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Oxygen Vacancy Substitution Linked to Ferric Iron in Bridgmanite at 27 GPa

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Abstract Ferric iron can be incorporated into the crystal structure of bridgmanite by either oxygen vacancy substitution (MgFeO2.5 component) or charge-coupled substitution (FeFeO3 component) mechanisms. We investigated the concentrations of MgFeO2.5 and FeFeO3 in bridgmanite in the MgO-SiO2-Fe2O3 system at 27 GPa and 1700–2300 K using a multianvil apparatus. The FeFeO3 content increases from 1.6 to 7.6 mol.% and from 5.7 to 17.9 mol.% with and without coexistence of (Mg,Fe)O, respectively, with increasing temperature from 1700 to 2300 K. In contrast, the MgFeO2.5 content does not show clear temperature dependence, that is, ~2–3 and < 2 mol.% with and without the coexistence of (Mg,Fe)O, respectively. Therefore, the presence of (Mg,Fe)O enhances the oxygen vacancy substitution for Fe3+ in bridgmanite. It is predicted that Fe3+ is predominantly substituted following the oxygen vacancy mechanism in (Mg,Fe)O-saturated Al-free bridgmanite when Fe3+ is below ~0.025 pfu, whereas the charge-coupled mechanism occurs when Fe3+ > 0.025 pfu.

Plain Language Summary Bridgmanite, the most abundant mineral of the Earth’s lower mantle, can contain Fe3+ although the valence of iron is 2+ in general. An important question is how Fe3+ is substituted in the crystal structure of bridgmanite. It may form the MgFeO2.5 component, in which oxygen vacancies are partly missing. Or it may form the FeFeO3 component, which has no missing cations or anions. Since bridgmanite is present in the lower mantle together with (Mg,Fe)O, we investigated the MgFeO2.5 and FeFeO3 contents in Al-free bridgmanite that coexists with and without (Mg,Fe)O under the topmost lower mantle conditions. The results show that the presence of (Mg,Fe)O enhances the formation of MgFeO2.5. The solubility of MgFeO2.5 component is about 2.5 mol.% in bridgmanite that coexists with (Mg,Fe)O, whereas it is nearly zero when (Mg,Fe)O is absent.

1. Introduction

Bridgmanite, with typical chemical formula (Mg,Fe)SiO3, comprises around 80 vol.% of the lower mantle (e.g., Frost, 2008; Ringwood, 1991; Tschauner et al., 2014), and therefore, it dominates the physical and chemical processes of the lower mantle. Due to its variable valence, iron is the most important element that affects the properties of bridgmanite and thus mantle dynamics (e.g., Ismailova et al., 2016). Generally, iron is incorporated into the crystal structure of bridgmanite in the Mg site (A site) with the ferrous valence state. However, it is also known that a large amount of iron in bridgmanite can be in the ferric valence state even under reduced lower mantle conditions (Frost et al., 2004; Frost & McCammon, 2008; Grocholski et al., 2009; Jackson et al., 2005; Lauterbach et al., 2000; McCammon, 1997; McCammon et al., 2004). By analogy with aluminum substitutions in bridgmanite through the formation of MgAlO2.5 and AlAI03 components (Brodoht, 2000; Kojitani et al., 2007; Liu, Ishii, & Katsura, 2017; Liu, Nishi, et al., 2017; Liu, Akaogi, & Katsura, 2019; Liu, Bofa-Ballaran, et al., 2019; Navrotsky, 1999), it is expected that the MgFeO2.5 (oxygen vacancy mechanism) and FeFeO3 (charge-coupled mechanism) components, respectively, will be formed by Fe3+ incorporation through the following reactions (e.g., Hummer & Fei, 2012; Navrotsky, 1999):
Table 1: Run Conditions, Phase Assemblage, and Composition of Bridgmanite in Run Products

| Composition of bridgmanite | Total | SiO2 | Fe2O3 | MgFeO2.5 |
|----------------------------|-------|------|-------|---------|
| MgO                        | 29.72 | 43.57 | 26.67 | 99.96   |
| SiO2                       | 74.23 | 46.78 | 23.40 | 101.63  |
| Fe2O3                      | 94.85 | 31.45 | 23.40 | 101.63  |
| MgO-rich                    | 2000  | 31.45 | 23.40 | 101.63  |

