Article

Effect of Climate Variability on Green and Blue Water Resources in a Temperate Monsoon Watershed, Northeastern China

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Abstract: Over 80% of global grain production relies on green water, water from precipitation that is stored in unsaturated soil and supports plant growth. Blue water, precipitation that turns into surface water and groundwater, is also a vital surface water resource, and it can be directly utilized. The Tanghe River Basin is a typical temperate continental monsoon watershed in Northern China where residents and crops rely on blue and green water resources. In this study, the spatiotemporal distributions of water resources in the Tanghe River Basin were simulated using the soil and water assessment tool (SWAT) model for the period between 1970 and 2015. The results demonstrate that the Nash–Sutcliffe efficiency and coefficient of determination were both higher than 0.64 during the calibration and validation periods at all hydrological stations, indicating high simulation accuracy. The average annual water resources of the Tanghe River Basin are 759.37 mm. Green and blue water account for 68% and 32% of the total water resources, respectively. The study period was divided into the reference period (1970–1976) and the variation period (1977–2015), to explore the impact of climate change on the green and blue water resources of the Tanghe River Basin water resources. Compared with the reference period, the average green and blue water resources in the variation period decreased by 78.48 and 35.94 mm/year, and their rate changes were −13.45% and −13.17%, respectively. The water resource relative change rates were high in the south and low in the north, and they were predominantly affected by precipitation. This study improves our understanding of the hydrological processes as well as the availability of blue and green water in the study region, and can prove beneficial in promoting the sustainable development of small basins and the integrated watershed management in areas with similar climatic conditions.

Keywords: SWAT model; blue water resources; green water resources; Tanghe River Basin

1. Introduction

Water is vital to people and ecosystems. As the global economy and population have grown, human demand for water resources has also increased, affecting the health of the global ecosystem [1–4]. Due to this ongoing competition between humans and the environment for water resources, the availability of water resources must be comprehensively assessed to deepen our understanding of renewable water sources and realize sustainable, effective, and fair water resource management [5,6]. In the current literature, considerable attention has been paid to blue water resources, which are generated by surface and underground runoff. However, green water resources have been gaining attention in recent years. The concept of green water was first proposed by Falkenmark in 1995 [7]. Green water resources refer to the water from precipitation that is stored in unsaturated soil and supports plant growth. Green water can be classified into two components: green water
flow, which refers to actual evapotranspiration, and green water storage, which refers
to water stored in the soil [8]. It plays a critical role in preserving the production and
service functions of land ecosystems. Green water resources account for the vast majority
of water resource utilization; over 80% of the water used by global food production and
forest grasslands is green water [9,10]. Climate change greatly impacts the hydrological
cycle and green water resources in basins [11], and it is primarily reflected by changes
in rainfall, temperature, and other meteorological elements. Therefore, determining the
relationships between the ecological and hydrological processes of green and blue water
resources by studying their spatiotemporal changes, mitigating the limitations of available
water resources [12], and evaluating water resources in a comprehensive scientific manner
is significant for water resource planning and management.

There are currently three primary evaluation method types for green water resources:
hydrological, biological, and bio-hydrological approaches [13]. The hydrological model
established by the hydrology method evaluates the spatial distribution changes of green
and blue water resources in a basin; therefore, it has been widely utilized by scholars
worldwide. The predominant global hydrological models include the hydrological land-
use change, agricultural catchment research unit, and soil and water assessment tool
(SWAT) models.

Modeling can be used to represent natural conditions and as a reference for addi-
tional. Models can also inform researchers and policymakers about the original state of a
river basin for decision-making. The SWAT model is particularly utilized due to its high
efficiency, high precision, long simulation cycles, and ease of operation for simulating
watershed hydrology [14]. Thus, we selected this model for our study. Studies on blue
and green water resources using the SWAT have been extensively conducted by various
authors worldwide, including in Africa [15–18], Europe [19], and Asia [20–22]. These study
types can provide a scientific basis for the planning and management of water resources.

