The different characteristics of the recent eruptions of Fernandina and Sierra Negra volcanoes (Galápagos, Ecuador)

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Abstract

After eight years of quiescence, Fernandina volcano experienced two short-lived eruptions, on 4 September 2017 and 16 June 2018. The eruptions were characterized by very short periods of unrest that started a few hours before the initiation of the eruptive activity. On the other hand, Sierra Negra volcano (Isabela Island) began a new eruptive period on 26 June 2018, after almost one year of persistent unrest characterized by an increase in the magnitude and number of seismic events and more than 5 meters of uplift since its last eruption in 2005. The Sierra Negra and Fernandina eruptions were located in remote zones where access is extremely complex. Thus, satellite images complement the continuous monitoring data of the Instituto Geofísico (IG-EPN) with remote observations and allow rapid response mapping in order to identify the areas affected by the lava flows. Finally, the aim of this Report is to encourage other scientists to investigate the behaviors of both pre-eruptive and eruptive periods registered during these eruptions.

Keywords: Galapagos; Sierra Negra; Fernandina; Remote sensing; Monitoring; Hazard Assessment

1 Introduction

The Galápagos Islands are a group of basaltic shield volcanoes related to a hotspot that currently lies 170 km south of the Galápagos Spreading Center [Allan and Simkin 2000]. The western volcanoes located on Isabela and Fernandina islands show frequent activity with approximately 69 eruptions reported since the early 1800s [Global Volcanism Program 2013c]. For instance, Fernandina volcano has had at least 28 historical eruptions [Global Volcanism Program 2013a] including one in 2017 and 2018, and Sierra Negra had 13 [Global Volcanism Program 2013b] including one in 2018. Nonetheless, geophysical studies have been limited due to the islands’ remote location [Rowland et al. 2003; Chadwick et al. 2011] and also because the only permanent human settlement in the area is the small town of Puerto Villamil at the south of Isabela. Thus, according to Rowland et al. 2003, p. 314 "many of the eruptions have gone unnoticed […]. Although the human population has grown, access to eruption sites is still difficult". Only since the 1990s have eruptions been systematically observed and instrumentally monitored [e.g. Allan and Simkin 2000; Rowland et al. 2003; Chadwick et al. 2011; Geist et al. 2008; Bagnardi et al. 2013; Bernard et al. 2015; Xu et al. 2016].

In 2014, Instituto Geofísico (IG-EPN) installed new broadband seismometers on Fernandina (2) and Isabela (4), allowing a much better follow-up of the activity and for the first time allowing the study of seismic unrest.
in detail. This Report presents an overview of the time series (seismic, ground deformation, SO$_2$, thermal) before and during recent eruptions in Galápagos (2017 and 2018 Fernandina eruptions and 2018 Sierra Negra). This information allows us to illustrate significant differences between those volcanic systems and their behavior.

2 Summary of the eruptions

The following summary is set in Coordinated Universal Time (UTC) which for Galápagos Archipelago means UTC-6 (local time).

2.1 Fernandina (Fernandina Island)

After 8 years of quiescence, Fernandina (0.37°S, 91.55°W, 1476 m a.s.l.), the westernmost volcano of the Galápagos Archipelago began a new eruptive phase on 4 September 2017. At 11h34 the IG-EPN monitoring network detected an increase in hybrid seismic events, which presented an energy peak at 16h20. Seismic signals changed at 17h25 from hybrid to low-frequency events (LF) and finally at 18h25 volcanic tremor was recorded. This tremor was associated with the beginning of the eruption and extrusion of lava flows (Figure 1A). Additionally, Interferometric Synthetic Aperture Radar (InSAR) detected uplift of 17 centimeters from March 2015 to September 2017 on the floor of the caldera (31 km$^2$), the final five centimeters of which occurred in the last two months before the eruption [IG-EPN 2017].