Note: All experiments were performed at 27 GPa. T: temperature; t: annealing duration; N: number of analyzed points by EPMA. aMgO-rich: mixture of oxides with bulk composition of MgSiO3 + MgFe2O4 + MgO (60:20:20 atomic ratio). bFe2O3-rich: mixture of oxides with bulk composition of MgSiO3 + MgFe2O4 + Fe2O3 (60:20:20 in mole proportion, respectively). cMgAlO2.5 decreases with decreasing Mg/Si ratio in the MgO-Al2O3 system because of the reaction 2MgO + Al2O3 = 2MgAlO2.5. Therefore, the phase relations of bridgmanite with MgFeO2.5 and FeFeO3 components, characterized by B site Fe3+, may depend on whether periclase (or ferropericlase) coexists or not according to the reaction 2MgO + Fe2O3 = 2MgFeO2.5. To clarify this further, we investigated the defect chemistry of bridgmanite in assemblage with a separate MgFe2O4 phase and either presence or absence of a (Mg,Fe)-rich phase, using a multianvil apparatus.

2. Experimental and Analytical Methods

2.1. Starting Material and High-Pressure Experiments

MgO, SiO2, and Fe2O3 (purity > 99.9%, natural isotopic ratios) from Sigma-Aldrich Corporation were used as starting materials. MgO and SiO2 were dried at 1273 K, whereas Fe2O3 was dried at 773 K prior to weighing. Fine-mixed powders with bulk compositions of MgSiO3 + MgFe2O4 + MgO (60:20:20 in mole proportion, hereafter MgO-rich sample) and MgSiO3 + MgFe2O4 + Fe2O3 (60:20:20 in mole proportion, hereafter Fe2O3-rich sample) were obtained by repeatedly grinding in ethanol for homogeneity.

After drying in a vacuum furnace at 420 K, the two mixtures (both MgO-rich and Fe2O3-rich separated by a piece of Pt foil) were filled into Pt tube capsules (OD = 1.0 mm, ID = 0.8 mm). Each capsule was loaded into an MgO sleeve within carbide anvils with edge length of 26 mm and truncation edge length of 3.0 mm using the multianvil press at Bayerisches Geoinstitut, IRIS-15 (Ishii et al., 2016). After reaching the target pressure of 27 GPa, the assembly was heated to a target temperature (1700–2300 K as listed in Table 1) at a ramping rate of 100 K/min measured by a W/Re (D-type) thermocouple. The annealing durations were 9–40 hr.
Table 1, which is sufficiently long to achieve chemical equilibrium (Frost & Langenhorst, 2002) as confirmed by the run products. The samples were quenched to room temperature by switching off the heating power and decompressed to ambient conditions over a duration exceeding 15 hr. The recovered run products were mounted in epoxy resin and their cross sections were prepared by polishing with emery papers and diamond pastes.

2.2. Sample Characterization

Scanning electron microscope (SEM). SEM analysis was performed on each sample, using a backscattering detector and an acceleration voltage of 20 kV associated with an energy dispersive detector. Run products appeared as coexisting bridgmanite (gray) + MgFe₂O₄ (bright) + (Mg,Fe)O (dark) phases and bridgmanite (gray) + MgFe₂O₄ (bright) phases, respectively (Table 1), as observed in backscattering detector images (Figure 1) and SEM-EDS point analysis. No unreacted particles (SiO₂ or Fe₂O₃ grains) were found in any of the sample capsules.

Microfocus X-ray diffraction (XRD). XRD analysis was performed on the recovered samples using a micro focused X-ray diffractometer (Brucker AXS D8 Discover) equipped with a two-dimensional solid-state

Figure 1. Backscattered SEM images of run products at 27 GPa and 1700–2300 K for both (a, c, and e) MgO-rich and (b, d, and f) Fe₂O₃-rich samples, which appear as coexisting bridgmanite (gray) + MgFe₂O₄ (bright) + (Mg,Fe)O (dark) phases and bridgmanite (gray) + MgFe₂O₄ (bright) phases, respectively.
detector and a microfocus source of Co-Kα radiation operated at 40 kV and 500 μA. The exposure time is about 4 hr for each sample with beam size of ~100 μm. Bridgmanite and ferropericlase phases were clearly identified, and additional peaks were characterized as a CaMn2O4-structured MgFe2O4 phase. This is consistent with the observation by Andrault and Bolfan-Casanova (2001), although a CaTi2O4-type structure was proposed by a more recent study (Greenberg et al., 2017). Pt peaks from sample capsules also appeared in the sample diffraction patterns due to the limited spatial resolution of the diffractometer (Figure 2).

