Detecting Ocean Glint on Exoplanets Using Multiphase Mapping

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Abstract

Rotational mapping and specular reflection (glint) are two proposed methods to directly detect liquid water on the surface of habitable exoplanets. However, false positives for both methods may prevent the unambiguous detection of exoplanet oceans. We use simulations of Earth as an exoplanet to introduce a combination of multiwavelength, multiphase, time-series direct-imaging observations and accompanying analyses that may improve the robustness of exoplanet ocean detection by spatially mapping the ocean glint signal. As the planet rotates, the glint spot appears to “blink” as Lambertian scattering continents interrupt the specular reflection from the ocean. This manifests itself as a strong source of periodic variability in crescent-phase disk-integrated reflected light curves. We invert these light curves to constrain the longitudinal slice maps and apparent albedo of multiple surfaces at both quadrature and crescent phase. At crescent phase, the retrieved apparent albedo of ocean-bearing longitudinal slices is increased by a factor of 5, compared to the albedo at quadrature phase, due to the contribution from glint.

The land-bearing slices exhibit no significant change in apparent albedo with phase. The presence of forward-scattering clouds in our simulated observation increases the overall reflectivity toward crescent, but we find that clouds do not correlate with any specific surfaces, thereby allowing for the phase-dependent glint effect to be interpreted as distinct from cloud scattering. Retrieving the same longitudinal map at quadrature and crescent phases may be used to tie changes in the apparent albedo with phase back to specific geographic surfaces (or long-standing atmospheric features), although this requires ideal geometries. We estimate that crescent-phase time-dependent glint measurements are feasible for between 1 and 10 habitable zone exoplanets orbiting the nearest G, K, and M dwarfs using a space-based, high-contrast, direct-imaging telescope with a diameter between 6 and 15 m.

Key words: astrobio(1)logy – planetary systems – planets and satellites: oceans – planets and satellites: surfaces – planets and satellites: terrestrial planets

1. Introduction

Finding and characterizing habitable exoplanets is an emerging frontier, with the detection of liquid surface water as a fundamental goal on the path toward discovering life beyond the solar system. A promising few small exoplanets have been discovered in the habitable zones of their parent stars (Anglada-Escudé et al. 2016; Gillon et al. 2016, 2017; Dittmann et al. 2017). These potentially habitable exoplanets are likely high priorities for upcoming observations to characterize their environments and search for signs of habitability. However, characterizing the nature of a planet’s habitability will likely require many different lines of evidence synthesized to assess the likelihood of surface liquid water (for a recent review, see Meadows & Barnes 2018; Robinson 2018).

Detecting an ocean on the surface of an exoplanet is the most straightforward way to confirm habitability, and high-contrast direct imaging is the most promising type of observation to accomplish this. Unlike transmission spectroscopy, which cannot probe the surface and near-surface atmosphere (García Muñoz et al. 2012; Bétrémieux & Kaltenegger 2014; Misra et al. 2014), direct imaging can probe all the way to the planetary surface, and is less susceptible to obscuration by atmospheric extinction (Fortney 2005), such as by aerosols (Kreidberg et al. 2014). As a result, only direct imaging offers a feasible path to exoplanet ocean discovery. There are two ways in which exoplanet oceans could be directly detected using direct imaging of planetary reflected light: spatial mapping using the planet’s rotational light curves, and ocean glint using phase-dependent and/or polarization observations (Zugger et al. 2010).

Although any given snapshot of a directly imaged exoplanet is disk integrated and contains no spatial information, as the planet rotates on its axis, time-series observations permit spatial mapping of heterogeneities in the atmosphere and on the surface (Ford et al. 2001; Pallé et al. 2008; Oakley & Cash 2009). Multiwavelength time-series observations can be used to infer the longitudinal map and color of multiple surfaces on a planet (Cowan et al. 2009; Cowan & Strait 2013), and multi-epoch observations can leverage the planet’s
obliquity and/or inclination to map latitudes and construct two-dimensional color surface maps of exoplanets (Kawahara & Fujii 2010, 2011; Fujii & Kawahara 2012; Berdyugina & Kuhn 2017). Such color maps could be used to identify large oceans on exoplanets.

An alternative approach to detecting exoplanet oceans is to observe glint—the specular reflection typical of a flat surface or liquid when viewed at an oblique angle relative to the illumination source—a remotely detectable characteristic of the Earth ocean (Sagan et al. 1993). Glint has been suggested as a promising habitability marker (Williams & Gaidos 2008; Robinson et al. 2010), and produces a substantial relative increase in the brightness of Earth when viewed at crescent phases (Robinson et al. 2014). Targeted crescent-phase observations could be used to identify the reflected light glint signature of an ocean-bearing exoplanet.

Although spatial mapping and glint are promising, both have degeneracies in interpretation that may limit the confidence with which each technique may be used to directly detect oceans. Light-curve inversion techniques to identify the albedo and geography of multiple surfaces, including oceans (e.g., Cowan et al. 2009), are under-constrained inference problems that suffer from mathematical degeneracies (Fujii et al. 2017) and have traditionally required a priori assumptions about the surface composition to find unique solutions (e.g., Fujii et al. 2010). More recently, Fujii et al. (2017) imposed the simple physical prior that the surface albedo be bounded between zero and one. This was insufficient to fully break this “rotational unmixing” degeneracy, but showed promise for achieving simultaneous constraints on both the color and geography of Earth-like oceans because of their low albedo and vast global distribution. Furthermore, simply inferring a dark blue surface is not a definitive ocean detection, as Rayleigh scattering above a gray surface can masquerade as “blue” in color (Krissansen-Totton et al. 2016). Glint also has potential false positives. The non-Lambertian increase in reflectivity toward crescent phase that is characteristic of ocean glint can occur independent of a glinting ocean as a result of forward-scattering from clouds and hazes, as observed for Venus (e.g., Mallama et al. 2006) and Earth (e.g., Robinson et al. 2010). Furthermore, high latitudes, which are more likely to be covered in ice, contribute more to the disk-integrated flux at crescent phase than more gibbous phases. This can naturally cause an increase in reflectivity toward crescent phase (Cowan et al. 2012). Additional strategies and techniques are needed to help increase the robustness of ocean detection.

Here we introduce a sequence of direct-imaging observations that combine longitudinal mapping with phase-dependent glint measurements to increase the robustness of ocean detection. Specifically, multiwavelength time-series observations at multiple phases and wavelengths can enable spectral unmixing as a function of illumination. An exoplanet in a nearly edge-on inclination orbit, with a negligible obliquity, will have rotational light curves that probe the same latitudes at all orbital phases, and therefore the two-dimensional mapping problem (e.g., Fujii & Kawahara 2012) will be intractable as longitudinal maps will appear the same at all observed phases. However, recognizing the same longitudinal map at another phase angle would offer a second look at the same surface, thereby allowing observers to trace the reflective properties of a given geographical surface as a function of illumination angle. Thus, the reoccurring map may lend confidence to any assumption that the same surface is being viewed, while the surface albedo necessary to reproduce the observed light curves would reveal its scattering phase function. In this way, it may be possible to infer that the same surface that appears dark blue at gibbous phases is significantly more reflective at crescent phase, thereby proving the non-isotropic scattering behavior of the surface.

In this paper we demonstrate the potential utility of multiphase longitudinal mapping as it may be applied to detect oceans on directly imaged exoplanets. In Section 2 we describe the methods we used to generate synthetic light curves of Earth and to invert them to map the planet. In Section 3 we present results on multiphase longitudinal mapping and observational feasibility tests. We offer a discussion in Section 4 and conclude in Section 5.

2. Methods

To investigate and develop a more robust technique for exoplanet ocean detection, we use a forward model for the multiwavelength time-series photometry of the disk-integrated Earth, apply realistic noise for direct-imaging exoplanet observations, and then invert the problem to solve for the geographical maps and colors of dominant surface types. The methods for generating synthetic light curves of Earth and fitting maps to the light curves are described in the following subsections.

