Assessing the Potential of Using Satellite Data on Radiation in the Analysis of Earth’s Climate System to Complement the Scarce Radiation Data Measured from the Ground Stations

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Authors’ contributions

This work was carried out in collaboration between all authors. Author COO designed the study, performed the statistical analysis, wrote the protocol, and wrote the first draft of the manuscript. Authors SWW and RS managed the analyses of the study. Author SWW managed the literature searches. All authors read and approved the final manuscript.

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Abstract

High quality solar radiation data is required for the appropriate monitoring and analysis of the Earth’s climate system as well as efficient planning and operation of solar energy systems. However, well maintained radiation measurements are rare in many regions of the world. Therefore, satellite-derived radiation estimates are an alternative to these scarce solar radiation measurements from the weather stations. Satellite estimates of solar radiation have an advantage over solar radiation measurements from weather stations because of their high spatial and temporal resolutions. These satellite radiation estimates at approximately 5-6 Km resolution derived from geostationary Meteosat satellites are available through the EUMETSAT Satellite Application Facilities (SAFs). CM-SAF (SAF on Climate Monitoring).

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provides consistent dataset of hourly, daily and monthly solar radiation from 1983 to 2013. In this study, we examined the potential of using satellite estimates of solar radiation to fill in the data gaps in records from the weather stations as well as the areas where radiation data is not available. The analysis carried out showed that the satellite data had fewer missing values than the ground data, and that they are both similar in distribution. The average correlation between the two data sets was found to be 0.71 for both monthly and daily analysis. However, the month of September showed a very low correlation of 0.21. Mean percentage error, mean bias error and mean absolute deviation were found to be 2.46, 18.84, 50.32 and 3.08, 559.87, 1135.93 for daily and monthly analysis, respectively.

The solar radiation distribution in Dodoma was found to follow Weibull distribution throughout the year.

Keywords: Weibull Distribution; CM-SAF; EUMETSAT; geostationary Meteosat satellites.

1 Introduction

Most of the biochemical processes that occur on the surface of the earth, such as photosynthesis and plant development, derive their energy from the sun [1]. Solar radiation is also important in determining the performance and monitoring the working of solar devices such as solar furnaces, solar collectors, and photovoltaics. It is therefore important to get solar radiation data of high quality at every place where these applications are applied for appropriate monitoring.

Solar radiation data are important for many application fields such as ecology, biodiversity, hydrology, agriculture, meteorology, and climatology. For example,

1. Hydrology: Most of the hydrological models require weather variables such as precipitation, maximum and minimum temperatures and solar radiation [2].

2. Agriculture: Farmers are faced with a challenge of choosing which crop to plant in a given area. This is normally influenced by a number of factors such as soil moisture of the area and plant’s rate of evapotranspiration. Soil moisture depends on the rate of evaporation in the soil, which to a large extent, depends on solar radiation. To understand this basic physiology of crop growth and development, many crop models have been developed [3]. These models are of great importance in quantifying the environmental limits to specific crop production at a given area. They are also applicable in prediction of crop yields. According to [4], solar radiation is an important factor in crop growth in Europe and other areas of the world (Fig. 1:1). This justifies the importance of accurate measurement of radiation data.

![Fig. 1.1. Geographic distribution of potential climatic constraints, temperature, radiation and water, to plant growth derived from long-term climate statistics [4]](image-url)
In the recent past, there has been a lot of work on examining the feasibility of solar cooking and solar water purification. Development of these technologies requires accurate information on hourly solar radiation as this will help in knowing when to use these appliances. The knowledge of the hourly solar radiation is also vital in determining how much of the energy can be stored for use when the sun goes down. According to [5], “it is anticipated that solar cooking technology will be demanded by a huge group of people in the near future because of its outstanding features”. Therefore, accurate measurement of solar radiation in most parts of the world will be of great importance to help in designing the solar cooker and solar water purifiers that can be can be used even in areas that receive low solar irradiance.

However, location specific and accurate solar radiation data is still a great challenge for most of the African countries, even though the problem has been discussed at many local and global meetings including the World Meteorological Organization (WMO) Regional Association (Africa) sessions that are held every four years [6]. This is mainly because there are few weather stations in Africa. On 6th Nov, 2006, Michel Jarraud said in a press conference that was held in Nairobi, Kenya, that Africa still needs approximately 200 automatic weather stations as a step towards rescuing historical data and also improving capacity building on climate and weather reporting. According to Medany et al. [7], there are 1152 World Weather Watch (WWW) stations in Africa, giving a station density of 1 station per 26,000Km² which is still 8 times below the minimum recommended level by WMO. In the working paper 3 of United Nations Economic Commission for African Climate Policy Centre [8], it is documented that the shortage of data in Africa is exacerbated by the uneven distribution of available weather stations leaving large areas unmonitored.

