Ecosystem carbon balance in the Hawaiian Islands under different scenarios of future climate and land use change

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Abstract

The State of Hawai'i passed legislation to be carbon neutral by 2045, a goal that will partly depend on carbon sequestration by terrestrial ecosystems. However, there is considerable uncertainty surrounding the future direction and magnitude of the land carbon sink in the Hawaiian Islands. We used the Land Use and Carbon Scenario Simulator (LUCAS), a spatially explicit stochastic simulation model that integrates landscape change and carbon gain-loss, to assess how projected future changes in climate and land use will influence ecosystem carbon balance in the Hawaiian Islands under all combinations of two radiative forcing scenarios (RCPs 4.5 and 8.5) and two land use scenarios (low and high) over a 90 year timespan from 2010 to 2100. Collectively, terrestrial ecosystems of the Hawaiian Islands acted as a net carbon sink under low radiative forcing (RCP 4.5) for the entire 90 year simulation period, with low land use change further enhancing carbon sink strength. In contrast, Hawaiian terrestrial ecosystems transitioned from a net carbon sink under low radiative forcing (RCP 4.5) to the atmosphere under high radiative forcing (RCP 8.5), with high land use accelerating this transition and exacerbating net carbon loss. A sensitivity test of the CO\textsubscript{2} fertilization effect on plant productivity revealed it to be a major source of uncertainty in projections of ecosystem carbon balance, highlighting the need for greater mechanistic understanding of plant productivity responses to rising atmospheric CO\textsubscript{2}. Long-term model projections such as ours that incorporate the interactive effects of land use and climate change on regional ecosystem carbon balance will be critical to evaluating the potential of ecosystem-based climate mitigation strategies.

1. Introduction

Terrestrial ecosystems are a major sink for atmospheric carbon dioxide (CO\textsubscript{2}), removing ~30% of human emissions on an annual basis and reducing the rate of increase in atmospheric CO\textsubscript{2} (Keenan and Williams 2018, Friedlingstein et al \textit{et al} 2019). There is increasing recognition among policymakers that natural and agricultural ecosystems can contribute to climate mitigation, which has given rise to the popularity of ‘natural climate solutions’ (Cameron \textit{et al} 2017). Defined as conservation and land management efforts aimed at enhancing ecosystem carbon storage (Griscom \textit{et al} 2017), natural climate solutions are appealing because they are seen as cost-efficient and readily available (Cameron \textit{et al} 2017, Galarraga \textit{et al} 2017, Fargione \textit{et al} 2018). However, effective implementation is complicated by the uncertainty surrounding the future direction and magnitude of the land carbon sink, especially at the regional scale. Despite this uncertainty, evidence indicates that both interannual and long-term variability in carbon uptake by land ecosystems is driven primarily by fluctuations in climate, land use, and land cover.
change (Ahlström et al 2015, Prestele et al 2017, Friedlingstein et al 2019). Incorporating the interactive effects of land use and climate into spatially explicit future projections of ecosystem carbon balance could therefore provide a reference point to evaluate the effectiveness of land-based mitigation. Although a complex challenge, a growing number of sub-national jurisdictions plan to incorporate land-based mitigation strategies into their emissions reduction efforts. These jurisdictions would benefit from understanding how future land use and climate-biosphere feedbacks will affect ecosystem carbon balance in their respective regions (Sleeter et al 2019).

The State of Hawai‘i exemplifies the challenges associated with projecting the interactive effects of future climate and land use change on ecosystem carbon balance at a regional scale. Hawai‘i was the first U.S. state to enact legislation committing to full carbon neutrality, requiring the state to account for and offset all of its greenhouse gas emissions by 2045 (State of Hawai‘i Acts 15 and 16). This legislation emphasizes the mitigation potential of natural ecosystems as a key component to emissions reduction, necessitating baseline estimates and future projections of land carbon sink strength. However, Hawai‘i’s challenging terrain complicates these assessment efforts. The main Hawaiian Islands are a complex mosaic of natural and human-dominated landscapes overlain by steep climate gradients across relatively short distances, with mean annual temperature ranging from ~4 °C to 24 °C (Giambelluca et al 2014) and mean annual rainfall ranging from ~200 to 10 200 mm (Giambelluca et al 2013). Temperatures have risen rapidly in the Hawaiian Islands since the mid 1970s (Giambelluca et al 2008, McKenzie et al 2019) and a long-term drying trend has persisted since the early 1920s (Frazier and Giambelluca 2017), resulting in reduced forest biomass and productivity (Barbosa and Asner 2017). These same drying and warming trends have increased the frequency and intensity of wildland fire (Trauernicht et al 2015, Trauernicht 2019) with predictable negative effects on ecosystem carbon balance (Selmants et al 2017). Ecosystem carbon stocks across the main Hawaiian Islands have also been strongly influenced by the legacy of past land use change (Osher et al 2003, Asner et al 2011). Thousands of hectares were deforested beginning in the late 19th century to clear land for sugar plantations and cattle pasture (Cuddihy and Stone 1990). Since the mid-20th century, much of this agricultural land has been steadily converted to urban areas, commercial forestry plantations, or simply abandoned and colonized by non-native grass and shrub species (Suryanata 2009, Perroy et al 2016). Although these past trends surely inform the future impact of climate and land use change on ecosystem carbon balance, high spatial and temporal heterogeneity complicates realistic projection efforts. To date only one study has attempted to integrate land use, climate, and natural disturbances into future projections of Hawaiian ecosystem carbon balance, with projections limited to the mid-21st century under a single land use change scenario and a single moderate radiative forcing scenario (Special Report on Emission Scenarios (SRES) A1B, equivalent to Representative Concentration pathway (RCP) 6; Selmants et al 2017).