The objective of this study was to establish the natural hydrological cycle process
and quantify the spatiotemporal dynamics of green and blue water under the natural
conditions of the Tanghe River Basin using the SWAT model. The Tanghe River Basin
exhibits a temperate continental monsoon climate and is an important grain-producing
area in Liaoning Province in Northern China. The Tanghe Reservoir in the basin is an
essential source of drinking water for Liaoyang and Anshan. Many studies have been
conducted on the water quality and surface water resources of the Tanghe River Basin;
however, few studies have focused on the green water resources, which account for the
greatest proportion of farmland water consumption. Therefore, it is necessary to study
blue and green water in the Tanghe River basin to provide a scientific basis for the rational
allocation and management of water resources in the basin.

2. Materials and Methods

2.1. Study Area

The Tanghe River is the largest tributary of the Taizi River Basin, and it is located in
Liaoning Province, China at 40°43′–41°16′ N, 123°03′–123°31′ E. Upstream of the river
basin is divided into east and west tributaries, which flow into the Tanghe Reservoir
through Anping and Gengjiatun and then into the Taizi River at Xishuangmiao Village
in Xiaotunzhen. The river is 90.7 km in length, and the river basin covers an area of
approximately 1466 km² (Figure 1). The Tanghe River Basin is located in a temperate
monsoon climate zone, with two distinct seasons: hot and rainy in the summer, and cold
and dry in the winter. The annual temperature ranges from −17 to 29 °C, and the annual
precipitation is approximately 710 mm. The basin terrain is low in the north and high in
the south; the upstream is primarily mountains and hills, whereas the downstream has
more plains. The land-use type is predominantly forest land, and downstream contains
more farmland and construction land.
2.2. Data

The data required to construct the SWAT model primarily included spatial and hydrometeorological data. The spatial data included land-use, soil, and digital elevation model (DEM) data. Of these data types, the land-use data (Figure 2) were obtained from the 2016 Landsat OLI remote sensing image set, which was downloaded from the interpretation of geospatial data cloud. The soil data (1:100,000) were obtained from the Harmonized World Soil Database [23]. The DEM data exhibited a spatial resolution of 30 m and were derived from the SRTM1v3.0 data of the U.S. Geological Survey website in February 2020 (http://glovis.usgs.gov/ (accessed on 17 February 2021)). Hydrological data were obtained and verified using daily inflow data from two hydrological stations in the Tanghe River Basin—Erdaohezi and Haojiadian—from 1970 to 2015. The meteorological data included the daily rainfall, temperature, solar radiation, wind speed, and relative humidity from 1970 to 2015, and they were derived from the Liaoyang meteorological station and national meteorological science data sharing service platform (http://data.cma.cn/en (accessed on 17 February 2021)).

2.3. SWAT Model Data Preparation

The SWAT model is a widely used model that can simulate hydrological processes [24], including runoff [25], sedimentation [26], and water pollution [27], in large and complex watersheds over long periods. Hydrological cycle SWAT model simulations can be divided into land and routing phases. The land phase of the hydrological cycle simulates the water yield in each hydrological response unit (HRU), and the routing phase models the movement of water through the channel network and other water bodies to the outlet [28].
First, the SWAT divides the watershed into sub-basins according to topography and river course. Then, it divides the sub-basin into multiple HRUs according to the soil and land-use types. Finally, it calculates the runoff of each HRU to obtain the total meridional flow [29,30]. The hydrologic cycle process of the basin is simulated according to the water balance principle. The corresponding equation is as follows [31,32]:

\[ SW_t = SW_0 + \sum_{i=1}^{t} \left( R_{day} + Q_{surf} - E_a - w_{seep} - Q_{gw} \right) \]  

(1)

where \( SW_t \) is the final soil water content (mm), i.e., the amount of water in the soil profile at the end of the period; \( SW_0 \) is the initial soil water content on day \( i \); \( R_{day} \) is the amount of precipitation on day \( i \); \( Q_{surf} \) is the amount of surface runoff on day \( i \); \( E_a \) is the amount of evapotranspiration on day \( i \); \( w_{seep} \) is the amount of seepage and bypass water from the soil profile on day \( i \); and \( Q_{gw} \) is the amount of return flow on day \( i \).

Based on the DEM, land-use data, and soil data, the spatial database of the SWAT model was established. The Tanghe River Basin was divided into 44 sub-basins and 655 HRUs. Monthly runoff data from the Erdaohezi and Haojiadian hydrological stations from 1970 to 1990 were used to calibrate the model. Monthly runoff data from 1991 to 2015 were used to verify the model, and the warm-up period was set to two years in order to ensure the establishment of basic flow conditions for the simulations [33].

