Implications of the interstellar object 1I/'Oumuamua for planetary dynamics and planetesimal formation

Sean N. Raymond,1,⋆ Philip J. Armitage,2,3 Dimitri Veras,4,5† Elisa V. Quintana6 and Thomas Barclay6,7

1Laboratoire d’Astrophysique de Bordeaux, CNRS and Université de Bordeaux, Allée Geoffroy St. Hilaire, F-33165 Pessac, France
2JILA, University of Colorado and NIST, 440 UCB, Boulder, CO 80309-0440, USA
3Department of Astrophysical & Planetary Sciences, University of Colorado, Boulder, CO 80309-0391, USA
4Department of Physics, University of Warwick, Coventry CV4 7AL, UK
5Centre for Exoplanets and Habitability, University of Warwick, Coventry CV4 7AL, UK
6NASA Goddard Space Flight Center, 8800 Greenbelt Rd, Greenbelt, MD 20771, USA
7University of Maryland, Baltimore County, 1000 Hilltop Cir, Baltimore, MD 21250, USA

Accepted 2018 February 8. Received 2018 January 11; in original form 2017 November 21

ABSTRACT
'Oumuamua, the first bona fide interstellar planetesimal, was discovered passing through our Solar system on a hyperbolic orbit. This object was likely dynamically ejected from an extrasolar planetary system after a series of close encounters with gas giant planets. To account for 'Oumuamua’s detection, simple arguments suggest that \( \sim 1 M_\oplus \) of planetesimals are ejected per solar mass of Galactic stars. However, that value assumes mono-sized planetesimals. If the planetesimal mass distribution is instead top-heavy, the inferred mass in interstellar planetesimals increases to an implausibly high value. The tension between theoretical expectations for the planetesimal mass function and the observation of 'Oumuamua can be relieved if a small fraction (\( \sim 0.1 - 1 \) per cent) of planetesimals are tidally disrupted on the pathway to ejection into 'Oumuamua-sized fragments. Using a large suite of simulations of giant planet dynamics including planetesimals, we confirm that 0.1–1 per cent of planetesimals pass within the tidal disruption radius of a gas giant on their pathway to ejection. 'Oumuamua may thus represent a surviving fragment of a disrupted planetesimal. Finally, we argue that an asteroidal composition is dynamically disfavoured for 'Oumuamua, as asteroidal planetesimals are both less abundant and ejected at a lower efficiency than cometary planetesimals.

Key words: comets: general – minor planets, asteroids: individual: 1I/'Oumuamua – planets and satellites: dynamical evolution and stability – planets and satellites: formation – protoplanetary discs.

1 INTRODUCTION

1I/'Oumuamua was discovered by the Panoramic Survey Telescope And Rapid Response System (Pan-STARRS) survey (Chambers et al. 2016) on 2017 October 19 (Meech et al. 2017; Williams 2017). The object’s orbit was determined to be inclined by 123° with respect to the ecliptic plane and to have an eccentricity of 1.2, placing it on a hyperbolic orbit that is not aligned with any known cluster of comets (de la Fuente Marcos & de la Fuente Marcos 2017). Given the lack of Solar system objects capable of imparting a large enough impulse to explain its anomalous velocity (e.g. Portegies Zwart et al. 2017), 'Oumuamua is likely of extra-solar origin and in the process of passing through the Solar system on its way back to interstellar space.

Interstellar planetesimals are likely to represent by-products of planet formation. Their existence is not a surprise as planet formation is well known to be less than 100 per cent efficient. The detectability of interstellar asteroidal and cometary planetesimals has been studied by a number of authors over the past several decades (e.g. McGlynn & Chapman 1989; Kresak 1992; Jewitt 2003; Moro-Martín, Turner & Loeb 2009; Jura 2011; Cook et al. 2016). The absence of detections prior to 'Oumuamua led Engelhardt et al. (2017) to place an upper limit on the abundance of interstellar planetesimals.

Planetesimals are thought to form via gas-assisted concentration of drifting particles in planet-forming discs (see the review by Johansen et al. 2014). A fraction of planetesimals can subsequently be ejected by receiving gravitational kicks from planets. A planet’s
ejection potential can be quantified with the Safronov number $\Theta$, defined as the ratio of the escape speed from a planet’s surface to the local escape speed from the star (see e.g. Ford & Rasio 2008; Raymond, Armitage & Gorelick 2010; Wyatt et al. 2017). Ejections are likely for $\Theta \gg 1$, while collisions are favoured over ejections if $\Theta < 1$. In the Solar system, although all the giant planets have $\Theta > 1$, Jupiter ejects the vast majority of planetesimals. Given the planets’ orbital architecture, outer Solar system planetesimals are generally passed inward from the ice giants to Saturn before being ejected after repeated encounters with Jupiter (Fernandez & Ip 1984). The resulting orbital migration is the basis for understanding Pluto’s orbit (Malhotra 1995), the Kuiper Belt (e.g. Nesvorny 2015), and the current architecture of the outer Solar system (Tsiganis et al. 2005). Contemporary models for early outer Solar system evolution invoke planetesimal discs with masses of $\approx 15–65 M_\oplus$ (Deienno et al. 2017), most of which is ultimately ejected.

Dynamical arguments show that the observation of one interstellar object during the duration of the Pan-STARRS survey implies that every star ejects about $1 M_\oplus$ of planetesimals (Laughlin & Batygin 2017; Trilling et al. 2017). By comparing this to the $>10 M_\oplus$ ejected from the Solar system, one might immediately conclude that the observation of ‘Oumuamua is expected and possibly overdue. However, this expectation is not obvious. Most of the known exoplanets are too close to their stars to be good candidates for ejecting planetesimals (Laughlin & Batygin 2017), and the known population of gas giants that probably dominate ejections are present around only a small fraction of stars. Planetesimals, moreover, have a distribution of masses, and if that distribution is skewed towards large masses, inferences from seeing a single small body need to be corrected for the much larger mass of the population.

