Implications of “peak oil” for atmospheric CO$_2$ and climate

Pushker A. Kharecha$^1$ and James E. Hansen$^1$

Received 26 November 2007; revised 11 March 2008; accepted 18 March 2008; published 5 August 2008.

[1] Unconstrained CO$_2$ emission from fossil fuel burning has been the dominant cause of observed anthropogenic global warming. The amounts of “proven” and potential fossil fuel reserves are uncertain and debated. Regardless of the true values, society has flexibility in the degree to which it chooses to exploit these reserves, especially unconventional fossil fuels and those located in extreme or pristine environments. If conventional oil production peaks within the next few decades, it may have a large effect on future atmospheric CO$_2$ and climate change, depending upon subsequent energy choices. Assuming that proven oil and gas reserves do not greatly exceed estimates of the Energy Information Administration, and recent trends are toward lower estimates, we show that it is feasible to keep atmospheric CO$_2$ from exceeding about 450 ppm by 2100, provided that emissions from coal, unconventional fossil fuels, and land use are constrained. Coal-fired power plants without sequestration must be phased out before midcentury to achieve this CO$_2$ limit. It is also important to “stretch” conventional oil reserves via energy conservation and efficiency, thus averting strong pressures to extract liquid fuels from coal or unconventional fossil fuels while clean technologies are being developed for the era “beyond fossil fuels”. We argue that a rising price on carbon emissions is needed to discourage conversion of the vast fossil resources into usable reserves, and to keep CO$_2$ beneath the 450 ppm ceiling.

Citation: Kharecha, P. A., and J. E. Hansen (2008), Implications of “peak oil” for atmospheric CO$_2$ and climate, Global Biogeochem. Cycles, 22, GB3012, doi:10.1029/2007GB003142.

1. Introduction

[2] M. King Hubbert, the late petroleum geologist and Shell oil company consultant, articulated the notion that oil production would peak when about half of the economically recoverable resource had been exploited. His successful prediction of peak oil production in the continental United States [Hubbert, 1956] has encouraged numerous analysts to subsequently apply his model or variations thereof to global oil production. The concept of peak extraction of a finite nonrenewable resource constrained by geology and geography has received support from similar patterns of growth, peak production, and decline of mineral resources [van der Veen, 2006], natural gas [Lam, 1998], and coal [Milici and Campbell, 1997] in specific regions.

[3] There is intense disagreement about when global “peak oil” might occur, but it is widely accepted that it will occur at some point this century [Wood et al., 2003; Kerr, 2005]. Despite the obvious relevance of peak oil to future climate change, it has received little attention in projections of future climate change.

[4] In this paper we emphasize the estimated magnitudes of fossil fuel resources (oil, natural gas, and coal), and the relevance of these limitations to the question of how practical it may be to avoid “dangerous anthropogenic interference” with global climate as outlined in the United Nations Framework Convention on Climate Change [UNFCCC, 1992]. We are motivated by the conclusion of Hansen et al. [2007a, 2007b] that “dangerous” climatic consequences are expected at atmospheric CO$_2$ levels exceeding 450 ppm and possibly at even lower levels. Thus we investigate whether atmospheric CO$_2$ can be kept to 450 ppm or less via constraints on the use of coal and unconventional fossil fuel resources.

[5] In estimating atmospheric CO$_2$ levels for given emission scenarios we employ both a linear atmospheric pulse response function (PRF) fit to the Bern carbon cycle model [Joos et al., 1996] and a nonlinear mixed-layer PRF fit to the same Bern carbon cycle model. The mixed-layer (or dynamic-sink) PRF allows nonlinear ocean carbonate chemistry, which is not included in the atmospheric (static-sink) PRF. The static-sink PRF underestimates CO$_2$ levels for large emission cases, but we include results for this PRF because our main interest is cases that keep atmospheric CO$_2$ at ~450 ppm, this PRF has the advantage of simplicity and transparency, and we have used this function in other studies [Hansen et al., 2007a, 2007b]. We also include results for the dynamic-sink PRF.

[6] We do not attempt to resolve the debate about the timing of peak oil or the magnitude of fossil fuel resources; rather, we consider reasonable alternative assumptions. We recognize that the magnitude of recoverable oil and gas resources depends on economic incentives and penalties.

---

$^1$NASA Goddard Institute for Space Studies and Columbia University Earth Institute, New York, New York, USA.

Copyright 2008 by the American Geophysical Union.

0886-6236/08/2007GB003142
units of gigatons of carbon (1 Gt C = 1 Pg C = 10^{15} g C) for the sake of uniformity, as well as in units of atmospheric CO$_2$ equivalent (1 ppm CO$_2$ = 2.12 Gt C).

[10] “Proven reserves” are the amounts of these fuels that are estimated to be economically recoverable under current economic and environmental conditions with existing technology. “Resources” are fossil occurrences whose existence is well-known but whose recoverable magnitudes are less certain; however, they are widely believed to contain immense amounts of fossil energy and carbon [e.g., see IPCC, 2001a]. “Reserve growth” is defined by EIA [2006] as expected additions to proven reserves (from resources) based on realistic expected improvements in extraction technologies.

[11] Reservoir estimates tend to be larger if they are made under the assumption of very high fuel prices, or if it is assumed that greater technology advances will allow recovery of a much higher percentage of fossil fuels in existing fields. On the other hand, if a substantial carbon price is applied to CO$_2$ emissions in the future, the available reservoir may decrease, as it becomes unprofitable to extract resources from remote locations or to squeeze hard-to-extract resources from existing fields. Because of these uncertainties, we also consider a range of estimates of fossil fuel reservoirs.