Figure 2. Tanghe River Basin land-use map.

2.4. SWAT Model Evaluation Index

In this study, the sequential uncertainty analysis method (SUFI-2) provided by the SWAT-CUP software was used to calibrate and verify the established SWAT model, and the SWAT performance was evaluated according to the statistical relationship between the observations and simulation results. The coefficient of determination \( (R^2) \) and Nash–Sutcliffe efficiency \( (E_{NS}) \) are the most commonly used statistical indicators for calibration.
and verification. $R^2$ represents the degree of coincidence between the simulated and measured values, and the value range is 0 to 1. $E_{NS}$ is used to judge the simulation effect of the hydrological model, and its value range is $-\infty$ to 1. Model results are considered credible when $R^2 \geq 0.6$ and $E_{NS} \geq 0.5$ [34–37]. The formulas for $R^2$ and $E_{NS}$ are expressed as follows:

$$R^2 = \frac{\sum_{i=1}^{n} (Q_{m,i} - \overline{Q}_m)(Q_{s,i} - \overline{Q}_s)^2}{\sum_{i=1}^{n} (Q_{m,i} - \overline{Q}_m)^2 \sum_{i=1}^{n} (Q_{s,i} - \overline{Q}_s)^2}$$ (2)

$$E_{NS} = 1 - \frac{\sum_{i=1}^{n} (Q_m - Q_s)^2}{\sum_{i=1}^{n} (Q_{m,i} - \overline{Q}_m)^2}$$ (3)

where $Q_m$ is the measured value over many years, $Q_s$ is the simulation value, $Q_m$ is the measured average value over many years, and $Q_s$ is the simulated average value.

### 2.5. Water Resource Statistics

The SWAT model can directly estimate the blue and green water components and quantify their available amounts. The water components include the subwatershed yield ($WYLD$), deep aquifer supply ($DA_RCHG$), actual evapotranspiration ($ET$), and soil moisture content ($SW$). Blue water is the sum of the subwatershed yield and deep aquifer recharge. According to the concept of green water and output results of the SWAT model, green water flow is the actual evapotranspiration ($ET$), and green water storage is soil water ($SW$). Thus, the total amount of green water resources is the sum of the actual evapotranspiration and soil moisture, i.e., the sum of the green water flow and reservoir [11]. The green water coefficient is the proportion of green water resources of the total water resources of a certain area, which can be used to evaluate the distributions of blue and green water resources in the basin [5,38,39]. The formulas for blue water, green water, and the green water coefficient are expressed by Equations (4)–(6), respectively.

$$B = WYLD + DA_RCHG$$ (4)

$$G = ET + SW$$ (5)

$$I_{GWC} = \frac{G}{G + B}$$ (6)

where $I_{GWC}$ is the green water coefficient, $B$ is the amount of blue water resources, and $G$ is the amount of green water resources.

In this study, the relative rate of change ($RCR$) was used to represent the changes in the green and blue water flows during different periods. The $RCR$ formula is expressed as follows:

$$RCR = \frac{V_i - V_0}{V_0} \times 100\%$$ (7)

where $V$ represents the green or blue water flow, $i$ represents the later period, and 0 represents the initial period.

### 2.6. Abrupt Change Point Detection

The Mann–Kendall test, which is widely used in hydro-meteorological time series, is a nonparametric statistical test method [40,41]. Given that $x_1, x_2 \ldots, x_n$ is a series that denotes the studied time series, $n$ is the sequence length, and the order column $S_k$ is constructed using the following equations:

$$S_k = \sum_{i=1}^{k} \sum_{j=i+1}^{n} \alpha_{ij} (k = 2, 3 \ldots n)$$ (8)

$$\alpha_{ij} = \begin{cases} 
1 & x_i > x_j \\
0 & x_i \leq x_j
\end{cases} \quad 1 \leq j \leq I$$ (9)
Statistical variables are defined as

\[ UF_K = \frac{|S_k - E(S_k)|}{\sqrt{Var(S_k)}} (K = 1, 2, \ldots, n) \]  

