Drainage Ditches Contribute Considerably to the CH4 Budget of a Drained and a Rewetted Temperate Fen

Daniel Köhn (✉ j.daniel.koehn@gmail.com)  
University of Rostock  https://orcid.org/0000-0001-5435-8831

Carla Welpelo  
Johann Heinrich von Thünen -Institut Institut für Agrarklimaschutz: Johann Heinrich von Thunen - Institut Institut fur Agrarklimaschutz

Anke Günther  
University of Rostock: Universitat Rostock

Gerald Jurasinski  
University of Rostock

Research Article

Keywords: Methane (CH4), peatland rewetting, restoration, ditches, methane budget, ebullition

Posted Date: February 19th, 2021

DOI: https://doi.org/10.21203/rs.3.rs-223145/v1

License: ☺ This work is licensed under a Creative Commons Attribution 4.0 International License.  
Read Full License
Abstract

Small water bodies including (former) drainage ditches can be hotspots for methane (CH$_4$) emissions from peatlands. We assessed the CH$_4$ emissions of a drained and a rewetted temperate fen including emissions of active and former drainage ditches over the course of 2.5 years, covering three vegetation periods. Ditch CH$_4$ emissions in the rewetted fen were significantly higher than in the drained fen. In the rewetted fen ditches contributed up to 91 % of the annual CH$_4$ budget, despite covering only 1.5 % of the area. In the drained fen CH$_4$ emissions were solely made up of ditch emissions. When including CH$_4$ uptake by the peat soil, the CH$_4$ balance of the drained fen was neutral. Dissolved organic carbon concentrations likely had an enhancing effect on CH$_4$ emissions while nitrate and sulphate in the ditch water seem to have had an inhibitory effect. Air and water temperature controlled seasonal variability of ebullitive as well as diffusive CH$_4$ emissions. Ebullition contributed less than 10 % to the overall CH$_4$ budget in the ditches. Drainage ditches represent a hotspot of CH$_4$ emissions and need therefore be taken into account when assessing the success of rewetting projects of peatlands.

Introduction

Peatlands are a globally important carbon store (Treat et al. 2019) that is turned into a strong source of greenhouse gases (GHGs) when drained, and faces other threats, for instance, from global warming (Loisel et al. 2020). Peatland rewetting represents an efficient way to reduce or stop GHG emissions. Due to the high availability of organic substrate in the soil, water-logged areas in drained or rewetted peatlands can become hotspots for emissions of the GHG methane (CH$_4$).

Small water bodies play an important role in the global carbon cycle (Bastviken et al. 2011; DelSontro et al. 2016; Holgerson and Raymond 2016). However, only few studies have so far examined the importance of CH$_4$ emissions from drainage ditches in peatlands. Drainage ditches can be important hotspots for CH$_4$ emissions in wetlands (Schrier-Uijl et al. 2011, 2010), sometimes contributing a major part of the total regional CH$_4$ budget (Schrier-Uijl et al. 2010). Also in agricultural landscapes drainage ditches may contribute significantly to the landscape carbon budget via high CH$_4$ emissions (Peacock et al. 2017). In this context, ebullition is often mentioned as an important pathway of CH$_4$ emissions in various aquatic ecosystems (Baulch et al. 2011; Bastviken et al. 2004; Repo et al. 2007; Yang et al. 2020).

The major biotic factor driving high CH$_4$ emissions is thought to be the trophic state of the water body (Schrier-Uijl et al. 2011). Phosphate (PO$_4$) and reduced iron in the ditch water are indicators for anaerobic conditions and can explain a large proportion of the variance in CH$_4$ emissions (Schrier-Uijl et al. 2011). In connection with the trophic status of water bodies, the oxygen concentration in the water column is a good indicator for CH$_4$ emissions (Liikanen et al. 2003). Since methanogenesis depends on small organic carbon molecules, either carbon dioxide (CO$_2$), hydrogen (H) or acetate (H$_3$CO) as a substrate (Kelly and
Chynoweth 1981), the concentrations of dissolved organic matter (DOM) or dissolved organic carbon (DOC) are important drivers of CH\textsubscript{4} emissions in small water bodies (Bastviken et al. 2004; Zhou et al. 2019). However, it is often unclear whether the organic matter in ditches mainly derives from high biomass production within the ditch or from allochthonous DOC that was potentially leached at high rates from surrounding decomposing peat as was shown in a mesocosm experiment (Laine et al. 2014).

High nutrient inputs from surrounding agriculturally-used peat soils can cause eutrophication in the ditches and thereby enhance plant and algal biomass production and subsequent depletion of oxygen from biomass decomposition (Zhou et al. 2019). Increased nutrient concentrations and experimental warming showed an increase in CH\textsubscript{4} emissions from small water bodies in a study on CH\textsubscript{4} ebullition from lake mesocosms (Davidson et al. 2018). This relationship was also shown in natural northern lakes and ponds (DelSontro et al. 2016). Shallow water bodies such as most ditches are highly susceptible to warming and eutrophication because of their small water volume and climate warming is expected to globally increase the CH\textsubscript{4} emissions via ebullition by up to 51 % (Aben et al. 2017).

Here, we study the importance of ditch CH\textsubscript{4} emissions in regional GHG budgets and the drivers for temporal variation in two peatlands with differing land use. We determine the effects of climatic (air temperature, water temperature, air pressure), biotic (DOC, nutrients) and morphological (water depth, orientation) variables on CH\textsubscript{4} emissions from ditches and evaluate the importance of ebullitive CH\textsubscript{4} fluxes in relation to diffusive fluxes. Using a 2.5 year time series of floating chamber measurements and closed chamber measurements in the adjacent peatland, we assess the interannual variability of CH\textsubscript{4} fluxes and seasonal CH\textsubscript{4} budgets.

**Materials And Methods**

**Site description**

The two studied fens are located 8 km apart in the valleys of the two rivers Recknitz (drained fen) and Trebel (rewetted fen) in north-eastern Germany. The average annual mean temperature is 9.1°C (DWD raster data, Krähenmann et al. 2016). The drained fen (PD, 54.13194° N, 12.62889° E, elevation a.s.l = 20 m) is an extensively used grassland which is harvested once a year for fodder production. The rewetted fen (PW, 54.10111° N, 12.73944° E, elevation a.s.l. = 2 m) has been rewetted in 1997 after being used as intensive grassland for decades. After rewetting, the water table in PW now fluctuates around the soil surface. Peat thickness is around 5 m in PD and approx. 6 m in PW. The peat in both sites is mainly of sedge and reed origin (Jurasiński et al. 2020). The vegetation at PD can be characterised as a uniform grassland dominated by *Ranunculus repens* L. and *Deschampsia cespitosa* (L.) P. Beauv. PW is dominated by sedges (*Carex acutiformis* Ehrh.) and occasional great willowherb (*Epilobium hirsutum* L.) and grey sallow (*Salix cinerea* L.). Especially around ditches and former peat cuttings large areas of reed (*Phragmites australis* Trin. ex Steud.) and occasional cattail (*Typha latifolia* L.) can be found.

**Study setup**
At each study site (PD and PW) a soil measurement site was established inside a fenced area (12 x 30 m) between April and June 2017 (Fig. 1). Five collars for the measurement of soil CH$_4$ exchange were installed along a boardwalk at both sites. The soil collars also included vegetation. Weather stations inside both fenced areas recorded air temperature, humidity, photosynthetic photon flux density (PPFD), wind speed, wind direction and precipitation (logged with CR300, Campbell Scientific, Bremen, Germany). Additionally, air pressure, vapour pressure and sunshine duration was obtained from three different weather stations in proximity of the soil sampling site (Warnemünde − 40 km NW, Barth − 30 km N, Greifswald, 40 km E), run by the german weather service (DWD). For analyses, the values of all three weather stations were averaged.

