Determinants of soil respiration in a semi-arid savanna ecosystem, Kruger National Park, South Africa

Soil respiration, which is a combination of root respiration and microbial respiration, represents one of the main carbon fluxes in savannas. However, it is remarkable how little is known about these components—regarding either process-level mechanisms or quantitative estimates, especially in savanna ecosystems. Given the extensive area of savannas worldwide, this limits our ability to understand and predict the critical changes in the global carbon budget that underlie the phenomenon of global climate change. From May 2000 to April 2001, bi-weekly soil respiration measurements from two savanna types were made in 14 sampling collars (diameter = 100 mm), using a PP Systems EGM-2 respirometer. Results indicated that there was a difference in the rate of respiration between the more clayey Acacia and sandier Combretum savanna soils ($p = 0.028$). The mean (± s.d.) soil respiration in the Acacia savanna was 0.540 g/m²/h ± 0.419 g/m²/h, whilst it was 0.484 g/m²/h ± 0.383 g/m²/h in the Combretum savanna. We also found that soil respiration was sensitive to both sites rose to a maximum soil temperature was at 28 °C and declined at higher temperatures, despite different temperature sensitivities. Soil respiration increased approximately linearly with an increase of soil moisture. In both savanna sites soil is subject to a combination of high temperature and water stress, which controls the fluxes of soil carbon dioxide. We found that the two sites differed significantly in their soil moisture characteristics ($p < 0.0001$) but not with regard to temperature ($p = 0.141$), which implies that soil moisture is the main factor responsible for the differences in respiration between Acacia and Combretum savannas.

**Conservation implications:** It is argued for many protected areas that they perform a climate change buffering function. Knowing the soil respiration rate and determining its controlling factors contribute to improved understanding of whether protected areas will be net sources or sinks of carbon in the future.

**Introduction**

The sun’s energy, captured by plants through photosynthesis, is used to combine atmospheric carbon dioxide (CO₂) and water from the soil into carbohydrates. At ecosystem scale, this process is called gross primary production (GPP) (Mäkelä et al. 2008; Roxburgh et al. 2005). The organic molecules, in their various forms, are subsequently oxidised during the process of respiration, releasing CO₂, water and energy. Respiration can take several forms: the plant itself respires (autotrophic respiration) in order to grow and metabolise, the microbes in the soil respire as they break down soil organic matter (SOM), some plant parts are eaten by herbivores and some plant residues are consumed by fire. It is important to be able to measure or estimate Rₑ in order to calculate net primary productivity (NPP), the carbon that appears as an increase in plant mass and the energy that is available to power the rest of the ecosystem, as shown in Equation 1. The exudates and volatile organic carbon (VOC) terms are typically much smaller than autotrophic respiration but in savannas both may be considerable. Savannas regularly experience a combination of high temperatures, high solar radiation and water stress, which leads to large effluxes of isoprene and monoterpenes—typically about 5% of GPP, but under exceptional circumstances as much as 40% (Otter, Guenther & Greenberg 2002).

\[
GPP = NPP + Rₑ + \text{exudates} + \text{VOC},
\]

[Eqn 1]

with \( Rₑ \) representing autotrophic respiration.

Globally, terrestrial GPP is estimated at 120 Pg C yr⁻¹ and NPP at 60 Pg C yr⁻¹ (Houghton 1995; IPCC 2001). In the long term, in a steady-state system, respiration from the entire ecosystem (which includes autotrophic respiration, as well as the respiration components due to herbivore activity and fire) should be approximately equal to GPP; that is, net biome production should...
be approximately zero. The global terrestrial ecosystem as a whole is currently not in steady state: it represents a net sink of 1 Pg C yr\(^{-1}\) – 3 Pg C yr\(^{-1}\), for reasons thought to relate to rising atmospheric CO\(_2\) levels, elevated nitrogen availability, climate change and ongoing change in disturbance regimes.

Autotrophic respiration is widely estimated to amount to about half of the GPP (Prentice et al. 2001; Watson et al. 2000). A large part of this occurs belowground from the roots themselves or from root exudates, which have been variously estimated as constituting 10% – 20% of the GPP. However, in whole soil it is technically very difficult to differentiate between the respiration from the roots themselves versus that of heterotrophic respiration of soil microbes, a portion of which comes from the oxidation of root exudates.

