Diversity and paleoenvironmental implications of an elasmobranch assemblage from the Oligocene–Miocene boundary of Ecuador

Jorge D. Carrillo-Briceño1, Jaime A. Villafaña2,3, Carlos Gracia2,4, F. Fernando Flores-Alcivar5, René Kindlimann1 and Juan Abella6,7

1 Palaeontological Institute and Museum, University of Zurich, Zurich, Switzerland
2 University of Vienna, Department of Paleontology, Vienna, Austria
3 Centro de Investigación en Recursos Naturales y Sustentabilidad, Universidad Bernardo O’Higgins, Santiago, Chile
4 Center of Tropical Paleocology and Archaeology, Smithsonian Tropical Research Institute, Panama, Panama
5 Universidad Estatal de la Península de Santa Elena, La Libertad, Santa Elena, Ecuador
6 Institut Català de Paleontologia Miquel Crusafont, Universitat Autònoma de Barcelona, Cerdanyola del Vallès, Barcelona, Spain
7 Instituto Nacional de Biodiversidad, (Parque La Carolina) Quito, Ecuador

ABSTRACT

The occurrence and diversity of elasmobranchs from the Oligocene–Miocene boundary from Tropical America is poorly known in comparison with the paleodiversity from younger Neogene intervals of the region. Here we describe a new elasmobranch assemblage from the rich fossil site of Montañita-Olón (Dos Bocas Formation, Santa Elena, Ecuador), where other vertebrates have already been described: for example, sea turtles and cetaceans. We report a total of 27 elasmobranch taxa, 19 of which are new fossil records for Ecuador, 10 new records for the Central Eastern Pacific and four new records for South America. Additionally, in order to reconstruct the environment where these marine remains were deposited, we performed abundance, paleobathymetric and habitat preference analyses, concluding that they were likely deposited in an outer neritic (open shelf) environment. The study of Oligocene and early Miocene marine elasmobranch faunas in Tropical America is key to addressing the issues in the evolutionary history of this group.

INTRODUCTION

The Oligocene–Miocene transition (OMT) was an important period for the evolutionary history of the marine biota in the northern margins of South America, especially due to the significant changes that affected marine biota in the Pacific and proto-Caribbean region (Johnson, Sánchez-Villagra & Aguilera, 2009). At that time, large scale geological processes like the closure of the Central American Seaway (CAS) and the rise of the Panamanian
Isthmus had not yet been completed (Coates & Stallard, 2013; Jaramillo et al., 2017). The CAS was a deep oceanic connection along the tectonic boundary between the Caribbean and South American plates that connected the Eastern Central Pacific (ECP) and Western Central Atlantic (WCA) (Jaramillo et al., 2017). This marine corridor allowed the flow of species (Leigh, O’Dea & Vermeij, 2014) until the end of the Miocene (Coates & Stallard, 2013; Bacon et al., 2015; Jaramillo et al., 2017). In this context, the evolutionary history of the marine fish fauna in Tropical America, according to the fossil record, has been linked with the geographic changes of the oceanic pathway that connected the ECP and WCA (e.g., Aguilera Socorro et al., 2011; Aguilera et al., 2017b; Schwarzhans & Aguilera, 2013; Carrillo-Briceño et al., 2018, and references therein). The study of Oligocene and early Miocene marine elasmobranch faunas is a proxy that allows us to address issues in evolutionary history in Tropical America, offering new light on factors that drove changes in biogeographic patterns of elasmobranchs before the closure of the CAS (Carrillo-Briceño et al., 2018). This oceanographic event resulted in a barrier that isolated the marine biota in both oceanic regions (ECP and WCA) (Lessios, 2008; Coates & Stallard, 2013; Leigh, O’Dea & Vermeij, 2014). Despite this, the Oligocene and early Miocene elasmobranch diversity from Tropical America is poorly known in comparison with the paleodiversity from younger Neogene intervals of the region (Carrillo-Briceño et al., 2018, fig. 5, tables S3 and S4). Oligocene elasmobranchs from Tropical America include a few isolated reports from the Caribbean region (Casier, 1958, 1966; Kruckow & Thies, 1990). In contrast, early Miocene assemblages from Tropical America are well known from the Caribbean and other WCA basins (Leriche, 1938; Casier, 1958; Santos & Travassos, 1960; Santos & Salgado, 1971; Iturralde-Vinent, Hubbell & Rojas, 1996; Sánchez-Villagra et al., 2000; Aguilera, 2010; Carrillo-Briceño et al., 2016a, 2016b, 2019; Aguilera et al., 2017a), as well as from the ECP in Peru (De la Cruz, 2008; Shimada et al., 2017; Landini et al., 2019). In this study, we present a revision of a new elasmobranch assemblage from the Oligocene–Miocene boundary (Dos Bocas Formation) of the Santa Elena Province, Ecuador, on the margin of the ECP (Fig. 1). The assemblage composition was studied, and we present a taxonomic list of sharks and rays with a comprehensive paleoenvironmental interpretation based on their bathymetric affinities. Additionally, the significance of the fossil assemblage for chronostratigraphic inferences is also discussed. The new elasmobranch assemblage, among other marine vertebrates found in the Dos Bocas Formation, including sea turtles (Cadena, Abella & Gregori, 2018), a new genus dolphin (Tanaka et al., 2017), and actinopterygian remains, increases the fossil record of the region and represents a critical window into marine tropical vertebrate faunas in the ECP during the OMT.

GEOLOGICAL SETTING

The outcrops of the Dos Bocas Formation studied herein correspond to a cliff of approximately 800 m in length along the coastline between the towns of Montañita and Olón (Fig. 1). Whittaker (1988) and Tanaka et al. (2017) provisionally identified the Montañita-Olón outcrops as part of the Zapotal Member of the Dos Bocas Formation. The fossil sharks and rays studied herein are from several points along the Montañita-Olón cliff, in localities that can only be accessed during low tides. For this study, a general
A stratigraphic column for the Montañita-Olón outcrops with an approximate thickness of 14 m was elaborated (Figs. 2A–2C). The geological section is dominated by strata composed of a moderately sorted fine to medium-grained sandstone with angular quartz-feldspatic clasts and rounded green grains (probably glauconite, although berthierine cannot be dismissed), with a micritic and volcanogenic matrix (Tanaka et al., 2017). Concretions are abundant in the section, even forming well-defined layers (Fig. 2B). In these concretions, other fossil vertebrates such as cetaceans, sea turtles and bony fishes have been collected (Tanaka et al., 2017; Cadena, Abella & Gregori, 2018).

A well-defined thin bioturbated layer, with abundant fossil teeth is present underlyng the above-mentioned concretion layer (Figs. 2B and 2C). Our observations confirm that bivalves and gastropods with poor preservation are also present, occurring in greater abundance in the upper layers of the section (Figs. 2A and 2C). It has been suggested as “general interpretation” that the Dos Bocas Formation was deposited in a shallow protected environment (see Tanaka et al., 2017; and references therein); although no clear evidence has been presented to justify this hypothesis for the Montañita-Olón outcrops.
In contrast, an upper platform environment based on micropaleontological evidence was suggested for the Dos Bocas Formations by Ordoñez, Jiménez & Suárez (2006) and Witt et al. (2019, fig. 3C). A late Oligocene age has been inferred for the Dos Bocas Formation based on its faunal composition (Olsson, 1931; Bristow, 1975; Tanaka et al., 2017; Cadena, Abella & Gregori, 2018). A recent U-Pb zircon dating for the Montañita-Olón outcrops confirmed an age of 23.5 ± 0.4 Ma, with a younger cluster average of 22.9 ± 0.6 Ma (Witt et al., 2019). The sample used for the above mentioned U-Pb zircon dating, and referred by Witt et al. (2019) as “sample CP705”, was collected by one of the authors (JA) in the outcrops of Montañita area (Figs. 2B and 2C).