As a way to reduce the problem of insufficient climatic data in Africa, United Nations Development programme rolled out a number of support programs that encourages climate resilient economic development in Africa (CIRDA) [9]. The main component of these programs is to provide credible and appropriate information on both weather and climate to help in decision making in most of the sectors that require these data [9]. To achieve this, Trans-African Hydro-meteorological Observatory (TAHMO) has been commissioned with the task of installing 20,000 on-the-ground automatic stations across Africa. These stations are to provide rainfall, temperature, and other critical data at high temporal resolution, which TAHMO plans to make freely available in their website. TAHMO is currently piloting this in several parts of Chad, Kenya, DRC, Uganda, among others [10].

Trans-African Hydro meteorological Observatory (TAHMO [10]) has the task of installing 20,000 stations in Africa to help solve the problem of insufficient climatic data in the continent.

In the absence of solar radiation measurements, a number of methods have been developed to derive estimates of solar radiation for places where it is not measured directly. The simplest of these methods is assigning the measured values from the nearest station to these places [11] or using spatial interpolation.
methods [12]. Solar radiation, at times is also estimated using other meteorological observations that are available, such as sunshine duration, cloud cover and air temperature using empirical models [13]. The estimated radiation calculated from these models varies in accuracy depending on variables used [14]. The model that uses the sunshine duration has been found to be more accurate, since it gives a measure of the time per day of direct solar radiation. But since the number of weather stations is low, solar radiation data recorded or estimated may not be enough to create a time series of solar radiation that has a high spatial resolution [15]. Thus, estimated solar radiation data from sources that provide spatially continuous information at high resolution such as data provided by the satellite have to be considered to supplement the station-specific estimates.

Satellite data on radiation provides continuous estimates of these radiation data with a very high spatial and temporal resolution of $0.05^\circ \times 0.05^\circ$ and 30 min, respectively, for over 30 years. These estimates from the satellite data are therefore worth investigating to see if they correlate well with ground truth, and hence can be used to fill the large data gaps.

2 Solar Radiation Data

Data as well as preparation of the data used in this work is presented here.

2.1 Description of the data

2.1.1 Satellite data on radiation

The data set that was used is called the Surface Solar Radiation Data Set-Heliosat (SARAH). It is a collection of satellite records of the solar surface irradiance (SIS), the surface direct normalized irradiance (DNI) and the effective cloud albedo (CAL) [16]. These records are available as hourly, monthly and yearly averages from 1983 – 01 to 2013 – 12, which have started to change as more recent data are being added to the data base, with a spatial resolution of $0.05^\circ \times 0.05^\circ$ and covering a region of $\pm 65^\circ$ longitude and $\pm 65^\circ$ latitude [16].

The SARAH data are derived from the Meteosat Visible and Infrared Imager (MVIRI) and Spanning Enhanced Visible and Infrared Imager (SEVIRI) instruments on the geostationary Meteosat satellites [17]. This is called the Heliosat Method.

The first step in this method is to retrieve CAL by using the normalized relation between all sky and clear sky reflection in the visible channel of the Meteosat instruments. CAL is then used to derive the cloud index (a measure for the impact of clouds on the clear-sky irradiance. The clear - sky irradiance is calculated using the all sky model SPECMAGIC [18]). The combination of cloud index and clear sky irradiance gives SIS. The Surface Direct Irradiance ($SID$) is then derived using the diffuse model of Skartveit et al., [19] and the cloud index. The direct normalized irradiance ($DNI$) is derived from normalization of $SID$ with the cosine of the solar zenith angle (SZA):

According to Mueller et al. [20],

$$DNI = \frac{SID}{\cos(SZA)},$$

(2.1.1)

**Effective Cloud Albedo (CAL)** is the normalized difference between the reflected irradiance for all sky and reflected irradiance for clear sky. It is dimensionless. CAL is a measure of the effect of cloud on the Earth’s solar radiation budget, since it measures the cooling effect in the solar spectrum caused by clouds.

According to Mueller et al. [20],

$$CAL = \frac{\rho - \rho_{cs}}{\rho_{max} - \rho_{cs}},$$

(2.1.2)
where, $\rho$ is the observed reflection for each pixel and time, $\rho_{cs}$ is the clear sky reflection and $\rho_{\text{max}}$ is the maximum reflection determined by the 95% percentile of all reflection values at local noon in a target region. From Posselt et al. [21], the monthly, daily and hourly averages of $\text{CAL}$ are calculated as

$$\text{CAL}_{\text{mean}} = \frac{1}{n} \left( \sum_{i=1}^{n} \text{CAL}_i \right)$$  \hspace{1cm} (2.1.3)

where $i$ is the index of the slots per hour for hourly average, index of hourly means for daily means, index for daily means for monthly mean calculations.

