METHODOLOGY

Improved volume variable cluster model method for crystal-lattice optimization: effect on isotope fractionation factor

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Abstract

The isotopic fractionation factor and element partition coefficient can be calculated only after the geometric optimization of the molecular clusters is completed. Optimization directly affects the accuracy of some parameters, such as the average bond length, molecular volume, harmonic vibrational frequency, and other thermodynamic parameters. Here, we used the improved volume variable cluster model (VVCM) method to optimize the molecular clusters of a typical oxide, quartz. We documented the average bond length and relative volume change. Finally, we extracted the harmonic vibrational frequencies and calculated the equilibrium fractionation factor of the silicon and oxygen isotopes. Given its performance in geometrical optimization and isotope fractionation factor calculation, we further applied the improved VVCM method to calculate isotope equilibrium fractionation factors of Cd and Zn between the hydroxide (Zn–Al layered double hydroxide), carbonate (cadmium-containing calcite) and their aqueous solutions under superficial conditions. We summarized a detailed procedure and used it to re-evaluate published theoretical results for cadmium-containing hydroxyapatite, emphasizing the relative volume change for all clusters and confirming the optimal point charge arrangement (PCA). The results showed that the average bond length and isotope fractionation factor are consistent with those published in previous studies, and the relative volume changes are considerably lower than the results calculated using the periodic boundary method. Specifically, the average Si–O bond length of quartz was 1.63 Å, and the relative volume change of quartz centered on silicon atoms was −0.39%. The average Zn–O bond length in the Zn–Al-layered double hydroxide was 2.10 Å, with a relative volume change of 1.96%. Cadmium-containing calcite had an average Cd–O bond length of 2.28 Å, with a relative volume change of 0.45%. At 298 K, the equilibrium fractionation factors between quartz, Zn–Al-layered double hydroxide, cadmium-containing calcite, and their corresponding aqueous solutions were

\[ \delta^{30}\text{Si}_{\text{Qtz}} = 2.20\% \]
\[ \delta^{18}\text{O}_{\text{Qtz}} = 36.05\% \]
\[ \Delta^{66/64}\text{Zn}_{\text{Zn–Al LDH}} = 1.12\% \]
\[ \Delta^{114/110}\text{Cd}_{\text{Cd–Cal}} = -0.26\% \]

respectively. These results strongly support the reliability of the improved VVCM method for geometric optimization of molecular clusters.

Keywords: Molecular cluster, VVCM, Geometric optimization, Relative volume change, Isotopic equilibrium fractionation factor

Introduction

In nature, the isotopic composition of an element varies with rock type, source, and age, especially those determined by high-resolution mass spectroscopy since the 2000s. This has considerably expanded the usefulness of
Research on solid earth science involves the geometric optimization of solid structures. The periodic boundary method is the most commonly used method for simulating solids. Mineral cells are used to construct periodic systems. In general, supercells are required to ensure that the constructed system is comparable to the real mineral lattice, but the total number of unit cells is limited by computational power. In addition, improper handling of atoms at the boundary of the supercells may considerably affect the accuracy of the calculated results. Given the shortcomings of the periodic boundary method, Rustad and Dixon [17], Rustad and Yin [18], and Rustad et al. [19, 20] introduced an embedded cluster model with a Pauling bond strength (PBS)-conserving termination of the atoms in their rigid shell. In this model, a large molecular cluster with three shells is employed to represent the minerals. A small core and a second shell are optimized at high and low levels of theory and basis sets, respectively. The third shell is fixed at the measured lattice positions. Chemical bonds entering the third shell are clipped and replaced by “Pauling bond strength-conserving quasi-atoms” along the clipped bonds, which make the core charge of each quasi-atom equal to the Pauling bond valence, that is, the charge of the removed cation divided by its coordination number. The distance between the quasi-atom and the third shell is maintained at 1 Å. Finally, the frequency calculations are performed only for the small core. As an alternative, the volume variable cluster model (VVCM) method was proposed by Liu [21], which was developed during the optimization of silicates, carbonates, oxides, hydroxides, and sulfides [22–28]. In contrast to the embedded cluster model, in the VVCM model all atoms of the cluster are geometrically optimized, and all or partial atoms are involved in the calculations of the vibrational frequency according to the computational cost. The total number of background point charges amounts to several hundreds, which is much higher than the number of quasi-atoms in the embedded cluster model. This ensures the coverage of point charges on the entire molecular cluster. Additionally, the positions of the background point charges are adjusted during the construction of the initial molecular clusters. More importantly, geometric optimization and harmonic vibrational frequency calculations can be performed for both periodic systems (minerals) and nonperiodic systems (solutions) with identical exchange–correlation functionals and basis sets. Although the main two limitations of the molecular cluster method are the same as those of the periodic boundary method (the size of the system and the treatment of atoms on the boundary) and its representation of the periodic mineral lattice is not as accurate as that of the periodic boundary method, it is able to precisely treat the local configuration around the atom of interest, which is the dominant factor controlling isotope
fractionation. In recent years, static quantum chemistry calculations based on molecular clusters have accurately calculated the stable isotope equilibrium fractionation factors [22–32]. The calculated results are in agreement with previous experimental and theoretical calculations. However, the critical question is to treat the local configuration around the atom of interest. In previous studies, a detailed procedure to obtain an optimized local configuration was absent. In addition, evaluation proxies for the optimized local configuration are not well-established. The energy of a solid (E) is a function of the molecular volume (V) [33], total number of atoms, and interactions among the atoms. Additionally, the relationship between energy and volume is linear and/or nonlinear [34]. For example, the nonlinear four-parameter Vinet equation of state [35, 36] can be expressed as follows:

\[
E(V) = a - \frac{4B_0V_0}{(B_0' - 1)^2} \left\{ 1 - \frac{3}{2}(B_0' - 1) \left( 1 + \frac{\Delta V}{V_0} \right)^{1/3} \right\} \times \exp \left\{ \frac{3}{2}(B_0' - 1) \left[ 1 - \left( 1 + \frac{\Delta V}{V_0} \right)^{1/3} \right] \right\}
\]

where the fitting parameter \(a = E_0 + 4B_0V_0/(B_0' - 1)^2\), \(V_0\), \(B_0\), and \(B_0'\) represent the equilibrium volume, bulk modulus, and its first derivative with respect to pressure, respectively; and \(\Delta V\) is the deviation from the equilibrium volume.

