Future sea level change under CMIP5 and CMIP6 scenarios from the Greenland and Antarctic ice sheets

Article

Published Version

Creative Commons: Attribution 4.0 (CC-BY)

Open Access

Payne, Antony J., Nowicki, Sophie, Abe-Ouchi, Ayako, Agosta, Cecile, Alexander, Patrick, Albrecht, Torsten, Asay-Davis, Xylar, Aschwanden, Andy, Barthel, Alice, Bracegirdle, Thomas J., Calov, Reinhard, Chambers, Christopher, Choi, Youngmin, Cullather, Richard, Cuzzzone, Joshua, Dumas, Christophe, Edwards, Tamsin L., Felikson, Denis, Fettweis, Xavier, Galton-Fenzi, Benjamin K., Goelzer, Heiko, Gladstone, Rupert, Golledge, Nicholas R., Gregory, Jonathan M., Greve, Ralf, Hattermann, Tore, Hoffman, Matthew J., Humbert, Angelika, Huybrechts, Philippe, Jourdain, Nicolas C., Kleiner, Thomas, Munneke, Peter Kuipers, Larour, Eric, Le clec'h, Sebastien, Lee, Victoria, Leguy, Gunter, Lipscomb, William H., Little, Christopher M., Lowry, Daniel P., Morlighem, Mathieu, Nias, Isabel, Pattyn, Frank, Pelle, Tyler, Price, Stephen F., Quiquet, Aurelien, Reese, Ronja, Ruckamp, Martin, Schlegel, Nicole-Jeanne, Seroussi, Helene, Shepherd, Andrew, Simon, Erika, Slater, Donald, Smith, Robin S., Straneo, Fiammetta, Sun, Sainan, Tarasov, Lev, Trusel, Luke D., Van Breedam, Jonas, van de Wal, Roderik, van den Broeke, Michiel, Winkelmann, Ricarda, Zhao, Chen, Zhang, Tong and Zwinger, Thomas (2021) Future sea level change under CMIP5 and CMIP6 scenarios from the Greenland and Antarctic ice sheets.
Geophysical Research Letters

RESEARCH LETTER
10.1029/2020GL091741

Key Points:
- We compare results from an ice sheet model inter-comparison forced using Coupled Model Intercomparison Project phase 5 and phase 6 climate projections.
- Projected sea level at 2100 is higher under CMIP6 scenarios than CMIP5, but similar for Antarctica under both scenarios.
- CMIP6 warmer climate results in increased Greenland surface melt while increased snowfall mitigates loss from ocean warming for Antarctica.

Supporting Information:
Supporting Information may be found in the online version of this article.

Correspondence to:
A. J. Payne and S. Nowicki, a.j.payne@bristol.ac.uk; sophien@buffalo.edu

Citation:
Payne, A. J., Nowicki, S., Abe-Ouchi, A., Agosta, C., Alexander, P., Albrecht, T., et al. (2021). Future sea level change under coupled model intercomparison project phase 5 and phase 6 scenarios from the Greenland and Antarctic Ice Sheets. Geophysical Research Letters, 48, e2020GL091741. https://doi.org/10.1029/2020GL091741

Received 20 NOV 2020
Accepted 2 FEB 2021

© 2021, The Authors.
This is an open access article under the terms of the Creative Commons Attribution License, which permits use, distribution and reproduction in any medium, provided the original work is properly cited.

Future Sea Level Change Under Coupled Model Intercomparison Project Phase 5 and Phase 6 Scenarios From the Greenland and Antarctic Ice Sheets