Surface Incoming Solar Radiation (SIS). This is the solar radiation flux on a horizontal Earth surface, expressed in $Wm^{-2}$. According to Mueller et al. [22], it is calculated from CAL as;

$$SIS = SIS_{\text{CLS}} (1 - \text{CAL}),$$  \hspace{1cm} (2.1.4)

where $SIS$ is the solar surface irradiance, and $SIS_{\text{CLS}}$ is clear sky irradiance. The SIS daily mean is

$$SIS_{\text{DA}} = \frac{SIS_{\text{CLS}}}{n} \left( \frac{\sum_{i=1}^{n} SIS_i}{\sum_{i=1}^{n} SIS_{\text{CLS}}} \right).$$  \hspace{1cm} (2.1.5)

where $SIS_{\text{DA}}$ is daily average of $SIS$, $SIS_{\text{CLS}}$ is the daily averaged $SIS$ is for clear sky, $SIS_i$ calculated $SIS$ for satellite image $i$, $SIS_{\text{CLS}}$ is the calculated clear sky, $SIS$ for image $i$ and $n$ is the total number of images available during the day.

**Figure 2.1:** SIS plot for satellite data  
**Figure 2.2:** Average DNI over Tanzania in 2013

Direct Normalized Irradiance. This is the radiation from the Sun directly to the Earth. It has a fixed direction. It is expressed in $Wm^{-2}$. In sunny regions or during the summer, it accounts for 70% – 80% of the total radiation.
Fig. 2.1 and Fig. 2.2 shows examples of the average SIS and DNI derived from the satellite over Tanzania for March 2013.

The satellite estimates of radiation used in this work was gotten from EUMETSAT Satellite Application Facility on Climate Monitoring (CM SAF) [23]. This website has hourly, daily and monthly averages of DNI, SIS and CAL from 1983/01/01 to 2013/12/31. For the sake of this work, hourly SIS was selected. The data ranged from 1983/01/01 to 2013/12/31 and covering an area bounded by longitude 35.7° E to 35.8° E and latitude 6.2° S to 6.1°S. This was selected because it is the area that bounds the area from where ground data were derived. The data is provided as a zipped folder.

2.1.2 Dodoma data on radiation.

Solar radiation data at the ground stations can be measured directly using instruments such as pyranometers and pyrheliometers or indirectly from the number of sunshine hours or maximum and minimum temperature.

Pyrheliometers are used to measure direct solar radiation. The working of pyrheliometer is such that sunlight enters it through a window and is directed onto a thermopile which converts heat to an electrical signal that can be recorded. A calibration factor is then applied to convert the millivolt signal to an equivalent radiant energy flux, measured in W m⁻². It is sensitive to wavelengths in the band from 0.28 μm to 3 μm. A pyranometer is used to measure global solar radiation. It gives the reading in W m⁻² and is also sensitive to wavelengths in the band from 0.28 μm to 3 μm. Diffuse radiation is measured using a pyranometer with a shading ring.

In Tanzania, the Tanzania Meteorological Agency (TMA) organizes and manages climatic condition records of the country. These records are derived from the weather stations that make hourly observations on variables such as solar radiation, hours of bright sunshine, maximum and minimum temperature and humidity.

The ground data used in this work was obtained from a synaptic station in Dodoma, Tanzania. These data are readily available in R-Instat through being made freely available by TMA. R - Instat is an open source front end to the statistics language R.

The data had other climatic records, rainfall, maximum and minimum temperatures. Figs. 2.3 and 2.4 show the first and the last few observations of sunshine duration data.

| Year | Month | Day | Rain (mm) | Tmax (°C) | Tmin (°C) | Sunh (h) |
|------|-------|-----|-----------|-----------|-----------|---------|
| 1    | 1974  | 1   | 7.7       | 31.2      | 19.5      | 10.9    |
| 2    | 1974  | 2   | 0.0       | 31.6      | 18.4      | 11.3    |
| 3    | 1974  | 3   | 0.0       | 29.9      | 17.8      | 10.2    |
| 4    | 1974  | 4   | 0.0       | 31.0      | 17.9      | 11.5    |
| 5    | 1974  | 5   | 0.0       | 31.7      | 18.6      | 11.3    |
| 6    | 1974  | 6   | 0.0       | 31.9      | 18.7      | 11.6    |

Fig. 2.3. The first 6 records of Dodoma sunshine duration hours

| Year | Month | Day | Rain (mm) | Tmax (°C) | Tmin (°C) | Sunh (h) |
|------|-------|-----|-----------|-----------|-----------|---------|
| 14240 | Dec   | 26  | 0.0       | 30.6      | 18.1      | 11.2    |
| 14241 | Dec   | 27  | 0.0       | 30.7      | 19.6      | 11.5    |
| 14242 | Dec   | 28  | 0.0       | 27.0      | 20.0      | 10.0    |
| 14243 | Dec   | 29  | 0.0       | 29.9      | 21.2      | 4.1     |
| 14244 | Dec   | 30  | 8.5       | 30.5      | 20.5      | 8.2     |
| 14245 | Dec   | 31  | 0.0       | 31.0      | 19.4      | 6.6     |

Fig. 2.4. The last 6 records of Dodoma sunshine duration hours
The column of interest for this work is the sunshine duration.

