Olistostrome and the mesozoic tectonic of the bantimala complex, South Sulawesi

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Abstract. The study aims to determine the presences and spreads of the olistostrome as one components of the tectonic complex of Bantimala area. The basement rocks of Bantimala area is a metamorphic rocks, which are unconformably overlain by Balangbaru Formation and radiolarian chert. Beneath in between cherts and basement rocks are presence breccia schists which give rise to various presumptions and interpretations of the environment and conditions of formation as submarine deposit, ideally cherts in the area underlain by oceanic crust.Schist breccias presence underneath of cherts in the Bantimala Complexes were suggested an olistostrome deposit. It was characterized by poorly sorting, unfoliated, shows deformed textures and composed olistoliths which are embedded in sandy matrices, and in the cherts are presence layer sandstones and schist fragments. Olistostrome is sedimentary deposit as preserved in the trench, they will give us an interpretation that prior to formed cherts in the Bantimala area, initially tectonic subduction activity which are deforming and brecciation of the basement rocks subsequently as the constituent material of olistostrome.

1 Introduction

The olistostrome of Bantimala Tectonic Complexes is likely related to Western Pacific plate has subducted to Asian Continent during Mesozoic, the interaction both of the plates were caused deformation, metamorphism and olistotrome deposits. These situation is reasonably allow to occur a debris fall due to steep slope in the part of trenches as suggested by [1], [2], their fragments possibly are eroded debris materials from continent and oceanic plates, they have been mixed as a chaotic sedimentary deposits, its immediately covered by deep marine sediments of radiolarian cherts.

The oldest rocks were exposed in the Bantimala Complexes are Triassic metamorphic rocks composed of glaukofan schists, mica-hornblende schists, eclogites, granulits, quartz phylite and quartzite [3]. On the top of this basement rocks is overlain by breccias, schists (olistostrome), sandstones and radiolarian cherts indicating Jurassic - Cretaceous age [4].

Our side observations that schist breccias in the study area is an olistostrome deposits characterized by present vary of olistolith schists, stratigraphically it overlain by radiolarian cherts. So far, the study of olistotrome in the area has never been studied, but previous studies it was indicated a mélange [5] thought a chaotic sedimentary processing which is more appropriately considered as olistostrome.

2 Methods

This study emphasizes field observations and petrographic analysis. Field observations was performed stratigraphic detail by measuring section and cross section along Patteteyang river, observing the relationship between the lithologic unit and indication deformation structures in the field and rocks sampling. Further observations were made followed by petrographic analysis to determine the texture and rock composition for all components of rocks, especially for olistolith and matrices.

3 Regional geology and tectonic setting

Tectonically the Bantimala Complexes is still contentious, including presence of rocks faragments and interrelationship each other such as presence of mélanges, schist basement rocks, rocks continent affinity, radiolarian cherts and Paleogene volcanics.

3.1 Geological setting

The basement rocks were exposed in the Bantimala area is part of shelf margin of East Kalimantan which was separated since Miocene coincided with the occurrence of Makassar Strait. The basement rocks is called the Bantimala Tectonic Complex where composed of: Triassic metamorphic rocks namely glaucohpane schists,
hornblende-mica schists, eclogites, granulites, phyllites and metaquarzites [3]; Jurassic-Cretaceous of mélangé composed of quartz schists, metacherts, metabasalts. Cretaceous sedimentary rocks which composed of fine-siliceous shales, sanstone, mudstone and radiolarians.

Ophiolite blocks consisting of harzburgites and serpentinites where occur through out tectonic obduction which covered or structural contact with the Tertiary rocks in the area, whereas sedimentary type of continental shelf margin are Cretaceous of typical flysch sedimentary rocks of Balangbaru-Paremba Formations. They were unconformity overlain by quartz sandstones of Mallawa Formation and Paleocene-Eocene volcanics. They are covered by Eocene-Miocene limestone of Tonasa Formation and Upper Miocene-Pliocene breccias and tuffs of Camba Formation [6] Fig 1.

3.2 Tectonic setting

The Bantimala Complex is formed in two tectonics phases, the first is subduction system of oceanic plate that took place during the Mesozoic to Tertiary. The second is ophiolite obduction system that took place during Tertiary to Quaternary. The presence of high grade of metamorphic rocks associated with system during pre-Cretaceous (Triassic - Jurassic), it may cause by the West Pacific plate subducted to the East Kalimantan continent, which was produced accretion complex on the eastern shelf of Asian Continental plates. During Jurassic to Cretaceous the accretion complex has deformed to produce a mélange rocks [7-8]. During the Early Cretaceous occurred brecciation in the accretion complexes formed breccia - schist [5] or here called olistostrome [9], while continuing deposited chert and contemporaneously with deposition in the trench of terrigenic sediment of Balangbaru (Figure 2).