MATERIALS AND METHODS

The fossil elasmobranch fauna from the Montañita-Olón site consists of 424 cranial (teeth and rostral spines) and postcranial (vertebrae and caudal spines) elements (Figs. 2–8; Table 1; Data S1; Tables S1–S3); a few actinopterygian remains were also found. All specimens were collected in situ from different points along the Montañita-Olón cliff section (Fig. 2C), during several field trips conducted by the authors (JA, JDCB and FFA).
and other collaborators between 2016 and 2018. The Ecuadorian Instituto Nacional de Patrimonio Cultural (INPC) excavation permit, Code: No 0039-DR5.INPC.2015 supported the field activities. The Montañita-Olón localities (Fig. 1) are located in a coastal cliff of around 800 m long, between the towns of Montañita and Olón (Santa Elena Province, Ecuador, coordinates: 1°48’9.3″S, 80°45’24,10W and 1°49’09.66″S, 80°45’30,50W). The fossil specimens are housed at the “Museo Paleontológico Megaterio”

Figure 3 Hexanchiformes and Squaliformes of the Montañita-Olón site (Dos Bocas Formation). (A–H) *Heptranchias* cf. *H. howellii* ((A–H): MPM-1365; lower lateral teeth (A–F); upper lateral tooth (G and H)). (I–O) *Hexanchus* cf. *H. griseus* ((I–O): MPM-1359; lower lateral teeth (I–M); upper tooth (N and O)). (P–W) *Centrophorus* cf. *C. granulosus* ((P–W): MPM-1367; lower antero-lateral teeth (P–U); upper antero-lateral tooth (V–W)). (X and Y) *Dalatias* sp. (lower tooth (MPM-1366)). View: labial (B, C, H, J, K, O, Q, S, U, W and Y) and lingual (A, D–G, I, L–N, P, R, T, V and X).

Full-size DOI: 10.7717/peerj.9051/fig-3
(MPM-) at the Universidad Estatal Peninsula de Santa Elena, Ecuador (Table S3).

Photographs were taken with a Leica MZ16F multifocal stereomicroscope and Scanning Electronic Microscope (SEM) for small teeth. Tooth measurements including total height, width and length (see Fig. S1), were taken for the specimens and are listed in Table S1.

**Figure 4** Echinorhiniformes, Pristiophoriformes, Orectolobiformes and Lamniformes of the Montañita-Olón site (Dos Bocas Formation). (A and B) cf. *Echinorhinus* sp. (antero-lateral tooth (MPM-1368)). (C and D) *Paraechinorhinus* cf. *P. barnesi* (antero-lateral tooth (MPM-1369)). (E–H) *Pristiophorus* sp. (rostral spines (MPM-1361)). (I–K) *Rhincodon* sp. (upper antero-lateral tooth (MPM-1370)). (L–S) *Isourus* cf. *I. oxyrinchus* ((L–S): MPM-1364; lower anterior teeth (L–O); lower tooth (P and Q); upper lateral tooth (R and S)); (T–Z) *Mitsukurina* cf. *M. lineata* ((T–Z): MPM-1371; lower anterior tooth (T–V); lower lateral tooth (W and X); upper anterior tooth (Y and Z)). View: labial (B, C, I, M, O, Q, S, V, X and Y), lingual (A, D, L, N, P, R, T, W and Z), dorsal (E–H) and profile (J and U), basal (K).
Dental terminology and the systematics utilized herein follow Cappetta (2012).

We identified all fossil elasmobranch remains to the lowest possible taxonomic level. Taxonomic identifications are based on literature review and comparative analysis between fossil and extant specimens from several collections including: the Fossil Vertebrate Section of the Museum für Naturkunde, Berlin, Germany (MB.Ma.); Natural History Museum of Basel (NMB), Switzerland; Natural History Museum of Vienna (NHMW), Austria; Paleontological collection of the Alcaldía del Municipio Urumaco.
(AMU-CURS); Paleontological collection of the Institut des Sciences de l’Evolution, University of Montpellier (UM), France; Palaeontological Institute and Museum at the University of Zurich (PIMUZ) and the René Kindlimann private collection with public access, Aathal, Switzerland; paleontological collections of the Mapuka Museum of Universidad del Norte (MUN), Barranquilla, Colombia.

Figure 6 Lamniformes and Carcharhiniformes of the Montañita-Olón site (Dos Bocas Formation). (A–C) Lamnidae indet. (Lower anterior tooth (MPM-1377)). (D–H) Alopias cf. A. exigua (lower anterior teeth (MPM-1374)). (I–M) Alopias latidens (anterior lateral teeth (MPM-1375)). (N–S) Carcharhinus gibbesii (upper antero-lateral teeth (MPM-1376)). (T–W) Galeocerdo aduncus (antero-lateral teeth (MPM-1379)). View: labial (C, F, H, K, M, N, P, T and V), lingual (A, D, G, I, L, O, Q–S, U and W) and profile (B, E and J).
An analysis of abundance for the Montañita-Olón site assemblage was carried out using percentages of specimens by family, genera and species. In addition, we performed a paleobathymetric analysis following the methodology of Nolf & Brzobohatý (1994), adapted to fossil sharks (Carrillo-Briceño et al., 2015a, 2016a, 2016b). A second paleobathymetric analysis following the methodology of Perez et al. (2017) was performed using R (R Core Team, 2019). This depth estimator uses weighted bootstrap analysis to estimate the depth distribution of the population sampled. To apply this technique, we calculated the mean for the total depth distribution range and the mean for the common distribution range for each taxon. A bootstrap was performed with the mean differences and using the relative abundance of each taxa as weight estimator of the mean. The data were resampled 10,000 times and plotted as a histogram that provides the mean depth. The 95% confidence interval was obtained using a percentile method. The raw data and

**Figure 7** Carcharhiniformes of the Montañita-Olón site (Dos Bocas Formation). (A–F) Physogaleus contortus ((A–F): MPM-1380; lower antero-lateral tooth (A–D); upper antero-lateral tooth (E and F)). (G–J) Hemipristis serra (upper lateral teeth (MPM-1354)). (K–N) Sphyra sp. (upper antero-lateral teeth (MPM-1381)). (O and P) Carcharhiniformes indet. vertebrae (MPM-1382). View: labial (B, D, F, H, J, L and M) and lingual (A, C, E, G, I, K and N).
script for the paleobathymetric analysis were included in Data S2. The average, minimum and maximum depth estimates were plotted in Fig. 10. The minimum and maximum depth are the lower and upper limits of the mean depth range in the distribution. For both analyses, we included only species/genera with closely related extant taxa. Extinct taxa without clear identification to the generic level were removed. We analyzed a total of 19 (out of 24) taxa that represent a total of 373 individuals for which ecological information is available based on closely related extant taxa. The bathymetric (depth range) and habitat preferences utilized for these analyses are available in Table S2, compilation following Compagno (1984a, 1984b), Compagno, Dando & Fowler (2005), Musick, Harbin & Compagno (2004), Ebert & Stehmann (2013) and the FishBase website (Froese & Pauly, 2019). We use the term “Tropical America” (Neotropics) to refer to the geographic area of

Figure 8 Myliobatiformes and Osteichthyes remains of the Montañita-Olón site (Dos Bocas Formation). (A–K) Mobula fragilis ((A–K): female antero-lateral teeth (MPM-1383)). (L–R) Mobula sp. ((L–R): female indet. position (MPM-1384)). (S and T) Myliobatiformes indet. (caudal spine fragments (MPM-1385)). (U) Caudal peduncle of Eoceleopoma sp. (MPM-1391). (V and W) Upper pharyngeal element of Hypsigenyini. (MPM-1392). (X) Isolated tooth of Labridae indet. (MPM-1393). View: labial (A, E, L, O, W), lingual (I, P), occlusal (C, G, J, M, Q, V), profile (B, F, K, N, R), basal (D, H), dorsal (T), ventral (S) and indet. (X), lateral (U).
the Western Hemisphere located between the Tropic of Cancer (23°27′ N) and the Tropic of Capricorn (23°7′ S). The ECP and the WCA are here referred to as the oceanic areas of Tropical America. The Eastern Pacific is referred to as the oceanic area from North America to the most southern point of South America.

