Research Article

The 4-Parameter Ionospheric Channel Model: Part 1—Theory and Simulation

You cannot significantly improve modeling unless you can significantly improve the methodology of modeling.

—Ilir Progri

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The 1-parameter ionospheric channel model that is based on estimating the ionospheric delay is insufficient to enable the GNSS receiver to reliably acquire and track GNSS signals under severe ionospheric conditions.

In order to solve this problem Giftet Inc. is proposing the 4-parameter ionospheric model. The 4-parameter ionospheric model is based on estimating four parameters: ionospheric delay, Doppler, amplitude, and phase scintillation. The 4-parameter ionospheric model presents a revolutionary advancement of our understanding of the ionospheric effects in the RF signals. In order to accurately estimate the ionospheric Doppler we have performed some of the most amazing mathematical calculations.

Afterwards, the theory and simulation, presented in this paper, are completely verified and accurately tested, the implementation thereof in a single frequency GNSS receiver will enable it to perform just as good as or even better than a differential GNSS receiver.

Index Terms—4-parameter, ionospheric, channel, model, delay, Doppler, amplitude, phase, scintillation, MATLAB.

1 Introduction

Ionosphere: is the part of the atmosphere that: (1) stretches from 50 to 1,000 km (31.06855 to 621.371 mi; 164,042 to 3,280,840 ft), (2) is ionized (i.e.; contains ions and electrons) by solar radiation, typically overlaps both the exosphere and the thermosphere; (3) forms the inner edge of the magnetosphere; (4) greatly influences, for example, radio propagation on the Earth; and (5) is responsible for auroras (see Figs. 1 and 2) [1], [2].

The radio window consists of frequencies which range from
about 5 MHz (5 million Hz) to 30 GHz (30 billion Hz). The low-frequency end of the window is limited by signals being reflected by the ionosphere [2] back into space or Earth, while the upper limit is caused by absorption of the radio waves by water vapor [3] and carbon dioxide in the atmosphere. As atmospheric conditions change the radio window can expand or shrink. On clear days with perfect conditions signals as high as 300 GHz have been detected (see Figs. 1-3) [1].

Normally radio “ground waves” propagate along the surface as creeping waves; i.e., they are only diffracted around the curvature of the earth, which is the reason why early long distance radio communications used long wavelengths with the best known exception of the HF (3-30 MHz) waves being reflected by the ionosphere [2], [4].

The reduced refractive index [5] due to lower densities at the higher altitudes in the Earth’s atmosphere bends the signals back toward the Earth. When signals encounter a higher refractive index [5] layer duct, tend to remain in that layer because of the reflection [6] and refraction [6] encountered at the boundary with a lower refractive index material and vice versa. In some weather conditions, such as inversion layers, density changes so rapidly that waves are guided around the curvature of the earth at a constant [8] altitude [4].

**Ionospheric reflection** : is a bending, through a complex process involving reflection and refraction [6], of electromagnetic waves propagating in the ionosphere [2] back toward the Earth [6]. The amount of bending depends on the extent of the penetration (which is a function of frequency), the angle of incidence, the polarization of the wave, and ionospheric conditions, such as the ionization density. It is negatively affected by incidents of ionospheric absorption [6].

Ionospheric absorption/ dispersion (or ISAB/ISD) is the scientific name for absorption/ dispersion occurring as a result of the interaction between various types of electromagnetic waves and the free electrons in the ionosphere [2], which can
interfere with radio transmissions which results in a significant loss/dispersion of the signal power [9].

Ionospheric absorption/dispersion is of critical importance when radio networks, telecommunication systems or interlinked radio systems are being planned, particularly when trying to determine propagation conditions [9].

The ionosphere [2] can be described as an area of the atmosphere in which radio waves on shortwave bands are refracted or reflected back to earth (see Figs. 1-4). As a result of this reflection, which is often the key in the long-distance propagation of radio waves, some of the shortwave signal strength is decreased. In this regard, ISAB/ISD is the primary limiting factor in radio propagation [9].

ISAB/ISD is only a factor in the period of the day where radio signals travel through the portion of the ionosphere [2] facing the sun. The solar wind and radiation cause the ionosphere [2] to become charged with electrons in the first place. At night, the atmosphere becomes drained of its charge, and radio signals can go much further with less loss of signal. In particular, low wavelength signals that would be attenuated to nothing during the day will be received much further away at night [9]. The specific amount of attenuation [10] can be derived as a function of the inverse-square law. The lower the frequency is, the greater the attenuation becomes [9], [10].

Ionospheric errors are caused by delays or changes in the RF signals’ time of arrival, Doppler shift, amplitude, and phase traversing the ionosphere [2], whose electron content is a function of the amount of incident solar radiation. Therefore, the ionospheric delay [11]-[13], Doppler shift, amplitude, and phase change with latitude, longitude, time of day, season (or day) of the year and, year of the millennium, following the 11.1 years solar cycle, with higher values by day (at around 14h00m local time), during the summer and at the peaks of the solar cycle. From these three periodic changes, the dominant one is diurnal variability, following the variation in incident solar radiation [13], [14].

Superimposed on these periodic changes, severe magnetic storms occur a few times (generally not more than four times) during each 11.1 years solar cycle, causing extreme delays on GPS signals, in addition to amplitude fading and scintillation. These magnetic storms affect mainly the auroral latitudes (around the geomagnetic poles) and “for some unknown reason (…) occur more frequently during the declining phase of the solar cycle” (see Figs. 1-4) [13], [14].

Ionospheric propagation vertical delays are typically 20-30 m during the day and 3-6 m at night [16]. These errors are entirely compensated by DGPS and their lack of correlation with time, for periods of tens of minutes, is very low, because the ionospheric delay [11] does not change significantly on such timescales—except in the case of major magnetic storms during which the TEC changes rapidly. Dusk [17] and dawn [18] are the periods when the temporal lack of correlation of the ionospheric error is generally higher, because the ionosphere [2] re-configures itself, but even then the de-correlation is not significant. Although the de-correlation of the ionospheric errors with time has a much smaller effect than its de-correlation with distance, this is the error which contributes the most to the growth of DGPS errors with age of corrections [13], [14].

Naturally, one might ask, why does not a single frequency GNSS receiver perform just as good as a DGPS receiver? I.e., why is not a single frequency GNSS receiver capable of performing the same corrections as a DGPS receiver? The answer lies in this fact that the channel model of a single frequency GNSS receiver is currently based on a 1-parameter ionospheric channel model. The 1-parameter ionospheric channel model that is based on estimating the ionospheric delay is insufficient to enable the GNSS receiver to reliably acquire and track GNSS signals under severe ionospheric conditions.

In order to solve this problem Giftet Inc. is proposing the 4-parameter ionospheric model. The 4-parameter ionospheric model is based on estimating four parameters: ionospheric delay, Doppler, amplitude, and phase scintillation. The 4-parameter ionospheric model presents a revolutionary advancement of our understanding of the ionospheric effects in the RF signals. In order to accurately estimate the ionospheric Doppler, we have performed some of the most amazing mathematical calculations.

This paper presents one of the most daunting tasks ever undertaken by a single scholar to very accurately model, analyze, and simulate the entire description of the ionospheric channel effects.

Afterwards the theory and simulation, presented in this paper, are completely verified and accurately tested, the implementation thereof in a single frequency GNSS receiver will enable it to perform just as good as or even better than a differential GNSS receiver (Progr 2018, [19]).
2 Description of RF Geolocation Propagation Channel Model

When an RF signal propagates from an RF transmitter (typically an RF (or GNSS) satellite orbiting the Earth at approximately 20,200 km altitude from the surface of the Earth) to a RF receiver antenna (typically most receivers are located from 0 to 30 km from the surface of the Earth), it suffers from degradation effects, such as, free space path loss, absorption and refraction from the atmosphere, reflection and masking from surrounding objects such as trees and buildings, jamming, and environmental noise as shown in Figs. 5-8.

2.1 Description of an ideal RF geolocation transmitted signal

Initially, we start with the description of the ideal RF geolocation propagation channel model.

The analytical expression of the ideal transmitted RF signal (or IS) can be written as the product of two components: (1) ideal baseband signal (or IBS), \( s = S_0 e^{j\phi_0} \), and (2) ideal carrier signal (or ICS), \( s_2(t; 1, 0, \omega_c) = 1 \times e^{j(0 + \omega_c t)} \).

\[
s(t) = S_0 e^{j(\phi_0 + \omega_c t)} = S_0 e^{j\phi_0} 1 \times e^{j(0 + \omega_c t)}
\]

\[
= S_0 e^{j\phi_0} 1 \times e^{j(0 + \omega_c t)} = s_1 s_2(t; 1, 0, \omega_c)
\]

The unknowns in (1) are as follows: (1) \( t \) is the time in [s]; (2) \( S_0 \) is amplitude in [V]; (3) \( \phi_0 \) is phase in [rad]; (4) \( \omega_c \) is the radiant carrier frequency in [rad/s] of the transmitted GPS (or GNSS) signal respectively.

The first main assumption is as follows: IBS is a time-invariant quantity; i.e., it only contains the desired signal amplitude, \( S_0 \), and (initial) phase, \( \phi_0 \); ICS contains the desired signal carrier [radiant] frequency, \( \omega_c \), and time (or time stamp), \( t \).

The second main assumption is that, in the description of IBS itself the desired signal amplitude, \( S_0 \), and phase, \( \phi_0 \), are also independent; in the description of ICS the desired signal carrier [radiant] frequency, \( \omega_c \), and time stamp, \( t \), are also independent; however, their product produces the accumulated carrier phase, \( \phi \); and since, \( \omega_c \) is constant the product (or correlated) information in the accumulated carrier phase, \( \phi \), gives precise information about the time stamp, \( t \).

Third, in an ideal environment the ideal RF geolocation propagation channel model can be completely characterized by means of four main parameters: (1) desired signal amplitude, \( S_0 \), (2) initial phase, \( \phi_0 \), (3) carrier [radiant] frequency, \( \omega_c \), and (4) time stamp, \( t \).

Fourth, in an ideal (i.e., linear and homogenous) environment it is straightforward to separate these four independent parameters; namely, the signal amplitude, \( S_0 \), initial phase, \( \phi_0 \), frequency, \( \omega_c \), and time stamp (or just time), \( t \); the first and the most important attribute of modeling the received RF signal is to come up with a novel, original, methodical, and useful method of separating, these four independent parameters, \( S_0 \), \( \phi_0 \), \( \omega_c \), and \( t \), from other nonlinear and heterogeneous parameters of the environment; hence, the description of the a realistic RF geolocation propagation channel model.

2.2 Description of free-space path loss and ideal propagation channel model

Before we consider the realistic propagation, channel let us first present the description of free-space path loss channel model.

The free space path loss, \( Q_0 \) , [20] (or the lossless environment parameter) provides the relation due to inverse square law [21] as given by

\[
Q_0 = \frac{P_r}{P_t}
= \frac{s}{S4\pi R^2 4\pi}
= \left(\frac{\lambda}{4\pi R}\right)^2
= \left(\frac{c}{4\pi R\omega_c}\right)^2
= \left(\frac{c}{4\pi Rf}\right)^2
\]

where (1) \( P_t \) is the transmitted signal power in [W] [22]; (2) \( P_r \) is the received signal power in [W]; (3) \( S \) is the spreading of the electromagnetic energy in [W/m²]; (4) \( \lambda \) is the wavelength [m]; (5) \( R \) is the transmitter receiver distance (or the line-of-sight (LOS) vector magnitude) in [m]; (6) \( c \) is the speed of light (SOL) in vacuum (299,792,458 m/s [22]); (7) \( \omega \) is the radiant frequency \( \omega = 2\pi f \) in [rad/s]; and (8) \( f \) is the frequency in [Hz].

In an ideal free space environment, the free space path loss, \( Q_0 \), is the only quantity that completely characterizes or describes the physics of the propagation of the channel model;
i.e., an ideal free space environment is in fact a one channel parameter.

2.3 Description of an ideal RF received signal

Before we consider the realistic received signal let us present the description of an ideal RF received signal (or IRS).

