Changes in particle sizes and geochemistry of Siyah Keshim lagoon sediment of Gilan province to determine origin and tectonic position of sediment

Saba Saanei¹, Khalil Rezaei², Mehran Arian³, Mohsen Aleali⁴, Pantea Giyahchi⁵

¹Islamic Azad University, Tehran, Iran
²Kharazmi University of Tehran, Iran, drkhalilrezaei@gmail.com
³Giyahchi Center of Geoscience, Tehran, Iran, pgiyahchi@gmail.com

Abstract. Useful information was obtained about the environmental condition of this region such as energy, sediment location, origin, sediment carrying path, pond evolutionary process, and tectonic conditions and origins of sediments by studying physical and geochemical sediment changes processes in place and time, distribution of sediments and elements in lagoon bed, and also identification the effective factors on sediment distribution model. In this regard, 59 sediment samples were taken from Siyah Keshim lagoon and were analyzed chemically and aggregation by XRF and ICP-MS technic. Adaptation of data by Folk diagrams showed that this region is placed in a range of sand, muddy sand, sand with a little gravel, muddy sand with a little gravel, and silty sand. Weak to medium sorting and negative tilting shows sediment in a coastal area. Geochemical evidence shows that SiO₂/Al₂O₃ ratio in these sediments is 2.6-3.7. In addition, Na₂O value shows relatively low sediment investigation for sediments of this lagoon. Moreover, determination of the weathering effect in origin place by the chemical index of alteration (CIA) and chemical weathering index (CWI) shows intensive chemical alternation on sediments. Index of combined variety was averagely 1.54 in the studied sediments and show that sediments resulted from the first cycle sediment. Using tectonic separating diagrams based on the primary and secondary oxidents percentage show the sediment in the active continent margin (ACM), continent-island arc (CIA), and oceanic island arc (OIA) and shows that the studied sediments are related to subduction margin.

Key words: Siyah Keshim, lagoon, sorting, sediment, geochemistry

Використання змін розмірів частинок та геохімії осадів лагуни Сія Кешим провінції Гілан для визначення походження та тектонічного положення осадів

Саба Санаеї¹, Халил Резаєї², Мехран Ариан³, Мохсен Альялі⁴, Пантея Джахчі⁵

¹Ісламський університет Азод, Тегеран, Іран
²Університет імені Хорезмі, Тегеран, Іран, drkhalilrezaei@gmail.com
³Гуйлянський Центр наук про Землю, Тегеран, Іран, pgiyahchi@gmail.com

Анотація. Отримано корисну інформацію про стан навколишнього середовища дослідженої регіону, зокрема щодо енергії, поширення осадів, їх походження, шляхів перенесення відкладів, процесу еволюції водоймища, тектонічних умов та походження відкладів шляхом вивчення фізичних та геохімічних процесів зміни осадів у місці та часі, розподілу відкладів та елементів у річці лагуни, а також виявлення ефективних факторів на моделі розподілу осадів. Для цього із лагуни Сія Кешим було взято 59 зразків осаду, які були проаналізовані хімічно та агрегатно за допомогою технік XRF та ICP-MS. Адаптація даних за діаграмами Фолка показала, що ця область розміщена в діапазоні піску, мулистого піску, піску з невеликою кількістю гравію, мулистого піску з невеликою кількістю гравію та мулистого піску. Від слабкого до середнього сортування та негативного націлив свідчить, що осад формувався у прибережній зоні. Геохімічні дані показують, що співвідношення SiO₂/Al₂O₃ у цих відкладах становить 2.6-3.7. Крім того, значення Na₂O свідчить про відносно низьке дослідження осадів для відкладів цієї лагуни. Більше того, визначення ефекту звітної в місці походження за хімічним індексом зміни (CIA) та індексом хімічного звітної (CWI) показує інтенсивне хімічне змінення відкладів. Індекс комбінованого різноманіття становить у середньому 1.54 у досліджуваних відкладах і показує, що відклади були результатом осаду першого циклу. Використовуючи тектонічні розділові діаграми на основі процентного вмісту первинних та вторинних окисників, можна побачити наявність осадів на ділянці активного континенту (ACM), дуже континент-острів (CIA) та острів-океанічний дію (OIA) та показує, що дослідженні відклади пов’язані із субдукційною границею.