RESULTS
Elasmobranch paleodiversity
The elasmobranch assemblage described herein from the Montañita-Olón site comprises 27 taxa (including three of indeterminate taxonomy) of squalomorphs, galeomorphs and batoids (Figs. 3–8; Table 1; Tables S1 and S2). A general descriptive taxonomy of the dental elements of key species (first and rare records in Ecuador and Tropical America) is presented in Data S1.

Table 1 Elasmobranch paleodiversity of the Montañita-Olón site (Dos Bocas Formation).

| Superorder       | Order               | Family           | Genus         | Taxon                                      |
|------------------|---------------------|------------------|---------------|--------------------------------------------|
| Squalomorphii    | Hexanchiformes      | Heptranchidae    | Heptranchias   | Heptranchias cf. †H. howellii*             |
|                  | Hexanchidae         |                  | Hexanchus     | Hexanchus cf. H. griseus*                 |
| Squaliforms      | Centrophoridae      | Centrophorus     | Centrophorus cf. C. granulosus*             |
|                  | Dalatiidae          | Dalatias         | Dalatias sp.*|                                            |
| Echinorhiniforms | Echinorhinidae      | Echinorhinus     |               | cf. Echinorhinus sp.                        |
|                  |                     | †Paracarcharinus  |               | †Paracarcharinus cf. †P. barnesi*         |
| Galeomorphii     | Priophoriiformes    | Priophorus       | Priophorus sp.|
|                  | Orectolobiformes    | Rhincodontidae   | Rhincodon     | Rhincodon sp.*                             |
|                  | Lamniformes         | Lamnidae         | †Isurus       | †Isurus cf. I. oxyrinchus*                 |
|                  | Mitsukurinidae      | Mitsukurina      | Mitsukurina cf. †M. lineata*               |
|                  | Odontaspididae      | Carcharias       | Carcharias sp.|
|                  |                     | Odontaspis       | Odontaspis sp.*                            |
|                  | †Otodontidae        | †Otodus          | †Otodus (Carcharocles) cf. †O. angustidens* |
|                  |                     | †Parotodus        | †Parotodus benedeni*                        |
|                  |                     | †Megalolamna      | †Megalolamna paradoxodon*                  |
|                  | Alopiidae           | Alopia           | Alopia cf. †A. exigua*                     |
|                  |                     |                 | †Alopia latidens*                          |
| Carcharhiniforms | Carcharhinidae      | Carcharhinus     | †Carcharhinus gibbesii*                    |
|                  |                     | Galeocerdo       | †Galeocerdo aduncus*                       |
|                  |                     | †Physogaleus     | †Physogaleus contortus*                    |
|                  | Hemigaleidae        | Hemipristis      | †Hemipristis serra                         |
|                  | Sphyrnidae          | Sphyrna          | Sphyrna sp.                                |
|                  | Indet.              | Indet.           | Indet.                                      |
| Batomorphii      | Myliobatiformes     | Mobulidae        | Mobula†                                    |
|                  |                     |                 | †Mobula fragilis*                          |
|                  |                     |                 | Mobula sp. (this morphotype*)               |
|                  | Indet.              | Indet.           | Indet.                                      |

Note: * First fossil record in Ecuador.
Squalomorphs, with seven species, are the second most diverse and abundant group of sharks in the Montañita-Olón assemblage (Fig. 9; Table 1), and their remains are represented exclusively by isolated teeth (Table S1). The identified taxa include the hexanchiforms *Heptranchias* cf. *H. howellii* (Reed, 1946) (Figs. 3A–3H; Data S1) and *Hexanchus* cf. *H. griseus* (Bonnaterre, 1788) (Figs. 3I–3O; Data S1), the squaliforms *Centrophorus* cf. *C. granulosus* (Bloch & Schneider, 1801) (Figs. 3P–3W; Data S1) and *Dalatias* sp. (Figs. 3X and 3Y; Data S1), the echinorhiniforms cf. *Echinorhinus* sp. (Figs. 4A and 4B; Data S1), and *Paraechinorhinus* cf. *P. barnesi* Welton in Pfeil, 1983 (Figs. 4C and 4D; Data S1), and the pristiophoriform *Pristiphorus* sp. (Figs. 4E–4H). With the exception of *Pristiphorus* sp., all the above-mentioned species are reported for the first time in the fossil record of Ecuador. *Heptranchias* cf. *H. howellii*, *Centrophorus* cf. *C. granulosus*, *Dalatias* sp., cf. *Echinorhinus* sp. and *Paraechinorhinus* cf. *P. barnesi* are reported for the first time in the fossil record of the ECP.

Galeomorphs (isolated teeth as well as vertebral remains) represent the most diverse and abundant group in the elasmobranch assemblage from the Montañita-Olón site (Fig. 9). This group of sharks is characterized by a total of 15 species within 14 genera and nine families of Orectolobiformes, Lamniformes and Carcharhiniformes (Table 1). Lamniforms represent the most diverse group in the overall assemblage of the Montañita-Olón site. With 10 taxa, it includes: *Isurus* cf. *I. oxyrinchus* Rafinesque, 1810 (Figs. 4L–4S; Table 1).
Data S1), an indeterminate lamnid species (represented by and isolated tooth, see Figs. 6A–6C; Data S1), *Mitsukurina* cf. *M. lineata* (Probst, 1879) (Figs. 4T–4Z; Data S1), *Carcharias* sp. (Figs. 5A and 5B; Data S1), *Odontaspis* sp. (Figs. 5C–5K; Data S1), *Otodus* (*Carcharocles*) cf. *O. angustidens* (Agassiz, 1833–1843) (Figs. 5L–5Q; Data S1), *Parotodus benedenii* (Le Hon, 1871) (Figs. 5R–5T; Data S1), *Megalolamna paradoxodon* (Shimada et al., 2017) (Figs. 5U–5X; Data S1), and *Alopias* cf. *A. exigua* (Probst, 1879) (Figs. 6D–6H). Carchariniforms are represented by *Carcharhinus gibbesii* (Woodward, 1889) (Figs. 6N–6S; Data S1), the most abundant taxon in the assemblage (Table S2), as well as *Galeocerdo aduncus* (Agassiz, 1833–1843) (Figs. 6T–6W), *Physogaleus contortus* (Gibbes, 1849) (Figs. 7A–7F), *Hemipristis serra* (Agassiz, 1833–1843) (Figs. 7G–7J), and *Sphyrna* sp. (Figs. 7K–7N). A group of 25 semi articulated vertebrae (Figs. 7O and 7P) was also collected in the Montañita-Olón site inside a concretion (Fig. 2C); however, due to the lack of diagnostic elements, a more accurate identification than “Carcharhiniformes indet.” is not possible. The presence of an incomplete tooth of the whale shark *Rhincodon* sp. (Figs. 4I–4K; Data S1) represents the only record of an orectolobiform shark in the Montañita-Olón site. With the exception of *Carcharias* sp. and *H. serra*, all the above-mentioned galeomorph species are reported for the first time in the fossil record of Ecuador.

*Rhincodon* sp., *Mitsukurina* cf. *M. lineata*, *Odontaspis* sp., *Otodus* (Carcharocles) cf. *O. angustidens*, *Alopias* cf. *A. exigua*, *Alopias latidens* and *Carcharhinus gibbesii* are reported for the first time in the fossil record of the ECP. The batoids, with only one genus and two species, are the least diverse group from the Montañita-Olón assemblage (Fig. 9; Table 1). Only a few isolated teeth of the mobulids *Mobula fragilis* (Cappetta, 1970) (Figs. 8A–8K; Data S1) and *Mobula* sp. (Figs. 8L–8R; Data S1) are herein reported. *Mobula fragilis* is here reported for the first time in the fossil record of Ecuador and the ECP. Although some isolated mobulid teeth have been reported before from Neogene deposits of the Central Eastern Pacific of Ecuador (Carrillo-Briceño, Aguilera & Rodríguez, 2014) and Panama (Carrillo-Briceño et al., 2018), the morphological pattern of the *Mobula* sp. specimens from the Montañita-Olón site has not been recognized before in the ECP. In addition, two eroded, broken and non-diagnostic caudal spine fragments (Figs. 8S and 8T) are referred herein to Myliobatiformes indet.