We selected two ditches to measure diffusive and ebullitive CH$_4$ exchange from the water surface in close proximity to each soil sampling site (~ 300–400 m distance, Fig. 1). At each site, one of the selected ditches runs parallel to the drainage direction (PD-p, PW-p) and one ditch runs orthogonal to the drainage direction (PD-o, PW-o) towards the main river. In all four ditches five sampling spots were established at approximately 10 m from each other (20 ditch sampling locations in total, Fig. 1).

The ditches at PD are relatively uniform with a width of approximately 2 m and are regularly excavated in summer (own observations). Accordingly, the depth of the ditches varies throughout the year, ranging from 10 to 70 cm. During summer, the ditches are often covered by common duckweed (**Lemna minor** L.) Further, water starwort (**Callitriche palustris** L.) was abundant. At PW the ditches are not managed and, thus, do not vary in depth over the year. PW-o, however, is significantly deeper than PW-d with average depths of 104 cm and 38 cm, respectively. Also, PW-o is much wider than PW-d with approximately 4 m and 2 m, respectively. The ditches at PW are often covered entirely with vegetation during the summer months, with **Stratiotes aloides** L. being dominant in PW-o and **Typha latifolia** L. and **Lemna minor** L. being dominant in PW-p. The banks of both ditches in PW are dominated by **Phragmites australis** (Cav.) Trin. ex Steud.

Flux measurements

Diffusive CH$_4$ fluxes

Diffusive emissions of CH$_4$ from the ditches were measured with a floating chamber. The floating chamber was constructed using a bucket (diameter = 20 cm, height = 25 cm), coated with reflective material to reduce heating inside the chamber (Fig. 2). The chamber was equipped with a temperature and humidity sensor as well as with a fan powered by a 9 V battery mounted inside the chamber lid. The chamber was placed inside a float (square 50 x 40 cm, Styrodur, BASF, Ludwigshafen am Rhein, Germany) and connected to a 1.5 m long handle. CH$_4$ concentration measurements were carried out in-situ with laser spectrometers (‘Ultra-Portable Greenhouse Gas Analyzer’, Los Gatos Research, Mountain
View, USA and ‘GasScouter’, Picarro, Santa Clara, USA) connected to the chamber with flexible polyurethane tubes (inner diameter: 4 mm). Measurements lasted 180 s.

Diffusive CH$_4$ flux measurements on the soil surface at PD and PW were carried out with circular flexible chambers constructed out of polyurethane walls varying in height between 0.9 and 1.4 m. The diameter of the soil chamber was 0.65 m. The soil chamber was also equipped with three fans at the chamber top ensuring constant mixing of the air inside the chamber. Diffusive fluxes from the ditches and the soil surface at both sites were performed between April 1st 2018 and September 29th 2020.

Diffusive fluxes were estimated using the $flux$ function of the package $flux$ (Jurasinski et al. 2014) for R (R development core team 2020). The slope between all concentration points was calculated and the median slope was used for flux estimation (median-based regression, Siegel 1982). All diffusive flux measurements were visually checked for signs of ebullition (i.e. strong, sudden increase in CH$_4$ concentrations). If an ebullition event was identified during a diffusive flux measurement, it was excluded from the calculation of annual CH$_4$ balances (157 fluxes excluded, 302 fluxes remaining at ditches in PD and 182 fluxes excluded, 374 fluxes remaining at ditches in PW).

**Ebullitive CH$_4$ fluxes**

Ebullitive CH$_4$ emissions were assessed during the vegetation period of 2018. Bubble traps were installed floating in the middle of the ditches (five measurement points at each ditch). The bubble traps were constructed from inverted polypropylene funnels (15 cm diameter opening) connected to a 120 ml syringe that functioned as the gas reservoir, similar to the approaches of Molongoski & Klug (1980) and Baulch et al. (2011). The funnel and the syringe were attached to each other with an insoluble adhesive sealant and a three-way stop cock allowed sampling at the top of the trap (Fig. 3). To prevent large water insects such as water scavenger beetles (Hydrophilidae) from entering the bubble trap we covered the opening of the funnel with a net (polyvinyl chloride, net width 5 mm). The traps were provided with a 20×30×5 cm cuboid float (Styrodur, BASF, Ludwigshafen am Rhein, Germany). The bubble traps were fixed in place by cables running between the float and both banks of the ditch to prevent any disturbance to the sediment.

To prepare for gas collection all bubble traps were filled with water completely. During the time in which the trap is deployed, rising bubbles are trapped in the funnel and replace the water inside the trap. After approximately two weeks (11–14 days) the volume of the accumulated gas in the trap was noted by reading the printed scales on the syringes. Gas samples were taken from the headspace collected in the syringe without disturbing the bubble trap by laying a portable aluminium footbridge across the ditch. Because PW-o was too wide to reach both banks, the bridge was instead placed onto a small, permanently-installed wooden platform inside the ditch. Gas samples were taken with a 60 ml syringe and immediately transferred to 12 ml exetainers (Labco, Lampeter, UK). The final sample volume was

Page 5/31
approximately 35 ml, thus the sample was stored with overpressure. After sampling we refilled the bubble trap completely with water.

Due to the long deployment times, CH$_4$ concentrations of air caught inside the bubble traps may have decreased due to equilibration with the water column or CH$_4$ oxidation (McGinnis et al. 2006). To quantify this potential error, fresh bubbles were collected using a mobile bubble trap by intentionally disturbing the sediment to induce ebullition at random locations within the ditch. Fresh bubbles were always collected after the permanent bubble traps had been sampled.

Gas analyses were performed within one week using a gas chromatograph (Shimadzu GC, Kyoto, Japan) with a flame ionization detector. As concentrations of CH$_4$ varied strongly, the samples had to be diluted up to a factor of 1000 and measured in different sensitivity ranges of the gas chromatograph.

Final ebullition fluxes were calculated as such: We assumed that bubbles caught in the traps originated from an area of sediment that corresponded to the area of the funnel opening (~ 0.0176 m$^2$). Thus, we normalized the recorded gas volumes in the bubble traps to 1 m$^2$ and divided by the number of days since the last sampling (ml m$^{-2}$ d$^{-1}$). Then, we multiplied this value with the CH$_4$ concentration measured inside the gas samples (ppb). When there was no gas sample taken the arithmetic mean of the CH$_4$ concentration from all gas samples was taken to estimate the ebullitive CH$_4$ emission. Every bubble was sampled for CH$_4$ concentration every four weeks, meaning that every second week the arithmetic mean was taken as an estimate for the CH$_4$ concentration.

The final ebullition flux $F_e$ was estimated by firstly estimating the CH$_4$ bubble rate in moles according to Eq. 1:

$$F_e = \frac{p \times v}{R \times T} \times c \times m$$  \hspace{1cm} (1)

with $P$ the atmospheric pressure, $v$ the volume of gas measured inside the bubble trap, $R$ the gas constant ($R = 0.0821$), $T$ the temperature in the laboratory during analyses (298 K), $c$ the concentration of CH$_4$ in the gas sampled (% by volume) and $m$ the molar weight of CH$_4$ (16.04 g mol$^{-1}$)

Greenhouse gas budgets

We used a combination of bootstrap, jackknife and linear interpolation of the fluxes to calculate seasonal budgets (Günther et al. 2017). For each measurement day, one flux value per flux subset (ditch or soil) and each site (PD or PW) was randomly chosen. This was repeated 100 times to obtain 100 different flux time series. Then, the area-under-curve (auc.mc function from the R package flux, Jurasinski et al. 2014) was calculated 100 times for each flux time series each time leaving out one flux value, leading to a total of 10 000 different CH$_4$ balances. For the final CH$_4$ balances per site and flux subset we calculated the
average and standard deviation of all balances. Using this procedure yields a more robust estimate of the seasonal CH₄ budgets as it is more sensible for temporal variation than the simple average of all flux measurements. The CH₄ balances were calculated per season (i.e. vegetation period (April – September) and non-vegetation period (October – March).

