3mm

Correspondence:
Author responses to reviews to "The Community Inversion Framework v1.0: a unified system for atmospheric inversion studies"

Antoine Berchet1,*, Espen Sollum2, Rona L. Thompson2, Isabelle Pison1, Joël Thanwerdas1, Grégoire Broquet1, Frédéric Chevallier1, Tuula Aalto3, Adrien Berchet14, Peter Bergamaschi4, Dominik Brunner5, Richard Engelen6, Audrey Fortems-Cheiney1, Christoph Gerbig7, Christine D. Groot Zwaartink2, Jean-Matthieu Haussaire5, Stephan Henne5, Sander Houweling8, Ute Kartens9, Werner L. Kutsch10, Ingrid T. Luijkx11, Guillaume Monteil9, Paul I. Palmer12, Jacob C. A. van Peet8, Wouter Peters11,13, Philippe Peylin1, Elise Potier1, Christian Rödenbeck7, Marielle Saunois1, Marko Scholze9, Aki Tsuruta3 and Yuanhong Zhao1

1Laboratoire des Sciences du Climat et de l’Environnement, CEA-CNRS-UVSQ, Gif-sur-Yvette, France
2Norwegian Institute for Air Research (NILU), Kjeller, Norway
3Finnish Meteorological Institute (FMI), Helsinki, Finland
4European Commission Joint Research Centre, Ispra (Va), Italy
5Swiss Federal Laboratories for Materials Science and Technology (Empa), Dübendorf, Switzerland
6European Centre for Medium-Range Weather Forecasts, Reading, RG2 9AX, UK
7Max Planck Institute for Biogeochemistry, Jena, Germany
8Vrije Universiteit Amsterdam, Department of Earth Sciences, Earth and Climate Cluster, Amsterdam, the Netherlands
9Dep. of Physical Geography and Ecosystem Science, Lund University, Sweden
10Integrated Carbon Observation System (ICOS-ERIC), Helsinki, Finland
11Meteorology and Air Quality Group, Wageningen University and Research, Wageningen, the Netherlands
12School of GeoSciences, University of Edinburgh, Edinburgh, EH9 3FF, UK
13Centre for Isotope Research, University of Groningen, Groningen, the Netherlands
14Institut Pprime (UPR 3346, CNRS-Université de Poitiers-ENSMA), Bat H2, 11 Boulevard Marie et Pierre Curie, TSA 51124, 86073, Poitiers Cedex 9, France

Correspondence: antoine.berchet@lsce.ipsl.fr

We thank the referees for their detailed and fruitful comments on our manuscript. We reproduce below their reviews and embed our responses in bold blue text in their comments. Revised section of the manuscript are reproduced in italic blue. We also provide a track-change manuscript at the end of the present document.

1 Referee #1

The submitted paper presents the Community Inversion Framework (CIF) to help rationalise development efforts and leverage the strengths of individual inversion systems into a comprehensive framework. The CIF is primarily a programming protocol to allow various inversion bricks to be exchanged among researchers. The system is supposed to allow running different atmospheric transport models, different observation streams and different data assimilation approaches. The paper describes a
system that will bring a major software advances for future inversion frameworks. The presentation is clear and well written. I have few main major concerns detailed in the main comments below.

1.1 Main comments

The paper advertises and the multi-CTM or model capability, but this is not shown in the paper only mentioning a future paper. This capability needs to be showcased in the paper or such claims should be removed or rephrased.

The CIF is a multi-purpose systems. The two main purposes are to enable the comparisons of different inversion methods with a given inversion case, and the comparison of different transport models for a given inversion configuration. We decided to split the two aspects into two different papers for clarity. We clarified the text accordingly

p4. l. 14: In the present paper, we lay out the basis of the CIF, giving details on its underlying principles and overall implementation. The proof-of-concept focuses on the implementation of several inversion methods, illustrated with a test case. We will dedicate a future paper to the evaluation of the system on a real-life problem with a number of interfaced atmospheric (chemistry-)transport models.

It is a bit limiting to narrow the scope only on GHG as inversion of reactive species and primary aerosols is quite relevant as well. Especially if CIF plans to, or already includes the CHIMERE CTM which is originally designed for air quality applications.

The CIF is primarily designed by members of the GHG community, thus the bias in presentation. However all the conceptual environment of the CIF can be applied directly to chemistry, as it is the case in CHIMERE, which solves complex chemical schemes reproducing the full chemistry of O3, NO2, VOC, CO, etc., as well as in LMDZ e.g. with simple chemical schemes around methane, CO or nitrous oxide. We clarified the text accordingly

p3 l34: In particular, although primarily designed for GHGs applications, the CIF is based on a general structure that will allow applications to air quality data assimilation.

Air quality applications are also mentioned throughout the introduction more explicitly.

The ensemble and analytical results point out possible problems in the implementation of the EnKF and in the analytical solution using correlation length scales.

There was indeed a difference in the way the correlations were computed in the analytical and the ensemble method, versus the variational method. We fixed the issue and also updated the test cases to better highlight the consistency of the results.

It is not entirely clear on how the system uses the ensemble information to perform the inversion. This ensemble-based inversion should not be called an EnKF as it is not sequential. This needs to be revised or diagnosed or explained.

We agree that in this application there is no time-stepping in the EnKF solution and hence the word Filter is not appropriate. However, the implementation of the solution methods allows it, and algorithmically follows the Ensemble Square Root Filter originally described in Whitaker and Hamil, (2002) and later used in Peters et al., (2005). The important distinction from the other methods is the use of the statistics of the ensemble to solve the Kalman equations,
and the possibility to solve non-linear problems. We think the name EnSRF is therefore suited, and will use it in the revised version. We also include the note about the time-stepping.

The section on Ensemble methods (p8-10) has been extensively rewritten to clarify the difference between EnSRF and other ensemble methods. See track-change manuscript

Finally, the testing of system using the Gaussian plumes on my local machine (Mac) was impossible for the time I gave in. I regularly use python and successfully setup conda on my local machine. But it seems that the CIF is dependent on many specific packages that bring conflict issues and are not straightforward to install (specifically GDAL seems problematic). It required some research to find different commands to install the packages from what is indicated on the online documentation. After trying for hours, I had to give up as my time to revise this paper is limited.

We are sorry to hear about difficulties to install our system. We will give more details on the installation in the documentation web pages. To make the testing of our system easier, we have now set up a docker image (publicly available here: https://hub.docker.com/repository/docker/pycif/pycif-ubuntu) with all packages correctly installed and demonstration data available as well. Thus, new users can simply download the docker image and run it on their system, avoiding complicated configuration of libraries.

The docker image is provided with the corresponding configuration command lines to be run on a standard Ubuntu system to obtain the same environment. These lines can be used to adapt to other OS and distributions.

More specifically, we agree that the GDAL library can create some issues when not correctly installed. However, GDAL is standard in geoinformation systems and such a library is essential for all geographical operations, such as reprojecting from one grid to another, or finding in which grid cell an observation point is. In very regular grids, such a library is not necessary, but the CIF is design to accommodate many transport models and input data, which are not necessarily always on regular lon/lat grid.

The docker image is now mentioned in Sect 3.1 p.17, l19: To foster portability and dissemination, a dedicated Docker image is distributed with pyCIF, providing a stable environment to run the system and enabling full reproducibility of the results from one machine to the other.

1.2 Specific comments

– P3, L8: Define better what a unified system would be here.

  Rephrased p3 l11 as: A unified system, as a community platform running multiple transport models, with diverse inversion methods, would provide ...

– P3, L12: OOPS main goals are on NWP. The CAMS system will run on OOPS (which is based on IFS) which will run inversions in the future. Maybe it is worth mentioning the CoCO2 projects (formerly CHE). The authors should also mention the Joint Effort for Data Integration led by UCAR/JCSDA. https://www.jcsda.org/jcsda-project-jedi.

The CAMS operational system will indeed run with OOPS, to be adapted for flux inversion purposes. The CIF is part of the CoCO2 project and is used as an interface between the operational community of CAMS and the
research community. The CIF will also be used for the design of operational systems at the national and regional scales in Europe. We clarified this aspect and also mention the JEDI effort as follows:

p3 l15: "Collaborative efforts towards unified systems are already happening in other data assimilation communities, with, e.g., the Object-Oriented Prediction System (OOPS; coordinated by the European Centre for Medium-range Weather Forecast, UK), or the Joint Effort for Data Integration (led by UCAR/JCSDA; https://www.jcsda.org/jcsda-project-jedi). The Data Assimilation Research Testbed (DART) is also an example of collective effort proposing common data assimilation scripts for diverse applications (e.g., Earth system, chemical species inversions)"

– P3, L13-14: DART is an EnKF (EAKF to be more precise) that allows to run on the CESM earth system model. There is also a WRF mode. I would not call this unified or modular system compared to OOPS, CIF or JEDI. Also, DART can run atmospheric composition and recently chemical species inversions as well (see Gaubert et al., 2020).

We thank the reviewer for these insights and have updated the text accordingly (see reply above).

– P5, L10: Some of the models mentioned here use semi-lagrangian advection scheme. Maybe remove Eulerian and use a different word.

We rephrased as follows: p5 119:

[... ] by numerical (chemistry-)transport models. These can be of various types e.g. global circulation models (e.g., LMDZ, Chevallier et al. 2010; TM5, Houweling et al. 2014; GEOS-Chem, van der Laan-Luijkx et al. 2017; Liu et al. 2015; Palmer et al. 2019; Feng et al. 2017; NICAM, Niwa et al. 2017), regional Eulerian chemistry-transport models (e.g., CHIMERE, Broquet et al. 2011; Fortems-Chiney et al. 2019; WRF-CHEM, Zheng et al. 2018; COSMO-GHG, Kuhlmann et al. 2019; LOTOS-EUROS, Curier et al. 2012) or Lagrangian particle dispersion models (e.g., FLEXPART, Thompson and Stohl 2014; STILT, Bagley et al. 2017; Brioude et al. 2013; Trusilova et al. 2010).

– P5, L11: Mizzi et al., 2016 is a WRF CHEM paper and also do not perform any source inversion but only concentration assimilation. Please consider removing this reference.

There was a mistake in our references indeed. We have now replaced Mizzi by Kuhlman et al. 2019.

– Equation 2: Does N indicates a normal distribution? Please clarified. Also, in the second brace element, is this zero in the first moment? If yes, I guess this assumes unbiased observation and prior. I know this is a requirement for data assimilation formulation. But often not the case in practice in atmospheric composition. Maybe you can detail and develop this in the text.

N indeed indicates a normal distribution. This is now stated in the text. We also clarified the underlying assumption of unbiased errors. This reads as follow now:

" is to assume prior and observation spaces to be normal distributions, noted $\mathcal{N}(\cdot, \cdot)$ below, the first argument representing the average of the distribution and the second argument the covariance matrix. In addition, when assuming that the distributions in the state vector space and the observation space are independent from each other, and that errors in
the observation and the state vector spaces have Gaussian, unbiased distributions, it is possible mathematically derive
the Bayes theorem and to represent the distributions of Eq. (2) as follows:

\[ \text{EQUATION 2} \]

with \( B \) and \( A \) the prior and posterior covariance matrix of uncertainties in the target vector, \( x^b \) and \( x^a \) the prior and posterior target vectors and \( R \) the covariance matrix of uncertainties in the observation vector and the observation operator."

The assumption that errors are unbiased, which makes it possible to write normal distributions in Eq. (1) with means
\( x^b \), \( 0 \) and \( x^a \) respectively, is needed to simplify the formulation of the Bayesian problem in atmospheric inversions. Neglecting error biases have known impacts on inversion results (Masarie et al., 2011); they can be accounted for
online as an unknown to be solved by the inversion (Zammit-Mangion et al., 2021), but are often treated offline from the
inversion, either through pre-processing of inputs or post-processing of outputs.

-- P6, L22-25: This is an important point, maybe the authors could expand more on this for the non-data assimilation expert?

**We give some more details, especially in terms of linearisation around a given point to approximate the observation operator by its Jacobian.**

p7 111: Analytical inversions can also be used on slightly non-linear problems, by linearizing the observation operator
around a given reference point using the tangent linear of the observation operator. It formulates as follows:

\[ \mathcal{H}(x^b+\delta x) \approx \mathcal{H}(x^b) + d\mathcal{H}_{x^b}(\delta x) = \mathcal{H}(x^b) + H_{x^b}\delta x \] (1)

with \( \delta x \) a small deviation from \( x^b \) within a domain where the linear assumption is valid, \( d\mathcal{H}_{x^b} \) the tangent-linear of \( \mathcal{H} \) at \( x^b \) and \( H_{x^b} \) the Jacobian matrix of \( \mathcal{H} \) at \( x^b \).

