Adaptability, resilience and environmental buffering in European Refugia during the Late Pleistocene: Insights from La Riera Cave (Asturias, Cantabria, Spain)

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The Upper Palaeolithic in Europe was a time of extensive climatic changes that impacted on the survival and distribution of human populations. During the Late Glacial Maximum (LGM), southern European peninsulas were refugia for flora, fauna, and human groups. One of these refugia, the Cantabrian region (northern Atlantic Spain), was intensively occupied throughout the Upper Palaeolithic. Characterising how climatic events were expressed in local environments is crucial to understand human and animal survival. La Riera Cave (Asturias) has a rich geo-cultural sequence dating between 20.5kyr BP to 6.5kyr BP and represents an ideal location in which to explore this. Stable isotope analysis of red deer and ibex is used alongside other environmental and climatic proxies to reconstruct Late Upper Palaeolithic conditions. Results show that during the LGM, ibex adapted their niche to survive, and became a major prey species for humans. The diverse environmental opportunities offered in the high-relief and coastal environs of La Riera may help to explain the high human population levels in the Cantabrian Region throughout the Late Upper Palaeolithic. Despite fluctuating conditions, herbivores and humans had the flexibility and resilience to adapt, demonstrating the importance of southern European refugia for the survival of different species.
are key to understanding why this region functioned as a refugium throughout the Late Pleistocene, enabling the survival of humanity throughout these climatically challenging conditions.

Ice core evidence from Greenland\(^2\) provides information on larger climatic events that occurred throughout MIS2/1, including the Last Glacial Maximum, Heinrich 1 event, Greenland Interstadial 1, Younger Dryas, and eventually the onset of the Holocene. The integration of ice and marine cores with terrestrial records has helped to refine the stratigraphy of these events\(^2\)\(^-\)\(^2\)\(^7\). These larger scale climatic records are physically far removed from conditions experienced on regional and local levels in northern Spain. Speleothem records\(^2\)\(^8\), lake sediments\(^2\)\(^9\),\(^3\(^0\), ostracod evidence from the continental shelf\(^3\(^1\)) reflect data closer to study areas, enabling regional-scale environmental reconstructions, but these are still distant from the sites occupied by human groups. More localised environmental indicators such as pollen, found within cave sites, can inform on past vegetation\(^3\(^2\),\(^3\(^3\),\(^3\(^4\)) and have been analysed in Asturian cave sites\(^3\(^3\),\(^3\(^4\)) but must be used with caution as they can be subject to taphonomic and diagenetic alterations\(^3\(^5\),\(^3\(^6\)). Both macro- and micro- mammal remains are ubiquitously found on these sites\(^3\(^7\)-\(^4\(^3\)) and are often less affected by taphonomic and sampling biases than other environmental proxies.

Bone collagen $\delta^{13}C$ and $\delta^{15}N$ analyses have been abundantly applied in European settings to explore palaeoclimates and palaeoenvironments\(^4\(^4\)-\(^4\(^7\)) and has yielded successful results for the Middle and Upper Palaeolithic of northern Spain\(^2\)\(^6\),\(^4\(^8\)-\(^5\(^2\)). Analysis of $\delta^{13}C$ and $\delta^{15}N$ in bone collagen is a powerful tool for reconstructing palaeoenvironments, especially when combined with other available environmental proxies\(^5\(^3\). Carbon and nitrogen in animal bones derives directly from consumed diet\(^5\(^4\), with a fractionation effect\(^5\(^5\), and shows long-term dietary behaviour of animals during the period of bone growth, typically representing the lifetime of each individual animal\(^5\(^6\),\(^5\(^7\).

A multitude of climatic and environmental factors can affect the $\delta^{13}C$ and $\delta^{15}N$ values of plants, and thus the animals that consume them. Regarding $\delta^{13}C$ values, temperature has been seen to affect values both positively and negatively\(^5\(^8\), although the $\delta^{13}C$ values of plants appear to have a greater correlation with mean annual precipitation than mean annual temperature, and typically higher levels of precipitation result in lower plant $\delta^{13}C$ values\(^5\(^9\),\(^6\(^0\)). Increased tree cover can result in lower $\delta^{13}C$ values in relation to more open environments linked to a combination of closed carbon cycling of CO$_2$ under the forest's canopy limiting atmospheric exchange and recycling of CO$_2$ from decomposing leaf matter\(^6\(^1\)-\(^6\(^4\)). Plants growing at higher altitude can typically produce elevated $\delta^{13}C$ values linked to decreased temperatures and changing CO$_2$ partial pressure\(^6\(^5\), and has been observed archaeologically to produce elevated $\delta^{13}C$ in the fauna that reside at higher altitudes\(^6\(^6\)). Global scale conditions such as a change in atmospheric CO$_2$ levels can also impact on $\delta^{13}C$ values of plants\(^6\(^7\),\(^6\(^8\), and thus be expressed in fauna. Larger-scale drivers and climatic events such as this would be expected to be observable in multiple species within a site, and typically should be observed across different geographical locations.

A negative correlation between plant $\delta^{15}N$ values and annual precipitation has been observed in a variety of ecosystems\(^6\(^9\)-\(^7\(^1\)). Higher temperatures and lower water availability are thought to produce higher animal $\delta^{15}N$ values in hot/arid potentially due to physiological stress due to heat and lack of water\(^2\)\(^2\)-\(^2\(^4\), although research has suggested that this changes could be linked to dietary shifts associated with these conditions\(^5\(^5\),\(^5\(^7\). Colder temperatures

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**Figure 1.** Location of La Riera Cave within the Cantabrian Region, Northern Spain (Panel A) with other contemporary archaeological sites, together with available environmental and climatic proxy records, referred in the text (Panel B). Panel C shows the outside view of La Riera Cave in the summer of 1976 and Panel D displays the stratigraphy sequence of squares G/H at the site (Photos by L.G. Straus). Panel E is a re-drawing of the stratigraphy at La Riera (based on the original version by G.A. Clarke\(^8\(^2\)).
are frequently associated with lower δ15N values due to reduced activity of nitrogen fixing mycorrhizae. There is a negative relationship between soil δ15N values and altitude. Different geographical locations can exhibit variation in baseline δ15N due to the nature of soil types within a given region, which affects plant communities, and nitrogen-fixing mycorrhizae present.

In addition to environmental and atmospheric drivers affecting plants (and their consumers), changing environmental conditions can also affect faunal behaviour, in terms of the dietary and habitat choices of animals. Animal diets may shift to include different plant species consumed, or parts of plants, with differing isotopic signatures resulting from plant physiological variations. Due to the complex processes influencing how localised and regional conditions are expressed within the isotopic values of plants, and the animals that consume them, temporal trends in δ13C and δ15N values can be expressed differently between regions, and changing conditions do not always affect every species in the same way. Wider environmental indicators can be used to facilitate interpretations of isotopic results to understand the trends that are visible.

When applied to bones with evidence of anthropogenic modifications (e.g., exhibiting cut marks, marrow cracking, hammer blows), bone collagen δ13C and δ15N analysis can provide direct evidence of past environments during the times when these animals were consumed and is one of the few proxies that can inform on how climatic events were experienced by humans and animals on a local scale. This study combines newly generated δ13C and δ15N evidence from La Riera, one of the most important, and well-published Upper Palaeolithic sites in the Cantabrian region with wider climatic and environmental proxies from local, regional and global records to explore environmental changes throughout the archaeological sequence to understand how humans and animals adapted and became resilient to the changing conditions that they faced.

**La Riera Cave**

La Riera (Posada de Llanes, Asturias) (Fig. 1) was discovered and first excavated by Conde de la Vega del Sella in 1916 and was re-excavated by L.G. Straus and G.A. Clark in 1976–79. All materials discussed in this paper relate to the findings from the 70s excavations. The site contains a long sequence of archaeologically rich levels dating from the Solutrean to the Asturian (Mesolithic) cultural periods, C14-dated between 20–6.5kyr BP. The site is located c. 30 m above sea level on the south-facing edge of a low ridge on the narrow coastal plain facing the steep, nearby cliffs of the 800 m a.s.l. Sierra de Cuera, a coastal mountain range north of the massive Picos de Europa chain. Currently located 2 km from the shore, during the LGM, when sea level was approximately 120 m lower, the site was around 10 km from the shore, roughly two hours walk from the Late Glacial littoral.

| Level | Cultural attribution | Lab number | Material | Dating method | 14C date ± | Red deer n | Ibex n | Total n |
|-------|----------------------|------------|----------|---------------|------------|-----------|--------|--------|
| 29    | Asturian             | GaK-3046   | Charcoal | C14           | 6500 200   | 7         | 4      | 11     |
|       |                      | RIE-199    | Shell    | AAR           | 7516 588   | 7         |        |        |
|       |                      | GaK-2909   | Charcoal | C14           | 8650 300   |           |        |        |
| 28    | Azilian              | Q-2933     | Shell    | C14           | 9125 90    | 7         | 3      | 10     |
| 27    |                      | BM-1494    | Bone     | C14           | 10630 120  |           |        |        |
|       |                      | Q-2935     | Shell    | C14           | 11285 70   | 7         | 7      | 14     |
| 24    | Upper Magdalenian    | GaK-6982   | Charcoal | C14           | 10890 430  | 7         | 3      | 10     |
|       |                      | Q-2926     | Shell    | C14           | 11775 75   |           |        |        |
|       |                      | UGAMS-2752 | Bone     | AMS           | 13530 35   |           |        |        |
| 23    |                      | Q-2932     | Shell    | C14           | 13095 110  | 7         | 4      | 11     |
|       |                      | UCR-1275D  | Bone     | C14           | 12620 300  |           |        |        |
| 19    | Lower Magdalenian    | Q-2116     | Charcoal | C14           | 15230 300  | 7         | 7      | 14     |
|       |                      | Q-2110     | Charcoal | C14           | 15520 350  | 7         | 7      | 14     |
|       |                      | GaK-6448   | Charcoal | C14           | 16420 430  |           |        |        |
| 18    | Solutrean            | Q2936      | Shell    | C14           | 15885 150  | 7         | 7      | 14     |
| 15    |                      | UCR-1272A  | Bone     | C14           | 17225 350  |           |        |        |
| 14    |                      | Q2927      | Shell    | C14           | 16705 105  |           |        |        |
| 13    |                      |           |          |               |            |           |        |        |
| 12    |                      | GaK-6446   | Charcoal | C14           | 17210 350  | 8         | 7      | 15     |
| 11    |                      |           |          |               |            |           |        |        |
| 10    |                      | GaK-6447   | Charcoal | C14           | 19820 390  | 7         | 7      | 14     |
| 7     |                      | Q2934      | Shell    | C14           | 20295 210  |           |        |        |

Table 1. Radiocarbon (and AAR) dates of the archaeological levels sampled in this study, including the number of specimens per species analysed for bone collagen δ13C and δ15N values. Where direct dates for a level were not available, dating evidence from adjacent levels has been included. Shell C14 dates adjusted upward by 105 years for the Holocene marine reservoir effect, per Soares et al. 143.
location of the site means that hunting areas available included the montane habitats of the Sierra, the low, rolling coastal plain, the valleys of the small Calabres and Bedón rivers, and their estuaries offering a selection of wild terrestrial, marine, and riverine resources32.