To estimate the contribution of CH₄ emitted from ditches to total ecosystem emissions we manually determined the area covered by ditches by digitizing them within a randomly-chosen 1 km² area around the soil sampling site using aerial imagery. The area share of ditches was approximately 1.52 % and 1.49 % in PD and PW, respectively. To derive the total contribution of ditch CH₄ emissions to the overall ecosystem CH₄ budget, the ditch budgets were weighed using their relative spatial share within the 1 km² area. The rest of the area was assumed to emit on average as much CH₄ as the soil sample locations.

Water and sediment characteristics

Water samples of ditch water and groundwater at the soil sampling site were taken to assess potential influences of chemical properties (i.e. nutrients, DOC) on CH₄ emissions. Ditch water samplings took place at irregular intervals, however, covering all seasons between April 2018 and March 2019. On each sampling occasion, one sample was taken for every flux measurement location in the ditch (n = 20). Groundwater samples were taken at the central site every four weeks between April 2018 and September 2020. The groundwater samples were obtained from three water gauges per site which were located close to the soil surface measurement plots. All water samples were directly filtered in the field with syringe filter units (pore size 0.45 µm, Sartorius, Göttingen, Germany) and afterwards stored cool (~5°C) or frozen until analysis. The water samples were analysed for DOC, dissolved inorganic carbon (DIC), phosphate (PO₄³⁻), nitrate (NO₃⁻), ammonium (NH₄⁺) and total nitrogen (TN). DOC, DIC and TN were analysed using a Dimatoc 2100 (Dimatec, Essen, Germany). All other nutrients were analysed with an AA3 SEAL Auto Analyzer 3HR continuous flow analyzer (SEAL Analytical, Norderstedt, Germany). Further, a Multiprobe AP 2000 (Aquaread, Bridge House, UK) was used to measure pH, O₂ saturation [%], water temperature [°C], electrical conductivity [S m⁻¹], redox potential [mv] and salinity [µS] directly in the ditches at a depth of approximately 15 cm.

In June 2018 sediment samples were taken in two depths (0–5 cm, 10–20 cm) from the sediment surface of each ditch. The samples were dried for 24 hours at 105°C and subsequently ground for three minutes. Carbon, nitrogen and sulphur concentrations in the sediment samples were analysed on a vario EL cube CNS analyser of elementar, Hanau, Germany.

Statistical analyses

All statistical analyses and visualizations were carried out with R 4.0.2 (R development core team 2020). The entire dataset and any subsets were tested for normality and homogeneity of variance using Shapiro-Wilk tests and Levene's tests, respectively. Where data was non-normally distributed or the variance was not homogeneous, Kruskal-Wallis tests were used to detect significant differences between
subpopulations within the dataset. In order to assess relationships between water chemical variables (pH, \(O_2\) saturation, water temperature, electrical conductivity, redox potential and salinity) and \(CH_4\) fluxes all values were averaged by date, because not every flux measurement had an associated measurement of water chemical variables. Diffusive \(CH_4\) fluxes were log-transformed in order to achieve a near normal distribution. Transformed diffusive flux values were subsequently directly linked to environmental variables from the weather station with multiple linear regressions (wind direction, wind speed, air temperature, air pressure, vapour pressure and pressure change over different time intervals). Additional variables, such as concentration values of nutrients and DOC that were either only available at certain dates were merged with daily average ebullitive and average non-transformed diffusive \(CH_4\) fluxes.

**Results**

**Environmental variables and ditch characteristics**

According to the weather station data air temperature differed only slightly between PD and PW. However, during the two-year study period PD was much drier than PW, receiving only 973 mm in comparison to 1173 mm at PW. Nonetheless, there were pronounced dry spells in both peatlands during the summer months of 2018 and 2019 (Fig. 4).

In PD water levels remained close to the soil surface in winter and were very low in summer (overall mean = -28 cm, overall minimum = -78 cm) whereas in PW water levels were more stable (overall mean = -0.5 cm, overall minimum = -28 cm) (Fig. 4). Soil temperatures were higher at PD than at PW, reflecting the overall drier conditions (10.3°C at PD and 9.1°C at PW). Due to extremely dry conditions in the summers of 2018 and 2019, the water levels in the ditches varied strongly over the seasons. Ditches at PD repeatedly fell dry in late summer. Generally, the amplitude of water table fluctuations in the ditches was lower in PW.

Nutrient loads in the ditches of PD and PW differed strongly. The two ditches at PD showed significantly higher loads of nitrate (Table 2, \(\chi^2 = 39.95, \text{d.f.}: 1, p < 0.01\)). DOC concentrations in the ditch water were significantly higher in PW compared to PD (Table 2, \(\chi^2 = 38.78, \text{d.f.}: 1, p < 0.01\)). Regardless of the site, ditches with an orthogonal orientation to the drainage direction (PW-o and PD-o) showed DOC concentrations that were almost double those of the ditches with parallel orientation to drainage direction (Table 2). Also, the sediment samples taken in the orthogonal ditches (PD-o, PW-o) showed much higher concentrations of carbon, nitrogen and sulphur (Table 1). In PW the concentrations of phosphate differed significantly between the ditches (\(\chi^2 = 13.35, \text{DF}: 1, p < 0.01\)) with PW-o having higher values than PW-p. DOC concentrations also differed significantly across sites (\(\chi^2 = 33.05, \text{DF}: 1, p < 0.01\)) with ditches at PW having higher values and also significantly higher diffusive \(CH_4\) fluxes (\(\chi^2 = 52.19, \text{DF}: 1, p < 0.01\)). Concentrations of DOC in the groundwater differed significantly between the two sites (\(\chi^2 = 131.26, \text{d.f.}: 1, p < 0.01\)). Average groundwater DOC concentrations at PD were more than four times higher than at PW.
(96.6 ± 44.1 vs. 20.7 ± 19.5 mg l⁻¹). Concentration values of other nutrients in the groundwater can be found in supplementary information (Table T1).

| Table 1 |
| Nutrient contents [% dry weight] of sediment samples in ditches (± denotes one standard deviation, n = 4 per ditch) |
| PD-p | PD-o | PW-p | PW-o |
| C 9.7 ± 4.9 | 29 ± 1.7 | 7.7 ± 3.2 | 38 ± 2.4 |
| N 0.7 ± 0.4 | 1.9 ± 0.1 | 0.6 ± 0.3 | 2.5 ± 0.1 |
| S 0.5 ± 0.4 | 3.8 ± 0.2 | 0.3 ± 0.1 | 1.6 ± 0.2 |

| Table 2 |
| Nutrient concentrations [mg l⁻¹] in the ditch water. (± denotes one standard deviation) |
| PD-p | PD-o | PW-p | PW-o |
| PO₄ 0.1 ± 0.1 | 0.1 ± 0.1 | 0.2 ± 0.1 | 0.5 ± 0.6 |
| NH₄ 0.1 ± 0.1 | 0.4 ± 0.6 | 0.2 ± 0.6 | 0.2 ± 0.4 |
| NO₂ 1.2 ± 1.4 | 0.5 ± 0.5 | 0.1 ± 0.0 | 0.0 ± 0.0 |
| NO₃ 21.3 ± 12.0 | 6.8 ± 11.3 | 0.3 ± 1.00 | 0.5 ± 1.3 |
| SO₄ 120.0 ± 11.6 | 79.8 ± 27.2 | 123.7 ± 48.2 | 41.3 ± 33.5 |
| Br 0.2 ± 0.0 | 0.2 ± 0.0 | 0.2 ± 0.0 | 0.2 ± 0.0 |
| Fl 0.2 ± 0.0 | 0.3 ± 0.1 | 0.3 ± 0.0 | 0.3 ± 0.0 |
| Cl 48.4 ± 2.2 | 47.4 ± 7.3 | 42.3 ± 10.0 | 43.4 ± 13.8 |
| DOC 7.7 ± 5.2 | 12.0 ± 6.3 | 11.4 ± 7.3 | 24.4 ± 5.9 |
| DIC 46.4 ± 8.6 | 36.7 ± 7.7 | 44.3 ± 8.5 | 60.2 ± 13.2 |