Given the role of molecular volume in the energy of the solid, the expansion or contraction of molecular clusters will influence the accuracy of harmonic vibrational frequencies and the isotope fractionation factor. However, the degree of expansion or contraction of molecular clusters was not considered in the previous molecular cluster modeling method.

In this study, we chose the typical oxide quartz (Qtz) present on the Earth’s surface [37, 38] as an example. By geometrically optimizing the Qtz molecular cluster, extracting its parameters (converged electronic energy, average bond length, and relative volume change), and calculating the equilibrium oxygen and silicon isotopic fractionation factors between molecular clusters and the aqueous solution, we aimed to verify the usefulness of the improved VVCM method after comparison of methods developed in previous studies. Based on the performance of the improved VVCM method for Qtz, we extended the method to the calculation of other nontraditional stable isotopes (Zn and/or Cd) between layered double hydroxide (Zn–Al LDH), carbonate (Cd–containing calcite, Cd–Cal), and their corresponding solutions. After rechecking previous theoretical results for cadmium-containing hydroxyapatite (Cd–HAp), emphasizing the relative volume change for all clusters, and confirming the optimal point charge arrangement (PCA), we constructed a detailed procedure to implement the improved VVCM method.

**Theory and methods**

To obtain an optimal geometric structure, we systematically used the optimization process based on the improved VVCM method, as shown in Fig. 1.

**Constructing molecular cluster and adding background point charges**

We reconstructed the three-dimensional crystal lattices of the minerals from X-ray single-crystal diffraction or neutron diffraction data obtained from previous studies [39–43]. The N–N–N principle must be followed when cutting molecular clusters. This requires the atoms (nearest neighbor) to form chemical bonds with the central atom (atom of interest), and the atoms (the second neighbor atom) forming chemical bonds with the nearest neighbor must be retained. In specific cases, the size of the molecular clusters depends on the computational cost and accuracy. Generally, the outermost layer of a cluster is composed of anions or anionic groups, while the metal ions and/or protons connected to these anions or anionic groups are deleted. The hanging bonds are typically saturated with positive point charges to fix the outermost atoms and to simulate the electron-neutral environment of the crystal. Existence of free interlayer anions makes LDH different from others. The general formula of LDH is \([\text{M}^{2+}_{1-x}\text{M}^{3+}_x(\text{OH})_2\text{x}^{-}\text{An}^{n-}_x\text{mH}_2\text{O}\), where \(\text{M}^{2+}\) and \(\text{M}^{3+}\) represent divalent and trivalent metal cations, respectively; \(\text{An}^{n-}\) is an interlayer anion including \(\text{CO}_3^{2-}\), \(\text{Cl}^-\), \(\text{SO}_4^{2-}\), \(\text{NO}_3^-\); \(x\) represents the molar ratio \(\text{M}^{3+}/(\text{M}^{2+} + \text{M}^{3+})\); and \(m\) is the number of water molecules [44, 45]. In nature, the \(\text{M}^{2+}/\text{M}^{3+}\) molar ratio is usually two [46]. Therefore, we focused on a typical Zn-containing LDH with a Zn:Al ratio of two. We constructed two kinds of molecular clusters for LDH without and with interlayer anions.

Oxygen atom can strongly attract the electron of hydrogen atom due to its strong electronegativity. It means that each hydrogen atom of a water molecule can accept a non-shared electron pair of the oxygen atom. Meanwhile, oxygen atom can contribute its two non-shared electron pairs to two hydrogen atoms. Therefore, water molecules can form 4 H bonds with each other in solution [47], which constructs a tetrahedron structure found in
ice crystals (Additional file 1: Figure S1). Inspired by this observation, we used point charges and connected chemical bonds to encompass the terminal anions. Different terminal anions have various arrangements of positive point charges. In a previous study, three PCAs were proposed [25], and for the terminated oxygen (η\textsuperscript{1}-O) capped by five positive point charges (5 \times), two positive point charges were used as two wings (2 \times) distributed on the bridge oxygen (\mu_2\textsuperscript{-}O) or the terminated hydroxyl (\eta_1\textsuperscript{-}OH), and one positive point charge was set to the apex (1 \times) of the bridge oxygen or hydroxyl (\mu_3\textsuperscript{-}O or \mu_2\textsuperscript{-}OH) (Fig. 2). Considering \eta_1\textsuperscript{-}O as the predominant anion on the surface of the cluster, we added point charges in four different ways to evaluate their effect on the theoretical results. The one was to add one point charge in the direction of the deleted atoms [17–20]. In other three ways, three, four, and five point charges were added to \eta_1\textsuperscript{-}O, respectively. However, point charges were not fixed at the directions of chipped bonds. Alternatively, these point charges were averagely distributed in the space surrounding the \eta_1\textsuperscript{-}O. Obviously, \eta_1\textsuperscript{-}O is more tightly capped by five point charges than by three or four point charges (Additional file 1: Figure S1). Furthermore, we performed geometrical optimizations for the Qtz cluster considering 1 \times, 3 \times, and 5 \times PCAs (Table 1). Although the average Si–O bond length optimized in three PCAs was slightly larger than that in the experimental data (1.61 Å [39] and 1.62 Å [48]), the more accurate relative volume change (\Delta V/V_0) and \Delta^{30/28}\text{Si}_{Qtz-H_2SiO_4} indicated that 5 \times was the best PCA for \eta_1\textsuperscript{-}O. Similarly, we safely capped \mu_2\textsuperscript{-}O and \mu_3\textsuperscript{-}O with two and one point charges, respectively, according to a previous study [25].