2.1. Generating Synthetic Light Curves

We use the VPL Earth Model to create high spectral resolution, time-dependent, disk-averaged observations of Earth at multiple phases (Robinson et al. 2011). The VPL Earth Model gathers observed cloud and ice coverage, surface wind speed, and atmospheric thermal structure and gas-mixing ratio profiles from Earth-observing satellites. These Earth science data products are used as inputs to the VPL Earth Model, which maps them to a HEALPix11 model (Górski et al. 2005) projection of Earth. Multi-stream, multi-scattering, line-by-line radiative transfer calculations are then performed across the grid of cloud coverages and illumination angles for each HEALPix pixel using the Spectral Mapping and Atmospheric Radiative Transfer (SMART) model (Meadows & Crisp 1996, developed by D. Crisp). SMART, and by extension the VPL Earth Model, includes the spectral reflection of the Sun or star off the ocean using the Cox-Munk glint model (Cox & Munk 1954). Performing SMART calculations yields a database of radiances that contains contributions emerging from different regions of Earth at any given time. Spatially and spectrally resolved data cubes of Earth are constructed according to user-specified sub-solar and sub-observer longitudes and latitudes (Robinson et al. 2011). An example VPL Earth model image of the crescent-phase Earth with a clearly visible glint spot is provided in Figure 1 of Robinson et al. (2010). These data cubes are then integrated over the visible disk to give a single composite disk-integrated spectrum.

The VPL Earth Model has been validated against spacecraft observations of Earth, including spatially resolved and time-dependent observations made by the EPOXI mission (Robinson et al. 2011), and phase-dependent measurements of Earth showing the glint effect that were taken by the Lunar Crater Observing and Sensing Satellite (LCROSS) (Robinson et al. 2011). An example VPL Earth model image of the crescent-phase Earth with a clearly visible glint spot is provided in Figure 1 of Robinson et al. (2010). These data cubes are then integrated over the visible disk to give a single composite disk-integrated spectrum.

11 http://healpix.sf.net/
We refer to Robinson et al. (2011, 2014) and Schwieterman et al. (2015, 2016) for a more complete description of the VPL Earth Model and its validation.

The VPL Earth Model is used to generate multiple different Earth and Earth-like spectral models that enable longitudinal mapping investigations at multiple phases. We make a simplifying assumption that the sub-solar and sub-observer latitudes are fixed at the equator to simulate a planet with zero obliquity in an edge-on orbit. We compute time-dependent spectra at three phase angles: 45°, 90°, and 135°, which we refer to as gibbous, quadrature, and crescent hereafter. At each phase we simulate 120 consecutive hours (5 days) of high-resolution, disk-integrated spectra at a 1 hr cadence. We neglect the small (<5°) phase change of the planet over the 120 hr of Earth modeling. The computational expense of the VPL Earth Model places limits on the observational duration and observational cadence that we may consider in this work, and in Section 3.1 we focus on experiments with 100 hr of data collection at a 1 hr cadence. We simulate Earth during northern spring. Data sets are generated both with true time-varying cloud coverage and without clouds.

To simulate realistic observations of our VPL Earth Model spectra as an analog for a directly imaged exoplanet, we use an open-source Python coronagraph noise model, originally developed by Robinson et al. (2016) for the WFIRST exoplanet direct-imaging mission and expanded upon in subsequent studies (Luger et al. 2017; Meadows et al. 2018a). The model simulates planetary and background photon count rates as a function of wavelength using analytic parameterizations for telescope, instrumental, and astrophysical noise sources. These sources include Poisson noise, read noise, dark current, coronagraph speckles, scattering off of zodiacal and exozodiacal dust, and thermal emission from the mirror. The telescope is assumed to perform a roll maneuver to allow for background subtraction (see Brown 2005), which increases the background by a factor of two (Robinson et al. 2016).

We perform coronagraph simulations for various telescope designs that encompass telescope mission concepts currently under study by NASA Science & Technology Definition Teams (STDTs) in advance of the 2020 decadal survey. Many of the telescope, instrument, and system parameters used here are the same as those used in Robinson et al. (2016, see their Table 2 and Table 3), with the exceptions listed in Table 1. We adopt a raw contrast of $C = 10^{-10}$ to pursue Earth-like exoplanets orbiting Sun-like stars. We consider telescope diameters between 4 and 15 m, and end-to-end throughputs between 0.02 and 0.5. These ranges generally encapsulate those relevant to the Large UV/Optical/IR (LUVIOR) (e.g., Dalcanton et al. 2015; Bolcar et al. 2017) Surveyor and Habitable Exoplanet Imaging Mission (HabEx) (e.g., Mennesson et al. 2016) telescope mission concepts. We adopt a highly optimistic inner working angle (IWA) of 1A/D and outer working angle (OWA) of 60A/D so that our simulated light curves solely represent the photon counting statistics. However, we separately consider constraints on the IWA in Section 3.2.1. We note that, like Robinson et al. (2016), the IWA and OWA are modeled as strict cutoffs—a step function in throughput from 100% to 0%—whereas formally they are defined as the 50% off-axis throughput point. As a result, our coronagraph modeling neglects throughput effects as the IWA and OWA are approached and exceeded.

We construct simultaneous multiwavelength light curves by simulating coronagraph spectroscopy with a resolving power of $R = \lambda/\Delta \lambda = 200$ at a one-hour cadence over a modest 18% bandpass. This gives a spectrum with a low signal-to-noise ratio (S/N) each hour, and from hour to hour, the spectrophotometry samples Earth’s rotational variability. While multiple hours of spectra can be coadded into a spectrum with a higher S/N that may be used to retrieve atmospheric gas abundances (Feng et al. 2018), here we instead coadd neighboring spectral resolution elements to increase the S/N on each hour of photometry. Since current coronagraph designs can only null over relatively modest spectral bandpasses (typically $\Delta \lambda/\lambda \approx 10\%-20\%$) in a single exposure, we focused on synthetic observations in an 18% bandwidth wavelength range between 0.67 and 0.8 µm. In this wavelength range, the Earth exhibits large-amplitude rotational variability owing to surface features such as ocean, soil, and vegetation. This sensitivity is due in large part to the Earth’s atmosphere being optically thin, with very few molecular absorption features and reduced Rayleigh scattering at most of these wavelengths. The one major exception is the opacity due to the 0.76 µm O2 absorption band, which is likely to be one of the most sought-after atmospheric diagnostics for Earth-like planets (Meadows 2017; Meadows et al. 2018b). Consequently, the same multiwavelength time-resolved data that are used to map the planet and search for oceans can be coadded to characterize the planet’s atmosphere and search for signs of life. We bin the spectrophotometry into two photometric bands, centered at $\lambda_1 = 0.70 \mu m$ and $\lambda_2 = 0.77 \mu m$, with widths $\Delta \lambda_1 = 0.08 \mu m$ and $\Delta \lambda_2 = 0.07 \mu m$, sampled over many consecutive hours and with corresponding noise applied, which yield light curves for studying the time-variability of the exoplanet. We assume no read time, no reset time, and no down time between consecutive exposures, thereby simulating a 100% efficient duty cycle.

### 2.2. Light-curve Fitting

Time-resolved multiwavelength planetary light curves can be used to constrain the longitudinal map of a planetary
surface, using techniques of “spectral-rotational unmixing”, which solve for the longitudinal map of $N$ different unique surface types and their geometric albedo spectrum (Cowan et al. 2009, 2012; Fujii et al. 2017). As described in detail in Fujii et al. (2017) and summarized here for completeness, the observed apparent albedo of the planet, $A_{ij}$, at wavelength $j$ and at time $i$ can be decomposed into

$$A_{ij} = \sum_{i,k} \tilde{f}_{ik} a_{ij},$$

(1)

where $\tilde{f}_{ik}$ is the apparent covering fraction of the $k$th surface at the $i$th time and $a_{ij}$ is the apparent albedo spectrum of the $k$th surface at the $j$th wavelength.