### Paleobathymetric analysis

Two paleobathymetric methods were applied to the elasmobranch fauna from the Montañita-Olón site: (1) an adaptation of Nolf’s method which calculates the percentage of species that share an assigned depth range (Nolf & Brzobohatý, 1994) and (2) the weighted method (Perez et al., 2017) by resampling the mean differences and using the relative abundance as weight estimator of the mean. Nineteen of the 24 total taxa, which were represented by 373 specimens (88.39% of total studied material), were included in the analysis. We used only taxa that have comparable modern analogs that allow us to estimate their depth ranges (Carrillo-Briceño et al., 2015a). This allows us to obtain more accurate estimations. For taxa without modern representatives, their presence is assumed in the estimated depth range.
Both results were comparable. For the Nolf technique, 73.7% of the studied taxa are represented in a depth range of 100–200 m (Fig. 10B). For the weighted paleobathymetry, a range between 87.8 and 378.0 m with a mean of 192 m was estimated (Fig. 10C). Our results suggest that the fossils were most likely deposited in an outer neritic (open shelf) environment. Of these two applied techniques, the weighted method provides more precision, but the Nolf method allows a better visualization of the depth distribution for each taxa in the assemblage (Fig. 10). The results obtained using these two techniques, improves the data interpretation in our study.
DISCUSSION

Diversity composition and biostratigraphy significance

Elasmobranchs from the Oligocene and early Miocene have not been reported from Ecuador, and their fossil record has been restricted to younger strata (Longbottom, 1979; Carrillo-Briceño, Aguilera & Rodríguez, 2014). The elasmobranch assemblage described here from the Montañita-Olón site includes at least 27 taxa (Table 1). This represents the most diverse elasmobranch assemblage known from Ecuador and for the Oligocene–Miocene boundary of Tropical America. The assemblage includes 13 extinct taxa (e.g., *Heptranchias* cf. *H. howelli*, *Paraechinorhinus* cf. *P. barnesi*, *Mitsukurina* cf. *M. lineata*, *Otodus* (Carcharocles) cf. *O. angustidens*, *P. benedenii*, *M. paradoxodon*, *Alopias* cf. *A. exigua*, *A. latidens*, *C. gibbesii*, *G. aduncus*, *P. contortus*, *H. serra* and *M. fragilis*) with a worldwide paleodistribution (Cappetta, 2012; Cappetta, Gregorová & Adnet, 2016; Szabó & Kocsis, 2016a; Shimada et al., 2017). From the overall elasmobranch assemblage at Montañita-Olón site, 19 species are reported here for the first time in the fossil record of Ecuador (Table 1). Ten species, including the squalomorphs *Heptranchias* cf. *H. howelli*, *Centrophorus* cf. *C. granulosus*, *Dalatias* sp., cf. *Echinorhinus* sp., the galeomorphs *Rhincodon* sp., *Mitsukurina* cf. *M. lineata*, *Odontaspis* sp., *Carcharhinus* *gibbesii* and the batoids *Mobula fragilis and Mobula* sp., are reported for the first time in the fossil record of the ECP. Records of *Paraechinorhinus* cf. *P. barnesi*, *Otodus* (Carcharocles) cf. *O. angustidens*, *Alopias* cf. *A. exigua*, *A. latidens*, are reported for the first time in the fossil record of South America. The presence of *Rhincodon* sp. from Ecuador, together with the record from the late Oligocene (Chattian) of Eastern USA (Cicimurri & Knight, 2009), represent the oldest records for this taxon. The *Heptranchias* cf. *H. howelli* specimens from the Montañita-Olón site (see Data S1) clearly resemble teeth of *H. howelli* from the Oligocene of North America (Welton, 1974), early Miocene of Colombia (Carrillo-Briceño et al., 2016b) and other regions (Cappetta, Gregorová & Adnet, 2016). The presence of *Paraechinorhinus* cf. *P. barnesi* in the Oligocene–Miocene of Ecuador increases its paleobiogeographic distribution and represents the oldest record of the genus in the Americas, as *P. barnesi* was known only from the middle Miocene of California (USA) (Pfeil, 1983; Cappetta, 2012).

As mentioned above in the results section, mobulid teeth identified as *Mobula* sp. have been reported before in the ECP from the Middle Miocene-early Pliocene of Ecuador (Carrillo-Briceño, Aguilera & Rodríguez, 2014, fig. 5) and late Miocene-Pliocene of Panama (see Perez et al., 2017; Carrillo-Briceño et al., 2018, table S6 and references therein). However, the morphological pattern of the *Mobula* sp. specimens from the Montañita-Olón site (Figs. 8L–8R) differ from those teeth reported from Ecuador and Panama. The two *Mobula* sp. specimens referred to here (MPM-1384), as well as those specimens illustrated by Carrillo-Briceño et al. (2016a, fig. 11.19–21) from the early Miocene of Venezuela, resemble teeth of the living species *Mobula munkiana* Notarbartolo Di Sciara (1987) (Adnet et al., 2012, fig. 4). Nevertheless, due to the scarcity of fossil and comparative
material, a more accurate specific identification is not possible at this time. The batoids are the less diverse group from the Montañita-Olón assemblage with only benthopelagic representatives (Table S2). The absence of other batoids, especially those associated with benthonic habitats, could likely be a result of bias in sampling. However, the benthonic batoids could also be associated with the ecological and environmental conditions that prevailed during the deposition of the Dos Bocas Formation, (e.g., Cappetta, Gregorová & Adnet, 2016).

To date, the elasmobranch assemblage from the Montañita-Olón site is the only one known from Oligocene–Miocene boundary of the ECP. Other Oligocene elasmobranchs from the region are unknown, and only a few early Miocene assemblages are restricted to southern Peru (Landini et al., 2019). The elasmobranch assemblage reported by Landini et al. (2019) from the Chilcatay Formation which is characterized by 22 taxa, clearly contains faunal differences in comparison with those from the Montañita-Olón site, where only a few galeomorphs (e.g., Isurus oxyrinchus, Carcharias, P. benedenii, M. paradoxodon, G. aduncus, P. contortus and H. serra) were present in both geological units. A few early Miocene assemblages from the southernmost areas of the Eastern Pacific in Chile have also been described (Villafaña & Rivadeneira, 2018; Villafaña et al., 2019). As in the Montañita-Olón site, early Miocene assemblages from Chile include taxa such as Carcharias, Odontaspis, Isurus (Villafaña et al., 2019), including a new species of sawshark †Pristiophorus humboldti, Villafaña et al. (2019) from the Navidad Formation. The presence of Pristiophorus sp. in the Montañita-Olón site represents the oldest record for this taxon in the ECP, although the taxon has been reported from the Eocene of southernmost Magallanes Region (Otero et al., 2013). The isolated rostral spines reported here show the typical characters described for the genus Pristiophorus from the Eastern Pacific (Carrillo-Briceño et al., 2013; Staig et al., 2015; Villafaña et al., 2019). However, the use of rostral spines is not recommended for species identification due to their high variability and insufficiency as a diagnostic character (Underwood & Schlögl, 2013; Engelbrecht et al., 2017; Villafaña et al., 2019). By contrast, the teeth display enough characters to assign specimens to lower taxonomic levels. Therefore, we prefer to identify to the genus level until more material is available. In addition, two coprolites with abundant bony fish remains and some isolated rostral spines of Pristiophorus were found in the Montañita-Olón site (Fig. S2). Although it is difficult to identify the possible producer of these coprolites, there is no doubt that sawfishes were part of the predator’s diet.