Then Eq. (3) can be easily adapted by replacing \( (y^o - Hx^b) \) by \( (y^o - \mathcal{H}(x^b)) \) and \( H \) by \( H_{x^b} \).

The computation of an analytical inversion faces two main computational limitations. First, the matrix \( H \) representing
the observation operator \( \mathcal{H} \) must be built explicitly. This can be done either column by column, in the so-called
response function method, or row by row, in the so-called footprint method. The two approaches require \( \dim(X) \), the
dimension of the target space and \( \dim(Y) \), the dimension of the observation space, independent simulations respectively.
In the response function method, each column is built by computing \( \{d\mathcal{H}_{x^b}(\delta x_i) \mid \forall \delta x_i \in B_X\} \) with \( B_X \) the canonical basis of the target space. For a given increment \( \delta x_i \), the corresponding column gives the sensitivity of observations to changes in the corresponding component of the target space. In the footprint method, each row is built by computing \( \{\mathcal{H}_{x^b}(\delta y_i) \mid \forall \delta y_i \in B_Y\} \) with \( B_Y \) the canonical basis of the observation space. For a given perturbation of \( \delta y_i \) of a component of the observation vector, the corresponding row of \( H \) gives the sensitivity of the inputs to that perturbation.
P7, L8: Are control space and target space the same thing or different? Please be consistent (I would recommend control space as it is more generally used in atmospheric DA) or explain the differences.

We think that the use of "control vector" can be misleading in some context and prefer "target vector" which is also widely used in the community. We have checked for the consistency of the use of target v control vector over all the manuscript.

P7, L12: This part of the sentence is not very clear. Replace "... problems such as the variational and the ensemble Kalman filter methods which are described below."

We have replaced by:

For this reason, smart adaptations of the inversion framework (including approximations and numerical solvers) are often necessary to tackle problems even when they are linear; in the following, we choose to elaborate on some of the most frequent approaches used in the atmospheric inversion community: the variational approach and one ensemble method, the Ensemble Square Root Filter.

P7, L18: “limited non-linearity”: This is not necessarily true, or it is misleading. EnKFs are frequently applied successfully on chemical transport models which can be significantly non-linear systems in the polluted boundary layer for example.

Indeed, the ensemble methods can in principle handle non-linear systems, although we found in a CO2-application of the EnSRF that non-linearity in the observation operator quickly degrades the analysis (Tolk et al., 2011). We have revised the text.

In the current version, only the EnSRF approach is implemented in the CIF. One should note that the EnSRF, as a direct approximation of the analytical solution, can be very sensitive to non-linearity in the observation operator (e.g., Tolk et al., 2011). It can generally cope only with slight non-linearity over the assimilation window, thus, the assimilation window length has to be chosen appropriately, contrary to other ensemble methods which are usually not sensitive to non-linearity.

P7, L25: “running computation window”: I believe the authors refer to what is commonly called assimilation window? (Note that inversion is part of data assimilation techniques, so it is fine to call it data assimilation window in the inversion framework). Please be consistent in the terminology in this paragraph. “Smaller”, smaller than what? I understand that for GHG inversions especially CO2 you need long window given the observation network and the CO2 atmospheric lifetime, but this is not general to all atmospheric composition inversion.

We have clarified the terminology. For a given period of interest, EnSRF computes smaller running assimilation windows. The length of the assimilation windows can vary depending on the scale, on the species of interest, the process of interest, the observations used, etc. All the ensemble methods section has been rewritten.
P7, L29-31: “for very dense...may not be sufficient”: This is not necessarily true, or the statement is misleading. This also needs to be proven. If the reference exists please add it. One can foresee that with a better coverage and higher satellite sensitivity towards the surface the window length could be reduced hence the number of observations to go through sequentially is reduced. Please correct or remove this statement or justify this more clearly.

We agree with the reviewer that this statement was too vague. The assimilation window must be shortened as much as possible to limit the number of observations, but too short windows generate their own problems of independence of windows and propagation of information between windows. We have clarified to:

for very dense observations, such as datasets from new-generation high-resolution satellites, the sequential assimilation of observations may not be sufficient, or at least methods may be needed to make the observation errors between sequential assimilation windows independent, for example by applying a whitening transformation to the observations to form a new set with uncorrelated errors as suggested by Tippett et al., 2003. The challenge is exacerbated for long-lived species such as CO2, for which assimilation windows must be long enough to maintain the propagation of information on the fluxes on long distances through transport. Propagating an covariance matrix from assimilation windows to assimilation windows as accurate as possible could in principle limit the later issue, as suggested in Kang et al., 2011 and 2012, but could still prove hard to apply in very high resolution problems.

P8, L11: “under-estimating”: This is not necessarily true. Small ensemble size can over and underestimate the uncertainty. More importantly they introduce spuriousness. Please clarified the statement.

We have clarified the statement. Small ensembles can cause the posterior ensemble to collapse, i.e., the posterior distribution is dominated by one or a very small number of members, which do not allow for a reliable assessment of the posterior uncertainties (Morzfeld et al., 2017); moreover, small ensembles introduce spuriousness in the posterior uncertainties, with unrealistic correlations being artificially generated.

p10: By using random sampling, ensemble methods are able to approximate large dimensional matrices at a reduced cost without using the adjoint of the observation operator (see variational inversion below) that can be challenging to implement. Small ensembles generally cause the posterior ensemble to collapse, i.e., the posterior distribution is dominated by one or a very small number of members, which does not allow for a reliable assessment of the posterior uncertainties (Morzfeld et al., 2017); moreover, small ensembles introduce spuriousness in the posterior uncertainties, with unrealistic correlations being artificially generated. In the EnSRF, small ensembles rather cause a misrepresentation of uncertainty structures, which limits the accuracy of the computed solution, and may require fixes as described in, e.g., Bocquet (2011). In any case, the level of approximation necessary for this approach to work is strongly different for each problem, which requires preliminary studies before consistent application. In particular, the so-called localization of the ensemble can be used to improve the consistency of the inversion outputs (e.g., Zupanski et al., 2007; Babenhauserheide et al., 2015).

In the current version, only the EnSRF approach is implemented in the CIF. One should note that the EnSRF, as a direct approximation of the analytical solution, can be very sensitive to non-linearity in the observation operator (e.g., Tolk
et al., 2011). It can generally cope only with slight non-linearity over the assimilation window, thus, the assimilation window length has to be chosen appropriately, contrary to other ensemble methods which are usually not sensitive to non-linearity.

- P8, L13-14: Could you possibly provide few examples of different problems having different approximations?

> We give two examples of past studies who discussed the use of the localization of the Monte Carlo ensemble to help the filter keeping consistent results.

- P8 L15-28: I recommend revising this introduction of the variational method. Add a sentence to present the cost function. Put the equation right after. Explain in few sentences that the aim is to minimize this cost function and that to achieve this the gradient is calculated. Then put the gradient formula.

> We have re-arranged the variational section to clarify it. See track-change for detailed modifications. In particular, we give details on how the posterior uncertainties are computed in variational inversions in the new version of the manuscript.

- P8, L25: I am not sure this is true. If the state vector is large, i.e. increasing resolution or augmenting the control variables, or increasing the number of observations, the minimization is expected to be slower.

> The limiting factor in the analytical inversion approach is the size of H which cannot be broken down (unlike B). The variational approach still has limitations but memory is unlikely to be a problem since B can be broken down, only the computational time could be a limiting factor.

> Technically, crazily huge target vector will impede the capability of variational inversions to be computed, if only due to memory limitations, and convergence time. However, these limits come for much bigger problems than for analytical and ensemble methods. Also, the minimization algorithms such as CONGRAD and M1QN3 are extremely efficient and manage a relatively good level of convergence with only a few tens of iterations even in very large dimensional problems (when correctly set up). For very large inversion problems (especially in the case of non-linearity), variational inversions still need thorough set up to guarantee the convergence.

- P12, Equation 12: Please, be explicit about what the index represents. "... of elementary interchangeable N transformations ..."

> We now give more details about this equation. This generalization of the decomposition of the observation operators allow us to simplify the configuration and development of inversions. In particular, the development of new functions is well identified and makes the development of the adjoint simpler.

p15 l15: In such a formalism, all intermediate transformations have the same conceptual level in the code. They are functions ranging from spatial reprojection, to temporal interpolations, to more complex operations such as the reconstruction of satellite total columns from concentrations simulated at individual levels in the transport model. In the CIF, all these transformations have the same input and output structure and, thus, their order can be changed seamlessly
to execute a given configuration. Please note that the commutative property of elementary transformations as pieces of
code does not guarantee the commutative property of the corresponding physical operators.

- P13, L10: It is a bit of pity that air quality application is only mentioned here. I think a system such as CIF could greatly
benefit the air quality community as much as the GHG community, which are getting more connected lately.

We agree. As mentioned above, the CIF is being developed by members of the GHG community primarily.
However, air quality applications are also included in the purposes of the CIF, as e.g., the chemistry transport
CHIMERE is implemented in the CIF with complex chemical schemes. AQ applications are now mentioned from
the introduction.

- P13, L13: I am not sure what the authors meant by spatial gradient as observations? Please clarified or remove.

Some inversion studies assimilate differences between concentrations at different locations in the observation
vector (Breon et al. 2015), instead of individual observations. This approach is meant to limit the influence of
background concentrations in the inversion of local fluxes.

p15 l7: authors implemented differences between observation sites and time in the observation vector instead of ob-
servations from individual sites in order to focus on spatial/temporal gradients, thus allowing to limit the influence of
background concentrations in the computation of local

- P13, L14-15: Maybe the authors could mention here that the main idea behind “super-obing” is to lower the discrepancies
of the representativeness between model and observations.

We thank the reviewer for helping to clarify our text. We have added statements about the representativeness of
the observations

p16 l4: In the case of super-observations (satellites retrievals, images, spatial gradients, etc.) in the observation vector,
it is often necessary to combine multiple simulated point observations in given grid cells and time stamps into a single
super-observation, to limit redundant observations, hence the size of the observation vector.

- P13, L15-16: “This is the case for...time and locations,” This sentence is very unclear. Please rephrase or explain better.

We agree with the reviewer. We have split and clarified our statement.

p16 l8: Super-observations are commonly used in the case for satellite observations being compared to all the model
levels above a given location; concentration gradients comparing observations at different time and locations (see e.g.,
Bréon et al., 2015; Stauffer et al., 2016) are another example of observation aggregation to reduce representativeness
errors. Isotopic ratios are also super-observations as they require to simulate separate isotopologues and recombine
them after the simulation (as done in e.g., van der Velde et al., 2018; Peters et al., 2018)

- P15, L21-31: What about satellite observations? I think the authors should add statement about this.

9
At the moment (in a version of the CIF more advanced than the one described in the present paper), satellite observations are included in the observation operator, but the user must format original files manually to a standard observation format readable by the CIF, with column, averaging kernels, etc. in a specific form. In the future, there will be a need to allow raw satellite formats automatically in the CIF to be ingested without manual formatting.

Satellite products, in particular the formatting of averaging kernels and other metadata, should also be included in the CIF in the near future as they play a growing role in the community.

Those studies use Gaussian plume technique but do not provide a continuous inversion constrain globally. Please correct the statement as it makes the reader believes that the Gaussian plume technique would allow a global inversion similar to a CTM based inversion.

We fully agree with the reviewer and have clarified the sentence.

To the larger scales with satellite measurements to optimize individual clusters of industrial or urban emissions.

The purpose of this set-up is to demonstrate that very large correlation lengths are equivalent to having only one uniform scaling factor. In the example set-up we had, the high number of observation sites counter-balanced the large correlation lengths. We changed our set-up, with a smaller domain and smaller number of observation sites to really demonstrate the equivalence of large correlation lengths and one single scaling factor for the whole domain. The corresponding figures are shown in Supplement now.

P19, L28: 10000m is 4 times the size of the domain here. No data assimilation system would use such long length scale. Is this on purpose? If yes, please justify.

The purpose of this set-up is to demonstrate that very large correlation lengths are equivalent to having only one uniform scaling factor. In the example set-up we had, the high number of observation sites counter-balanced the large correlation lengths. We changed our set-up, with a smaller domain and smaller number of observation sites to really demonstrate the equivalence of large correlation lengths and one single scaling factor for the whole domain. The corresponding figures are shown in Supplement now.