Before optimizing the configuration, we set the charge quantity and distance of the point charges of the
outermost oxygen atoms. The quantity of charge was equal to the amount of charge provided by the outermost anions or anionic groups of the cluster. Specifically, the total amount of point charge on each terminal anion was equal to the sum of the contributions of all “cut” metal ions and/or protons. The distance between the point charge and the outermost oxygen atom was set empirically. Upon adjustment of the distance between the point charge and the outermost oxygen atom, the electronic energy of the system also changed, and each atom of the inner part of the cluster relaxed under its three freedoms (X, Y, and Z dimensions). When the inner atoms amounted to N, 3 N freedoms were observed. The outermost layers of the Qtz clusters centered on silicon or oxygen were all η 1-O (Fig. 2). In the Qtz cluster, the total electrical charge of the point charges on each η 1-O was equal to 1/4 of that of the removed Si$^{4+}$, that is, $1/4$. We set the electric quantity of each point charge to $1/30$. The total charge on µ 2-OH was equal to the removed charge of one hexa-coordinated Al$^{3+}$ ($3/6$). The electric quantity of each point charge was $1/2$. Furthermore, for Cd–Cal, the two PCA modes were 5 × and 2 × (Fig. 2). The total charge on η 1-O was equal to the total amount of charge of the two removed hexa-coordinated Ca$^{2+}$ ($2/6 + 2/6$), so the amount of charge divided by each point charge was $+(2/6 + 2/6)/5$. The total amount of charge on µ 2-O was +2/6, and the electric quantity of each point charge was $+ (2/6)/2$.

The configuration of Cd-containing hydroxyapatite is further complicated by the simultaneous existence of η 1-O, µ 2-O, and µ 3-O in its outermost layer (Fig. 2). Two ways can be used to set the electric quantity for each point charge in the 5 × mode. One fraction of η 1-O obtains the total charge of two nine-coordinated Ca$^{2+}$ and one seven-coordinated Ca$^{2+}$ ($2 	imes 2/9 + 2/7$), setting the electric quantity of each point charge to $+(2 	imes 2/9 + 2/7)/5$. The other fraction of η 1-O obtains the total charge ($2 	imes 2/7 + 2/9$) of two seven-coordinated Ca$^{2+}$ and one nine-coordinated Ca$^{2+}$, with the electric quantity of each point charge set to $+(2 	imes 2/7 + 2/9)/5$. The three ways to set the electric quantity for each point charge when adding two points of charge are as follows: (1) the µ 2-O

| Optimal PCA (pm) | Si–O bond length (Å) | ΔV/V₀ (%) | $\Delta^{30/28}{Si_{Qtz} - H_4SiO_4}$ (298 K, %) |
|------------------|---------------------|-----------|-----------------------------------------------|
| 1 × 94           | 1.630               | 2.03      | 0.78                                          |
| 3 × 111          | 1.629               | 1.00      | 1.76                                          |
| 5 × 116          | 1.625               | −0.39     | 2.02                                          |
Table 2  PCA on the anion or anion group in the outermost layer of a molecular cluster

| Cluster      | Molecular formula | Total valence | 1 × | 2 × | 5 × | Net charge |
|--------------|-------------------|---------------|-----|-----|-----|------------|
| Si–Qtz       | Si24O66           | – 36          | –   | –   | –   | 180 × (4/4) × 1/5 | 0 |
| O–Qtz        | Si29O77           | – 38          | –   | –   | –   | 190 × (4/4) × 1/5 | 0 |
| Zn–Al LDH    | Zn13Al12O60C4H30  | – 39          | –   | –   | –   | 90 × (1 + 3/6 + 2/6) × 1/5 | 0 |
| Cd–Cal       | Ca6Cd(CO3)26      | – 60          | 12 × 3/6 | –   | –   | 210 × (2/6 + 2/6) × 1/5 | 0 |
| Cd–HAp       | Ca12Cd(PO4)30(OH)3 | – 67          | 2 × 2/7 | 10 × 2/9 | –   | 150 × (2 × 2/9 + 2/7) × 1/5 | 0 |

The outermost layer of the Si–Qtz molecular cluster has 5 × point charges (180). The outermost layer of the O–Qtz molecular cluster has 5 × point charges (190). The outermost layer of the Zn–Al LDH molecular cluster has 1 × point charges (12) and 5 × point charges (90). The outermost layer of the Cd–Cal molecular cluster has 2 × point charges (60) and 5 × point charges (210). The outermost layer of the HAp molecular cluster has 5 × (2), b1 × (10), b2 × (48), b5 × (40), b5 × (15), and b5 × (40) point charges.

Si–Qtz net charge (0) = total cluster valent (-36) + total point charge (+36). O–Qtz net charge (0) = total cluster valent (-38) + total point charge (+38). Zn–Al LDH net charge (0) = total cluster valent (-39) + total point charge (+39). Cd–Cal net charge (0) = total cluster valent (-38) + total charge of point charge (+38). Cd–HAp net charge (0) = total cluster valent (-67) + total charge of point charge (+67).
obtains the total charge of two nine-coordinated \( \text{Ca}^{2+} \) \((2/9 + 2/9)/2\); (2) the \( \mu_2\)-O obtains the total charge of two seven-coordinated \( \text{Ca}^{2+} \) \((2/7 + 2/7)/2\) and the electric quantity of each point is set to \((2/7 + 2/7)/2\); and (3) the \( \mu_2\)-O obtains a total charge of a seven-coordinated \( \text{Ca}^{2+} \) and a nine-coordinated \( \text{Ca}^{2+} \) \((2/7 + 2/9)/2\) and the electric quantity of each point is set to \((2/7 + 2/9)/2\). The two types of electric quantities for each point charge in the \(1 \times \) mode are as follows: the electric quantity of each point charge on \( \mu_3\)-O is set to \((2/7)\) when the seven-coordinated \( \text{Ca}^{2+} \) is removed; and the point charge on \( \mu_3\)-O obtains a charge of \((2/9)\) when the nine-coordinated \( \text{Ca}^{2+} \) is removed.

