Abstract

Tidally locked gas giants, which exhibit a novel regime of day–night thermal forcing and extreme stellar irradiation, are typically in several-day orbits, implying a modest role for rotation in the atmospheric circulation. Nevertheless, there exist a class of gas-giant, highly irradiated objects—brown dwarfs orbiting white dwarfs in extremely tight orbits—whose orbital and hence rotation periods are as short as 1–2 hr. Phase curves and other observations have already been obtained for this class of objects, raising fundamental questions about the role of an increasing planetary rotation rate in controlling the circulation. So far, most modeling studies have investigated rotation periods exceeding a day, as appropriate for typical hot Jupiters. In this work, we investigate atmospheric circulation of tidally locked atmospheres with decreasing rotation periods down to 2.5 hr. With a decreasing rotation period, we show that the width of the equatorial eastward jet decreases, consistent with the narrowing of the equatorial waveguide due to a decrease of the equatorial deformation radius. The eastward-shifted equatorial hot-spot offset decreases accordingly, and the off-equatorial westward-shifted hot areas become increasingly distinctive. At high latitudes, winds become weaker and more rotationally dominated. The day–night temperature contrast becomes larger due to the stronger influence of rotation. Our simulated atmospheres exhibit variability, presumably caused by instabilities and wave interactions. Unlike typical hot Jupiter models, the thermal phase curves of rapidly rotating models show a near alignment of peak flux to secondary eclipse. This result helps to explain why, unlike hot Jupiters, brown dwarfs closely orbiting white dwarfs tend to exhibit IR flux peaks nearly aligned with secondary eclipse. Our results have important implications for understanding fast-rotating, tidally locked atmospheres.

Unified Astronomy Thesaurus concepts: Brown dwarfs (185); Exoplanet atmospheres (487); White dwarf stars (1799); Hydrodynamics (1963); Hot Jupiters (753)

1. Introduction

1.1. Observational Motivation

Since the first transiting observation of HD 209458 b (Charbonneau et al. 2000), gas-giant planets with small semimajor axes (aka the hot Jupiters) remain the best-characterized group of exoplanets. Due to their close-in distances, they experience extreme stellar irradiation and, because they are expected to be tidally locked, exhibit a novel regime of permanent day–night thermal forcing. Thermal phase curves and secondary eclipse measurements have been obtained for a variety of hot Jupiters, providing constraints on their three-dimensional (3D) thermal structure and circulation pattern (Heng & Showman 2015). The flood of observations has stimulated a growing body of modeling studies on the 3D atmospheric circulation of tidally locked gas giants (e.g., Showman & Guillot 2002; Cooper & Showman 2005; Dobbs-Dixon & Lin 2008; Showman et al. 2009; Menou & Rauscher 2010; Heng et al. 2011a; Perna et al. 2012; Rauscher & Menou 2013; Mayne et al. 2014; Mendonça et al. 2016; Roman & Rauscher 2019; Steinrueck et al. 2019). Yet most observations and theoretical models to date have focused on planets that are typically in several-day orbits, implying a low rotation rate and a modest role for rotation in the atmospheric circulation.

Nevertheless, there exists a class of gas-giant, highly irradiated objects—brown dwarfs (BDs) orbiting white dwarfs (WDs) in extremely tight orbits—whose orbital and hence rotation periods can be as short as 1–2 hr. These extremely close-in gas giants are expected to be synchronously rotating, as the tidal spin-down time is short compared to system ages (e.g., Guillot et al. 1996). Thus, their orbital period provides a good estimate of the rotation period. These systems are often survivors of binary evolution—when Sun-like stars become old and expand, engulfing their close-by companions, the orbits of their smaller companions shrink due to friction. The outer layer of the evolved star eventually puffs off, leaving a remnant WD with a closely orbiting low-mass (i.e., BD) companion (Hellier 2001; Percy 2007). Many secondary BDs of such systems are donors; i.e., materials from their atmospheres are accreted onto the WDs, and their observations are often complicated by contributions from the accretion disks, as well as the distinctly nonspherical shape of the BDs (e.g., Santitubol et al. 2016). Fortunately, some of them are detached systems in which there are no outflows from the secondary BDs, making them the ideal targets to understand the 3D atmospheric structure and circulation of gas giants that are extremely fast-rotating and highly irradiated. An observational advantage of this class of objects is that near-IR photons emitted from the companion BDs are generally well separated in wavelength from those emitted from the host WD (whose flux peak is typically in the UV); typically, the BD emits far greater IR energy than the WD, greatly aiding observational characterization.4 Phase curves and other observations have been
obtained for several such systems, including NLTT 5306 (orbital period 102 minutes; Steele et al. 2013), WD 0137–349 (116 minutes; Casewell et al. 2015; Longstaff et al. 2017), EPIC 21223532 (68 minutes; Casewell et al. 2018a), WD 1202–024 (71 minutes; Rappaport et al. 2017), and SDSS J141126.20+200911.1 (122 minutes; Liddlefair et al. 2014; Casewell et al. 2018b). In particular, phase curves of WD 0137–349 and SDSS J141126.20+200911.1 were taken with multiple wavelengths. Several candidate detached WD+BD systems should be observable in the near future (see a summary in Casewell et al. 2015). These phase curves often exhibit large day–night temperature differences and nearly zero phase offsets between the peak flux and phase 0.5, where a secondary eclipse would occur if the system were eclipsing. In some cases, the flux peaks slightly after phase 0.5, which is in contrast to most hot Jupiter near-IR phase curves that show robust peak flux before the secondary eclipse. This raises fundamental questions about the role of strong rotation in controlling the global circulation and day–night heat transport in tidally locked atmospheres.

In between the extremely close-in BDs around WDs and the canonical hot Jupiters that are typically in several-day orbits, a handful of giant planets and BDs with orbital periods close to or less than a day around Sun-like stars have been discovered or characterized (e.g., WASP-12b, orbital period 1.1 days, Swain et al. 2013; WASP-103b, 0.93 day, Cartier et al. 2016; Kreidberg et al. 2018; WASP-18b, 0.94 day, Sheppard et al. 2017; Arcangeli et al. 2018; WASP-19b, 0.78 day, Espinoza et al. 2019; NGTS-7Ab, 0.68 day, Jackman et al. 2019; TOI 263.01, 0.56 day, Parvainen et al. 2020). There are significant observational motivations to investigate atmospheric physics for this emerging class of exoplanets (e.g., Bell & Cowan 2018; Kitzmann et al. 2018; Komacek & Tan 2018; Lothringer et al. 2018; Parmentier et al. 2018; Tan & Komacek 2019). In the near future, global atmospheric properties and circulation of these short-period planetary companions will be better constrained via phase curves and secondary eclipse measurements, and therefore a better understanding of the stronger role of rotation in atmospheric dynamics is needed.

### 1.2. Theoretical Motivation

The dominant dynamical response to day–night radiative forcing in the atmosphere of a canonical hot Jupiter is the excitation of standing, global-scale equatorial Rossby and Kelvin waves. The superposition of the equatorial Rossby and Kelvin waves (the so-called Matsuno–Gill pattern; Matsuno 1966; Gill 1980) causes phase tilts in the eddy velocities that pump eastward momentum from high latitudes to the equator, thereby inducing a fast eastward equatorial jet, or superrotation. Combined with advection due to the equatorial superrotation, the eastward propagation of the Kelvin waves also helps displace the hot spot eastward of the substellar point. Analytical wave solutions were given by Showman & Polvani (2011) in a linearized shallow-water system with finite radiative damping and frictional drag. Subsequently, Tsai et al. (2014) investigated solutions in a continuously stratified atmosphere with a uniform zonal wind in the regime of an equal radiative and drag timescale, Heng & Workman (2014) extended the shallow-water solutions to include magnetic field and spherical geometry, Hammond & Pierrehumbert (2018) investigated the effects of the general profiles of zonal-mean zonal wind on the wave-mean flow interactions using the shallow-water system, and Debras et al. (2020) explored the time-dependent wave solutions under day–night forcing. An example of the steady wave solution with moderate nondimensionalized radiative and drag timescales that are appropriate for typical hot Jupiters is shown in the upper panel of Figure 1. In this regime, the northwest–southeast (northeast–southwest) eddy tilt in the northern (southern) hemisphere pumps eastward angular momentum from the off-equatorial region to the equatorial region. A general understanding of the emergence of this eddy tilt is the three-way horizontal force balance between pressure gradient, Coriolis, and frictional drag force. Even when the explicit frictional drag is weak, as may be the case in many hot Jupiters, a similar three-way force balance can still be achieved by nonlinear advection terms, allowing the emergence of such an eddy tilt and driving the equatorial superrotation (Showman & Polvani 2011; for a brief review, see Showman et al. 2013).

This Matsuno–Gill pattern is typical in slow rotators in which the Rossby number \( \text{Ro} \sim 1 \) and the equatorial Rossby deformation radius \( L_D \) is comparable to the planetary radius. The deformation radius is defined as \( L_D = \sqrt{c/\beta} \), where \( c \) is the gravity wave speed and \( \beta = df/dy \) is the meridional gradient of the Coriolis parameter \( f \), where \( y \) is distance, increasing northward. The Coriolis parameter is \( f = 2\Omega \sin \phi \), where \( \phi \) is latitude and \( \Omega \) is the planetary rotation rate, and therefore \( \beta = 2\Omega \cos \phi /a \), where \( a \) is the planetary radius, implying \( \beta = 2\Omega /a \) at the equator. A good estimate of the horizontal phase speed of long vertical wavelength gravity waves is \( c \sim 2NH \), where \( N \) is the Brunt–Väisälä frequency and \( H \) is the scale height. One can then express the equatorial deformation radius as \( L_D = \sqrt{NH\Omega}/\Omega \). The theory of Showman & Polvani (2011) predicts that the meridional half-width of the equatorial jet on synchronously rotating planets is approximately the equatorial deformation radius.