We used a stochastic, spatially explicit simulation model to estimate ecosystem carbon balance for Hawai‘i’s natural and agricultural lands on an annual basis for the period 2010–2100 under a range of assumptions about future climate, land use, land cover, disturbance, and global CO₂ emissions (Daniel et al 2016, 2018, Sleeter et al 2019). We developed four unique scenarios that explored different pathways, or future trajectories, of land use and climate change. These four scenarios represent all combinations of two land use change pathways (low and high) and two radiative forcing pathways (RCP 4.5 and RCP 8.5). In addition to these four scenarios, we conducted a separate series of simulations to examine how ecosystem carbon balance estimates vary in response to different levels of a CO₂ fertilization effect (CFE) on net primary productivity (NPP; Sleeter et al 2019). Our goals were to estimate statewide changes in Hawaiian ecosystem carbon balance and their uncertainties under a range of plausible future scenarios, quantify the relative impact of major controlling processes such as land use change, disturbance, and climate change, and assess the sensitivity of modeled ecosystem carbon balance estimates to varying levels of a CFE on NPP.

2. Methods

We used the Land Use and Carbon Scenario Simulator (LUCAS), an integrated landscape change and carbon gain-loss model, to project changes in ecosystem carbon balance for the seven main Hawaiian Islands. The landscape change portion of LUCAS is a state-and-transition model, where ‘state’ is the pre-defined information specific to each simulation cell (e.g. land cover class, climate zone) and ‘transition’ is the change in state of each simulation cell over time (Daniel et al 2016). Landscape change occurs by applying a Monte Carlo approach to track the state and age of each simulation cell in response to a pre-determined set of transition pathways (Daniel et al 2016). Transition pathways define which state types can be converted to any other state type. The carbon gain-loss portion tracks carbon stocks within each simulation cell over time as continuous state variables, along with a pre-defined set of continuous flows specifying rates of change in stock levels over time (Daniel et al 2018, Sleeter et al 2019). We parameterized the Hawai‘i LUCAS model to estimate annual changes in carbon stocks and fluxes in response to land use, land use change, wildland fire,
and long-term climate variability for the time period 2010–2100.

2.1. Study area
The spatial extent of this study was the terrestrial portion of the seven main Hawaiian Islands (figure 1), a total land area of 16,554 km$^2$. We subdivided this landscape into a grid of 264,870 simulation cells, each of which was 250 × 250 m in size. Each simulation cell was assigned to one of 210 possible state types based on the unique combination of seven islands, three moisture zones (dry, mesic, and wet; supplemental figure 1 (available online at stacks.iop.org/ERL/16/104020/mmedia)), and ten discrete land cover classes (figure 1). We also tracked the age and time since transition for each simulation cell as continuous state variables (Daniel et al. 2016).

2.2. States and transitions
We developed two land use scenarios (low and high) with transition pathways modified from Daniel et al. (2016). Transition pathways were pre-defined to represent urbanization, agricultural contraction, agricultural expansion, harvesting of tree plantations, and wildfire. Agriculture, forest, grassland, shrubland, and tree plantation land cover classes each had multiple transition pathways, i.e. they could each be converted into a variety of other land cover classes. The barren land cover class could only transition to developed, so its sole transition pathway was urbanization. Although most land cover classes had an urbanization transition pathway, there was no transition pathway out of an urbanized (developed) state. Water and wetland land cover classes remained static throughout the simulation period.

Transition targets set the area to be transitioned over time. These targets were based on either recent historical trends of land use change in the Hawaiian Islands from 1992 to 2011 (NOAA 2020) or on population projections for the State of Hawai‘i (Kim and Bai 2018). For the high land use scenario, rates of agricultural contraction, agricultural expansion, and urbanization for each timestep and Monte Carlo realization were sampled from uniform distributions bounded by the median and maximum recent historical rates for each island. For the low land use scenario, rates of agricultural contraction and expansion were sampled from uniform distributions bounded by zero and the minimum recent historical rates for each island, which shifts the balance toward agricultural contraction and leads to a reduction in land area under active cultivation over time. Urbanization rates in the low land use scenario were based on island-level population estimates and projections at five year intervals from 2010 to 2045 (Kim and Bai 2018).

We converted population projections into urbanization transition targets following Sleeter et al. (2017) by calculating population density for each island and then projecting future developed area based on the

Figure 1. Land cover classification of the seven main Hawaiian Islands. Agriculture in this map combines herbaceous and woody crops, but these two crop types are treated as separate land cover classes in the simulation model. Water and Wetland land cover classes are not shown. Reproduced from Jacobi et al. (2017). Credit: US Geological Survey.
five-year incremental change in island population. The spatial extent of agricultural contraction, agricultural expansion, and urbanization was constrained in both land use scenarios based on existing zoning maps (Daniel et al 2016). Transition targets for tree plantation harvest were set at ~75% of recent historical rates in the high land use scenario and ~40% of recent historical rates in the low land use scenario (Daniel et al 2016). Tree plantation forestry in the State of Hawai‘i is primarily short-rotation (5–7 years) Eucalyptus spp., harvested at a rate of approximately 2 km² yr⁻¹ over the past decade. In both the high and low land use scenarios, approximately 60% of tree plantation harvests were replacement harvests resulting in conversion to grassland or agriculture. The remaining 40% were rotation harvests replanted to Eucalyptus spp. There was no transition pathway from any land cover class into tree plantation in either land use scenario, which is consistent with recent historical trends of stable or declining tree plantation land area since the year 2000. For the contemporary period (2010–2020), transition targets in both the high and low land use scenarios were set at low land use scenario rates and diverged after 2020.

