Trend analysis of climatic variables in an arid and semi-arid region of the Ajmer District, Rajasthan, India

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Abstract

In the present study, trends and variations in climatic variables (i.e. rainfall, wet day frequency, surface temperature, diurnal temperature, cloud cover, and reference and potential evapotranspiration) were analyzed on seasonal (monsoon and non-monsoon) and annual time scales for the Ajmer District of Rajasthan, India. This was done using non-parametric statistical techniques, i.e. the Mann–Kendall (MK) and Modified Mann–Kendall (MMK) tests, over a period of 100 years. The MK test with prewhitening (MK–PW) of climatic series was also applied to climatic variables and the results were compared to those obtained through the MK and MMK tests in order to assess the performance of trend detection methods. The Pettitt–Mann–Whitney (PMW) test was applied to detect the temporal shift in climatic series. The trend analysis revealed that annual and seasonal rainfall did not show any statistically significant trend at a 10% significance level. A noticeable trend increase was found in wet day frequency, surface temperature and reference evapotranspiration (ET) during the non-monsoon season from the three non-parametric statistical tests at a 10% significance level. A statistically significant decrease in maximum temperature was found during the non-monsoon season by the MK–PW test alone. This analysis of several climatic variables at the district scale is helpful for the planning and management of water resources and the development of adaptation strategies in adverse climatic conditions.

Key words: Mann–Kendall test, modified Mann–Kendall test, Pettitt–Mann–Whitney test, Rajasthan, Sen’s slope, trends

INTRODUCTION

The effects of climate change on precipitation occurrence, distribution, intensity, quality and quantity are resulting in noticeable changes in the hydrologic cycle. While the International Panel on Climate Change (IPCC) has projected an increase in global precipitation due to the effects of climate change, both increases and decreases in precipitation have been projected at the regional scale [IPCC 2007]. The later part of the 20th century experienced an average temperature increase of 0.6°C and by the end of the 21st century, temperatures are expected to increase dramatically by 1.4°C to 5.4°C based on various climate
and that long term data is often required. However, it was also evident that a small data set might not provide an accurate portrayal of climate change and that long term data is often required. However, the study was based solely on statistical analyses and trend analysis techniques (e.g. YOON, LEE [2003]; DE, RAO 2004; ARORA et al. 2005; BASISTHA et al. 2007; PINGALE et al. 2014) and only a few studies assessed different climatic variables at the district scale [GOWDA et al. 2008]. GOWDA et al. [2008] analyzed the climatic parameters (i.e. rainfall, relative humidity, maximum temperature, minimum temperature, sunshine hours and wind speed) for the Devangere district of India to assess local scale climate change over a period of 32 years. While the results showed a mild trend in climate change in and around the Devangere region, it was also evident that a small data set might not provide an accurate portrayal of climate change and that long term data is often required. However, the study was based solely on statistical analyses and trends and shifts in climatic variables were not assessed.

Ghossi et al. [2009] observed a varied trend in Indian summer monsoon season rainfall that was not only affected by global warming but also by local changes arising from rapid urbanization, industrialization, and deforestation. DUHAN and PANDEY [2013] explored the spatial and temporal variability of precipitation in 45 districts in Madhya Pradesh (MP), India on an annual and seasonal basis. The Mann–Kendall (MK) test and Sen's slope estimator test were used for trend detection. An increase and decrease in the precipitation trend was found in the districts of MP on an annual and seasonal basis, respectively. However, the study only assessed precipitation and a comparative study of different climatic variables using different statistical techniques was not performed.

There is need to analyze additional parameters within the political boundaries of countries, to clearly assess and understand climate change [ADAMOWSKI et al. 2012; ARAGHI et al. 2015; BELAYNEH et al. 2014; CAMPISI et al. 2012; HADJARY et al. 2013; TIWARI, ADAMOWSKI 2014]. Micro-scale studies are required to assess climate change and to identify its real causes for the proper planning and management of water resources [ADAMOWSKI et al. 2012; ADAMOWSKI, PROKOPH 2013; HALBE et al. 2013; SAADAT et al. 2011]. No studies have been reported in the literature that have investigated climate in terms of inter-seasonal and inter-annual variation in temperature, precipitation, number of rainy days, humidity, cloud cover, evaporation and evapotranspiration along with seasonal shifts for the Ajmer District for the State of Rajasthan, India. Ajmer is an important holy place in the central part of the State of Rajasthan. The effects of climate change are evident in the uneven rainfall pattern and temperature increases, and may due to increased population and developmental activities in Ajmer [PINGALE et al. 2015]. Therefore, the statistical analysis should be extended to analyze additional climatic parameters and their relationships with water resources, land use/cover change, urbanization, etc. Given this, the present study was undertaken with the objective to analyze representative meteorological parameters (i.e. rainfall, wet day frequency, temperature, cloud cover, reference and potential evapotranspiration) using non-parametric statistical tests (i.e. Mann–Kendall (MK) test, modified Mann–Kendall (MMK) test, MK test with prewhitening of series (MK–PW) and Pettitt–Mann–Whitney (PMW) test) to observe the trends and shifts in these climatic variables in the Ajmer District of Rajasthan, India. Also, a comparative assessment of the different trend detection techniques was performed to analyze trends and shifts in the selected climatic variables. Such information can be useful to allow stakeholders to plan and manage water resources in a more sustainable manner [BUTLER, ADAMOWSKI 2015; HALBE et al. 2014; INAM et al. 2015; KOLJIVADY et al. 2014a, b; STRAITH et al. 2014].