As the rotation period decreases (rotation rate increases), the deformation radius decreases relative to the planetary radius, and the meridional extent of the associated standing waves are expected to be smaller and smaller, confining the wave structure closer to the equator. At high latitudes—beyond the equatorial waveguide—the force balance is expected to be primarily between the pressure-gradient force and the Coriolis

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**Figure 1.** Height fields (with arrows representing wind vectors overplotted) of the wave solutions triggered by day–night forcing for a shallow-water system from Showman & Polvani (2011). The solution adopts finite \( \tau_{\text{drag}} \) and \( \tau_{\text{rad}} \) which are appropriate for typical hot Jupiters. At the high day–night forcing amplitudes relevant to hot Jupiters, nonlinear effects are important, and in that case, advection can play the qualitative role of drag, leading to a similar structure even when there is no drag in the active layer. Note that for typical hot Jupiters, the structure extends all the way from the equator to high latitudes. At a faster rotation rate, a similar structure occurs but becomes confined closer to the equator due to the decrease in equatorial deformation radius with increasing rotation rate.
force if the frictional drag is weak, which is the so-called geostrophic regime.

Based on the above framework, we expect to see an emergence of two distinct behaviors with decreasing rotation period: standing waves are triggered but confined closer to the equator, and the flow becomes increasingly geostrophic at high latitudes. Equatorial superrotation may still be prevalent, but its meridional extent would decrease due to the smaller equatorial deformation radius. These changes in the circulation pattern will directly influence the thermal phase curves and other observables. In the slow-rotating regime, the eastward-shifted hot spot dominates and results in a peak thermal deformation radius. These changes in the circulation pattern would decrease due to the smaller equatorial force if the frictional drag is weak, which is the so-called geostrophic regime.

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To gain a comprehensive understanding of the circulation on rapidly rotating, tidally locked gas-giant planets, it is best to seek solutions using well-controlled numerical experiments. In this study, utilizing a general circulation model (GCM) with idealized day–night thermal forcing, we investigate the effects of increasingly strong rotation on the atmospheric circulation of tidally locked atmospheres. This will help to provide a foundation for understanding current and future observations of ultraclose-in, rapidly rotating BDs and exoplanets. To isolate the dynamical effects of rotation in a clean, systematic manner, we intentionally simplify our model by excluding other potentially important effects on the circulation, such as feedback involving radiatively active clouds (e.g., Lee et al. 2016; Lines et al. 2018; Roman & Rauscher 2019), magnetic drag (e.g., Perna et al. 2010; Batygin et al. 2013; Rauscher & Menou 2013; Rogers & Komacek 2014; Rogers 2017), and heat transport by hydrogen dissociation and recombinations (Bell & Cowan 2018; Komacek & Tan 2018; Tan & Komacek 2019). In addition, vigorous atmospheric circulation of field BDs has been inferred by numerous observations (see a recent review by Biller 2017). The circulation on these field BDs is forced by the internal heat flux, which presumably does not have day–night or equator–pole variation (Showman & Kaspi 2013). Yet these objects likewise exhibit fast rotation, high temperature, and high surface gravity. The use of a simplified GCM here offers a clean comparison to investigations of isolated BDs using idealized global models (Zhang & Showman 2014; Tan & Showman 2017; Showman et al. 2019) whose atmospheric circulation is forced in a very different manner.

This paper is organized as follows. We introduce the numerical model in Section 2. In Section 3, we first describe the general results for models with a short radiative timescale in Section 3.1 and models with a long radiative timescale in Section 3.2. We explore the effects of drag on rapid rotators in Section 3.3 and present synthetic phase curves of models in the grid in Section 3.4. Finally, we discuss the implications of our results in Section 4 and highlight key conclusions in Section 5.

2. Model

We solve the standard global hydrostatic primitive equations in pressure coordinates, which include the horizontal momentum, hydrostatic equilibrium, continuity, and thermodynamic energy equations. The ideal gas law is assumed for the equation of state of the atmosphere. Our physical model is almost identical to those used in Liu & Showman (2013) and Komacek

Becker & Showman 2013; Komacek & Showman 2016).

Previous work has explored the influence of rotation rate on the atmospheric circulation of hot Jupiters. Showman et al. (2008) varied the rotation rate under the assumption of synchronous rotation, while Showman et al. (2009) and Rauscher & Kempton (2014) explored the rotation rate varying simultaneously with allowing the heating pattern to migrate nonsynchronously, all within a factor of 2 or so variations in rotation rate. Showman et al. (2015) widened the range of rotation rate variations, including nonsynchronous rotation. Komacek et al. (2017) varied the rotation rate and stellar insolation in a self-consistent manner. Kataria et al. (2013) performed simulations for quasi-fast rotators in the context of eccentric planets rotating pseudosynchronously. Penn & Vallis (2017) explored the effects of a different rotation rate and nonsynchronization on the thermal phase-curve offset using shallow-water models. Flowers et al. (2019) systematically compared the observed high-resolution transmission spectrum of HD 189733b to models with a wide range of rotation periods. In general, these studies tend to support the picture that faster rotation leads to a narrower equatorial jet. But in most of these studies, the complicating effects of nonsynchronous rotation, varying stellar insolation, and eccentric orbits complicate an understanding of the effect of varying rotation by itself. More importantly, none of these studies reach the ultrafast-rotation regime relevant to WD–BD binaries, as we want to do here.

Time-dependent processes, including large-scale instability and free propagating waves, have been given little attention in existing theoretical work for hot Jupiters, partly because, in both observations and numerical simulations of hot Jupiters, the dominant atmospheric features are almost time-invariant. A well-known source of time variability in rapidly rotating atmospheres is baroclinic instability, which dominates the large-scale midlatitude dynamics in Earth’s troposphere (for reviews, see, e.g., Showman et al. 2013; Vallis 2017) and may contribute to driving the multiple zonal jets in Jovian and Saturnian atmospheres (see a recent review by Showman et al. 2019). Conditions on the close-in gas giants differ from low-temperature terrestrial atmospheres in several ways, including a much shorter radiative timescale near the photospheres, possibly strongly stratified photospheres, the lack of a surface temperature gradient, and, finally, day–night thermal forcing instead of the equator-to-pole forcing more typical on planets like Earth, Mars, and the solar system’s giant planets. At present, it is therefore unclear whether (and to what extent) baroclinic instability plays an important role in maintaining the atmospheric circulation of close-in gas giants. Other types of large-scale instability—for example, barotropic instability, which is relevant in Jovian and Saturnian atmospheres (Ingersoll & Pollard 1982)—are also worth examining in the context of day–night-forced, rapidly rotating atmospheres.

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Radiative heating and cooling are represented using a Newtonian scheme in which the temperature field is relaxed toward a prescribed equilibrium temperature profile over a characteristic radiative timescale \( \tau_{\text{rad}}(p) \). The radiative timescale \( \tau_{\text{rad}} \) is a function of pressure only. The radiative-equilibrium temperature structure \( T_{\text{eq}}(\lambda, \phi, p) \), which is hot on the dayside and cold on the nightside, is taken to be the following function of longitude \( \lambda \), latitude \( \phi \), and pressure \( p \):

\[
T_{\text{eq}}(\lambda, \phi, p) = \begin{cases} 
T_{\text{night,eq}}(p) + \Delta T_{\text{eq}}(p) \cos \lambda \cos \phi & \text{dayside}, \vspace{1mm} \\
T_{\text{night,eq}}(p) & \text{nightside}.
\end{cases}
\]  

Here \( T_{\text{night,eq}}(p) + \Delta T_{\text{eq}}(p) \) is the equilibrium profile on the nightside and \( T_{\text{night,eq}}(p) \) is that at the substellar point. The nightside profile is acquired by subtracting our chosen \( \Delta T_{\text{eq}}(p)/2 \) from a temperature profile of HD 209458b from Iro et al. (2005). The day–night difference in radiative-equilibrium temperature, \( \Delta T_{\text{eq}}(p) \), is set to be a constant \( \Delta T_{\text{eq,top}} \) at a pressure lower than \( p_{\text{eq,top}} \) zero at a pressure larger than \( p_{\text{bot}} \), and varying linearly with log-pressure in between. We take \( p_{\text{eq,top}} = 10^{-3} \) bars, \( p_{\text{bot}} = 10 \) bars, and \( \Delta T_{\text{eq,top}} = 1000 \) K for most of the models. The radiative timescale is generally short in the upper atmosphere and long at depth (e.g., Iro et al. 2005; Showman et al. 2008). To qualitatively capture this behavior, \( \tau_{\text{rad}} \) is set to a small constant \( \tau_{\text{rad,bot}} \) at pressures less than \( p_{\text{rad,bot}} \), and a large constant \( \tau_{\text{rad,bot}} \) at pressures greater than \( p_{\text{bot}} \); then it varies linearly in logarithmic space in between. We set \( p_{\text{rad,bot}} = 10^{-2} \) and \( p_{\text{bot}} = 10 \) bars. This forcing setup is nearly identical to that in Liu & Showman (2013) and Komacek & Showman (2016). Because the BDs in BD+WD systems experience a wide range of radiative time constants, and because there is a strong theoretical motivation for investigating how the atmospheric circulation responds to varying radiative timescales (Showman & Polvani 2011; Perez-Becker & Showman 2013; Komacek & Showman 2016), we systematically explore values of \( \tau_{\text{rad,bot}} \) ranging from \( 10^4 \) to \( 10^7 \) s. The rotation period is likewise varied over a wide range.

The frictional drag is represented by a linear damping of horizontal velocities, which represents missing physics, for instance, turbulent mixing or the Lorentz force on the velocity (see Komacek & Showman 2016 for a detailed discussion). Kinetic energy dissipated by the drag is returned to the system as a heating term in the thermodynamic energy equation. As in Komacek & Showman (2016), there are two components in the drag. The first one is a basal drag applied only near the bottom of the model domain, which crudely parameterizes interactions between the weather layer and a relatively quiescent planetary interior. The basal drag is zero (meaning no drag) at pressure less than \( p_{\text{drag,bot}} \), and the drag coefficient linearly increases with increasing pressure from zero at pressure \( p_{\text{drag,bot}} \) to a maximum value \( 1/\tau_{\text{drag,bot}} \) at the bottom of the domain (~200 bars). At the bottom of the computational domain, the basal drag strength is characterized by a drag timescale \( \tau_{\text{drag,bot}} \). We set \( p_{\text{drag,bot}} = 10 \) bars and \( \tau_{\text{drag,bot}} = 10^6 \) s for all models. The basal drag does not directly influence the dynamics at pressures less than 10 bars. The second component, which we only include in some models, is a spatially independent drag—applied everywhere in the domain—that is characterized by a constant drag timescale \( \tau_{\text{drag}} \). It is intended primarily as a very crude representation for Lorenz forces that might be associated with partial ionization on particularly hot planets (e.g., see Perna et al. 2010; Rauscher & Menou 2013; Komacek & Showman 2016 for discussion). Note that most of our models lack the second component; however, we will present some models that include it. The total frictional drag at any grid point takes the maximum of the two drag components.

