Arpu Kuilpu: An H5 from the outer main belt

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Abstract—On June 1, 2019, just before 7:30 p.m. local time, the Desert Fireball Network (DFN) detected a −9.3 magnitude fireball over South Australia near the Western Australia border. The event was observed by six fireball observatories, and lasted for 5 s. One station was nearly directly underneath the trajectory, greatly constraining the trajectory solution. This trajectory’s backward numerical integrations indicate that the object originated from the outer main belt with a semimajor axis of 2.75 au. A light curve was also extracted and showed that the body experienced very little fragmentation during its atmospheric passage. A search campaign was conducted with several DFN team members and other volunteers. One 42 g fragment was recovered within the predicted fall area based on the dark flight model. Based on measurements of short-lived radionuclides, the fragment was confirmed to be a fresh fall. The meteorite, Arpu Kuilpu, has been classified as an H5 ordinary chondrite. This marks the fifth fall recovered in Australia by the DFN, and the smallest meteoroid (∼2 kg) to ever survive entry and be recovered as a meteorite.

INTRODUCTION

The Desert Fireball Network (DFN) is a system of automated photographic all-sky fireball observatories covering over 2.5 million km² of Australian outback (Howie, Paxman, Bland, Towner, Cupak, et al., 2017). It is the largest single photographic fireball network in the world. The DFN’s primary objective is to recover meteorite falls using fireball observations. These meteorites with precise orbital information can help us better understand small bodies and the debris they generate in the solar system. The DFN is a partner network of the Global Fireball Observatory, a collaborative project consisting of 18 international partners from all over the world (Devillepoix et al., 2020). Our observatories are optimized to precisely observe bright meteors (“fireballs”) when they impact the atmosphere, providing invaluable information about the asteroidal debris in the inner solar system (Devillepoix et al., 2019).

Over 40 meteorite falls in total have been recovered with the assistance of fireball observations (Colas et al., 2020; Devillepoix et al., 2020; Gardiol et al., 2021; Granvik & Brown, 2018). The most numerous of the meteorite fall types recovered are H chondrites. This is somewhat unexpected, as the most common Antarctic meteorite type are L chondrites, followed by H chondrites (Binzel et al., 2015). More meteorite fall recoveries can clarify whether this trend is significant. Within this study, we discuss the DFN fireball observations and data reduction, which led to the recovery of Arpu Kuilpu, an H5 chondrite. This is the fifth meteorite recovered in Australia by the DFN (Bland et al., 2009; Devillepoix et al., 2018; Dyl et al., 2016; Sansom et al., 2020).
FIREBALL OBSERVATION AND TRAJECTORY

In total, six DFN camera systems detected the event internally referenced as DN190601_01 (Table 1), using the on-board algorithm of Towner et al. (2020). The location of these is mapped in Fig. 1 along with the fireball as observed by each system.

The DFN captures long exposure images and obtains timing information from de Bruijn sequence dashes (Howie, Paxman, Bland, Towner, Cupak, et al., 2017; Howie, Paxman, Bland, Towner, Sansom, et al., 2017; Howie et al., 2020). This methodology of encoding a de Bruijn sequence into the shutter frequency allows us to obtain absolute timing of the fireballs without the need of an additional subsystem—drastically reducing the size, cost, and power requirements.

Observing conditions during the fireball were ideal, enabling three cameras to observe from over 400 km away. Encoded timing was acquired for five sites, but was not resolvable for DFNEXT051. The best convergence angle between the four observations was a reasonable 36°. However, one camera, DFNEXT041 Hughes, was nearly directly underneath the fireball trajectory. This observation significantly helped to constrain the fit and was instrumental in recovering the small 42 g stone meteorite. The observations from the six DFN fireball observatories provided 256 data points to fit the trajectory (Fig. 1).

Astrometry from Photographs

After time decoding and centroiding the fireball positions, we use the same method as described by Devillepoix et al. (2018) to convert the pixel positions in the image to true horizontal coordinates. The calibration frame used for most observatories was taken at 2019-06-01T10:00:30, 7 min after the fireball, using ≃1000–1500 reference stars for close cameras, and 2000+ using background subtraction on farther observatories (Mt Barry and Kanandah) in order to maximize the number of stars low on the horizon. We propagate uncertainties all the way through, taking into account both pixel picking uncertainties and residuals on the star fit. This yields formal astrometric uncertainties on the order of 1.5′ (1σ) for most records. Note that the same angular precision is achieved throughout the frame, thanks to a change of pixel pitch due to the stereographic lens projection, from 120″ at the zenith, down to 70″ at 5° on the horizon.