Soil respiration, the combination of root and microbial respiration, typically represents the single largest carbon flux from savanna ecosystems to the atmosphere. At a global scale, soil respiration is the second-largest terrestrial carbon flux, estimated at 75 Pg C yr\(^{-1}\) – 100 Pg C yr\(^{-1}\) (IPCC 2001; Raich & Potter 1995), 79.3 Pg C yr\(^{-1}\) – 81.8 Pg C yr\(^{-1}\) (Raich, Potter & Bhagawati 2002) or 98 Pg C ± 12 Pg C, increasing at 0.1 Pg C yr\(^{-1}\) (Bond-Lamberty & Thomson 2010). High uncertainty in soil respiration estimates is an indication of the imprecision with which it is known globally. Such variation is due to the spatial and temporal variability of soil respiration and the inability to measure the parameter remotely remains a significant constraint to regional and global evaluations of the feedback potential that soil processes have on the global carbon cycle (Reichstein & Beer 2008).

Soil respiration responds to a suite of drivers, including temperature, moisture and vegetation productivity, all at different spatial and temporal scales (Ågren et al. 1996; Reichstein et al. 2003; Vincent et al. 2006; Webster et al. 2008). Many studies show that soil respiration is sensitive to climatic variability, expressed through soil moisture and temperature (Anderson 1995; Conant et al. 2004; Davidson, Belk & Boone 1998; Fan et al. 1995; Jarvis & Leverenz 1983; Lavigne et al. 1997; Raich & Schlesinger 1992; Veenendaal, Kolle & Lloyd 2004; Wang et al. 2010). Global warming is generally predicted to reduce ecosystem carbon storage by increasing respiration, especially of soils (Conant et al. 2004; Houghton & Woodwell 1989; Scholes & Walker 1993). Cooling in the past increased carbon storage (Schimel 1995), and cool areas generally have greater soil carbon stores than warm areas. Various studies have shown that the temperature sensitivity of respiration (conventionally expressed as Q\(_{10}\), the ratio of the rate at a given temperature to the rate ten degrees cooler) declines from the boreal to the temperate zone (e.g. Chen & Tian 2005; Peng et al. 2009; Raich & Schlesinger 1992; Reichstein et al. 2003), but it is still unclear if this represents a long-term evolutionary adaptation, a short-term acclimation or a difference in substrate quality (Kirschbaum 1995; Townsend, Vitousek & Holland 1992; Zheng et al. 2009; Wang et al. 2010). Soil respiration is broadly proportional to the amount of oxidisable substrate that is present (SOM, including humified material, finely divided soil litter such as dead roots, exudates and recently dead microbes), but is also affected by the proportions of SOM in various forms, because they differ with regard to the ease with which they can be oxidised. In a warm, semi-arid environment, availability of soil moisture is the controlling factor for soil respiration (e.g. Xu, Baldocchi & Tang 2004; Zhang et al. 2003).

Given that the various processes of respiration are just as important as photosynthesis in determining the carbon balance in an ecosystem, it is remarkable how little is known about them, with reference to either process-level mechanisms or quantitative estimates, especially in savanna ecosystems. Ecosystem warming experiments (Luo, Wan & Hui 2001), modelling analysis (Raich et al. 2002) and fundamental biokinetics (Davidson & Janssens 2006) all suggest that soil respiration should change with climate. The relative contributions of these mechanisms have been difficult to confirm observationally because of the high spatial variability of soil respiration, inaccessibility of the soil medium and the difficulty of measuring respiration continuously and at large scales, for instance by using remote sensing (Bond-Lamberty & Thomson 2010; Jones & Cox 2001). Modelling efforts that link observations of soil respiration at different scales are making important contributions to understanding the variability of the parameter at local, regional and global scales (Reichstein & Beer 2008).

This paper reports the results of a year-long series of measurements of soil respiration conducted at the flux site at Skukuza, Kruger National Park, South Africa. Data such as these can help construct models of respiration in semi-arid savannas. The objectives of the study were to:

- determine the effect of soil temperature and moisture on soil respiration
- determine whether there is a difference in soil respiration between fine-leaved *Acacia* and broad-leaved *Combretum* savannas
- determine whether there is a relationship between soil respiration and measures of carbon assimilation in the savannas, such as the fraction of absorbed photosynthetically active radiation (FAPAR).