To date, La Riera remains one of the most extensively studied sites in the Cantabrian Region, and was a landmark study, as it provided detailed scientific analysis, including extensive analyses of palaeoenvironmental remains including: pollen records, deposited sediments, bird bones, macro- and micro-fauna, and even a pioneering analysis of mollusc shell isotopes34,37,83–86. Several series of radiocarbon dates82,87,88 were undertaken both by the excavators and subsequent researchers, and two new dates for Levels 23 and 2490 run for this study have resolved some previous dating discrepancies in chronologies at the site (Table 1).

Although not permanently occupied, La Riera was repeatedly visited and inhabited, during different seasons, by human groups throughout the Upper Palaeolithic and Mesolithic and was an important hunting and gathering base32. La Riera was as part of a network of Upper Palaeolithic sites across the Cantabrian region. The Solutrean (Levels 4 to 17) dates to between 20.5–17kyr BP, and the earliest Solutrean occupations are interpreted as representing mainly a hunting camp for the procurement of red deer and mainly ibex90. The site then became a multi-purpose residential camp, still exploiting large numbers of red deer and ibex with fish and shellfish also being collected90. The Lower Magdalenian (Levels 18–20) date to between 17–14kyr BP and the Middle and Upper Magdalenian (Levels 21–26) span between 14.5–13kyr BP. The Azilian, a form of local Mesolithic industry, (Levels 27–28) falls between around 10.5–9kyr BP, and at the end of the sequence, Level 29 (attributed to the Asturian, also a local Mesolithic industry) is a shell midden (‘conchero’) formed between roughly 9-6.5kyr BP. Towards the end of the sequence, faunal evidence suggests stricter bi-seasonality, with occupation focussed on the summer and winter90. The faunal record for La Riera shows evidence of intensive carcass processing, including marrow extraction from the low-yield phalanges and mandibles of red deer, and large numbers of individuals per level were recorded37 despite the relative small portion of the site excavated, reflecting intensive hunting strategies. Most notably from the Upper Magdalenian levels onwards (Level 24–29) there is increasing evidence of slaughtering Cervus hinds and their young, alongside a more diverse faunal suite, including greater frequencies of birds, fish, and marine molluscs37, suggesting increasing levels of dietary diversification and further intensification of diet throughout the Upper Palaeolithic.

La Riera is 26 km west of coastal El Pindal cave, famous for its Magdalenian cave art and yielding a speleothem record now used as a palaeoenvironmental proxy28,91,92 (Fig. 1). Lake Enol, situated at c. 1000 m above present sea level about 20 km WSW in the Picos de Europa of Asturias (Fig. 1) contains valuable lake sediment information regarding temperatures throughout the Late Pleistocene27,30. The relative proximity of La Riera to sites with additional environmental proxies, facilitates a more nuanced interpretation of the past conditions experienced by the human groups that occupied La Riera.

Results and Discussion

Overview of the δ13C and δ15N data. In total 128 specimens (72 red deer and 56 ibex) were analysed for δ13C and δ15N in bone collagen from 10 different archaeological levels at La Riera, from the Solutrean to the Asturian period, covering a temporal span of almost 15kyr BP (Table 1). Summary statistics of these results are shown in Table 2. Preservation at the site was exceptional and collagen was successfully extracted and analysed from 121 specimens (94.5%). All specimens analysed had atomic C:N ratios between 3.2–3.5, falling within the accepted ranges for well-preserved collagen90, with 110 specimens of these having ratios between 3.2–3.4, adhering to more rigid criteria of Van Klinken94. The vast majority of specimens had >35%C and %N values between 10–16%, and 109 of the specimens collagen yields were >1% (based only on the >30Ka fraction, rather than total collagen), indicative of well-preserved in vivo collagen90, as detailed in the Supplementary Information (SI1).

A series of trends were observable throughout the sequence due to changing δ13C and δ15N values of the species analysed. To facilitate discussion the sequence has been divided into interpretative zones (Z1-25) based on key characteristics of the data including; changing statistical relationships of the δ13C or δ15N values of each species, increases/or decreases in the isotopic values of both species, observably by statistically significant differences between archaeological levels, and in one instance notably large ranges in isotopic values of both species observed. A thorough explanation of the key characteristics of each interpretative zone is as follows: Zone (Z1), at the bottom of the stratigraphic sequence, dating to the Solutrean, is characterised by partitioning of the red deer and the ibex, with the latter having 15-N depleted values. Zone 2 (Z2, Middle into Upper Magdalenian) is marked by an increase in the δ13C value of the ibex, followed by stability, in both δ13C and δ15N values. Zone 3 (Z3, Upper Magdalenian) is differentiated by a small increase in the δ13C values of both species observed between Levels 23 and 24. Zone 4 (Z4, Azilian) is marked by a dramatic decrease in the δ13C values in both species. Finally, Zone 5 (Z5, Asturian) exhibits similar δ13C and δ15N values to the preceding levels (27 and 28), but is marked as being different due to a larger range in the δ13C values of both the red deer and ibex in comparison to the rest of the sequence. The isotope data are correlated with environmental proxies on a local, regional and global scale to determine how trends observed relate to wider environmental and climatic trends. Raw data are provided in SI1–10.

The results and interpretation are framed by exploring the patterns in the stable isotope record at La Riera, outlining the conditions experienced continentally within the region and the implications that this had for human groups, before being drawn together in the final discussion.

Zone 1: The Solutrean (Levels 7-13: 20- 17kyr BP). The Solutrean levels (7, 11 and 13) exhibit a marked difference in the δ15N values of ibex and red deer, with red deer having consistently higher δ15N values by at least 1‰, and no overlap in standard deviation is seen between the two species (Fig. 2). A statistically significant difference in in the δ15N was seen between the two populations for each of these levels (Table 3). From Level
18 onwards, no difference in the δ15N values of the two species was observed (Fig. 2, Table 3). The ibex have the lowest range in δ15N values seen within the sequence at La Riera during this time (Table 2).

The δ13C values show little difference between the two species. Only in Level 11 were the mean ibex values 0.7‰ lower than those seen at La Riera (Table 1, Fig. 2), and the populations were statistically different (p = 0.01) (Table 3). The mean δ13C value of ibex was also slightly higher within Level 13 (Table 2), although there was no statistically significant difference (p = 0.07) (Table 3). There is a decrease in δ13C values of the deer between levels 7 and 11, which was statistically significant (p = 0.02) (Table 4).

The difference in δ15N values indicates niche partitioning of ibex and red deer. Pollen records available from Levels 6–15 indicate a predominance of *Ericacea* (heather/heath plants), accounting for between 60–90% of the pollen recorded, and little tree cover, with arboreal pollen accounting for <5% (Table 5). *Microtus oeconomus*, a species adapted to cold environments, is identified in Level 7 (Fig. 3). Montane species (ibex) are exploited in greater quantities (74% of NISP in Level 7, declining to 30% and 15% throughout this zone), although red deer are still present in great numbers (25% in Level 7, increasing to 80–90% in Levels 11 and 13 this zone) (Figs. 3 and 4). The ostracod *Microcyprides armatus* (Capra pyrenaicus) was also identified in Level 7 (Fig. 3) which could indicate colder/drier conditions.

Within the Solutrean Z1 lower δ15N values in the red deer were noted at El Mirón in comparison to values within Z2 from the site (p = 0.00) (Table 5; Fig. 3), which could suggest colder, drier conditions during this time. Typically El Mirón δ15N values are lower than those seen at La Riera (mean values range between 2.4–2.8‰ in Z1–4), indicative of wider, baseline differences between these two geographically distinct locations (Table 5; Fig. 3). Antoliñako Koba δ15N values were higher in Z1, compared to those seen later on in the sequence (Table 5; Fig. 3) which could indicate colder/drier conditions.

The water temperature estimates for Lake Enol are calculated at 1 °C between 20-17kyr BP (Fig. 3), with no change registered in the Ca precipitation record, although during cooler times, this record becomes less precise (Fig. 3). The ostracod evidence from El Pindal cave shows some of the lowest δ13C values and a hiatus in growth between 18.2-15.4kyr BP (Fig. 3). This interruption is interpreted as belonging to one of the coldest and potentially relatively driest parts of the sequence (Fig. 3). Lower δ18O values in the NGRIP record at 20kyr BP indicate that this part of the sequence corresponds to the end of the LGM, which concluded at c. 19kyr BP, when the onset of Heinrich 1 event marks a climatic improvement (Fig. 3). The ostracod evidence, from the Basque continental shelf, also attests to generally cooler sea temperatures (Fig. 3).