With all the fluxes and fluxes differences plots it would be better to use python colour maps that are diverging and centred around 0. This will improve the readability of the figures.

We have used centered color scales for the uncertainty matrix as plotted now. But as fluxes are not centered, we find that our color scales are the best compromise for readability.

P20, L20: I am unsure if I understand correctly but I believe the authors are speaking about iterations in the variational sense and members in the EnKF sense. If yes, please have few sentences to explain what the word ‘simulation’ means. Alternatively, I would replace the word ‘simulations’ by’iterations or members.’

The reviewer understood correctly. We clarified. We use this measure in the x-axis to be comparable between variational and ensemble methods, as one iteration of a variational algorithm requires two simulations (one forward and one adjoint). Details are given about the number of simulations in Sect. 4.2.2

More precisely, users and developers of adjoint transport models choose the number of forward re-computations to be carried out, based on a space-speed trade off: by saving all forward intermediate states, the adjoint is as costly as the forward, but the disk space burden can be extremely challenging to manage, thus increasing the overall computation time in return.
P20, L32: 10000m is 4 times the size of the domain here. See the related comment above. Such long length scales would provide very uniform structure on the increments. Structure similar to the 500m increments is seen on the result using the 10000m length scale. I expect there is typo through the text a 0 should be removed to 1000m?

This was indeed 10000m, but the high number of observation sites counter-balanced the long correlation lengths. We have changed the set-up with less observation sites to have very uniform increments when having very large correlation scales. This is a simple test to demonstrate the consistency of the definition of the correlations in the system. To make this test clearer, we now prescribe a correlation of 200 km, hence 100 times the size of the domain.

P20,L33-P21,L2: This look to me more like a bug rather noise as vertical lines or “chessboard” like pattern appear in the increments results. Such pattern also occurs in the analytical solution. This not looking like sampling noise to me. I think this points out possible problems in the implementation of the EnKF and of the systems other than the variational methods. It is not entirely clear if the system uses the ensemble information to derive jacobians for the calculation of H and perform a minimization? In this case this inversion cannot be called EnKF. This needs to be revised or diagnosed or explained.

Thanks for pointing this out. There was a small difference in the definition of the B matrix between analytical and ensemble methods on one side, and variational inversions in the other site. We have removed it and have correspondingly re-computed the inversions.

1.3 Technical comments

P2, L27-29: an order in all those references, chronological or alphabetical.

We have re-ordered those references chronologically

P7, L16: replace "... linear, simple cases." by "... linear and simple cases."

Corrected

P7, L19: change “characterized” to “characterize”.

Corrected

P9, L13: Chevallier et al., 2005: put this reference to the previous parenthesis for clarity.

Corrected

P10, L5: There is a colon here but nothing after. Should we expect equations?

An equation was missing in the submitted version of the manuscript. We detail how posterior matrices can be deduced in variational inversions: either with Monte Carlo experiments or with truncated eigen vectors and values of the Hessian matrix
The approximation of the posterior uncertainty matrix $A$ in the case of CONGRAD reads as follows:

$$A = \text{Hess}(J)^{-1} \approx V_x^T \Lambda_x^{-1} V_x$$

with $V_x$ the dominant eigenvectors of the Hessian matrix at the point $x^a$ and $\Lambda_x$ the matrix of the dominant eigenvalues of the Hessian matrix. Please note that the dominant eigenvalues of the Hessian matrix correspond to components with low posterior uncertainties in $A$.

Another approach to quantify the posterior uncertainty matrix $A$, valid for both linear and non-linear cases, is to carry out a Monte Carlo ensemble of independent inversions with sampled prior vectors from the prior distribution $N(x^b, B)$ (e.g., [?]). An ensemble of posterior vectors are inferred and used to compute the posterior matrix as follows:

$$A \approx \frac{1}{N-1} (x^a_1 - x^a_{\text{ref}}, x^a_2 - x^a_{\text{ref}}, \ldots, x^a_N - x^a_{\text{ref}}) . (x^a_1 - x^a_{\text{ref}}, x^a_2 - x^a_{\text{ref}}, \ldots, x^a_N - x^T)$$

with $N$ the size of the Monte Carlo ensemble, $x^a_i$ the posterior vector corresponding to the prior $x^b_i$ of the Monte Carlo ensemble and $x^a_{\text{ref}}$ the average over sampled posterior vectors.

– P15, L18: I understand what the authors means but this need to be reformulated. “Meteos” is not really correct English.

This referred to the keyword used in the CIF, which is abbreviated for the sake of readability; we replaced the word in the body of the manuscript by "meteo-data".

– P19, L3-10: Maybe add the equations in the text where they are mentioned.

We re-arranged this section accordingly.

– P22, L6: This is not a fair statement considering on how the “EnKF” has been implemented here. It doesn’t seem to be strictly an EnKF as the sequential assimilation is not performed. Consider changing or removing this in the conclusion. Also, this is not the scope of GMD paper here to compare methods but to describe and showcase model/software developments for geoscience.

We agree that a GMD "description paper" should not go too far in interpreting demonstration results. We have reformulated and reduced the comparison of methods. See track-change for modifications in the result section.

– P22, L10-12: The authors could add some sentences about the challenges in implementing actual CTMs in the CIF framework. For example, among many other challenges, I/O is a limiting factor for data assimilation especially with CTMs.

We agree that I/O can be a limiting factor. With its very general formalism, the CIF may not be able to accommodate very specific tweaks necessary for some models and some very large applications. However, using
very efficient numerical libraries in python (which uses C in the background), one should be able to keep up the pace with systems with compiled codes (Fortran mainly in the community). We have added a paragraph in the conclusion to discuss this issue:

The next step of the CIF is the implementation of a large variety of CTMs. The implementation of new CTMs already interfaced with other inversion systems should not bring particular conceptual challenges as all interface operations are already written in their original inversion system; in most cases, re-arranging existing routines is sufficient to interface a model to the CIF. One particular challenge concerns I/O optimizations: the generation of inputs and the processing of outputs can be time consuming and in some very heavy applications requires specific numerical and coding optimizations. The very general formalism of the CIF may hamper the ability of applying these particular optimizations for some models. Best efforts will have to be deployed to take full advantage of advanced I/O and data manipulation libraries in python to limit this weakness.

2 Referee #2: Peter Rayner

2.1 General Comments

This paper introduces a software system to facilitate ensembles of inverse calculations. The likely focus is atmospheric inversions in which model inputs such as surface fluxes are inferred from combinations of prior information and observations. Such a system is clearly useful since ensembles based on choices of ingredients (such as models or input covariances) provide a better basis for both the means and uncertainties of posterior estimates. The paper is certainly within scope for GMD since it describes a system which might well be widely used in the future. It is also well written though some final proofreading, especially around pluralisation, will be needed.

I note I am not reviewing the software itself. There is only one major critique of the paper. I think the benefits of the system are slightly over-sold (but what software release doesn’t do this). I believe, for example, that CIF doesn’t provide a framework for comparing estimates where the target spaces differ such as ENKF with short windows vs long-term mean approaches. That’s an intellectual/mathematical task which one could instantiate in software. This would probably involve classes representing PDFs which might be internally represented by ensembles or parameters and functions then defining transformations among these PDFs. That quibble aside there is no doubt the system will help such comparisons.

We agree that the CIF does not provide a plug’n’play solution to the problem of comparing totally different inversion set-ups in a statistically consistent way. Doing so would require further mathematical considerations that may prove necessary in the future to compare methods/models that cannot be computed on, e.g., exactly the same inversion window (i.e., when comparing inversion results for a given region, one computed with a regional model and the other with a global one). However, having a consistent framework with comparable configuration parameters and identical output format will definitely help to develop such a comparison.
On the positive side the ability to configure the computational pipeline is a marvellous innovation. Building our own, less ambitious system, we have struggled to make it agnostic about computation.

Perhaps more important than supporting intercomparisons, the tools provided as part of the system will lower the entry barrier for new groups entering the field. This was a surprising spin-off from the TRANSCOM effort 20 years ago. Most systems since then have been too closely tied to a given model for universal use so CIF can provide a major boost to the field. The fact that it is well documented will only help this.

There is one aspect, though, that seems seriously deficient. If I read correctly, methods for calculating uncertainty in the Gaussian posterior are not provided. One important role of a new framework like CIF is to educate, to encourage good practice. That includes calculating and engaging with the uncertainty in the results. Certainly any of the analytic methods under-estimate this term but it is still true that under-sampling of concentration is a major source of uncertainty for inversions at all scales. I request that the authors at least detail a road map for including posterior uncertainty.

We agree: this was missing. Of course the posterior uncertainty calculation depends on the approach used (analytical, CONGRAD, M1QN3, or EnKF), as well as how the error matrices are computed in the system. We have implemented the computation of the posterior error matrices for all methods and have re-arranged the manuscript to demonstrate the capability to provide uncertainty reduction and posterior error matrices. In particular, posterior matrices are shown in the Result section.

2.2 Specific Comments

– P7 bottom this correlation of errors is a common critique of ensemble methods but I’ve rarely seen it implemented in a non-ensemble method either and it’s perfectly possible to do in ensemble methods, one just adds some extra state variables carrying observational corrections and build the correct autocorrelations into them instead

We agree most variational inversions assume diagonal observation error matrices, or block diagonals, which is somehow equivalent to the running assimilation window of the Kalman Filter.

– P8 L22saying "the maximum probability" and "the mode" is tautological I think

We have re-arranged the variational section following comments from reviewer #1 and have removed this tautology.

– P9 congratulations for explicitly writing out the regularisation/reduction step which is so often assumed. It is called pre-conditioning so often that you should add the term for clarity

We now mention the term pre-conditioning for consistency with the terminology used in the community

– P9 it’s worth commenting that leading eigen-vectors of the Hessian correspond to the low uncertainty parts of the solution which might or might not be what you want

Thank you for pointing out this aspect. We now give details on the formula used to deduce the posterior matrix from the eigen vector of the Hessian and mention this.
The approximation of the posterior uncertainty matrix $A$ in the case of CONGRAD reads as follows:

$$A = Hess(J)_{x^a}^{-1} \approx V_{x^a}^T \Lambda_{x^a}^{-1} V_{x^a}$$

with $V_{x^a}$ the dominant eigenvectors of the Hessian matrix at the point $x^a$ and $\Lambda_{x^a}$ the matrix of the dominant eigenvalues of the Hessian matrix. Please note that the dominant eigenvalues of the Hessian matrix correspond to components with low posterior uncertainties in $A$.

Another approach to quantify the posterior uncertainty matrix $A$, valid for both linear and non-linear cases, is to carry out a Monte Carlo ensemble of independent inversions with sampled prior vectors from the prior distribution $N(x^b, B)$ (e.g., ?). An ensemble of posterior vectors are inferred and used to compute the posterior matrix as follows:

$$A \approx \frac{1}{N-1} (x_1^a - x_{ref}^a, x_2^a - x_{ref}^a, \ldots, x_N^a - x_{ref}^a) \cdot (x_1^a - x_{ref}^a, x_2^a - x_{ref}^a, \ldots, x_N^a - x^T)$$

with $N$ the size of the Monte Carlo ensemble, $x_i^a$ the posterior vector corresponding to the prior $x_i^b$ of the Monte Carlo ensemble and $x_{ref}^a$ the average over sampled posterior vectors.

– P11 citing the Bousserez and Henze paper is a good advertisement for your approach, you might want to say in the abstract or intro that many new methods are hybrid and this combined framework simplifies implementing such methods.

We insist on the advantage for modern hybrid methods of such a flexible structure. The decomposition of elementary operations makes it easier to recompose them in various ways. Besides, the CIF allows new methods to be tested with multiple models, hence inquiring their strengths and weaknesses in various contexts.

p25 l14: With modern inversion methods moving towards an hybrid paradigm of variational and ensemble methods, the flexibility of the CIF will be a valuable asset as abstract methods can be easily mixed with each other.

– P15 a technical question, how do you handle target variables which aren’t needed at various stages, e.g. bias corrections in observations which aren’t inputs to the CTM? the pipe lining structure doesn’t lend itself well to this problem.

