Shear strength of shallow expansive soil varies along with the depth under the freeze-thaw effect. This work investigates shear strength characteristics of shallow expansive soil by simulating the actual freeze boundary conditions of seasonal frozen areas with water supplement. An integrated approach incorporating the freeze-thaw test and direct shear test was adopted. Firstly, uni-directional freezing tests for expansive soil columns under three different freezing temperature gradients were carried out. Secondly, direct shear tests under low vertical stress were performed on the standard samples, which were prepared by using cutting rings cut the thawed expansive soil columns into nine segments along with the depth. Temperature, water content, and dry density at different depths were also investigated after the freeze-thaw process. The test results showed that, after the freeze-thaw process, the shear strength of expansive soil columns showed significant differences along with the depth and highly correlated with water content, specifically the higher water content and the lower shear strength. The minimum shear strength in the expansive soil columns occurred at the soil layer below the frozen and unfrozen zones interface. The expansive soil column’s shear strength changed most under the moderate freezing temperature gradient corresponding to the most considerable shear strength reduction. Moreover, the significant decrease in cohesion was the main reason for the shear strength reduction of expansive soil after the freeze-thaw process. These results indicate significant depth variability in shear strength of expansive soil under the freeze-thaw effect.

1. Introduction

Expansive soil is a highly plastic soil with many hydrophilic clay minerals like montmorillonite and illite. Expansive soil exhibits significant swell-shrink potential and is more prominent shear strength degradation upon climate variations, such as freeze-thaw effect [1–3]. Large areas of expansive soils have been discovered in seasonal frozen regions accompanying the engineering construction. For example, nearly, 380 km expansive soil open channel in the middle route of South-to-North Water Transfer Project (SNWTP) and more than 260 km of expansive soil cutting slope in the Jilin-Tumen-Hunchun high-speed railway in China were both located in seasonal frozen regions [4–6]. In these hydraulic engineering and traffic engineering projects, slope instability of expansive soil frequently occurs after the freeze-thaw process. For slope stability analysis, expansive soil’s shear strength under the freeze-thaw effect is the necessary mechanical parameter.

When the atmosphere temperature drops below the freezing point under field conditions, the shallow soil will be frozen unidirectional from top to bottom. A typical soil profile for shallow soil in the freezing process includes three parts: the frozen zone, the frozen fringe, and the unfrozen zone [7]. During the freezing process, soil pore water migrates from the unfrozen area to the frozen area under cryogenic suction action [8]. When the atmospheric temperature increases, the upper soil layers begin to thaw first. However, it is difficult for the melted water to infiltrate into the lower layer due to the compacted soil’s low permeability.
Thus, a layer with high water content formed along the plane of the most massive ice lens in the thawed soil, which allows the formation of slurry layers and a consequential weak plane in the soil [9]. Usually, this weak soil layer is the main sliding surface of a soil slope landslide [10, 11]. It is reasonable to believe that the reduction of shear strength varies with depth after the freeze-thaw process. However, there is a minimal understanding of shear strength characteristics of shallow expansive soil after freeze-thaw.

Laboratory experiments offer a possibility to explore the shear strength variation characteristics of soil under the freeze-thaw effect [12–19]. It needs to be emphasized that the soil samples’ freezing and thawing boundary conditions in the laboratory experiments should be consistent with the actual situation on-site. There are two situations for underground water: a high water table and a low water table. If the underground water table is relatively low, it is difficult for water to enter the upper soil. If the groundwater is relatively shallow, water may be induced to migrate upwards and accumulate under cryogenic suction action [20]. An open system and a closed system can be used to simulate high and low underground water tables [21]. Li and Wang [22] have reported shearing strength degradation characteristics of expansive soil due to a closed system’s freeze-thaw effect. However, an open system considering external water supplement is a lack of research.