Astrometric data files for all cameras are available in the supporting information, and documentation associated with these data are available at https://dfn.gfo.rocks/data_documentation.html.

Photometry from Video Records

The latest generation observatories of type DFNEXT are equipped with a monochrome digital...
video camera (Point Grey/FLIR BFLY-U3-23S6M-C, Fujinon Fisheye 1:1.4/1.8 mm). The effective field of view is an all-sky circle that fits in a square of \(\approx 1080\) pixels. This addition in parallel to the still high-resolution imager was introduced to not only observe fainter meteors than in the photographs, provide observational coverage during the dead time (3 s = 10\%) between long exposures, yield better photometry, and also observe during the daytime. These systems run a modified version of Freture (Colas et al., 2020) to detect and save fireballs (code available at https://github.com/desertfireballnetwork/freeture_DFN). Frames are saved as 8-bit lossless compressed FITS files.

At the time of the event, the recording software deployed on the observatories was still in development: it was not recording long exposure calibration frames for astrometry and also failed to record the beginning of the fireball on DFNEXT041 (Table 1). Although the record was saturated in the brightest part of the fireball, we used a point spread function (PSF) fitting photometry technique to estimate the true luminosity. Compared to aperture photometry, PSF photometry mitigates the saturation issue to some extent, but still cannot guarantee the accuracy of the results. To calibrate the records, we also perform PSF fitting on \(\alpha\) Crucis to use it as a reference luminosity source \(V_{\text{mag}} = 1.25\) after atmospheric extinction taken into account). Finally, we use the triangulation solution (see the Trajectory Modeling section) to determine the range of each observation in order to obtain the distance-corrected (normalized to 100 km distance) absolute magnitude (Fig. 2). We also track the brightness of a secondary fragment for about half a second (Fig. 2), \(\approx 2\) magnitudes fainter than the main mass just after it broke apart. This trailing fragment is also visible in the still image (Fig. 3), in which a slight lateral deviation is visible. We note that in some frames, up to five fragments are visible in the video (main mass, aforementioned fragment, and three smaller ones; see the middle frame in Fig. 3).

### TRAJECTORY MODELING

#### Initial Triangulation

To determine a trajectory from the fireball observations, we utilized a modified version of the straight line least squares method of Borovička (1990) and Sansom et al. (2015). The velocity and mass information is then extracted by using an extended Kalman smoother (EKS; Sansom et al., 2015). The velocity and mass information is then extracted by using an extended Kalman smoother (EKS; Sansom et al., 2015). Finally, camera observations are weighted according to their distance from the observed fireball, as the increased distance increases the cross-track uncertainties (Fig. 4).

The fitted trajectory was mostly constrained by the DFNEXT041—Hughes, which was nearly directly underneath the trajectory (Table 1). The fireball was seen for 5 s from an altitude of 86.4–28.9 km. The meteoroid impacted the upper atmosphere at a modest angle (\(\sim 50^\circ\)). For further details about the fireball trajectory and dynamic information, please refer to Table 2.

The relatively featureless light curve for this fireball (Fig. 2) indicates few large fragmentation events. The increase in brightness at 2.07 s occurs at an atmospheric ram pressure of 0.11 MPa, while the event forming the largest secondary fragment at 4.6 s corresponds to pressures of 0.82 MPa. This is similar to the pressures experienced during fragmentation of other H5 chondrite falls such as the Grimsby, Kosice, and Ejby (Borovička et al., 2013; Brown et al., 2011; Spurný et al., 2017). Interestingly, the progenitor meteoroids for these falls span three orders of magnitude in mass, but each has initial fragmentations.

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Table 1. Locations of Desert Fireball Network observatories that obtained photographic records of DN190601_01, and nature of data obtained. Times are relative to first fireball observation at 09:53:04.326 UTC on June 1, 2019 (from the Hughes camera). Timing information was unable to be extracted from the observations made from DFNEXT051—Kanandah.