[12] “Unconventional” fossil fuels are those that exist in a physical state other than conventional oil, gas and coal. The contribution of unconventional fossil fuels to CO$_2$ emissions is negligible to date [IPCC, 2001a]. We do not include unconventional fossil fuels in the scenarios that we illustrate, because we are interested primarily in scenarios that cap atmospheric CO$_2$ at 450 ppm or less. However, it should be borne in mind that unconventional fossil fuels could contribute huge amounts of atmospheric CO$_2$ if the world should follow an unconstrained “business-as-usual” scenario of fossil fuel use.

2.2. CO$_2$ Emissions Scenarios

[13] We illustrate five CO$_2$ emissions scenarios for the period 1850–2100. The first case, Business-As-Usual (BAU), assumes continuation of the ~2% annual growth of fossil fuel CO$_2$ emissions that has occurred in recent decades [EIA, 2006; Marland et al., 2006]. This 2% annual growth is assumed to continue for each of the three conventional fuels until ~half of each total reservoir (historic + remaining) has been exploited, after which emissions are assumed to decline 2% annually.

[14] The second scenario, labeled Coal Phase-out, is meant to approximate a situation in which developed countries freeze their CO$_2$ emissions from coal by 2012 and a decade later developing countries similarly halt increases in coal emissions. Between 2025 and 2050 it is assumed that both developed and developing countries will linearly phase out emissions of CO$_2$ from coal usage. Thus in Coal Phase-out we have global CO$_2$ emissions from coal increasing 2% per year until 2012, 1% per year growth of coal emissions between 2013 and 2022, flat coal emissions for 2023–2025, and finally a linear decrease to zero CO$_2$ emissions from coal in 2050. These rates refer to emissions to the atmosphere and do not constrain consumption of coal,
provided the \( \text{CO}_2 \) is captured and sequestered. Oil and gas emissions are assumed to be the same as in the BAU scenario.

[15] The third, fourth, and fifth scenarios include the same phase-out of coal but investigate the effect of uncertainties in global oil usage and supply. The Fast Oil Use scenario adopts an alternative approach for calculating the peak in global oil emissions/usage, following the method of Wood et al. [2003]. It assumes that 2% annual growth in oil use continues past the midpoint of oil supplies, until the ratio of remaining reserves to emissions decreases to 10 years from the current value of \( \sim 50 \) years. This scenario causes “peak oil” to be delayed \( \sim 21 \) years to 2037. The fourth scenario, Less Oil Reserves, uses the same trends as in Coal Phase-out but omits the oil “reserve growth” term. This fourth scenario may be most relevant to a situation in which oil companies have been too optimistic about reserves and/or a high price on carbon emissions discourages exploration for oil in remote locations. Finally, the fifth scenario, Peak Oil Plateau, assumes that oil emissions exhibit a sustained peak from 2020–2040, using supply and 21st-century usage estimates of R. Nehring [Kerr, 2007]. It assumes the global oil reserve base is \( \sim 50 \) Gt C larger than in our other scenarios (R. Nehring, personal communication, 14 June 2007; http://www.aapg.org/explorer/2007/05may/nehring.cfm). This scenario reflects the possibility that the peak productivity of major oil fields may occur at different times over the next several decades, leading to an extended, rather than abrupt, global oil peak.

[16] In addition to fossil fuel \( \text{CO}_2 \) emissions, we also include historical and projected estimates for net emissions from land use in each of the above scenarios. We use historical (1850–2000) land use emissions estimates from CDIAC [Houghton and Hackler, 2002], and projections from the midrange IPCC Special Report on Emissions Scenarios (SRES) A1T scenario [values from IPCC, 2001b].

[17] These illustrative scenarios cover a broad range of fossil fuel reserves, but more extreme estimates do exist. We quantitatively investigate the effect of very low estimates of natural gas and crude oil reserves [IPCC, 2001a] and coal reserves [WEC, 2007]. Although some analysts estimate reserves exceeding those in the illustrated scenarios, we do not study those because the cases we consider already reach far into the range of “dangerous” atmospheric \( \text{CO}_2 \) levels.

### 2.3. Atmospheric \( \text{CO}_2 \) Projections

[18] For each of the \( \text{CO}_2 \) emissions scenarios we generate a time series of atmospheric \( \text{CO}_2 \), using the following parameterization of the Bern carbon cycle model of Joos et al. [1996]:

\[
\text{\text{CO}_2}(t) = 18 + 14\exp(-t/420) + 18\exp(-t/70) + 24\exp(-t/21) + 26\exp(-t/3.4)
\]

(1)

where \( \text{CO}_2 (t) \) is the percentage of emitted \( \text{CO}_2 \) remaining in the atmosphere after \( t \) years, and the coefficients of each term are rounded from those provided by Shine et al. [2005]. Note that equation (1) implies that about one-third of anthropogenic \( \text{CO}_2 \) emissions remain in the atmosphere after 100 years and one-fifth after 1000 years.

[19] Equation (1) is a static-sink PRF for anthropogenic \( \text{CO}_2 \) emissions, i.e., it is an approximation for the proportion (percent) of \( \text{CO}_2 \) remaining airborne \( t \) years following an emission pulse. The time evolution of atmospheric \( \text{CO}_2 \) is obtained by taking the 1850 \( \text{CO}_2 \) concentration as 285.2 ppm [Etheridge et al., 1998], recursively applying equation (1) to the emission scenario, and integrating the results from 1850 to year \( t \).

[20] We also investigate the same scenarios using a dynamic-sink PRF that incorporates some carbon cycle feedbacks. Both approximations are based on the Bern carbon cycle model with the HILDA and 4-box biosphere models [equations (3)–(6), (16)–(17), A.2.2, and A.3 of Joos et al., 1996]. The dynamic-sink PRF includes simplified nonlinear ocean carbonate chemistry as well as biospheric carbon uptake and respiration (both for a fixed climate). We demonstrate that the dynamic-sink PRF yields slightly different quantitative results than the static-sink PRF (equation (1)), but the differences are small enough to only reinforce our conclusions.