Similarly, as a result of the relationship between time variability and surface heterogeneity for a rotating planet, the observed apparent albedos can be written in terms of a longitudinal map. In this case, $\tilde{f}_{ik}$ may be approximated by

$$\tilde{f}_{ik} \approx \sum_{l} W_{il} f_{lk},$$

(2)

where $f_{lk}$ is the area fraction of the $k$th surface in the $l$th longitudinal slice and $W_{il}$ is the weight of the $l$th longitudinal slice at the $i$th time, which is a function of the observational geometry. The approximation in Equation (2) is only strictly valid when the true area fraction within each longitudinal slice does not change or changes little across the slice for all surfaces. Plugging Equation (2) into Equation (1), we reach the following expression in terms of longitudinal slices:

$$A_{ij} \approx \sum_{i,k} W_{il} f_{lk} a_{ij}.$$  (3)

Equations (1) and (3) both form the two modular forward models in the Surface Albedo Mapping Using Rotational Inference (samurai)\(^{16}\) Python code developed in this work. samurai attempts to solve the inverse problem for the map (\(\tilde{f}_{i} f_{k}\) or \(f_{i} f_{k}\)) and albedo (\(a_{ij}\)) of $N$ different surfaces using multiband (\(j\)), time-series (\(i\)) photometric observations \(y(\lambda)\) and their observational uncertainties \(\sigma(\lambda)\), subject to the following constraints on the model \(A_{ij}\):

$$\tilde{f}_{i} f_{k} \geq 0 \quad \text{for any } i, l, k$$  (4a)

$$\sum_{k} \tilde{f}_{i} f_{k} = \sum_{k} f_{lk} = 1 \quad \text{for any } i, l$$  (4b)

$$0 \leq a_{ij} \leq 1 \quad \text{for any } k, j.$$  (4c)

Equation (4a) ensures that the area fraction of any surface $k$ at any time $i$ or in any longitudinal slice $l$ is non-negative, while Equation (4b) enforces that the area fractions for all surface types at any time or in any longitudinal slice sum to unity. Equation (4c) imposes that the albedo of a surface $k$ at any wavelength $j$ does not exceed 1.

The hard boundaries in Equations (4a)–(4c) make it difficult to efficiently sample the likelihood space using a Markov chain Monte Carlo (MCMC) algorithm. One way around this is to reparameterize so that the domain of our parameters extends from $-\infty$ to $+\infty$. This can be done by introducing the transformed parameters

$$b_{ij} = \ln \left( \frac{a_{ij}}{1 - a_{ij}} \right) \quad \text{for any } k, j$$  (5a)

$$g_{ik} = \ln \left( \frac{f_{ik}}{1 - \sum_{m=1}^{M} f_{im}} \right) \quad \text{for } k = 1, ..., N - 1,$$  (5b)

where $N$ is the number of surface types included in the model. Note that conditions (4a)–(4c) are implicitly satisfied by this reparameterization, and there is a one-to-one correspondence between albedo matrices $a$ and $b$ and between area-covering fraction matrices $f$ and $g$. Similar reparameterizations have been applied to gas-mixing ratios in exoplanet atmospheric retrieval problems (see Benneke & Seager 2012).

It is important to note that when reparameterizing our model, we must also transform the prior probability of each model parameter. Since we have no a priori information on the albedos and area fractions of individual surfaces, we enforce a flat prior for these variables. This can be obtained by imposing the prior probabilities

$$P(b_{ij}(a_{ij})) \propto \frac{e^{b_{ij}}}{(1 + e^{b_{ij}})^2}$$  (6)

$$P(g_{ik}(f_{ik})) \propto \frac{e^{g_{ik}}}{(1 + e^{g_{ik}})^2}$$  (7)

on our transformed variables $b$ and $g$, where $P$ is a probability distribution function (PDF). Note that Equation (7) does not contain the covariance term that arises due to the sum in Equation 5(b). As a result, Equation (7) does not guarantee a flat prior for three or more surfaces, and thus all investigations in this paper are performed with only two surfaces. We impose one additional prior $\mathcal{O}(g_{ik})$ to ensure that surface-covering fractions are ordered from largest to smallest to break a surface-labeling degeneracy. Considering all elements in $b$ and $g$, and $\mathcal{O}$, our full prior is

$$\pi(b_{ij}, g_{ik}) = P(b_{ij}) P(g_{ik}) \mathcal{O}(g_{ik})$$  (8)

$$\propto \prod_{k,j} \frac{e^{b_{ij}}}{(1 + e^{b_{ij}})^2} \prod_{i,k} \frac{e^{g_{ik}}}{(1 + e^{g_{ik}})^2} \mathcal{O}(g_{ik}).$$  (9)

The aim of our parameter estimation is to evaluate the probability of the albedo and map parameters—the posterior PDF—given multiwavelength time-series data. By Bayes’ theorem, the posterior PDF is

$$P(b_{ij}, g_{ik} | y_{ij}) \propto \mathcal{L}(y_{ij} | b_{ij}, g_{ik}) \times \pi(b_{ij}, g_{ik}),$$  (10)

where the likelihood $\mathcal{L}(y_{ij} | b_{ij}, g_{ik})$ is easily calculated from the model and data,

$$\ln \mathcal{L}(y_{ij} | b_{ij}, g_{ik}) = -\frac{1}{2} \sum_{i,j} \left( \frac{y_{ij} - W_{ij} f_{ik} a_{ij}}{\sigma_{ij}} \right)^2.$$  (11)

Unless otherwise stated, we sample the log-probability (log of Equation (10)) using the emcee\(^{17}\) implementation of MCMC (Foreman-Mackey et al. 2013). We use a linear combination of 15% affine-invariant (Goodman &

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\(^{16}\) https://github.com/jlustig/yama

\(^{17}\) emcee https://github.com/dmb/emcee.
3. Results

We performed spectral unmixing experiments at quadrature, crescent, and gibbous phase to determine if a combination of phase-dependent surface-mapping and glint observations could provide more robust constraints on the presence of an ocean for an Earth-analog exoplanet. We then quantified the observational feasibility and telescope requirements for this technique with respect to coronagraph IWA and the observed photometric S/N.

3.1. Multiphase Longitudinal Mapping

Our first multiphase mapping experiment focused on models without clouds to provide an ideal end-member to isolate the signal from the ocean. This also allowed us to assess the unique impact of clouds once they were included.

The upper panel of Figure 1 shows the quadrature ($\alpha = 90^\circ$; solid lines) and crescent ($\alpha = 135^\circ$; dotted lines) phase multiwavelength light curves and median model fits to these data. The light curves were simulated for an optimistic case of Earth at 5 pc, observed for 100 hr at each phase, at a 1 hr cadence, with a 15 m, 50% throughput space telescope. The spectrophotometry was binned into two wavelength intervals to allow for slopes in the albedo of each surface. This yielded light curves at 0.70 ± 0.04 $\mu$m and 0.77 ± 0.03 $\mu$m, with an average 1 hr photometric uncertainty of 5.0% (S/N = 20.1) and 4.0% (S/N = 25.4) at quadrature, and 6.8% (S/N = 14.8) and 6.1% (S/N = 16.5) at crescent. The decrease in photometric precision seen for the crescent light curves in Figure 1 is due to the decrease in the illuminated portion of the planet’s disk for observations using the same telescope.

The lower panels of Figure 1 show the retrieved apparent albedo and longitudinal slice maps of each surface for a two-surface 16-slice fit to the synthetic multiwavelength light curves shown in the upper panel. The choice of 16 slices was made so that the angular slice width ($22.5^\circ$) divides the change in phase (45°) by an integer, thereby allowing the quadrature and crescent-phase maps to be shifted and aligned. The slice maps are displayed on a Robinson map projection, and use color intensity to indicate the median retrieved area-covering fraction in each slice. The retrieved surface albedos are shown with a 1$\sigma$ uncertainty envelope about the median posterior PDFs derived from the MCMC. The apparent albedo of Surface 1 (blue) exhibited a strong increase from quadrature to crescent phase, becoming $\sim16\times$ brighter, while the albedo of Surface 2 (orange) increased by only $\sim15\%$, which was within the derived uncertainties.

Figure 2 explores the statistical significance of the increase in apparent albedo of Surface 1 and Surface 2 between the retrievals at quadrature and crescent phase. Histograms showing the change in apparent albedo of each surface were produced by taking 5000 random samples from the wavelength-averaged posterior PDFs at each phase and then subtracting the quadrature phase result from the crescent-phase result. The apparent albedo of Surface 1 increased by 0.47 ± 0.02, while that of Surface 2 remained constant with phase (0.02 ± 0.02).

