Single particle diversity and mixing state measurements

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Abstract. A newly developed framework for quantifying aerosol particle diversity and mixing state based on information-theoretic entropy is applied for the first time to single particle mass spectrometry field data. Single particle mass fraction estimates for black carbon, organic aerosol, ammonium, nitrate and sulfate, derived using single particle mass spectrometer, aerosol mass spectrometer and multi-angle absorption photometer measurements are used to calculate single particle species diversity ($D_s$). The average single particle species diversity ($D_s$) is then related to the species diversity of the bulk population ($D_{\alpha}$) to derive a mixing state index value ($\chi$) at hourly resolution. The mixing state index is a single parameter representation of how internally/externally mixed a particle population is at a given time. The index describes a continuum, with values of 0 and 100% representing fully external and internal mixing, respectively. This framework was applied to data collected as part of the MEGAPOLI winter campaign in Paris, France, 2010. $D_s$ values are low (~2) for fresh traffic and wood-burning particles that contain high mass fractions of black carbon and organic aerosol but low mass fractions of inorganic ions. Conversely, $D_s$ values are higher (~4) for aged carbonaceous particles containing similar mass fractions of black carbon, organic aerosol, ammonium, nitrate and sulfate. Aerosol in Paris is estimated to be 59 % internally mixed in the size range 150–1067 nm, and mixing state is dependent both upon time of day and air mass origin. Daytime primary emissions associated with vehicular traffic and wood-burning result in low $\chi$ values, while enhanced condensation of ammonium nitrate on existing particles at night leads to higher $\chi$ values. Advection of particles from continental Europe containing ammonium, nitrate and sulfate leads to increases in $D_s$, $D_\alpha$ and $\chi$. The mixing state index represents a useful metric by which to compare and contrast ambient particle mixing state at other locations globally.

1 Introduction

Determination of the chemical composition of ambient aerosol at the single particle level is advantageous for several reasons. Single particle composition information can be used to associate certain single particle types or classes with their respective sources in urban environments (Reinard et al., 2007; Moffet et al., 2008; Snyder et al., 2009; Ault et al., 2010; Dall’Osto et al., 2013). Single particle chemical composition can also be linked directly to climate-relevant aerosol physical properties such as cloud condensation nuclei activity, hygroscopicity, optical absorption and scattering (Furutani et al., 2008; Herich et al., 2009; Moffet and Prather, 2009; Zelenyuk et al., 2010). The composition of single particles can vary significantly across an ambient particle population at a given time, depending upon local and regional sources and the extent of atmospheric processing (Whiteaker et al., 2002; Guazzotti et al., 2003; Ault et al., 2009; Pratt et al., 2011; Cahill et al., 2012; Zauscher et al., 2013). A term that is frequently used in this context is “aerosol mixing state”. In this work, we use this term to
describe how chemical species are distributed amongst the particles present in an aerosol population. In many cases, however, aerosol mixing state is described loosely as “internally mixed” or “externally mixed”, even though the reality typically lies somewhere between these two descriptions.

Riemer and West (2013) recently developed a framework for relating single particle chemical composition and bulk aerosol chemical composition in order to derive a quantitative mixing state index ($\gamma$). This approach is based on information-theoretic diversity measures, whereby the Shannon entropy of the chemical species distribution is used to determine both the average single particle species diversity ($D_\alpha$) and the bulk population species diversity ($D_\chi$). The affine ratio of these terms is then used to derive the mixing state index, $\gamma$. Such diversity measures have traditionally been applied in the field of ecology to quantify ecosystem biodiversity (Whittaker, 1965), but the approach has also been applied in several other fields to date, including economics and genetics (Attaran, 1986; Rosenberg et al., 2002).

