Coral Energy Reserves and Calcification in a High-CO2 World at Two Temperatures

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Abstract

Rising atmospheric CO2 concentrations threaten coral reefs globally by causing ocean acidification (OA) and warming. Yet, the combined effects of elevated pCO2 and temperature on coral physiology and resilience remain poorly understood. While coral calcification and energy reserves are important health indicators, no studies to date have measured energy reserve pools (i.e., lipid, protein, and carbohydrate) together with calcification under OA conditions under different temperature scenarios. Four coral species, Acropora millepora, Montipora monasteriata, Pocillopora damicornis, Turbinaria reniformis, were reared under a total of six conditions for 3.5 weeks, representing three pCO2 levels (382, 607, 741 μatm), and two temperature regimes (26.5, 29.0 °C) within each pCO2 level. After one month under experimental conditions, only A. millepora decreased calcification (~53%) in response to seawater pCO2 expected by the end of this century, whereas the other three species maintained calcification rates even when both pCO2 and temperature were elevated. Coral energy reserves showed mixed responses to elevated pCO2 and temperature, and were either unaffected or displayed nonlinear responses with both the lowest and highest concentrations often observed at the mid-pCO2 level of 607 μatm. Biweekly feeding may have helped corals maintain calcification rates and energy reserves under these conditions. Temperature often modulated the response of many aspects of coral physiology to OA, and both mitigated and worsened pCO2 effects. This demonstrates for the first time that coral energy reserves are generally not metabolized to sustain calcification under OA, which has important implications for coral health and bleaching resilience in a high-CO2 world. Overall, these findings suggest that some corals could be more resistant to simultaneously warming and acidifying oceans than previously expected.

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Introduction

Anthropogenic climate change threatens many marine ecosystems today, and coral reefs are among the most sensitive to current changes in ocean biogeochemistry [1,2]. Rising atmospheric carbon dioxide (CO2) concentrations have already caused an increase of 0.6 °C in the average temperature of the upper layers of the ocean over the past 100 years [3], and about one third of all anthropogenic CO2 has been absorbed by the ocean, causing ocean acidification (OA) [4,5]. Since scleractinian corals are calcifying organisms that already live close to their upper thermal tolerance limits [6], both ocean warming and acidification severely threaten their survival and role as reef ecosystem engineers [1,7].

The uptake of anthropogenic CO2 by the ocean changes the carbonate chemistry of seawater by increasing proton (H+), and bicarbonate (HCO3-) concentrations, while at the same time decreasing the concentration of carbonate (CO32-). Consequently, seawater pH (i.e., -log[H+]) and the saturation state with respect to aragonite [Ωarag = [Ca2+][CO32-]/Ksp with Ksp being the ionic product of [Ca2+] and [CO32-] under solution-mineral equilibrium) decreases. As aragonite is the form of calcium carbonate (CaCO3) precipitated by modern corals, this process compromises marine calcification [8–10]. Over the past century, Ωarag in the tropics has decreased from 4.6 to 4.0 [8] and is expected to decrease to 2.5–3.0 by the year 2100 [1,8,11]. Further, it has been estimated that scleractinian calcification rates may drop by up to 35%–40% by the end of this century [8,12].
Coral calcification typically decreases in response to experimentally reduced seawater pH [13–24] but not always [14,15,25–31]. Seawater temperature also influences calcification [32–36], resulting in potentially interactive effects of temperature and OA on coral calcification. For example, negative effects of elevated seawater pCO₂ on calcification are often exacerbated when temperature is simultaneously increased [20,27,30], suggesting a synergistic interactive effect. However, this is not always observed [19,28,31] and in one study even the opposite was shown [37]. Clearly, further studies are required to gain a better understanding of the interactive effects of elevated temperature and pCO₂ on coral calcification and its resistance to OA.

Much less is known about how combined OA and warming will influence other aspects of coral physiology such as energy reserves and tissue biomass. If calcification becomes energetically more costly under elevated pCO₂ due to a decreased aragonite saturation state [38–40], then the extra energy needed to maintain calcification might be drawn from one or more of the following sources: 1) Coral energy reserves (i.e., lipids, protein, carbohydrates), 2) Enhanced endosymbiotic algal production due to CO₂ fertilization [41], and 3) Increased heterotrophy (i.e., zooplankton, particulate and/or dissolved organic carbon) [28,42]. These responses may be even more extreme with the simultaneous increases in seawater temperature because tissue biomass, energy reserves, and endosymbiotic algal density are typically lowest when temperature (and irradiance) is highest on seasonal timescales [43,44,45] and under bleaching scenarios [46,47,48].

Although tissue biomass and energy reserves are important indicators of coral health [46,47] and play a significant role in promoting resilience to bleaching [49], no studies to date have measured all three energy reserve pools (i.e., lipid, protein, and carbohydrate) under OA conditions at elevated temperature. While protein concentrations were either unaffected [30,31] or increased in response to elevated pCO₂ alone [13,50], the effects of OA, or OA plus elevated temperature, on coral lipids and carbohydrates are unknown. Studies specifically addressing all three energy reserve pools are needed to get a better understanding of how OA affects coral energetics and their overall resistance to future climate change.

