Optimally coherent sets in geophysical flows: A new approach to delimiting the stratospheric polar vortex

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The “edge” of the Antarctic polar vortex is known to behave as a barrier to the meridional (poleward) transport of ozone during the austral winter. This chemical isolation of the polar vortex from the middle and low latitudes produces an ozone minimum in the vortex region, intensifying the ozone hole relative to that which would be produced by photochemical processes alone. Observational determination of the vortex edge remains an active field of research. In this letter, we obtain objective estimates of the structure of the polar vortex by introducing a new technique based on transfer operators that aims to find regions with minimal external transport. Applying this new technique to European Centre for Medium-Range Weather Forecasts (ECMWF) ERA-40 three-dimensional velocity data we produce an improved three-dimensional estimate of the vortex location in the upper stratosphere where the vortex is most pronounced. This novel computational approach has wide potential application in detecting and analysing mixing structures in a variety of atmospheric, oceanographic, and general fluid dynamical settings.

I. INTRODUCTION

Stratospheric ozone in the Southern Hemisphere high latitudes has decreased dramatically since the early 1970s. This long-term trend has been attributed to a combination of natural and anthropogenic factors\textsuperscript{[1–4]}. In particular, it has been discovered that the ozone depletion in the lower stratosphere of the Southern Hemisphere is particularly pronounced, due in part to a strong barrier to meridional transport between middle and high latitudes during the austral winter and early spring\textsuperscript{[1]}. Barriers such as these, which often coexist with turbulent mixing, play a major role in the dynamics of the stratosphere. The polar vortex is a known strong barrier to transport, enclosing a persistent, non-dispersive, coherent region over the high latitudes. Our aim in this letter is to precisely determine the spatial location and movements of this coherent region, improving significantly over existing methods of estimation. Our study focuses on the upper stratosphere where the polar vortex is best developed.

It is common meteorological practice to diagnose the polar vortex edge at the position of maximum meridional gradient of potential vorticity (PV). Potential vorticity is a quantity combining measures of circulation and stratification which is materially conserved for adiabatic, inviscid flow (both of which are good approximations in stratospheric flow over timescales of a week or two). It can be shown that strong PV gradients produce a “restoring force” inhibiting meridional motion of air parcels\textsuperscript{[5]}. Nevertheless, PV gradients alone provide only indirect measures of mixing barriers. In contrast, the present study characterises regions of minimal mixing directly in terms of the transport properties of the observed stratospheric flow. We present an innovative new mathematical technique to determine the polar vortex location at different times, directly as coherent structures in observed velocity fields. Lagrangian PV-based measures of the vortex such as those presented in\textsuperscript{[16]} are complicated by the fact that PV is generally a noisy field (as vorticity is the curl of the velocity field). The velocity field is generally much smoother; barriers to mixing estimated from the velocity field can be expected to be less sensitive to (poorly-observed) small-scale features of the flow.

Our new mathematical approach for detecting minimal transport structures with high accuracy has a broad range of potential applications to geophysical fluids. For example, transport properties in other long-lived atmospheric coherent structures such as blocking highs are of interest. There is also increasing interest in the transport properties of mesoscale (on the order of 10 to 100 km in diameter) ocean eddies and their influence on biological processes within the upper, sunlit part of the water column\textsuperscript{[6, 7]}.

II. INPUT DATA AND NON-AUTONOMOUS FLOW

Our input data consists of three-dimensional velocity fields obtained from the ECMWF ERA-40 data set\textsuperscript{[http://data.ecmwf.int/data/index.html]}. The data is on a three-dimensional grid with 2.5 degree resolution in the latitude and longitude direction (144 by 73 grid points over the Southern Hemisphere). Vertical coordinates are in units of hPa, with data provided at 5 pressure levels (3, 5, 7, 10, 20 hPa). We use 62 days of 6-hourly velocity fields from August 1 to September 31 in 1999. The velocity fields will be interpolated linearly in space and in time; thus we can only aim to detect features at the resolution of the data provided. While we recognize that there may be biases in the reanalysis data, particularly in the upper pressure levels near the model’s upper boundary, the purpose of this study is to demonstrate the
ability of the transfer operator approach to characterize coherent sets in highly unsteady flows. A climatology of the polar vortex would require a more careful consideration of the dataset under consideration.