Antony J. Payne1, Sophie Nowicki2, Ayako Abe-Ouchi4, Cécile Agosta5, Patrick Alexander6,7, Torsten Albrecht8, Xylar Asay-Davis9, Andy Aschwanden10, Alice Barthel10, Thomas J. Bracegirdle11, Reinhard Calov8, Christopher Chambers12, Youngmin Choi13, Richard Cullather2, Joshua Cuozzo14, Christophe Dumas5, Tamsin L. Edwards15, Denis Feliskin16, Xavier Fettweis7, Benjamin K. Galton-Fenzi18,19, Heiko Goelzer20,21,22, Rupert Gladstone23, Nicholas R. Golledge24, Jonathan M. Gregory25,26, Ralf Greve11,27, Tore Hattermann28,29, Matthew J. Hoffman10, Angelika Humbert30,31, Philippe Huybrechts32,33, Nicolas C. Jourdain34, Thomas Kleiner10, Peter Kuipers Munneke20,34, Eric Larour14, Sebastien Le clec'h32, Victoria Lee1, Gunter Leguy34, William H. Lipscomb34, Christopher M. Little15, Daniel P. Lowry36, Mathieu Morlighem13, Isabel Nias2,37, Frank Pattyn11, Tyler Pelle13, Stephen F. Price3, Aurélien Quietuk3, Ronja Reese3, Martin Rücker38, Nicole-Jeanne Schlegel14, Hélène Seroussi14, Andrew Shepherd13, Erika Simon2, Donald Slater39, Robin S. Smith32, Fiammetta Straneo37, Sainan Sun21, Lev Tarasov40, Luke D. Trusel41, Jonas Van Breemad32, Roderik van de Wal20,42, Michiel van den Broeke39, Ricarda Winkelmann8,43, Chen Zhao14, Tong Zhang9, and Thomas Zwinger44

1Centre for Polar Observation and Modelling, University of Bristol, Bristol, UK, 2Cryospheric Sciences Laboratory, NASA Goddard Space Flight Center, Greenbelt, MD, USA, 3Geology Department and RENEW Institute, University at Buffalo, Buffalo, NY, USA, 4Atmosphere and Ocean Research Institute, The University of Tokyo, Kashiwa-shi, Japan, 5Laboratoire des Sciences du Climat et de l’Environnement, LSCE-IPSL, CEA-CNRS-UVSQ, Université Paris-Saclay, Gif-sur-Yvette, France, 6Lamont-Doherty Earth Observatory, Columbia University, Palisades, NY, USA, 7NASA Goddard Institute for Space Studies, New York, NY, USA, 8Potsdam Institute for Climate Impact Research (PIK), Member of the Leibniz Association, Potsdam, Germany, 9Theoretical Division, Los Alamos National Laboratory, Los Alamos, NM, USA, 10Geophysical Institute, University of Alaska, Fairbanks, AK, USA, 11British Antarctic Survey, Cambridge, UK, 12Institute of Low Temperature Science, Hokkaido University, Sapporo, Japan, 13Department of Earth System Science, University of California Irvine, Irvine, USA, 14Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA, 15Department of Geography, King’s College London, London, UK, 16Universities Space Research Association, Goddard Earth Sciences Technology and Research Studies and Investigations, Columbia, MD, USA, 17Laboratory of Climatology, Department of Geography, University of Liège, Liège, Belgium, 18Australian Antarctic Division, Kingston, TAS, Australia, 19Australian Antarctic Program Partnership, Institute for Marine and Antarctic Studies, University of Tasmania, Hobart, TAS, Australia, 20Institute for Marine and Atmospheric research Utrecht, Utrecht University, Utrecht, The Netherlands, 21Laboratoire de Glaciologie, Université Libre de Bruxelles, Brussels, Belgium, 22NORCE Norwegian Research Centre, Bjørknes Centre for Climate Research, Bergen, Norway, 23Arctic Centre, University of Lapland, Finland, 24Antarctic Research Centre, Victoria University of Wellington, New Zealand, 25National Center for Atmospheric Science, University of Reading, Reading, UK, 26Met Office, Hadley Centre, Exeter, UK, 27Arctic Research Center, Hokkaido University, Sapporo, Japan, 28Norwegian Polar Institute, Tromso, Norway, 29Department of Physics and Technology, The Arctic University of Tromsø, Norway, 30Alfred Wegener Institute for Polar and Marine Research, Bremerhaven, Germany, 31Department of Geoscience, University of Bremen, Bremen, Germany, 32Earth System Science and Department Geographie, Vrije Universiteit Brussel, Brussels, Belgium, 33Univ. Grenoble Alpes/ CNRS/IRD/G-INP, Institut des Géosciences de l’Environnement, France, 34Climate and Global Dynamics Laboratory, National Center for Atmospheric Research, Boulder, CO, USA, 35Atmospheric and Environmental Research, Inc., Lexington, MA, USA, 36GNS Science, Lower Hutt, New Zealand, 37School of Environmental Sciences, University of Liverpool, Liverpool, UK, 38Centre for Polar Observation and Modelling, University of Leeds, UK, 39Scripps Institution of Oceanography, University of California San Diego, La Jolla, CA, USA, 40Department of Physics and Physical Oceanography, Memorial University of Newfoundland, Canada, 41Department of Geography, Pennsylvania State University, University Park, PA, USA, 42Geosciences, Physical Geography, Utrecht University, Utrecht, The Netherlands, 43Department of Physics and Astronomy, University of Potsdam, Potsdam, Germany, 44CSC-IT Center for Science, Espoo, Finland
Abstract  Projections of the sea level contribution from the Greenland and Antarctic ice sheets (GrIS and AIS) rely on atmospheric and oceanic drivers obtained from climate models. The Earth System Models participating in the Coupled Model Intercomparison Project phase 6 (CMIP6) generally project greater future warming compared with the previous Coupled Model Intercomparison Project phase 5 (CMIP5) effort. Here we use four CMIP6 models and a selection of CMIP5 models to force multiple ice sheet models as part of the Ice Sheet Model Intercomparison Project for CMIP6 (ISMIP6). We find that the projected sea level contribution at 2100 from the ice sheet model ensemble under the CMIP6 scenarios falls within the CMIP5 range for the Antarctic ice sheet but is significantly increased for Greenland. Warmer atmosphere in CMIP6 models results in higher Greenland mass loss due to surface melt. For Antarctica, CMIP6 forcing is similar to CMIP5 and mass gain from increased snowfall counteracts increased loss due to ocean warming.