Ключові слова: Сія Кешим, лагуна, сортування, осад, геохімія
Introduction. Lagoons have many advantages for having hydrological characteristics. Lagoons not only supply a part of underground waters, but also provide an environment as the best ecosystem to supply many birds and aquatics foods (Nuri, 2009). In addition, lagoon prevents penetration of salty waters and floods prevalence by setting water flow and has a significant role in preventing desert (Nuri, 2009). In addition to the mentioned cases, lagoons have a significant role in preventing erosion of coastal lines, stabilization of the regions air based on rainfall and ambient temperature, livestock feed, agricultural usages, minerals extraction such as potassium and phosphorus, and air purification of adjacent areas, etc. (Mohammadi et al., 2012). Anzali lagoon is one of the international protected lagoons that were studied by researchers (Amirn-Ranjbar, 1998; Riyahi et al., 2005, Sharifi, 2006). This lagoon is an ecosystem of many endangered species and has a more effective role in preventing flood and acts such as sediment traps (Kazanci et al, 2004).

Anzali lagoon was made because of coastal sediment growth as coastal tabs and dams. The expansion of sand stripe masses separated a part of coastal waters that finally formed coastal back lakes after enclosure. Alzali lagoon is one of these coastal back lakes that formed a part of Caspian Sea coastal waters in the past. This lagoon developed a lot in the past but was filled gradually by alluvial-delta sediment of Sefidrud branches and rivers of Rasht, Fuman and Masal region. Anzali lagoon is the tectonic resulted lethargy that became separated from Caspian seawater recession and formed by the sand blade (the area between Anzali and Kapurchal) caused by sea movements and its waves (Ranjbar, 2012).

Siyah Keshim area that is in the southwest of lagoon has the average depth of 1m and us mostly covered by vegetable plants, especially reed, Cattail, and Azolla. The relationship between Anzali lagoon and Siyah Keshim area is possible through the narrow strait in its northeast (Mohammadi et al., 2005).

Sediment granulation is used as the most principal and important characteristics of sediment particles to discern and analyze particle size and classification of sediment environment, sediment origin, transmission history, and erosion conditions (Bui et al, 1990; Folk & Ward, 1957; Friedman, 1979). Since the sediment particles are in varied sizes and can be observed in various sizes, their granulation is based on their highest diameter that was produced by Adon and Wentworth for the first time. Wentworth scale is a logarithmic scale in which each grade is twice bigger than the next smaller grade. Phi scale is shown by $\Phi$ is the changed form of Wentworth scale that was invented by Chrombin in 1937, and scaling boundaries must change to $\Phi$ values.

$\Phi = - \log_{2}d$ where $d$ – particle diameter (mm)

Sediment was geochemically studied in the next step to fill and confirm the obtained information from granulation. Using the geochemical element to determine the old environment conditions is significantly important, and distribution of the main and secondary elements in sediment depends on mineralogical composition, temperature, fluid composition, oxidation, and reduction conditions (Adabi, 2004). Studies show that elements amounts and compositions are sensitive to environmental conditions, sai. Analysis of the main (Mg, Ca) and secondary (Sr, Mn, Na, Fe) elements of sediments open a way of understanding sedimentary environmental time by indicating their dispersion and distribution. The element geochemistry for the present elements in each sample is a good guidance for climate changes because some elements in the soil are sensitive to environmental conditions in a way that some elements report the certain conditions; for example, high amounts of Fe, Mn shows reduction environment, (Nekukhu, 2003). The elements geochemistry can show temperature and raining conditions well. Elements of Mn, Sr, Mn replace with elements such as Ca under the specific climate conditions in crystal network. Changes in Mg/Ca, Sr/Ca, Mg/Ca ratios along cores show changes in lake conditions chronically (Taghavi et al., 2013).

Methods and materials. Geographical conditions and geometry of the studied region. Siyah Keshim lagoon of Gilan province (Fig. 1) is limited to Anzali city from north to Sowme’eh Sara city from south, to Rasht city from east, and to Kapurchal and Anzali waterfront from the west.

Based on the geographical map studies of Bandar-e Anzali in 1:100000 from publications of Geological Survey & Mineral Explorations of Iran (GSI) Tehran (Khabaznia et al., 2005) of Anzali lagoon region contains rocks of Paleozoic and Caucasian periods and has specific characteristics, but Siyah Keshim lagoon is very new geologically and its creation can be known for about 7 centuries ago.

