THE ATMOSPHERIC CIRCULATION AND CLIMATE OF TERRESTRIAL PLANETS ORBITING SUN-LIKE AND M-DWARF STARS OVER A BROAD RANGE OF PLANETARY PARAMETERS

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ABSTRACT

The recent detections of temperate terrestrial planets orbiting nearby stars and the promise of characterizing their atmospheres motivates a need to understand how the diversity of possible planetary parameters affects the climate of terrestrial planets. In this work, we investigate the atmospheric circulation and climate of terrestrial exoplanets orbiting both Sun-like and M-dwarf stars over a wide swath of possible planetary parameters, including the planetary rotation period, surface pressure, incident stellar flux, surface gravity, planetary radius, and cloud particle size. We do so using a general circulation model (GCM) that includes non-grey radiative transfer and the effects of clouds.

The results from this suite of simulations generally show qualitatively similar dependencies of circulation and climate on planetary parameters as idealized GCMs, with quantitative differences due to the inclusion of additional model physics. Notably, we find that the effective cloud particle size is a key unknown parameter that can greatly affect the climate of terrestrial exoplanets. We confirm a transition between low and high dayside cloud coverage of synchronously rotating terrestrial planets with increasing rotation period. We determine that this cloud transition is due to eddy-driven convergence near the substellar point and should not be parameterization-dependent. Finally, we compute full-phase light curves from our simulations of planets orbiting M-dwarf stars, finding that changing incident stellar flux and rotation period affect observable properties of terrestrial exoplanets. Our GCM results can guide expectations for planetary climate over the broad range of possible terrestrial exoplanets that will be observed with future space telescopes.

Subject headings: hydrodynamics - methods: numerical - planets and satellites: terrestrial planets - planets and satellites: atmospheres

1. INTRODUCTION

We are approaching an era in which the atmospheres of terrestrial exoplanets can be characterized by astronomical observations. This is enabled by the recent detection of a variety of nearby stellar systems with temperate terrestrial exoplanets, including TRAPPIST-1 (Gillon et al. 2017), Proxima Centauri (Anglada-Escudé et al. 2016), and GJ 1132 (Berta-Thompson et al. 2015). These discoveries are timely, as the James Webb Space Telescope will be launched in the coming years and will have the capability to detect whether or not terrestrial planets orbiting M-dwarf stars have atmospheres (Seager & Deming 2009) and could potentially characterize the atmospheres of these planets (Kreidberg & Loeb 2016, Morley et al. 2017, Diamond-Lowe et al. 2018, de Wit et al. 2018). Additionally, planned future space-based telescopes (e.g. LUVOIR, HabEx) would have the ability to detect reflected light spectra of temperate terrestrial planets orbiting Sun-like stars and constrain their atmospheric properties (Feng et al. 2018).

Motivated by the prospect of future observations, it is crucial to develop a theoretical understanding of the atmospheres of terrestrial planets. This includes knowing how properties of exoplanet atmospheres vary over the wide potential phase space and how physical changes in atmospheric properties manifest in observations. In principle, there are a variety of a priori unknown planetary parameters that could affect the atmospheric dynamics. For example, the rotation period affects the width of the tropical circulation (Held 2000, Kaspi & Showman 2015), the size and distribution of extratropical eddies (Eady 1949, Charney 1967, Kaspi & Showman 2015), and the number of extratropical jets (Williams 1978, Cho & Polvani 1996, Chemke & Kaspi 2015b). Meanwhile, the atmospheric mass affects the strength of the atmospheric circulation and the surface temperature distribution (Chemke et al. 2016, Chemke & Kaspi 2017). Additionally, the spin state of terrestrial exoplanets orbiting M-dwarf stars varies substantially from that of planets orbiting Sun-like stars, as the spins of planets orbiting M-dwarfs are affected by tidal dissipation due to their close-in orbits (Lee & et al. 2015).

How the atmospheric dynamics and climate of planets orbiting Sun-like stars vary with planetary parameters has been explored in a wide variety of circulation models. As a first step, many modeling studies utilize idealized general circulation models (GCMs) which have reduced complexity in radiative transfer and cloud treatment relative to full physics simulations in order to understand the basic physics controlling atmospheric properties (Schneider 2006, Kaspi & Showman 2015, Chemke & Kaspi 2017). Notably, Kaspi & Showman (2015) varied a wide swath of planetary parameters (rotation period, incident stellar flux, surface pressure, surface gravity, and radius) and found that each parameter can significantly alter the atmospheric circulation of exoplanets. Yang et al. (2014) used the more complex Community Atmosphere Model GCM, which includes non-grey radiative effects and cloud parameterizations. Yang et al. found that in the Venus-like case of very long rotation...
periods high cloud coverage builds up on the dayside of the planet, significantly increasing the planetary albedo and therefore the incident stellar flux at which the planet can stay habitable (i.e., the inner edge of the habitable zone, Kasting et al. 1993). Similarly, Jansen et al. (2018) found using the ROCKE-3D GCM (Way et al. 2017) that the surface temperature of Earth-like planets decreases with increasing rotation period. The inner edge of the habitable zone for planets around Sun-like stars has been studied using a variety of complex GCMs (Leconte et al. 2013a, Wolf & Toon 2015, Popp et al. 2016, Way et al. 2016b, Wolf et al. 2017), with all models finding inner edges that are at higher incident stellar flux relative to 1D models (e.g., Kopparapu et al. 2013) due to the effects of clouds and reduced relative humidity.

The atmospheric circulation of planets orbiting M-dwarf stars is vastly different from that of planets orbiting Sun-like stars due to the likely spin-synchronized or otherwise tidally damped nature of their spin states. A wide variety of GCM work has been performed to study planets orbiting M-dwarfs (Joshi et al. 1997, Merlis & Schneider 2010, Selsis et al. 2011, Leconte et al. 2013b, Yang et al. 2013, Hu & Yang 2014, Wang et al. 2014, Koll & Abbot 2015, Carone et al. 2015, Koll & Abbot 2016, Kopparapu et al. 2016, Turbet et al. 2016, Boutle et al. 2017, Del Genio et al. 2018, Kopparapu et al. 2017, Noda et al. 2017, Wolf 2017, Haqq-Misra et al. 2018, Lewis et al. 2018), studying both synchronous rotators and planets in a 3:2 spin-orbit resonance. As for planets orbiting Sun-like stars, the rotation period has been found to play a key role in the atmospheric circulation of planets orbiting M-dwarf stars. The effect of rotation period on the circulation of planets orbiting M-dwarf stars has been analyzed using models with a hierarchy of radiative schemes, from those with radiative relaxation (Carone et al. 2015) to a double-grey radiation scheme without clouds (Noda et al. 2017), and models including sophisticated radiative transfer (Yang et al. 2013, Kopparapu et al. 2017). The latter simulations have found that, as for Venus-like planets, dayside cloud coverage on slowly rotating M-dwarf planets is large enough to significantly increase the planetary albedo. This increased albedo moves the inner edge of the habitable zone closer in for planets orbiting M-dwarf stars. In general, slowly rotating planets tend to have larger-scale atmospheric features due to the broader latitudinal scale of their mean circulation and larger eddy length scales, which provides a heuristic explanation why slowly rotating M-dwarf planets have greater dayside cloud coverage (Leconte et al. 2013b, Haqq-Misra et al. 2018). This effect is especially strong for planets orbiting earlier-type (hotter) M-dwarfs, as these planets have wider orbits for a given incident stellar flux, with correspondingly longer rotation periods (assuming synchronous rotation).

In this work, we study how the atmospheric circulation and climate of planets orbiting both Sun-like stars and M-dwarfs vary with planetary parameters. We utilize a complex GCM (ExoCAM, a modified version of the Community Atmosphere Model with an improved radiation scheme for exoplanet atmospheres and cloud and sea ice parameterizations) and conduct simulations over a wide region of planetary parameter space, varying the incident stellar flux, rotation period, surface pressure, planetary radius, surface gravity, and effective liquid water cloud particle size. Our simulations of planets around Sun-like stars can hence be thought of as a more complex rendition of the simulations of Kaspi & Showman (2015), now including the effects of clouds, non-grey radiative transfer, and sea ice. Additionally, we vary the effective cloud particle size, as the cloud particle size is a priori unknown in exoplanet atmospheres and can greatly modify the cloud radiative effect. Our simulations for synchronously rotating planets orbiting M-dwarf stars with varying rotation period are comparable to the idealized works of Carone et al. (2015) and Noda et al. (2017), and explore how the conclusions of Kaspi & Showman (2015) extend to synchronously rotating planets. We also calculate model phase curves from our simulations of planets orbiting M-dwarfs to show how varying planetary parameters can affect astronomically observable properties of the planet.