Like Pristiophorus, other elasmobranch taxa such as Dalatias, Carcharias, and Isogomphodon, were at the end of the Neogene regionally extirpated from the Eastern Pacific, but with extant representatives still inhabiting the Western Atlantic (Carrillo-Briceño et al., 2018). The presence of Dalatias sp. in the Montañita-Olón site is the oldest record of this taxon in the Eastern Pacific. This record from Ecuador together with the one known from the late Miocene–Pliocene of the Atacama Region in Chile (Villafaña & Rivadeneira, 2018), suggest that Dalatias was present in the Eastern Pacific from the late Oligocene to the end of the Neogene, when it became regionally extinct. The presence of
Mitsukurina cf. *M. lineata* in the Montañita-Olón site also suggests the extirpation of the genus from the Eastern Pacific (see Carrillo-Briceño et al., 2018, table S4).

A late Oligocene–early Miocene age has been suggested for the Dos Bocas Formation on the basis of radiometric dating (Witt et al., 2019). According to Witt et al. (2019), the U-Pb zircon dating for the Montañita-Olón outcrop in the Montañita area (Fig. 2B) yielded an age of 23.5 ± 0.4 Ma, with a younger cluster average of 22.9 ± 0.6 Ma. These results agree with a late Oligocene age proposed for the unit using faunal composition (Olsson, 1931; Bristow, 1975; Tanaka et al., 2017). In addition, the age of the Dos Bocas Formation (Montañita-Olón site) proposed here is also supported by the presence of the typical Oligocene megatooth species *Otodus* (*Carcharocles*) cf. *O. angustidens* from other localities around the world (Baut & Génault, 1999; Gottfried & Fordyce, 2001; Reinecke, Stapf & Raisch, 2001; Szabó & Kocsis, 2016b).

**Paleoenvironmental inferences**

A shallow protected environment has been suggested as the most plausible depositional environment for the Dos Bocas Formation (see Tanaka et al., 2017; Witt et al., 2019, and references therein). The base of the fossiliferous outcrop consists of massive, moderately sorted, fine to medium-grained sandstone with angular quartz-feldspathic clasts and probably glauconite rounded green grains (Fig. 2). The matrix is micritic and volcanogenic, possibly bentonitic. According to Tanaka et al. (2017), bedding is massive to indistinct, suggesting little influence by traction currents or storm waves, and in turn implying a quiet setting; estuarine or mid-shelf is possible. Witt et al. (2019, fig. 8A) reported for the outcrops of the Dos Bocas Formation in the Montañita-Olón area, local large arthropod burrows and some evidence of patch reefs of bivalves and worms, suggesting that these beds were deposited in a protected shallow-water environment, likely under a period of strong tectonic deformation. Nevertheless, arthropod burrows are not exclusively from shallow water environments, and our field observations suggest that the thick layer of large arthropod burrows and the probable evidence of patch reefs with fragmentary mollusks reported by Witt et al. (2019) for the Montañita-Olón area are represented towards the top of the section (Fig. 2A). In contrast, micropaleontological evidence suggests for the Dos Bocas Formations, by Ordoñez, Jiménez & Suárez (2006) and Witt et al. (2019, fig. 3C), an upper platform environment. In our bathymetric analysis of the Montañita-Olón site (Dos Bocas Formation), 73.7% of the studied elasmobranch taxa are represented in a depth range of 100–200 m (Fig. 10). For the weighted paleobathymetry, a range between 87.8 and 378.0 m with a mean of 192 m was estimated. Our results suggest that the fossils were most likely deposited in an outer neritic (open shelf) environment, which coincides with the upper platform environment suggested by Witt et al. (2019, fig. 3C). Additional support for this estimation is derived from the presence of the benthopelagic squalomorph sharks *Heptranchias* cf. *H. howellii*, *Hexanchus* cf. *H. griseus*, *Centrophorus* cf. *C. granulosus*, *Dalatias* sp. and cf. *Echinorhinus* sp., whose extant representatives usually prefer deep-water environments near the continental slope (see Table S2). The extinct *Heptranchias* cf. *H. howellii* has been reported from other localities of the Americas and Europe as a species with deep environment preferences (Cappetta, Gregorová & Adnet,
Paraechinorhinus cf. P. barnesi is another extinct taxon associated with deep-water paleoenvironments (Welton, 1979; Pfeil, 1983). In addition, the presence of “Goblin shark” Mitsukurina cf. M. lineata in the Montaña-Ólón assemblage could support our bathymetric estimation. The extant Mitsukurina bottom-dwelling species is usually living in deep waters, on the outer continental shelves and upper slopes (Compagno, 1984a; Cappetta, Gregorová & Adnet, 2016); although it should not be ruled out that the species could also have occasional visits to the neritic areas (Ebert & Stehmann, 2013). As an example, an isolated tooth assigned to Mitsukurina was referred from an infralittoral environment during the early Miocene of Costa Rica (Laurito et al., 2014).

Some of the benthopelagic and pelagic galeomorph taxa of the Montaña-Ólón assemblage (Table S2) have been associated with both coastal and open-water environments in the fossil record, as well as for their living representatives (Laurito, 1999; Aguilera, 2010; Cappetta, 2012; Compagno, 1984b, Purdy et al., 2001; Compagno, Dando & Fowler, 2005; Ebert & Stehmann, 2013; Carrillo-Briceño et al., 2013, 2015b, 2016a, 2018, 2019; Perez et al., 2017; Landini et al., 2019, and references there in). Taxa such as Isurus cf. I. oxyrinchus, Carcharias sp., Odontaspis sp., Otodus (Carcharocles) cf. O. angustidens, Parotodus benedenii, M. paradoxodon, Alopias cf. A. exigua, G. aduncus, P. contortus and H. serra have been reported with a wide global distribution during the Cenozoic (Gottfried & Fordyce, 2001; Cappetta, 2012; Reinecke et al., 2014; Carrillo-Briceño, Aguilera & Rodríguez, 2014; Carrillo-Briceño et al., 2015a), suggesting significant distances over oceanic basins and wide environmental ranges. The extinct Carcharhinus gibbesii, with 146 isolated teeth in the total sample, is the most abundant taxon in the elasmobranch assemblage from the Montaña-Ólón site. Like the above-mentioned sharks, C. gibbesii was an oceanodromous species with a wide distribution in North America, Europe and Tropical America during the Oligocene–early Miocene (Cicimurri & Knight, 2009; Reinecke et al., 2014; Carrillo-Briceño et al., 2016b, 2019). In the early Miocene of Colombia, C. gibbesii was reported in both shallow and deep-water environments (Carrillo-Briceño et al., 2016b, 2019).

Other vertebrates reported for the Montaña-Ólón outcrops include the dolphin Urkudelphis chawpipacha Tanaka et al., 2017, and a Pan-Cheloniidae sea turtle (Cadena, Abella & Gregori, 2018). However, no environmental conditions were inferred on the basis of these taxa. Bony fishes, including the tail of a Scombridae (Fig. 8U), and dental battery and isolated tooth of Labridae (Figs. 8V–8X), were collected in the Montaña-Ólón outcrops (Fig. 2C). The caudal peduncle of Scombridae in the Montaña-Ólón is assigned to †Eocoelopoma Woodward, 1901 by having preural vertebrae 2–4 abruptly shortened and a hypural plate formed by hypurals 1–4 (−5) (Monsch & Bannikov, 2011). Eocoelopoma is a primitive scombrid closely related to the Sardini + Thunnini clade and cannot be confused with Thunnini because the reduction of the preural vertebrae in this tribe is extremely pronounced (Monsch & Bannikov, 2011). The genus Eocoelopoma is known from the Paleocene of Turkmenistan, early Eocene of England, and from Equatorial Africa without specific age. This report for the Montaña-Ólón represents the first record of the genus for Tropical America (Bannikov, 1985; Monsch, 2004;
The presence of *Eocoelopoma* sp. in the Montañita-Olón locality suggests open to oceanic marine environments, due to the paleoenvironmental inferences suggested in other localities where this taxon was reported (Friedman et al., 2016). The caudal skeleton of *Eocoelopoma* presents an typical interlocking of vertebrae, which are well adapted to a pelagic mode of life (Fierstine & Walters, 1968). A second specimen from the Montañita-Olón locality represented by upper pharyngeal bones with phylloidont condition, where the teeth are developed in distinct stacks, is diagnostic of Hypsigenyini (Bellwood et al., 2019), and the multiple oblique tooth rows (Figs. 8V and 8W) are diagnostic characters for †Trigonodon (Bellwood et al., 2019). Hypsigenyini are restricted to the Miocene-Pliocene of Europe and Africa (Schultz & Bellwood, 2004; Bellwood et al., 2009), and *Trigonodon* in only known from the Miocene of Europe by the type species †Trigonodon oweni Sismonda in Michelotti (1847). Although *Trigonodon* has been referred from the early Miocene of Costa Rica (Laurito et al., 2014), a future detailed taxonomic revision for our specimens from Ecuador and those from Costa Rica could help give new insights on the taxonomy of these fossil Hypsigenyini fishes from Tropical America. According to Schultz & Bellwood (2004), *Trigonodon* was a shallow-water fish possibly living primarily on coral reefs and associated hard-grounds. The presence of shallow-water components in the Montañita-Olón site could be explained by elements being washed into deeper water by turbidity currents or slumping (Vialle, Adnet & Cappetta, 2011).