The wildfire transition sub-model was modified from Daniel et al (2016) by incorporating a new 21 year historical wildfire spatial database of the Hawaiian Islands (supplemental figure 2). We used this new spatial database to calculate historical wildfire size distribution and ignition probabilities for each unique combination of moisture zone (supplemental figure 1), island, and land cover class (figure 1) for the years 1999–2019. Starting in 2020, the number and size of fires was randomly drawn from one of these historical year-sets for each timestep and Monte Carlo realization, using burn severity probabilities from Selments et al (2017). Wildfire in the low land use scenario was sampled from the subset of historical fire years at or below the median area burned statewide from 1999–2019. The high land use scenario sampled from historical fire years above the median area burned over the same 21 year period (supplemental figure 2(a)). The vast majority of wildfire in Hawai‘i is human-caused and initiates in non-native grasslands and shrublands (Trauernicht et al 2015), but can spread into nearby forest areas. Moderate to high severity fires that spread into mesic and wet forests can convert these areas into grasslands about half the time. However, moderate to high severity fires in mesic and wet forests account for <5% of the total annual area burned, on average (Selments et al 2017, Daniel et al 2018).

2.3. Carbon stocks and flows
The fate of carbon stocks was tracked for each simulation cell as continuous state variables for each simulation cell, including live biomass, standing dead wood, down dead wood, litter, and soil organic carbon. We also included and tracked carbon in atmospheric, aquatic, and harvest product pools to enforce carbon mass balance (Daniel et al 2018). To transfer carbon between stocks, we defined baseline carbon flows as continuous variables resulting from growth, mortality, deadfall, woody decay, litter decomposition, and leaching (which includes runoff). We also defined carbon flows resulting from land use, land use change, and wildfire (Selments et al 2017, Daniel et al 2018).

Initial carbon stocks and baseline carbon flows were derived from the Integrated Biosphere Simulator (IBIS; Foley et al 1996, Liu et al 2020), a process-based dynamic global vegetation model. We initiated IBIS with minimal vegetation and simulated forward for 110 years using 30 year climate normals for the Hawaiian Islands (supplemental figure 3; Giambelluca et al 2013, 2014). We calibrated IBIS carbon stocks with statewide gridded datasets of soil organic carbon (Soil Survey Staff 2016) and total forest live biomass. The total forest live biomass data was derived from Asner et al (2016) estimates of aboveground forest live biomass by applying a power function to calculate the belowground portion (Mokany et al 2006) as described in Selments et al (2017). We calibrated gross photosynthesis in IBIS using a Hawai‘i-specific gridded dataset of gross primary production (GPP) derived from MODIS satellite imagery (Kimbball et al 2017). Mean annual GPP from Kimball et al (2017) ranged from 0.123 to 3.88 kg C m⁻² yr⁻¹ across the Hawaiian Islands, resulting in a NPP-to-GPP ratio ranging from 0.37 to 0.56.

We initiated NPP in LUCAS for each simulation cell as the mean IBIS-derived value for each combination of moisture zone and land cover class (supplemental table 1) adjusted with a set of spatially explicit stationary NPP multipliers to reflect local variation driven by microclimate (supplemental figure 5; Sleeter et al 2018, 2019). The stationary spatial NPP multipliers were the NPP anomaly for each simulation cell relative to the mean empirically-derived NPP values for each state type (supplemental figure 5). Soil carbon flux to the atmosphere (heterotrophic respiration, $R_h$) and aquatic soil carbon losses (leaching and overland flow) were estimated as the ratio of the IBIS-derived flux for each state type to the microclimate-adjusted NPP value for each simulation cell. We derived soil carbon loss as a fraction of NPP because photosynthesis is the dominant factor driving variation in ecosystem carbon fluxes, including $R_h$, and the vast majority of annual carbon loss from soils was fixed by photosynthesis within that year (Kuzyakov and Gavrichkova 2010, Baldocchi et al 2018). All other carbon flow rates were estimated as the ratio of the mean IBIS-derived flux.
for each state type to the size of the originating carbon stock at each age (Daniel et al. 2018, Sleeter et al. 2018). IBIS-derived carbon stocks for each state type were allowed to equilibrate with spatially variable NPP (supplemental figure 5) via a 100 year LUCAS spinup model run with no fire or land cover change and initiated with SSURGO soil carbon (Soil Survey Staff 2016) and IBIS values for live biomass, standing dead wood, down dead wood, and litter carbon stocks. Spatially variable carbon stocks from this spinup model run were used to initiate the main LUCAS model run.

Climate change impacts on carbon flows were represented by temporal growth and decay multipliers applied to each simulation cell based on mid-century (2049–2069) and end-of-century (2070–2099) statistically downscaled CMIP5 temperature and rainfall projections for the Hawaiian Islands under the RCP 4.5 and RCP 8.5 radiative forcing scenarios (supplemental figure 4; Elison Timm et al. 2015, Elison Timm 2017). Annual increments in temporal growth and decay multipliers were calculated by dividing mid-century or end-of-century estimates by the number of intervening years. Temporal growth multipliers used to represent the impact of future changes in rainfall and temperature on NPP were calculated using empirical NPP equations (Schuur et al. 2003, Del Grosso et al. 2008) and climate model projections of temperature and rainfall for each radiative forcing scenario (supplemental figure 6; Sleeter et al. 2019). Temporal decay multipliers used to represent the effect of future warming on turnover rates of dead organic matter were calculated using $Q_{10}$ temperature coefficients based on climate model temperature projections for each radiative forcing scenario (supplemental figure 4; Elison Timm 2017). The $Q_{10}$ temperature coefficients represented the proportional change in detrital carbon turnover due to a 10 °C change in mean annual temperature. We applied a $Q_{10}$ of 2.0 for wood and soil organic matter decay flows (Kurz et al. 2009, Sleeter et al. 2019) and a $Q_{10}$ of 2.17 for litter decay flows (Bothwell et al. 2014). Transition-triggered carbon flows resulting from disturbances associated with land use change, timber harvesting, and wildfire were based on values from Don et al. (2011), Selmants et al. (2017), and Daniel et al. (2018).