Finally, the algal endosymbiont (Symbiodinium sp.) provides healthy corals with up to 100% of their daily metabolic energy demand via photosynthesis [51]. If algal productivity is enhanced under OA due to CO₂ fertilization [41], this might help maintain calcification rates and/or energy reserves under OA as energetic costs for calcification increase. Further, Symbiodinium sp. exhibit high sensitivity to elevated seawater temperature [52]. Thus, it is important to monitor endosymbiont and chlorophyll a concentrations in studies manipulating both pCO₂ and temperature.

Here, we studied the single and interactive effects of pCO₂ (382, 607, 741 μatm) and temperature (26.5 and 29.0°C) on coral calcification, energy reserves (i.e., lipid, protein, and carbohydrate), chlorophyll a, and endosymbiont concentrations in 4 species of Pacific coral with different growth morphologies. It was hypothesized that 1) calcification and energy reserves decrease in response to elevated pCO₂ and elevated temperature, 2) decreases are larger when pCO₂ and temperature are elevated simultaneously, and 3) that physiological responses are species-specific.

We show that only one of the four coral species studied here decreased calcification in response to average ocean acidification levels expected by the second half of this century (741 μatm), even when combined with elevated temperature (+2.5°C). Further, we show for the first time that energy reserves were largely not metabolized in order to sustain calcification under elevated pCO₂ and temperature, suggesting that some coral species will be more resistant to combined ocean acidification and warming than previously expected.

### Materials and Methods

#### Experiment

Six parent colonies of Acropora millepora, Porites damicornis, Montipora monasterioa, and Turbinaria reniformis were purchased from Reef Systems Coral Farm (New Albany, Ohio, USA) which is a CITES permit holder. The parent colonies were specifically collected for this experiment from 3–10 m in northwest Fiji (17°29′19″S, 177°23′39″E) in April 2011. Colonies of the same species were collected at least 10 m apart to increase the probability that different genotypes of the same species were selected. All colonies were shipped to Reef Systems Coral Farm and maintained in recirculating indoor aquaria with natural light (greenhouse, 700–1000 μmol quanta m⁻² s⁻¹) and commercially available artificial seawater (Instant Ocean Reef Crystals) for 2.5 months until the start of the experiment.

From April 22 - May 19, 2011, six fragments were collected from each parent colony and mounted on PVC tiles for a total of 144 fragments (4 species × 6 colonies × 6 fragments; Fig. 1). Starting on June 19, 2011, corals were gradually acclimated to a custom-made artificial seawater (ESV Aquarium Products Inc.), which was designed to mimic the chemical composition and alkalinity of natural reef seawater. On July 8 and 9, 2011, all 144 fragments were transferred to the experimental recirculating indoor aquaria with artificial light (Tek Light T5 actinic lights, 275 μmol quanta m⁻² s⁻¹; 9:15 hrs light:dark cycle) and allowed to acclimate to the artificial light conditions for 10 days under ambient seawater conditions (i.e., 26.5°C and 382 μatm). Photosynthesis to irradiance (P/E) curves performed on Acropora millepora showed that photosynthesis was fully saturated at these light levels. Due to logistical reasons, P/E curves were not performed on the other species.

For each of the 6 treatments, the recirculating tank system consisted of one 905 L sump and six aquaria of 57 L each. One fragment per parent colony per species was put in one of the 6 aquaria in each system such that there were a total of 4 fragments (one of each species) in each aquarium, and each parent colony of each species was represented in each system. By placing the same genotypes in each treatment, genotypic variation between treatments was minimized and our ability to detect treatment effects was optimized. Replication of treatments and independent tanks within treatments was not possible due to the complexity and cost of operating tanks under modified pH conditions. While this is, strictly speaking, a pseudo-replicated design [53], the disadvantages of this design are outweighed by the advantages of being able to simultaneously manipulate six combinations of temperature and pH. To optimize the experimental design conditions, coral fragments were rotated daily within tanks and every 3 days among tanks within each system to minimize any tank or positional effects within each system. Further, tanks were cleaned every three days, and great care was taken to ensure similar conditions across treatments except for carbonate chemistry and temperature.

Experimental treatments were assigned to each system as follows: 26.5°C and 382 μatm, 26.5°C and 607 μatm, 26.5°C and 741 μatm, 29.0°C and 382 μatm, 29.0°C and 607 μatm, and 29.0°C and 741 μatm (Fig 1). The three pCO₂ levels–382, 607, and 741 μatm – were designed to represent present day pCO₂, and two μatm levels expected by the second half of the 21st century, respectively. The control temperature (26.5°C) represents the current average annual temperatures in Fiji (http://www.ospo.org).
noaa.gov/Products/ocean/index.html), whereas 29.0 °C represents the upper limit of current summer temperatures but is still below the bleaching threshold at that location. Therefore, the 26.5 °C and 382 μatm treatment served as control. The experiment lasted for 24 days from July 19-August 12, 2011.