### 3. Results

#### 3.1. Historical \( \text{CO}_2 \) Emissions and Concentrations

[21] Integration of the product of equation (1) and fossil fuel emissions over the period 1850–2007 yields an airborne fossil fuel \( \text{CO}_2 \) amount of \( \sim 85 \) ppm in 2007. The land use net emissions estimates of Houghton and Hackler [2002], with SRES A1T estimates for 2001–2007, yield an additional contribution of \( \sim 35 \) ppm in 2007. The \( \text{CO}_2 \) amount in 2007 is thus overestimated by \( \sim 20 \) ppm using equation (1), or by \( \sim 13 \) ppm when the dynamic-sink PRF is used (Figure 2). We suggest that, rather than a model deficiency, this overestimate is probably due to overestimate of net land use emissions (particularly for 1950–2000), based on the large uncertainties inherent in those estimates [Houghton, 2003], in contrast with the relatively high certainty of fossil fuel emissions estimates. Part of this discrepancy also may be due to the carbon cycle model. When the land use emissions are reduced by 50%, as supported by other studies [e.g., see IPCC, 2007, Ch. 7], the model-data differences amount to at most \( \sim 2 \%) from 1850–2007 (Figure 2). Although part of the discrepancy could be a result of “fertilization” of the biosphere, via anthropogenic \( \text{CO}_2 \) and nitrogen emissions, we show in a paper in preparation that incorporation of these effects in the carbon cycle model, as an alternative to reducing the land use source, has negligible impact on our present investigation. All of our scenarios therefore assume the above reduction in the Houghton and Hackler [2002] estimates.

[22] This calculation for 1850–2007 provides a check on the reasonableness of the carbon cycle approximation (equation (1)) for \( \text{CO}_2 \) in the range 280–385 ppm. We infer that the model may continue to provide useful estimates for scenarios with moderate fossil fuel emissions, i.e., for the scenarios of special interest that keep atmospheric \( \text{CO}_2 \) less than or approximately 450 ppm.

[23] As mentioned above and discussed in section 3.4, for the larger \( \text{CO}_2 \) emissions of BAU scenarios, equation (1) may begin to underestimate airborne \( \text{CO}_2 \), as it excludes the
nonlinearity of the ocean carbon cycle as well as anticipated climate feedbacks on atmospheric CO$_2$ and CH$_4$, the latter being eventually oxidized to CO$_2$.

### 3.2. Projected CO$_2$ Emissions

Table 1 provides an overview of the fossil fuel emissions and resulting atmospheric CO$_2$ amounts in our five emissions scenarios. We list peak fossil fuel emission years as ranges where necessary, to reflect minor differences in historical emissions estimates. (For example, there are differences between the historic emissions of CDIAC [Marland et al., 2006] and EIA [2006]. Also, relatively minor changes arise from year to year in the CDIAC fossil fuel data due to retroactive updates to the UNSTAT database (T. Boden, personal communication, 5 June 2007.).

**Table 1.** Approximate Peak Fossil Fuel CO$_2$ Emissions and Atmospheric CO$_2$ Levels in Each Scenario

| Scenario          | Peak Fuel Emission | Year of Peak | Peak Fuel CO$_2$ Level | Year of Peak |
|-------------------|--------------------|--------------|------------------------|--------------|
| BAU               | 14 Gt C a$^{-1}$   | 2077 ± 2 a   | 563 ppm                | 2100         |
| Coal phase-out    | 10 Gt C a$^{-1}$   | 2016 ± 2 a   | 428 ppm                | 2047         |
| Fast oil use      | 11 Gt C a$^{-1}$   | 2025 ± 2 a   | 446 ppm                | 2046         |
| Less oil reserves | 9 Gt C a$^{-1}$    | 2022 ± 2 a   | 422 ppm                | 2045         |
| Peak oil plateau  | 10 Gt C a$^{-1}$   | 2025 ± 2 a   | 440 ppm                | 2060         |

Additional features of our four mitigation scenarios are listed in Table 2. Figure 3 shows CO$_2$ emissions over time for the five main scenarios and Figure 4 shows the resulting atmospheric CO$_2$ concentrations.

**Table 2.** Salient Features and Metrics of Mitigation Scenarios

| Scenario          | 2007–2050 Coal Emissions | Total 2007–2050 Fuel Emissions | Reduction in 2050 Versus 2007 Fuel Emissions |
|-------------------|--------------------------|--------------------------------|---------------------------------------------|
| Coal phase-out    | ~110 Gt C               | ~330 Gt C                      | 57%                                         |
| Fast oil use      | ~110 Gt C               | ~390 Gt C                      | 54%                                         |
| Less oil reserves | ~110 Gt C               | ~300 Gt C                      | 66%                                         |
| Peak oil plateau  | ~110 Gt C               | ~360 Gt C                      | 40%                                         |

[25] Peak oil emission in the BAU scenario occurs in 2016 ± 2 a, peak gas in 2026 ± 2 a, and peak coal in 2077 ± 2 a (Figure 3a). Coal Phase-out moves peak coal up to 2022 (Figure 3b). Fast Oil Use causes peak oil to be delayed until 2037 [Wood et al., 2003], but oil use then crashes rapidly (Figure 3c). Less Oil Reserves results in peak oil moving to 2010 ± 2 a (Figure 3d), under the assumption that usage approximates the near symmetrical shape of the classical Hubbert curve. In the Peak Oil Plateau case, oil emissions peak in 2020 and remain at that level until 2040 [Kerr, 2007], thereafter decreasing approximately linearly (Figure 3e).