2.4. CO$_2$ fertilization effect

Increasing atmospheric CO$_2$ concentrations stimulate leaf-level photosynthesis, potentially increasing NPP as well (Franks et al. 2013). However, the magnitude and persistence of this effect is highly uncertain, particularly across a range of climatic conditions and over long time spans (Walker et al. 2020). Following Sleeter et al. (2019), we developed a separate set of scenarios designed to test the sensitivity of LUCAS model projections of ecosystem carbon balance to different rates of a CFE. We incorporated a CFE multiplier for NPP that represented the percent increase in NPP for every 100 ppm increase in atmospheric CO$_2$ concentration under the high land use and high radiative forcing (RCP 8.5) scenario. We tested five CFE levels ranging from 5% to 15%, which is within the range of CFEs observed in free air CO$_2$ enrichment (FACE) experiments. For all CFE levels, we assumed a saturation point at an atmospheric CO$_2$ concentration of 600 ppm, with no further stimulation of NPP despite a continued increase in CO$_2$ concentration to 930 ppm by 2100. This 600 ppm threshold generally coincides with the upper limit from FACE experiments and is reached by the year 2060 under RCP 8.5.

2.5. Scenario simulations and analysis

Each of the four unique scenarios were run for 90 years at an annual timestep and repeated for 30 Monte Carlo realizations, using initial conditions corresponding to the year 2010. All simulations were performed within the SyncroSim (version 2.2.4) software framework with ST-Sim (version 3.2.13) and SF (version 3.2.10) add-on modules (Daniel et al. 2016, 2018). Model input data and output summaries were prepared with the R statistical computing platform (R Core Team 2019) using the tidyverse (Wickham et al. 2019), raster (Hijmans 2020), and rsyncrosim (Daniel et al. 2020) packages. Carbon stocks and fluxes for the seven main Hawaiian Islands were calculated for each scenario by summing within each Monte Carlo realization on an annual basis and then calculating annual means as well as the annual upper and lower limits of the 30 Monte Carlo realizations. Carbon balance for the seven main Hawaiian Islands was calculated on annual basis for each scenario and Monte Carlo realization as net biome productivity (NBP), which was equal to annual carbon input in the form of NPP minus the annual sum of all carbon losses from terrestrial ecosystems, including heterotrophic respiration ($R_b$) from litter and soil, carbon fluxes to the atmosphere triggered by land use and land use change, wildfire emissions, and aquatic carbon losses through leaching and overland flow. Positive NBP values indicated ecosystems of the seven main Hawaiian Islands were acting as a net sink for atmospheric CO$_2$, while negative NBP values indicated that these ecosystems were acting as a net carbon source to the atmosphere (Chapin et al. 2006).

3. Results

3.1. Carbon stocks and fluxes

Terrestrial ecosystems of the seven main Hawaiian Islands stored an estimated 316 Tg of carbon at the beginning of the simulation period in 2010 (figure 2(a)), with 58% in soil organic matter, 22% in living biomass, and 20% in surface dead organic matter (litter and dead wood; figure 2(b)). Ecosystems accumulated carbon in all scenarios but at different rates, with trajectories shaped primarily by climate change and to a lesser extent by land use change.
Figure 2. Projected changes in total ecosystem carbon storage (a) and individual carbon stocks (b) for the seven main Hawaiian Islands. Solid lines indicate the mean of 30 Monte Carlo realizations for each scenario, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations.

The highest and most consistent projected accumulation of ecosystem carbon occurred under the combination of low radiative forcing and low land use change, yielding a ~15% increase in ecosystem carbon to an average of 363 Tg by 2100 (figure 2(a)). In contrast, high radiative forcing and high land use change resulted in the lowest ecosystem carbon gain, reaching a peak of ~332 Tg in 2063 and a decline to 327 Tg in 2100, resulting in a net increase of only 3% by the end of the simulation period (figure 2(a)). Ecosystem carbon accumulation was driven exclusively by increasing soil organic carbon across all four scenarios, all other stocks declined over time (figure 2(b)).

Net primary production (NPP) for the seven main Hawaiian Islands was ~8.1 Tg yr⁻¹ at the beginning of the simulation period in 2010 (figure 3) with land use change driving an approximate 2% decline during the contemporary period (2010–2020). NPP continued to decline throughout the rest of the simulation period (2020–2100) across all four scenarios, but this long-term decline was driven primarily by climate change and to a lesser extent by land use change (figure 3). The combination of high radiative forcing (RCP 8.5) and high land use change led to the steepest decline in NPP over time, driven by intense long-term drying on the leeward sides of islands under RCP 8.5 (supplemental figures 4 and 6) and sustained losses of forest and shrubland land area in the high land use scenario (supplemental figure 7). In contrast, intense warming under RCP 8.5 (supplemental figure 4) led to an increase in heterotrophic respiration (Rₘ) in the latter half of the 21st century under the low land use scenario (figure 3), and Rₘ was 3% higher on average by 2100 in the RCP 8.5 radiative forcing scenario than in the RCP 4.5 radiative forcing scenario. Heterotrophic respiration declined substantially over time in the high land use scenario (figure 3) because of long-term reductions in forest and shrubland land area (supplemental figure 7), similar to trends in NPP. Transition-triggered carbon fluxes to the atmosphere from land use, land use change, and wildfire were largely independent of changes in climate, stabilizing by mid-century at an average of ~0.4 Tg yr⁻¹ in the high land use scenario and ~0.2 Tg yr⁻¹ in the low land use scenario (figure 3). Uncertainty around transition-triggered carbon fluxes were higher in the high land use scenario, driven primarily by greater variability in wildland fire probabilities.