This work aims to investigate the shear strength characteristics of shallow expansive soil along with the depth due to the freeze-thaw effect by simulating the actual freezing boundary conditions in an open system. An integrated approach was to incorporate unidirectional freezing tests, and direct shear tests were adopted. This work is expected to contribute to a clearer understanding of the shallow expansive soil’s shear strength characteristics under the freeze-thaw effect considering water supplement.

2. Materials and Methods

2.1. Material Properties. The expansive soil sample was collected from a depth of 1.0 m on a construction site in Nanyang, Henan province, China. The mineral composition of the expansive soil sample was determined by an X-ray diffractometer (Table 1). The total proportion of expansive clay minerals (montmorillonite and illite) is 58.6%. The expansive soil sample’s physical properties were tested according to the Standard for Soil Test Method (GB/T 50123-2019) (Table 2). The specific gravity of the expansive soil was 2.73, which was typical clay. Liquid limit ($w_L = 69.1\%$) and plastic limit ($w_P = 21.3\%$) indicated the expansive soil sample possessing high water adsorption capacity: the plasticity index ($I_p$) was calculated to be 47.8%. Furthermore, the expansive soil sample can be classified as CH (high plasticity clay). According to the Technical Code for Building in Expansive Soil Regions, the free swelling ratio of 103% shows that the expansive soil sample has strong expansion potential (GB/T 50112-2013). The maximum dry density was 1.68 g/cm³ at an optimum water content of 23%. The freezing point of $-1.64°C$ was obtained from the expansive soil samples at the state of 18% water content and 1.55 g/cm³ dry density.

### Table 1: Mineral composition of the expansive soil sample.

| Mineral composition     | Content (%) |
|-------------------------|-------------|
| Montmorillonite         | 48.8        |
| Illite                  | 9.8         |
| Chlorite                | 3.4         |
| Kaolinite               | 1.2         |
| Quartz                  | 32.9        |
| Feldspar                | 2.8         |
| Dolomite                | 1.1         |

### Table 2: Physical properties of expansive soil samples.

| Physical index          | Value     |
|-------------------------|-----------|
| Specific gravity, $G_s$ | 2.73      |
| Liquid limit, $w_L$ (%) | 69.1      |
| Plastic limit, $w_P$ (%)| 21.3      |
| Plasticity index, $I_p$ | 47.8      |
| Free swelling ratio (%) | 103       |
| Optimum water content (%)| 23.0     |
| Maximum dry density (g/cm³)| 1.68 |
| Freezing point (°C)     | -1.64     |

2.2. Experiment Equipment. The freeze-thaw device consists of the soil column, temperature control parts, and data measurement and acquisition parts (Figure 1). Soil columns were prepared in a plexiglass tube (height of 36 cm and diameter of 10 cm), which was wrapped by an insulation material to isolate the expansive soil column from the test’s external temperature disturbance. A water valve was set at the bottom of the soil columns to control the external water supplement.

Temperature control parts contain two sets of cold baths and a temperature control chamber. The two sets of cold baths can separately control the heat exchangers’ temperatures, installed on the top and bottom of the soil columns. A particular freeze environment with different temperature gradients was achieved during the test by applying different temperature combinations on the two heat exchangers. The temperature control chamber with an internal dimension of $80cm \times 60cm \times 60cm$ (height $\times$ width $\times$ depth) was used to control the temperature environment around the soil column. The two heat exchangers’ temperature and the temperature control chamber can be controlled with a range from $-60$ to $40°C$ with an accuracy of $\pm0.2°C$.

Data measurement and acquisition parts contain nine thermocouples, one displacement transducer, and one data logger. Nine thermocouples (accuracy of $\pm0.1°C$) were horizontally installed with an equal interval of 4 cm at a depth of 2, 6, 10, 14, 18, 22, 26, 30, and 34 cm in the expansive soil columns to monitor the internal temperature dynamic change during freeze-thaw process. The displacement transducer (accuracy of $\pm0.01 mm$) was placed on the top heat exchanger to monitor the soil columns’ vertical deformation during the test.