Mössbauer Spectroscopy. Mössbauer spectra were collected over 500 μm diameter spots on samples of 120 μm thickness at room temperature in transmission mode on a constant acceleration Mössbauer spectrometer with a nominal 370 MBq 57Co high specific activity (point) source in a 12 μm Rh matrix. The velocity scale was calibrated relative to 25 μm thick α-Fe foil using the positions certified for standard reference material no. 1541 of the (former) National Bureau of Standards. Line widths of 0.36 mm/s for the outer lines of α-Fe were obtained at room temperature. Measurement times for each spectrum varied from 4 to 8 hr. The effective Mössbauer thickness of samples varied between 5 and 10 mg Fe/cm². Spectra were fitted by the program MossA (Prescher et al., 2012) using the full transmission integral to multiple doublets with pseudo-Voigt line shape to account for next nearest neighbor effects. There is no detectable Fe²⁺ in bridgmanite from the run products within experimental uncertainty (Figure 3), and we assume that Fe³⁺/ΣFe ≈ 100%.

Electron probe microanalysis (EPMA). Concentrations of major elements (Mg, Si, Fe) in run products were obtained using a JEOL JXA-8200 electron microprobe with a wavelength-dispersive spectrometer operated with an acceleration voltage of 15 kV, a beam current of 5 nA, and a counting time of 20 s. Focused point analysis with an excitation region of ~1 μm was applied. An enstatite single crystal was used as the standard for Mg and Si, metallic Fe was used for Fe, whereas O was calculated by stoichiometry. Since the Pt capsule may absorb Fe from the samples, grains near the capsule wall (within a few microns) with slightly lower Fe% were avoided. The compositions of bridgmanite are listed in Table 1, whereas (Mg,Fe)O and MgFe2O₄ phases, which have Fe/Mg ratios of 0.02–0.07 and 1.94–2.38, respectively, are given in the supporting information.

Molar concentrations of MgSiO₃, MgFeO₂·₅, and FeFeO₃ components in bridgmanite were calculated from

![Figure 2. Representative X-ray diffraction spectra of run products (Run I567). (top) Fe₂O₃-rich sample. (bottom) MgO-rich sample. Bridgmanite (bdg), CaMn₂O₄-type MgFe₂O₄ (CM), and (Mg,Fe)O-periclase (pc) phases were identified, as well as the Pt from the sample capsules.](image-url)
\[
\text{Mg}_x\text{Fe}_y\text{Si}_z\text{O}_{x+y+z} = z\text{MgSiO}_3 + (x-z)\text{MgFeO}_{2.5} + \frac{y-x+z}{2} \text{FeFeO}_3
\]

where \(x, y, \text{ and } z\) are the atomic numbers per formula unit (pfu) from EPMA analysis. The \(\text{Fe}^{2+}\text{SiO}_3\) component was excluded since \(\text{Fe}^{3+}/\Sigma\text{Fe} \approx 100\%\) based on Mössbauer spectroscopy analysis, which is also confirmed by the low Fe/Mg ratio in ferropericlase.

### 3. Results and Discussion

#### 3.1. Total \(\text{Fe}^{3+}\) Concentration in Bridgmanite

The concentration of \(\text{Fe}^{3+}\) in bridgmanite systematically increases from 0.06 to 0.17 pfu and from 0.12 to 0.37 pfu in MgO-rich and Fe\(_2\text{O}_3\)-rich samples, respectively, with increasing temperature from 1700 to 2300 K (Figure 4a). It is noted that only two phases, bridgmanite and MgFe\(_2\text{O}_4\), coexist in run products of Fe\(_2\text{O}_3\)-rich samples (Table 1). According to the phase rule, there is one more degree of freedom in addition to pressure and temperature in the MgO-SiO\(_2\)-Fe\(_2\text{O}_3\) three-component system, and as a result, the \(\text{Fe}^{3+}\) concentration in bridgmanite should be correlated with the bulk Fe\(_2\text{O}_3\) content in the starting material.