Molecular cluster optimization

We systematically adjusted the distance between the point charge and outermost oxygen atom to change the distribution of the background charge and to change the molecular cluster volume and geometry.

Qtz

Only one PCA exists in the Qtz molecular cluster, that is, \(5 \times \) PCA. Here, we considered the Qtz molecular cluster centered on a silicon atom as an example to provide a detailed description. According to previous studies, the range of distances between the \(5 \times \) point charge and \( \eta_1\)-O \((d_{5x})\) is approximately 100–120 pm \([22]\). Thus, we adjusted \(d_{5x}\) to 100 pm. Each time, we placed the point charge with two sets of distances, with a difference of only 1 pm, for instance, 100 and 101 pm. We determined the direction of further distance adjustment by determining the lower electronic energy (corresponding to 101 pm) in the two groups. Therefore, increasing the distance between the \(5 \times \) point charge and \( \eta_1\)-O decreased the electronic energy of the cluster (Fig. 3). Subsequently, we successively adjusted \(d_{5x}\) until we determined the optimal value. For the Zn–Al LDH and Cd–Cal molecular clusters, we determined the optimal \(d_{1x}\) or \(d_{2x}\) after fixing \(d_{5x}\). However, confirming the most suitable distance \(d_{2x}\) \((d_{1x})\) in the Cd–HAp molecular cluster requires \(d_{1x}\) \((d_{2x})\) and \(d_{5x}\) to be both fixed.

In this study, we used the exchange–correlation functional B3LYP for Qtz and Zn–Al LDH, and BP86 for Cd–Cal and Cd–HAp. We used the all-electronic basis set 6–311G \((d)\) \([49, 50]\) to describe the wavefunction of the Qtz molecular cluster and a mixed basis set to describe the wavefunction of other systems: LanL2DZ for Zn \([51–53]\); LanL2MB for Cd \([52]\); 6–311 + G \((d, p)\) for Al, O, and H in Zn–Al LDH; and 6–311G \((d)\) for C, Ca, P, O, and H in Cd–Cal and Cd–HAp. We also performed harmonic vibrational frequency analyses to check whether there was an imaginary frequency and to ensure that the most stable configuration corresponded to at least the local minima on the potential electronic energy surface. We calculated electronic energy and performed geometrical optimization and harmonic vibrational frequency analyses using Gaussian09 code \([54]\).

Relative volume change

We calculated the relative volume change \((\Delta V/V_0)\) to quantify the expansion or contraction of the molecular
clusters during geometric optimization. By comparing the relative volume change (relative to the volume of the unoptimized molecular cluster) and the system electronic energy, we screened the configuration with little volume change with respect to the natural mineral lattice and the lowest electronic energy. Only this method could ensure that the simulation matched the real scenario to the largest extent. Because a single molecular cluster is not a unit cell or supercell of a mineral, accurately calculating the molecular cluster volume using cell parameters is impossible. The wavefunction or electronic density of a molecule is a function of space points. Moreover, the Monte Carlo method can be used to calculate the molecular cluster volume (molecular van der Waals volume), which is the space enclosed by an isosurface with an electron density of $x \text{e/Bohr}^3$, where $x$ represents the electron density value of the isosurface. We set up a rectangular box with volume $V$ for molecular volume integration. Subsequently, we placed many random points ($m$) in the box for sampling and tested the electron densities at these space points to determine if they were larger than $x \text{e/Bohr}^3$. Finally, we counted the number of points ($n$) with electron densities larger than $x \text{e/Bohr}^3$, for which the van der Waals volume of the molecule was $n/m*V$ \([55, 56]\). Bader et al. \([57]\) proposed a more accurate definition of the van der Waals volume, stating that if a molecule is in the gas phase, an isosurface with an electron density of 0.001 e/Bohr$^3$ is considered a van der Waals surface. Such surfaces typically contain more than 98% the electron density of a molecule. For molecules in a condensed state, intermolecular extrusion and various forms of interaction cause the van der Waals surface to penetrate. Using an isosurface with an electron density of 0.002 e/Bohr$^3$ as the van der Waals surface is generally recommended. Among the molecular clusters constructed in this study, Cd–HAp had the largest number of atoms (no more than 169 atoms). Hence, we treated the molecular clusters as small molecules in the gas phase and used an isosurface with an electron density of 0.001 e/Bohr$^3$ as the van der Waals surface of the molecular clusters. Cd–HAp was studied by He et al. \([25]\), but they did not select the optimal PCA, by strictly adhering to our procedure. We rechecked the optimal PCA for this mineral, emphasizing the relative volume changes for all clusters.

Calculating the van der Waals volume requires the molecular wavefunctions. We obtained the wavefunction files using Gaussian09 code \([54]\). By inputting $i$, $x$, and $k$ values, we calculated the volume using the Multiwfn program \([58]\), where $i$ is a proxy for the number of random points distributed in the box and $k$ is the space reserved by the box around the molecule. Alternatively, the shortest distance between the outermost atom and the boundary of the box is $k$ times the van der Waals radius of the atom. For specific calculations, we distributed $100 \times 2^i$ random points in a box and set the value of $k$ to 1.7. We used an isosurface with an electron density of 0.001 e/Bohr$^3$ to envelop the molecule. Normally, to obtain a more accurate molecular van der Waals volume, the density of random points in the box must not be less than 2000 points/Bohr$^3$. We used the following formula for these calculations:

$$\Delta V/V_0 = \frac{V_{\text{opt}} - V_0}{V_0} \times 100\%$$  \(2\)

where $V_{\text{opt}}$ denotes the optimized volume and $V_0$ is the unoptimized volume. A positive $\Delta V/V_0$ indicates expansion of the molecular cluster, whereas a negative $\Delta V/V_0$ suggests contraction of the molecular cluster.

