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Authors
Kendall, Brian
Creaser, Robert A
Reinhard, Christopher T
et al.

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Transient episodes of mild environmental oxygenation and oxidative continental weathering during the late Archean

Brian Kendall,1* Robert A. Creaser,2 Christopher T. Reinhard,3 Timothy W. Lyons,4 Ariel D. Anbar5,6

IT IS NOT KNOWN WHETHER ENVIRONMENTAL O2 LEVELS INCREASED IN A LINEAR FASHION OR FLUCTUATED DYNAMICALLY BETWEEN THE EVOLUTION OF OXYGENIC PHOTOSYNTHESIS AND THE LATER GREAT OXIDATION EVENT. NEW RHENIUM–OSMIUM ISOTOPE DATA FROM THE LATE ARCHEAN MOUNT MCRAE SHALE, WESTERN AUSTRALIA, REVEAL A TRANSIENT EPISODE OF OXIDATIVE CONTINENTAL WEATHERING MORE THAN 70 MILLION YEARS BEFORE THE ONSET OF THE GREAT OXIDATION EVENT. A DEPOSITIONAL AGE OF 2495 ± 14 MILLION YEARS AND AN INITIAL 187Os/188Os OF 0.34 ± 0.19 WERE OBTAINED FOR RHENIUM- AND MOLOYDENUM-RICH BLACK SHALES. THE INITIAL 187Os/188Os IS HIGHER THAN THE MANTLE/EXTRATERRESTRIAL VALUE OF 0.11, POINTING TO MILD ENVIRONMENTAL OXYGENATION AND OXIDATIVE MOBILIZATION OF RHENIUM, MOLOYDENUM, AND RADIogenic OSMIUM FROM THE UPPER CONTINENTAL CRUST AND TO CONTEMPORANEOUS TRANSPORT OF THESE METALS TO SEAWATER. BY CONTRAST, STRATIGRAPHICALLY OVERLAYERING BLACK SHALES ARE RHENIUM- AND MOLOYDENUM-POOR AND HAVE A MANTLE-LIKE INITIAL 187Os/188Os OF 0.06 ± 0.09, INDICATING A REDUCED CONTINENTAL FLUX OF RHENIUM, MOLOYDENUM, AND OSMIUM TO SEAWATER BECAUSE OF A DROP IN ENVIRONMENTAL O2 LEVELS. TRANSIENT OXYGENATION EVENTS, LIKE THE ONE CAPTURED BY THE MOUNT MCRAE SHALE, PROBABLY SEPARATED INTERVALS OF LESS OXYGENATED CONDITIONS DURING THE LATE ARCHEAN.

INTRODUCTION

MULTIPLE LINES OF GEOCHEMICAL EVIDENCE FROM SEDIMENTARY ROCKS POINT TO THE PRODUCTION AND ACCUMULATION OF PHOTOSYNTHETIC O2 IN SURFACE ENVIRONMENTS SINCE AT LEAST 3 BILLION YEARS AGO (Ga) (1–7). IN CONTRAST, SULFUR MASS-INDEPENDENT FRACTIONATION (S-MIF) INDICATES THAT ARCHEAN ATMOSPHERIC O2 LEVELS WERE PROMINENTLY BELOW 0.001% OF PRESENT AT-FUR MASS-INDEPENDENT FRACTIONATION (S-MIF) INDICATES THAT ARCHEAN AND AN INITIAL 187Os/188Os OF 0.34 ± 0.19 WERE OBTAINED FOR RHENIUM- AND MOLOYDENUM-RICH BLACK SHALES. THE INITIAL 187Os/188Os IS HIGHER THAN THE MANTLE/EXTRATERRESTRIAL VALUE OF 0.11, POINTING TO MILD ENVIRONMENTAL OXYGENATION AND OXIDATIVE MOBILIZATION OF RHENIUM, MOLOYDENUM, AND RADIogenic OSMIUM FROM THE UPPER CONTINENTAL CRUST AND TO CONTEMPORANEOUS TRANSPORT OF THESE METALS TO SEAWATER. BY CONTRAST, STRATIGRAPHICALLY OVERLAYERING BLACK SHALES ARE RHENIUM- AND MOLOYDENUM-POOR AND HAVE A MANTLE-LIKE INITIAL 187Os/188Os OF 0.06 ± 0.09, INDICATING A REDUCED CONTINENTAL FLUX OF RHENIUM, MOLOYDENUM, AND OSMIUM TO SEAWATER BECAUSE OF A DROP IN ENVIRONMENTAL O2 LEVELS. TRANSIENT OXYGENATION EVENTS, LIKE THE ONE CAPTURED BY THE MOUNT MCRAE SHALE, PROBABLY SEPARATED INTERVALS OF LESS OXYGENATED CONDITIONS DURING THE LATE ARCHEAN.

