An experimental and theoretical investigation of HCN production in the Hadean Earth atmosphere

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ABSTRACT

A critical early stage for the origin of life on Earth may have involved the production of hydrogen cyanide (HCN) in a reducing, predominantly H2 atmosphere. HCN is crucial for the origin of life as it is a possible precursor to several biomolecules that make up RNA and proteins including nucleobases, nucleotides, amino acids, and ribose. In this work, we perform an in depth experimental and theoretical investigation of HCN production in reducing atmospheric conditions (89–95% H2) possibly representing the earliest stages of the Hadean eon, ∼4.5–4.3 billion years ago. We make use of cold plasma discharges—a laboratory analog to shortwave UV radiation—to simulate HCN production in the upper layers of the atmosphere for CH4 abundances ranging from 0.1–6.5%. We then combine experimental mass spectrum measurements with our theoretical plasma models to estimate the HCN concentrations produced in our experiments. We find that upper atmospheric HCN production scales linearly with CH4 abundance with the relation [HCN] = 0.13±0.01[CH4]. Concentrations of HCN near the surface of the Hadean Earth are expected to be about 2–3 orders of magnitude lower. The addition of 1% water to our experiments results in a ∼50% reduction in HCN production. We find that four reactions are primarily responsible for HCN production in our experiments: (i) 4N + CH3 → H2CN + H → HCN + H2, (ii) 4N + CH → CN + H followed by CN + CH4 → HCN + CH3, (iii) C2H4 + 4N → HCN + CH3, and (iv) 4N + 3CH2 → HCN + H. The most prebiotically favorable Hadean atmosphere would have been very rich in CH4 (>5%), and as a result of greenhouse effects the surface would be likely very hot. In such a prebiotic scenario, it may have been important to incorporate HCN into organic hazes that could later release biomolecules and precursors into the first ponds.

Keywords: early Earth — astrobiology — atmospheric chemistry — origin of life — HCN production

1. INTRODUCTION

The origin of life on Earth occurred in the Hadean eon, roughly between 4.5 and 3.7 billion years ago (bya) (e.g., Pearce et al. (1); Joyce (2)). At this time, comets and asteroids were impacting the surface of the Earth at a rate 100 thousand to 100 million times greater than today3,4. It is expected that a large fraction of the impactors were similar in composition to enstatite meteorites, and contained a significant amount of iron5. This is based on the similarities of isotopes and relative proportions of lithophile (O, Ca, Ti, and Nd), moderately siderophile (Cr, Ni, and Mo), and highly siderophile (Ru) elements in the Earth’s mantle compared to enstatite meteorites6,5. Furthermore, zircon analyses suggest the Hadean Earth had surface water by 4.4 bya7,8, indicating that warm little pond environments were present for the emergence of life9. The oxidation of this iron by water (i.e. Fe + H2O → FeO + H2) released large amounts of H2 into the atmosphere (e.g. 60–90%10,11), providing reducing conditions favorable for the production of hydrogen cyanide (HCN): a key biomolecule precursor. As the impact rate steadily decreased from 4.5 to 3.7 bya, the hydrogen content of the atmosphere dropped as a result of hydrodynamic escape from the upper atmosphere. By ∼4.3 bya, the bombardment rate was likely too low to maintain H2-dominant conditions10, and an atmosphere dominated by N2/CO2 from outgassing would follow12,13. HCN, and its precursor CH4 are rapidly destroyed in oxidizing (CO2-rich) environments, e.g., via the OH radical10.
The famous Miller-Urey experiment was the first demonstration of abiotic biomolecule production in simulated H$_2$-dominated early Earth conditions. Using electrodes attached to a glass flask filled with reducing gases (H$_2$, CH$_4$, NH$_3$ and H$_2$O), Miller simulated lightning strikes to produce atmospheric HCN. Attached below the flask was a condenser, where this HCN (along with other water soluble species) rained out into a reservoir and reacted in aqueous solution to produce biomolecules. Many have repeated this famous experiment to demonstrate the production of dozens of amino acids as well as the four canonical RNA nucleobases in atmospheres containing various reducing and oxidizing gas mixtures (e.g., Ring et al., Wolman et al., Miller and Schlesinger, Schlesinger and Miller, Miyakawa et al., Cleaves et al., Ferus et al.).

In general, reduced carbon is required as a precursor to produce atmospheric HCN. In contrast, the pathway to HCN from oxidized carbon (i.e., CO$_2$ and CO) is unfavorable, as there is a large energetic barrier to remove oxygen from these species. CH$_4$ is considered the dominant source of reduced carbon in planetary atmospheres, as it is typically the most abundant compared to other reduced carbon species (e.g., C$_2$H$_6$, C$_2$H$_4$). Equilibrium models suggest that early hydrothermal systems on Earth could produce up to ~2.5 ppm of atmospheric CH$_4$. On the other hand, very large impactors similar in size to Vesta (~400 km) have been proposed as a source of ~0.1–10% atmospheric CH$_4$. Abundant CH$_4$ is unlikely to have come from volcanism. However, molecular nitrogen would have been present as a result of volcanic outgassing, supplying nitrogen atoms for HCN.

Several atmospheric photochemical models have identified a link between HCN production and atmospheric CH$_4$ abundance (e.g., Pearce et al.; Zahnle et al.; Rimmer and Rugheimer; Tian et al.; Zahnle). However, a relation between these two species based on models and experiments has not been calculated.

Miller and Schlesinger are the only ones that have performed spark-discharge experiments to investigate the effects of CH$_4$ abundance on HCN production in reducing atmospheres. For CH$_4$ abundances ranging from 17–50%, they found relatively constant ~10% yields of HCN. However, such high atmospheric CH$_4$ concentrations would be unlikely on early Earth, even in typical large impact scenarios.

Considering the importance of HCN as a precursor for biomolecule production, it is of great interest to understand how HCN production scales with CH$_4$ abundance for plausible Hadean Earth atmospheres. A novel and comprehensive approach to exploring this relation is to couple laboratory experiments with theoretical chemical kinetics simulations. The advantages to this approach are: 1) ground-truth experiments provide a means to validate theory, 2) theoretical simulations provide a way to estimate the HCN concentrations produced in experiments, without the difficulty and danger of having to run pure HCN gas standards, and 3) theoretical simulations provide a pathway to discover which reactions are responsible for HCN production in the studied atmospheres.

Atmospheric HCN production requires an energy source to dissociate N$_2$ and CH$_4$ into reactive radicals. For example, $^4$N + CH$_3$ $\rightarrow$ HCN + H$_2$ is one of the dominant reactions for HCN production in Titan’s atmosphere (e.g., Pearce et al.). Recent models suggest that UV radiation was the main energy source for HCN production in the Hadean atmosphere. In contrast, lightning’s contribution to the global HCN concentration was negligible—even when considering volcanic electrical storms. Photodissociation of N$_2$ requires very short UV wavelengths ($\lesssim$ 110 nm) that are outside the spectrum of most UV lamps. For this reason, cold plasma discharges (i.e., electron/ion collisions) can be used as a laboratory analog to excite and dissociate N$_2$ similarly to short wavelength UV radiation (e.g., Cable et al.). We note that HCN and biomolecules can also be produced in high-velocity impacts, and delivered to early Earth via comets.

In this paper, we perform laboratory experiments to simulate the production of atmospheric HCN as function of CH$_4$ abundance on the Hadean Earth. We make use of the Planetary Haze Research (PHAZER) experimental setup, which allows us to apply cold plasma discharges to different gas mixtures and detect changes in mass spectra resulting from HCN production with a Residual Gas Analyzer (RGA). We then perform 0D chemical kinetic calculations using three different chemical networks and compare the resulting theoretical HCN concentrations with the increases to the experimental 26 peak from the plasma experiments. From this we are able to calculate a relation, based on experiment and theory, between HCN production and CH$_4$ abundance for our reducing atmospheric experiments. This relation can also be applied to theoretical HCN production via shortwave UV radiation in the upper atmospheric layers of the Hadean Earth. Lastly, we perform sensitivity analyses on the theoretical simulations to determine which reactions are responsible for HCN production and destruction for each simulated experiment.
2. METHODS

2.1. PHAZER Experiments

In Figure 1, we display a schematic of the PHAZER experimental setup used for our Hadean Earth experiments. PHAZER is a flow system with a stainless steel chamber that is used to simulate the chemistry occurring in planetary atmospheres. Within the chamber is the option to attach two electrodes that produce a cold plasma discharge when a voltage differential is applied. The electrons and ions produced by the AC glow discharge are roughly in the 1.5 to 18.5 eV range, and are energetic enough to directly dissociate all four of our input species (H$_2$, H$_2$O, CH$_4$, and N$_2$). The total gas temperature is not significantly altered by the plasma, and remains at roughly room temperature. Although the Hadean upper atmosphere was likely somewhat cooler than this, we do not adjust the chamber temperature for simplicity. When decreasing the temperature by 50°C in our CRAHCN-O chemical kinetic model of Experiment 1, we see only a ~14 % decrease in HCN production. Such a setup has been used in the past as a laboratory analog for planetary UV and lightning chemistry (e.g., Hörst et al. (35); Ferus et al. (36, 22)). The molecules produced in the cold plasma discharge flow via stainless steel piping to the RGA quadrupole mass spectrometer for identification and measurement.

Our experimental procedure is similar to previous PHAZER studies. First, we prepare the initial gas mixtures using high-purity gases (H$_2$-99.9999%, N$_2$-99.9997%, CH$_4$-99.999%; Airgas) and high performance liquid chromatography-grade water (Fisher Chemical). For the gas mixture with ~1% water, the partial pressure of vapor is roughly controlled by putting a cylinder of water in a dry ice-methanol-water cooling bath maintained in the -42 to -46°C range. We wait at least 2 hours for the water temperature to equilibrate with the cooling bath before turning on the gas flow.

We obtain two background spectra (30-scan averages) of the RGA chamber prior to each experiment and subtract the average of these spectra from the mass spectra measurements of the gas mixtures during each experiment. The RGA is set to the standard 70 eV ionization energy with a scanning mass range of 1–100 AMU.

To begin each experiment, we flow the gas mixture continuously through the system at 5 standard cubic meters per minute. We then obtain two histogram spectra and one analog spectrum of the input gas mixtures prior to turning on the plasma source. All spectra measurements are an average of 30 consecutive scans. We wait 20 minutes after turning on the 170 W/m$^2$ plasma source before measuring the spectra with the RGA in order to ensure the experimental composition reaches a steady state. Then, we obtain 8–9 histogram spectra and one analog spectra with the AC glow discharge turned on, over the course of 4–5 hours.

In Table 1, we display the gas compositions for our five Hadean Earth experiments. These compositions are H$_2$-dominant - in line with the expected reducing conditions produced by the oxidation of iron from enstatite impactors at ~4.5–4.3 bya. We maintain the nitrogen content at 5% for all experiments, which is expected to be present from volcanic outgassing. We also include water at 1% in one experiment to see whether a large atmospheric water abundance affects our results. We note that no unexpected or unusually high safety hazards were encountered in performing these experiments.