**Method**

**Study area**

The study site was located at the Skukuza flux tower, in a semi-arid savanna ecosystem. The Skukuza flux tower is found in the southern part of the Kruger National Park. The exact location was 25°01.184’S, 31°29.813’E and oriented true north at an altitude of 365 m above sea level (Scholes et al. 2001). As is common in semi-arid savannas on granitic geologies, the undulating landscape results in a catena with sandy crests and clay-rich bottomlands. The flux tower is located at the ecotone between the broad-leaved (*Combretum*) savanna, which occupies the crests, and the fine-leaved (*Acacia*) savanna on the more clayey bottomlands (Scholes et al. 2001).

The area is characterised by shallow soils, seldom deeper than 50 cm and underlain by granite rocks of the Basement complex (Nelspruit Suite) intruded by mafic dykes of both pre-Karoo and post-Karoo age (Venter 1990). This gives rise
to a landscape of moderately undulating plains. The crests of the slopes are well-drained interfluvial areas with slope angles seldom more than 10%. The soil here is shallow and stony, presenting as a brownish, coarse sandy loam. Rock outcrops make up a small percentage of the landscape. At mid-slope there are moderately to poorly drained slopes with slope angles of between 1% and 13%. Here one finds shallow, grey, eutrophic, coarse sand over weathered rock or greyish-brown clay. At the foot of the slope there are poorly to moderately drained slopes with angles of between 1% and 9%. The soils here are paraduplex sandy clay, often associated with mafic dykes. They are shallow to moderately deep red and brown eutrophic loams. Pedocutanic clays also occur (Venter 1990).

The site has a subtropical climate, with warm to hot summers and mild winters. Temperatures can reach 44 °C in the summer, but seldom fall below freezing point in winter. Mean annual rainfall is 547 mm, with an interannual coefficient variation of 30% (Scholes et al. 2001). Rainfall occurs in the warm summer months. The first rainfall event generally occurs in late September or early October and the last event is usually in April or early May. Fire is a common phenomenon during the six months of the dry winter period, occurring on average once every five years (although the site itself had three fires in the period between 2001 and 2006).

**Data collection**

A total of 14 white PVC cylinders (collars), each 100 mm wide and 70 mm high, were inserted vertically into the soil to a depth of 5 mm in February 2000. Half the number of collars (i.e. seven) were placed in the *Acacia* savanna site, whilst the remaining seven were placed in the *Combretum* savanna site. Individual collars in each cluster were approximately 3 m apart.

Soil respiration measurements were performed in each collar bi-weekly over one year, from May 2000 to April 2001, using a closed-system soil respirometer (PP Systems EGM-2 IRGA, Amesbury, MA), with a ventilated, insulated SRC-1 soil chamber (volume = 150 cm³). Corresponding soil temperature and moisture measurements were also made for each soil respiration measurement. Both soil temperature and soil moisture were measured at a depth of 100 mm, using a thermocouple probe (Campbell Scientific) and a time-domain reflectometer instrument (Hydrosense, Campbell Scientific), respectively.

Data regarding FAPAR, an index of aboveground plant activity, were derived from the NASA SeaWifs satellite for the period May 2000 to April 2001. SeaWifs data were used because of their high spatial resolution (4.5 km in global mode).

**Data analysis**

Data were captured in a spreadsheet, which was imported to STATISTICA (version 7.1) for statistical analysis. Two-dimensional scatterplots fitted with polynomial regression curves were used to analyse the effect of soil temperature and moisture on soil respiration. Two-dimensional mean plots (box-and-whisker) were used to analyse the mean monthly respiration and FAPAR. Three-dimensional scatterplots were used to analyse the relationship between soil respiration, temperature and moisture. A multiple regression (z/zy) function was used to compute the effect of soil temperature and moisture on respiration. A single-factor analysis of variance (ANOVA) was used to examine statistically significant differences between means. Temperature sensitivity (Q₁₀) of soil respiration was determined according to Equation 2:

\[
Q_{10} = \left( \frac{R_i}{R_j} \right)^{\frac{10}{T_j - T_i}}
\]

where \(Q_{10}\) is the factor by which the reaction rate increases when the temperature is raised by 10 °C, \(R_i\) is the measured reaction rate at temperature \(T_i\) and \(R_j\) is the measured reaction rate at temperature \(T_j\).

**Results**

**Effects of soil temperature on the rate of soil respiration**

The effect of soil temperature on the rate of respiration is shown in Figures 1a and 1b. No soil temperature data below 10 °C were recorded at either sampling site. The small number of records between 11 °C and 12 °C meant that the increase in the rate of soil respiration between these temperatures in both the *Acacia* and *Combretum* savannas was not statistically significant (p = 0.141; see Table 1). The highest mean monthly respiration rates in the savannas were measured between October and December, when the mean monthly soil temperature was between 24 °C and 28 °C (Table 1).