Our paper is structured as follows. We first point out a fundamental shortcoming of the simple estimates for the abundance of interstellar objects (Section 2). Assuming that the detected object was drawn from a reasonable size distribution implies an order of magnitude increase in this abundance. We then (Section 3) use existing simulations of extrasolar planetary dynamics (Raymond et al. 2011) to develop a framework for the possible sources of ejected planetesimals. We show that this naturally matches the nominal (lower) estimate for the abundance of interstellar objects. We argue on dynamical grounds that it is much easier to explain ‘Oumuamua as a dormant cometary nucleus rather than an asteroid. In Section 4 we show that a fraction $f \approx 1$ per cent of planetesimals in our simulations underwent extremely close encounters with one or more giant planets before being ejected. These encounters should lead to tidal disruption. We demonstrate that this solves the problem pointed out in Section 2 if the population of ejected planetesimals is characterized by a standard size distribution with a small contamination of fragments, which contain a small fraction of the mass but dominate by number. In Section 5 we compare our model with an alternate one that assumes ‘Oumuamua to have originated in a binary star system, and conclude with a discussion of the implications for the birth size distribution of planetesimals in Section 6.

2 ESTIMATING THE PLANETESIMAL MASS EJECTED PER STAR

The observation of a flux of interstellar objects passing through the inner Solar system directly constrains the interstellar number density $n$ of such bodies. With an estimate of the size and density of the intruders, we can convert the number density to a mass density and infer the average amount of mass that must be ejected for each Solar mass of local stellar population. This estimate has been made by several authors (Laughlin & Batygin 2017; Meech et al. 2017; Portegies Zwart et al. 2017; Trilling et al. 2017; Do, Tucker & Tonry 2018; Rafikov 2018). It is necessarily uncertain for so long as we have only a single detected object.

Recapping the elementary argument, the cross-section for an object with relative velocity at infinity $v_\infty$ to approach to a peri-centre distance $q$ is

$$\sigma = \pi q^2 \left[ 1 + \left( \frac{v_\infty}{v_\infty'} \right)^2 \right],$$

where $v_\infty$ is the escape velocity from $q$. The observation of a rate of such encounters $\Gamma$ determines the number density $n = \Gamma/v_\infty$, and an interstellar mass density,

$$\rho_{ej} = \frac{4\pi \Gamma r_p}{3 \sigma v_\infty},$$

where $r$ is the radius and $\rho_m$ the material density of the objects. ‘Oumuamua has $v_\infty = 26 \text{ km s}^{-1}$ and $q = 0.25 \text{ au}$, and hence we adopt $\sigma \sim 3 \text{ au}$. Large variations in the absolute magnitude suggest a highly elongated ellipsoidal figure, with an equivalent spherical radius of $\approx 70 \text{ m}$ (Bannister et al. 2017; Jewitt et al. 2017; Meech et al. 2017; Bolin et al. 2018). The lack of evidence for outgassing means that the purely observational prior favours an asteroid-like composition, with $\rho_m \approx 3 \text{ g cm}^{-3}$. The rate is highly uncertain in the absence of a detailed study of survey completeness, but is of the order of $\Gamma \sim 0.1 \text{ yr}^{-1}$.

We can compare $\rho_{ej}$ to the local mass density of stars. McKee, Parravano & Hollenbach (2015) infer a local surface density of stars and stellar remnants $\Sigma_* = 33.4 \pm 3 M_\odot \text{ pc}^{-2}$, and a volume density $\rho_v = 0.060 M_\odot \text{ pc}^{-3}$. If we make the simple (and probably incorrect) assumption that the amount of mass ejected scales linearly with $M_\star$, we find that the mass ejected per Solar mass of the local stellar population is

$$\epsilon_{ej} \approx 0.65 \left( \frac{\Gamma}{0.1 \text{ yr}^{-1}} \right) \left( \frac{\sigma}{3 \text{ au}} \right)^{-1} \left( \frac{v_\infty}{26 \text{ km s}^{-1}} \right)^{-1} \times \left( \frac{r}{70 \text{ m}} \right)^3 \left( \frac{\rho_m}{3 \text{ g cm}^{-3}} \right) M_\odot M_\odot^{-1}. \tag{3}$$

This estimate drops to $\epsilon_{ej} \approx 0.2 M_\odot M_\odot^{-1}$ for an icy (lower density) composition, and has a large uncertainty inherent in deriving a rate from a single object in an incompletely characterized survey. We take $1 M_\odot M_\odot^{-1}$ as a central estimate, and consider 0.1–10 $M_\odot M_\odot^{-1}$ as a reasonable range. This is somewhat lower than the values inferred by Trilling et al. (2017) and Do et al. (2018), who found roughly 20 and 40 $M_\odot$ ejected per star, respectively. As we show below and in Section 4, any inferred values (including ours) are unlikely to represent the true value but rather may be underestimated or overestimated by a significant amount, subject to unconstrained parameters of the population of interstellar objects.

The mass estimate in equation (3) is valid if the population of interstellar planetesimals is mono-disperse. If the mass distribution of the ejected population is instead top heavy – as is observed for the asteroid and Kuiper Belts, and expected theoretically in some planetesimal formation scenarios (see Johansen et al. 2014) – then the inferred $\epsilon_{ej}$ increases.

