Alpine ice and the annual political economy of the Angevin Empire, from the death of Thomas Becket to Magna Carta, c. AD 1170–1216

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High-resolution analysis of the ice core from Colle Gnifetti, Switzerland, allows yearly and sub-annual measurement of pollution for the period of highest lead production in the European Middle Ages, c. AD 1170–1220. Here, the authors use atmospheric circulation analysis and other geoarchaeological records to establish that Britain was the principal source of that lead pollution. The comparison of annual lead deposition at Colle Gnifetti displays a strong similarity to trends in lead production documented in the English historical accounts. This research provides unique new insight into the yearly political economy and environmental impact of the Angevin Empire of Kings Henry II, Richard the Lionheart and John.

Introduction

Twenty years ago, Brännvall et al. (1999) published the first high-resolution evidence demonstrating that the largest-scale lead pollution in Northern Europe prior to the modern era occurred between c. AD 1150 and 1200. Their evidence came from concentration and...
isotope analysis of lead deposited in annual varve sediments from Swedish lakes (Brännvall et al. 1999: 4392–93). This peak in lead pollution almost certainly reflected later twelfth-century urban population and economic growth in Western and Central Europe. Brännvall et al.’s (1999: 4394) application of heavy metal aerosol-transport modelling suggested that the largest quantities of lead pollution had been delivered, in order of magnitude, from England, Germany, Wales and Poland.

Our study uses ultra-high-resolution evidence from the new Colle Gnifetti (CG) ice core (Switzerland), along with textual and archaeological sources, to affirm Brännvall et al.’s (1999) conclusion that England was the major contributor of pollution in the later twelfth and early thirteenth centuries AD. Detailed yearly analysis of the different sources of evidence indicates that, for the mid 1160s to 1216, the CG ice core provides a measurable, continuous record of the environmental impact and economic history of silver and lead production in England under the Angevin kings, Henry II, Richard the Lionheart and John. As a result of the way in which they affected production, key events linked to the political economy of the Angevin Empire, from warfare and castle and monastery construction to the deaths of monarchs, are inscribed into the pollution record.

The Colle Gnifetti ice-core chronology and annual lead-pollution record

The new CG ice core was extracted in 2013 from the north-facing slope of the glacier at Monte Rosa (45°55′ N, 7°52′ E), 4450m asl (Figure 1). The chronological framework for analysis of the ice core was established in two stages: firstly, through continuous-flow analysis, high-resolution discrete inductively coupled plasma mass spectrometry (ICP-MS) and ion chromatography; and secondly by annual layer-counting using ultra-high-resolution laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) (see Sneed et al. 2015; More et al. 2017; Bohleber et al. 2018; Loveluck et al. 2018). Discrete liquid-based ICP-MS measurement of the concentration of lead and other elements was also conducted along the small melted portion of the core, at an average resolution of 45mm from a water stream generated by the continuous melter system (Bohleber et al. 2018).

The LA-ICP-MS annual-layer chronology and lead-intensity record was tied to absolute chronological markers to constrain the methodological error margin. For this study, covering the mid twelfth to early thirteenth centuries, the initial absolute-dating ties were provided by recorded major Saharan dust events (Bohleber et al. 2018), although the most important absolute markers come from the AD 1349–1353 Black Death pandemic (More et al. 2017, 2018), and the eruption of the Icelandic volcano Eldgjá, in AD 939 (Oppenheimer et al. 2018). Theoretically, while there could have been a 35- to 40-year error at the point in the ice dated to 1349 by layer-counting, in practice the lead intensity plummeted to levels barely measurable precisely at the layer-counted ice dated to 1349. This independently demonstrates that there was no actual error in the annual layer-counting in the mid fourteenth century. The CG Eldgjá marker was dated by layer-counting to 945, giving a maximum error of six years from the eruption date in 939. Therefore, a six-year error margin had accumulated over the 410-year interval between 1349 and 939. Assuming an equal
Figure 1. Location of environmental records referred to in the text (figure by S. Troadec).
build-up of error, the error margin envelope encompassing this study would be 1.8 years at 1220 (±1 year), and 2.9 years (±1.5 years) at 1150.

Figure 2 shows both the LA-ICP-MS ultra-high-resolution yearly mean lead-intensity measurements (based on 100–102 measurements per year; see the online supplementary material (OSM) 1) and the lower-resolution lead concentration measurements from the discrete ICP-MS analysis, with one sample for every three to seven years. Despite the disparity in resolution, the plotted trends from the concentration data are very similar to the intensity data, although with an expected discrepancy of up to two years between the concentration measurements and the mean yearly intensity readings.