More over the Cretaceous of Bantimala Complexes have been situated by some of previous study such as [10] and [11], according to them that the establishment of a high-pressure metamorphic rocks associated low-grade metamorphic rocks, mélange and ultramafic in this region are the results of system subduction of oceanic crust into the shelf margin of continental plates in the Jurassic to the Early Cretaceous (114 to 132 Ma). Based on calculations of pressure - temperature of rocks garnet - glaucophane resulted a temperature of 580 – 640 °C and pressures 18-24 kbar, occurs at depth 65-85 km [5]. Furthermore the breccia of Bantimala Complexes unconformably overlap by the Radiolarian chert which ranging from Albian to Cenomanian in age (± 100 Ma) [12].

4. Physical properties of Olistostrome

Lithologic characterized of olistostrome of Bantimala Complex could be described into two aspects i.e. based on components and sedimentary types.

4.1. Component Fabrics

Olistostrome component in the Bantimala Complex shows various clast of material component which is indicated difference origin (polymictic), it composed by deformed and reworked materials derived from subduction activity which consisting of schist, quartzite and gneiss, serpentinite, metachert, the components are vary in grain size ranging from centimeter to several meter, poorly sorting and angular to subangular. Fragments floating in reddish chert matrices suggesting retransported material within a trench, deformed textures indicated by lenses, fracturing, shows pseudofoliation with slipped surface (Figure 9 and 10).

The following is results of petrographic analysis of olistoliths of olistostrome, schists, cherts, and sandstone as follows:

1. Schists and Quartzite Olistoliths:
   - Schist olistoliths of schist muscovite quartz (OL, 5C), lepidoblastic textures, composed by muscovite 40%, quartz 30%, chlorite 20%, actinolite 5%, mineral opak 5% (Fig. 3).
   - Quartzite Olistolith (OL, 5D), composed by quartz 75%, muscovite 12%, and biotite 10%, mineral opak 5%, granoblastic textures (Figure 4).
2. Schist and greenschist
- Olistolith schist (OL, 2B), lepidoblastic textures composed by muscovite 65%, quartz 20%, chlorite 5%, albite 6%, and ore mineral 4% (Fig. 5).

- Green schist/ actinolite schist (ML, 3B), lepidoblastic textures, composed by actinolite 90%, quartz 5%, chlorite 3%, ore mineral (Fig. 6).

3. Chert (C, 4E)
Chert composed by radiolaria 40%, quartz 56%, ore mineral 4%, (Fig. 7).

4 Sandstone (ST. 13A)
Sandstone composed by muscovite 70%, quartz 25%, ore mineral 5%, (Fig. 8).
4.2 Type and physiography sediment

4.2.1 Deposition type Olistostrome

Quartz muscovite schist in one complex with glauophane schists, granulits, eclogites and radiolarian cherts, as an indication of sediment avalanche or slope slumping at a relatively steep slope suggested rocks that preserve from subduction between the continental plates with oceanic plates, where olistostrome interbedded in part the bottom of the radiolarian cherts [13]. To mixing materials in chert clast and fining upwards structures, cobble – granule in grain sizes, indicating that they were preserve in a deep marine or trenches on subducting tectonic conditions, based on the above determination so that the olistostrome of Bantimala Complex is categorized as subduction type.

4.2.2 Sediment Physiography

The appearance of interbedded between the sandstones and radiolarian chert as the top of the olistostrome, gradational structure components in the western part of the study area including fluxo deposit. Based on deposit type in the eastern part and the contact between the radiolarian chert coarse grains of olistolith indicated proksima type deposit (Fig. 9 and 10). Based on the spread and characteristics of layering can be interpreted the depositional direction of olistostrome from west to east. The presence of serpentinite – jadeite blocks of as exotic block in olistostrome is separate phenomenon to be determine in the future work. Body of olistostrome in Bantimala Complex suggested as a lensis in radiolarian cherts unit.

5. The occurrences of olistotrome

The process of subduction of the Pacific plate along continental margin of Kalimantan during Jurassic, initially the process of forming olistostrome of Bantimala. During that period was preserved tectonic deformation, brecciation and metamorphism of the two plates subducted each other, which is accompanied by the formation of deep trenches as the depositional environment. Plate continental and oceanic plates that have undergone at low grade metamorphism and brecciated forming blocks as olistoliths that shows deformation textures.