(10)

in which

\[ E(S_k) = \frac{k(k + 1)}{4} \]  

(11)

\[ Var(S_k) = \frac{k(k - 1)(2k + 5)}{72} \]  

(12)

The data sequence \( x \) is arranged in reverse order. Then, according to Equation (10),

\[
\begin{cases}
UB_k = -UF_k' \\
k' = n + 1 - k 
\end{cases} (k = 1, 2, \ldots, n)
\]  

(13)

where \( UF_k \) and \( UB_k \) are the forward and backward sequences, respectively. If \( UF_k \) is positive, the sequence is increasing, and when it is negative, it is decreasing [42]. For a given significance level, \(|UF_k| \) or \(|UB_k| > U_{a/2}\) indicates that the sequence exhibits a significant change in trend; otherwise, it is not significant. The significance level in this study was 0.05, and the critical values were \( U_{0.05} = \pm 1.96 \). The intersection of the \( UF_k \) and \( UB_k \) curves is the abrupt change point.

3. Results

3.1. SWAT Model Construction, Calibration, and Validation

Parameter sensitivity analyses were conducted using the SWAT-CUP software [43,44], and 10 highly sensitive parameters were selected (Table 1). Parameter selection and initial ranges were based on a literature review and prior knowledge of similar watersheds [45,46]. Then, the parameters were automatically calibrated for streamflow using SWAT-CUP and 1000 simulations. The results are displayed in Table 1 and Figure 3.

The \( R^2 \) of the Erdaohezi hydrological station was 0.79 for the calibration period and 0.72 for the validation period, and the \( E_{NS} \) was 0.79 and 0.72, respectively. The \( R^2 \) of the Haojiadian hydrological station was 0.81 for the calibration period and 0.74 for the validation period, and the \( E_{NS} \) was 0.74 and 0.64, respectively. The SWAT model performed well with high accuracy [36,47], indicating that it can credibly describe the Tanghe River Basin hydrological cycle.

3.2. Variance Tendency of Green and Blue Water Resources from 1970 to 2015

As illustrated in Figure 4, the average annual water resources of the Tanghe River Basin were 759.37 mm, with an average annual green water resource of 516.83 mm, which included 176.50 mm of green water flow and 340.33 mm of green water storage. The blue water resource constituted 242.55 mm, and the average annual green water coefficient was 68%. These results indicated that the water resources in the Tanghe River Basin were primarily green water resources, which were 2.13 times the amount of blue water resources.

According to the unitary regression analysis results, the precipitation, blue water, and green water, including green water flow and water storage, in the Tanghe River Basin did not significantly change from 1970 to 2015. In these 46 years, the precipitation, blue water, and green water change rates were \(-11.48\), \(-7.91\), and \(-14.96\) mm/decade, respectively. In addition, the change rates of the green water flow and storage were 0.07 and \(-15.03\) mm/decade, respectively.
Figure 3. Observed and simulated monthly runoff rates at (a) Erdaohezi and (b) Haojiadian hydrological station in Tanghe River Basin during calibration and validation periods.

Table 1. SWAT calibration parameters.

| Parameter Name and Document Type | Description                                                       | Initial Range | Calibrated Value |
|----------------------------------|-------------------------------------------------------------------|---------------|-----------------|
| v_SLSUBBSN.hru                   | Average slope length (m)                                          | −20% to +20%  | Increased to 11.971% |
| v_CANMX.hru                      | Maximum canopy storage                                            | 0–100         | 18.780          |
| r_CN2.mgt                        | SCS curve number for moisture condition II                        | −20% to +20%  | Decreased to 0.027% |
| r_SOL_K().sol                    | Saturated hydraulic conductivity (mm h⁻¹)                         | −20% to +20%  | Increased to 0.173% for the first layer |
| v_ESCO.hru                       | Soil evaporation compensation factor                             | 0–1           | 0.952           |
| v_EPCO.hru                       | Plant uptake compensation factor                                  | 0–1           | 0.341           |
| v_CH_N2.rte                      | Manning “n” value for the primary channel                        | 0.01–0.3      | 0.146           |
| v_RCHRG_DP.gw                    | Deep aquifer percolation fraction                                 | 0–1           | 0.132           |
| v_GWQMN.gw                       | Threshold depth of water in the shallow aquifer required for return flow to occur (mm H₂O) | 0–5000 | 1.669 |
| r_SOL_AWC().sol                  | Available soil water capacity (mm H₂O mm⁻¹ soil)                 | −20% to +20%  | Increased to 0.084% for the first layer |