[26] Total fossil fuel CO$_2$ emissions peak in 2077 ± 2 a at ~14 Gt C a$^{-1}$ in the BAU scenario, almost double the current level, decreasing to ~9 Gt C a$^{-1}$ in 2100 (Figure 3a). Fuel emissions peak in 2016 ± 2 a at ~10 Gt C a$^{-1}$ in the Coal Phase-out scenario, decreasing to ~1 Gt C a$^{-1}$ by 2100 (Figure 3b). Cumulative 21st-century fossil fuel emissions are ~1100 Gt C in the BAU scenario, ~500 Gt C in the Coal Phase-out scenario, ~520 Gt C in the Fast Oil use scenario, ~430 Gt C in the Less Oil Reserves scenario, and ~550 Gt C in the Peak Oil Plateau scenario.

### 3.3. Projected Atmospheric CO$_2$ Concentrations

[27] Figure 4 shows atmospheric CO$_2$ concentrations resulting from our fossil fuel emissions scenarios. All of our scenarios include land use CO$_2$ emissions from the midrange SRES marker scenario (A1T). Additional calculations (not shown here) reveal that the land use emissions projections of the high-end scenario (A2) would add ~5 ppm to all of our CO$_2$ projections, while the low-end projections (B1) would cause a reduction by about the same amount.

[28] Peak CO$_2$ in the BAU scenario is ~575 ppm in 2100, with fuel emissions alone raising CO$_2$ to over 560 ppm (Figure 4a). This is more than double the preindustrial CO$_2$ amount of ~280 ppm and already far past the 450 ppm threshold under consideration. Likely nonlinearities in the carbon cycle with such large CO$_2$ amounts would make the real-world peak CO$_2$ even greater, as would any contribution from unconventional fossil fuels.

[29] Our interest is primarily in scenarios that limit atmospheric CO$_2$ to ~450 ppm or less. Therefore all scenarios other than BAU include phase-out of coal emissions at least as rapidly as in the standard Coal Phase-out scenario.

Figure 2. Computed versus observed time evolution of industrial-era atmospheric CO$_2$ from 1850–2007. (CDIAC LU = Houghton and Hackler [2002] land use emissions for 1850–2000). Inclusion of the CDIAC LU emissions causes increasing overestimation of pCO$_2$ by the model between 1950–2000 for both the static-sink and dynamic-sink PRFs, suggesting that those LU estimates may be overestimates (see section 3.1). When the CDIAC LU estimates are reduced by 50%, both PRFs produce very good agreement with observed CO$_2$. Observations prior to 1958 are based on Law Dome ice core data [Etheridge et al., 1998], and from 1958 onwards based on high-precision flask and in situ measurements [Keeling and Whorf, 2005; Conway et al., 2007; Thoning et al., 2007], with the specific data series as concatenated and adjusted to global means by Hansen and Sato [2004].
scenario, which has moderate continued growth of global coal emissions until 2025, followed by linear phase-out of global emissions from coal between 2025 and 2050. This standard Coal Phase-out scenario has peak atmospheric CO$_2$ at ~445 ppm in 2046; fossil fuels alone raise CO$_2$ to ~428 ppm in 2047 (Figure 4b). 

The Fast Oil Use scenario (Figure 4c) yields a peak atmospheric CO$_2$ level of ~463 ppm with fossil fuels raising CO$_2$ to ~446 ppm, i.e., faster use of the same oil amount increases the peak atmospheric amount by about 18 ppm. However, in the absence of carbon feedbacks, this difference decreases with time, practically disappearing by 2100.

The Less Oil Reserves scenario (Figure 4d) yields a peak atmospheric CO$_2$ level of ~439 ppm, with fossil fuels raising CO$_2$ to ~422 ppm. Thus omission of oil reserve growth (Figure 1) reduces the peak atmospheric CO$_2$ amount by ~6 ppm from the baseline Coal Phase-out scenario.

Last, the Peak Oil Plateau scenario (Figure 4e) yields a peak atmospheric CO$_2$ level of ~456 ppm, with fossil fuels raising CO$_2$ to ~440 ppm. The sustained 20-year peak in global oil emissions leads to an increase of ~10 ppm from the baseline Coal Phase-out case.

### Additional Scenarios Considered

Effects of using the dynamic-sink PRF of Joos et al. [1996] are shown in Figures 5 and 6. The mean 1960–2007 total CO$_2$ airborne fraction of ~50% implied by the...
dynamic-sink PRF is lower than the ~53% based on the static-sink PRF, equation (1), and both are higher than the ~48% from observed airborne CO₂ and assumed fossil fuel and land use emissions (Figure 5). The difference in the responses of the two PRFs to a pulse of 5 ppm CO₂ (~10 Gt C) is not major, although it does persist for centuries (Figure 6a). Despite this, there is little difference (less than ~3%) between results from the two functions for the entire 21st century, regardless of whether high or moderate emissions are assumed (Figures 6b and 6c). However, the BAU scenario with unrestrained emissions diminishes the buffering capacity of the ocean, leading to a ~20 ppm increase in peak CO₂ (Figure 6b). On the other hand, with the lower emissions of the Coal Phase-out scenario, the oceanic/biospheric CO₂ uptake is similar for the two response functions (Figure 6c). [34] We also computed time series for emissions and atmospheric CO₂ levels for several alternative sets of conventional oil, gas, and coal reserve estimates, including one from the World Energy Council [WEC, 2007] and one from IPCC [2001a, Table 3.28b]. These estimates are lower than the EIA [2006] estimates (Figure 1), and therefore, even assuming BAU growth and decline (section 2.2), they yield earlier emissions peaks and lower peak atmospheric CO₂ levels (Figure 7). Specifically, WEC [2007] coal reserves (~450 Gt C) yield peak coal emissions in ~2040, and IPCC [2001a] oil and gas reserves (118 Gt C and 82 Gt C, respectively) yield peak oil emissions in ~2004 and peak gas emissions in ~2009 (Figure 7a).