**All ditches were strong sources of CH₄ during the measurement period. Diffusive CH₄ fluxes from ditches were generally significantly higher than soil fluxes from the adjacent peatlands. This holds true for both PD and PW (PD: \( \chi^2 = 358.59, \) d.f.: 1, \( p < 0.01 \); PW: \( \chi^2 = 259.66, \) d.f.: 1, \( p < 0.01 \)). Maximum CH₄ fluxes from the ditches were reached during summer months with up to 1469.5 mg m⁻² h⁻¹ for PW and 464.7 mg m⁻² h⁻¹ for PD (Fig. 5). The average diffusive CH₄ fluxes from the ditches were higher in summer than in winter by many orders of magnitude and they differed strongly between different ditches (Fig. 5). Ditches at PW showed much higher diffusive CH₄ fluxes than at PD (75.7 ± 213.3 mg m⁻² h⁻¹ vs. 13.8 ± 37.4 mg..."
\( m^{-2} h^{-1}, \chi^2 = 52.19, \text{d.f.:} 1, p < 0.01, \text{Fig. 5} \). Also, within one site there were large differences between ditches. Orthogonal ditches (PD-o and PW-o) emitted much more CH\(_4\) than parallel ones (PD-p and PW-p) in both sites \( (\chi^2 = 181, \text{d.f.:} 1, p < 0.01, \text{Fig. 6}) \). CH\(_4\) fluxes from soils ranged around or slightly below 0 in PD, while soils in PW were on average a weak source of CH\(_4\) \( (0.5 \pm 1.6 \text{ mg } m^{-2} h^{-1}) \) with a maximum CH\(_4\) flux \( (15.4 \text{ mg } m^{-2} h^{-1}) \) recorded on May 3rd 2018.

**Ebullitive fluxes**

Ebullitive fluxes showed a seasonal pattern with the tendency of high fluxes in summer (maximum CH\(_4\) flux: 23.4 mg m\(^{-2}\) h\(^{-1}\) at PW-p on August 21st 2018) (Fig. 7). The average ebullitive CH\(_4\) flux was significantly higher at PW than at PD \( (7.0 \pm 4.5 \text{ mg } m^{-2} h^{-1} \text{ vs. } 2.7 \pm 2.7 \text{ mg } m^{-2} h^{-1}, \chi^2 = 64.15, \text{d.f.:} 1, p < 0.01) \). However, there were no significant differences between the two ditches within one site \( \text{(PD: } \chi^2 = 2.57, \text{d.f.:} 1, p = 0.11; \text{PW: } \chi^2 = 1.01, \text{d.f.:} 1, p = 0.31; \text{Fig. 8}) \). The bubble rate, estimated with the bubble traps differed significantly only between PW-p and PW-o \( (\chi^2 = 7.83, \text{d.f.:} 1, p < 0.01) \). However, CH\(_4\) concentrations in the gas samples taken from the bubble traps were significantly higher in the ditches at PW than at PD \( (\chi^2 = 86.37, \text{d.f.:} 1, p < 0.01, \text{Fig. 8}) \).

**CH\(_4\) budgets**

Seasonal CH\(_4\) budgets were roughly ten times lower in winter than in summer. In both sites, highest seasonal CH\(_4\) emissions were estimated for ditches in summer 2018 (Table 3). Non-ditch CH\(_4\) emissions in PD were negligible while PW was a weak source. In summer 2018 approximately 9.1 % and 2.5 % of the total ditch CH\(_4\) emissions were transported via ebullition in PD and PW, respectively.

Seasonal CH\(_4\) emissions differed strongly from year to year for both ditches and adjacent peatlands. Especially seasonal soil CH\(_4\) emissions at PW decreased by approximately 90 %, when comparing summer 2018 and summer 2019. Also, the ditch CH\(_4\) emissions declined by 68 %, comparing summer 2018 and summer 2019 at PW while they were comparably stable at PD. Winter CH\(_4\) emissions from the ditches roughly made up between 7.5 and 15 % of the annual ditch CH\(_4\) budgets in PD and PW, respectively.
Table 3

CH\(_4\) budgets in g m\(^{-2}\) per season for each flux subset (ditch and soil). Summer season denotes the period between April 1st and September 30th and the winter season from October to March (± denotes 1 standard deviation).

| Site | type       | Apr. — Sept. 2018 | Oct. — Mar. 2018/19 | Apr. — Sept. 2019 | Oct. — Mar. 2019/20 | Apr. — Sept. 2020 |
|------|------------|-------------------|---------------------|-------------------|---------------------|-------------------|
| PD   | ditch      | 102.5 ± 19        | 8.2 ± 7.5           | 108.8 ± 24        | 14.0 ± 4.3          | 75.6 ± 33         |
| PD   | diffusive  |                   |                     |                   |                     |                   |
| PD   | ebullition  | 10.3 ± 1.5        |                     |                   |                     |                   |
| PD   | soil       | -0.1 ± 0.02       | 0.0 ± 0             | -0.1 ± 0.01       | 0.4 ± 0.2           | 0 ± 0             |
| PD   | diffusive  |                   |                     |                   |                     |                   |
| PW   | ditch      | 919.6 ± 147       | 65.1 ± 17.0         | 293.5 ± 75        | 29.6 ± 14.5         | 123.2 ± 28        |
| PW   | diffusive  |                   |                     |                   |                     |                   |
| PW   | ditch      | 23.0 ± 1.8        |                     |                   |                     |                   |
| PW   | ebulition  |                   |                     |                   |                     |                   |
| PW   | soil       | 4.0 ± 0.7         | 0.04 ± 0.0          | 0.4 ± 0.1         | 0.11 ± 0.0          | 1.3 ± 0.3         |
| PW   | diffusive  |                   |                     |                   |                     |                   |

Ditches in PD and PW covered only 1.52 and 1.49 % of the area, respectively. Still, CH\(_4\) emissions from ditches were of high relevance for the total ecosystem CH\(_4\) budgets (Table 4). Because emissions from ditches greatly exceeded the weak sink or source from peat soils in PD and PW, ditches dominated the total ecosystem CH\(_4\) budgets in both sites.
Table 4

CH$_4$ budgets [g m$^{-2}$ season$^{-1}$] from ditches and soils weighed with their respective spatial share and the relative contributions [%] of the ditch emissions to the combined ecosystem CH$_4$ balance. Weighed seasonal balances of CH$_4$ emissions were multiplied with the spatial share of the respective landscape element (ditch, soil) estimated from an area of 1 km$^2$ around the soil sample locations. Summer denotes the period between April 1st and September 30th and winter the period from October to March.