3.2. Ecosystem carbon balance

Net biome productivity (NBP) averaged approximately 0.6 Tg C yr⁻¹ at the start of the simulation period and declined over time in all four scenarios (figure 4). On average, terrestrial ecosystems of the seven main Hawaiian Islands collectively acted as a net carbon sink under the RCP 4.5 radiative forcing scenario throughout the simulation period, but carbon sink strength was ~40% lower in the high land use scenario compared to the low land use scenario by the end of the simulation period (figure 4). In contrast, ecosystems of the Hawaiian Islands acted as a net carbon source to the atmosphere toward the latter half of the simulation period under RCP 8.5, with the transition from sink to source occurring 15 years earlier on average in the high land use scenario than in the low land use scenario (figure 4). The high land use scenario under RCP 8.5 represented a ~40% larger net source of carbon to the atmosphere by the year 2100 than the low-land use scenario under the same radiative forcing. Over the entire simulation period, both global emissions reductions and local avoided land conversion resulted in substantial increases in cumulative NBP (figure 5). However, switching from RCP 8.5 to RCP 4.5 increased cumulative NBP in the Hawaiian Islands more than twice as much as reducing emissions from local land use change and wildfire disturbance (figure 5). Switching from RCP 8.5 to RCP 4.5 under the low land use scenario yielded the greatest cumulative increase in NBP, resulting in
Figure 3. Projected changes in net primary production (NPP), heterotrophic respiration ($R_h$) and carbon fluxes induced by land use and land use change for the seven main Hawaiian Islands. Solid lines indicate the mean of 30 Monte Carlo realizations for each scenario, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations.

A median gain of 26.5 Tg of carbon over the entire 90 year simulation period.

3.3. CO$_2$ fertilization effect
Projected estimates of both total ecosystem carbon storage and ecosystem carbon balance were highly sensitive to differing rates of a CFE on plant productivity. Under the high radiative forcing (RCP 8.5) and high land use scenario, the inclusion of a CFE ranging from 5% to 15% led to $\sim$33–98 Tg of additional carbon storage in ecosystems by the end of the century, a $\sim$10%–30% increase (figure 6(a)). Compared to the reference scenario (0% CFE), a 5% CFE was sufficient to transform Hawaiian terrestrial ecosystems from a net carbon source to the atmosphere during the latter half of the 21st century (figure 4(b)) to a net carbon sink for the entire simulation period (figure 6(b)), completely offsetting all other carbon losses induced by high radiative forcing and high land use. Net carbon sink strength was further enhanced at higher CFE rates, with NBP increasing by an average of 0.07 Tg C yr$^{-1}$ for each 1% increase in CFE (figure 6(b)). When compared to other scenarios, applying a 5% CFE to the high radiative forcing and high land use scenario resulted in a mean annual NBP of $0.46 \pm 0.3$ Tg C yr$^{-1}$, roughly equivalent to mean annual NBP in the low radiative forcing and low land use scenario with no CFE (0.52 ± 0.12). A 15% CFE applied to the high radiative forcing and high land use scenario resulted in a mean annual NBP of $1.18 \pm 0.29$ Tg C yr$^{-1}$, more than double that of the low radiative forcing and low land use scenario with no CFE.

4. Discussion
We estimated that terrestrial ecosystems of the Hawaiian Islands have been a consistent net sink for atmospheric carbon over the last decade (figure 4).
For the time period 2011–2019, NBP averaged 0.64 Tg C yr\(^{-1}\) and ranged from 0.46 to 0.88 Tg C yr\(^{-1}\) across all scenarios. Based on this mean annual NBP estimate, Hawaiian terrestrial ecosystems offset approximately 13% of 2015 statewide CO\(_2\) emissions from energy production and transportation (5.04 Tg C), the State of Hawai‘i’s largest source of greenhouse gas emissions (State of Hawai‘i 2019). Future projections indicate Hawaiian terrestrial ecosystems will continue to be a net sink for atmospheric carbon if global CO\(_2\) emissions peak around 2040 and then decline (RCP 4.5), and that carbon sink strength can be further enhanced by reducing the intensity and extent of future land use change. If, however, global CO\(_2\) emissions continue to rise throughout the 21st century (RCP 8.5), our projections indicate Hawaiian ecosystems will transition from a net sink to a net source of CO\(_2\) to the atmosphere, with high levels of land use change accelerating this transition and exacerbating net carbon loss. Our model results also indicate that projections of ecosystem carbon balance are highly sensitive to the introduction of a CFE. Even a 5% increase in NPP for every 100 ppm increase in atmospheric CO\(_2\) was sufficient to completely offset all other carbon losses induced by the high radiative forcing and high land use scenario, maintaining Hawaiian Island ecosystems as a net carbon sink for the entire simulation period instead of transitioning to a net carbon source by mid-century. Reconciling the high uncertainty surrounding the response of net photosynthesis to rising atmospheric CO\(_2\) is essential to more realistically constrain model projections of ecosystem carbon balance.

4.1. Impact of different climate and land use pathways
By comparing ecosystem carbon balance estimates under different scenario combinations, we were able to assess the relative impact of both global emissions reductions and regional actions to reduce emissions.
Figure 5. Projected changes in cumulative net biome productivity (NBP) for the seven main Hawaiian Islands when switching from the high to low land use change scenario under each radiative forcing scenario (top panel) and when switching from the high (RCP 8.5) to low (RCP 4.5) radiative forcing scenario under each land use scenario (bottom panel). Box plots indicate the median (vertical black line), 25th and 75th percentiles (colored boxes), 10th and 90th percentiles (thin horizontal lines), and values outside of this range (black circles). Note the different x-axis scales in each panel.