Figure 5. CO₂ “airborne fraction” (AF) of 1960–2007 anthropogenic CO₂ emissions, computed as the mean measured atmospheric CO₂ concentration of a given year minus the amount in the previous year divided by either the fossil fuel emissions in the given year (AF1) or the sum of fossil fuel and land use emissions in the given year (AF2). The 1960–2007 mean derived from observed CO₂ is ~57% for the former (AF1 obs.) and ~48% for the latter (AF2 obs.). For the static-sink and dynamic-sink PRFs, the 1960–2007 model mean AF2 values are ~53% and ~50%, respectively. (See Figure 2 caption for CO₂ data sources.)

Figure 6. Effects of using the static-sink PRF versus the dynamic-sink PRF. (a) Annual fraction of CO₂ remaining airborne after a pulse emission of 5 ppm (~10 Gt C), with differences highlighted at 100, 500, and 1000 years. Atmospheric CO₂ concentrations for the baseline (b) BAU and (c) Coal Phase-out scenarios. The difference is generally negligible (less than ~3% throughout the 21st century), but in the high-emission BAU scenario the dynamic-sink PRF yields ~20 ppm greater peak CO₂ in 2100.
Figure 7. (a) Alternate “low-end” BAU emissions scenario assuming conventional oil and gas reserves from IPCC (2001a) with no reserve growth, and coal reserves from World Energy Council [WEC, 2007] (LU = land use). (b) Resulting atmospheric CO$_2$ from this scenario, compared with the baseline (“high-end”) BAU scenario from Figure 4a.

[35] Assuming the IPCC [2001a] oil and gas reserves along with WEC [2007] coal reserves yields a peak atmospheric CO$_2$ level of ~457 ppm in 2076, with fossil fuels alone raising CO$_2$ to ~442 ppm (Figure 7b). Although this latter scenario represents relatively moderate “BAU” cases, it relies heavily on the assumption that carbon-positive substitute fuels cannot or will not be developed in the future to replace declining conventional fuel reserves, e.g., due to a rising price on carbon emissions (see Discussion).

3.5. Comparison With IPCC-SRES and EMF-21 Scenarios

[36] In contrast with all of the above scenarios, peak total emissions in the four SRES scenario families [IPCC, 2000] range from ~12 Gt C a$^{-1}$ in 2040 (B1 marker scenario) to a staggering ~28 and 29 Gt C a$^{-1}$ in 2100 (A2 and A1F1 marker scenarios, respectively). Time-integrated 21st-century emissions for these SRES marker scenarios range from ~970 Gt C (B1) to ~1900 and 2100 Gt C (A2 and A1F1). Thus it is clear that the high-end SRES scenarios implicitly assume that, in the absence of climate mitigation policies, massive amounts of unconventional or “undiscovered” resources will become viable substitutes for dwindling conventional reserves.

[37] Resulting atmospheric CO$_2$ amounts in 2100 in the SRES scenarios range from ~540 ppm to ~970 ppm, excluding carbon cycle feedbacks [IPCC, 2001b]. Model simulations suggest that carbon cycle feedbacks under a high-end emissions scenario (A2) can yield an additional ~20–200 ppm of CO$_2$ by 2100 [Friedlingstein et al., 2006]. Note that our four mitigation scenarios, however, are consistent with current assessments of the cumulative 21st-century emissions needed to stabilize atmospheric CO$_2$ at 450 ppm even after factoring in carbon cycle feedbacks [e.g., see IPCC, 2007].

[38] Our scenarios also differ from those of EMF-21 [Weyant et al., 2006] in several ways. First, their scenarios have a long-term global radiative forcing target of ~4.5 W m$^{-2}$ above the preindustrial level, which yields a target global temperature rise of over 3°C, for nominal estimates of climate sensitivity [Hansen et al., 2007b; IPCC, 2007]. Also, their mean model ensemble 21st-century CO$_2$ emissions are assumed to continually increase from ~14 Gt C a$^{-1}$ in 2025 to ~25 Gt C a$^{-1}$ by 2100. Thus like the SRES scenarios, the EMF-21 scenarios reflect climate change that we would classify as “dangerous” and therefore highly undesirable.

4. Discussion

4.1. Avoidance of “Dangerous” Anthropogenic Climate Change

[39] Practically all nations of the world have agreed that a “dangerous” increase of atmospheric greenhouse gases should be avoided [UNFCCC, 1992], but the dangerous level of gases is not well defined. Hansen et al. [2007a, 2007b] have argued that additional global warming above that in 2000 must be kept less than 1°C, and that, therefore, the dangerous CO$_2$ level is at most about 450 ppm, and likely less than that. Although moderate trade-off with non-CO$_2$ gases is possible [Hansen and Sato, 2004], CO$_2$ is the most important climate forcing because a considerable fraction of fossil fuel CO$_2$ emissions remains in the atmosphere for many centuries [Archer, 2005]. Given that CO$_2$ has already increased during the industrial era from ~280 ppm to ~385 ppm, there is some urgency in determining what steps are practical to limit further growth of atmospheric CO$_2$. Indeed, the scenarios used in climate projections by the IPCC [2000, 2001a, 2001b, 2007] all have CO$_2$ increasing well beyond 450 ppm.

[40] On the other hand, fossil fuel reservoirs are finite, and existing deposits do not have to be fully exploited. Also, it may be practical to capture and sequester much of the CO$_2$ emitted in burning coal at power plants. Thus it is important to estimate expected atmospheric CO$_2$ levels for realistic estimates of fossil fuel reserves and to determine how the CO$_2$ level depends upon possible constraints on coal use.

[41] We suggest that, for the sake of simplicity and transparency, it is useful to make such estimates with simple pulse response functions for airborne CO$_2$, although similar studies should also be made with comprehensive carbon cycle models. We view the pulse response functions that we have employed as providing an approximate lower bound
for the proportion of fossil fuel CO₂ emissions that remain airborne. The uptake capacity of the carbon sinks may decrease if the CO₂ source increases, and there are potential climate feedbacks that could add CO₂ to the atmosphere, e.g., carbon emissions from forest dieback [Cox et al., 2000], melting permafrost [Walter et al., 2006; Zimov et al., 2006], and warming ocean floor [Archer, 2007]. In addition, land use change and deforestation may remain a significant source of positive anthropogenic climate forcing this century, e.g., Gruber et al. [2004] assert that land use-related human activities could lead to the release of up to ~40 Pg C over the next ~20 years and ~100 Pg C this century due to the alteration of live biomass pools in tropical and subtropical ecosystems.