We adopted planetary parameters relevant for gas giants, including specific heat \( c_p = 1.3 \times 10^4 \) J kg\(^{-1}\) K\(^{-1}\), specific gas constant \( R = 3714 \) J kg\(^{-1}\) K\(^{-1}\), and planetary radius \( 7 \times 10^7 \) m (similar to that of Jupiter). Because the major motivation of this study comes from observations of BDs around WDs, we adopted the surface gravity \( g = 1000 \) m s\(^{-2}\) relevant for BDs. However, as discussed in Showman et al. (2019), the entire system is independent of the value of gravity for our particular model formulation (i.e., the equations are solved in pressure coordinates with a prescribed Newtonian cooling scheme that is also specified to be a function of pressure rather than height). Thus, the qualitative implications from our numerical results should be applicable to both hot Jupiters and BDs. We systematically vary the rotation period from 80 to 2.5 hr and \( \tau_{\text{rad,bot}} = 10^4 \) to \( 10^7 \) s and assume a drag-free atmosphere for the main suite of simulations (in other words, our main suite of simulations has only the basal drag scheme at pressures greater than 10 bars; the main atmosphere at \( p < 10 \) bars lacks any frictional drag). In some rapidly rotating experiments, we include the finite drag timescale \( \tau_{\text{drag}} \) throughout the atmosphere to explore dynamics with increasing drag strength.

We use the MITgcm to solve the global hydrostatic primitive equations using the cubed-sphere grid (Adcroft et al. 2004). A standard fourth-order Shapiro filter is applied to the horizontal momentum and thermodynamics equations to maintain numerical stability, and it has a minimal effect on the large-scale structures. The pressure domain is between 200 and \( 2 \times 10^{-4} \) bars and evenly divided into 40 levels in log-pressure. Because the dynamical length scales (e.g., deformation radius) are small at fast rotation, properly resolving the global-scale dynamics at fast rotation requires greater horizontal resolution than at slow rotation. As is standard in the field, we adopt a resolution for our slowly rotating, 80 and 40 hr models of \( C_{32} \) (meaning \( 32 \times 32 \) points on each cubed-sphere face, equivalent to a global resolution of \( 128 \times 64 \) in longitude and latitude) but increase the resolution up to \( C_{192} \) for models at a 2.5 hr rotation period. The latter implies that we use \( 192 \times 192 \) points on each cubed-sphere face, equivalent to a global resolution of approximately \( 768 \times 384 \) in longitude and latitude, or \( 0.47 \)° of longitude or latitude per grid cell. These latter models are among the highest-resolution models of tidally locked giant planets that have ever been performed.\(^5\)

\(^5\) Liu & Showman (2013) showed, for a canonical hot Jupiter with a rotation period of 3.5 days and a forcing setup similar to that here, that the statistically equilibrated final state obtained using a horizontal resolution of \( C_{32} \) is almost identical to that obtained at a much higher resolution of \( C_{128} \), implying that \( C_{32} \) is adequate for understanding many aspects of the system under the slowly rotating conditions of canonical hot Jupiters. Menou (2020) also showed that for canonical hot Jupiters, modest resolution is sufficient to resolve the main dynamical structure. However, \( C_{32} \) would be insufficient for the faster-rotation models considered in this paper.
All models are integrated to a statistical steady state. The initial state is at rest with a globally uniform temperature profile set to the Iro et al. (2005) profile. Note that the equilibrated final state is insensitive to initial conditions (Liu & Showman 2013).

3. Results

We performed simulations in a grid with varying rotation periods of 80, 40, 20, 10, 5, and 2.5 hr and radiative timescales at the top $\tau_{\text{rad, top}} = 10^2$, $10^3$, $10^6$, and $10^7$ s. Simulations in this grid are drag-free at pressure lower than 10 bars, and a basal drag at pressures higher than 10 bars is included. Although the $\tau_{\text{rad, top}}$ of some models are higher than expected for real objects, understanding dynamics in the high-$\tau_{\text{rad}}$ regime is fruitful to obtain a full dynamical picture, as has been shown by previous studies of canonical hot Jupiters (Perez-Becker & Showman 2013; Komacek & Showman 2016; Komacek et al. 2017). Our grid connects to that in Komacek & Showman (2016), in which they explored the circulation with a different drag strength and radiative timescale but a fixed rotation period appropriate for canonical hot Jupiters.

Snapshots of a horizontal temperature structure with superposed wind vectors for models in the grid are shown in Figure 2 at a pressure of 7 mbars, Figure 3 at 80 mbars, and Figure 4 at 1.3 bars. The time-averaged zonal-mean zonal wind structures are shown in Figure 5. In these plots, the radiative timescale increases from left to right, and the rotation period decreases from top to bottom. We first describe the general outcomes of this grid. Circulation in the canonical hot Jupiter regime (those with rotation periods of 80 and 40 hr) has been extensively discussed in Komacek & Showman (2016), and we focus more on the dynamics of the rapidly rotating, short radiative timescale models, as they are in a novel dynamical regime and relevant for light-curve observations of the WD+BD systems. Some simulations with spatially independent
drag will also be discussed. Finally, we present synthetic phase curves for models with a short radiative timescale.

### 3.1. Circulation with Short Radiative Timescales \( \tau_{\text{rad, top}} = 10^4 \) and \( 10^5 \) s

In the observable atmosphere (pressures less than a few hundred mbars), models with a short radiative time constant exhibit a global-scale, large-amplitude day–night temperature difference across a wide range of rotation periods. In general, the characteristic day–night temperature difference increases with decreasing rotation period, a trend clearly evident in the left two columns of Figures 2 and 3. At \( \tau_{\text{rad, top}} = 10^4 \) s, the radiative timescale is so short at low pressure that the dynamics are inefficient at redistributing heat from day to night, and the temperature structure is not far from radiative equilibrium at all rotation rates explored. Thus, at 7 mbars, the day–night temperature difference increases only mildly with rotation rate (left column of Figure 2). Interestingly, though, the trend is much more drastic at 80 mbars (left column of Figure 3). This occurs because the dynamics efficiently redistributes the heat from day to night when the rotation is slow but is suppressed from doing so when the rotation is fast, leading to a very strong dependence of the day–night temperature difference on rotation rate. For models with \( \tau_{\text{rad, top}} = 10^5 \) s, the systematic increase of the day–night temperature difference with decreasing rotation period is strong throughout the observable atmosphere and prominent at both 7 and 80 mbars in Figures 2 and 3. In this case, the longer radiative timescale permits significant dynamical readjustment of the temperature structure when rotation is slow, but this dynamical readjustment mechanism becomes much less efficient when rotation is fast.

The day–night temperature difference becomes visibly weaker or even indistinguishable at pressures exceeding 1 bar, where the radiative timescale is longer and the day–night forcing is weaker. The global circulation is instead dominated by a zonally symmetric configuration in which eddies of various horizontal length scales are embedded. The temperature structure at 1.3 bars is shown as an example in...
Figure 4. The tendency of the deep circulation to become increasingly zonally banded is not surprising, since when the radiative timescale becomes longer than the relevant dynamical timescales at depth, latitude-dependent dynamical processes such as interactions between Rossby waves and turbulence could overcome the direct day–night forcing and drive a zonally banded circulation (e.g., Showman et al. 2013, 2015).

With decreasing rotation period, the global circulation transitions from the canonical hot Jupiter regime to a rotation-dominated, geostrophic regime. At relatively low pressure, the global-scale pattern of slow rotators with 40 and 80 hr rotation periods is dominated by a broad equatorial eastward jet and the eastward-shifted hot spot. The standing Rossby waves extend all the way to the poles and are Doppler-shifted toward the east by the strong equatorial jet. This regime has been extensively explored by previous work (see a review by, e.g., Heng & Showman 2015). In the case with a 20 hr rotation period, the off-equatorial Rossby waves are more robustly developed than those with 80 and 40 hr rotation periods. As the rotation period decreases, the equatorial deformation radius becomes smaller and—for rotation periods less than about 10 hr—is only a small fraction of the planetary radius. The Matsuno–Gill wave pattern excited by the day–night forcing is thus confined closer to the equator, and the poleward edges of the standing Rossby waves shift to lower latitudes as the rotation period decreases. Poleward of the Matsuno–Gill pattern, horizontal winds are much weaker than those at low latitudes, and wind vectors follow the isotherms more closely (Figures 2 and 3). The Rossby number at mid-to-high latitude is much less than 1 over the global scale, corresponding to the geostrophic circulation regime. At pressure levels where slowly rotating models exhibit a systematic eastward offset of the global-scale dayside hot regions, this offset visibly decreases with decreasing rotation period. This is particularly evident at $\tau_{\text{rad,top}} = 10^5$ s in Figure 2 and both $\tau_{\text{rad,top}} = 10^4$ and $10^5$ s in Figure 3.

Off the equator, Rossby numbers can be as small as a few percent in our rapidly rotating models. For example, in our
most rapidly rotating model with a 2.5 hr period, the rotation rate is $\Omega = 7 \times 10^{-4} \text{ s}^{-1}$, implying a typical value of the Coriolis parameter in midlatitudes of $f \approx 10^{-3} \text{ s}^{-1}$. Adopting a wind speed of $200 \text{ m s}^{-1}$ and a typical meridional length scale of $10^7 \text{ m}$ yields $Ro = U / fL \approx 0.02$.