The outline of this paper is as follows. Section 2 describes our model setup and the parameter space studied by our large suite of numerical circulation models. In Section 3, we show our results for simulations of planets orbiting Sun-like stars and how their climate varies with rotation period, surface pressure, and effective liquid cloud particle size. In Section 4, we show how the climate of spin-synchronized planets orbiting M-dwarf stars varies as the incident stellar flux and rotation period are consistently changed, the rotation period alone is changed, and the surface pressure alone is changed. Note that we show further results from our suite of simulations in the Appendix. Then, in Section 5 we compare our results to previous work, discuss the transition between small to large dayside cloud coverage in our simulations of planets orbiting M-dwarf stars, and compute model full-phase light curves from our simulations of planets orbiting M-dwarf stars. Lastly, we summarize the main points of our work in Section 6.

2. METHODS

2.1. Model setup

To simulate the atmospheres of terrestrial exoplanets, we use ExoCAM1 (Wolf & Toon 2015, Kopparapu et al. 2016, 2017, Wolf et al. 2017, Wolf 2017). ExoCAM is a modified version of the Community Atmosphere Model version 4 (CAM4) with correlated-k radiative transfer, updated spectral coefficients using the HITRAN 2012 database, and a novel treatment of water vapor continuum absorption. ExoCAM has been used to simulate the climates of Earth (Wolf & Toon 2015), the TRAPPIST-1 system (Wolf 2017), and exoplanets near the inner edge of the habitable zone (Kopparapu et al. 2017, Haqq-Misra et al. 2018).

In this work, we use the same version of ExoCAM as in Kopparapu et al. (2017) and Haqq-Misra et al. (2018), which has a finite volume dynamical core with a horizontal resolution of 4° × 5° and 40 vertical levels from the surface to a top pressure of 1 mbar. This resolution is the same as that used in previously published ExoCAM simulations (Kopparapu et al. 2017, Wolf et al. 2017, Haqq-Misra et al. 2018), but note that it is necessarily lower than the resolution of some idealized models that have been applied to study exoplanet climate (e.g.,

1 https://github.com/storyofthewolf/ExoCAM
Kaspi & Showman 2015, Koll & Abbot 2015, Chemke & Kaspi 2017) due to the complexity of the ExoCAM model. We assume an atmosphere that is comprised purely of N$_2$ and H$_2$O overlying an aquaplanet slab ocean with a depth of 50 m. As in Kopparapu et al., we neglect ocean heat transport and allow sea ice to form. We use sub-grid parameterizations for clouds from Rasch & Kristjánsson (1998) and for convection from Zhang & McFarlane (1995). Note that our results may quantitatively differ from those of other GCMs with different cloud parameterizations. As a result, we focus on broad qualitative changes in how the climate depends on planetary parameters due to robust physical mechanisms. We assume that planets have zero obliquity and that their orbit has zero eccentricity. For all of the simulations presented here, we use a time step of 30 minutes, with 16 dynamical substeps per time step and the radiative transfer calculated every 3 time steps. We run simulations until they reach a statistically steady state. For all of the simulations presented here, we use a time step of 30 minutes, with 16 dynamical substeps per time step and the radiative transfer calculated every 3 time steps. We run simulations until they reach a statistically steady state, typically 45 – 50 Earth years. All results shown in this work are time-averaged over the last ten years of model time.

### 3. CIRCULATION AND CLIMATE OF TERRRESTRIAL PLANETS ORBITING SUN-LIKE STARS

#### 3.1. Zonally averaged temperature and circulation strength

In the following sections, we show how varying the rotation period, surface pressure, effective liquid cloud particle size, incident stellar flux, planetary radius, and surface gravity affects the atmospheric circulation and climate of planets orbiting Sun-like stars.

###### 3.1.1. Rotation period

We show the zonal-mean temperature, potential temperature, zonal wind, and mass streamfunction for planets orbiting Sun-like stars with varying rotation period from 0.5-16 Earth days$^2$ in Figure 1. We find that the equator-to-pole temperature difference decreases with increasing rotation period (decreasing rotation rate), indicative of enhanced latitudinal heat transport at higher rotation period. The reduced equator-to-pole temperature difference at higher rotation period is due to the

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| Planetary Parameter | Unit | Parameter Values |
|---------------------|------|------------------|
| Rotation period     | Earth Days | 0.5, 1, 2, 4, 8, 16 |
| Surface pressure    | Bars  | 0.25, 0.5, 1, 2, 4 |
| Incident stellar flux | Earth flux | 0.544, 0.667, 0.816, 1 |
| Planetary radius    | Earth radii | 0.5, 0.707, 1, 1.414, 2 |
| Liquid cloud effective radius | Earth surface gravity | 0.5, 0.707, 1, 1.414 |

**Table 1**

Values of planetary parameters used in the suite of general circulation model experiments of planets orbiting both Sun-like and M-dwarf stars.
Fig. 1.— Zonal-mean temperature, potential temperature, zonal wind, and streamfunction for simulated planets orbiting a Sun-like star with varying rotation periods from 0.5 - 16 Earth days. The left-hand panels show zonal-mean temperature (colors) and potential temperature (contours). All panels share a color scale, and the potential temperature contours run from 200-500 K, in 30 K intervals. The y-axis of each panel is normalized pressure $\sigma = p/p_s$, and the x-axis is latitude. The right-hand panels show mass streamfunction (colors) and zonal-mean zonal wind (contours). The zonal-mean zonal wind is shown in 10 m s$^{-1}$ intervals with a maximum of $\pm 100$ m s$^{-1}$, and with solid contours representing positive values and dashed contours representing negative values. The left-hand panels show that the width of the warm equatorial region increases with increasing rotation period, leading to a decrease in the equator-to-pole temperature contrast. The right-hand panels show that the Hadley circulation increases in width with increasing rotation period, leading to extratropical jets that lie at higher latitudes.

increase in eddy length scale with increasing rotation period (Eady 1949, Pedlosky 1987, Chemke & Kaspi 2015a), which more efficiently carry heat poleward than smaller eddies (Kaspi & Showman 2015), along with the expansion of the Hadley cell with increasing rotation period (Held & Hou 1980). Additionally, the peak zonal wind speed moves to higher latitudes with increased rotation period due to the increased size of the Hadley cell, with the peak zonal wind occurring poleward of the Hadley cell edge. Lastly, the speed of the subtropical zonal jets does not increase with increasing mass streamfunction, instead reaching a maximum wind speed at a rotation period of 1 day.

As discussed in Kaspi & Showman (2015), the non-monotonic trend in subtropical jet speed with rotation period is due to a trade-off between the increase in rotation rate leading to an increase in angular momentum of poleward-moving air in the Hadley cell and the narrowing of the Hadley cell causing the jet to acquire less mean angular momentum. Ignoring the effect of eddies, air moving poleward on slowly rotating planets does not have as much angular momentum as on faster rotating planets. As a result, the angular momentum-conserving jet speed is slower on more slowly rotating planets, since it scales with the inverse of the rotation period (Held & Hou 1980). However, for very fast rotating planets the jet speed is inhibited by eddies, leading to a jet speed that is weaker than the angular-momentum conserving limit (Kaspi & Showman 2015). As a result, there is a transition between the slowly rotating and fast rotating regimes in which the jet speeds go from increasing to decreasing in strength with increasing rotation period.

The findings discussed above all agree qualitatively with the results of Kaspi & Showman (2015), who used an idealized GCM with double-grey radiative transfer and no clouds to determine how the atmospheric dynamics of planets orbiting Sun-like stars are affected by varying planetary parameters. We will quantitatively compare our results for planets with varying rotation period to those of Kaspi & Showman (2015) in Section 5.1.