We model the mass function of interstellar planetesimals as a power law,

$$\frac{dN}{dn} = km^n,$$  \tag{4}
between $m_{\min}$ and $m_{\max} \gg m_{\min}$. We assume that the mass function is dominated by number by small objects, and by mass by large ones ($1 < p < 2$). The expected mass for the first object to be seen from this distribution is given by setting the number of objects with a lower mass equal to the number with a higher mass, $m_{1/2} = 2^{(p-1)/p} m_{\min}$. Proceeding as above we would then estimate a total mass for the population of $M_{\text{est}} = N m_{1/2}$, where $N$ is the total number of bodies. The true mass, however, is given by integrating over the mass function as $M_{\text{final}} \approx k m_{\text{max}}^p / (2 - p)$. The ratio between the true mass of the population and the one estimated using a single size is then,

$$b = 2^{p-1} \left( \frac{p - 1}{2 - p} \right) \left( \frac{m_{\max}}{m_{\min}} \right)^{(3/2) - (p)},$$

written in terms of the minimum and maximum radii.

This leads to a discrepancy with models in which planetesimals form via the streaming instability (Youdin & Goodman 2005). Simulations show that a single burst of planetesimal formation from initial conditions unstable to streaming yields a power-law mass function with $p \approx 1.6$ (Johansen et al. 2015; Simon et al. 2016, 2017; Schäfer, Yang & Johansen 2017). Taking $m_{\min} = 100$ m and $m_{\max} = 100$ km (a conservative upper size limit, as forming a belt with a total mass of $\sim 10^5$ M$_{\oplus}$ in one episode would probably yield a maximum size closer to 1000 km), we find $b \sim 10^3$. A still substantial bias $b > 10^2$ occurs for steeper but still top-heavy mass functions, for example for $p = 11/6$ as appropriate for a steady-state collisional cascade (Dohnanyi 1969).

If ejected planetesimals sample a streaming mass function ($p = 1.6$), then the estimate of 1 M$_{\oplus}$ M$_{\odot}^{-1}$ in ejected planetesimals is underestimated by a factor of $b > 100$. The inferred mass is hundreds of Earth masses, larger than the total solid budget of most planet-forming discs (see Section 3.2).

We conclude this section with a conundrum. An estimate for the mass in interstellar planetesimals based on an unrealistic, mono-disperse size distribution yields a reasonable value. Yet an estimate based on a plausible size distribution produces an unrealistically large value.

### 3 DYNAMICAL SOURCES OF EJECTED PLANETESIMALS

The mass of ejected planetesimals is determined by the initial configuration of a planetary system and by the location and masses of planetesimal belts. For gas giants we have reasonable estimates of planetary frequency out to 5–10 au (Cumming et al. 2008; Mayor et al. 2011; Wittenmyer et al. 2016), and circumstantial evidence from orbital eccentricities for their dynamical history (e.g. Adams & Laughlin 2003; Chatterjee et al. 2008). There is no evidence from direct imaging surveys or microlensing for a substantial population of gas giants in wide orbits (Meshkat et al. 2017; Mróz et al. 2017). We know from the existence of debris discs (Wyatt 2008) that extrasolar planetesimal belts exist, but have only limited knowledge of their initial masses.

The above constraints do not allow us to uniquely predict a stellar population averaged mass ejection efficiency. What we can do, however, is to define a set of dynamical models that fit the basic known properties of the giant extrasolar planet population, using assumptions for the masses of planetesimal belts that are in the range inferred for the Solar system. We first ask whether such a baseline model, developed previously (Raymond et al. 2010), is consistent with the observation of 'Oumuamua. We assume the following:

(i) A fraction $f_{\text{giant}}$ of stars host giant planets. Within this subset, a fraction $f_{\text{unstable}}$ form planets in an initial configuration that is subject to subsequent dynamical instability (leading to orbit crossing, planetary ejections, and efficient ejection of neighbouring planetesimals).

(ii) The remaining stars do not host giant planets. These systems might host ice giants, but without secure knowledge of the ice giant populations it is reasonable to assume that they are less efficient at ejecting planetesimals on average.

The mean mass in planetesimals ejected per star is then

$$M_{\text{interstellar}} = \left(1 - f_{\text{giant}}\right) \times M_{\text{unstable}}^{\text{ejected}}$$

$$+ f_{\text{giant}} \times \left(1 - f_{\text{unstable}}\right) \times M_{\text{giant}}^{\text{stable}, \text{ejected}}$$

$$+ f_{\text{giant}} \times f_{\text{unstable}} \times M_{\text{giant}}^{\text{unstable}, \text{ejected}}$$

where $M_{\text{ejected}}$ is the average mass ejected from systems without a gas giant (nogiant), and in stable (giant, stable) and unstable (giant,unstable) giant planet systems.

Long-duration radial velocity surveys constrain the occurrence rate of giant planets $f_{\text{giant}}$ to be $\sim 10–20$ per cent for Sun-like stars (Cumming et al. 2008; Mayor et al. 2011; Foreman-Mackey et al. 2016; Rowan et al. 2016; Wittenmyer et al. 2016). Most gas giants are located beyond 0.5–1 au at radii where they could plausibly eject planetesimal belts. Low-mass stars have fewer gas giants than high-mass stars (Johnson et al. 2007; Lovis & Mayor 2007; Dressing & Charbonneau 2015) so the population-averaged rate is likely smaller than the value for Sun-like stars.