The altitudes of the catchment area of this wetland are related to Cretaceous to Jurassic courses and also the third period in the context of limestone, conglomerate, and sandstone. Based on the structural geology of the studied region, a resulted face of the earth performance shear-stress structure is made by the rock of diatomaceous deposits from the Paleozoic era and other sedimentary deposits and Azar Caves.
related to the Mesozoic era. Strain-slip faults are one of the most vivid structural elements seen in this region. Anticlines and syntaxes in the north-east axis of the Southwest were seen in the southwest of the studied region that changes in center and northwest to west-east and northwest-southeast line. Most of these buildings were unidirectional and classified in conical faults, Aghanabati, (2006).

**Sampling methods and data analysis.** Sampling was conducted in August 2016. Therefore, two stations were selected in various lagoons parts (Fig. 3). And the sedimentary cores were conducted in the Egher machine in the sediment of the lagoon. Thus, the first core section had 675 cm depth with 34 samples and the second core section had 486 cm depth with 25 samples (Fig. 2).

Samples were transferred to the laboratory to prepare for granulation, studying sedimentary characteristics, mineralogy nature, determination of the major and trace elements. After hydrometric tests, the related calculations to correct data were conducted by calculating laboratory temperature and silt and clay weight percentages were obtained (Feyznia, 2008).

To calculate these parameters, first, the accumulative frequency diagrams of each sample were drawn after performing granulometric tests by Gradistat and Excel software. Then, total skewness index (SKI) and total standard deviation index (SDI), mean (Mz), kurtosis (KU) was calculated as following using relations Folk (1980) and Ward (1957) using values of Ø84, Ø75, Ø50, Ø25, Ø16, Ø5. Since these diameters are based on mm and the obtained formula for statistical parameters are based on $\phi$, the related percentage diameters changed from mm to $\phi$ and then were inserted in formulas (Table 1).

These samples were sent to calculate the major elements (Mg, Ca) and secondary (Sr, Mn, Na, Fe) to Geochemical Laboratory of Geological Survey of
Iran. The range of the major element (Mg, Ca) and secondary ones (Sr, Mn, Na, Fe) was calculated mg/g by OES, XRF ICP-MS, ICP- methods. Plots of major oxides and sub-elements on different diagrams (Basu et al., 1975; Bhatia., 1983; Dickinson et al., 1983; Roser and Korsch., 1986; Bhatia and Crook., 1986; Suttner and Duta., 1986) helped to obtain the related results to the origin, tectonic position, and climatic conditions of these sediments.

**Discussion and conclusion.** Folk diagrams (1954) were used to determine the name of sediments of this region based on their constitutional particles that showed the sediments of this region were placed around sand, mud sands, gravel sand with slightly muddy sandstone, and silty sand (Fig. 3).

Based on the obtained information from the statistical calculation, and the related tables and diagrams using Excel and Gradistat indicated that sediment sorting of this area is 0.7-1.9, sediment kurtosis is 0.1-0.3, and kurtosis is in 0.6-0.9 (Folk, 1980).

The sediments of this area are wide based on kurtosis, negative skewness (toward big particles), and medium to bad sorting (Feyznia, 2008). The based on geochemistry of the major elements. The tectonic position has two characteristics of studying the origin places including continental blocks, volcanic arc system, and collision belts and examining the boundary type among sheets including rift or inactive continental margins, active or orogeny continental margins, or striped fault margins (Dickinson and Suczek.,1979; Dickinson et al, 1983; Garzanti et al, 2003; Garzanti et al, 2007).

The similar results were obtained in the determination of sediments tectonic position by drawing Rusar and Kurosh tectonic separating diagrams (Roser and Korsch., 1986) and 2d and Bhatia functional separating diagrams (Bhatia, 1983) (Fig. 4&5). These diagrams based on the logarithmic ratio of $\frac{K_2O}{Na_2O}$ versus $SiO_2$ percent (Fig. 4) and $TiO_2$ and $Al_2O_3/SiO_2$ values versus $Fe_2O_3+MgO$ were drawn. Their formation in Arctic islands is confirmed for sedimentary samples. As it is observed in these diagrams, it can be claimed that $TiO_2$ and total values of $Al_2O_3+Fe_2O_3+MgO$ in arctic islands reduced to inactive margins (Fig. 6).
Fig. 4 Tectonic separating diagrams of sediments based on K2O/Na2O versus SiO2 percent logarithmic ratio.