This data was from 1983 to make its comparison with the satellite data easy. The subset data had 10958 entries from Jan 1983 to Dec 2012. Out of these, a total of 3429 observations were missing values, with completely missing values for the years 1994, 1995, 1996, 1997 and 1998. Leaving out these years gave a total of 9131 observations with 1599 missing values. This makes approximately 18% of the values missing. Fig. 2.5 shows the missing sections of the data.

![Fig. 2.5. Missing values of sunshine hours](image)

2.2 Data preparation

2.2.1 Satellite data

CMSAF provides, other than the data, R scripts that can be used to unzip the data folder. The script Prep.Data.R was used to uncompress the data. This gave several data files in Network Common Data Format (netCDF) with a single file containing all the files merged. To get the data in csv format, the open NetCDF option from climate menu in R-Instat was used. Fig. 2.6 shows a snapshot of the climate menu of R-Instat 0.2.3.

![Fig. 2.6. Climate menu of R-Instat version 0.2.3](image)

Clicking on open NetCDF opens a dialogue box that allows one to select the file to be opened. The opened file can then be exported as a CSV file.
Because the ground data had missing values from 1994 to 1998, the satellite data was split into two, the lower data from 1983 to 1992 and the upper data ranging from 1999 to 2012. Fig. 2.7 and Fig. 2.8 show the first 6 elements of the lower and upper datasets, respectively.

| period | lat  | lon  | station SIS |
|--------|------|------|-------------|
| 1      | 0    | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | NA |
| 2      | 0    | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | NA |
| 3      | 1    | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | 0  |
| 4      | 1    | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | 0  |
| 5      | 2    | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | 0  |
| 6      | 2    | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | 0  |

**Fig. 2.7. The first 6 elements of the lower data**

| period | lat  | lon  | station SIS |
|--------|------|------|-------------|
| 1      | 140256 | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | NA |
| 2      | 140256 | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | NA |
| 3      | 140257 | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | 0  |
| 4      | 140257 | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | 0  |
| 5      | 140258 | -6.149998 | 35.70 | S6.1499981880188_E35.7000045776367 | 0  |
| 6      | 140258 | -6.149998 | 35.75 | S6.1499981880188_E35.7000045776367 | 0  |

**Fig. 2.8. The first 6 elements of the upper data**

The period columns in both datasets were converted to respective dates indicating the year, month, date and hour. The data was then subset by removing the night hours (7 PM – 5 AM) because no radiation is received. Figs 2.9 and 2.10 show the lower and upper data with the dates after sub-setting.

| lat1 | lon1 | station1 | period | split | date_year | month | day | hour | SIS |
|------|------|----------|--------|-------|-----------|-------|-----|------|-----|
| 15   | -6.149998 | 35.70   | 1      | 4     | 1983-01-01 | 1983  | 1   | 7   | 4   |
| 16   | -6.149998 | 35.75   | 2      | 4     | 1983-01-01 | 1983  | 1   | 7   | 5   |
| 17   | -6.149998 | 35.70   | 1      | 5     | 1983-01-01 | 1983  | 1   | 8   | NA  |
| 18   | -6.149998 | 35.75   | 2      | 5     | 1983-01-01 | 1983  | 1   | 8   | NA  |
| 19   | -6.149998 | 35.70   | 1      | 6     | 1983-01-01 | 1983  | 1   | 9   | 48  |
| 20   | -6.149998 | 35.75   | 2      | 6     | 1983-01-01 | 1983  | 1   | 9   | 44  |

**Fig. 2.9. First 6 elements of the lower dataset after subsetting**

| lat1 | lon1 | station1 | period | split | date_year | month | day | hour | SIS |
|------|------|----------|--------|-------|-----------|-------|-----|------|-----|
| 15   | -6.149998 | 35.70   | 1      | 140261 | 1999-01-01 | 1999  | 1   | 7   | 167 |
| 16   | -6.149998 | 35.75   | 2      | 140261 | 1999-01-01 | 1999  | 1   | 7   | 237 |
| 17   | -6.149998 | 35.70   | 1      | 140262 | 1999-01-01 | 1999  | 1   | 8   | 393 |
| 18   | -6.149998 | 35.75   | 2      | 140262 | 1999-01-01 | 1999  | 1   | 8   | 431 |
| 19   | -6.149998 | 35.70   | 1      | 140263 | 1999-01-01 | 1999  | 1   | 9   | 708 |
| 20   | -6.149998 | 35.75   | 2      | 140263 | 1999-01-01 | 1999  | 1   | 9   | 699 |

**Fig. 2.10. First 6 elements of the upper dataset after subsetting**
The values were then aggregated to daily data for easy comparison with the daily sunshine hours. The lower dataset had a total of 3652 observations with 147 missing, while the upper, out of 5479 observations, 80 were missing. This gives a total of 227 missing values out of 9131 observations, translating to 2.5% of the satellite observations missing. Fig. 2.11 shows the missing values in SIS derived from the satellite.