Interestingly, outside the equatorial regions, the zonal-mean temperature of rapid rotators oscillates nonmonotonically as a function of latitude. This “fingering” phenomenon is prominent on the dayside at 7 and 80 mbars, where zonally aligned tongues of colder and warmer air interleave as a function of latitude (Figures 2 and 3). It is also prominent on both the dayside and nightside at deeper levels of $\sim$1 bar (Figure 4). The meridional extent of these zonal bands decreases, and the number of zonal bands increases, with decreasing rotation period. The meridional positioning of the zonal bands is consistent between high and low pressures.

These zonal temperature bands are accompanied by numerous zonal jets, alternating in sign as a function of latitude (see panels in lower left corner of Figure 5). As can be seen from Figures 2–4, the off-equatorial eastward jets are positioned where the zonal-mean temperature decreases poleward, while the westward jets are positioned where the zonal-mean temperature increases poleward. This relationship implies, through thermal-wind balance, that the zonal jets gradually decay with increasing pressure. Indeed, just such behavior can be seen in Figure 5.

Interestingly, this nonmonotonic zonal-mean temperature structure of our rapidly rotating models differs drastically from Earth’s troposphere and many GCMs for rapidly rotating giant planets (e.g., Williams 2003; Lian & Showman 2008; Schneider & Liu 2009), all of which are under zonally symmetric, equator-to-pole thermal forcing and exhibit a monotonic decrease of zonal-mean temperature (on isobars) with increasing latitude. The nonmonotonic temperature structure seen in our models likely results from the presence of day–night thermal forcing. One possibility is that the numerous off-equatorial eastward and westward zonal jets induce the

Figure 5. Time-averaged zonal-mean zonal wind as a function of latitude and pressure of models over the grid with a different rotation period in the vertical direction and a different radiative timescale $\tau_{\text{rad, top}}$ in the horizontal direction. All models are drag-free at pressures less than 10 bars.
temperature fingering by advecting cold nightside air onto the dayside in meridionally localized tongues of air (a mechanism that cannot occur when the forcing is zonally symmetric, as it is in many Jupiter models). Still, the zonal advection timescale for this process is several $\times 10^5$ s, which may be too long—relative to the radiative time constant—to maintain these temperature variations. Another possibility is vertical motion associated with meridional circulations occurring on the small meridional scale of the off-equatorial zonal jets; such motions would advect entropy vertically and induce horizontal temperature contrasts. Indeed, this mechanism would seem to be preferred at pressures exceeding $\sim 1$ bar, where day–night temperature differences are minimal (Figure 4).

There exist interesting off-equatorial coherent structures and small-scale eddies in rapid rotators with a rotation period of 10 hr or less. Eddies with a small horizontal length scale propagate westward at mid-to-high latitudes, an effect that is especially prominent at 1.3 bars. These eddies have a long vertical wavelength and affect the temperature structure at pressures lower than 80 mbars, which is partly responsible for the hot “tails” west of the dayside meridionally stacked structures there. The off-equatorial small-scale eddies likely originate from baroclinic instability, triggering Rossby waves that then propagate westward on the zonal jets. In the quasi-geostrophic limit, one of the necessary conditions for baroclinic instability to occur is that the meridional gradient of the zonal-mean potential vorticity (PV) changes sign in the vertical direction (e.g., Pierrehumbert & Swenson 1995; Chapter 6 of Vallis 2006). The statistically equilibrated zonal-mean configurations of our rapidly rotating models satisfy such a condition. In Figure 6, we show the zonal-mean PV as a function of latitude at pressures of 80 mbars and 0.93 bar for the model with a rotation period of 2.5 hr and $\tau_{\text{rad, top}} = 10^4$ s. The full hydrostatic PV in pressure coordinates shown in Figure 6 is written as (e.g., Vallis 2006)

$$Q = -g \left( f + \zeta \right) \frac{\partial \theta}{\partial p} - \frac{\partial v}{\partial p} \frac{\partial \theta}{\partial x} + \frac{\partial u}{\partial p} \frac{\partial \theta}{\partial y},$$

where $\zeta = \nabla \times \mathbf{v}$ is the vertical component of the relative vorticity, $\mathbf{v}$ is the isobaric velocity vector, and $\theta$ is the potential temperature. We have estimated that the Rossby number at midlatitudes is about $R \sim 0.02$, and thus the quasi-geostrophic limit holds well at midlatitudes. Poleward of about $\pm 20^\circ$, there are regions wherein the magnitude of PV first increases then decreases poleward (see regions in between the eastward jet cores marked by the dashed lines). However, at 0.93 bar, the magnitude of PV monotonically increases poleward in most latitudes except regions within $\pm 10^\circ$. This results in a change of sign for $\partial Q/\partial y$ in the vertical direction between these layers, which likely promotes the baroclinic instability at midlatitudes.

At low latitudes, transient eddies are also active and exhibit a wide range of horizontal length scales. The most obvious feature would be the meandering of the equatorial jets, as visibly seen in the temperature maps. All models with $\tau_{\text{rad, top}} = 10^4$ and $10^5$ s show more or less meandering of their equatorial jets. Even smaller-scale eddies exhibiting high-frequency oscillations are present on the flanks of the equatorial jet, but they do not play a major role in maintaining the circulation. The PV structure near the flanks of the equatorial superrotating jet shown in Figure 6 exhibits a strong poleward decrease in amplitude (in regions between about $2^\circ$ and $7^\circ$ and between $-2^\circ$ and $-7^\circ$) at both 80 mbars and 0.93 bar, indicating a violation of the barotropic stability criteria $\partial^2 \Pi/\partial y^2 < 0$ (e.g., Ingersoll & Pollard 1982). This suggests that the equatorial jet is barotropically unstable in the model with a rotation period of 2.5 hr, which is likely responsible for the meandering of the jet and small-scale transient eddies at the flanks of the equatorial jet. Whether barotropic instability would occur in hot Jupiter atmospheres and its effect on limiting the jet speed have been discussed in the literature (Menou & Rauscher 2009; Heng et al. 2011b; Fromang et al. 2016; Menou 2020). All of our models with $\tau_{\text{rad, top}} = 10^4$ s strongly violate the barotropic stability criterion, and apparently barotropic instability occurs in these models. Models with $\tau_{\text{rad, top}} = 10^5$ s also slightly violate the barotropic stability criterion. However, such instabilities may not be a critical factor in limiting the equatorial jet speed in our models; otherwise, the instability would efficiently remove the kinetic energy associated with the strong wind shear by forcing the jet profile to be neutral to barotropic stability.

The transition of the dynamical regime with respect to rotation impacts the phase-curve amplitude and offset. With decreasing rotation period, the global day–night temperature difference visibly increases, and straightforwardly, the day–night flux difference is expected to increase. For a relatively short rotation period at a low latitude within the Matsuno–Gill pattern, the hottest area is always east of the substellar point due to the strong equatorial superrotating jet and the equatorial Kelvin wave. The off-equatorial anticyclonic Rossby gyres are west of the substellar point and are associated hot regions. As the rotation period decreases, the projected area of the Rossby gyres increases and more thermal flux from west of the substellar point contributes to the disk-integrated flux, while the meridional extent of the equatorial region decreases.
combination of the two reduces or even reverses the phase-curve offset that is typically observed to peak before the secondary eclipse for canonical hot Jupiters. In addition, the off-equatorial Rossby gyres can sometimes be much hotter than the equatorial region and exhibit fluctuations, and there exist hot tails west of the Rossby gyres that can extend to the nightside. We expect that these features will influence the phase-curve offset and induce time variability. Synthetic phase curves will be shown in Section 3.4.

We now discuss the structure of the zonal jets, the relation to the relevant dynamical length scales, and the jet-driving mechanism in our models. We start off with the equatorial superrotating jets. In Section 1.2 we summarized the mechanism that drives the equatorial superrotation on tidally locked exoplanets, namely, the interactions of the standing tropical waves with the mean flow, and this theory predicts that the equatorial jet has a half-width approximately equal to the equatorial deformation radius, \( \sqrt{NH_0/\Omega} \) (Showman & Polvani 2011). We measure the half-widths of the equatorial jets in our GCM simulations to compare them to this scaling theory. The measured half-widths of the equatorial superrotating jets from GCMs scale quite well with our analytic prediction, showing a close correlation with \( \Omega^{-1/2} \). The half-width of the equatorial jet is defined as the distance between the equator and where the jet profile first reaches a minimum. The jet profiles used here are vertically averaged between 1 and 10\(^{−3}\) bars. In the top panel of Figure 7, we show the jet half-widths measured from GCMs with \( \tau_{\text{rad,top}} = 10^4 \) and 10\(^{−1}\) s and comparisons to the equatorial deformation radius (scaled up by a constant 1.8) as a function of rotation period assuming an isothermal atmosphere with \( T \sim 1400 \) K. For models with \( \tau_{\text{rad,top}} = 10^4 \) s, the scaling is particularly good for rotation periods \( \sim 40 \) hr. For rotation periods \( \sim 40 \) hr, the measured jet half-width flattens out with increasing rotation period. For models with \( \tau_{\text{rad,top}} = 10^5 \) s, the flattening starts at 20 hr. The mismatch at a long rotation period is partly because the wave theory and the corresponding scaling (Showman & Polvani 2011) are established on the equatorial \( \beta \) plane, a local approximation that could overestimate the jet width when the waves and wave-mean flow interactions can take place at mid-to-high latitudes where the full spherical geometry should be taken into account. In the case of slow rotators, the deformation radius is comparable to the planetary radius, and thus the accuracy of our scaling is affected. It is common in practice that GCM simulated jets differ from the theoretical scaling by an O(1) constant (e.g., Schneider & Walker 2006; Schneider & Liu 2009). In reality, several factors may also affect this comparison. First, the scaling is based on a theory that is linearized about a rest state, and our measured jet width is based on the fully equilibrated, nonlinear calculations. Second, the wave structure and wave-mean flow interactions would be altered by the presence of a strong equatorial jet (Tsi et al. 2014; Hammond & Pierrehumbert 2018). Third, the nonlinear effects may not be precisely quantified as the frictional drag. Finally, the dynamics of a continuously stratified atmosphere is considerably more complicated than that of the shallow-water system. Given all the complexities, the scaling of the width of the equatorial superrotating jet works surprisingly well in explaining our GCM results and supports a uniform mechanism that drives the equatorial superrotation from slow-rotating regime to rapidly rotating regime.