Unstable giant planet systems are extremely efficient at ejecting planetesimals. Planet–planet scattering is the leading model to explain the eccentric orbits of giant exoplanets (e.g. Lin & Ida 1997; Adams & Laughlin 2003; Chatterjee et al. 2008) and the observed distribution can be matched if $\sim 75–95$ per cent of giant exoplanets are the survivors of dynamical instabilities (Juric & Tremaine 2008; Raymond et al. 2011). Simulations show that instabilities typically result in the ejection of one or more giant planets and the gravitational perturbations during this process often eject the entire outer planetesimal disc out to $\sim 30$ au (Raymond et al. 2010, 2011, 2012; Raymond & Armitage 2013; Marzari 2014).

A baseline model for the source population of interstellar objects invokes efficient ejection from a small fraction of stars with unstable giant planet systems, less efficient ejection for stars with stable giants, and no ejection from the bulk of stars without massive planets (based on the Safronov number argument). Given that we infer a high unstable fraction $f_{\text{unstable}} \approx 90$ per cent from dynamical arguments, the key unknown is the mass ejected per unstable system ($M_{\text{giant}}^{\text{unstable}}$). The ejection of $10–100$ M$_{\oplus}$ per star hosting unstable gas giants suffices to match the central estimate (of 1 M$_{\oplus}$ ejected per solar mass of stars) for the population averaged ejection efficiency assuming a stellar mass-averaged $f_{\text{giant}}$ of 1–10 per cent. This is consistent with Solar system expectations if 'Oumuamua formed as an icy body (see Section 3.1).

It is worth noting that equation (6) likely underestimates planetesimal ejection. For example, in binary star systems, ejection may be more efficient than in single-star systems (Smullen, Kratter & Shannon 2016; Sutherland & Fabrycky 2016; Gong & Ji 2017). Planetesimal ejection may also be efficient after the main sequence (see Veras 2016, for a review of the effects of stellar evolution on orbital dynamics). Extrasolar Oort Cloud analogues – formed partly during planet formation and partly from continuous interchange with the interstellar medium – are almost entirely ejected during
3.1 An interstellar asteroid or comet?
Observations of 'Oumuamua have not conclusively determined its composition. Despite its very close approach to the Sun (perihelion of 0.25 au), 'Oumuamua has shown no signs of sublimed water in the form of a coma (Jewitt et al. 2017; Meech et al. 2017; Ye et al. 2017). However, Fitzsimmons et al. (2018) showed that 'Oumuamua's perihelion passage would only have exposed volcanoes down to a depth of roughly 40 cm. 'Oumuamua's colours fall within the distribution of primitive, organic-rich Solar system objects (Bannister et al. 2017; Bolin et al. 2018; Fitzsimmons et al. 2018; Jewitt et al. 2017). It has a red spectral slope with no obvious absorption features (Fitzsimmons et al. 2018; Masiero 2017; Meech et al. 2017; Ye et al. 2017), similar to some primitive Solar system bodies (e.g. Bus & Binzel 2002; Carry et al. 2016). Put together, it seems plausible to imagine that 'Oumuamua is either rocky (asteroidal) in nature or an extinct cometary nucleus (as we discuss further below). The Damocloids, with orbits similar to Halley-type and long-period comets but no detectable activity (Asher et al. 1994; Jewitt 2005), may offer a suitable analogy. On the other hand, Cook et al. (2016) found that surveys may be biased towards finding asteroids if small \((r < 1 \text{ km})\) comets are very rare.

Given the inconclusive direct evidence we use a suite of dynamical simulations to address the question of whether asteroidal or cometary planetesimals are more likely to be ejected from systems with gas giants. The Safronov number for a planet of given mass increases with orbital radius, so from a dynamical standpoint it is easier to eject icy planetesimals from beyond the snow line than rocky bodies that form closer to the star (e.g. Gaidos, Williams & Kraus 2017)\(^2\).

Fig. 1 shows the fraction of planetesimals that were ejected or hit the central star in simulations from Raymond et al. (2011, 2012). These simulations included an inner disc of terrestrial-planet-forming (rocky asteroidal) material, three giant planets (with masses chosen to match the observed planetary mass function), and an outer disc of presumably comet-like, volatile-rich planetesimals. In these simulations, dynamical instabilities in the (dominant) subset of initially unstable systems were almost 100\% efficient in ejecting icy planetesimals across a broad range of radii. In contrast, a substantial fraction of rocky planetesimals were excited to high enough eccentricities to impact the star. The same trends are seen in the stable subset (see also Barclay et al. 2017), but these systems were only able to eject nearby planetesimals in the outer belt.

The above arguments do not rule out the possibility that the majority of interstellar bodies are asteroidal in nature. This would require, however, that primordial rocky planetesimal belts are typically substantially more massive than icy ones at larger radii. Moreover, the frequency of massive planets that could eject asteroids from close-in orbits is already constrained. Roughly, we would require typical asteroid belts of the order of \(10 \text{M}_{\oplus}\) to meet the central estimate for ejection efficiency. This is not impossible, since aerodynamically coupled pebbles may drift inward and preferentially produce planetesimals close-in (Drążkowska, Alibert & Moore 2016). We would be forced to consider the possibility that asteroid belts are typically more massive than belts of icy planetesimals if future observations suggest that most interstellar bodies are rocky.