**Interpretation.** The Solutrean was a cold period, representing the latter stages of the LGM and start of the Heinrich 1 event, as seen continentally and in ice core and offshore marine records. The consistently lower δ15N values of ibex during this time, in comparison to the red deer (and in relation to the ibex analysed from later in the sequence) is distinctive at La Riera. Higher altitudes are associated with depleted δ15N values of soils and plants, and it has been proposed that animals feeding at altitudes may have produced lower δ15N values of faunal, as seen in Northern Europe, and at the Cataluña site of Balma Guilanya during the 20–10kyr BP. Typically higher altitudes are also associated with elevated δ13C values in plants, which has been archaeologically observable within ibex at the site of Valdegoba in central Northern Spain, and is not evident at La Riera. There is no

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**Table 2. Summary statistics of red deer and ibex δ13C and δ15N values from levels sampled at La Riera.**

| Level | δ13C n | Mean | Min. | Max. | Range | 1σ | δ15N n | Mean | Min. | Max. | Range | 1σ |
|-------|--------|------|------|------|-------|----|--------|------|------|------|-------|----|
| 7     | 7      | −20.2| −20.8| −19.6| 1.2   | 3.4| 5.8    | 2.4  | 1.0  |
| 11    | 7      | −20.9| −21.2| −20.4| 0.8   | 3.7| 4.3    | 3.2  | 0.4  |
| 13    | 7      | −20.6| −21.2| −20.2| 1.0   | 3.5| 4.3    | 3.4  | 0.5  |
| 18    | 7      | −20.7| −21.1| −20.5| 0.6   | 4.0| 4.4    | 0.8  | 0.3  |
| 19    | 7      | −20.8| −21.3| −20.7| 0.6   | 4.2| 3.7    | 4.8  | 1.1  | 0.4  |
| 23    | 7      | −20.8| −21.0| −20.0| 1.0   | 3.9| 2.8    | 5.3  | 2.5  | 0.7  |
| 24    | 7      | −20.2| −20.8| −19.6| 1.2   | 4.3| 3.4    | 5.8  | 2.4  | 1.0  |
| 27    | 6      | −21.2| −21.5| −20.9| 0.6   | 4.5| 3.8    | 5.1  | 1.3  | 0.5  |
| 28    | 5      | −21.0| −21.3| −20.7| 0.6   | 3.8| 3.5    | 4.3  | 0.8  | 0.3  |
| 29    | 7      | −21.1| −21.6| −20.1| 1.5   | 4.1| 2.4    | 5.1  | 2.7  | 1.0  |

**Spanish ibex (Capra pyrenaicus)**

| Level | δ13C n | Mean | Min. | Max. | Range | 1σ | δ15N n | Mean | Min. | Max. | Range | 1σ |
|-------|--------|------|------|------|-------|----|--------|------|------|------|-------|----|
| 7     | 5      | −20.5| −21.3| −19.9| 1.4   | 3.0| 3.5    | 1.2  | 0.5  |
| 11    | 5      | −20.1| −20.5| −19.5| 1.0   | 2.5| 2.0    | 3.0  | 1.0  | 0.4  |
| 13    | 6      | −20.2| −20.4| −20.0| 0.4   | 2.7| 2.1    | 3.8  | 0.6  |
| 18    | 7      | −20.6| −21.4| −20.0| 1.4   | 4.0| 2.6    | 5.3  | 2.7  | 0.8  |
| 19    | 6      | −20.6| −21.7| −19.9| 1.8   | 4.0| 3.6    | 4.4  | 0.8  | 0.3  |
| 23    | 4      | −20.6| −21.4| −20.3| 1.1  | 3.6| 2.9    | 4.5  | 1.6  | 0.7  |
| 24    | 3      | −19.9| −19.9| −19.8| 0.1   | 4.6| 4.0    | 5.2  | 1.2  | 0.6  |
| 27    | 7      | −21.4| −21.9| −20.2| 1.7   | 4.4| 3.1    | 4.9  | 1.8  | 0.7  |
| 28    | 3      | −20.7| −21.1| −20.4| 0.7   | 3.5| 3.2    | 3.9  | 0.7  | 0.4  |
| 29    | 4      | −20.9| −21.7| −20.3| 1.4   | 3.8| 2.7    | 5.2  | 2.5  | 1.0  |

**Interpretation.** The Solutrean was a cold period, representing the latter stages of the LGM and start of the Heinrich 1 event, as seen continentally and in ice core and offshore marine records. The consistently lower δ15N values of ibex during this time, in comparison to the red deer (and in relation to the ibex analysed from later in the sequence) is distinctive at La Riera. Higher altitudes are associated with depleted δ15N values of soils and plants, and it has been proposed that animals feeding at altitudes may have produced lower δ15N values of faunal, as seen in Northern Europe, and at the Cataluña site of Balma Guilanya during the 20–10kyr BP. Typically higher altitudes are also associated with elevated δ13C values in plants, which has been archaeologically observable within ibex at the site of Valdegoba in central Northern Spain, and is not evident at La Riera. There is no
Figure 2. Mean δ¹³C values (left) and δ¹⁵N values (right) of red deer (denoted by black circles) and ibex (denoted by grey circles) from the Upper Palaeolithic levels sampled throughout the sequence at La Riera. Error bars show 1σ.

Table 3. Comparisons of red deer and Spanish ibex δ¹³C values and δ¹⁵N values, within each archaeological level sampled, to determine similarities and differences between the dietary niches of both species. Numbers in bold denote statistically significant results.

| Species | Levels Compared | 7&11 | 11&13 | 13&18 | 18&19 | 19&23 | 23&24 | 24&27 | 27&28 | 28&29 |
|---------|----------------|------|-------|-------|-------|-------|-------|-------|-------|-------|
| Red deer | δ¹³C | 0.02 | 0.25 | 0.61 | 0.34 | 0.70 | 0.03 | 0.00 | 0.23 | 0.62 |
|         | δ¹⁵N | 0.48 | 0.37 | 0.07 | 0.37 | 0.20 | 0.61 | 0.43 | 0.03 | 0.33 |
| Ibex    | δ¹³C | 0.40 | 0.58 | 0.37 | 0.93 | 0.51 | 0.04 | 0.03 | 0.15 | 0.86 |
|         | δ¹⁵N | 0.17 | 0.71 | 0.02 | 0.58 | 0.24 | 0.11 | 0.69 | 0.15 | 0.86 |

Table 4. Independent comparisons of red deer (upper rows) and ibex (lower rows) δ¹³C values and δ¹⁵N values between consecutive archaeological levels sampled to explore temporal changes. Numbers in bold denote statistically significant results.
difference between the ibex and the red deer δ¹³C values throughout the sequence, suggesting that both species are feeding at isotopically similar altitudes, making this less likely to be the cause for the lower values of the ibex. Furthermore, studies of Mediterranean vegetation indicate that both species density and diversity declines with altitude⁹⁶, suggesting that scaling higher altitudes would be less profitable for ibex, especially during the hostile conditions of the Solutrean. Thus, alternative drivers are responsible for the difference in δ¹⁵N values of the two species during the Solutrean.

The lower δ¹⁵N may be linked to changing patterns of vegetation consumption, resulting in the niche of the ibex being altered to adapt to the cold conditions of the Solutrean. The lower ibex δ¹⁵N values occur alongside increased quantities of Ericaceaeous (heath) plants³⁴ (Fig. 3). Ericaceaeous plants typically occupy regions with lower quality, acidic soils and have mycorrhiza which enhance their ability to absorb nitrogen from soils⁹⁷. The larger proportions of ericaceous plants suggest more hostile environmental conditions were experienced during these levels. Studies of recent red deer and ovicaprids in Scotland have shown that red deer can digest heather better than ovicaprids⁹⁸. Ecological research into the feeding behaviour of ovicaprids and red deer in mosaic heath landscapes showed that the dietary selection of ovicaprids was heavily affected by the presence of heather, resulting in the two species consuming different diets⁹⁹, and could explain the observed niche partitioning in these archaeological specimens during the Solutrean. Also, ecological studies of ibex have demonstrated that they are pushed into alternative niches when competing with livestock¹⁰⁰, and a similar effect might be observed when faced with red deer as a competitor. Red deer are an ecologically flexible species, thriving in open grasslands, in addition to forested environments¹⁰¹–¹⁰³. Their combination of grazing and browsing behaviour means that they have great flexibility in their dietary behaviour, already observed archaeologically¹⁰⁴, and thus may have been able to continue existing the ericaceous landscape that proved challenging to the ibex. Ibex are known to have a high plasticity in the feeding strategies depending on ecological constraints¹⁰⁵, and it is possible that the ibex populations were being pushed away from the ericaceous vegetation communities of the δ¹⁵N depleted heathlands and were consuming a different diet. The mobility of ibex reduces when food sources are scarcer and encounter

Figure 3. Multi-scale environmental and climatic comparisons throughout the Upper Palaeolithic on a local, regional and global level. Stable isotope δ¹³C and δ¹⁵N values of red deer and ibex sampled at La Riera are shown alongside additional environmental proxies from the site²⁶,²⁷, and isotopic values from red deer across the Cantabrian Region²⁶,⁴⁸,⁴⁹,⁵². Wider continental environmental indicators²⁸,⁹², offshore proxies³¹ and NGRIP ice core evidence³⁵, and climatic events¹⁴¹ are also displayed to further contextualise trends observed through time.

Figure 4. The %NISP of each of key prey species exploited within the sampled levels in this study showing the variation of each species through time, based on data produced by Altuna³⁷.
deviations (Table 2), and the diet of both species becomes isotopically indistinguishable. An increase in mean $\delta^{13}$C values of ibex in Lower Magdalenian Level 18 are 1.3‰, higher than in Level 13 (Fig. 2, Table 2), and the values of red deer and ibex track each other relatively closely in mean values, ranges and standard deviations (Table 2), and the diet of both species becomes isotopically indistinguishable. An increase in mean $\delta^{13}$C values of red deer is seen between Levels 13 and 18, although this was not statistically significant (p = 0.07) (Table 4). Throughout the Lower Magdalenian and into the Upper Magdalenian (Levels 18, 19, 23) (16-15kyr BP), the $\delta^{15}$N and $\delta^{13}$C values of both the red deer and the ibex remain stable (Table 2, Fig. 2). Red deer and ibex populations within Levels 18–19 and Levels 19–23 were statistically similar (Table 4).

With the conclusion of the Solutrean, from Level 16 onwards, there was a large decrease in Ericaceae and heath was replaced by steppic vegetation (including Liguliflorae, Tubuliflorae, Chenopodiaceae, and Artemisia) (Fig. 3). Arboreal pollen increased slightly from before, but remains low, accounting for less than 10% of the total pollen record between Levels 16–23. Lowland species (notably red deer) are the most commonly exploited game (70%), but with montane (steep, rocky habitat-preferring) species (predominantly ibex) also representing a valuable resource within these levels (Fig. 4).

The wider isotopic data from the Cantabrian Region red deer also show little temporal change in $\delta^{13}$C values in the Lower Magdalenian to the Upper Magdalenian (17-12kyr BP) (Table 5; Fig. 3). Within Z2, the El Mirón $\delta^{13}$C values indistinguishable from those seen in Z1 (p = 0.798) (Table 5). The mean $\delta^{15}$N value at El Mirón is 2.8‰ which is 0.4‰ higher than observed in Z1 (Table 5). Lake Enol temperatures of 2 °C are estimated during the 1b event, followed by Greenland Interstadial 1, which represents a warmer climate (Fig. 3). A decrease in the “Northern Guest” benthic foraminifera between 16-15kyr BP, indicate warmer sea bottom temperatures off the Basque shelf, and warmer seawater surface temperatures are suggested by the lower $\delta^{18}$O values and decline in N. pachyderma sin. % % (Table 5).