Figure 6. Sensitivity of projected changes in total ecosystem carbon storage (a) and mean annual net biome productivity (b) to different rates of a carbon dioxide fertilization effect (CFE) in the seven main Hawaiian Islands under the RCP 8.5 radiative forcing and high land use scenario. The CFE is the percent change in net primary productivity (NPP) for every 100 ppm increase in atmospheric carbon dioxide. Solid lines in (a) indicate mean total ecosystem carbon storage across 30 Monte Carlo realizations for each CFE rate, with shaded areas indicating the minimum and maximum range of Monte Carlo realizations. Solid circles in (b) represent mean annual net biome productivity averaged across all years and Monte Carlo realizations for each CFE rate, with vertical lines indicating the standard deviation of the mean. The dashed horizontal line in (b) represents the boundary between ecosystems acting as a net carbon sink (positive values) and a net carbon source (negative values).

from land use, land use change, and wildland fire (figure 5). Global adherence to a lower emissions trajectory (i.e. switching from RCP 8.5 to RCP 4.5) had the largest impact, resulting in a median cumulative increase of 26 Tg C sequestered by Hawaiian ecosystems over the 90 year simulation period. Long-term reductions in the intensity of land use change also consistently led to an increase in ecosystem carbon sequestration, but to a lesser degree than global emissions reductions. Switching from the high to the low land use scenario resulted in a median cumulative retention of an additional 11.6 Tg C in
Hawaiian ecosystems by 2100. The combination of global climate mitigation and local reductions in land use conversion had the largest potential benefit to ecosystem carbon sequestration, reducing cumulative net losses by over 400% (37.7 Tg C). Notably, the relative impact of reducing emissions from land use change was much greater under the high radiative forcing pathway (RCP 8.5). Cumulative NBP increased by 130% when switching from the high to low land use scenario under RCP 8.5, as opposed to a 37% cumulative increase in NBP when switching from high to low land use under RCP 4.5. These results demonstrate that reducing ecosystem carbon losses from land use change, harvest, and wildland fire can be an important component of greenhouse gas reduction efforts by sub-national jurisdictions like the State of Hawai’i, regardless of the global emissions trajectory. These results also highlight the utility of Hawai’i’s multi-pronged approach of participating in global climate mitigation efforts by reducing emissions from the energy and transportation sectors while also reducing land use emissions to minimize positive feedbacks to the climate system.

4.2. Comparison to other studies
There are few estimates of contemporary ecosystem carbon balance for the main Hawaiian Islands, and even fewer model projections of future ecosystem carbon balance in response to climate and land use change. Our mean annual NBP estimate of 0.64 Tg C yr⁻¹ for the period 2011–2020 agrees well with a recent State of Hawai’i Greenhouse Gas Inventory report, which estimated an annual net carbon sink of 0.66 Tg C in 2015 from agriculture, forestry, and other land uses (State of Hawai’i 2019). In contrast, our NBP estimate for the past decade was ~88% higher than a previous statewide LUCAS model estimate covering the same time period (0.341 Tg C yr⁻¹; Selmants et al 2017). This discrepancy was likely driven by modifications in how we calculated NPP, soil Rs, and soil aquatic carbon loss compared to previous versions of the LUCAS model, as well our model’s finer spatial resolution (Selmants et al 2017, Daniel et al 2018). Previous versions of a Hawai’i LUCAS model were run at 1 km spatial resolution and simulation cells within each unique combination of moisture zone and state type all had the same mean IBIS-derived NPP value applied to them at the beginning of the simulation period. In contrast, our NPP estimates at 250 m spatial resolution were adjusted on a cell-by-cell basis using Hawai’i-specific climate data (supplementary figure 3; Gambelluca et al 2013, 2014) as described in section 2.3. As a result, our statewide NPP estimates from 2011 to 2020 were 9.5% lower on average than previous LUCAS model estimates for Hawai’i during the same time period (Selmants et al 2017), likely because of the greater influence of more arid simulation cells. Soil carbon losses via Rs, leaching, and overland flow in previous versions of the LUCAS model were calculated as the ratio of the IBIS-derived flux to the size of the originating carbon stock, in this case soil organic carbon to 1 m depth (Daniel et al 2018). Here we calculated soil Rs and aquatic carbon losses as the ratio of the mean IBIS-derived flux to the microclimate-adjusted NPP value of each simulation cell, which is a more realistic driver than stock size (Kuzyakov and Gavrichkova 2010, Jackson et al 2017, Baldocchi et al 2018). Compared to previous Hawai’i LUCAS model estimates (Selmants et al 2017), soil Rs and aquatic carbon losses from 2011 to 2020 were reduced by an average of 15% and 21%, respectively, which widened the gap between carbon gain (NPP) and carbon losses and accounted for the overall increase in NBP estimates for this time period.