3.2 Limits on ejected masses
A planetary system cannot eject an arbitrarily large amount of planetesimals because the process is self-limiting. A growing gas giant ejects a significant fraction of nearby planetesimals (Raymond & Izidoro 2017) and, after dispersal of the gaseous disc, planetesimals from within a few Hill radii are perturbed on to giant planet-crossing orbits and ejected (Duncan, Quinn & Tremaine 1987; Charnoz & Morbidelli 2003). At this point ejection is slowed because the reservoir of unstable planetesimals is cut off (e.g. Bonsor, Augereau & Thébault 2012). Additional planetesimals can be ejected if a giant planet is perturbed on to an eccentric orbit, usually by planet–planet scattering. However, if the mass in outer disc of planetesimals is within a factor of a few of the scattered planet’s mass, angular momentum exchange damps the planet’s eccentricity and lifts its pericentre, quenching the giant planet instability (Raymond, Armitage & Gorenck 2009, 2010). For this reason we do not expect a system to eject more than its planetary budget in planetesimals (at least while its host star is on the main sequence; see Veras et al. 2014b, Stone et al. 2015; Rafikov 2018). Planetesimal formation could in principle start early, when the disc mass is high. For a gas disc mass of \(0.1 \text{M}_{\odot}\), the mass in icy material for a standard metallicity is approximately \(300 \text{M}_{\oplus}\). If extremely massive planetesimal belts formed frequently, the easiest way to eject them without violating observational constraints on the number of planets at large orbital radii would be if the same systems hosted multiple ice giants. In such a model of the order of \(100 \text{M}_{\oplus}\) of icy planetesimals could be ejected from most stars. This scenario appears contrived, though it shares some similarities with models which seek to explain the dust rings in systems such as HL Tau (ALMA Partnership 2015) as a consequence of early-forming ice giants.

---

1. Two new studies explore a post-main sequence ejection origin for ‘Oumuamua (Hansen & Zuckerman 2017; Rafikov 2018).
2. A few per cent of rocky bodies may be scattered into outer icy regions where they did not form (Weissman & Levison 1997; Shannon et al. 2015; Meech et al. 2016), but this does not affect our argument.
4 CONSEQUENCES OF TIDAL DISRUPTION DURING EJECTION

During the scattering phase, planetesimals undergo close encounters during which they pass within a gas giant’s Hill sphere. Close encounters on the pathway to ejection can tear planetesimals apart. It is worth considering whether 'Oumuamua could itself be a fragment of a larger planetesimal that was tidally disrupted. If disrupted ejected planetesimals are to be common, they must walk a tightrope: too close an approach to a giant planet will result in a collision, and many distant approaches will not cause disruption.

We investigated the encounters undergone by planetesimals on the path to ejection. We used the same batch of simulations as presented in Fig. 1 (from Raymond et al. 2011, 2012). For each of the >10⁵ planetesimals ejected in those 200 simulations, we tabulated the parameters of the close encounters undergone with each of the three giant planets before the planetesimal’s ejection. We divided up the planetesimals by origin; planetesimals closer to the star than the giant planets (between 0.5 and 4 au) are referred to as ‘asteroidal’ and planetesimals originating beyond the giant planets (typically past 10 au; see Fig. 1) are ‘cometary’. In contrast with Fig. 1, we now combine both stable and unstable simulations into a single set.

While comets were ejected at a higher rate than asteroids (Fig. 1), ejected asteroids typically underwent more encounters than ejected comets. The median number of close encounters with a giant planet before ejection was 54 for asteroids and 19 for comets. There is a clear correlation between the number of encounters and the closest close encounter: planetesimals that underwent more encounters had a statistically closer closest encounter. The critical radius for tidal disruption $R_t$ is defined as

$$R_t \approx r \left( \frac{M_{\text{giant}}}{m} \right)^{1/3},$$

where $m$ and $r$ are the planetesimal’s mass and radius, respectively, and $M_{\text{giant}}$ is the giant planet mass. It is clear that $R_t \propto \rho^{-1}$, where $\rho$ is the planetesimal’s bulk density. For a planetesimal with density of 1 g cm⁻³ undergoing a close encounter with Jupiter, $R_t \approx 76$ 800 km, just ~10 per cent wide of Jupiter’s actual radius of ~70 000 km.

Fig. 2 shows the cumulative distribution of the closest encounter for ejected asteroids (in red) and comets (in blue), scaled to the tidal disruption radius $R_t$ of the relevant planet. The median closest encounter was 35$R_t$ for asteroids and 87$R_t$ for comets. The fraction of planetesimals that underwent encounters within 1, 3, and 10 $R_t$ was, respectively, 0.2 per cent, 3.9 per cent, and 16 per cent for asteroids and 0.05 per cent, 1.9 per cent, and 7 per cent for comets. Among planetesimals with approaches closer than $R_t$, 18 per cent/6 per cent/1.4 per cent had 2/3/4 or more passages within $R_t$, i.e. multiple likely disrupting encounters. See Raymond, Armitage & Veras (2018) for more on tidal disruption of ejected planetesimals in a range of dynamical settings.

The kink in the distribution of closest encounters at $d_{\text{min}} \leq R_t$ is due to the tidal radius approaching the planet’s physical radius such that many encounters within $R_t$ simply lead to planetary collisions, not ejections. Given that giant planets compress and shrink in radius as a function of time (e.g. Baraffe et al. 2003; Fortney, Baraffe & Militzer 2010), the parameter space available for tidal disruption followed by ejection (rather than a simple collision) increases at later times. In our simulations each of the giant planets – with masses drawn from the observed $dN/dM \propto M^{-1.1}$ distribution (Butler et al. 2006; Udry & Santos 2007) between Saturn’s mass and three Jupiter masses – was given Jupiter’s bulk density of 1.3 g cm⁻³. This may underestimate the radius of some planets and modestly overestimate the contribution of tidally disrupted ejected planetesimals.