Plain Language Summary  The melting of the Greenland and Antarctic ice sheets (GrIS and AIS) will result in higher sea level in the future. How sea level will change depends in part on how the atmosphere and ocean warm and how this affects the ice sheets. We use multiple ice sheet models to estimate possible future sea levels under climate scenarios from the models participating in the new Coupled Model Intercomparison Project phase 6 (CMIP6), which generally indicate a warmer world that the previous effort (CMIP5). Our results show that the possible future sea level change due Antarctica is similar for CMIP5 and CMIP6, but the warmer atmosphere in CMIP6 models leads to higher sea-level contributions from Greenland by the end of the century.

1. Introduction

The overall aim of this paper is to assess whether the stronger future warming shown by many Coupled Model Intercomparison Project phase 6 (CMIP6) models (Forster et al., 2019; Meehl et al., 2020) compared with Coupled Model Intercomparison Project phase 5 (CMIP5) has a significant impact on future Global Mean Sea Level Rise (GMSLR). We compare projections for the sea-level contribution of the Greenland and Antarctic ice sheets (GrIS and AIS) under climate forcing from a small group of models from the CMIP6 ensemble (Eyring et al., 2016) with that of models using forcing from the CMIP5 model ensemble (Taylor et al., 2012). Goelzer, Nowicki, et al. (2020) and Seroussi et al. (2020) present detailed analyses of the latter set of experiments for GrIS and AIS, respectively. In both cases, a great deal of attention was paid to sampling the CMIP5 ensemble effectively, so that the CMIP5 models used to provide climate forcing both represented the present-day climate of the ice sheets well and sampled the range of future projections of the overall ensemble. Details of this procedure can be found in Barthel et al. (2020).

Global warming as manifested in regional atmospheric and oceanic change can impact the ice sheet mass budget, and hence contribution to GMSLR, in a number of ways. Warming of the atmosphere over the ice sheet promotes increased melt from snow and ice surfaces leading to increased mass loss in the form of runoff to the oceans. It may also be associated with increased precipitation because of the increased moisture-carrying capacity of warmer air. The relationship between global warming and the warmth of polar ocean water masses impinging on the ice sheets is likely to be more complex. The warming of these water masses is expected to increase GMSLR by increasing mass loss from the marine-terminating outlet glaciers of the GrIS, and by processes associated with Marine Ice Sheet Instability (Schoof, 2007) for the AIS. An additional complexity for GrIS is that marine mass loss is partly controlled by freshwater fluxes from the surface melt (Slater et al., 2019). Finally, Marine Ice Sheet Instability could also be triggered by atmospheric warming leading to the fracture and collapse of floating ice shelves (Trusel et al., 2015). This process may in turn lead to the subsequent rapid retreat of the exposed marine ice cliffs (DeConto & Pollard, 2016). In summary, the range and complexity of the ways in which climate affects ice-sheet mass budget suggests that the greater global warming found in CMIP6 models may not necessarily lead to increased GMSLR.