3.1.2. Surface pressure

In Figure 2 we show how the climate and circulation of planets orbiting Sun-like stars depends on varying the surface pressure from 0.25 – 4 bars. We find that planets with lower surface pressures are hotter, and that planets with surface pressures $\geq$ 2 bars are so cold that they
are in a fully glaciated (snowball) state. This is because the Rayleigh scattering of N$_2$ increases with increasing atmospheric mass, as was shown to affect the inner edge of the habitable zone by Kopparapu et al. (2014). As a result, here we show that Rayleigh scattering also affects the outer edge of the habitable zone by potentially leading planets with no CO$_2$ and zero obliquity to be fully ice-covered at large surface pressures.

Note that our result of surface cooling with increasing atmospheric mass differs from that of Goldblatt et al. (2009), Li et al. (2009), Charnay et al. (2013), Wolf & Toon (2014), Chemke et al. (2016), who found that higher atmospheric masses early in Earth’s history would increase the surface temperature. However, we are not considering an Earth-like setup and ignore CO$_2$, so even though we include the pressure broadening effect of H$_2$O the combined scattering effect of clouds and the background N$_2$ gas leads to cooling. This also agrees with the expectation that pressure broadening of CO$_2$ is dominant over the pressure broadening of H$_2$O for Earth (Li et al. 2009). Note that we do not expect that our results are generally applicable due to the lack of CO$_2$ in our simulations. However, we show that without the pressure broadening effect of CO$_2$ the climate of terrestrial planets should not warm with increasing surface pressure, the opposite of what is expected to occur on early Earth.

We also find that the Hadley cell strength is non-monotonic with increasing surface pressure, due to the greatly reduced Hadley cell strength in our fully glaciated simulations with surface pressures of 2 and 4 bars. Though as in Chemke & Kaspi (2017) we find that above a pressure of 1 bar, more massive atmospheres have reduced Hadley cell strengths, we also find evidence for an increase in the Hadley cell strength from surface pressures of 0.25 - 1 bar. Lastly, note that our simulations include the effect of pressure broadening while those in Chemke & Kaspi (2017) do not. The inclusion of pressure broadening in our simulations may be the cause of the warmer equator and relatively larger equator-to-pole temperature contrasts in our simulations relative to those of Chemke & Kaspi (2017).

### 3.1.3. Liquid cloud particle size

The effect of varying the effective liquid cloud particle radius from 7 – 21 µm on the circulation and climate of planets orbiting Sun-like stars is shown in Figure 3. The liquid water cloud particle size is a fixed parameter in our suite of simulations, with no assumed distribution, with the standard value of 14 µm taken to tune the model to an Earth-like climate. One might expect the effective cloud particle size to change the albedo of the planet due to the dependence of particle scattering properties on particle size and the wavelength of incident radiation. We find that increasing the effective cloud particle size from the canonical value of 14 µm to 21 µm causes the planet to warm by ≈ 7.5 K in the global mean, as the planet receives more incident stellar light. Given that the cloud coverage is greater at the equator than the pole the equator warms up relative to the polar regions causing a larger equator-to-pole temperature contrast, with a correspondingly stronger Hadley circulation and subtropical jet. Note that our simulations are not warm enough for moist effects to cause a reduction in the strength of the Hadley cell due to the amount of saturated water vapor increasing faster than it can be precipitated (Held & Soden 2006).

Conversely, decreasing the particle size from 14 µm to 7 µm leads to a ≈ 19 K cooler global-mean surface, with a significantly reduced Hadley cell strength and subtropical jet speed. This is because smaller effective particle sizes scatter more incident stellar light, leading to much less downward visible flux reaching the surface. Note that the exact value of the warming and cooling due to varying cloud particle size depends on the base state of the atmosphere, and in a warmer base climate the cloud particle effect may be larger in magnitude.

Based on the results shown in Figure 3, we find that the effective cloud particle size is a key unknown in the study of terrestrial exoplanet atmospheres. The effective cloud particle size directly affects how much incident radiation reaches the surface to drive circulation, which then feeds back on the cloud distribution. In reality, the cloud particle size distribution itself may also be affected by the circulation, including the spatial pattern of vertical mixing and heating/cooling. As a result, future work on the atmospheric circulation of terrestrial exoplanets should consider the effect of cloud parameterizations on climate and work toward building models that couple cloud microphysics and the radiative effect of clouds with atmospheric circulation.

### 3.1.4. Incident stellar flux

Figure 4 shows the zonal-mean circulation and climate for planets orbiting Sun-like stars with varying incident stellar flux from 0.544 – 1 $F_{\odot}$. We find that the surface temperature increases sharply with increasing incident flux, with a corresponding increase in the strength of the mass streamfunction and peak zonal wind speed. Kaspi & Showman (2015) also found an increase in zonal wind speed with increasing incident stellar flux, due to the greater energy input to drive atmospheric circulation. Additionally, we find a decrease in the pressure level at which the subtropical jet reaches its maximum with increasing incident stellar flux, also found by Kaspi & Showman (2015).
Fig. 4.— Zonal-mean temperature, potential temperature, zonal wind, and streamfunction for simulated planets orbiting a Sun-like star with varying incident stellar flux from $0.544 - 1 F_\odot$. The potential temperature and zonal-mean zonal wind contour intervals are the same as in Figure 1. The climate warms and the circulation strengthens with increasing incident stellar flux.

Fig. 5.— Zonal-mean temperature, potential temperature, zonal wind, and streamfunction for simulated planets orbiting a Sun-like star with varying planetary radius from $0.5 - 2 R_\oplus$. The potential temperature and zonal-mean zonal wind contour intervals are the same as in Figure 1. The equatorial temperature and equator-to-pole temperature contrast increase with increasing radius due to reduced eddy heat transport.

3.1.5. Planetary radius

Figure 5 shows how the circulation of planets orbiting Sun-like stars varies with changing planetary radius from $0.5 - 2 R_\oplus$. We find that the equatorial temperature increases with increasing radius, indicative of reduced equator-to-pole heat transport in the atmospheres of larger planets. We also find that the peak magnitude of the subtropical jet is more equatorward for higher radius, with a corresponding narrowing of the Hadley cell. Note that we find a reduction in the strength of the Hadley cell and subtropical jets at a radius of $0.707 R_\oplus$, where the circulation transitions from being characterized by subtropical jets that peak in strength at approximately $\pm 60^\circ$ latitude to being more like the Earth with subtropical jets that peak near $\pm 30^\circ$ latitude. In general, the trend we find of decreased equator-pole heat transport with increasing radius agrees with the simulations of Kaspi & Showman (2015), who found that the equator-to-pole temperature contrast of larger planets is enhanced due to the reduced eddy length scale relative to the radius of the planet.

3.1.6. Surface gravity

Figure 6 shows how the climate of simulated planets orbiting a Sun-like star depends on varying surface gravity from $0.5 - 1.414 g_\oplus$. We find that we keep surface pressure fixed in this set of experiments and hence allow the atmospheric mass to vary. However, this suite of simulations differs from the simulations with varying pressure in Section 3.1.2 due to the inclusion of pressure broadening in our radiative transfer scheme and the effects of gravity on the vertical static stability and convection due to the increase of convective available potential energy with gravity. We find that the equator-to-pole temperature contrast increases with increasing surface gravity. As a result, our simulated planets with larger surface gravities have stronger Hadley cells and correspondingly stronger subtropical jets. This is consistent with the simulations of Kaspi & Showman (2015) who found that the Hadley...
cell intensity increases with increasing surface gravity, qualitatively in accord with the theory of Held & Hou (1980). Note that our simulations do not cover as broad a range as Kaspi & Showman (2015), who found that this increase in Hadley cell strength with gravity continues up to a surface gravity 16 times that of Earth. Additionally, simulations with reduced gravity are hotter, in accord with the simulations of Chenke et al. (2016) who found that a larger atmospheric mass on early Earth would lead to warmer surface temperatures.