Fig 2 and Eq. 11-12-13 may give the impression that all elementary transformations are computed in a serialized way. One should keep in mind that $x$ in Figure 2 includes all target variables, not only those regarding fluxes (as assumed in the figure). Therefore, if some variables are related to observation biases, they are kept identical along all transformations before and after the models, and they are only used at the last step when reconstructing the simulated observation vector equivalent. We have modified the figure to highlight this aspect. The text describing the figure has been extended:

p16 l13: The case of Fig. 2 also include background concentrations in the target vector: a background is often used to fix some biases in initial and lateral concentrations in limited-area models, and in observations (mostly satellites); the background variables are processed at the very end of the pipe when re-constructing the observations vector.
The mathematical formalism of Eq. (15) and (16) suggests that transformations are necessarily computed in a serialized way, thus limiting applications to simple target variables upstream the transport model. However, each elementary transformation handles components of the inputs it is concerned with, leaving the rest identical and forwarding it to later transformations. Typically, it does not actually limit applications to simple target variables upstream the CTM. For instance, in the case of target variables optimizing biases in the observations, the corresponding components of the target vector $\mathbf{x}$ are forwarded unchanged by all transformations in Fig. 2 until the very last operation, where they are used for the final comparison to the observation vector.

- P18 I think Eq. 17, do you consider sources as points or area integrals, if so do you need to integrate over the area of the source, especially for nearby observations?

The formulas used in Eq. 16 and 17 are based on point sources. The word "gridded" in the text is therefore misleading. We have clarified this statement by:

p21 l18: we use a grid of point surface fluxes to simulate concentrations using the Gaussian plume equation.

- P20 it’s worth adding that the amount of intermediate calculation needed is at user discretion; a space-speed trade-off. CMAQ, for example needs almost none, having check pointed what it needs on the forward run it had to do anyway at the start of the iteration.

We thank the reviewer for highlighting this aspect. We have extended the paragraph as follows to mention this aspect:

p23 l29: More precisely, users and developers of adjoint transport models choose the number of forward re-computations to be carried out, based on a space-speed trade off: by saving all forward intermediate states, the adjoint is as costly as the forward, but the disk space burden can be impossible to manage, thus increasing the overall computation time in return.
The Community Inversion Framework v1.0: a unified system for atmospheric inversion studies

Antoine Berchet\textsuperscript{1,}\textsuperscript{*}, Espen Sollum\textsuperscript{2}, Rona L. Thompson\textsuperscript{2}, Isabelle Pison\textsuperscript{1}, Joël Thanwerdas\textsuperscript{1}, Grégoire Broquet\textsuperscript{1}, Frédéric Chevallier\textsuperscript{1}, Tuula Aalto\textsuperscript{3}, Adrien Berchet\textsuperscript{14}, Peter Bergamaschi\textsuperscript{4}, Dominik Brunner\textsuperscript{5}, Richard Engelen\textsuperscript{6}, Audrey Fortems-Cheiney\textsuperscript{1}, Christoph Gerbig\textsuperscript{7}, Christine D. Groot Zwaaftink\textsuperscript{2}, Jean-Matthieu Haussaire\textsuperscript{5}, Stephan Henne\textsuperscript{5}, Sander Houlweling\textsuperscript{8}, Ute Kartens\textsuperscript{9}, Werner L. Kutsch\textsuperscript{10}, Ingrid T. Luijkx\textsuperscript{11}, Guillaume Monteil\textsuperscript{9}, Paul I. Palmer\textsuperscript{12}, Jacob C. A. van Peet\textsuperscript{8}, Wouter Peters\textsuperscript{11,13}, Philippe Peylin\textsuperscript{1}, Elise Potier\textsuperscript{1}, Christian Rödenbeck\textsuperscript{7}, Marielle Saunois\textsuperscript{1}, Marko Scholze\textsuperscript{9}, Aki Tsuruta\textsuperscript{3} and Yuanhong Zhao\textsuperscript{1}

\textsuperscript{1}Laboratoire des Sciences du Climat et de l’Environnement, CEA-CNRS-UVSQ, Gif-sur-Yvette, France
\textsuperscript{2}Norwegian Institute for Air Research (NILU), Kjeller, Norway
\textsuperscript{3}Finnish Meteorological Institute (FMI), Helsinki, Finland
\textsuperscript{4}European Commission Joint Research Centre, Ispra (Va), Italy
\textsuperscript{5}Swiss Federal Laboratories for Materials Science and Technology (Empa), Dübendorf, Switzerland
\textsuperscript{6}European Centre for Medium-Range Weather Forecasts, Reading, RG2 9AX, UK
\textsuperscript{7}Max Planck Institute for Biogeochemistry, Jena, Germany
\textsuperscript{8}Vrije Universiteit Amsterdam, Department of Earth Sciences, Earth and Climate Cluster, Amsterdam, the Netherlands
\textsuperscript{9}Dep. of Physical Geography and Ecosystem Science, Lund University, Sweden
\textsuperscript{10}Integrated Carbon Observation System (ICOS-ERIC), Helsinki, Finland
\textsuperscript{11}Meteorology and Air Quality Group, Wageningen University and Research, Wageningen, the Netherlands
\textsuperscript{12}School of GeoSciences, University of Edinburgh, Edinburgh, EH9 3FF, UK
\textsuperscript{13}Centre for Isotope Research, University of Groningen, Groningen, the Netherlands
\textsuperscript{14}Institut Pprime (UPR 3346, CNRS-Université de Poitiers-ENSMA), Bat H2, 11 Boulevard Marie et Pierre Curie, TSA 51124, 86073, Poitiers Cedex 9, France

\textbf{Correspondence:} antoine.berchet@lsce.ipsl.fr

\textbf{Abstract.} Atmospheric inversion approaches are expected to play a critical role in future observation-based monitoring systems for surface \textit{fluxes of} greenhouse gas (GHG)\textit{fluxes, pollutants and other trace gases}. In the past decade, the research community has developed various inversion softwares, mainly using variational or ensemble Bayesian optimization methods, with various assumptions on uncertainty structures and prior information and with various atmospheric chemistry-transport models. Each of them can assimilate some or all of the available observation streams for its domain area of interest: flask samples, in-situ measurements or satellite observations. Although referenced in peer-reviewed publications and usually accessible across the research community, most systems are not at the level of transparency, flexibility and accessibility needed to provide the scientific community and policy makers with a comprehensive and robust view of the uncertainties associated with the inverse estimation of GHG \textit{and reactive species} fluxes. Furthermore, their
development, usually carried out by individual research institutes, may in the future not keep pace with the increasing scientific needs and technical possibilities. We present here a **Community Inversion Framework** (CIF) to help rationalize development efforts and leverage the strengths of individual inversion systems into a comprehensive framework. The CIF is primarily a programming protocol to allow various inversion bricks to be exchanged among researchers. In practice, the ensemble of bricks makes a flexible, transparent and open-source python-based tool to estimate the fluxes of various GHGs both at and reactive species both at the global and regional scales. It will allow running different atmospheric transport models, different observation streams and different data assimilation approaches. This adaptability will allow a comprehensively comprehensive assessment of uncertainty in a fully consistent framework. We present here the main structure and functionalities of the system, and demonstrate how it operates in a simple academic case.

1 Introduction

The role of greenhouse gases (GHGs) in climate change has motivated an exceptional effort over the last couple of decades to densify the observations of GHGs around the world (Ciais et al., 2014): from the ground, e.g., with the European Integrated Carbon Observation System (ICOS, https://www.icos-cp.eu/), from mobile platforms (e.g., from aircrafts, or balloons equipped with Aircore sampling; Filges et al., 2016; Karion et al., 2010), and from space (e.g., Crisp et al., 2018; Janssens-Maenhout et al., 2020), despite occasional budgetary difficulties (Houweling et al., 2012). These observations quantify the effect of exchange between the surface and the atmosphere on GHG concentrations (e.g., Ramonet et al., 2020) and can thus be used to determine the surface fluxes of GHGs through the inversion of atmospheric chemistry and transport (e.g., Peylin et al., 2013; Houweling et al., 2017). Alongside improved observation capabilities, national and international initiatives pave the way towards an operational use of atmospheric inversions to support emissions reporting to the United Nations Framework Convention on Climate Change (UNFCCC; e.g., Say et al., 2016; Henne et al., 2016; Bergamaschi et al., 2018a; Janssens-Maenhout et al., 2020, or the EU projects CHE – CO₂ Human emissions; che-project.eu – or VERIFY – verify.lsce.ipsl.fr).

In the past, research groups have developed various atmospheric inversion systems based on different techniques and atmospheric transport models, targeting specific trace gases or types of observations, and at various spatial and temporal scales, according to the particular scientific objectives of the study. All these systems have their own strengths and weaknesses and help explore the range of systematic uncertainty in the surface to atmosphere fluxes. Inter-comparison exercises are regularly conducted to assess the strengths and weaknesses of various inversion systems (e.g., Locatelli et al., 2013; Monteil et al., 2019; Bergamaschi et al., 2018b; Gurney et al., 2003; Saunois et al., 2003; Peylin et al., 2013; Babenhauserheide et al., 2015; Brunner et al., 2017; Bergamaschi et al., 2018b). Inter-comparisons also provide an assessment of the systematic uncertainty on final flux estimates induced
by the variety of options and choices in different inversion systems. However, although the inversion systems are referenced in peer-reviewed literature, and are usually accessible to the research community, they are typically not at the level of transparency, documentation, flexibility and accessibility required to provide both the scientific community and policy makers with a comprehensive and robust view of the uncertainties associated with the inverse estimation of GHG trace gases (primarily GHGs and reactive species) fluxes. In particular, the differences between inversion systems (such as the atmospheric transport model, prior and observation space uncertainties, and inversion algorithm) make comparing their results particularly challenging, even when they are applied to the same problem. Moreover, research inversion systems are so far not ready for operational use, and their development, usually carried out by individual research institutes or limited consortia, may not keep pace with the scientific and technical needs to come, such as those linked to the increasing availability of high resolution satellite GHG and reactive species observations (Janssens-Maenhout et al., 2020). A unified system, as a community platform running multiple transport models, with diverse inversion methods, would provide new possibilities to effectively and comprehensively assess GHG and various reactive species budgets, trends, and their uncertainties and quantify limitations and development needs related to different approaches, all which is needed in order to properly support emission reporting. Collaborative efforts towards unified systems are already happening in other data assimilation communities, with, e.g., the Object-Oriented Prediction System (OOPS; coordinated by the European Centre for Medium-range Weather Forecast, UK), or the Joint Effort for Data Integration (led by UCAR/JCSDA: www.jcsda.org/jcsda-project-jedi), The Data Assimilation Research Testbed (DART; Anderson et al., 2009) dedicated to weather and marine forecasts. The Community Inversion Framework is also an example of collective effort proposing common data assimilation scripts for diverse applications (e.g., Earth system, or reactive species inversions; Gaubert et al., 2020). The Community Inversion Framework (CIF) is an initiative by members of the GHG atmospheric inversion community to bring together the different inversion systems used in the community, and is supported by the European Commission H2020 project VERIFY. The CIF will also support operational applications of atmospheric inversions in the CoCO2 project (coco2-project.eu), that will design an operational inversion system based on OOPS and interfaced with the research community through the CIF.

Despite their differences in methodology, application and implementation, almost all inversion systems rely on the same conceptual and practical bases: in particular, they use model-observation mismatches in a statistical optimization framework (most of the time based on Bayes’ Theorem), and numerical atmospheric tracer transport and chemistry models to simulate mixing ratios of GHGs and trace gases based on surface fluxes. The objectives of the CIF are to develop a consistent input/model interface, to pool development efforts, and to have an inversion tool that is well-documented, open-source, and ready for implementation in an operational framework. The CIF is designed to be a flexible and transparent tool to estimate the fluxes of different GHGs (e.g. carbon dioxide CO₂, methane CH₄, nitrous oxide N₂O, or halocarbons) and other
species, such as reactive species (e.g., CO, NO₂, HCHO), based on atmospheric measurements. In particular, although primarily designed for GHGs applications, the CIF is based on a general structure that will allow applications to air quality data assimilation. It is also designed to run at any spatial and temporal scales and with different atmospheric (chemistry-)transport models (global and regional, Eulerian and Lagrangian), with various observation data streams (ground-based, remote sensing, etc.), and a variety of data assimilation techniques (variational, analytical, ensemble methods, etc.). It will be possible to run it on multiple computing environments and corresponding set-ups and tutorials will be well documented. Community development will help in tackling the challenges in set-up and running, and accelerate adoption of the tool into wider use. One of the main foreseen advantages of the CIF is the capability to quantify and compare the errors due to the modeling of atmospheric transport and the errors due to the choice of a given inversion approach and set-up to solve a specific problem, in a fully consistent framework. The CIF will provide a common platform for quickly developing and testing new inversion techniques with several transport models, and it is hoped that with the combined community effort, it will be continuously improved and revised, keeping it state-of-the-art.