2. The CMIP6 Ensemble

We compare a small ensemble of four Earth System Models (ESMs) submitted to the CMIP6 exercise. These models are UKESM1-0-LL, CESM2, CNRM-CM6-1, and CNRM-ESM2-1, which were the only ones available for downscaling at the time. Because the sample is small and based on availability only, it is important to
understand the difference between the selected models and the larger CMIP6 model ensemble. Effective Climate Sensitivity (ECS) (IPCC, 2013) is a convenient measure of this. ECS estimates the global mean temperature response to doubled atmospheric carbon dioxide concentration (Flato et al., 2013). The four selected models all have ECS at the upper end of the CMIP6 ensemble (CESM2, CNRM-CM6-1, CNRM-ESM2-1 and UKESM1-0-LL have ECS of 5.2°C, 4.8°C, 4.8°C, and 5.3°C, respectively). Roughly half of the CMIP6 ensemble has an ECS of between 4.6°C and 5.6°C, while there is a second similarly sized group with markedly lower ECS in the range 2.3°C–3.2°C (Meehl et al., 2020). In contrast, the CMIP5 ensemble exhibited a fairly continuous range of ECS between 2.1°C and 4.7°C (Flato et al., 2013). The CMIP5 models used in Goelzer, Nowicki, et al. (2020) and Seroussi et al. (2020) were typically drawn from the upper end of this distribution (e.g., MIROC-ESM, HadGEM2-ES, CSIRO-Mk3-6-0 and IPSL-CM5A-LR with ECS of 4.7°C, 4.6°C, 4.1°C, and 4.1°C, respectively) or lay close to the median (e.g., CCSM4, NorESM1-M and MIROC5 with ECS of 2.9°C, 2.8°C and 2.7°C, respectively).

Summaries of the atmospheric and ocean forcing for the two ice sheets are shown in Figures 1 and 2, respectively. Surface warming exhibited over the AIS in CMIP6 lies at or above the high end of the CMIP5 range. A similar pattern is evident in projected changes in Surface Mass Balance (SMB, the annual difference between mass addition, such as snowfall and refrozen rainfall, and mass loss, such as melt and subsequent runoff) over the ice sheet. Neither quantity is, however, significantly higher than the CMIP5 range. For

Figure 1. Atmospheric forcing used in Coupled Model Intercomparison Project phase 6 (CMIP6)-forced experiments. (a) and (b) mean annual surface air temperature and Surface Mass Balance (SMB) anomalies over Antarctic ice sheet (AIS). (c) and (d) mean annual surface air temperature and SMB anomaly for Greenland ice sheet (GrIS). Individual CMIP6 experiments are as shown as colored lines (legend in panel (d)). Gray shading reflects range of CMIP5 forcing encompassed by all of the CMIP5 experiments used by ISMIP6 (i.e., highest and lowest CMIP5 forcing for each year).
GrIS, SMB was derived by forcing the MAR regional climate model of Greenland (Fettweis et al., 2013) with CMIP6-derived boundary conditions. In this case, the CMIP6-forced SMB is significantly more negative (i.e., higher GMSLR rise) than is the case for CMIP5 forcing. Indeed, all four SSP585 ESMs fall outside the CMIP5 range and, by 2100, anomalies from UKESM1-0-LL and CESM2 approach twice that of largest CMIP5 ESM. The oceanic forcing of the AIS is described in detail by Jourdain et al. (2020) and for the GrIS by Slater et al. (2020). The thermal forcing derived from the CMIP6 models for both ice sheets lies within the range of the CMIP5 models with the exception of UKESM1-0-LL SSP585, which is occasionally higher. In many cases, the forcing lies towards the center of the CMIP5 range despite the higher ECS of the CMIP6 models. As would be expected thermal forcing from CNRM-CM6-1 SSP126 is less than that from CNRM-CM6-1 SSP585, however the difference is similar to the difference between the four SSP585 models.