4. CIRCULATION AND CLIMATE OF SPIN-SYNCHRONIZED TERRESTRIAL PLANETS ORBITING M-DWARF STARS

4.1. Temperature & cloud coverage

In the following sections, we show how varying the rotation period and incident stellar flux together, rotation period alone, and surface pressure affect the climate of planets orbiting M-dwarf stars. We defer a discussion of how the incident stellar flux, planetary radius, surface gravity, and effective cloud particle size affect the climate of planets orbiting M-dwarf stars to Appendix A, as they are broadly similar to their impact on the climate of planets orbiting Sun-like stars.

4.1.1. Consistently varying rotation period and incident stellar flux

To analyze the climate of simulations of planets orbiting M-dwarf stars with consistently varying incident stellar flux and rotation period, we show maps of the temperature, total cloud fraction, and integrated cloud water (liquid and ice) path in Figures 7 and 8. Figure 7 shows results for planets orbiting an early type M-dwarf with an effective temperature of 4000 K, while Figure 8 shows results from simulations for planets orbiting a late-type M-dwarf star with an effective temperature of 2600 K. In both cases, we find that the temperature increases monotonically with increasing incident flux, as expected. However, in the case of planets orbiting a late M-dwarf star, we find that there is a transition between large sub-stellar cloud cover for our cases with lower incident flux and slower rotation and reduced cloud cover at higher flux and faster rotation.

This decrease in dayside cloud coverage with faster rotation corresponds to a dynamical transition of the atmosphere from a Rhines-rotator (with the Rossby deformation radius greater than the planetary radius, but the Rhines scale less than the planetary radius) regime to a rapidly rotating regime (Haqq-Misra et al. 2018). In the rapidly rotating regime, both the Rhines length scale and Rossby deformation radius are smaller than the planetary radius. Given that the Rhines length scale determines the length scale of the turbulent energy cascade, turbulence can lead to the formation of mid-latitude zonal jets and departure from day-night symmetry in this rapidly rotating regime. The Rossby deformation radius sets the characteristic length scale of the mean zonal circulation (Haqq-Misra et al. 2018). Consequently if the Rossby deformation radius is smaller than the planetary radius, the equatorial zonal circulation is confined to an equatorial strip, the length scale of which is set by the Rossby deformation radius. The shift in the dynamical state of the atmosphere from Rhines to rapidly rotating affects the dayside cloud distribution by reducing the latitudinal extent of the zonal circulation, leading to a reduction in the equatorward fluxes of momentum and moisture that cause vigorous convection at the substellar point, which is needed to form dayside clouds (Yang et al. 2013). The transition from Rhines rotating to rapidly rotating planets is not seen in our more slowly rotating simulations of planets orbiting an early M-dwarf star, as all of those simulations are in a slowly rotating regime (Haqq-Misra et al. 2018), where both the Rhines length scale and Rossby deformation radius are larger than the planetary radius. Note that all of our simulations of planets orbiting an early type M-dwarf star show significant dayside cloud coverage due to their slow rotation periods of 74.3-117.4 days. We will discuss the effects of rotation on dayside cloud coverage further in Section 5.3.

4.1.2. Rotation period

To analyze the effects of varying rotation alone, in Figure 9 we show maps of surface temperature and cloud coverage from simulations of planets orbiting a late M-dwarf star with varying rotation period from 0.5-16 Earth days. These simulations all use the same incident stellar flux of 1 \( F_{\odot} \). Note that these experiments are unrealistic, as the rotation period is not varied consistently with the incident stellar flux. In Section 4.1.1, we took into account the decrease in rotation period with increasing incident stellar flux, but in this section we vary the rotation period alone to isolate its effect on the atmospheric circulation. We can see that as the rotation period is decreased, the length scales of the circulation correspondingly decrease. With slow rotation, dynamical structures are large-scale, with a hot dayside hemisphere and cold nightside hemisphere. With faster rotation, the dayside temperature (and cloud) pattern is sheared out at high latitudes due to the effect of rotation, resulting in mid-latitude jets for Earth-like rotation periods (Noda et al. 2017).

As in the case with consistently varying rotation period and incident stellar flux, we see a transition in the dayside cloud coverage with increasing rotation period. For fast rotating planets, the dayside is relatively cloud-free, causing the planet to have a smaller top-of-atmosphere albedo and a hotter surface. As the rotation period is increased, the planet surface cools due to the enhanced dayside cloud coverage. Then, between a rotation period of 8 and 16 days, the dayside cloud coverage increases drastically, with clouds centered on the substellar point for the 16 day rotation period case. The centering of the cloud pattern on the substellar point increases the top-of-atmosphere albedo, leading to a significantly cooler surface. We will discuss the dynamics of this cloud transition in more detail in Section 5.3.

4.1.3. Surface pressure

We show the effects of varying surface pressure on the climate of M-dwarf planets in Figure 10 for simulations of planets orbiting a late M-dwarf planet with an incident stellar flux and rotation period set equal to that of Earth. We use this unrealistically short rotation period in order to contrast our simulations with those of planets orbiting Sun-like stars (see Section 5.5). However, even in this fast-rotating regime with mid-latitude jets identified by Noda et al. (2017) and Haqq-Misra et al. (2018),
Fig. 7.— Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting an early M-dwarf star with varying incident stellar flux from $0.544 - 1 F_\odot$, corresponding to varying rotation periods from 117.4 – 74.3 days. The center of each map corresponds to the substellar point of the planet. All simulations are in the slowly rotating regime, with a hot dayside and cold nightside and large dayside cloud coverage.

Fig. 8.— Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying incident stellar flux from $0.544 - 1 F_\odot$, corresponding to varying rotation periods from 6.49 – 4.11 days. The center of each map corresponds to the substellar point of the planet. A transition between reduced cloud coverage at the substellar point with faster rotation (larger incident fluxes) and enhanced cloud coverage with slower rotation (lower incident fluxes) can be seen.
Fig. 9.—Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying rotation periods from 0.5 - 16 days. The center of each map corresponds to the substellar point of the planet. Faster-rotating planets have smaller dynamical lengthscales, while slow rotators have large dynamical lengthscales and a large day-night temperature contrast. The dayside cloud cover increases with increasing rotation period, with an abrupt transition in the substellar cloud cover between a rotation period of 8 and 16 days.
Fig. 10.—Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying surface pressure from 0.25−4 bars. The center of each map corresponds to the substellar point of the planet. Increasing surface pressure leads to larger dayside cloud coverage, with a corresponding decrease in the global-average temperature for more massive atmospheres.

we can see the effect of atmospheric pressure on the M-dwarf cloud transition.

In our suite of simulations, planets with lower surface pressures have reduced dayside cloud coverage relative to those with higher surface pressures. We find that increasing the surface pressure (and hence atmospheric mass) leads to enhanced dayside convective fluxes. Additionally, we find that the surface relative humidity increases with increasing surface pressure, leading to more available condensate. Note that the enhanced convective fluxes and increase in surface relative humidity are linked, as convection moistens the atmosphere in the limit of weak subsidence (Romps 2014). The larger cloud coverage on the dayside of M-dwarf planets with higher surface pressures is a result of the increased dayside convective fluxes and relative humidity in our simulations with large surface pressures. Conversely, we find that the combination of reduced dayside cloud coverage along with the reduced Rayleigh scattering from the background atmospheric N$_2$ (Kopparapu et al. 2014) causes the surfaces of M-dwarf planets with thinner atmospheres to be hotter relative to those with thicker atmospheres.

5. DISCUSSION

5.1. Comparison with Kaspi & Showman (2015)

In this section, we compare our simulation results for how the Hadley circulation and equator-to-pole temperature contrast of terrestrial exoplanets varies with planetary parameters to the results of Kaspi & Showman (2015). Kaspi & Showman used the GCM of Frierson et al. (2006), Frierson et al. (2007) to investigate how varying planetary parameters affects the atmospheric dynamics of terrestrial exoplanets. This GCM has an idealized radiative transfer scheme with a double-grey opacity, that is, one visible and one infrared band, parameterized moist convection, and no clouds (Betts 1986, Betts & Miller 1986). Our simulations are therefore a more complex version of those performed by Kaspi & Showman.

In Figure 11 we show that the latitude of maximum cloud coverage along with the reduced Rayleigh scattering from the background atmospheric N$_2$ (Kopparapu et al. 2014) causes the surfaces of M-dwarf planets with thinner atmospheres to be hotter relative to those with thicker atmospheres.