The first thermal anomaly detected by GOES-16 satellite was at 18h30 close to the caldera rim, five minutes after the onset of the eruption based on volcanic tremor (Figure 1A). The eruption generated a column that reached 4000 m a.s.l., composed mainly of volcanic gases with low ash content drifting to the west. The eruption originated from a three-kilometer-long arcuate fissure located on the southwestern upper-flank of the volcano, at the same location as the 2005 eruption [Chadwick et al. 2011; Bagnardi et al. 2013]. The fissure emitted several lava flows that covered an area of 6.5 km$^2$ (Figure 2). The maximum run-out of the flows was four kilometers downslope. These lava flows did not reach the ocean. After two and a half days, a decrease in the seismic energy back to background levels indicated that the eruption concluded (Figure 1A). SO$_2$ degassing was clearly seen by OMI and OMPS satellite instruments. The highest SO$_2$ mass observed by OMPS was on September 5, reaching 8.2×10$^{10}$ kg.

The most dangerous hazard associated with this eruption was the series of wildfires triggered by the high temperatures of the lava flows and the prevailing surface wind direction. The fires burned an area of 16 km$^2$ on the western flank where endemic fauna and flora were present (Figure 2). The wildfires ended one month after the eruption concluded.

On 16 June 2018, nine months after the previous eruption, a new period of unrest began. The seismic monitoring network of the IG-EPN detected nine earth-Quakes larger than 2.5 M$_{LV}$, starting at 14h37, with the largest event of 4.1 M$_{LV}$ occurring at 15h22. At 17h15 volcanic tremor was observed indicating the beginning of the eruption and lava emission (Figure 1B). The eruption occurred on the northern flank of the volcano 1.5–2 km downslope of the caldera rim (Figure 2). It produced an eruptive column of volcanic gases and low ash content drifting to the southwest that reached 3000 m a.s.l. The radial fissures were 100–200 meters long and extruded lava flows that reached the ocean (4 km run-out) and covered 1.58 km$^2$ (Figure 2). Similar to the previous eruption, this ended two days after the seismic activity began (Figure 1B). SO$_2$ degassing reached a minimum of 32×10$^6$ kg according to the OMPS image of July 17. InSAR data did not show significant deformation from September 2017 to June 2018 [Paul Lundgren personal comm.].

2.2 Sierra Negra (Isabela Island)

Sierra Negra volcano (0.81°S, 91.13°W, 1124 m a.s.l.) is located at the southern end of Isabela Island. Its last eruption, in 2005, was characterized by the extrusion of 150 million cubic meters (150 Mm$^3$) of lava flows that covered a large part of the caldera floor and a smaller area to the north of the caldera [Geist et al. 2008]. According to Global Volcanism Program 2013b, Sierra Negra averages one eruptive period every 11–12 years with most of the recent activity located on the northern flank and inside its vast caldera (82 km$^2$).

Since July 2017, persistent signals of unrest were detected by the IG-EPN. Thousands of earthquakes (>15,000) located inside and at the rim of the caldera were recorded at shallow depths (1–15 km). Many of them were larger than 3 M$_{LV}$ and some even larger than 4 M$_{LV}$ (11 seismic events). Galápagos National Park rangers housed at the touristic rim entrance, called El Cura, regularly felt the largest earthquakes. These events also produced small rockfalls at the walls of the caldera. Most of these seismic events were volcAno-tectonic (VT) events, but low-frequency (LF), very low-frequency (VLF) and hybrid events were also recorded. The number and magnitude of the seismic events increased significantly until the beginning of the eruption in June 2018.

A high-precision GPS (cGPS) network installed by UNAVCO, University of Idaho, and Oregon State University, and now maintained by Penn State University and UNAVCO, detected more than 4.5 meters of vertical uplift since 2006 (Figure 3A). InSAR images processed at the Rosenstiel School of Marine and Atmospheric Science (RSMAS, https://insarmaps.miami.edu/) confirm this result and detected a cumulative uplift of 2.46 meters on
Figure 1: Median Seismic Amplitude data in 10-minute windows during two Fernandina volcano eruptions. Traces are corrected for instrument response and filtered between 1 – 8 Hz.: [A] 2017, seismic station nearest FER2 and [B] 2018, seismic station nearest FER1. See Figure 2 for the location of the seismic stations.