In contrast, three phases, bridgmanite, (Mg, Fe)O, and MgFe\(_2\text{O}_4\), coexist in the MgO-rich samples (Table 1). With fixed pressure and temperature, the composition of each phase should also be fixed and independent of the bulk composition of the starting material in the three-component system. Therefore, the \(\text{Fe}^{3+}\) contents in the MgO-rich samples should represent the solubility of \(\text{Fe}^{3+}\) in bridgmanite with saturation of (Mg,Fe)O under the given pressure and temperature conditions, that is, from 0.06 to 0.17 pfu with temperatures from 1700 to 2300 K (Figure 4a).

#### 3.2. MgFe\(_{2.5}\) and FeFe\(_3\) Concentrations in Bridgmanite by EPMA

The molar fractions of the bridgmanite components, MgSiO\(_3\), MgFe\(_{2.5}\), and FeFe\(_3\) were calculated sequentially based on Si, remaining Mg, and remaining Fe, respectively (Table 1). Figure 4b illustrates the proportions of ferric iron components versus temperature and shows that the FeFe\(_3\) component increases systematically for both MgO-rich and Fe\(_2\text{O}_3\)-rich samples with increasing temperature from 1700 to 2300 K (Figure 4a).

The concentration of bridgmanite component MgFe\(_{2.5}\) is well below 2 mol.% in Fe\(_2\text{O}_3\)-rich samples although the uncertainty is relatively large. MgO-rich samples have a MgFe\(_{2.5}\) concentration of 2 to 3 mol.% at 1700-2300 K (Figure 4b). Therefore, the presence of MgO enhances the formation of the MgFe\(_{2.5}\) component in bridgmanite, but only slightly in the presence of MgFe\(_2\text{O}_4\). A similar relationship was also observed in the Fe-free and Al-bearing bridgmanite coexisting with MgO and MgAl\(_2\text{O}_4\) (Liu, Boffa-Ballaran, et al., 2019). In contrast to the temperature-induced increase in the FeFe\(_3\) component, the concentration of the MgFe\(_{2.5}\) component is insensitive to temperature. This could be because either the temperature dependence is hidden by the uncertainty of data points, or because the MgFe\(_{2.5}\) reaches the maximum solubility even at 1700 K. Thus, the increase of total \(\text{Fe}^{3+}\) with temperature is dominated by increasing FeFe\(_3\) component, rather than increasing MgFe\(_{2.5}\) component.
3.3. Fe$^{3+}$ Substitution Mechanisms in Bridgmanite

It has been previously proposed that oxygen vacancy substitution by Fe$^{3+}$ operates in bridgmanite in addition to the charge-coupled substitution mechanism (e.g., Frost & Langenhorst, 2002; Frost & McCammon, 2008; Walter et al., 2004); however, the concentration and formation condition of the MgFeO$_{2.5}$ component have not been well constrained. Andrault and Bolfan-Casanova (2001) and Catalli et al. (2010) found roughly equal distribution of Fe$^{3+}$ on Mg and Si sites; nevertheless, sample compositions in these diamond anvil experiments were uncertain due to small sample size. Hummer and Fei (2012) investigated Fe$^{3+}$ substitution mechanisms based on well-constrained chemical compositions of bridgmanite samples recovered from large-volume multianvil experiments. They concluded that Fe$^{3+}$ substitutes into bridgmanite by a combination of oxygen vacancy (equation 1) and charge-coupled (equation 2) mechanisms when Fe content is low (<0.05 pfu) because Mg content is either higher or

![Figure 4](image-url). Atomic concentrations of MgFeO$_{2.5}$, FeFeO$_3$, and total Fe$^{3+}$ per formula unit (oxygen = 3) in bridgmanite, assuming MgFeO$_{2.5}$ + FeFeO$_3$ + MgSiO$_3$ = 100%. (upper panel) Total Fe$^{3+}$ concentration. (lower panel) Proportions of the MgFeO$_{2.5}$ and FeFeO$_3$ components.
comparable with Si. When Fe content is higher (0.074 pfu), charge-coupled (equation 2) and Mg vacancy mechanisms ($3\text{Mg}^\times \rightarrow 2\text{Fe}^\times + \text{V}_{A}^\times$ with Fe$^{3+}$SiO$_3$ component) occur because Si > Mg. However, the relative magnitudes of Mg and Si that they obtained show no systematic variation with Fe content (Figure 5 in this study and Figure 3 in Hummer & Fei, 2012), and therefore are insufficient for interpretation of Fe$^{3+}$ substitution mechanisms.