| Observatory name       | Latitude (°) | Longitude (°) | Altitude (m) | Instrument record | Range* (km) | Start time observed | End time observed |
|-------------------------|--------------|---------------|--------------|-------------------|-------------|---------------------|------------------|
| DFNEXT041—Hughes        | 30.652867 S  | 129.700608 E  | 144          | P, V              | 70          | 0.0                 | 5.0              |
| DFNEXT027—Forrest       | 30.858058 S  | 128.115032 E  | 166          | P                 | 190         | 0.67                | 4.47             |
| DFNEXT043—Kybo          | 31.024935 S  | 126.591345 E  | 176          | P                 | 332         | 0.87                | 4.35             |
| DFNSMALL26—Rawlina      | 29.742252 S  | 125.750340 E  | 209          | P                 | 422         | 1.67                | 2.87             |
| DFNEXT032—Mount Barry   | 28.516433 S  | 134.886256 E  | 169          | P                 | 538         | 1.74                | 3.64             |
| DFNEXT051—Kanandah      | 30.901091 S  | 124.884540 E  | 195          | P                 | 495         | –                   | –                |

*P: Photographic record (long-exposure high resolution image, see the Astrometry from Photographs section). V: lossless compressed digital video (30 frames per second, see the Photometry from Video Records section). Ranges are from when the meteoroid was at 65 km altitude. *Distance from the observatory to the meteoroid at 65 km altitude.
around ~0.1 MPa. This low fragmentation pressure range was proposed to be associated with reassembled and cemented material by Borovička et al. (2020), meaning this feature may be constant within the $10^0$–$10^3$ kg range for ordinary chondrites.

Estimating Initial and Terminal Masses

We initially use the $\alpha$–$\beta$ methods of Gritsevich et al. (2012) and Sansom et al. (2020) to determine a first approximation of final mass and velocity to initiate Bayesian filtering techniques. The dimensionless ballistic ($\alpha$) and mass loss ($\beta$) parameters calculated for this fireball are $\alpha = 36.16$ and $\beta = 1.13$ (Fig. 5). Assuming meteoroid properties, such as a brick-like shape ($A = 1.55$; Gritsevich, 2009), a bulk density of 3500 kg m$^{-3}$, and a shape change parameter of 2/3 (see Sansom et al. [2019] and references therein), a terminal mass of 60 g is predicted, with an estimated initial mass of 1.4 kg. Assuming a spherical shape gives a minimum initial mass of 680 g, with a meteorite mass of 30 g. The fitted final velocity is 5.1 km s$^{-1}$.

The EKS filter predicts changes to the state (position, velocity, and mass) using the single body aerodynamic equations (Sansom et al., 2015). The state is initiated at $t_f = 5.0$ s, with state values taken from initial fitting above, with 10% uncertainties in position and velocity, and 100% in mass; for velocity $5.1 \pm 0.5$ km s$^{-1}$ and mass $0.06 \pm 0.06$ kg. Based on the deceleration of the object and a shape parameter once again of $A = 1.55$, the final velocity was $4.74 \pm 0.19$ km s$^{-1}$, and the initial and final masses were estimated to be $2.2 \pm 0.1$ kg and $0.04 \pm 0.01$ kg, respectively (Sansom et al., 2015). This is very consistent with the 42 g meteorite sample recovered.

In order to get the best estimate of initial velocity for orbit calculation, we need to account for any variation in the trajectory from a straight line due to

Table 2. Atmospheric trajectory information for event DN190601.01. Apparent entry radiant and velocities are in an inertial.

| Time (UTC) | Beginning | Terminal |
|------------|-----------|----------|
| Longitude (WGS84; °E) | 130.132155 | 129.631139 |
| Latitude (WGS84; °S) | 30.610298 | 30.623636 |
| Height (km) | 86.44 ± 0.03 | 28.90 ± 0.01 |
| Velocity (km s$^{-1}$) | 17.58 ± 0.04 | 4.74 ± 0.19 |
| Angle from local horizontal (°) | 50.0 | 49.7 |
| Apparent entry radiant (RA °) | 210.48 ± 0.01 | – |
| Apparent entry radiant (Dec °) | -22.08 ± 0.01 | – |
| Calculated mass* (kg) | 2.2 ± 0.1 | 0.04 ± 0.01 |
| Total duration | 5.0 s | |
| Number of data points | 256 | |
| Number of observing stations | 6 (5 with timing) | |
| Recovery location (WGS84) | 129.631274°E 30.623689°S | |
| Recovered mass | 42 g | |

ECI = earth-centered reference frame.

*Assumptions listed in the Estimating Initial and Terminal Masses section.
the atmosphere. This involves recalculating the trajectory and EKS modeling for the observations above 60 km altitude. In doing so, the initial velocity is calculated to be $17.58 \pm 0.04 \text{ km s}^{-1}$.