What happens when a planetesimal passes inside the tidal disruption radius? One would expect the object to be torn apart down to the size scale at which tensile strength dominates over self-gravity, orders of magnitude smaller than the typically assumed planetesimal size of ~100 km (e.g. Benz & Asphaug 1999; O’Brien & Greenberg 2003; Leinhardt & Stewart 2009). This would create a huge number of tidally distorted fragments, possibly with large axis ratios like ‘Oumuamua’ (Bannister et al. 2017; Fitzsimmons et al. 2018; Fraser et al. 2017; Jewitt et al. 2017; Meech et al. 2017; Bolin et al. 2018). After tidal disruption, the resulting debris could recollapse into larger bound objects (Hahn & Reitgl 1998; Veras et al. 2014a; Coughlin et al. 2016). Detailed modelling of this process is beyond the scope of this paper, yet it seems reasonable to assume that a very large number of small fragments may be created by the disruption of a single large planetesimal. Indeed, comet Shoemaker-Levy 9’s close encounter with Jupiter in 1992 caused it to disrupt into a chain of 21 detectable fragments (Weaver et al. 1995; Noll, Weaver & Feldman 1996; Movshovitz, Asphaug & Korycansky 2012), many of which were similar in size to ‘Oumuamua.

We now explore the consequences of a bimodal population of interstellar planetesimals. We naively assume that a small fraction $f$ from a $dN/dM = k m^{-p}$ profile has been tidally disrupted and transformed into mono-sized fragments with size $r_{\text{tag}}$ and mass $m_{\text{tag}}$ (assuming a constant physical density across all sizes). The total number of fragments $N_{\text{frag}}$ is

$$N_{\text{frag}} = f \left( \frac{k}{2 - p} \right) \left( \frac{m_{\text{max}}^{2-p}}{m_{\text{tag}}^{2-p}} \right),$$

assuming that $1 < p < 2$ such that the mass is dominated by the largest bodies. Given that the constant $k$ can be written in terms of the total number of bodes in the initial distribution $N$ (as $k \approx \frac{N}{m_{\text{min}}^{2-p}}$)

\[3\] We have deliberately kept our formulation for $R_t$ simple; more complex treatments would take into account the effect of other physical properties of planetesimals including tensile strength (e.g. Cordes & Shannon 2008; Veras et al. 2014a; Bear & Soker 2015; Veras 2016).
the ratio of the total number of fragments to the bodies in the original distribution is
\[ \frac{N_{\text{frag}}}{N} \approx f \left( \frac{p - 1}{2 - p} \right) \left( \frac{m_{\text{max}}}{m_{\text{min}}} \right) \left( \frac{1}{m_{\text{frag}}} \right) . \] (9)

For fiducial values let us assume an initial planetesimal population that follows a streaming instability distribution with \( p = 1.6 \) from \( r_{\text{min}} = 1 \text{ km} \) up to \( r_{\text{max}} = 100 \text{ km} \) with bulk densities \( \rho = 1 \text{ g cm}^{-3} \). We further assume that, as in our simulations (Fig. 2), \( f = 1 \text{ per cent} \) of planetesimals are disrupted and produce fragments of \( r_{\text{frag}} = 100 \text{ m} \) in size (similar to 'Oumuamua). With those assumed values, the number of fragments dominates the total distribution, with \( N_{\text{frag}}/N \approx 3 \text{ per cent} \). If we instead assume that the fragments created match the minimum in the already-established size distribution, i.e. that \( r_{\text{frag}} = r_{\text{min}} \), then \( N_{\text{frag}}/N \approx 4 \).

In Section 2 we showed that assuming a mono-size distribution for all interstellar objects yielded an underestimate of the true mass. This was quantified in equation (5) using the bias \( b \), which was typically two or more orders of magnitude. This discrepancy is alleviated in a natural fashion if we invoke binary systems.

\[ \frac{N_{\text{frag}}/N}{b} = \left( \frac{r_{\text{min}}}{r_{\text{frag}}} \right)^3 . \] (10)

If \( N_{\text{frag}}/N > b \) then fragments dominate the population of interstellar objects to a sufficient extent that the mono-disperse value from equation (3) overestimates the true value. If \( N_{\text{frag}}/N < b \) the mono-disperse value still underestimates the true value.

Both an underestimate and an overestimate are entirely within reason. For plausible choices of parameters \( p = 1.6 \), \( f = 1 \text{ per cent} \), \( r_{\text{min}} = 1 \text{ km} \), and \( r_{\text{frag}} = 100 \text{ m} \) – the mass inferred from a mono-sized distribution overestimates the true one by a factor of 6.6. This would reduce Trilling et al. (2017)’s and Do et al. (2018)’s estimate to \( 3 \text{ M}_\oplus \) and \( 6 \text{ M}_\oplus \) in planetesimals ejected per star, respectively. For only slightly different parameters – keeping \( p = 1.6 \) and \( r_{\text{frag}} = 1 \text{ km} \) but decreasing \( f \) to 0.5 per cent and increasing \( r_{\text{frag}} \) to 200 m – the mass inferred from a mono-sized distribution underestimates the real one by a factor of 2.4. This would increase Do et al. (2018)’s estimate to \( \sim 10 \text{ M}_\oplus \) ejected per star. Given the large uncertainty in parameters it is premature to overinterpret these inferred masses.

It is interesting that our nominal parameters come close to entirely cancelling out the bias. If we naively equate \( N_{\text{frag}}/N \) and \( b \) we can solve for the fraction of planetesimals that must disrupt on the way to ejection to perfectly cancel each other out:

\[ f = 2^{p-1} \left( \frac{r_{\text{frag}}}{r_{\text{min}}} \right)^3 . \] (11)

Assuming that fragments are much smaller than the smallest planetesimals that form by the streaming instability, i.e. \( r_{\text{frag}} = 0.1 \text{ r}_{\text{min}} \), only 0.15 per cent of ejected planetesimals need to perfectly cancel out the bias. If fragments are slightly larger \( r_{\text{frag}} = 0.2 \text{ r}_{\text{min}} \) – then \( f \approx 1 \text{ per cent} \), as we found in our simulations (Fig. 2).