First, IRS or $u_0(t; A, \varphi, \tau, w_c)$ represents the received signal in vacuum; i.e., as if there were no ionosphere, or troposphere, or multipath or anything else; i.e., including only the inverse square-law propagation \[21\] in vacuum, which can be expressed as a product of the ideal received BS (or IRBS), $u_1$, which is only a function of $(A_1, \varphi_1)$ and ideal received CS (or IRCS), $u_2(t)$, which is a function of $(A_2, \varphi_2, \tau, w_c)$

$$u_0(t) = \frac{u_1(t)}{\text{IRBS}} \frac{u_2(t)}{\text{IRCS}}$$

$$= A(t; \tau, w_c) e^{j\varphi_0 + \omega_c (t-\tau)}$$

$$= A(t; \tau, \omega_c) e^{j[\varphi_0 + \omega_c (t-\tau)]}$$

$$= A(t; \tau, \omega_c) e^{j[\varphi_0]} e^{j\omega_c (t-\tau)}$$

$$= A(t; \tau, \omega_c) e^{j[\varphi_0]} s_2(t-\tau; 1, 0, \omega_c)$$

$$= A(t; \tau, \omega_c) e^{j[\varphi_0]} s_2(t-\tau; 1, 0, \omega_c)$$

$$= A(t; \tau, \omega_c) e^{j[\varphi_0]} \delta A(t; \tau, \omega_c) s_2(t-\tau; 1, 0, \omega_c)$$

(3)
where $A(t; \tau, \omega_c)$ can be computed as the square root of the transmitter-receiver antenna gain, $G_{tr}$, and free space path loss, $Q_0$, given by (2)

$$A(t; \tau, \omega_c) = \sqrt{G_{tr}}Q_0$$

$$= \sqrt{G_{tr}} \frac{c}{2\omega_c \tau}$$

$$= \sqrt{G_{tr}} \frac{c}{2\omega_c \tau(t)}$$

$$= \left[ A_0 \equiv \frac{\sqrt{G_{tr}}}{2} \right] ^{\delta A_0(t) \equiv \frac{1}{\omega_c \tau}}$$

Finally, equating both sides of (8), (8)-(10) we finally obtain

$$A_z(t; \tau, \omega_c) = \delta A_0(t; \tau, \omega_c) = (\phi_c = \omega_c \tau)^{-1}$$

In an ideal RF propagation channel, we should be able to precisely know (or estimate) four “independent” signal parameters $A(t)$, $\varphi(t)$, $\tau(t)$, and $\omega_c$ as follows:

$$A(t) = \frac{A_0 \equiv A_z(\tau(t), \omega_c)}{\text{IRBS}}$$

$$\equiv \frac{[\varphi_c(t) = \omega_c \tau(t)]^{-1}}{\text{IRCS}}$$

$$\equiv \frac{1 \equiv 1 \times 1}{\text{ERS} \equiv \text{IRBS} \times \text{ERCS}}$$

$$\varphi(t) = \frac{\varphi_0 \equiv \varphi_1 + \delta \varphi_0 = \varphi_2(t(t), \omega_c) + 0 \equiv 0 + 0}{\text{IRBS}}$$

$$\equiv \frac{\omega_c = 0}{\text{ERCS}}$$

$$\equiv \frac{0 \equiv 0 + 0}{\text{ERS} \equiv \text{IRBS} + \text{ERCS}}$$

$$\equiv \frac{\tau(t) = 0}{\text{ICs}}$$

$$\equiv \frac{0 \equiv 0 + 0}{\text{ERS}}$$

$$\equiv \frac{0 \equiv 0}{\text{IRCS}}$$

$$\equiv \frac{\omega_c = 0}{\text{ERS}}$$

$$\equiv \frac{0 + 0}{\text{IRCS}}$$

In an ideal RF propagation channel, the received signal amplitude is inversely proportional with the phase velocity $\phi_c(t)$ or the product of $\omega_c \tau(t)$. The ideal received signal phase decreases linearly with the phase velocity $\phi_c(t)$ or the product of $\omega_c \tau(t)$. We can clearly see that to make a correct estimation of the signal amplitude and phase depends on a correct estimation of both the signal time delay (i.e., or distance) and frequency.

Next, let us describe what happens to these parameters when we consider a realistic RF geolocation propagation channel model.

### 2.4 Description of a realistic RF geolocation propagation channel model

The analytical expression of the received RF GPS (or GNSS) signal, $u^z(t)$, which is a function of $(A^z, \varphi^z, \tau^z, \omega^z)$ and has undergone propagation through the ionosphere, troposphere (all of which are included, later on, and of the multipath model) can be written as a product of two independent signals: received BS (or RBS), $u^z_1 = A^z_1 e^{j \phi^z_1}$, and received CS or (RCS), $u^z_2 = A^z_2 e^{j \phi^z_2}$, and received CS or (RCS),

$$u^z(t) = A^z_1 e^{j \phi^z_1} + A^z_2 e^{j \phi^z_2}$$

$$= u^z_1 + u^z_2$$

Where
On the other hand, the received RF GPS (or GNSS) signal, \( u_s(t) \), can also be expressed as a product of signal amplitude, \( A_s(t) \), and exponential phase, \( e^{j[\phi_s(t) + \omega_s(t)\tau - \frac{\omega_s}{c}(t - \tau) + \omega_s(t)\tau]} \), as follows

\[
u_s(t) = A_s(t) e^{j[\phi_s(t) + \omega_s(t)\tau - \frac{\omega_s}{c}(t - \tau)]}
\]  

(19)

Since (17) and (19) really represent the same thing equating both sides we obtain

\[
A^s(t) = A^i(t) e^{j[\phi^i(t) + \omega^i(t)\tau]}
\]  

(20)

\[
\phi^s(t) = \phi^i + \phi^e(t)
\]  

(21)

\[
\rho = R + cD^e
\]  

(22)

Although (20) and (21) are an accurate description of the received RF GPS (or GNSS) signal, \( u^s(t) \), amplitude, \( A^s(t) \), and phase, \( \phi^s(t) \), they do not show the relationship between the ideal transmitted signal, \( s(t) \), and the received signal, \( u^s(t) \). In order to accomplish that we need to express the received RF GPS (or GNSS) signal, \( u^s(t) \), as a product of the ideal received signal (or IRS), \( u_0(t) \), which is a function of \( (A, \phi, \tau, \omega_e) \), and environmental received signal (or ERS), \( \delta u_e(t) \), which is a function of \( (\delta A_e^i, \delta \phi_e^i, D_e^i, \omega_e^i) \), as follows
\[ u^s(t) = u_d(t) \frac{\delta u^s_\text{ERS}(t)}{\delta u^s_\text{IRS}} \]  

(26)

Where

\[ \delta u^s_\text{ERS}(t) = \delta A^s_\text{ERS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) \]  

(27)

\[ \delta A^s_\text{ERS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) = A^s(t; \tau^s, \omega^s) A(t; \tau, \omega_c)^{-1} \]  

(28)

\[ \delta \phi^s_\text{ERS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) + \omega^s_\text{ERS}(t - D^s_\text{ERS}) = \left[ \phi^s(t; \tau^s, \omega^s) + \omega^s \left( t - \frac{D^s_\text{ERS}}{c} \right) \right] - \phi_0 - \omega_c (t - \tau) \]  

(29)

The description of the ERS is still not enough so we will need to further separate the ERS, \( \delta u^s_\text{ERS}(t) \), into two components as follows:

\[ \frac{\delta u^s_\text{ERS}(t)}{\delta u^s_\text{IRS}} = \frac{\delta u^s_\text{IRS}(t)}{\delta u^s_\text{ERBS}} \]  

(30)

where

\[ \frac{\delta u^s_\text{IRS}(t)}{\delta u^s_\text{ERBS}} = \frac{\delta A^s_\text{IRS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) e^{i \delta \phi^s_\text{IRS}(t)}}{\delta A^s_\text{ERBS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) e^{i \delta \phi^s_\text{ERBS}(t)}} \]  

(31)

\[ \frac{\delta u^s_\text{ERBS}(t)}{\delta u^s_\text{ERCS}} = \frac{\delta A^s_\text{ERBS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) e^{i \delta \phi^s_\text{ERBS}(t)}}{\delta A^s_\text{ERCS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) e^{i \delta \phi^s_\text{ERCS}(t)}} \]  

(32)

Finally, in a realistic RF channel we should be able to precisely know (or estimate) the four “independent” signal parameters: \( A^s(t), \phi^s(t), \tau^s(t), \) and \( \omega^s(t) \) as follows:

\[ A^s(t) = A(t) \frac{\delta A^s_\text{ERS}(t; D^s_\text{ERS}, \omega^s_\text{ERS})}{\delta A^s_\text{IRS}(t; D^s_\text{ERS}, \omega^s_\text{ERS})} \]  

(33)

\[ \phi^s(t) = \phi(t) + \phi^s_\text{IRS}(t; D^s_\text{ERS}, \omega^s_\text{ERS}) \]  

(34)

\[ \tau^s(t) = \frac{\rho(t)}{c} = \frac{R(t) + D^s_\text{IRS}(t)}{c} \]  

(26)

\[ \omega^s(t) = \frac{0 + 0 + \frac{R(t)}{c} + D^s_\text{IRS}(t)}{\text{ERCS}} \]  

(35)

Equations (33) through (36) provide the entire description of the four most important “independent” signal parameters, \( A^s(t), \phi^s(t), \tau^s(t), \) and \( \omega^s(t) \), of a realistic RF propagation channel model.

In a realistic RF propagation channel, the realistic received signal amplitude is inversely proportional with the phase velocity \( \phi_c(t) \) or the product of \( \omega_c \tau(t) \) and it is a complicated function of the ERCS or \( (D^s_\text{ERS}, \omega^s_\text{ERS}) \).

The received signal phase decreases linearly with the phase velocity \( \phi_c(t) \) and it is a complicated function of ERCS or \( (D^s_\text{ERS}, \omega^s_\text{ERS}) \). We can clearly see that a correct estimation of the signal amplitude and phase depends on a correct and complicated estimation of both the signal delay and frequency or their respective errors. It is not only important to estimate the signal delay error, \( D^s_\text{ERS}, \omega^s_\text{ERS} \), but it is just as important to estimate the signal frequency error, \( \omega^s_\text{ERS} \), which is the derivative of the signal delay error.

We should also mention that this is a first order approximation; i.e., \( f^s_\text{IRCS}(t) = dD^s_\text{IRS}(t)/dt \) is the only correction applied to the received carrier frequency. If this approximation is not good enough, based on verification of the fidelity of the model from real data, then later on the model can be enhanced to include second or higher order approximations. However, as we are going to see throughout the derivations of this paper, because the computation of \( f^s_\text{IRCS}(t) \) is in general very laborious, for the purposes of this paper, we are going to restrict all derivations based on this main assumption of the first order approximation. An example of higher order ionosphere errors is given by [24].
Next, we consider the description of ionospheric RF geolocation propagation channel model.

3 Description of Ionospheric RF Geolocation Propagation Channel Model

The description of the ionospheric RF geolocation propagation channel model contains the description of the ionospheric refractive index and refractivity, delay, Doppler frequency, and amplitude and phase scintillation.

3.1 Ionospheric refractive index and refractivity

Refractive index: In RF signal propagation, the refractive index or index of refraction, \( n \), of the medium (or media or propagation channel or environment in our case the ionosphere) is a dimensionless number that describes how RF signals propagate through that medium.