Temperature was controlled by titanium aquarium heaters submerged in each system sump (Aqua Medic) and connected to a digital control system (Neptune Systems Apex AquaController). Temperature loggers (Onset Hobo Pro v2) were placed in each sump and recorded temperature every 5 minutes. Seawater CO₂ was controlled by bubbling in pure CO₂, CO₂-free air, or ambient air delivered by an outdoor air pump (Sweetwater, Aquatic Eco-Systems Inc.) into each system sump. CO₂-free air was achieved by moving ambient air through CO₂-scrubbers consisting of a 1.5 m long tube (10 cm diameter) filled with soda lime (SodaSorb HP). Supply of all gases was controlled via a pH stat system using custom designed software (KSgrowstat, written by K. Oxborough, University of Essex). Seawater pH was measured every 5 seconds by microelectrodes (Thermo Scientific Orion Ross Ultra pH glass electrode), which were calibrated daily.

For the elevated temperature (29.0 °C) treatments, temperature was gradually increased over several days until the desired temperature was reached. For the medium (607 μatm) and high (741 μatm) CO₂ treatments, CO₂ was gradually increased over several days starting from 382 μatm until the final CO₂ was achieved. Recirculating seawater flow rate was 210–230 l/hour and little pumps (Accela Powerheads) created additional water circulation within each aquarium. A quarter of the entire water volume of each treatment system was exchanged every 3 days. Non-carbonate ceramic filter media (MarinePure High Performance Biofilter Media, CerMedia) were placed in the sumps to filter the water. Tanks were cleaned every 3 days or as needed.

Figure 1. Photos of representative coral fragments from (a) Acropora millepora, (b) Pocillopora damicornis, (c) Montipora monasteriata, and (d) Turbinaria reniformis. Rectangles indicate subsamples taken from each fragment for lipid, protein/carbohydrate, and tissue biomass analyses. The remaining tissue was airbrushed for chlorophyll a and endosymbiont density measurements.

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Monitoring of seawater chemistry during the experiment

Temperature and salinity were measured daily (YSI 63), and salinity was adjusted daily to 35 ppt. Daily water samples were taken using screw-top high-density polyethylene bottles for pH and alkalinity analyses. After equilibration at 25 °C in a recirculating water bath (30 min), sample pH_NBS was measured with an Orion® Ross glass electrode (precision 0.01 pH units) [56], which was calibrated daily at 25 °C. Total alkalinity (TA) was titrated with HCl on the same samples using an AS-ALK2 (Apollo SciTech Inc.) alkalinity titrator [57] (precision 0.1%). The HCl solution was calibrated with Certified Reference Material (CRM) from A.G. Dickson (Scripps).

Treatment xCO₂ (dry air), aragonite saturation state (Ω_{arag}), and pH₇ were calculated using the program CO2SYS [58] based on measured pH_NBS and alkalinity at the respective temperature. xCO₂ was converted to pCO₂ using the equation in Weiss et al. [59]. Carbonate dissociation constants were taken from Millero et al. [60]. In addition, a custom-made CO₂ analyzer based on a LI-COR 820 was used weekly to crosscheck with calculated sump xCO₂ values according to methods by Wang & Cai [56], and indicated good agreement of measured and calculated values (r² = 0.97, n = 66).

Since healthy corals in situ can acquire up to 46% of their daily metabolic energy demands by feeding on zooplankton [48,54], corals were fed every three days with 48 h old brine shrimp nauplii (Artemia sp., Carolina Biological Supply). Corals were allowed to acclimate to the dark for 30 min before feeding was conducted. They were fed for one hour in separate, partially submerged plastic containers containing water from their respective treatment, and at a concentration of approximately 1 brine shrimp ml⁻¹ which is representative of zooplankton concentrations on natural Pacific reefs [55]. At the end of the hour, brine shrimp nauplii remained in the feeding chambers indicating that the corals had not captured all brine shrimp nauplii available to them. Following feeding, the corals were placed back in their respective aquaria and the feeding container water discarded so as not to introduce brine shrimp into the recirculating systems.
Laboratory analyses

**Calcification.** Net calcification was determined using the buoyant weight technique [61]. Each coral fragment was buoyantly weighted at the beginning, middle (after 11 experimental days), and at the end of the experiment (after 23 experimental days). As such, it was possible to assess if calcification rates varied during the experiment. Daily calcification rates were calculated as the difference between initial, middle, and final weights, divided by the respective number of days elapsed, and standardized to surface area (see below).

For tissue analyses, corals were frozen at −80°C and a total of three branch tips or growing edge pieces were saved from each fragment for lipid, protein/carbohydrate, and tissue biomass analyses, respectively (Fig. 1). The remaining tissue was airbrushed for chlorophyll a and endosymbiont density measurements.