Figure 2: Map of the eruptions of Fernandina volcano in September 2017 and June 2018. Note color lines showing the area burnt by wildfires triggered by the lava flows and the prevailing wind during the 2017 eruption. Green dots are the location of the seismic stations FER1 and FER2. Data provided by: Sentinel-2: 21 June, 16 July 2018 and 9, 19 and 29 Sept 2017. Landsat-8 12 Sept 2017. Hillshade provided by JAXA. Inset shows the main islands of Galápagos Archipelago and the red rectangle the location of Fernandina volcano.

Figure 3: [A] Time-series of the vertical component for three cGPS stations on the rim (GV01) and within the caldera (GV02, GV04) at Sierra Negra provided by Penn State (see Figure 5 for locations). The cumulative uplift exceeds 4.5 meters. The cGPS stations provide high-temporal sampling of the deformation signal. [B] Green box shows the time-series inversion assessed by InSAR since 2014 on GV02’s location using Sentinel-1 (128-path) taken from RSMAS (https://insarmaps.miami.edu/). Note that cGPS and InSAR trends looks similar in the same period of time. Dashed lines show important events in the volcano.

On Tuesday, 26 June 2018 at 09h15, a 5.3 M\textsubscript{L}V earthquake with a focal mechanism indicating reverse slip occurred near the southwestern caldera rim at 5.3 km depth (Figure 4A, Figure 5). This earthquake was large enough to be felt and reported by residents in nearby Puerto Villamil located 19 km southeast of the epicenter (see inset in Figure 5). It produced an uplift of the center of the caldera floor since December 2014 (Figure 3B). According to the cGPS network, since January 2018 the uplift rate accelerated and was close to 11 centimeters per month.
the center of the caldera floor before the eruption began. After a few hours of relative seismic quiescence, at 17h17, an intense seismic swarm commenced and included a 4.6 M_{LV} earthquake and several VLF, VT and LF events embedded in the sequence. Finally, at 19h40 volcanic tremor began to dominate the seismic record indicating the beginning of the eruption (Figure 4A).

GOES-16 satellite images identified an eruptive column of volcanic gases at 10.5 km a.s.l. drifting to the Northwest in addition to thermal anomalies at the northern flank of the caldera.

Two main phases were recognized during the eruption, which occurred along five eruptive fissures:

1. Phase one lasted less than one day and was the most intense. During this phase, fissures 1, 2, 3, 4 and 5 were active and lava flows covered an area greater than 17 km² (see Figure 5).

2. Phase two lasted 57 days and was focused only on fissure 4 (Figure 5). This fissure extruded several lava flows which in total covered an area of 13 km².

**Figure 5** shows the five fissures mentioned. They will be briefly described below:

- **Fissure 1 (1070 m a.s.l.)** is a tangential discontinuous fissure of four kilometers in length, aligned WNW to the caldera rim, and located in the *Volcán Chico* sector. This fissure produced a lava fountain that extruded several lava flows covering 14.7 km² in one day of activity. The maximum run-out of these flows was seven kilometers downslope.

- **Fissures 2 (700 m a.s.l.)** and **3 (550 m a.s.l.)** are located to the northwest and west of the caldera, three and four kilometers downslope of the rim, respectively. These fissures were approximately 250 meters long and produced lava flows that covered 2.3 km² and 0.4 km², respectively, in a single day of eruptive activity.

- **Fissure 4 (100 m a.s.l.),** is located on the northwestern flank at 7–8 km north of the caldera rim, and was the most active and the only vent with ongoing activity during all the eruptive period (58 days). The most recent data indicates that flows covered an area of 13.3 km². Most of the volume of the eruption is concentrated in this flat area. According to Sentinel-1 data collected on 6 July, lava flows reached the ocean and began to change the Ecuadorian coastline. Important hazards associated with this phenomenon were explosions due to rapid evaporation as the hot lavas enter the cold ocean. Until 25 August, the Ecuadorian territory increased about 1.5 km².

- **Finally, fissure 5 (840 m a.s.l.),** is located to the west 1.5 km downslope of the caldera rim [Marco Bagnardi personal comm.]. It is 170 meters long and produced 0.026 km² of lava flows at the beginning of the eruption.

Until 25 August, we estimated that the lava emitted by this eruption covered a total area of 30.6 km², which is equivalent to the area occupied by Cuenca, the third largest city in Ecuador.