The direct shear apparatus is controlled by a servo control system, which can automatically record the horizontal displacement and shear stress. The main technical parameters of the direct shear apparatus are as follows:
300 kN maximum shear stress with a precision of ±5 N, five strain rates (0.02, 0.04, 0.08, 1.00, and 2.00 mm/min), 50 mm maximum horizontal displacement with ±0.01 mm precision, and 20 mm vertical displacement transducer with 0.01 mm precision.

2.3. Experiment Process. The experiment process consists of three procedures: expansive soil columns preparation, freeze-thaw tests, and direct shear tests. Predetermined amounts of water and air-dried soil samples after screening from a 2.0 mm sieve were mixed and allowed to equilibrate for 24 h in sealed bags to achieve an initial water content of 18% and then dynamically compacted in six layers in the plexiglass tube. The dry densities of all soil columns were controlled to be 1.6 g/cm³. Vaseline was smeared on the plexiglass tube’s internal face before soil compaction to reduce boundary friction.

Three different freezing schemes (F1, F2, and F3) were adopted in unidirectional freezing tests (Table 3). Moreover, the different freezing schemes had different freezing temperature gradients by setting different temperature combinations to the two heat exchangers. The top heat exchangers’ temperatures were set as −5, −10, and −15°C, corresponding to freezing schemes F1, F2, and F3, respectively. Furthermore, the bottom heat exchangers were always controlled to be +1°C during the freezing process. Therefore, the three freezing temperature gradients were 0.17, 0.31, and 0.44°C·cm⁻¹, respectively. All the expansive soil columns were precooled at +1°C in the temperature control chamber for 24 h before freezing tests, and then exposed at different freezing schemes until the Markov bottle’s water level is no longer falling, which means water migration is adequate [23]. After the freezing process, the expansive soil columns thawed at room temperature (+15°C) until reaching a stable thermal state. The displacement transducer was placed on the top heat exchanger to monitor the soil columns’ vertical deformation all through the freezing-thawing test. The water valve was opened during the whole freeze-thaw process to allow external water supplement [24].

After thawing, cut the thawed expansive soil columns into nine segments (4 cm thick) along with the depth. Cutting rings are used to prepare the standard direct shear sample (20 mm height and 61.8 mm diameter) in each segment. After standard direct shear sample preparation, each segment’s water content was determined by the drying method on the remaining expansive soils. Each segment’s dry density was calculated using phase relationships based on the measured water content and the measured cutting sample weight.

A quick direct shear test with a constant strain rate of 0.08 mm/min was adopted during the test to avoid soil structure change in the freeze-thaw process. A series of low normal stresses (25, 50, 75, and 100 kPa) were utilized in the direct shear test to reflect the actual stress state of shallow expansive soil [25]. The direct shear test terminated as a shear displacement was 4.0 mm if a peak reading was shown.

Table 3: Freezing temperature gradients and corresponding temperature combinations.

| Scheme no. | Top (°C) | Bottom (°C) | Temperature gradient (°C·cm⁻¹) |
|------------|----------|-------------|---------------------------------|
| F1         | −5       | 1           | 0.17                            |
| F2         | −10      | 1           | 0.31                            |
| F3         | −15      | 1           | 0.44                            |
in the dynamometer; otherwise, it is stopped at the shear displacement of 6.0 mm. The relationship curves between shear stress and shear displacement were plotted when the direct shear tests were finished. The peak value of shear stress on the relationship curve between shear stress and shear displacement was appointed as the shear strength of each sample. If there was no peak value, the shear stress corresponding to the shear displacement of 6 mm was taken as shear strength. The shear strength parameters (cohesion and internal friction angle) were obtained from the strength envelopes obtained based on the stress-strain curves. The direct shear test under the same conditions for identical expansive soil samples without freeze-thaw effect was also carried out for comparison.