**Chlorophyll a and endosymbiont density.** Coral tissue was stripped off the coral skeleton with a waterpik [62] containing 40 ml of synthetic seawater (Instant Ocean). The endosymbionts were isolated from the host tissue via centrifugation and then resuspended in 10 ml of synthetic seawater. For chlorophyll a concentrations, 1 ml of this algal suspension was pelleted and the cells lysed in 1 ml of 4°C methanol using a bead-beater for 60 seconds. Samples were then immediately placed on ice and allowed to extract for one hour in the dark. Samples were centrifuged to remove cellular debris and measured spectrophotometrically (λ = 652, 665 & 750) on a 96-well plate reader. The equations for chlorophyll a in methanol described by Porra et al. [63], along with path length correction [64], were used to calculate chlorophyll a concentrations (pg/cell), and were then standardized to surface area (see below). Another 1 ml subsample of the algal suspension was preserved with 10 µl of 1% glutaraldehyde solution for endosymbiont quantification, which was calculated using 6 independent replicate counts on a hemocytometer, using a Nikon microphot-FXA epifluorescent microscope at 100x magnification. Photographs were analyzed through Image J using the analyze particles function.

**Energy reserves and tissue biomass.** For all energy reserve and tissue biomass measurements, only branch tips or samples with a growing edge were used. While tissue composition may vary across the surface of a coral [65], this approach was used to allow for comparison with previously published studies [46,47,66]. Soluble lipids (referred to hereafter simply as lipids) were extracted from a whole, ground coral sample (skeleton + animal tissue + algal endosymbiont) in a 2:1 chloroform:methanol solution for 1 hour [46,66] washed in 0.88% KCl followed by 100% chloroform and another wash with 0.88% KCl. The extract was dried to constant weight under a stream of pure nitrogen 100% chloroform and another wash with 0.88% KCl. The extract was washed in 0.88% KCl followed by allowed to extract for one hour in the dark. Samples were centrifuged to remove cellular debris and measured spectrophotometrically (λ = 652, 665 & 750) on a 96-well plate reader. The equations for chlorophyll a in methanol described by Porra et al. [63], along with path length correction [64], were used to calculate chlorophyll a concentrations (pg/cell), and were then standardized to surface area (see below). Another 1 ml subsample of the algal suspension was preserved with 10 µl of 1% glutaraldehyde solution for endosymbiont quantification, which was calculated using 6 independent replicate counts on a hemocytometer, using a Nikon microphot-FXA epifluorescent microscope at 100x magnification. Photographs were analyzed through Image J using the analyze particles function.

Tissue biomass was measured by drying a third branch tip of whole coral sample (skeleton + animal tissue + algal endosymbiont) to constant dry weight (24 hours, 60°C) and burning it (6 hours, 450°C). The difference between dry and burned weight was the ash-free dry weight which was standardized to the surface area of this branch tip.

**Surface area.** Surface area of plating *M. monasteriata* and *T. reniformis* fragments was determined using the aluminum foil technique [69], whereas surface area of branching *A. millepora* and *P. damicornis* fragments was determined using the single wax dipping technique [70,71] after the tissue had been removed. Natural wooden blocks of varying sizes and shapes were used as calibration standards [71]. Wax dipping was conducted using household paraffin wax (Gulf Wax, Royal Oak Enterprises) heated to 65°C. Dried coral skeletons and wooden calibration standards were maintained at room temperature prior to weighing.

**Statistical analyses**

Three-way mixed-model analyses of variance (ANOVA) tested the effects of pCO2, temperature, and parent colony on calcification rates in the first and second half of the experiment, chlorophyll a, algal endosymbiont density, lipid, protein, carbohydrate, and tissue biomass. Temperature and pCO2 were fixed and fully crossed, whereas parent colony was a random factor. The ANOVAs were run for each species separately. All data were normally distributed according to plots of residuals versus predicted values for each variable, or transformed to meet the condition of normality. Outlier values greater than 3 times the interquartile range were excluded. Post hoc Tukey tests were performed when main effects were significant (p≤0.05). A posteriori slice tests (e.g., tests of simple effects) [72] determined if the ambient (26.5°C) and elevated (29.0°C) temperature treatment averages significantly differed within each pCO2 level. Bonferroni corrections were not applied [73,74], therefore significant model p-values >0.0016 (0.05/32 tests) should be interpreted with caution. Statistical analyses were performed using SAS software, Version 9.2 of the SAS System for Windows.

**Results**

All corals appeared healthy throughout the experiment. No visible paling and no mortality occurred. The average seawater temperature, pH, pCO2, saturation state, and total alkalinity for all six treatments are summarized in Table 1.

**Calcification**

In *Acropora millepora*, calcification rates during the first half of the experiment were overall unaffected by both temperature (p = 0.36) and pCO2 (p = 0.79) (Fig. 2a; Table S1). However, at the highest pCO2 level (741 µatm) calcification was 43% lower at 29.0°C than at 26.5°C (Fig. 2a). During the second half of the experiment, calcification rates were significantly affected by pCO2 (p = 0.0011) but not temperature (p = 0.42), and were lower by 53% at the highest compared to the lowest pCO2 level (Fig. 2b; Table S1).

In *Pocillopora damicornis*, a significant interaction of pCO2 and temperature (p<0.001) was observed for calcification rates during the first half of the experiment (Fig. 2c, Table S1). During the second half of the experiment, calcification rates were generally unaffected by temperature (p = 0.06) and pCO2 (p = 0.07). However, at ambient seawater pCO2 (382 µatm) corals kept at elevated temperature (29.0°C) calcified 91% more compared to those kept at 26.5°C (Fig. 2d, Table S1).