![Fig. 2.11. Missing values of SIS](image)

2.2.2 Dodoma data on radiation

Other than sub-setting, there were no further preparations that were done on it, since it was in a format that can be analyzed.

3 Methodology

Here we elucidate the methods used to address the main objective of this work. The methods used to compare satellite data and ground data are described. The chapter also describes the probability distributions functions that were used to describe the distribution of solar radiation.

3.1 Comparison of radiation data sets

The comparison of the two radiation datasets (Satellite and Ground Data) is to be done in two ways;

(i) Drawing of comparative plots of surface solar irradiance from the satellite and the sunshine hours from Dodoma.

(ii) Calculation of statistical parameters, mean percentage error (MPE), correlation (r), mean absolute deviation (MAD) and mean bias error (MBE) were calculated from the mean monthly SIS data as follows [24];

Let \( D_{gi} \) be measurements from the ground and \( D_{si} \) be measurements from the satellite;
The statistical parameters described are to be used to give an idea of whether satellite-derived measurements can be used to fill the gap in the ground measurements.

3.2 Distribution of solar irradiance

Because of their reliability when it comes to solving modern technology systems problems whose functionality depends on the reliability of its components, Log-normal, Gamma, and Weibull distributions are tested to get the best fit for solar radiation in a given month.

(i) Log-normal distribution. If a random variable \( Y = \ln(X) \) has a normal distribution with mean \( \mu \) and standard deviation \( \sigma \), the continuous random variable \( X \) is said to have a log normal distribution, and its density function is

\[
f(x|\mu, \sigma) = \frac{1}{\sigma \sqrt{2\pi x}} \exp \left( -\frac{1}{2\sigma^2} \ln(x) - \mu \right) \quad \text{for } x \geq 0.
\]  

The mean and variance of the distribution are

\[
\text{mean} = \exp \left( \mu + \frac{1}{2} \sigma^2 \right)
\]

\[
\text{variance} = (\exp (2\mu + \sigma^2)) \exp (\sigma^2) - 1.
\]

(ii) Gamma distribution. If a random variable \( X \) has a Gamma distribution, its density function is

\[
f(x|a, b) = \frac{1}{b^a \Gamma(a)} x^{a-1} e^{-x/b}, \quad X > 0.
\]

With \( \Gamma(a) = \int_0^\infty x^{a-1} e^{-x} \, dx \).

Where \( a \) and \( b \) are parameters of the Gamma distribution. The mean and variance of Gamma distribution are \( \mu = ab \), and \( \sigma^2 = ab^2 \).
Weibull distribution. This distribution has become more popular due to its ability to fit the data from many fields such as engineering, weather, and life data, to mention but a few. The general Weibull distribution has the density function

$$f(x|a, b, c) = \frac{a}{b} \left(\frac{x-a}{b}\right)^{a-1} e^{-\left(\frac{x-a}{b}\right)^a}, \quad x \geq a, a \in \mathbb{R}, \text{ and } b, c \in \mathbb{R}^+, \quad \text{(3.2.3)}$$

where $a$ is the location parameter, $b$ is the scale parameter and $c$ the shape parameter. The Weibull distribution used in this work assumes that a failure can occur at point zero. Therefore, the location parameter is zero. This gives a distribution that has the density distribution

$$f(x|a, b) = \frac{a}{b} \left(\frac{x}{b}\right)^{a-1} e^{-\left(\frac{x}{b}\right)^a}. \quad \text{(3.2.4)}$$

The mean and variance of the Weibull distribution are

$$\mu = b \Gamma \left(\frac{a}{a} + 1\right), \text{and } \sigma^2 = b^2 \left\{ \Gamma \left(\frac{a}{a} + 1\right) - \Gamma^2 \left(\frac{a}{a} + 1\right) \right\}.$$ 

The probability distribution function that best describes the distribution of the solar irradiance in a given month, in Dodoma, was found to be Weibull distribution.

## 4 Analysis and Results

The results of the analysis are presented. The data were processed and subsequently analyzed in response to the main objective of this work. The findings presented in this chapter demonstrate the potential of using satellite-derived radiation data as a substitute of scarce radiation data measured from the ground. We start with the analysis of the ground data followed by satellite data and finally the comparison of the two datasets.

### 4.1 Fitted PDFs for solar radiation

When the PDFs of lognormal, Weibull and gamma distributions were fitted, Weibull was found to describe the data much better than the other two.

![Fig. 4.1. Fitted distribution for April](image1)

![Fig. 4.2. Fitted distribution for July](image2)

### 4.2 Sunshine Hours

Here, the analysis of the ground data is presented. Yearly plots are first presented. These include the box-plots and a time series plot showing the daily variation of sunshine hours over the years. The monthly plots and finally a histogram are plotted.
Fig. 4.4. Daily variation of sunshine hours over the years

The following observations were made from the yearly plots for the ground data that were drawn.