2.4. Data Analysis Method. A dimensionless parameter, vertical deformation rate ($\varepsilon$), was defined to reflect the vertical deformation ratio after freezing or thawing and calculated by

$$\varepsilon = \frac{z_p - z_0}{h_0} \times 100\%, \quad (1)$$

where $\varepsilon$ is the vertical deformation ratio, $z_p$ is the reading of the displacement sensor after freezing or thawing process, $z_0$ is the initial reading of the displacement sensor at the beginning of the testing, and $h_0$ is the original height of the soil sample. When $\varepsilon$ is positive, it refers to volume expansion; if contrary, it represents volume contraction.

The shear strength ratio ($S$-ratio), cohesion ratio ($C$-ratio), internal friction angle ratio ($A$-ratio), water content ratio ($W$-ratio), and dry density ratio ($D$-ratio), respectively, were calculated by equation (2). The ratio greater than 1.0 indicates the parameter increase, and correspondingly, the ratio less than 1.0 indicates a decrease in the parameter due to freeze-thaw:

$$S, C, A, W, D = \frac{\text{data after freeze - thaw process}}{\text{data of the unfrozen soil}}. \quad (2)$$

3. Results

3.1. Frozen Characteristics of Expansive Soil Columns. During the freezing process, the temperature dynamic characteristics of each expansive soil column were different (Figure 2). After freezing for 24 h, 48 h, and 72 h, expansive soil columns' temperature profiles corresponding to schemes F1, F2, and F3 were all linearly distributed along with the depth, which indicates that the thermal states of the soil columns have been reached stable. The expansive soil columns' frozen depths reached 21 cm, 26 cm, and 29 cm, corresponding to freezing schemes F1, F2, and F3, respectively. All the expansive soil columns were divided into the frozen zone and unfrozen zone according to the freezing point of $-1.64^\circ$C, considering the thickness of the frozen fringe is very small, usually ranging from 1.5 mm to 8 mm [26, 27]. Each expansive soil column's temperature dynamic characteristics were also varied with freezing schemes (Figure 3). After thawing for 18 h, 32 h, and 48 h, expansive soil columns' temperature profiles corresponding to scheme F1, F2, and F3 were all reached stability.

The expansive soil columns' water content profiles exhibited significant differences along with depth after thawing and varied with different freezing schemes (Figure 4). The water content of each soil columns all increased at different depths. The water content increased significantly in the frozen soil area while it increased little in the unfrozen area. The free water was sucked from the Markov bottle to the bottom of the soil columns; then, soil pore water migrated from the unfrozen zone to the frozen zone under the temperature gradients during the freezing process. The expansive soil column's average water content was 21.7%, 23.6%, and 22.3%, corresponding to freezing schemes F1, F2, and F3, respectively. The water content at the frozen and unfrozen zone interface was highest except for the bottom of the soil columns and that was 23.3%, 28.2%, and 26.9%, corresponding to freezing schemes F1, F2, and F3, respectively. The amount of water migration was from scheme F2, F1, and F3 in the descending order.

The expansive soil columns' dry density profiles also exhibited significant differences along with the depth after thawing and varied with different freezing schemes (Figure 5). The dry densities were generally decreased and less than the initial dry density (1.60 g/cm$^3$) at different depths. The mean dry densities were 1.57 g/cm$^3$, 1.55 g/cm$^3$, and 1.56 g/cm$^3$, corresponding to freezing schemes F1, F2, and F3. The minimum dry densities were 1.52 g/cm$^3$, 1.50 g/cm$^3$, and 1.51 g/cm$^3$, corresponding to freezing scheme F1, F2, and F3, obtained at the interfaces of the frozen and unfrozen zones for each specimen. Changes in dry density mean changes in soil structure, and the decrease in dry density means that the soil structure becomes loose.