Calcification rates of *Montipora monasteriata* were affected by temperature (p = 0.04) but not pCO2 (p = 0.42) during the first half...
of the experiment (Fig. 2e–f, Table S1), with corals calcifying 18% more at elevated compared to ambient temperature. This was largely driven by significant temperature differences at both 382 and 607 µatm but not 741 µatm. During the second half of the experiment, calcification rates were unaffected by both temperature ($p = 0.82$) and $p$CO$_2$ ($p = 0.14$).

In contrast, calcification rates of *Turbinaria reniformis* during both first and second half of the experiment (Fig. 2g–h, Table S1) did not respond to changes in seawater temperature ($p = 0.45$ and 0.17) or $p$CO$_2$ ($p = 0.36$ and 0.09). Notably, the two plating species (*M. monasteriata* and *T. reniformis*) calcified more than twice as fast as the two branching species (*A. millepora* and *P. damicornis*).

**Chlorophyll a and endosymbiont density**

The chlorophyll a concentrations of *A. millepora* were significantly affected by $p$CO$_2$ ($p<0.001$) but not temperature ($p = 0.054$), with concentrations being 31% lower at 607 µatm than at either 382 or 741 µatm (Fig. 3a, Table S2). Endosymbiont densities were not affected by either seawater temperature ($p = 0.054$) or $p$CO$_2$ ($p = 0.03$ but overall model $p = 0.24$) (Fig. 3b, Table S2).

In *P. damicornis*, a significant interaction of temperature and $p$CO$_2$ was observed for both chlorophyll a concentrations and endosymbiont densities ($p<0.001$ and $p = 0.02$, respectively) (Fig. 3c–d, Table S2). When temperature was elevated, chlorophyll a concentrations were higher by 19% and 67%, respectively, at both 382 and 741 µatm. At 607 µatm, endosymbiont densities decreased by 36% at 29°C compared to concentrations at ambient temperature.

In *M. monasteriata*, a significant interaction of seawater temperature and $p$CO$_2$ was observed for chlorophyll a concentrations ($p = 0.01$) (Fig. 3e, Table S2), with concentrations being 45% and 30% lower at elevated compared to ambient temperature, respectively, under both 382 and 741 µatm conditions. Endosymbiont densities were significantly affected by both temperature ($p<0.001$) and $p$CO$_2$ ($p = 0.01$) but the interaction term was not significant ($p = 0.38$) (Fig. 3f, Table S2). Densities were 32% lower at elevated compared to ambient temperature, and were lowest

| 400 ppm target | elevated temp. | 600 ppm target | elevated temp. | 800 ppm target | elevated temp. |
|----------------|----------------|----------------|----------------|----------------|----------------|
| **Temp.(°C)** | 26.45±0.01 | 29.31±0.02 | 26.37±0.01 | 28.53±0.02 | 26.61±0.01 | 28.93±0.02 |
| **pH** | 8.07±0.01 | 8.04±0.01 | 7.90±0.01 | 7.89±0.01 | 7.83±0.01 | 7.81±0.01 |
| **pCO$_2$ (µatm)** | 364.31±9.69 | 400.62±16.83 | 598.37±18.50 | 616.08±24.24 | 732.04±22.37 | 749.63±26.21 |
| **TA (µmol kg$^{-1}$)** | 2269.4±10.84 | 2270.1±11.15 | 2303.8±9.34 | 2288.3±10.43 | 2306.3±10.64 | 2304.5±9.08 |
| **$P_{max}$** | 3.69±0.07 | 3.79±0.09 | 2.75±0.05 | 2.91±0.05 | 2.40±0.06 | 2.52±0.06 |

Mean ± 1 SE are shown. Sample size was 25 for all measurements. Temp. = Temperature. doi:10.1371/journal.pone.0075049.t001

**Table 1.** Average conditions for each of the 6 treatments representing three $p$CO$_2$ levels at two temperature regimes (ambient, elevated = ambient + 2.5°C).

![Figure 2](https://example.com/figure2.png)
overall at 607 μatm (−25%) compared to the other two pCO2 levels.

Chlorophyll a concentrations of *T. reniformis* were affected by pCO2 (\(p = 0.03\)) but not temperature (\(p = 0.11\)), with concentrations being 38% lower at 607 compared to 382 μatm (Fig. 3g, Table S2). Endosymbiont densities were not affected by either temperature (\(p = 0.90\)) or pCO2 (\(p = 0.21\)) but were lower by 55% at elevated compared to ambient temperature under 741 μatm conditions (Fig. 3h, Table S2).

### Energy reserves and tissue biomass

Lipid concentrations of *A. millepora* were affected by seawater pCO2 (\(p = 0.01\)) but not temperature (\(p = 0.053\)), with concentrations being 28% and 21% higher at 607 and 741 μatm, respectively, compared to concentrations at ambient pCO2 (Fig. 4a, Table S3). A significant interaction of seawater pCO2 and temperature was observed for protein concentrations (\(p = 0.01\)) (Fig. 4b, Table S3). Carbohydrate concentrations were affected by both temperature (\(p = 0.02\)) and pCO2 (\(p = 0.01\)) but the interaction term was not significant (\(p = 0.85\)) (Fig. 4c, Table S3). Across all pCO2 treatments, carbohydrate concentrations were 18% lower at 29.0°C compared to 26.5°C, and 41% lower at 607 μatm than at 741 μatm. Tissue biomass was unaffected by changes in seawater temperature (\(p = 0.99\)) and pCO2 (\(p = 0.07\)) (Fig. 4d, Table S3).