**CONCLUSIONS**

We report 27 elasmobranch taxa, of which 19 are new fossil records for Ecuador, 10 new records for the Central Eastern Pacific and four new records for South America. This elasmobranch fossil assemblage represents the most diverse known from Ecuador and for the Oligocene–Miocene boundary of Tropical America. The elasmobranch habitat preferences and paleobathymetric analyses support the hypothesis that the Montañita-Olón site was likely deposited in an outer neritic (open shelf) environment. Future work and new findings could help improve and refine our interpretations. The assemblage from the Montañita-Olón site increases the fossil record of the region and represents a critical window into marine tropical vertebrate faunas in the ECP during the OMT, a critical moment in the evolutionary history of the marine biota of Tropical America.

**ACKNOWLEDGEMENTS**

We thank students and University (UPSE) staff involved in the fieldwork, especially Xavier Guncay. Thanks to the curators of the paleontological collections referred to in the methods section for their valuable assistance and permission to review fossil collections. Thanks to Marcelo Sánchez-Villagra and the Evolutionary Morphology and Palaeobiology group at the Palaeontological Institute and Museum at the University of Zurich for their support. Special thanks to the Center for Microscopy and Image Analysis of the University of Zurich for their assistance and support performing the scanning electron microscopy analysis. To Ana Balcarcel for the revision of the preliminary version of the
manuscript. We greatly appreciate comments and suggestions from Editor Andrew A. Farke and the two reviewers David Cicimurri and Victor Perez.

ADDITIONAL INFORMATION AND DECLARATIONS

Funding
This study was supported by the project from the Universidad Estatal Península de Santa Elena, (Instituto de investigación Científica y Desarrollo Tecnológico) code 91870000.0000.381019. The Spanish Research Projects CGL 2015-68333-P and CGL2016-76431-P (Agencia Estatal de Investigación (AEI), European Regional Development Fund and the European Union (UE)), the Research Groups CSIC 64 1538 and CAM-UCM 910607, and the Generalitat de Catalunya (Fundación Catalana para la Investigación (CERCA)) and Beatriu de Pinós contract 2017 BP 00223 from the Agencia de Gestión de Ayudas Universitarias y de Investigación funded Juan Abella. The Becas-Chile Scholarship, Advanced Human Capital Program of the National Comission for Scientific and Technological Research (CONICYT, Chile) was granted to Jaime Andres Villafaña Navea. The funders had no role in study design, data collection and analysis, decision to publish, or preparation of the manuscript.

Grant Disclosures
The following grant information was disclosed by the authors:
Universidad Estatal Península de Santa Elena (Instituto de investigación Científica y Desarrollo Tecnológico): 91870000.0000.381019.
The Spanish Research Projects (Agencia Estatal de Investigación (AEI), European Regional Development Fund and the European Union (UE)): CGL2015-68333-P and CGL2016-76431-P.
Research Groups: CSIC 64 1538 and CAM-UCM 910607.
Generalitat de Catalunya (Fundación Catalana para la Investigación (CERCA)).
Agencia de Gestión de Ayudas Universitarias y de Investigación: 2017 BP 00223.
The Becas-Chile Scholarship, Advanced Human Capital Program of the National Commission for Scientific and Technological Research (CONICYT, Chile).

Competing Interests
The authors declare that they have no competing interests.

Author Contributions
- Jorge D. Carrillo-Briceño conceived and designed the experiments, performed the experiments, analyzed the data, prepared figures and/or tables, authored or reviewed drafts of the paper, fieldwork, and approved the final draft.
- Jaime A. Villafañe conceived and designed the experiments, performed the experiments, analyzed the data, prepared figures and/or tables, authored or reviewed drafts of the paper, and approved the final draft.
Carlos De Gracia conceived and designed the experiments, performed the experiments, analyzed the data, prepared figures and/or tables, authored or reviewed drafts of the paper, and approved the final draft.

F. Fernando Flores-Alcívar analyzed the data, authored or reviewed drafts of the paper, fieldwork, and approved the final draft.

René Kindlimann analyzed the data, authored or reviewed drafts of the paper, and approved the final draft.

Juan Abella conceived and designed the experiments, performed the experiments, analyzed the data, authored or reviewed drafts of the paper, fieldwork, and approved the final draft.

**Field Study Permissions**

The following information was supplied relating to field study approvals (i.e., approving body and any reference numbers):

Ecuadorian Instituto Nacional de Patrimonio Cultural (INPC) approved this field study (Excavation permit, Code: No 0039-DR5.INPC.2015).

**Data Availability**

The following information was supplied regarding data availability:

Raw data and script are available in Data S2.

All the fossil specimens described here are available at the “Museo Paleontológico Megaterio” (MPM-) at the Universidad Estatal Peninsula de Santa Elena, Ecuador. Every specimen or group of specimens is assigned with a catalog number using the acronyms “MPM-”. Data on fossil specimens with their respective catalog numbers are available in Table S3.

**Supplemental Information**

Supplemental information for this article can be found online at http://dx.doi.org/10.7717/peerj.9051#supplemental-information.

**REFERENCES**

Adnet S, Cappetta H, Guinot G, Di Sciara GN. 2012. Evolutionary history of the devilrays (Chondrichthyes: Myliobatiformes) from fossil and morphological inference. Zoological Journal of the Linnean Society 166(1):132–159 DOI 10.1111/j.1096-3642.2012.00844.x.

Agassiz JLR. 1833–1843. Recherches sur les Poissons fossiles. A Neuchâtel: Imprimerie Petitpierre.

Aguilera OA. 2010. Peces fósiles del caribe de venezuela. Washington, D.C.: Gorham Printing.

Aguilera Socorro OA, Ramos MIF, Paes ET, Costa SARF, Sánchez-Villagra MR. 2011. The neogene tropical America fish assemblage and the paleobiogeography of the Caribbean region. Swiss Journal of Palaeontology 130(2):217–240 DOI 10.1007/s13358-011-0020-9.

Aguilera O, Luz Z, Carrillo-Briceño JD, Kocsis L, Vennemann TW, De Toledo PM, Nogueira A, Amorim KB, Moraes-Santos Hisa, Polck MR, De Lourdes Ruivo M, Linhares AP, Monteiro-Neto C. 2017a. Neogene sharks and rays from the Brazilian 'Blue Amazon'. PLOS ONE 12(8):e0182740 DOI 10.1371/journal.pone.0182740.

Aguilera O, Silva GOA, Lopes RT, Machado AS, Dos Santos TM, Marques G, Bertucci T, Aguiar T, Carrillo-Briceño J, Rodrigue F, Jaramillo C, Vermeij GJ. 2017b. Neogene
proto-caribbean porcupinefishes (Diodontidae). *PLOS ONE* **12**(7):e0181670 DOI 10.1371/journal.pone.0181670.

Bacon CD, Silvestro D, Jaramillo C, Smith BT, Chakrabarty P, Antonelli A. 2015. Biological evidence supports an early and complex emergence of the Isthmus of Panama. *Proceedings of the National Academy of Sciences of the United States of America* **112**(19):6110–6115 DOI 10.1073/pnas.1423853112.