Calculating the species diversity of the bulk population, $D_\chi$, is relatively straightforward, and can be performed by obtaining mass concentrations for the species of interest over a chosen time period using routine aerosol chemical speciation measurements. However, calculating the diversity of a single particle ($D_\alpha$), a requirement for calculating $D_\chi$, is more challenging, because the relative mass fraction of each chemical species present in that particle must first be determined (Riemer and West, 2013). Thus, measurements of the mass fractions of each chemical species present in each single particle are required in order to determine the mixing state index at a given site.

Single particle mass spectrometers have proven to be suitable for source identification and source apportionment of particulate matter in various environments (Pekney et al., 2006; Eatough et al., 2008; Snyder et al., 2009; Healy et al., 2010). However, attempting quantitative chemical speciation at the single particle level remains problematic due to the high uncertainty associated with the required scaling procedures. Particle counts must be scaled to account for size-dependent detection efficiency, often using estimated effective density values and assuming spherical shape (Qin et al., 2006; Pratt and Prather, 2009). Composition-dependent desorption/ionisation efficiencies and therefore detection efficiencies have also been reported (Kane and Johnston, 2000; Wenzel et al., 2003). At the mass spectral level, composition-dependent matrix effects can lead to differing instrumental sensitivities for the same chemical species depending on the presence or relative abundance of other constituent species within the same particle (Liu et al., 2000). Furthermore, variability in the desorption/ionisation laser power density associated with each pulse can also cause variations in mass spectral ion intensities (Gross et al., 2000; Wenzel and Prather, 2004; Reinard and Johnston, 2008). Despite these significant uncertainties, recent attempts to quantify chemical species based on single particle mass spectral data have proven reasonably successful (Ferge et al., 2006; Pratt et al., 2009; Froyd et al., 2010; Hatch et al., 2011; Jeong et al., 2011; Healy et al., 2013).

The aim of this work was to use a combination of single particle mass spectrometer, aerosol mass spectrometer and multi-angle absorption photometer measurements to estimate the mass fractions of black carbon, organic aerosol, ammonium, nitrate and sulfate present in ambient single particles detected in Paris, France, during the MEGAPOLI winter campaign 2010. These estimates were then used to calculate single particle species diversity ($D_\alpha$) and hourly average single particle species diversity ($D_\chi$) values. Finally, the mixing state index ($\gamma$) was determined for each hour of the measurement period. The methodology described here is expected to be applicable for the determination of aerosol mixing state in other locations globally.

2 Methods

2.1 Instrumentation

The sampling site and instrumentation used during the MEGAPOLI winter campaign have been described in detail previously (Healy et al., 2012, 2013). Briefly, measurements were performed at the Laboratoire d’Hygiène de la Ville de Paris (LHVP), Paris (48.75° N, 2.36° E), from 15 January–11 February 2010. An aerosol time-of-flight mass spectrometer (ATOFMS, TSI model 3800) (Gard et al., 1997) fitted with an aerodynamic focusing lens (TSI, model APL100) (Su et al., 2004) was used to collect single particle mass spectra in the size range 150–1067 nm (aerodynamic diameter, $d_{eq}$). Single particles are sampled through a critical orifice and focused in the aerodynamic lens before transmission to the sizing region. Here, $d_{eq}$ for each particle is measured based on its time-of-flight between two sizing lasers (Nd : YAG, 532 nm). Particles are subsequently desorbed/ionised using a Q-switched Nd : YAG laser (266 nm, operated at 1.1–1.3 mJ per pulse), and the resulting positive and negative ions are detected using two time-of-flight mass spectrometers. The instrument was located in an air-conditioned van and sampled aerosol through a stainless steel sampling line at a height of 4 m above ground level.