**RADIANT AND HELIOCENTRIC ORBIT DETERMINATION**

As stated in the Trajectory Modeling section, the straight-line trajectory may not be relied upon to provide representative entry vectors of the meteoroid from which to calculate orbits. Thus, the initial radiant and velocity were calculated using only the top of the trajectory (>60 km) before major deceleration. This apparent radiant and velocity was used to calculate the preatmospheric orbit of the meteoroid using the integration method of Jansen-Sturgeon et al. (2019). The Monte Carlo results are illustrated in Fig. 6, with orbital values given in Table 3.

A Monte Carlo simulation was performed to characterize the orbital characteristics and history for Arpu Kuilpu. One thousand test particles were generated within the formal triangulation uncertainties assuming a Gaussian distribution (Table 3). These particles were then propagated backward in time 10 Myr using the IAS15 integrator implemented through the Python-based REBOUND module, taking into account all relevant planetary perturbations (all planets, the Sun, and the Moon; Rein & Liu, 2012; Rein & Spiegel, 2015). We found that 35.2 ± 4.2% of the particles were unstable over the previous 10 kyr, having numerous close encounters with Jupiter and terrestrial planets. As previously noted by Tancredi (2014), the unpredictability of the evolution of an object's orbit on a 10 kyr time scale is a diagnostic feature of Jupiter-family comets. However, given our simulations, the Arpu Kuilpu meteoroid's orbital evolution is more likely consistent with debris from the main asteroid belt. This is consistent with the findings of Shober, Jansen-Sturgeon, Bland, et al. (2020), Shober, Jansen-Sturgeon, Sansom, et al. (2020), and Shober et al. (2021), which showed that nearly all of the centimeter- to meter-sized debris on Jupiter-family comet-like orbits are initially from the main asteroid belt. After ~20 kyr, the orbital history of Arpu Kuilpu becomes too challenging to characterize. A significant number of close encounters for all the particles removes the ability to gain further information through backward integration.

The orbit for Arpu Kuilpu (Table 3) is very similar to Košice, Hamburg, and Ejby, other H chondrite falls recovered with orbits elements of $a \sim 2.7 \text{ au}$, $e \sim 0.65$, $\text{inc} < 2^\circ$. Granvik and Brown (2018) determined the likely escape routes from the main belt for Košice and Ejby based on a debiased Near-Earth Object (NEO)
orbital model (Granvik et al., 2018). They found that these meteorites’ most likely transport mechanisms were the 3:1 or 5:2 mean motion resonances. Thus, Arpu Kuilpu should be expected to have a similar history. Based on the NEO model, the meteorites also had a very high probability of originating from the Jupiter family.

Fig. 5. Trajectory data with velocities normalized to the velocity at the top of the atmosphere \( V_0 = 17.58 \text{ km s}^{-1}; \) Table 2) and altitudes normalized to the atmospheric scale height, \( h_0 = 7160 \text{ m}. \) The best fit to equation 10 of Gritsevich (2009) results in \( \alpha = 36.16 \) and \( \beta = 1.13 \) and is shown by the black line. (Color figure can be viewed at wileyonlinelibrary.com.)
comet (JFC) population, but, as previously stated, this is expected for meteorites coming from comet-like orbits (Shober, Janson-Sturgeon, Bland, et al., 2020; Shober, Janson-Sturgeon, Sansom, et al., 2020, 2021).

Wind Modeling and Dark Flight

We use the Weather Research and Forecasting (WRF) model version 4.0 with dynamic solver ARW (Advanced Research WRF; Skamarock et al., 2019) for atmospheric situation numerical modeling. The weather profile (Fig. 7) includes wind speed and direction at heights ranging up to 30 km, fully covering the dark flight altitudes (<28.9 km). The dark flight trajectory of the meteorite was substantially affected by the atmospheric winds, dominated by the subtropical jet stream at altitudes approximately 10–15 km. However, the winds during the Arpu Kuilpu fall were not as strong as in the case of Murrili (Sansom et al., 2020) or Dingle Dell (Devillepoix et al., 2018). From the 3D data produced by the WRF/ARW modeling software, we extract weather parameters relevant for the dark flight modeling (wind speed, wind direction, pressure, temperature, and relative humidity). We then interpolate these conditions in time and 3D coordinates.