Sourcing the origins of the CG lead pollution record between AD 1150 and 1220

The decades from c. 1150–1220 constitute the period of the highest quantifiable levels of lead pollution at CG between c. AD 150 and 1650, as identified by discrete ICP-MS measurement of lead concentration (More et al. 2017: fig. 1). This peak in lead pollution c. AD 1200 was previously identified at a lower resolution in an earlier ice core from CG (Gabrieli &
Barbante 2014: 76), and in North-western and Central European peat-core records (see OSM 2). In order to exploit the new yearly lead-pollution record and go beyond existing observations of general coincidence with historic phenomena, however, we needed to identify the most probable sources of the pollution. This was achieved using textual and archaeological evidence, lower-resolution pollution records from European mining regions, and Climate Reanalyzer™ visualisation, with NOAA-CIRES climate data from 1979–2014 (Figures 1–5).

The Harz Mountains represent the mining region least likely to have contributed to the CG lead record from c. 1150–1220. Prior to early modern exploitation, the production of silver-lead concentrated in two periods in this region: from c. AD 960–1050 and c. AD 1280–1330 (Böhme 1978; Spufford 1988). The first large-scale silver-lead mining in the 960s–990s facilitated the minting of the ‘Otto-Adelheid’ pennies, with silver mainly from the Rammelsberg mine, near Goslar (Figure 3; Ehlers 1997: 48–49; Kluge 2002: 93–101). By c. 1040, these pennies were no longer struck, and Harz silver-lead production was minimal from 1050–1300 (Spufford 1988: 93–94). Archaeological evidence, however, suggests that copper- and iron-mining activities overall expanded in the Upper Harz, between c. 1100 and 1300 (Klappauf et al. 2008; Knapp et al. 2015: 52–54). A peat core from the Sonnenberger Moor in the Upper Harz shows no significant lead pollution from c. 950–1050, but does present copper pollution—and lead as a by-product—from c. 1100–1300, reflecting that localised expansion (Kempter & Frenzel 2000: 100).
The largest continental European region that produced silver and lead from the mid twelfth century was the Erzgebirge (Ore Mountains) region around Freiberg, Saxony (Figure 3). Probably discovered in 1168, this ore-field was the paramount supplier of European silver until c. 1250 (Böhme 1978: 75; Spufford 1988: 112–13). Silver production in the Erzgebirge was large-scale and continuous through the first half of the thirteenth century, contrasting with reduced pollution at CG between c. 1190 and 1250 (Figure 2; More et al. 2017).

The other major lead-silver-producing regions in Europe between c. 1150 and 1220 were located in Britain, and were generally administered under the label ‘mine’. By the 1150s, the most important of these were the mine of Carlisle, and the mines of the Derbyshire Peak District, administered from Castleton and Wirksworth. Others included the mine of Yorkshire, focused on Boroughbridge; the mine of Shropshire, administered from Shrewsbury; and the mines of the Bishops of Durham, in the northern Pennines (Figure 4). The production of lead is well documented in the audit records of royal taxation, known as the Pipe rolls (Great Rolls of the Pipe). These records were collated yearly by the Royal Treasury/Exchequer of the Angevin Kings of England from 1156, and in great detail from 1167 onwards (Barratt 2007: 248–54; Allen 2011). They record annual taxation owed and paid—known as ‘farm’ payments—for holding the franchises of the ‘mines’, as well as royal requisitions/purchases of lead by right of ‘pre-emption’ in ‘cart-loads’. The cart-load was a unit of account, averaging approximately one imperial ton, although its weight-measure varied slightly in different ore-fields (Latimer 1999: 49–50).

Continuous mining into the modern era has destroyed much of the medieval evidence at British ore-fields. This has included the systematic re-smelting of medieval slag heaps, particularly in the Peak District (Barnatt et al. 2013). Nevertheless, peat pollution records at Lindow Moss (Cheshire) and Raeburn Flow (Dumfries and Galloway) confirm that the nearby mines of Shropshire and Carlisle were producing large quantities of lead between c. 1150 and 1200, with a decline after 1200 (within their error margins) (Le Roux et al. 2004: 504–507; Küttner et al. 2014: fig 3c; see OSM 2).

Atmospheric circulation visualisation using Climate Reanalyzer™ has revealed a consistent pattern of wind delivery to the Western Alps from the north-west, especially during the summer, which coincided with the medieval lead-producing season. Previous analyses of textual and archaeological evidence for the decades AD 640–670 and 1330–1370 suggest that the pollution at CG originated mainly from western France and Britain, respectively, at those times (More et al. 2017: 214–17; Loveluck et al. 2018: 6–10). Figure 5 shows a Climate Reanalyzer plot for June, July and August, with the largest silver-lead-producing regions of Europe from c. 1150–1220 (triangle sizes on Figure 5 equate to the scale of production). The mines of Britain are indicated as the most probable sources of the pollution. Any lead aerosol from the ore-fields of the Harz and the Erzgebirge travelled east and south-east, to Poland from the Harz, and to Austria and the Balkans from the Erzgebirge. This long-term north-west to south-east wind direction, driven by the Iceland Low, is confirmed consistently by studies on the transport of Icelandic tephra across North-west Europe to Italy and Slovenia from the onset of the Holocene to the present; and by analysis of historic volcanic plume transport from Icelandic eruptions (see OSM 3). For this study, the possible influences of climatic- and wind-strength variations on quantities of lead aerosol delivered to CG, within