Figure 3 examines the difference in the retrieved longitudinal maps with phase. A qualitative comparison shows that the maps at quadrature and crescent phase in Figure 1 are very similar, but a direct comparison of the maps and their uncertainties at each phase reveals both regions of agreement and regions of disagreement. The two longitudinal maps in Figure 3 show the median (left) and standard deviation (right) of the difference between the maps at each phase in each slice. This is simply the difference between the two Surface 1 slice maps in Figure 1, with uncertainties propagated by randomly sampling the MCMC posterior distributions. Figure 3 is shown with respect to Surface 1, although for a two-surface longitudinal model, Surface 2 is simply one minus the map of Surface 1 (see condition 4(b)). The median difference between the maps is less than 0.5 for all slices. The histogram in the bottom panel of Figure 3 shows the slice-averaged (or total) fractional area increase with the change in phase. The crescent-phase results include slightly more contribution from Surface 1 (0.10 ± 0.06) than the quadrature results. However, the longitudinal maps agree best for slices containing majority open ocean, with the marginally significant change in the maps with phase primarily driven by disagreements in the land-bearing slices. This suggests that even despite the ideal system geometry, telescope noise properties, and lack of clouds, the inferred maps have intrinsic differences.

Simulations performed at gibbous phase (45°) show similar results to those at quadrature phase and are therefore not shown. At gibbous phase, the recovered apparent surface albedos are consistent with the recovered albedos at quadrature, with one very dark surface and another more reflective surface that increases in albedo from about 0.1 to 0.2 over the wavelength range of interest. However, the retrieved maps at gibbous phase are less precise than both the quadrature and crescent-phase maps, although they have a higher 1 hr photometric S/N. This difference is due to the longer phase, which reflects more total light to the observer, but this light is integrated over a larger portion of the planetary disk, which lowers the observer’s sensitivity to each individual longitude, thereby lowering the resolution of the recovered maps (Fujii et al. 2017).

We performed the same experiment on simulations of Earth with real time-varying cloud coverage. Figure 4 shows results from multiple surface-mapping retrievals performed for Earth at 90° (quadrature; solid lines) and 135° (crescent; dotted lines), where each row of plots corresponds to data generated at a different S/N. The upper panel of Figure 4 shows longitudinal mapping results from light curves simulated for the optimistic case of Earth at 5 pc, observed for 100 hr at each phase, at a 1 hr cadence, with a 15 m space telescope with 50% throughput (same setup as Figure 1). The middle and lower panels of Figure 4 show similar results to the upper panel, except that they correspond to data simulated for a 15 m
telescope with 10% throughput and to a 6 m telescope with 10% throughput, respectively. These three data sets have a 1 hr photometric S/N at quadrature phase of 23 (4.3%), 10 (10%), and 2.4 (42%), from the top to bottom panels. As in the cloudless case, the spectrophotometry was binned into two wavelength intervals, and the maps were fit with a two-surface model with 16 longitudinal slices. The envelope about each line in Figure 4 shows the 1σ uncertainty about the median retrieved albedo spectrum parameters derived from the MCMC.

Comparing results at each phase for the data with the highest S/N, we find that the albedo of Surface 1 exhibits the same large increase at crescent as seen for the cloudless case, even in the presence of realistic clouds, and the longitudinal maps agree reasonably well, particularly at longitudes between −180° and 0°. By the same test as shown in Figure 2, the apparent albedo of Surface 1 increased by 0.57 ± 0.02 with the change in phase, while that of Surface 2 increased by a slight margin of 0.08 ± 0.02. Similarly, the crescent-phase results include more contribution from Surface 1 (0.14 ± 0.06) than

Figure 1. Multiphase inverse problem example for the cloudless Earth. Top: Synthetic light curves for two wavelength intervals with error bars for the cloudless Earth, and retrieved fits to the observations (lines) at quadrature (solid lines) and crescent (dotted lines) phase. Data were simulated for the cloudless Earth at 5 pc observed with a 15 m telescope with 50% throughput, which yielded a 1 hr photometric precision of 4.5% at quadrature and 6.5% at crescent. Bottom left: Retrieved surface albedos from a two-surface fit to rotational light curves of the cloudless Earth at 90° phase and 135° phase. The color of each line identifies the surface, and the line style denotes the phase. The envelope about each line shows the 1σ interval about the median posterior PDF derived from the MCMC. From quadrature to crescent phase, the average albedo of Surface 1 increases by a factor of ∼16 with 15σ confidence, while the average albedo of Surface 2 increase by ∼15% at <1σ. Bottom right: Median retrieved longitudinal slice maps. Each longitudinal slice is colored proportional to the retrieved area-covering fraction of the given surface, with white meaning that none of the given surface is found within the slice. Maps are shown in the Robinson projection.
the quadrature results, and again, the longitudinal maps agree best for slices containing majority open ocean, with the change in the maps with phase driven primarily by disagreements in land-bearing slices.

The longitudinal mapping results from the data sets with a lower S/N in Figure 4 (lower panels) demonstrate the photometric limitations of multiphase mapping. The middle panel with S/N = 10 is quite similar to the case with S/N = 23, but with slightly larger uncertainties on the crescent-phase surface albedos and longitudinal maps. The lower panel with S/N = 2.4 no longer captures the variability at crescent phase: the albedo and longitudinal map of both surfaces are unconstrained and overlap significantly. The convergence of the two surfaces implies that a single-surface model (mean observed albedo and no map) is warranted at crescent phase. At quadrature phase, the S/N = 2.4 result still shows success in extracting the albedos of both surfaces and mapping them to the ocean and land. In the following subsection, we further explore the limitations and feasibility of multiphase longitudinal mapping, particularly with respect to difficult crescent-phase observations.

3.2. Observational Feasibility

We now assess the feasibility of using this phase-dependent technique to detect oceans on extrasolar planets, considering angular separation, photometric sensitivity, and target yields. To obtain direct-imaging observations of an Earth-analog exoplanet at a crescent phase of 135°, the planet must fall outside the coronagraph IWA. Additionally, the crescent-phase light curves must be observed with sufficient photometric precision for the relatively short exposure times needed for mapping to be able to retrieve multiple surfaces. Below, we describe these observing considerations, and estimate the number of nearby planetary systems where this technique
could be applied to search for oceans with upcoming telescopes.

### 3.2.1. IWA Constraints

The IWA is the minimum angular separation, \( \theta_{IWA} \), for which a coronagraph can occult a star without blocking the light of an exoplanet in orbit about it. Since \( \theta_{IWA} \propto \lambda/D \), it is convenient to define

\[
\theta_{IWA} = N \lambda / D, \tag{12}
\]

where \( N \) is a factor set by the specific coronagraph design. Lower values of \( N \) enable planets at smaller angular separations from their host star to be imaged. In the limit that the distance to the observed system, \( d \), is much farther away than the projected planet–star separation, \( r \), then

\[
\theta_{IWA} \approx r / d, \tag{13}
\]

We can solve for the coronagraph specific IWA parameter \( N \) in terms of key telescope and system parameters,

\[
N \approx 4.8 \left( \frac{r}{1 \text{ au}} \right) \left( \frac{D}{10 \text{ m}} \right) \left( \frac{10 \text{ pc}}{d} \right) \left( \frac{1 \mu \text{m}}{\lambda} \right). \tag{14}
\]

We initially consider observations of habitable Earth-like exoplanets orbiting Sun-like stars for which \( r \approx 1 \text{ au} \), at wavelengths near \( \lambda \approx 0.8 \mu \text{m} \), but using a telescope with a currently unspecified mirror diameter, and for systems that have yet to be discovered.