3. Summary of ISMIP6 Experimental Procedure

The procedures used to convert the climate information summarized in Figures 1 and 2 into forcing imposed on ice sheet models are summarized in a series of papers for Antarctic ocean (Favier et al., 2019; Jourdain et al., 2020), Greenland ocean (Slater et al., 2019, 2020) and Greenland atmosphere (Fettweis et al., 2013).
Table 1
Overview of Experiments and Modeling Groups Participating in the CMIP6-Forced Exercise for AIS

| Group  | Model  | Open | Standard | Symbol |
|--------|--------|------|----------|--------|
| AWI    | PISM   | 1–5  | 1–5      | ⚖      |
| ILTS_PIK | SICOPOLIS | 1–5  |           | ⊹       |
| JPL    | ISSM   | 1–5  |          | ▲       |
| NCAR   | CISM   | 1–5  | 1–5      | △       |
| LSCE   | GRISLI | 1–5  | □        |        |
| UCIJPL | ISSM   | 1–5  | ▼        |        |
| VUB    | AISMPALEO | 1–3  | ♦        |        |

Note. Please refer to Seroussi et al. (2020) for model and group details. Symbols are those used in Figure 3.

Abbreviations: AIS, Antarctic ice sheet; CMIP6, Coupled Model Intercomparison Project phase 6.

Table 2
Overview of Experiments and Modeling Groups Participating in the CMIP6-Forced Exercise for GrIS

| Group  | Model  | Open | Standard | Symbol |
|--------|--------|------|----------|--------|
| AWI    | ISSM1  | 1–5  |          |       |
| AWI    | ISSM2  | 1–5  |          | ▼      |
| AWI    | ISSM3  | 1–5  |          | △      |
| BGC    | BISICLES | 1–3 |          |        |
| GSFC   | ISSM   | 1–2  | □        |        |
| ILTS_PIK | SICOPOLISI | 1–5 |          | ▲      |
| ILTS_PIK | SICOPOLIS2 | 1–5 |          | △      |
| IMAU   | IMAUICE2 | 1–3,5 |        | ♦      |
| JPL    | ISSM   | 1–5  | ⊹        |        |
| JPL    | ISSMPALEO | 1–3,5 |         | ▼      |
| LSCE   | GRISLI | 1–5  | △        |        |
| NCAR   | CISM   | 1–5  | ▲        |        |
| UAF    | PISM1  | 1–3,5 |          | △      |
| UAF    | PISM2  | 1–3,5 |          | △      |
| UCIJPL | ISSM1  | 1–3  | ♦        |        |
| VUB    | GISM   | 1–5  | +        |        |

Note. Please refer to Goelzer, Nowicki, et al. (2020) for model and group details. Symbols are those used in Figure 4. “f” refers to filled symbol. Abbreviations: CMIP5, Coupled Model Intercomparison Project phase 6; GrIS, Greenland ice sheet.

4. GMSLR Projections

Figure 3 shows projections for the AIS from the seven participating ice sheet models for each CMIP6-forced experiment along with ranges from the equivalent CMIP5-forced experiments (Seroussi et al., 2020). Figure 3b–3d compares these projections with ranges derived for the CMIP5 ensemble at 2100. The equivalent ranges for the whole AIS are −14–155 mm for RCP2.6, and −76–300 mm for RCP8.5. The regional contributions from West and East AIS are within or below the ranges reported for CMIP5 forcing. In many cases, they sit in the lower half of this range. This, however, is likely to reflect the high GMSLR associated with one ESM in CMIP5 ensemble of six (HadGEM2-ES), whose projected GMSLR was typically much higher (roughly twice that of the other ESMs for West AIS and positive rather than negative for East AIS). The projected GMSLR for all three AIS regions for CMIP6 and CMIP5 is very compatible if HadGEM2-ES is excluded from the latter.

Comparing projections for SSP126 (one ESM only) and SSP585 (four ESMs) suggests that there is little impact of emission scenario on projected GMSLR for AIS. This is, again, most likely to be related to the contrasting impacts for global warming on the ice sheet’s mass budget through increases in both mass loss by ice-sheet discharge and gain by snow accumulation.

The relationship between forcing and GMSLR for each CMIP6 ESM is complicated. For instance, ocean thermal forcing (Figure 2), air temperature anomalies (Figure 2) tend to be larger for UKESM1-0-LL; however, this is not reflected in their projected GMSLR. This is most likely to be associated with the compensatory effect of increased precipitation (Figure 2) in these ESMs.