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Figure 4. Location of places in Britain discussed in the text (figure by S. Troadec).
its predominantly spring-to-summer record (Bohleber et al. 2018), have been addressed by taking the mean value of the 100–102 measurements per year, between AD 1150 and 1220. Smaller-scale lead production in southern France could also have fed the CG lead profile. The Narses Mortes peat core (Mont Lozère, Cévennes), for example, has a lead peak from c. 1156–1264 (at 95.4 per cent confidence), and nearby smelting sites have a similar date range (Baron et al. 2005, 2006). Unlike the CG record, however, this localised production did not decline from 1200–1250. Lead pollution also increased from c. AD 1000–1200 at Étang-de-la-Gruère (Jura), and origins were suggested from sources in Germany (but not the Harz), using lead isotopes (Shotyk et al. 1998: 1637–40). Production, however, was not local; non-German sources were not considered by Shotyk et al. (1998), and lead isotope characteristics often overlap between ore-fields (Blichert-Toft et al. 2016: S1). Copper and lead-silver mining/smelting sites in the Chassezac Valley (Ardèche) also date from the eleventh to thirteenth centuries, but were small in scale compared to British and Erzgebirge production; its pollution would have travelled south-east over the Mediterranean (Bailly-Maître

Figure 5. Climate Reanalyzer™ plot for June, July and August 1984, showing predominant directions of wind delivery to Colle Gnifetti and principal lead-producing regions, AD 1150–1220 (figure by A. More & the Climate Change Institute, University of Maine).
et al. 2013; see OSM 3). Contemporaneous evidence for lead-silver extraction is also absent in French regions that became major producers from the later Middle Ages, such as Sainte-Marie-aux-Mines, Vosges, and Castel-Minier, Ariège (Téreygeol 2009: 84–85; Flamant et al. 2019). Environmental records from other areas (e.g. Lorraine, the Eifel region, the Pyrenees or Spain), dating from 1150–1200, have yielded either no pollution, or very limited localised activity (see OSM 2; Figures 1 & 3).

A yearly history of the political economy of the Angevin Kings, AD 1167–1216

The contextual evidence all suggests that the ore-fields of Britain were the principal source of the lead pollution at CG from 1150–1220. The yearly mean LA-ICP-MS lead record from CG was therefore compared with the only other detailed yearly lead record from this period: the English Pipe rolls from 1167–1216.

Two principal proxies were selected from the Pipe rolls as probable reflections of the production of lead, and also silver in the case of the ‘mine of Carlisle’. Firstly, the ‘farm’ franchise payments in silver were quantified for the two main lead-producing regions: the ‘mine of Carlisle’ and the mines of the Derbyshire Peak District. As demands were often far beyond what franchise holders could pay, farm payments were quantified based on what was actually paid to the Treasury each year. Arrears payments for earlier years were quantified in the year that they were accounted in the rolls. The aim was to examine the pattern of farm payments against levels of pollution per year; logically, larger yearly silver payments might reflect larger-scale silver and lead production (Claughton 2003; Allen 2011: 116). Secondly, the number of cart-loads of lead were quantified by year as cumulative totals sent from all the mines combined (individual farm and cart-load totals by year for different ‘mines’ are recorded in OSM 4–8; all numerical data are presented in OSM 9). These could then be compared with the CG yearly mean pollution record.

Henry II obtained the mine of Carlisle from the King of Scotland in 1157, after it had been ceded to Scotland in 1136 (Carpenter 2003: 195; Mattinson & Cherry 2013). The mine’s principal silver and lead source was in the northern Pennines on Alston Moor, near Nenthead, and Pipe roll evidence and surviving coinage suggest that operations expanded considerably under Henry. Opinions have differed on the importance of silver from the mine of Carlisle to overall coinage supply in later twelfth-century England, with major and lesser contributions suggested (see Claughton 2003; Allen 2011). Erzgebirge silver would certainly have arrived via trade in commodities such as English wool and tin (Bartlett 2000: 367–68). The Carlisle mine, however, was unquestionably the largest indigenous supplier of silver between 1158 and 1216, and a desire for increased silver extraction from larger quantities of its low-grade ore would have produced much larger quantities of both lead aerosol and metal as by-products.