Maximum respiration rates were observed in November, with a mean of 1.053 g CO₂/m²/h ± 0.575 g CO₂/m²/h at the *Acacia* site and 0.831 g CO₂/m²/h ± 0.329 g CO₂/m²/h at the *Combretum* site. The mean monthly soil temperature for November in both the *Acacia* and *Combretum* savannas was about 24 °C (see Table 1). The mean \(Q_{10}\) values were 2.22 for soil temperatures between 11.4 °C and 21.4 °C, 2.65 for soil temperatures between 21.4 °C and 31.4 °C, and 0.59 for soil temperatures between 31.4 °C and 44.8 °C at the *Acacia* site. For the *Combretum* site, mean \(Q_{10}\) values were 3.07 for soil temperature between 12.6 °C and 22.6 °C, 1.64 for temperatures between 22.6 °C and 32.6 °C, and 1.25 for temperatures between 32.4 °C and 37.1 °C. The decrease in \(Q_{10}\) values with increasing temperature at both sites shows that soil respiration does not follow a single, exponential response to temperature over the entire range (as is often assumed in respiration models). The rate of increase of respiration with temperature declined when soil temperature rose above 28 °C (Figures 1a and 1b).

**Effects of soil moisture on the rate of soil respiration**

We found that an increase in volumetric soil moisture content had a positive effect on the rate of respiration, as shown in Figures 2a and 2b. The highest rate of soil respiration was
TABLE 1: Mean (± s.d.) monthly effect of soil temperature, moisture and fraction of absorbed photosynthetically active radiation on the rate of soil respiration measured over 12 consecutive months at Skukuza, Kruger National Park, South Africa.

| Month | Soil temperature (°C) | Soil moisture (%) | Soil CO₂ respiration (g/m²/h) | FAPAR (μm) |
|-------|-----------------------|-------------------|-------------------------------|------------|
|       | Acacia savanna | Combretum savanna | Acacia savanna | Combretum savanna | Acacia savanna | Combretum savanna | All sites |
| Jan   | 37.329 ± 4.187 | 33.343 ± 2.841 | 1.357 ± 0.745 | 1.000 ± 0.000 | 0.338* ± 0.131 | 0.531* ± 0.191 | 0.342 ± 0.039 |
| Feb   | 25.914 ± 1.427 | 25.010 ± 1.468 | 8.537 ± 4.997 | 4.357 ± 2.432 | 0.538* ± 0.357 | 0.558* ± 0.399 | 0.317 ± 0.037 |
| Mar   | 26.371 ± 2.388 | 26.371 ± 2.388 | 3.714 ± 2.761 | 2.429 ± 1.484 | 0.739* ± 0.245 | 0.712* ± 0.276 | 0.433 ± 0.031 |
| Apr   | 23.507 ± 1.053 | 23.045 ± 1.122 | 5.571 ± 3.630 | 3.619 ± 1.794 | 0.617* ± 0.227 | 0.559* ± 0.214 | 0.414 ± 0.023 |
| May   | 18.311 ± 1.706 | 19.130 ± 1.189 | 4.357 ± 2.938 | 4.821 ± 0.993 | 0.568* ± 0.261 | 0.496* ± 0.458 | 0.330 ± 0.043 |
| Jun   | 18.021 ± 1.354 | 17.307 ± 0.554 | 3.571 ± 1.016 | 4.714 ± 0.726 | 0.375* ± 0.217 | 0.305* ± 0.176 | 0.239 ± 0.017 |
| Jul   | 16.689 ± 2.206 | 16.015 ± 1.831 | 2.804 ± 1.458 | 3.048 ± 1.066 | 0.223* ± 0.121 | 0.220* ± 0.154 | 0.180 ± 0.016 |
| Aug   | 19.939 ± 1.734 | 19.516 ± 1.812 | 1.125 ± 0.384 | 1.018 ± 0.134 | 0.162* ± 0.105 | 0.181* ± 0.190 | 0.151 ± 0.014 |
| Sep   | 22.883 ± 1.681 | 23.390 ± 1.850 | 1.071 ± 0.261 | 1.048 ± 0.216 | 0.089* ± 0.067 | 0.078* ± 0.052 | 0.146 ± 0.023 |
| Oct   | 24.493 ± 1.311 | 25.257 ± 1.609 | 7.286 ± 2.351 | 5.405 ± 1.326 | 0.834* ± 0.362 | 0.693* ± 0.296 | 0.176 ± 0.038 |
| Nov   | 24.188 ± 1.835 | 23.789 ± 1.938 | 7.375 ± 2.976 | 5.089 ± 1.761 | 1.053 ± 0.575 | 0.830* ± 0.329 | 0.407 ± 0.087 |
| Dec   | 28.364 ± 3.520 | 27.332 ± 3.296 | 4.679 ± 2.639 | 3.607 ± 1.873 | 0.851* ± 0.213 | 0.831* ± 0.134 | 0.436 ± 0.041 |
| Year mean | 22.831 ± 4.763 | 22.406 ± 4.441 | 4.169 ± 3.410 | 3.294 ± 2.069 | 0.540 ± 0.419 | 0.484 ± 0.383 | 0.297 ± 0.027 |
| Year maximum | 44.800 ± - | 37.100 ± - | 17.000 ± - | 9.000 ± - | 2.770 ± - | 2.770 ± - | 0.551 ± - |
| Year minimum | 11.400 ± - | 12.600 ± - | 1.000 ± - | 1.000 ± - | 0.010 ± - | 0.010 ± - | 0.115 ± - |