Fig. 11.—Comparison of dependence of the latitude of the maximum in the latitude of the maximum of zonal wind speed between this work and that of Kaspi & Showman (2015). We generally find good agreement between the two models over a broad range of rotation rates.
The latitude of maximum zonal wind speed corresponds to the location of the subtropical jets that form at the flanks of the Hadley cell. Hence, the overall width of the Hadley cell and its dependence with rotation rate are similar between the two models.

In Figure 12 we show how the equator-to-pole temperature contrast depends on rotation rate, surface pressure, planetary radius, and incident stellar flux in our simulations and those of Kaspi & Showman. Because our simulations are not tuned to Earth, include sea ice, have zero obliquity, and do not include CO$_2$, the absolute equator-to-pole temperature contrasts in our simulations are up to $\approx 100$ K larger than those in Kaspi & Showman. As a result, we normalize the equator-to-pole temperature contrast to the average of the equator-to-pole temperature contrast for each model parameter variation in order to directly compare results from the two models. We generally find similar trends in the equator-to-pole temperature contrast with varying planetary parameters: the equator-to-pole temperature contrast increases with increasing rotation rate, decreasing surface pressure, increasing radius, and increasing flux. However, as discussed in Section 3.1.2 our simulations with surface pressures of 2 and 4 bars are fully ice-covered, leading to greatly reduced equator-to-pole temperature contrasts (due to the cooler equatorial region) relative to our simulations with lower pressures. We find very good agreement with Kaspi & Showman that planets with larger radii have larger equator-to-pole temperature contrasts due to a decreased ratio of the eddy length scale to the planetary radius. Similarly to Kaspi & Showman, we find a general trend of increasing equator-to-pole temperature contrast with increasing incident flux (due to the decreasing radiative time constant) until a critical flux at which the equator-to-pole temperature contrast begins to decrease with increasing flux (due to moist energy transport). By running a simulation with higher incident flux ($1.225 F_{\oplus}$, see the last panel of Figure 12), we find that the decrease in the equator-to-pole temperature contrast begins at higher values of incident stellar flux than was found by Kaspi & Showman due to the significant glaciation in our model with incident flux equal to that of Earth.

5.2. Sea ice coverage

All simulations of planets orbiting Sun-like stars shown in this work have significant sea ice cover due to the idealized assumptions of zero obliquity, no atmospheric CO$_2$, and no ocean heat transport. As a result, we find that the planet reaches a snowball in our simulations below an incident flux of $0.816 F_{\oplus}$. Note that this is lower than the snowball limit for modern Earth, which is $0.91 - 0.96 F_{\oplus}$ (Voigt & Marotzke 2010, Voigt et al. 2011). The reduced critical incident flux to reach a snowball in our simulations is due to the assumption of zero obliquity, which causes the equatorial regions to be ice-free at a lower incident flux. Very few of our simulations with an Earth value of incident stellar flux have sea ice that encompasses the entire planet (i.e., reach a snowball), as this only occurs with surface pressures greater than that of Earth.

Though the majority of simulations with Earth’s value of incident flux are not snowballs, only the near-equatorial region of the planet is ice-free. Figure 13 shows the zonal-mean ice and snow thickness for our
simulation with all parameters set equal to that of Earth. The planet is only ice-free in a small latitudinal strip near the equator, with up to $\sim 45$ m of sea ice cover at the poles. We find that the snow cover peaks poleward of the ice edge, and aligns with a region where the sea ice cover does not change much with latitude (at $\pm 30^\circ$). Snow has a much higher albedo than sea ice in CAM, which protects the ice sheet from melting in the mid-latitudes. The snow cover decreases to zero toward the equator, which reduces the top-of-atmosphere albedo and leads to melting. This climate state is the exoplanet equivalent to the “Jormungand” global climate state proposed for Earth (Abbot et al. 2011), and here we show that similar feedback processes as found for Earth-like simulations can occur in the zero-obliquity state. As a result, the reason that we find a lower incident flux limit for the snowball state in our simulations relative to modern Earth is because of the reduced snow-albedo feedback in equatorial regions in the low obliquity regime. We find that this mechanism is the same that may have kept the equatorial region of Earth ice-free during Neoproterozoic glaciations.

5.3. M-dwarf cloud transition

In our simulations of planets orbiting late M-dwarf stars, we find that there is a clear transition between low and high dayside cloud coverage between orbital periods of 8 and 16 days (see Figure 9), with further analysis showing that this transition occurs at $\approx 10$ days. For simulations of planets orbiting early M-dwarf stars, we find that the transition occurs between orbital periods of 4 and 8 days. To analyze this cloud transition in more detail, in Figure 14 we show the total cloud fraction, vertical wind speed, zonal wind speed, and meridional flux of zonal momentum ($\overline{u'v'}$) from our simulations of planets orbiting late M-dwarfs with orbital periods of 8 and 16 days.

Figure 14 shows that the faster rotating simulation with an orbital period of 8 days (left hand panels) has relatively low dayside cloud coverage, while the more slowly rotating simulation with an orbital period of 16 days (right hand panels) has larger dayside cloud coverage and an associated strong dayside updraft that occurs near the substellar point of the planet. The dayside updraft in the simulation with a rotation period of 16 days corresponds to the location where there is strong zonal convergence between eastward winds west of the substellar point and westward winds east of the substellar point. The strong zonal convergence is due to meridional transport of eddy momentum from higher latitudes toward the equator. This can be seen in the strong phase tilts in the meridional flux of eddy momentum that are northwest-southeast in the northern hemisphere and southwest-northeast in the southern hemisphere. The eddy phase tilts lead to negative eddy momentum flux north of the equator and positive eddy momentum flux south of the equator, combining to transport eastward momentum toward the equator (Held 2000, Vallis 2006, Showman & Polvani 2011). This eastward eddy momentum then accelerates the flow west of the substellar point, leading to enhanced convergence at the substellar point. This strong convergence implies transport of moisture and momentum to the substellar point, which leads to a strong updraft and dayside cloud cover.

The eddy phase tilts that transport momentum toward the equator in the 16 day period case are not seen in the 8 day period case, leading to much weaker dayside cloud coverage. Additionally, the region west of the substellar point (where there are strong updrafts) is very dry in the 8 day period model. In general, a positive correlation between upward vertical velocity and moisture is needed for net upward moisture transport (Zhang & Showman 2018). This correlation is not found in our 8 day period simulation, which is instead moist east of the substellar point where there are weak downdrafts. Note that our simulation with a period of 8 days lies in the rapidly rotating regime demarcated by Haqq-Misra et al. (2018) (where the Rhines and Rossby deformation scales are smaller than the planetary radius), while the 16 day period case is in a slowly rotating regime, where the relevant dynamical length scales are similar to or larger than the planetary radius. As a result, we find that the dynamical transition found by Haqq-Misra et al. (2018) between rapid and slow rotators leads to a transition in substellar cloud coverage due to eddy-mean flow interactions.

Previous authors have shown that the dayside cloud transition for M-dwarf planets is not affected significantly by changing planetary parameters (Yang et al. 2013) or by changing the model convective parameterizations (Way et al. 2018). Notably, Way et al. (2018) showed that dayside cloud coverage of slowly rotating planets is still large even when all upwelling condensate is assumed to rain out, showing that continuous convective transport of condensate is not required for large dayside cloud coverage. Our finding that the cloud transition is due to a shift in the dynamical state of the atmosphere with increasing rotation period is further evidence that the M-dwarf cloud transition is not parameterization dependent. Instead, the cloud transition of slowly rotating planets orbiting M-dwarfs is due to a robust dynamical mechanism that should not be dependent on the details of model cloud and convective parameterizations.

5.4. Observational consequences

To analyze the observational consequences of our simulations of planets orbiting M-dwarf stars, we show full-phase light curves (“phase curves”) in Figure 15. These phase curves are bolometric, using the upwelling top-of-atmosphere longwave flux (i.e., the OLR) from our simulations of planets orbiting M-dwarf stars with varying effective temperature and insolation, using the correct rotation period for a given value of insolation (Kopparapu et al. 2017). The flux is then weighted by the cosine of the angle between the grid point and the observer to take into account planetary limb darkening (Cowan & Agol 2008).

