Introduction

In Morocco, the continual decrease in water resources is the current major problem, since water is a determining factor for sustainable development and improving the quality of life. The quantitative and qualitative management of groundwater and surface water resources is currently a necessity, given the continuous increase in water demand, on the one hand, and the degradation of water quality observed during the past few years on the other hand. Water resources are limited and drought affects most of the country and becomes seriously threatening.

In the south-east of Morocco, or more precisely in the province of Errachidia, the scarcity and randomness of precipitation, the increasingly worrying desertification, the overexploitation of groundwater, linked to the strong demographic pressure and agriculture and animal husbandry activities have a great influence on the piezometric levels of groundwater (Amharref, 1991). The salt water from the source of Ain Al Atti has a negative impact on the environment by inducing direct chemical contamination on the aquifers in place. This therefore requires urgent intervention to curb the polluting flow from this source (Ammary, 2007). The problem addressed by this study concerns the potential groundwater in the locality of Douira north of the Ain Al Atti source. This article helps assess this potential by identifying fracturing zones. Faults represent a special target for hydrogeological exploration and can play an important role in the supply of groundwater (Reinhard, 2006). To achieve the objective of this article to highlight the litho-structural parameters, the working procedure will be based on electrical soundings, in order to determine the vertical variation of the resistivity obtained at the level of Douira and at the level of Ain Al...
Atti. On the other hand, it will be based on the Very Low Frequency Electromagnetic technique (VLF-EM) in order to locate the alignment of the fractured areas.

**Geography and Geological Setting**

The study area is located about 50 km south of Errachidia and 20 km north of Erfoud Fig. 1. It is a part of cretaceous basin of Errachidia (South-eastern Morocco). This basin is characterized by a stratigraphic series that ranges from paleozoic to quaternary. In the outcrop Fig. 2, the basin is generally composed of carbonate deposits of Turonian (Choubert and Faure-Muret, 1960), sandstone and sand with gypsum intercalation attributed to Infracenomanian and sand clay with gypsum and anhydrite of Senonian (Choubert, 1948). Locally, from drilling data, the Infracenomanian deposits overlie the Paleozoic (angular discordance). The quaternary is presented by alluvium and conglomerates. It shows varying thickness layers between 5 and 40 m (Amharref, 1991).

**Lithostratigraphy**

The stratigraphic log in Fig. 3 shows the formations constituting the basin range from the Paleozoic to the Quaternary.

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**Fig. 1:** Geographic situation of the study area

**Fig. 2:** Geological section of the Errachidia basin by (Margat, 1977)
**Fig. 3:** Synthetic stratigraphic log of the Cretaceous basin of Errachidia (Amharref, 1991)

| Age          | Lith                    | Description                                      | Thickness (m) |
|--------------|-------------------------|--------------------------------------------------|---------------|
| Quaternary   | Lake limestones, silts and conglomerates            |                                                  | 5-40          |
| Cretaceous   | Red sandstone sands with intercalation of marl and gypsum |                                                  |               |
| Senonien     | Clay sands                                                          |                                                  | 150-600       |
|              | Lumachellic massive limestones                            |                                                  |               |
|              | Bed limestone with marl intercalation                      |                                                  | 60-150        |
|              | Green manes anhydrite                                       |                                                  | 40-50         |
|              | Red sandstone                                               |                                                  |               |
|              | Sandy clays                                                 |                                                  |               |
|              | Anhydrites                                                  |                                                  |               |
|              | Fine sands                                                  |                                                  | 80-400        |
| Jurassic     | Coarse sandstone                                           |                                                  | 150-280       |
|              | Red clays                                                   |                                                  |               |
|              | Conglomerate                                               |                                                  |               |
| Bajocien     | Dolimitic limstones                                         |                                                  | 100-320       |
|              | Alternation of limstones and marls                          |                                                  |               |
| Aalenien     | Alternation of marls and dolomitic limestones               |                                                  | 100-320       |
| Toarcien     | Fossiliferous marls                                         |                                                  | 50-250        |
| Domerien     | More or less marly dolomitic limestone                      |                                                  | 150-200       |
| Trias        | Red marls                                                  |                                                  | 500           |
|              | Basal doleritic                                             |                                                  |               |
|              | Anhydritic red marls                                        |                                                  |               |
| Paléozoïc    | Schist, quartzite and limestone                            |                                                  |               |
Materials and Methods