Additional instruments were located in an adjacent container. These included a multi-angle absorption photometer (MAAP, Model 5012, Thermo Scientific) (Petzold and Schönlinner, 2004), a high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS, Aerodyne Research Inc.) (DeCarlo et al., 2006), and a twin differential mobility particle sizer (TDMS) (Birmili et al., 1999). The HR-ToF-AMS measures mass concentrations of non-refractory aerosol species including ammonium, nitrate, sulfate and organic aerosol. Here the latter refers to the mixture of primary and secondary organic compounds present in the particle phase. A collection efficiency of 0.4 was calculated.
for the HR-ToF-AMS as discussed elsewhere (Crippa et al., 2013). The uncertainty associated with the HR-ToF-AMS and MAAP mass concentration measurements is 30 and 12 %, respectively (Bahreini et al., 2009; Crippa et al., 2013; Petzold and Schönlinner, 2004). All instruments located in the container sampled aerosol through a PM$_{10}$ inlet at 6 m above ground level. An aerosol diffusion dryer system was used to ensure relative humidity in the sampling line remained below 30 % (Tuch et al., 2009).

2.2 ATOFMS data analysis

Approximately 1.75 million single particle mass spectra were collected during the MEGAPOLI winter campaign, and imported into ENCHILADA (Gross et al., 2010) for further analysis. In this work, an adaptation of the quantitative approach described by Healy et al. (2013) has been employed. Healy et al. (2013) previously derived campaign-averaged ATOFMS relative sensitivity factors (RSFs) for organic aerosol, elemental carbon (EC), ammonium, nitrate, sulfate and potassium using a combination of mass spectral marker ion intensities and concurrent bulk aerosol composition measurements. Healy et al. (2013) then applied these RSFs to 10 discrete particle “types” assumed to have static chemical composition.

In this work, RSFs have been calculated at hourly resolution, and applied at the single particle level instead, in an effort to reduce the uncertainty associated with single particle chemical speciation estimates. Hourly average black carbon (BC) mass concentrations were measured using the MAAP and hourly average mass concentrations for ammonium, nitrate, sulfate and organic aerosol (OA) were measured using the HR-ToF-AMS. These data were used to constrain the ATOFMS RSF values as described in the Supplement. The five species chosen represent > 90 % of the composition of PM$_{2.5}$ measured at the site by mass (Bressi et al., 2013). The average contribution of potassium, sodium and chloride to the measured PM$_{2.5}$ mass concentration was relatively low (∼ 3 %) (Bressi et al., 2013), and therefore these species have not been included in the quantification procedure or in the diversity analysis here. The ATOFMS marker ions used for OA, BC, ammonium, nitrate and sulfate are the same as those employed previously (Healy et al., 2013). ATOFMS relative peak areas for mass spectral ions were normalised separately for positive and negative ion mass spectra, although combined dual ion mass spectral normalisation produced highly similar results. The RSF values derived for each hour of the campaign were applied to each single particle mass spectrum in that hour to estimate the mass fraction of each chemical species in each single particle. RSF values were found to be reasonably stable across the entire campaign (Fig. S1 in the Supplement). Good agreement was also observed between ATOFMS-derived mass concentration estimates of each chemical species and those measured directly by the HR-ToF-AMS and MAAP instruments, although this is expected because RSF values were calculated at hourly resolution (Figs. S2–S4 in the Supplement).

There are several assumptions involved in the estimation of the mass fraction of each chemical species at the single particle level using this approach that are likely to introduce significant uncertainty in the speciation estimates. First, it is assumed that all particles are composed exclusively of OA, BC, ammonium, nitrate and sulfate. This is a reasonable simplification in this case, since these species account for > 90 % of the average mass of particulate matter smaller than 2.5 μm in diameter (PM$_{2.5}$) in Paris (Bressi et al., 2013). The relative contribution of these five species is expected to be even higher for the size range investigated here (150–1067 nm), due to a lower contribution from sea salt and crustal material at smaller sizes. Second, it is assumed that the ATOFMS RSF values derived every hour are not significantly dependent upon either particle size or chemical mixing state. While the reasonably low variability of calculated RSF values is encouraging, error is certainly introduced through this assumption (Fig. S1 in the Supplement). Third, when scaling ATOFMS number size distributions, it is assumed that all particles are spherical, with an effective density of 1.5 g cm$^{-3}$ (Healy et al., 2013). A detailed discussion of the hourly RSF values obtained, the marker ions chosen and estimates of uncertainty associated with ATOFMS-derived particle speciation are provided in the Supplement. Briefly, the uncertainty associated with hourly mass fraction estimates for BC, OA, ammonium, sulfate and nitrate are 40, 44, 76, 41 and 41 %, respectively. The highest uncertainty is associated with ammonium, the species for which the ATOFMS also exhibits the poorest sensitivity. It is important to note, however, that uncertainty at the single particle level may be higher, but in the absence of accurate single particle composition measurements for comparison, this cannot be assessed for the ambient data set reported here.