We find that slower-rotating planets that receive less incident flux have phase curves that peak near secondary eclipse and have large dayside-to-nightside flux contrasts, evident in the large amplitude of the phase curve. Hotter planets that are faster rotating have larger offsets of the maximum of the phase curve from secondary eclipse, with the phase curve maximum even occurring near the eastern limb for planets orbiting earlier-type M-dwarfs that receive the same incident flux as Earth. The phase curve offset is likely determined by a combination of wave dynamics and cloud coverage, as wave-mean flow inter-
Fig. 14.— A shift in the dynamical state of the atmosphere causes the transition in dayside cloud coverage between rotation periods of 8 (left) and 16 (right) days for planets orbiting M-dwarfs. The top panel shows the total cloud fraction, the second panel shows the vertical wind speed at the lowest vertical level (in pressure units of milli-Pascals per second), the third panel shows the zonal wind speed at the lowest vertical level, and the bottom panel shows the meridional flux of zonal momentum $\nabla'$ at the lowest vertical level. Over-plotted contours range from the minimum to maximum value shown on the colorbar, and panels on the same row share a color scale. In the simulation with a rotation period of 16 days, there are strong phase tilts of the meridional flux of zonal eddy momentum west of the substellar point. These phase tilts are northwest-southeast in the northern hemisphere and southwest-northeast in the southern hemisphere and therefore flux zonal momentum from higher latitudes toward the equator. This equatorward momentum flux causes strong eastward winds west of the substellar point, which leads to zonal convergence at the substellar point. The zonal convergence forces a strong updraft centered on the substellar point, leading to high cloud coverage on the dayside.
actions provide an explanation of the phase curve offsets of larger tidally locked planets (Zhang & Showman 2017, Komacek et al. 2017, Hammond & Pierrehumbert 2018), and clouds can affect the outgoing infrared flux. However, we leave a detailed understanding of the phase curve offset of tidally locked terrestrial planets to future work. Additionally, we find that faster rotating planets have smaller relative day-to-night flux contrasts. As a result, we agree with the findings of Haqq-Misra et al. (2018) that phase curves can be used to infer the dynamical state of the atmosphere for terrestrial planets orbiting M-dwarf stars. In future work, we will investigate how transmission, emission and reflectance spectra vary over the broad suite of our simulations, with implications for future observations with the James Webb Space Telescope and LUVOIR/HabEx.

5.5. Comparison between the climates of planets orbiting Sun-like stars and M-dwarfs

Our suite of simulations is unique in that we utilized the same GCM setup to study planets orbiting both Sun-like stars and M-dwarf stars. We used the same model parameters in the two suites, but assumed a different spin-state (spin-synchronized for M-dwarf planets, zero obliquity and Earth’s rotation period for Sun-like star planets) and used a different incident stellar spectrum.

We find that the spin state greatly changes the dynamics and resulting climate. Increasing the rotation period causes a sharp decrease in the globally averaged surface temperature for simulations of planets orbiting M-dwarf stars (~45 K with increasing rotation period from 0.5-16 days for planets orbiting late M-dwarfs), whereas the surface temperature for planets orbiting Sun-like stars varies by ≲10 K with varying rotation period. This is due to the dayside cloud transition that enhances the top-of-atmosphere albedo of slowly rotating M-dwarf planets, but does not occur in our suite of simulations for planets orbiting Sun-like stars. Note that in our suite of simulations we did not exceed a rotation period of 16 days for simulations of Earth-like planets orbiting Sun-like stars. However, planets orbiting Sun-like stars will reach a similar cloud transition at very long rotation periods relevant to early Venus (Yang et al. 2014, Way et al. 2016a).

We also find that M-dwarf planets are uniformly warmer than planets orbiting Sun-like stars. This corresponds with a lower top-of-atmosphere albedo for planets orbiting M-dwarf stars relative to planets orbiting Sun-like stars. This lowered albedo for planets orbiting M-dwarf stars is due to both the greater absorption of incident near-infrared radiation by atmospheric water vapor and significantly reduced dayside sea ice coverage in these simulations. This reduced dayside sea ice coverage for planets orbiting M-dwarf stars is due to two effects: first, the decrease in the ice-albedo feedback due to the relatively red incident stellar radiation at which sea ice absorbs more incoming radiation (Joshi & Haberle 2012), and second, the pattern of incident flux ensures that the daysides of M-dwarf planets that receive similar flux to Earth will have an ice-free region centered on the substellar point (Pierrehumbert 2011). Note that we do not include a dynamical ocean in these simulations, which would act to increase the amount of ice-free ocean due to heat transport (Hu & Yang 2014). Hence, the conclusion that M-dwarf planets are uniformly hotter than
planets orbiting Sun-like stars is likely to be robust to changing our assumption of zero ocean heat transport.

6. CONCLUSIONS

We find that the climates of exoplanets are sensitive to planetary parameters, with the incident stellar flux, rotation period, and surface pressure all able to greatly affect the surface temperature and determine whether or not a planet is habitable. Planetary parameters impact climate differently for planets around Sun-like stars and synchronously rotating planets orbiting M-dwarf stars. Most notably, increasing the rotation period leads to significant cooling for planets orbiting M-dwarf stars but does not greatly affect the surface temperature of planets orbiting Sun-like stars. Here we list important results from this work for planets orbiting Sun-like and M-dwarf stars.

The key takeaways from our simulations of planets orbiting Sun-like stars are:

1. The scaling of Hadley cell width and equator-to-pole temperature differences from our sophisticated general circulation modeling broadly match those of Kaspi & Showman (2015) using an idealized GCM with double-grey radiative transfer and no clouds. The width of the Hadley circulation of Earth-like planets increases with increasing rotation period, leading to an increase in the latitude at which the magnitude of the subtropical jet peaks with increasing rotation period. However, we find large differences in the absolute value of equator-to-pole temperature differences between our simulations and those of Kaspi & Showman (2015) because we include sea ice and do not tune our simulations to match Earth.

2. The globally averaged temperature of terrestrial planets increases with increasing incident stellar flux, planetary radius, and effective liquid cloud particle size and decreases with increasing surface pressure and surface gravity. The equator-to-pole temperature contrast correlates with the globally averaged temperature except for varying rotation period and surface gravity, in which case the equator-to-pole temperature contrast increases with decreasing rotation period and increasing surface gravity. Our findings broadly agree with previous theoretical and numerical work which finds that the equator-to-pole temperature contrast should increase with increasing planetary radius and gravity and decrease with increasing surface pressure and rotation period.

3. We find that the effective cloud particle size is a key unknown in the study of the climate of terrestrial exoplanets. With the assumption of uniform effective cloud particle sizes, larger particle sizes lead to warmer climates, with a correspondingly more vigorous atmospheric circulation. Future work is necessary to understand the expected cloud particle size distribution in the atmospheres of terrestrial exoplanets and how it affects climate.

The key points from our simulations of planets orbiting M-dwarf stars are:

1. We find that a transition between small and large dayside cloud coverage of synchronously rotating planets occurs with increasing rotation period. For planets that receive the same incident flux as Earth and orbit a late M-dwarf star with an effective temperature of 2600 K, the transition between low and high dayside cloud coverage occurs at a rotation period of ≈ 10 days. We find a similar transition between low and high dayside cloud coverage for planets orbiting late-type M-dwarfs with consistently varying rotation period and incident stellar flux, but in this case the transition occurs at a smaller rotation period of ≈ 5 days. For planets orbiting early type M-dwarf stars with an effective temperature of 4000 K, we find extensive dayside cloud coverage for all combinations of rotation period and incident stellar flux that we consider.

2. We find that the transition between small and large dayside cloud coverage is due to a shift in the dynamical state of the atmosphere with increasing rotation period. As the rotation period is increased, the meridional eddy flux of zonal momentum shows northwest-southeast phase tilts in the northern hemisphere (southwest-northeast phase tilts in the southern hemisphere) that flux zonal momentum from higher latitudes toward the equator. This equatorward momentum flux causes strong eastward flow west of the substellar point that leads to convergence of the zonal winds near the substellar point. This convergence causes a strong updraft on the dayside of the planet, leading to large dayside cloud coverage. As a result, we find that the transition in dayside cloud coverage is due to a largely dynamical mechanism and should not be parameterization-dependent. This agrees with the previous results of Yang et al. (2013) and Way et al. (2018), who have shown that dayside cloud coverage is not significantly affected when varying model parameters and convective schemes.