The establishment of potential proxies for Peak District production is more complex, as they could be hidden under different farm payment labels. Between the 1150s and early 1200s, only lead production is recorded in the Pipe rolls. Most of the mining/smelting seems to have taken place within the Wapentake (administrative territory) of Wirksworth in the Low Peak. Farm payments for the ‘mines of Derbyshire’ were returned from 1156–
1158, and for the ‘mine of the Peak’ in 1172. Cart-loads of lead from Derbyshire are listed between 1172 and 1178, again presumably from Wirksworth Wapentake, as High Peak cart-load production, administered from Castleton, was specifically differentiated (e.g. in 1188; Figure 6 & OSM 5).

From 1169 onwards, regular farm payments were made by the land-holders of Wirksworth Wapentake (see OSM 5). According to the customary rights of lead mining in Derbyshire, administered from Wirksworth and first written down in 1288, anyone could mine for lead, but they paid every thirteenth ‘dish’ of ore to landowners/farm holders (Rieuwerts 1988: 15). The paramount social organisation of medieval lead-mining across England was for free or tenant farmers to mine and smelt lead in the summer months, between the two harvests of winter and spring crops, with ore renders and cash payments in coin made to farm-holders (Blanchard 1972). Although the latter payments were generated predominantly from trade in lead, which could create a slight delay between mining/smelting and generation of silver coin, the importance of lead-related income for peasant-farmer economies kept delays minimal (Blanchard 1972: 99–101; Latimer 1999). Given the importance of lead in Wirksworth Wapentake, it seems logical to suggest that levels of lead production may be reflected in the collective farm payments made for the Wapentake. Hence, these payments, and those of the mines of Derbyshire, the mine of the Peak, and High Peak cart-loads, were quantified together under ‘Wirksworth Wapentake’, as a proxy for Peak District production.

To attempt an overall proxy representation of lead production from 1167–1216, Figure 7 combines the yearly cart-load totals of lead from all British mines and farm payments in silver from Carlisle and Wirksworth Wapentake from 1167–1216. These data are plotted alongside the CG yearly mean lead-pollution record, and reflect the year when farm payments and cart-loads were sent to their destinations. The payments for 1182 and 1183 from Carlisle were both recorded in 1183 (see OSM 4), but have been plotted separately in their respective years. When large cart-load shipments were recorded in a single year (in lieu of farm payments for one or more years), the totals have been averaged over the years of production: in 1180–1181, for cart-loads from all British mines and for Carlisle cart-loads; and from 1208–1211 for Durham cart-loads. Indeed, production for Durham is explicitly identified in the 1211 roll (see OSM 8–9). There were no ice-core measurements for 1171, as the core was cut at
that point for transport, and no Pipe rolls survive for 1213 and 1216. Mid-point proxies between values from adjacent years were therefore allocated at these points.

Figure 8a–b shows the combined annual farm payments from Carlisle and Wirksworth and total cart-loads separately. Figure 9a–b shows the individual trends from the mine of Carlisle and Wirksworth Wapentake.

Statistical correlations are inappropriate for this type of interdisciplinary study. This is because of the comparison of combined quantified totals of raw numbers from very different types of data (‘tower pounds’ of silver and ‘tons’ of lead) against mean average totals of 100–102 lead-intensity readings per year, and the occasional missing and combined records. Instead, we use a qualitative comparative approach, employed previously for comparing volcanic eruption events to records of cold episodes in textual sources (Ludlow et al. 2013), and for comparing temperature and precipitation proxies from dendro-chronological data with annual historical records (Huhtamaa 2015: 570–75). Unlike with scatter-plot correlative analysis, this comparative approach takes into account the different forms of data and the influences at work on them. It has been the standard approach for recent studies comparing ice-core pollution studies from Alpine and Greenland glaciers with historical and archaeological records (McConnell et al. 2019; Preunkert et al. 2019). Our comparison of the
Figure 8.  a) Sum of farm payments of the mine of Carlisle and Wirksworth Wapentake combined; b) total production of lead cart-loads per year, and the Colle Gnifetti yearly lead record, AD 1167–1216 (figure by C. Loveluck & A. More).
Figure 9. Farm payments and cart-loads, AD 1167–1216, from the mine of Carlisle (a) and Wirksworth Wapentake (b), alongside the Colle Gnifetti yearly lead record (figure by C. Loveluck & A. More).
similarity in rises and falls (peaks and troughs) in the CG lead record, farm payments and cart-loads on a yearly basis between AD 1167 and 1216 (see OSM 10) shows an 89 per cent match between British lead production and the CG pollution record, with individual similarities of 72 per cent with Wirksworth and 57 per cent with Carlisle (see OSM 10a).