In *P. damicornis*, lipid concentrations were affected by seawater pCO2 (\(p = 0.01\)) but not temperature (\(p = 0.53\)) (Fig. 4e, Table S3), with concentrations being 41% and 18% higher at 607 and 741 μatm, respectively, compared to concentrations at ambient pCO2 (Fig. 4e, Table S3). Neither protein, nor carbohydrate concentrations or tissue biomass were affected by seawater temperature (\(p = 0.63\), 0.88, 0.33, respectively) and pCO2 (\(p = 0.52\), 0.35, 0.41, respectively) (Fig. 4f–h, Table S3).

The lipid concentrations of *M. monasteriata* were unaffected by both seawater temperature (\(p = 0.38\)) and pCO2 (\(p = 0.23\)) (Fig. 4i, Table S3). A significant interaction of temperature and pCO2 was observed for protein concentrations (\(p < 0.001\)): at 382 μatm, they decreased (−27%) at elevated compared to ambient temperature, whereas at 607 and 741 μatm, they increased (+36% and +60%, respectively) (Fig. 4j, Table S3). Carbohydrate concentrations were affected by seawater temperature (\(p = 0.02\)) but not pCO2 (\(p = 0.36\)) (Fig. 4k, Table S3), and the concentrations were 25% higher at elevated than at ambient temperature under 741 μatm conditions. Tissue biomass was also unaffected by both seawater temperature (\(p = 0.70\)) and pCO2 (\(p = 0.12\)) (Fig. 4l, Table S3).

In *T. reniformis*, none of the measured energy reserve pools responded to changes in seawater temperature and pCO2 (Fig. 4m–o, Table S3). Tissue biomass was also unaffected by both temperature (\(p = 0.62\)) and pCO2 (\(p = 0.58\)), but was 21% lower at 29.0°C compared to 26.5°C at 382 μatm (Fig. 4p, Table S3).

### Effects of parent colony

Parent colony was a significant effect in many of the measured variables, but no single parent colony or group of specific parent colonies was consistently different from all other parent colonies in any of the species studied (Tables S1–S3).

### Discussion

Coral calcification has been predicted to decrease dramatically by the end of this century, thus threatening the existence of coral reefs in the future. Although the response of coral calcification is not uniform across species, most studies have found that calcification decreases with increasing seawater pCO2 [75–77]. Here, we show that only one of the four Pacific coral species studied here decreased calcification in response to average ocean acidification levels expected by the second half of this century (741 μatm), even when combined with elevated temperature (+2.5°C). Further, we investigated for the first time the effects of OA on coral energy reserves and show that they were largely not metabolized in order to sustain calcification under elevated pCO2 and temperature.

*Acropora millepora* was the only coral out of the four species studied here that decreased calcification rates in response to OA (Fig. 2b). While calcification rates were not affected by elevated pCO2 and/or temperature during the first half of the experiment, they declined by 53% during the second half of the experiment.
due to acidification alone. As the second half of the experiment is more likely to reflect the long term response of corals to ocean acidification, this negative response to OA is consistent with other studies on *Acropora* sp. [14,20,24,78], although the amount of decline differs between species. The absence of any change in calcification of *Pocillopora damicornis* is consistent with another study [14], whereas declines in calcification of 50% in *P. meandrina* were reported [37]. The lack of any change in calcification rates of *Montipora monasteriata* and *Turbinaria reniformis* due to acidification (Fig. 2d, f, h) is in contrast to other studies which reported a 15–20% decline in *Montipora capitata* [16], and a 13% decline in *T. reniformis* (albeit at pCO2 levels that were considerably higher than those in the present study) [78]. Although a significant interaction of pCO2 and temperature was observed in *P. damicornis* during the first half of the experiment (Fig. 2c), this was not observed during the second half. Similarly, *M. monasteriata* calcified more at elevated compared to ambient temperature during the first half (Fig. 2c), but not during the second half of the experiment. Thus, it appears that with the exception of *A. millepora*, these species may be resistant to changes in pCO2 and temperature within the parameter ranges investigated in this study.

In the current study, elevated temperature did not exacerbate or counteract the negative effects of OA on calcification in *A. millepora*, and did not have an overall negative affect on calcification in the other three species. This is in contrast to other studies where elevated temperature was found to mitigate negative OA effects. For example, Anthony *et al.* [20] found that elevated temperature (28–29°C vs. 25–26°C) prevented a decline of calcification in *A. intermedia* at elevated pCO2 (520–705 μatm). Muehllehner and Edmunds [37] showed that the negative effects of elevated pCO2 (720 μatm) were fully mediated in *P. meandrina* when OA was combined with elevated temperature (29°C vs. 27°C). Overall, these findings add to the growing body of evidence that the response of coral calcification to OA is highly species specific, and that some coral species may maintain calcification under combined ocean acidification and warming in the future. Although the current study was conducted using artificial seawater, it is unlikely that this influenced the observed responses of calcification to ocean acidification. The carbonate chemistry of the custom-made seawater mimicked natural conditions very well (Table 1), and calcification rates – as well as chlorophyll a concentrations, endosymbiont densities, energy reserves, and tissue biomass – were within the range observed in the field and/or other experimental studies using natural seawater [13,28,43,45,46,66,78–89].