In order to derive the refractive index for the ionospheric channel we make the following fundamental assumption. If we treat ionosphere as ionized gas as a perfect dielectric with relative permittivity \( \varepsilon_r \) and relative permeability \( \mu_r = 1 \) its refractive index \( n \), for a signal frequency, \( \pi k \), is defined as the ratio of the SOL \( \frac{c}{v} \) in vacuum (or in medium with refractive index \( n = 1 \)), \( \frac{c}{v} \), and the phase velocity or the SOL (or RF signals) in the medium (or ionosphere) with refractive index \( n \neq 1 \), \( \frac{c}{v} \).

\[
n = \frac{c}{v} = \frac{\varepsilon_r}{\varepsilon} = \frac{\varepsilon_r}{\varepsilon_r + j \varepsilon_{rl}} = \frac{1 - j \frac{\sigma}{\varepsilon \omega_0}}{\varepsilon_r + j \varepsilon_{rl}} \quad (37)
\]

\[
\sigma = \frac{\varepsilon_n}{m_e \omega_0} \quad (38)
\]

\[
\omega_0 = 2\pi f_k \quad (39)
\]

where (1) \( \sigma_d \) is the dielectric conductivity of the ionosphere; (2) \( e \) is the elementary charge \([30]\); (3) \( m_e \) is the electron mass \([30]\); (4) \( \varepsilon_0 \) is the dielectric permittivity of vacuum \([31]\); and (5) \( n_e \) the plasma \([32]\) electron density (PED) in \([\text{el} \cdot \text{m}^{-3}]\).
\( n(f_k \to 0) = -\infty; \quad n(f_k \to \infty) = 1 \)  \hspace{1cm} (43)

Which results in the following values of the refractivity:
\[ N(f_k \to 0) = -\infty; \quad N(f_k \to \infty) = 0 \]  \hspace{1cm} (44)

Since, RF GNSS L5, L2, L3, and L1 signal frequency fall in the RF L-band range of \( \forall f_k \in [f_{c0} = 1 \text{ GHz,} \ f_{c1} = 2 \text{ GHz}] \), then for most practical purposes, typical values of refractive index and refractivity are as follows,
\[ n(f_{c0} = 1 \text{ GHz}) = n_{c0} \approx 1 + N_{c0} \times 10^{-6} \]  \hspace{1cm} (45)
\[ n(f_{c1} = 2 \text{ GHz}) \approx n_{c1} \approx 1 + N_{c1} \times 10^{-6} \]  \hspace{1cm} (46)
\[ N(f_{c0} = 1 \text{ GHz}) \approx N_{c0} \approx -40 \]  \hspace{1cm} (47)
\[ N(f_{c1} = 2 \text{ GHz}) \approx N_{c1} \approx -10 \]  \hspace{1cm} (48)

Below a certain frequency (\( f_{c1} < 1 \text{ GHz} \)), called the critical frequency \( f_{c0} = [5-898.026]^{10} \text{ MHz} \), the ray cannot cross the ionosphere [2], [44] for any elevation angle. Hence, for most practical purposes it is safe to assume that the GNSS RF L-band signals will be able to penetrate the ionosphere and the received GNSS RF L-band signals will be bent as a result of their propagation through the ionosphere by refractive index and refractivity values given by (42)-(48) respectively.

This concludes the description of the ionospheric refractive index and refractivity. Next, we consider the description of the Ionospheric delay.

### 3.2 Description of the Ionospheric delay

**Ionospheric delay** : It is the amount of additional transmission (or propagation) time a signal incurs as it passes through the ionosphere [11], [12]. The amount of delay through the ionosphere varies with the frequency of the signal that passes through the ionosphere, solar activity which influences the electron (and/or ions) content or PED of the ionosphere, time of day, day of the year, year of the millennium, elevation or zenith angles etc.

In general, this delay, \( D_{\text{lt}}(t) \), is defined as the difference between the actual path of the carrier \( S \) and the straight-line path in a vacuum \( L^n \):
\[ D_{\text{lt}}(t) = \int_S \frac{ds}{c} - \int_L \frac{dl}{c} \rightarrow \]  \hspace{1cm} (49)
\[ cD_{\text{lt}} = \int_S n ds - \int_L dl \]
\[
I_s = \int_f(n - 1)ds + \left\{ \int_s ds - \int_f dl \right\}
\]

In terms of distance, after multiplying by the SOL [2], \( c \),
gives two components: (1) \textit{change of refractive delay along path length} and (2) \textit{change of path length} [45].

Because the change of path length must be independent of speed then the following must hold
\[
\{ \int_s ds - \int_f dl \} = 0
\]

Hence,
\[
cD_{lt}^x = \int_s nds - \int_f dl = \int_f(n - 1)ds = 10^{-6} \int_s Nds
\]

Up until now the atmosphere was considered to be horizontally layered and azimuthally symmetric; i.e., refractive index (or refractivity) of the ionosphere is constant (or does not vary) as a function of horizontal position; it only varies as a function of height. Since the refractive index (or refractivity) of the ionosphere is not constant, but varies gradually with both height and horizontal position, ionosphere is known as a gradient-index or GRIN medium and is described by gradient index or azimuthally asymmetric mathematics (see Figs. 9-12) [5], [48], [49]. Although for most applications this assumption is appropriate, the azimuthally asymmetric condition may introduce significant errors in geodetic measurements where high precision is required to eliminate deep signal fading effects due to ionospheric scintillation [46] and ionospheric plasma bubble. The azimuthally asymmetric condition is described as a tilted atmosphere; whose refractivity \( N \) as function of height \( h \) and horizontal position \( x \) reads: (where the gradient vector (or tensor [48], [49]), \( \nabla N \), is defined as below (see Figs. 9-12)):
\[
N(x, s) = N(0,s) + \left[ \nabla N(s) \equiv \frac{\partial N(x,s)}{\partial x(s)} \right] x = 0 \cdot x(s)^{\text{**}}
\]

The total ionospheric delay [11] in an azimuthally asymmetric atmosphere is given by integrating both sides of (54) as follows
\[
D_{lt}^{ax} (\zeta) = 10^{-6} \int_0^\infty [N(x,s) \equiv N(0,s) + \nabla N(s) \cdot x(s)] ds
\]

Where the component of (54) can be written as: \( D_{lt}^x (\zeta) \), is the azimuthally symmetric component of the total ionospheric delay
\[
D_{lt}^x (\zeta) = 10^{-6} \int_0^\infty N(0,s)ds
\]

And, \( D_{lt}^{ax} (\zeta) \), is the azimuthally asymmetric component of the same
\[
D_{lt}^{ax} (\zeta) = 10^{-6} \int_0^\infty \nabla N(s) \cdot x(s)ds
\]

Hence,
\[
D_{lt}^{ax} (\zeta) = D_{lt}^x (\zeta) + D_{lt}^{ax} (\zeta)
\]

The vector \( x(s) \) is in the projected slant direction (projection on the surface) (see Figs. 9-12). Because we have the slant vector equal to
\[
s = m_i(90° - \zeta) x(s) = m_i(\theta) x(s)
\]

And the differential of the slant as
\[
ds = m_i(\zeta) dh
\]

hence, the absolute value of \( x(s) \) is given by
\[
| x(h) | = | x(s) |
\]

\[
\approx \frac{m_i(\zeta)h}{m_i(90°-\zeta)}
\]

\[
= \tan(\zeta')h
\]

Hence, vector \( x(s) \) is approximately given by
\[
\hat{x}(s) \approx \frac{m_i(\zeta)}{m_i(90°-\zeta)} \hat{x}(h) h
\]

First, substituting (59) into (55) yields,
\[
D_{lt}^x (\zeta) = 10^{-6} m_i(\zeta) \int_0^\infty N(0,h)dh
\]

Second, substituting, (61) and (59) into (56) yields,
\[
D_{lt}^{ax} (\zeta) = 10^{-6} \frac{m_i(\zeta)}{m_i(90°-\zeta)} \int_0^\infty \nabla N(h) \cdot x(h)dh
\]

The typical range of variation for the ionospheric delay is very dependent on the frequency but it can move from less than one meter to more than one hundred meters. In general, the \( D_{lt}^x (\zeta) \) integrated along a path joining a satellite \( k \) and a receiver \( A \) can be split up into two contributions: vertical TEC and slant function, in the following way
\[
D_{lt}^x (\zeta) = m_i(\zeta) D_{lt}^x = D_{lt}^x + \Delta_{lt}^{ax}(\zeta)
\]

where \( D_{lt}^x \) is the delay due to vertical TEC
\[
D_{lt}^x = 10^{-6} \int_0^\infty N(0,h)dh \approx \pm \alpha_{TEC} \frac{m_i}{l_k}
\]

and \( \Delta_{lt}^{ax}(\zeta) \) is the slant component
\[
\Delta_{lt}^{ax}(\zeta) = D_{lt}^x [m_i(\zeta) - 1]
\]

The dot in (63) denotes an inner product [48]. The
ionospheric delay [11] in an azimuthally asymmetric atmosphere can be given as the sum of the delay we would have in a symmetric atmosphere and a correction term due to the azimuthally asymmetric ionosphere.

The ionospheric obliquity factor \( m_d[\theta(t)] \) (or \( m_d[\zeta(t)] \)) given by (85) or (86), on the other hand, is independent of the frequency but it is monotonic decreasing/increasing with elevation, zenith \( \theta(t), \zeta(t) \in [0^\circ, 90^\circ] \) (or \( \zeta(t) = 90^\circ - \theta(t) \)) as shown in Fig. 15. The ionospheric azimuth obliquity factor \( m_{\Delta}(\zeta(t)) \) (or \( m_{\Delta}(\theta(t)) \)) given by (83) or (84), on the other hand, is independent of the frequency but it is monotonic decreasing/increasing with elevation, zenith \( \theta(t), \zeta(t) \in [0^\circ, 90^\circ] \) (or \( \zeta(t) = 90^\circ - \theta(t) \)) as shown in Fig. 16.

The correction term follows from (63) as

\[
D_{\Delta}'(\zeta) \approx \frac{m_{\Delta}^2(\zeta)}{m_{\theta}(\theta)} \left[ \int_{h_0}^{h_{\infty}} \nabla n(h) \times (h) d h \right] \frac{10^{-6}}{l_k} = G_l \cdot e \equiv G_{l}(\phi)
\]

(67)

and \( D_{\Delta}'(\zeta) \) can be approximated as in the right-hand side of (62); \( e = [\cos \phi; \sin \phi] \) is a unit vector in the direction of \( x(h) \); hence, we obtain

\[
G_l = 10^{-6} \int_{h_0}^{h_{\infty}} \nabla n(h) h d h
\]

(68)

Fig. 17. Refractive index, \( n(f) \), of the entire RF L band 1-2 GHz.

Fig. 18. Refractivity, \( N(f) \), of the entire RF L band 1-2 GHz.

Fig. 19. Ionospheric delay, \( D_{\Delta}(f) \) (m), of the entire RF L band 1-2 GHz.

Fig. 20. Ionospheric critical delay, \( D_{\Delta}^{\text{crit}}(f) \) (m), of the entire RF L band 1-2 GHz.

Which is equivalent with

\[
G_l = -\frac{d\left[ \int_{h_0}^{h_{\infty}} \nabla n(h) h d h \cdot e \right]}{d e_l} \bigg|_{e_l=0}, \quad l = \{1 \equiv NS, 2 \equiv EW\}
\]

(69)

where the expression of \( v_l(h) \) can be obtained from

\[
v_l(h) = \frac{\partial \mathbf{n}_l(h)}{\partial e_l} \bigg|_{e_l=0}, \quad l = \{1 \equiv NS, 2 \equiv EW\},
\]

(70)

where \( \mathbf{v}(h) = \{v_l(h)\}, \quad l = \{1 \equiv NS, 2 \equiv EW\} \), is the velocity of the PED bubble with components in the North-South and East-West directions and \( \nabla \mathbf{n}_l(h) \) is the gradient vector of the \( n_l(h) \) in the opposite direction of the projected normal with components

\[
\nabla \mathbf{n}_l(h) \equiv \mathbf{v}(h)
\]

(71)

If we further define

\[
e_l = \frac{\partial \Theta_{EC}}{\partial e_l} \bigg|_{e_l=0}
\]
The first part of (67), azimuth normal with components is the gradient vector in the opposite direction of the projected azimuth zenith delay. Therefore, the azimuth delay can be computed from either the gradient or the critical azimuth zenith delay. Hence, the final formula of the gradient, \( G_i \), (or the critical azimuth zenith delay, \( D_i^{az} \)) can be computed from

\[
D_i^{az}(\zeta) = m_i^{az}(\zeta) G_i(\phi); \quad m_i^{az}(\zeta) = \frac{m_i(\zeta)}{m_i(\theta)}
\]

as shown in Figs. 9-12

In analogy to the symmetric case, the “asymmetric” delay in the slant direction is also a product of a MF and, in this case, the delay in the direction of the normal \( D^n \approx D^f \)

\[
D_i^{as}(\phi, \zeta) = m_i(\zeta - \epsilon \cos(\beta))D^n
\]

where \( \beta \) is the angle between the projected normal and the projected slant direction. The approximation is allowed because \( \epsilon \cos(\beta) \) is small. From (57) and (81) now follows

\[
D_i^{as}(\phi, \zeta) = -\left[ \frac{d m_i(\zeta)}{d \zeta} \approx \sec(\zeta)\tan(\zeta) \right] \epsilon \cos(\beta) D_i^f
\]

So, if the RF signal is received from the same direction as the normal, the total delay is the smallest, caused by a thinner atmosphere. Because the MF is approximately \( m_i(\zeta) \approx \sec(\zeta) \), and \( d m_i(\zeta)/d \zeta \approx \sec(\zeta)\tan(\zeta), \) from (67), (72), and (74) can be seen that \( G_i \cdot e = |G_i| \cos(\beta) \approx -D_i^f \epsilon \cos(\beta). \) Therefore, the length of the gradient vector, \(|G_i| = D_i^f |e|\) is about which is an upper bound of the gradient delay for a particular azimuth as shown in Figs. 9-12.