Bannikov AF. 1985. Fossil scombrids of the USSR. *Trudy Paleontologicheskogo Instituta: Akademiya Nauk SSSR* **210**:1–111 [in Russian].

Baut JP, Génault B. 1999. Les Elasmobranches des Sables de Kerniel (Rupélien), à Gellik, Nord Est de la Belgique. *Memoirs of the Geological Survey of Belgium* **45**:1–61.

Bellwood DR, Schultz O, Sisqueira AC, Cowman PF. 2019. A review of the fossil record of the Labridae. *Annalen des Naturhistorischen Museums in Wien, Serie A* **121**:125–193.

Bloch ME, Schneider JG. 1801. M.E. Blochii Systema Ichthyologiae iconibus ex illustratum: post obitum auctoris opus inchoatum absolvit, corriguit, interpolavit. Saxo: J.G. Schneide.

Bonnaterre JP. 1788. *Tableau encyclopédique et methodique des trois règnes de la nature Ichtyologie*. Paris: Chez Panckoucke.

Bristow CR. 1975. On the age of the Zapotal sands of Southwest Ecuador. *Newsletters on Stratigraphy* **42**(2):119–134 DOI 10.1127/nos/4/1975/119.

Cadena E, Abella J, Gregori M. 2018. The first Oligocene sea turtle (Pan-Cheloniidae) record of South America. *PeerJ* **6**:e4554 DOI 10.7717/peerj.4554.

Cappetta H. 1970. Les Sélaciens du Miocène de la région de Montpellier. *Palaeovertebrata, Mémoire Extraordinaire* **1970**:1–39.

Cappetta H. 2012. *Chondrichthyes—mesozoic and cenozoic elasmobranchii: teeth*. Munich: Dr. Friedrich Pfeil.

Cappetta H, Gregorová R, Adnet S. 2016. New selachian assemblages from the Oligocene of Moravia (Czech Republic). *Neues Jahrbuch für Geologie und Paläontologie-Abhandlungen* **280**(3):259–284 DOI 10.1127/njgpa/2016/0579.

Carrillo-Briceño JD, González-Barba G, Landaeta MF, Nielsen SN. 2013. Condritios fósiles del Plioceno Superior de la Formación Horcón, Región de Valparaíso, Chile central. *Revista Chilena de Historia Natural* **86**(2):191–206 DOI 10.4067/S0716-078X2013000200008.

Carrillo-Briceño JD, Aguilera OA, Rodríguez F. 2014. Fossil Chondrichthyes from the central eastern Pacific Ocean and their paleoceanographic significance. *Journal of South American Earth Sciences* **51**:76–90 DOI 10.1016/j.jsames.2014.01.001.

Carrillo-Briceño JD, De Gracia C, Pimiento C, Aguilera OA, Kindlimann R, Santamarina P, Jaramillo C. 2015a. A new late Miocene chondrichthyan assemblage from the Chagres Formation, Panama. *Journal of South American Earth Sciences* **60**:56–70 DOI 10.1016/j.jsames.2015.02.001.

Carrillo-Briceño JD, Maxwell E, Aguilera OA, Sánchez R, Sánchez-Villagra MR. 2015b. Sawfishes and other elasmobranch assemblages from the Mio-Pliocene of the South Caribbean (Urumaco Sequence, Northwestern Venezuela). *PLOS ONE* **10**(10):e0139230 DOI 10.1371/journal.pone.0139230.

Carrillo-Briceño JD, Aguilera OA, De Gracia C, Aguirre-Fernández G, Kindlimann R, Sánchez-Villagra MR. 2016a. An early Neogene Elasmobranch fauna from the southern Caribbean (Western Venezuela). *Palaeontologia Electronica* **19**(2):1–32 DOI 10.26879/664.
Gibbes RW. 1849. Monograph of the fossil Squalidae of the United States. *Journal of the Academy of Natural Sciences of Philadelphia* 1(2):191–206.

Gottfried MD, Fordyce RR. 2001. An associated specimen of *Carcharodon angustidens* (Chondrichthyes, Lamnidae) from the late Oligocene of New Zealand, with comments on *Carcharodon* interrelationships. *Journal of Vertebrate Paleontology* 21(4):730–739 DOI 10.1671/0272-4634(2001)021[0730:AASOCA]2.0.CO;2.

Iturralde-Vinent MA, Hubbell G, Rojas R. 1996. Catalogue of Cuban fossil Elasmobranchii (Paleocene to Pliocene) and paleogeographic implications of their lower to middle Miocene occurrence. *Boletín de la Sociedad Jamaicana de Geología* 31:7–21.

Jaramillo C, Montes C, Cardona A, Silvestro D, Antonelli A, Bacon CD. 2017. Comment (1) on “Formation of the Isthmus of Panama” by O’Dea et al. *Science Advances* 3(6):e1602321 DOI 10.1126/sciadv.1602321.

Johnson KG, Sánchez-Villagra MR, Aguilera OS. 2009. The Oligocene-Miocene transition on coral reefs in the falcon basin (NW Venezuela). *PALAIOS* 24(1):59–69 DOI 10.2110/palo.2008.p08-004r.

Kruckow T, Thies D. 1990. Die neoselachier der paleokaribik (Pisces: Elasmobranchii). *Courier Forschungsinstitut Senckenberg* 119:1–102.

Landini W, Collareta A, Di Celma C, Malinverno E, Urbina M, Bianucci G. 2019. The early Miocene elasmobranch assemblage from Zamaca (Chilcatay Formation, Peru). *Journal of South American Earth Sciences* 91:352–371 DOI 10.1016/j.jsames.2018.08.004.

Laurito C. 1999. Los seláceos fósiles de la localidad de Alto Guayacán (y otros ictiolitos asociados), Mioceno Superior-Plioceno Inferior de la Formación uscari, provincia de Limón, Costa Rica. San José: C.A. Laurito.

Laurito MCA, Calvo C, Valero Al, Calvo A, Chacón R. 2014. Ictiofauna del Mioceno inferior de la localidad de Pacuare de Tres Equis, Formación Río Banano, provincia de Cartago, Costa Rica, y descripción de un nuevo género y una nueva especie de scaridae. In: *Revista Geológica de América Central*. Vol. 50. 153–192.

Le Hon H. 1871. *Préliminaires d’un mémoire sur les poissons tertiaires de Belgique*. Brussels: H. Merzbach.

Leigh EG, O’Dea A, Vermeij GJ. 2014. Historical biogeography of the Isthmus of Panama. *Biological Reviews* 89(1):148–172 DOI 10.1111/brv.12048.

Leriche M. 1938. Contribution à L’étude des Poissons fossiles des pays riverains de la Méditerranée américaine, Venezuela, Trinité, Antilles. *Mexique Mémoires de la Sociéte Paléontologique du Suisse* 61:1–52.

Lessios HA. 2008. The great American Schism: divergence of marine organisms after the rise of the central American Isthmus. *Annual Review of Ecology, Evolution, and Systematics* 39(1):63–91 DOI 10.1146/annurev.ecolsys.38.091206.095815.

Longbottom AE. 1979. Miocene shark’s teeth from Ecuador. *Bulletin of the British Museum (Natural History)*, *Geology* 32:57–70.

Michelotti G. 1847. *Description des fossils des terrains miocènes de l’Italie septentrionale*. Leiden: A. Arnz & Comp.

Monsch K. 2004. Revision of the scombroid fishes from the Cenozoic of England. *Earth and Environmental Science Transactions of the Royal Society of Edinburgh* 95(3–4):445–489 DOI 10.1017/S0265359300001164.

Monsch KA, Bannikov AF. 2011. New taxonomic synopses and revision of the scombroid fishes (Scombroidi, Perciformes), including billfishes, from the Cenozoic of territories of the former...
USSR. Earth and Environmental Science Transactions of the Royal Society of Edinburgh 102(4):253–300 DOI 10.1017/S1755691011010085.