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**Figure 4.** Average lipid, protein, carbohydrate concentrations, and tissue biomass of (a–d) *Acropora millepora*, (e–h) *Pocillopora damicornis*, (i–l) *Montipora monasteriata*, and (m–p) *Turbinaria reniformis*. Averages ± 1 SE are shown for three pCO2 levels and two temperature regimes (26.5, 29.0°C). Asterisks indicate significant differences between 26.5 and 29.0°C within a specific pCO2 level (determined by a posteriori slice tests). The letters a and b indicate results of the post hoc Tukey tests when there was a significant pCO2 effect. Sample sizes ranged between 4 and 6. Statistical details can be found in Table S3.

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While many studies note a decline in coral calcification with increasing \( \mu \text{CO}_2 \) [75–77], there is considerable among-study variation [75,76], and some species are more resistant than others [15,90]. Such differences may be due to experimental duration, how seawater carbonate chemistry is altered (i.e., bubbling \( \text{CO}_2 \) vs. acid addition), and how calcification is measured (i.e., buoyant weight vs. total alkalinity anomaly technique). Meta-analyses have shown that experimental duration or the method of carbonate chemistry manipulation did not explain the large variability of responses observed among studies [75,76]. While this study suggests that experimental duration can influence the response of calcification to OA in some species (i.e., calcification of A. millepora decreased only during the second half of the second half), it is likely that biological aspects have a stronger influence on the sensitivity of coral calcification to OA than differences in methodology. Important biological aspects include energetic status and feeding [28,38], enhanced algal production [41,91], and cellular \( \text{pH} \) control [92–95].

Despite the assumption that calcification becomes energetically more costly under OA [38–40], energy reserves did not decline with increasing \( \mu \text{CO}_2 \) (Fig. 4). Lipid concentrations increased under OA conditions in both A. millepora and P. damicornis, and were fully maintained in M. monasteriata and T. reniformis. Protein, carbohydrate, and tissue biomass were overall maintained under OA conditions in all species. Further, temperature did not negatively affect energy reserves and tissue biomass except for carbohydrate concentrations in A. millepora, which were lower at elevated compared to ambient temperature. Importantly, energy reserves and tissue biomass were fully maintained or even increased at the highest \( \mu \text{CO}_2 \) level in A. millepora despite dramatic decreases in calcification rates. These findings suggest that (1) energy reserves are generally not metabolized under OA conditions or OA at elevated temperature, and (2) that either energy reserves do not play a role in sustaining calcification under OA conditions, or that the increased energetic costs of maintaining calcification under OA are relatively insignificant. This is consistent with other work showing that calcification likely does not become energetically more costly under OA conditions [81], and that the extra energy required to up-regulate \( \text{pH} \) at the site of calcification under OA conditions is <1% of that produced by photosynthesis [92].

Further, from an energetic standpoint of view, the total amount of energy reserves present in a coral species did not seem to be related to their calcification response to OA. The energetic content of lipid, protein, and carbohydrates is better assessed from an energetic point of view [96,97], as specific enthalpies of combustion differ significantly among energy reserve pools: -39.5 kJ g\(^{-1}\) for lipid, -23.9 kJ g\(^{-1}\) for protein, and -17.5 kJ g\(^{-1}\) for carbohydrate [96]. When the total amount of energy available to each species was calculated (i.e., the sum of lipid, protein, and carbohydrate expressed in kJ g\(^{-1}\) tissue biomass), A. millepora had the lowest amount of all species in the control treatment (6.6 vs. up to 8.1 kJ g\(^{-1}\) in P. damicornis), but the highest amount in the high-CO\(_2\) treatment (9.0 vs. 6.6 kJ g\(^{-1}\) in M. monasteriata), and a similar amount as both M. monasteriata and T. reniformis in the high-CO\(_2\)– high temperature treatment (7.3 kJ g\(^{-1}\) vs. 7.2 and 7.8 kJ g\(^{-1}\), respectively). It is therefore unlikely that high levels of energy reserves per se help corals maintain calcification rates under OA conditions.

However, maintaining energy reserves and tissue biomass under ocean acidification does have crucial implications for other aspects of coral health and resistance to stressors such as coral bleaching. For example, maintenance of lipid concentrations may enable corals to maintain their reproductive output [98], even under future OA and warming. This may be critical considering that many other processes involved in coral reproduction such as fertilization, settlement success, and metamorphosis are compromised under OA [99,100]. Furthermore, maintenance of energy reserves has been shown to be associated with higher resistance to coral bleaching and to promote recovery from bleaching [46,49], which could prove critical as bleaching events will increase in frequency over the coming decades [101].