**Table 1.** Input molar compositions (in %) for six Hadean Earth experiments. Each dry experiment is run twice and Experiment 6 is run once with water to see how it affects HCN production.

| Gas | Exp. 1 | Exp. 2 | Exp. 3 | Exp. 4 | Exp. 5 | Exp. 6 |
|-----|--------|--------|--------|--------|--------|--------|
| CH$_4$ | 6.5 | 3.6 | 1.5 | 0.5 | 0.1 | 3.6 |
| N$_2$ | 5 | 5 | 5 | 5 | 5 | 5 |
| H$_2$ | 89 | 91 | 94 | 95 | 95 | 90 |
| H$_2$O | 0 | 0 | 0 | 0 | 0 | 1 |

2.2. Chemical Kinetics Models

We run 0D closed system chemical kinetic models at room temperature to simulate our experimental conditions using the KINTECUS software package. We use three different chemical networks that are commonly used to determine HCN chemistry as input into our models to see if the results are network dependent. These networks are: 1) CRAHCN-O, 2) Hébrard et al. (37;38;39;40;41;42;43), and 3) Venot et al. (36, 22)). We add 53 new reactions to CRAHCN-O in order to better estimate the production of hydrocarbons and other potential species of interest such as NO and NH$_3$. We also include 12 reactions from the previous versions of CRAHCN-O that produced species that had no sink reactions in our network and led to erroneous build up of these species (e.g., OH + H$_2$CN + M $\rightarrow$ H$_2$CNOH + M and OH + NH + M $\rightarrow$ HNOH + M). The new reactions added to CRAHCN-O are displayed in Table S1 in the Support.
CRAHCN-O was developed to model HCN and H$_2$CO chemistry in planetary atmospheres dominated by any of H$_2$, N$_2$, CH$_4$, CO$_2$, and H$_2$O. It is a reduced network containing experimental rate coefficients when available, and consistently calculated quantum chemical rate coefficients otherwise.$^{28,46,10}$ CRAHCN-O has been used to model HCN production in Titan’s atmosphere and produces results that agree with the observations of HCN taken by the Cassini spacecraft.$^{28}$ CRAHCN-O has also been used to model HCN and H$_2$CO production in the Hadean Earth atmosphere.$^{10}$

The Hébrard12 network was developed to model HCN and HNC production in Titan’s atmosphere$^{47}$, and also produces results that agree with the observational measurements of HCN by the Cassini spacecraft.$^{28}$ CRAHCN-O has also been used to model HCN and H$_2$CO production in the Hadean Earth atmosphere.$^{10}$

The Venot15 network was developed to study the chemistry of warm carbon-rich exoplanet atmospheres.$^{48}$ It is the largest of the networks used in this study, containing experimental, theoretical, and estimated rate coefficients for 2002 gas phase reactions. The Venot15 network is complete for carbon species up to C $\leq$ 14.

Electron impact dissociation rate coefficients from our plasma source take the form

$$A + e^- \rightarrow B + C + e^-$$

(1)

We use the following standard rate coefficient equation for electron-induced processes$^{49}$ to estimate the dissociation rate coefficients for each of the species for which electron impact dissociation data was available.

$$k = \sqrt{2q/m_e} \int_{1.5eV}^{18.5eV} \epsilon^{1/2} \sigma(\epsilon) f(\epsilon) d\epsilon$$

(2)

where $k$ is the dissociation rate coefficient for a particular electron energy in m$^3$ s$^{-1}$, $q$ is used to convert J to eV ($1.602 \times 10^{-19}$ J/eV), $m_e$ is the mass of an electron ($9.11 \times 10^{-31}$ kg), $\epsilon$ is the electron energy (eV), $\sigma(\epsilon)$ is the electron impact dissociation cross-section as a function of electron energy in m$^{-2}$, and $f(\epsilon)$ is the electron energy distribution function (EEDF) (eV$^{-1}$).

We run simulations using both Maxwellian and Druyvesteyn EEDFs to understand the importance of
Table 2. Summary of the three chemical networks input in our chemical kinetic models.

| Network     | Species | Gas Phase Reactions | Plasma Dissociation Reactions | Source                        |
|-------------|---------|---------------------|-------------------------------|-------------------------------|
| CRAHNC-O    | 49      | 300                 | 17                            | Pearce et al. (28, 46, 10) + this work |
| Hébrard12   | 136     | 788                 | 16                            | Hébrard et al. (47) + this work |
| Venot15     | 238     | 2002                | 16                            | Venot et al. (48) + this work  |

EEDF choice on our chemical kinetic results. If the electrons are in thermodynamic equilibrium with the surrounding plasma, their distribution function is expected to be Maxwellian. Maxwellian EEDFs are conventionally used in plasma theory; however, low-pressure discharges such as those in our experiments are generally non-Maxwellian. At a pressure of 0.3 Torr, one study found ionized plasmas to be Druyvesteyn-like. This is within a factor of 6 of our typical chamber pressure (≈1.8 Torr). These two EEDFs take the form:

\[ f(\epsilon) = c_1 \epsilon^{1/2} e^{-c_2 \epsilon^x} \]  

where \( c_1 \) and \( c_2 \) are defined as:

\[ c_1 = \frac{x}{<\epsilon>^{3/2}} \frac{\Gamma(\frac{5}{2} + x/2)}{\Gamma(\frac{5}{2})} \]

\[ c_2 = \frac{1}{<\epsilon>^x} \frac{\Gamma(\frac{5}{2} + x/2)}{\Gamma(\frac{5}{2})} \]

where \( <\epsilon> \) is the average electron energy (eV), and \( \Gamma \) is the Gamma function.

In Figure 2, we plot the analytical EEDFs, assuming an average electron energy in the center of the range for our plasma source, i.e., 10 eV. Studies of EEDFs from DC glow discharge in H\(_2\)-dominant atmospheres suggest the average negative glow region to be roughly 6 and 8 eV for \( H_2 \) abundances of 70 and 80%, respectively. Given our AC glow discharge experiments are roughly \( H_2 = 90\% \), 10 eV may be a reasonable choice for an average electron energy.

We reduce Equation 2.2 to a first order rate coefficient (s\(^{-1}\)) by multiplying \( k \) by the electron number density \( n \) (m\(^{-3}\)). This is reasonable under the assumption that the plasma discharge maintains a constant density of electrons. We estimate the electron number density assuming that the plasma energy density (170 W m\(^{-2}\)) is primarily electrons, and the average electron energy \( <\epsilon> = 10 \text{ eV} \). The electron number density is calculated using the equation below.

\[ n = \frac{\omega}{<\epsilon> q v_{th}} \]

where \( \omega \) is the plasma energy density (170 W m\(^{-2}\)), and \( v_{th} \) is the most probable thermal velocity of a free electron \( (v_{th} = \sqrt{2k_B T/m_e} = \sqrt{4k_B T/3m_e}) \), where \( k_B \) is the Boltzmann constant and \( m_e \) is the mass of the electron.

In Table 3, we display the calculated dissociation rate coefficients used in our study along with the sources for electron impact dissociation cross-sections \( \sigma(\epsilon) \).

Figure 2. Analytical electron energy distribution functions (EEDFs) for the PHAZER plasma source for the two distributions explored in this work.

3. RESULTS - EXPERIMENTS

In Figure 3, we display the analog RGA mass spectra for Experiments 1–5. Analog plots are valuable for verifying the presence or absence of distinct peaks at each mass-to-charge ratio. Both the initial gas mixtures and the compositions after the plasma source is turned on are displayed.

In all plots, we can see that the peaks corresponding to the mass spectrometry fragment pattern of CH\(_4\), i.e., m/z = 16, 15, 14, 13, decrease after plasma is turned on, suggesting methane has been converted into other carbon species. We also see that the 26 peak, corresponding to HCN or hydrocarbons (e.g., C\(_2\)H\(_6\), C\(_2\)H\(_4\), C\(_2\)H\(_2\)) increases after plasma is turned on, and that the increase correlates with the abundance of CH\(_4\) in the experiment. The 27 peak, which is the main detection
fragment for HCN, is completely washed out by the tail of the intense 28 peak of N₂ and CO. Given the mass spectrometry fragmentation peak for HCN at m/z = 27 is 5.9 times greater than the HCN fragmentation peak at m/z = 27\(^{66}\), we can only detect a HCN peak at m/z = 27 when the tail of the 28 peak from N₂ and CO bleeding into m/z = 27 is < 5.9 times greater than the HCN peak at m/z = 26. In Figure 3, we see that the signal at m/z = 27 from the N₂/CO tail is 7.2, 6.8, 10, 23, and 29 times higher than the signal at m/z = 26 for experiments 1–5, respectively.

However, the 26 peak alone cannot be used to determine HCN production, given that it can also represent hydrocarbon production. The increase in the peaks at 24 and 25 after plasma is turned on suggest that hydrocarbons (e.g., C₂H₂ and C₂H₄) are produced in these experiments. For this reason, we run chemical kinetic models for each experiment to deduce the mostly likely source of the increase in 26 peaks from cold plasma discharge.

In Figures 4 and S1, we plot the histogram RGA mass spectra of Runs 1 & 2 of Experiments 1–5, and Experiment 6 with 1% water. Histogram plots are valuable for performing quantitative analyses.

Similar to the analog plots, we see a clear increase in the 26 peaks after plasma is turned on. The only exception is for the weakest signal in Run 1 of Experiment 5, for which we can consider the 26/28 peak ratio to see the increase. This ratio is the more critical factor for HCN detection in histogram scans, because for small “shoulder” peak increases at m/z = 26 as we see for Experiment 5 in Figure 3, the tail from the 28 peak makes up a significant portion of the signal. Therefore, if the 28 peak decreases after plasma is turned on, then the component of the signal at m/z = 26 that is due to the 28 tail decreases. Considering the 26/28 ratio accounts for this and allows us to measure the weaker signals of Experiment 5.

Because we know that the signals at m/z = 27 are completely due to the 28 tails, it is not surprising to see that the 27 peaks in Figures 4 and S1 decrease after plasma is turned on, along with the decrease in the 28 peaks.

We see increases to the 24 and 25 peaks, or at least the 24/28 and 25/28 ratios, due to the cold plasma discharge in all experiments. These signals are most likely due to hydrocarbon production. C₂H₂, C₂H₄, C₂H₆, and C₂H₈ have fragment peaks for m/z = 24 and 25 of 0.05 and 0.19, 0.02 and 0.08, 0.005 and 0.035, 0.001 and 0.005, respectively\(^{66}\), we might expect C₂H₂ and C₂H₄ to be the main contributors to these peaks. However, again,
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Figure 3. Analog mass spectra for each of the five Hadean Earth experimental compositions, normalized to the 2 (H\textsubscript{2}) peak. Included are scans of the input gas mixtures, as well as the compositions after cold plasma discharge has been turned on. Each mass spectrum is an average of 30 consecutive scans. The 26 peak is circled in each plot as this is the peak with which HCN can be detected in our experiments (via the CN\textsuperscript{+} fragment). Conversely, any HCN in the 27 peak is completely washed out by the tail from the 28 peak of N\textsubscript{2}.

we look to our chemical kinetics models in the following section to deduce the most likely source of these signals.