Fig. 5 Tectonic separating diagrams of sediments based on Al2O3/SiO2 percent versus Fe2O3+MgO logarithmic ratio.

Therefore, the secondary elements are significantly mentioned in the determination of tectonic origin and place (Bhatia and Crook, 1986; McLennan, 2001; Eriksson et al., 1994) which cause many studies on the present secondary elements in the sedimentary rocks to determine the tectonic position and their origin (Bahlburg, 1998; Burnett and Quirk, 2001; McLennan et al., 1993; Zimmermann and Bahlbarg, 2003).

Therefore, 3d diagrams for the secondary elements were drawn to determine the tectonic position and the obtained results of them in confirmation the resulted diagrams for the major elements showed that the tectonic position of Quaternary sediments of Siyah Keshim lagoon is an arctic continental island (Fig. 7). The related area to the firmed sediments in active and inactive continental margins was overlapped in the La-Th-Sc 3d drawing, while these two environments are completely separated in Th-Sc-Zr/10 3d drawing (Fig. 7) (Adabi, 2004).
Based on drawing the paired Ti/Zr elements versus La/Sc that is shown in Fig. 8, sediments tectonic position in Suyah Keshim area of Gilan province shows arctic continental margin (ACM) and arctic continental island (CIA). Moreover, the mentioned rocks tectonic environments by diagram (Schandl & Gorton, 2002) are shown in Fig. 9 and 10 based on the secondary elements. All the studied samples are around ACM based on this diagram.

The determined areas include A: oceanic arctic islands, B: arctic continental island, C: active continental margin, D: inactive margins

Based on the obtained results from drawing the high geochemical diagrams based on the oxidants percentage of the major and secondary elements that show the tectonic position of this region sediments of continental and oceanic arctic island, continental active margin, and inactive continental margin. It can be concluded based on citing studies (Asiabanha & Foden, 2012) that the related studied sediments are related to subduction margin.

**Interpretation of origin area weathering.** The mobility of the major elements during weathering, transportation, and processes after sediment can be used to determine the chemical maturity of sediments, (McLennan et al., 1993).

The very low concentration of Na₂O in sediments shows high sedimentary maturity (Fig. 11). Moreover, SiO₂/Al₂O₃ ratio is the usable index to determine sediment maturity (Potter, 1978). SiO₂/Al₂O₃ ratio higher than 5-6 in sedimentary rocks show their high sedimentary maturity (Roser et al., 1996). SiO₂/Al₂O₃
Fig. 8 Tectonic separating diagram based on Ti/Zr versus La/Sc ratio for the studied samples (Bhatia & Crook, 1986)

Fig. 9 The diagram of tectonic environment determination based on Th/Ta versus Yb ratio (Schandl & Gorton, 2002)

Fig. 10 The diagram of tectonic environment determination based on Th/Hf versus Ta/Hf (Schandl & Gorton, 2002)

The ratio in samples is varied in 2.6-3.7. These numbers show relatively low maturity in the studied region sediments.

ICV combined variety can be used to determine the first cycle sediment or the obtained sediments from the second cycle (Cox et al., 1995) that is obtained from the following formula:

\[
\text{ICV} = \frac{(\text{Fe}_2\text{O}_3 + \text{K}_2\text{O} + \text{Na}_2\text{O} + \text{CaO} + \text{MgO} + \text{MnO} + \text{TiO}_2)}{\text{Al}_2\text{O}_3}
\]

Samples with higher ICV than 1 are probably for the first cycle sediments, and which with smaller...
ICV than 1 may be from the second cycle sediments or the severely weathered sediments from the first cycle sediments (Cullers & Podkovyrov, 2002). The calculated values from ICV in quaternary sediments of this region are 0.9-5.35 with an average of 1.54. Thus, it can be stated that the most sediments of the studied lagoon are related to the first cycle sediments.