The box-plots in Fig. 4.3 show the daily sunshine hour split by years. It shows that over the years, the sunshine hours has been stationary, there is no pattern from one year to another. This is confirmed by the time series plot of the sunshine hours as shown in Fig. 4.4. The Ljung – Box test confirmed this from its small p-values which suggest that the series is stationary. The box-plots also show unusual features in the data like gaps and outliers. The years 1994 to 1998 had no records and thus were omitted from the analysis. Other than yearly plots, monthly box-plots of the sunshine hours was also plotted.
The box-plots of sunshine hours split by months (Fig. 4.5) showed some possibility of seasonality in sunshine hours; the minimum values are recorded in the months of March and April, and the maximum in the months of June to October. The month of September shows less variability in sunshine hours received. This is depicted by the smaller length of the boxplot. The opposite is seen in the months of December, January and February which show high variability in the sunshine hours.

In an attempt to know the distribution of the sunshine hours (where the majority of values falls and how much variation is in the data), a histogram was also plotted.

The distribution of sunshine hours in Fig. 4.6 shows heavy skew to the left. This is an indication that majority of values fall between 5 and 11, an evidence that Dodoma receives a lot of solar radiation. This section presented the analysis that were done using the ground data. For easy comparison, the same analysis were done for the satellite derived estimates as shown in section 4.3.
4.3 Satellite estimates of solar radiation

This section describes the analysis done using the satellite-derived estimates. The yearly plots followed by the monthly box-plots and finally the histogram of the satellite-derived measurements are presented.

The box-plots for daily SIS split by years, like those for sunshine hours, show that the amount of solar radiation received over the years is stationary. The box-plots also depicts unusual features, outliers and gaps, in the data. Most of these outliers are lower outliers, apart from a few especially in the year 1990, which shows very extreme values and thus are worth investigation. These outliers were investigated further in the comparison of the two datasets. The years 1994 to 1998 were omitted to allow easy comparison with the sunshine hours.

Fig. 4.7. SIS radiation by years

Fig. 4.8. Time series plot of the SIS
These time series plots also depict SIS stationarity as the SIS variations does not show any pattern over the years. It also shows the gap between 1994 to 1998, the years that were left out for easy comparison with the ground data. These are indications that the area under study receives on average, the same sunshine durations each year. The monthly variations were also investigated to get an insight of the distribution of the sunshine hours in each month.

![Box-plot of SIS radiation by months](image)

**Fig. 4.9. SIS radiation by months**

The box-plots of SIS split by months shows that the months of January, February, March, August, September, October, November and December record larger measurements of solar radiation, with the month of September and October recording the highest amount. It also shows that the month of September receive almost the same amount of radiation throughout. This is shown by the small length of the box-plots, which is an indicator of less variability. The months of April to July receives lower surface solar irradiance, with the month of May recording the lowest readings. This may be because these are wet months.

![Histogram of SIS](image)

**Fig. 4.10. Histogram of SIS**

The histogram of SIS (Fig. 4.10), just like Fig. 4.6, shows that solar surface radiation is heavily skewed to the left. This is an indication that Dodoma receives a lot of solar radiation, with the majority of values falling between 400 and 750.
To find the extent of the difference between the datasets, other statistical parameters, Correlation \( r \), MPE (\%), MAD (\( WM^{-2} \)), and MBE (\( WM^{-2} \)) were calculated.

### 4.4 Comparison of sunshine hours and solar radiation estimates

The exploratory analysis conducted shows that the two datasets, sunshine hours and the solar radiation estimates from satellite, are similar in distribution. Even though the two datasets are measured in different units, the similarity may be attributed to the fact that the underlying process generating the two datasets is the same.

The comparison of the two datasets was done in two levels, daily data and monthly data.

For easy comparison, the sunshine hours were converted to the same unit as the radiation data from satellite using Angstrom’s formulae\[25\],

\[
R_s = \left( a_s + b_s \frac{n}{N} \right) R_a, \tag{4.4.1}
\]

Where, \( R_s \) is the solar radiation in \( MJm^{-2}day^{-1} \), \( n \) is the actual duration of sunshine hours, \( N \) is the maximum possible duration of sunshine, \( a_s \) is the regression constant showing the fraction of extraterrestrial radiation reaching the earth on overcast days \((n = 0)\), \( a_s + b_s \) is the fraction of the extraterrestrial radiation reaching the surface on a clear day and \( R_a \)

\[
R_a = \frac{24(6.0)}{\pi} G_{sc} d_r [w_s \sin(f) \sin(\delta) + \cos(f) \cos(\delta) \sin(w_e)], \tag{4.4.2}
\]

Where is the solar constant, \( f \) is latitude in radians \( d_r = 1 + 0.0033 \cos \left( \frac{2\pi}{36} f \right) \), \( \delta = 0.409 \sin \left( \frac{2\pi}{36} f - 1.39 \right) \) and \( w_s = \arccos \left( -\tan(f) \tan(\delta) \right) \).