The reason that coupling chemical kinetics to these experiments is so valuable, is that there are often multiple combinations of molecular species and concentrations that can produce identical signals in the mass spectra. Without chemical kinetics, it is not always possible to determine the correct combination. In the past, authors have demonstrated the use of Monte Carlo methods to obtain a statistical distribution of potential compositions\textsuperscript{67}. Here, we demonstrate the use of plasma modeling with chemical kinetics in order to obtain a single most-likely solution. Experimentally, we find the increase in the 25 peak relative to the 26 peak (normalized to the 28 peak), to range from 0.08:1 to 0.37:1, with an average of (0.16±0.07):1. If we remove the 0.37:1 outlier, corresponding to Experiment 5 Run 1 which had the weakest 25 and 26 peaks, the average
Figure 4. Histogram mass spectra for Run 1 of Experiments 1–5, as well as a single Experiment 6 run with 1% water vapor. Run 2 data can be found in the Supporting Information. The 26 peak is circled in each plot to direct the reader’s eye to the location where HCN can be detected in our experiments. Experiment 5 also has peak 28 circled to emphasize that the 26/28 ratio increases after plasma is turned on.

is (0.14±0.02):1. Given Experiment 5 (0.1% CH₄) is at the detection limit for HCN production using our instruments, we suggest the removal of this outlier may provide a more reasonable estimate. Regardless, the following exercise to show that multiple solutions can produce this ratio works similarly whether the ratio is 0.14:1 or 0.16:1. We can estimate possible solutions for this signal with the equation:

\[
\frac{\sum_{m/z=25} n_i \sigma_i f_i}{\sum_{m/z=26} n_i \sigma_i f_i}
\]

where \(n_i\) is the molar concentration of species \(i\), \(\sigma_i\) is the ionization cross-section for species \(i\) at 70 eV, and \(f_i\) is the fragmentation fraction for species \(i\) at the specified m/z peak. The ionization cross-sections for HCN, C₂H₂, and C₃H₄ at 70 eV are 3.6 Å², 4.4 Å², and 5.1 Å²,
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respectively. The fragmentation fractions for theses species at m/z=25 and m/z=26 are 0 and 0.17, 0.19 and 1, and 0.08 and 0.53, respectively.

Three possible solutions for an increase in the 25 peak relative to the 26 peak that match our experimental average of 0.14 are (in molar percentages HCN/C₂H₂/C₂H₄): 60/20/20, 20/0/80, and 70/30/0. By modeling our experiments using chemical kinetics, we can calculate the expected concentrations for each of theses species and determine the most likely solution.

In the mass spectra of Experiment 6 with water in Figure 4, we can see a small reduction to the 24, 25, and 26 peaks, suggesting water has some effects on the production of HCN and hydrocarbons. The reduction in the 26/28 ratio for Experiment 6 when water is introduced is approximately a factor of 2.

Finally, we see an increase in the 41 peak in approximately half of our experimental runs. The 41 peak increase is often, but not always associated with increases in the 38 and 39 peaks. The associated peak increases can be seen in the analog spectra for Experiments 1–3 in Figure 3 and the histogram spectra for Experiments 2, 3, and 5 in Figure S1. The increases in the 38, 39, and 41 peaks may suggest propane (C₃H₈) or acetonitrile (CH₃CN) was produced. However, given the large range of experimental relative increases for peak 38 to 41 (0.07–0.44:1) and peak 39 to 41 (0.04–5.81), we cannot use the fragmentation patterns for propane or acetonitrile to deduce which product is most likely producing these signals. We again rely on our chemical kinetic models to try to determine which of these species is the most likely source of the 41 peak increases due to the cold plasma discharge in our experiments.

4. RESULTS - CHEMICAL KINETICS

In Figure 5, we display the chemical kinetics models of Experiment 2 (3.6% CH₄) from t = 0–5 seconds using each of the three chemical networks in Table 2, and the Druyvesteyn EEDF. In Figure S2 in the Supporting Information, we display the same three models using the Maxwellian EEDF, and find that the results do not greatly differ. Perhaps the most interesting result of these models is that the concentration and temporal evolution of HCN agrees across all three networks. HCN molar mixing ratios at t = 1 second differ by at most a factor of 4 across all simulations, ranging from 0.37–1.5×10⁻².

The value of t for our experimental setup corresponds to the length of time the gas mixture is in contact with the plasma before it flows out of the chamber and towards the RGA. Given the experimental flow rate of 5 sccm, the ~2 Torr chamber pressure, and the plasma discharge volume of roughly 30 cm³, using p₁v₁ = p₂v₂ we expect t to be roughly one second. At t = 1 second, we see that the second most abundant product ranges from C₂H₄ and C₂H₂ for the CRAHCN-O network simulation, C₂H₆ for the Hébrard12 network simulation and
$C_2H_4$, $C_3H_8$, and $C_2H_2$ for the Venot15 network simulation. The concentrations of these products range from a factor of seven to three orders of magnitude lower than the concentration of HCN.

If we compare the Hébrard network to our experimental results, the dominance of the HCN concentration compared to hydrocarbon species suggests that essentially 100% of the 26 peak would be due to HCN. However, this network solution does not accurately account for the increase in the 25 peak relative to the 26 peak, as described in the previous section.

For the CRAHCN-O and Venot15 networks, we can roughly estimate the contribution from other species to the 26 peak by multiplying the species concentration by the species theoretical mass spectrometry fragment fraction at $m/z = 26$, and the species theoretical ionization cross-section at 70 eV. Then, we can divide this value for each species by the sum of this multiplier for all contributing species. The mass spectrometry fragment fractions for peak 26 are 0.17 for HCN, 1.0 for $C_2H_2$, 0.23 for $C_2H_6$, 0.09 for $C_3H_8$, and 0.53 for $C_2H_4$.66 The ionization cross-sections at 70 eV are 3.6 $\AA^2$ for HCN, 4.4 $\AA^2$ for $C_2H_2$, 6.4 $\AA^2$ for $C_2H_6$, 8.6 $\AA^2$ for $C_3H_8$, and 5.1 $\AA^2$ for $C_2H_4$.6869 Considering these two networks, we calculate that $C_2H_4$, $C_2H_2$, and $C_3H_8$ could contribute up to $\sim$23–47%, 9–15%, and 1–4% to peak 26, respectively.

Depending on the choice of EEDF, HCN could contribute as low as 56% and as high as 60% to the 26 peak increase when using the CRAHCN-O network simulations, and as low as 33% and as high as 65% to the 26 peak increase when using the Venot15 network simulations. To sum up these results, with the exception of the Maxwellian Venot15 simulation, our chemical kinetic models suggest that HCN is the main contributor to the 26 peak increase in our experiments due to the cold plasma discharge. We note again that low-pressure discharges are generally non-Maxwellian; therefore, we expect our Druyvesteyn EEDF results to be slightly more accurate.

The hydrocarbons that are most likely to contribute to the increase in the 24 and 25 peaks when plasma is turned on depends strongly on the chemical network. Similar to our calculations above, using CRAHCN-O network simulations, the experimental fragment patterns66 and ionization cross sections68, we calculate that 69–76% of the 25 peak increase is due to $C_2H_4$ with the remainder coming from $C_2H_2$. For the Venot15 network, the contribution from $C_2H_4$ is still the highest, at 62–70%; however, both $C_2H_2$ and $C_3H_8$ contribute 28–37% and 1–2% to the signal, respectively. Finally, using the Hébrard12 network simulations, we calculate that 61–62% of the 25 peak increase is due to $C_2H_5$, 25–26% is due to $C_2H_2$, 11–12% is due to $C_2H_4$ and 1–2% is due to $C_3H_8$. These calculations are consistent to within 1% when calculating the hydrocarbon contributions to the 24 peak increase for the CRAHCN-O and Venot15 network simulations, and within 14% for the Hébrard12 network simulation. Given that the Venot15 network is the most extensive network for hydrocarbon species, and there is some agreement with the CRAHCN-O network, we suggest that $C_2H_4$ and $C_2H_2$ are most likely the main contributors to the increases in the 24 and 25 peaks due to the cold plasma discharge.

We also aim to shed light on the species that produces an increase to the 41 peak in approximately half of our experiments after plasma is turned on. The two main considerations are $C_3H_8$ and CH$_3$CN. It is worth noting that $C_3H_8$ is not an available product in CRAHCN-O, and CH$_3$CN is not an available product in Venot15; therefore only the Hébrard12 network can be used to compare both concentrations. Given the over an order of magnitude difference between $C_3H_8$ and CH$_3$CN in the Hébrard12 network simulations, and the considerably high $C_3H_8$ mixing ratio for the Venot15 network simulations, we expect that the 41 peak increase due to the cold plasma discharge in our simulations is most likely due to propane rather than acetonitrile.

We find that the main difference in our models when switching from the Maxwellian and Druyvesteyn EEDF is that concentrations decrease, increase and reach steady state more rapidly.

In Figure 6, we plot the HCN abundances from our chemical kinetics models of Experiments 1–5 as well as Experiment 6 which contains water vapor. For this figure, we use the CRAHCN-O network and the Druyvesteyn EEDF. However, we note that differences in HCN abundances at $t = 1$ second are modest across all networks and EEDFs, varying by at most a factor of 4.

The average relation between HCN and CH$_4$ from the theoretical results of Experiments 1–5 at $t = 1$ second are [HCN] = 0.13$\pm$0.01 [CH$_4$]. These data are used for comparison with our experimental results in the following section.

5. RESULTS - EXPERIMENT/THEORY COMPARISON

In Figure 7, we display the 26/28 peak increase from cold plasma discharge, multiplied by the theoretical HCN contribution of 0.6 from our CRAHCN-O model, for our experimental Hadean atmospheres as a function of initial CH$_4$ abundance. Then, we fit the data to a linear function to find the experimental relation between
26/28 peak increase and CH$_4$ abundance. A linear function was chosen because, A) past spark discharge experiments suggested a fairly constant yield for HCN production for CH$_4$ mixing ratios from 17–50%\textsuperscript{15}, which corresponds to a linear relation, and B) a linear relation fit the data reasonably well. The calculated relation is $[(26/28)_{\text{plasma}} - (26/28)_{\text{gas}}] \times 0.6 = 0.02[\text{CH}_4]$.