2.3 Diversity and mixing state index calculations

The following terminology and formulae relating to information-theoretic entropy, diversity and aerosol mixing state have been described in detail by Riemer and West (2013), but are discussed again here for clarity. For a population of $N$ aerosol particles, each consisting of specific amounts of $A$ distinct aerosol species, the mass of species $a$ in particle $i$ is denoted as $\mu_{ia}$ for $i = 1, \ldots, N$ and $a = 1, \ldots, A$. The total mass of particle $i$, termed $\mu_i$, is equal to the sum of the mass of each species present in that particle:

$$\mu_i = \sum_{a=1}^{A} \mu_{ia}^a$$

and the total mass of species $a$ in the total particle population, termed $\mu^a$, is given by
The average per-particle Shannon entropy ($H_a$) is given by

$$H_a = \sum_{i=1}^{N} p_i H_i.$$  \hspace{1cm} (8)

The Shannon entropy of the entire bulk population ($H_r$) is given by

$$H_r = \sum_{a=1}^{A} -p^a \ln p^a.$$  \hspace{1cm} (9)

$H_l$ and $H_a$ are thus calculated using ATOFMS-derived single particle composition estimates, while $H_r$ is calculated using the HR-ToF-AMS/MAAP measurements of the bulk population composition. The particle diversity ($D_l$), or effective number of species in particle $i$, is calculated as follows:

$$D_l = e^{H_l}.$$  \hspace{1cm} (10)

and the average single particle diversity ($D_a$) is calculated as follows:

$$D_a = e^{H_a}.$$  \hspace{1cm} (11)

Finally, the bulk population diversity ($D_r$), or effective number of species in the bulk population, is given by

$$D_r = e^{H_r}.$$  \hspace{1cm} (12)

These terms can be related to derive the mixing state index ($\chi$):

$$\chi = \frac{D_a - 1}{D_r - 1}.$$  \hspace{1cm} (13)

Thus, values for single particle diversity are estimated at the single particle level ($D_l$), averaged across the entire particle population ($D_a$), and related to the bulk population diversity ($D_r$) to derive mixing state index values ($\chi$) at hourly temporal resolution. For a pure external mixture, $D_l = D_a = 1$, so $\chi = 0 \%$, while for a pure internal mixture, all particles have the same composition as the bulk aerosol and thus $D_a = D_r$, so $\chi = 100 \%$.

3 Results and discussion

3.1 Relating single particle composition and diversity

Estimating the mass fractions of BC, OA, ammonium, nitrate and sulfate present at the single particle level allows for the classification of particles based on their composition and size. The ATOFMS data set was queried based on size and BC mass fraction for each hour of the campaign and particle counts were scaled using concurrent TDMS number-size distribution data. The campaign-average of these hourly
two-dimensional number distributions is shown in Fig. 1. At the smallest detectable sizes (150–250 nm, \(d_{\text{a}}\)), the majority of particles are characterised by BC mass fractions higher than 0.5. This is consistent with the fresh, locally emitted combustion particles previously identified and apportioned to vehicular traffic and wood-burning activities in Paris (Healy et al., 2013). At larger sizes (>250 nm), particles are characterised by much lower BC mass fractions (<0.1), and higher inorganic ion content. Condensation of ammonium nitrate on locally emitted combustion particles and transport of ammonium, nitrate, sulfate and secondary organic aerosol from outside Paris have been demonstrated to contribute significantly to the mass concentration of particles >250 nm detected at the site (Healy et al., 2013; Crippa et al., 2013).