MATERIALS AND METHODOLOGY

STUDY AREA AND DATA USED

Ajmer District is located between 25°38' to 26°58’N latitude and 73°54’ to 75°22’E longitude (Fig. 1) in India. It is situated almost in the heart of Rajasthan, and it is bordered by the Nagaur district to the north, the Jaipur and Tonk districts to the east, the Bhiwara district to the south, and the Pali district to the west. To the north of Ajmer city is a large artificial lake called Anasagar, which is adorned with a marble structure called Baradari. The Ajmer District has an area of 8,481 km² and a population of 2,584,913 [Registrar General, India 2011]. It is situated 486 m above mean sea level (msl), has a hot climate characterized by extremely hot summers and receives a fairly large amount of rainfall.

In the present study, 0.5°×0.5° gridded meteorological data (i.e. monthly rainfall, wet day frequency, surface temperature, cloud cover, reference evapotranspiration and potential evapotranspiration) from the Ajmer District of Rajasthan were utilized from the year 1903 to 2002. These data sets were obtained from the Indian water portal web, developed from the Climate Research Unit TS 2.1 dataset [MITCHELL, JONES 2005], UK and GIS software GRASS (Geographic Resources Analytical Support System). The meteorological data of 100 years (1903 to 2002) was selected to assess climate change trends and variations.
on seasonal and annual scales for the Ajmer District. The monsoon season runs for the months of June to September, and the non-monsoon season represents the months of October to May.

**METHODOLOGY**

Assessment of climate change for the Ajmer District was carried out on seasonal (monsoon and non-monsoon) and annual time scales using the non-parametric (MK, MK–PW, MKK and PMW test), statistical analyses. The derived reference and potential evapotranspiration from the FAO-56 manual [Allen et al. 1998] were also analyzed for the Ajmer District over the period of 100 years.

**Normalisation and autocorrelation analysis of time series**

Normalised climatic time series were used to test for the presence of outliers, which were obtained from the following relationship [Pingale et al. 2014; 2015; Rai et al. 2010]:

\[
X_i = (x_i - \bar{x}) / \sigma
\]

Where \(X_i\) is the normalized anomaly of the series, \(x_i\) is the observed time series, \(\bar{x}\) and \(\sigma\) are the long-term mean and standard deviation of annual/seasonal time series.
The autocorrelation (serial coefficients) test was performed to check the randomness and periodicity in the time series [MODARRES, SILVA 2007]. If lag-1 serial coefficients were not statistically significant then the MK test was applied to the original time series [KARPOUZOS et al. 2010; LUO et al. 2008; PINGALE et al. 2015]. The MMK test was applied to the statistically significant time series after removing the effect of serial correlation. The serial correlation coefficients of normalized climatic series were computed for lags $L = 0$ to $k$, where $k$ is the maximum lag (i.e. $k = n/3$), and $n$ is the length of the series. The autocorrelation coefficient $r_k$ of a discrete time series for lag-$k$ was estimated as follows:

$$r_k = \frac{\sum_{i=1}^{n-k}(X_i - \bar{X})(X_{i+k} - \bar{X})}{\sum_{k=1}^{n-k}(X_i - \bar{X})^2(X_{i+k} - \bar{X})^2}^{0.5} \tag{2}$$

Where $r_k$ is the lag-$k$ serial correlation coefficient. The hypothesis of serial independence was then tested by the lag-1 autocorrelation coefficient as $H_0 : r_1 = 0$ against $H_1 : |r_1| \geq 0$ using the test of significance of serial correlation [RAI et al. 2010; YEVEVICH 1972]:

$$(r_1)_{ij} = \frac{1 \pm t_{\alpha/2}(n-k)^{0.5}}{n-k} \tag{3}$$

Where $(r_1)_{ij}$ is the normally distributed value of $r_1$, $t_{\alpha/2}$ is the normally distributed statistic at ‘$\alpha$’ level of significance. The values of $t_{\alpha/2}$ are 1.645, 1.965 and 2.326 at significance levels of 0.10, 0.05 and 0.01, respectively. If $|r_1| \geq (r_1)_{ij}$, the null hypothesis about serial independence was rejected at significance level $\alpha$ (0.05 in this study). For non-normal series, the MK test is an appropriate choice for trend analysis [BASISTHA et al. 2009; YUE, PIOLON 2004]. Therefore, the MK test was used in the present study where autocorrelation was found to be non-significant at the 5% level of significance.