Notably, falls in pollution coincide with major political changes in Britain (Figure 7). Troughs in lead deposition, for example, correspond to the deaths and successions of Angevin monarchs, from Henry II to Richard I the Lionheart in 1189, Richard to John in 1199, and John to Henry III in 1216. The lack of royal demand during interregna and occasional uncertainty over succession resulted in reduced production and minimal farm payments in those years. Sizable payments were made in 1190 and 1200, once the successions were secure, and these payments match increases in lead pollution. Lead deposition also fell at CG during political crises or royal absences. Examples include 1170, during Henry II’s dispute with the Church, leading to the murder of Thomas Becket, Archbishop of Canterbury; 1173–1174, during civil war and conflict with Scotland (see Figure 9a–b); 1191–1192, when Richard the Lionheart was co-leading the third crusade; and 1215–1216, during the civil war when John signed Magna Carta, and the French invaded England.

The farm payments mirror the CG lead record during these political upheavals. No farm payment was made in 1170 for the mine of Carlisle, although it did pay a large sum in 1171. It also made no payments in 1174–1175, following the King of Scots lands invasion of Northumberland and the region around Carlisle in 1173 and 1174 (Figure 9a). Payments were also minimal in the years that Richard the Lionheart was away on crusade from 1191–1193, and ceased in 1215–1216, following Carlisle’s annexation by the Scots (Carpenter 2003). Wirksworth farm payments were also minimal from 1196–1199, when Richard was at war in Normandy; and no payments were made during the civil war and French invasion of 1215–1216 (Figure 9b).

The mine of Carlisle made larger annual farm payments during the reign of Henry II, although they had declined by the reign of John. Wirksworth payments were generally smaller until the reign of Richard (1190s), but by the reign of John, they were consistently higher than those of Carlisle. This suggests that the mine of Carlisle was the largest consistent lead and silver producer between the mid 1160s and 1190s, while the Peak District became the paramount producer by the mid 1190s. Surviving records of the Tideslow mine in the Peak District reflect this trend, with annual production levels in 1195 double those of the period from 1216–1249 (Rieuwerts 1988: 16; Ford & Rieuwerts 2000: 20). CG lead pollution seems to reflect the combined lead production illustrated by both sets of farm payments, certainly from 1169 onwards (Figure 8a), albeit with a closer correspondence with Wirksworth production overall. Farm payment peaks usually align directly with pollution increases, as in 1169, 1200 and 1214. During extended peaks in lead pollution, such as from 1193–1197, however, farm payment peaks were not always continuous. There was a Wirksworth farm peak in 1194–1195 and a Carlisle farm peak in 1197. For the 1207–1209 lead-pollution peak, there is a Carlisle farm peak in 1207 and a Wirksworth peak in 1209, as Carlisle payments stopped during warfare with Scotland in 1208–1209 (Figure 9a–b).

The cart-load data reflect lead requisitioned, often in lieu of farm payments, from all the mining regions from 1168 onwards. Cart-loads were frequently sent to their destinations at a moment that coincides with the end of CG lead-pollution spikes, or during a slight drop in
pollution the following year (Figure 8b). This possibly reflects a time lag between mining/smelting and dispatch, due to limitations of productive capacity and transport infrastructure to meet large-scale demand. The cart-loads were ordered for major construction projects in England and France, sometimes beyond Angevin territories. Henry II, for example, sent cart-load shipments to Caen (Basse-Normandie) in 1168; to Clairvaux Abbey (Aube), from 1179–1188; to Winchester Palace (Hampshire) in 1172; to Woodstock Palace (Oxfordshire) in 1176; to Waltham Abbey (Essex) in 1181–1183; to Amesbury Abbey (Wiltshire) in 1181–1184; and to Gisors Castle (Vexin) in 1184 (see OSM 4–7). In 1176, Henry purchased lead worth 40 pounds for Grandmont Abbey (Limousin) from Carlisle. All the shipments to the Continent came from the mines of Carlisle, Yorkshire and High Peak, and were transported via the Rivers Tyne and Humber, and the North Sea, Channel and Biscay coasts to Rouen and La Rochelle (Figure 3; see OSM 4–6).

It is sometimes possible to specify the time lag between production and arrival at continental destinations. The surviving Norman Pipe rolls show that the 100 cart-loads produced at Carlisle in 1179 for Clairvaux Abbey had arrived in Rouen in 1180, for onward transport along the River Seine via Paris (Moss 2004: 51). The 20 cart-loads sent from Yorkshire to Gisors Castle in 1184 are recorded as having arrived in Rouen in the same year (Moss 2004: 79). The time lag between production and arrival of these shipments reflects the capacity of what could be produced and transported within one year. Although there were no cart-load shipments recorded in the reign of Richard, an unspecified quantity of lead was procured for castle construction, for example, at Château Gaillard (Eure) in 1197–1198 (Moss & Everard 2016: 144–46). The CG pollution rise from 1194–1196, and the large farm payments from Wirksworth in 1194 and Carlisle in 1197, may reflect the production phase of the lead used in Richard’s castle-building in Normandy (Figure 9a–b). Overall, however, it is notable that no cart-loads of lead are recorded as travelling from England to continental Europe during the wars of 1173–1174, 1194–1199 and 1203–1207 (Figure 8b).