The 2.5-Ga Mount McRae Shale (Hamersley Basin, Western Australia) highlights the difficulties inherent in identifying changes in ancient environmental O2 levels. Black shales from the upper Mount McRae Shale record an episode of high Mo and Re concentrations (Fig. 1) that suggest oxidative mobilization of these metals from the upper continental crust, followed by their transport as oxyanions to seawater (1). Increased Mo and Re concentrations may capture a transient increase in oxidative continental weathering, possibly associated with a rise in atmospheric O2 levels, and the subsequent decrease in Mo and Re concentrations may reflect a return to a lower redox state (1). However, sedimentary Fe speciation shows that the Mo and Re concentrations rise in phase with a shift in local bottom water redox conditions from anoxic and Fe2+-rich (ferruginous) to anoxic and H2S-rich (euxinic) (15). Because the magnitude of Mo and Re enrichment in marine sediments is influenced by local bottom water redox conditions, organic carbon export fluxes, and sedimentation rates (16, 17), the elevated Mo and Re concentrations may reflect changes in local depositional conditions against a backdrop of mild environmental O2 levels (1). Both scenarios are consistent with stratigraphic trends in S, N, Mo, U, and Se isotope data in the Mount McRae Shale (1, 2, 15, 18–21). Distinguishing between these two possibilities is critical in understanding the controls on possible O2 production and accumulation before the GOE.

To determine whether the chemostratigraphic trends in the Mount McRae Shale capture continuous or transient oxidative continental weathering, and to directly test recent challenges to the geochemical evidence for pre-GOE oxygenesis (22), we turned to the Re-Os isotope system. An isochron regression of Re-Os data from temporally related black shales yields a depositional age, and the initial 187Os/188Os records the local seawater 187Os/188Os ratio at the time of deposition (23, 24). Seawater 187Os/188Os, in turn, reflects a time-varying balance between radiogenic and unradiogenic Os fluxes to the ocean. Modern seawater has a highly radiogenic 187Os/188Os of ca. 1.06 (25, 26) due to riverine runoff of radiogenic Os derived from oxidative continental weathering (~1.5) (27). A subordinate flux of unradiogenic Os originates from dissolution of cosmic dust and hydrothermal alteration of ultramafic oceanic rocks (~0.13) (28–30). On continental margins, organic-rich sediments are an Os sink and directly record local seawater 187Os/188Os (31). Hence, the initial 187Os/188Os from ancient black
shales can be used to infer temporal changes in the Os fluxes to seawater from mantle/extraterrestrial sources versus those tied to oxidative continental weathering.

RESULTS

We measured Re and Os isotope data (table S1) for eight Re- and Mo-rich black shale samples of the Mount McRae Shale from 148.09- to 148.15-m depth in the ABDP-9 core (Archean Biosphere Drilling Project core #9), which previously yielded geochemical evidence for oxidative continental weathering and possible surface ocean oxygenation (1, 2, 15, 18–21). The analytical methods are described in Materials and Methods. Regression of the eight analyses (plus four replicate analyses) together with four previous Re-Os analyses from 145.22, 146.08, 147.10, and 148.32 m in ABDP-9 (1) yields a 16-point Re-Os age of 2495 ± 14 million years ago (Ma) (the 2σ age uncertainty includes a 0.31% uncertainty on the 187Re decay constant) and an initial 187Os/188Os of 0.34 ± 0.19 (Fig. 2).