The deformation rates of each expansive soil column were positive both in the freezing and thawing processes, and they were also varied with freezing schemes (Figure 6). For schemes F1, F2, and F3, the deformation rates were 4.1%, 5.3%, and 2.8% after the freezing process and 2.9%, 3.6%, and 2.4% after the thawing process. Expansive soil columns showed a frost heave phenomenon during the freezing process. The expansive soil columns' deformation showed a slight decrease during the thawing process than the freezing process. The expansion deformation that existed after the thawing process indicates the expansive soil columns' structure has been changed.

3.2. Shear Strength Characteristics. Shear strength of the expansive soil columns at different depths, under different freezing schemes, and different normal stress after the freeze-thaw process are listed in Table 4; besides, the expansive soil samples' shear strength without freeze-thaw effect is 32.9 kPa, 58.3 kPa, 91.3 kPa, and 119.3 kPa, corresponding to the normal stress of 25 kPa, 50 kPa, 75 kPa, and 100 kPa, respectively. The shear strength ratio ($S$-ratio), water content ratio ($W$-ratio), and dry density ratio ($D$-ratio) were calculated by equation (2).

The $S$-ratios showed significant differences along with each expansive soil column's depth and were less than 1.0 (Figure 7). The $S$-ratios decreased rapidly at the bottom of
the expansive soil column. The minimum $S$-ratios under different freezing schemes were all obtained at the interfaces of the frozen zones and unfrozen zones. Taking the normal stress of 50 kPa as an example, the average $S$-ratios of the expansive soil in the frozen zone was 0.82, 0.75, and 0.77 and those in the unfrozen zone were 0.65, 0.64, and 0.66, and the minimum $S$-ratios were 0.64, 0.63, and 0.65 corresponding to freezing schemes F1, F2, and F3, respectively. The results indicate that the shear strength reduction in the frozen zone is smaller than that in the unfrozen zone. Another phenomenon was that the average decrease in $S$-ratios of the expansive soil in frozen zones was F2, F1, and F3 in the descending order.

The relationship curves between $S$-ratio (normal press 25 kPa and 100 kPa) and $W$-ratio and $D$-ratio were plotted in Figures 8 and 9. There is a high correlation between the $S$-
ratios and $W$-ratios in the variation law. All the $S$-ratios’ reduction in the frozen zone was relatively small, corresponding to the small increase in the $W$-ratios’. In the unfrozen zone, the $S$-ratios showed a substantial reduction, corresponding to the $W$-ratios’ substantial increase. In particular, both the $S$-ratios and $W$-ratios achieved minimum and maximum values at the freeze-thaw interface of the frozen and unfrozen zones, respectively. However, the correlation between the $S$-ratios and $D$-ratios in the frozen zone was low (Figure 9). All the average $D$-ratios in the unfrozen zones was 0.96, and in the frozen zones, it was 0.98, corresponding to freezing schemes F1, F2, and F3. Nevertheless, the average $S$-ratios in the frozen zone was 0.82, 0.75, and 0.77 and those in the unfrozen zone were 0.65, 0.64, and 0.66.

3.3. Shear Strength Parameter Characteristics. The cohesion ratio ($C$-ratio) and the internal friction angle ($A$-ratio) at the expansive soil columns’ different depths after the freeze-thaw process were calculated by equation (2) and plotted in Figures 10 and 11. Note that the expansive soil samples’ cohesion and internal friction angle without freeze-thaw effect was 34.1 kPa and 22.1°, respectively. Both of the $C$-ratios and $A$-ratios were less than 1.0 at the whole expansive soil columns, but they were varied with freezing schemes. For freezing schemes F1, F2, and F3, the average $C$-ratios were 0.72, 0.64, and 0.69, while the average $A$-ratios were 0.97, 0.96, and 0.96, respectively. The reduction
Table 4: Shear strength of expansive soil columns at different depths under different normal stress after the freeze-thaw process.