Table 5. Summary statistics to compare $\delta^{13}$C and $\delta^{15}$N values of red deer for each isotopic zone at La Riera with contemporary results from the wider the wider Cantabrian Region sites of El Mirón (Cantabria)52, Kiputz IX (Gipuzkoa)48, Antoliñako Koba (Bizkaia)26 and La Paloma (Asturias)49, where data are available (see Supplementary SI6).

| Archaeological site | $\delta^{13}$C values | $\delta^{15}$N values |
|---------------------|----------------------|----------------------|
|                     | Isotope Zone | N | Mean | Min | Max | Range | 1σ | Mean | Min | Max | Range | 1σ |
| La Riera            | 1           | 20 | −20.7 | −21.2 | −20.2 | 1.0 | 0.3 | 3.7 | 3.0 | 3.0 | 1.3 | 0.4 |
| El Mirón            | 11          | 21 | −20.8 | −21.3 | −20.0 | 1.2 | 0.3 | 3.9 | 2.5 | 5.7 | 3.2 | 1.2 |
| La Paloma           | 6           | 26 | −20.4 | −21.0 | −19.6 | 1.6 | 0.6 | 3.3 | 2.8 | 5.3 | 2.5 | 0.5 |
| Antoliñako Koba     | 3           | 5   | −20.9 | −21.1 | −20.3 | 1.2 | 0.5 | 3.4 | 2.4 | 3.8 | 0.8 | 0.7 |
| Kiputz IX           | 1           | 1   | −21.3 | −21.3 | −21.3 | 1.2 | 0.3 | 3.9 | 2.5 | 3.2 | 0.8 | 0.5 |
| Kiputz IX           | 2           | 1   | −20.9 | −20.8 | −19.7 | 1.2 | 0.5 | 3.4 | 2.4 | 3.8 | 0.8 | 0.7 |
| Kiputz IX           | 3           | 7   | −20.2 | −20.8 | −19.7 | 1.2 | 0.5 | 3.4 | 2.4 | 3.8 | 0.8 | 0.7 |
| Kiputz IX           | 4           | 3   | −20.6 | −21.0 | −20.4 | 0.6 | 0.3 | 3.7 | 2.1 | 3.2 | 1.1 | 0.6 |
| Kiputz IX           | 5           | 5   | −20.9 | −21.0 | −20.7 | 0.3 | 0.1 | 3.4 | 2.8 | 3.5 | 0.9 | 0.7 |
| Kiputz IX           | 6           | 1   | −21.1 | −21.6 | −20.1 | 1.5 | 0.5 | 4.1 | 2.4 | 5.1 | 2.7 | 1.0 |
| Kiputz IX           | 7           | 3   | −21.4 | −22.0 | −20.5 | 1.5 | 0.8 | 4.0 | 3.5 | 4.5 | 1.0 | 0.5 |

Interpretation. Climatic amelioration is attested within the La Riera sequence by a shift in vegetation, that occurs with the change to G1 signalled by a decline in heath plants and an increase in steppic plants, then suggesting a warmer and slightly drier environment. The $\delta^{15}$N values of the ibex increase in this zone, and they become isotopically indistinguishable from the red deer. Warmer conditions which may explain the increased $\delta^{15}$N values of the ibex, as a product of increased soil microbial activity 77,78, although this would be expected to affect the red deer values also. The lack of covariance in the $\delta^{15}$N values of the La Riera red deer suggests that drivers behind this change are specific to the niche of the ibex. The shift in the ibex $\delta^{15}$N values may be linked to dietary change of the ibex related to the reduction in Ericaceae, allowing the ibex could expand into regions that were previously less desirable, adapting their niches to suit these new environmental conditions. The consistency in the stable isotope values of both species throughout Levels 18–23 indicates that the environment was relatively...
stable within these levels, as is mirrored in continental stable isotope results across the region. The increase in red deer bones within the zooarchaeological assemblage suggests that there is a shift in hunting focus, associated within these milder conditions.

Zone 3: Upper Magdalenian (Level 24: 13.5kyr BP). During the Upper Magdalenian occupation of the site, between Levels 23 and 24, an increase in $\delta^{13}C$ is observed, and mean values in both red deer and ibex are higher within Level 24 (Table 2, Fig. 2), which was statistically significant for both species (Table 4). A slight increase in $\delta^{15}N$ values is visible for both species (Table 2, Fig. 2), although this was not statistically significant for red deer ($p = 0.61$) or ibex ($p = 0.11$) (Table 4). A stratigraphic hiatus was recorded between these two levels and new dating of ungulate bones from these levels (Table 1), revealed a 2,000-year gap between them, with Level 24 marking a restart in the archaeological sequence.

The lack of thermophilic species in Level 24 in the palynological record suggests a cooler environment. There is a small increase in arboreal plants, steppic taxa continue to be represented, and generally tree cover remains low (Fig. 3) with ferns present only in the upper parts of level 24. Previous interpretations of the level suggested that it was formed during a cold spell within the Allerød (Fig. 4) and sedimentary evidence indicated a colder environment than in the previous level. There is evidence of a greater exploitation of montane species (mostly represented by ibex, accounting for 40% of the NISP) (Fig. 4).

On a wider regional level, continental records at Lake Enol show a slight temperature increase of 2–4 °C at around 14-13kyr BP (Fig. 3). There are few comparative isotopic results from this period, with only 3 specimens from El Mirón (in the montane interior of Cantabria, c. 160 km to the east) dating to c.13kyr BP, slightly later than the specimens from La Riera. The $\delta^{13}C$ values are very similar to those observed previously in the sequence, although the mean $\delta^{15}N$ value is slightly lower at this time (Table 5, Fig. 3). At 14kyr BP, the El Pindal Cave $\delta^{13}C$ and $\delta^{18}O$ show lower values, but increase again at 13kyr BP (Fig. 5), indicative of changing conditions. The NGRIP record shows the end of GI1, where a slight decrease in $\delta^{18}O$ values, indicative of a cooler environment (Fig. 3). There are no "Northern Guest" benthic foraminifera between 15-12.5kyr BP (Fig. 3), suggesting warmer sea bottom temperatures, and this phase is part of a longer trend of declining N. pachyderma sin percentages.

Interpretation. The shift in the $\delta^{13}C$ values of the red deer between Levels 23 and 24 appears to be a product of colder, and possibly drier environmental conditions. Positive and negative correlations are seen between temperature and $\delta^{13}C$ values, and elevated $\delta^{13}C$ can be indicative of drier conditions (also suggested by the small increase in $\delta^{15}N$ values, which typically can become elevated when exposed to more arid conditions). Given that this shift coincides with a cooler period, as marked by the NGRIP $\delta^{18}O$ at the end of GI1, and wider environmental evidence from the site suggests colder environments, a decreasing in temperature, and possible drier environments are potentially responsible for the increase in $\delta^{13}C$ values. The increased exploitation of ibex in this level (42% of NISP in Level 24, compared to 13% of NISP in Level 23, Fig. 4), a characteristic also seen in the Solutrean levels, indicate a tendency to hunt greater quantities of montane animals during these more rigorous conditions. This zone reflects a short duration of time, with few contemporary isotopic specimens from other sites available to date (Table 5, Fig. 3). It is possible that La Riera was occupied in favour of other sites due to the prevalence of ibex that had found a niche in order to succeed in the area, even during periods of environmental stress.

Zone 4: Azilian (Levels 27 and 28: 12-9kyr BP). With the transition to the Azilian (Level 27) a dramatic decrease in the $\delta^{13}C$ values is evident for both main game species, with mean values becoming more negative at this point, by 1‰ for red deer and 1.5‰ for ibex, and stay at these lower values throughout the Azilian and Asturian (Levels 27–29) (Table 2, Fig. 2). The $\delta^{15}N$ values of both species remain at a consistent level in comparison to the previous levels sampled (Table 2, Fig. 2).

Between Levels 27 and 28 (c. 12-9kyr BP) a shift in $\delta^{15}N$ was observed in red deer. The mean $\delta^{15}N$ value in Level 28 is 0.8‰ lower than Level 27 (Table 2, Fig. 2), and the populations show a statistically significant difference ($p = 0.02$, Table 3). Values remain at this lower level within Level 29. A similar effect is also observed within the ibex, with the mean $\delta^{15}N$ value being 0.9‰ lower, although this is not statistically significant ($p = 0.15$), potentially due to the smaller sample size available for Level 28.

The palynological record shows a small increase in arboreal woodland species such as oak, elm and hazel, in addition to ferns (Polypodium vulgare), suggesting a warmer, more humid environment during Levels 27 and 28 (Fig. 3). Dormouse (Glis glis), which inhabits more temperate and moist environments, was present (Fig. 3). Macroinvertebrate assemblage continue to show dominance in the expalntial and woodland environments, notably both red deer and the appearance of wilde, for the first time, in the sequence within Levels 27 and 28. An increase in animals exploited from open landscapes is evident (notably bovines accounting for 10% of the NISP in Level 28) (Figs. 3 and 4).

Other isotopic records from the Cantabrian Region are more limited for this temporal period. The mean value of specimens from Antolínako Koba (c. 190 km east on a mountaintop near the coast) is 0.7‰ lower in Z4 than the mean values seen in Z1, the only other zone for which data are available from this site (Table 5). At El Mirón within Z4, the mean $\delta^{13}C$ values and $\delta^{15}N$ values are indistinguishable from those seen within Z3 and no statistically significant differences were observed between populations within these zones, and it is only within Z5 when decreased $\delta^{13}C$ values and increased $\delta^{15}N$ values are evident at El Mirón (Table 5), indicative of changing Holocene conditions. Between 12-9kyr BP, Lake Enol temperature estimates average temperature of 10 °C, which is warmer than observed earlier in the sequence. At El Pindal Cave from 12kyr BP onwards, there is a continual increase in the $\delta^{18}O$ record.

The NGRIP data for this period is marked by the brief lowering of $\delta^{18}O$ values marking the Younger Dryas, followed by the large increase in $\delta^{18}O$ values that indicate the onset of the Holocene (Fig. 3). No "Northern guests" are
seen in the Basque continental shelf cores, indicating warmer sea bottom temperatures. The slow, but steady decline in the N. pachyderma sin taxon and the steady increase in the δ18Olob record is also indicative of warmer sea temperatures.