a v indicates parameter changes by the specified value; r displays parameter changes by the original value, 1+ specified value; hru indicates HRU conventional data entry; mgt represents HRU management; sol represents soil data entry; rte indicates river hydrology; and gw represents groundwater hydrology [46]. b Source: Neitsch et al. [48]. c Ranges used for parameter sensitivity analysis to maximize E₉ [49].
Table 2. Average annual precipitation and green and blue water resource statistics in reference and variation periods

|                        | Precipitation | Green Water | Blue Water |
|------------------------|---------------|-------------|------------|
| Reference Period (1970–1976) (mm/year) | 763.71        | 583.36      | 273.02     |
| Variation Period (1977–2015) (mm/year) | 699.70        | 504.89      | 237.08     |
| Value of Change (mm)    | −64.01        | −78.48      | −35.94     |
| Rate of Change (%)      | −8.38         | −13.45      | −13.17     |

Figure 4. Amounts of blue and green water resources in Tanghe River Basin.

3.3. Effect of Climate Variability on Green and Blue Water Resources in Tanghe River Basin

The Mann–Kendall test was used to explore changes in precipitation and green and blue water resources. As illustrated in Figure 5, the UF and UB of the precipitation (Figure 5a), blue water resources (Figure 5b), and green water resources (Figure 5c) all intersected in 1976 and were all within a significance level of \( p < 0.05 \). The results demonstrated that 1976 is the transition point of the hydrological series of the Tanghe River, and the change in precipitation decisively influenced the green and blue water resources.

Figure 5. Mann–Kendall test results for (a) precipitation, (b) green water resources, and (c) blue water resources.

Considering the abrupt changes in precipitation, the entire period was divided into two parts, a reference period (1970–1976) and variation period (1977–2015), to explore the impact of climate change on the green and blue water resources of the Tanghe River Basin. The period results are displayed in Table 2.
After the abrupt change in 1976, the average annual precipitation decreased by 64.01 mm/year, and the rate of change was −8.38%. As a result, the average annual green water resources decreased by 78.48 mm, with a change rate of −13.17%. The average annual blue water resources decreased by 35.94 mm, with a change rate of −13.45%. The sum of the blue and green water was less than the precipitation amount. This result occurred because green water storage primarily refers to water that is stored in the soil, which includes groundwater recharge. The change rates of green and blue water resources were much higher than those of precipitation, indicating that the water system was particularly dependent on the climate and similar to those of climate changes, but it was more sensitive than the climate system.

From the spatial perspective, the precipitation RCR was high in the south and low in the north, and the RCRs were negative throughout the basin. The highest sub-basin RCR was upstream of the west tributary of the Tanghe River Basin, which was as high as −7% to −10% (Figure 6a). The spatial distribution of the RCR of green water resources (Figure 6b) was similar to that of precipitation. In all the sub-basins, green water resources decreased by −26% to 0%. The sub-basins with the greatest RCR range (−26% to −20%) were located in the eastern Tanghe River Basin, while the sub-basins with the smallest RCR range (−3% to −2%) were located in the east and upstream of the Tanghe River Basin. Compared with the green water resource RCRs, those of the blue water resources both increased and decreased (Figure 6c). In most of the sub-basins, the blue water resource RCRs decreased by −33% to 0%. The sub-basins with the largest RCR range (−33% to −17%) were located in the west branch of the Tanghe River Basin. In a few sub-basins, blue water resources increased by 0% to 37% in the upstream of the west branch of the Tanghe River Basin.

![Spatial distributions of the precipitation, blue water, and green water RCRs in the Tanghe River Basin.](image_url)

Figure 6. Spatial distributions of the (a) precipitation, (b) blue water, and (c) green water RCRs in the Tanghe River Basin.

Regarding the spatial distributions, the differences between precipitation and the green and blue water demonstrated that the variations in green and blue water resources were predominantly affected by rainfall; they may also be affected by other factors, such as human activities and the underlying surface characteristics. Thus, the reasons for changes in green and blue water resources are complex.