DISCUSSION

Osmium isotope evidence for oxidative continental weathering at 2.5 Ga

Recently, a sophisticated set of petrographic analyses was used to suggest that the purported geochemical evidence for oxidative continental weathering recorded by the Mount McRae Shale is instead entirely a product of metasomatic overprints (22). However, the new Re-Os black shale age of 2495 ± 14 Ma for 145.22 to 148.32 m is statistically identical to a Re-Os black shale age of 2495 ± 20 Ma from 128.71 to 129.85 m in the ABDP-9 core (1) and to an U-Pb zircon age of 2504 ± 5 Ma from a tuffaceous bed within the Mount McRae Shale (32). The agreement between the Re-Os and U-Pb ages is an important observation, given the possibility that sedimentary strata of the Pilbara Craton have been chemically altered by post depositional hydrothermal fluid flow (33). Our new data confirm that metasomatic processes have not noticeably perturbed the Re-Os system in the Mount McRae Shale subsequent to deposition at ~2.5 Ga.

Given that Mo is a redox-sensitive, siderophilic, and chalcophilic trace metal that should be characterized by postdepositional chemical mobility and overprints similar to those of Re and Os, it is likely that Mo has also escaped significant metasomatic alteration. In any case, Re enrichment alone provides strong evidence for oxidative continental weathering, and the likely preservation of its primary (depositional) signal is directly linked to the demonstrably unperturbed Re-Os system. A primary depositional signal is also supported by strong correlations between total organic carbon and both Mo and Re concentrations, an expected signature of oxidative mobilization and transport of soluble MoO₄²⁻ and ReO₄⁻, and subsequent removal of these metals to anoxic sediments (1).

Most significantly, the initial 187Os/188Os value of the Re-Os regression is statistically higher than the mantle/extraterrestrial value of 0.11 at 2.5 Ga (the modern value of 0.13 is higher because of 187Re decay to 187Os since 2.5 Ga). This observation indicates that local paleo-seawater 187Os/188Os was higher than the mantle value and hence that oxidative continental weathering supplied radiogenic Os from the upper crust to Hamersley Basin seawater. Previously, the oldest known example of radiogenic Os delivery to seawater was captured by sedimentary rocks of the Huronian Supergroup, deposited in the aftermath of a widespread Paleoproterozoic glaciation at ca. 2.3 Ga (34). All other black shales between 2.7 and 2.0 Ga (n = 5) yield initial 187Os/188Os values from Re-Os isochron regressions that are indistinguishable from the mantle/extraterrestrial value (1, 35–38). The Mount McRae Shale thus captures the oldest known example to date of a radiogenic continental Os contribution.
to seawater and represents the most direct evidence for chemical interaction between photosynthetic O$_2$ and crustal rocks at Earth’s surface.

Continental sulfide minerals probably supplied radiogenic Os to Hamersley Basin seawater. The similar ionic size and charge of Re$^{4+}$ and Mo$^{4+}$ enable significant uptake of Re into molybdenites during crystallization. The average Re concentration of 71 ppm in Archean molybdenite is four orders of magnitude higher than in bulk silicate crust (39). Hence, molybdenites contain high $^{187}$Os concentrations generated by in situ radioactive decay of $^{187}$Re. Iron-bearing sulfide minerals (for example, pyrite, arsenopyrite, and pyrrhotite) have Re concentrations (parts-per-billion levels) that are orders of magnitude lower than those of molybdenite but are relatively more abundant in the continental crust (24, 35, 40, 41). Relationships between dissolved SO$_4^{2-}$, Re, and Mo concentrations in modern rivers suggest that oxidative continental weathering of iron-bearing sulfide minerals is an important source of Re and Mo to modern seawater (42). The oxidative weathering of black shales (which hosts Re and Os in organic matter and pyrite) may be a major supplier of Re but is a relatively minor source of Os and Mo to modern seawater (42, 43).

Other continental minerals (for example, silicates) were not likely to be major sources of Re, Mo, and Os to late Archean seawater. Uranium concentrations in the Mount McRae Shale are consistently low (indicating limited oxidation of U$^{4+}$ to U$^{6+}$ and thus low seawater U concentrations), and U is primarily held in nonsulfide minerals (for example, silicates, phosphates, zircon, and uraninite) that oxidize more slowly relative to sulfides (1, 44, 45). Hence, oxidative mobilization of Re, Mo, and radiogenic Os from continental sulfide minerals followed by their transport to seawater together as metal oxyanions provides a coherent and compelling explanation for the high Re and Mo enrichments and elevated seawater $^{187}$Os/$^{188}$Os captured by the black shales at 145.22 to 148.32 m in the ABDP-9 drill core (Fig. 1).