Electrical Survey

The electrical sounding consists of the analysis of the apparent resistivity of the subsoil, measured by a symmetrical quadrupole device (AMNB) for a succession of spacing of the electrodes AB. The investigation depth is adjusted by varying the distance between the current injection electrodes A and B. The difference in the potential measured between M and N reflects an apparent resistivity linked to the true resistivity and to the thicknesses of all the layers interested in the AMNB device. This method is widely used and generally provides a good initial approach (Fig. 4). The apparent resistivity \( \rho_a \) is calculated from the current I and the potential difference \( \Delta V \) (Equation 1). The coefficient \( K \) is called geometric factor. For Schlemberger configuration, the factor \( K \) can be calculated from the electrode spacing by (Equation 2):

\[ \rho_a = K \frac{\Delta V_{MN}}{I_{AB}} \]  

(1)

\[ K = 2\pi a \]  

(2)

VLF Electromagnetic Survey

The Very Low Frequency Electromagnetic (VLF-EM) is based on the use of radio waves in the range of 15 to 30 kHz. The signal primary magnetic field \( (H_p) \) emitted by the VLF stations, can be captured in the field by the VLF instruments. When a conductor (e.g., a fracture zone) is crossed by the \( H_p \) electromagnetic field, an induced current (Current of Foucault) flows through it and produces a secondary magnetic field \( (H_s) \) out-of-phase with primary magnetic field oriented in any direction (McNeill and Labson, 1991). In this case the conductive body acts as a second source (Kaya et al., 2007). The resulting field from the sum of primary magnetic field and secondary magnetic field is elliptically polarized. This ellipse of polarisation has two components with the same frequency, but different amplitude and phase, (Eze et al., 2004). The in-phase \( H_p \) is the Real component \( (R_p) \) proportional to the tilt \( \tau \) (inclination of the major axis of the ellipse), while the out of phase \( H_p \) is the Imaginary component \( (I_m) \) proportional to the ellipticity \( \varepsilon \) (the ratio between the minor and the major axis b/a). These two components \( \tau \) and \( \varepsilon \) are described by the Equations (3 and 4) below (Saydam, 1981):

\[ \tau = R_p / H_p \]  

(3)

\[ \varepsilon = I_m / H_p \]  

(4)

During our study, the survey was carried out using the Receiver T-VLF Iris Instruments, operating in tilt angle mode, in order to measure the parameters of the ellipse of polarization, which are the tilt \( \tau \) and the ellipticity \( \varepsilon \) Fig. 5. In this mode, it is convenient to operate with a transmitter (VLF station) which is located in the supposed strike (±45°) of the prospected target for a maximum coupling. For detecting the supposed fractures in the study area, the GBR station located in Rugby (England) has been chosen, with a power of 750 KW, which emits a signal with a frequency of 16 KHz. On the fieldwork, 6 VLF EM profiles were conducted, with profile length reaches 1600 m. Readings were taken respecting a spacing of 20 m. The profiles lines were oriented NNW-SSE and N-S directions.

![Fig. 4: Representation of the schlumberger type device](image-url)
For our measurements, in one hand, the Karous and Hjelt (KH) filter has been applied to the real component. This filter permits the draw of apparent current density cross-sections, which show the response of the conductor in depth (Karous and Hjelt, 1983). Qualitatively, it is possible to discriminate between conductive and resistive structures using apparent current density cross-section (Karous and Hjelt, 1977), where a high positive value corresponds to conductive structure and low negative values are related to resistive one (Benson et al., 1997; Sharma and Baranwal, 2005). In the other hand, the Fraser filter has been applied to the real component too and presented in the form of a contour map. Therefore, the filtered real component will always show a positive peak above an anomalous zone (Fraser, 1969). In order to perform Karous-Hjelt and Fraser filtering on VLF EM data, the software KHFFILT is used in the interpretation of the measurements along VLF traverses.

Results

In order to obtain better recognition of the study site, we used two geophysical techniques; the method of electrical and electromagnetic exploration. In fact, two Schlumberger-type electrical soundings were carried out (AB = 600 m long) and six electromagnetic profiles (Fig. 6).

Electrical Prospecting

The electrical survey N°1 is carried out next to the village of Douira on the left bank of the national road N°13 (about 100 m) which leads to Erfoud. Its direction is N-S (Fig. 7a). While the survey N°2 was carried out at 250 m from the source of Ain El Atti in direction NW-SE Fig. 7b.

Electromagnetic Prospecting

We focus on the analysis of positive Karous-Hjelt anomalies, for semi-quantitative interpretation and visualization of the target. For the data filtered by KH, the cross-sectional curves of the apparent current density were produced (Figs. 9 to 11). Here, the traces KH are represented for three types of lines: 2, 3 and 4.