Heterotrophy is known to promote energy storage, tissue synthesis, and skeletal growth in healthy and bleached corals [47,48,102] as well as corals subjected to OA [28,42]. Therefore, bivweekly feeding in this study (intended to mimic zooplankton contribution to the coral diet on the reef) may have helped corals to sustain energy reserves and tissue biomass under these conditions. It has further been suggested that coral tissue reacts to availability of such resources faster than skeletal growth [103,104], which could explain why tissue biomass - but not necessarily calcification – was maintained or even increased in all four species irrespective of \( \mu \text{CO}_2 \) or temperature conditions. As feeding rates and heterotrophic plasticity are highly species-specific [48,54,105], it is likely that heterotrophic carbon intake differed significantly among the species studied here, potentially contributing to their differential responses to OA.

Enhanced algal productivity due to \( \text{CO}_2 \)-fertilization may help corals to maintain calcification under OA conditions. Although chlorophyll a concentrations and endosymbiont density were unaffected at the highest \( \mu \text{CO}_2 \) level (except for chlorophyll in T. reniformis), \( \text{CO}_2 \)-fertilization may nevertheless have played a role in helping corals to maintain energy reserves and/or calcification. Increased availability of \( \text{CO}_2_{a} \) under OA conditions may enhance algal productivity, especially in Symbiodinium phylotypes with less efficient carbon-concentrating mechanisms, which rely to a greater extent on the passive, diffusive uptake of \( \text{CO}_2_{a} \) [41]. Thus, a potentially increased translocation of autotrophic carbon to the animal host may have contributed to the maintenance of energy reserves and tissue biomass observed here.

Interestingly, both chlorophyll a concentrations and endosymbiont density were often lowest at 607 \( \mu \text{m} \), showing a non-linear relationship with increasing \( \mu \text{CO}_2 \). Nevertheless, the lack of any significant difference in chlorophyll a and/or symbiont density at 741 \( \mu \text{m} \) versus ambient \( \mu \text{CO}_2 \) concentrations (except for chlorophyll in T. reniformis) is consistent with other studies [23,28,29,106]. The reason for the observed minima at ~600 \( \mu \text{m} \) is unknown. Similar non-linear responses were not observed for calculation rates, tissue biomass, and most energy reserve pools, suggesting that this did not translate into a decreased performance of the animal host. Edmunds [81] also observed a non-linear \( \mu \text{CO}_2 \) threshold between 756 and 861 \( \mu \text{m} \) affecting photosynthesis and respiration in massive Porites corals, thus highlighting the importance of studying multiple \( \mu \text{CO}_2 \) levels in OA experiments in order to assess non-linear physiological responses and to better forecast physiological responses over the coming century as the oceans continue to warm and acidify.

In addition to energetic status and enhanced algal productivity due to \( \text{CO}_2 \) fertilization, other factors such as the amount of control over the carbonate chemistry at the site of calcification may explain the observed differences in susceptibility of calcification to OA. Corals have the ability to significantly up-regulate the \( \text{pH} \) at the site of calcification compared to ambient seawater, even under OA conditions [92–95,107,108]. Yet, the degree to which corals are able to control the \( \text{pH} \) at the site of calcification likely varies among species [92,95,109]. Acropora spp. may have the lowest capacity to up-regulate \( \text{pH} \) at the site of calcification based
on boron isotopic measurements [92,110]. Further, crystallization under OA was most compromised in *Acropora millepora* and least compromised in *T. reniformis* [76]. Thus, we hypothesize that *A. millepora* has a weaker proton pump than the other coral species studied here, making its calcification rate more sensitive to future OA. Although pH up-regulation has not been studied in *Acropora *branching *Acropora* corals, which are important reef builders, are likely to be “losers” on future coral reefs because they are highly susceptible to both bleaching [111] and OA [20,24,90,99]. This can be expected to have severe impacts on reef diversity, structural complexity, and overall reef functioning. Nevertheless, some corals could be more resistant to combined ocean acidification and warming expected by the end of this century than previously thought, as three of the four species fully maintained calcification under elevated pCO2 and temperature without compromising overall energy reserves or biomass. Further, the immediate effects of rising seawater temperature and ocean acidification may be tolerable for some species.

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**Supporting Information**

**Table S1** Results of 8 two-way ANOVAs for average calcification rate during the first and second half of the experiment. (DOCX)

**Table S2** Results of 8 two-way ANOVAs for average chlorophyll a concentrations and symbiont density. (DOCX)

**Table S3** Results of 16 two-way ANOVAs for average soluble lipid, animal soluble protein, animal soluble carbohydrate concentrations, and tissue biomass. (DOCX)

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**Author Contributions**

Conceived and designed the experiments: AGG MEW WJC. Performed the experiments: VS AGG MEW WJC TF M KDH DTP XH QL HX YW YMY JHB. Analyzed the data: VS. Contributed reagents/materials/analysis tools: VS AGG MEW WJC TF M KDH DTP XH QL HX YW YMY JHB. Wrote the paper: VS AGG MEW WJC TF M KDH DTP XH QL HX YW YMY JHB.
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