Now we look at the off-equatorial jets. It is widely accepted that the off-equatorial jet spacing in simulations of rapidly rotating, thin atmospheres subject to small-scale turbulent forcing is typically close to the Rhines scale (see a recent review by Showman et al. 2019), a scale over which two-dimensional turbulence interaction becomes anisotropic due to planetary rotation (Rhines 1975; Vallis & Maltrud 1993). Here we compare the measured off-equatorial jet spacing to the Rhines scale in the middle panel of Figure 7, which shows the ratio between the Rhines scale and jet spacing as a function of the Coriolis parameter \( f \) at jet cores. The measured jet spacing is based on jet profiles vertically averaged between 1 and 10\(^{−3}\) bars for GCMs with \( \tau_{\text{rad,top}} = 10^4 \) and 10\(^5\) s. Middle panel: the ratio between the Rhines scale defined using vertically mass-weighted eddy velocity and the off-equatorial jet spacing (circles) and the ratio between the Rhines scale defined using eastward jet velocity and the off-equatorial jet spacing (triangles). These ratios are shown as a function of the Coriolis parameter \( f \) at jet cores. The measured jet spacing is based on jet profiles vertically averaged between 1 and 10\(^{−3}\) bars for GCMs with \( \tau_{\text{rad,top}} = 10^4 \) and 10\(^5\) s. Bottom panel: ratio between the extratropical deformation radius (assuming an isothermal atmosphere and \( T \sim 1400 \) K) and the off-equatorial jet spacing.
are two flavors of the Rhines scale that are commonly used in the literature. The first one is expressed using a characteristic eddy velocity \( L_R \sim 2\pi U_{\text{edd}}/f \) (e.g., Rhines 1975; Schneider & Walker 2006; Chemke & Kaspi 2015), where \( U_{\text{edd}} \) is a vertically mass-weighted eddy velocity between 1 and \( 10^{-3} \) bars following Schneider & Liu (2009). The other one is expressed using a characteristic jet velocity \( L_R \sim \pi \sqrt{2U_{\text{jet}}/f} \) (e.g., Williams 1978; Lian & Showman 2008), where \( U_{\text{jet}} \) is the vertically averaged eastward jet core speed. The eddy Rhines scale ratios are represented as circles in the middle panel of Figure 7, and those scaled using zonal-mean jet speed spacing scales well with the Rhines scale. The Rhines scales based on both the eddy and jet velocities are equally well fitted to our jet spacing. This is perhaps because at the mid-to-high latitudes of our rapid rotators, kinetic energy associated with eddies is comparable to that associated with zonal jets. It is interesting to note that kinetic energy near the equator is dominated by the strong equatorial superrotating jet, which contains more than 95% of the total kinetic energy, and this is the case for all models with a different rotation period and \( \tau_{\text{rad, top}} = 10^4 \) and \( 10^5 \) s.

The extratropical deformation radius \( L_D \sim 2\pi c/f \), where \( c \) is an appropriate gravity phase speed, is a relevant energy injection scale, especially for jets that are driven by baroclinic instability. The bottom panel in Figure 7 shows the ratio between the extratropical deformation radius (with \( c \sim 2NH \) assuming an isothermal atmosphere with \( T \sim 1400 \) K) and the jet spacing as a function of \( f \). Likewise, the deformation radius also exhibits relatively good agreement with the jet spacing, with a maximum deviation of about 1.3. There is little scale separation between the Rhines scale and the deformation radius shown in Figure 7, suggesting that nonlinear eddy–eddy interactions are weak in our fast-rotating models, similar to the extratropical regime in Earth’s atmosphere (e.g., Held & Larichev 1996; Schneider & Walker 2006). The zonal jets are maintained mainly by interactions with eddies generated from baroclinic instability and day–night forcing, and nonlinear eddy–eddy interactions appear not to be essential in determining the jet structure. There is an emerging trend that \( L_D/j_c \) is larger than 1 at small \( f \) and systematically decreases to below 1 at large \( f \). This trend reflects a tendency of scale separation as a function of the local Coriolis parameter, which has been documented in previous studies (e.g., Chemke & Kaspi 2015).

Comparison of the zonal jets between models with \( \tau_{\text{rad, top}} = 10^4 \) and \( 10^5 \) s (Figure 5) is interesting. The spacing and strength of the off-equatorial zonal jets, as well as the associated zonal-mean meridional temperature gradients, are quantitatively similar between models with \( \tau_{\text{rad, top}} = 10^4 \) and \( 10^5 \) s. In contrast, for the equatorial jets, despite their similar jet width, the jet speed with \( \tau_{\text{rad, top}} = 10^5 \) s is almost half that with \( \tau_{\text{rad, top}} = 10^4 \) s. The fact that, given a fixed \( \tau_{\text{rad, top}} \), the equatorial jet speed is almost the same among models with different rotation rates is also worth noticing. Recall that frictional drag is absent at pressures less than 10 bars in models shown in Figure 5. The above comparisons indicate that, in the relatively short-\( \tau_{\text{rad, top}} \) regime, the scaling of the off-equatorial jet strength with rotation rate differs from that for the equatorial jet. No theory currently exists to predict the equatorial jet speed of tidally locked planets, and theories predicting the off-equatorial jet speed often apply only under conditions of zonally symmetric, equator-to-pole forcing. Zhang & Showman (2017) showed that the scaling of eddy velocity using the theory of Komacek & Showman (2016) provides a good match to the equatorial jet speed, indicating the intimate relation of eddy fluxes and the equatorial jet speed. These issues are challenging, and we do not attempt to resolve them here.

Finally, we briefly present diagnostics that further illustrate the role of eddies in maintaining the zonal-mean circulation and the transition from slow to rapid rotation. The zonal-mean zonal momentum equation of primitive equations in pressure coordinates can be written as

\[
\frac{\partial \pi}{\partial t} = \tau_f - \frac{v}{a \cos \phi} \frac{\partial (\pi \cos \phi)}{\partial \phi} - \omega \frac{\partial \pi}{\partial p} - \frac{1}{a \cos^2 \phi} \frac{\partial (u'v' \cos^2 \phi)}{\partial \phi} - \frac{\partial (u'\omega')}{\partial p},
\]

where \( f \) is the Coriolis parameter; \( \phi \) is latitude; \( \pi, v, \) and \( \omega \) are zonal-mean zonal, meridional, and vertical velocities in pressure coordinates; and \( u', v', \) and \( \omega' \) are their deviations from the zonal average, respectively. The terms on the right-hand side represent Coriolis acceleration associated with the mean meridional flow, meridional momentum advection by the mean flow, vertical momentum advection by the mean flow, horizontal convergence of eddy momentum, and vertical convergence of eddy momentum.

Accumulation on the zonal-mean zonal wind by various terms on the right-hand side of Equation (3) for a slowly rotating model with a rotation period of 40 hr and \( \tau_{\text{rad, top}} = 10^4 \) s are shown in the left column of Figure 8. In agreement with previous hot Jupiter studies (e.g., Showman & Polvani 2011; Showman et al. 2015; Mayne et al. 2017), the standing Matsuno–Gill pattern driven by the day–night thermal forcing generates horizontal eddy convergence of momentum that is responsible for maintaining the equatorial superrotation and mainly balanced by vertical eddy convergence at low latitudes. The acceleration is strongest near the flanks of the equatorial jet rather than at the center of the jet, which is slightly different from that in the spin-up phase (see, for example, Showman & Polvani 2011; Showman et al. 2015). This is presumably caused by modification of the standing wave pattern by the strong equatorial jet and has been similarly shown in the GCM diagnostics of Mayne et al. (2017). At high latitudes, the Coriolis acceleration associated with the mean meridional circulation becomes important in the balance of zonal angular momentum. Overall, the momentum advection by mean flows is not dominant at all latitudes and pressures.

The diagnostics of the rapidly rotating model with a rotation period of 2.5 hr and \( \tau_{\text{rad, top}} = 10^4 \) s are shown in the right column of Figure 8. Near the equator (within \( \pm 10^5 \)), despite being confined at much lower latitudes, excitation of the Matsuno–Gill pattern by the day–night forcing is responsible for maintaining the equatorial superrotation, similar to that in the slowly rotating model. At mid-to-high latitudes (poleward of \( \pm 10^5 \)), the maintenance of subtropical eastward jets is due to
the horizontal eddy convergence of momentum by the baroclinic eddies (see positive accelerations by horizontal eddy transport at about ±30°, ±45°, and ±60° in panel (f), which correlates well to the eastward jets shown in Figures 5 and 6). However, the major balance in the zonal momentum at mid-to-high latitudes is between the Coriolis acceleration and horizontal eddy convergence (see a comparison between panels (f) and (h) in Figure 8), namely,

\[ f \sigma \approx \frac{1}{a \cos \phi} \frac{\partial (u'v' \cos^2 \phi)}{\partial \phi}. \] (4)

This is consistent with the fact that the dynamics is quasi-geostrophic at mid-to-high latitudes, and the above balance in the statistically steady state is well known in this regime (e.g., Schneider 2006; Vallis 2006; Holton & Hakim 2012).

