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CFD convective flow simulation of the varying properties of CO$_2$-H$_2$O mixtures in geothermal systems

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Abstract

Numerical simulation of a geothermal reservoir, modelled as a bottom-heated square box, filled with water-CO$_2$ mixture is presented in this work. Furthermore, results for two limiting cases of a reservoir filled with either pure water or CO$_2$ are presented. Effects of different parameters including CO$_2$ concentration as well as reservoir pressure and temperature on the overall performance of the system are investigated. It has been noted that, with a fixed reservoir pressure and temperature, any increase in CO$_2$ concentration leads to better performance, i.e. stronger convection and higher heat transfer rates. With a fixed CO$_2$ concentration, however, the reservoir pressure and temperature can significantly affect the overall heat transfer and flow rate from the reservoir. Details of such variations are documented and discussed in the present paper.

Keywords: Geothermal, reservoir, CO$_2$, water, CFD

Introduction

Simultaneous power generation and geo-sequestration makes CO$_2$ a very attractive choice for geothermal power plants. As such, carbon-dioxide-based engineered geothermal systems (CO$_2$-EGS) have been previously proposed as an alternative to water-based EGS systems (Brown 2000). Subsequent studies
added further details and reported possibility of improved energy extraction (Pruess 2006; Pruess 2008; Atrens et al. 2009; Nago and Nieto 2011; Fard et al. 2011; Hooman et al. 2014). Interestingly, the ‘dry-out period’, or transition from an initially water-filled EGS system to a CO$_2$-rich one (Atrens et al. 2011a-c; Borgia et al. 2012) has received a lot of attention mainly because of possible mineral dissolution and precipitation as a result of changes in the reservoir fluid composition and reservoir permeability alteration (Kneafsey and Pruess 2010; Spycher and Pruess 2010; Spycher and Pruess 2011). Additionally, Atrens and Gurgenci (2013) investigated the effects of CO$_2$-rich phase compositions on the production flow rate and the heat extraction from the reservoir. What is yet to be reported in the literature is a detailed numerical simulation of a water-CO$_2$ mixture filling a reservoir. Simple thermodynamic analysis of a reservoir shows that more heat can be extracted (compared to a water-saturated reservoir) mainly because a CO$_2$-water mixture is more buoyant than pure water. This could significantly affect the energy extraction from a water-saturated reservoir. Most of these reservoirs, suffer from permeability drops with depths. As such, having a more buoyant fluid which could move up against gravity, at least partially, is not only desirable but also going to significantly affect the drilling cost which is expected to grow exponentially with the well depth (Atrens et al., 2010; Augustine et al., 2006).

The role of fluid migration in the build-up of heat in underground geothermal systems is not well-understood. It is known that degassing of CO$_2$ with isotopic composition indicating mantle-sources in regions of tectonic activity is associated with locally elevated geothermal temperatures (Italiano et al., 2011; Uysal et al., 2009 and 2011). Convective fluid plumes may play a role in enhancing heat flows from the mantle to geothermal reservoirs, and within the reservoirs themselves. Additionally, convection within a geothermal reservoir may enhance the productive life-time of geothermal reservoir systems by enhancing heat supply from underlying strata, and by ensuring a more even distribution of thermal energy throughout the reservoir – i.e. by off-setting localized cooling along major flow paths.

This could also imply that for a water-saturated reservoir which is classified as non-productive, mainly due to lack of convection currents, injection of CO$_2$ can lead to formation of convection cells and thus facilitate heat extraction. In what follows, a numerical analysis of this problem is presented to quantitatively measure the improved convective flow patterns and enhanced heat transfer from the reservoir. This is intended to provide insight as to the possible mechanisms by which CO$_2$ presence in or addition to underground reservoirs could lead to enhancement of convective heat transfer.

**Modelling**
The reservoir is modelled as a bottom-heated square box with adiabatic lateral boundaries and a cold top wall as Fig. 1 shows. The cold and hot temperatures are varied from 331-431 K and 416-516 K, respectively, in a way that the hot-cold temperature difference remains at 85 K for each case. For constant properties, and of course with the same reservoir size, porosity and permeability, one would expect the results to be the same as long as the temperature difference is not altered. It will, however, be shown in the forthcoming sections that this is not the case in our problem as properties significantly vary with both temperature and pressure. The reservoir porosity is fixed at 0.05 and the permeability-length product is kept constant at $10^{11}$ m$^3$ similar to Fard et al. (2010) with no through-flow. The reservoir pressure is varied from 20 to 60 MPa (equivalent to the hydrostatic pressure of approximately 2 to 6 km of water) to cover a wide range of practical applications for geothermal development. The flow is modelled using Darcy flow model with the single-phase fluid properties obtained as linear superposition of those of individual fluids weighted with their respective fraction, i.e.