**Atmospheric O$_2$ concentrations at 2.5 Ga**

Mass-balance modeling places constraints on the relative importance of the continental Os flux during deposition of the Mo- and Re-rich shales at 145.22 to 148.32 m. Compared with today, the Archean continental crust was smaller in volume (46) and had lower $^{187}$Os/$^{188}$Os because less radiogenic $^{187}$Os had accumulated from $^{187}$Re decay. Early Paleoproterozoic (ca. 2.3 Ga) sandstones and siltstones have initial $^{187}$Os/$^{188}$Os as high as 1.1 (34), which may represent an approximate upper limit for the $^{187}$Os/$^{188}$Os of the eroding upper continental crust at 2.5 Ga. The Os isotope mass balance equation for seawater can be represented by

$$f_{\text{continental}} \times f_{\text{mantle/extraterrestrial}} = f_{\text{mantle/extraterrestrial}}$$

where the $f$ terms represent the relative fraction of Os input to seawater and $f_{\text{continental}} + f_{\text{mantle/extraterrestrial}} = 1$. The mantle and extraterrestrial (chondritic) contributions have similar $^{187}$Os/$^{188}$Os. This equation can be rearranged to obtain the continental flux of Os as follows

$$f_{\text{continental}} = \frac{^{187}\text{Os}_{\text{seawater}}}{^{187}\text{Os}_{\text{continental}} + ^{187}\text{Os}_{\text{mantle/extraterrestrial}}}$$

Assuming a seawater $^{187}$Os/$^{188}$Os of 0.34 ± 0.19 (based on our data from the Mount McRae Shale), a mantle/extraterrestrial $^{187}$Os/$^{188}$Os of 0.11, and a continental $^{187}$Os/$^{188}$Os of 1.10, the continental flux represented ~5 to 40% of Os in Hamersley Basin seawater. A higher ratio of continental versus mantle/extraterrestrial inputs would occur if riverine $^{187}$Os/$^{188}$Os was less than 1.10.

A significant continental Os flux to seawater does not require high atmospheric O$_2$ concentrations because Os is a siderophile and chalcophile metal that resides primarily in easily weathered sulfide minerals. Subaerial oxidation of continental sulfide minerals is likely to be more efficient than submarine oxidative weathering within oxygenated surface waters along ocean margins (47). It has also been shown that oxidative dissolution of pyrite and molybdenite can potentially occur on geologically short time scales (tens of thousands of years or less) beneath an atmosphere containing sufficiently low O$_2$ concentrations (~0.001% PAL) (9) to allow preservation of S-MIF signatures in sedimentary rocks (15, 48). However, atmospheric O$_2$ concentrations would likely need to reach ~0.03% PAL (this estimate has an order-of-magnitude uncertainty), at least locally, to allow efficient riverine transport of Os to seawater; otherwise, dissolved Os would have been reduced and removed to riverbeds and floodplains via redox reactions with Fe(II) (5, 49). Above this
atmospheric O$_2$ threshold, theoretical calculations indicate that Fe(II) should be efficiently oxidized to Fe(III) during continental weathering (5).

Recently, it has been hypothesized that localized (sub-meter scale) oxidative weathering environments associated with microbial mats in soils, rock surfaces, and sediments in lakes, rivers, and estuaries could permit the oxidative release of redox-sensitive trace metals and their transport to Archean oceans without any appreciable increase in atmospheric O$_2$ concentrations (50). Our data do not firmly rule out this hypothesis. However, because of the reactivity of Os with Fe(II), long-range transport of Os from a locally oxidizing weathering environment to an organic-rich sediment sink should require sufficiently high background atmospheric O$_2$ levels to allow for pervasive oxidation and removal of Fe(II) from meteoric waters. Thus, we suggest that the radiogenic seawater $^{187}$Os/$^{188}$Os captured by the black shales from 145.22 to 148.32 m is more easily explained if nontrivial amounts of O$_2$ ($\geq 0.03\%$ PAL) were present in the atmosphere because this allows efficient riverine transport of continental Os to seawater.