The azimuthal MF (AMF), \( m_i^{az}(\zeta) \) (or \( m_i^{az}(\theta) \)), is given as

\[
m_i^{az}(\zeta) = \frac{m_i^2(\zeta)}{m_i(\theta)} = \cos^2(\theta') = \frac{\sin^2(\zeta')}{\cos(\zeta')}
\]

\[
m_i^{az}(\theta) = \frac{m_i(\theta)}{m_i(\zeta)} = \frac{\sin^2(\zeta')}{\sin(\theta')} = \cos(\theta')
\]

Where

\[
m_i(\zeta) = \frac{1}{\sqrt{1 - q^2}}
\]

\[
= \frac{1}{\cos(\sin^{-1}q)}
\]

\[
= \frac{1}{\cos^{-1}(\zeta')}
\]

\[
= \frac{1}{\sec(\zeta')}
\]

\[
\approx \frac{1}{\sqrt{1 - 0.948\sin^2(\zeta)}}
\]

\[
m_i(\theta) = \frac{1}{\sin(\theta')}
\]

\[
= \csc(\theta')
\]
1 \approx 1 - 0.948 \cos^2(\theta) \quad (86)

\begin{align}
q &= r \frac{\cos(\theta)}{\sin(\zeta)}; \\
r &= \frac{R_e}{R_e + h_i}
\end{align} \quad (87)

\[
D_{i}^{az}(\zeta) = 10^{-6} \int_{h_i}^{\infty} \nabla N(h_0) \cdot d\mathbf{s}
\]

\[
\approx \frac{\sin^2(\zeta')}{\cos(\zeta')} \mathbf{G}_i \cdot \mathbf{e}
\]

\[
= m_i^{az}(\zeta) D_{i}^{az,xc}(t)
\]

As can be seen from Figs. 13-16, gradients become increasingly influential at lower elevation angles.

Another equivalent expression of the effects of azimuthally asymmetric, atmospheric delay is primarily caused by inhomogeneous PED distribution. \( D_{i}^{az}(t) \) can be expressed as

\[
D_{i}^{az}(t) = m_i^{az}(\theta) \left( G_{1,1} \cos \phi + G_{1,2} \sin \phi \right)
\]

\[
= m_i^{az}(\theta) D_{i}^{az,xc}(t)
\]

where \( D_{i}^{az,xc} \) is the critical azimuth zenith delay corresponding to \( \cos^2(\theta')/\sin(\theta') = 1 \) or \( \theta^c = 41.8227^\circ \); i.e., for \( 0^\circ \leq \theta \leq 41.8227^\circ \) \( D_{i}^{az}(\theta) \geq D_{i}^{az,xc} \) and for \( \theta > \theta^c = 41.8227^\circ \) we have that \( D_{i}^{az}(\theta) < D_{i}^{az,xc} \).

Up to this point the derivation of the total ionospheric delay included only the influence of the azimuth, elevation (or zenith) and height of the ionosphere. Since experimental data has shown that PED varies considerably as a function of GNSS receivers’ geodetic latitude, \( \theta_h \), longitude, \( \phi_h \), and time, \( t \), then the following expression of the zenith/slant ionospheric delays and the ionospheric gradients along with MFs can be used to estimate the total asymmetrical ionospheric delay as [11], [62]-[66].

\[
D_{i}^{az}(t) = D_{i}^{yv}(t) + D_{i}^{az}(t)
\]

\[
D_{i}^{az}(t) = D_{i}^{yv}(t) + D_{i}^{az}(t)
\]
Where
\[ D_i(t) = m_i[\zeta(t)]D_i^z(t) \]  \hspace{1cm} (92)
\[ D_i^y(t) = m_i[\theta(t)]D_i^z(t) \]  \hspace{1cm} (93)
\[ D_i^{xy}(t) = m_i[\zeta(t)]D_i^{az,xc}(t) \]  \hspace{1cm} (94)
\[ D_i^{xz}(t) = m_i[\theta(t)]D_i^{az,xc}(t) \]  \hspace{1cm} (95)

where \( D_i^{az} / D_i^{az} \) is the azimuth delay as a function of \( \zeta(t) / \theta(t) \), \( \varphi_h, \varphi_b \) and \( t \) and \( D_i^{az,xc} \) is the critical zenith ionospheric horizontal gradient and \( \zeta(t) / \theta(t) \) are the zenith and elevation angles of the IP point (see Figs. 9-12) [64], [67], [68].

The differential ranging error due to ionospheric spatial gradient was also modeled using the associated LAAS model [63], [65], [69],[75] or L2C [76]
\[
D_i^{az}(\zeta) = \frac{\alpha \frac{d\zeta}{dx}}{\sqrt{1 - \zeta^2}} \equiv \left( \frac{\alpha d\zeta}{dx} + 2r v_{air} \right) m_i(\zeta)^{1/2} \]  \hspace{1cm} (96)
\[
D_i^{az}(\theta) \equiv \left( \frac{\alpha d\zeta}{dx} + 2r v_{air} \right) m_i(\theta) \]  \hspace{1cm} (97)

where \( D_i^{az}(\theta), D_i^{az}(\theta) \) is the [\( L_1 \) for example] ionospheric error (or azimuth delay, negative for carrier), \( dD_i^z/dx \) is the vertical ionospheric gradient; and \( x_r \) (or \( x_{air} \)) is the distance (slant range) of the airplane from the ship (or the GBAS reference point [m]); \( r \) is the service type dependent time constant [8] of the smoothing filter and \( v_{air} \) is the horizontal speed of the aircraft [m/s].

Equations (96) or (97) provides the formula that is able to reconcile the expression not found anywhere else
\[
\frac{d\zeta}{dx} x = \cot(\theta') \left[g_{ix} \cos(\phi) + g_{iz} \sin(\phi)\right] = \Delta_i^{az}(\theta, \phi, t) \]  \hspace{1cm} (98)

And
\[
\frac{d\zeta}{dx} x m_i(\zeta) = m_i^{az}(\zeta)D_i^{az,xc} \rightarrow \]  \hspace{1cm} (99)

\[
\frac{d\zeta}{dx} x = \frac{m_i^{az}(\zeta)D_i^{az,xc}(t)}{x_{air} m_i(\zeta)} \]  \hspace{1cm} (100)

Substituting (80) into (100) the final expression for \( D_i^{az,xc} \) is obtained
\[
D_i^{az,xc}(t) = \frac{m_i(\theta)}{m_i(\zeta)} \frac{d\zeta}{dx} x \]  \hspace{1cm} (101)

This concludes the description of the ionospheric delay.

### 3.3 Description of the Ionospheric Doppler Frequency

There appears to be at least three good references: Nickisch et al. in 2006, [77] in “Range rate – Doppler correlation for HF propagation in traveling ionospheric disturbance environments,” and Zhou et al., 2012 [78] “Evidence of low-latitude daytime large-scale traveling ionospheric disturbances observed by high-frequency multistatic backscatter sounding system during a geomagnetically quiet period,” who are directly linked with the ionospheric Doppler frequency estimation and Cham et al. 2012 [79] “Statistical investigation of horizontal propagation of gravity waves in the ionosphere over Europe and South Africa” which is indirectly liked with the ionospheric Doppler frequency estimation. Nickisch et al. in 2006, [77] claims: “Using ionospheric sounding together with fast computational inverse processing, it is now possible to obtain good real-time ionospheric models for use in geolocation for over-the-horizon (OTH) radar.” Zhou et al. 2012 [78] claims: “Vertical phase velocity was also evaluated to be < 42 m/s through the Doppler measurements.” Cham et al. 2012 [79] claims: “Simultaneous measurements from nearby ionosonde made it possible to estimate the height of the ground waves (GW) observations and show that the analyzed GWs propagated at altitudes from ~150 to ~250 km. The analyzed waves were mainly observed after sunrise and around sunset. Our statistical study shows that the analyzed GWs propagated with horizontal velocities from ~70 to 250 m/s. The average observed horizontal velocities were ~100 m/s in the local summer and 125-250 m/s in the local winter.” According to Jandieri et al. 2009 [80], “Experimental investigations of Doppler shifts of ionospheric signal (basic frequency was ~ 17 MHz) reflected from the ionosphere were carried out in Moscow State University, Moscow (Lat. 0 N, 0 55.5, Long. 0 E, 0 37.3) USSR in 1979. It was established that electron density fluctuations have a power-law spectrum with spectral index 3.8 ≤ p ≤ 4.6, (p) = 4.2 . Drift velocity of large-scale ionospheric irregularities is about ~170 m/s and corresponds to linear scales from several hundred to several tens of kilometers.” Furthermore, horizontal mean drift velocity is 257 m/s (day) and 185 m/s (night); more frequently 136 ~ 170 m/s translate to ionospheric Doppler frequencies in the range of 40-
However, the derivations in this section are original and unique neither based on any analytical treatise of Nickisch et al. in 2006, [77], nor Zhou et al. 2012 [78], and nor Chum et al. 2012 [79] which are specifically used for radar signal processing. Moreover, it is important to know that depending on the solar activity the Doppler frequency estimated from the theory presented in this section is in total agreement with 4-parameter model theory presented in this paper.

The ionospheric Doppler frequency, $\pi D_D$, is a subcomponent of the $\pi_s(t)$ or assuming that all other channel effects are ignored then

$$f_D(t) = f_i^{D,as}(t) + f_i^{D,as}(t): t \in [0,T]$$

(102)

In this paper we are going to ignore the other channel components such as the tropospheric and multipath effects.

Nevertheless, before we can discuss the ionospheric Doppler frequency, we need to do some preliminary calculations.

Typically, a GPS satellite orbital period is one-half a sidereal day, i.e., 11 hours and 58 minutes or $T = 718$ min or $T = 718 \times 60$ s. Hence, the angle $\alpha'(t)$ as a function of time is as follows:

$$\alpha'(t) = \frac{2\pi}{2\pi} \equiv k_{\alpha'} t: t \in [0,T)$$

(103)

Next, we compute the angle $\beta(t)$ as follows:

$$\beta(t) = \tan^{-1} \left[ \frac{R_s \sin \alpha'(t)}{R_s - R_e \cos \alpha'(t)} \right]: t \in [0,T)$$

(104)

where $R_e = 26,600$ km is the GPS satellite orbital radius.