| PD          | Apr. – Sept. 2018 | Oct. – Mar. 2018/19 | Apr. – Sept. 2019 | Oct. – Mar. 2019/20 | Apr. – Sept. 2020 |
|-------------|-------------------|---------------------|-------------------|---------------------|-------------------|
| Ditch diffusive | 1.6 ± 0.3         | 0.1 ± 0.1           | 1.7 ± 0.4         | 0.2 ± 0.1           | 1.2 ± 0.5         |
| Ditch ebullition | 0.2 ± 0          |                     |                   |                     |                   |
| Soil diffusive   | 0 ± 0             | 0 ± 0               | 0 ± 0             | 0.4 ± 0.2           | 0 ± 0             |
| Combined ecosystem balance | 1.8 ± 0.3       | 0.1 ± 0.1           | 1.7 ± 0.4         | 0.6 ± 0.3           | 1.2 ± 0.5         |
| Ditch contribution | 100              | 100                 | 92                | 33                  | 100               |

| PW          | Apr. – Sept. 2018 | Oct. – Mar. 2018/19 | Apr. – Sept. 2019 | Oct. – Mar. 2019/20 | Apr. – Sept. 2020 |
|-------------|-------------------|---------------------|-------------------|---------------------|-------------------|
| Ditch diffusive | 13.7 ± 1.7        | 1.0 ± 0.3           | 4.4 ± 1.1         | 0.4 ± 0.2           | 1.8 ± 0.4         |
| Ditch ebullition | 3.5 ± 0          |                     |                   |                     |                   |
| Soil diffusive   | 3.9 ± 0.7         | 0 ± 0               | 0.4 ± 0.1         | 0.1 ± 0             | 1.3 ± 0.3         |
| Combined ecosystem balance | 21.1 ± 2.4      | 1 ± 0.3             | 4.8 ± 1.2         | 0.5 ± 0.2           | 3.1 ± 0.7         |
| Ditch contribution | 82 ± 12          | 100 ± 30            | 92 ± 25           | 80 ± 50             | 58 ± 22           |

Drivers of diffusive and ebullitive CH$_4$ emissions

Both approaches of linking the diffusive CH$_4$ fluxes to water chemical parameters or weather variables — using daily averages and direct flux values — revealed that air temperature was the most important factor for explaining seasonal variation at both sites. Looking only at climatic variables that were available for every diffusive flux measurement, air temperature and water depth were significantly correlated with the diffusive CH$_4$ flux at PW but still could only explain 14 % of the overall variability ($R^2 = 0.14$, $F = 24.45$, DF: 3 and 422, $p < 0.01$). At PD air temperature and solar radiation were significantly correlated with diffusive CH$_4$ flux. However, the explanatory power of the multiple linear regression was very low ($R^2 = 0.09$, $F = 12.52$, DF: 3 and 378, $p < 0.01$). Considering all averaged diffusive CH$_4$ fluxes and additional water chemical variables as well as other daily climatic variables regardless of the site, the explanatory power of a multiple linear regression increased to 45 % ($R^2 = 0.45$, $F = 6.12$, DF = 4 and 30, $p < 0.01$). Again, air temperature showed the strongest relationship with diffusive CH$_4$ fluxes. However, also water depth, wind speed and vapour pressure were significantly related. The relationships between the explanatory variables and CH$_4$ emissions held true for PW. At PD none of the explanatory variables was correlated.
with CH$_4$ emissions. Thus, it is likely that single high fluxes at PW dominated the multiple linear regression models. Summary statistics of the regression models can be seen in supplementary information (Table T2).

Nutrient concentrations varied among ditches of different orientation at both sites (Table 2). The much higher nitrate concentrations in PD-p compared to PD-o corresponded to significantly lower diffusive CH$_4$ fluxes in PD-p than in PD-o ($\chi^2 = 82.19$, DF: 1, p < 0.01). In connection with significantly higher diffusive CH$_4$ fluxes ($\chi^2 = 82.19$, DF: 1, p < 0.01). Hence, ditches with higher DOC concentrations consistently exhibited higher diffusive CH$_4$ fluxes. This is the only apparent connection between nutrient loads and diffusive CH$_4$ fluxes. Linear regressions between averaged nutrient concentrations (including DOC) by ditch failed to produce significant relationships (supplementary information, Table T2). Daily diffusive CH$_4$ fluxes averaged across all ditches exhibited a relatively strong relation with vapour pressure (Fig. 9).

Daily ebullitive CH$_4$ fluxes were significantly correlated with air temperature, wind speed and wind direction ($R^2 = 0.51$, $F = 8.29$, DF: 4 and 32, p < 0.01). When averaged by day and across both sites and all four ditches, ebullitive CH$_4$ fluxes exhibited a strong correlation with vapour pressure ($R^2 = 0.53$, $F = 10.15$, DF: 1 and 9, p = 0.01) (Fig. 9). Air pressure as such was not significantly correlated with ebullitive CH$_4$ flux. Correlations between all explanatory variables and ebullitive or diffusive CH$_4$ fluxes are shown in supplementary information (SF1 and SF2).

**Discussion**

**CH$_4$ emissions from ditches and soils**

In general, diffusive CH$_4$ emissions from ditches were much higher than those from the peat soils in both the drained (PD) and the rewetted peatland (PW). This is in line with other studies who find ditches to be important sources of CH$_4$ emissions in peatlands (Schrier-Uijl et al. 2011). The much higher ditch CH$_4$ emissions at PW compared to PD highlight the importance of (former) drainage ditches also for the greenhouse gas balance of rewetted peatlands. Overall, the magnitude of the diffusive ditch CH$_4$ emissions in our study compare well to the few other studies that exist, depending on whether the surrounding peatland was in a rather natural state (PW) or drained for agriculture (PD). Generally, emission rates and maxima of the diffusive fluxes in PW were similar to values reported in the only study on diffusive CH$_4$ emissions from ditches in temperate fens that we are aware of (Peacock et al. 2017), but higher than values from infilled and vegetated ditches in a blanket bog (Cooper et al. 2014). Furthermore, emission rates from PD compare well with Crawford et al. (2016) that studied stream CH$_4$ emissions in an agriculturally used landscape. Generally, seasonal variation of diffusive ditch CH$_4$ emissions in our study was very high with distinct maxima in July and August. Average emissions in summer were 5 to 10 times higher than average winter emissions for PD and PW. Yet, winter emissions were consistent and not negligible.
Ebullitive fluxes showed maxima at both sites in summer 2018, which is in line with the known temperature dependence of CH$_4$ ebullitions (Davidson et al., 2018; DelSontro et al., 2016, Wik et al. 2013). Average CH$_4$ ebullitive fluxes were much lower than values reported from temperate ponds with high nutrient loading (Yang et al., 2020) and small lakes and ponds in the boreal region (DelSontro et al., 2016). Also, CH$_4$ concentrations in the air samples from the bubble traps were much lower compared to studies from temperate or subtropical regions (Maeck et al. 2013; Martinez and Anderson 2013). Instead, CH$_4$ concentrations in the bubbles and the associated ebullitive fluxes were rather comparable with values from subarctic peatlands (Wik et al., 2013).

Unlike the PD ditches with consistent CH$_4$ emissions, PD soil was a small CH$_4$ sink during summer and a small CH$_4$ source during winter 2019/2020 which compares well with comparable studies (van den Pol-van Dasselaar et al. 1998; Nykänen et al. 1995). The soils at PW showed lower average fluxes than found by a previous study at this site (0.6 mg m$^{-2}$ h$^{-1}$, our study; 5–20 mg m$^{-2}$ h$^{-1}$, Huth et al. 2013) and also lower than average values reported for temperate peatlands (4.5 mg m$^{-2}$ h$^{-1}$, Turetsky et al. 2014). It is possible that methanogenesis in the soils was reduced in our study due to the drought conditions in the summers of 2018 and 2019 (Jurasinski et al. 2020).

Drivers of ditch CH$_4$ emissions

Diffusion as well as ebullition fluxes were explained best by temperature, as was also found in other small water bodies recently (Audet et al. 2020). Increased CH$_4$ production is often associated with an increase in temperatures of aquatic environments (Kelly and Chynoweth 1981; Duc et al. 2010). Thus, higher temperatures likely caused the increased CH$_4$ emission rates in summer.