Next, we place the results of our coupled theoretical models on the same plot, by calibrating to the central point on the linear relation. In other words, we assume that the 26/28 peak increase at [CH$_4$] = 0.015 is due to the theoretical HCN concentration produced in our coupled chemical kinetics model in Figure 6. Each point on the theoretical 26/28 peak increase curve is calculated from the HCN concentration from the corresponding chemical kinetics model. The relation between HCN production and CH$_4$ mixing ratio from these theoretical models—calculated in Section 4—is $[\text{HCN}] = 0.13 \pm 0.01[\text{CH}_4]$.

Next, on the right side of the y-axis, we label the corresponding HCN concentrations for each point on the theoretical 26/28 peak increase curve. In doing this, we have created a relation between our experimental 26/28 peak increases and theoretical HCN production, linked by a conversion factor of 5.37.

The validity of using the theoretical relation for the experimental results can be analyzed by looking at the agreement between the shapes of the two curves. We find good agreement between the theoretical HCN curve and the experimental 26/28 peak increase curve, with a maximum difference of 15% at [CH$_4$] = 0.001. Given all possible calibration points from [CH$_4$] = 0.001–0.065, the maximum difference can increase up to 22%. These results suggest that roughly 13 ± 1% of CH$_4$ is converted into HCN in our experiments.

We emphasize that the relation between HCN production and CH$_4$ mixing ratio derived from our experiments and theory is based on plasma modeling results using the CRAHCN-O network and Druyvesteyn EEDF. We think this network is the most accurate of the three networks for modeling HCN production, as it contains nearly 100 reactions previously undiscovered in the literature that were calculated using consistent quantum chemistry and statistical mechanics methods that have gone through rigorous validations in comparison to experimental data\textsuperscript{45;28;46;10}. This network has also been validated with observational data when modeling HCN production in the atmosphere of Titan\textsuperscript{28}. When using the Hébrard\textsuperscript{12} and Venot networks and the Druyvesteyn EEDF, the HCN relations increase to [HCN] = 0.43 ± 0.06 and 0.41 ± 0.02[CH$_4$], respectively. These yields are quite high compared to past experiments (0.09–0.0135\textsuperscript{15}), which may suggest these networks are less accurate for modeling HCN production and destruction. In the sensitivity analysis in Section 6 we analyze these networks in detail and suggest potential reasons for the inaccuracies.

The largest discrepancy between the two runs of each experiment are for Experiment 5, where there is approximately a factor of 2.5 difference between Runs 1 and 2. For Experiment 5, we made the CH$_4$ mixing ratio as low as possible while still remaining within the detection limit for changes to the 26 peak. Given the m/z = 26 signal was just above the background for these runs, we expect the uncertainty of the Experiment 5 results to be the highest of all experiments.

Lastly, we plot the experiment with 1% water and find that the addition of water leads to a 50% reduction in HCN production. This experimental reduction in HCN due to water is about 34% larger than the theoretical reduction in our chemical kinetic models.

6. RESULTS - SENSITIVITY ANALYSES

In order to understand which reactions are critical for the production and destruction of HCN in our Hadean Earth experiments, we run sensitivity analyses for Experiment 2 and Experiment 6 (which only differ by the absence or presence of 1% water), for each of the three chemical networks employed in this study. The sensitivity analyses involve removing one reaction at a time, running the chemical kinetic simulation, and comparing the difference to the HCN abundance at $t = 1$ second with the reference simulation. If the HCN abundance
The key direct source of HCN varies across network simulations. For CRAHCN-O, the number one direct route to HCN is $^4\text{N} + \text{CH}_3 \rightarrow \text{H}_2\text{CN} + \text{H}$—which proceeds most rapidly to HCN via $\text{H}_2\text{CN} + \text{H} \rightarrow \text{HCN} + \text{H}_2$. The second step of this reaction does not appear in the Table 4 because although it is the dominant subsequent reaction, there are other efficient replacement reactions when it is removed, e.g., $^4\text{N} + \text{H}_3\text{CN} \rightarrow \text{HCN} + \text{NH}$. Both $\text{C}_2\text{H}_4 + ^4\text{N} \rightarrow \text{HCN} + \text{CH}_3$ and $^4\text{N} + ^3\text{CH}_2 \rightarrow \text{HCN} + \text{H}$ are also important contributors to HCN production for this network. For Hébrard12, the most efficient direct route to HCN is also $^4\text{N} + \text{CH}_3 \rightarrow \text{H}_2\text{CN} + \text{H} \rightarrow \text{HCN} + \text{H}_2$. However, $\text{C}_2\text{N} + \text{H} \rightarrow \text{HCN} + \text{C}$ contributes a significant amount to HCN production as well. In the Venot15 network simulation, the dominant direct route to HCN is $^4\text{N} + \text{CH} \rightarrow \text{CN} + \text{H}$—which proceeds to HCN via $\text{CN} + \text{CH}_4 \rightarrow \text{HCN} + \text{CH}_3$. The second dominant source of HCN for this network is $^4\text{N} + ^3\text{CH}_2 \rightarrow \text{HCN} + \text{H}$. We note that the rate coefficient for the dominant HCN production reaction in the Venot15 simulation is approx-
Table 4. The results of the HCN sensitivity analyses for Experiment 2 for the three reaction networks employed in this study, using the Druyvesteyn EEDF for plasma chemistry. The top 10 dominant sources and sinks are included for percent increases/decreases > 3%. In bold are reactions dominant in all three network simulations. In purple are reactions dominant in two of the three network simulations. In orange are reactions that are a dominant source in one network simulation and a dominant sink in another network simulation.

| Network      | Source Reaction         | Percent Dec. | Sink Reaction          | Percent Inc. |
|--------------|-------------------------|--------------|------------------------|--------------|
| CRAHCN-O     | N₂ + e⁻ → 4N + 4N + e⁻  | 100%         | CH₃ + H + M → CH₄ + M  | 303%         |
|              | CH₄ + e⁻ → 3CH₂ + H₂ + e⁻| 50%          | H₂ + e⁻ → H + H + e⁻   | 202%         |
|              | C + H₂ + M → 3CH₂ + M    | 51%          | 4N + NH → N₂ + H       | 119%         |
|              | 4N + CH₃ → H₂CN + H      | 49%          | CH₃ + H₂ + M → CH₃ + M  | 34%          |
|              | NH + H → N + H₂          | 46%          | CH₃ + H₂ + M → CH₃ + M  | 34%          |
|              | CH₄ + e⁻ → CH₃ + H + e⁻   | 35%          | 3CH₂ + H → CH + H₂      | 24%          |
|              | H + H + M → H₂ + M       | 33%          | 3CH₂ + H + M → CH₃ + M  | 5%           |
|              | C₂H₄ + 3N → HCN + CH₃    | 30%          |                         |              |
|              | CH₄ + CH → C₂H₄ + H      | 16%          |                         |              |
|              | 4N + 3CH₂ → HCN + H      | 11%          |                         |              |
| Hébrard12    | N₂ + e⁻ → 4N + 4N + e⁻   | 100%         | CH₃ + H + M → CH₄ + M   | 35%          |
|              | CH₄ + e⁻ → CH₃ + H + e⁻   | 63%          | CH₃ + H + M → CH₄ + M   | 27%          |
|              | 4N + CH₃ → H₂CN + H       | 45%          | CH₃ + H → C + H₂        | 22%          |
|              | H + H + M → H₂ + M       | 23%          | HCN + C₂ → C₃N + H      | 21%          |
|              | C₂H₂ + H + M → C₂H₄ + M   | 16%          | C + C₂H₃ → C₂H₂ + H₂    | 21%          |
|              | C₂N + H → HCN + C        | 13%          | CN + 4N → N₂ + C       | 20%          |
|              | C + H₂ + M → 3CH₂ + M     | 10%          | C + C₂H₄ → C₂H₃ + H    | 14%          |
|              | H + C₂H₅ + M → C₂H₄ + M  | 67%          | 4N + C₃ → CN + C₂      | 14%          |
|              | H₂CN + H → HCN + H₂      | 7%           | H₂ + e⁻ → H + H + e⁻   | 14%          |
|              | H + CH₃C₂H + M → C₂H₅ + M| 6%           | C + C₂H₂ → C₃ + H₂     | 10%          |
| Venot15      | N₂ + e⁻ → 4N + 4N + e⁻   | 100%         | 3CH₂ + H → CH + H₂      | 25%          |
|              | CH₄ + e⁻ → CH₃ + H + e⁻   | 58%          | H₂ + e⁻ → H + H + e⁻   | 18%          |
|              | CH₄ + e⁻ → 3CH₂ + H₂ + e⁻ | 22%          | CH₄ + CH → C₂H₄ + H     | 18%          |
|              | 4N + CH → CN + H         | 25%          | H₂CN + 4N → N₂ + 3CH₂  | 13%          |
|              | C₂H₂ + H → C₂H₃          | 20%          | CN + H₂ → HCN + H      | 8%           |
|              | C₂H₃ + H → C₂H₂ + H₂     | 20%          | CH + H → C + H₂        | 8%           |
|              | CH + C₂H₂ → C₂H₂ + H     | 20%          |                         |              |
|              | 4N + 3CH₂ → HCN + H      | 14%          |                         |              |
|              | CN + CH₄ → HCN + CH₃     | 6%           |                         |              |
|              | CH₃ + C → C₂H₂ + H       | 3%           |                         |              |

approximately a factor of two higher than the value suggested by the average of experiments.70,71,72

Given the lack of experimental studies for the C₂N + H → HCN + C pathway for HCN production in the Hébrard12 network simulation, we analyze this reaction in detail using quantum chemistry. Hébrard et al.47 were the first to estimate a rate coefficient for this reaction for use in the chemical modeling of Titan’s atmosphere. They used capture rate theory to estimate a barrierless upper limit rate coefficient of 2.0×10⁻¹⁰ cm³ s⁻¹. Takahashi and Takayanagi73 studied this reaction with quantum chemistry, and found the pathway to involve four transition states, i.e., C₂N + H → ³HCCN → ³HCCN_isomer → ³HCNC → HCN + C. Their calculations show that transition states 2, 3, and 4 are found to have barriers of 58.2, 4.9, and 66.7 kcal/mol at the B3LYP/6-311++G(d,p) level of theory, respectively. We go one step further to analyze this reaction, by 1) remodeling the transition state chemistry for this reaction using the Gaussian
we conclude that C
hashi and Takayanagi (73) - which can be considered negligible. Furthermore, if mental values always within about an order of magnitude experimental values, and are coefficients calculated using this method most frequently land with experiments show that reaction rate coefficients calculated using this method most frequently land within a factor of two of experimental values, and are always within about an order of magnitude of experimental values. Contrary to the barrierless estimate by Hébrard et al., we find that this reaction proceeds through multiple barriers, consistent with the quantum chemistry results by Takahashi and Takayanagi (73). Our calculated rate coefficient is $2.6 \times 10^{-50}$ cm$^3$ s$^{-1}$ which can be considered negligible. Furthermore, if we instead use the energy barriers calculated by Takahashi and Takayanagi (73) for our calculations, which use a density functional known to underestimate barrier heights, the calculated rate coefficient is still negligible, at $8.2 \times 10^{-45}$ cm$^3$ s$^{-1}$. Given these results, we conclude that $C_2N + H \rightarrow HCN + C$ is not a dominant source of HCN in our Hadean Earth experiments.