The zonal-mean stream function \( \Psi \), defined by \( \sigma = g (2\pi a \cos \phi)^{-1} \partial \Psi / \partial \rho \) and \( \varpi = -g (2\pi a \cos \phi)^{-1} \partial \Psi / \partial \phi \), where \( a \) is the planetary radius, is shown in the bottom panels of Figure 8 for the slow-rotating model with a rotation period of 40 hr and \( \tau_{\text{rad,}0} = 10^4 \) s in the left column and the model with a rotation period of 2.5 hr and \( \tau_{\text{rad,}0} = 10^4 \) s in the right column. Panel (a) shows the horizontal eddy acceleration (the term \( \frac{1}{a \cos \phi} \frac{\partial (u' v' \cos^2 \phi)}{\partial \rho} \) in Equation (3)) on the zonal-mean zonal wind, panel (b) shows the vertical eddy acceleration \( -\frac{\partial \Psi}{\partial \rho} \) on the zonal-mean zonal wind, panel (c) shows the Coriolis acceleration associated with the zonal-mean meridional circulation \( (\Psi) \), panel (d) shows the acceleration by zonal-mean momentum transport \( \frac{1}{a \cos \phi} \frac{\partial (u' v' \cos^2 \phi)}{\partial \rho} \) on the zonal-mean zonal wind, and panel (e) shows the zonal-mean stream function. In the bottom panels, red lines are clockwise with log-spaced contours between \( 3 \times 10^8 \) and \( 3 \times 10^{10} \) kg s\(^{-1}\), and blue lines are counterclockwise with log-spaced contours between \( -3 \times 10^8 \) and \( -3 \times 10^{10} \) kg s\(^{-1}\). Panels (f)-(j) are similar but for the model with a rotation period of 2.5 hr and \( \tau_{\text{rad,}0} = 10^4 \) s. of Figure 8 for the slow-rotating model with a rotation period of 40 hr on the left and the rapidly rotating model with a rotation period of 2.5 hr on the right. Red lines are clockwise, and blue lines are counterclockwise. There are mainly two zonal-mean meridional circulation cells above 0.5 bar in each hemisphere. The air associated with the mean meridional circulation descends at the equator and ascends near ±40°, indicating a thermally indirect zonal-mean meridional circulation near the equator. Interestingly, there are two thermally direct circulation cells within about ±10° at a deeper pressure than the thermally indirect cells. The zonal-mean stream function of the rapid rotator (panel (j)) exhibits up to nine cells in each hemisphere, which is not surprising given the multiple zonal banded structure in the simulation. At low latitudes, air associated with the mean meridional circulation ascends at about ±10° and
descends near the equator and is similarly a thermally indirect circulation. Similar to the slow-rotating case, there are small thermally direct cells within about ±3° at a pressure deeper than 50 mbars. At mid-to-high latitudes, as indicated by the balance in Equation (4), the mean meridional circulation is eddy-driven, and the strength is constrained by magnitude of the eddy momentum convergence. Indeed, the position and strength of the off-equatorial cells correlate well with the horizontal eddy acceleration shown in panel (f). The eddy-driven circulation at mid-to-high latitudes is similar to the picture that has been proposed for the off-equatorial jets on Jupiter and Saturn near the cloud level (Del Genio et al. 2007; Showman 2007; Schneider & Liu 2009).

3.2. Circulation with Long Radiative Timescale

Models with long radiative timescales ($\tau_{\text{rad,top}} = 10^6$ and $10^7$ s) show qualitatively different horizontal structure from those with short radiative timescales, analogous to the behavior found in Perez-Becker & Showman (2013) and Komacek & Showman (2016). As straightforwardly seen in Figures 2 and 3, given any rotation period, the day–night temperature differences at low pressure are much smaller in models with long radiative timescales than those with short radiative timescales. In slow rotators with 40 and 80 hr rotation period, the horizontal temperature structure is rather zonal even at low pressure. Surprisingly, in models with a rotation period of 20 hr or less, a global-scale wave pattern can still be visible at 7 and 80 mbars, albeit with much weaker winds and temperature perturbations than those with shorter $\tau_{\text{rad,top}}$. For models with 5 and 2.5 hr rotation, the wave patterns are still vaguely seen even at 1.3 bars. For rapid rotators, in contrast to those with $\tau_{\text{rad,top}} = 10^4$ and $10^5$ s, the circulation with long $\tau_{\text{rad,top}}$ is characterized overall by meridionally broad wave structures, showing no obvious multiple zonal banded structures. It is possible that with a long radiative timescale, global-scale waves are the major response to the day–night forcing, and the waves significantly redistribute heat around the globe, eliminating the large equator-to-pole temperature difference. As a result, instability and small-scale eddies do not easily occur in these models.

Equatorial superrotation exists in all models with $\tau_{\text{rad,top}} = 10^6$ s, although with weaker wind speeds than those with $\tau_{\text{rad,top}} = 10^5$ s (see Figure 5). In rapid rotators, the depth of the equatorial jets at $\tau_{\text{rad,top}} = 10^6$ s is shallower than that with $\tau_{\text{rad,top}} = 10^4$ and $10^5$ s by more than a scale height; the off-equatorial zonal flows are much weaker, and their meridional shapes are less well developed than those with $\tau_{\text{rad,top}} = 10^4$ and $10^5$ s. These features are not surprising because in general, the magnitude of eddies gets weaker with longer $\tau_{\text{rad,top}}$, and the zonal jets driven by them become weaker. Interestingly, for models with 40 and 80 hr rotation, although there is equatorial superrotation, the jet profile actually peaks off the equator, indicating a stronger role of meridional circulation in shaping the overall jet structure than that with shorter $\tau_{\text{rad,top}}$. For models with $\tau_{\text{rad,top}} = 10^5$ s, slow rotators exhibit qualitatively similar jet structures to those with $\tau_{\text{rad,top}} = 10^6$ s. However, models with a 20 hr or less rotation period have little or no equatorial superrotation, and their overall circulation is yet much weaker than those with $\tau_{\text{rad,top}} = 10^6$ s.

We summarize the results over the entire grid by discussing the fractional day–night temperature differences. The day–night temperature difference is mainly regulated by wave adjustment. This process competes with radiative damping, frictional drag damping, and planetary rotation (Perez-Becker & Showman 2013; Komacek & Showman 2016). The fractional day–night temperature differences quantified using the nondimensional quantity $A = \Delta T/\Delta T_{\text{eq}}$ at 80 mbars (as defined in Komacek & Showman 2016) for drag-free models with a different rotation period and radiative timescale $\tau_{\text{rad,top}}$. Solid lines are predictions using the scaling theory of Komacek & Showman (2016) and Zhang & Showman (2017). Here $A \sim 0$ represents a homogeneous longitudinal temperature distribution, and $A \sim 1$ represents a temperature distribution close to radiative equilibrium. Overall, the day–night temperature difference increases with decreasing rotation period and radiative timescale. A long radiative timescale results in a negligible day–night temperature difference despite a short rotation period.

The scaling theory captures the overall trends of the GCM results, consistent with the findings in Komacek & Showman (2016) and Zhang & Showman (2017), but systematically predicts lower fractional day–night temperature differences than that from our GCMs (Figure 9). As shown in Komacek & Showman (2016), the analytic theory works best in the strong-drag regime, which typically has a laminar day–night flow, but relatively less well in the drag-free regime, where the strong equatorial jet is driven. The Komacek & Showman (2016) scaling theory has no explicit representation of the equatorial jet or the Matsuno–Gill standing wave pattern; the theory assumes that the wind flows straight from day to night and that the temperature structure comprises a simple day–night difference, which are assumptions that (for a given wind speed and day–night temperature difference) maximize the day–night heat transport. In reality, Figures 2 and 3 make clear that winds outside the equatorial jet often follow nearly parallel to
isotherms (this is true not only for very rapid rotators but also to some degree even for periods of 10, 20, and 40 hr). Only the component of wind perpendicular to isotherms will advect temperature, and this component is weaker than the total wind amplitude. This phenomenon suggests that, for a given rms wind speed and day–night temperature difference, the actual day–night heat transport will be smaller than predicted by the theory, and therefore the equilibrated day–night temperature difference will be larger than predicted by the theory. It may be possible to improve the theory to account for some of these effects, a task we leave for the future.

3.3. The Role of Frictional Drag

The thermal ionization fraction could be significant in hot atmospheres of extremely close-in gas giants in the sense that the atmospheric flows can be strongly coupled to the intrinsic magnetic field and experience significant magnetohydrodynamic drag. Some GCMS parameterize this dissipation as simple frictional drag and explore its consequences on the flows (e.g., Perna et al. 2010, 2012; Komacek & Showman 2016). There have been self-consistent magnetohydrodynamic models but with simplified thermal forcing schemes (e.g., Rogers & Komacek 2014; Rogers 2017). Both models predict slower winds due to the Lorentz force. Other possible dissipation mechanisms include turbulent dissipation (e.g., Li & Goodman 2010). To crudely represent the effects of magnetic or turbulent effects on reducing the wind speed and shaping the circulation pattern, we include a pressure-independent frictional drag in a set of models with a 2.5 hr rotation period, \( \tau_{\text{drag, top}} = 10^4 \) s, and two different drag timescales, \( \tau_{\text{drag}} = 10^6 \) and \( 10^4 \) s.

Figure 10 shows horizontal temperature maps with overplotted winds at three different pressure levels for models with \( \tau_{\text{drag}} = 10^6 \) s on the left and \( \tau_{\text{drag}} = 10^4 \) s on the right. The overall circulation pattern of the models with \( \tau_{\text{drag}} = 10^6 \) s at a relatively low pressure qualitatively resembles that of the drag-free model, showing a strong, robust equatorial superrotation, characteristic Matsuno–Gill wave patterns at low latitudes, and nearly geostrophic flow at high latitudes. The magnitude of wind speed is lower than in our drag-free models due to the drag. Small-scale vortices and waves are visibly seen at both low and high latitudes, which are similarly caused by baroclinic and barotropic instability. A quantitative difference is that there is no obvious zonal banding even at high pressure (1.3 bars shown in Figure 10), despite the existence of eddies.

The drag timescale is likely shorter than the dynamical timescale on which eddies can accelerate the zonal jets and over which turbulence can reorganize into a zonalized configuration. Thus, when the drag timescale is sufficiently short, it plays a critical role in preventing off-equatorial jet formation. These results are consistent with those of Perna et al. (2010), Showman et al. (2013), and others, who described the role drag can play in suppressing jet formation in the deeper layers of the troposphere.