**Mann–Kendall test**

The MK test is a non-parametric test used to detect trends in a time series [MANN 1945]. The non-linear trend and the turning point can be derived from Kendall test statistics [KENDALL 1975]. This method searches for a trend in a time series without specifying whether the trend is linear or nonlinear. It has been found to be an excellent tool for trend detection and many researchers have used this test to assess the significance of trends in hydro-climatic time series such as water quality, stream flow, temperature and precipitation (e.g. LUDWIG et al. [2004]; ZHANG et al. [2004]; McBEAN, MOTIEE [2008]; ADAMOWSKI et al. [2009]; BASISTHA et al. [2009]; ADAMOWSKI et al. [2010]; RAI et al. [2010]; NALLEY et al. [2012]; PASTRA et al. [2012]; NALLEY et al. [2013]; PINGALE et al. 2015]. The MK test can be applied to a time series $x_i$ ranked from $i = 1, 2, \ldots n$ and $x_j$ ranked from $j = i + 1, 2, \ldots n$. Each data point $x_i$ is used as a reference point and is compared with all other data points $x_j$ such that:

$$\text{sgn}(x_j - x_i) = \begin{cases} 1 & \text{if } (x_j - x_i) > 0 \\ 0 & \text{if } (x_j - x_i) = 0 \\ -1 & \text{if } (x_j - x_i) < 0 \end{cases} \tag{4}$$

The Kendall test statistic $(S)$ is given as:

$$S = \sum_{i=1}^{n-1} \text{sgn}(x_j - x_i) \tag{5}$$

Where $\text{sgn}(x_j - x_i)$ is the signum function. The test statistic $(S)$ has been assumed to be asymptotically normal, with $E(S) = 0$ for sample size $n \geq 8$ and variance as follows:

$$V(S) = \frac{[n(n-1)(2n+5) - \sum_i (t_i - 1)(2t_i + 5)]}{18} \tag{6}$$

Where $t_i$ denotes number of ties up to sample $i$. The standardized MK test statistics $(Z_{mk})$ can be estimated as follows:

$$Z_{mk} = \frac{S - 1}{\sqrt{V(S)}} \text{ if } S > 0 \tag{7}$$

$$Z_{mk} = \frac{0}{\sqrt{V(S)}} \text{ if } S = 0$$

$$Z_{mk} = \frac{S + 1}{\sqrt{V(S)}} \text{ if } S < 0$$

The standardized MK test statistics $(Z_{mk})$ follows the standard normal distribution with a mean of zero and variance of one. If $|Z_{mk}| \leq Z_{0.25}$ (here $\alpha = 0.1$), then the null hypothesis for no trend is accepted in a two sided test for trend and the null hypothesis for no trend is rejected if $|Z_{mk}| > Z_{0.25}$. Failing to reject $H_0$ (i.e. null hypothesis) does not mean that there is no trend. Rather, it is a statement that the evidence available is not sufficient to conclude that there is a trend [HELSEL, HIRSCH 2002]. A positive value of $Z_{mk}$ indicates an ‘upward trend’ and a negative value indicates a ‘downward trend’.

**Mann–Kendall test with pre-whitening of series**

The prewhitening of time series involves the computation of serial correlation and removing the correlation if the calculated serial correlation is significant at a significance level of 0.05 [BURN, ELNUR 2002]. The pre-whitening of the time series is accomplished as follows:

$$X'_i = (X_{i+1} - rx_k) \tag{8}$$
Where \( x_i \) is the original time series with autocorrelation for time interval \( k \); \( X^* \) is the pre-whitened time series; and \( r \) is the lag-1 autocorrelation coefficient. This pre-whitened series is then subjected to the MK test for detecting the trend.

**Modified Mann–Kendall test**

The MK–PW test is used to detect trends in a time series in the presence of autocorrelation [CUNDERLIK, BURN 2004]. However, pre-whitening reduces the detection rate of significant trends in the MK test [YUE, HASHINO 2003]. Therefore, the MMK test is employed for trend detection of an auto-correlated series [HAMED, RAO 1998]. Only significant values of \( \rho_k \) are used to calculate the variance correction factor \( n/n^*_s \) (equation 9) as the variance of \( S \) is underestimated when the data are positively autocorrelated.

\[
n/n^*_s = 1 + \\
+ \frac{2}{n(n-1)(n-2)} \sum_{i=1}^{n-1} (n-i)(n-i-1)(n-i-2) \rho_i(i) \tag{9}
\]

Where \( n \) is the actual number of observations, \( n^*_s \) is considered as the ‘effective’ number of observations to account for autocorrelation in data, and \( \rho_i(i) \) is the autocorrelation function of the ranks of the observations which is given as follows [KENDALL 1955]:

\[
\rho_i(i) = \sin^{-1} \left( \rho(i) \right) \tag{10}
\]

Where \( \rho(i) \) is the parent autocorrelation function of rank of the observation. The corrected variance is then computed as:

\[
V^*(S) = V(S) \cdot n/n^*_s \tag{11}
\]

Where \( V(S) \) is obtained from equation (6). The remaining process of trend analysis is included in the MK test by incorporating the corrected variance in equation (7). A significance level of 10% was used for the autocorrelation of rank \( \rho_i(i) \), which produces the best overall empirical significance level. The advantage of using corrected variance is that there is no need to either normalize data or their autocorrelation function [RAI et al. 2010].