Considering the CRAHCN-O and Venot15 networks, there is a discrepancy in the number one direct pathway to HCN. The main reason for this is that the Venot15 network has an extensive set of reactions that build larger hydrocarbons using smaller hydrocarbons (e.g., $C_2H_4$) and methane fragments (e.g., CH$_3$ and $^3$CH$_2$). The competition between this extensive set of reactions and HCN reactions for hydrocarbons and methane fragments shifts the dominance of direct HCN production to cyanide routes. One caveat to using the extensive Venot15 network is that, due to lack of data, we do not include electron impact dissociation reactions for the hundreds of large hydrocarbons in these chemical kinetics simulations. It may be the case that including such data would recycle more CH$_3$ and $^3$CH$_2$ fragments back into the gas mixture, and the methane fragment routes for HCN production would become more dominant as with the CRAHCN-O network simulations. This, however, remains uncertain.

Overall, we see that the four main direct routes to HCN in the dry simulations are $^4N + CH_3 \rightarrow HCN + H$, $^2H_2 + C_2H_4 + ^4N \rightarrow HCN + CH_3$, $^4N + CH \rightarrow CN + H$ followed by $CN + CH_4 \rightarrow HCN + CH_3$, and $^4N + ^3CH_2 \rightarrow HCN + H$. Ironically, $CN + CH_2 \rightarrow HCN + H$ is actually a slight sink reaction in the Venot15 simulation. The reason is because this reaction produces HCN at the expense of producing H atoms, which destroy HCN precursor $^3$CH$_2$. The dominant cyanide route to HCN, i.e., $CN + CH_3 \rightarrow HCN + CH_3$, produces an HCN precursor ($CH_3$) rather than H atoms.

The most critical sink reactions are $CH_3 + H + M \rightarrow CH_4 + M$ and $^3CH_2 + H \rightarrow CH + H_2$. The first sink removes CH$_3$, which is a key reactant for $^4N + CH_3 \rightarrow HCN + H$. The second sink removes $^3$CH$_2$, which is a reactant for $^4N + ^3$CH$_2 \rightarrow HCN + H$. Another key sink across all network simulations is $H_2 + e^- \rightarrow H + H + e^-$. One reason for this is that the source reaction $C + H_2 + M \rightarrow ^3$CH$_2 + M$ requires $H_2$ to make a key precursor to HCN. Another reason is that H atoms efficiently destroy $^3$CH$_2$ via the sink reaction $^3$CH$_2 + H \rightarrow CH + H_2$. Other key sinks in the list are involved in the destruction of precursors to HCN, such as CN, $^4N$, and H$_2$CN. Interestingly, $CH_4 + CH \rightarrow C_2H_4 + H$ is a key sink reaction in the Venot15 simulation, and a key source reaction in the CRAHCN-O simulation. This is due to the differences in the key direct sources of HCN between these simulations. $C_2H_4$ is a key direct source of HCN in the CRAHCN-O simulation via $C_2H_4 + ^4N \rightarrow HCN + CH_3$, whereas CH$_4$ is a direct source of HCN in the Venot15 simulation via $CN + CH_4 \rightarrow HCN + CH_3$.

In Table 5, we display the dominant source and sink reactions for HCN from a sensitivity analysis of Experiment 6 - which contains a 1% molar mixing ratio of water. Labelled in blue are all the reactions that were not previously dominant in the dry sensitivity analyses of Experiment 2 in Table 4.

Only in the CRAHCN-O and Venot15 network simulations are there differences to the dominant source and/or sink reactions with the addition of water.

In the CRAHCN-O network, the three direct sources of HCN remain the same when water is added. For sinks, two main new reactions create an approximate 16% reduction in HCN by removing $^4N$ atoms that would otherwise be used to produce HCN. These reactions are $H_2O + ^4N \rightarrow OH + H + e^-$ followed by $OH + ^4N \rightarrow NO + H$.

There are three new dominant source reactions and four new dominant sink reactions for the Venot15 net-
Table 5. The results of the HCN sensitivity analyses for Experiment 6 for the three reaction networks employed in this study, using the Druyvesteyn EEDF for plasma chemistry. The top 10 dominant sources and sinks are included for percent increases/decreases > 3%. In bold are reactions dominant in all three network simulations. In blue are reactions that are not dominant in the corresponding simulation without water in Table 4.

| Network    | Source Reaction          | Percent Dec. | Sink Reaction          | Percent Inc. |
|------------|--------------------------|--------------|------------------------|--------------|
| CRAHCN-O   | N₂ + e⁻ → ⁴N + ⁴N + e⁻   | 100%         | CH₃ + H + M → CH₄ + M   | 354%         |
|            | CH₄ + e⁻ → ³CH₂ + H₂ + e⁻ | 58%          | H₂ + e⁻ → H + H + e⁻   | 227%         |
|            | ⁴N + CH₃ → H₂CN + H       | 56%          | ³N + H + M → NH + M     | 105%         |
|            | C + H₂ + M → ³CH₂ + M     | 45%          | ⁴N + NH → N₂ + H       | 98%          |
|            | NH + H → ⁴N + H₂          | 45%          | CH + H₂ + M → CH₃ + M   | 41%          |
|            | CH₄ + e⁻ → CH₃ + H + e⁻   | 40%          | ³CH₂ + H → CH + H₂      | 28%          |
|            | H + H + M → H₂ + M        | 36%          | H₂O + e⁻ → OH + H + e⁻  | 17%          |
|            | CH₄ + ³N → HCN + CH₃     | 16%          | OH + ³N → NO + H       | 17%          |
|            | ⁴N + ³CH₂ → HCN + H      | 14%          | ³CH₂ + H + M → CH₃ + M  | 6%           |
|            | CH₄ + CH → C₂H₄ + H      | 13%          |                         |              |
| Hébrard12  | N₂ + e⁻ → ⁴N + ⁴N + e⁻   | 100%         | ³CH₂ + H → CH + H₂      | 35%          |
|            | CH₄ + e⁻ → CH₃ + H + e⁻   | 64%          | CH₃ + H + M → CH₄ + M   | 26%          |
|            | ⁴N + CH₃ → H₂CN + H       | 45%          | CH + H → C + H₂         | 20%          |
|            | H + H + M → H₂ + M        | 22%          | HCN + C₂ → C₃N + H     | 20%          |
|            | C₂H₃ + H + M → C₂H₄ + M   | 14%          | CN + ⁴N → N₂ + C       | 18%          |
|            | C₂N + H → HCN + C        | 13%          | H + C₂H₃ → C₂H₂ + H₂   | 18%          |
|            | C + H₂ + M → ³CH₂ + M     | 9%           | H₂ + e⁻ → H + H + e⁻   | 15%          |
|            | H₂CN + H → HCN + H₂      | 7%           | C + C₂H₄ → C₃H₃ + H    | 13%          |
|            | H + C₂H₅ + M → C₂H₆ + M   | 7%           | ⁴N + C₃ → CN + C₂      | 12%          |
|            | H + CH₃C₂H + M → C₃H₅ + M | 6%           | C + C₂H₂ → C₃ + H₂     | 9%           |
| Venot15    | N₂ + e⁻ → ⁴N + ⁴N + e⁻   | 100%         | ³CH₂ + H → CH + H₂      | 38%          |
|            | CH₄ + e⁻ → CH₃ + H + e⁻   | 47%          | H₂ + e⁻ → H + H + e⁻   | 31%          |
|            | ⁴N + CH → CN + H         | 37%          | H₂CN + ⁴N → N₂ + ³CH₂  | 29%          |
|            | H + H + M → H₂ + M       | 23%          | C + NO → CO + ⁴N       | 20%          |
|            | CH₄ + e⁻ → ³CH₂ + H₂ + e⁻ | 15%          | CH + H → C + H₂        | 18%          |
|            | ⁴N + ³CH₂ → HCN + H      | 13%          | C₃H₃ + ³O → C₂H + H₂CO | 15%          |
|            | H₂CN + H → HCN + H₂      | 8%           | CN + H₂ → HCN + H      | 10%          |
|            | CN + CH₄ → HCN + CH₃     | 7%           | CH₄ + CH → C₂H₄ + H    | 10%          |
|            | ⁴N + CH₃ → HCN + H₂      | 3%           | H₂O + e⁻ → OH + H + e⁻ | 7%           |
|            |                         |             | ⁴N + OH → NO + H       | 4%           |

In summary, when water is introduced to the simulation, the four main direct routes to HCN remain dominant, i.e., ⁴N + CH₃ → H₂CN + H → HCN + H₂, ⁴N + CH → CN + H followed by CN + CH₄ →
HCN + CH₃, C₂H₄ + ^4N → HCN + CH₃, and ^4N + ^3CH₂ → HCN + H.

7. DISCUSSION AND CONCLUSIONS

Several early Earth models have revealed the connection between atmospheric HCN production and CH₄ abundance^10;11;25;26;27. Moreover, experiments simulating lightning-based chemistry in the Hadean Earth atmosphere have found nearly constant 9–13.5% yields for very high (17–50%) atmospheric CH₄ mixing ratios. We are the first to use simulating upper atmospheric UV chemistry for a range of plausible Hadean CH₄ abundances, to find a relation between atmospheric CH₄ content and HCN production at the earliest habitable stages of our planet.

For our experimental setup, we see a relation between HCN production and CH₄ molar mixing ratio that follows the function [HCN] = 0.13 ± 0.01[CH₄]. Our relation, which extends from [CH₄] = 0.001–0.065, agrees well with spark discharge experiments performed by Miller and Schlesinger^15 for [CH₄] = 0.17–0.5, which they find to be [HCN] = 0.09–0.135[CH₄]. How do these experimental relations relate to HCN production from shortwave UV chemistry on the Hadean Earth?

In our previous study Pearce et al.,^10 we performed 1D Hadean Earth atmospheric photochemical simulations for H₂-dominant atmospheres, with methane abundances in the ppm-range. We saw peak HCN production from UV chemistry occurred at altitudes of 500 km and pressures of ~0.0015 Torr. Peak HCN production in these 1D atmospheric models follows the relation [HCN] = 0.73–0.91[CH₄] after 30–50 million years. The main difference between the 1D atmospheric models and our experiments is gas exposure time to UV/plasma. We see in our 0D chemical kinetics models in Figure 6 that HCN has not reached steady state by 1 second. Thus, we expect that longer exposure times such as what would occur in the Hadean atmosphere, would result in higher HCN yields.

HCN produced in the upper Hadean atmosphere is turbulently mixed downwards to the troposphere where it can rain out into favorable environments for the emergence of life such as warm little ponds^10. Early Earth photochemical models suggest the HCN abundances near the surface are approximately 2–3 orders of magnitude lower than the peak abundances produced in the upper atmosphere^10;26;27.