Additionally, the seismic record showed five eruptive pulses on 1-2, 7-8, 31 July and 4 and 9 August (Figure 4A). These seismic peaks were associated with the SO₂ emission peaks detected by OMI and OMPS satellite instruments. The highest values were registered at the beginning of the eruption 29×10⁶ kg, and on July 2 and 8, with values of 46×10⁶ kg and 50×10⁶ kg respectively (Figure 4B). Daily values were typically higher than 3×10⁵ kg. These pulses were also recorded by the MODVOLC system, which detected the evolution of thermal anomalies (Figure 4C). The pulses could be related to an increase in the eruptive emission rate of lava flows at fissure 4, which formed a scoria cone (200 meters width).
Figure 5: Map of 2018 Sierra Negra eruption (updated until 19 September 2018). Green dots are the main reference locations, which include cGPS and seismic permanent stations, epicenter of the earthquake M\text{LV} 5.3 and the Amarradero de caballos which is a place where tourists have a panoramic view of the caldera during quiescence periods. Access to this area was forbidden during pre-eruptive and eruptive period. Data provided by: Landsat-8: 27 June, 13 July and Sentinel-2: 6, 11, 16 and 31 July, 5, 10 and 25 August and 9 and 19 September 2018. Inset shows the main islands of the Galápagos Archipelago and the red rectangle the location of Sierra Negra volcano.

### 3 Size of the eruptions

The two principal quantities to define the size of volcanic eruptions are the magnitude (the mass of the material erupted) and the intensity (the mass eruption rate "MER"). According to Pyle 2015, magnitude and intensity are logarithmic-scales to describe and compare historic, pre-historic, and geological events. Since it is very difficult to quantify the intensity fluctuation of pre-historic and geological eruptions of which neither reports nor direct observations are available, intensity is likely to be a time-averaged estimate based on total erupted mass and duration [Pyle 2015].

Table 1 shows the parameters calculated for the size of the eruptions of Fernandina and Sierra Negra since 1979 to present. In 2018 the Fernandina eruption extruded a bulk volume of 7.92±4 million m$^3$ assuming 5±2.5 meters thickness and had a magnitude of 3.2±0.2 and intensity of 8±0.2. On the other hand, the 2017 eruption extruded a volume of 13±6.5 million m$^3$ assuming 2±1 meters thickness based on the similarities of area, fissure location and eruptive behavior of the 2005 Fernandina’s eruption [Chadwick et al. 2011]. Thus, the 2017 Fernandina eruption had a magnitude of 3.4±0.2 and intensity of 8.1±0.2. In 2018 Sierra Negra erupted 189±94 million m$^3$ (bulk volume), assuming 3±1.5 meters thickness on fissures 1, 2, and 3, 1±0.5 meters thickness in fissure 5, and 10±5 meters on fissure 4 based on [Geist et al. 2008; Reynolds et al. 1995]. This would correspond to a magnitude of 4.6±0.2 and intensity of 7.8±0.2.

Finally, taking into account the two phases described for the 2018 Sierra Negra eruption, it was possible to quantify the intensity of each phase. Thus, phase one had an intensity of 9.1±0.2 and phase two of 7.8±0.2. The higher intensity of the first phase is compatible with the higher pulses of energy detected by the seismic stations.

### 4 Conclusions

Fernandina and Sierra Negra are both shield volcanoes fed by magma from the same hot spot, which showed similar eruptive style. Nonetheless, their unrest periods reported here were quite different.

It seems that Sierra Negra usually shows a clear and long-lasting unrest period characterized by an increase in the number and magnitude of volcanic earthquakes and is also accompanied by significant ground deformation (up to few meters) of the caldera floor (i.e., 2005 and 2018 eruptions). Moreover, the M\text{LV} 5.3 earthquake was a clear precursor of the eruption, which was also the case in 2005 [Geist et al. 2008]. On the other hand, Fernandina volcano displayed short-lasting seismic unrest (few hours) and less impressive ground deformation (few centimeters) in the last two eruptions (2017 and 2018). Also, vents and fissures open up on different flanks of the volcano, which is more difficult to anticipate for the hazard assessment.