The coin reforms of Henry II and John are not clearly reflected in the CG lead-pollution record, or in the lead-production proxies from the Pipe rolls. Henry’s reforms in 1158–1159, 1180–1182 and 1188 (Allen 2007; Bolton 2012: 146–47) would have involved melting down old coinage and re-striking new currency, releasing lead-aerosol pollution. The 1158–1159 and 1188 reforms are not reflected at CG, while the 1180–1182 reform might have contributed to the 1180–1181 lead peak. This, however, also coincides with Henry’s period of maximum acquisition of metal for construction. John conducted a re-coinage from 1205–1207 (Allen 2011: 117), but, again, this is not clearly reflected in the CG core. All of these examples suggest that the correspondence between the CG lead record and the English Pipe rolls predominantly reflects primary metal production and the linked effects of political events.

Conclusions

The correspondence between the CG lead-pollution profile and English lead production under the Angevin kings is striking. Detailed analysis of textual, archaeological and geo-archaeological records across Europe, together with analysis of atmospheric circulation, using Climate Reanalyzer™ and tephra, confirms Britain as the principal source of the lead
pollution between c. 1167 and 1216. The 89 per cent correspondence between the CG lead record and annual British lead production recorded in the Pipe rolls, including markers for royal successions, wars and major construction projects, provides further historical anchors for the LA-ICP-MS-layer-counted chronology at CG, in addition to Saharan dust events, tephra and the 1349–1353 Black Death lead record that also had a British source. The interdisciplinary results presented in this study at annual resolution are much less likely to represent chance coincidence than pollution records from peat cores, ice cores with wider radiocarbon-dating errors, or ice-core pollution records not yet subjected to detailed analysis of pollution sourcing and atmospheric circulation incorporating tephra studies (see OSM 3).

The current study provides a unique insight into the relationship of the Angevin royal government to metal production and the wider economy in Britain and western France, at a critical moment for European state formation. The impact of political events, such as warfare, the deaths and successions of monarchs and disputes between Church and State, are all reflected in the CG glacier. The results illustrate with remarkable clarity the impact of political (in)stability and transport security on specialist commodity production and movement in one of Europe’s most powerful political hegemonies during the later twelfth century.

Acknowledgements

This is contribution 13 of the ‘Historical Ice Core’ project funded by Arcadia (grants AC3862 & AC4190), a charitable fund of Lisbet Rausing and Peter Baldwin. We gratefully acknowledge their support. C.P.L. would like to thank the History Department and the Initiative for the Science of the Human Past at Harvard University, Dumbarton Oaks and the University of Nottingham. All ice-core analyses were conducted at the W.M. Keck Laser Ice Facility, Climate Change Institute, funded by the W.M. Keck foundation and the National Science Foundation (PLR-1042883, PLR-1203640, PLR-1142007). Thanks are also due to Solenn Troadec (for Figures 1, 3 & 4) and the Peak District National Park Authority (for Figure 6).

Supplementary material

To view supplementary material for this article, please visit https://doi.org/10.15184/aqy.2019.202

References

ALLEN, M. 2007. Henry II and the English coinage, in C. Harper-Bill & N. Vincent (ed.) Henry II: new interpretations: 257–77. Woodbridge: Boydell.

2011. Silver production and the money supply in England and Wales, 1086–c. 1500. Economic History Review 64: 114–31. https://doi.org/10.1111/j.1468-0289.2010.00552.x

BAILLY-MAÎTRE, M.-C., N. MINIVIELLE-LAROUSSE, E. KAMMENTHALER, T. GONO & G. GUIONOVA. 2013. L’exploitation minière dans la vallée du Chassezac (Ardèche): le plomb, l’argent et le cuivre au Moyen Âge (XIe–XIIIe siècle). Archéologie Médiévale 43: 47–76. https://doi.org/10.4000/archeomed.9655

BARNATT, J., K. HUSTON, D. MALLON, R. NEWMAN, R. PENNY & R. SHAW. 2013. The lead legacy: an updated inventory of important metal and gangue mining sites in the Peak District. Mining History: Bulletin of the Peak District Mines Historical Society 18(6): 1–112.
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BARON, S., M. LAVOIE, A. PLOQUIN, J. CARIGNAN, M. PULIDO & J.-L. DE BEAULIEU. 2005. Record of metal workshops in peat deposits: history and environmental impact on the Mont Lozère Massif, France. Environmental Science and Technology 39: 5131–40. https://doi.org/10.1021/es048165l

BARON, S., J., CARIGNAN, S. LAURENT & A. PLOQUIN. 2006. Medieval lead making on Mont Lozère Massif (Cévennes, France); tracing ore sources using Pb isotopes. Applied Geochemistry 21: 241–52. https://doi.org/10.1016/j.apgeochem.2005.09.005

BARRATT, N. 2007. Finance and the economy in the reign of Henry II, in C. Harper-Bill & N. Vincent (ed.) Henry II: new interpretations: 242–56. Woodbridge: Boydell.