$$\rho_g = \omega_f \rho_f + \omega_e \rho_e$$  \hfill (1)

$$\beta_g = \omega_f \beta_f + \omega_e \beta_e$$  \hfill (2)

$$c_{pg} = \omega_f c_{pf} + \omega_e c_{pe}$$  \hfill (3)

$$\mu_g = \omega_f \mu_f + \omega_e \mu_e$$  \hfill (4)

The real (i.e. non-ideal) properties are determined for pure CO$_2$ using a Helmholtz-free-energy based equation of state (Span and Wagner, 1996), and for pure H$_2$O using the International Association for the Properties of Water and Steam equation of state (Span et al, 2000). The use of a single-phase basis for modelling is to enable examination of the effect of changing properties without additional complexities of two-phase flow separation behaviour. Further comment is provided in the discussion. The reservoir properties are then obtained similarly using the porosity and solidity as the weight-functions as follows

$$\lambda_R = \phi(\omega_f \lambda_f + \omega_e \lambda_e) + (1 - \phi)\lambda_m$$  \hfill (5)

One notes that in the above formulation changes in the reservoir porosity and permeability are not taken into account while fluid properties are updated in each iteration.

An in-house code, used by Hooman and Gurgenci (2008), was cross-validated with the commercially available software ANSYS-FLUENT and used to create and mesh the geometry and finally solve the governing equations. The governing equations were derived as for standard two-dimensional heat transfer conditions, based on the assumptions of adiabatic lateral walls and constant temperature for the top and bottom boundaries, and are as follows:
\[
\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = 0 \quad (6)
\]

\[
\frac{\partial u}{\partial y} - \frac{\partial v}{\partial x} = -\frac{\rho g \beta K}{\mu} \frac{\partial T}{\partial x} \quad (7)
\]

\[
u \frac{\partial T}{\partial x} + v \frac{\partial T}{\partial y} = \alpha \left( \frac{\partial^2 T}{\partial x^2} + \frac{\partial^2 T}{\partial y^2} \right) \quad (8)
\]

Subject to the boundary conditions illustrated in Fig. 1.

Fig. 1 Schematic view of the computational domain

**Numerical details**

Grid independence was verified by running the software on different combination of grid sizes. It was observed that the results changed less than 2% when a 100×100 mesh system is used instead of a finer mesh with 200X200 grid points. Results are also verified for constant property free convection of water in a porous cavity, i.e. the Darcy-Benard problem. It was noted that the correlation between Nusselt number \((Nu)\), and Rayleigh number \((Ra)Nu = Ra/40\) best fits out numerical data, as Fig. 2 shows. \(Ra\) in this instance is determined as per equation 9:

\[
Ra = \frac{g \beta \rho}{\mu a} (T_h - T_c) KH \quad (9)
\]
As a further check on the accuracy of our results, variable property Darcy-Benard free convection of pure water in a porous cavity was investigated to observe that using the reference temperature approach of Hooman and Gurgenci (2008) the above correlation can still be used within 5%.

Results and discussion

In what follows we focus on free convection heat and fluid flow of a water-CO$_2$ mixture in a porous cavity. We use $Nu$ and maximum flow rate as our metrics to evaluate the strength of convective flow patterns. $Nu$ is the total heat transfer divided by that of pure conduction through the same cavity (no convective flow patterns). As such, any $Nu$ value in excess of unity shows some degree of convection. Obviously, higher $Nu$ values mark stronger convective cells. The flow rate reported here is the one induced by free convection only, i.e. without a well-head pump or any other suction/injection mechanisms. We systematically change the CO$_2$ mass fraction from zero (pure water) to unity (pure CO$_2$) over a range of reservoir pressure and temperature in a way that the hot-cold temperature difference
remains the same. For a constant property sub-critical fluid flow, one would expect that with the same temperature difference, and hence, the same Rayleigh number \((Ra)\), the overall heat transfer and fluid flow will not alter. However, as \(CO_2\) is supercritical within the range of conditions of underground reservoir systems, that is not the case for mixtures of \(CO_2\) and \(H_2O\), as demonstrated by Fig. 3. This figure shows \(Nu\) versus \(CO_2\) mass fraction at 20MPa with the same hot-cold temperature difference but with different hot and cold temperatures as denoted on the plots. As seen, the heat transfer increases with \(CO_2\) mass fraction for any given \(T_h\) and \(T_c\) combination. Furthermore, moving from pure water to pure \(CO_2\), the increase in heat transfer is significant; about one order of magnitude is the minimal heat transfer augmentation. More interestingly, however, is the fact that \(Nu\) is the highest with the lowest \(T_c\) (and obviously lowest \(T_h\) to remain the same \(\Delta T\) of 85K) mainly because the lower temperatures are closer to those of pseudocritical conditions where \(Ra\) is expected to reach a maximum value; see also Foroozhi et al. (2012-2014). This is obviously in favour of low temperature geothermal reservoirs which may not be productive when pure water is the working fluid.