**Sen’s estimator of slope and percentage change over a period**

If a linear trend is present in a time series, then the true slope of trend can be estimated using a simple non-parametric procedure [SEN 1968; THEIL 1950] that is computed as:

\[
Q_i = \text{median} \left( \frac{x_j - x_k}{j-k} \right) \quad \forall k \leq j \tag{12}
\]

Where \( x_j \) and \( x_k \) are data values at times \( j \) and \( k \) \((j > k)\), respectively. The median of \( N \) values of \( Q_i \) is Sen’s estimator of slope. If \( N \) is odd, then Sen’s estimator is computed by \( Q_{med} = Q_{(N+1)/2} \) and if \( N \) is even, then Sen’s estimator is computed by \( Q_{med} = (Q_{(N/2)} + Q_{(N/2+1)}/2) \). Finally, \( Q_{med} \) is tested by a two-sided test at a 100(1 – \( \alpha \))% confidence interval.

The percentage change (%) is estimated assuming a linear trend in which magnitude by Theil and Sen’s median slope and mean are used [BASISTHA et al. 2009; YUE, HASHINO 2003]. The %changes over a period is expressed as follows:

\[
\% \text{change} = \left( \frac{\text{median slope} \cdot \text{length of period}}{\text{mean}} \right) \tag{13}
\]

**Pettitt–Mann–Whitney (PMW) test for shift detection**

The PMW test is used for the determination of shift in climatological time series [PETTITT 1979]. This test was performed using the evaluation version of XLstat 2011 software. This test can be briefly described using PMW statistics [BASISTHA et al. 2009; KIELY et al. 1998] where \( T \) is the length of the time series and \( \tau \) is the year of the most likely change point. Considering the time series as two samples represented by \( X_1 \ldots X_T \) and \( X_{r+1} \ldots X_T \), the index \( V_\tau \) is defined as:

\[
V_\tau = \sum_{j=1}^{T} \text{sgn}(X_j - X_j) \quad \text{for any } \tau \tag{14}
\]

Let a further index \( U_\tau \) be defined as:

\[
U_\tau = \sum_{j=1}^{T} \sum_{\tau} \text{sgn}(X_j - X_j) \tag{15}
\]

A plot of \( U_\tau \) against \( \tau \) for a time series with no change point would result in a continually increasing value of \( |U_\tau| \). However, if there is a change point (even a local change point), then \( |U_\tau| \) would increase up to the change point and then begin to decrease. The most significant change point \( \tau \) can be identified as the point where the value of \( |U_\tau| \) is a maximum and can be found using equation (16).

\[
K_\tau = \max_{1 \leq \tau \leq T} |U_\tau| \tag{16}
\]

The probability of a change point being at a year where \( |U_\tau| \) is the maximum, is approximated by:

\[
p = 1 - \exp \left( -\frac{6K_\tau^2}{T^3 + T^2} \right) \tag{17}
\]

Further for \( 1 \leq \tau \leq T \), the series...
is introduced and defined as:

\[ p(\tau) = 1 - \exp\left( -\frac{6U(\tau)^2}{T_1^3 + T_2^3} \right) \]  

(19)

In this way, series consisting of probabilities of the change point at each year are obtained for the shift detection in the time series of annual and seasonal rainfall and temperature over a period of time.

### RESULTS AND DISCUSSION

In the present study, assessment of climate change (i.e., trends, shifts and variability) for the Ajmer District was carried out at seasonal (monsoon and non-monsoon) and annual time scales through the MK, MK–PW, MMK and PMW non-parametric statistical tests. Sen’s slope estimator and percentage confidence interval were decided by the 95% confidence level. The rainfall, wet day frequency, surface temperature (minimum, average and maximum), cloud cover, reference evapotranspiration (ET) and potential evapotranspiration (PET) were assessed for the Ajmer District over the period of 100 years.

### NORMALISATION AND AUTOCORRELATION ANALYSIS

Initially autocorrelation analysis was performed to identify the suitable trend analysis method for the original and normalized series of climatic variables. The results of autocorrelation analysis up to the 20th lag are shown in Table 1 and Fig. 2 to 4 for the climatic variables (i.e. rainfall, wet day frequency, temperature (minimum, average and maximum), cloud cover, reference ET and PET) for the monsoon season, non-monsoon season and annual scale. The upper and lower boundaries were decided by the 95% confidence interval to test the limits of the autocorrelation coefficient [Pingale et al. 2015]. The significant autocorrelation in climatic variables at lag-1 was observed for the monsoon season, non-monsoon season and annual scale at the 0.05 significance level. For auto-correlated series at lag-1, the MK test was applied for non-significant climatic variables while the MMK test was used for the significant climatic variables. The autocorrelation coefficient at lag-1 was selected because Wagesho et al. [2013] found that the dependence of physical systems on past values is likely to be strongest for the most recent past.