The zenith angle, $\zeta(t)$, is the sum of two angles $\alpha'(t)$ with $\beta(t)$ (see Figs. 9-12) as follows:

$$\zeta(t) = \alpha'(t) + \beta(t) = k_{\alpha'} t + \beta(t)$$

(105)

The angle, $\zeta(t)$, can be computed from

$$\zeta'(t) = \sin^{-1} [\sin \zeta(t)]: t \in [0,T)$$

(106)

Next, we compute the derivative of the zenith angle, $\frac{d\zeta(t)}{dt}$, as follows:

$$\frac{d\zeta(t)}{dt} = k_{\alpha'} + \left( \frac{R_e \sin (k_{\alpha'} t)}{R_e - R_e \cos (k_{\alpha'} t)} \right) \frac{d\alpha'(t)}{dt}$$

(107)
\[ \frac{df(t)}{dt} = R_e k' \cos(k' \cdot t) \]

Substituting (109) and (110) into (108) yields

\[ \frac{d\delta(t)}{dt} = k' + \frac{R_e k' \cos(k' \cdot t) - R_e \cdot k'}{R_e^2 - 2 R_e R_a \cos(k' \cdot t) + R_a^2} \]

Next, let us find the derivative of \( \frac{dm_i(t)}{dt} \) as follows

\[ \frac{dm_i(t)}{dt} = \sec[\delta(t)] \tan[\delta(t)] \frac{du(t)}{dt} - \frac{\sec[u(t)] \tan[u(t)] \frac{du(t)}{dt}}{\sqrt{1 - u^2(t)}} \]

\[ \frac{dm_i(t)}{dt} = \sec[\delta(t)] \frac{\cos(\delta(t)) - \sin(\delta(t))}{\cos(\delta(t))} \frac{du(t)}{dt} - \frac{\sec[u(t)] \tan[u(t)] \frac{du(t)}{dt}}{\sqrt{1 - u^2(t)}} \]

\[ \frac{dm_i(t)}{dt} = m_i(t) \frac{d\delta(t)}{dt} \]

Where \( \frac{d\delta(t)}{dt} \) is already given previously by (110).

If we wanted to include the gradient component in the SID, \( D^{\delta \gamma}_1(t) \) (or \( D^{\delta \alpha}_1(t) \)) then we get

\[ D^{\delta \gamma}_1(t) = D^{\delta \gamma}_1(t) + D^{\delta \alpha}_1(t) \]

\[ = m_i(t) D^{\delta \gamma}_1(t) + m_i^{\delta \gamma}(\theta) D^{\delta \alpha \gamma \kappa \chi}_1(t) \]

where

\[ D^{\delta \gamma}_1(t) = D^{\delta \gamma}_1(t) + D^{\delta \alpha}_1(t) \]

\[ = m_i(t) D^{\delta \gamma}_1(t) + m_i^{\delta \gamma}(\theta) D^{\delta \alpha \gamma \kappa \chi}_1(t) \]
Finally substituting (118) and (119) or (120) into (117) we obtain Doppler frequency, \( f_i^D \) (or \( f_i^D \)): (1), as a result of the change in ionospheric zenith delay, \( D_i^z(t) \) (see (65)), and (2) the other as a result of the change in the obliquity factor \( m_i(\xi^k) \) (or \( m_i(\theta^b) \)) (see (85) or (86)); i.e.,

\[
f_i^{D,xy}(t) = - \frac{d TEC(x,t)}{c f} - \frac{d TEC(y,t)}{c f} \frac{d \theta(t)}{dt}
\]

(122)

\[
f_i^{D,zy}(t) = - \frac{a TEC(x,t)}{c f} - \frac{a TEC(y,t)}{c f} \frac{d \theta(t)}{dt}
\]

(123)

Since, the (or symmetric) Doppler component \( f_i^{D,xy} \) (or \( f_i^{D,zy} \)) is already computed in (122) or (123); the azimuth

\[
f_i^{D,az}(t) = \frac{d \theta(t)}{dt}
\]

(117)

Doppler component from \( f_i^{D,az} \) (or \( f_i^{D,az} \)) can be computed from

\[
f_i^{D,az}(t) = \frac{d \theta(t)}{dt}
\]

(124)

where

\[
m_i^z(t) = \frac{-r^2 \cos[\theta(t)] \sin[\theta(t)]}{\sin^3[\theta(t)]}
\]

(121)
Since, $\zeta'(t)$ is given by (106); hence,

$$\frac{d\zeta'(t)}{dt} = \frac{\cos\theta(t)}{\cos^{2}[\zeta(t)]} \frac{d\zeta(t)}{dt}$$

(126)

Similarly,

$$\theta'(t) = \cos^{-1}[\cos\theta(t)]$$

(127)

$$\frac{d\theta'(t)}{dt} = \frac{\sin\theta(t) d\theta(t)}{\sin\theta'(t)} \frac{d\theta(t)}{dt}$$

(128)

However, since $\theta(t) = 90^\circ - \zeta(t)$ then, hence,

$$\frac{d\theta(t)}{dt} = -\frac{d\zeta(t)}{dt}$$

(129)

Finally, substituting (126), (128)/(129), into (124)/(125) we obtain

$$f_i^{D,ax}(t) = -\frac{m_i^{ax}(t) D_i^{ax,xc} d\theta(t)}{\lambda} m_i^{ax}(t) \frac{d\theta^{ax,xc}}{dt}$$

(130)

$$f_i^{D,ax}(t) = -\frac{m_i^{ax}(t) D_i^{ax,xc} d\theta(t)}{\lambda} m_i^{ax}(t) \frac{d\theta^{ax,xc}}{dt}$$

(131)

Where

$$m_i^{ax}(t) = \frac{\sin[\theta'(t)] [1 + \sec^2[\theta'(t)]]}{\cos[\theta'(t)]} \frac{d\sin\theta(t)}{dt}$$

(132)

$$m_i^{ax}(t) = \frac{\cos[\theta'(t)] [1 + \sec^2[\theta'(t)]]}{\sin[\theta'(t)]} \frac{d\sin\theta(t)}{dt}$$

(133)

Next, let us compute $\frac{dB_i^{ax,xc}(t)}{dt}$ by differentiating (78); therefore, first we compute

$$\frac{d\zeta}{\sqrt{1 + c^2}} = \frac{d\zeta}{\sqrt{1 + c^2}}$$

(134)

$$\frac{d\zeta}{\sqrt{1 + c^2}} = \frac{d\zeta}{\sqrt{1 + c^2}}$$

$$\frac{d\zeta}{\sqrt{1 + c^2}} = \frac{d\zeta}{\sqrt{1 + c^2}}$$

$$\frac{d\zeta}{\sqrt{1 + c^2}} = \frac{d\zeta}{\sqrt{1 + c^2}}$$

Finally, substituting (122)/(123), (120)/(131) into (114)/(115) we obtain the final Doppler expression due to total ionospheric delay [11]

$$f_i^{D,as}(t) = -\frac{a m_i(t) (D_i^{as,xc}(t))}{\lambda}$$

(138)

$$f_i^{D,as}(t) = -\frac{a m_i(t) (D_i^{as,xc}(t))}{\lambda}$$

(139)

Where

$$D_i^{D,ax}(t) = m_i^{ax}(t) D_i^{ax,xc}(t)$$

(140)

$$D_i^{D,ax}(t) = m_i^{ax}(t) D_i^{ax,xc}(t)$$

(141)

Special cases: There are four special cases that we should take into consideration.

(1) Static component. If we assume that $\frac{df(t)}{dt} = \frac{d\theta(t)}{dt} = 0$

then $A(t) = B(t) = 0$.

$$f_i^{D,as}(t) = -\frac{a m_i(t) (D_i^{as,xc}(t))}{\lambda}$$

(142)

$$f_i^{D,as}(t) = -\frac{a m_i(t) (D_i^{as,xc}(t))}{\lambda}$$

(143)

If $m_i(t) = m_i(t) = 1$ then the static component becomes the same as the zenith component; i.e., zenith component is a subcomponent of the static component.

(a) Zenith component with no gradient component. On the other hand, another effect known as the phase advance $\Delta \phi_k$ affects the carrier frequency of the transmitted signal, $f_k$. In this case the signal is affected by the following factor in cycles (or in radians) [81]

$$\Delta \phi_k = \frac{a f_k}{c f_k} = f_k \frac{d\phi}{c}$$

(144)

$$\Delta \phi_k = 2\pi \Delta \phi_k = \omega_k \frac{d\phi}{c}$$

(145)

where (1) $D_i^2$ is the ionospheric zenith delay in [m]; (2) $c$ is the SOL in vacuum [2] [m/s].

Therefore, if one was to assume the ionosphere [2] to be a
layer between the satellite and the user, the TEC encountered by the signal directly depends on the satellite elevation and change as the satellite is moving. This change induces a change in the phase advance observed and as such a frequency shift known as the Doppler frequency induced by the ionosphere [2] given by [81], [82]

\[ \Delta \phi_k = \frac{d \Delta \phi_k}{dt} \]

\[ = f_D(t) \]

\[ = - \frac{a \cdot dT\text{EC}(t)}{c \tau_k} \]

\[ = \left[ \frac{1.344536593729647 \times 10^{-7}}{\pi \lambda_k} , \text{cycles} \right] \frac{dT\text{EC}}{dt} \]

\[ = - f_k \frac{d \psi(t)}{c \cdot dt} \]

(146)

Where

\[ a = 1.344536593729647 \times 10^{-7} \text{els}^{-1} \cdot \text{m}^2 \cdot \text{s}^{-1} \cdot \text{c} \]

\[ = \frac{1.344536593729647}{10^7} \times \frac{m^2}{\text{els} \cdot \text{s} \cdot \text{c}} \times 299,729,458 \frac{m}{s} \]

\[ = 40.308193030515824 \text{ el} \cdot \text{s}^{-1} \cdot \text{m}^2 \cdot \text{s}^{-2} \]

(147)

where TEC is the integral of electron density along the path of the radio signal from transmitter to receiver, where \( f_{1,2} \) is evaluated in units [83] Hz, \( d\text{TEC}/dt \) is evaluated in units TEC/s, and where TECN = 1 el \cdot m^{-2} [81].

(2) **Slant (symmetric) component with and no change in TEC and no gradient component.** With dual frequency measurements, we can compute the ionospheric delay [11] and numerically differentiate to produce an additional correction term to the predicted model Doppler. However, without this information, we can assume that the path delay is a function of the zenith delay and the obliquity factor (itself a function of the satellite elevation). Differentiating the path slant delay, \( D_{i}^{\text{SY}}(\theta) \), with respect to time, we find that

\[ \frac{dD_{i}^{\text{SY}}(\theta,t)}{dt} = -D_{i}^{0}(t) \frac{m_{i}^{\text{z}2}(t)}{\lambda} \frac{d \theta(t)}{dt} \]  

\[ = f_{i}^{D}(t) \]

\[ = - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{SY}}(\theta)}{dt} = \frac{d \psi(t)}{\lambda_{i}} m_{i}^{\text{z}2}(t) \frac{d \theta(t)}{dt} \] (see (121))  

(149)

where \( \theta \) is the elevation angle of the lowest layer, \( D_{i}^{0} \) is the zenith delay in meters. The Doppler shift from the ionosphere [2] is the negative of the time rate of change of the path delay over the wavelength [84].

The negative sign is due to the phase advance being opposite in sign from the path delay. The new predicted frequency is thus

\[ f_{i}^{D} = f_{\text{model}}^{0} + f_{i}^{D} \]  

(150)

Leveraging dual-frequency code-based pseudoranges from the SATRACK data, we calculated an approximate zenith delay for each space vehicle (SV), which varied between 1.5 m and 3 m. The total effect of \( f_{i}^{D} \) is small (thousandths of Hz) but can correct some of the trends observed in the residual phase measurements [84].

The same result in (149) can be obtained from (139) if

\[ \frac{dT\text{EC}(t)}{dt} = 0 \text{ and } D_{i}^{\text{az},\text{zc}}(t) = 0 \text{ then } \frac{dD_{i}^{\text{az},\text{zc}}(t)}{dt} = 0. \]

(3) **Slant (symmetric) component with no gradient component.**

If we assume \( D_{i}^{\text{az},\text{zc}} = 0 \) then \( \frac{dD_{i}^{\text{az},\text{zc}}}{dt} = 0 \) then from (137)/(138) the following is obtained

\[ f_{i}^{0,\text{sy}}(t) = - \frac{a m_{i}(t) D_{i}^{\text{az},\text{zc}}(t)}{c f} \frac{D_{i}^{0}(t)m_{i}^{\text{z}2}(t) d \xi(t)}{\lambda} \]

\[ = - \frac{a m_{i}(t) D_{i}^{\text{az},\text{zc}}(t)}{c f} \frac{D_{i}^{0}(t)m_{i}^{\text{z}2}(t) d \theta(t)}{\lambda} \]

(151)

(152)

This is the slant (symmetric) Doppler component with no gradient component same as (122)/(123).