Isotopic equilibrium fractionation factor

The Bigeleisen–Mayer equation and Urey model \([59, 60]\) provide the equilibrium constants for the isotopic exchange reaction, laying the foundation for the theory and calculation of stable isotope geochemistry. Consider the following isotope exchange reaction as an example:

$$AX' + X = AX + X'$$  \(3\)

where $AX$ and $AX'$ represent compounds with heavy and light isotopes, respectively; $X$ and $X'$ denote the ideal single-atom gas with heavy and light isotopes, respectively. The reaction equilibrium constant $K$ can be obtained as follows from the ratio of the partition functions of the products and reactants:

$$K = \frac{Q_{\text{trans}}^{AX}Q_{\text{rot}}^{AX}Q_{\text{vib}}^{AX}Q_{\text{elec}}^{AX}}{Q_{\text{trans}}^{AX'}Q_{\text{rot}}^{AX'}Q_{\text{vib}}^{AX'}Q_{\text{elec}}^{AX'}}$$  \(4\)

where $Q_{\text{trans}}$ is the transitional partition function, $Q_{\text{rot}}$ is the rotational partition function, $Q_{\text{vib}}$ is the vibrational partition function, and $Q_{\text{elec}}$ is the electronic partition function \([61]\). For light elements, the difference in the ground-state electronic energy of the compound before and after isotope substitution is indistinguishable, so $Q_{\text{elec}}$ can be ignored. The translational, rotational, and vibrational partition functions can be respectively written as follows:

$$Q_{\text{trans}} = V \left(\frac{2\pi MkT}{\hbar^2}\right)^{3/2}$$  \(5\)

$$Q_{\text{rot}} = \frac{\pi^{1/2}(8\pi^2 kT)^{3/2}(I_AI_BI_C)^{1/2}}{\sigma \hbar^3}$$  \(6\)
where $V$ is the volume, $M$ is the mass, $I_A$ is the moment of inertia around the axis $A$ of rotation, $\sigma$ is the symmetry number of the molecule, and $u_i = \hbar c \omega_i / kT$, where $\hbar$ represents the Planck constant, $k$ is the Boltzmann constant, $T$ is the temperature in degrees Kelvin, $c$ is the speed of light, and $\omega$ is the harmonic frequency in cm$^{-1}$. The Teller-Redlich [62] approximation is done as follows:

$$Q_{vib} = \prod_i \frac{e^{-u_i/2}}{1 - e^{-u_i}}$$  \hspace{1cm} (7)

Therefore, the ratio of the translational partition function and the ratio of the rotational partition function can be expressed by the vibrational frequency. Through algorithm transformation, the reduced partition function ratio (RPFR($AX/AX'$)) is defined as follows:

$$\text{RPFR}(AX/AX') = \frac{\sigma}{\sigma'} K = \prod_i \frac{u_i}{u'_i} \left( \frac{e^{-u_i/2}}{e^{-u'_i/2}} \right) \left( \frac{1 - e^{-u_i}}{1 - e^{-u'_i}} \right)$$  \hspace{1cm} (8)

When only one atom is to be exchanged in the molecular cluster, the RPFR is equivalent to the $\beta$ factor [61], which is the isotopic fractionation factor between a compound and the ideal atomic gas [63]. Given the $\beta$ factors of a pair of compounds or minerals in equilibrium, the isotope fractionation factor $\alpha$ can be derived from the ratio of their $\beta$ factors: $\alpha = \beta_A / \beta_B$, where $A$ and $B$ denote different compounds or minerals. The isotopic equilibrium fractionation between the two phases can then be obtained as follows:

$$\Delta_{A-B} \approx 10^3 \ln \alpha_{A-B} = 10^3 (\ln \beta_A - \ln \beta_B)$$  \hspace{1cm} (10)

**Results**

**Influence of $d_{5x}$, $d_{2x}$, and $d_{1x}$ on system electronic energy**

We collected the electronic energy values of the four systems and performed polynomial fitting. The results showed that the system electronic energy had a good parabolic relationship with $d_{5x}$, $d_{2x}$, and $d_{1x}$, and the $R^2$ value ranged from 0.60 to 0.99 (Figs. 3, 4, 5, 6). The PCA modes corresponding to the most stable configurations of the molecular clusters were $5 \times 116$, $5 \times 116$, $5 \times 111/1 \times 76$, $5 \times 123/2 \times 122$, and $5 \times 123/2 \times 119/1 \times 115$ pm for Si–Qtz, O–Qtz, Zn–Al LDH, Cd–Cal, and Cd–HAp, respectively.

**Effect of $d_{5x}$, $d_{2x}$, and $d_{1x}$ on relative volume change**

During the geometric optimization process, the volume of the Si–Qtz molecular cluster decreased, with $\Delta V/V_0$ ranging from $-2.05$ to $-0.11\%$ (Fig. 7). We observed an extremely strong linear correlation between $\Delta V/V_0$ and $d_{5x}$ ($R^2 = 0.9895$). $\Delta V/V_0$ gradually increased with an increase in $d_{5x}$. We recorded the $\Delta V/V_0$ value as $-0.39\%$ when the Si–Qtz molecular cluster reached the optimal structure. However, the volume of the Zn–Al LDH molecular cluster without interlayer anions
expanded during the optimization process, with $\Delta V/V_0$ ranging from 3.17 to 3.54% (Fig. 8). The relative volume change was 3.54% when the system reached its lowest electronic energy configuration ($5 \times 111/1 \times 76$ pm). But when interlayer anions were introduced, the relative volume change reached 1.96% for the lowest electronic energy configuration ($5 \times 111/1 \times 76$ pm). As shown in Fig. 9, the volume of Cd–Cal expanded during the optimization process ranging from 0.24 to 0.50%. When the structure reached the maximum stability ($5 \times 123/2 \times 122$ pm), the volume expansion was 0.45% (Table 4). The volume of the Cd–HAp molecular clusters also expanded during the optimization process, with a volume expansion of 0.64–0.92%. We found...
that $\Delta V/V_0$ was 0.74% when the system reached the most stable configuration ($5 \times 123/2 \times 119/1 \times 115$ pm) (Fig. 10), which is identical to the 0.7% reported by He et al. [25].