We find that water has a slight impact on HCN production. Experimentally, we see a reduction of about 50% in HCN after adding a molar mixing ratio of ~1% water in Experiment 6. Using the CRAHCN-O network, we see a theoretical 16% reduction in HCN due to the addition of water. The discrepancy is possibly due to missing reactions in our network, such as ion-based reactions. Although we worked hard to discover all the previously unknown neutral reactions related to H₂O, its dissociation products, and HCN in Pearce et al.,^16, we still do not include ion chemistry, and likely several unstudied electron impact dissociation reactions. In any case, a 34% discrepancy is not very large and hints that the CRAHCN-O network does a reasonable job at simulating HCN chemistry in the presence of H₂O. We also note that the upper atmosphere is unlikely to have water vapor mixing ratios as high as 1%; therefore the reduction of HCN production due to the presence of water in the upper Hadean atmosphere is likely much less than 50%.

We find that the key direct sources of HCN are the reactions (A) ^4N + CH₃ → H₂CN + H → HCN + H₂, (B) ^4N + CH → CN + H followed by CN + CH₄ → HCN + CH₃, (C) C₂H₄ + ^4N → HCN + CH₃, and (D) ^4N + ^3CH₂ → HCN + H. The first two reactions are consistent with two of the four dominant sources of HCN in Titan’s atmosphere^28. The dominant sink ^3CH₂ + H → CH + H₂ is also one of the dominant sinks in models of Titan’s atmosphere^28. We find that the addition of water to our simulations does not have a great impact on the dominant sources and sinks for HCN. The main source of the 16% theoretical reduction in HCN is due to the two reactions H₂O + e⁻ → OH + H + e⁻ followed by the ^4N removal reaction, OH + ^4N → NO + H.

Since HCN production scales roughly linearly with CH₄ abundance in the upper Hadean atmosphere, a rich CH₄ (e.g., >5%) reducing atmosphere is the most favorable for prebiotic chemistry. Paradoxically, an atmosphere too rich in CH₄ and H₂ may also be uninhabitable, given the greenhouse effects maintained by H₂–H₂ collisional induced absorption and CH₄ absorption in the infrared^10. The atmosphere requires the depletion of CH₄ to cool, which reduces HCN production. If HCN is best produced in the upper atmosphere when the surface is too hot for liquid water, then how might it persist until the surface cools to habitable temperatures?

One solution might be for HCN to be incorporated into solids during this hot CH₄-rich phase. For example, organic hazes produced from discharge experiments of gas mixtures with 2–5% CH₄ have been shown to contain several biomolecules including all five canonical nucleobases and several proteinogenic amino acids^35;81. Perhaps HCN and biomolecules could have survived in organic hazes while the Hadean atmosphere was hot, and later been incorporated into surface ponds after sufficient cooling. Organic hazes also have a cooling effect.
on the atmosphere, and may increase the cooling rate of the surface. In future experiments, we aim to explore the production and organic composition of organic hazes and exposing them to the hot conditions present in a CH₄-rich Hadean troposphere. Such solids may have been a key source of biomolecule feedstock to warm little ponds on early Earth.

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| Reaction Equation | β | γ | Source(s) |
|-------------------|---|---|-----------|
| C₂H + H₂ →→ C₂H₂ + H | 2.5×10⁻¹¹ | 0 | 1560 | Baulch et al. (83) |
| C₂H + OH →→ C₂H₂ + ³O | 3.0×10⁻¹¹ | 0 | 0 | Tsang and Hampson (84) |
| C₂H + OH →→ CO + ³CH₂ | 3.0×10⁻¹¹ | 0 | 0 | Tsang and Hampson (84) |
| C₂H + HO₂ →→ OH + HCCO | 3.0×10⁻¹¹ | 0 | 0 | Tsang and Hampson (84) |
| C₂H + HCN →→ HC₃N + H | 5.3×10⁻¹² | 0 | 769 | Hoobler and Leone (85), Hébrard et al. (86), Cvetanović (87), Gannon et al. (88), Sayah et al. (89), Westenberg and DeHaas (90), Avramenko and Krasnen'kov (91) |
| C₂H₂ + ³O →→ CO + ³CH₂ | 3.5×10⁻¹² | 1.5 | 850 | Tsang and Hampson (84) |
| C₂H₂ + CN →→ HC₃N + H | 2.3×10⁻¹⁰ | 0 | 0 | Gannon et al. (88) |
| C₂H₂ + CN →→ HCN + C₂H | 2.2×10⁻¹⁰ | 0 | 0 | Sayah et al. (89) |
| C₂H₂ + ³O →→ HOC + CH₃ | 8.9×10⁻¹³ | 1.55 | 216 | Tsang and Hampson (84) |
| C₂H₂ + ³O →→ H₂CO + ³CH₂ | 8.3×10⁻¹² | 0 | 754 | Gannon et al. (88) |
| C₂H₂ + ³N →→ HCN + CH₃ | 2.1×10⁻¹³ | 0 | 754 | Sayah et al. (89) |
| C₂H₂ + OH →→ C₂H₃ + H₂O →→ C₂H₂ + H₂O + H | 1.7×10⁻¹³ | 2.75 | 2100 | Tsang and Hampson (84), Baulch et al. (83) |
| C₂H₂ + CH₃ →→ C₂H₄ + CH₄ →→ C₂H₂ + CH₄ + H | 6.9×10⁻¹² | 0 | 5600 | Gannon et al. (88) |
| C₂H₂ + CN →→ CH₂CHCN + H | 3.2×10⁻¹⁰ | 0 | 0 | Sayah et al. (89) |
| C₂H₂ + CN →→ HCN + C₂H₃ →→ HCN + C₂H₂ + H | 2.1×10⁻¹⁰ | 0 | 0 | Hébrard et al. (86), Balucani et al. (92) |
| C₂H₂ + ³N →→ CH₃CN + H | 2.2×10⁻¹⁰ | 0 | 500 | Baulch et al. (83) |
| C₂H₆ + ³CH₂ →→ C₂H₅ + CH₃ →→ C₂H₄ + CH₃ + H | 1.1×10⁻¹¹ | 0 | 3980 | Baulch et al. (83) |
| C₂H₆ + ³O →→ C₂H₅ + OH →→ C₂H₄ + OH + H | 8.6×10⁻¹² | 1.5 | 2920 | Baulch et al. (83) |
| C₂H₆ + H →→ C₂H₅ + H₂ →→ C₂H₄ + H₂ + H | 1.2×10⁻¹¹ | 1.5 | 3730 | Baulch et al. (83) |
| C₂H₆ + OH →→ C₂H₅ + H₂O →→ C₂H₄ + H₂O + H | 1.1×10⁻¹² | 2.0 | 435 | Baulch et al. (83) |
| C₂H₆ + CH →→ C₂H₄ + CH₃ | 1.3×10⁻¹⁰ | 0 | 0 | Galland et al. (94) |
| C₂H₆ + CH →→ CH₂CHCH₂ + H | 3.0×10⁻¹¹ | 0 | 0 | Galland et al. (94) |
| C₂H₆ + CH₃ →→ CH₄ + C₂H₅ →→ CH₄ + C₂H₄ + H | 1.8×10⁻¹⁶ | 6.0 | 3040 | Baulch et al. (83) |
| C₂H₆ + C₂H →→ C₂H₄ + C₂H₂ →→ C₂H₂ + C₂H₄ + H | 6.0×10⁻¹² | 0 | 0 | Baulch et al. (83) |
| C₂H₆ + CN →→ HCN + C₂H₅ →→ HCN + C₂H₄ + H | 3.5×10⁻¹² | 2.16 | 624 | Baulch et al. (83) |
| C₂H₆ + ³O →→ C₂H₅ + OH →→ C₂H₄ + OH + H | 6.3×10⁻¹⁰ | 0 | 0 | Matsumi et al. (95) |
| C₂H₆ + ³O →→ C₂H₅ + ³O | 7.3×10⁻¹⁰ | 0 | 0 | Fletcher and Husain (96) |

SUPPORTING INFORMATION

In Table S1, we display the 53 chemical reactions added to the CRAHCN-O chemical network for our 0D chemical kinetics models of the Hadean Earth. CRAHCN-O was originally developed to accurately model HCN and H₂CO in atmospheres dominated by any of H₂, CH₄, H₂O, CO₂, and N₂. We added these new reactions in order to better estimate the production of hydrocarbons and a few other potential molecules of interest including NH₃ and NO.
CH$_3$OH + $^4$N $\rightarrow$ CH$_3$ + HNO
CH$_3$OH + OH $\rightarrow$ H$_2$CO + H$_2$O + H
CH$_3$CHO + $^4$N $\rightarrow$ HCN + H$_2$ + HCO
CH$_3$CHO + H $\rightarrow$ CO + H$_2$ + CH$_3$
CH$_3$CHO + H $\rightarrow$ CH$_4$ + HCO
NH$_2$ + $^3$O $\rightarrow$ H + HNO
NH$_2$ + $^3$O $\rightarrow$ OH + NH
NH$_2$ + $^3$O $\rightarrow$ H$_2$ + NO
NH$_2$ + NO $\rightarrow$ N$_2$ + H$_2$O
NH$_2$ + OH $\rightarrow$ H$_2$O + NH
NH$_2$ + OH $\rightarrow$ NH$_3$ + $^3$O
NH$_2$ + HO$_2$ $\rightarrow$ H$_2$O + HNO
NH$_2$ + HO$_2$ $\rightarrow$ NH$_3$ + O$_2$
NH$_2$ + H$_2$ $\rightarrow$ NH$_3$ + H
NH$_2$ + C$_2$H$_6$ $\rightarrow$ NH$_3$ + C$_2$H$_5$ $\rightarrow$ NH$_3$ + C$_2$H$_4$ + H
NH$_2$ + CH$_4$ $\rightarrow$ NH$_3$ + CH$_3$
NH$_3$ + $^3$O $\rightarrow$ NH$_2$ + OH
NH$_3$ + OH $\rightarrow$ NH$_2$ + H$_2$O
NH$_3$ + CN $\rightarrow$ HCN + NH$_2$
HNO + H $\rightarrow$ H$_2$ + NO
HNO + $^3$O $\rightarrow$ OH + NO
HNO + OH $\rightarrow$ H$_2$O + NO
HNO + CN $\rightarrow$ HCN + NO
HNO + HCO $\rightarrow$ H$_2$CO + NO

$4.0 \times 10^{-10}$ 0 4330 Roscoe and Roscoe (98)
$1.1 \times 10^{-12}$ 1.4 56 Srinivasan et al. (99)
$1.0 \times 10^{-14}$ 0 0 Lambert et al. (100)
$5.0 \times 10^{-13}$ 2.75 486 Sivaramakrishnan et al. (101)
$8.8 \times 10^{-14}$ 0 0 Lambert et al. (102)
$7.5 \times 10^{-11}$ 0 0 Cohen and Westberg (103)
$1.2 \times 10^{-11}$ 0 0 Cohen and Westberg (103)
$5.9 \times 10^{-11}$ 2.37 437 Cohen and Westberg (103)
$7.7 \times 10^{-13}$ 1.5 230 Cohen and Westberg (103)
$5.0 \times 10^{-15}$ 2.6 870 Cohen and Westberg (103)
$6.1 \times 10^{-16}$ 0.55 265 Sumathi and Peyerimhoff (105)
$1.9 \times 10^{-16}$ 1.5 1020 Sumathi and Peyerimhoff (105)
$2.1 \times 10^{-12}$ 0 4281 Demissy and Lesclaux (106)
$6.1 \times 10^{-13}$ 0 3600 Demissy and Lesclaux (106)
$7.8 \times 10^{-12}$ 0 4680 Möller and Wagner (107)
$1.6 \times 10^{-11}$ 0 3670 Baulch et al. (83)
$3.5 \times 10^{-12}$ 0 925 Atkinson et al. (108)
$1.5 \times 10^{-11}$ 0 181 Sims and Smith (109)
$3.0 \times 10^{-11}$ 0 500 Tsang and Herron (110)
$3.8 \times 10^{-11}$ 0 0 Inomata and Washida (111)
$8.0 \times 10^{-11}$ 0 500 Tsang and Herron (110)
$3.0 \times 10^{-11}$ 0 0 Tsang and Hampson (84)
$1.0 \times 10^{-12}$ 0 1000 Tsang and Herron (110)

In Figure S1, we plot Run 2 of Experiments 1–5. Run 1 data is displayed in the main text.