Other variables that have been found to influence CH$_4$ emissions of small water bodies are water depth (Vermaat et al. 2011; West et al. 2016), pH (Ye et al. 2012), or trophic status of the water body (e.g., Peacock et al. 2019). In our study, the deepest ditch showed the highest diffusive and ebullitive fluxes. However, given that the deepest ditch had the highest DOC concentrations it seems more likely that the DOC as a substrate was a more important driver of CH$_4$ emission. Further, we could not find a relationship between pH and CH$_4$ emissions. The range of pH in the ditch water in our study was very small.

Chlorophyll A content seems to also be a good proxy for eutrophication and, thus, for CH$_4$ production and/or emission (e.g. West et al. 2016; DelSontro et al. 2018; Beaulieu et al. 2019). In our study, nitrate, phosphate and DOC contents can be seen as indicators for the trophic status of the ditches. Regression models between averaged CH$_4$ emissions and nutrient/DOC contents failed to produce significant results. However, individual ditches in our study differed substantially in their chemical properties, morphology and in their CH$_4$ emissions. Ditches with increased DOC concentrations (PD-o, PW-p, PW-o, Table 2) showed significantly higher diffusive and ebullitive CH$_4$ emissions, with PW-o standing out with uniquely high diffusive and ebullitive fluxes and DOC concentrations. This has also been shown in previous studies where TOC/DOC concentrations were related with CH$_4$ emissions in ponds and ditches (Crawford and Stanley 2016; Peacock et al. 2019). Interestingly, DOC concentrations in the groundwater were higher
at PD than at PW (supplementary information, Table T1), which is characteristic for drained sites and indicates carbon leaching (Hyvönen et al. 2013). However, the opposite pattern is observed in the ditches (Table 2), which suggests fast carbon turnover in the ditch water.

Our analyses of nutrient concentrations in the ditches allowed for analyzing potential impacts of nutrient loads on CH$_4$ production and emission. For instance, nitrate is known to have an inhibitory effect on CH$_4$ production because it acts as a more favourable electron acceptor when organic substrate is limited (Watson and Nedwell 1998; Audet et al. 2020). The low CH$_4$ emissions in PD-p and PD-o may be explainable by the comparably high nitrate concentrations in these ditches. Diffusive CH$_4$ emissions from ditches in our study, however, were of similar magnitude in a recent study with a comparable nitrate load (Crawford et al. 2016). Sulphate is also known to inhibit methanogenesis (Lovley and Klug 1983, Dean et al. 2018, Zak et al. 2020). Ditches parallel to the general drainage direction (PD-p, PW-p) showed roughly double the sulphate concentration of the ditches running orthogonally to drainage direction. Accordingly, these ditches showed lower diffusive and ebullitive CH$_4$ emissions. It is possible that due to a potentially lower water flow velocity and a longer residence time of the water in the orthogonal ditches, sulphate reduction can be more efficient and thus leads to lower concentrations and less inhibition of CH$_4$ production. However, lower water flow velocity could also lead to lower concentration of oxygen in the water, which is again favourable for methanogenesis. Finally, phosphorus is an indicator for eutrophication and, thus, for production, and therefore seems to be strongly related to increased methanogenesis and/or ebullition (DelSontro et al. 2016). Across both sites, the ditches with higher phosphate concentrations showed higher diffusive and ebullitive CH$_4$ emissions. Generally, it is important to note that the nutrient status can also indirectly influence CH$_4$ emissions through its effects on dominant vegetation (Davidson et al. 2015; Audet et al. 2020).

Other studies have found that water chemical properties are not always good predictors for CH$_4$ concentrations or fluxes in aquatic systems (Ortega et al. 2019). Apart from nutrient and substrate availability in the water column, the nutrient status of the sediment is important for methanogenesis. For instance, sediment accumulation rates are thought to be directly linked to the rate of methanogenesis (Maeck et al. 2013). Both PD-o and PW-o showed higher carbon contents in the sediment than PD-p and PW-p (Table 1). This difference may be driven by higher sediment accumulation rates from slower water flow velocities orthogonally to the general drainage direction. PD-o and PW-o also showed higher diffusive CH$_4$ emissions. Concludingly, nitrate likely inhibited methanogenesis in PD ditches and higher DOC concentrations in PW ditches may have led to higher CH$_4$ emissions (Fig. 10). Here it is interesting to note that DOC concentrations in the groundwater were higher at PD than at PW (supplementary information, Table T1) which is characteristic for drained sites and indicates carbon leaching (Hyvönen et al. 2013). Here, DOC concentrations could have been higher due to the drought conditions in the soil, potentially leading to higher percolation rates. However, the opposite pattern is observable for the ditches (Table 2) for which we have no potential explanation.
Since the residence time of organic matter is likely higher in the orthogonal ditches, the time that is available for matter decomposition and subsequent cycling is also increased. Generally, ditches in PW had higher macrophyte abundance which may also have provided an increased amount of organic matter as substrate for methanogenesis (Davidson et al. 2015).

Importance of ditch emissions for total ecosystem budgets

Since the soils at PD were almost neutral with respect to CH$_4$ emissions, the ecosystem CH$_4$ budget was dominated by emissions from ditches at the drained site. The annual CH$_4$ budgets reported for PWp and PWo were larger than values from other studies (Peacock et al. 2017, Schrier-Uijl et al. 2010). Annual budgets from PDp and PDo compared well to fluxes reported from streams in agriculturally used landscapes (Crawford et al. 2016).

The relative importance of ditch emissions for total ecosystem CH$_4$ budgets was higher than values presented in the scarce studies that exist (Hyvönen et al. 2013). Since the soils at PD were around neutral with respect to CH$_4$ emissions, the ecosystem CH$_4$ budget was dominated almost entirely by emissions from ditches at the drained site. At the rewetted peatland PW, relative contributions by ditches to ecosystem CH$_4$ budgets were still >50%.

Weather conditions were very dry during the study period, especially during the summer months. Thus, CH$_4$ emissions from the soils were likely lower than the long-term mean due to low water levels. Our results support the finding that ditches are hotspots of CH$_4$ emissions and, since extreme weather events are likely to become more frequent, could even gain in importance in the future.

Ebullition contributed <10 % to the overall CH$_4$ budgets of ditches in summer 2018. Many studies report contributions of over 50 % by ebullition to total CH$_4$ emissions (Tokida et al. 2007; Wilcock and Sorrell 2008; Baulch et al. 2011; Vermaat et al. 2011; Martinez-Cruz et al. 2017). However, few studies reported low importance of ebullition of between 10 and 38 % (Minkkinen et al. 1997; Higgins et al. 2008). Moreover, CH$_4$ concentrations in the samples obtained from the bubble traps in our study were low compared to other studies (e.g. Maeck et al. 2013; Martinez and Anderson 2013), while fresh bubbles from the sediment showed higher CH$_4$ concentrations. This could be an indicator for measurement error due to relatively long residence times of the gas inside the traps prior to sampling and equilibration with the water in the trap. If the concentrations of the fresh bubbles only would have been used for the calculation of the ebullition, the contribution of ebullition to the overall CH$_4$ budget would have increased to 14 % and 4 % at PD and PW, respectively. Thus, ebullition would still be of minor importance.

CH$_4$ emissions from ditches and soils continued during winter, although at a much lower rate. Overall, winter emissions contributed between 6 and 11 % of the total annual ditch CH$_4$ emissions. When just looking at the winter season, CH$_4$ budgets consisted almost entirely of emissions from ditches in both sites. With data on this subject being so scarce, future studies should ideally involve measurements of winter emissions.
Conclusions

Ditches can play an important role in the overall GHG budgets of peatlands. Here, we showed that active and inactive ditches in drained and rewetted temperate fens can act as hotspots for CH$_4$ emissions. Despite ditches covering only a small part of the peatlands, the total ecosystem CH$_4$ budget was periodically determined entirely by ditch emissions even in the rewetted fen. Emissions from ditches in the rewetted fen were much higher than from ditches at the drained site. High nitrate concentrations in ditches of the drained fen seemed to reduce CH$_4$ emissions, while increased DOC concentrations in the ditch water seemed to foster CH$_4$ emissions at the rewetted fen. As inactive ditches tend to have higher macrophyte abundance, they can be particularly strong hotspots for CH$_4$ emissions. This must be considered in rewetting projects and filling these ditches must be taken into consideration, although data on the emissions of infilled ditches in temperate fens is lacking. In drained fens high CH$_4$ emissions from ditches add to the high CO$_2$ emissions from the drained peat soils. There, emissions from drainage ditches need to be included in the calculations of GHG budgets.