**Discussion**

**Geometrically optimized configuration**

In this study, we obtained an average Si–O bond length in the Qtz cluster with the lowest electronic energy of 1.63 Å (Table 3). Purton et al. [64] used ab initio Local Density Functional (LDF) theory to study the structure of α–quartz, and obtained an average Si–O bond length of 1.62 Å. He and Liu [22] calculated an average Si–O bond length of 1.61 Å on the 6–311G (2df) level. Hazen et al. [39] and Kihara [48] obtained the average Si–O bond lengths of 1.61 Å and 1.62 Å, respectively, using X-ray diffraction. Our results are consistent with those of previous theoretical calculations and experimental studies.

The central Zn$^{2+}$ in the Zn–Al LDH cluster had a coordination number of six, with an average Zn–O bond length of 2.10 Å. Gou et al. [3], by X-ray diffraction and EXAFS experiments, obtained an average Zn–O bond length of 2.08 Å (coordination number = 6.5) in the Zn–Al layer-like double hydroxides. Thus, the findings of these two studies are in good agreement with each other. For Cd–Cal, the average Cd–O (2.28 Å) bond length calculated in this study is consistent with the results obtained using EXAFS data (2.30 Å), as measured by Bailey et al. [65]. According to our procedure, the average Cd–O bond length in the Cd–HAp cluster was confirmed to be 2.39 Å, which is exactly the same as the average Cd–O (2.39 Å) bond length measured by Hata et al. [66] and the average Cd–O (2.38 Å) bond length calculated by He et al. [25].
Relative volume changes

The relative volume change of mineral lattices during geometrical optimization has only been studied by a few researchers. Wang et al. [4] calculated the ΔV/V₀ values of Ba-containing species by using the periodic boundary method. Except for Ba(OH)₂(H₂O)₈, the relative volume change of which was -0.67%, the ΔV/V₀ of the other species ranged from 1.26% to 6.30%, with an average of 4.46%. Ducher et al. [67] also used periodic boundary method to calculate the ΔV/V₀ of various Zn-containing minerals (sulfide, carbonate, oxide, silicate, sulfate, and arsenate), which ranged between 1.40% and 5.80%, with an average of 3.56%. Thus, we concluded that for the majority of the studied systems, the ΔV/V₀ values calculated using the molecular cluster method are much smaller (Table 4), approximately one-tenth of those obtained by the periodic boundary method.

Table 3 Isotopic equilibrium fractionation between molecular clusters and their corresponding aqueous solutions

| Cluster     | Calculated bond length (Å) | Experimental bond length (Å) | 1000 lnβ (cluster, 298 K, ‰) | 1000 lnβ (aquocomplex, 298 K, ‰) | Δ cluster-aquocomplex (298 K, ‰) |
|-------------|----------------------------|-----------------------------|-------------------------------|----------------------------------|----------------------------------|
| Si–Qtz      | 1.63, 1.62ab, 1.61b        | 1.61ab, 1.62b               | 73.65                         | 71.63                            | 2.02 (Si isotopes)               |
| O–Qtz       | 1.63                       | 1.61c                       | 110.31                        | 74.26                            | 36.05 (O isotopes)               |
| Zn–Al LDH   | 2.10                       | 2.08a                       | 4.87                          | 3.75                             | 1.12 (Zn isotopes)               |
| Cd–Cal      | 2.28                       | 2.30f                       | 2.62                          | 2.88g                            | −0.26 (Cd isotopes)              |

* Si–O bond length (α–quartz, Ab initio LDF theory, Purton et al. [64])
* Si–O bond length (quartz, DFT calculations, He and Liu [22])
* Si–O bond length (quartz, X-ray diffraction, Hazen et al. [39])
* Si–O bond length (quartz, X-ray diffraction at 848 K, Kihara [48])
* Zn–O bond length (X-ray diffraction and EXAFS, Gou et al. [3])
* Cd–O bond length (EXAFS, Bailey et al. [65])
* DFT calculations (He et al. [25])

Table 4 Relative volume change of solid molecular clusters

| Cluster     | ΔV/V₀ (%) | This study | Previous studies |
|-------------|-----------|------------|------------------|
| Si–Qtz      | −0.39     | 7.17a, −2.66b, 4.61c, −2.75d |
| Zn–Al LDH   | 1.96      | −          | 4.87             |
| Cd–Cal      | 0.45      | 3.54e, 3.47f |
| Cd–HAp      | 0.74      | 0.79, 3.43h, 4.37i |

* PBE [68]
* LDA [68]
* GGA(PW91) [69]
* LDA [69]
* CaCO₃ [70]
* [71]
* DFT calculations [25]
* GGA–PBE [72]
* GGA–PW91 [73]

Relative volume changes

We found that the ΔV/V₀ of quartz using the periodic boundary method calculated by the PBE and LDA functionals, primarily used by Méheut et al. [68], were 7.17% and -2.66%, respectively. Hamann [69] obtained the ΔV/V₀ of quartz using GGA (4.61%, PW91) and LDA (-2.75%). The ΔV/V₀ obtained in this study was -0.39%, which is substantially less than that obtained in previous studies. The Zn–Al LDH has a slightly larger expansion (1.96%) during geometric optimization, compared to that of other minerals in our study. But it is still smaller than those obtained by the periodic boundary method. The volume of the calculated Cd-containing calcite also increased (0.45%). However, the relative volume change of the Mn-containing calcite (3.54%) [70] and tetrahedral AO₄²⁻ group (A=S, Cr, Se)-doped calcite (3.47%) [71] calculated by the periodic boundary method were larger than that obtained during this study. For HAp without impurities, Sailuam et al. [72] calculated a ΔV/V₀ of 3.43% using GGA–PBE as the exchange–correlation functional. Bhat et al. [73] obtained ΔV/V₀ (4.37%) using the exchange–correlation functional GGA–PW91. In contrast, the ΔV/V₀ of Cd–HAp (Cd/(Cd + Ca) = 1/13) calculated in this study was 0.74%, which is much smaller than the previously reported values.