Our interest is focused on the Lagrangian dynamics in the higher latitudes of Southern Hemisphere. Therefore, we will work on the phase space $X = S^1 \times [-\theta^\circ, -30^\circ] \times D$, where $S^1$ is a circle parameterized from $0^\circ$ to $360^\circ$ and $D = [3, 20]$ denotes the range of pressure in hPa. The Lagrangian motion of passive tracers is represented by their trajectories $x(t) := \Phi(x, t; \tau)$, where the flow map $\Phi : X \times \mathbb{R} \to X$ is a function of time $\tau$ and gives the terminal point $\Phi(x, t; \tau)$ in $X$ of a particle initially located at $x \in X$ at time $t$ and flowing for $\tau$ time units. The flow map $\Phi(x_0, t_0; \tau)$ is obtained as a solution of the nonautonomous ODE $\frac{dx}{dt} = f(x(t), t)$ with initial condition $x(t_0) = \Phi(x_0, t_0; 0)$, where $f(x(t), t)$ in this report is the prescribed velocity data.

III. ALMOST-IN Variant SETS, COHERENT PAIRS, AND TRANSFER OPERATORS

We shall be interested in finding a pair of sets $A_t, A_{t+\tau}$ at times $t$ and $t + \tau$ so that $\Phi(A_{t+\tau}, t + \tau; -\tau) \approx A_t$. Moreover, this pair of sets should retain this property even when some diffusion is added to the system. Let $\mu$ be a probability measure that is preserved by the flow at all times. We call $A_t, A_{t+\tau}$ a $(\rho_0, t, \tau)-$coherent pair if

$$\rho_\mu(A_t, A_{t+\tau}) := \mu(A_t \cap \Phi(A_{t+\tau}, t + \tau; -\tau))/\mu(A_t) \geq \rho_0,$$

and $\mu(A_t) = \mu(A_{t+\tau})$. The requirement that diffusion is crucial. Clearly, there are many $(\rho_0, t, \tau)-$coherent pairs according to the above definition without diffusion. One may simply select any set $A_t \subset X$ and define $A_{t+\tau} = \Phi(A_t, t; \tau)$ to produce a $(1, t, \tau)$-coherent pair. In chaotic systems, such an image set $A_{t+\tau}$ is likely to be significantly less regular than $A_t$ because of stretching and folding. We are seeking $(\rho_0, t, \tau)$-coherent pairs with both sets regular. The requirement that hold even under diffusion acts as a selection principle, removing irregular sets, and selecting pairs that are robust to perturbation. At a certain level of diffusion, we may ask to find a coherent pair that maximises $\rho_0$, and may expect a unique such pair.

To identify sets satisfying, we use a transfer operator $P_{t, \tau} : L^1(X, m) \circ \Phi$ defined by

$$P_{t, \tau} g(x) := \frac{g(\Phi(x, t + \tau; -\tau))}{|\det D\Phi(\Phi(x, t + \tau; -\tau), t; \tau)|}$$

where $m$ is the normalized Lebesgue measure on $X$. In particular, if $g(x)$ is a density of passive tracers at time $t$, $P_{t, \tau} g(x)$ provides their density at time $t + \tau$ induced by the flow $\Phi$.

In the autonomous setting, eigenfunctions $f$ of $P_{t, \tau}$ ($= P_{t}$ for all $t$) corresponding to positive eigenvalues $\Lambda \approx 1$ were used to find almost-invariant sets [8,11]. The key point of difference between these prior studies and the present work is that the sets studied previously do not move significantly over the time duration studied, while our present work seeks highly mobile coherent sets, which are far from being almost-invariant. The new theory and numerics introduced in the next section are specifically designed for nonautonomous or time-dependent systems.

IV. NUMERICAL APPROACH

We partition $X$ into a grid of boxes $\{B_1, \ldots, B_n\}$. The pressure extents of the boxes are either 3-5, 5-7, 7-10, or 10-20 hPa. Each pressure layer consists of approximately equal cross-sectional area in the latitude/longitude directions, leading to $n = 26420$ boxes in total. To numerically approximate the transfer operator $P_{t, \tau}$, we construct a finite-dimensional approximation based on Ulam’s approach [12]:

$$P^{(\tau)}(t)_{i,j} = \frac{m(B_i \cap \Phi(B_j, t + \tau; -\tau))}{m(B_i)},$$

where $m$ is a normalised volume measure in (lat, lon, pressure) coordinates. The entry $P^{(\tau)}(t)_{i,j}$ may be interpreted as the probability that a point selected uniformly at random in $B_i$ at time $t$ will be in $B_j$ at time $t + \tau$. The discretisation naturally produces a diffusion at the level of box diameters. As our boxes are of approximately the same dimensions as the distances between neighbouring ERA-40 data points, it is unnecessary to impose additional diffusion. If our boxes were significantly smaller the distances between neighbouring data points, it is possible that spurious fine features below the resolution supported by the data could appear; in such a situation, additional diffusion would be required to remove spurious fine features.