In the present paper, we lay out the basis of the CIF, giving details on its underlying principles and overall implementation. The proof-of-concept focuses on the implementation of several inversion methods, illustrated with a test case. We will dedicate a future paper to the evaluation of the system on a real-life problem with a number of interfaced atmospheric (chemistry-)transport models. At the time of writing the present article, the following models are interfaced with the CIF: the Global Circulation Models LMDZ (Chevallier et al., 2005) and TM5 (Krol et al., 2005; van der Laan-Luijkx et al., 2017), the regional chemistry-transport Eulerian model CHIMERE (Fortems-Cheiney et al., 2021) and the Lagrangian Particle Dispersion models FLEXPART (Pisso et al., 2019) and STILT (Trusilova et al., 2010). For the sake of simplicity, we refer to all types of (chemistry-)transport models generically as CTMs in the following. In Section 2, we describe the general theoretical framework for atmospheric inversions and how the CIF will include the theory in a flexible and general way. In Section 3, the practical implementation of the general design rules is explained, with details on the python implementation of the CIF. In Section 4, we demonstrate the capabilities of the CIF in a simple test case, applying various inversion techniques in parallel.

2 General principle

The version of the CIF presented here is implemented around Bayesian data assimilation methods with Gaussian assumptions, which constitute the main framework used in the atmospheric inversion systems for GHG fluxes and other trace gases (e.g., Enting, 2002; Bocquet et al., 2015). However, some studies have proposed possible extensions to more general probability density functions beyond the classical Gaussian case (e.g., truncated Gaussian densities, log-normal distributions, etc.; Michalak and Kitanidis, 2005; Berga-
Therefore, we propose here a general and flexible structure for our system that will be independent of limiting assumptions, as described in Sect 2.3, to allow future extensions to more general theoretical frameworks. In the following, mathematical formulas are written following notations based on Ide et al. (1997) and Rayner et al. (2019). We present the theoretical basis and several inversion methods that are implemented in the CIF as demonstrators.

2.1 General Bayesian data assimilation framework

The Bayesian approach consists in estimating the following conditional probability density function (pdf):

$$ p^a(x) = p(x \mid y^o - \mathcal{H}(x^b), x^b) \propto p(y^o - \mathcal{H}(x) \mid x) p^b(x) \tag{1} $$

with $x$ the target vector, $p^a(x)$ the posterior distribution of the target vector, $p^b(x)$ the prior knowledge of the target vector, characterized by its mode $x^b$, $y^o$ the observation vector gathering all observations implemented in the inversion and $\mathcal{H}$ the observation operator linking the target vector to the observation vector. In the following, we also refer to $\mathcal{X}$ and $\mathcal{Y}$ as the target and observation spaces, respectively, from where the target and observation vectors are sampled. Classically, for atmospheric inversions, the observation vector $y^o$ includes ground-based measurements of GHG trace gases mixing ratios on fixed or mobile platforms, and remote sensing observations such as satellite observations. The target vector $x$ includes the variables to be optimized by the inversion; it includes the main variables of interest, such as the surface fluxes, but also variables related to atmospheric chemical sources and sinks, background concentrations in the case of limited-area transport models, model parameters, etc., which are required to make the inversion physically consistent. The observation operator $\mathcal{H}$ mainly includes the computation of atmospheric transport and chemistry (if relevant) by numerical Eulerian-globale circulations (chemistry-) transport models. These can be of various types: e.g., global transport models (e.g., LMDZ, Chevallier et al. 2010; TM5, Houweling et al. 2014; GEOS-Chem, van der Laan-Luijkx et al. 2017; Liu et al. 2015; Palmer et al. 2019; Feng et al. 2017; NICAM, Niwa et al. 2017), regional Eulerian chemistry-transport models (e.g., CHIMERE, Broquet et al. 2011; Fortems-Cheiney et al. 2021; WRF-CHEM, Zheng et al. 2018; COSMO-GHG, Mizzi et al. 2016; Kuhlmann et al. 2019; LOTOS-EUROS, Curier et al. 2012) or Lagrangian Particle Dispersion models (e.g., FLEXPART, Thompson and Stohl 2014; STILT, Bagley et al. 2017; Brioude et al. 2013; Trusilova et al. 2010). It also includes pre- and post-processing operations required to project the target vector to a format compatible with the model input and the model outputs to the observation vector; these operations can be the applications of e.g., averaging kernels in the case of satellite operations, or interpolation of the target vector to higher resolution model inputs.

As errors in inversion systems come from a large variety of independent causes superimposing on each other, it is often assumed that the most relevant way of representing the distributions in Eq. (1) is to assume...
prior and observation spaces to be normal distributions, noted $\mathcal{N}(\cdot, \cdot)$ below, the first argument representing the average of the distribution and the second argument the covariance matrix. In addition, when assuming that the distributions in the state vector space and the observation space are independent from each other, and that errors in the observation and the state vector spaces have Gaussian, unbiased distributions, it is possible mathematically derive the Bayes theorem and to represent the distributions of Eq. \((1)\) as follows:

\[
\begin{align*}
    p^b(x) & \sim \mathcal{N}(x^b, B) \\
    p(y^o - \mathcal{H}(x^b)) & \sim \mathcal{N}(0, R) \\
    p^a(x) & \sim \mathcal{N}(x^a, A)
\end{align*}
\]

with $B$ and $A$ the prior and posterior covariance matrix of uncertainties in the target vector, $x^b$ and $x^a$ the prior and posterior target vectors and $R$ the covariance matrix of uncertainties in the observation vector and the observation operator.

The assumption that errors are unbiased, which makes it possible to write normal distributions in Eq. \((1)\) with means $x^b$, $0$ and $x^a$, respectively, is needed to simplify the formulation of the Bayesian problem in atmospheric inversions. Neglecting error biases have known impacts on inversion results (e.g., Masarie et al., 2011); they can be accounted for online as an unknown to be solved by the inversion (e.g., Zammit-Mangion et al., 2021), but are often treated offline from the inversion, either through pre-processing of inputs or post-processing of outputs.

### 2.2 Computation modes in the CIF

The present version of the CIF includes three main categories of inversion methods: 1) analytical, i.e. algebraic solution of the unbiased Gaussian Bayesian problem, 2) ensemble methods with the Ensemble Kalman Filter-Square Root Filter (EnSRF), and 3) variational with two examples of minimizing algorithms (MICE and CONGRAD). Other types of data assimilation methods (e.g. direct sampling of probability density functions through Monte Carlo approaches) are also used by the community. The choice of implementing the three aforementioned methods first is motivated by their dominant use, and because these three use very different approaches for solving the Bayesian inversion problem: that is, with/without random sampling of probability distributions, and with/without the use of the adjoint of the observation operator. The adjoint of the observation operator, noted $\mathcal{H}^*$, is built following the mathematical definition of the adjoint; heuristically, it operates backwards compared to the observation operator (e.g., Errico, 1997) in the sense that it determines the sensitivity to inputs (e.g. fluxes) given an incremental perturbation to outputs (e.g. concentrations) (e.g., Errico, 1997). In addition to the mentioned data assimilation methods, the CIF also includes the possibility to run forward simulations and to test the adjoint and the tangent linear of the observation operator for given inversion configurations. In the following we call all inversion methods and other types of computation in the CIF "computation modes". With these computation modes implemented in a flexible and
computing in the so-called response function method, it is anticipated that other inversion methods could be easily added to the CIF in the future (see Sect. 2.3).

2.2.1 Data assimilation methods

Analytical inversions

Analytical inversions compute the algebraic solution of the Gaussian Bayesian problem when it is linear and are used extensively at all scales (e.g., Stohl et al., 2009; Turner and Jacob, 2015; Kopacz et al., 2009; Bousquet et al., 2011; Wang et al., 2018; Palmer et al., 2006). When the observation operator is linear, $\mathcal{H}$ equals its Jacobian matrix $\mathbf{H}$, and conversely its adjoint $\mathcal{H}^*$ is the transpose of the Jacobian $\mathbf{H}^T$. In that case, $\mathbf{x}^*$ and $\mathbf{A}$ can be explicitly written as matrix products. There are two equivalent formulations of the matrix products for the solution of the problem (e.g., Tarantola and Valette, 1982):

\[
\begin{align*}
\mathbf{x}^* &= \mathbf{x}^b + \mathbf{K} (\mathbf{y}^o - \mathbf{H} \mathbf{x}^b) \quad \text{or} \quad \mathbf{x}^* = \mathbf{x}^b + (\mathbf{H}^T \mathbf{R}^{-1} \mathbf{H} + \mathbf{B}^{-1})^{-1} \mathbf{H}^T \mathbf{R}^{-1} (\mathbf{y}^o - \mathbf{H} \mathbf{x}^b) \\
\mathbf{A} &= \mathbf{B} - \mathbf{K} \mathbf{H} \mathbf{B}
\end{align*}
\]

with $\mathbf{K}$ the Kalman gain matrix: $\mathbf{K} = \mathbf{B} \mathbf{H}^T (\mathbf{R} + \mathbf{H} \mathbf{B} \mathbf{H}^T)^{-1}$

Analytical inversions can also be used on slightly non-linear problems, by linearizing the observation operator around a given reference point using the tangent linear of the observation operator. It formulates as follows:

\[
\mathcal{H}(\mathbf{x}^b + \delta \mathbf{x}) \approx \mathcal{H}(\mathbf{x}^b) + d\mathcal{H}_{\mathbf{x}^b}(\delta \mathbf{x}) = \mathcal{H}(\mathbf{x}^b) + \mathbf{H}_{\mathbf{x}^b} \delta \mathbf{x}
\]

with $\delta \mathbf{x}$ a small deviation from $\mathbf{x}^b$ within a domain where the linear assumption is valid, $d\mathcal{H}_{\mathbf{x}^b}$ the tangent-linear of $\mathcal{H}$ at $\mathbf{x}^b$ and $\mathbf{H}_{\mathbf{x}^b}$ the Jacobian matrix of $\mathcal{H}$ at $\mathbf{x}^b$.

Then Eq. (3) can be easily adapted by replacing $(\mathbf{y}^o - \mathbf{H} \mathbf{x}^b)$ by $(\mathbf{y}^o - \mathcal{H}(\mathbf{x}^b))$ and $\mathbf{H}$ by $\mathbf{H}_{\mathbf{x}^b}$.

The computation of an analytical inversion faces two main computational limitations. First, the matrix $\mathbf{H}$ representing the observation operator $\mathcal{H}$ must be built explicitly. This can be done either column by column, in the so-called response function method, by computing $\{\mathcal{H}(\mathbf{x}_i) \ | \ \forall \mathbf{x}_i \in B_X\}$ or row by row, in the so-called footprint method. The two approaches require $\dim(Y)$, the dimension of the target space and $\dim(Y)$, the dimension of the observation space, independent simulations respectively. In the response function method, each column is built by computing $\{d\mathcal{H}_{\mathbf{x}^b}(\delta \mathbf{x}_i) \ | \ \forall \delta \mathbf{x}_i \in B_X\}$ with $B_X$ the canonical basis of the target space, or row by row, in the so-called footprint method. For a given increment $\delta \mathbf{x}_i$, the corresponding column gives the sensitivity of observations to changes in the corresponding component of the target space. In the footprint method, by computing $\{\mathcal{H}^*\delta \mathbf{y}_i \ | \ \forall \delta \mathbf{y}_i \in B_Y\}$ each row is built by computing $\{\mathcal{H}^*_{\mathbf{x}^b}(\delta \mathbf{y}_i) \ | \ \forall \delta \mathbf{y}_i \in B_Y\}$ with $B_Y$ the
canonical basis of the observation space. For a given perturbation of $\delta y_i$ of a component of the observation vector, the corresponding row of $H$ gives the sensitivity of the inputs to that perturbation.

Depending on the number of available observations or the size of the target vector, one of the two is preferred to limit the number of observation operator computations to be carried out explicitly. When the dimension of the target vector is relatively small, the response function is generally preferred, and conversely, when the observation vector is small, the footprint approach is preferred. The type of transport model used to compute the matrix $H$ also plays a role in the choice of the approach: for Eulerian models, the response function approach is preferred for multiple reasons: (i) their adjoint is often much more costly than their forward, (ii) the adjoint may not be available for some models or is difficult to generate, and (iii) the computation time of the forward is constant no matter how numerous the observations; for Lagrangian models, the footprint approach is preferred as they often compute backward transport simulations for each observation, allowing a straightforward computation of the adjoint (Seibert and Frank, 2004). In both cases, the explicit construction of the matrix $H$ requires numerous independent simulations, which can be an insurmountable computational challenge.