If roughly 1 per cent of planetesimals undergo tidal disruption events on their way to ejection, and if the typical fragment size is modestly smaller than the smallest bodies in the initial distribution, then the large difference (quantified by \( b \) in equation 5) between estimates of the mass in interstellar planetesimals for mono-disperse and realistic size distributions decreases significantly and may even disappear. This would imply that the distribution of interstellar objects is dominated by fragments by number and by the largest planetesimals by mass. If this interpretation is correct we would expect future interstellar objects to be dominated by objects similar in size to ‘Oumuamua and also to show signs of tidal distortion via their stretched-out shapes. Alternately, if this interpretation is wrong and the true distribution is similar to the streaming one, then we would expect a much lower rate of future discoveries than predicted (e.g. by Trilling et al. 2017).

Finally, at face value Fig. 2 shows that asteroidal planetesimals are disrupted before ejection at a higher rate than cometary planetesimals. Given the potential for a large number of fragments per disruption event, a relatively small difference in efficiency could create a large discrepancy by number. This might tilt the scales in favour of detecting asteroidal interstellar planetesimals rather than cometary ones, in contrast with the arguments presented in Section 3.1. However, this interpretation remains premature. Asteroids have higher physical densities than comets (Carr 2012), with correspondingly smaller tidal disruption radii. The shrinking value of \( R \), would shift the asteroidal curve to the right in Fig. 2, towards the cometary curve. For factor of two higher physical density for asteroids, the difference in disruption efficiency between ejected asteroids and comets vanishes. In fact, it may be that disruption is extremely rare for asteroids.

5 COMPARISON WITH OTHER ORIGIN HYPOTHESES FOR ‘OUMUAMUA

We have argued that single stars with gas giants can account for the abundance and properties of ‘Oumuamua. In our scenario, ‘Oumuamua is likely to be the fragment of an extinct comet that was born beyond the snow line of its parent star.4 Planetesimals will also be ejected (and may be tidally disrupted) from binary star systems. Binary systems are abundant in the Galaxy (Duchêne & Kraus 2013), and a binary companion provides stronger dynamical perturbations than a gas giant. Typically, however, we would expect a binary to interact with a much smaller fraction of surrounding planetesimals than an (initially unstable) giant planet system, and hence there is a trade-off which must be assessed quantitatively in order to determine if binary ejections are significant. For completeness, we briefly discuss two other origin hypotheses for ‘Oumuamua that invoke binary systems.

Jackson et al. (2017) proposed that most interstellar planetesimals form in circumbinary discs. Close binary stars have well-defined dynamical stability limits, as objects that enter within a critical orbital radius are destabilized (Holman & Wiegert 1999). Jackson et al. (2017) assume that planetesimals form beyond the stability limit and systematically drift inward across the limit under the effect of aerodynamic gas drag. They performed simulations to show that all planetesimals that drift inside the stability limit are ejected. For a significant fraction of close binaries the snow line is exterior to the stability limit such that some ejected planetesimals may be volatile-poor. This process could dominate the population of interstellar planetesimals if two pre-conditions are met. First, a sufficiently large fraction of the entire solid disc mass must drift interior to the binary stability limit [Jackson et al. (2017) assume this fraction is 10 per cent]. If the drift occurs while the solids are in the form of

4 Several studies have tried to pinpoint the Galactic origins of ‘Oumuamua but there is currently insufficient evidence to clearly link its trajectory to a specific star or star-forming region (Gaidos et al. 2017; Mamajek 2017; Zuluaga et al. 2017; Zhang 2018).
planetesimals, this requires relatively small planetesimals (r ≤ 1 km) which can drift significantly under aerodynamic forces during the disc lifetime (e.g. Adachi, Hayashi & Nakazawa 1976; Thommes, Duncan & Levison 2003). Secondly, the drifting planetesimals need to enter the dynamically unstable zone close to the binary, rather than piling up at the pressure maximum in the gaseous circumbinary disc which may be at modestly greater orbital radii (Artymowicz & Lubow 1994; Pierens & Nelson 2007; Marzari, Thébault & Scholl 2008; Lines et al. 2016).

Čuk (2018) proposed that ‘Oumuamua is a fragment of a planet that formed in a binary system and was tidally disrupted after passing too close to one of the stars, which itself must be sufficiently dense to allow for tidal disruption rather than collision (this favours M dwarf stars). This would represent a low-probability event but one that creates enough fragments, if ejected, to flood the distribution of interstellar planetesimals. Čuk (2018) suggests that these events are possible in S-type binary systems, in which the binary companion is on a wider orbit than the planet. It is likely, however, that planet–planet scattering provides one of the dominant channels for triggering planet–star near-collisions, and hence the ‘direct’ planetesimal ejection / tidal disruption mechanism that we have discussed would necessarily accompany planetary disruptions. It is also unclear whether planetary tidal disruptions would yield fragments of the size inferred for ‘Oumuamua.

6 CONCLUSIONS

The observation of ‘Oumuamua provides a first crude estimate of the number density of interstellar planetesimals. Converting that number density to a mass is model-dependent and highly uncertain. By assuming a mono-disperse source population, we calculated a seemingly reasonable value of 1 M⊙ in ejected planetesimals per stellar mass in stars (Section 2). However, for a mass function that is dominated by small bodies, but by mass by large ones, the first observation will almost always be of an object close to the minimum size. Inferring an interstellar mass density from the derived number density then leads to an underestimate of the typical ejected mass, typically by several orders of magnitude. Future observations of interstellar objects therefore have the potential to place unique constraints on the mass distribution of planetesimals, with the caveat that the ejected bodies do not necessarily fairly sample the primordial distribution.

