Application of Thermal Remote Sensing technique for mapping of ultramafic, carbonate and siliceous rocks using ASTER data in Southern Rajasthan, India

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In the present study, thermal remote sensing technique and ASTER data have been used to delineate ultramafic, carbonate and siliceous rocks. The study gains importance as mineralized carbonate and ultramafic rocks are present in the southern region of Rajasthan, India between Udaipur and Dungarpur districts. The rocks in the study area include phyllites, mica schist, chlorite schist, quartzite, dolomite, granite, granitoids, gneiss and intrusive serpentinite. ASTER thermal bands were used to map ultramafics, siliceous and carbonate rocks on a scale of 1 : 380,000. Delineation of ultramafics was done using MRI-AV and MI-N indices, however the former provided a more informative map compared to the latter. QRI-AV, QI-N and QI-RH indices were used for mapping siliceous rock. QI-RH provided a more informative map compared to QRI-AV and QI-N. The index used for carbonate rocks was CI-N, but this did not provide a satisfactory map.

Keywords: ASTER TIR, carbonate and siliceous rocks, thermal indices, thermal remote sensing, ultramafics.

MAJORITY of oxides–hydroxides, clay and carbonate minerals reflect spectral absorption in the visible near infrared (VNIR) and shortwave infrared (SWIR) regions1,2, whereas major rock-forming minerals such as quartz (silicates) and orthoclase/plagioclase (feldspar) do not represent any kind of spectral absorption features. The thermal infrared (TIR) region is beneficial for the discrimination of such kinds of minerals based on variances in the emissivity spectra. Atomic and molecular vibrations are the main factors responsible for changes in the emissivity spectra of different minerals. In the TIR spectral domain, vibrational spectroscopy provides ample amount of information about minerals due to variation in emissivity, because minerals have a different chemical composition and molecular structure2,3. Majority of the silicate minerals show variation in vibrational energy pattern of the Si–O bond4,5, which reflects variation in the emissivity spectra of these minerals and is responsible for shifting of emissivity minima of silicate-bearing rocks in TIR spectra6,5,6. TIR spectroscopy is useful for mapping of geological materials which exhibit vibrational absorption in the spectral range 3–50 μm (refs 2, 3 and 5) like silicate as well as carbonate-rich rocks with diagnostic emissivity features in the TIR region5,6,5–11. TIR spectroscopy is less common compared to reflectance spectroscopy to map and delineate various kinds of rocks and minerals. This is due to the fact that very few hyperspectral (spectral observation in contiguous channels) and multispectral space-borne sensors are operative in the TIR domain. This is, in turn, related to poor signal-to-noise (SNR) ratio of data recorded by the thermal channels within the TIR domain12. This is another hindrance for using contiguous spectral bands with finer spectral resolution to detect subtle variations in thermal emissivity within the TIR wavelength range.

The present study aims to evaluate the thermal remote sensing technique for identification and demarcation of ultramafic, carbonate and siliceous rocks in the study area using different thermal mineral indices by analysing ASTER data of Terra satellite. It also aims to compare better suitability of the indices for the region. Different varieties of rock like quartzite, carbonate, phyllite, schist and serpentinite are present in the region. So it is assumed that thermal spectroscopy will provide a better understanding of the distribution pattern in this region.

Geology

The study area is in the southern region of Rajasthan, extending between Udaipur and Dungarpur districts (Figure 1). The litho-units consists of the Archaean basement rocks of the Banded Gneissic Complex (BGC), Proterozoic supracrustals of the Aravalli Supergroup and intrusives13–17 (Table 1). The BGC comprises of Sarara Inlier and Kherwara Inlier granitoids and the Aravalli Supergroup comprises of conglomerate, dolomite, quartzite, mica schist and phyllite rocks. Intrusives are mainly
Figure 1. Map showing location of the study area in Rajasthan, India, and imagery extended from south of Udaipur to Dungarpur.

Table 1. Stratigraphic distribution of geological litho-units of the study area (modified after Gupta et al.13)

| Age       | Supergroup                  | Group                  | Lithology                                                                 |
|-----------|-----------------------------|------------------------|---------------------------------------------------------------------------|
| Proterozoic| Udaipur granites and gneisses (intrusion) | Granites, Gneiss        | Serpentinites                                                             |
|           | Rakhabdev Ultramafic Suite (intrusion) |                        |                                                                            |
| Aravalli  | Udaipur                     |                        | Phyllite, slates with interbands of quartzite, dolomite and meta-dolerite, quartzite and variant facies of dolomite, pebbly greywacke, phyllite, carbonaceous phyllite with minor intercalations of quartzite and dolomite |
|           | Debari                       |                        | Dolomite, mica schist and phyllite with subordinate calcareous and carbonaceous rocks, metaconglomerate, meta-arkose and quartzite, basic volcanic–sedimentary rocks (hornblende–actinolite schist and chlorite-actinolite schist), gritty and conglomeratic quartzite |
| Archean   | Bhilwara                    |                        | Banded gneissic rocks, migmatites and schist                                |

The study area is known for lead–zinc mineralization hosted by dolomites of the Middle Aravalli Group15.