Weathering history of the clastic rocks is mostly estimated by calculating the mobile oxides ratio of K₂O, Na₂O, and CaO than non-mobile oxide Al₂O₃ (Nesbitt & Young, 1982). The most used index in this formula is alternation chemical index (Nesbitt & Young, 1982). This index is obtained by the following relation and oxides in it was stated in mole.

\[
\text{CIA} = \frac{\text{Al}_2\text{O}_3}{\text{Al}_2\text{O}_3 + \text{CaO} + \text{Na}_2\text{O} + \text{K}_2\text{O}} \times 100
\]

CaO is the present calcium in rock silicate components, and this value must be modified in samples that high CaO is related to diagenesis cement. CIA range may be low, medium, and high varied from 50 to 100. Increasing CIA from down to up is related to alternation chemical degree. Low CIA shows no alternation or very low alternation and is a reflection of cold and dry climate conditions, while medium and high CIA with mobile cations transmission such as (K+, Na+, Ca²⁺) and remaining the constituters (Ti⁺ and Al⁺) is related to less mobility (Nesbitt & Young, 1982). Samples were considered with higher CaO than 5% to determine CIA precisely and CaO removal from carbonate cement (Batumike et al., 2006, Nesbitt, 2003, Garcia et al., 2004).

The calculated CIA for the studied samples was 0.78-63 and is 72 averagely in samples which show severe chemical alternation in sediments. Chemical index of weathering (CIW) is extensively used to determined rock weathering degree and is obtained by the following formula (Harnois, 1988).

\[
\text{CIW} = \frac{\text{Al}_2\text{O}_3}{\text{Al}_2\text{O}_3 + \text{CaO} + \text{Na}_2\text{O}} \times 100
\]

The mean of this index for the sediments of Siyah Keshim area is 80.73. Using CIA and CIW indexes in samples with high CaO changes doesn’t show interesting results (Cullers, 2000). In this regard (Cullers, 2000), another weathering index can be offered for samples with high CaO. This index is stated as follows:

\[
\text{CWI'} = \frac{\text{Al}_2\text{O}_3}{\text{Al}_2\text{O}_3 + \text{Na}_2\text{O}} \times 100
\]

In which, oxides are considered in molecular ratio. The mean of this index in the studied samples is 95.47 that show the severe weathering of these samples.

This result was obtained finally based on the calculated indexes in which the sediments of Siyah Keshim area is for the first sedimentary cycle with relatively low sedimentary maturity, and has severe chemical weathering on then and alternate them.

**Conclusions.** Sediments are considered as data in environmental studies and its conditions are the most important evidence particularly in the previous environmental conditions. Many other pieces of

![Fig. 11 Drawing Na₂O versus K₂O](image)
evidence such as paleontology, botanical, and biological effects generally and even hum civilization effects are evaluated by sedimentary data.

In other words, the real history of the earth and its conditions are hidden in the body of the earth sediments. Sediments can be examined through various views such as particles size and diameter, ration of the formed particles of a sediment, sorting, particles kurtosis, skewness, rounding, and some other statistical parameters such as mean, mode, standard deviation, etc. each of these cases shows genre, origin, and environmental conditions that formed sediments.

Studying quaternary sediments for two studied sections of Siyah Keshim lagoon in Gilan province showed that this area has a sandy texture, mud sands, sand with slightly gravel, sandy mud with slightly gravel, and silty sand through sedimentology view.

Sediments sorting are medium to weak, negative skewness, and many big particles and plate kurtic through kurtosis view which show sediment in the coastal area. Regarding the obtained results from drawing the geochemical diagrams based on the major and secondary elements oxides, it can be concluded that the studied sediments are related to a subduction margin. Moreover, geochemical evidence shows that SiO₂/Al₂O₃ ratio in these sediments is 2.6-3.7 and also Na₂O has the relatively low sedimentary maturity for the sediments of this lagoon. The mean coefficient weathering indexes such as ICV, CIW, CWI, and CIA also show that the mentioned sediments are mainly related to the first sediment cycle and tolerate high chemical weathering.

References

Adabi, M. H., 2004, Sedimentary geochemistry, first edition, Tehran: Aryanazim publication, 448 p.

Aghanabati, S. A., 2006, Geology of Iran. 2nd edition. Tehran: Publications of Geological Survey of Iran, 590 p.

Amini Ranjbar, Gh., 1998. Heavy metal concentration in surficial sediments from Anzali Wetland, Iran. Water, Air, and Soil Pollution, 104:305-312.

Asiabeha, A. Foden, J., 2012. Post-collisional transition from an extensional volcano-sedimentary basin to a continental arc in the Alborz Ranges, N-Iran. Journal of Lithos 148: 98-111.