![Graph showing Nusselt number versus CO2 mass fraction](image-url)

**Fig 3** Nusselt number versus CO2 mass fraction for different \(T_h\) and \(T_c\) combinations with \(\Delta T=85K\)
Figure 4a-c are presented to demonstrate $Nu$ versus CO$_2$ mass fraction for different reservoir pressures and hot-cold temperature combinations. $Nu$ increases with mass fraction for a fixed pressure and hot-cold temperature combination. Comparing the relationship between any of the figure 4a-c for a fixed pressure will result in conclusions similar to what were drawn based on close examination of figure 3. That is, heat transfer increases for temperatures close to pseudocritical conditions. Moreover, based on plots in the same chart, increasing the pressure leads to lower heat transfer rates for a fixed CO$_2$ mass fraction and temperature. This could be explained as the obvious decrease in compressibility and increase in the fluid density with higher pressures, with a fixed fluid temperature, which will lead to lower thermal expansion coefficients. As a result, at the same temperature, either of the two fluids will be less buoyant at higher pressures when compared to lower ones; so will be the mixture in the absence of any phase transitions.
Fig. 4 Nusselt number versus CO$_2$ mass fraction for different reservoir pressures and temperature combinations: a) $T_h=416$ K; $T_c=331$K, b) $T_h=466$ K; $T_c=381$K, and c) $T_h=516$ K; $T_c=431$K

The convective flow rates also reflect a dependence on compressibility, as demonstrated by Fig. 5. The dimensionless flow rate (normalized stream function on the vertical axis) is obtained by normalizing the actual flow rate with appropriate scales for velocity, area and density:

$$\xi = \frac{\dot{m}}{\rho u A}$$

Mathematically, it means that we used the group $\rho A u$ to non-dimensionalize the flow rate. It needs to be mentioned that the choice of these parameters is optional but we tried to use constant values for density and length to make it easy for the reader to generate estimates, based on our calculation, for expected flow rates through a given reservoir. Moreover, what we are more interested in is the trend of the flow rate plot against the mass fraction than the actual flow rate values. In doing so, the (constant) density of water at atmospheric condition is used where the unit area is used defined as the length of the cavity multiplied by unity (1 m). The flow velocity, for single-phase constant property case, is assumed to be linearly
proportional to the product of the thermal diffusivity and \( Ra^{1/2} \) and inverse linearly proportional to the cavity length, e.g.

\[
u \sim \frac{\alpha}{H} \sqrt{Ra}
\]

(10)

The flow rate is given by

\[
\dot{m} = \rho Hu \sim \rho \alpha \sqrt{Ra}
\]

(11)

The product of thermal diffusivity and density is independent of density and leaves us with a group \( \lambda/c_p \).

Consequently the mass flow rate scale is represented as

\[
\dot{m} = \rho Hu \sim \frac{k}{c_p} \sqrt{Ra}
\]

(12)

While \( \lambda \) and \( c_p \) are calculated under standard atmospheric conditions, \( Ra \) is affected by fluid property variation and following the use of Eqns. (1-5). This flow rate here is the buoyancy-induced flow rate due to changes in fluid density. The buoyancy induced flow leads to an upward movement of hot fluid toward the top wall, where it is cooled and then displaced by other rising hot fluid. Results of mass flow rate normalized by using Eq. (12) above are presented in Fig. 5 for different mass fractions, pressure and temperature combinations. Similar to \( Nu \) plots, one notes that the mass flow rate is not only sensitive to the temperature difference but also to the actual wall temperature values. Furthermore, higher \( CO_2 \) mass fraction leads to higher flow rates. It can be noted, moving from Fig. 5a to 5c, that flow rates are less sensitive to pressure as \( T_h \) is increased. With a fixed \( T_h \) and \( T_c \), one notes different trends in flow rate when pressure changes. Depending on the temperature values, an increase in pressure can either increase (Fig. 5a) or decrease the flow rate (Fig. 5b).
This work is an initial analysis of the role of CO$_2$ enhancement of convective heat transfer within geothermal reservoirs. It deliberately assesses the behaviour of a single-phase mixture of the two components. Further work is necessary to extend this to account for multiple phases. There are three particular qualitative effects through which multiphase flow is expected to alter the results presented here:

1. Transient exsolution of dissolved CO$_2$ as bubbles within the two phase region should lead to local enhancement of convective flow around the bubble due to its upwards buoyancy-driven motion. One expectation of this would be an increase in the gradient of convective heat flux with mass fraction (i.e. $d\xi/dx$) at the bubble line, where CO$_2$ begins to exsolve from the H$_2$O phase.