Materials and method

Advanced spaceborne thermal emission and reflection radiometer (ASTER) launched on-board by the earth observation satellite (EOS), Terra, is a multi-spectral sensor system18,19. ASTER has three bands in the VNIR domain, six bands in the SWIR domain and five bands in the TIR domain (band nos 10–12). ASTER data have been extensively used for mapping few significant types of surface mineralization signatures associated with hydrothermal alteration zones and oxidation capping of supergene enrichment deposits based on the diagnostic absorption features of different clay, carbonate, aluminium–iron hydroxide minerals in the VNIR–SWIR domain11,18–20.
Figure 2. Overall methodology adopted to process the ASTER-TIR data for preparation of different index maps.

Table 2. Technical specifications of the ASTER L1T data

| Granule ID  | Sensor    | Band number | Spectral width (μm) | Spatial resolution (m) | Radiometric resolution (bits) | Valid range |
|------------|-----------|-------------|---------------------|------------------------|-------------------------------|-------------|
| AST_L1T_0  | ASTER-TIR | 10          | 8.125–8.475         | 90                     | 12                            | 0–65535     |
| 0304222003 | 11        | 8.475–8.825 | 90                  | 12                     | 0–65535                       |
| 055021_201 | 12        | 8.925–9.275 | 90                  | 12                     | 0–65535                       |
| 5042803151 | 13        | 10.25–10.95 | 90                  | 12                     | 0–65535                       |
| 0_40583    | 14        | 10.95–11.65 | 90                  | 12                     | 0–65535                       |

The VNIR–SWIR bands of ASTER were also used for delineating few economically important rock types with spectrally sensitive minerals.\(^6,^{10,11,18}\)

ASTER L1T data used in this study were downloaded from the United States Geological Survey (USGS) Earth–Explorer portal (https://earthexplorer.usgs.gov/). L1T are radiometrically and terrain corrected data (https://lpdaac.usgs.gov/documents/71/AST_L1T_User_Guide_V3.pdf) and 16-bit integer data are present in the thermal bands (bands 10–14) at 90 m spatial resolution. Table 2 provides more technical specifications. Band ratio and relative band depth formulation were also applied on radiance data to identify the distribution of various lithounits (rocks) in the region.\(^3,^{6,8,10,21–26}\) Figure 2 illustrates the methodology adopted.

Each pixel of the L1T data carries a digital number (DN), which is converted into calibrated radiance, i.e. \(L_{\text{sen}}\) using eq. (1)\(^{24,27–29}\)

\[
L_{\text{sen}} = \text{cof}^i \times (\text{DN}^i - 1),
\]

where

\[
\begin{align*}
\text{cof}^{10} & = 0.006882, \\
\text{cof}^{11} & = 0.006780, \\
\text{cof}^{12} & = 0.006590, \\
\text{cof}^{13} & = 0.005693, \\
\text{cof}^{14} & = 0.005224.
\end{align*}
\]

Guha and Vinod Kumar\(^6\) (eq. (2)), Rockwell and Hofstra\(^10\) (eq. (3)), and Ninomiya et al.\(^7\) (eq. (4)) have given the indices for the mapping of silicate rocks. These three different thermal indices have been applied on radiance data for the identification and mapping of silicate rocks.

Quartz-bearing rock index (QRI-AV)

\[
\text{QRI-AV} = (\text{Band 10}/\text{band 12}) \times (\text{band 13}/\text{band 12}),
\]

Quartz index (QI-RH) \(= (\text{Band 11}/(\text{band 10 + band 12})) \times (\text{band 13}/\text{band 12})\)

Quartz index (QI-N) \(= (\text{Band 11} \times \text{band 11})/(\text{band 10} \times \text{band 12})\)
Guha and Vinod Kumar\(^6\) and Ninomiya \(et\ al.\)\(^7\) have also given indices for mapping of mafic minerals (eq. (5) and (6) respectively); these have been applied to map the ultramafics.

**Mafic-bearing rock index (MRI-AV)**

\[ \text{MRI-AV} = \left( \frac{\text{Band 12}}{\text{band 13}} \right) \times \left( \frac{\text{band 14}}{\text{band 13}} \right), \]

(5)

**Mafic index (MI-N)**

\[ \text{MI-N} = \frac{(\text{band 12} \times (\text{band 14}^3))}{(\text{band 13}^4)}. \]

(6)

Ninomiya \(et\ al.\)\(^7\), and Rockwell and Hofstra\(^30\) have given an equation for mapping of carbonate rocks (eq. (7)). This equation has been applied to determine the availability and for mapping of carbonate rocks in the region without applying any kind of correction, i.e. atmospheric correction on the radiance data\(^7\).