**Interpretation.** This zone corresponds with the start of the Holocene and is marked by more negative δ13C values of both species at La Riera. This pattern appears to extend further and is seen in the herbivore bone collagen record within wider Europe datasets, where after 14kyr BP there is a trend towards decreased δ13C values. This change may in part be related to increased concentration in atmospheric CO2 alongside a possible increased water availability, as further evidenced by the increased presence of arboreal pollen, and forest dwelling species within the wider environmental records at the site. The lower δ15N values within Level 28 support this interpretation, as there is a correlation between lower δ15N values and wetter environments, which are also reflected in El Mirón later, in Z5. This indicates a possible delayed response in nitrogen to the changing climate.

The increased suite of vegetation present in the pollen record, indicates greater habitat diversity, which in turn is reflected in the change in the suite of animals exploited in Level 28. In Level 28, a shift towards exploiting larger areas of the landscape was observed, with the inclusion of plain species (bison) which are infrequently present in assemblages before this point (Fig. 4). Diet in general becomes broader, with larger quantities of marine shells appear in the malacological record. The Holocene in Asturias, and in the whole Cantabrian region, is marked by a change in the suite of macrofauna fauna present, with reinder, bison and cave bear disappearing and increasing numbers of forest-dwelling species such as wild boar and roe deer, which are more suited to these temperate environments.

**Zone 5: Asturian (Level 29: 8.6-6.5 kyr BP).** Within Level 29 (a shell midden deposit), the δ15N ranges for both species are larger than seen in any other period (red deer 2.5‰, ibex 2.7‰, Table 2). This indicates a greater inter-individual variability within both red deer and ibex populations. The δ13C values remain at a similar level to those observed within Z4.

The pollen record for Level 29 is similar to that of Level 28, with the presence of abundant tree pollen, dominated by hazel, with oak, elm, alder and birch also represented. In addition, higher, instances of ferns and slightly higher proportions of arboreal pollen are observed, indicative of a moist and temperate environment (Fig. 3). Sediment composition also suggests that the level was formed under temperate and humid conditions.

Red deer of El Mirón (9kyr BP) have higher δ13C with mean value of 4‰, which is 1.2‰ than seen within Z3, and δ15N are 1‰ lower (Table 5). Lake Enol has records of water temperature of 12–13°C at 8.5 to 8kyr BP, before dropping slightly to 10–11°C (Moreno et al. 2014). In the El Pindal Cave speleothem record, there is a decrease in δ13C and δ18O during the last part of the record at c. 8kyr BP indicative of a cooler environment than seen previously in the sequence and may be reflecting the 8.2kyr BP event.

The higher δ18O values in the NGRIP record are part of the continued trend for warmer Holocene climates (Fig. 3). The marine core influx of benthic “Northern Guest” species c. 9.3kyr BP and 7.8kyr BP indicate cooler sea bottom waters and the sharp decrease in δ18Olob is thought to relate to the early Holocene cooling events.

**Interpretation.** This level corresponds to the shell midden (‘conchero’) that capped the infilling of the site. The larger ranges in the δ15N values might indicate that animals were consuming plants from a broader environmental area, encompassing a range of isotopically different locations. These larger ranges could be related to the accumulation of midden deposits over a long timescale, resulting in a wider variety in values observed. The longer spans of time represented in the midden layers, in comparison to the lower levels, could have resulted in a higher variation in values, representing a series of environmental fluctuations within levels. An alternative explanation for the diversity in stable isotope values is that the animals hunted from this level came from isotopically diverse areas of the landscape and could be indicative of procurement of resources from larger catchment areas. The concept of base-camps being used, with hunting parties likely travelling within a 15–40 km walking distance of the site has been proposed as a model of site use, with coastal-inland movements aimed at maximising exploitation of local resources. Animals gradually accumulated, from different parts of the landscape at La Riera, could reflect exploitation from several isotopically distinct locations.

**Discussion**

Throughout the sequence at La Riera a series of changes in environment were detected in the bone collagen δ13C and δ15N isotopic records, and other environmental proxies showed shifts in both animal niches and human hunting habits, which has global significance for enhancing our understanding of the importance of European refugia during the Late Pleistocene.

At the beginning of the sequence, during the Solutrean (20.5-17kyr BP, Zone 1), conditions appeared to be cold and arid, with ericaceous plants dominating the landscape. Throughout Z1, during relatively hostile Late Glacial conditions, ibex were able to adapt their niche and survive. The high-relief environments surrounding La Riera enabled them to successfully alter their feeding behaviour in response to those adverse environmental conditions. Niche partitioning of herbivores is a characteristic observed throughout the Pleistocene, both isotopically and using dental microwear. Archaeological studies have indicated that niche partitioning may not necessarily be predictable through time. Adapting feeding habits in relation to changing environmental conditions can be a valuable survival strategy, reducing competition for resources, potentially reflecting mosaic environments. Despite the pronounced effects of the climatic severity, the suite of environments surrounding La Riera provided alternative habitats for ibex, outside of the niche of the red deer, enabling them to thrive and thus to be available to be hunted extensively by humans near the site. The predictable presence of ibex surrounding La Riera environs may have been one of the major attractions of hunter-gatherer populations during times of hardship.

Studies of Late Glacial macromammals imply that local hunting strategies were adapted to the local conditions, and at La Riera the results indicate that the availability of ibex during this time of climatic turmoil may...
have allowed hunters to adapt their strategies. This may, in part explain, the success of this region for human and animal populations during the LGM as attested by genetic studies in validation of the classic archaeo-

genetic refugium theory. Both red deer and ibex were crucial to the survival of Solutrean populations of Cantabrian Spain, although with fluctuations as to which species was dominant . The environments surrounding La Riera cave would have offered a reliable source of game, a highly favourable characteristic for Late Glacial populations and may have acted as a hub or base camp for Solutrean bands inhabiting the area.

The climatic improvement that occurred at the end of the LGM throughout the Lower–Upper Magdalenian (16–15kyr BP, Zone 2) is represented by a shift in the niche of the ibex, causing them to have more similar δ13C and δ15N values to those of the red deer. This change in values also coincides with a change in vegetation away from ericaceous plants, towards more steppe conditions, and suggests that plants in the landscape could successfully support both species. The consistency in both the δ13C and δ15N isotopic values throughout this period suggests that, the Cantabrian region offered environmental predictability which may explain why this area was so densely occupied with a remarkable increase in the number of archaeological sites, estimated at 10–15 per millennium. Stability in the resource availability, during this newly improved climate and environment, would have been a major draw of the region to human populations.

Colder conditions observed in Level 24 (13.5kyr BP, Zone 3) coincide with a shift towards greater consumption of ibex, a trait observed during the Solutrean at La Riera when the environment was also characterised as being cold. This suggests that La Riera inhabitants were able to adapt their hunting strategies towards targeting different prey when conditions were harsher, indicating human resilience. This cool spell was not reflected in three other sites in the Cantabrian region, where stable isotope records were available and may reflect series of micro-environ-
ments or resource ‘patches’ facilitating survival of different species in around the region effectively, buffering the varying climatic conditions. This environmental buffering, enabled by differential expressions of climate on a local level across the region, may have allowed animals and humans to shift locations depending on resource availability. The movement of populations and the expansion of wider territories into both, coastal and inland areas, has long been discussed as a model of occupation in the region. By having the flexibility to change location, even during times of environmental hardship, the occupants of the Cantabrian Region refugium were able to survive, and even thrive. La Riera was a location that offered abundant game, even during colder more challenging conditions and demonstrates its importance in the network of sites in the local and regional framework.

The final parts of the sequence at La Riera show two key trends. During the Azilian, in Zone 4 (12–9kyr BP), warmer and wetter Holocene conditions are recorded. These conditions continue during the Asturian (Mesolithic) (8.6–6.5kyr BP, Zone 5) with another trend, that is, an increase in the ranges of δ15N values in both animal species. Hunting animals from a wider catchment area, causing a larger spread in the data, and may indicate increased hunting ranges. Expanded hunting ranges could be due to a need to feed larger numbers of people, to maximise food returns. An increase in human demography is seen at this time, supporting this hypoth-

esis. Other changes in diet are also seen, with a wider dietary breadth in the region throughout the Late Upper Magdalenian, Azilian and Asturian, with a greater number of marine species being exploited, along with lower ranked ungulate prey (such as roe deer), species that are more dangerous and difficult to hunt (e.g. wild boar) and with high percentages of juvenile animals. Additionally, more intensive processing of carcasses, such as marrow cracking of small bones, such as phalanges is seen in La Riera from the Upper Magdalenian onwards, and at the site of Las Caldas in central Asturias at this time. In the wider region, stone boiling to maximise nutritional output from carcasses is observed. There have been questions as to whether this dietary change is result of environmental change or population pressures and over-hunting. The evidence from earlier in the La Riera sequence suggests that the major ungulate species (particularly ibex) being exploited by humans at the site were quite resilient and able to adapt to changing conditions, even during times of harsh climatic condi-
tions. This suggests that human population pressures on ungulate populations in Upper Magdalenian–Azilian times may have been responsible for this dietary change. The increase in number of sites per millennia recorded in the archaeological record, throughout the Late Pleistocene, attests to the opportunities offered by the region, even in the face of increased population size. Whilst the environment played a role in the suite of species on offer during the Pleniglacial, the move towards hunting lower ranked, less sustainable prey is likely a product of an increased human population, causing pressure on resource availability.

During the Mesolithic (Asturian) in the Cantabrian region, populations start to abandon much of the interior, moving away from the densely forested inland areas with limited ungulate prey, likely due to over-exploitation, and towards the coastal ecotone strip with newly formed estuaries and inlets, something which is seen at La Riera, where coastal resources (i.e., fish, molluscs, crustaceans) became major and easily collected food resources. This phenomenon has also been observed across the wider Cantabrian region, as attested by the zooarchaeological record and by technological specialisation, and a marked increase in the number of coastal sites emerging in the Mesolithic. A decrease in the exploitation of larger mammals is evident in the region. It is thought that demographic pressures, more than (or at least in combination with) environmental conditions, were likely ultimately responsible for the shift in human diet. Intensive carcass processing, and increased exploitation of smaller species and aquatic resources has been identified across Europe during the latter stages of the Upper Palaeolithic and is often attested to population growth, suggesting that the observations at La Riera are consistent with trends occurring happening on a pan-European scale at this time.