On the other hand, if fossil fuel emissions of CO₂ decrease, concerns about possible nonlinear positive feedbacks are diminished. Indeed, we suggest a possible dichotomy of scenarios: if CO₂ emissions decrease, the proportion of CO₂ taken up by sinks could increase, with a resulting climate forcing that is much less than that in scenarios with continually increasing CO₂ emissions.

Given these basic considerations, we have focused on scenarios in which coal use is phased out except where the CO₂ is captured. We find that, with such an assumption, it is possible to keep maximum 21st-century atmospheric CO₂ less than 450 ppm, provided that the EIA estimates of oil and gas reserves and reserve growth are not significant underestimates. This limit on CO₂ is achieved in our scenarios only if cumulative global emissions from coal between the present and 2050 amount to ~100 Gt C or less. Thus even if coal reserves are much lower than historically assumed [e.g., NRC, 2007], there is surely enough coal to take the world past 450 ppm CO₂ without mitigation efforts such as those described here. On the bright side, our findings indicate that a feasible timescale for reductions can keep CO₂ below 450 ppm.

### 4.2. Future Use of Fossil Fuels

Goals for atmospheric CO₂ amount surely must be adjusted as knowledge about climate change and its impacts improves. Recent evidence of sea ice loss in the Arctic and accelerating net mass loss from the West Antarctic and Greenland ice sheets suggest that the allowable level of warming is likely less than 1°C above the 2000 global temperature and the CO₂ limit is likely less than 450 ppm. Thus details about the magnitude of fossil fuel reserves and the rate at which the reserves are exploited may be important. We find that the maximum 21st-century atmospheric CO₂ level varies by ~18 ppm depending upon the rate at which given oil and gas reserves are consumed. This variation decreases with time, however, so the size of the exploited oil and gas reservoirs is a more important consideration.

The size of economically recoverable oil and gas resources is flexible, depending upon the degree to which fossil fuels are priced to cover their environmental costs. Thus we have argued [Hansen et al., 2007a; Hansen, 2007] in favor of placing a significant rising price on CO₂ emissions. One effect of a rising carbon price would be to slow the rate at which fossil fuel resources are exploited, thus reducing the maximum atmospheric CO₂ amount, as illustrated above. More importantly, a carbon price would result in some of the oil, gas, and coal being left in the ground, primarily deposits at great depths or in extreme environmental locations. Given that the world must move beyond fossil fuels for its energy needs eventually, it is appropriate to encourage that transition soon, and thus minimize anthropogenic climate change. We note that there are various ways of placing a price on carbon emissions, such as a progressive carbon tax, industry “cap and trade” measures, or individual “ration and trade” measures. The pricing scheme should be chosen based on economic effectiveness and fairness.

Hirsch et al. [2005] note that it requires decades to remake energy infrastructure, and thus peaking of oil and gas production has the potential for severe economic disruption if steps are not taken to encourage technology development and implementation. This consideration adds to the need for prompt actions to conserve readily available oil and gas, thus stretching out these conventional supplies, while encouraging innovations in energy efficiency and alternative (nonfossil) energies. Stretching of supplies is a principal function of an increasing carbon price. Nuclear power could be one viable alternative option, if strict provisions are followed for public safety, waste disposal, and elimination of potential weapons-grade by-products; adoption of an international nuclear environmental treaty could be a significant step toward this end [Robock et al., 2007].

### 4.3. Additional Climate Change Mitigation Measures

Finally, we note that, as understanding of climate change and its impacts improves, it is possible that even lower limits on atmospheric CO₂ and the net anthropogenic climate forcing than discussed here may prove to be highly desirable. It has been suggested that to buy extra time to enact such large-scale mitigation, societies should adopt an approach that incorporates both emissions reductions and geoengineering options like periodic, sustained stratospheric sulfate aerosol injection [Crutzen, 2006; Wigley, 2006]. However, a geoengineering “quick fix”, if not sustained precisely to the degree and length of time needed, could do more harm than good [Matthews and Caldeira, 2007], and it is difficult to define how much “fix” is needed. Thus while geoengineering might provide some benefit, the potential gain must be weighed against long-term risks to climate and oceanic/stratospheric chemistry. Especially given the existence of low-cost and no-cost methods to reduce CO₂ emissions [Lovins, 2005], slowing of fossil fuel CO₂ emissions warrants highest priority.

Further reductions of anthropogenic climate forcing, beyond the 2025–2050 coal phase-out strategy that we quantified here, could be achieved as follows.

1. A freeze on new construction of traditional coal-fired power plants (without CO₂ sequestration) by 2010, with a linear phase-out of all such existing plants between 2010 and 2030. This action reduces the maximum atmospheric CO₂ from ~445 ppm in our standard Coal Phase-out scenario to ~400–430 ppm (depending on oil and gas...
reserve size). Fossil fuel contribution to atmospheric CO₂ level decreases from ~428 ppm to ~390–410 ppm.

[50] Intensive efforts to reduce non-CO₂ anthropogenic climate forcings, especially methane, tropospheric ozone, and black carbon. Hansen and Sato [2004] estimate that realistic potential savings from such reductions are equivalent to 25–50 ppm of CO₂.