\[ \Delta^{30/28}\text{Si}_{\text{Qtz}} - \text{H}_2\text{SiO}_3 \text{ and } \Delta^{18/16}\text{O}_{\text{Qtz}} - (\text{H}_2\text{O})_n \]

With different exchange–correlation functionals and basis sets, 1000 lnβ may be different. At 298 K, Méheut et al. [74] calculated the silicon β-factor of quartz as 69‰ and the oxygen β-factor as 102‰. At the same temperature, our silicon β-factor (73.65‰) and oxygen β-factors (110.31‰) were slightly larger than those reported by Méheut et al. [74]. However, when calculating the fractionation factor α, the difference in the β-factor for the same mineral can be offset. No substantial difference
was observed between the $\Delta^{18/16}\text{O}_{\text{quartz}} - (\text{H}_2\text{O})$ value reported by Méheut et al. [68] and that obtained in this study (31.37‰ vs. 36.05‰). We compared the silicon and oxygen fractionation factors calculated in this study with previous experimental and theoretical data.

He and Liu [22] calculated the Si isotopic equilibrium fractionation between quartz and H$_4$SiO$_4$ solution at 273–673 K using B3LYP/6–311G(2df). They thought that a large Si isotopic equilibrium fractionation factor of approximately 3.3‰ exists between quartz and H$_4$SiO$_4$ solution at 298 K. Dupuis et al. [75] calculated the fractionation factors of Si isotopes between quartz and aqueous H$_4$SiO$_4$ at 273–323 K using first-principles methods, and found that at 300 K, the fractionation between quartz and H$_2$SiO$_4$ was $+2.1 \pm 0.2$‰. Figure 11 shows that our calculation results are in good agreement with those of Dupuis et al. [75]. Under neutral and acidic conditions, Si often exists in the form of aqueous H$_2$SiO$_4$ solutions. We considered the experimentally determined $\Delta^{30/28}\text{Si}_{\text{quartz-aqueous solution}}$ to be $\Delta^{30/28}\text{Si}_{\text{quartz-H}_2\text{SiO}_4}$. From this perspective, the available experimental data [5, 7, 76, 77] match the existing theoretical data.

To verify the reliability of the VVCM method, we further calculated the equilibrium oxygen isotope fractionation factor between quartz and the aqueous solution and then compared our result with previous experimental and theoretical results (Fig. 12). In the low-temperature range, compared to the theoretical results of the linear fitting reported by Méheut et al. [68], our calculated results are more consistent with the experimental values [6, 83] and the fitting equation based on experimental data [84]. However, at high temperatures, both theoretical results are less than the experimental value [85–91]. For example, at 673 K, our calculated oxygen isotope equilibrium fractionation factor and the experimental values reported by Méheut et al. [68] and Sharp et al. [84] were 1.79‰, 2.77‰, and 4.25‰, respectively.

In conclusion, the quartz geometric optimization and calculation of traditional/nontraditional stable isotope fractionation factors was accurate using the improved VVCM method. Finally, we can reliably apply this method to other systems such as Zn–Al LDH and Cd–Cal.

$\Delta^{66/64}\text{Zn}_{\text{Zn–Al LDH}} - (\text{ZnH}_2\text{O}_4)^{2+}$ and $\Delta^{114/110}\text{Cd}_{(\text{Cd–Cal})} - (\text{CdH}_2\text{O}_4)^{2+}$

Because no experimental or theoretical data on Zn isotope fractionation between Zn–Al LDH and the solution have been published, we chose the Zn isotope fractionation produced during Zn$^{2+}$ adsorption on the aluminum oxide/hydroxide surface for comparison. As reported by Komárek et al. [92] and Gou et al. [93], at higher surface coverage, Zn–Al LDH was detected by XRD on the surface of γ-Al$_2$O$_3$ with a fractionation factor of $\Delta^{66/64}\text{Zn}_{\text{solid-aq}} = 0.02 \pm 0.07$‰. Pokrovsky et al. [94] also reported lower fractionation on corundum ($\Delta^{66/64}\text{Zn}_{\text{solid-aq}} = 0.19 \pm 0.05$‰) and gibbsite ($\Delta^{66/64}\text{Zn}_{\text{solid-aq}} = 0.10 \pm 0.05$‰) surfaces, suggesting that the potential formation of Zn–Al LDH on the surface of gibbsite and corundum is similar to that of γ-Al$_2$O$_3$. Although the size of calculated isotope fractionation (1.12‰) was larger than that obtained experimentally, the fractionation directions were consistent with each other. In fact, experimentally determined Zn isotope fractionation factor for adsorption on aluminum oxides/hydroxides may be complicated by mixed adsorption and precipitation. In future, we hope
that Zn isotope fractionation can be experimentally determined for the individual Zn–Al LDH precipitation.

The calculated isotopic equilibrium fractionation between Cd–Cal and Cd-containing aqueous solution was -0.26‰. Previous studies on precipitated calcite from artificial seawater solutions reported the entry of Cd$^{2+}$ into the calcite from seawater, with the Cd equilibrium isotope fractionation, $\Delta^{114/110}$Cd$_{calcite-seawater} = -0.45 \pm 0.12$‰ \[95\]. Furthermore, Xie et al. \[96\] revealed apparent Cd isotope fractionation in the coprecipitation of Cd–calcite, with $\Delta^{114/110}$Cd$_{calcite-aqueous} = -0.38 \pm 0.18$‰ to -0.20 ± 0.12‰. Our theoretical calculation results are consistent with previous experimental results.