The S-MIF signatures at 145.22 to 148.32 m (2) do not disappear as expected if atmospheric O$_2$ concentrations were high enough to permit quantitative oxidative transport of Fe(II) in meteoric waters and thus efficient transport of dissolved Os by rivers. We suggest that the persistence of S-MIF signatures may result from propagation of a relic crustal signature (10). Alternatively, the atmospheric O$_2$ level required for widespread removal of dissolved Fe(II) from continental meteoric waters is overestimated. In any case, our data robustly point to oxidative mobilization of Os from Earth’s crust followed by large-scale riverine transport of Os to the Hamersley Basin.

**Transient oxidative continental weathering at 2.5 Ga**

In contrast to the 145.22- to 148.32-m interval, the black shales at 128.71 to 129.85 m yield an initial $^{187}$Os/$^{188}$Os of 0.06 ± 0.09 (Fig. 2) that is statistically indistinguishable from the mantle/extra-terrestrial value of ~0.11 (1), pointing to weak riverine transport of continental Os and associated oxidative continental weathering. The extreme opposite ends of the $2\sigma$ uncertainties of the initial $^{187}$Os/$^{188}$Os from 128.71 to 129.85 m (0.06 ± 0.09) and 145.22 to 148.32 m (0.34 ± 0.19) overlap, but an unpaired $t$ test indicates that the initial $^{187}$Os/$^{188}$Os values from the two intervals are most likely different ($P < 0.0001$, two-tailed).

Stratigraphic trends in the initial $^{187}$Os/$^{188}$Os calculated for individual samples within the two stratigraphic intervals cannot be evaluated because the uncertainties of individual sample ratios (as determined by numerical error propagation of age, $^{187}$Re decay constant, and present-day $^{187}$Re/$^{188}$Os and $^{187}$Os/$^{188}$Os uncertainties) are large and thus statistically overlap (table S1). This is a consequence of the old (late Archean) age of the Mount McRae Shale and its high present-day $^{187}$Os/$^{188}$Os ratios. In contrast, the initial $^{187}$Os/$^{188}$Os of Re-Os isochron regressions, defined by multiple samples, provide a more precise average estimate of seawater $^{187}$Os/$^{188}$Os for each stratigraphic interval. The statistically higher initial $^{187}$Os/$^{188}$Os for 145.22 to 148.32 m (compared with 128.71 to 129.85 m) indicates that the flux of radiogenic riverine Os to seawater became more prominent relative to the radiogenic mantle/extraterrestrial flux at this time.

An alternative explanation for the higher seawater $^{187}$Os/$^{188}$Os at 145.22 to 148.32 m is a decreased flux of mantle/extraterrestrial Os to seawater, but we consider this scenario to be highly unlikely. The upper Mount McRae Shale (125 to 150 m in ABDP-9) was predominantly deposited from euxinic bottom waters based on Fe speciation data (Fig. 1) (15). Hence, the stratigraphic variations in Re and Mo concentrations likely reflect changes in seawater metal concentrations rather than variations in local bottom water redox conditions. The highest Re and Mo concentrations are observed at 140 to 150 m in ABDP-9 (1) and cannot be explained by a decline in the mantle/extra-terrestrial fluxes because such decreases should result in lower Re and Mo concentrations in black shales. Similarly, an increase in the mantle/extra-terrestrial flux during deposition of the black shales above 140 m cannot account for the lower seawater $^{187}$Os/$^{188}$Os at 128.71 to 129.85 m because this scenario should result in an increase rather than a decrease in the Re and Mo concentrations. Hence, the decline in Re and Mo enrichments above 140 m in ABDP-9 is consistent with decreased oxidative weathering rates (1). Together, the high Re and Mo enrichments and elevated seawater $^{187}$Os/$^{188}$Os at 145.22 to 148.32 m point to a transient increase in oxidative continental weathering and likely atmospheric and surface ocean O$_2$ levels.