Declarations

Acknowledgements

The authors would like to acknowledge the support of several student assistants and field technicians for their support in field measurements and laboratory analyses. We would also like to thank Jacob Mulder from the agricultural cooperative Schulenberg for letting us carry out our measurements on his ground. Further we would like to thank the funding from the european social fund for the WETSCAPES project (ESF/14-BM-A55-0030/16).

Funding

Funding was provided by the European Social Fund in the WETSCAPES Project (grant number: ESF/14-BM-A55-0030/16).

Conflicts of interests

The authors declare no conflicts of interest.

Availability of data and material

Data is available via PANGAEA link:

Code availability
R code for analyses is available via PANGAEA link:

**Author contributions**

DK, CW, AG and GJ contributed to the study design. DK and CW collected the data in the field and analysed the data and did laboratory work. DK wrote the first version of the manuscript. AG, CW and GJ added comments which were incorporated by DK. All authors read and approved the final manuscript.

**Ethics approval**

Not applicable.

**Consent to participate**

Not applicable.

**Consent for publication**

Not applicable.

**References**

Aben RCH, Barros N, Van Donk E, et al. (2017) Cross continental increase in methane ebullition under climate change. Nature Communications 8:1–8. doi: 10.1038/s41467-017-01535-y

Audet J, Carstensen MV, Hoffmann CC, et al. (2020) Greenhouse gas emissions from urban ponds in Denmark. Inland waters. doi: 10.1080/20442041.2020.1730680

Bastviken D, Cole J, Pace M, Tranvik L (2004) Methane emissions from lakes: Dependence of lake characteristics, two regional assessments, and a global estimate. Global Biogeochemical Cycles 18:1–12. doi: 10.1029/2004GB002238

Bastviken D, Tranvik LJ, Downing JA, et al. (2011) Freshwater Methane Emissions Offset the Continental Carbon Sink. Science. doi: 10.1126/science.1196808

Baulch HM, Dillon PJ, Maranger R, Schiff SL (2011) Diffusive and ebullitive transport of methane and nitrous oxide from streams: Are bubble-mediated fluxes important? Journal of Geophysical Research: Biogeosciences. doi: 10.1029/2011JG001656

Beaulieu JJ, DelSontro T, Downing JA (2019) Eutrophication will increase methane emissions from lakes and impoundments during the 21st century. Nature Communications 10:1–6. doi: 10.1038/s41467-019-
Crawford JT, Stanley EH (2016) Controls on methane concentrations and fluxes in streams draining human-dominated landscapes. Ecological Applications 26:1581–1591. doi: 10.1890/15-1330

Crawford JT, Stanley EH, Spawn SA, et al. (2014) Ebullitive methane emissions from oxygenated wetland streams. Global Change Biology 20:3408–3422. doi: 10.1111/gcb.12614

Davidson TA, Audet J, Jeppesen E, Landkildehus F (2018) Synergy between nutrients and warming enhances methane ebullition from experimental lakes. Nature Climate Change. doi: 10.1038/s41558-017-0063-z

Davidson TA, Audet J, Svenning JC, et al. (2015) Eutrophication effects on greenhouse gas fluxes from shallow-lake mesocosms override those of climate warming. Global Change Biology 21:4449–4463. doi: 10.1111/gcb.13062

DelSontro T, Beaulieu JJ, Downing JA (2018) Greenhouse gas emissions from lakes and impoundments: Upscaling in the face of global change. Limnology and Oceanography Letters 3:64–75. doi: 10.1002/lol2.10073

DelSontro T, Boutet L, St-Pierre A, et al. (2016) Methane ebullition and diffusion from northern ponds and lakes regulated by the interaction between temperature and system productivity. Limnology and Oceanography 61:S62–S77. doi: 10.1002/lno.10335

Duc NT, Crill P, Bastviken D (2010) Implications of temperature and sediment characteristics on methane formation and oxidation in lake sediments. Biogeochemistry 100:185–196. doi: 10.1007/s10533-010-9415-8

Günther A, Jurasinski G, Albrecht K, et al. (2017) Greenhouse gas balance of an establishing Sphagnum culture on a former bog grassland in Germany. Mires and Peat 20:Article 02. doi: 10.19189/MaP.2015.0MB.210

Higgins TM, McCutchan JH, Lewis WM (2008) Nitrogen ebullition in a Colorado plains river. Biogeochemistry 89:367–377. doi: 10.1007/s10533-008-9225-4

Holgerson MA, Raymond PA (2016) Large contribution to inland water CO₂ and CH₄ emissions from very small ponds. Nature Geoscience 9:222–226. doi: 10.1038/ngeo2654

Huth V, Günther A, Jurasinski G, Glatzel S (2013) The effect of an exceptionally wet summer on methane effluxes from a 15-year re-wetted fen in north-east Germany. Mires and Peat 13:1–7.

Hyvönen NP, Huttunen JT, Shurpali NJ, et al. (2013) The role of drainage ditches in greenhouse gas emissions and surface leaching losses from a cutaway peatland cultivated with a perennial bioenergy crop. Boreal Environment Research 18:109–126.
Jurasinski G, Ahmad S, Anadon-rosell A, et al. (2020) From understanding to sustainable use of peatlands: The WETSCAPES approach. 1–28. doi: 10.20944/preprints202001.0250.v1

Jurasinski G, Koebsch F, Günther A, Beetz S (2014) R package Flux.

Kelly CA, Chynoweth DP (1981) The contributions of temperature and of the input of organic matter in controlling rates of sediment methanogenesis. Limnology and Oceanography 26:891–897.

Laine MPP, Strömmer R, Arvola L (2014) DOC and CO$_2$-C releases from pristine and drained peat soils in response to water table fluctuations: A mesocosm experiment. Applied and Environmental Soil Science. doi: 10.1155/2014/912816

Liikanen A, Martikainen PJ (2003) Effect of ammonium and oxygen on methane and nitrous oxide fluxes across sediment-water interface in a eutrophic lake. Chemosphere 52:1287–1293. doi: 10.1016/S0045-6535(03)00224-8

Lovley DR, Klug MJ (1983) Sulfate reducers can outcompete methanogens at freshwater sulfate concentrations. Applied and Environmental Microbiology 45:187–192. doi: 10.1128/aem.45.1.187-192.1983

Maeck A, Delsontro T, McGinnis DF, et al. (2013) Sediment trapping by dams creates methane emission hot spots. Environmental Science and Technology 47:8130–8137. doi: 10.1021/es4003907

Martinez-Cruz K, Gonzalez-Valencia R, Sepulveda-Jauregui A, et al. (2017) Methane emission from aquatic ecosystems of Mexico City. Aquatic Sciences 79:159–169. doi: 10.1007/s00027-016-0487-y

Martinez D, Anderson MA (2013) Methane production and ebullition in a shallow, artificially aerated, eutrophic temperate lake (Lake Elsinore, CA). Science of the Total Environment 454–455:457–465. doi: 10.1016/j.scitotenv.2013.03.040

McGinnis DF, Greinert J, Artemov Y, et al. (2006) Fate of rising methane bubbles in stratified waters: How much methane reaches the atmosphere? Journal of Geophysical Research: Oceans 111:1–15. doi: 10.1029/2005JC003183

Minkkinen K, Laine J, Nykänen H, Martikainen PJ (1997) Importance of drainage ditches in emissions of methane from mires drained for forestry. Canadian Journal of Forest Research 27:949–952. doi: 10.1139/cjfr-27-6-949

Molongoski J, Klug M (1980) Anaerobic metabolism of particulate organic matter in the sediments of a hypereutrophic lake. Freshwater Biology 10:507–518.