FAPAR, fraction of absorbed photosynthetically available radiation; s.d., standard deviation.

Symbols represent the seasonal effects of soil temperature, moisture and FAPAR on the rate of soil respiration:

- Low respiration.
- Moderate respiration.
- High respiration.

Combined effects of soil temperature and moisture on soil respiration

Respiration responses to increases in soil temperature have been shown to be constrained by soil moisture (e.g. Conant et al. 2004). Multiple regressions, considering both moisture and temperature effects and their interactions, account for nearly two-thirds of the variance in the data at both the Acacia and Combretum savanna sites (Figures 3a and b). The two sites differed in their soil moisture characteristics ($p < 0.0001$) but not with regard to temperature ($p = 0.141$), suggesting that soil moisture is the key factor controlling soil respiration in these savannas.

Relationship between soil respiration and the fraction of absorbed photosynthetically available radiation

The results showed a strong seasonal variation in the rate of soil respiration in both the Acacia and Combretum savannas ($p = 0.028$) (see Table 1 and Figure 4a). Soil respiration rate and FAPAR both peaked during the wet season (summer months) and were lower in the dry season (winter months). The observed trend is likely due to the combined effects of measured in November, with means (± s.d.) measured at 1.053 g/m²/h ± 0.575 g/m²/h in the Acacia savanna and 0.5830 g/m²/h ± 0.329 g/m²/h in the Combretum savanna.

For the Acacia savanna data, a linear increase with increasing soil moisture content, reaching a plateau above 10% moisture content. For the Combretum savanna data, a more-or-less linear increase with increasing soil moisture content, without a plateau.
seasonal patterns of soil moisture and temperature. In our study period, soil respiration (Figure 4a) and FAPAR (Figure 4b) both began to increase after September, when the rain began, and not July, the coldest month. Respiration reached a maximum in November, whilst FAPAR response lagged by a month, reaching its maximum value in December (Figure 4b).

**Discussion**

**Effects of soil temperature and moisture on the rate of soil respiration**

Our study corroborates findings of other studies, which have shown that soil respiration is greatest at around 25 °C (e.g. Conant *et al.* 2004). At this temperature microbes are active but the temperature is not too high to retard their activity (Swift, Heal & Anderson 1979). Our study found that when soil temperatures rose above 28 °C, soil respiration declined, but this observation is confounded by the low soil moisture, which always coincided with high temperatures. Temperature dependence of respiration therefore remains unclear at the site, as also reported by Williams *et al.* (2009), although results from Kutsch *et al.* (2008) suggest that respiration suppression at high temperatures may be occurring. \( Q_{10} \) is defined, by convention, as the factor by which respiration increases over a 10-degree temperature range, but can be calculated over any temperature difference, and can be thought of as the sensitivity of respiration to temperature at any given temperature. We calculated it for several temperature ranges. Similar to the results of Wang *et al.* (2010), we show a \( Q_{10} \) value of more than 1 in both sites at all temperature ranges analysed, but which declines at higher temperature. The decline of \( Q_{10} \) at high temperatures has also been observed by several other researchers (e.g. Allison, Wallenstein & Bradford 2010; Bradford *et al.* 2008; Chen & Tian 2005). This decline at high temperature may be due to a temperature–moisture relationship (Williams *et al.* 2009; Xu & Baldocchi 2004; Xu *et al.* 2004), metabolic processes (White 2003) and ecosystem self-adjustment on a long time scale (Davidson & Janssens 2006).