4. Discussion

Green water resources accounted for 68% of the water resources in the Tanghe River Basin. this result is similar to that of our previous research where 69% of water resources were green water in the Xihe River Basin [39], which is also located in the temperate monsoon climate zone of Northeastern China. The runoff yield mechanism under this climate
is primarily excess infiltration; thus, more precipitation was converted to green water. The same results have been found for other basins, such as the Wuyuer River Basin [50], where the average annual green water coefficient was 90%. Under different climatic conditions such as abundant precipitation, the runoff yield mechanism is predominantly excess storage. Blue water resources in basins such as the Dongjiang River Basin [38] and Xixi Basin [51] are substantially greater than the green water resources; the average annual green water coefficients of these basins were 39% and 36%, respectively. In combination with other research results presented in Table 3, we postulate that the precipitation magnitude and underlying surface properties determine the runoff yield mechanism and subsequently the allocation of blue and green water resources.

Table 3. Distributions of green water resources in various river basins

| River Basin            | Climate Type                        | Precipitation (mm) | Green Water Coefficient (%) |
|------------------------|-------------------------------------|--------------------|----------------------------|
| Wuyuer River Basin     | Temperate monsoon climate           | 496.7              | 90                         |
| Huangshui Basin        | Plateau continental climate         | 400–600            | 79                         |
| Xihe River Basin       | Temperate monsoon climate           | 700–900            | 69                         |
| Tanghe River Basin     | Temperate monsoon climate           | 763.71             | 68                         |

From the spatial distribution perspective, precipitation is the primary factor affecting the spatial distribution of blue water. However, it has no apparent effect on the spatial distribution of green water. Underlying surface properties and human activities (e.g., the conversion of farmland to forest or grassland, construction of reservoirs, agricultural irrigation, etc.) can affect the spatial distribution pattern of green water. This study contributes to our understanding of hydrological processes and the availability of blue and green water in the Tanghe River Basin. This knowledge is essential for promoting the sustainable development of small basins and the integrated management of watersheds under similar climatic conditions. The impacts of human activities such as agricultural irrigation and reservoir construction on blue and green water were not considered in the simulation of blue and green water in this study. Additionally, the conversion mechanism of blue and green water that can relieve the pressure on water resources and the basin environment remains unclear and requires further study in the future.

5. Conclusions

This study investigated the long-term spatiotemporal changes in the Tanghe River Basin from 1970 to 2015 by establishing a SWAT model, simulating the hydrological cycle processes within the basin, and exploring the spatiotemporal changes in green and blue water resources in the temperate monsoon climate zone of North China. The primary conclusions are as follows.

1. At the Erdaohezi hydrologic station, the $E_{NS}$ was 0.79 and the $R^2$ was 0.79 in the calibration period and 0.72 and 0.72 in the validation period, respectively. At the Haojiadian hydrologic station, the $E_{NS}$ was 0.74 and the $R^2$ was 0.81 in the calibration period and 0.64 and 0.74 in the validation period, respectively. The model simulations of the two hydrological stations were extremely precise, indicating that the model can accurately describe basic watershed hydrologic cycle processes.

2. From 1970 to 2015, the average amount of annual water resources in the Tanghe River Basin was 759.37 mm, that of green water resources was 516.83 mm, and that of blue water resources was 242.55 mm. The average annual green water coefficient was 68%, and the average amount of annual green water resources was 2.13 times that of the blue water resources. The amount of green water storage, green water resources, and blue water resources all demonstrated an insignificant downward trend.

3. From 1970 to 2015, the precipitation, green water resources, and blue water resources in the Tanghe Basin all experienced an abrupt change in 1976. Compared with that of the reference period, the average annual precipitation and green and blue water
resources in the variation period decreased by 64.01, 78.48, and 35.94 mm/year, and their change rates were −8.38%, −13.45%, and −13.17%, respectively.

(4) The spatial distributions of green and blue water resources and the precipitation RCR were similar, with high values in the south and low values in the north. Additionally, the RCRs of the green and blue water were all negative except for those in a small part upstream of the Tanghe Basin. These results indicate that precipitation greatly impacts green and blue water resources, and the influence of climate change in the region was significant.

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