[51] Anthropogenic draw-down of atmospheric CO₂. Farming and forestry practices that enhance carbon retention and storage in the soil and biosphere should be supported [McCarr and Sands, 2007], as should large-scale reforestation. Direct removal of CO₂ from the air through expedited carbonate formation also holds great potential [Lackner, 2003; Keith et al., 2006]. In addition, burning biofuels in power plants with carbon capture and sequestration can draw down atmospheric CO₂ [Hansen, 2007], in effect putting anthropogenic CO₂ back underground where it came from. However, careful measures must be taken to ensure that biofuel production does not occur at the expense of food crops and tropical forests that are not converted to biofuel farms. For instance, agricultural waste, natural grasses and other cellulosic material can be used [e.g., Tilman et al., 2006]. Fertilizers used in their production should minimize emission of non-CO₂ greenhouse gases as well. CO₂ sequestered beneath ocean sediments is inherently stable [House et al., 2006], and other safe geologic sites are also available.

[52] Acknowledgments. We thank Makiko Sato for providing the extrapolated historic CO₂ emissions data for 2004 and 2005 as well as the compiled historical CO₂ concentration data. We also thank Giosetta Petracic, Fortunat Joos, Jackson Harper, Dave Rutledge, and two anonymous reviewers for helpful comments on the manuscript. Research support was provided by Hal Harvey of the Hewlett Foundation, Gerry Lenfest, and NASA Earth Science Research Division managers Jack Kaye and Don Anderson.

References

Archer, D. (2005), Fate of fossil fuel CO₂ in geologic time, J. Geophys. Res., 110, C09S05, doi:10.1029/2004JC002625.

Archer, D. (2007), Methane hydrate stability and anthropogenic climate change, Biogeosci. Disc., 4, 993–1057.

British Petroleum (BP) (2006), Putting energy in the spotlight: BP Statistical Review of World Energy, June 2006 (Available at http://www.bp.com/pdf/statistical.review.of.world.energy.full.report2006.pdf).

Conway, T. J., P. M. Lang, and K. A. Masarie (2007), Atmospheric carbon dioxide dry air mole fractions from the NOAA ESRL carbon cycle cooperative global air sampling network, 1968–2006, Version: 2007-09-19 (Available at ftp://ftp.cmdl.noaa.gov/ccg/co2/flash/event/).

Cox, P. M., R. A. Betts, C. D. Jones, S. A. Spall, and J. J. Totterdell (2000), Acceleration of global warming due to carbon-cycle feedbacks in a coupled climate model, Nature, 408, 184–187.

Crutzen, P. (2006), Albedo enhancement by stratospheric sulfur injections: A contribution to resolve a policy dilemma?, Clim. Change, 77, 211–219.

Energy Information Administration (EIA), U.S. Dept. of Energy (2006), International Energy Outlook 2006 (Available at http://www.eia.doe.gov/oiaf/archive/eio06/index.html).

Etheridge, D. M., L. P. Steele, R. L. Langenfelds, R. J. Franckey, J.-M. Barnola and V. I. Morgan (1998), Historical CO₂ records from the Law Dome DE08, DE08-02, and DSS ice cores, in Trends: A Compendium of Data on Global Change, Carbon Dioxide Inf. Anal. Cen., Oak Ridge Natl. Lab., Dept. of Energy, U.S., Oak Ridge, Tenn.

Friedlingstein, P. et al. (2006), Climate-carbon cycle feedback analysis: Results from the C-MIP model intercomparison, Clim. J., 19, 3337–3353.

Gruber, N., P. Friedlingstein, C. B. Field, R. Valentini, M. Heimann, J. E. Richy, P. R. Lankao, E-D. Schulze, and C. Chen (2004), The vulnerability of the carbon cycle in the 21st century: An assessment of carbon-climate-human interactions, in The Global Carbon Cycle, edited by C. B. Field and M. R. Raupach, pp. 45–76, Island Press, Washington.

Hansen, J. E. (2007), Dangerous human-made interference with climate, Testimony to Select Committee on Energy Independence and Global Warming, United States House of Representatives, 26 April 2007 (Available at http://globalwarming.house.gov/list/hearing/global_warming_hearing_070423.shtml; updated version available at http://arxiv.org/abs/0706.3720).

Hansen, J. E., and M. Sato (2004), Greenhouse gas growth rates, Proc. Natl. Acad. Sci., 101, 16,109–16,114.

Hansen, J. E., et al. (2007a), Dangerous human-made interference with climate: A GISS modelE study, Atmos. Chem. Phys., 7, 2287–2312.

Hansen, J. E., M. Sato, P. Kharecha, G. Russell, D. W. Lea, and M. Siddall (2007b), Trace gases and climate change, Philos. Trans. R. Soc. Ser. A, 365, 1925–1954, doi:10.1098/rsta.2007.2052.

Hirsch, R. L., R. Bezdkek, and R. Wendling (2005), Peaking of world oil production: impacts, mitigation, and risk management, report to U.S. Dept. of Energy - Natl. Energy Techn. Lab. (Available at http://www.doe.gov/publications/others/pdf/Oil.Peaking.NETL.pdf).

Houghton, R. A. (2003), Revised estimates of the annual net flux of carbon to the atmosphere from changes in land use and land management 1850–2000, Tellus, Ser. A and Ser. B, 55B, 378–390.

Houghton, R. A., and J. L. Hackler (2002), Fluxes of carbon to the atmosphere from Land-Use Changes, in Trends: A Compendium of Data on Global Change, Carbon Dioxide Inf. Anal. Cent., Oak Ridge Natl. Lab., Dept. of Energy, U.S., Oak Ridge, Tenn.

House, K. Z., D. P. Schrag, C. F. Harvey, and K. S. Lackner (2006), Permanent carbon dioxide storage in deep-sea sediments, Proc. Natl. Acad. Sci., 103(33), 12291, doi:10.1073/pnas.0605318103.

Hubbert, M. K. (1956), Nuclear energy and the fossil fuels, Publication No. 95, 40 pp., Shell Development Company, Houston, Tex.