BARTLETT, R. 2000. England under the Norman and Angevin Kings 1075–1225. Oxford: Clarendon.

BLANCHARD, I. 1972. The miner and the agricultural community in late medieval England. Agricultural History Review 20: 93–106.

BLICHERT-TOFT, J., H. DELILE, C.-T. LEE, Z. STOS-GALE, K. BILLSTRÖM, T. ANDERSEN, H. HANNU & F. ALBARÈDE. 2016. Large-scale tectonic cycles in Europe revealed by distinct Pb isotope provinces. Geochemistry, Geophysics, Geosystems 17: 3854–64. https://doi.org/10.1002/2016GC006524

BOHLEBER, P., T. ERHARDT, N. SPAULDING, H. HOFFMANN, H. FISCHER & P.A. MAYEKOLKI. 2018. Temperature and mineral dust variability recorded in two low accumulation Alpine ice cores over the last millennium. Climate of the Past 14: 21–37. https://doi.org/10.5194/cp-14-21-2018

BÖHME, H.W. 1978. Der Erzbergbau im Westharz und die Besiedlung des Oberharzes seit dem frühen Mittelalter, in Westlicher Harz. Clausthal-Zellerfeld, Osterode, Seesen (Führer zu vor- und frühgeschichtlichen Denkmälern 36): 59–126. Mainz: von Zabern.

BOLTON, J.L. 2012. Money in the medieval English economy 973–1489. Manchester: Manchester University Press.

BRÄNNVALL, M.-L., R. BINDLER, I. RENBERG, O. EMTERYD, J. BARTNICKI & K. BILLSTRÖM. 1999. The medieval metal industry was the cradle of modern large-scale atmospheric lead pollution in Northern Europe. Environmental Science and Technology 33: 4391–95. https://doi.org/10.1021/es990279n

CARPENTER, D. 2003. The struggle for mastery: Britain 1066–1284. London: Allen Lane/Penguin.

CLAUGHTON, P. 2003. Production and economic impact: northern Pennine (English) silver in the 12th century, in Proceedings of the 6th International Mining History Congress, September 2003, Akabira City, Hokkaido, Japan: 146–49. Hokkaido: International Mining History.

EHLERS, C. 1997. Die Anfänge Goslars und das Reich im elften Jahrhundert. Deutsches Archiv für Erforschung des Mittelalters 53: 45–79.

FLAMENT, J., F. MERCIER, C. DUBOIS & F. TÉREVGEOL. 2019. Mining archaeology and micro-raman analysis associated with ESEM-EDX: toward a chrono-spatial definition of ore consumption in a Pyrenean medieval workshop. Archaeometry 61: 99–116. https://doi.org/10.1111/arcm.12404

FORD, T.D. & J.H. RIEUWERTS. 2000. Lead mining in the Peak District. Matlock: Peak District Mines Historical Society.

GABRIELI, J. & C. BARRANTE. 2014. The Alps in the age of the Anthropocene. Rendiconti Lincei: Scienze fisiche e naturali 25: 71–83. https://doi.org/10.1007/s12210-014-0292-2

HUHTAMAA, H. 2015. Climate anomalies, food systems, and subsistence crises in medieval Novgorod and Ladoga. Scandinavian Journal of History 40: 562–90. https://doi.org/10.1080/03468755.2015.1063081

KEMPTER, H. & B. FRENZEL. 2000. The impact of early mining and smelting on local tropospheric aerosol detected in ombrotrophic peat bogs in the Harz, Germany. Water, Air and Soil Pollution 121: 93–108. https://doi.org/10.1023/A:1005253716497

KLAPPAU, L., C. BARTELS, F.-A. LINKE & B. ASMUS. 2008. Das Montanwesen am Rammelsberg und im Westharz – Historische und archäologische Quellen zum ’12.’ und ’13.’ Jahrhundert, in M. BRANDT (ed.) Bild und Bestie: Hildesheimer Bronzen der Stauferzeit: 65–76. Hildesheim: Schnell & Steiner.

KLUGE, B. 2002. OTTO REX/OTTO IMP. Zur Bestandsaufnahme der ottonischen Münzprägung, in B. SCHNEIDMÜLLER & S. WEINFURTER (ed.) Ottonische Neuanfänge: 85–112. Mainz: von Zabern.