4.3. Limitations of this study
There is ample evidence that increasing atmospheric CO₂ concentrations can stimulate NPP (Norby et al 2005, Zhu et al 2016), but the magnitude and persistence of this effect remains highly uncertain (Frankis et al 2013, Walker et al 2020). Our results demonstrate that long-term projections of ecosystem carbon balance are highly sensitive to uncertainty in CFE strength. With no CFE, Hawaiian ecosystems became a net source of CO₂ to the atmosphere beginning in the latter half of the 21st century under the high land use and high radiative forcing scenario. However, a CFE equivalent to a 5% increase in NPP for every 100 ppm increase in atmospheric CO₂ applied to the same scenario resulted in Hawaiian ecosystems remaining a net carbon sink throughout the entire simulation period. A 15% CFE applied to the high land use and high radiative forcing scenario resulted in a nearly 5-fold increase in mean annual NBP averaged across all years and Monte Carlo realizations. Despite this demonstrated sensitivity to a CFE, several potentially attenuating factors complicate the selection of a realistic CFE value with any degree of confidence (Walker et al 2020). Nitrogen and phosphorus limitation can reduce or eliminate a CFE (Reich et al 2006, Norby et al 2010, He et al 2017, Terrer et al 2019), as can water limitation and heat stress (Obermeier et al 2017, Birami et al 2020). Forest age may also be a factor, with young aggrading forests showing a strong positive growth response to CO₂ fertilization (Walker et al 2019), while old-growth forests show little to no response (Jiang et al 2020, Yang et al 2020). This evidence indicates that a CFE may be highly variable across space and time, suggesting it may be unrealistic to apply a single CFE rate value across an entire region over several decades. Until mechanistic understanding is improved, the most conservative approach when projecting future ecosystem balance in the context of climate mitigation planning may be to assume no CFE, with the knowledge that any realized CFE will only enhance ecosystem carbon sequestration.
Our model does not currently differentiate between forests dominated by native versus non-native tree species, which could influence estimates of ecosystem carbon balance. Native forests in the Hawaiian Islands are dominated by ‘ohi’a (Metrosideros polymorpha Gauditch), an endemic foundational tree species found across a broad range of climatic and edaphic conditions (Ziegler 2002). Beginning in 2010, a fungal disease termed rapid ‘Ohi’a death (ROD) caused by Ceratocystis spp. has emerged and caused widespread mortality to mature ‘ohi’a trees across a range of size classes, primarily on Hawai‘i Island (Mortenson et al 2016, Fortini et al 2019). The distribution and potential range of this emerging threat to Hawai‘i’s dominant native tree species have only recently been mapped (Vaughn et al 2018, Fortini et al 2019), but the pulse of newly dead organic matter and reduction in photosynthetic capacity induced by widespread tree mortality could significantly alter ecosystem carbon balance over the long-term (Sleeter et al 2019). ROD-affected forests could also undergo a replacement of canopy dominant stress-tolerating ‘ohi’a trees by non-native tree species with faster relative growth rates, lower wood density, and faster tissue turnover, potentially altering the long-term trajectory of carbon cycling in Hawaiian forest ecosystems. New model projections for Hawai‘i that incorporate ROD spread rates and forest restoration scenarios will therefore require differentiation among forest ecosystems dominated by native and non-native tree species.

Interannual climate variability is a primary factor influencing spatial and temporal patterns of global wildland fire activity (Abatzoglou et al 2018), with climate warming expected to increase wildland fire frequency and wildfire season length across a wide range of biomes (Sun et al 2019). Although our model projections capture the spatial and temporal variation in ignition probability and fire extent by sampling from previous fire years (1999–2019), we did not incorporate the effects of projected future climate change on the frequency and extent of wildland fire. Recent fire probability modeling for the northwest portion of Hawai‘i Island indicated that projected drying and warming trends under RCP 8.5 could increase maximum fire probability values more than three-fold and shift areas of peak flammability to higher elevation by mid-century (Trauernicht 2019). Extending this probability fire modeling approach statewide would provide a quantitative, spatially explicit assessment of wildland fire probability for the main Hawaiian Islands as predicted by climate, land cover, and ignition density, which is highly correlated with population density (Trauernicht et al 2015, Trauernicht 2019). This approach would provide future simulation model projections of Hawaiian ecosystem carbon balance with more realistic scenarios of expected annual area burned based on the integrated effects of future climate and land use change.

5. Conclusion

Although terrestrial ecosystems are currently an important sink for atmospheric CO$_2$, the future direction and magnitude of the land carbon sink are highly uncertain, especially at regional scales. Our simulation modeling results indicated that projected climate change, dictated by long-term trajectories in global greenhouse gas emissions, was the primary factor influencing terrestrial ecosystem carbon balance in the Hawaiian Islands. Long-term reductions in the intensity of land use change and wildland fire also consistently led to an increase in ecosystem carbon sequestration, but to a lesser degree than global emissions reductions. CO$_2$ fertilization of NPP was the largest source of uncertainty in long-term projections of ecosystem carbon balance in the Hawaiian Islands, highlighting the need for greater mechanistic understanding of the cascading effects of rising atmospheric CO$_2$ on ecosystem carbon sequestration. By incorporating the interactive effects of land use and climate change into future projections of ecosystem carbon balance, our model results could be used as a set of baseline projections for the State of Hawai‘i to evaluate different ecosystem-based climate mitigation strategies. Studies like ours that incorporate stochasticity into spatially explicit simulation models could also provide a framework for the growing number of sub-national jurisdictions that plan to incorporate ecosystem carbon sequestration into their emissions reduction efforts. These long-term projections will be critical to assessing the impact of future land use change and climate-biosphere feedbacks on meeting climate mitigation goals.

Data availability statement

The data that support the findings of this study are openly available at the following URL/DOI: https://doi.org/10.5066/P9AWLFKZ.

LUCAS model input data and R code used to format input data, summarize output data and compile this manuscript are available from a GitHub repository (https://github.com/selmants/HI_Model) archived at https://doi.org/10.5281/zenodo.5198072.