Based on the autocorrelation analysis, the MMK test was applied for the diurnal temperature for the monsoon season, non-monsoon season and annual scale. For the annual scale the MMK test was applied for PET and cloud cover while for the non-monsoon season the MK test was applied for the temperature, cloud cover, reference ET and PET. For the remaining variables, the MK test was applied for the monsoon season, non-monsoon season and annual scale (Tab. 1). The MK and MMK tests were applied at the 10% significance level for the trend analysis of the monsoon season, non-monsoon season and annual time scale. However, the MK, MK–PW and MMK test were also applied for all climatic variables for the monsoon season, non-monsoon season and annual scale. This was done in order to compare the results and accurately estimate the trends in climatic variables over the period of 100 years.

### TREND ANALYSIS

The climatic variables were analyzed for the monsoon season, non-monsoon season and annual scale using the MK, MK–PW and MMK trend tests. The MMK and MK–PW test were used for the series that had significant lag-1 autocorrelation at the 0.05 level of significance. The results of the three trend tests were derived for all climatic variables and compared at the 10% significance level. The test statistics \( S \) and \( Z_{mk} \) derived for the monsoon season, non-monsoon season and annual scale are presented in Tables 2 to 4 for both the significant and non-significant climatic variables. The positive values of \( Z_{mk} \) statistics indicated increasing trends in climatic variables while negative values of \( Z_{mk} \) showed decreasing trends for the significant and non-significant climatic variables.

After the MK test, the Sen’s estimator of slope was employed to find out the change per unit time of the trends observed in the time series of all climatic variables. The corresponding results are presented in Table 5. The Sen’s slope estimates with upper and lower bounds for statistically significant upward slopes for wet day frequency, surface temperature and downward slope of reference ET are shown in Fig. 5. The percentage changes in climatic variables over the 100 year period of study for the statistically significant climatic variables are presented in Table 5 for monsoon season, non-monsoon season and annual scale. The mean and standard deviation of all the nine

### Table 1. Summary of autocorrelation analysis at lag-1 for climatic variables

| Parameters         | Monsoon season | Non-monsoon season | Annual     |
|--------------------|----------------|--------------------|------------|
| Rainfall           | 0.060          | 0.008              | 0.041      |
| Wet day frequency  | 0.007          | 0.123              | 0.024      |
| Minimum temperature| 0.024          | 0.229              | 0.166      |
| Average temperature| 0.016          | 0.229              | 0.165      |
| Maximum temperature| 0.008          | 0.235              | 0.165      |
| Diurnal temperature| 0.311          | 0.298              | 0.384      |
| Cloud cover        | 0.102          | 0.490              | 0.549      |
| Reference ET       | -0.007         | 0.206              | 0.150      |
| PET                | 0.048          | 0.311              | 0.384      |

Explanations: ET = evapotranspiration; PET = potential evapotranspiration; highlighted bold values indicates significant autocorrelation at the 0.05 significance level.

Source: own study.
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Fig. 2. Autocorrelation analysis of climatic variables for monsoon season: a) rainfall, b) wet day, c) minimum temperature, d) average temperature, e) maximum temperature, f) diurnal temperature, g) ET, h) PET, i) cloud cover; source: own study
Fig. 3. Autocorrelation analysis of climatic variables for non-monsoon season: a) rainfall, b) wet day, c) minimum temperature, d) average temperature, e) maximum temperature, f) diurnal temperature, g) ET, h) PET, i) cloud cover; source: own study
Fig. 4. Autocorrelation analysis of climatic variables for annual scale: a) rainfall, b) wet day, c) minimum temperature, d) average temperature, e) maximum temperature, f) diurnal temperature, g) ET, h) PET, i) cloud cover; source: own study
Table 2. Trend analysis of climatic variables for monsoon season at Ajmer District

| Parameters            | MK test    | MK test with prewhitening | MMK test |
|-----------------------|------------|---------------------------|----------|
|                       | $S$ $Z_{mk}$ | $S$ $Z_{mk}$              | $S$ $Z_{mk}$ |
| Rainfall              | 250 0.74   | 245 0.73                  | 250 0.74 |
| Wet day frequency     | 406 1.21   | 409 1.22                  | 406 1.20 |
| Minimum temperature   | -501 -1.49 | -498 -1.48                | -501 -2.04 |
| Average temperature   | -505 -1.50 | -505 -1.50                | -505 -2.58 |
| Maximum temperature   | -501 -1.49 | -502 -1.49                | -501 -2.04 |
| Diurnal temperature   | -247 -0.80 | -302 -0.97                | -247 -0.25 |
| Cloud cover           | -8 -0.02   | -4 -0.01                  | -8 -0.01 |
| Reference ET          | -613 -1.82 | -613 -1.82                | -613 -2.63 |
| PET                   | -284 -0.84 | -284 -0.84                | -284 -0.83 |

Explanations: $S = $ Kendall test statistic, $Z_{mk} = $ standardized MK test statistics, $ET, PET$ as in Table 1; trends are highlighted in bold at the 10% level of significance.