Bahlburg, H., 1998. The geochemistry and provenance of Ordovician turbidites in the Argentine Puna. In: Pankhurst, R.J., Rapela, C.W. (Eds.), The Proto-Andean Margin of Gondwana Special Publication 142. Geological Society of London, 127-142.

Batumilke, M. J., Kampunzu, A. B. & Cailleux, J. H., 2006. Petrology and geochemistry of the Neoproterozoic Nguba and Kundelungu Groups, Katangan Supergroup, southeast Congo: Implications for provenance, paleoweathering and geotectonic setting. Journal of African.

Bhatia, M R. Crook, K A W. 1986. Trace element characteristics of greywackes and tectonic setting discrimination of sedimentary basins. Contributions to Mineralogy and Petrology, 92: 181-193.

Bhatia, M R. 1983. Plate tectonics and geochemical composition of sandstones. Journal of Geology, 91: 611–627.

Bui, EN. Mazullo, J. Wilding, LP. “Using quartz grain size and shape analysis to distinguish between Aeolian and fluvial deposits in the Dallol Bosso of Niger (West Africa)”, Earth Surface Processes and Landforms. 14. 157–166.1990.

Burnett, D J. Quirk, D G. 2001. Turbidite provenance in the Lower Paleozoic Manx Group, Isle of man; implications for the tectonic setting of Eastern Avalonia. Journal of Geological Society of London, 158: 913–924.

Cox, R., Low, D. R. & Cullers, R. L., 1995. The influence of sediment recycling and basement composition on evolution of mudrock chemistry.

Cullers, R. L. & Podkovyrov, V. N., 2002. The source and origin of terrigenous sedimentary rock in the Mesoproterozoic UI group, southeaste Russia. Journal of Precambrian Research 117: 157-183.

Cullers, R. L., 2000. The geochemistry of shales, siltstones and sandstones of Pennsylvanian-Permian, Colorado, USA: Implication for provenance and metamorphic studies. Lithos 51: 181-203.

Dickinson, W.R. and Suczek, C.A., 1970. Plate tectonics and sandstone compositions. American Association of Petroleum Geologist, Bulletin 63:2164-2182.

Dickinson, W.R. Beard, L.S. Brakenridge, G.R. Evjavec, J.L. Ferguson, R.C Inman, K.F. Knepp, R.A. Lindberg, F.A. and Ryberg, P.T., 1983. Provenience of North American Phanerozoic sandstones in relation to tectonic setting. Geological Society of America, Bulletin 94: 222-235.

Earth Sciences 44: 97-115.

Eriksson, P. G., Schreiber, U. M., Reczko, B. F., Snyman, C. P., 1994. Petrography and geochemistry of sandstones interbedded with the Rooiberg Felsite Group (Transvaal sequence, south Africa): implication for provenance and tectonic setting. Journal of Sedimentary Research, A64: 836-846.

Feyznia, S., 2008. Applied sedimentology with emphasis on soil erosion and sediment production, Gorgan University of agricultural sciences and natural resources, 356.

Folk, R.L. and W.C. Ward, 1957. Brazos River bar: a study in the significance of grain size.
Folk, R.L., 1954. The distinction between grain size and mineral composition in sedimentary rock nomenclature, Jour. Geol. V.62, 344-359.

Folk, R.L., 1980. Petrology of sedimentary rocks, Hamphill Food and Agriculture Organization (FAO), 1980, Human induced erosion maps. http://www.FAO.com.

Friedman, G.M., 1979. Differences in size distributions of populations of particles among sands of various origins, Sediment logy. 26, 3-32.

Garcia, D., Ravenne, C., Marechal, B. & Moutte, J., 2004. Geochemical variability induced by entrainment sorting: quantified signals for provenance analysis. Sediment Geol 171: 113-128.

Garzanti, E., Vezzoli, G. and Ando S., 2002. Modern sand from subducted ophiolite belts (Sultanate of Oman and United Arab Emirates). Geology, 110(4): 371-391.

Garzanti, E., Vezzoli, G., Ando, S.Lave, J., Attal, H., France – Lanord C. and De Celles P., 2007. Quantifying sand provenance and erosion (Marsyandi River, Nepal Himalaya). Earth and Planetary Science Letters, 258(3): 500-515.

Gondwana Special Publication 142. Geological Society of London, 127-142.

Harnois, L., 1988. The CIW index: a new chemical index of weathering. Journal of Sedimentary Geology 55: 219-227.