Conclusions
The large-scale climatic changes that dominated the Late Pleistocene have been detected within the bone collagen record δ13C and δ15N correlating with the wider proxies within the archaeological sequence at La Riera, thus providing insights into the conditions experienced by humans occupying this refuge region. The isotopic study has shown that the climatic conditions were responsible for some of the affecting changes in vegetation, and thus plant and faunal δ13C and δ15N values. Behavioural changes associated with those fluctuating conditions were also
integral in producing many of the trends observed in the isotopic record at La Riera, influencing both animal feeding behaviour and human hunting strategies. Humans occupying the site and animals living in the locality of La Riera showed the capacity to adapt to these new environments, and to change their feeding behaviour, demonstrating their resilience. Animals might have altered their ecological niches by feeding in different environmental zones, that were readily available for human exploitation, buffering the effects of harsher climatic conditions. The predictability of resources in the environs of the site, even during more environmentally unstable times, may explain why this region was routinely occupied throughout the Upper Palaeolithic. La Riera was just one site in a dense network of Upper Palaeolithic sites in the Cantabrian Region, and by moving between sites and exploiting locally available sources during different seasons and environmental conditions, humans succeeded to survive, and even thrive during the harsh Late Glacial conditions. During the end of the Upper Palaeolithic, in the Azilian and in the Asturian, environments become wetter and increased hunting ranges are seen, in addition to more intensive carcass processing and dietary diversification. Human and animal populations were able to cope with adapting to more hostile conditions earlier in the sequence, suggesting that these changes during early Holocene may be linked instead to pressure on resources, due to an increased human population size. The multi-proxy approach used here shows the value of integrating lines of evidence to characterise how global climatic changes were expressed on continental and local levels and how these affected the Late Glacial human populations, and has wider applications for understanding the importance of European refugia during these times of climatic turbulence.

Materials and Methods

Bone sampling strategy. Specimens were sampled from the archive collections held at the Museo Arqueológico de Asturias in Oviedo. The archaeological material was organised relating to the original numerical level attributions assigned during excavation which are inverse to the levels discussed and explained in the site monograph. In this study, archaeology excavation levels have been correlated to the levels as discussed throughout the original publication to facilitate comparison with the wider palaeoenvironmental datasets available from the site. Levels with secure stratigraphy, pertaining to discrete temporal units, spanning the Solutrean to the Asturian period, were selected to provide insights into changes in environment through time.

Red deer (Cervus elaphus) and Spanish ibex (Capra pyrenaica), the most consistently consumed species throughout the sequence at the site, and within the wider Cantabrian region during Late Pleistocene were selected for analysis. Mature individuals, as identified from state of bone fusion were selected for analysis to prevent any animals with elevated δ15N nursing signatures, which can mask environmental signatures. Bones with evidence of anthropogenic alteration were targeted, including specimens exhibiting cut marks and green fractures indicative of marrow extraction, to ensure that the palaeoenvironmental evidence was directly linked to periods of human activity.

Analytical methods. Collagen extraction was undertaken following a modified Longin method with an ultra-filtration step. Samples ranging between 0.7–0.9 g were drilled and cleaned using abrasion. Samples were demineralised in 0.5 M HCl at 6–8 °C for between 3–10 days, and then washed three times using de-ionised water before being gelatinised in a weak acidic solution (pH3 HCl) at 70 °C for 48 hours. Samples were filtered using 5–8 μm Ezee® mesh filters (Elkay Laboratory Products), before being ultra-filtered to separate out the larger >30k collagen chains. The >30k fraction was then frozen and lyophilized for 48 hours. Between 0.2–0.45 mg of collagen was weighed into tin capsules for analysis of stable carbon (δ13C) and nitrogen (δ15N) isotope ratios. Specimens were analysed in duplicate, using a Delta XP mass spectrometer coupled to a Flash EA 2112 elemental analyser. The δ13C values and δ15N values are reported relative to the V-PDB standard and AIR standards. A series of international and internal standards were used to calculate analytical error which was ± 0.1‰ (1σ) or better. The mean difference observed between duplicate measurements was 0.03 for δ13C, and 0.01 for δ15N. Data were not normally distributed, meaning that the nonparametric Mann-Whitney U test, with a post-hoc Holm-Bonferroni correction was used. A p-value of < 0.05 or less was deemed to be indicative of a statistically significant result.

Comparative environmental and climatic proxies. Macromammalian faunal information from La Riera was compiled and plotted as %NISP of the major prey species exploited to create Fig. 4, excluding minor species such as carnivores represented by occasional fragments. The full faunal spectrum from La Riera is provided in Supplementary SI2. For the creation of Fig. 3 species were sorted by habitat type for the creation of Fig. 3. Red deer and wild boar were included in the mixed/forest category. Montane species were, predominantly, represented by ibex, which preferentially occupy rocky cliffs and slopes such as in the nearby Sierra de Cuera. Species attributed to open landscapes included horse and bovids (aurochs and bison which are often not possible to specify further). Microfaunal remains at the site were scarce, likely linked to sampling strategies used to recover materials (i.e., hand-picking from fine-mesh water-screen residues, rather than from flotation or dedicated stratigraphic column sampling), meaning that the assemblage may not be entirely representative of the full suite of species inhabiting the landscape. Environmental indicators species were identified including Glis glis (dormouse), which is characteristic of more temperate environments and Microtus oeconomus (root or tundra vole), indicative of cold (and humid) environments, and only their presence is indicated in Fig. 3. The microfaunal data from La Riera are also provided in Supplementary SI3. Pollen data available from la Riera were collated into categories that have been combined into Mediterranean forest, Atlantic forest, heath and steppic environments (see Supplementary SI4). A summary of evidence and interpretations of the sedimentary record is included in SI5.
throughout the study period27 (see Supplementary SI7). Speleothem δ18O and δ13C records from El Pindal cave (also located in Asturias) were collated28 (see Supplementary SI8). Additional bone collagen δ13C and δ15N studies from across the region have been included for comparative purposes. Data was available from La Paloma, roughly 100 km southeast of La Riera40, El Mirón Cave40 (on the northern edge of the Cantabrian Cordillera in eastern Cantabria, about 160 km east of La Riera and 20 km from the present shoreline at 260 m a.s.l.). Two more sites from the Basque region were also included: Antolihakoa Koba49 and the palaeontological site of Kiputz IX48 (all shown in Fig. 1). Data from red deer was included to enable cross-comparisons, but ibex data were not available (see Supplementary SI6). Isotopic data were plotted according to the radiocarbon ages rather than archaeologcal levels, to enable direct comparisons with results generated from La Riera.

Records available from the continental shelf, including the percentages of Neogloboquadrina pachyderma sin. Ostracods, δ18O values from Globigerina bulloides and interpretations were available31 using data from cores KS05-10, KS05-05, and KS04-16 (see Supplementary SI9). This enables a regional, offshore perspective of environment for comparison, and are the closet available records to this study region. These provide an indication of regional temperature changes.

Finally, NGRIP data δ18O values available25 were plotted at 500-year intervals in Fig. 3 to show general climatic trends throughout the study period (see Supplementary SI10) facilitating correlations of site based, and local data. Global climatic events42 are also marked within Fig. 3.

Data availability
All data generated or analysed during this study are included in this published article, and within the associated Supplementary Information Files SI1–10).

Received: 10 May 2019; Accepted: 31 December 2019; Published online: 27 January 2020