Figure 4 shows projections for the GrIS from the 14 participating ice-sheet models for each CMIP6-forced experiment along with ranges from the equivalent CMIP5-forced experiments (Goelzer, Nowicki, et al., 2020). Projected GMSLR is either at the upper end of the CMIP5-forced range.
5. Discussion

We present the first comparison between CMIP5 and CMIP6-based projections of the contribution of ice sheets to future GMSLR up to 2100. This comparison is particularly interesting because many CMIP6 ESMs have higher climate sensitivity than their CMIP5 counterparts (Forster et al., 2019; Meehl et al., 2020) and their projections of future global warming are therefore higher. The comparison is hampered by the use of a relatively small ensemble of available CMIP6 ESMs, which are all at the upper end of CMIP6's range of climate sensitivity.

The comparison between CMIP5 and CMIP6 is markedly different for the two ice sheets, reflecting the very different ways in which the ice sheets are impacted by and respond to changes in the global climate system.

Figure 3. Global Mean Sea Level Rise (GMSLR) contribution from the Antarctic ice sheet (AIS) to 2100. (a) Time series of contribution between 2015 and 2100 (in mm) for whole ice sheet as a function of ice sheet model (symbol) and experiment (see legend). Contribution at 2100 for (b) West AIS, (c) East AIS and (d) Antarctic Peninsula. Symbols refer to ice sheet models and are given in Table 1. Filled symbols refer to "open" experiments and unfilled for "standard". Boxes in panels (b)–(d) refer to ranges from equivalent CMIP5-forced experiments (see Seroussi et al. (2020)).
For the GrIS, our results suggest that GMSLR contributions under CMIP6 are much higher than for CMIP5 perhaps by a factor of two. They also suggest a significant difference between SSP585 and SSP126, with the former experiencing accelerating rates of mass loss in marked contrast to the tendency towards stabilization of the latter.

Goelzer, Nowicki, et al. (2020) demonstrate that in excess of 80% of GrIS' contribution to GMSLR can be explained by changing SMB (primarily by surface melt and subsequent runoff), which is mostly controlled by atmospheric processes. The link between global warming and mass loss from the ice sheet is therefore fairly direct and a strong relationship between the two should be expected. The higher climate sensitivity of the sampled CMIP6 ESMs will therefore manifest itself as a larger GMSLR contribution in comparison to CMIP5. It should also be noted that for GrIS (in contrast to AIS), global warming is likely to favor increased mass loss by both atmospheric (i.e., SMB) and ocean forcing (i.e., discharge). However it appears that, at least within the ISMIP6 experimental design, ocean forcing plays a secondary role to the atmosphere.

For AIS, our results up to 2100 suggest little difference between CMIP6 and CMIP5-forced projections. This reflects the more complex interactions between this ice sheet and the global climate system. Global warming is likely to favor mass loss through changes in discharge resulting from increased ocean thermal forcing; however, the opposite is expected of the atmospheric forcing where warming is likely to favor mass gain (as a consequence of increased snow accumulation). The higher climate sensitivity of the sampled CMIP6 ESMs is therefore associated with both increased mass gain (snowfall) and mass loss (discharge) resulting in little net change in comparison to CMIP5 forcing. The complicated regional nature of interactions between ocean thermal forcing and AIS' discharge (e.g., Jenkins et al., 2018) is also likely to weaken any link between global warming and AIS mass loss.

The experimental design of the CMIP6-forced experiments reported here does not include the fracture and collapse of AIS' floating ice shelves resulting from meltwater ponding due to significant atmospheric warming (Trusel et al., 2015). This process has been cited as a necessary precursor to rapid ice loss by the retreat of marine ice cliffs (DeConto & Pollard, 2016). As ice shelf fracture was included in the CMIP5-forced experiments, an initial assessment can be made by comparing the amount of atmospheric warming projected to occur under CMIP5 and CMIP6. The higher climate sensitivity of the sampled CMIP6 ESMs is therefore associated with both increased mass gain (snowfall) and mass loss (discharge) resulting in little net change in comparison to CMIP5 forcing. The complicated regional nature of interactions between ocean thermal forcing and AIS' discharge (e.g., Jenkins et al., 2018) is also likely to weaken any link between global warming and AIS mass loss.