Kazanci, N., et al. 2004. Sedimentary and environmental characteristics of the Gilan Mazenderan plain, northern Iran: influence of long- and short-term Caspian water level fluctuations on geomorphology, Journal of marine systems.46:145-168.

Khabaznia et al., 2005. Bandar Anzali geometrical map, 1:100000, Geological Survey of the country.

Mansuri, N., Khorasani, N., 1985. Lagoons and their values and importance for human, journal of ecology, 12th year, No. 13, Tehran University publication.

McLennan, S. M., Hemming S., McDaniel, D. K., Hanson, G. N., 1993. Geochemical approaches to sedimentation, provenance and tectonics. In: Johnson, M.J. and Basu, A.,(Eds). Geological Society of America, Special Papers 285: 21–40.

McLennan, S. M., 2001. Relationships between the trace element composition of sedimentary rocks and upper continental crust. Geochemistry Geophysics Geosystems, 2: C000109.

Mohammadi, A., Amani, M., Navernaviri, M., 2012. Sedimentology report of surface sediments of Anzali lagoon, geological survey of Iran, 139.

Nekukhu, M., 2003. Geochemistry and sedimentary environments of permain carbonate sediments in southeast of Shahreza city, MSc thesis, Faculty of earth sciences, Shahid Beheshti University, 27-108.

Nesbitt, H. W. & Young, G. M., 1982. Early Proterozoic climates and plate motions inferred from major element chemistry of lutites. Nature 717-715: 299.

Nesbitt, H. W., 2003. Petrogenesis of siliciclastic sediments and sedimentary rocks. In: Lenz, D.R. (Ed.), Geochemistry of Sediments and Sedimentary Rocks, Geotext 4. Geological Association of Canada, Newfoundland: 39–51

Nuri, Gh., 2009. Lagoon importance, adopted from Iran deserts electronic website: www.irandesert.com parameters, Sedimentary Petrology, 27: 3-26.

Potter, P.E., 1978. Petrology and chemistry of modern big river sands. Journal of Geology, 86: 423-449.

Ranjar, M., 2012. Anzali lagoon changes and its morphological effects in ground application, land geographical quarterly, 9th year, no. 34

Riyahi, A.R., et al., 2005. Determination of Heavy Metals Content in Astacus Leptodactylus Caspicus of Anzali: In CEST2005: 784-790.

Roser, B. P., Korsch R. J., 1986. Determination of tectonic setting of sandstone– mudstone suites using SiO₂ content and K₂O/Na₂O ratio. Journal of Geology, 94: 635–650.

Roser, B. P., Korsch R. J., 1988. Provenance signatures of sandstone–mudstone suites determined using discriminant function analysis of major-element data. Chemical Geology, 67: 119–139.

Roser, B. P., Cooper, R. A., Nathan, S. & Tulloch, A. J., 1996. Reconnaissance sandstone geochemistry, provenance and tectonic setting of the lower Palaeozoic terranes of the West Coast and Nelson, New Zealand. New Zealand Journal of Geology and Geophysics 39: 1–16.

Sai, K., 2004. Geochemistry of lake sediments as a Record of Environmental Change in a High Arctic Watershed, Chemie der Erde 64, 257-275.

Sanzivala, Z., Pakzad, H., Makkizadeh, M.A., Kamali, M., 2011. Investigation of sedimentology and sediment of sandy sediments of southern part of Anzali lagoon, 15th Iranian Geological Society.

Schandl, E. S. & Gorton, M. P., 2002. Application of high field strength elements to discriminate tectonic setting in VMS environments. Journal of Economic Geology 97: 629-642.

Sharifi, M., 2006. The Pattern of Caspian Sea Water Penetration into Anzali Wetland: Introduction of a Salt Wage.Caspian J. Env. Sci. Vol. 4. No.1:77-81.

Taghavi, L., Tayebi, S., Karimiyani, B., 2013. Long-term analysis of Northern part of Gavkhouni lagoon using geochemistry of the major and secondary elements, Journal of Oceanology, Wetland ecology research, Islamic Azad University of Ahvaz, 5th year, No.16.

Zimmermann, U, Bahlburg, H., 2003. Provenance analysis and tectonic setting of the Ordovician clastic deposits in the southern Puna Basin, NW Argentina. Sedimentology, 50: 1079–1104.