(4) **Slant azimuth Doppler component.** If we assume

\[ \frac{dT\text{EC}(t)}{dt} = 0 \text{ then from (138)/(139) the following is obtained } \]

\[ f_{i}^{\text{D,as}} = - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} m_{i}^{\text{z}2}(t) \frac{d \theta(t)}{dt} \]

\[ = - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} m_{i}^{\text{z}2}(t) \frac{d \theta(t)}{dt} \]

(153)

(154)

where \( D_{i}^{\text{m}}(t) \) and \( D_{i}^{\text{m},\text{zc}}(t) \) are given by (140) and (141).

This is the slant azimuth Doppler component; i.e., gradient and derivative of the gradient only component with no change in TEC, neither zenit nor elevation component. Furthermore, slant azimuth Doppler component can be decomposed into two subcomponents:

(a) **gradient only with no change in TEC** if \( \frac{dD_{i}^{\text{az},\text{zc}}(t)}{dt} = 0 \) or \( D_{i}^{\text{az},\text{zc}} = \text{const.} \)

\[ f_{i}^{\text{D,as}}(t) = - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} \frac{d \xi(t)}{\lambda} - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} \frac{d \theta(t)}{\lambda} \]

\[ = - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} \frac{d \xi(t)}{\lambda} - \frac{1}{\lambda_{i}} \frac{dD_{i}^{\text{m}}(t)}{dt} \frac{d \theta(t)}{\lambda} \]

(155)

(156)

and (b) **derivative of the gradient only with no change in TEC**
if \( \frac{dt}{dt} = \frac{d\theta(t)}{dt} = 0 \) then \( m^{az}(t) = m^{az}(t) = 0 \) hence

\[
\begin{align*}
 f_i^{D,az}(t) &= -\frac{m^{az}(t) d\theta^{az}(t)}{\lambda} \\
 f_i^{D,az}(t) &= -\frac{m^{az}(t) d\theta^{az}(t)}{\lambda}
\end{align*}
\]

This concludes the description of the ionospheric Doppler frequency calculations.

### 3.4 Description of the amplitude and phase scintillation

Amplitude scintillation is a measure of the signal amplitude degradation due to scintillation and is measured by an index known as \( S_4 \).

The \( S_4 \) amplitude scintillation index [85] is computed as the standard deviation of normalized signal intensity [86] (or irradiance, \( I \)) (expressed in linear units [83]) or power \( P \).

\[
S^2_4 = \frac{(\langle P^2 \rangle - \langle P \rangle^2)}{\langle P \rangle^2} = \frac{(\langle P_{\text{ave}} \rangle)^2}{\langle P \rangle^2}
\]

If the signal is measured in power units (or Watts [22]) then \( Y \) : Power \( E[Y] = \langle P \rangle ; \sigma_{|Y|} = S_4(P) \) else if the signal is measured in amplitude units then \( X \) : Amplitude \( E[X] = \langle A \rangle = \sqrt{2\langle P \rangle} ; \sigma_{|X|} = \sqrt{2S_4(P)} = \sqrt{2S_4(A)} = \sqrt{2\sigma_A} \).

Here the angle brackets \( \langle \alpha \rangle \equiv E[\alpha] \) formally denote an ensemble average of \( \alpha \) or \( E[\alpha] \) but in practice it indicates spatial or temporal averages.

One school of thought believes that the \( S_4 \) index used as a sole measure of scintillation does neither parameterize phase fluctuation index nor provides detailed characterization of fade depth or duration. According to this school of thought, since \( S_4 \) reaches the neighborhood of unity (focusing on saturation regime), scintillation characteristics may vary even as \( S_4 \) remains relatively constant. Nevertheless, \( S_4 \) remains a worthwhile means of characterizing scintillation due to several factors: (1) large, long term activities of \( S_4 \) measurements exist and morphological studies use these archives the natural scintillation environment (Aarons, 1993 [87]); (2) a good number of scintillation detecting systems are largely based on \( S_4 \) measurements (Groves et al, 1997, [88]; Thomas et al., 2001, [89]); (3) additionally scintillation at L-band (1-2 GHz) or higher frequencies often does not approach saturation except in limited circumstances (Fremouw, et al, 1980 [90]; Yeh and Liu, 1982 [91]); (4) finally, empirical relationships between \( S_4 \) and properties like signal fade statistics allow historical or climatological \( S_4 \) values to be used in system design [85]. As we are going to see from the derivations and simulation analyses, the conclusions provided by this school of thought are erroneous.

Another school of thought, as we shall see in the further development of this section, believes that there exists a well-defined relationship between the phase fluctuations and the amplitude scintillation index, \( S_4 \). Therefore, we are first going to produce an understanding about the phase fluctuations or phase scintillation index or phase variance due to scintillation and then derive the amplitude scintillation index, \( S_4 \) [80], [92]-[94].

In Jandieri et al. 2009 [80] phase fluctuations are suggested as being responsible for fluctuations of the angle-of-arrival (AOA). AOA measurements have been performed by interferometers; whereas a part of radar propagation effects program at the Millstone Hill radar facility [1], AOAs has been measured with a single mono-pulse tracking system. According to Jandieri et al. 2009 [80] the study of AOA could provide useful information about the structure of the ionospheric irregularities.

Let as consider a simplified model of the magnetized plasma in order to visualize the influence of medium absorption and anisotropy on the AOA of scattered waves [80].

Let us assume a 3-layer approximation between a Satellite Transmitter and a Receiver with respective thickness, \( h_i, i = \{1,2,3\} \) as shown in Fig. 12; where the middle, scattered layer contains random magnetized plasma and irradiated high-frequency electromagnetic waves.

The LOS distance between the Satellite Transmitter and a Receiver is given by

\[
H_0 = \sum_{i=1}^{2} h_i
\]

Suppose a magnetic field with constant flux density vector, \( \textbf{B}_0, \) and magnitude of \( B_0 \) and angle of \( \zeta_0 \) from the z axis is given by

\[
\textbf{B}_0 = B_0 e^{j\zeta_0}
\]

Suppose that the wavenumber vector, \( \mathbf{k} \), of the signal in vacuum is given by

\[
\mathbf{k} = k_0 e^{j(\gamma+\beta)}
\]

where the wavenumber magnitude of the signal in vacuum, whose radiant frequency is \( \omega \), is given by

\[
k_0 = \omega/c
\]
Suppose that the plasma density in the layer, \( N(r) \), is the sum of two terms: the constant term, \( N_0 \), and the random function of the spatial coordinates describing density fluctuations, \( N_1(r) \), then
\[
N(r) = N_0 + N_1(r)
\]  
(164)

Assuming a Gaussian distribution of the random function of the spatial coordinates plasma density fluctuations, \( N_1(r) \), its variance is given by \( \langle N_1^2 \rangle \); hence, the variance of the plasma density in the layer, \( N(r) \), is denoted by \( \sigma_N^2 \) and it can be computed from
\[
\sigma_N^2 = \langle N_1^2 \rangle / N_0^2
\]  
(165)

Let us assume that \( v_{\text{eff}} \) is the effective collision frequency of electrons with other plasma particles; the plasma frequency can be computed from [80]
\[
\omega_p^2 = \frac{e^2 n_e}{\varepsilon_0 m_e} = 4\pi^2 f_p^2
\]  
(166)

The angular gyro-frequency for the magnetic field, \( \omega_B \), is defined as [80]
\[
\omega_B = \frac{e B_0}{m_c}
\]  
(167)

Suppose that the following conditions are fulfilled [80]
\[
\omega_B \gg \omega; \ \omega_B \gg \omega_p; \ \omega_B \gg v_{\text{eff}}
\]  
(168)

We define the following non-dimensional plasma parameters \( \nu \) and \( s \) as follows
\[
\nu = \omega_p^2 / \omega^2
\]  
(169)
\[
s = v_{\text{eff}} / \omega
\]  
(170)

Anisotropic Gaussian correlation function of electron density fluctuations in F-region of ionospheric plasma, when prolate (or elongate along the polar diameter) irregularities produce an angle, \( \gamma \), with respect to the external magnetic field.

Assuming that the irregularities are elongated along a direction of the geomagnetic field
\[
\alpha = k_0 \left[ \alpha_0 \equiv \frac{1}{2N_0} \left( \sin^2 \theta_0 + (3\cos^2 \theta_0 - 5/2) \right) \right]^2
\]  
(171)

Then it can be easily shown that \( \alpha > 1 \) [80].

---

1 The plasma frequency, \( f_p \), was introduced earlier in (40).

2 Please do not confuse \( \alpha \) with \( \alpha' \).
Let $h_\parallel$ and $h_\perp$ indicate the relative characteristic spatial scale sizes in the field aligned, $\parallel$, and field-perpendicular, $\perp$, directions; then $\chi$ is the anisotropic parameter (see Fig. 12) given by

$$\chi = h_\parallel/h_\perp$$  \hspace{1cm} (172)

$$\bar{h}^2 = \frac{h_\parallel^2}{\sin^2\gamma + \chi^2 \cos^2\gamma}$$  \hspace{1cm} (173)

Next $a$, $b$, and $d$ are defined as non-dimensional spatial parameters normalizing on thickness of a plasma layer

$$a = h_\parallel/h_2$$  \hspace{1cm} (174)

$$b = h_1/h_2$$  \hspace{1cm} (175)

$$d = h_3/h_2$$  \hspace{1cm} (176)

Next, we define the following non dimensional plasma parameters

$$A = v \frac{b+d}{1+b+d} \sin \zeta_0 \cdot \cos \zeta_0 = \nu \gamma$$  \hspace{1cm} (177)

$$J_0 = v \frac{b}{1+b+d} \sin \zeta_0 \cdot \cos \zeta_0 = \nu \delta$$  \hspace{1cm} (178)

Next, we define the following remaining parameters as follows:

$$a_1 = \frac{(1-\chi^2)^2 \sin^2 \alpha \cos^2 \alpha}{\chi^2}$$  \hspace{1cm} (179)

$$a_2 = \frac{\sin^2 \alpha + \chi^2 \cos^2 \alpha}{\chi^2}$$  \hspace{1cm} (180)

$$a_3 = \frac{(1-\chi^2) \sin \alpha \cos \alpha}{\chi^2}$$  \hspace{1cm} (181)

$$G = a_1 + a_2 A^2 - 2a_3 A = a_1 + a_2 \nu^2 \gamma^2 - 2a_3 \nu \gamma$$  \hspace{1cm} (182)

$$J = 2s J_0$$  \hspace{1cm} (183)

Assume that distance between two receivers normal to the direction of wave propagation in which phase correlation is measured is $x$ then the AOA and frequency fluctuations intensity for anisotropic Gaussian correlation function at $G_0 \equiv G(\chi = 0)$ has the following form [80]

$$\sigma_0^2 = \frac{\sqrt{\pi}}{4} \frac{s_A^2}{a_A} \frac{k_0^2 h_\parallel^2 \sin \zeta_0}{a_X} \left[ \frac{\nu^2}{(1-\sin^2 \zeta_0)} \right] \equiv \beta e^{ \left( \frac{J}{\omega_{p}^2} \right)^2}$$  \hspace{1cm} (184)

This expression is valid for arbitrary location of a point source and the receiver with respect to scattered layer [80].