For the area under study, \( f = -0.0342 \pi, \ G_{sc} = 1369 Wm^{-2}, \ N = 12.1 \), the maximum possible sunshine duration. Because there were no actual solar radiation data, no calibration was done to change, \( a_s \) and \( b_s \), therefore, the default values of \( a_s = 0.25 \) and \( b_s = 0.5 \) were used.

#### 4.4.1 Evaluation of monthly data

The comparison of monthly ground data calculated from sunshine hours and monthly satellite data was done using the monthly sums. It showed a bias of 727.2, mean percentage error of 4.04 and mean absolute deviation of 1194.40, as shown in Table 1 This small MPE indicates that the two datasets are similar. The correlation between satellite data and ground data is strong with a value of 0.687. However, the months of March, April and September showed lower values of 0.402, 0.572 and 0.25, respectively, possibly because these are wet months. The months of May, June, July and August recorded the highest bias between the two datasets, probably because of the higher number of sunshine hours during this period.
Table 1. Bias, MAD, MPE and correlation ($r$) for the comparison of Monthly ground data and satellite data

|       | JAN | FEB | MAR | APRIL | MAY | JUN | JUL | AUG | SEP | OCT | NOV | DEC | MEAN |
|-------|-----|-----|-----|-------|-----|-----|-----|-----|-----|-----|-----|-----|------|
| $r$   | 0.858 | 0.8734 | 0.572 | 0.735 | 0.6211 | 0.867 | 0.71 | 0.25 | 0.7167 | 0.779 | 0.855 | 0.687 |
| MPE   | -2.85 | -3.78 | -2.81 | 4.37 | 15.94 | 18.07 | 14.24 | 7.62 | 0.34 | -2.09 | 1.51 | -2.10 | 4.04 |
| MBE   | -406.5 | -569.2 | -402.3 | 722.5 | 2577.1 | 3178.4 | 2542.5 | 1357.8 | 64.7 | -351.1 | 274.0 | -261.1 | 727.2 |
| MAD   | 617.0 | 668.8 | 841.3 | 835.2 | 2577.1 | 3178.4 | 2542.5 | 1357.8 | 303.8 | 357.2 | 542.1 | 511.66 | 1194.40 |

Table 2. Bias, MAD, MPE and correlation for the comparison of daily ground data and satellite data

|       | JAN | FEB | MAR | APR | MAY | JUN | JUL | AUG | SEP | OCT | NOV | DEC | MEAN |
|-------|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|------|
| $r$   | 0.872 | 0.818 | 0.708 | 0.648 | 0.651 | 0.633 | 0.454 | 0.384 | 0.757 | 0.877 | 0.877 | 0.683 |
| MPE   | -4.91 | -4.09 | -4.90 | 2.44 | 14.0 | 16.4 | 12.10 | 6.76 | -0.21 | -3.64 | -0.11 | -3.90 | 2.495 |
| MBE   | -15.28 | -15.27 | -13.47 | 23.12 | 80.72 | 97.23 | 73.99 | 42.60 | 1.19 | -10.42 | 7.22 | -8.24 | 21.95 |
| MAD   | 44.38 | 41.63 | 56.70 | 50.98 | 89.12 | 98.31 | 77.64 | 49.67 | 30.74 | 37.57 | 41.23 | 42.40 | 55.03 |

Table 3. Monthly statistical parameters after removing outliers

|       | JAN | FEB | MAR | APR | MAY | JUN | JUL | AUG | SEP | OCT | NOV | DEC | MEAN |
|-------|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|------|
| $r$   | 0.86 | 0.87 | 0.44 | 0.77 | 0.74 | 0.68 | 0.87 | 0.69 | 0.21 | 0.718 | 0.79 | 0.86 | 0.71 |
| MPE   | -4.30 | -5.37 | -2.76 | 4.49 | 14.93 | 17.30 | 13.29 | 6.57 | -0.69 | -2.92 | -0.48 | -3.13 | 3.08 |
| MBE   | -628.22 | -771.76 | -394.79 | 733.04 | 2405.14 | 3008.21 | 2338.72 | 1157.39 | -120.92 | -496.87 | -440.44 | 559.87 |
| MAD   | 711.73 | 786.57 | 747.38 | 738.29 | 2405.14 | 3008.21 | 2338.72 | 1157.39 | 317.64 | 496.87 | 454.05 | 469.15 | 1135.93 |