In the model with \( \tau_{\text{drag}} = 10^4 \) s, the drag timescale is much closer to the rotation timescale; thus, the overall horizontal force balance transitions from the geostrophic regime to a three-way balance regime. At mid-to-high latitude, this can be recognized due to the fact that the wind vectors do not necessarily follow isotherms on isobaric surfaces (right column in Figure 10). The equatorial jet is suppressed. The off-equatorial Rossby waves are still seen but much weakened. The pattern is again quite similar to predictions made by using linearized analytic models. With an even stronger drag, \( \tau_{\text{drag}} = 10^5 \) s (not shown), the circulation pattern is dominated by the pure day-to-night flow, with a weak flow speed and a temperature pattern nearly in radiative equilibrium. In this case, the drag timescale is shorter than any other relevant dynamical timescales (especially the rotation timescale), and the force balance is simply between pressure gradient and drag. The weak wind speeds allow only relatively weak day–night heat transport, helping to preserve the temperature structure close to radiative equilibrium. The existence of strong drag has important consequences for the thermal phase curve due to the influence on the horizontal temperature structure. The phase curve at different wave-lengths can provide a diagnostic tool to discriminate the extent to which the drag shapes the circulation. Still, Lorentz forces will have a far more complex pattern than allowed by our simple Rayleigh-drag scheme, so these strong-drag models should be viewed more as experiments to understand how the circulation responds to idealized drag of various strengths, rather than as predictions for how actual atmospheres affected by Lorentz forces or strong turbulent mixing will behave.

3.4. Synthetic Phase Curves

The change of global temperature structure with decreasing rotation period has major implications for thermal phase curves. We show normalized synthetic thermal phase curves in Figure 11 sampled from temperatures at different pressures for models with a different rotation period, a short radiative timescale \( \tau_{\text{rad, top}} = 10^4 \) s, and without frictional drag, using the following relation (Zhang & Showman 2017):

\[
F(\delta, \phi) = \int_{\pi/2 - \delta}^{\pi/2 + \delta} d\lambda \int_{-\pi/2}^{\pi/2} d\phi \sigma T^4(\lambda, \phi, p) R_p^2 \cos^2 \phi \cos \lambda d\phi. \tag{5}
\]

Here we assume that the orbit of the BD is aligned with the line of sight of the observer (i.e., we assume the BD is a transiting object), which maximizes the day–night flux contrast. The temperature structure used for the phase curves shown in Figure 11 is time-averaged and static as the phase rotates. A few features are obvious. First, at the same pressure level, the amplitude of the phase curves increases with decreasing rotation period, consistent with the aforementioned increasing day–night temperature contrast with decreasing rotation period. Second, with the same rotation period, the phase-curve amplitude is generally greater for phase curves sampling lower pressure, due to the larger day–night temperature differences that tend to occur at lower pressure relative to higher pressure. Finally, for slow rotators, the thermal fluxes sampling several pressure levels shown in Figure 11 all peak before the secondary eclipse, which is attributed to the eastward-shifted equatorial hot spot. As the rotation period decreases, the flux peak moves closer to the secondary eclipse. At low pressures (80 and 7 mbars), the thermal flux peak is aligned with the secondary flux for a rotation period of less than 5 hr. At a slightly higher pressure (231 mbars), the flux peaks after the secondary eclipse for a rotation period of less than 10 hr. As we
have seen from the horizontal map in Figures 2 and 3, for faster rotators, the westward-shifted hot areas associated with the Rossby waves move closer to the equator, contributing more to the total flux. Meanwhile, the eastward-shifted hot spot at the equator shrinks, contributing less to the total flux. The two effects tend to compensate for each other, resulting in almost no shift between the peak flux and the secondary eclipse at low pressures. Note that similar effects on the phase curve have been shown in Carone et al. (2020), who simulated the rapidly rotating hot Jupiter WASP-43b with a rotation of 0.8 day. These results—showing that strong rotation suppresses the offset of the flux peak from the secondary eclipse—provide an explanation for the absence of a flux-peak offset in many phase curves of BD+WD systems (e.g., Casewell et al. 2015). Moreover, the possibility that in some cases the westward-shifted contribution can slightly dominate may explain why phase curves of some BD+WD systems appear to have a “westward-shifted hot spot” (for example, the Kepler broadband phase curve for WD 1202–024; Rappaport et al. 2017).

The contribution of spatial temperature variability causes time variability of the synthetic phase curves. In typical hot Jupiter simulations, such variability is typically small, around the 1% level (Showman et al. 2009; Fromang et al. 2016; Komacek & Showman 2019), and here we concentrate on the variability of the rapidly rotating models. Figure 12 shows the normalized dayside flux and phase offsets as a function of time for models with four different rotation periods but all with $\tau_{\text{rad, top}} = 10^4$ s and drag-free at pressures less than 10 bars. In each case, the upper panel shows the dayside flux at the secondary eclipse (normalized by the time-mean value) as a function of time for three pressure levels, and the lower panel shows the phase offsets of the peak thermal flux relative to the secondary eclipse. In general, the dayside flux emitted from 10 mbars varies by less than 1%, and the phase-curve offsets show almost no variation for all rotation periods. At deeper levels of 80 and 330 mbars, the peak-to-peak dayside flux variation can reach up to about 2.5%, or slightly more in cases with a 2.5, 10, or 20 hr rotation period and less than 1% in the 5 hr case. The phase-curve offsets at 80 mbars can reach a few

Figure 10. Snapshots of horizontal temperature maps with superposed wind vectors at 7 mbars in the first row, 80 mbars in the second row, and 1.3 bars in the third row for models with a 2.5 hr rotation period, $\tau_{\text{drag}} = 10^6$ s, and different frictional drag timescales $\tau_{\text{drag}} = 10^6$ (left column) and $10^4$ s (right column).
Figure 11. Normalized synthetic thermal phase curves sampling different pressure levels for drag-free models with a different rotation period from 80 to 2.5 hr and a short radiative timescale $t_{\text{rad}} = 10^4$ s. These phase curves are obtained by rotating the modeled atmosphere using time-mean temperature structures. Although the phase curves peak before secondary eclipse in the canonical hot Jupiter regime (black and dark blue curves), the maximum flux tends to occur right at secondary eclipse for very rapid rotators (light blue curves). This phenomenon can help explain the absence of such offsets in phase-curve observations of WD+BD systems (e.g., Casewell et al. 2015).

Figure 12. These panels show the time variability of the synthetic light curves in terms of the dayside flux at the secondary eclipse and the phase offsets of the maximum peak flux for four models with different rotation periods, as indicated by the titles above each panel. In each case, we show the evolution of the normalized dayside flux (normalized around the time-mean value) and phase offsets at three different pressure levels of 10, 80, and 330 mbars. These models are drag-free at pressures less than 10 bars.
degrees, and those at 330 mbars are even larger, which, in the case with a 20 hr rotation period, can reach more than 10° of peak-to-peak variation. The typical variability timescale ranges from several to more than 10 days. Interestingly, the dayside flux variability generally exhibits obvious phase offsets at 80 mbars relative to 330 mbars for all rotation periods. The off-equatorial propagating eddies are mainly responsible for the flux and phase-offset variability. To prove this, taking the 2.5 hr case as an example, we have compared synthetic phase curves from the full model, the model with the equatorial region (within ±10° latitude) replaced with the time-mean temperature, the model with the off-equatorial region (within ±[15° 30°] latitude) replaced with the time-mean temperature, and the model with the high-latitude region (poleward of ±30° latitude) replaced with the time-mean temperature. Phase curves from models with the equatorial and high-latitude regions replaced show very similar results to the full model, but phase curves from the model with only the midlatitude region replaced show much less variation, as well as a dissimilar evolutionary shape of the light curve. The phase difference of the dayside flux at different pressure levels is caused by the horizontal phase difference of the propagating waves at different pressures.

On the other hand, if the circulation suffers from a sufficiently strong drag, the phase-curve amplitude is generally larger than that of drag-free models, and there is little phase offset due to the suppression of strong equatorial jets and off-equatorial eddies. Comparisons of phase curves sampling different pressures for models with a 2.5 hr rotation period and various drag strength (no drag, $\tau_{\text{drag}} = 10^3$ and $10^4$ s) are shown in Figure 13. The effect of drag is small at 10 mbars but becomes increasingly important at larger pressures. In particular, for phase curves sampling 330 mbars, the day–night flux amplitude increases and the phase offset decreases with decreasing drag timescale. The dayside flux variability from the model with $\tau_{\text{drag}} = 10^3$ s is much less than 1%, and the phase-curve offset variation is negligible. The drag timescale is comparable to or shorter than the timescale that takes the waves in the drag-free model to propagate across the day–night distance, which apparently helps to suppress the temperature variation caused by these waves. Future phase-curve observations probing both low and high pressures would be helpful to disentangle the role of drag in shaping the circulation in these atmospheres.

4. Discussion

This study made one of the first efforts to understand the atmospheric circulation relevant to BDs in extremely tight orbits around WDs and thoroughly investigate the role of rotation in shaping the atmospheric circulation of tidally locked planetary companions. In a parallel study, Lee et al. (2020) simulated atmospheric circulation specifically for the BD orbiting WD 0137–349 with an orbital period of about 2 hr using a GCM coupled with a dual-band gray radiative transfer scheme. Their results are qualitatively similar to our most rapidly rotating case, showing a narrowing equatorial superrotating jet, nearly geostrophic flows at mid-to-high latitudes, and a large day–night temperature difference. Although using a different GCM and radiative forcing setup, the agreement between the two studies is encouraging. A direct implication of our study is to understand and interpolate the phase-curve observations of BDs in extremely close-in orbits around WDs. Because these atmospheres might be coupled to the intrinsic magnetic field, it is interesting to constrain the role of possible magnetic drag by using multiwavelength phase curves and long-time monitoring to measure the dayside flux variability. An alternative way may be the Doppler effect by winds on the high-resolution transmission spectroscopy, and this may be able to constrain whether the flow is predominantly day-to-night or zonal (Showman et al. 2013; Louden & Wheatley 2015; Brogi et al. 2016; Flowers et al. 2019).