Nykänen H, Alm J, Lang K, et al. (1995) Emissions of CH$_4$ , N$_2$O and CO$_2$ from a Virgin Fen and a Fen Drained for Grassland in Finland. Journal of Biogeography 22:351. doi: 10.2307/2845930
Ortega SH, Romero C, Quijano G, et al. (2019) Methane emissions from contrasting urban freshwaters: Rates, drivers, and a whole-city footprint. Global Change Biology 25:4234–4243. doi: 10.1111/gcb.14799

Peacock M, Audet J, Jordan S, et al. (2019) Greenhouse gas emissions from urban ponds are driven by nutrient status and hydrology. Ecosphere. doi: 10.1002/ecs2.2643

Peacock M, Ridley LM, Evans CD, Gauci V (2017) Management effects on greenhouse gas dynamics in fen ditches. Science of the Total Environment 578:601–612. doi: 10.1023/A:1010372914805

R Core Team (2020). R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. URL https://www.R-project.org/.

Repo ME, Huttunen JT, Naumov A V., et al. (2007) Release of CO$_2$ and CH$_4$ from small wetland lakes in western Siberia. Tellus, Series B: Chemical and Physical Meteorology 59:788–796. doi: 10.1111/j.1600-0889.2007.00301.x

Schrier-Uijl AP, Kroon PS, Leffelaar PA, et al. (2010) Methane emissions in two drained peat agro-ecosystems with high and low agricultural intensity. Plant and Soil 329:509–520. doi: 10.1007/s11104-009-0180-1

Schrier-Uijl AP, Veraart AJ, Leffelaar PA, et al. (2011) Release of CO$_2$ and CH$_4$ from lakes and drainage ditches in temperate wetlands. Biogeochemistry 102:265–279. doi: 10.1007/s10533-010-9440-7

Siegel AF (1982) Robust regression using repeated medians. Biometrika 69:242–244.

Tokida T, Miyazaki T, Mizoguchi M, et al. (2007) Falling atmospheric pressure as a trigger for methane ebullition from peatland. Global Biogeochemical Cycles 21:1–8. doi: 10.1029/2006GB002790

Treat CC, Kleinen T, Broothaerts N, et al. (2019) Widespread global peatland establishment and persistence over the last 130,000 y. Proceedings of the National Academy of Sciences of the United States of America 116:4822–4827. doi: 10.1073/pnas.1813305116

Turetsky MR, Kotowska A, Bubier J, et al. (2014) A synthesis of methane emissions from 71 northern, temperate, and subtropical wetlands. Global Change Biology 20:2183–2197. doi: 10.1111/gcb.12580

van den Pol-van Dasselaar A, Corré WJ, Priemé A, et al. (1998) Spatial Variability of Methane, Nitrous Oxide, and Carbon Dioxide Emissions from Drained Grasslands. Soil Science Society of America Journal 62:810–817. doi: 10.2136/sssaj1998.03615995006200030039x

Vermaat JE, Hellmann F, Dias ATC, et al. (2011) Greenhouse gas fluxes from dutch peatland water bodies: Importance of the surrounding landscape. Wetlands 31:493–498. doi: 10.1007/s13157-011-0170-y
Watson A, Nedwell DB (1998) Methane production and emission from peat: The influence of anions (sulphate, nitrate) from acid rain. Atmospheric Environment 32:3239–3245. doi: 10.1016/s1352-2310(97)00501-3

West WE, Creamer KP, Jones SE (2016) Productivity and depth regulate lake contributions to atmospheric methane. Limnology and Oceanography 61:S51–S61. doi: 10.1002/lno.10247

Wik M, Crill PM, Varner RK, Bastviken D (2013) Multiyear measurements of ebullitive methane flux from three subarctic lakes. Journal of Geophysical Research: Biogeosciences 118:1307–1321. doi: 10.1002/jgrg.20103

Wilcock RJ, Sorrell BK (2008) Emissions of greenhouse gases CH4 and N₂O from low-gradient streams in agriculturally developed catchments. Water, Air, and Soil Pollution 188:155–170. doi: 10.1007/s11270-007-9532-8

Yang P, Zhang Y, Yang H, et al. (2020) Ebullition was a major pathway of methane emissions from the aquaculture ponds in southeast China. Water Research 184:116176. doi: 10.1016/j.watres.2020.116176

Ye R, Jin Q, Bohannan B, et al. (2012) pH controls over anaerobic carbon mineralization, the efficiency of methane production, and methanogenic pathways in peatlands across an ombrotrophic-minerotrophic gradient. Soil Biology and Biochemistry 54:36–47. doi: 10.1016/j.soilbio.2012.05.015

Zak D, Hupfer M, Cabezas A, et al. (2021) Sulphate in freshwater ecosystems: A review of sources, biogeochemical cycles, ecotoxicological effects and bioremediation. Earth-Science Reviews. doi: 10.1016/j.earscirev.2020.103446

Zhou Y, Zhou L, Zhang Y, et al. (2019) Autochthonous dissolved organic matter potentially fuels methane ebullition from experimental lakes. Water Research 166:1–12. doi: 10.1016/j.watres.2019.115048

**Figures**
Figure 1

Location of the study area in Germany and measurement locations for soil and ditch emissions at the drained (PD) and the rewetted fen (PW). Note: The designations employed and the presentation of the material on this map do not imply the expression of any opinion whatsoever on the part of Research Square concerning the legal status of any country, territory, city or area or of its authorities, or concerning the delimitation of its frontiers or boundaries. This map has been provided by the authors.
Figure 2

Setup of the floating chamber which was used to measure diffusive CH4 emissions.
**Figure 3**

Setup of the bubble trap for measurements of ebullitive CH4 emissions.

**Figure 4**

Seasonal course of a) daily precipitation, b) daily mean air temperature and c) groundwater level at PD and PW. At c) the solid line depicts PD and the dashed line depicts PW. Daily precipitation and daily mean air temperature were averaged between PD and PW, since differences were not observable due to the low distance of the sites.
Figure 5

Seasonal course of diffusive CH4 fluxes from ditches at PD (a) and PW (b), and soil surface at PD (c) and soil surface at PW (d). Note the differing y-axes between upper and lower panels.
Boxplot of log-transformed diffusive CH4 emissions by ditch. Kruskal-Wallis tests were used to test for differences between the ditches.
Figure 7

Bubble rate recorded at the bubble traps at PD (a) and PW (b) and CH4 concentrations in air samples from the bubble traps at PD (c) and PW (d). Crosses indicate the CH4 concentrations of individual fresh bubble samples.
Boxplot of ebullitive CH4 emissions by ditch. Kruskal-Wallis tests were used to test for differences between the ditches.
Figure 9

Linear regression between vapour pressure and a) daily averaged diffusive CH4 fluxes and b) daily averaged ebullitive CH4 fluxes. Vapour pressure data was obtained from three weather stations of the German weather service (DWD) located within 40 km to the NW, N and E. Vapour pressure values from all weather stations were averaged.
Figure 10

Qualitative diagram of the most important factors influencing CH4 emissions in ditches.

Supplementary Files

This is a list of supplementary files associated with this preprint. Click to download.

- SupplementaryinformationSF1correlationMatrix.pdf
- SupplementaryinformationSF2correlationMatrix.pdf
- SupplementaryinformationT1.docx
- SupplementaryinformationT2.docx