To better understand the coronagraphic requirements needed for ocean detection, we explored contours of constant \( N \) while varying the telescope diameter and system distance. Figure 5 shows coronagraph IWA requirements to directly image an Earth-analog exoplanet at quadrature phase (top panel) and at crescent phase (bottom panel). At quadrature phase, the projected planet–star separation is the semimajor axis, \( a = 1 \text{ au} \), but at crescent phase with \( \alpha = 135^\circ \), the projected separation for an edge-on inclination system is \( a \sin(\alpha) \approx 0.707 \text{ au} \). This tighter planet–star separation imposes a stricter constraint on the coronagraph IWA that ultimately reduces the distance out to which exoplanet systems may be observed at \( 135^\circ \) phase relative to quadrature phase for any given coronagraph design. For instance, a 10 m telescope with \( \theta_{IWA} = 3 \lambda / D \) can image Earth-analog exoplanets out to \( \sim 20 \text{ pc} \) at quadrature, but only \( \sim 14 \text{ pc} \) at \( 135^\circ \) degrees.

### 3.2.2. Photometric Constraints

We also considered the feasibility of detecting large changes in the albedo of a particular surface with phase as a function of observational precision, and the effects of telescope design and target distance on scientific yields. We repeated the same multiphase longitudinal mapping experiment as in Section 3.1 for the true cloudy Earth observed for 100 hr with a 1 hr cadence, but now for varying levels of photometric precision. However, instead of fixing the S/N on each light curve, we fixed the underlying telescope architectural parameters used to simulate the light curves to enforce self-consistency between observations at quadrature and crescent phase. We then determine for what data quality or telescope design statistically significant increases in the albedo of Surface 1 may be retrieved concurrent with statistically similar longitudinal maps.

Figure 6 shows both changes in retrieved apparent albedo of the two surfaces and changes in retrieved longitudinal maps as a function of data quality. The lower x-axis corresponds to the 1 hr photometric S/N for the Earth at quadrature phase, while the upper x-axis corresponds to the 1 hr S/N at crescent phase for observations of the same exoplanet at the same distance with the same model telescope configuration and identical integration time. Thus, fixing all parameters except for the phase, crescent-phase observations have approximately half the S/N of quadrature phase observations. The upper panel of Figure 6 shows the phase-dependent albedo increase of Surface 1 (blue) and Surface 2 (orange) with 1σ error bars derived by randomly sampling the MCMC posterior distributions. This visualization presents posterior PDFs, analogous to those shown in Figure 2, as points with errors along an x-axis of increasing data quality. The lower panel shows a similarly derived result, but for changes in longitudinal area-covering fraction, which is analogous to results from Figure 3, but now as a function of S/N.
observer latitude. We propose that this discrepancy is due to the effect of a multicomponent scattering phase function on the inferred longitudinal maps, which we discuss further in Section 4.1.

The asymptotic behavior of the apparent albedo increase in Surface 1 for \( S/N > 10 \) suggests a natural lower limit of the precision below which the difference in apparent albedo with phase will be lost in the noise. To further investigate this limit and determine the photometric limits of mapping—particularly at crescent phase—we find the minimum number of surfaces that are justified in fits to data with different \( S/N \). We use the Bayesian information criterion (BIC; Schwarz 1978) to penalize overly complicated model fits. We take the “optimal” number of surfaces to be the model with the minimum BIC. We define a convention that the optimal number of surfaces determined by the BIC is the number of detected surfaces, where a one-surface fit is a homogeneous map and a two-surface fit is the bare minimum for performing any meaningful mapping. This is similar to using principal component analysis to quantify detected surfaces (Cowan et al. 2009; Fujii et al. 2017), except that we find it to be significantly more robust against observational uncertainties.

We used the BIC to find the number of detected surfaces across a finely sampled grid in photometric \( S/N \) at both quadrature and crescent phase. For each \( S/N \), we calculate the BIC for a one-surface model by finding the best fit from 100 randomly initialized optimizations using the L-BFGS-B algorithm (Zhu et al. 1997). We then increment the number of surfaces and repeat the procedure until we discover the minimum BIC. The two panels of Figure 6 are shaded red to denote the crescent-phase photometric precision for which only one surface is detected. The boundary occurs at a photometric precision of \( \sim 15\% \) (\( S/N = 6.7 \)) for observations at crescent phase, which we take to be the minimum precision capable of performing phase-dependent mapping on a partially cloudy Earth-like exoplanet. We also find that the boundary between one and two surfaces requires a slightly improved photometric precision of \( \sim 10\% \) (\( S/N = 10 \)) for observations at quadrature phase. This indicates that the increased color contrast between the two surfaces at crescent aids in multi-surface mapping, although higher precision measurements are easier to acquire at quadrature phase. We also note that three surfaces are detected at photometric precision \( > 2.8\% \) (\( S/N = 35 \)) at quadrature.

Adopting the two-surface threshold at crescent phase as the lower limit to the photometric precision, we determine the ability for different possible future telescopes to meet this requirement with respect to scientific yields. Figure 7 shows the distance out to which an Earth-like exoplanet could have crescent-phase (\( \alpha = 135^\circ \)) observations that meet the \( S/N \) threshold for two-surface longitudinal mapping as a function of telescope diameter and end-to-end telescope throughput. For example, a 15 m telescope with 20\% throughput could perform crescent-phase mapping out to 10 pc, while an 8 m telescope with 20\% throughput could only go out to 5.4 pc, and a 4 m telescope might reach only 2.7 pc.

3.2.3. Extending to Low-mass Stars

The forward models used here were developed for Earth–Sun analog systems. However, we can broaden our scope to consider observations of planets with surface oceans orbiting later-type stars, using the Earth as a proxy for these planets. Cooler stars have compact habitable zones and therefore
of different diameters, assuming $\theta_{\text{IWA}} = 3\lambda/D$, to determine if a planet in such an orbit is observable. Additionally, to meet the photometric requirements for mapping, we also check to see if an Earth-size and albedo planet at the inner edge of the HZ exceeds our 1 hr photometric S/N threshold of 6.7, assuming the same telescope and instrumental specifications provided above and a 20% throughput. If a given stellar system satisfies both the IWA and the photometric conditions, then we count that system as observable.

However, not all stars will have an HZ planet, and not all HZ planets will be optimally inclined so that they pass through a crescent phase of $135^\circ$. These are questions on the occurrence rate of HZ exoplanets and the random distribution of system inclinations, both of which are fundamental to exoplanetary target yield estimates. Numerous studies have attempted to infer the occurrence rate of Earth-like planets in the HZ of different stellar types (e.g., Dressing & Charbonneau 2013; Kopparapu 2013; Foreman-Mackey et al. 2014; Burke et al. 2015), but large uncertainties persist for temperate rocky worlds. We assume an occurrence rate of 20% for G, K, and M dwarfs, which is consistent with published values. Although trends in occurrence rate with stellar type may exist (Mulders et al. 2015), we elect to keep our calculations simple so that they may be easily scaled as occurrence rates are refined. The prior probability of randomly oriented system inclinations is $\sin(i)d_i$ (Luger et al. 2017). This means that 70% of randomly oriented systems will have planets tracing orbital paths that pass through a phase of $135^\circ$ ($i > 45^\circ$). We use the 20% occurrence rate and the 70% random inclination prior to scale our potential stellar targets into an expected number of exoplanets for which oceans may be searched for.

The left y-axis of Figure 9 shows the number of stars for which an assumed optimally aligned planet ($i = 45^\circ–90^\circ$) in the HZ of G (yellow), K (orange), and M (red) dwarfs could be observed at phase angles of $90^\circ$ (left panel) and $135^\circ$ (right panel) as a function of telescope diameter. The right y-axis of Figure 9 shows the expected number of HZ exoplanets that could potentially be searched of oceans using our crescent-phase glint-mapping technique. These observations are assumed to be at $0.8 \mu m$, using a coronagraph with an IWA of $3\lambda/D$. The shaded regions in Figure 9 are bounded below by the inner edge of the HZ and above by the outer edge. The dashed lines correspond to the sum of all partially observable HZs, where a partial HZ is calculated as the radial fraction of the HZ that is observable beyond the IWA. Thus the dashed lines fall between the corresponding shaded region because 100% of the HZ is observable if the inner edge falls outside the IWA, while 0% of the HZ is observable if the outer edge is equal to the IWA.