2. Relative permeability within the two-phase region would act to reduce enhancement of flows, as the reservoir permeability to the minor phase within a two-phase flow is typically substantially reduced.
3. Under steady-state conditions, there would be an expectation of phase separation into two horizontal phases based on relative density, i.e. an upper CO₂-rich phase, and a lower H₂O-rich phase. The upper phase would experience significantly enhanced convective heat transfer rates, as it would have internal heat transfer characteristics similar to the right sides of Figs. 4 & 5. Additionally, the heat transfer would be further enhanced by the temperature-dependent solubility of H₂O in the CO₂-rich phase, leading to additional H₂O evaporating into the CO₂ phase at the boundary between the two phases, and condensing at the upper surface of the reservoir. The H₂O-rich phase would experience the converse effect, depressing the rate of heat transfer, although CO₂ solubility in H₂O is far less dependent on temperature than that of H₂O in CO₂ (Atrens et al., 2014).

These expected qualitative behaviours require further analysis accounting for multiphase flow behaviour to determine their relative contribution to overall convective heat flow enhancement.

However, the sum of these changes is not expected to reverse the overall trend demonstrated here, of increased convective heat flux as CO₂ is added to the reservoir system. Considering that the results presented here indicate that CO₂ may enhance flow rates by up to a factor of 2.67, we conclude that this is a potentially important mode of heat transport within geothermal reservoirs and warrants further study. We anticipate the next steps to be consideration of the additional flow behaviours when multiple phases are present.

Conclusion

The effect on convective heat transport within a closed reservoir system of varying fluid composition was analysed by CFD modelling of a single-phase fluid with properties derived from a composition-dependent average of pure CO₂ and pure H₂O. As compositional properties were varied from H₂O toward that of CO₂, substantial increases were observed in Nusselt number (by a factor of up to 10) and normalised stream function (by a factor of up to 2.67). We conclude that this indicates substantial increase in convective heat transport.

Convective heat transport may be further modified by multiphase flow behaviours, and we conclude that due to the potential magnitude of heat flow enhancement by the addition of CO₂, further research exploring the effect of these behaviours on heat transport is warranted.

We anticipate this finding to have implications for the study of natural geothermal reservoirs, where the role of dissolved gas exsolution on heat transfer enhancement remains unquantified. Additionally, these
findings may be of potential interest with regards to CO₂ injection into geothermal reservoirs, as it may lead to improved productivity through the mechanisms elucidated in this work.

The magnitude of this effect on the thermal productivity of geothermal power plants is difficult at this stage to quantify, and probably not meaningful to speculate on due to the limitation of assessing only single-phase flow behaviours within the reservoir. These findings do demonstrate, however, that reservoirs with elevated CO₂ content will experience greater convective heat transfer, and therefore be of comparatively higher temperature (and therefore of greater resource value). Additionally, it can be concluded that any increase in CO₂ content of an existing reservoir will enhance convective flow behaviours (although the true magnitude of this effect will depend on two-phase flow behaviour as well as the particulars of the reservoir), and consequently will enhance the productivity and/or the longevity of geothermal energy extraction.

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**Glossary**

\( \rho \) – Density, kg m\(^{-3}\)

\( \omega \) – Mass fraction, dimensionless

\( \beta \) – Thermal expansion coefficient, K\(^{-1}\)

\( \mu \) – Dynamic Viscosity, Pa s

\( \lambda \) – Thermal conductivity W m\(^{-1}\) K\(^{-1}\)

\( \phi \) – Reservoir porosity, dimensionless

\( c_p \) – Constant-pressure heat capacity, kJ kg\(^{-1}\) K\(^{-1}\)

\( g \) – Gravitational acceleration, 9.81 m s\(^{-2}\)

\( u \) – Horizontal component of fluid velocity, m s\(^{-1}\)

\( v \) – Vertical component of fluid velocity, m s\(^{-1}\)
Nu – Nusselt Number, dimensionless

Ra – Rayleigh Number, dimensionless

T – Temperature, K

ξ – Normalised flow rate, dimensionless

m – Mass flow rate, kg s\(^{-1}\)

A – Area, m\(^2\)

H – Reservoir height, m

K – Reservoir permeability, m\(^2\)

α – Thermal diffusivity, m s\(^{-2}\)

**Subscripts**

γ – Carbon Dioxide

e – Water

g – Fluid mixture

R – Reservoir

m – Rock

h – Highest system temperature, at bottom boundary

c – Lowest system temperature, at top boundary

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