Source: own study.

Table 3. Trend analysis of climatic variables for non-monsoon season at Ajmer District

| Parameters            | MK test    | MK test with prewhitening | MMK test |
|-----------------------|------------|---------------------------|----------|
|                       | $S$ $Z_{mk}$ | $S$ $Z_{mk}$              | $S$ $Z_{mk}$ |
| Rainfall              | 323 0.96   | 315 0.94                  | 323 1.14 |
| Wet day frequency     | 669 1.99   | 671 2.00                  | 669 1.99 |
| Minimum temperature   | 1050 3.12  | 1046 3.11                 | 1050 3.82 |
| Average temperature   | 1053 3.13  | 1053 3.15                 | 1053 3.13 |
| Maximum temperature   | 1055 3.14  | 1053 3.13                 | 1055 3.39 |
| Diurnal temperature   | 27 0.08    | 36 0.11                   | 27 0.03  |
| Cloud cover           | 501 1.49   | 497 1.48                  | 501 0.85 |
| Reference ET          | 998 2.97   | 998 2.97                  | 998 2.97 |
| PET                   | 550 1.64   | 550 1.64                  | 550 1.40 |

Explanations as in Table 2.
Source: own study.

Table 4. Trend analysis of climatic variables for annual level at Ajmer District

| Parameters            | MK test    | MK test with prewhitening | MMK test |
|-----------------------|------------|---------------------------|----------|
|                       | $S$ $Z_{mk}$ | $S$ $Z_{mk}$              | $S$ $Z_{mk}$ |
| Rainfall              | 344 1.02   | 343 1.02                  | 344 1.02 |
| Wet day frequency     | 667 1.98   | 674 2.00                  | 667 2.52 |
| Minimum temperature   | 558 1.66   | 557 1.66                  | 558 2.26 |
| Average temperature   | 525 1.56   | 534 1.59                  | 525 1.22 |
| Maximum temperature   | 517 1.54   | 515 1.53                  | 517 1.66 |
| Diurnal temperature   | -333 -1.04 | -377 -1.15                | -333 -0.50 |
| Cloud cover           | 425 1.26   | 430 1.28                  | 425 0.76 |
| Reference ET          | 443 1.32   | 443 1.32                  | 443 1.37 |
| PET                   | 247 0.73   | 247 0.73                  | 247 1.23 |

Explanations as in Table 2.
Source: own study.

Table 5. Sen’s slope and percentage change over 100 years of climatic variables

| Parameter             | Monsoon season | Non-monsoon season | Annual |
|-----------------------|----------------|-------------------|--------|
|                       | Sen's slope % change | Sen's slope % change | Sen's slope % change |
| Rainfall              | 0.397 8.03 | 0.079 17.15 | 0.675 12.46 |
| Wet day frequency     | 0.020 8.09 | 0.111 17.83 | 0.033 10.77 |
| Minimum temperature   | -0.004 -1.48 | 0.007 4.25 | 0.003 1.53 |
| Average temperature   | -0.004 -1.24 | 0.007 2.86 | 0.003 1.07 |
| Maximum temperature   | -0.004 -1.09 | 0.007 2.20 | 0.003 0.89 |
| Diurnal temperature   | 0.000 0.00 | 0.000 0.00 | 0.000 0.00 |
| Cloud cover           | 0.000 0.00 | 0.004 2.57 | 0.003 1.22 |
| Reference ET          | -0.002 -0.94 | 0.006 1.64 | 0.003 0.53 |
| PET                   | -0.001 -0.29 | 0.003 0.51 | 0.001 0.18 |

Explanations: $ET, PET$ as in Table 1; trends in Table 1; bold shows statistically significant Sen’s slope at the 10% level of significance.
Source: own study.

climatic variables are given in Table 6. These represent the changes and variation in all the climatic variables over the 100 year period for the above mentioned time scales in the Ajmer District.