**Carbonate index (CI-N)**

\[ \text{CI-N} = \left( \frac{\text{Band 13}}{\text{band 14}} \right). \]

(7)

**Results and discussion**

The ultramafic map was developed using mafic indices of Guha and Vinod Kumar\(^6\) and Ninomiya \(et\ al.\)\(^7\). MRI-AV mapped the ultramafic rock outcrop with very high values in Rakhabdev region. However, in the southern zone, these indices also defined a water body as mafic outcrop on the upper right side of the area (Figure 3\(a\)). Another minor parallel belt of the ultramafic outcrop was also identified as a major rock belt in the map. MI-N mapped the ultramafic outcrop with high values confined to Rakhabdev area, while at a few places it also mapped the ultramafic outcrops with lower values (Figure 3\(b\)). Minor parallel belts of the ultramafics were not identified using this index. Ultramafic rock outcrop is more precisely mapped by MRI-AV compared to MI-N.

Quartz/siliceous rock map was developed using QRI-AV, QI-RH and QI-N. All the three indices gave diverse kinds of pattern. The QRI-AV index defined siliceous rock outcrops in the region. Siliceous outcrops were marked in the outer periphery of the granitoid gneissic terrain of the Sarara Inlier. Another siliceous outcrop was identified as intrusive granitic body near Rakhabdev. This index suppresses the major mafic and carbonate rocks with certainty but low values are present in the entire region, which is a major drawback of the index (Figure 4\(a\)). QI-N index produced noisy imagery. It demarked the presence of ample siliceous rocks in the region, which is in all probability a wrong interpretation regarding the abundance of these rocks (Figure 4\(b\)). The QI-RH index demarked the siliceous rocks in the region quite precisely. In Figure 4\(b\), siliceous rocks are mapped near the granitoid rock outcrops as well as close to the intrusive granite outcrops in the right lower region. This pattern of siliceous outcrops is more precise and correlatable to the published geological map of Gupta \(et\ al.\)\(^13\). QI-RH decreased the values of carbonate and ultramafic outcrops and is useful to determine the distribution of siliceous outcrops (Figure 4\(c\)).

The CI-N technique used in Figure 5 to map carbonate outcrops resulted in a large amount of noise. Consequently, distinct identification of the actual carbonate outcrop was
not possible. Carbonate outcrops gave poor interpretation due to the presence of banding noise and poor signal received at band 14 (refs 3, 31 and 32). Ultramafics and silicates also gave lower values in the carbonate map. Therefore, this was a better option for the interpretation of ultramafic and siliceous rocks (Figure 5).

A lithological map was prepared using a combination of siliceous, ultramafic and carbonate-generated maps. The map has an RGB (red, green and blue) combination of the siliceous, ultramafic and carbonate maps (Figure 6). Siliceous rock bodies (shown in dark red colour in the figure) are present as peripheral to the basement bodies. They are also parallel to the carbonate rock bodies. Few silicate rock outcrops are scattered near the ultramafics at the lower centred part (Figure 6). Ultramafic outcrops are present in bright green colour from middle to lower right. The mapped ultramafic body is highly correlatable with the published map of Gupta et al.\textsuperscript{13}. Carbonate bodies are visible on the map (cyan colour, Figure 6). However, majority of the study area is covered with cyan in the figure, which is not an accurate interpretation of the carbonate outcrop pattern. Nevertheless, the known carbonate occurrences are mapped accurately giving a high tone of cyan compared to others.

**Validation**

The prepared lithological map was validated with the help of published geological map of the region, matching of spectral signature of the rocks with library spectra and also with ground-collected data. The geological map is highly correlatable with the mapped ultramafic and siliceous rock outcrops (Figure 7\textsuperscript{a} and \textsuperscript{b}), and also with the known carbonate occurrences. Spectral signatures matched with JHU mineral library, representing the
Figure 5. Carbonate map using the CI-N thermal index.

Figure 6. Lithological map of the study area generated by a combination of developed rock maps. The combination of bands are QI-RH in red, MRI-AV in green and CI-N in blue.
matching of dolomite spectra in the carbonate region and olivine spectra in the ultramafic (Figure 7 c and d). Field-collected ground data clearly demarked the presence of ultramafic, siliceous and carbonate bodies in the thermally identified regions (Figure 7 e–h).

**Conclusion**

ASTER data of Terra satellite have been successfully used to develop the regional scale outcrop pattern of ultramafics with the help of MRI-AV index. Besides this, the known occurrences of carbonate rocks hosting lead–zinc mineralization have also been mapped and extrapolated, though the cyan colour distribution was not so accurate with the CI-N index. Siliceous rocks identified in the study are also important as they are suggestive of distribution of granitoids. The QRI-AV, QI-RH and QI-N indices applied for the identification of these rocks played a unique role as thin bands parallel to the carbonates were also mapped. RGB image of the study area was generated using the siliceous, ultramafic and carbonate outcrop patterns.
The authors declare that there is no conflict of interest.

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