Showman et al. (2015) also explored the effects of increasing rotation (but nonsynchronously rotating) in the circulation of hot Jupiters. Their fast-rotating models exhibit weak or no equatorial superrotation, while strong jets exist at high latitudes due to baroclinic instability. This is different from our models with short $\tau_{\text{rad,top}}$, in which strong equatorial superrotation is prevalent no matter how fast they rotate. This is partly because the forcing mechanisms are different. In the fast-rotating, nonsynchronous models of Showman et al. (2015), the rapid westward migration of the irradiated pattern may disrupt the formation of a standing Matsuno–Gill pattern, therefore weakening the driving force for the equatorial superrotation. In addition, the equator-to-pole temperature gradient as a result of nonsynchronous rotation generates baroclinic instability, and Rossby waves from high latitudes propagate to low latitudes and contribute to the westward acceleration on the zonal wind. Together, these imply much weaker (or even no) equatorial superrotation and strong high-latitude jets in the fast-rotation models of Showman et al. (2015), whereas in our models, the steady day–night forcing exists no matter how fast the model rotates. Therefore, there is always a standing Matsuno–Gill pattern that drives strong equatorial superrotating jets no matter how fast the model rotates.

The atmospheric temperature of BDs in tight orbits around WDs may be sufficiently high to thermally dissociate some fraction of the molecular hydrogen, especially at low pressure, where the gas absorbs the UV flux from the WD. The dynamical effects of hydrogen dissociation and recombination, including chemical heat release/absorption and change of mean molecular weight (Tan & Komacek 2019), would play an additional role in shaping the atmospheric circulation. The eddies associated with the instabilities could be more energetic, analogous to the effects of latent heat on baroclinic eddies in the midlatitudes of Earth’s atmosphere (e.g., Lapeyre & Held 2004). The size of the eddies might be sufficiently large and the associated temperature perturbations sufficiently high for these effects to have observable consequences.

Due to the inefficient day–night heat transport, the temperature structure on the nightside of BDs in tight orbits around WDs may be similar to that of field BDs, and optically thick silicate or iron clouds could form there. As was demonstrated in Tan & Showman (2019), the intrinsic variability driven by cloud radiative feedback may occur over the nightside of BDs and can induce high-amplitude temperature and cloud structure variation. The heating/cooling rate associated with the variability could even be comparable to that associated with the day–night forcing. Interactions between the global-scale, day–night driven dynamics and the small-scale intrinsic variability driven by cloud radiative feedback are still an exotic regime that deserves future exploration.

In this work, the circulation of tidally locked atmospheres with increasing rotation rate has been investigated in the
context of gas giants, i.e., fluid hydrogen planets with no solid surface and with a radius close to that of Jupiter. Still, our study also has important implications for the extremely close-in terrestrial planets observed to date. Understanding the climate dynamics of these will be complementary to those in or near the habitable zones around low-mass stars, which are synchronously rotating but with a rotation period of several more than 10 days (e.g., Merlis & Schneider 2010; Yang et al. 2014; Carone et al. 2015; Haqq-Misra et al. 2018). Several ultrashort-period terrestrial planets have been reported: KOI 1843.03 (4.245 hr of orbital period and hence the rotation period if it is tidally locked; Rappaport et al. 2013), K2-137 b (4.3 hr; Smith et al. 2018), Kepler-78 (8.5 hr; Winn et al. 2018), LHS 3844 b (11 hr; Kreidberg et al. 2019; Vanderspeck et al. 2019), 55 Cancri e (17.7 hr; Demory et al. 2016), and Kepler 10b (20 hr; Rouan et al. 2011). Their primary atmospheres are likely blown away by stellar irradiation or stellar wind due to their extreme proximity, but the vigorous outgassing due to the high surface temperature may help them to retain secondary atmospheres. Although their measured radii are much less than Jupiter’s radius, the equatorial deformation radius may still be a small fraction of their planetary radius depending on the thickness and stratification of their atmospheres. Effects of much lower surface gravity and the presence of a solid surface likely do not affect the qualitative dynamical regime transition discussed in this study. As long as the equatorial Rossby deformation radius is much smaller than the planetary radius and \( Ro \ll 1 \) at midlatitudes, the quasi-geostrophic regime naturally emerges. Thus, based on our results, we expect that the thermal phase curves may be aligned with the secondary eclipse, and the day–night temperature differences should be large for these extremely close-in terrestrial worlds. If the major atmospheric composition is rock or iron vapor, the atmospheres may be able to condense on the much cooler nightside, and the dynamical properties of nondilute condensable atmospheres (Ding & Pierrehumbert 2016; Pierrehumbert & Ding 2016; Hammond & Pierrehumbert 2017) would pose further questions about the nature of the atmospheric circulation of this kind. Note that for lava planets, rotation and tidal dissipation could affect the phase curve in addition to atmospheric flows (e.g., Selsis et al. 2013).

Carone et al. (2015) pointed out that the climate regime transition from a single equatorial jet to more jets of tidally locked rocky planets occurs when the equatorial deformation radius is smaller than half of the planetary radius. Haqq-Misra et al. (2018) suggested another interesting regime transition that could occur when the Rhines scale becomes smaller than the planetary radius. The smallest orbital (and hence the rotation) period of these studies is still comparable to a day. As mentioned above, it would be interesting to examine the regime transition further to a rotation period down to several hours for tidally locked rocky planets.

The BDs orbiting WDs may host an intrinsic heat flux much larger than those of typical hot Jupiters, which could result in thermal perturbations caused by convection impinging against the base of the atmosphere near the radiative-convective boundary (Showman & Kaspi 2013; Zhang & Showman 2014; Showman et al. 2019). These perturbations are presumably small-scale, and they can drive large-scale zonal jets and bands via wave-mean flow interactions (Showman & Kaspi 2013; Showman et al. 2019). In this study, we do not include such effects in order to keep a clean modeling framework. However, the irradiated flux of these extremely close-in gas giants still dominates over the intrinsic flux. For example, the irradiated equilibrium temperature of WD 0137–349b (Casewell et al. 2015) is on the order of 2000 K, but the intrinsic flux may be on the order of only 1000 K in terms of effective temperature. Moreover, the efficiency driving atmospheric circulation by external day–night forcing is expected to be much larger than that driven by the internal flux. The former can be thought of as a heat engine powered by the day–night contrast with a typical estimated thermodynamic efficiency on the order of 10% for hot Jupiter–like atmospheres (Koll & Komacek 2018). The circulation is more indirectly forced by intrinsic heat flux, which has to turn its convected interior heat flux first into atmospheric waves and turbulence and then into large-scale flow being forced by these eddies. Its efficiency is therefore significantly reduced by the two processes. So we expect that the circulation near the photosphere driven by day–night forcing would dominate over that by thermal perturbations in typical extremely close-in BDs around WDs. At a depth where the radiative timescale is long and the day–night thermal forcing is weak, zonal jets and banded structures can still

![Figure 13. Normalized synthetic thermal phase curves sampling different pressure levels for models with a rotation period of 2.5 hr, a short radiative timescale \( \tau_{\text{rad}} = 10^5 \) s, and different frictional drag (no drag, \( \tau_{\text{drag}} = 10^2 \) and \( 10^4 \) s).](image-url)
emerge, although the exact driving mechanism may differ from that in isolated BD models due to possible interactions between the interior convective flux and the external day–night forcing.

5. Conclusions

To summarize, we presented idealized GCMs of synchronously rotating, strongly irradiated gas-giant planets spanning a wide range of rotation periods from 80 to 2.5 hr to explore the properties of the circulation with increasing rotation rate. Our motivation is to understand the dynamics of BDs in extremely tight, close-in (1–4 hr period) orbits around WDs. Our primary results are as follows.

1. A decreasing rotation period (increasing rotation rate) causes a narrowing in the meridional width of the superrotating equatorial jet and the associated eastward-shifted hot spots, as well as the confinement to low latitudes of the standing Rossby wave (Matsumo–Gill) pattern. At fast rotation rates, regions poleward of a few degrees latitude enter a regime of small Rossby number and geostrophic balance, which exhibits very different dynamics than the canonical, slow-rotating hot Jupiter regime (where the Rossby number is order unity).

2. We showed that, on global scales, the day–night temperature difference increases with decreasing rotation period. Likewise, on global scales, the eastward offset of the dayside hot spot decreases as the rotation period decreases.

3. The thermal phase curves of our fast-rotating models show a close alignment of the peak flux to phase 0.5, where a secondary eclipse would occur if the system were eclipsing, unlike the situation with typical hot Jupiters. This result helps explain the small offset of the flux peak from phase 0.5 that has been observed in rapidly rotating BD+WD systems (e.g., Casewell et al. 2015). Generally, as compared to hot Jupiters at similar irradiation levels, we expect BDs orbiting WDs on short-period orbits to exhibit phase curves with larger day–night flux contrasts but smaller offsets of the flux peak from the secondary eclipse. Atmospheric drag (if any) exerts a significant effect on the phase curves, although more realistic models are needed for a proper representation of the Lorentz force our drag is intended to parameterize.

4. On rapid rotators, there exists a wealth of small-scale eddies and associated time variability at both low and high latitudes that is likely due to barotropic and baroclinic instability. Both types of instability may be important in shaping the mean atmospheric state, including the structure of the zonal jets and temperature. Meandering equatorial features are prevalent on the superrotating jets of all models, probably due to barotropic instability.

5. The equatorial superrotating jet is prevalent in models at all rotation periods. The meridional half-width of the equatorial jet scales closely with the equatorial deformation radius, as predicted by Showman & Polvani (2011). In rapidly rotating models, away from the equator, strong day–night forcing causes the generation of numerous alternating eastward and westward zonal jets. These are analogous to the zonal jets caused by convection or baroclinic instabilities in more familiar Jovian settings (e.g., Vasavada & Showman 2005), but here these jets form in response to a very different type of forcing. These off-equatorial jets scale well with the Rhines scale.

6. We also explore models with a long radiative timescale. In contrast to those with a short radiative timescale, these models show a much smaller horizontal temperature difference. They do exhibit modest equator-to-pole temperature contrasts and winds that are predominantly zonally aligned, but they do not exhibit the small-scale alternating jet structures prevalent on rapid rotators at short radiative time constants.

7. Our results have major implications for understanding the atmospheric circulation of BDs around WDs in extremely tight orbits, fast-rotating hot Jupiters, and terrestrial planets in ultrashort orbits.

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