Based on experiments [14] and [15] on the slopes of the critical condition allow submarine landslide occurred in the trench, the material falling as debris flow or slumping and widespread deposited on the ocean floor which its mixture with clast of continental and oceanic components, [16] and [17], (Fig. 11).

Based on their appearance in the field, olistostrome layer at Pateteyang River (Bantimurung Village), there are at least four cycle of avalanche or debris deposition of material inter bedded with cherts, namely:

- The first is preserve debris avalanches and slumping, forming as rock blocks with a thickness of ± 340 m olistolith sizes between 1-150 cm, angular and boudinage. Components consist of chlorite schist, mica schist, amphibolite schist, and quartzite gneiss.
- The second, before the second deposition is preceded by the deposition of a thin layer of chert 20 cm thick and mixed with coarse grains of chert material followa debris avalanche of gravel-blockschist clasts (2-40 cm), relatively smoother than the first avalanche, with a thickness ± 150 cm.
- The third is also inter bedded by chert layers 60 cm in thick. The third layer is composed of schist clast with 25 cm thick.
- The fourth, deposition above of layer chert (120 cm thick), a thin layer of chert clast ± 20 cm thick and
covered by layers cherts that show the body olistostrome of Bantimala (Fig. 12). The composition and structure of the sediment grain size olistostrome shows fining upwards as debris flow sediment or turbidity (Table 1).

Fig. 11. Cross section of formation of olistotrome of Bantimala Complex, subduksi lempeng Oceanic Pasific plate (black) subducted to Asean Continental plate (orange) during Early Cretaceous.

Fig. 12. Outcrop olistostrome (X) with radiolarian chert (Y) in the Pateteyang River, Bantimala

6 Discussion

Synsedimentary deformation and tectonic implications, field data shows deformation of soft-sediment on cherts with showing small folds (antiform and synform) (Fig. 13), brecciation, dragfold, microthust, boudinage olistolith surrounded by a fine grained schist matrix, slump fold and slump ball (Fig. 14), indicated that during sedimentation contemporaneous with active tectonic condition.

The stratigraphy position of olistostrome by Wakita et al. (1996) is above the wedge with Radiolarian chert and intercalated with Balangbaru sandstone, and there are also schists interbedded chert on the bottom. Breccia schist by Wakita et al. (1996) and the authors include as olistostrome deposit. Stratigraphy position hardly visible in the field but by outcopping thin layer of ± 50 cm (Figure 10) olistostrome interbedded in chert in the eastern part of the area of research, it can be interpreted that the (1996) that is deposited in the same time with the slump sedimentation system on the bottom layer of chert.

Fig. 13. Tabel of stratigraphic column of olisostrom in Bantimala Complex

The stratigraphy relationship breccia schist in this case olistostrome is unconformable overlaying cherts (Wakita, et al., 1994), conversely based on the age of the fossil in radiolarian chert, breccia or sandstone is conformable. It is also supported by the appearance of the field that showed gradational from bottom-top with fining upward structures and almost breccias composed by schist clast indicated same environment and sedimentary sources suggested that the relationship between radiolarian chert and breccia schists is conformable.

7 Conclusion

a. Geological conditions and tectonic Bantimala Complex composed by pre-Tertiary components (Schists, olistostromes, radiolarian cherts and sandstones of Balangbaru Formation) and Tertiary components (Volcanic Paleocene, sandstones of Mallawa Formation, limestones of Tonasa Formation, Camba volcanic and basic-intermediate-acid igneous intrusions. Some of them bounded by thrust faults.

b. Characteristics of olistostrome in the Bantimala Complex demonstrated by heterogeneous components with polymictic breccias consisting of schist, gneiss,
quartzite and metachert components, composed of three layers, namely breccia polymictic, coarse sand -granule and sand. They including in olistostrome subduction type that formed in Lower Cretaceous (Albian).

c. Stratigraphically the position of olistostrome or breccia schist conformably overlaying by radiolarian chert.

d. The olistostrome component is characterized by deformation textures and interrelationship with radiolarian chert where it was deformed during of post-sedimentation of the Jurassic – Cretaceous.

Fig. 14. Antiform (X) and synform (Y) structures in the cherts, Pateteyang River, Bantimala

Fig. 15. Outcrop of radiolarian cherts deformed, Pateteyang River, Bantimala

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