We estimate $P^{(\tau)}(t)_{i,j}$ by

$$P^{(\tau)}(t)_{i,j} \approx \#(\ell : y_{i,\ell} \in B_i, \Phi(y_{i,\ell}, t; \tau) \in B_j)/Q,$$

where $y_{i,\ell}, \ell = 1, \ldots, Q$ are uniformly distributed test points in $B_i$ and $\Phi(y_{i,\ell}, t; \tau)$ is obtained via a numerical integration. We set $Q = 147$ in our experiments and calculate $\Phi(y_{i,\ell}, t; \tau)$ using the standard Runge-Kutta method with step size of 3/4 hours. Linear interpolation is used to evaluate the velocity vector of a tracer lying between the data grid points in the longitude-latitude-pressure coordinate. In the temporal direction the data is affinely interpolated independently in the longitude, latitude and pressure level directions. The step size of 3/4 hours is small enough to guarantee that a tracer will usually not flow to a neighboring data grid set; this limits the numerical integration error.

We assume that the mass density of particles in $X$ is at equilibrium and denote the fractional mass of particles contained in $B_i$ by $p_i$. Thus $\sum_{i=1}^n p_i = 1$. We construct a reverse time transition matrix from time $t + \tau$ to $t$ denoted $P^{(\tau)}(t)$ as $P^{(\tau)}(t)_{i,j} = P^{(\tau)}(t)_{j,i} p_j/p_i$. 

Introduce a weighted inner product \( \langle x, y \rangle_p := \sum_{i=1}^{n} x_i^p y_i^p \). One has \( \langle x \mathbf{P}^{(\tau)}(t), y \rangle_p = \langle x, y \mathbf{P}^{(\tau)}(t) \rangle_p \) for all \( x, y \in \mathbb{R}^n \).

Our new approach to finding a coherent pair is intuitively based upon seeking a solution to

\[
\max_{w \in \{\pm 1\}^n} \frac{\langle w \mathbf{P}^{(\tau)}(t), w \mathbf{P}^{(\tau)}(t) \rangle_p}{\langle w, w \rangle_p}. \tag{5}
\]

We think of \( A_t := \cup_{i:t_i=1} B_t \) and \( A_t^c := \cup_{i:t_i=-1} B_t \) as a coherent partition of \( X \). The numerator in (5) represents the size of the forward image of the vector \( w \). If there is little transport from \( A_t \) to \( \Phi(A_t^c, t; \tau) \) and from \( A_t^c \) to \( \Phi(A_t, t; \tau) \) (so \( A_t, \Phi(A_t, t; \tau) \) and \( A_t^c, \Phi(A_t^c, t; \tau) \) are both coherent pairs), this numerator will be large. To produce non-trivial partitions, we may need to place lower bounds on the masses of both \( A_t \) and \( A_t^c \). Such a balanced bisection problem is combinatorially hard to solve. Therefore we remove the discrete condition \( \rho_n \). The maximizing \( y \) Denote the corresponding left singular vector by \( \tau \mathbf{y} \).

Let \( \mathbf{D}_{ij} = \delta_{ij} p_i \) and noting \( \langle w, v \rangle_p = \langle w \mathbf{D}^{1/2}, v \mathbf{D}^{1/2} \rangle \), this is easily solved by computing the second largest singular value \( s \) of \( D^{-1/2} \mathbf{P}^{(\tau)}(t) D^{1/2} \).

Denote the corresponding left singular vector by \( \mathbf{y} \) (under multiplication on the right). The maximizing \( w = w(t) \) is constructed as \( w(t) = y D^{-1/2} \). The minimizing \( v \) turns out to be \( v D^{-1/2} \) where \( u \) is the leading left singular vector of \( D^{-1/2} \mathbf{P}^{(\tau)}(t) D^{1/2} \). We also construct \( z \) as the corresponding right singular vector and set \( w(t+\tau) = z D^{-1/2} \). We assume that \( w(t), w(t+\tau) \) are normalised so that \( \langle w(t), w(t) \rangle_p = 1 \) and \( \langle w(t+\tau), w(t+\tau) \rangle_p = 1 \).

One now has:

1. \( w(t) \mathbf{P}^{(\tau)}(t) = s w(t+\tau) \),
2. \( w(t+\tau) \mathbf{P}^{(\tau)}(t) = s w(t) \),
3. \( \langle w(t) \mathbf{P}^{(\tau)}(t), w(t) \mathbf{P}^{(\tau)}(t) \rangle_p = s^2 \).