3. We computed full-phase light curves from our numerical circulation models of terrestrial exoplanets orbiting M-dwarf stars with varying effective temperatures from 2600 − 4000 K and with varying incident stellar flux from 0.544 − 1 $F_\oplus$. We find that the phase curve amplitude increases with decreasing incident stellar flux due to the larger day-to-night temperature differences of cooler planets. We additionally find that faster-rotating planets that receive a greater amount of incident stellar flux have larger phase curve offsets. Hence, as proposed by Haqq-Misra et al. (2018), we find that phase curves can be used to infer the dynamical state of terrestrial exoplanet atmospheres.

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APPENDIX

A. PLANETS ORBITING M-DWARF STARS: FURTHER RESULTS

A.1. Incident stellar flux

Figure 16 shows how the surface temperature and cloud coverage for synchronously rotating planets around M-dwarf stars vary with increasing incident stellar flux from $0.544 - 1 F_\text{E}$. As expected, the surface temperatures increase with increasing incident flux, with a relatively larger increase in the nightside temperature than dayside temperature. The cloud pattern also changes drastically, from broad equatorial cloud coverage with low flux to large dayside cloud coverage with moderate flux and then reduced dayside and enhanced nightside cloud coverage at an Earth-like value of flux. Comparing these results to the simulations with consistent incident stellar flux and rotation period (Figures 7 and 8), one can see visually the importance of slow rotation for the habitability of planets orbiting M-dwarf stars. In the fast-rotating simulations shown in Figure 16, the dayside cloud albedo mechanism is not present to reduce the amount of incident flux reaching the surface. This leads these simulations with the rotation period fixed at 1 Earth day to be much hotter than those with consistent (slower) rotation.
Fig. 17.— Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying planetary radius from $0.5 - 2 \, R_\oplus$. The center of each map corresponds to the substellar point of the planet.

A.2. Planetary radius

Figure 17 shows how the surface temperature and cloud coverage for planets orbiting M-dwarf stars varies with increasing radius from $0.5 - 2 \, R_\oplus$. We find that the surface temperature increases with increasing planetary radius, largely due to decreasing dayside cloud cover (and increasing nightside cloud cover). In Section 5.3, we showed that the large dayside cloud coverage of M-dwarf planets is due to a wave-mean flow interaction that leads to zonal convergence near the substellar point. Given that the relative ratio of the Rossby deformation radius (which controls the length scale over which gravity waves can interact with the flow) and the planetary radius decreases with increasing planetary radius, one might expect that the wave-mean flow interactions driving this convergence decrease in efficacy with increasing radius.

A.3. Surface gravity

Figure 18 shows the effect of varying surface gravity from $0.707 - 1.414 \, g_\oplus$ on the climate of planets orbiting M-dwarf stars. We find that simulations of planets with smaller surface gravity are hotter, the same result found for planets orbiting Sun-like stars. This is because reduced surface gravity increases the atmospheric mass, leading to a larger total atmospheric heat capacity. Additionally, we find that the nightside cloud coverage is much larger in simulations with lower gravity due to their hotter nightsides causing more water vapor to be present in the nightside atmosphere.

A.4. Liquid cloud particle size

Figure 19 shows the effect of varying the effective liquid cloud radius from $7 - 21 \, \mu m$ on the climate of planets orbiting M-dwarf stars. As for planets orbiting Sun-like stars, larger effective cloud particle sizes lead to hotter surfaces due to the decreased scattering of the incoming stellar radiation (see Figure 3). The cloud particle size also significantly affects the cloud distribution, with reduced dayside cloud coverage and enhanced nightside cloud coverage for simulations with larger cloud particle size.

B. TABLES OF SIMULATION OUTPUT

Here we show tables of output from all of our simulations. In Table 2 we show the globally averaged surface temperature, equator-to-pole temperature contrast, top-of-atmosphere net shortwave flux, top-of-atmosphere net longwave...
Fig. 18.— Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying surface gravity from 0.707 to 1.414 g. The center of each map corresponds to the substellar point of the planet.

Fig. 19.— Maps of surface temperature, integrated cloud coverage, and integrated cloud water path from simulations of spin-synchronized planets orbiting a late M-dwarf star with varying liquid cloud particle radius from 7 to 21 µm. The center of each map corresponds to the substellar point of the planet.

flux, and albedo from our simulations of planets orbiting Sun-like stars. In Tables 3, 4, and 5 we show the globally averaged temperature, day-night temperature difference, top-of-atmosphere net shortwave flux, top-of-atmosphere net longwave flux, and albedo for our simulations of tidally locked planets orbiting M-dwarf stars. We also show the equivalent orbital distance from the host star, in AU, for all of our assumed incident stellar fluxes for a given stellar type. Table 3 shows the results from our simulations of planets orbiting an M-dwarf star with $T_{\text{eff}} = 2600$ K, Table 4 shows data for planets orbiting an M-dwarf with $T_{\text{eff}} = 3300$ K, and Table 5 shows results for planets orbiting an M-dwarf with $T_{\text{eff}} = 4000$ K.
### Table 2

Key simulation quantities for the simulations of planets orbiting sun-like stars, averaged over the last ten years of model time.

| Simulation parameters | Global-average T [K] | Eq-pole ∆T [K] | TOA net SW [Wm⁻²] | TOA net LW [Wm⁻²] | Albedo |
|-----------------------|----------------------|----------------|-------------------|-------------------|--------|
| Reference:            | Earth parameters (1 AU) | 231.79 | 139.37 | 151.72 | 154.96 | 0.55 |
| **Varying rotation period:** | | | | | | |
| 0.5 days | 220.85 | 128.37 | 123.29 | 129.30 | 0.64 |
| 2 days | 218.86 | 109.45 | 116.61 | 121.50 | 0.66 |
| 4 days | 217.63 | 97.76 | 114.74 | 119.82 | 0.66 |
| 8 days | 220.89 | 93.93 | 118.13 | 123.00 | 0.65 |
| 16 days | 222.41 | 75.04 | 118.96 | 123.96 | 0.65 |
| **Varying surface pressure:** | | | | | | |
| 0.25 bars | 234.89 | 140.20 | 164.29 | 167.07 | 0.52 |
| 0.5 bars | 235.19 | 147.19 | 162.10 | 164.44 | 0.52 |
| 2 bars | 211.24 | 96.06 | 106.43 | 111.27 | 0.69 |
| 4 bars | 211.49 | 85.36 | 102.07 | 108.28 | 0.70 |
| **Varying incident stellar flux:** | | | | | | |
| 0.544 F⊙ (1.36 AU) | 182.42 | 94.29 | 56.09 | 63.04 | 0.69 |
| 0.667 F⊙ (1.22 AU) | 191.93 | 104.31 | 70.81 | 75.68 | 0.69 |
| 0.816 F⊙ (1.11 AU) | 200.81 | 108.22 | 86.24 | 91.02 | 0.69 |
| 1.25 F⊙ (0.984 AU) | 268.85 | 110.37 | 235.97 | 277.70 | 0.46 |
| **Varying planetary radius:** | | | | | | |
| 0.5 R⊕ | 239.06 | 112.42 | 168.21 | 171.52 | 0.51 |
| 0.707 R⊕ | 216.89 | 104.15 | 113.50 | 118.27 | 0.67 |
| 1.414 R⊕ | 236.09 | 147.65 | 158.47 | 161.42 | 0.53 |
| 2 R⊕ | 244.65 | 161.95 | 171.02 | 173.62 | 0.50 |
| **Varying surface gravity:** | | | | | | |
| 0.5 g⊕ | 236.72 | 118.03 | 148.98 | 152.10 | 0.56 |
| 0.707 g⊕ | 232.49 | 130.18 | 147.36 | 151.02 | 0.57 |
| 1.414 g⊕ | 231.23 | 142.14 | 153.37 | 156.94 | 0.55 |
| **Varying liquid cloud particle radius:** | | | | | | |
| 7 μm | 213.05 | 110.93 | 107.31 | 113.41 | 0.69 |
| 21 μm | 239.22 | 138.15 | 167.46 | 170.33 | 0.51 |