These eruptions show that long-lasting unrest periods could be related with long-lasting eruptive periods (58 days in the case of Sierra Negra volcano). On the other hand, short-lasting unrest in Fernandina volcano produced short eruptive periods (two days). This observation has a direct correlation with the volume of magma erupted, the intensity of the eruption and the risk assessment. Nonetheless, this observation requires more detailed studies, specially given that Wolf volcano (Galápagos) showed a short unrest period and long-lasting eruption [Bernard et al. 2015].
The recent eruptions of Fernandina and Sierra Negra

Table 1 – Comparison of the size of the eruptions of Fernandina and Sierra Negra volcanoes. The parameters were taken from Bourquin et al. 2009 (2009-eruption), Chadwick et al. 2011 (2005), Rowland et al. 2003 (1995), Chadwick et al. 1991 (1988) in the case of Fernandina and from Geist et al. 2008 (2005-eruption) and Reynolds et al. 1995 (1979) in the case of Sierra Negra. Moreover, based on past literature an average thickness (±50% error), a rock density of 3000 kg m$^{-3}$ and deposit density (25% void) were assumed.

| Parameter                  | Fernandina | Sierra Negra |
|----------------------------|------------|--------------|
| Eruption date              | 2018       | 2017         |
| Area (10^6 m$^2$)          | 1.58       | 6.5          |
| Average thickness (m)      | 5±2.5      | 2±1          |
| Volume (10^6 m$^3$)        | 7.92±4     | 13±6.5       |
| DRE (10^6 m$^3$)           | 5.9±3      | 9.7±4.9      |
| Erupted Mass (10^10 kg)    | 1.78±0.9   | 2.9±1.5      |
| Duration (days)            | 2          | 2.5          |
| MER (m$^3$ s$^{-1}$)       | 34.4±17    | 45.1±22.6    |
| Magnitude                  | 3.2±0.2    | 3.4±0.2      |
| Intensity                  | 8±0.2      | 8.1±0.2      |

The areas covered by lava flows during 2018 Sierra Negra eruption was of approximately 30.6 km$^2$ in 58 days while in Fernandina eruptions the covered areas were of 1.58 km$^2$ (2018) and 6.5 km$^2$ (2017) in two days each.

Galápagos Archipelago is a fragile ecosystem with unique native flora and fauna. Both can be highly threatened by lava flows and wildfires as occurred during the Fernandina eruption in 2017.

Satellite images provided critical information for hazard assessment and unrest and eruption monitoring in such remote areas. Freely available satellite data complement the monitoring networks of volcanic observatories and help to improve the communication with the stakeholders.

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Author contributions

Francisco Vasconez, Patricio Ramón, Silvana Hidalgo and Benjamin Bernard used satellite images to monitor the evolution of the eruption, which include thermal anomalies, lava flows displacement, ash columns, degassing patterns, ground deformation and secondary hazards. Stephen Hernandez, Mario Ruiz and Alexandra Alvarado analyzed and interpreted the seismic data of the unrest and eruptive period of these eruptions. Peter La Femina and Gorki Ruiz collected and analyzed the cGPS data.

Francisco Vasconez, Silvana Hidalgo, Benjamin Bernard, Stephen Hernandez, Peter La Femina, Patricio Ramón, Mario Ruiz and Alexandra Alvarado led the writing of the manuscript, with contributions from all authors.

Data availability

Thermal anomalies (http://modis.higp.hawaii.edu/, http://www.mirovaweb.it/), multispectral satellite images (Sentinel-2 and Landsat-8 https://www.sentinel-hub.com/, https://eos.com/landviewer/, https://earthexplorer.usgs.gov/), radar images (Sentinel-1, https://vertex.daac.asf.alaska.edu/), degassing data (https://so2.gsfc.nasa.gov/), and InSAR ground-deformation (https://insarmaps.miami.edu/) were provided by several free-online platforms. Seismic and cGPS data recorded by IG-EPN and Penn
University and UNAVCO can be requested for further analysis and future work at https://www.igepn.edu.ec/ and http://www.unavco.org/ respectively.

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