Because we are looking forward to obtaining the phase variance as a function of the frequency we need to make the following simplification of $\alpha$ by substituting (163) and (169) into (171)

$$\alpha = k_0^2 \left[ \alpha_0 \equiv \frac{1}{2N_0} \frac{\sin^2 \zeta_0}{(1-\nu \cos \zeta_0)^2} \right] \equiv \frac{1}{2N_0} \frac{\omega_p^2 \sin \zeta_0}{c^2}$$  \hspace{1cm} (185)

Similarly, we simplify $\beta$ and $J/G_0$ by substituting (169) into (182) and then (169) and (182) into (184) and assuming that $\omega \gg \omega_p$ yields

$$\beta = \frac{\omega_p^3/\omega^3}{\left(1-\frac{2\omega_p^2 \cos \theta_0}{\omega^2}\right) \left(a_1 + a_2 \frac{\omega_p^4}{\omega^6} \gamma^2 - 2a_3 \frac{\omega_p^2}{\omega^4} \gamma \right)} \cong \frac{1}{\omega^4} \left( \frac{\omega_p^2}{\alpha_1} \right)$$  \hspace{1cm} (186)

$$\frac{J}{G_0} = \frac{2s J_0}{a_1 + a_2 A^2 - 2a_3 A} = \frac{2\nu \omega_p^2 \beta}{a_1 + a_2 \frac{\omega_p^4}{\omega^6} \gamma^2 - 2a_3 \frac{\omega_p^2}{\omega^4} \gamma} \cong \frac{2\nu \omega_p^2 \beta}{\alpha_1}$$  \hspace{1cm} (187)

Hence, substituting (186) and (187) into (184) yields an approximation equal to

$$\sigma_0^2 \approx \frac{\nu \omega_p^2}{\alpha_1} e^{\frac{x^2}{\omega_x^2}}$$  \hspace{1cm} (188)

For the isotropic case $\chi = 1$ and ignoring the influence of the geometric parameters we obtain the main component of the phase variance as a function of the signal frequency

$$\sigma_0^2 \approx \frac{\nu \omega_p^2}{\alpha_1} e^{\frac{x^2}{\omega_x^2}}$$  \hspace{1cm} (189)

If the Satellite Transmitter and the Receiver are located on the upper and lower boundaries of plane plasma layer then we have $b = d = 0$, in isotropic case, $\chi = 1$, at $\zeta_0 = 90^\circ$, then (184) becomes

$$\sigma_0^2 = \sqrt{\pi} \sigma_0^2 h_\parallel h_2 \left( \frac{k_0^2 \nu^2}{4} \cong \frac{\omega_p^2}{\omega_x^2} \right) \cong \left( \frac{\omega_p^2}{\omega_x^2} \right)^2$$  \hspace{1cm} (190)

This is an upper bound of the phase variance as it is invariant of the frequency.

In Jandieri et al, 2007 [92] and 2008, [93] the phase variance $\langle \phi_z^2 \rangle$ and intensity scintillation index $S_4$ are related as follows

$$\langle \phi_z^2 \rangle = \pi \sigma_0^2 (D^2 + E^2) \frac{k_0^2 h_\parallel}{\omega_x^2}$$  \hspace{1cm} (191)

$$S_4^2 = 2 \langle \phi_z^2 \rangle G_{\text{ovl}}$$  \hspace{1cm} (192)

$$G = 1 - \frac{\cos \left[ \tan^{-1} \left( \frac{2k_0^2}{\omega_x^2} \right) + \tan^{-1} \left( \frac{\omega_p^2}{\omega_x^2} \left( q \chi x^2 k_0^2 \right) \right) \right]}{\left( 1 + \chi x^2 k_0^2 \right)^{1/2} \left( 1 + \chi x^2 k_0^2 \right)^{1/2}}$$  \hspace{1cm} (193)

$$Q^2(\alpha, \chi) = 1 + \frac{\chi^2 w^2}{x^2} + b^2 \frac{x^2}{\chi^2} - 2b w^2$$  \hspace{1cm} (194)

$$D = \frac{\alpha_0^2}{\alpha_1}$$
\[ D = D_1 + D_2 + D_3 \]
\[ D_1 = \varepsilon'(n'-n')[(e'-n_2^2 \sin^2 \zeta_1)] \]
\[ D_2 = \frac{n_2^2 (n'-n') n_z (e'-\eta \sin \zeta_1)}{2N_z [2n_e (n'-n') + n_z (e'-\eta \sin \zeta_1)]} \]
\[ D_3 = \frac{[n_2^2 (n'-n') n_z (e'-\eta \sin \zeta_1)] (e'-n_2^2 \sin^2 \zeta_1)}{2N_z [2n_e (n'-n') + n_z (e'-\eta \sin \zeta_1)]} \]
\[ b = \frac{\alpha''}{\alpha_x} \]
\[ w^2 = \frac{x^2-1}{x^2} \sin^2 \alpha \cos \alpha \]
\[ \eta' = 1 - \frac{u(1+s^2-u)}{1-s^2+4u^2} \]
\[ \eta'' = \frac{u(1+s^2-u)}{1-s^2+4u^2} \]
\[ \varepsilon' = 1 - \frac{u}{1-s^2} \]
\[ \varepsilon'' = \frac{su}{1+s^2} \]
\[ \mu' = \frac{2su}{(1+s^2-u)^2+4s^2} \]
\[ \mu'' = \frac{su}{(1+s^2-u)^2+4s^2} \]
\[ N_2 = \frac{n_2^2 + n_2^2 - n_2^2}{2} \]
\[ n_2^2 = 1 - 2v \frac{(1-v)[2(1-u-x^2)-\sin^2 \zeta_1 \pm \overline{\rho}]{x}[2(2u-v^2) \pm 2]}{[2(1-u-x^2)-\sin^2 \zeta_1 \pm \overline{\rho}]{2}[2(2u-v^2) \pm 2]} \]
\[ n_2^2 = -2v \frac{(1-v)[2(1-u-x^2)-\sin^2 \zeta_1 \pm \overline{\rho}]{x}[2(2u-v^2) \pm 2]}{[2(1-u-x^2)-\sin^2 \zeta_1 \pm \overline{\rho}]{2}[2(2u-v^2) \pm 2]} \]
\[ P = \sqrt{0.5(\sqrt{p^2 + q^2} + p)} \]
where a sum of two integrals:

\[ \int v_1 dh = \int v_2 dh \]

Example 1: In the first example we assume that \( v_i(h) \) is given by

\[ v_i(h) = \begin{cases} \frac{v_i}{2}(1 - e^{-\alpha_i(h-h_0)^2}), & h_0 \leq h < h_i \\ \frac{v_i}{2}e^{-\beta_i(h-h_i)^2} - 1, & h_i \leq h \leq h_i \end{cases} \]

Where \( i = 1 \equiv \text{NS} \) and \( i = 2 \equiv \text{EW} \).

First, let us compute the integral \( I \)

\[ I = \int_{h_0}^{\infty} v_i(h) dh \]

as a sum of two integrals: \( I_1 \) and \( I_2 \)

\[ I_1 = \int_{h_0}^{h_i} v_i(1 - e^{-\alpha_i(h-h_0)^2}) dh \]

\[ I_2 = \int_{h_0}^{h_i} v_i(\frac{h^2}{2} + e^{-\alpha_i(h-h_0)^2} - 1) dh \]

This concludes the description of the amplitude and phase scintillation indexes.

### 4 Description of a simplified ionospheric channel model

The exact computation of the ionospheric channel model based on the equations presented previously is a daunting task, because the results will depend on the profile of \( v_i(h) \) which is obtained from measurements; hence, we consider two simplified examples of \( v_i(h) \) profiles.

#### 4.1 Examples of description of simplified ionospheric delay

*Example 1:* In the first example we assume that \( v_i(h) \) is given by

\[ v_i(h) = \begin{cases} \frac{v_i}{2}(1 - e^{-\alpha_i(h-h_0)^2}), & h_0 \leq h < h_i \\ \frac{v_i}{2}e^{-\beta_i(h-h_i)^2} - 1, & h_i \leq h \leq h_i \end{cases} \]

Where \( i = 1 \equiv \text{NS} \) and \( i = 2 \equiv \text{EW} \).

First, let us compute the integral \( I \)

\[ I = \int_{h_0}^{\infty} v_i(h) dh \]

as a sum of two integrals: \( I_1 \) and \( I_2 \)

\[ I_1 = \int_{h_0}^{h_i} v_i(1 - e^{-\alpha_i(h-h_0)^2}) dh \]

\[ I_2 = \int_{h_0}^{h_i} v_i(\frac{h^2}{2} + e^{-\alpha_i(h-h_0)^2} - 1) dh \]

This concludes the derivations of Example 1.

*Example 2:* In the second example we assume that \( v_i(h) \) is given by

\[ v_i(h) = \begin{cases} \frac{v_i}{2}(1 - e^{-\alpha_i(h-h_0)^2}), & h_0 \leq h < h_i \\ \frac{v_i}{2}e^{-\beta_i(h-h_i)^2} - 1, & h_i \leq h \leq h_i \end{cases} \]

Where \( i = 1 \equiv \text{NS} \) and \( i = 2 \equiv \text{EW} \).

First, let us compute the integral \( I \)

\[ I = \int_{h_0}^{\infty} v_i(h) dh \]

as a sum of two integrals: \( I_1 \) and \( I_2 \) (same as (241))
One then can utilize these calculations to compute a simplified total ionospheric delay based on the equations that were previously presented in this paper.

4.2 Examples of description of the simplified ionospheric Doppler frequency

In the following two examples we will utilize exactly the same simplified profile that we utilized in simplified ionospheric delay.

Example 1: Assuming the first profile of the azimuth gradient is given by (246)-(251); hence, \( \frac{dc_{i1}(t)}{dt} \) can be obtained from the derivative of three expressions:

\[
\frac{dc_{i1}(t)}{dt} = \frac{dc_{i2}(t)}{dt} + \frac{dc_{i3}(t)}{dt}
\]

The first derivative, \( \frac{dc_{i1}(t)}{dt} \), can be computed from

\[
\frac{dc_{i1}(t)}{dt} = \sum_{j=1}^{2} \left[ a_{ij} v_{ij}(t) + a_{mj} \frac{dv_{mi}(t)}{dt} \right]
\]

Where

\[
a_{v11} = h_i(t) \frac{dh_i(t)}{dt} - h_0(t) \frac{dh_0(t)}{dt}
\]

\[
a_{v21} = h_i(t) \frac{dh_i(t)}{dt} - h_1(t) \frac{dh_1(t)}{dt}
\]

\[
a_{v22} = \frac{h_i^2(t) - h_0^2(t)}{2}
\]

The second derivative, \( \frac{dc_{i2}(t)}{dt} \), can be computed from

\[
\frac{dc_{i2}(t)}{dt} = a_{vml} v_{mi}(t) + a_{vmi} \frac{dv_{mi}(t)}{dt}
\]

Where

\[
a_{vml} = \frac{1}{a_i(t)^2} - \frac{1}{\beta_i(t)^2}
\]

The third derivative, \( \frac{dc_{i3}(t)}{dt} \), can be computed from

\[
\frac{dc_{i3}(t)}{dt} = a_{\alpha_1} + a_{\alpha_1} \frac{d\alpha_1(t)}{dt} + a_{\alpha_1} \frac{d\alpha_1(t)}{dt}
\]

Where

\[
\alpha_1 = \frac{2}{a_i(t)^2} \left[ \frac{d\beta_i(t)}{dt} \frac{1}{\beta_i(t)^2} - \frac{d\alpha_1(t)}{dt} \frac{1}{a_i(t)^2} \right]
\]

\[
a_{\phi mi} = \frac{1}{a_i(t)^2} - \frac{1}{\beta_i(t)^2}
\]

The third derivative, \( \frac{dc_{i3}(t)}{dt} \), can be computed from

\[
\frac{dc_{i3}(t)}{dt} = a_{\beta_1} + a_{\beta_1} \frac{d\beta_1(t)}{dt} + a_{\alpha_1} \frac{d\alpha_1(t)}{dt}
\]

Where
Example 1: In Figs. 13-16 the following are depicted: (1) the ionospheric gradient PED profile vs. the height of the ionosphere and (2) the obliquity factor $m_1(t)$ (or $m_1(t)$) and the azimuth obliquity factor $m_2^R(t)$ (or $m_2^R(t)$) as a function of elevation $\theta(t)$ (or zenith $\zeta(t)$) angle. First, the ionospheric gradient PED profile of the first or the second scenario increases monotonically from 0/0 els/m$^2$ to 60/96 Mels/m$^2$ as a function of height from 50 km to 350 km in EastWest (EW)/NorthSouth (NS) and then decreases monotonically from 60/96 Mels/m$^2$ to 0/0 els/m$^2$ as a function of height from 350 km to 1000 km from EW/NS. In either profile the maximum ionospheric gradient PED profile 60/96 Mels/m$^2$ in EW/NS corresponds to the ionospheric height of 350 km. Now this height can change and so can the maximum value of the ionospheric gradient PED profile in the simulation setting.