The second obstacle consists in that the computation of the Kalman gain matrix in Eq. 3 (left) requires inverting of a matrix of the dimension of the observation vector space, $\text{dim}(\mathcal{Y})$, while for the other formulation (Eq. 3 right) a matrix is of dimension $\text{dim}(\mathcal{X})$, the dimension of the control target space. If the dimensions of both the observation and the control-target spaces are very high, as in many inversion applications, the explicit computation of Eq. 3 with matrix products and inverses is not computationally feasible. For this reason, smart adaptations of the inversion framework (including approximations and numerical solvers) are often necessary to tackle even linear problems even when they are linear; in the following, we choose to elaborate on some of the most frequent approaches used in the atmospheric inversion community: the variational approach and the Ensemble Kalman filter are described below one ensemble method, the Ensemble Square Root Filter (EnSRF). Less frequently, intermediate adaptations of the analytical inversion also include sequential applications (e.g., Michalak, 2008; Bruhwiler et al., 2005; Brunner et al., 2012), that are a compromise between tackling the above-mentioned computational obstacles while maintaining the simplicity of the analytical inversion; however, such sequential analytical inversions are limited to specific linear and simple cases.

**Ensemble methods**

Ensemble methods are commonly used to tackle high-dimensional problems with limited non-linearity and to approximately characterize the optimal solution. In ensemble methods, such as Ensemble Kalman filters (EnKFs) or smoothers (e.g., Peters et al., 2005; Zupanski et al., 2007; Zupanski, 2005; Feng et al., 2009; Chatterjee et al., 2012) (e.g., Whitaker and Hamill, 2002; Peters et al., 2005; Zupanski et al., 2007; Zupanski, 2005; Feng et al., 2009; Chatterjee et al., 2012)
The issue of high dimensions in the system of Equations 3, Eq. (3) is avoided using the two following main procedures:

- observations are first assimilated sequentially in the system to reduce the dimension of the observation space, making it possible to explicitly compute matrix products and inverses, and thus propagating information from the target space to the observation space; the overall inversion time window period is processed incrementally using a smaller running computation assimilation window including a manageable number of observations; intermediate inversions are solved on the smaller running window that is gradually moved from the beginning to the end of the overall inversion data assimilation window; the running assimilation window with so-called analysis and forecast steps introduces complex technical challenges to rigorously propagate errors from one iteration of the running window to the next one; moreover, the sequential assimilation of observations is valid only under the assumption that observations errors are not correlated between assimilation windows, which may prove incorrect for high-density data sets, but is an assumption also done in, e.g., variational inversions. For very dense observations, such as datasets from new-generation high-resolution satellites, the sequential assimilation of observations may not be sufficient; moreover, the sequential assimilation of observations is valid only under the assumption that observations for each assimilation window are not correlated with each others, which may prove incorrect for high-density data sets, or at least methods may be needed to make the observation errors between sequential assimilation windows independent, for example by applying a whitening transformation to the observations to form a new set with uncorrelated errors as suggested by Tippett et al. (2003). The challenge is exacerbated for long-lived species such as CO2, for which assimilation windows must be long enough to maintain the propagation of information on the fluxes on long distances through transport; propagating a covariance matrix from assimilation windows to assimilation windows as accurate as possible could in principle limit the later issue, as suggested in Kang et al. (2011, 2012), but could still prove hard to apply in very high resolution problems.

- matrices; the posterior distribution at a given step of the filter is then characterized explicitly by applying Eq. (1) on each member of the ensemble; the new intermediate posterior distribution is then sampled and propagated to the next data assimilation window.

In the atmospheric inversion community, another ensemble method is widely used, based on the CarbonTracker system (Peters et al., 2005), which uses an Ensemble Square Root Filter (EnSRF; Whitaker and Hamill, 2002). In that approach, the observations are split using running data assimilation windows as for other ensemble methods, but instead of directly characterizing the posterior distribution from the ensemble, the statistics of the ensemble is used to solve the analytical equation, Eq. (3), approximately. Thus, the EnSRF method is less general as EnKFs methods, as it relies on the Gaussian assumption, as well as limited non-linearity
in the inversion problem, but proves very efficient at computing an approximated solution of the inversion problem. Matrix products in Eq. 1 (3) involving the target vector covariance matrix $B$ ($BH^T$ and $BH^T$) are approximated by reducing the space of uncertainties to a low-rank representation; this is done in practice by using a Monte Carlo ensemble of possible control target vectors sampling the distribution $\mathcal{N}(x^b, B)$; with such an approximation, matrix products can be written as follows:

\[
\begin{align*}
BH^T & \approx \frac{1}{N-1} \left( \mathcal{H}(x_1), \mathcal{H}(x_2), \ldots, \mathcal{H}(x_N) \right)^T \left( \mathcal{H}(x_1), \mathcal{H}(x_2), \ldots, \mathcal{H}(x_N) \right) \\
BH^T & \approx \frac{1}{N-1} (x_1, x_2, \ldots, x_N)^T \left( \mathcal{H}(x_1), \mathcal{H}(x_2), \ldots, \mathcal{H}(x_N) \right)^T
\end{align*}
\] (5)

where $N$ is the size of the ensemble.

From there, Eq. 1 (1) is solved analytically by replacing $BH^T$ and $BH^T$ by their respective approximations. Using random sampling, ensemble methods are able to approximate large dimensional matrices at a reduced cost without using the adjoint of the observation operator (see variational inversion below) that can be challenging to implement. However, using too small ensembles causes degenerate approximations (under estimating the uncertainty magnitude, or misrepresenting uncertainty structures). Small ensembles generally cause the posterior ensemble to collapse, i.e., the posterior distribution is dominated by one or a very small number of members, which does not allow for a reliable assessment of the posterior uncertainties (Morzfeld et al., 2017); moreover, small ensembles introduce spuriousness in the posterior uncertainties, with unrealistic correlations being artificially generated. In the EnSRF, small ensembles rather cause a misrepresentation of uncertainty structures, which limits the accuracy of the computed solution, and may require fixes as described in, e.g., Bocquet (2011). In any case, the level of approximation necessary for this approach to work is strongly different for each problem, which requires preliminary studies before consistent application. In particular, the so-called localization of the ensemble can be used to improve the consistency of the inversion outputs (e.g., Zupanski et al., 2007; Babenhauserheide et al., 2015) in the current version, only the EnSRF approach is implemented in the CIF. One should note that the EnSRF, as a direct approximation of the analytical solution, can be very sensitive to non-linearity in the observation operator (e.g., Tolk et al., 2011). It can generally cope only with slight non-linearity over the assimilation window, thus, the assimilation window length has to be chosen appropriately, contrary to other ensemble methods which are usually not sensitive to non-linearity.

**Variational inversions**

Variational inversions are a numerical approximation to the solution. Variational inversions use the fact that finding the mode of the inversion problem: they involve the gradient of the cost function in Eq. 5 and require to run forward and adjoint simulations iteratively (e.g., Meirink et al., 2008; Bergamaschi et al., 2010; Houweling et al., 2016, 2014; ?; Chevallier et al. ...). In variational inversions, the solution $x$ is defined as being that with maximum posterior probability. In
the case of Gaussian assumptions, it is equivalent to computing the mode $\mathbf{x}^a$ of the normal distribution. Computing $\mathbf{x}^a$—posterior Gaussian distribution $p^a(\mathbf{x}) \sim \mathcal{N}(\mathbf{x}^a, \mathbf{A})$ in Eq. 1 (2) is equivalent to finding the minimum $\mathbf{x}^a$ of the cost function $J$:

$$J(\mathbf{x}) = \frac{1}{2}(\mathbf{x} - \mathbf{x}^b)^T \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}^b) + \frac{1}{2}(\mathbf{y}^o - \mathbf{H}(\mathbf{x}))^T \mathbf{R}^{-1}(\mathbf{y}^o - \mathbf{H}(\mathbf{x}))$$ (6)

The variational formulation does not require calculation of complex matrix products and inversions, contrary to the analytical inversion, and is thus not limited by vector dimensions. Still, the inverses of the uncertainty matrices $\mathbf{B}$ and $\mathbf{R}$ need to be computed, potentially prohibiting the use of very large and/or complex general matrices; this challenge is often overcome by reducing $\mathbf{B}$ and $\mathbf{R}$ to manageable combinations of simple matrices (e.g., Kronecker products of simple shape covariance matrices; see Sect. 2.3.1).

In variational inversions, the minimum of the cost function in Eq. 6 (6) is numerically estimated iteratively using quasi-Newtonian algorithms based on the gradient of the cost function:

$$\nabla J_\mathbf{x} = \mathbf{B}^{-1}.(\mathbf{x} - \mathbf{x}^b) + \mathbf{H}^*(\mathbf{R}^{-1}.(\mathbf{y}^o - \mathbf{H}(\mathbf{x})))$$ (7)

Quasi-Newtonian methods are a group of algorithms designed to compute the minimum of a function iteratively. It should be noted that in high-dimension problems, it can take a very large number of iterations to reach the minimum of the cost function $J$, forcing the user to stop the algorithm before convergence, thus reaching only an approximation of $\mathbf{x}^a$; in addition, iterative algorithms can reach local minima without ever reaching the global minimum, making it essential to thoroughly verify variational inversion results; this can happen in non-linear non-linear cases, but also, due to numerical artefacts, in linear cases (some points in the cost function have gradients so close to zero that the algorithm sees them as convergence points, whereas the only unique global minimum is somewhere else). In the community, examples of quasi-Newtonian algorithms commonly used are M1QN3 (Gilbert and Lemaréchal, 1989), the conjugate gradient CONGRAD (applicable only to linear or linearized problems; Fisher, 1998) algorithm (e.g., Chevallier et al., 2005) or the Broyden–Fletcher–Goldfarb–Shanno (BFGS) algorithm (Zheng et al., 2018; Bousserez et al., 2015), M1QN3 (Gilbert and Lemaréchal, 1989), and the CONGRAD algorithm (applicable only to linear or linearized problems; Fisher, 1998; Chevallier et al., 2005) based on the Lanczos method, which iterates to find the eigenvalues and eigenvectors of the Hessian matrix, which is then used (in a single step) to calculate the analysis vector, $\mathbf{x}^a$. In general, quasi-Newtonian methods require an initial regularization, or "pre-conditioning" of $\mathbf{x}$, the vector to be optimized, for better efficiency. In atmospheric inversions, such a regularization is generally made by optimizing $\chi = \mathbf{B}^{-1/2}.(\mathbf{x} - \mathbf{x}^b)$ in-
stead of $x$; we note $\mathfrak{A}$ the regularization space: $\chi \in \mathfrak{A}$. This transformation translates in Eq. 6 and 7 as follows:

$$\nabla J_{\chi} = \chi + B^{1/2} \mathcal{H}^* R^{-1} (y^o - \mathcal{H}(x))$$

$$\nabla J_{\chi} = \chi + B^{1/2} \mathcal{H}^* (R^{-1} (y^o - \mathcal{H}(B^{1/2} \chi + x^b)))$$

(8)

Solving Eq. 6 and 7 in the target vector space or Eq. 8 in the regularization space is mathematically fully equivalent, but the solution in the regularization space is often reached in fewer iterations. Moreover, in the regularization space, one can force the algorithm to focus on the main modes of the target vector space by filtering the smallest eigenvalues of the matrix $B$. This reduces the dimension of $\chi$ and accelerates further the rate of convergence, although the solution of the reduced problem is only an approximation of the solution of the full problem. In the following we thus prefer calling the "regularization space" the "reduction space". The link between the two can be written as follows:

$$\chi_{\text{full}} = QA^{-1/2} (x - x^b)$$

$$\chi_{\text{reduced}} = Q' \Lambda'^{-1/2} (x - x^b)$$

(9)

with $B^{1/2} = QA^{1/2} Q^{-1} B^{1/2} = QA^{1/2} Q^T$, $Q$ and $\Lambda$ being the matrices of the eigenvector and the matrix of the corresponding eigenvalues of the matrix $B$. $Q'$ and $\Lambda'$ are the reduced matrices of eigenvalues and eigenvectors with a given number of dominant eigenvalues.