In our MgO-rich samples, both Mg and Si contents deviate slightly from theoretical calculations of a pure charge-coupled substitution mechanism (Figure 5). The Mg content is systematically higher than Si within the investigated Fe content range from 0.06 to 0.17 pfu, indicating an oxygen vacancy substitution with formation of a MgFeO$_{2.5}$ component. On the other hand, Fe$_2$O$_3$-rich samples have Mg and Si contents that match a pure charge-coupled mechanism without detectable oxygen vacancy substitution (Figure 5). Therefore, the Fe$^{3+}$ substitution mechanism in bridgmanite is dominantly controlled by the saturation condition of MgO, rather than by Fe content. The Mg vacancy substitution mechanism proposed by Hummer and Fei (2012) is not observed in this study, although the exact reason for this discrepancy is unclear. One possible reason is that ferric iron in their iron-rich sample (Fe = 0.074 pfu) might be partially reduced, leading to the formation of a Fe$^{2+}$SiO$_3$ component.

In Figure 5, the lines of decreasing Mg and Si with increasing Fe$^{3+}$ in bridgmanite from MgO-rich samples intersect the corresponding Si- and Mg-vectors for pure MgFeO$_{2.5}$ substitution at about 0.025 pfu, suggesting that oxygen vacancy substitution is preferred relative to charge-coupled substitution for low Fe$^{3+}$. When Fe$^{3+} \leq 0.025$ pfu, the majority of Fe$^{3+}$ should follow the oxygen vacancy substitution mechanism. With increasing Fe$^{3+}$, MgFeO$_{2.5}$ content increases in this compositional range. When the MgFeO$_{2.5}$ content reaches the solubility limit of 0.025 pfu, additional Fe$^{3+}$ will follow the charge-coupled mechanism and thus the amount of MgFeO$_{2.5}$ remains constant, which appears as nearly parallel curves of Mg or Si contents and pure charge-coupled substitution in Figure 5.
3.4. Implications for Bridgmanite Chemistry

The Earth’s lower mantle is mainly composed of bridgmanite, ferropericlase, and CaSiO₃ perovskite (Frost, 2008; Ringwood, 1991; Tschauner et al., 2014), meaning that bridgmanite is under MgO-saturated conditions. Since a significant amount of iron in bridgmanite is ferric (Frost et al., 2004; Grocholski et al., 2009; Jackson et al., 2005; Lauterbach et al., 2000; McCammon, 1997; McCammon et al., 2004), the concentration of the MgFeO₂.₅ component by oxygen vacancy substitution could be significant. On the other hand, bridgmanite in the lower mantle contains some Al (e.g., Irifune et al., 2010), which could affect the Fe³⁺ substitution mechanism via the formation of the Fe₅O₉ component with Fe³⁺ and Al³⁺ occupying A and B sites, respectively (e.g., Richmond & Brodholt, 1998; Zhang & Oganov, 2006 as reviewed in Frost & McCammon, 2008). Similarly, Fe³⁺ may also affect formation of the MgAlO₂.₅ component. The Al/Fe³⁺ ratio in Al-bearing bridgmanite under lower mantle conditions is generally larger than 1.0 in both peridotitic and basaltic lithologies (e.g., Mohn & Trønnes, 2016; Nakajima et al., 2012; Prescher et al., 2014), the FeAlO₃ component may thus prohibit oxygen vacancy substitution, although some regions such as harzburgite layers that was brought into the lower mantle by subduction have Al/Fe³⁺ ratios smaller than 1.0 (Liu et al., 2018). Studies of iron spin state indicate almost no Fe³⁺ (<1%) in the B site because the spin transition for the B site Fe³⁺ is observed in Al-free bridgmanite at ~40–60 GPa, but not found in Al-bearing samples (e.g., Liu et al., 2018; Mohn & Trønnes, 2016). Nevertheless, oxygen vacancy substitution is significantly suppressed by pressure (Liu, Nishi, et al., 2017); therefore, it is still unclear how much of the oxygen vacancy component is formed in Al-bearing bridgmanite, especially near the topmost lower mantle. Further studies at ~24–40 GPa in the MgO-SiO₂-Al₂O₃-Fe₂O₃ system are required.

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References

Andrautl, D., & Bolfan-Casanova, N. (2001). High-pressure phase transformations in the MgFe₂O₄ and Fe₂O₃-MgSiO₃ systems. *Physics and Chemistry of Minerals*, 28, 211–217.