The particle population can also be classified according to the diversity value \(D_i\) associated with each single particle as shown in Fig. 2. This value is a representation of the effective number of species present in each single particle (Eq. 10). A particle composed of 100% BC, for example, will have a \(D_i\) value of 1, while a particle consisting of 50% OA and 50% BC will have a \(D_i\) value of 2. A particle consisting of BC and OA in unequal proportions will have a diversity value between 1 and 2. Five species were chosen for this analysis and therefore calculated \(D_i\) values will lie in the range 1–5. The majority of particles detected in Paris are characterised by \(D_i\) values of approximately 2, consistent with combustion particles composed predominantly of BC and OA (Fig. 2). However, as shown in Fig. 3 (top panel), when total particle counts in each size bin are normalised to 1, an increase in \(D_i\) with particle diameter is apparent. This is due to the higher inorganic ion content of larger accumulation mode particles (Fig. 3, bottom panel), which results in a more equally weighted contribution from each chemical species at the single particle level, and thus a higher value for \(D_i\). It follows that single particle diversity is thus expected to increase for carbonaceous particles with time as they accumulate secondary inorganic ions through atmospheric processing (Riemer and West, 2013). Typical single particle mass spectra, associated chemical composition estimates, and calculated \(D_i\) values for three particles detected during the campaign are given in Fig. 4.

### 3.2 Single particle diversity, bulk population diversity and mixing state

The average single particle diversity, \(D_a\) (Eq. 11), a representation of the average effective number of species present at the single particle level, was also calculated at hourly resolution using the ATOFMS particle composition estimates. This value can potentially range from 1 (all single particles are pure and contain only 1 species) to 5 (all single particles contain identical mass fractions of all five species). The species diversity of the bulk population, \(D_y\), was calculated directly from the HR-ToF-AMS and MAAP mass concentration measurements of BC, OA, ammonium, nitrate and sulfate for each hour of the measurement period (Eq. 12). \(D_y\) therefore represents the effective number of species present in the bulk population. This value can also range from 1 (only one species is present in the bulk population) to 5 (all five species are present in equal ratios in the bulk population). Thus, \(D_y\) ignores heterogeneity at the single particle level and describes the diversity of the bulk composition only.

The affine ratio of \(D_a\) and \(D_y\) is used to derive the mixing state index \((\chi)\) at hourly resolution (Eq. 13). The mixing state index can potentially range from 0 % when all particles are pure and contain one chemical species (fully externally mixed) to 100 % when all particles have identical mass fractions of each chemical species (fully internally mixed). The relationship between these three terms for the Paris data set is shown in Fig. 5. Mixing state index values range from 37 to 72 %, with a mean value of 59 %, thus indicating that the particle population measured during the MEGAPOLI winter campaign can never be described as fully internally mixed. It is also apparent that while the bulk population diversity \((D_y)\), at times, approaches the maximum value of 5, the average single particle diversity \((D_a)\) is rarely greater.
3.3 Dependence of mixing state upon time of day

$D_a$, $D_f$ and the mixing state index ($\chi$), were observed to be dependent upon time of day, as shown in Fig. 6. Higher values for $D_a$ are observed in the early hours of the morning (00:00–04:00 LT). During this period, local emissions associated with vehicular traffic and wood-burning activities are low, and the single particle mass fractions of ammonium and nitrate are increasing through enhanced condensation at lower temperatures (Crippa et al., 2013). The relationship between diurnal ammonium and nitrate mass concentrations and temperature is shown in Fig. S5 in the Supplement. The increase in the mass fractions of ammonium and nitrate present at the single particle level leads to an increase in $D_a$. Condensation of ammonium nitrate is also reflected in the simultaneous increase in the bulk population diversity ($D_f$). Overall, this phenomenon leads to stable, but relatively higher, mixing state index ($\chi$) values from 00:00–04:00. This behaviour is consistent with that observed for a complex urban plume scenario generated using the particle resolved PartMC-MOSAIC model (Riemer and West, 2013). In that simulation, $D_a$ and $D_f$ also increased due to condensation of ammonium nitrate on pre-existing carbonaceous particles.