The experimental design of the CMIP6-forced experiments reported here does not include the fracture and collapse of AIS' floating ice shelves resulting from meltwater ponding due to significant atmospheric warming (Trusel et al., 2015). This process has been cited as a necessary precursor to rapid ice loss by the retreat of marine ice cliffs (DeConto & Pollard, 2016). As ice shelf fracture was included in the CMIP5-forced experiments, an initial assessment can be made by comparing the amount of atmospheric warming projected to occur under CMIP5 and CMIP6. Figure 2 suggests that CMIP6 ESMs lie close to or above the maximum CMIP5 surface temperature warming for AIS. For CMIP5 forcing, this process is limited to the Antarctic Peninsula and areas around George VI ice shelf and Totten glacier and its impact on GMSLR is ~28 mm (Seroussi et al., 2020). Ice-shelf fracture and associated processes may therefore be important under some CMIP6 forcing, in particular for CESM2 and UKESM1-0-LL, and likely be enhanced beyond 2100.
Data Availability Statement

The data analyzed in this study are available from Zenodo with digital object identifier https://doi.org/10.5281/zenodo.4498331.

References

Barthel, A., Agosta, C., Little, C. M., Hattermann, T., Jourdain, N. C., Goelzer, H., et al. (2020). CMIP5 model selection for ISMIP6 ice sheet model forcing: Greenland and Antarctica. *The Cryosphere*, 14, 855–879. https://doi.org/10.5194/tc-14-855-2020

DeConto, R. M., & Pollard, D. (2016). Contribution of Antarctica to past and future sea-level rise. Nature, 531(7596), 591–597. https://doi.org/10.1038/nature17145

Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., & Taylor, K. E. (2016). Overview of the coupled model intercomparison project phase 6 (cmip6) experimental design and organization. *Geoscientific Model Development*, 9(5), 1937–1958. https://doi.org/10.5194/gmd-9-1937-2016

Favier, L., Jourdain, N. C., Jenkins, A., Merino, N., Durand, G., Gagliardini, O., et al. (2019). Assessment of sub-shelf melting parameterizations using the ocean-ice sheet coupled model NEMO/v3.6-Elmer/Ice/v8.3. *Geoscientific Model Development*, 12, 2255–2283. https://doi.org/10.5194/gmd-12-2255-2019

Fettweis, X., Franco, B., Tedesco, M., van Angelen, J. H., Lenaerts, J. T. M., van den Broeke, M. R., & Gallée, H. (2013). Estimating the Greenland ice sheet surface mass balance contribution to future sea level rise using the regional atmospheric climate model mar. *The Cryosphere*, 7(2), 469–489. https://doi.org/10.5194/tc-7-469-2013

Flato, G., Marotzke, J., Aebi, D., Braconnot, P., Chou, S., Collins, W., & Rummukainen, M. (2013). Evaluation of climate models. In Stocker, T., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., & Midgley, P. M. (Eds.), *Climate change 2013: The physical science basis*. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change (pp. 741–866). Cambridge and New York, NY: Cambridge University Press.

Forster, P., Maycock, A., McKechnie, C., & Smith, C. (2019). Latest climate models confirm need for urgent mitigation. *Nature Climate Change*, 10, 7–10. https://doi.org/10.1038/s41558-019-0660-0

Goelzer, H., Noël, B. P. Y., Edwards, T. L., Fettweis, X., Gregory, J. M., Lipscomb, W. H., et al. (2020). Remapping of Greenland ice sheet surface mass balance anomalies for large ensemble sea-level change projections. *The Cryosphere*, 14(6), 1747–1762. https://doi.org/10.5194/tc-14-1747-2020

Goelzer, H., Nowicki, S., Payne, A., Larour, E., Seroussi, H., Lipscomb, W., & van den Broeke, M. (2020). The future sea-level contribution of the Greenland ice sheet: A multi-model ensemble study of ISMIP6. *The Cryosphere*, 14(9), 3071–3096. https://doi.org/10.5194/tc-14-3071-2020

IPCC (2013). Annex III. Glossary. In Stocker, T., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., & Midgley, P. M. (Eds.), *Climate change 2013: The physical science basis*. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change (p. 1447–1466). Cambridge and New York, NY: Cambridge University Press. https://doi.org/10.1017/CBO9781107415324.031