Figure 9 suggests that K and M dwarfs will make up the bulk of the stellar hosts for surface ocean assessments. Low-mass K and particularly M dwarfs make up the bulk of the nearby stellar population. However, these stars have compact HZs that are difficult to observe in direct imaging. Additionally, the planet-to-star contrast ratio increases toward fainter stars with lower mass. This decreases the noise because of unocculted stellar light (speckles), particularly for systems with contrast ratios that are lower than the design contrast of the direct-imaging mission ($10^{-10}$ assumed here). Figure 9 suggests that K dwarfs strike a balance between these effects for the smaller telescope diameters considered, and may be optimal stellar hosts for direct-imaging studies of HZ planets. G dwarfs have
appreciably separated HZs, so that if the systems are near enough to achieve a sufficient photometric precision, then the IWA will not be a restricting factor for crescent-phase observations. This is in agreement with our earlier findings. M dwarfs occupy a different domain: their favorable planet-to-star contrast ratios more easily satisfy S/N requirements, so that IWA is instead more limiting on both the stellar distance out to which M dwarfs may have their HZs observed and on crescent or gibbous phase observations of any HZ planet.

Considering the IWA and S/N constraints, we find that 1 to nearly 10 HZ planets could be assessed for signs of time-varying ocean glint signals with telescopes in the 6–15 m diameter range. The nearby single stars Tau Ceti, Epsilon Eridani, and Lacaille 8760 (AX Microscopii) are the respective G, K, and M dwarfs that have the largest angular separation to the inner edge of the HZ, making them quite optimistic targets, considering IWA constraints. Telescopes smaller than 12 m may have the potential for crescent-phase observations around more K dwarfs than either G or M dwarfs. Telescopes larger than 12 m may have more M dwarf targets than K dwarf targets. A 15 m telescope with an IWA of $3 \lambda/D$ could perform crescent-phase 0.8 μm direct imaging in the HZ (including the sum of partially observable HZs) of approximately 9 G, 25 K, and 20 M stars. A 12 m telescope could do the same for 5 G, 16 K, and 8 M stars; a 9 m telescope for 2 G, 8 K, and 3 M stars; a 6 m telescope for 1 G, 4 K, and 0 M stars; and a 4 m telescope for 1 G, 1 K, and 0 M stars. Considering random system geometries and assuming an occurrence rate as discussed above, we find that a 15 m, 12 m, 9 m, 6 m, and 4 m telescope will have approximately 8, 4, 2, 1, and 0 expected planet(s), respectively, in the HZ amenable to crescent-phase mapping investigations.

4. Discussion

We investigated the unique time- and phase-dependent observables of Earth oceans in the context of direct habitability assessments for terrestrial exoplanets. In the following subsections we discuss how our findings may be used to increase the robustness of ocean detection on exoplanets (Section 4.1) and assess the potential yield and required observational conditions for ocean searches on extrasolar planets (Section 4.2).

4.1. Increasing the Robustness of Exoplanet Ocean Detection

We find that the robustness of exo-ocean detection can be increased via solutions to the apparent albedo and longitudinal map of multiple surfaces observed at multiple orbital phase angles. We retrieved the longitudinal map and apparent albedo of two unique surfaces from simulated light curves of Earth at quadrature, crescent, and gibbous phase. At quadrature and gibbous phase, one low-albedo surface is mapped to longitudinal slices that we know from Earth to be ocean bearing, while the other more reflective surface is mapped to the continents. At crescent phase, the maps of the dominant surfaces appeared quite similar to the quadrature phase solutions and the albedo of the land-containing slices was nearly the same as at quadrature, but the apparent albedo of the ocean-dominated slices exhibited a large and discernible increase in albedo of $\sim$0.5. This sizable increase in apparent albedo of the ocean-bearing longitudinal slices seen at crescent phase comes from the ocean glint spot. Since our inverse model assumes surfaces are Lambertian scatterers, deviations from Lambertian are encoded in the inferred apparent albedo of the ocean-bearing longitudinal slices seen at crescent phase. This helps to attribute changes in surface reflectivity to specific surface components. The planetary surface that had appeared dark at quadrature must reflect light more efficiently at crescent phase. Such a defined deviation from Lambertian behavior is strongly suggestive of ocean glint.

Although our cloudless-Earth tests demonstrated that glint effects can be extracted from crescent-phase longitudinal mapping, our tests with realistic time-varying clouds revealed limitations and caveats for the technique. Clouds made it much more difficult to extract the same map at both quadrature and crescent phases. Crescent-phase observations are more
sensitive to clouds than more nadir-viewing geometries. Clouds therefore inhibit, but do not fully prevent, accurate surface longitudinal mapping at crescent phase. We were still able to detect the presence of large landmasses, like Africa, which effectively interrupted the glint spot.

However, clouds had minimal impact on the surface albedo results with phase. Despite the added difficulty for accurately detecting the same map, we did confidently detect the increase in albedo of the ocean-bearing slices, along with the lack of albedo increase of the land-bearing slices. Both retrieved surfaces exhibited a slight increase in albedo going from quadrature to crescent phase when clouds were included. Without clouds, Surface 2 (land) did not increase in inferred albedo (0.02 ± 0.02) with phase, while with clouds, it increased by about 0.1 (0.08 ± 0.02), which is a 30% increase. Surface 1 (ocean) without clouds increased in albedo by 0.5 (0.47 ± 0.02) with phase, which is a factor of ~20 increase, while with clouds, it increased by 0.6 (0.57 ± 0.02), which is a factor of ~5 increase with phase. This added increase of 0.1 in the apparent albedo of both surfaces toward crescent is due to forward-scattering by clouds, which we find imparts a significantly smaller phase-dependent effect than ocean glint (see Robinson et al. 2010), and the effect of glint strongly correlates with only one of the surface components. Cowan et al. (2009) also noted that clouds appeared to increase the albedo of the two principle components used in their study, rather than appearing as a unique principle component (or surface). As a result of decomposing the light curves into two surfaces, we see a much stronger effect from glint (~5×) than that noted by Robinson et al. (2010; ~2×), which could make this technique a more sensitive tool for ocean detection than attempting to observe phase-dependent brightness in the disk average.

Longitudinal mapping at crescent phase naturally breaks the latitude-albedo false positive for ocean glint (Cowan et al. 2012) by attributing the global albedo increase observed at crescent phases to a specific surface on the planet. Reflected light observations with high S/N and high cadence at crescent and quadrature (or gibbous) phase enable the longitudinal map and albedo to be retrieved. The time-variability of a rotating planet’s disk-integrated light is used to deconvolve the relative albedo increase seen toward crescent phase into specific surfaces that are responsible for the relative increase. This may increase the confidence that high-latitude ice is not responsible for the relative albedo increase. Furthermore, the boosted strength of the glint signal found after the deconvolution (~5×) exceeds the likely increase in disk-integrated albedo that is expected from the latitude-albedo effect for an Earth-like planet with polar ice caps (~2×; Cowan et al. 2012).

We find that phase-dependent deviations in the longitudinal maps of the cloudless Earth (our ideal case) further implicate non-Lambertian surface-scattering behavior. The difference between the quadrature and crescent-phase maps in Figure 1 can be attributed to the footprint on the planet that contributes the most flux at each phase and therefore drives our inferences. Cowan et al. (2013) refer to this as the kernel of convolution, which has a longitudinal and latitudinal extent determined by the system geometry and phase function of the reflecting surface (Schwartz et al. 2016). At quadrature phase, the planetary surface is well described by a Lambertian surface with a kernel that extends well above and below the equator (for our case here, with a 90° inclination and 0° obliquity). However, at crescent-phase, the ocean deviates from diffuse Lambertian scattering as light reflected off the ocean is dominated by the glint spot. The effective size of the glint spot is significantly more concentrated in both longitude and latitude than the Lambertian kernel of convolution. Consequently, the ocean glint probes a relatively small latitudinal extent. Thus, inferred glint maps are highly sensitive to interruptions (e.g., land, ice, and clouds) along the direct path that the glint spot takes across the surface of the planet. In our
test case of Earth with 90° inclination and 0° obliquity, the dominant latitude and path of the glint spot is the equator.