Musick JA, Harbin MM, Compagno LJV. 2004. Historical zoogeography of the Selachii. In: Carrier JC, Musick JA, Heithaus MR, eds. Biology of Sharks and their Relatives. Washington, D.C.: CRC Press, 33–78.

Nolf D, Brzobohatý R. 1994. Fish otoliths as paleobathymetric indicators. Paleontologia y Evolutió 24–25:255–264.

Olsson AA. 1931. Contributions to the tertiary paleontology of northern Peru: part 4—the Peruvian Oligocene. Bulletins of American paleontology 17(63):97–260.

Perez VJ, Pimiento C, Hendy A, González-Barba G, Hubbell G, MacFadden BJ. 2017. Late Miocene chondrichthyans from Lago Bayano, Panama: functional diversity, environment and biogeography. Journal of Paleontology 91(3):512–547 DOI 10.1017/jpa.2017.5.

Pfeil FH. 1983. Zahnmorphologische Untersuchungen an rezenten und fossilen Hainen der Ordnungen Chlamydoselachiformes und Echinorhiniformes. Palaeo Ichthyologica 1:1–315.

Probst J. 1879. Beiträge zur Kenntniss der fossilen Fische aus der Molasse von Baltringen. Hayeshefte des Vereins für vaterländische Naturkunde in Württemberg 35:127–191.

Purdy R, Clellan JHM, Schneider VP, Applegate SP, Meyer R, Slaughter R. 2001. The Neogene sharks, rays and bony fishes from Lee Creek Mine, Aurora, North Carolina. Smithsonian Contributions to Paleobiology 90:71–202.

R Core Team. 2019. R: a language and environment for statistical computing. Vienna: The R Foundation for Statistical Computing. Available at https://www.R-project.org/.

Rafinesque CS. 1810. Caratteri di alcuni nuovi generi e noue specie di Animali e Piante della Sicilia con varie osservazioni sopra I medesimi. Palermo: Per le stampe di Sanfilippo.

Reed D. 1946. New species of fossil shark from New Jersey. Notulae Naturae of the Academy of Natural Sciences of Philadelphia 172:1–3.

Reinecke T, Stapf H, Raisch M. 2001. Die selachier und chimären des unteren meeressandes und schleichsandes im mainzer becken (alzey- und stadecken-formation, rupelium, unteres oligozän). Palaeontos 1:1–73.

Reinecke T, Balsberger M, Beaurry B, Pollerspoeck J. 2014. The elasmobranch fauna of the Thalberg Beds, early Egerian ( Chattian, Oligocene), in the Subalpine Molasse Basin near Siegsdorf, Bavaria, Germany. Palaeontos 26:1–127.

Sánchez-Villagra MR, Burnham RJ, Campbell DC, Feldmann RM, Gaffney ES, Kay RF, Lozsán R, Purdy R, Thewissen JGM. 2000. A new near-shore marine fauna and flora from the early neogene of northwestern Venezuela. Journal of Paleontology 74(5):957–968 DOI 10.1017/S002233600003126.

Santos RS, Travassos H. 1960. Contribuição à Paleontologia do estado do Pará. Peixes fósseis da Formação Pirabas. Rio de Janeiro: Ministério da Agricultura, Departamento Nacional da Produção Mineral, Divisao de Geologia e Mineralogia Monografia.
Santos RS, Salgado MS. 1971. Contribuição à paleontologia do estado do Pará: Novos restos de peixes da Formação Pirabas. Boletim do Museu Paraense Emílio Goeldi 16:1–13.

Shimada K, Chandler RE, Lam OLT, Tanaka T, Ward DJ. 2017. A new elusive otodontid shark (Lamniformes: Otodontidae) from the lower Miocene, and comments on the taxonomy of otodontid genera, including the 'megatoothed' clade. Historical Biology 29(5):704–714 DOI 10.1080/08912963.2016.1236795.

Schultz O, Bellwood DR. 2004. Trigonodon oweni and Asima jugleri are different parts of the same species Trigonodon jugleri, a Chiseltooth Wrasse for the Lower and Middle Miocene in Central Europe (Osteichthyes, Labridae, Trigonodontinae). Annalen des Naturhistorischen Museums in Wien Serie A 105:287–305.

Schwarzhans W, Aguilera O. 2013. Otoliths of the Myctophidae from the Neogene of tropical America. Palaeo Ichthyologica 13:83–150.

Staig F, Hernández S, López P, Villafañ a JA, Varas C, Soto LP, Carrillo-Briceño JD. 2015. Late Neogene elasmobranch fauna from the Coquimbo Formation. Chile Revista Brasileira de Paleontologia 18(2):261–272 DOI 10.4072/rbp.2015.2.07.

Szabó M, Kocsis L. 2016a. A new middle Miocene selachian assemblage (Chondrichthyes, Elasmobranchii) from the Central Paratethys (Nyírad, Hungary): implications for temporal turnover and biogeography. Geologica Carpathica 67(6):573–594 DOI 10.1515/geoca-2016-0036.

Szabó M, Kocsis L. 2016b. A preliminary report on the early Oligocene (Rupelian, Kiscellian) selachians from the Kiscell Formation (Buda Mts, Hungary), with the re-discovery of Wilhelm Weiler’s shark teeth. Fragmenta Palaeontologica Hungarica 33:31–64 DOI 10.17111/FragmPalHung.2016.33.31.

Tanaka Y, Abella J, Aguirre-Fernández G, Gregori M, Fordyce RE. 2017. A new tropical Oligocene dolphin from Montañita/Olón, Santa Elena, Ecuador. PLOS ONE 12(12):e0188380 DOI 10.1371/journal.pone.0188380.

Underwood CJ, Schlögl S. 2013. Deep-water chondrichthyans from the early miocene of the Vienna basin (central paratethys, Slovakia). Acta Palaeontologica Polonica 58(3):487–509.

Vialle N, Adnet S, Cappetta H. 2011. A new shark and ray fauna from the middle Miocene of Mazan, Vaucluse (southern France) and its importance in interpreting the paleoenvironment of marine deposits in the southern Rhodanian Basin. Swiss Journal of Palaeontology 130(2):241–258 DOI 10.1007/s13358-011-0025-4.

Villafañ a JA, Rivadeneira MM. 2018. The modulating role of traits on the biogeographic dynamics of chondrichthyans from the Neogene to the present. Paleobiology 44(2):251–262 DOI 10.1017/pab.2018.7.

Villafañ a JA, Nielsen SN, Klug S, Kriwet J. 2019. Early Miocene cartilaginous fishes (Chondrichthyes: Holocephali, Elasmobranchii) from Chile: diversity and paleobiogeographic implications. Journal of South American Earth Sciences 96:102317 DOI 10.1016/j.jsames.2019.102317.

Welton BJ. 1974. Heptanchias Howellii (Reed, 1946) (Selachii, Hexanchidae) in the Eocene of the United States and British Columbia. Paleobios 17:1–15.

Welton BJ. 1979. Late Cretaceous and Cenozoic squalomorphii of the Northw est Pacific Ocean. Ph.D. thesis, University of California.

Whittaker JE. 1988. Benthic Cenozoic foraminifera from Ecuador: taxonomy and distribution of smaller benthic foraminifera from Coastal Ecuador (late Oligocene-late Pliocene). London: British Museum (Natural History).
Witt C, Reynaud JY, Barba D, Poujol M, Aizprua C, Rivadeneira M, Amberg C. 2019. From accretion to forearc basin initiation: the case of SW Ecuador, Northern Andes. *Sedimentary Geology* 379:138–157 DOI 10.1016/j.sedgeo.2018.11.009.

**Woodward AS. 1889.** *Catalogue of the fossil fishes in the British Museum: part. I.* Vol. I. London: British Museum (Natural History).

**Woodward AS. 1901.** *Catalogue of the fossil fishes in the British Museum (Natural History), Vol. 4, containing the Actinopterygian Teleostomi of the suborders Isospondyli (in part), Ostariophysii, Apodes, Peresoces, Hemibranchii, Acanthopterygii, and Anacanthini.* London: British Museum (Natural History).