Table 4. Daily statistical parameters after removing outliers

|       | JAN | FEB | MAR | APR | MAY | JUN | JUL | AUG | SEP | OCT | NOV | DEC | MEAN |
|-------|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|------|
| $r$   | 0.86 | 0.87 | 0.44 | 0.77 | 0.74 | 0.68 | 0.87 | 0.69 | 0.21 | 0.718 | 0.79 | 0.86 | 0.71 |
| MPE   | -4.32 | -4.03 | -3.18 | 3.40 | 13.79 | 15.94 | 11.52 | 5.99 | -1.06 | -4.03 | -0.98 | -3.55 | 2.46 |
| MBE   | -16.63 | -17.0 | -9.91 | 24.47 | 78.04 | 93.01 | 68.62 | 36.54 | -5.22 | -15.24 | 0.07 | -10.62 | 18.84 |
| MAD   | 39.36 | 33.34 | 43.02 | 43.78 | 84.75 | 94.05 | 71.61 | 43.44 | 30.23 | 38.36 | 38.0 | 39.0 | 50.32 |

Fig. 3.9. SIS variation by month
Combined line graphs of ground and satellite data. The red lines represent the data collected from the ground while the blue lines represent the satellite data. Other than the monthly analysis, daily analysis was also done. This is shown in subsection 4.4.2.

4.4.2 Evaluation of daily data

The comparison of ground data calculated from sunshine hours and satellite data showed a bias of 21.95, Mean percentage error of 2.495 and mean absolute deviation of 55.03, as shown in Table 2. This is an indication that there is no big difference between the two datasets and therefore, the satellite data are a good approximation of the ground truth.

Also, the correlation between satellite data and ground data is strong with a value of 0.683, but is lower than the correlation from the monthly data. However, the months of September and October showed lower values of 0.454 and 0.384, respectively, an indication that the satellite data and ground data correlate well in some months and not others, probably because of outliers or the season. This was further investigated.

As stated earlier, there were a number of outliers with very extreme values in 1990, the upper outliers of this year, which start at about 820, corresponds to very high sunshine hours, 15.1, that cannot be reached at any particular time in the area under study. These outliers, together with those that are very low, below 150, were left out, and the statistical parameters calculated again.

The monthly and daily statistical parameters after removing outliers are shown in Tables 3 and 4. The new average correlation increased to 0.71, and the other statistical parameters reduced considerably. However, the months of January, February and December consistently showed high correlation, and the month of September consistently low correlations. Because of this, the months of February and September were investigated further to get an insight of what can be the cause of these differences.

For the month of February, it was found that there is linear spread of sunshine hours from minimum value of 0.0 to the maximum value of 12 as depicted by the mean 7.9. Fig. 4.12 shows the spread of the radiation data from the two sources for the month of February. The solar radiation estimates from the satellite were also found to have a good spread with a mean of 533.9. The trend of the two sources of radiation were found to be similar throughout the month Fig. 4.14. The high correlation can therefore be attributed to the linear trend of data in this month.

In the month of September, out of 688 observations, with the maximum record as 11.4, only 267 are less than 10. This is a clear indication of clustering of observations. High average record of 10.07 justify the clustering observed. The same pattern was seen in solar radiation estimates from satellite, where only 218 observations out of 688 had values less than 600. Fig. 4.13 shows the spread of ra-diation data from the two sources for the month of September. The low correlation can be explained by the non-uniform distribution in the recorded data. In spite of the low correlation and the uneven spread of records, some general trends were observed. The records from the two data sources showed some general trend over the years (Fig. 4.11) and the month Fig. 4.15.

In this chapter, we compared the radiation data from satellite and data from the ground. Some months showed very high correlation, while others recorded low correlations. The high correlation was found to be due to the good spread and similar general trends. Low correlations were found to be mainly due to clustered observations. However, the records in the months with low correlation showed similar general trend over the years as well as daily observations within the months. Our general feeling therefore is that the two data source shows a good similarity, and satellite data can be used for places where the data is not available to provide continuous data.
5 Conclusion

The satellite data from EUMETSAT is the first freely available radiation data record for hourly, daily and monthly surface solar irradiance covering a region bounded between latitude $\pm 65^\circ$ and longitude $\pm 65^\circ$ for the time period 1983 to 2013. It has high quality, seen from fewer missing values and high spatial resolution $0.05^\circ \times 0.05^\circ$. Hence, it is qualified for many climate applications such as trend analysis and variation. This work has looked at the possibility of using satellite data to fill the gap of weather data shortage in Africa to help in making well informed climatic decisions.

In this work, we mainly considered solar radiation as a climatic element in Dodoma, Tanzania. The satellite data that was freely obtained from CMSAF was compared with the actual measurements from Dodoma.

Comparison showed that the two data sets are similar in distribution. The correlation of the two data sets was also found to be 0.71 on average for both monthly and daily records, which is an indication that the two data sets are similar, and satellite data can be used to fill in the gaps in ground measured records as well as be used in places where the records are not available.

From the work, a Weibull probability distribution best describes the distribution of solar radiation of all the months for the area under study. This can be used to estimate the energy output a given photovoltaic cell can produce and therefore it is possible to select a solar panel that is suitable for a particular region.
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Competing Interests

Authors have declared that no competing interests exist.

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