References
1. Müller, U. C. et al. The role of climate in the spread of modern humans into Europe. Quat. Sci. Rev. 30, 273–279 (2011).
2. Tallavaara, M., Luoto, M., Korhonen, N., Järvinen, H. & Seppä, H. Human population dynamics in Europe over the Last Glacial Maximum. Proc. Natl. Acad. Sci. 112, 8232–8237 (2015).
3. Verpoorte, A. Limiting factors on early modern human dispersals: the human biogeography of late Pleniglacial. Europe. Quat. Int. 201, 77–85 (2009).
4. Banks, W. E. et al. Human ecological niches and ranges during the LGM in Europe derived from an application of eco-cultural niche modeling. J. Archaeol. Sci. 35, 481–491 (2008).
5. Burke, A. et al. Exploring the impact of climate variability during the Last Glacial Maximum on the pattern of human occupation of Iberia. J. Hum. Evol. 73, 35–46 (2014).
6. Burke, A., Riel-Salvatore, J. & Barton, C. M. Human response to habitat suitability during the Last Glacial Maximum in Western Europe. J. Quat. Sci. 33, 335–345 (2018).
7. Gamble, C., Davies, W., Pettitt, P. & Richards, M. Climate change and evolving human diversity in Europe during the last glacial. Philos. Trans. R. Soc. B Biol. Sci. 359, 243–254 (2004).
8. Soares, P. et al. The Archaeoecogenetics of Europe. Current Biology 20 (2010).
9. Sommer, R. S. & Nadachowski, A. Glacial refugia of mammals in Europe: evidence from fossil records. Mamm. Rev. 36, 251–265 (2006).
10. Straus, L. G. Human Geography of the Late Upper Paleolithic in Western Europe: present State of the Question. J. Anthropol. Res. 47, 259–278 (1991).
11. Straus, L. G. Iberia before the Iberians: the Stone Age prehistory of Cantabrian Spain. (University of New Mexico Press, 1992).
12. Straus, L. G. A quarter-century of research on the Solutrean of Vasco-Cantabria, Iberia and beyond. J. Anthropol. Res. 56, 39–58 (2000).
13. Straus, L. G. Recent developments in the study of the Upper Paleolithic of Vasco-Cantabrian Spain. Quat. Int. 364, 255–271 (2015).
14. Consuegra, S. et al. Mitochondrial DNA variation in Pleistocene and modern Atlantic salmon from the Iberian glacial refugium. Mol. Ecol. 11, 2037–48 (2002).
15. Meiri, M. et al. Late-glacial recolonization and phylogeography of European red deer (Cervus elaphus L.). Mol. Ecol. 22, 4711–4722 (2013).
16. Achilli, A. et al. The Molecular Dissection of mtDNA Haplogroup H Confirms That the Franco-Cantabrian Glacial Refugia Was a Major Source for the European Gene Pool. Am. J. Hum. Genet. 75, 910–918 (2004).
17. Fu, Q. et al. The genetic history of Ice Age Europe. Nature 534, 200–205 (2016).
18. de la Rasilla Vives, M. & Straus, L. The Last Glacial Maximum Settlement of Cantabrian Spain: the Gravettian and Solutrean. Kult. (cercle Aneu) 8, 209–241 (2004).
19. Straus, L. G., Bicho, N. & Winegardner, A. C. The Upper Palaeolithic settlement of Iberia: First-generation maps. Antiquity 74, 553–566 (2000).
20. Straus, L. G., Bicho, N. & Winegardner, A. C. Mapping the Upper Paleolithic regions of Iberia. J. Iber. Archaeol. 2, 7–42 (2000).
21. Straus, L. G. The Upper Paleolithic of Iberia. Trab. Prehist. 75, 9 (2018).
22. Straus, L. G., González Morales, M. R., Martínez, M. A. F. & García-Gelabert, M. P. Last Glacial human settlement in eastern Cantabria (northern Spain). J. Archaeol. Sci. 29, 1403–1414 (2002).
23. Straus, L. G. Southwestern Europe at the Last Glacial Maximum. Carr. Anthropol. 32, 189–199 (2002).
24. Rasmussen, S. O., Vinther, B. M., Clausen, H. B. & Andersen, K. K. Early Holocene climate oscillations recorded in three Greenland ice cores. Quat. Sci. Rev. 26, 1907–1914 (2007).
25. Rasmussen, S. O. et al. A stratigraphic framework for abrupt climatic changes during the Last Glacial period based on three synchronized Greenland ice-core records: Refining and extending the INTIMATE event stratigraphy. Quat. Sci. Rev. 106, 14–28 (2014).
26. Roffe, J. et al. Combining Small-Vertebrate, Marine and Stable-Isotope Data to Reconstruct Past Environments. Sci. Rep. 5, 1–12 (2015).
27. Moreno, A. et al. A compilation of Western European terrestrial records 60–8kaBP: Towards an understanding of latitudinal climatic gradients. Quat. Sci. Rev. 106, 167–185 (2014).
28. Moreno, A. et al. A speleothem record of glacial (25–11.6 kyr BP) rapid climatic changes from northern Iberian Peninsula. Glob. Planet. Change 71, 218–231 (2010).
65. Körner, C., Farquhar, G. D. & Wong, S. C. Carbon isotope discrimination by plants follows latitudinal and altitudinal trends. Oecologia 88, 30–40 (1991).
66. Feranec, R., García, N., Díez, J. C. & Arsuaga, J. L. Understanding the ecology of mammalian carnivores and herbivores from Valdegoba cave (Burgos, northern Spain) through stable isotope analysis. Palaeogeogr. Palaeoclimatol. Palaeoecol. 297, 263–272 (2010).
67. Ehleringer, J. R. & Cerling, T. E. Atmospheric carbon dioxide and the ratio of intercellular to a. (1995).
68. Williams, D. G. et al. Global isotope discrimination by Sorghum bicolor under CO2 enrichment and drought.
69. Amundson, R. et al. Global patterns of the isotopic composition of soil and plant nitrogen. Glob. Biogeochim. Cycles 17, 1031 (2003).
70. Austin, A. T. & Vitousek, P. M. Nutrient dynamics on a precipitation gradient in Hawaiʻi. Oecologia 113, 519–529 (1998).
71. Arranž, J. N. et al. Nitrogen cycling in the soil-plant system along a precipitation gradient in the Kalahari sands. Glob. Chang. Biol. 10, 359–373 (2004).
72. Heaton, T. H. E., Vogel, J. C., Von La Chevallerie, G. & Collett, G. Climatic influence on the isotopic composition of bone nitrogen. Nature 322, 822–823 (1986).
73. Cormier, A. B. & Schwarz, H. P. Effects of climate on deer bone δ15N and δ13C: Lack of precipitation effects on δ15N for animals consuming low amounts of C4 plants. Geochim. Cosmochim. Acta 60, 4161–4166 (1996).
74. Gröcke, D. R., Bocherens, H. & Mariotti, A. Annual rainfall and nitrogen-isotope correlation in macropod collagen: application as a palaeoprecipitation indicator. Earth Planet. Sci. Lett. 153, 279–285 (1997).
75. Murphy, B. P. & Bowman, D. M. J. S. Kangaroo metabolism does not cause the relationship between bone collagen δ15N and water availability. Funct. Ecol. 20, 1062–1069 (2006).
76. Hartman, G. Are elevated δ15N values in herbivores in hot and arid environments caused by diet or animal physiology? Funct. Ecol. 25, 122–131 (2011).
77. Handley, L. L. et al. The 15N natural abundance (δ15N) of ecosystem samples reflects measures of water availability. Funct. Plant Biol. 26, 185 (1999).
78. Hobbie, E. A., Jumpson, A. & Trappe, J. Foliar and fungal 15 N:14 N ratios reflect development of mycorrhizae and nitrogen recycling in temperate and boreal ecosystems? Evidence from modern and ancient ungulates. Palaeogeogr. Palaeoclimatol. Palaeoecol. 266, 69–82 (2008).
79. Straus, L. & Clark, G. A. La Riera Cave: Stone Age hunter-gatherer adaptations in northern Spain. (Anthropological Research Papers 36, Arizona State University, 1986).
80. Eastham, A. The La Riera avifauna. In La Riera Cave: Stone Age hunter-gatherer adaptations in northern Spain (eds Straus, L. G. & Clark, G.) 274–284 (Anthropological Research Papers 36, Arizona State University, 1986).
81. Deith, M. & Shackleton, N. Seasonal exploitation of marine molluscs: Oxygen isotope analysis of shell from La Riera. In La Riera Cave: Stone age hunter-gatherer adaptations in northern Spain (eds. Straus, L. G. & Clark, G.) 299–313 (Anthropological Research Papers 36, Arizona State University, 1986).
82. Laville, H. Stratigraphy, sedimentology, and chronology of the La Riera Cave deposits. In La Riera Cave: Stone Age hunter-gatherer adaptations in northern Spain. (eds. Straus, L. G. & Clark, G.) 25–55 (Anthropological Research Papers 36, Arizona State University, 1986).
83. Ortega, J. The malacology of La Riera Cave. In La Riera Cave: Stone Age hunter-gatherer adaptations in northern Spain (eds. Straus, L. G. & Clark, G.) 289–298. (Anthropological Research Papers 36, Arizona State University, 1986).
84. Craighead, A. Climate change and patterns in the exploitation of economic resources (marine mollusca and ungulate fauna) in Cantabrian Spain at the end of the Pleistocene, ca. 21–6.5 kyr BP. In Zooarchaeology of the Pleistocene/ Holocene Boundary (ed. Driver, J. C.) 9–20 (British Archaeological Reports, 1999).
85. Ortiz, J. E. et al. The aminochronology of man-induced shell middens in caves in Northern Spain. Archaeometry 51, 123–139 (2009).
86. Straus, L. G. & González Morales, M. R. New Dates for the Solutrean and Magdalenian of Cantabrian Spain: El Miron and la Riera. Palaeogeogr. Palaeoclimatol. Palaeoecol. 189–201 (2001).
87. Bertran, A. B. & Schwarcz, H. P. Effects of climate on deer bone δ15N and δ13C: Lack of precipitation effects on δ15N for animals consuming low amounts of C4 plants. Geochim. Cosmochim. Acta 60, 4161–4166 (1996).
88. Ortiz, J. E. et al. A compilation of Western European terrestrial records 60–8kaBP: Towards an understanding of latitudinal climatic gradients. Quat. Sci. Rev., https://doi.org/10.1016/j.quascirev.2014.06.030 (2014).
89. DeNiro, M. J. Postmortem preservation and alteration of in vivo bone collagen isotope ratios in relation to palaeodietary reconstruction. Nature 317, 806–809 (1985).
90. Milne, J. A., Macrae, J. C., Spence, A. M. & Wilson, S. A comparison of the voluntary intake and digestion of a range of forages at different times of the year by the sheep and the red deer (Cervus elaphus). Br. J. Nutr. 30, 347–357 (1978).
91. Milne, J. A., Macrae, J. C., Spence, A. M. & Wilson, S. A. The biology of mycorrhiza in the Ericaceae XX. Plant and mycorrhizal nematodes as nitrogenous substrates for the ericoid mycorrhizal fungus Hymenoscyphus ericae and its host. New Phytol. 119, 353–360 (1998).
92. Milne, J. A., Macrae, J. C., Spence, A. M. & Wilson, S. A. A comparison of the voluntary intake and digestion of a range of forages at different times of the year by the sheep and the red deer (Cervus elaphus). Br. J. Nutr. 40, 347–357 (1978).
93. Milne, J. A., Macrae, J. C., Spence, A. M. & Wilson, S. A. The biology of mycorrhiza in the Ericaceae XX. Plant and mycorrhizal nematodes as nitrogenous substrates for the ericoid mycorrhizal fungus Hymenoscyphus ericae and its host. New Phytol. 119, 353–360 (1998).
94. van Klinken, G. J. Bone Collagen Quality Indicators for Palaeodiets and Radiocarbon Measurements. J. Archaeol. Sci. 26, 687–695 (1999).
95. García-Guixé, E., Martínez-Moreno, J., Mora, R., Núñez, M. & Richards, M. P. Stable isotope analysis of human and animal remains from the Late Upper Palaeolithic site of Balma Guilanyà, southeastern Pre-Pyrenees, Spain. J. Archaeol. Sci. 36, 1018–1026 (2009).
96. Montalvo, J., Casado, M. A., Levassor, C. & Pineda, F. D. Species diversity patterns in Mediterranean grasslands. J. Veg. Sci. 4, 213–222 (1993).
97. Kerley, S. J. & Read, D. J. The biology of mycorrhiza in the Ericaceae XX. Plant and mycorrhizal nematodes as nitrogenous substrates for the ericoid mycorrhizal fungus Hymenoscyphus ericae and its host. New Phytol. 139, 353–360 (1998).
98. Milne, J. A., Macrae, J. C., Spence, A. M. & Wilson, S. A comparison of the voluntary intake and digestion of a range of forages at different times of the year by the sheep and the red deer (Cervus elaphus). Br. J. Nutr. 30, 347–357 (1978).
99. Heiter, A., Gordon, I., Baille, G. J. & Tappin, E. Foraging behaviour of sheep and red deer within natural heather/grass mosaics. J. Appl. Ecol. 36, 133–146 (1999).
100. Azevedo, P., Cassinello, J. & Gortazar, C. The Iberian ibex is under an expansion trend but displaced to suboptimal habitats by the presence of extensive goat livestock in central Spain. Biodivers. Conserv. 16, 3361–3376 (2007).
101. Clutton-Brock, T. H. & Albon, S. D. Red deer in the Highlands. (Blackwell, 1989).
102. Geibert, C. & Verheyden-Tixier, H. Variations of diet composition of Red Deer (Cervus elaphus L.) in Europe. Mammm. Res. 31, 189–201 (2001).
103. Straus, L. G. On the habitat and diet of Cervus Elaphus. Munibe 33, 175–182 (1981).
104. Rivals, F., Schulz, E. & Kaiser, T. M. Late and middle Pleistocene ungulates dietary diversity in Western Europe indicate variations of Neanderthal paleoenvironments through time and space. Quat. Sci. Rev. 28, 3388–3400 (2009).
105. Acvedo, P. & Cassinello, J. Biology, ecology and status of Iberian ibex Capra pyrenaica: A critical review and research prospectus. *Mamm. Rev.* **39**, 17–32 (2009).
106. Alados, C. L. & Escos, J. Relationships between movement rate, agonistic displacements and forage availability in Spanish ibexes (Capra pyrenaica). *Biol. Behav.* **12**, 245–255 (1987).
107. García-González, R. & Cuartas, P. Food habits of Capra pyrenaica, Cervus elaphus and Dama dama in the Cazorla Sierra (Spain). *Mammalia* **56**, 195–202 (1992).
108. Hill, P. W., Handley, L. L. & Raven, J. A. Juniperus communis L. ssp. communis at Balnaguard, Scotland: Foliar carbon discrimination (? 13 C) and 15-N natural abundance (? 15 N) suggest gender-linked differences in water and N use. *Bot. J. Scotl.* **48**, 209–224 (1996).
109. O'Leary, O. H. Environmental effects on carbon isotope fractionation in terrestrial plants. In *Stable Isotopes in the Biosphere* (eds. Wada, E., Toneyama, T., Minagawa, M., Ando, T. & Fry, B. D.) 78–91 (Kyoto University Press, 1995).
110. Álvarez Lao, D. J. Macromamíferos fósiles del Pleistoceno de Asturias. *Nat. Cantab.* **2**, 11–24 (2003).
111. Strauss, L. G. Late Würm adaptive systems in Cantabrian Spain: The case of eastern Asturias. *J. Anthropol. Archaeol.* **5**, 330–368 (1986).
112. Feranec, R. S., Hadly, E. A. & Paytan, A. Stable isotopes reveal seasonal competition for resources between late Pleistocene bison (Bison) and horse (Equus) from Rancho La Brea, southern California. *Palaeogeogr., Palaeoclimatol., Palaeoecol.* **271**, 153–160 (2009).
113. Rivals, F. & Lister, A. M. Dietary flexibility and niche partitioning of large herbivores through the Pleistocene of Britain. *Quat. Sci. Rev.* **146**, 116–133 (2016).
114. Jones, E. L. Archaeofaunal evidence of human adaptation to climate change in Upper Paleolithic Iberia. *J. Archaeol. Sci. Reports* **2**, 257–263 (2015).
115. Strauss, L. G. Terminal Pleistocene faunal exploitation in Cantabria and Gascony. In *Animals and archaeology: Hunters and their prey* (eds. J. C.-B. & Grigson, C.) 209–225 (British Archaeological Reports, 1983).
116. Bailey, G. N. & Davidson, I. Site exploitation territories and topography: Two case studies from palaeolithic Spain. *J. Archaeol. Sci.* **10**, 87–115 (1983).
117. Butzer, K. Paleolithic adaptations and settlement in Cantabrian Spain. In *Advances in World Archaeology (Volume 5)* (eds. Wendendorf, F. & Close, A.I.) 201–252 (Academic Press, 1986).
118. Marín-Arroyo, Y., Strauss, L. G. Human response to holocene warming on the cantabrian coast (Northern Spain): An Unexpected Outcome. *Quat. Sci. Rev.* **81**, 1–11 (2013).
119. Mateos Cachorro, A. Apuntes sobre estacionalidad y subsistencia de los grupos humanos del Cantábrico occidental en torno al 12,000 bp. *Trab. Prehist.* **39**, 27–41 (2002).
120. Mateos Cachorro, A. Estudio de la fragmentación de falanges y mandíbulas en la secuencia temporal del 19000–13000 BP de la Cueva de las Caldas (Priorio, Oviedo). Implicaciones paleoecológicas. *Gall. Rev. Arqueol. e ant Igualdade* **22**, 9–20 (2003).
121. Nakazawa, Y., Strauss, L. G., González-Morales, M. R., Solana, D. C. & Sáiz, J. C. On stone-boiling technology in the Upper Paleolithic: behavioral implications from an Early Magdalennian hearth in El Mirón Cave, Cantabria, Spain. *J. Archaeol. Sci.* **36**, 684–693 (2009).
122. Freeman, L. G. The Significance of Mammalian Faunas From Paleolithic Occupations in Cantabrian Spain. *Am. Antq.* **38**, 3–44 (1973).
123. Strauss, L. G. Of deerslayers and mountain men: Paleolithic faunal exploitation in Cantabrian Spain. In *For theory building in archaeology* (ed. Binford, L. R.) 41–76 (Academic Press, 1977).
124. González Morales, M., Strauss, L. G., Díez, A. & Ruiz, J. Postglacial coast and Inland: The Epipaleolithic-Mesolithic-Neolithic Transitions in the Vasco-Cantabrian Region. *Mundo Antropol.* **56**, 61–78 (2004).
125. Strauss, L. G. A Quater-Century of Research on the Solutrean of Vasco-Cantabria. *J. Anthropol. Res.* **56**, 39–58 (2009).
126. Strauss, L. G. et al. Paleoequology at La Riera (Asturias, Spain). *Curr. Anthropol.* **22**, 655–682 (1981).
127. Adán, G. E., Álvarez-Lao, D., Turrero, P., Arbizu, M. & García-Vázquez, E. Fish as diet resource in North Spain during the Upper Paleolithic. *J. Archaeol. Sci.* **36**, 895–899 (2009).
128. Álvarez-Fernández, E. Humans and marine resource interaction reappraised: Archaeofauna remains during the late Pleistocene and Holocene in Cantabrian Spain. *J. Anthropol. Archaeol.* **30**, 327–343 (2011).
129. Clark, G. The Asturian of Cantabria: Subsistence Base and the Evidence for Post-Pleistocene Climatic Shifts. *Am. Anthropol.* **124**–1257 (1971).
130. Clark, G. A. Asturian of Cantabria: Early Holocene Hunter-Gatherers in Northern Spain. (University of Arizona Press, 1983).
131. Strauss, L. G. The late upper paleolithic-mesolithic-neolithic transitions in cantabrian Spain. *J. Anthropol. Res.* **65**, 287–298 (2009).
132. Clark, G. A. & Lerner, S. Prehistoric resource utilization in early Holocene Cantabrian Spain. *Anthropol. UCLA* **53**–96 (1980).
133. Strauss, L. G., Clark, G., Ágel, J. & Rato, O. Ice-Age Subsistence in Northern Spain. *Sci. Am.* **242**, 142–153 (1980).
134. Clark, G. Iberian mesolithic in the European contextTitle. In *The Mesolithic of the Atlantic Façade* (Eds. González Morales, M. R. & Clark, G.) 205–224 (Arizona State University, 2004).
135. Stiner, M. C. et al. The Tortoise and the Hare: Small-Game use, the Broad-Spectrum Revolution, and Paleolithic Demography. *Curr. Anthropol.* **41**, 39–79 (2000).
136. Starkovich, B. M., Munro, N. D. & Stiner, M. C. Terminal Pleistocene subsistence strategies and aquatic resource use in southern Greece. *Quat. Int.* **465**, 162–176 (2018).
137. Schurr, M. R. Stable Nitrogen Isotopes as Evidence for the Age of Weaning at the Angel Site: A Comparison of Isotopic and Demographic Measures of Weaning Age. *J. Archaeol. Sci.* **24**, 919–927 (1997).
138. Schurr, M. R. Using stable nitrogen-isotopes to study weaning behavior in past populations. *World Archaeol.* **30**, 327–342 (1998).
139. Longin, R. New Method of Collagen Extraction for Radiocarbon Dating. *Nature* **230**, 241–242 (1971).
140. Brown, T. A., Nelson, D. E., Vogel, J. & Southon, J. Improved collagen extraction by modified Longin method. *Radiocarbon* **30**, 171–177 (1988).
141. Holm, S. A Simple Sequentially Rejective Multiple Test Procedure. *Scand. J. Statist.* **6**, 65–70 (1979).
142. Seierstad, I. K. et al. Consistently dated records from the Greenland GRIP, GISP2 and NGRIP ice cores for the past 104ka reveal regional millennial-scale 818O gradients with possible Heinrich event imprint. *Quat. Sci. Rev.* **106**, 29–46 (2014).
143. António, M. M. S. et al. Marine Radiocarbon Reservoir Effect in Late Pleistocene and Early Holocene Coastal Waters off Northern Iberia. *Radiocarbon* **58** (4): 869–883 (2016).