Dynamical considerations suggest that systems with gas giants are responsible for ejecting the bulk of interstellar planetesimals (Section 3). Cometary planetesimals from beyond the gas giants’ orbits are ejected with a much higher efficiency than asteroidal planetesimals from closer-in. Ejection is most efficient in systems in which the gas giants themselves become unstable and undergo a phase of planet–planet scattering, thus ejecting the bulk of outer planetesimal discs (see Fig. 1 and Raymond et al. 2011, 2012).

During the process of ejection, some planetesimals may pass so close to giant planets that they are tidally disrupted (Section 4). This process requires planetary densities that are comparable to or higher than (present-day) Jupiter, and hence is sensitive to the abundance of super-Jovian planets and to the timing of planet-planet scattering. However, it appears possible that of the order of 1 per cent of planetesimals could be tidally disrupted during ejection. If those planetesimals are shredded into much smaller (‘Oumuamua-sized) fragments, then the fragments dominate by number (but not by mass). For reasonable parameter choices, the mass inferred for a mono-disperse interstellar population approaches the true mass of a more complex distribution, although the mono-disperse mass may still under- or overestimate the true mass by a large factor (see equation 10). If ‘Oumuamua represents a fragment of a disrupted planetesimal, then the disruption event itself may be the key to explaining its shape and tumbling rotation (Drahus et al. 2017; Fraser et al. 2017).

ACKNOWLEDGEMENTS

We thank the referee for a helpful report. We acknowledge discussions with Arnaud Pierens, Franck Selsis, Jean-Marc Huré, and Jeremy Leconte, and helpful comments from Karen Meech, Robert Jedicke, Darin Ragozzine, and Jorge Zuluaga. DV acknowledges the support of the STFC via an Ernest Rutherford Fellowship (grant ST/P003850/1). PJA acknowledges support from NASA through grant NNX16AB42G. SNR thanks Agence Nationale pour la Recherche grant ANR-13-BS05-0003-002 (grant MOJO) and NASA Astrobiology Institute’s Virtual Planetary Laboratory Lead Team, funded via the NASA Astrobiology Institute under solicitation NNH12ZDA002C and cooperative agreement no. NNA13AA93A.

REFERENCES

Adachi I., Hayashi C., Nakazawa K., 1976, Prog. Theor. Phys., 56, 1756
Adams F. C., Laughlin G., 2003, Icarus, 163, 290
ALMA Partnership, 2015, ApJ, 808, L3
Artymowicz P., Lubow S. H., 1994, ApJ, 421, 651
Asher D. J., Bailey M. E., Hahn G., Steel D. I., 1994, MNRAS, 267, 26
Barnes J. M. T. et al., 2017, ApJ, 851, L38
Baraffe I., Chabrier G., Barman T. S., Allard F., Hauschildt P. H., 2003, A&A, 402, 701
Barclay T., Quintana E. V., Raymond S. N., Penny M. T., 2017, ApJ, 841, 86
Bear E., Soker N., 2015, MNRAS, 450, 4233
Benz W., Asphaug E., 1999, Icarus, 142, 5
Bolin B. T. et al., 2018, ApJ, 852, L2
Bonsor A., Augereau J.-C., Thébault P., 2012, A&A, 548, A104
Bus S. J., Binzel R. P., 2002, Icarus, 158, 146
Butler R. P. et al., 2006, ApJ, 646, 505
Carry B., 2012, Planet. Space Sci., 73, 98
Carry B., Solano E., Eggl S., DeMeo F. E., 2016, Icarus, 268, 340
Chambers K. C. et al., 2016, preprint (arXiv:1612.05560)
Charnoz S., Morbidelli A., 2003, Icarus, 166, 141
Chatterjee S., Ford E. B., Matsumura S., Rasio F. A., 2008, ApJ, 686, 580
Cook N. V., Ragozzine D., Granvik M., Stephens D. C., 2016, ApJ, 825, 51
Cordes J. M., Shannon R. M., 2008, ApJ, 682, 1152
Coughlin E. R., Nixon C., Begelman M. C., Armitage P. J., Price D. J., 2016, MNRAS, 455, 3612
Cumming A., Butler R. P.,Marcy G. W., Vogt S. S., Wright T. J., Fischer D. A., 2008, PASP, 120, 531
Čuk M., 2018, ApJ, 852, L15
de la Fuente Marcos C., de la Fuente Marcos R., 2017, Res. Notes Am. Astron. Soc., 1, 5
Deienno R., Morbidelli A., Gomes R. S., Nesvorný D., 2017, AJ, 153, 153
Do A., Tucker M. A., Tonry J., 2018, ApJ, 855, 5
Dohanyi J. S., 1969, J. Geophys. Res., 74, 2531
Drahus M., Guzik P., Waniak W., Handzlik B., Kurowski S., Xu S., Xu S., 2017, preprint (arXiv:1712.00437)
Dressing C. D., Charbonneau D., 2015, ApJ, 807, 45
Dražkowska J., Alibert Y., Moore B., 2016, A&A, 594, A105
Duchêne G., Kraus A., 2013, ARA&A, 51, 269
Duncan M., Quinn T., Tremaine S., 1987, AJ, 94, 1330
Engelhardt T., Jedecie R., Vereš P., Fitzsimmons A., Denneau L., Beshore E., Meinke B., 2017, AJ, 153, 133
Fernandez J. A., Ip W., 1984, Icarus, 58, 109