Table 3. Pre-encounter orbital parameters expressed in the heliocentric ecliptic frame (\textit{J2000}) and associated \(1\sigma\) formal uncertainties.

| Epoch (TDB) | Semi-major axis \(a\) (AU) | Eccentricity \(e\) | Inclination \(i\) | Argument of perihelion \(\omega\) | Longitude of the ascending node \(\Omega\) | Perihelion distance \(q\) (AU) | Aphelion distance \(Q\) (AU) | Corrected radiant (RA) \(\alpha\) | Corrected radiant (Dec) \(\delta\) | Geocentric speed \(V_g\) (m \(s^{-1}\)) | Tisserand’s parameter \(T\) |
|-------------|-----------------------------|-------------------|-----------------|-------------------------------|---------------------------------|----------------------------|---------------------|-----------------|-----------------|-------------------|------------------|
| 2019-06-01  | 2.75 ± 0.03                 | 0.671 ± 0.003     | 2.03 ± 0.01     | 43.25 ± 0.02                  | 250.36 ± 0.01                   | 0.90 ± 0.01                | 4.59 ± 0.05         | 215.76 ± 0.01    | −20.04 ± 0.01    | 13,280 ± 50        | 2.97 ± 0.02       |

Fig. 6. Orbit calculated from observations for the \textit{DN190601_01} meteoroid (red). The orbit is displayed in the barycentric inertial reference frame (BCI), coinciding with that of the International Celestial Reference System (ICRS). Planets and the sun are displayed in gray. (Color figure can be viewed at wileyonlinelibrary.com.)
during dark flight modeling, following the meteorite fall trajectory at exact time and space (Towner et al., 2021).

From this wind profile and the mass and shape estimates above, we generate fall position predictions on the ground, showing hypothetical locations for a range of possible masses. Using a range of possible masses with all else equal, one can construct a ground fall line, to guide searching. However, this line is a construct of scenarios; it does not imply that pieces are present all along the line. If one further considers all possible scenarios, such as different wind model predictions and different shapes as well as possible masses, a heat map of possible fall positions can be made (see for example Moilanen et al., 2021). However, in Fig. 8 we plot lines for only three possible scenarios for clarity (white lines), as detailed in the caption. Other possible scenarios are easily generated (see supporting information). The environmental conditions at this fall were generally quite stable, and this can be seen in Fig. 8, where the fall lines for different wind models, but for the same hypothetical shape, are relatively close together.

SEARCH AND RECOVERY

Although the expected main mass was quite small, this particular fall was prioritized for searching thanks to several factors:

- Very clean data, both for trajectory and dynamics, thanks to a close-by viewpoint (Hughes station). This gave us confidence the search area would be very well constrained, and the mass would be well bracketed.
- Lack of recent rain and therefore vegetation likely sparse in the area.
- Relatively low-risk access to the site.
- Interesting JFC-like orbit with a larger semimajor axis.

A team of six people carried out a searching expedition for 2 weeks from July 17, 2019 onward. On site, the searching conditions were excellent, with most vegetation being sparse low bluebush and saltbush, with little to no grass, and sizable barren clay pan areas. The meteorite was recovered on the second day of searching, but searching continued for the entire trip. As a searching strategy, the spacing between party members was 2–4 m, even though the sightlines were excellent, based on the probable scenario of a relatively small meteorite. The meteorite was spotted about 1 m away from the searcher.

The meteorite was on the ground for approximately 6 weeks: The nearest weather station is Forrest airport, in Western Australia, and during the period of June 1 to July 20, records show 8.4 mm of rain. Nullarbor rainfall is relatively localized, mainly resulting from thunderstorms, so more significant rainfall at the fall site cannot be ruled out. During winter, it is also not unusual for ground fog and dew to occur in the Nullarbor.

The Meteorite Sample

The recovered meteorite sample consists of a single, rounded stone ~1–1.5 cm in diameter, with no noticeable terrestrial alteration. A portion of the meteorite was set in epoxy, polished, and carbon coated to enable classification via scanning electron microscope and electron probe microanalysis (Fig. 9). Arpu Kuilpu’s mineralogy and texture is typical of equilibrated ordinary chondrites, containing relict chondrules, metal, and troilite grains suspended in a recrystallized silicate matrix. The relative abundances of plagioclase, olivine, pyroxene, metal, and troilite within the sample, combined with the chemical compositions of its olivines, pyroxenes, and chromites, confirm that Arpu Kuilpu is an H5 chondrite.