| Depth (cm) | 25 kPa | 50 kPa | 75 kPa | 100 kPa |
|-----------|--------|--------|--------|--------|
|           | F1     | F2     | F3     | F1     | F2     | F3     | F1     | F2     | F3     | F1     | F2     | F3     |
| 2         | 27.9   | 25.4   | 27.1   | 51.7   | 47.4   | 49.1   | 82.5   | 76.8   | 80.0   | 111.2  | 104.1  | 108.4  |
| 6         | 27.9   | 24.6   | 26.9   | 51.7   | 47.0   | 49.4   | 82.9   | 76.4   | 80.1   | 110.8  | 103.7  | 107.5  |
| 10        | 27.9   | 24.6   | 26.7   | 50.2   | 46.0   | 48.3   | 81.3   | 75.7   | 79.2   | 109.6  | 102.7  | 105.8  |
| 14        | 27.4   | 23.0   | 25.3   | 49.9   | 43.6   | 45.6   | 81.7   | 74.7   | 77.8   | 109.4  | 101.1  | 103.9  |
| 18        | 25.7   | 22.4   | 23.6   | 47.1   | 42.4   | 44.6   | 79.5   | 72.9   | 74.4   | 106.9  | 98.6   | 101.5  |
| 22        | 20.8   | 22.1   | 23.2   | 37.8   | 40.7   | 43.2   | 62.4   | 66.5   | 72.5   | 84.9   | 89.0   | 99.3   |
| 26        | 20.4   | 20.0   | 22.2   | 38.0   | 37.2   | 42.0   | 63.3   | 61.0   | 71.0   | 86.5   | 84.0   | 95.3   |
| 30        | 20.5   | 19.8   | 20.7   | 38.1   | 37.6   | 38.0   | 63.6   | 61.5   | 62.2   | 86.9   | 86.4   | 86.1   |
| 34        | 20.4   | 20.0   | 20.6   | 38.1   | 37.3   | 38.4   | 63.7   | 61.4   | 63.0   | 87.0   | 85.2   | 86.2   |

in C-ratios was much higher than the A-ratios. The mean C-ratios in the frozen zones were 0.79, 0.74, and 0.66 but those in the unfrozen zone were 0.58, 0.54, and 0.50, corresponding to freezing schemes F1, F2, and F3. Three obvious inflection points between the frozen and non-frozen regions in the C-ratios curves can be seen in Figure 10. The A-ratios changes little in both frozen and unfrozen zones (Figure 11).

4. Discussion

4.1. Relationship between Shear Strength and Frozen Characteristics. The minimum shear strength in the expansive soil columns occurred at the soil layer below the frozen and unfrozen zone interface. The reduction of shear strength is attributed to the soil structure damage during the freezing process [28]. The formation of irregular pores and large cracks due to water migration and water phase change in soil particles and pores is the main reason for soil structure damage [29–31]. The more significant the water migration, the more the considerable changes in the soil structure, which induce more significant degradation of soil shear strength. The inhomogeneity of shear strength along depth is consistent with that of water content along with the depth. Changes in soil water potential in response to temperature gradients during the freezing process induce a situation in which soil water migrates from the unfrozen zone to the frozen zone [32, 33]. Moreover, the necessary condition for water migration is that the water migration rate must be higher than the freezing rate [34]. A higher freezing rate formed by a large temperature gradient can lead to an insufficient moisture migration due to pore water migration.
freezing rapidly. On the contrary, the lower freezing rate cannot provide enough driving force for pore water migration. The amount of water migration under moderate freezing temperature gradient (scheme F2) is most considerable than the low and high freezing temperature gradients (Figure 4), and similar observations have also been reported by Talamucci [35]. Thus, the shear strength degradation for freezing scheme F2 was more evident than that of F1 and F3 accordingly.