The effect of kernel size helps to explain why our cloud-free longitudinal maps at quadrature phase (Figure 1) show a sensitivity to continents in the mid-latitudes, such as North America and Asia, which does not appear in the crescent-phase maps. These continents never intersect the path of the latitudinally compact glint spot and are therefore not mapped at crescent phase. Thus, the latitudinal heterogeneity of the planet leads the inferred maps to slightly disagree between quadrature and crescent phase. This does not occur strictly due to the system geometry as explored in Schwartz et al. (2016), but instead because the scattering phase function of the planetary surface(s) affects the width of the convolution kernel. Conversely, the differences found between the quadrature and gibbous phase maps are driven more by the illumination geometry, which favors smoother longitudinal maps at gibbous phase because more illuminated longitudes are integrated together, thus decreasing spatial resolution in the maps (Fujii et al. 2017).

For demonstrative simplicity, we have considered planets with zero obliquity and in an edge-on inclined orbit. Since the sub-observer latitude does not change with phase, this is the optimal case for detecting the same longitudinal map at multiple phases. However, even for this optimistic case with the same latitude dominating the peak reflected flux, we found it difficult to recognize the same map at multiple phases due to latitudinal heterogeneity, as discussed above.

This presents the possibility of false negatives that may arise due to the difficulty of inferring the same longitudinal map at multiple phases. If the observed planet has a prominent north–south continental dichotomy and a nonzero obliquity, like Earth, or a non-edge-on inclination, then time-series observations at multiple phases will constrain the two-dimensional map of the planet (e.g., Kawahara & Fujii 2010, 2011; Fuji & Kawahara 2012; Berdyugina & Kuhn 2017) and the obliquity (Schwartz et al. 2016; Farr et al. 2018). For these cases it will be more difficult to distinguish between changes with phase and changes in surface type. As a result, multiphase longitudinal mapping will perform best for system geometries that are not amenable to two-dimensional spin–orbit mapping. However, exoplanets with nonzero obliquity may still exhibit strong light-curve variations at crescent phase that are best explained by the combination of a Lambertian and a non-Lambertian component. It may be that congruence of the multiphase maps is not necessary to place strong constraints on the detection of glint. Indeed, our finding that the glint signal is found in only one inferred surface and that the glint effect far exceeds forward-scattering by clouds may favor an ocean glint interpretation, albeit with less confidence than if the multiphase maps matched.

Even the detection of glint is not uniquely indicative of surface liquid water. Surface liquid reservoirs of any composition may cause a specular reflection due to their mirror-like smoothness. This is perhaps best exemplified by the glint detection from Titan’s hydrocarbon lakes (Stephan et al. 2010). Freshly formed smooth ice may also deviate from Lambertian scattering (Williams & Gaidos 2008). Furthermore, terrestrial glint has been identified from light scattering off oriented ice crystals in high-altitude cirrus clouds (Kolokolova et al. 2010; Marshak et al. 2017). Such signals could be distinguished from surface glint by their lack of reproducibility on subsequent rotations, or by their anticorrelation with water vapor absorption features in the spectrum that would indicate a cloud-truncated path length. These cases exemplify why additional context is necessary for conclusions to be drawn on the surface liquid composition, such as surface pressure and temperature constraints (Robinson 2018), or the phase-dependent relative change in the surface albedo as discussed in this work.

Because the VPL Earth Model only accurately reproduces Earth, our study only considered Earth’s true geography, rotation rate, and atmospheric composition. Future studies should generalize the technique of glint-mapping with randomly sampled rotation rates and geographies. Furthermore, we assumed that the rotation rate was precisely known. However, a simple periodogram analysis indicates that the time-series data we used for longitudinal mapping with a one-hour cadence and 100 hr duration is more than sufficient to confidently infer the rotation rate of Earth. In general, when the S/N is sufficient for longitudinal mapping to favor two or more surfaces and the observational duration exceeds the rotation rate, then the rotation rate can reliably be retrieved. This agrees with the findings of Pallé et al. (2008) and Fuji & Kawahara (2014). However, the rigorous treatment of directly imaged terrestrial exoplanet rotation rates is beyond the scope of this paper, and we leave it for follow-up work.

Additionally, closer-in planets are more likely to be in tidally locked rotation states, for which our assumption breaks down that the rotational period is much shorter than the orbital period. Further work should investigate the potential for glint detection in reflected light phase-curve measurements of tidally locked exoplanets in the HZs of late-K and M dwarfs. In these cases, the interruption of a glinting ocean by land may manifest itself as a strong asymmetry in the phase curve between the waxing and waning crescent phases.

Using the apparent albedo assumes that the illumination phase angle and planet radius are known, but in practice, both of these quantities will be difficult to precisely constrain for directly imaged exoplanets. This is due to uncertainties in orbit determinations and the radius-albedo-phase degeneracy (Nayak et al. 2017). Planet size cannot be measured unless the planet is transiting, and it is highly unlikely that the few planets amenable to glint-mapping observations have transiting orbits. Alternatively, constraints on the planet radius may be derived for non-transiting planets from a combination of photometric measurements and constraints on the planet albedo, radius, and atmospheric composition determined via spectrum retrieval (e.g., Nayak et al. 2017; Feng et al. 2018). Multi-epoch observations can be used to derive the planetary orbit, either prior to, or using the multi-epoch observations studied in this paper. The radius-albedo degeneracy could potentially be broken by relating the inferred surface composition to its laboratory-measured reflectance spectrum, particularly if phase-dependent glint were to provide a strong indication on the composition of the reflecting surface. Similarly, identifying the slight increase in reflectivity toward crescent for both surfaces could indicate forward-scattering by clouds and may be used to rule out a low-planet albedo, thereby helping to constrain the planet radius. In the event that the planetary radius is unconstrained or incorrectly assumed, multiphase observations may still reveal relative increases in the apparent albedo of multiple surfaces, with the caveat that the apparent
albedo and longitudinal area-covering fraction of each surface are not precisely known.

A number of improvements to the fitting procedures used in this work could be made in the future. We fit the light curves at each phase independently, but multiple epochs can instead be fit simultaneously (e.g., Farr et al. 2018) while allowing for the surface albedo to change with phase. A stronger prior could be enforced that only one of the surfaces is allowed to change albedo as a function of phase. This would turn the physical prior that land does not glint into a model constraint, thereby reducing the number of fitting parameters and helping with cases of lower S/N. Additionally, a Gaussian process could be used to enforce smoothness in the longitudinal maps and allow for the characteristic angular length scale to be inferred (e.g., Farr et al. 2018). This could provide enough flexibility in the mapping procedure to more accurately assess the congruence of maps at multiple phases.

Our work has extended rotational mapping to multiple observed planetary phases to combine specific observables of exoplanetary habitability. Previous studies have demonstrated the utility of multiwavelength light-curve inversions to map directly imaged exoplanets in one-dimension (Cowan et al. 2009; Fujii et al. 2010, 2011; Cowan & Straub 2013) and two-dimensions (Kawahara & Fujii 2010; Fujii & Kawahara 2012; Berdyugin & Kuhn 2017; Farr et al. 2018). While previous two-dimensional exoplanet mapping concepts have leveraged the planet’s obliquity and/or inclination to scan latitudes using observations at multiple orbital epochs (Schwartz et al. 2016), we have instead considered phase-dependent effects for observations at multiple epochs for cases where negligible obliquity and a nearly edge-on planet orbit would prevent two-dimensional mapping. Our crescent-phase simulations of Earth demonstrate a clear increase in reflectively relative to more illuminated phases due to specular reflection off the oceans, as seen in previous studies (Williams & Gaidos 2008; Robinson et al. 2010, 2014). The variability of Earth increases toward crescent phase as the illuminated crescent may be more completely dominated by a single end-member surface spectrum (Oakley & Cash 2009; Robinson et al. 2010). We demonstrated for the first time that crescent-phase variability can be leveraged to map ocean glint.