Cosmogenic radionuclide concentrations have been analyzed by means of nondestructive high purity germanium gamma spectroscopy. The counting efficiencies have been calculated using thoroughly tested Monte Carlo codes. One specimen of Arpu Kuilpu was
measured in the underground laboratories at the Laboratori Nazionali del Gran Sasso (LNGS; Laubenstein, 2017) for 40.63 days (183 days after the fall). Very low activity of $^{60}\text{Co}$ ($0.6 \pm 0.3$ dpm kg$^{-1}$) suggests that the preatmospheric size of the Arpu Kuilpu meteoroid was rather small and no significant production of secondary thermal neutrons took place within the meteoroid during its recent cosmic ray exposure in space. The measured $^{26}\text{Al}$ activity is consistent with that expected for a small-size H chondrite (Bhandari et al., 1989; Bonino et al., 2001; Leya & Masarik, 2009).

When we compare the radionuclide concentrations with cosmic ray production estimations for $^{26}\text{Al}$ (Leya

Fig. 8. The fall area lies within the South Australia area of the Nullarbor. Fall lines are shown with predicted masses for three scenarios; a hypothetical chondritic sphere (red circle) and chondritic cylinder (red diamond) using the preferred wind model, and a chondritic sphere with an alternative wind model (white circle). Meteorite recovery position shown as filled black star, with an inset image of the meteorite as found with GPS unit for scale. Blue tracks show the areas searched by the team. The location of the recovered meteorite is about 30 m from the chondritic sphere fall line, but corresponding to a mass of about 20 g, compared to a recovered mass of 42 g. (Color figure can be viewed at wileyonlinelibrary.com.)

Fig. 9. Mineral map of Arpu Kuilpu section constructed from Fe, Ni, Mg, Ca, S, Si, Cr element maps obtained by scanning electron microscopy (SEM) with energy dispersive X-ray spectrometry (EDS). Blue = olivine; teal = orthopyroxene; green/brown = clinopyroxene; yellow/green = plagioclase; orange = troilite; red = phosphate; pink = chromite; gray = FeNi metal. (Color figure can be viewed at wileyonlinelibrary.com.)
& Masarik, 2009), $^{60}$Co (Eberhardt et al., 1961; Spergel et al., 1986), $^{54}$Mn (Kohman & Bender, 1968), and $^{22}$Na (Bhandari et al., 1993), and assume the specimen is from the central part, the best agreement is obtained for radii of $<10$, $<20$, $<5$, and $5–8$ cm, respectively. Combining all results of these radionuclides, we infer a roughly spherical meteoroid with about 5–10 cm in radius. This estimate corresponds to an initial mass of $\sim$1.8–14.7 kg. The dynamic mass determined using the deceleration of the meteoroid in the atmosphere (2.2 kg; Table 2) favors the lower end of this estimate. Therefore, the true initial mass of the meteoroid was likely $\sim$2 kg. The small size of the progenitor body and minimal fragmentation observed indicates it was a single boulder with negligible macroscopic fractures.

For further information, please refer to the Meteoritical Bulletin Database (https://www.lpi.usra.edu/meteor/metbull.php?code=74013). A forthcoming paper will more thoroughly discuss the sample analysis.

CONCLUSIONS

Arpu Kuilpu is the fifth meteorite to be recovered by the DFN in Australia. The sample was found within close proximity to two of the three predicted fall lines, varying the shape and wind model. The predicted final mass (40 ± 10 g) is very close to the recovered stone’s mass (42 g). With an initial mass of $\sim$2 kg, the meteoroid is by far the smallest to ever survive entry and be recovered as a meteorite (the Cavezzo meteorite being second smallest at 3.5 kg [Gardiol et al., 2021]). The diminutive size of the progenitor body also demonstrates the robustness of our data reduction pipeline. Additionally, it supports the DFN’s original logic that desert regions are prime locations for fireball networks, where the likelihood of a successful meteorite recovery is greatest (Bland et al., 2012). Arpu Kuilpu is an H5 ordinary chondrite from the outer main belt (a $\sim$2.75 au). The orbital evolution over the previous 10 kyr is consistent with asteroidal debris, despite the proximity to a comet-like Tisserand’s parameter ($T_j = 2.97$). This is consistent with the results of Shoher et al. (2021), which demonstrated that nearly all sporadic comet-like debris seen by fireball networks are asteroidal in origin. Given the orbit, the most likely escape routes from the main belt are the 3:1 or 5:2 mean motion resonances (Granvik et al., 2018; Granvik & Brown, 2018).

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Data Availability Statement—Supplementary material has been uploaded as a Zenodo record at http://doi.org/10.5281/zenodo.2575775303. It contains all images (including calibration data), astrometry tables, the trajectory, wind profiles, as well as meteorite searching records.

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