Intergovernmental Panel on Climate Change (IPCC) (2000), Special Report on Emissions Scenarios, edited by N. Nakicenovic and R. Swart, Cambridge Univ. Press, Cambridge, U.K.

Intergovernmental Panel on Climate Change (IPCC) (2001a), Climate Change 2001: Mitigation, edited by B. Metz et al., Cambridge Univ. Press, Cambridge, U.K.

Intergovernmental Panel on Climate Change (IPCC) (2001b), Climate Change 2001: The Scientific Basis, edited by J. T. Houghton et al., Cambridge Univ. Press, Cambridge, U.K.

Intergovernmental Panel on Climate Change (IPCC) (2007), Climate Change 2007: The Physical Science Basis, Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change, edited by S. Solomon et al., Cambridge Univ. Press, Cambridge, U.K.

Joos, F., M. Bruno, R. Fink, T. F. Stocker, U. Siegenthaler, C. Le Quere, and J. J. L. Sarmiento (1996), An efficient and accurate representation of complex oceanic and biospheric models of anthropogenic carbon uptake, Tellus, Ser. A and Ser. B, 48B, 397–417.

Keeling, C. D. and T. P. Whorf (2005), Atmospheric CO₂ records from sites in the SIO air sampling network, in Trends: A Compendium of Data on Global Change, Carbon Dioxide Inf. Anal. Cent., Oak Ridge Natl. Lab., Dept. of Energy, U.S., Oak Ridge, Tenn.

Keith, D. W., M. H. Duong, and J. K. Stolaroff (2006), Climate strategy with CO₂ capture from the air, Clim. Change, 74, 17–45.

Kerr, R. A. (2005), Bumpy road ahead for world’s oil, Science, 310, 1106–1108.

Kerr, R. A. (2007), The looming oil crisis could arrive uncomfortably soon, Science, 316, 351.

Lackner, K. (2003), A guide to CO₂ sequestration, Science, 300, 1677–1678.

Lam, M. (1998), Louisiana short term oil and gas forecast, report to Louisiana Dept. of Natural Resources - Technol. Assessment Div. (Available at http://dnr.louisiana.gov/sec/execdiv/techasmt/oil_gas/forecasts/shortterm_1998/03-production.htm).

Lovins, A. B. (2005), More profit with less carbon, Macmillan, London.

Matthews, H. D., and K. Caldeira (2007), Transient climate-carbon simulations of planetary geoengineering, Proc. Natl. Acad. Sci., 104, 9949–9954, doi:10.1073/pnas.0700419104.

McCarl, B. A., and R. D. Sands (2007), Competitive ness of terrestrial greenhouse gas offsets: Are they a bridge to the future?, Clim. Change, 80, 109–126.
Milici, R. C., and E. V. M. Campbell (1997), A predictive production rate life-cycle model for southwestern Virginia coalfields, *USGS Circular, 1147* (Available at http://pubs.usgs.gov/circ/c1147/).

National Research Council (NRC) (2007), *Coal: Research and Development to Support National Energy Policy*, Elsevier, Washington.

Robock, A., O. B. Toon, R. B. Turco, L. Oman, G. L. Stenchikov, and C. Bardeen (2007), The continuing threat of nuclear weapons: Integrated policy responses, *Eos Trans. AGU, 88*(21), 228, doi:10.1029/2007EO210012.

Shine, K. P., J. S. Fuglestvedt, K. Hailenariani, and N. Stuber (2005), Alternatives to the global warming potential for comparing climate impacts of emissions of greenhouse gases, *Clim. Change, 68*, 281–302.

Thoning, K. W., D. R. Kitzis, and A. Crotwell (2007), Atmospheric Carbon Dioxide Dry Air Mole Fractions from quasi-continuous measurements at Barrow, Alaska; Mauna Loa, Hawaii; American Samoa; and South Pole, 1973–2006, Version: 2007-10-01 (Available at ftp://ftp.cmdl.noaa.gov/ccg/co2/in-situ).

Tilman, D., J. Hill, and C. Lehman (2006), Carbon-negative biofuels from low-input high-diversity grassland biomass, *Science, 314*, 1598–1600.

United Nations Framework Convention on Climate Change (UNFCCC) (1992), (Available at http://unfccc.int/essential_background/convention/background/items/1349.php).

van der Veen, C. J. (2006), Reevaluating Hubbert’s prediction of U.S. peak oil, *Eos Trans. AGU, 87*(20), 199, doi:10.1029/2006EO200003.

Walter, K. M., S. A. Zimov, J. P. Chanton, D. Verbyla, and F. S. Chapin (2006), Methane bubbling from Siberian thaw lakes as a positive feedback to climate, *Nature, 443*, 71–75.

Weyant, J. P., F. C. de la Chesnaye, and G. J. Blanford (2006), Overview of EMF-21: Multigas mitigation and climate policy, *Energy J., Special Issue 3*, 1–32.

Wigley, T. M. (2006), A combined mitigation/geoengineering approach to climate stabilization, *Science, 314*, 452–454.

Wood, J. H., G. Long, and D. Morehouse (2003), World conventional oil supply expected to peak in 21st century, *Offshore, 63*, 90/92/94/150 (online version available at http://www.eia.doe.gov/pub/oil_gas/petroleum/feature_articles/2004/worldoilsupply/oilsupply04.html).

World Energy Council (2007), *Survey of Energy Resources*, 21st ed., edited by J. Trinnaman and A. Clarke (Available at http://www.worldenergy.org/publications/survey_of_energy_resources_2007/default.asp).

Zimov, S. A., E. A. G. Schuur, and F. S. Chapin (2006), Permafrost and the global carbon budget, *Science, 312*, 1612–1613.

J. E. Hansen and P. A. Kharecha, NASA Goddard Institute for Space Studies and Columbia University Earth Institute, 2880 Broadway, New York, NY 10025, USA. (pushker@giss.nasa.gov)