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References

Abatzoglou J T, Williams A P, Boschetti L, Zubikova M and Kolden C A 2018 Global patterns of interannual climate–fire relationships Glob. Change Biol. 24 5164–75
Ahlström A et al 2015 The dominant role of semi-arid ecosystems in the trend and variability of the land CO2 sink Science 348 895–9
Asner G P, Hughes R F, Mascaro J, Uowolo A L, Knapp D E, Jacobson J, Kennedy-Boydowin T and Clark J K 2011 High-resolution carbon mapping on the million-hectare Island of Hawaii Front. Ecol. Environ. 9 434–9
Asner G P, Soussan S, Knapp D E, Selmants P C, Martin R E, Hughes R P and Giardina C P 2016 Rapid forest carbon assessments of oceanic islands: a case study of the Hawaiian archipelago Carbon Balance Manage. 11 1
Baldocchi D, Chu H and Reichstein M 2018 Inter-annual variability of net and gross ecosystem carbon fluxes: a review Agric. For. Meteorol. 249 520–33
Barboza J M and Asner G P 2017 Effects of long-term rainfall decline on the structure and functioning of Hawaiian forests Environ. Res. Lett. 12 094002
Birami B, Nägele T, Gattmann M, Preisler Y, Gast A, Arneth A and Barbosa J M and Asner G P 2016 State-and-transition simulation models: a framework for forecasting landscape change Methods Ecol. Evol. 7 1413–23
Daniel C J, Sleeter B M, Frid L and Fortin M-J 2018 Integrating continuous stocks and flows into state-and-transition simulation models of landscape change Methods Ecol. Evol. 9 1133–43
Daniel G, Hughes J, Embrey A, Frid L and Lucet V 2020 rsyncrosim: the R interface to SyncroSim (available at: https://github.com/rsyncrosim/rsyncrosim)
Del Grosso S, Parton W, Stohlgren T, Zheng D, Bachelet D, Prince S, Hibbard K and Olson R 2008 Global potential net primary production predicted from vegetation class, precipitation and temperature Ecology 89 2117–26
Don A, Schumacher J and Freibauer A 2011 Impact of tropical land-use change on soil organic carbon stocks—a meta-analysis Glob. Change Biol. 17 1658–70
Elison Timm O 2017 Future warming rates over the Hawaiian Islands based on elevation-dependent scaling factors Int. J. Climatol. 37 1093–104
Elison Timm O, Giambelluca T W and Diaz H F 2015 Statistical downscaling of rainfall changes in Hawaii based on the CMIP5 global model projections J. Geophys. Res.: Atmos. 120 92–112
Fargione J et al 2018 Natural climate solutions for the United States Sci. Adv. 4 eaat1869
Foley J A, Prentice I C, Ramankutty N, Levis S, Pollard D, Sitch S and Haxeltine A 1999 An integrated biosphere model of land surface processes, terrestrial carbon balance and vegetation dynamics Glob. Biogeochem. Cycles 10 603–28
Fortini I, B, Kaiser L R, Keith L M, Price J, Hughes R F, Jacobi J D and Friday J B 2019 The evolving threat of Rapid ‘ohia Death (ROD) to Hawaii’s native ecosystems and rare plant species For. Ecol. Manage. 448 376–85
Franks P J et al 2013 Sensitivity of plants to changing atmospheric CO2 concentration: from the geological past to the next century New Phytol. 197 1077–94
Frazier A G and Giambelluca T W 2017 Spatial trend analysis of Hawaiian rainfall from 1920 to 2012 Int. J. Climatol. 37 2523–31
Friedlingstein P et al 2019 Global carbon budget 2019 Earth Syst. Sci. Data 11 783–838
Galaraga I, de Murrieta E S and Franca J 2017 Climate policy at the sub-national level Trends in Climate Change Legislation ed A Averchenkova, S Fankhauser and M Nachmany (Northampton, MA: Edward Elgar Publishing) pp 143–73
Giambelluca T W, Chen Q, Frazier A G, Price J P, Chen Y-L, Chu P-S, Eisheild J K and Delporte D M 2013 Online rainfall atlas of Hawaii’s Bull. Am. Meteorol. Soc. 94 313–6
Giambelluca T W, Diaz H F and Luke M S A 2008 Secular temperature changes in Hawaii’s Geophys. Res. Lett. 35 L12702
Giambelluca T W, Shuai X, Barnes M L, Alliss R J, Longman R J, Miura T, Chen Q, Frazier A G, Mudd R G, Cao L and Businger A D 2014 Evapotranspiration of Hawaii’s Technical Report
Griscom B W et al 2017 Natural climate solutions Proc. Natl. Acad. Sci. 114 11645–50
He L, Chen J M, Croft H, Gonsamo A, Luo X, Liu J, Zheng T, Liu R and Liu Y 2017 Nitrogen availability dampens the positive impacts of CO2 fertilization on terrestrial ecosystem carbon and water cycles Geophys. Res. Lett. 44 11590–600
Hijmans R J 2020 Raster: geographic data analysis and modeling (available at: https://CRAN.R-project.org/package=raster)
Jackson R B, Lajtha K, Crow S E, Hugelius G, Kramer M G and Friday J B 2019 The fate of carbon in a mature forest under carbon dioxide enrichment Nature 580 227–31
Keenan T and Williams C 2018 The terrestrial carbon sink Proc. Natl. Acad. Sci. 115 2197–224
Kolden C A 2018 Global patterns of interannual climate–fire relationships Proc. Natl Acad. Sci. 114 12833–8
Chapin F S et al 2006 Reconciling carbon-cycle concepts, terminology and methods Ecosystems 9 1041–50
Cuddihy L W and Stone C P 1990 Alteration of Hawaiian Vegetation: Effects of Humans, Their Activities and Introductions (Honolulu, HI: University of Hawaii Press)
Daniel C J, Sleeter B M and Fortin M-J 2016 State-and-transition simulation models: a framework for forecasting landscape change Methods Ecol. Evol. 7 1413–23
Daniel C J, Sleeter B M, Frid L and Fortin M-J 2018 Integrating continuous stocks and flows into state-and-transition simulation models of landscape change Methods Ecol. Evol. 9 1133–43