Overall, variational inversions are a numerical approximation to the solution of the inversion problem; they involve the gradient of the cost function in Eq. 7 and require to run forward and adjoint simulations iteratively (e.g., Meirink et al., 2008; Bergamaschi et al., 2010; Houweling et al., 2016, 2014; Fortems-Cheiney et al., 2021; Chevallier et al., 2010).

The variational formulation does not require calculation of complex matrix products and inversions, contrary to the analytical inversion, and is thus not limited by vector dimensions. Still, the inverses of the uncertainty matrices $B$ and $R$ need to be computed, potentially prohibiting the use of very large and/or complex general matrices; this challenge is often overcome by reducing $B$ and $R$ to manageable combinations of simple matrices (e.g., Kronecker products of simple shape covariance matrices; see Sect. 2.3.1).

When the observation operator is linear, the posterior uncertainty matrix $A$ is equal to the inverse of the Hessian matrix at the minimum of the cost function. In most cases the Hessian cannot be computed explicitly, because of memory limitations, which is a major drawback of variational inversions. But some variational algorithms such as CONGRAD provide a coarse approximation of the Hessian: in the case of CONGRAD based on the Lanczos method, leading eigenvectors of the Hessian can be computed, together
with their eigenvalues (Fisher, 1998). The approximation of the posterior uncertainty matrix $A$ in the case of CONGRAD reads as follows:

$$A = Hess(J)_{x^a}^{-1} \approx V_{x^a}^T \Lambda_{x^a}^{-1} V_{x^a}$$

(10)

with $V_{x^a}$ the dominant eigenvectors of the Hessian matrix at the point $x^a$ and $\Lambda_{x^a}$ the matrix of the dominant eigenvalues of the Hessian matrix. Please note that the dominant eigenvalues of the Hessian matrix correspond to components with low posterior uncertainties in $A$.

Another approach to quantify the posterior uncertainty matrix $A$, valid for both linear and non-linear cases, is to carry out a Monte Carlo ensemble of independent inversions with sampled prior vectors from the prior distribution $\mathcal{N}(x^b, B)$ (e.g., Liu et al., 2017). An ensemble of posterior vectors are inferred and used to compute the posterior matrix as follows:

$$A \approx \frac{1}{N-1} (x^a_1 - x^a_{ref}, x^a_2 - x^a_{ref}, \ldots x^a_N - x^a_{ref}) . (x^a_1 - x^a_{ref}, x^a_2 - x^a_{ref}, \ldots x^a_N - x^T)$$

(11)

with $N$ the size of the Monte Carlo ensemble, $x^a$ the posterior vector corresponding to the prior $x^b$ of the Monte Carlo ensemble and $x^a_{ref}$ the average over sampled posterior vectors.

### 2.2.2 Auxiliary computation modes

#### Forward simulations

Forward simulations simply use the observation operator to compute simulated observation equivalents. It reads as:

$$(x^b, y^a) \rightarrow \mathcal{H}(x^b)$$

(12)

This mode is used to make quick comparisons between observations and simulations to check for inconsistencies before running a full inversion. It is also used by the analytical inversion mode to build response functions.

#### Test of the adjoint

The test of the adjoint is a crucial diagnostic for any inversion system making use of the adjoint of the observation operator. Such a test is typically required after making any edits to the code (to the forward observation operator or its adjoint) before running an inversion. Coding an adjoint is prone to errors and even small errors can have significant impacts on the computation of the gradient of the cost function in Eq. 3(7).
Thus, one needs to make sure that the adjoint rigorously corresponds to the forward. This test consists in checking the definition of the mathematical adjoint of the observation operator. It writes as follows for a given target vector \( x \) and incremental target perturbation \( \delta x \):

\[
< DdH_x(x, \delta x) | DdH_x(x, \delta x) > = < \delta x | (H^* \circ DdH_x)(x, \delta x) > \tag{13}
\]

\( D\mathcal{H}(x, \delta x) \cdot d\mathcal{H}_x(\delta x) \) is the linearization of the observation operator \( \mathcal{H} \) at the point \( x \) for a given increment \( \delta x \); it is computed with the tangent linear model, which is the numerical adaptation of \( D\mathcal{H}(x, \delta x) \cdot d\mathcal{H}_x(\delta x) \).

Then, \( (H^* \circ D\mathcal{H})(x, \delta x) \cdot (H^* \circ d\mathcal{H}_x)(\delta x) \) is calculated with the adjoint of the tangent-linear of \( \mathcal{H} \) at the point \( x \).

In practice, the two terms of the equation are rarely exactly equal. Nevertheless, the difference should never exceed a few times the machine epsilon. Besides, Eq. (13) should be verified for any given target vector and increment. In practice, it is not possible to explicitly verify all possible combinations; but as the result of the test is highly sensitive to any error in the code, it is assumed that a few typical couples \( (x, \delta x) \) are sufficient to certify the validity of the adjoint.

### 2.3 Identification of common elementary transformations

#### 2.3.1 General purpose operations

*Every inversion algorithm and computation mode mentioned above can be decomposed into a pipeline of elementary transformations. These transformations are listed in Tab. 1 and include: the observation operator and its adjoint (their matrix representations in linear cases), matrix products with control-target and observation error covariance matrices and corresponding adjoints, and random sampling of Normal distributions. To limit redundancy in the CIF as much as possible, these elementary transformations are included in the CIF as generic transformation blocks on the same conceptual level. Overall, the decomposition of computation modes presently implemented in the CIF into elementary transformations leads to the structure in Fig. 1.*

Avoiding redundancy makes the maintenance of the code much easier, and provides a clear framework for extensions to other inversion methods or features. For instance, inverse methods based on probability density functions other than Normal distributions could be easily implemented by updating the random ensemble generator, or by implementing new cost functions representing non-Gaussian distributions, while keeping the remaining code unmodified. In particular, non-Gaussian cost functions still rely on the computation of the observation operator. New combinations of elementary transformations can also directly lead to new methods. For instance, ensemble variational inversion (e.g., Bousserez and Henze, 2018) is a direct combination of the available variational pipeline and the random sampling pipeline. Inver-
sions estimating hyper-parameters through maximum-likelihood or hierarchical Bayesian techniques (e.g., Michalak et al., 2005; Berchet et al., 2014; Ganesan et al., 2014) could be integrated in the CIF by adapting the Gaussian cost function and by implementing a corresponding computation pipeline.

The complexity of the selected elementary transformations spans a wide range, from one-line straightforward codes to computationally expensive and complex code implementation. In small dimensional and/or linear problems, the computation of the observation operator using its Jacobian and matrix products may be computationally expensive, but is in principle rather straightforward to implement. For non-linear and/or high-dimensional problems, these transformations require simplifications and numerous intermediate steps. For instance, applying matrix products to the error covariance matrix \( R \) and \( B \) and computing their inverse is easy in small dimensions, but can be limiting in high dimensional problems. For that reason, the error covariance matrices are often reduced to particular decompositions; for instance, the error covariance matrix on the target vector \( B \) is often written as a Kronecker product of multiple spatial and/or temporal covariance matrices of lower dimensions, making matrix products manageable (e.g., Chevallier et al., 2005; Meirink et al., 2008; Yadav and Michalak, 2013).

In any case, the observation operator (see details in Sect. 2.3.2) appears as the center piece of any inversion method.

### 2.3.2 Observation operator

The observation operator is a key component of all inversion methods. It links the target space to the observation space, and conversely, its adjoint links the observation space to the target space. To do so, the observation operator projects its inputs through various intermediate spaces to the outputs. As atmospheric inversions need a representation of the atmospheric transport (and chemistry if relevant) to link the target vector (including surface fluxes, atmospheric sources and sinks, initial and boundary conditions for limited domains and time-windows, etc.) to the observation vector (including some form of atmospheric concentration measurements), the observation operator is built around a given (chemistry-)transport model CTM in most cases: Eq. (14) illustrates the various projections in the common case.

\[
x \xrightarrow{\Pi_x^\beta} \ f \xrightarrow{\Pi_{xF}^\beta} \text{inputs} \xrightarrow{\text{model}} \text{outputs} \xrightarrow{\Pi_{C}^{\Omega}} \ c \xrightarrow{\Pi_{\Omega}^{\gamma}} \mathcal{H}(x)
\]  

(14)

with \( f \) the target vector projected at the CTM’s resolution (includes fluxes, but also other types of inputs required by the CTM), \( c \) the raw outputs extracted from the run of the CTM’s executable (in general 4-dimensional concentration fields). \( \Pi \) operators are intermediate projectors: \( \Pi_x^\beta \) projects the target vector at the spatial and temporal resolutions of the CTM’s inputs, \( \Pi_{xF}^\beta \) dumps the input vector in files usable by the CTM’s executable, \( \Pi_{C}^{\Omega} \) reads the CTM’s outputs, \( \Pi_{\Omega}^{\gamma} \) reprojects the raw outputs at the observation vector.
resolution (mostly the temporal resolution as the model and the observation worlds do not follow the same time line).

The targeted structure of the CIF should allow a full flexibility of observation operators, from the straight-forward widely-used decomposition detailed in Eq. 14 (14) to more elaborated approaches including multiple transport models and/or complex super-observations (e.g., concentration gradients or aggregates: Bréon et al., 2015) (e.g., in Bréon et al., 2015; Staufer et al., 2016, authors implemented differences between observation sites and time in the observation vector instead) and hyper-parameters (e.g., emission factors and model parameters used to produce emission maps; Rayner et al., 2010; Asefi-Najafabady et al., 2014) (e.g., emission factors and model parameters used to produce emission maps; Rayner et al., 2010; Asefi-Najafabady et al., 2014) . Therefore, the observation operator is designed as a pipeline of elementary interchangeable transformations with standardized input and output formats such that:

\[ H = H_1 \circ H_2 \circ \cdots \circ H_N \]  

(15)

Such a design also In such a formalism, all intermediate transformations have the same conceptual level in the code. They are functions ranging from spatial reprojection, to temporal interpolations, to more complex operations such as the reconstruction of satellite total columns from concentrations simulated at individual levels in the transport model. In the CIF, all these transformations have the same input and output structure and, thus, their order can be changed seamlessly to execute a given configuration. Please note that the commutative property of elementary transformations as pieces of code does not guarantee the commutative property of the corresponding physical operators.

Such a transformation-based design allows us to rigorously separate transformations and thus to implement and test their respective adjoints more easily. Once adjoints for each individual operation are implemented, the construction of the general adjoint is straightforward by reversing the order of forward operations:

\[ H^* = H_N^* \circ H_{N-1}^* \circ \cdots \circ H_1^* \]  

(16)

Fig. 2 shows an example of a typically targeted observation operator. Operators from Eq. 14 (14) are reported for the illustration. It includes two numerical models chained with each other; they can be for instance a coarse global CTM and a finer resolution regional CTM, such as in Rödenbeck et al. (2009) or Belikov et al. (2016). The system applies a series of transformations to the target vector, including spatial deaggregation for the optimization of emissions by regions, sector deaggregation to separate different activity sectors, reprojection to the CTM’s resolution (a simple interpolation of mass-conserving regridding is available so far, with regular and irregular domains), application of temporal profiles (which is critical in air quality and anthropogenic CO2 applications), unit conversions to the required inputs for the CTMs. On the observation vector side, observations can span multiple model time-steps, requiring posterior temporal averages,
etc. In the case of super-observations (satellites retrievals, images, spatial gradients, etc.) in the observation vector, it is often necessary to combine multiple simulated point observations in given grid cells and time stamps into a single super-observation, to limit redundant observations, hence the size of the observation vector, but also to limit representativeness issues. Super-observations are commonly used in the case for satellite observations being compared to all the model levels above a given location; concentration gradients comparing observations at different time and locations (see e.g., Bréon et al., 2015; Staufer et al., 2016) are another example of observation aggregation to reduce representativeness errors; isotopic ratios are also super-observations as they require to simulate separate isotopologues and recombine them after the simulation (as done in e.g., van der Velde et al., 2018; Peters et al., 2018). This is the case for satellite observations being compared to all the model levels above a given location, or for concentration gradients comparing different time and locations, or also isotopic ratios that require to simulate separate isotopologues and recombine them after the simulation. The case of Fig. 2 also include background concentrations in the target vector: a background is often used to fix some biases in initial and lateral concentrations in limited-area models, and in observations (mostly satellites); the background variables are processed at the very end of the pipe when re-constructing the observations vector.