In Figure S2, we investigate the chemical kinetics models of Experiment 2 using a Maxwellian EEDF, for the three chemical networks in this study. These results do not greatly differ from the Druyvestyn EEDF results in Figure 5 in the main text.

REFERENCES

[1] Pearce, B. K. D.; Tupper, A. S.; Pudritz, R. E.; Higgs, P. G. Constraining the Time Interval for the Origin of Life on Earth. Astrobiology 2018, 18, 343–364.
[2] Joyce, G. F. The antiquity of RNA-based evolution. Nature 2002, 418, 214–221.
[3] Ferus, M.; Heays, A. N.; Knížek, A. Prebiotic Photochemistry: From Urey–Miller-like Experiments to Recent Findings; The Royal Society of Chemistry, 2021; pp 239–264.
[4] Chyba, C. F. Impact delivery and erosion of planetary oceans in the early inner solar system. Nature 1990, 343, 129–133.
[5] Catling, D. C.; Zahnle, K. J. The Archean Atmosphere. Sci. Adv. 2020, 6, eaax1420.
[6] Dauphas, N. The isotopic nature of the Earth’s accreting material through time. Nature 2017, 541, 521–524.
[7] Mojzsis, S. J.; Harrison, T. M.; Pidgeon, R. T. Oxygen-isotope evidence from ancient zircons for liquid water at the Earth’s surface 4,300Myr ago. Nature 2001, 409, 178–181.
[8] Peck, W. H.; Valley, J. W.; Wilde, S. A.; Graham, C. M. Oxygen isotope ratios and rare earth elements in 3.3 to 4.4 Ga zircons: Ion microprobe evidence for high $\delta^{18}$O continental crust and oceans in the Early Archean. Geochim Cosmochim Acta 2001, 65, 4215–4229.
[9] Damer, B.; Deamer, D. W. The Hot Spring Hypothesis for an Origin of Life. Astrobiology 2020, 20, 429–452.
[10] Pearce, B. K. D.; Molaverdikhani, K.; Pudritz, R. E.; Henning, T.; Cerrillo, K. E. Towards RNA life on Early Earth: From atmospheric HCN to biomolecule production in warm little ponds. Astrophys. J., in press 2022.
Figure S1. Histogram mass spectra for Run 2 of Experiments 1–5. The 26 peak is circled in each plot to direct the reader’s eye to the location where HCN can be detected in our experiments.

[11] Zahnle, K. J.; Lupu, R.; Catling, D. C.; Wogan, N. Creation and Evolution of Impact-generated Reduced Atmospheres of Early Earth. *Planet. Sci. J.* 2020, 1, 11.

[12] Trail, D.; Watson, E. B.; Tailby, N. D. The oxidation state of Hadean magmas and implications for early Earth’s atmosphere. *Nature* 2011, 480, 79–82.

[13] Levine, J. S. *The photochemistry of atmospheres: Earth, the other planets, and comets*; Academic Press, Inc.: Orlando, FL, 1985.

[14] Miller, S. L. A Production of Amino Acids Under Possible Primitive Earth Conditions. *Science* 1953, 117, 528–529.

[15] Miller, S. L.; Schlesinger, A. G. The atmosphere of the primitive earth and the prebiotic synthesis of organic compounds. *Adv. Space Res.* 1983, 3, 47–53.

[16] Bada, J. L. New insights into prebiotic chemistry from Stanley Miller’s spark discharge experiments. *Chem. Soc. Rev.* 2013, 42, 2186–2196.

[17] Ring, D.; Wolman, Y.; Friedmann, N.; Miller, S. L. Prebiotic synthesis of hydrophobic and protein amino acids. *Proc. Natl. Acad. Sci. U.S.A.* 1972, 69, 765–768.
Figure S2. Chemical kinetics simulations of Experiment 2 using three different chemical networks and the Maxwellian EEDF. The grey vertical dotted line represents the time step closest to $t = 1$ second that is used to compare molecule concentrations across simulations.

[18] Wolman, Y.; Haverland, W. J.; Miller, S. L. Nonprotein amino acids from spark discharges and their comparison with the murchison meteorite amino acids. *Proc. Natl. Acad. Sci. U.S.A.* 1972, 69, 809–811.

[19] Schlesinger, A. G.; Miller, S. L. Prebiotic synthesis in atmospheres containing CH$_4$, CO, and CO$_2$. *J. Mol. Evol.* 1983, 19, 376–382.

[20] Miyakawa, S.; Yamanashi, H.; Kobayashi, K.; Cleaves, H. J.; Miller, S. L. Prebiotic synthesis from CO atmospheres: Implications for the origins of life. *Proc. Natl. Acad. Sci. U.S.A.* 2002, 99, 14628–14631.

[21] Cleaves, H. J.; Chalmers, J. H.; Lazcano, A.; Miller, S. L.; Bada, J. L. A reassessment of prebiotic organic synthesis in neutral planetary atmospheres. *Orig. Life Evol. Biosph.* 2008, 38, 105–115.

[22] Ferus, M.; Pietrucci, F.; Saitta, A. M.; Knúžek, A.; Kubelík, P.; Ivanek, O.; Shestivska, V.; Civiš, S. Formation of nucleobases in a Miller–Urey reducing atmosphere. *Proc. Nat. Acad. Sci. U.S.A.* 2017, 114, 4306–4311.

[23] Guzmán-Marmolejo, A.; Segura, A.; Escobar-Briones, E. Abiotic Production of Methane in Terrestrial Planets. *Astrobiology* 2013, 13, 550–559.

[24] N., W.; Krissansen-Totton, J.; Catling, D. C. Abundant Atmospheric Methane from Volcanism on Terrestrial Planets Is Unlikely and Strengthens the Case for Methane as a Biosignature. *Planet. Sci. J.* 2020, 1, 58.

[25] Rimmer, P. B.; Rugheimer, S. Hydrogen cyanide in nitrogen-rich atmospheres of rocky exoplanets. *Icarus* 2019, 329, 124–131.

[26] Tian, F.; Kasting, J. F.; Zahnle, K. Revisiting HCN formation in Earth’s early atmosphere. *Earth Planet. Sci. Lett.* 2011, 308, 417–423.

[27] Zahnle, K. J. Photochemistry of Methane and the Formation of Hydrocyanic Acid (HCN) in the Earth’s Early Atmosphere. *J. Geophys. Res.* 1986, 91, 2819–2834.

[28] Pearce, B. K. D.; Molaverdikhani, K.; Pudritz, R. E.; Henning, T.; Hébrard, E. HCN production in Titan’s Atmosphere: Coupling quantum chemistry and disequilibrium atmospheric modeling. *Astrophys. J.* 2020, 901, 110.

[29] Resonance ltd., 143 Ferndale Drive North, Barrie, Ontario L4N 9V9, www.resonance.on.ca 2022.

[30] Cable, M. L.; Hörl, S. M.; Hodys, R.; Beauchamp, P. M.; Smith, M. A.; Willis, P. A. Titan Tholins: Simulating Titan Organic Chemistry in the Cassini-Huygens Era. *Chem. Rev.* 2012, 112, 1882–1909.

[31] Ferus, M. et al. One-Pot Hydrogen Cyanide-Based Prebiotic Synthesis of Canonical Nucleobases and Glycine Initiated by High-Velocity Impacts on Early Earth. *Astrobiology* 2020, 20, 1476–1488.
HCN production in the Hadean Earth atmosphere

[32] Todd, Z. R.; Öberg, K. I. Cometary Delivery of Hydrogen Cyanide to the Early Earth. Astrobiology 2020, 20, 1109–1120.

[33] Altwegg, K. et al. Prebiotic chemicals–amino acid and phosphorus–in the coma of comet 67P/Churyumov-Gerasimenko. Sci Adv 2016, 2, e1600285–e1600285.

[34] He, C.; Hörst, S. M.; Riemer, S.; Sebree, J. A.; Pauley, N.; Höst, S. M.; Yelle, R. V.; Buch, A.; Carrasco, N.; Ferus, M.; Kubelík, P.; Knížek, A.; Pastorek, A.; Lewis, N. K.; Moses, J. I.; Yu, X.; Höst, S. M.; Lewis, N. K.; Kempton, E. M.-R.; McGuiggan, P.; Marley, M. S.; Kempton, E. M.-R.; Morley, C. V.; Valenti, J. A.; Vuitton, V. Haze Formation in Warm H2-rich Exoplanet Atmospheres. Planet. Sci. J. 2020, 1, 51.

[35] Moran, S. E.; Höst, S. M.; Vuitton, V.; He, C.; Lewis, N. K.; Flandinet, L.; Moses, J. I.; North, N.; Orthous-Daunay, F.-R.; Sebree, J.; Wolters, C.; Kempton, E. M.-R.; Morley, M. S.; Morley, C. V.; Valenti, J. A. Chemistry of Temperate Super-Earth and Mini-Neptune Atmospheric Hazes from Laboratory Experiments. Planet. Sci. J. 2020, 1, 17.

[36] Ianni, J. C. Kintecus, Linux Version 6.80, www.kintecus.com 2019.

[37] Pearce, B. K. D.; Ayers, P. W.; Pudritz, R. E. A Consistent Reduced Network for HCN Chemistry in Early Earth and Titan Atmospheres: Quantum Calculations of Reaction Rate Coefficients. J. Phys. Chem. A 2019, 123, 1861–1873.