KNAPP, H., O. NELLE, W. KIRLEIS. 2015. Charcoal usage in medieval and modern times in the Harz
Mountains area, central Germany: wood selection and fast overexploitation of the woodlands. *Quaternary International* 366: 51–69. https://doi.org/10.1016/j.quaint.2015.01.053

Küttner, A., T.M. Mighall, F. de Vleeschouwer, D. Mauquoy, A. Martínez Cortizas, I.D.L. Foster & E. Krupp. 2014. A 3300-year atmospheric metal contamination record from Raeburn Flow raised bog, south-west Scotland. *Journal of Archaeological Science* 44: 1–11. https://doi.org/10.1016/j.jas.2014.01.011

Latimer, P. 1999. Early thirteenth-century prices, in S.D. Church (ed.) *King John: new interpretations*. 41–73. Woodbridge: Boydell.

Loveluck, C.P. et al. 2018. Alpine ice-core evidence for the transformation of the European monetary system, AD 640–1635. *Antiquity* 92: 1571–85. https://doi.org/10.15184/ajy.2018.110

Ludlow, F. et al. 2013. Medieval Irish chronicles reveal persistent volcanic forcing of severe winter cold events, 431–1649 CE. *Environmental Research Letters* 8: 024035. https://doi.org/10.1088/1748-9326/8/2/024035

Mattinson, J. & P. Cherry. 2013. The Carlisle mint coinages of Henry I, Stephen, David I and Earl Henry. *British Numismatic Journal* 83: 101–11.

 McConnell, J.R. et al. 2019. Pervasive Arctic lead pollution suggests substantial growth in medieval silver production modulated by plague, climate and conflict. *Proceedings of the National Academy of Sciences of the USA* 116: 14910–15. https://doi.org/10.1073/pnas.1904515116

More, A.F. et al. 2017. Next-generation ice core technology reveals true minimum natural levels of lead (Pb) in the atmosphere: insights from the Black Death. *GeoHealth* 1: 211–19. https://doi.org/10.1002/2017GH000064

– 2018. The role of historical context in understanding past climate, pollution and health data in trans-disciplinary studies: reply to comments on More et al. 2017. *GeoHealth* 2: 162–70. https://doi.org/10.1029/2017GH000121

Moss, V. (ed.). 2004. *Pipe rolls of the exchequer of Normandy for the reign of Henry II, 1180 and 1184* (New Series 53). London & Loughborough: Pipe Roll Society.

Moss, V. & J. Everard (ed.). 2016. *Pipe Rolls of the exchequer of Normandy for the reign of Richard I, 1194–95 & 1197–98* (New Series 60). Leicester: Pipe Roll Society.

Oppenheimer, C., A. Orchard, M. Stoffel, T.P. Newfield, S. Guillet, C. Corona, M. Sigl, N. di Cosmo & U. Büntgen. 2018. The Eldgjá eruption: timing, long-range impacts and influence on the Christianisation of Iceland. *Climatic Change* 147: 369–81. https://doi.org/10.1007/s10584-018-2171-9

Preunkert, S. et al. 2019. Lead and antimony in basal ice from Col du Dome (French Alps) dated with radiocarbon: a record of pollution during Antiquity. *Geophysical Research Letters* 46: 4953–61. https://doi.org/10.1029/2019GL082641

Rieuwerts, J.H. 1988. *A history of the laws and customs of the Derbyshire lead mines*. Sheffield: Derbyshire Barmote Court.

Le Roux, G., D. Weiss, J. Grattan, N. Givelet, M. Krachler, A. Cherbukin, N. Rausch, B. Kober & W. Shotyk. 2004. Identifying the sources and timing of ancient and medieval atmospheric lead pollution in England using a peat profile from Lindow Bog, Manchester. *Journal of Environmental Monitoring* 6: 502–10. https://doi.org/10.1039/B401500B

Shotyk, W., D. Weiss, P.G. Appleby, A.K. Cherbukin, R. Frei, M. Gloor, J.D. Kramers, S. Reese & W.O. van der Knaap. 1998. History of atmospheric lead deposition since 12,370 14C yr BP from a peat bog, Jura Mountains, Switzerland. *Science* 281: 1635–40. https://doi.org/10.1126/science.281.5383.1635

Snedc, S. et al. 2015. New LA-ICP-MS cryocell and calibration technique for sub-millimeter analysis of ice cores. *Journal of Glaciology* 61: 233–42. https://doi.org/10.3189/2015JoG14J139

Spufford, P. 1988. *Money and its use in medieval Europe*. Cambridge: Cambridge University Press. https://doi.org/10.1017/CBO9780511583544

Téreygeol, F. 2009. Aux sources des matériaux: les mines de plomb, in A. Timbert (ed.) *L’homme et la matière: l’emploi du plomb et du fer dans l’architecture gothique*. 83–88. Paris: Picard.