Brodholt, J. P. (2000). Pressure-induced changes in the compression mechanism of aluminous perovskite in the Earth’s mantle. *Nature*, 407, 620–622.

Catalli, K., Shi, S. H., Prakapenka, V. B., Zhao, J., Sturhahn, W., Chow, P., et al. (2010). Spin state of ferric iron in MgSiO₃ perovskite and its effect on elastic properties. *Earth and Planetary Science Letters*, 299, 68–75.

Frost, D. (2008). The upper mantle and transition zone. *Elements*, 4, 171–176.

Frost, D., & Langenhorst, P. (2002). The effect of Al₂O₃ on Fe-Mg partitioning between magnesiowüstite and silicate perovskite. *Earth and Planetary Science Letters*, 199, 227–241.

Frost, D., Liebske, C., Langenhorst, F., McCammon, C. A., Trønnes, R. G., & Rüdib, D. C. (2004). Experimental evidence for the existence of iron-rich metal in the Earth’s lower mantle. *Nature*, 428, 409–412.

Frost, D., & McCammon, C. A. (2008). The redox state of Earth’s mantle. *Annual Review of Earth and Planetary Sciences*, 36, 389–420.

Greenberg, E., Xu, W. M., Nikolaevsky, M., Bykova, E., Garbarino, G., Glazyrin, K., et al. (2017). High-pressure magnetic, electronic, and structural properties of Mg₂Fe₂O₅ (M=Mg,Fe) ferric spinels. *Physical Review B*, 95(19), 195150. https://doi.org/10.1103/PhysRevB.95.195150

Grocholski, B., Shi, S. H., Sturhahn, W., Zhao, J., Xiao, Y., & Chow, P. C. (2009). Spin and valence states of iron in (Mg₀.₆Fe₀.₄)₂SiO₃ perovskite. *Geophysical Research Letters*, 36, L24303. https://doi.org/10.1029/2009GL041262

Hummer, D. R., & Fei, Y. (2012). Synthesis and crystal chemistry of Fe⁺⁺⁺-bearing (Mg,Fex⁺⁺⁺-SiFe,Fe⁺⁺⁺)₂SiO₃ perovskite. *American Mineralogist*, 97, 1915–1921.

Irifune, T., Shimizu, T., McCammon, C. A., Miyajima, N., Rubie, D. C., & Frost, D. J. (2010). Iron partitioning and density changes of pyrolite in Earth’s lower mantle. *Science*, 327(5962), 193–195. https://doi.org/10.1126/science.1181443

Ishii, T., Shi, L., Huang, R., Tsujino, N., Druzhbin, D., Myhill, R., et al. (2016). Generation of pressures over 40 GPa using Kawai-type multi-anvil press with tungsten carbide anvils. *Review of Scientific Instruments*, 87, 024501. https://doi.org/10.1063/1.4941716

Ismailova, L., Bykova, E., Bykov, M., Cerantola, V., McCammon, C., Boffa-Ballaran, T., et al. (2016). Stability of Fe, Al-bearing bridgmanite in the lower mantle and synthesis of pure Fe-bearing bridgmanite. *Science Advances*, 2, e1600427. https://doi.org/10.1126/sciadv.1600427

Jackson, J. M., Sturhahn, W., Shen, G., Zhao, J., Hu, M. Y., Errandonea, D., et al. (2005). A synchrotron Mössbauer spectroscopy study of (Mg₂Fe₂O₅)SiO₃ perovskite up to 120 GPa. *American Mineralogist*, 90, 199–205.

Kojitani, H., Katsura, T., & Akaogi, M. (2007). Aluminum substitution mechanisms in perovskite-type MgSiO₃. An investigation by Rietveld analysis. *Physics and Chemistry of Minerals*, 34, 257–267.

Kröger, F. A., & Vink, H. J. (1956). Relations between the concentrations of imperfections in crystalline solids. *Solid State Physics*, 3, 307–435.

Lauterbach, S., McCammon, C. A., van Aken, P., Langenhorst, F., & Seifert, F. (2000). Mössbauer and ELNES spectroscopy of (Mg,Fe)₃Al₂O₅ perovskite: a highly oxidised component of the lower mantle. *Contributions to Mineralogy and Petrology*, 138, 17–26.