**Acknowledgements**

This research was funded by a Marie Skłodowska-Curie Individual Fellowship grant awarded to JRJ (H2020-MSCA-IF-2014-656122). ABMA was supported by a Marie Curie Career Integration Grant (FP7-PEOPLE-2012-CIG-322112), the Spanish Ministry of Economy and Competitiveness (HAR2012-33956 and Ramón y Cajal (RYC-2011-07690) and Santander Talent Attraction for Research (STAR-1). The new radiocarbon dates obtained by LGS for Levels 23 and 24 were kindly financed by the Stone Age Research Fund (University of New Mexico), I. and R. Auel principal donors. The authors would like to thank L. Agudo Pérez for helping to draw...
Figs. 1 and 3 and for her assistance sampling the remains from La Riera. Permission to recreate the stratigraphic section from the La Riera monograph was very kindly given by Geoffrey Clark (Arizona State University), editor of the Anthropological Research Papers series. Many thanks to C. Garcia de Castro at the Museum of Asturias for facilitating access to their collections. Special thanks to R. Macdonald and M. Wong (University of British Columbia) for their laboratory and technical assistance.

Author contributions
Data generation and analysis: J.J. Funding acquisition: J.J. Investigation: J.J., A.B.M.A., M.R. and L.G.S., Methodology: M.R. and J.J. Resources: A.B.M.A. and M.R., Writing – original draft: J.J. Writing – review & editing: J.J., A.B.M.A, L.G.S. and M.R.

Competing interests
The authors declare no competing interests.

Additional information
Supplementary information is available for this paper at https://doi.org/10.1038/s41598-020-57715-2.

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