Choosing \( v \) via the minimization in (5) ensures that \( w \) has the transformation properties 1. and 2. above, which are crucial to the definition of coherent sets.

We now extract a coherent pair \( A_t \) and \( A_{t+\tau} \) from a pair of vectors \( w(t) \) and \( w(t+\tau) \). We create sets that are unions of boxes with \( w \)-values above certain thresholds. Define \( \bar{A}_t^+((c) := \cup_{t: w(t) > c} B_t \) and \( \bar{A}_{t+\tau}^+(c) := \cup_{t: w(t+\tau) > c} B_t \), \( c \in \mathbb{R} \). Denote \( \mu_n(\bar{A}_t^+(c)) = \sum_{t: w(t) > c} p_i \) and \( \rho_n(\bar{A}_{t+\tau}^+(c)) = \sum_{t: w(t+\tau) > c} p_i \).

For \( A_t = \bigcup_{i \in I_t} B_i \) and \( A_{t+\tau} = \bigcup_{i \in I_{t+\tau}} B_i \), define

\[
\rho_n(A_t, A_{t+\tau}) = \sum_{i \in I_t, j \in I_{t+\tau}} p_i \mathbf{P}^{(\tau)}(t)_{ij} / \sum_{i \in I_t} p_i.
\]

The quantity \( \rho_n \) measures the discretised coherence for the pair \( A_t, A_{t+\tau} \). Our procedure is summarised below:

1. Let \( \tau(c) = \arg \min_{c \in \mathbb{R}} |\mu_n(\bar{A}_t^+(c)) - \rho_n(\bar{A}_{t+\tau}^+(\tau(c)))| \). This is to enforce \( \mu_n(A_t) = \mu_n(A_{t+\tau}) \).
2. Set \( c^* = \arg \max_{c \in \mathbb{R}} \rho_n(A_t^+(c), A_{t+\tau}^+(\tau(c))) \). The value of \( c^* \) is selected to maximize the coherence.
3. Define \( A_t := \bar{A}_t^+(c^*) \) and \( A_{t+\tau} := \bar{A}_{t+\tau}^+(\tau(c^*)) \).

We remark that one has to ensure that the sign of \( w(t) \) and \( w(t+\tau) \) manifest the same “parity”, i.e., the salient features of \( w(t) \) and \( w(t+\tau) \) to be extracted must have the same sign. It may thus be necessary to multiply one of \( w(t) \) or \( w(t+\tau) \) by \(-1\).

The major computational cost is the construction of \( \mathbf{P}^{(\tau)} \). The calculation of large singular values and corresponding singular vectors is relatively quick, as \( \mathbf{P}^{(\tau)} \) is very sparse and iterative methods for sparse matrices may be used. The construction of \( \mathbf{P}^{(\tau)} \) requires numerical integration of \( Q \cdot n \) trajectories for a flow duration of \( \tau \) time units. The trajectory computations are of course highly parallelisable, and further computational savings might be made by reusing already computed trajectory segments to link with new trajectories when the latter pass nearby.

**V. NUMERICAL RESULTS**

We computed the SVD of \( D^{-1/2} \mathbf{P}^{(t)}(\tau) D^{1/2} \) at \( t = 14 \) with \( \tau = 14 \) days to obtain the left (resp. right) singular vectors \( y \) (resp. \( z \)) and hence \( w(14) \) and \( w'(28) \). Figure 3 illustrates the vectors \( w(14) \) and \( w'(28) \) with the components monotonically rescaled to uniformly distributed values between 0 and 1. The highlighted part of these vectors describes the most coherent pair of sets. We now threshold \( w(14) \) and \( w'(28) \) using the algorithm described above to extract the corresponding pair of coherent sets; see Figure 4. We find the optimal coherence ratio is \( \rho_n(A_{14}, A_{28}) \approx 0.7902 \); this means that about 21% of the mass in \( A_{14} \) on August 14, 1999 falls outside \( A_{28} \) on August 28, 1999.

Interestingly, our coherent pair has a “hole” over the south pole. Further calculations have revealed that the reason for this is that around twice as many particles in this vertical hole on August 14 have exited the slice 3-20 hPa vertically by August 28 when compared to similar vertical exits of particles starting in the identified coherent set on August 14. Thus, this inner part of the vortex is less coherent and excluded from our coherent pair. This hole may be an artifact of the reanalysis data, although it is consistent with evidence of very strong polar descent in this region.
We now compare our coherent pair of sets to sets defined by contours of potential vorticity (PV). A common approach, developed in [14, 15] is to define the vortex boundary as the isoline of the largest gradient of PV w.r.t. the equivalent latitude. We employ this approach to define potential coherent pairs at $t = 14$ and $t = 28$. We additionally enforce the constraint that the mass enclosed by a PV isocontour at $t = 28$ is approximately equal to the mass of the set enclosed by the determined PV isocontour at $t = 14$. The computational cost of the PV approach is NARATIP, PLEASE ADD MATERIAL ON COMPUTATIONAL COST.