### Table 3

Key simulation quantities for the simulations of planets orbiting an M-dwarf star with an effective temperature of 2600 K, averaged over the last ten years of model time. Simulations marked “unstable” do not reach an equilibrated state, but are shown for comparison.

| Simulation parameters | Global-average T [K] | Day-night ∆T [K] | TOA net SW [Wm⁻²] | TOA net LW [Wm⁻²] | Albedo |
|-----------------------|----------------------|----------------|-------------------|-------------------|--------|
| Reference:            | Earth parameters (0.022 AU) | 281.25 | 18.25 | 263.59 | 263.63 | 0.23 |
| **Varying rotation period:** | | | | | | |
| 0.5 days | 308.26 | 9.76 | 307.19 | 298.47 | 0.10 |
| 2 days | 303.94 | 12.08 | 313.28 | 312.64 | 0.08 |
| 4 days | 298.12 | 13.36 | 301.09 | 300.44 | 0.12 |
| 8 days | 286.00 | 23.20 | 280.35 | 280.48 | 0.18 |
| 16 days | 262.24 | 24.18 | 229.78 | 230.78 | 0.33 |
| **Varying surface pressure:** | | | | | | |
| 0.25 bars (unstable) | 322.18 | 5.44 | 315.16 | 301.28 | 0.07 |
| 0.5 bars | 284.56 | 15.51 | 267.50 | 267.38 | 0.21 |
| 2 bars | 283.38 | 19.62 | 266.40 | 265.91 | 0.22 |
| 4 bars | 279.89 | 23.61 | 269.93 | 270.39 | 0.21 |
| **Varying incident stellar flux:** | | | | | | |
| 0.544 F⊙ (0.030 AU) | 209.55 | 48.05 | 120.31 | 124.31 | 0.35 |
| 0.667 F⊙ (0.027 AU) | 229.98 | 39.36 | 159.01 | 162.35 | 0.30 |
| 0.816 F⊙ (0.025 AU) | 250.26 | 35.20 | 203.67 | 206.10 | 0.27 |
| **Varying planetary radius:** | | | | | | |
| 0.5 R⊕ | 273.16 | 20.29 | 259.41 | 260.50 | 0.24 |
| 0.707 R⊕ | 275.99 | 19.45 | 259.99 | 260.79 | 0.24 |
| 1.414 R⊕ | 269.53 | 15.56 | 267.43 | 266.94 | 0.22 |
| 2 R⊕ | 317.08 | 10.42 | 304.15 | 302.90 | 0.11 |
| **Varying surface gravity:** | | | | | | |
| 0.5 g⊕ (unstable) | 362.44 | 0.89 | 313.33 | 294.64 | 0.08 |
| 0.707 g⊕ | 350.39 | 13.94 | 269.07 | 270.09 | 0.24 |
| 1.414 g⊕ | 274.11 | 23.42 | 256.37 | 257.28 | 0.25 |
| **Varying liquid cloud particle radius:** | | | | | | |
| 7 μm | 266.25 | 23.64 | 234.86 | 236.79 | 0.31 |
| 21 μm | 284.42 | 18.96 | 274.03 | 273.89 | 0.20 |
| **Varying incident stellar flux and rotation period:** | | | | | | |
| 0.544 F⊙, 6.49 days | 214.58 | 48.68 | 130.59 | 133.75 | 0.29 |
| 0.667 F⊙, 5.57 days | 227.32 | 42.96 | 157.22 | 160.14 | 0.31 |
| 0.816 F⊙, 4.79 days | 246.97 | 32.58 | 205.44 | 208.24 | 0.26 |
| 4 F⊙, 4.11 days | 300.16 | 14.12 | 304.57 | 304.55 | 0.11 |
TABLE 4
Key simulation quantities for the simulations of planets orbiting an M-dwarf star with an effective temperature of 3300 K with consistently varying incident stellar flux and rotation period, averaged over the last ten years of model time.

| Simulation parameters | Global-average T [K] | Day-night ∆T [K] | TOA net SW [Wm$^{-2}$] | TOA net LW [Wm$^{-2}$] | Albedo |
|-----------------------|----------------------|------------------|-------------------------|-------------------------|--------|
| 0.544 $F_0$, 33.8 days (0.134 AU) | 214.09 | 57.67 | 129.17 | 134.21 | 0.30 |
| 0.667 $F_0$, 30.7 days (0.121 AU) | 224.19 | 55.26 | 147.74 | 151.70 | 0.35 |
| 0.816 $F_0$, 26.4 days (0.109 AU) | 234.61 | 50.66 | 172.81 | 175.11 | 0.38 |
| 1 $F_0$, 22.7 days (0.099 AU) | 251.84 | 34.89 | 204.65 | 206.39 | 0.40 |

TABLE 5
Key simulation quantities for the simulations of planets orbiting an M-dwarf star with an effective temperature of 4000 K, averaged over the last ten years of model time.

| Simulation parameters | Global-average T [K] | Day-night ∆T [K] | TOA net SW [Wm$^{-2}$] | TOA net LW [Wm$^{-2}$] | Albedo |
|-----------------------|----------------------|------------------|-------------------------|-------------------------|--------|
| Reference: Earth parameters (0.296 AU) | 260.26 | 31.20 | 220.25 | 222.57 | 0.35 |
| Varying rotation period: | | | | | |
| 0.5 days | 267.81 | 29.28 | 230.82 | 232.68 | 0.33 |
| 2 days | 267.91 | 26.17 | 237.14 | 239.78 | 0.30 |
| 4 days | 260.65 | 28.62 | 229.19 | 231.79 | 0.33 |
| 8 days | 244.33 | 53.45 | 199.18 | 201.34 | 0.42 |
| 16 days | 242.89 | 48.47 | 191.96 | 193.94 | 0.44 |
| Varying surface pressure: | | | | | |
| 0.25 bars | 261.13 | 30.26 | 223.32 | 227.60 | 0.34 |
| 0.5 bars | 259.34 | 32.19 | 222.96 | 226.06 | 0.35 |
| 2 bars | 262.43 | 28.69 | 226.11 | 228.18 | 0.34 |
| 4 bars | 271.05 | 24.04 | 243.16 | 244.69 | 0.29 |
| Varying incident stellar flux: | | | | | |
| 0.544 $F_0$ (0.402 AU) | 201.58 | 51.87 | 103.05 | 108.31 | 0.44 |
| 0.667 $F_0$ (0.363 AU) | 220.22 | 43.49 | 136.31 | 140.70 | 0.40 |
| 0.816 $F_0$ (0.328 AU) | 246.17 | 36.81 | 187.88 | 191.10 | 0.32 |
| Varying planetary radius: | | | | | |
| 0.5 $R_0$ | 260.24 | 27.90 | 226.89 | 229.37 | 0.33 |
| 0.707 $R_0$ | 263.40 | 27.60 | 230.58 | 233.05 | 0.32 |
| 1.414 $R_0$ | 262.93 | 32.06 | 221.73 | 223.80 | 0.35 |
| 2 $R_0$ | 268.11 | 32.06 | 221.73 | 223.80 | 0.35 |
| Varying surface gravity: | | | | | |
| 0.5 $g_0$ | 277.23 | 18.73 | 240.67 | 241.56 | 0.29 |
| 0.707 $g_0$ | 260.13 | 23.10 | 230.16 | 232.00 | 0.32 |
| 1.414 $g_0$ | 255.80 | 36.71 | 215.68 | 218.13 | 0.37 |
| Varying incident stellar flux and rotation period: | | | | | |
| 0.544 $F_0$, 117.4 days | 210.58 | 65.36 | 122.20 | 126.78 | 0.34 |
| 0.667 $F_0$, 109.7 days | 220.61 | 61.92 | 139.92 | 143.30 | 0.38 |
| 0.816 $F_0$, 86.6 days | 226.19 | 54.78 | 160.48 | 163.16 | 0.42 |
| 1 $F_0$, 74.3 days | 240.57 | 47.26 | 185.13 | 186.79 | 0.46 |