[38] Pearce, B. K. D.; Ayers, P. W.; Pudritz, R. E. CRAHCN-O: A consistent reduced atmospheric hybrid chemical network oxygen extension for hydrogen cyanide and formaldehyde chemistry in CO2-, N2-, H2O-, CH4-, and H2-dominated atmospheres. J. Phys. Chem A. 2020, 124, 8594–8606.

[39] Hébrard, E.; Dobrijevic, M.; Loison, J. C.; Bergeat, A.; Hickson, K. M. Neutral production of hydrogen isocyanide (HNC) and hydrogen cyanide (HCN) in Titan’s upper atmosphere. Astron. Astrophys. 2012, 541, A21.

[40] Venot, O.; Hébrard, E.; Agúndez, M.; Decin, L.; Bounceur, R. New chemical scheme for studying carbon-rich exoplanet atmospheres. Astron. Astrophys. 2015, 577, A33.

[41] Alves, L. L.; Bogaerts, A.; Guerra, V.; Turner, M. M. Foundations of modelling of nonequilibrium low-temperature plasmas. Plasma Sources Sci. Technol. 2018, 27, 023002.

[42] Godyak, V. A.; Piejak, R. B.; Alexandrovich, B. M. Probe diagnostics of non-Maxwellian plasmas. J. Appl. Phys. 1993, 73, 3657–3664.

[43] Gudmundsson, J. T. On the effect of the electron energy distribution on the plasma parameters of an argon discharge: a global (volume-averaged) model study. Plasma Sources Sci. Technol. 2001, 10, 76–81.

[44] Khalilpour, H.; Foroutan, G. The effects of electron energy distribution function on the plasma sheath structure in the presence of charged nanoparticles. J. Plasma Phys. 2020, 86, 905860206.
Yoon, J.-S.; Song, M.-Y.; Han, J.-M.; Hwang, S. H.; Itikawa, Y.; Mason, N. Cross Sections for Electron Collisions with Hydrogen Molecules. *At. Data Nucl. Data Tables* 2006, 92, 375–406.

Tabata, T.; Shirai, T.; Sataka, M.; Kubo, H. Analytic cross sections for electron impact collisions with nitrogen molecules. *At. Data Nucl. Data Tables* 2008, 37, 913–931.

Chakrabarti, K.; Laporta, V.; Tennyson, J. Calculated cross sections for Electron Collisions with Hydrogen Molecules. *J. Phys. Chem. Ref. Data* 2008, 37, 913–931.

Cosby, P. Electron-impact dissociation of oxygen. *J. Chem. Phys.* 1993, 98, 9560–9569.

Shirai, T.; Tabata, T.; Tawara, H. Analytic cross sections for electron collisions with CO, CO₂, and H₂O relevant to edge plasma impurities. *At. Data Nucl. Data Tables* 2001, 79, 143–184.

Chakrabarti, K.; Laporta, V.; Tennyson, J. Calculated cross sections for low energy electron collision with OH. *Plasma Sources Sci. Technol.* 2019, 28, 045013.

Song, M.-Y.; Yoon, J.-S.; Cho, H.; Karwasz, G. P.; Kokoouline, V.; Nakamura, Y.; Tennyson, J. Cross Sections for Electron Collisions with NO, N₂O, and NO₂. *J. Phys. Chem. Ref. Data* 2019, 48, 043104.

Shirai, T.; Tabata, T.; Tawara, H.; Itikawa, Y. Analytic cross sections for electron collisions with hydrocarbons: CH₄, C₂H₆, C₂H₄, C₂H₂, C₃H₄, and C₅H₆. *At. Data Nucl. Data Tables* 2002, 80, 147–204.

May, O.; Kubala, D.; Allan, M. Absolute cross sections for dissociative electron attachment to HCN and DCN. *Phys. Rev. A* 2010, 82, 010701.

Song, M.-Y.; Yoon, J.-S.; Cho, H.; Karwasz, G. P.; Kokoouline, V.; Nakamura, Y.; Tennyson, J. Cross Sections for Electron Collisions with Acetylene. *J. Phys. Chem. Ref. Data* 2017, 46, 013106.

Saiber, W.; P elic, A.; Limão-Vieira, P.; Mason, N. J.; Limtrakul, J.; Scheier, P.; Probst, M.; Märk, T. D. Low energy electron attachment to CH₃CN. *Chem. Phys. Lett.* 2003, 381, 216–222.

Wallace, W. E. In *NIST Chemistry WebBook, NIST Standard Reference Database Number 69*; Linstrom, P. J., Mallard, W. G., Eds.; National Institute of Standards and Technology: Gaithersburg MD, 20899, 2021.

Gautier, T.; Serigano, J.; Bourgalais, J.; Hörst, S. M.; Trainer, M. G. Decomposition of electron ionization mass spectra for space application using a Monte-Carlo approach. *Rapid Commun. Mass Spectrom.* 2020, 34.

Kim, Y.-K.; Irikura, K. K.; Rudd, M. E.; Ali, M. A.; Stone, P. M.; Chang, J.; Coursey, J. S.; Dragoset, R. A.; Kishore, A. R.; Olsen, K. J.; Sansonetti, A. M.; Wiersma, G. G.; Zucker, D. S.; Zucker, M. A. *NIST Standard Reference Database 107*; National Institute of Standards and Technology: Gaithersburg MD, 20899, 2004.

Pandya, S. H.; Shelat, F. A.; Joshipura, K. N.; Vaishnav, B. G. Electron ionization of exotic molecular targets CN, C₂N₂, HCN, HNC and BF-Theoretical cross sections. *Int. J. Mass. Spectrom.* 2012, 323-324, 28–33.

Daranlot, J.; Hu, X.; Xie, C.; Loison, J.-C.; Caubet, P.; Costes, M.; Wakelam, V.; Xie, D.; Guo, H.; Hickson, K. Low temperature rate constants for the N(³S) + CH(X²Π) reaction. Implications for N₂ formation cycles in dense interstellar clouds. *Phys. Chem. Chem. Phys.* 2013, 15, 13888–13896.

Brownsword, R. A.; Gatensby, S. D.; Herbert, L. B.; Smith, I. W. M.; Stewart, D. W. A.; Symonds, A. C. Kinetics of reactions between neutral free radicals. Rate constants for the reaction of CH radicals with N atoms between 216 and 584 K. *J. Chem. Soc. Faraday Trans.* 1996, 92, 723–727.

Messing, I.; Filseth, S. V.; Sadowski, C. M.; Carrington, D. Absolute rate constants for the reactions of CH with O and N atoms. *J. Chem. Phys.* 1981, 74, 3874–3881.

Takahashi, K.; Takayanagi, T. Theoretical study on HNC(³Σ⁺) production from C₂(X³I₄) + NH(X²Σ⁺) reaction. *Chem. Phys. Lett.* 2006, 429, 399–404.

Frisch, M. J.; Trucks, G. W.; Schlegel, H. B.; Scuseria, G. E.; Robb, M. A.; Cheeseman, J. R.; Scalmani, G.; Barone, V.; Petersons, G. A.; Nakatsuji, H.; et al., 2009; *Gaussian 09*, Revision E01; Gaussian, Inc., Wallingford, CT, 2009.

Truhlar, D. G.; Garrett, B. C. Variational Transition State Theory. *Annu. Rev. Phys. Chem.* 1984, 35, 159–189.

Becke, A. D. A new mixing of Hartree-Fock and local density-functional theories. *J. Chem. Phys.* 1993, 98, 1372–1377.

Lee, C.; Yang, W.; Parr, R. G. Development of the Colle-Salvetti correlation-energy formula into a functional of the electron density. *Phys. Rev. B* 1988, 37, 785–789.
[100] Lambert, R. M.; Christie, M. I.; Golesworthy, R. C.; Linnett, J. W. Mass spectrometric study of the reaction of nitrogen atoms with acetaldehyde. *Proc. R. Soc. London A* 1968, 302, 167–183.

[101] Sivaramakrishnan, R.; Michael, J. V.; Klippenstein, S. J. Direct Observation of Roaming Radicals in the Thermal Decomposition of Acetaldehyde. *J. Phys. Chem. A* 2010, 114, 755–764.

[102] Lambert, R. M.; Christie, M. I.; Linnett, J. W. A novel reaction of hydrogen atoms. *Chem. Commun. (London)* 1967, 388–389.

[103] Cohen, N.; Westberg, K. R. Chemical Kinetic Data Sheets for High-Temperature Reactions. Part II. *J. Phys. Chem. Ref. Data* 1991, 20, 1211–1311.

[104] Song, S.; Hanson, R. K.; Bowman, C. T.; Golden, D. M. Shock Tube Determination of the Overall Rate of NH$_2$ + NO $\rightarrow$ Products in the Thermal De-NOx Temperature Window. *Int. J. Chem. Kinet.* 2001, 33, 715–721.

[105] Sumathi, R.; Peyerimhoff, S. D. A quantum statistical analysis of the rate constant for the HO$_2$ + NH$_2$ reaction. *Chem. Phys. Lett.* 1996, 263, 742–748.

[106] Demissy, M.; Lesclaux, R. Kinetics of Hydrogen Abstraction by NH$_2$ Radicals from Alkanes in the Gas Phase. A Flash Photolysis-Laser Resonance Absorption Study. *J. Am. Chem. Soc.* 1980, 102, 2897–2902.

[107] Möller, W.; Wagner, H. G. Messung der geschwindigkeitskonstanten der reaktion NH$_2$ + CH$_4$ $\rightarrow$ NH$_3$ + CH$_3$ hinter einfallenden stosswellen. *Z. Naturforsch. A* 1984, 39, 846–852.

[108] Atkinson, R.; Baulch, D. L.; Cox, R. A.; Crowley, J. N.; Hampson, R. F.; Hynes, R. G.; Jenkin, M. E.; Rossi, M. J.; Troe, J. Evaluated kinetic and photochemical data for atmospheric chemistry: Volume I - gas phase reactions of Ox, HOx, NOx and SOx species. *Atmos. Chem. Phys.* 2004, 4, 1461–1738.

[109] Sims, I. R.; Smith, I. W. M. Pulsed laser photolysis-laser-induced fluorescence measurements on the kinetics of CN(v = 0) and CN(v = 1) with O$_2$, NH$_3$ and NO between 294 and 761 K. *J. Chem. Soc. Faraday Trans. 2* 1988, 84, 527–539.

[110] Tsang, W.; Herron, J. T. Chemical kinetic data base for propellant combustion. I. Reactions involving NO, NO$_2$, HNO, HNO$_2$, HCN and N$_2$O. *J. Phys. Chem. Ref. Data* 1991, 20, 609–663.

[111] Inomata, S.; Washida, N. Rate Constants for the Reactions of NH$_2$ and HNO with Atomic Oxygen at Temperatures Between 242 and 473 K. *J. Phys. Chem. A* 1999, 103, 5023–5031.