The two-dimensional plots of PV-determined coherent pairs at $t = 14$ and $t = 28$ are compared with the coherent sets in Figure 2. To estimate the transport of particles from the inside the set at $t = 14$ to outside the set at $t = 28$, we use a method similar to the contour crossing method introduced in [16]. The tracer particle is considered to be outside the boundary if its potential vorticity is larger than that of the boundary. Note that the contour crossing method is originally developed to estimate the transport on the 2D isentropic surface but we would like to extend its utility to estimate the transport across the boundary surface. Therefore, we interpolate the PV at the final time ($t = 28$) to obtain the PV at the particle’s final position. We also interpolate the PV of the boundaries of the set at $t = 28$ along the pressure co-ordinate to determine the boundary at the pressure level the advected particle resides in. This calculation shows that the fraction of particles initially inside the surface at $t = 14$ remains inside the boundary surface at $t = 28$ is approximately 0.7204. Thus our transfer operator based approach yields coherent pairs with 9.69% greater coherence. Moreover, our approach is able to detect finer structures, including a vertical hole near the south pole.

VI. CONCLUSIONS

The Antarctic polar vortex is a well-known feature of the austral wintertime stratosphere separating polar and midlatitude air masses. The strong barrier to transport at the vortex edge plays an important role in ozone dynamics, particularly the development of the Southern Hemisphere ozone hole in austral spring. Diagnosis of the vortex edge from observations is a challenging problem that remains a subject of active research.

Previous approaches to this problem have been based on kinematic (following the advection of some tracer) or dynamic (considering gradients of PV) arguments. We presented a new kinematic method of accurately estimating the three-dimensional location of the vortex. This new method uses the velocity field to diagnose “optimally coherent pairs” and was able to determine a significantly more accurate estimate of transport barriers, with almost 10% less external transport from the identified vortex region than the PV-based estimate. Future, more detailed studies will include an investigation of the climatology of the polar vortex on isentropic surfaces throughout the stratosphere.

Our new computational approach for detecting minimal transport structures with high accuracy has a broad range of potential application to studies of transport and mixing in the atmosphere and ocean, and in general fluid dynamics settings.

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FIG. 2: Comparison between coherent pair and PV surface boundary. The optimal coherent set $A_{14}$ and $A_{28}$ obtained from thresholding the vectors $w(14)$ and $w'(28)$. The coherent ratio $\rho(A_{14}, A_{28}) \approx 0.7902$. The blue curve in each plot shows the PV surface boundary obtained from the maximum PV gradient w.r.t equivalent latitude.

[1] M. E. McIntyre, J. Atmos. Terr. Phys. 51, 29-43 (1989).
[2] C. R. Webster et al., Science 261, 1130-1134 (1993).
[3] T. G. Shepherd, J. Meteorol. Soc. Jpn. 85B, 165-191 (2007).
[4] T. G. Shepherd, Atmosphere-Ocean 46, 117-138 (2008).
[5] G. K. Vallis, Atmospheric and Oceanic Fluid Dynamics (Cambridge University Press, 2006).
[6] W. R. Crawford et al., Prog. Oceanogr. 75, 287-303 (2007).
[7] F. Nencioli, et al., Deep-Sea Res. II 55, 1252-1274 (2008).
[8] M. Dellnitz and O. Junge, SIAM J. Numer. Anal. 36, 491 (1999).
[9] G. Froyland and M. Dellnitz, SIAM J. Sci. Comput. 24, 1839 (2003).
[10] G. Froyland, Physica D 200, 205 (2005).
[11] G. Froyland, K. Padberg, M. H. England, and A. M. Treguier, Phys. Rev. Lett. 98(22):224503, (2007).
[12] S. Ulam, A collection of mathematical problems (Interscience Publishers, 1960).
[13] G.L. Manney and R.W. Zureck, and A. O’Neill, and R. Swinbank. J. Atmos. Sci. 51, 2973 (1994).
[14] M. E. McIntyre and T. N. Palmer, J. Atmos. Terr. Phys. 46, 825-849 (1983).
[15] E. R. Nash et al., J. Geophys. Res. Phys. 101, 9471-9476 (1996).
[16] A. H. Sobel et al., J. Atmos. Sci. 54, 2241-2260 (1997).
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