4.2. Observational Feasibility and Yield Estimates

The observations discussed above will require a telescope that can achieve simultaneous multiwavelength, time-series, direct-imaging observations of an Earth-like exoplanet at crescent phase. This will require a next-generation, high-contrast, space-based direct-imaging mission equipped with a coronagraph or a starshade. For glint-mapping observations to be feasible, light from the planet must fall outside the coronagraph or starshade IWA and at sufficiently high S/N to justify spectral unmixing inverse methods, without overfitting.

We find that the IWA constraints that must be overcome to directly image an Earth-analog at a crescent phase of 135° are less stringent than the constraints on the photometric precision. A 9 m, 12 m, and 15 m space-based coronagraph with an IWA < 3λ/D would be capable of observing an Earth–Sun analog at a 135° phase angle out to ~13, 17, and 21 pc, respectively. However, a 9 m, 12 m, and 15 m telescope with a 20% end-to-end throughput can only reach the requisite ~15% 1 hr photometric precision on multiwavelength light curves of Earth at 135° phase out to ~6, 8, and 10 pc, respectively.

We estimate that there will be potentially habitable exoplanets orbiting nearby stars that will be amenable to glint searches using a next-generation space-based telescope. The IWA, which is less problematic than photometric precision for G dwarfs, becomes a more serious concern for later-type stars where the HZ is more compact. On the other hand, the larger population of nearby K and M dwarfs compared to G dwarfs and the more favorable planet–star contrast ratios for these later-type stars both help to counteract the IWA concern. This allows for the possibility of quadrature and even crescent-phase observations of G, K, and M dwarf HZ planets using coronagraph-equipped telescopes with diameters in the 6–15 m range. Telescopes with diameters approximately ≥8 m may have more K and M dwarf targets than G dwarf targets, depending on how exoplanet occurrence rates scale with stellar type.

Since performing longitudinal mapping at multiple phases requires observations with relatively high S/N, larger mirror sizes and/or improvements in optical throughput and coronagraph spectral bandwidth could increase the potential number of oceans detected. Our derived 1 hr S/N requirement assumed 100 hr of consecutive observations at each phase, but a lower S/N with more exposures or a higher S/N with fewer exposures could still achieve qualitatively similar results. This could potentially open the door to glint-mapping observations for smaller telescopes, such as a 4 m instrument, but would require more than 100 hr of observation at each phase, and might only be able to search for oceans on planets orbiting in the HZ of Alpha Centauri A and/or B because of their proximity.

Although our study specifically modeled coronagraph-equipped telescopes, our results are also applicable to external occultors, such as a starshade. The main differences between coronagraph modeling and starshade modeling for the purposes of this study are the IWA, throughput, and spectral bandwidth. Although starshades can attain λ/D IWA, they are currently being considered for smaller telescopes (<4 m; Seager et al. 2015; Spergel et al. 2015; Mennesson et al. 2016) than those using only an internal coronagraph. Therefore, a 4 m class starshade-enabled telescope with an IWA of λ/D would be comparable to an 8 m telescope with an internal coronagraph and an IWA of 2λ/D. For equivalent IWAs, the higher yield of possible ocean observations will go to larger aperture and/or enhanced throughput, and external occultors more readily achieve higher throughput than internal occultors because they have fewer optical elements (Mennesson et al. 2016). Additionally, starshade technology permits larger spectral bandpasses to be simultaneously observed. This will allow more photons to be collected, effectively increasing the throughput, and will also provide a broader wavelength range over which to identify unique planetary surface albedo components. Figures 5 and 7 can easily be interpreted for a telescope with a starshade by considering IWAs, diameters, and throughputs that are more applicable to such an observatory. However, the multi-epoch approach discussed in this work will require two pointings at the same target, possibly months apart, which is more constrained by the formation-flying operations of the starshade.

Because of contrast considerations, large-aperture ground-based telescopes will be unlikely to observe habitable planets...
around G dwarfs in reflected light, but they may excel for K and M dwarfs. Ground-based 30 m class telescopes have an estimated coronagraph raw contrast of $10^{-7} - 10^{-8}$ (Guyon 2005; Pueyo & Norman 2013), compared with estimates for space-based observatories of $\sim 10^{-10}$. Since Earth–Sun analogs have a flux contrast in the visible of $\sim 10^{-10}$, ground-based observations will face $10^2 - 10^3$ times more noise from an inability to fully block stellar light, which will more than erase gains from mirrors with larger diameters. However, ground-based direct imaging of nearby HZ planets orbiting K and M dwarfs may be optimal for such 30 m class facilities owing to more favorable planet–star contrast ratios, and the larger aperture helps to ameliorate some IWA concerns for these compact systems.

While similar previous studies have focused on broad wavelength coverage (e.g., Cowan et al. 2009), we focused on a narrow wavelength region ($0.68 - 0.8 \mu m$) to remain relevant to current coronagraph design that is unable to effectively control the wavefront over broad wavelength ranges (Pueyo & Norman 2013). As a result of the narrow spectral coverage and the intrinsic faintness of target habitable exoplanets, we simulated multiwavelength time-series photometry by rebinning hourly spectroscopy. This allows for simultaneous multiwavelength time-series without the need to change photometric filters. This also means that the entire duration of the time series may be coadded into one spectrum that contains the disk- and time-averaged spectrum of the planet, which would be an excellent data product for atmospheric retrieval studies (e.g., Feng et al. 2018). Thus, we argue that time-series spectroscopy is more favorable than traditional filter photometry for the time-intensive observations discussed in this paper. Low-noise detector technology will enable this type of time-dependent spectroscopic observation. Microwave kinetic inductance detectors are a promising new technology that directly measures photon energy as they excite the detector, allowing for simultaneous photon counting and low-resolution spectroscopy (Meeker et al. 2018). Such advances would naturally allow for simultaneous multiwavelength time-series observations. Furthermore, the wavelength range that we considered ($0.68 - 0.8 \mu m$) contains the $0.76 \mu m$ O$_2$ A band and the vegetation red-edge—two potential biosignatures—and thus our IWA results are also directly applicable to observing biosignatures on potentially habitable nearby exoplanets. Ultimately, the atmospheric context necessary to more confidently characterize the habitability of an exoplanet target may be contained within the same data set used for the mapping results presented in this paper, but along the spectral rather than temporal dimension.

5. Conclusions

We successfully mapped Earth as an exoplanet using simulated multiwavelength time-series observations at multiple orbital phases, while including ocean glint and forward-scattering by clouds. We confirmed the findings of Cowan et al. (2009) that planet mapping in the specular reflection regime can identify two surface components that correspond to oceans and continents. However, our results extend this work to show that the specularly reflecting glint spot dominates the inferred crescent-phase maps, just as a dark blue surface dominates our quadrature maps for ocean-dominated worlds. Interruptions in the crescent-phase glint signal as Lambertian scattering continents rotate into view leads to a strong blinking effect, which allows for separate longitudinal maps to be constructed of Lambertian and non-Lambertian surfaces. We identified that the partitioning of Lambertian and non-Lambertian surfaces between retrieved surfaces is in stark contrast to the effects of forward-scattering by clouds, which introduces a global apparent albedo increase for all surfaces toward crescent phase. Such multi-epoch studies may be used to detect the strong deviations from Lambertian scattering that oceans exhibit and thereby increase the robustness of ocean detection on directly imaged exoplanets.

We found that recognizing the same longitudinal map at multiple phases will be difficult for a variety of reasons. Clouds introduce uncertainty into the surface maps, nonzero obliquities and inclinations $< 90^\circ$ cause latitudinal sensitivities with phase, and the effective latitudinal extent of the mapped longitudes is a function of phase. However, even cases with distinct longitudinal maps at each phase may exhibit a time-dependent rotationally induced glint signal, which differs significantly from forward-scattering by clouds.

We estimate that space-based high-contrast telescopes with diameters larger than 6 m will have at least one potentially habitable target amenable to glint-mapping via crescent-phase observations, with a 15 m telescope offering nearly 10 targets. Although the technique of glint-mapping discussed here is not applicable to all potentially habitable planets, it may be a key tool for increasing the robustness of habitability assessment for the most promising direct-imaging targets. The discovery of an ocean on an exoplanet will transform exoplanet science by unveiling a truly habitable planet beyond Earth, bringing us one step closer to answering fundamental questions on the prevalence of life in the universe.

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