Jenkins, A., Shosmith, D., Dutrieux, P., Jacobs, S., Kim, T. W., Lee, S. H., et al. (2018). West Antarctic Ice sheet retreat in the Amundsen sea driven by decadal oceanic variability. *Nature Geoscience*, 11(10), 733–738. https://doi.org/10.1038/s41561-018-0207-4

Jourdain, N. C., Assay-Davis, X., Hattermann, T., Straneo, F., Seroussi, H., Little, C. M., & Nowicki, S. (2020). A protocol for calculating basal melt rates in the ISMIP6 Antarctic ice sheet projections. *The Cryosphere*, 14(9), 3111–3134. https://doi.org/10.5194/tc-14-3111-2020

Meehi, G. A., Senior, C. A., Eyring, V., Flato, G., Lamarque, J.-F., Stouffer, R. J., et al. (2020). Context for interpreting equilibrium climate sensitivity and transient climate response from the CMIP6 earth system models. *Science Advances*, 6(26), eaba1981. https://doi.org/10.1126/sciadv.aba1981

Nowicki, S., Goelzer, H., Seroussi, H., Payne, A. J., Lipscomb, W. H., Abe-Ouchi, A., et al. (2020). Experimental protocol for sea level projections from ISMIP6 stand-alone ice sheet models. *The Cryosphere*, 14(7), 2331–2368. https://doi.org/10.5194/tc-14-2331-2020

Nowicki, S. M. J., Payne, A., Larour, E., Seroussi, H., Goelzer, H., Lipscomb, W., et al. (2016). Ice sheet model intercomparison project (IS- MIP6) contribution to CMIP6. *Geoscientific Model Development*, 9(12), 4521–4545. https://doi.org/10.5194/gmd-9-4521-2016

Schoof, C. (2007). Ice sheet grounding line dynamics: Steady states, stability, and hysteresis. *Journal of Geophysical Research, 112*(F3), F03S28. https://doi.org/10.1029/2006JF000664

Seroussi, H., Nowicki, S., Payne, A. J., Goelzer, H., Lipscomb, W. H., Abe-Ouchi, A., & Zwinger, T. (2020). ISMIP6 Antarctica: A multi-model ensemble of the Antarctic ice sheet evolution over the 21st century. *The Cryosphere*, 14(9), 3033–3070. https://doi.org/10.5194/tc-14-3033-2020

Slater, D. A., Felikson, D., Straneo, F., Goelzer, H., Little, C. M., Morlighem, M., et al. (2020). Twenty-first century ocean forcing of the Greenland ice sheet for modeling of sea level contribution. *The Cryosphere*, 14(3), 985–1008. https://doi.org/10.5194/tc-14-985-2020

Slater, D. A., Straneo, F., Felikson, D., Little, C. M., Goelzer, H., Fettweis, X., & Holte. J. (2019). Estimating Greenland tidewater glacier retreat driven by submarine melting. *The Cryosphere*, 13(3), 2498–2509. https://doi.org/10.5194/tc-13-2498-2019

Taylor, K. E., Stouffer, R. J., & Meehl, G. A. (2012). An overview of CMIP5 and the experiment design. *Bulletin of the American Meteorological Society*, 93(4), 485–498. https://doi.org/10.1175/BAMS-D-11-00094.1

Trusel, L. D., Frey, K. E., Das, S. B., Karnauskas, K. B., Kuipers Munneke, P., van Meijgaard, E., & van den Broeke, M. R. (2015). Divergent trajectories of Antarctic surface melt under two twenty-first-century climate scenarios. *Nature Geoscience*, 8(12), 927–932. https://doi.org/10.1038/ngeo2563

Acknowledgments

The authors would like to thank the Climate and Cryosphere (C2ic) project for providing support for ISMIP6 through the sponsorship of workshops and hosting the ISMIP6 website. We acknowledge the University at Buffalo for their help with ISMIP6 data distribution, and the multiple agencies that support CMIP5, CMIP6, and the Earth System Grid Federation. This is ISMIP6 contribution number 12. This project received funding from the European Union’s Horizon 2020 research and innovation program under Grant agreement No 689304, PROTECT contribution number 4.