One then can utilize these calculations to compute a simplified total ionospheric Doppler frequency based on the equations that were presented previously in this journal paper.
As can be seen from Figs. 13-16, gradients become increasingly influential at lower elevation angles.

Example 2: The second simulation example is the illustration of the RF L-band (1-2 GHz) ionospheric propagation channel model in Figs. 17-20. This basic simulation contains the description of: (1) refractive index \([5]\), \(n\); (2) refractivity, \(N\); (3) ionospheric zenith delay, \(D_z[m]\) and (4) ionospheric azimuth critical zenith delay, \(D_{\alpha\beta}[m]\) as a function of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\). By taking a look at the plots in Figs. 17-19 we observe that refractivity index and refractivity/ionospheric delay \([11]\) increase/decrease monotonically for values of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\).

The smallest refractive index \([5]\) 0.9999596918069695 or the smallest (largest in magnitude) refractivity \(\Delta n\) corresponds to values of \(f = 1\text{GHz}\) and the largest refractive index \([5]\) 0.999990229517423 or the largest (smallest in magnitude) refractivity \(\Delta n\) correspond to values of \(f = 2\text{GHz}\). The ionospheric zenith delay decreases from 16 m at 1 GHz to 4 m at 2 GHz and the ionospheric azimuth critical zenith delay decreases from 8.5 m at 1 GHz to 1 m at 2 GHz.

Example 3: The third simulation example is the illustration of the ionospheric Doppler frequency elements such as \(\delta\text{TEC}/\delta x_i\), \(dc(t)/dt\), \(\alpha'(t)\), and, \(\beta(t)\) vs. time, zenith angle, \(\zeta(t)\), is the sum of two angles \(\alpha'(t)\) with \(\beta(t)\), and \(\zeta'(t)\), (see Figs. 21-24).

Example 4: The fourth example of a simulation of the RF L-band (1-2 GHz) ionospheric propagation channel model 2 is illustrated in Figs. 25-28.

This basic simulation contains the description of: (1) total ionospheric delay \([11]\), \(D_z[m]\); (2) total ionospheric Doppler, \(D_{\alpha\beta}[\text{MHz}]\); as a function of time of day \([h]\) and of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\) using the first envelope of the PED profile (see Fig. 13). By taking a look at the plots in Figs. 25-28 we observe that total ionospheric delay \([11]\) and total ionospheric Doppler increase/decrease monotonically as a function of time/of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\).

The smallest values of the total ionospheric delay \([11]\) of 8.59 m \((f = 1,176.45\text{MHz})\) and 4.79 m \((f = 1,575.42\text{MHz})\) correspond to \(t = 0\) h at zenith or \((\zeta = 0)\) and largest values of the total ionospheric delay \([11]\) of 183 m \((f = 1,176.45\text{MHz})\) and 102 m \((f = 1,575.42\text{MHz})\) correspond to \(t = 2.52\) h at the horizon or \((\zeta = 90^\circ)\).

Example 5: The fifth example of a simulation of the RF L-band (1-2 GHz) ionospheric propagation channel model 3 is illustrated in Figs. 29-32. This basic simulation contains the description of: (1) total ionospheric delay \([11]\), \(D_{\alpha\beta}[m]\); (2) total ionospheric Doppler, \(D_{\alpha\beta}[\text{MHz}]\); as a function of time of day \([h]\) and of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\) using the second envelope of the PED profile (see Fig. 14).

By taking a look at the plots in Figs. 29-32 we observe that total ionospheric delay \([11]\) and total ionospheric Doppler increase/decrease monotonically as a function of time/of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\).

The smallest values of the total ionospheric delay \([11]\) of 8.59 m \((f = 1,176.45\text{MHz})\) and 4.79 m \((f = 1,575.42\text{MHz})\) correspond to \(t = 0\) h at zenith or \((\zeta = 0)\) and largest values of the total ionospheric Doppler of 121 Hz \((f = 1,176.45\text{MHz})\) and 165 Hz \((f = 1,575.42\text{MHz})\) correspond to \(t = 2.52\) h at the horizon or \((\zeta = 90^\circ)\).

Example 6: The sixth example of a simulation of the RF L-Band (1-2 GHz) ionospheric scintillation amplitude, \(S_{4}[\mu]\), and phase standard deviation, \(\sigma_p[\mu]\), is illustrated in Figs. 33-36. This basic simulation contains the description of: (1) ionospheric scintillation amplitude standard deviation, \(S_{4}[\mu]\); (2) phase standard deviation, \(\sigma_p[\mu]\), as a function of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\) having the PED, external magnetic field, and electron collision as parameters whose significance is given in sequential order.

By taking a look at the plots in Figs. 33-36 we observe that ionospheric scintillation amplitude, \(S_{4}[\mu]\), and phase, \(\sigma_p[\mu]\), standard deviation decreases monotonically as a function of frequency \(1\text{GHz} \leq f \leq 2\text{GHz}\) on the top two plots: the largest/smallest values of \(S_{4}\) are: \(1300/276[\mu]\) \((f = 1.1765\text{GHz})\) and 761/154 \([\mu]\) \((f = 1.547542\text{GHz})\) and \(\sigma_p[\mu]\) are: 27.2/3.55 \([\mu]\) \((f = 1.1765\text{GHz})\) and 21.4/2.65 \([\mu]\) \((f = 1.547542\text{GHz})\) correspond to \(n_e = [1 \times 10^{12}\text{el/m}^3]\), \(B = [0 \times 25 \times 10^{-4}\text{T}]\), \(\theta_0 = [0 \times 89]\), and \(v_{\text{eff}} = [0 \times 147.53]\text{Hz}\).
On the bottom two plots: the largest/smallest values of $S_4$ are: 1300/249 $[\mu \text{V}]$ ($f = 1.1765 \text{ GHz}$) and 761/139 $[\mu \text{V}]$ ($f = 1.547542 \text{ GHz}$) and $\sigma_\theta$ are: 27.2/3.19 $[^\circ]$ ($f = 1.1765 \text{ GHz}$) and 21.4/2.39 $[^\circ]$ ($f = 1.547542 \text{ GHz}$) correspond to $n_e = [0.9 \ 1] \times 10^{12} \text{el/m}^3$, $B = [0 \ 25 \times 10^{-4}] \text{T}$, $\theta_\theta = [0 \ 89]^\circ$, and $v_{\text{eff}} = [0 \ 147.53] \text{Hz}$.

An example amplitude variation and phase fluctuation of the basic scintillated component from the simulation are shown in [95]. The amplitude scintillation intensity is represented by the $S_4$ index which is defined as the normalized standard deviation of the signal power. The phase scintillation is characterized by the phase standard deviation. In the example shown in [95], the amplitude scintillation index $S_4$ is 1.01 and the phase scintillation standard deviation of 0.44 radians [96]. It is possible that some of the data is due to a 1-parameter ionospheric model used. The 4-parameter ionospheric model that is proposed in this paper should be able to present amplitude and phase scintillation data that are much lower (three to six orders of magnitude) than the ones in [95] or contrast with theoretical amplitude and phase scintillation depicted in Figs. 33-36.

I did however find a good reference from Abdullah et al. 2009 [97] that have employed a dual-frequency GPS data collected during the one-month ionospheric experimental campaign that was used for TEC and scintillation computation and analysis. The TEC with 15 sec interval were computed from combined L1 and L2 code-pseudorange and carrier phase measurements. Whereas, the scintillation parameter S4 index was computed as a standard deviation of the received signal power normalized to average signal power every 60 sec on L1. A corrected $S_4$ (without noise effects) was also computed and used in the analysis.

Abdullah et al. 2009 [97] have reported daily maxima VTEC for Parit Raja (PR) ranged from 38-100 TECU, which is generally higher than those of Sipitang, which ranged from 30-42 TECU and corrected amplitude scintillation index $S_4$ no greater than 0.1 and average on the order of 0.03-0.04 [97]. The dual frequency corrected scintillation index $S_4$ experimental results are in complete agreement with the 4-parameter ionospheric model presented in this journal paper.

## 6 Conclusions

In conclusion we have demonstrated in theory and simulations that the 1-parameter ionospheric channel model that is based on estimating the ionospheric delay is insufficient to enable the GNSS receiver to reliably acquire and track GNSS signals under severe ionospheric conditions.

In order to solve this problem Giftet Inc. has proposed a 4-parameter ionospheric model. The 4-parameter ionospheric model is based on estimating four parameters: ionospheric delay, Doppler, amplitude, and phase scintillation. The 4-parameter ionospheric model presents a revolutionary advancement of our understanding of the ionospheric effects in the RF signals.

In order to accurately estimate the ionospheric Doppler, amplitude, and phase scintillations we have performed some of the most amazing mathematical calculations.

This paper presents one of the most daunting tasks ever undertaken by a single scholar to accurately model, analyze, and simulate very accurately the entire description of the ionospheric channel effects.

After the theory and simulation presented in this paper are completely verified and accurately tested, the implementation thereof in a single frequency GNSS receiver will enable it to perform just as good as or even better than a differential GNSS receiver.

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This journal paper is dedicated to four special men in my life: my grandfather, Xhevdet Progri, my dear father, Fiqiri Progri, my father’s first cousin Dr. Peter Demir, and the forty-first President of the United States of America, George H.W. Bush.
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1 This paper presents one of the most daunting tasks ever undertaken by a single scholar to very accurately model, analyze, and simulate the entire description of the ionospheric channel effects.

2 These four parameters are not really independent of each other but they start our as independent parameters and they can be observed and measured independently of each other; hence, a more appropriate definition would be semi-independent.

3 Ditto i.

4 It is assumed that most of the conductivity comes as result of the presence of a magnetic field ignoring the average time interval between collisions as a result of the drift movements of the electrons and the effects from free ions [50].

5 It is possible to remove only the first order ionospheric refraction terms (up to the 99% of the total refractive effect) using their dependence on frequency [34].

6 Equation (9) in Matteo et al. 2009, [24] is an equivalent expression for \( \alpha = 0.5C_0 \) given by (41); however, there is an error in units in (9) in Matteo et al. 2009, [24].

7 Equation (41) uses the same units; i.e., \( \text{els}^{-1}\text{m}^{-2} \equiv \text{els}^{-1}\text{Wm}^{-3} \equiv \text{els}^{-1}\text{ms}^{-1} \equiv \text{els}^{-1}\text{s}^{-2} \) [22], [24].

8 The earliest calculation of the ionospheric critical frequency comes from [Mitra 1935, pg. 153 [44]]. Mitra gives a value of \( 5 \times 10^6 \) Hz \( \equiv 5.4 \) MHz for the ionospheric critical frequency in Calcutta (presently Kolkata, India).

9 The ionospheric delay, \( D_I(t) \), is a subcomponent of the \( D_e(t) \) assuming that all other channel effects are ignored then

\[
D_I(t) = D_{I,s}(t) : t \in [0, T)
\]

10 The detailed discussion of the Snell’s law, Fresnel’s equations, and ray’s propagation path that provide a complete derivation and interpretation of the ray path delay will be published in a separate journal paper.

11 Refractivity of ionosphere is actually a rank-2 (or second) tensor (a 3x3 matrix) or a mathematical object analogous to but more general than a vector, represented by an array of components that are functions of the coordinates of a space [5]. Tensors are geometric objects that describe linear relations between geometric vectors, scalars, and other tensors [48]. In this case the propagation of the RF signal cannot simply be described by refractive indices except for polarizations along principal axes [5].

12 Andrei et al. (2009, [41]) for \( h_1 \) a value in the range of 300-500 km is typical.

13 Blanch et al. (2004, [53]) provide a model to include two properties of the ionosphere: the vertical density profile (ignored in the thin shell models which will be subject of a separate journal paper) and the decorrelation of the ionospheric delay as a function of distance (ignored in the tomographic models.)

14 To this purpose, we generalize the two dimensional model that is used in Kriging [51]-[53], [55] and is based on multivariate Gaussian random fields [51], [52], [60] by introducing l layers instead of a single thin shell. Another approach is “the conical domain approach” proposed by Sparks et al. (2004, [57]).

15 There was a typo in (2) in [69] the ‘+’ should be ‘−’ as in (96) or (89).

16 The detailed discussion and the complete derivations and interpretation of this section will be published in a separate journal paper.