The reference $ET$ in the monsoon season showed a significant decreasing trend over the study period by the MK test, MK–PW test ($Z_{mk} = -1.82$) and MMK test ($Z_{mk} = -2.63$). However, these tests failed to detect a significant trend for the other climatic variables using the MK and MK–PW tests at a 10% level of significance. The MMK test failed only for rainfall, diurnal temperature and cloud cover in the monsoon season at the 10% significance level. The null hypothesis of no trend indicated that the evidence of trends was not sufficient on seasonal and annual time scales in some of the studied climatic variables at the 10% significance level. These results were in agreement with some studies in other regions of India that also found no trends in rainfall (e.g. Arora et al. [2005]; DASH et al. [2007]; KUMAR et al. [2010]; PINGALE et al. [2013]). However, increasing and decreasing trends in rainfall have been observed in other regions of the country (e.g. BASISTHA et al. [2007]; RAI et al. [2010]; PINGALE et al. [2014]). It is clear from Table 2 that no significant increasing trend existed in wet day frequency and minimum temperature in the monsoon season based on the MMK test. However, GOWDA et al. [2008] found decreasing trends in annual rainfall and expected high intensity rainfall or storm events for the Devangere district of Karnataka, India. Also, increasing trends in minimum temperature have been observed by other researchers in India (e.g. JAIN et al. [2013]; PINGALE et al. [2014]; PINGALE et al. [2015]). Although significant decreasing trends were observed in average and maximum temperature using the MMK test (Tab. 2), previous studies have reported increasing trends in average and maximum temperature for other regions of India (e.g. SINGH et al. [2008]).
Trend analysis of climatic variables in an arid and semi-arid region of the Ajmer District, Rajasthan, India

Table 6. Summary of statistics of climatic variables at Ajmer District

| Parameters                      | Monsoon mean | Monsoon SD | Non-monsoon mean | Non-monsoon SD | Annual mean | Annual SD |
|--------------------------------|--------------|------------|------------------|----------------|-------------|-----------|
| Rainfall, mm                   | 495.07       | 177.11     | 46.29            | 28.71          | 541.35      | 183.78    |
| Wet day frequency, days        | 24.87        | 4.81       | 5.99             | 1.65           | 30.86       | 5.18      |
| Minimum temperature, °C        | 24.73        | 0.68       | 15.66            | 0.57           | 18.68       | 0.49      |
| Average temperature, °C        | 29.33        | 0.68       | 23.34            | 0.58           | 25.33       | 0.49      |
| Maximum temperature, °C        | 33.95        | 0.68       | 31.04            | 0.58           | 32.01       | 0.49      |
| Diurnal temperature, °C        | 9.20         | 0.09       | 15.38            | 0.10           | 13.32       | 0.08      |
| Cloud cover, %                 | 49.93        | 2.54       | 23.34            | 1.98           | 25.33       | 0.98      |
| Reference ET, mm               | 3.21         | 0.32       | 3.57             | 0.55           | 3.79        | 0.67      |
| PET, mm                        | 26.43        | 0.28       | 53.04            | 0.47           | 79.47       | 0.60      |

Explanations: SD = standard deviation; ET, PET as in Table 1.

The significant increasing trends were observed in most of the climatic variables in the non-monsoon season by all three methods over the 100 year period with the exception of maximum temperature by the MK–PW test, which showed a significant decreasing trend in the non-monsoon season at the 10% significance level (Tab. 3). Significant increasing trends were observed in wet day frequency, minimum and average temperature and reference ET in the non-monsoon season from the three tests. However, these tests failed to detect significant trends for other climatic variables at the 10% level of significance (Tab. 3).

Significant increasing trends were observed in wet day frequency and minimum temperature on an annual scale using the three statistical methods over the study period. However, a significant increasing trend was only observed in annual maximum temperature using the MMK test (Tab. 4). These results are in good agreement with previous studies in India (e.g. KOTHYARI, SINGH [1996]; KOTHAWALE, KUMAR [2005]; JAIN et al. [2013]). A study conducted by PINGALE et al. [2015] similarly observed average temperature increases over time in the Ajmer District, which contrasted with findings of the IPCC that demonstrated both an average increase and decrease in temperature at the global level.

Significant increasing trends were observed in wet day frequency and minimum temperature on an annual scale using the three statistical methods over the study period. However, a significant increasing trend was only observed in annual maximum temperature using the MMK test (Tab. 4). These results are in good agreement with previous studies in India (e.g. KOTHYARI, SINGH [1996]; KOTHAWALE, KUMAR [2005]; JAIN et al. [2013]). A study conducted by PINGALE et al. [2015] similarly observed average temperature increases over time in the Ajmer District, which contrasted with findings of the IPCC that demonstrated both an average increase and decrease in temperature at the global level. All three trend tests failed to detect significant trends in other remaining climatic variables at the 10% level of significance (Tab. 3). However, statistically significant upward...
slopes were observed in wet day frequency on an annual scale (Tab. 5), which indicated an increase in annual wet day frequency and a higher possibility of intense storms in the Ajmer District.