Moreover, this is similar to the findings reported by Steiner et al. [9], who reported that the most substantial change in shear strength of an illite clay occurred between −5°C and −10°C due to the formation of large ice lenses at such moderate temperature range, and these lenses acted as
the primary failure plane in the thawed soil and caused a reduction in shear strength. During the freezing process, a violent water-ice phase transition occurs at the interface of frozen and unfrozen zones (freezing front) in which has the highest ice content. As the upper soil layers beg into thaw, the melted water does not easily infiltrate into the lower layer due to the compacted soil’s low permeability. Therefore, the lowest shear strength occurred at the frozen-unfrozen zones interface corresponding to the maximum water content. In an open system, due to water replenishment from the soil columns’ bottom, the water content at soil columns’ bottom was relatively high, and the bottom soil softens, so the shear strength at the bottom decreases significantly.

In addition to water content, dry density is generally considered another critical factor that has a significant effect on shear strength, and the shear strength increases with the increase of dry density under other fixed conditions [36, 37]. The dry density changes (the average $D$-ratio 0.98) were minimal in the frozen zones and unfrozen zones. Therefore, the effect of changes in dry density on shear strength is limited. The expansive soil columns’ volume has slightly expanded after thawing, consistent with the dry density’s slight decrease. Frost heave and thawing settlement are typical soil deformation characteristics during the freeze-thaw process [38, 39]. It is worth noting that the expansive soil columns showed a phenomenon of frost expansion and thawing expansion (Figure 6). Different observations have been found in a compacted expansive soil reported by Lu et al. [17], who found that, for expansive soils with lower water content, the overall deformation decreases upon freezing but increases thawing. Hamilton [40] reported that regardless of whether expansion or contraction during the freezing process, the volumes presented expansion for compacted soil at the end of the thawing process. Viklander [41] reported that the volume typically decreased for initially loose soil and increased for a dense soil after freeze-thaw, and independent of whether the initial soil structure was loose or dense, a constant “residual” void ratio could be obtained after 1–3 cycles. The residual deformation between the freezing and thawing stages indicates that the soil structure has been changed, directly affecting the thawed soil’s shear strength.

4.2. Shear Strength Parameter Characteristics after Freeze-Thaw. The cohesion of expansive soil that underwent freeze-thaw effect was significantly reduced, while the internal friction angle was slightly reduced and not noticeable. The decrease of cohesion is mainly induced by weakening of soil cementation, which results from cracks developed by ice lens growth during the freezing process [42, 43]. The cohesion decreased after freeze-thaw has been recognized by scholars...
[9, 16, 19]. The change of internal friction angle of the soils after freeze-thaw is still controversial. The increased results, decreased results, and consistent results all can be found in the literature [12, 14, 15, 18]. There are two opposite effects on the internal friction angle of the soils subjected to freeze-thaw. On the one hand, the increase of contact points between soil particles induced by the freeze-thaw effect can increase the internal friction angle; on the other hand, the lubrication effect caused by the increase of water content can reduce the internal friction angle [44]. The final results are the interaction of the two opposite effects mentioned above. As for expansive soil, Tang et al. [5] found that the internal friction angle decreased significantly after one freeze-thaw cycle, which is different from our results. The difference between the results may be attributed to the differences in test methods.

Only the initial water content of 18% was studied in the present work, and it is necessary to study the influence of initial water content and water replenishment on shear strength in future research. Besides, the influence of freeze-thaw cycles on shallow expansive soil’s shear strength also needs further investigation.

5. Conclusions

The shear strength characteristics of shallow expansive soil due to the freeze-thaw effect with water supplement were investigated using an integrated approach that incorporated the unidirectional freezing test and direct shear test. We found that shallow expansive soil’s shear strength after the freeze-thaw effect showed significant differences with depth and was highly correlated with water content. Furthermore, the most considerable reduction of the shear strength of expansive soil after freeze-thaw was under the moderate freezing temperature gradient. Moreover, the significant decrease in cohesion makes a prominent contribution to reducing the shear strength after freeze-thaw.

Data Availability

The data used to support the findings of this study are available from the corresponding author upon request.

Conflicts of Interest

The authors declare that there are no conflicts of interest regarding the publication of this paper.

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