We emphasize that oxidative weathering was not necessarily negligible before and after this transient event. Concentrations of Mo and Re and Mo isotope compositions in black shales above 140 m and below 150-m depth in ABDP-9 are slightly elevated above continental values, suggesting weak oxidative weathering (1, 19). The main point is that the initial $^{187}$Os/$^{188}$Os data from 145.22 to 148.32 m indicate a significant temporal shift in the scale of oxidative continental weathering during deposition of the Mount McRae Shale. Therefore, the new Os isotope data highlight that the onset of oxygenesis, oxidative continental weathering, and Earth surface oxygenation was most likely a complex and dynamic process (7, 10).

Hence, our new Os isotope data for the 2.5-Ga Mount McRae Shale have profound implications for the nature of photosynthetic O$_2$ accumulation in Archean surface environments. Rather than linear stepwise increases in environmental O$_2$ levels, it is likely that transient oxygenation events were separated by intervals of less oxygenated conditions. Oscillation in O$_2$ levels driven by the combined effects of complex biospheric feedbacks and tectonic processes may have been particularly characteristic of the late Archean Eon. The frequency, magnitude, and duration of these transient environmental oxygenation events may have increased through time—culminating in the early Paleoproterozoic Great Oxidation Event.

**MATERIALS AND METHODS**

**Rhenium-osmium isotope analyses**

The Re-Os analyses were carried out at the Canadian Centre for Isotopic Microanalysis, Department of Earth and Atmospheric Sciences, University of Alberta. Eight samples of finely laminated homogeneous black shale containing disseminated pyrite (but not macroscopic pyrite nodules) were selected from a depth of 148.09 to 148.15 m. Each sample (comprising 15 to 25 g) was ground to remove cutting and drilling marks, broken into chips using metal-free methods, and powdered in an agate mill. About ~0.2 g of sample powder and a known amount of $^{187}$Re-$^{186}$Os spike solution were digested in 8 ml of a Cr(VI)-H$_2$SO$_4$ solution for 48 hours at 200°C in sealed Carius tubes (51). Multiple studies have demonstrated that the Cr(VI)-H$_2$SO$_4$ solution minimizes the dissolution of detrital Os from silicate minerals, whereas the organic matter and sulfide minerals are digested, thus releasing hydrogensulfide and Os into the solution (38, 51–53).

After digestion, Os was separated from the Cr(VI)-H$_2$SO$_4$ solution by solvent extraction into chloroform and further purified by extraction into concentrated HBr, followed by double microdistillation (54, 55). An aliquot (2 ml) of the Cr(VI)-H$_2$SO$_4$ solution was taken and mixed with 15 ml of 5 N NaOH and 15 ml of acetone to separate Re by solvent
higher present-day 187Re/188Os and 187Os/188Os isotope ratios and yielded ABDP-9 (analyses from 128.71 to 129.85 m and four analyses from 145.22 to 148.32 m) which, along with the previous studies (38, 56), Total procedural blanks for Re and Os were 15 and 0.3 pg, respectively, with a blank 188Os/186Os of 0.20.

Statistical analysis
Regression of Re-Os isotope data was carried out using K. Ludwig’s Microsoft Excel add-in program Isoplot 4.15 (Berkeley Geochronology Center) using a 187Re decay constant of 1.666 × 10^-51 year^-1 (57, 58), calculated 2σ uncertainties for sample 187Re/188Os and 187Os/188Os as determined by numerical error propagation, and the correlation (p) between the isotope ratios. The 2σ age uncertainty includes the 187Re decay constant uncertainty of ±0.31% (57).

Further analysis of the Re-Os isotope systematics in the Mount McRae Shale
Previously, a Re-Os isochron age of 2501.1 ± 8.2 Ma (MSWD = 1.1; the calculated 2σ uncertainties for sample 187Re/188Os and 187Os/188Os as determined by numerical error propagation, and the correlation (p) between the isotope ratios. The 2σ age uncertainty includes the 187Re decay constant uncertainty of ±0.31% (57).

SUPPLEMENTARY MATERIALS
Supplementary material for this article is available at http://advances.sciencemag.org/cgi/suppl/1/10/1e1500777/DC1

Table S1. Re-Os abundance and isotope data for the Mount McRae Shale, Western Australia (drill core ABDP-9).

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