PETTIT–MANN–WHITNEY (PMW) TEST FOR SHIFT DETECTION

Shift detection is very important in identifying evidence of climate change and it provides valuable information for planning adaptation measures. Therefore, shift detection in the climatic time series was carried out using the PMW test on the seasonal and annual time scales for the climatic variables and a shift was observed in statistically significant climatic variables in the Ajmer District over the 100 year time period (Fig. 6 to 14). A shift in wet day frequency and average cloud cover variation (%) on an annual scale was observed during the years of 1941 and 1954, respectively (Fig. 6 and 14). However, a shift was not found in wet day frequency during the monsoon season and non-monsoon season and average cloud cover variation (%) during the monsoon season. No shift was observed in rainfall on either the seasonal or annual scale in the Ajmer District. However, this contrasted with Ajmer city where significant shifts have been observed in non-monsoon season rainfall in addition to an increasing trend in non-monsoon rainfall at the 10% significance level [PINGALE et al. 2015]. The surface temperature (minimum, average and maximum) in the non-monsoon season had a significant shift in 1937 (Fig. 6 to 9). There were no observablesignificant changes in surface temperature for the annual and monsoon season over the 100 year period. Significant shifts (increases) in both the reference evapotranspiration and PET were observed in the non-monsoon season in 1937 and 1930, respectively (Fig. 10 and 12). However, significant decreases in shift were observed for the PET during the monsoon season from 1952 onward (Fig. 11). There were no shifts in reference ET during the monsoon season and annual scale and no shifts in PET were observed at the annual time scale for the study period. The shifts in climatic variables were likely due to the global climate shift or factors such as the weakening global
monsoon circulation, the reduction in forest cover and increasing aerosol concentration in the atmosphere due to anthropogenic activities [Basistha et al. 2009].

The assessment of trends and shifts in different climatic variables on annual and seasonal scales is crucial for adaptation planning in arid and semi-arid regions of India at district scales. Similar studies like Pingale et al. [2014] that assess trends on local and regional levels at annual and seasonal time scales are useful for adaptation planning measures.

SUMMARY AND CONCLUSIONS

In this study, trend analysis of nine climatic variables during the monsoon season, non-monsoon season and annual scale was performed for the period of 1903 to 2002 in the Ajmer District using three non-parametric statistical tests (i.e., MK test, MK–PW test and MMK test). The PMW test was applied to detect temporal shifts in the climatic series over the study period. The Sen’s slope and percentage change in climatic variables were also estimated over the period of 100 years. The trend analysis revealed an overall upward precipitation trend even though no statistically significant trends were observed for the seasonal and annual scales. On the seasonal and annual scales, a noticeable increase was revealed in wet day frequency, surface temperature (i.e. minimum, average and maximum temperature) and reference ET during the non-monsoon season based on the results of the three non-parametric statistical tests at the 10% significance level, while maximum temperature experienced statistically significant decrease based on the MK–PW test). No change in average diurnal temperature was observed over the period of 100 years in the study area. This type of information should be helpful in facilitating a transition to more sustainable and adaptive water resources planning and management in the Ajmer District of India. Ultimately, this will help policy makers and scientists to focus on district scale planning measures for climate change adaptation and mitigation, by considering regional and local scale variability in trends compared to global climatic trends.

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Analiza trendu czynników klimatycznych na suchych i półsuchych obszarach dystryktu Ajmer w Radżasthanie, Indie

STRESZCZENIE

Słowa kluczowe: nachylenie Sena, test Manna–Kendalla (MK), test Pettitta–Manna–Whitneya (PMW), trendy, Radżasthan, zmodyfikowany test Manna–Kendalla (MMK)

W pracy analizowano trendy i zmienność czynników klimatycznych (opad, częstotliwość dni wilgotnych, temperaturę powierzchni ziemi, temperaturę dobową, zachmurzenie oraz ewapotranspirację wskaźnikową i potencjalną) w skali sezonowej i rocznej w dystrykcie Ajmer, w Radżasthanie (Indie). Analizę przeprowadzono za pomocą nieparametrycznych technik statystycznych Manna–Kendalla (MK) i zmodyfikowanej techniki MK (MMK) dla 100-letniego okresu. Test MK z eliminacją korelacji serii klimatycznych (prewhitening – MK–PW) zastosowano także do zmiennych klimatycznych, a wyniki porównano z uzyskanymi z użyciem testów MK i MMK, co pozwoliło na ocenę wiarygodności wykrywania trendu zmian w czasie.

W celu wykrycia czasowych przesunięć serii klimatycznych zastosowano test Pettitta–Manna–Whitneya (PMW). Na podstawie analizy trendu stwierdzono, że opady roczne i sezonowe nie wykazywały statystycznie istotnego trendu na poziomie istotności 10%. Wykorzystując trzy testy nieparametryczne, stwierdzono rosnący trend w przypadku częstotliwości występowania wilgotnych dni, temperatury powierzchni i ewapotranspiracji wskaźnikowej w okresie pozamonsunowym na poziomie istotności 10%. Statystycznie istotny spadek maksymalnej temperatury w tym okresie stwierdzono jedynie, gdy stosowano test MK–PW. Przedstawiona analiza kilku zmiennych klimatycznych w skali dystryktu może być pomocna w planowaniu i zarządzaniu zasobami wodnymi i w rozwoju strategii adaptacji do niekorzystnych warunków klimatycznych.