Liu, J., Dolfam, S. M., Zhu, F., Li, J., Wang, Y., Zhang, D., et al. (2018). Valence and spin states of iron are invisible in Earth’s lower mantle. *Nature Communications*, 9, 1284.

Liu, Z., Akaogi, M., & Katsura, T. (2019). Increase of the oxygen vacancy component in bridgmanite with temperature. *Earth and Planetary Science Letters*, 505, 141–151.

Liu, Z., Boffa-Ballaran, T., Huang, R., Frost, D., & Katsura, T. (2019). Strong correlation of oxygen vacancies in bridgmanite with Mg/Si ratio. *Earth and Planetary Science Letters*, 523, 115697. https://doi.org/10.1016/j.epsl.2019.06.037

Liu, Z., Ishii, T., & Katsura, T. (2017). Rapid decrease of MgAlO₂.₅ component in bridgmanite with pressure. *Geochemical Perspectives Letters*, 5, 12–18.
Liu, Z., Nishi, M., Ichii, T., Fei, H., Miyajima, N., Boffa-Ballaran, T., et al. (2017). Phase relations in the system MgSiO3-Al2O3 up to 2300 K at lower mantle pressures. *Journal of Geophysical Research: Solid Earth*, 122, 7775–7788. https://doi.org/10.1002/2017JB014579

McCammon, C. A. (1997). Perovskite as a possible sink for ferric iron in the lower mantle. *Nature*, 387, 694–696.

McCammon, C. A., Lauterbach, S., Seifert, F., Langenhorst, F., & van Aken, P. A. (2004). Iron oxidation state in lower mantle mineral assemblages I. Empirical relations derived from high-pressure experiments. *Earth and Planetary Science Letters*, 222, 435–449.

Mohn, C. E., & Trønnes, R. G. (2016). Iron spin state and site distribution in FeAlO3-bearing bridgmanite. *Earth and Planetary Science Letters*, 440, 178–186.

Nakajima, Y., Frost, D. J., & Rubie, D. C. (2012). Ferrous iron partitioning between magnesium silicate perovskite and ferropericlase and the composition of perovskite in the Earth’s lower mantle. *Journal of Geophysical Research*, 117, B08201. https://doi.org/10.1029/2012JB009151

Navrotsky, A. (1999). A lesson from ceramics. *Science*, 284, 1788–1789.

Prescher, C., Langenhorst, F., Dubrovinsky, L. S., Prakapenka, V. B., & Miyajima, N. (2014). The effect of Fe spincrossovers on its partitioning behavior and oxidation state in a pyrolitic Earth&apos; lower mantle system. *Earth and Planetary Science Letters*, 399, 86–91.

Prescher, C., McCammon, C., & Dubrovinsky, L. (2012). MossA: A program for analyzing energy domain Mössbauer spectra from conventional and synchrotron sources. *Journal of Applied Crystallography*, 45, 329–331.

Richmond, N. C., & Brodholt, J. P. (1998). Calculated role of aluminium in the incorporation of ferric iron into magnesium silicate perovskite. *American Mineralogist*, 83, 947–951.

Ringwood, A. E. (1991). Phase transformations and their bearing on the constitution and dynamics of the mantle. *Geochimica et Cosmochimica Acta*, 55, 2083–2110.

Sinnmyo, R., Nakajima, Y., McCammon, C. A., Miyajima, N., Petitgirard, S., Myhill, R., et al. (2019). Effect of Fe2+ on phase relations in the lower mantle: Implications for redox melting in stagnant slabs. *Journal of Geophysical Research: Solid Earth*, 124, 12,484–12,497. https://doi.org/10.1029/2019JB017704

Tschauner, O., Ma, C., Beckett, J. R., Prescher, C., Prakapenka, V. B., & Rossman, G. R. (2014). Discovery of bridgmanite, the most abundant mineral in Earth, in a shocked meteorite. *Science*, 346(6213), 1100–1102. https://doi.org/10.1126/science.1259369

Walter, M. J., Kubo, A., Yoshino, T., Brodholt, J., Koga, K. T., & Ohishi, Y. (2004). Phase relations and equation-of-state of aluminous Mg-silicate perovskite and implications for Earth’s lower mantle. *Earth and Planetary Science Letters*, 222, 501–516.

Zhang, F., & Oganov, A. R. (2006). Valence state and spin transitions of iron in Earth’s mantle silicates. *Earth and Planetary Science Letters*, 249, 436–443.