The moisture content of the soil is dependent on the rate of water lost through evapotranspiration and drainage. The temperature regime of soils is greatly affected by their moisture content. Skopp, Jawson & Doran (1990) found that microbial activity depends on soil water content to
function properly. Increases in soil moisture tend to optimise soil temperature, which, in turn, enhances the decomposition of organic matter by soil microorganisms and thus increases the rate of soil respiration. Increased respiration following wetting as reported here was also observed in the same site by Williams et al. (2009) and similar findings have been reported in other water-limited ecosystems (e.g. Jenerette, Scott & Huxman 2008; Xu & Baldocchi 2004; Xu et al. 2004).

This study concludes that soil moisture is the main factor responsible for the differences in respiration rates between Acacia and Combretum savannas, which concur with findings by Zhang et al. (2003) and Xu et al. (2004). However, it should also be noted that soil respiration is also influenced by the amount of soil carbon (i.e. the respiration substrate), which is typically low in savannas (Montgomery & Askew 1980; Scholes 1990). The typically sandy savanna soils have a low potential for stabilising carbon; the high temperatures common in savanna systems promote soil respiration. Furthermore, a significant fraction of the carbon is also lost to fire and herbivory before it can enter the soil (Scholes & Van der Merwe 1996).

**Relationship between soil respiration and the fraction of absorbed photosynthetically available radiation**

One way to tease out the autotrophic from the heterotrophic components of soil respiration is to look at the degree to which soil respiration is coupled to plant activity, indexed for instance by FAPAR. Our findings agree with those of Wang et al. (2010), who concluded from a number of studies that seasonal vegetation activity plays an important role in the seasonal variation of soil respiration. The variation of soil respiration in the growing season in, for instance, a temperate forest is mainly determined by variations of root respiration in response to seasonal temperature changes (Boone et al. 1998). Variations in plant phenology significantly contributed to the variations in calculated seasonal $Q_{10}$ values of soil respiration in a mixed temperate forest (Curiel Yuste et al. 2004), which were later attributed directly to seasonal differences in ecosystem photosynthesis (Sampson et al. 2007).

Soil moisture variability between Acacia and Combretum savannas is fundamentally due to the difference in soil types, given that both receive the same rainfall (there is some downslope movement of water from the sandy soils to the clays soils). The clayey Acacia soils retain moisture longer than the sandy soils at the Combretum site. Fine-textured soils tend to have a higher SOM content and decomposition rate than coarse-textured soils (Schimel, Stillwell & Woodmansee 1985). The result is a higher respiration rate in the clayey, broad-leaved Acacia savannas than in the sandy, fine-leaved Combretum savannas.

A recent study suggests that the effect of variations in seasonal amplitude of vegetation activity on the variation of seasonally derived $Q_{10}$ is higher than that of climatic variables such as temperature and precipitation, which highlights the linkage between plant physiological processes and soil processes (Wang et al. 2010). Vegetation activities may therefore substantially stimulate seasonal $Q_{10}$ of soil respiration (Wang et al. 2010).

**Conclusion**

This study showed that the highest rate of soil respiration in the Acacia and Combretum savannas was measured at a soil temperature of 28 °C. We observed that soil respiration is always low at much higher temperatures, but because high temperature was always associated with very low soil moisture, we cannot come to any definitive conclusion about the effect of high temperature on soil respiration from this field study alone – it will require a manipulative experiment. Because the soil temperature at this site is generally quite warm, even in winter, the availability of soil moisture is the primary control factor of soil respiration. Respiration had a positive linear response to increasing soil moisture in both the Acacia and Combretum savannas. Respiration responds more rapidly to soil wetting by rain than FAPAR does,
which seems to lag by about a month, suggesting that root respiration is a relatively small part of the observed flux. Respiration rate was 15% higher in the Acacia site than in the Combretum site, which corresponds to the difference in observed volumetric soil moisture content. Other factors that might have influenced respiration variation amongst the two sites include substrate inputs, soil texture and percentage SOM.

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Authors’ contributions

R.A.M. analysed data and collected supplementary data. R.J.S. designed and supervised the study.

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