When vehicular traffic emissions increase from 06:00–10:00, $D_a$ decreases due to the introduction of high numbers of particles with high BC and OA mass fractions and low $D_i$ values (Fig. 6). Simultaneously, however, $D_f$ increases because the increasing mass fraction of BC present at the bulk population level leads to a more evenly weighted contribution from all five chemical species. Overall, this phenomenon results in a more externally mixed population, and a corresponding decrease in $\chi$ is observed (Fig. 6).

A decrease in $D_f$, and to a lesser extent $D_a$, is observed at lunchtime (12:00–14:00), when primary cooking organic aerosol contributions are relatively high (Crippa et al., 2013).
Cooking-related particles are expected to have low \( D_i \) values because they are composed predominantly of OA, however the ATOFMS does not detect these particles efficiently (Healy et al., 2013), and therefore their number contribution is expected to be significantly under-represented in the ATOFMS data set. This may be either due to an absence of internally mixed BC or poor absorption efficiency at 266 nm for cooking-related aliphatic organics (Silva and Prather, 2000). Emissions of pure OA particles are expected to produce a more externally mixed population and therefore result in a decrease in \( \chi \). A minor increase in \( \chi \) is observed in this case from 12:00–14:00, although it is possible that this may be due to poor ATOFMS desorption/ionisation efficiency for cooking OA particles.

In the evening (16:00–20:00), emissions of fresh primary traffic and wood-burning particles with high BC and OA content and low \( D_i \) lead to a minor decrease in \( D_\alpha \). \( D_\gamma \) decreases to a greater extent during this period because the high OA mass contribution from wood-burning particles (Crippa et al., 2013; Healy et al., 2013) leads to a less evenly weighted contribution from all five species in the bulk population. From 20:00–23:00, however, primary emissions decrease and condensation of ammonium nitrate leads to an increase in both \( D_\alpha \) and \( D_\gamma \). Overall, these phenomena result in an increase in \( \chi \) from 16:00–23:00.

The relationship between the average \( D_\alpha \) and average \( D_\gamma \) values, sorted by hour of the day (\( N = 26 \) days), is shown in Fig. 7. At night, condensation of ammonium nitrate and a decrease in primary emissions lead to increases in both \( D_\alpha \) and \( D_\gamma \). Early morning vehicular emissions result in a decrease in \( D_\alpha \) initially, followed by a decrease in both \( D_\alpha \) and \( D_\gamma \). Despite this dependence upon time of day, the minimum and maximum hourly average \( \chi \) values are 56 and 63 %, respectively, indicating that the particle population in Paris remains relatively externally mixed throughout the day.
3.4 Dependence of mixing state upon air mass origin