Conclusions
We used the first-principles calculation method to geometrically optimize the molecular clusters of Qtz, Zn–Al LDH, and Cd–Cal and obtained relative volume changes of -0.39%, 1.96%, and 0.45%, respectively. Compared to the periodic boundary method, the improved VVCM method slightly changes the molecular cluster volume during optimization, which increased the accuracy of our simulations. We also calculated the isotopic equilibrium fractionation between Qtz, Zn–Al LDH, Cd–Cal, and their corresponding aqueous solutions at 298 K, which were as follows: $\Delta^{30/28}$Si$_{Qtz-H_2O} = 2.20$‰, $\Delta^{18/16}$O$_{Qtz-H_2O} = 36.05$‰, $\Delta^{65/64}$Zn$_{Zn-Al LDH-ZnH_2O_4^{2-}} = 1.12$‰ and $\Delta^{114/110}$Cd$_{Cd-Cal-CdH_2O_4^{2-}} = -0.26$‰. These results are consistent with those of previous studies, confirming the reliability of the improved VVCM method for crystal-lattice optimization.

Abbreviations
VVCM: Volume variable cluster model; DFT: Density Functional Theory; Si–Qtz: Layered double hydroxide; O–Qtz: Quartz centered on oxygen atoms; O–Qz: Quartz centered on oxygen atoms; LDH: Layered double hydroxide; Cd–Cal: Cd–containing calcite; Cd–HAp: Cd-containing hydroxyapatite; PCA: Point charge arrangement.

Supplementary Information
The online version contains supplementary material available at https://doi.org/10.1186/s12932-022-00078-6.

Additional file 1: Figure S1 Hydrogen bonds connected water molecules and scheme of point charge arrangements (PCA) for 1 x, 3 x, 4 x, and 5 x.

Acknowledgements
The authors are grateful for the financial support from the Hebei Natural Science Foundation, Hebei Education Department Key Program and the State Natural Sciences Foundation.

Author contributions
YFW carries out calculations and analyses, and prepares the first draft; PDW helps to collect data; XYL, LXJ, JL, TDZ and HNZ assist in analysis and discussion; YFW and HTH design the simulation frame, discuss and edit the manuscript. All authors read and approved the final manuscript.

Funding
This work was financially supported by the Hebei Natural Sciences Foundation (Nos. D2020402004 and D2021402020), Hebei Education Department Key Program (No. ZD2018086), and State Natural Sciences Foundation (No. 4160301).

Availability of data and materials
The data sets that support the conclusions of this study are included in the study.

Declarations

Competing interests
There is no conflict of interest.

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Received: 30 November 2021 Accepted: 31 March 2022
Published online: 22 May 2022

References
1. Zhu X-C, Di D-R, Ma Ming-G, Shi Wei-Y, (2019) Stable isotopes in green-house gases from soil: a review of theory and application. Atmosphere 10:377
2. Erkaev NV, Scherf M, Thaller SE, Lammer H, Mezentsev AV, Ivanov VA, Mandt KE (2020) Escape and evolution of Titan’s N$_2$ atmosphere constrained by $^{14}$N/$^{15}$N isootope ratios. Mon Not R Astron Soc 500:2020–2035
3. Gou WX, Siebecker MG, Wang ZM, Li W (2018) Competitive sorption of Ni and Zn at the aluminum oxide/water interface: an XAFS study. Geochim Trans 19:9
4. Wang WZ, Wu ZQ, Huang F (2021) Equilibrium barium isotope fractionation between minerals and aqueous solution from first-principles calculations. Geochim Cosmochim Acta 292:64–77
5. Wang W, Geilert S, Wei HZ, Jiang SY (2021) Competition of equilibrium and kinetic silicon isotope fractionation during silica precipitation from acidic to alkaline pH solutions in geothermal systems. Geochim Cosmochim Acta 306:44–62
6. Layebyre L (1974) New approach to surface seawater paleotemperatures using $^{18}$O/$^{16}$O ratios in silica of diatom frustules. Nature 248:40–42
7. Dai S, Li D, Ren D, Tang Y, Shao L, Song H (2004) Geochemistry of the late Permian No. 30 coal seam, Zhijin Coalfield of Southwest China: influence of a siliceous low-temperature hydrothermal fluid. Appl Geochem 19:1315–1330
8. Clayton RN, Epstein S (1961) The use of oxygen isotopes in high-temperature geological thermometry. J Geol 69:447–452
9. Klein RT, Loehmann KC, Thayer CW (1996) Bivalve skeletons record sea-surface temperature and $^{34}$S via Mg/Ca and $^{18}$O/$^{16}$O ratios. Geology 24:415
10. Mao JW, Zhang ZC, Zhang ZH, Du A (1999) Re–Os isotopic dating of molybdenites in the Xiaoliugou W (Mo) deposit in the northern Qilian mountains and its geological significance. Geochim Cosmochim Acta 63:1815–1818
11. Wan Y, Liu D, Wilde SA, Cao J, Chen B, Dong C, Song B, Du L (2010) Evolution of the Yunkai terrane, south China: evidence from shrimp zircon U-Pb dating, geochemistry and Nd isotope. J Asian Earth SCI 37:140–153
12. Zhang Z, Zuo R (2014) Sr–Nd–Pb isotope systematics of magnetite: implications for the genesis of Makeng Fe deposit, southern China. Ore Geol Rev 63:1815–1818
13. Ma YB, Bagas L, Xing SW, Zhang ST, Wang RJ, Li N, Zhang ZJ, Zou YF, Yang XQ, Wang Y (2016) Genesis of the stratiform Zhenzigou Pb-Zn deposit in the North China Craton: Pb–Sr–Nd–Pb isotope constraints. Ore Geol Rev 79:88–104
14. Zhang Y, Wen H, Zhu C, Fan H, Cloquet C (2018) Cadmium isotopic evidence for the evolution of marine primary productivity and the biological