Discussion

The survey in Fig. 6 suggests the presence of a conglomerate formation with gravel on the surface, of resistivity \( \rho = 2000 \text{ Ohm.m} \) and of thickness 1 m. The latter rests on a formation of weathered sandstone with clays, of resistivity \( \rho = 300 \text{ Ohm.m} \) and thickness 7 m, which overcomes a formation of limestone’s of resistivity \( \rho = 3060 \text{ Ohm.m} \) and thickness 64 m. In depth, we find sandy sandstones with clays of resistivity \( \rho = 750 \text{ Ohm.m} \) (Fig. 8a). For the survey in Fig. 7b highlights the sands and gravels in surface (resistivity \( \rho = 300 \text{ Ohm.m} \) and thickness 2.6 m) overcome a clay-marly formation with resistivity \( \rho = 9 \text{ Ohm.m} \) and thickness 16 m, which rests on the sandstones (\( \rho = 270 \text{ Ohm.m} \) and thickness 16 m). At the base, we find a marl formation of resistivity 8 Ohm-m (Fig. 8b).
Fig. 6: Location of electrical and electromagnetic surveys

Fig. 7: Electrical survey: a- Douira, b- Ain Al Atti
Taking into account the electrical logs (Fig. 8a and 8b), we observe that for the upper part of the field investigated, the formations encountered at Douira are relatively the same encountered at Ain Al Atti, however in depth we find a variation of facies (gray and limestone at the level of Douira and the marls at the level of Ain Al Atti), these marly formations with intercalation of gypsum encountered in depth, caused a significant decrease in the electrical resistivity in the vicinity of Ain Al Atti.

The line (Fig. 8a) represents the variation of the apparent section of the current density as a function of the distance, along line 2. The plot shows a prominent positive response between 900-1000 m, from the start of the profile, corresponding to a conductive ax, resulting in a fracture zone at a depth ranging from the surface to 170 m. The results of line 3 are represented as a cross section of apparent current density in Fig. 9a. The plot reveals a positive anomaly linked to the presence of a conductive target at 1580 m from the start of the profile. It is interpreted as a fracture zone at a depth ranging from surface to 140 m. The KH Filtering data from the rest of these lines does not show any pronounced anomalies. The result of line 4 (Fig. 10a) suggests the presence of a conductive anomaly at a distance of 350 m which corresponds to the same fractured zone. The filtered Fraser data reveals the presence of anomalies which are located and correlated to the anomalies obtained by Karous Hjelt filtration. Three positive anomalies were distinguished, where the apex of the anomalies was observed at locations L2 (Fig. 8b), L3 (Fig. 9b) and L4 (Fig. 10b). These are linked to the presence of conductive structures, interpreted as fracture zones. This fracture has a major NW-SE direction which confirms the work carried out at the level of the locality Zaouia Jdida by Ait (Bahammou et al., 2019) is schematized on the model of Fig. 12.

**Fig. 8:** Synthetic electric logs: a- Survey carried out in Douira, b- Survey carried out in Ain Al Atti
Fig. 9: Electromagnetic VLF-L2: a- apparent current density cross section, b- fraser filtering

Fig. 10: Electromagnetic VLF-L3: A- apparent current density cross section, b- fraser filtering
Fig. 11: Electromagnetic VLF-L4: a- apparent current density cross section, b- Fraser filtering

Fig. 12: Schematic model of the location of the fracturing zone
Conclusion

In this article, we have contributed to the litho-structural study at the level of the locality of Douira and near the artesian source of Ain Al Atti, by combining the electrical technique and the very low frequency electromagnetic method. The results of the electrical soundings showed the presence of very conductive grounds in the vicinity of the source which is due to the capillary rise of the salinity coming from marly formations with gypsum intercalations in depth. Data filtered by VLF revealed the presence of a fracture anomaly which reached 150 m in depth and which is oriented NW-SE. The identified fracture anomaly constitutes the potential groundwater circulation area that led to the supply of the Ain Al Atti source.

Given the importance of the results obtained at the level of the prospected area, it is advisable to carry out other electromagnetic profiles and electrical trails in the northern part of the locality of Douira and to carry out a mechanical survey at the level of these conductive anomalies. To block the loss of salt water from the Ain Al Atti source.

Acknowledgement

The authors are sincerely grateful to the Faculty of Science and Technology Errachidia (FSTE) for proving the geophysical machines used in this study.

Author’s Contributions

Dakir Ibrahim: I carried out the geophysical measurements, I contributed to the interpretation of the data, the writing of the article.

Benamara Ahmed: I am the professional supervisor of Mr. DAKIR. I helped the main author with the interpretation of geophysical data.

Aassoumi Habiba: I am the administrative supervisor of Mr. DAKIR. Helped the main author with manuscript writing and translation.

Ouallali Abdessalam: I contributed to the field level; we carried out together several geophysical companies.

Ait Bahammou Youssef: I contributed to the field level; we carried out together several geophysical companies.

Ethics

This article is original and contains unpublished data. The corresponding author confirms that all of the other authors have read and approved the manuscript and no ethical issues involved

Data Availability

The data used to support the findings of this study are available from the corresponding author upon request.

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