The MEGAPOLI winter campaign was also characterised by distinct periods influenced by either marine or continental air masses (Healy et al., 2012; Crippa et al., 2013). Local sources contributed most significantly to aerosol mass concentrations under marine air mass conditions from 28 January 2010–7 February 2010. Advection of significant additional aerosol mass from continental eastern and northwestern Europe was observed from 26–28 January 2010 and from 7–11 February 2010 (Healy et al., 2013). It is clear that air mass origin has a significant impact upon $D_\alpha$, $D_\gamma$, and $\chi$, as shown in Fig. 8. When marine air masses prevail, $D_\alpha$ is low due to the dominance of local emissions of carbonaceous particles. These particles have low associated $D_\i$ values due to their low inorganic ion content (Fig. 9). During periods influenced by continental air masses, however, additional contributions of larger, transported particles containing higher ammonium, nitrate and sulfate mass fractions are observed. These particles have much higher $D_\i$ values (Fig. 9), which in turn leads to higher overall $D_\alpha$ values. The larger size and mass of the transported particles also impacts the bulk population composition considerably, leading to higher $D_\gamma$ values (Fig. 8). The relationship between $D_\alpha$, $D_\gamma$, and air mass origin is shown in Fig. 10. During periods influenced by continental air masses, data points are shifted towards the upper-right of the $D_\alpha$–$D_\gamma$ space. The average $\chi$ values observed for marine and continental air mass conditions are 55 and 60%, respectively. The latter value remains relatively low, because locally emitted particles with low $D_\i$ values persist independent of air mass origin (Fig. 9), resulting in a relatively externally mixed particle population at all times. Transported, aged aerosol with high-inorganic ion content is thus expected to have a higher impact upon the aerosol mixing state index in remote environments, where contributions from local sources are minimal.

4 Conclusions

Single particle mass spectrometry, aerosol mass spectrometry and black carbon measurements have been used for the first time to investigate ambient aerosol mixing state using a newly developed framework based on information-theoretic entropy (Riemer and West, 2013). ATOFMS measurements, constrained using concurrent HR-ToF-AMS and MAAP data, were used to estimate the mass fractions of BC, OA, ammonium, nitrate and sulfate present in single particles detected in Paris, France, during the MEGAPOLI winter campaign. Species mass fractions were then used to calculate single particle species diversity ($D_\i$). Low $D_\i$ values (∼2) were determined for fresh, locally emitted vehicular traffic and wood-burning carbonaceous particles composed predominantly of BC and OA. Higher $D_\i$ values (∼4) were observed for larger, transported particles containing higher mass fractions of inorganic ions. The average diversity of single particles ($D_\alpha$), and the bulk population diversity ($D_\gamma$),

Figure 8. Temporal trends for average single particle diversity ($D_\alpha$), average bulk population diversity ($D_\gamma$), mixing state ($\chi$) and average bulk population composition. “M” and “C” correspond to periods influenced by marine and continental air masses, respectively.

Figure 9. Average dependence of particle number concentration on ATOFMS-derived single particle diversity ($D_\i$) and particle diameter for periods influenced by marine (top) and continental air masses (bottom).
were used to derive the mixing state index (χ) at hourly resolution. $D_\alpha$, $D_\gamma$ and χ were all found to be dependent upon time of day. $D_\alpha$ and $D_\gamma$ values were lowest during the daytime and early evening hours when local traffic and wood-burning emissions were highest, but increased at night due to a simultaneous decrease in local emissions and enhanced condensation of ammonium nitrate on existing particles. Advection of particles with high ammonium, nitrate and sulfate mass fractions from eastern and northwestern Europe led to significant increases in $D_\alpha$, $D_\gamma$ and, to a lesser extent, χ. The average value of χ determined for aerosol in the size range 150–1067 nm detected during the MEGAPOLI winter campaign is 59%, and is thus midway between externally and internally mixed.

The approach described here represents a new means by which to compare and contrast aerosol chemical mixing state in various environments globally. Potential applications include investigating changes in the mixing state of aerosol species as a function of plume age. Ambient aerosol diversity measurements are also expected to be useful for constraining or validating regional and global scale models that resolve aerosol chemical mixing state. In particular, this metric will be useful for assessing how differences between models depend on mixing state and how differences between model results and observations depend on mixing state. The description of quantitative diversity and mixing state measures presented here implicitly assumes that “species” should mean “chemical species”. Many other possibilities exist, however, including using species subsets, for example particles that do or do not contain BC. Alternative species definitions may be advantageous to relate mixing state to aerosol climate impacts, for example direct radiative forcing. Regarding future measurement efforts, comparisons of χ values derived using single particle mass spectrometry and off-line single particle analytical techniques will also be valuable.

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