The mathematical formalism of Eq. (15) and (16) suggests that transformations are necessarily computed in a serialized way, thus limiting applications to simple target variables upstream the transport model. However, each elementary transformation handles components of the inputs it is concerned with, leaving the rest identical and forwarding it to later transformations. Typically, it does not actually limit applications to simple target variables upstream the CTM. For instance, in the case of target variables optimizing biases in the observations, the corresponding components of the target vector \( x \) are forwarded unchanged by all transformations in Fig. 2 until the very last operation, where they are used for the final comparison to the observation vector.

3 Practical implementation

3.1 General rules

The Community Inversion Framework project follows the organisation scheme of Fig. 3. A centralized website is available at community-inversion.eu. The website includes all information given in the present paper, as well as further documentation details, practical installation instructions, tutorials and examples of possible set-ups. To foster the collaborative dynamics of our project, all scripts and codes are available in open-access on a GitLab server at git.nilu.no/VERIFY/CIF, where updates are published regularly. The frozen version of the code, documentation and data used for the present publication is available in Berchet et al. (2020). The repository includes the documentation, sources for the CTMs implemented in the CIF, as well as the Python library pyCIF. Our project is distributed as an
open-source project under the CECILL-C licence of the French law (cecill.info). The license grants full rights for the users to use, modify and redistribute the original version of the CIF, conditional to the obligation to make their modifications available to the community and to properly acknowledge the original authors of the code. The authors of modifications own intellectual property of their modifications, but under the same governing open license. Software that may be built around the CIF in the future can have different licensing, but all parts of the code originating from the CIF will be governed by the original CeCILL-C license, hence must remain open source. Similarly, some constituting pieces of the CIF can be adapted from other softwares governed by other licenses and simply interfaced to the CIF (e.g., transport models, minimizing algorithms, etc.); in such cases, the corresponding softwares keep their original license and their use and distribution in the CIF is subject to authorization by their owners (although open distribution and integration in the standard version of the CIF is encouraged). This is the case of the CONGRAD and M1QN3 algorithms which are used as minimizing algorithms in the variational inversions of the demonstration case in Sect. 4. The M1QN3 algorithm is distributed under the GNU General Public License, whereas CONGRAD is owned by ECMWF and is not open source; the latter was interfaced with the CIF but is not openly distributed.

The pyCIF library, written in Python 3, is the practical embodiment of the CIF project. All theoretical operations described in Sect. 2 are computed by this module. It includes inversion computations, pre- and post-processing of CTM inputs and outputs, as well as target and observation vector reprojections, aggregation, etc., as written in Eq. 15 (15). Python coding standards follow the community standards PEP-8 (python.org/dev/peps/pep-0008/).

Test cases (including the ones presented in Sect. 4) are distributed alongside the CIF codes and scripts. To foster portability and dissemination, a dedicated Docker image is distributed with pyCIF, providing a stable environment to run the system and enabling full reproducibility of the results from one machine to the other.

3.2 Plugin-based implementation

To reflect the theoretical flexibility required in the computation of various inversion methods and observation operators, we made the choice of implementing pyCIF following an abstract structure with a variety of so-called Python plugins, which are dynamically constructed and inter-connected depending on the set-up.

3.2.1 Objects and classes in pyCIF

General classes of objects emerge from the definition of the abstract structure of the inversion framework. These classes are defined by the data and metadata they carry, as well as by the methods they include and their interaction with other classes. The main classes are the following:
– computation modes: forward computations, the test of the adjoint, variational inversions, EnKF En-SRF and analytical inversions are available (see details in Sect. 2.2);

– models: interfaces to CTMs; includes generation of input files, executing the code and post-processing outputs; included are a Gaussian model described in Sect. 4 for the demonstration of the system, as well as CHIMERE, LMDZ, FLEXPART, TM5, and STILT, all of which will be described in a dedicated future publication;

– platforms: deal with specific configurations on different clusters; it includes a standard platform as well as two supercomputers where the CIF was tested;

– control target vectors: store and apply operations related to the control target vector, including spatial and temporal aggregation, deaggregation, regularization of the target vector;

– observation vectors: store and apply operations related to the observation vector, including application of observation errors;

– observation operators: drive CTMs and apply elementary operations between the control and observation vectors;

– transformations: elementary operations used to build the observation operator; includes temporal averaging or deaggregating of the target and observation vectors, projection of the target vector at the model input resolution, etc.;

– data vectors: store all information on inputs for pyCIF: this vector is used by the observation and control target vector classes to build themselves;

– minimizers: algorithms used to minimize cost functions, including M1QN3 and CONGRAD algorithms so far;

– simulators: cost functions to minimize in variational inversions; only includes the standard Gaussian cost function so far;

– domains: store information about the CTM’s grid, including coordinates of grid cell centers and corners, vertical levels, etc.;

– fluxes, fields, meteos meteo-data: fetch, read and write different formats of inputs for CTMs (surface fluxes, 3D fields and meteorological fields respectively); so far includes only inputs specific to included CTMs, but will ultimately include standard data streams, such as widely used emission inventories or meteorological fields such as those from ECMWF;
measurements: fetch, read and write different types of observation data streams; only include the World Data Center for Greenhouse Gases so far (https://gaw.kishou.go.jp/), but classical data providers such as ICOS (icos-cp.eu) or ObsPack Masarie et al. (2014) (Masarie et al., 2014) will also be implemented in the CIF; satellite products should also be included in the CIF in the near future as they play a growing role in the community.

Details on metadata and operations for each class are given in Supplements, Tab. ?? S1. Our objective was to design a code that is fully recursive in the sense that modifying some instance of a class does not require to update other classes calling or being called by the modified class. Thus, each class is built so that it only needs internal data, as well as data from the execution level just before and after it, in order to avoid complex dependencies while allowing proper recursive behaviour in building the transformation pipe. For instance, the observation operator applies a pipe of transformations from the target vector to the observation vector. Some transformations will use the model class to run the model, or the domain class to carry out reprojections, or the target vector to aggregate/de-aggregate target dimensions, etc. Despite the many complex transformations carried out under the umbrella of the observation operator, only the sub-transformations of the pipe are accessible at the observation operator level, which do not have to carry directly information about e.g., the model or other classes required at sub-levels. This makes the practical code of the observation operator much simpler and as easy to read as possible.

3.2.2 Automatic construction of the execution pipe

To translate the principle scheme of Fig. 1, pyCIF is not built in a sequential rigid manner. Plugins are interconnected dynamically at the initializing step of pyCIF depending on the chosen set-up (see Sect. 3.3 for details on the way to configure the CIF). The main strength of such a programming structure is the independence of all objects in pyCIF. They can be implemented separately in a clean manner. The developer only needs to specify what other objects are required to run the one being developed and pyCIF makes the links to the rest. It avoids unexpected impacts elsewhere in the code when modifying or implementing a feature in the system. In the following, we call this top-down relationship in the code a dependency.

For each plugin required in the configuration (primarily the computation mode), pyCIF initializes corresponding objects following simple rules. Following dependencies detailed in Tab. ?? S1, for every object to initialize, pyCIF will fetch and initialize required plugins and attach them to the original plugin. If the required plugin is explicitly defined in the configuration, pyCIF will fetch this one. In some cases, some plugins can be built on default dependencies that do not need to be defined explicitly in the configuration file. In that case, the required plugin can be retrieved using default plugin dependencies specified in the code itself and not needed in the configuration.
For instance, in the call graph in Fig. 1, "variational inversion" (inversion) is a "computation mode" object in pyCIF. To execute, it requires a "minimizer" object (CONGRAD, M1QN3, etc.) that is initialized and attached to it. The minimizer requires a "simulator" object (the cost function) that itself will call functions in the "control vector" object and the "observation operator" object. Then the "observation operator" will initialize a pipeline of transformations including running the "model", and so on and so forth.

3.3 Definition of configurations in the CIF

In practice, pyCIF is configured using a YAML configuration file (yaml.org). This file format was primarily chosen for its flexibility and intuitive implementation of hierarchical parameters. In the YAML language, key words are specified with associated values by the user. Indentations indicate sub-levels of parameters, which makes it a consistent tool with the coding language python.

To set-up a pyCIF computation, the user needs to define the computation mode and all related requirements in the YAML configuration file. Every plugin has mandatory and optional arguments. The absence of one mandatory argument raises an error at initialization. Optional arguments are replaced by corresponding default values if not specified. Examples of YAML configuration files used to carry out the demonstration cases are given in Supplement Section S3.

4 Demonstration case

In the following we describe a demonstration case based on a simple implementation of a Gaussian plume dispersion model and simple inversion set-ups. The purpose of this demonstration case is a proof-of-concept of the CIF, with various inversion methods. We comment and compare inversion set-ups and methods for the purpose of the exercise, but conclusions are not made to be generalized to any inversion case study due to the simplicity of our example. The test application with a simple Gaussian plume model allows users to quickly carry out the test cases themselves, even on desktop computers, to familiarize themselves with the system. Nevertheless, the Gaussian plume model is not only relevant for teaching purposes, but also for real applications, as it is used in many inversion studies from the scale of industrial sites with in-situ fixed or mobile measurements (e.g., Kumar et al., 2020; Foster-Wittig et al., 2015; Ars et al., 2017) to the global scale larger scales with satellite measurements to optimize individual clusters of industrial or urban emissions (e.g., Nassar et al., 2017; Wang et al., 2020). Other models implemented in the CIF will be presented in a future paper evaluating the differences when using different transport models with all other elements of the configuration identical. The purpose of such an evaluation is to produce a rigorous inter-comparison exercise identifying the effect of transport errors in inversion systems.
4.1 Gaussian plume model

Gaussian plume models approximate real turbulent transport by a stable average Gaussian state (Hanna et al., 1982). Such models are not always suitable to compare with continuous measurements but can be adapted when using observations averaged over time. In the following, we consider the Gaussian plume assumption to be valid for comparing to hourly averaged observations. A simple application of the Gaussian plume model was implemented in the CIF as a testing and training utility. It is computationally easy to run, even on desktop computers. It includes the most basic Gaussian plume equations. In that application, concentrations $C$ at location $(x_0, y_0, z_0)$ downwind from a source of intensity $f$ at $(x_1, y_1, z_1)$ are given by:

$$ C(x_0, y_0, z_0) = \frac{f}{2\pi \sigma_y \sigma_z \bar{u}} \exp\left(-\frac{y^2}{\sigma_y^2}\right) \cdot \exp\left(-\frac{z^2}{\sigma_z^2}\right) $$

(17)

with

$$
\begin{align*}
\sigma_z &= a \cdot x^b \\
\sigma_y &= |465.11628 \times x \cdot \tan(0.017653293 (c - d \cdot \ln x))|
\end{align*}
$$

(18)

$x$ is the downwind distance between the source and receptor points along the wind axis, $y$ is the distance between the wind axis and the receptor point, $v_{(source, receptor)}$ is the vector linking the source and the receptor point, $z$ is the difference between the source and the receptor altitudes, $\bar{u}$ is the average wind speed, with $\bar{u}$ is the average wind speed in the domain of simulation. $\langle \cdot | \cdot \rangle$ and $(\cdot \times \cdot)$ depict the scalar and the vector products respectively. $(a, b, c, d)$ are parameters depending on the Pasquill-Gifford atmospheric vertical stability classes. There are 7 Pasquill-Gifford stability classes, from A extremely unstable (mostly in summer during the afternoon) to G very stable (occurring mostly during nighttime in winter). As the purpose of the demonstration case is primarily to work on coarsely realistic concentration fields, with a computational cost as low as possible, our implementation of the Gaussian plume model does not include any representation of particle reflection on the ground or on the top of the planetary boundary layer.
To illustrate atmospheric inversions, we use gridded surface fluxes to simulate concentrations using the Gaussian plume equation. Thus, the concentration at a given point and time \( t \) is the sum of Gaussian plume contributions from all individual grid points:

\[
C(x_0, y_0, z_0, t) = \sum_{(x_1, y_1, z_1) \in \text{grid}} \frac{f(x_1, y_1, z_1, t)}{2 \pi \sigma_y(t) \sigma_z(t) \bar{u}(t)} \exp \left( -\frac{y^2}{\sigma_y(t)^2} \right) \exp \left( -\frac{z^2}{\sigma_z(t)^2} \right)
\]

This formulation highlights the linear relationship between concentrations and fluxes. As the concentrations can be expressed as a matrix product, the computation of the adjoint of the Gaussian model is straightforward and does not require extra developments:

\[
\delta f(x_1, y_1, z_1, t) = \sum_{(x_0, y_0, z_0) \in \text{obs}} \frac{\delta C(x_0, y_0, z_0, t)}{2 \pi \sigma_y(t) \sigma_z(t) \bar{u}(t)} \exp \left( -\frac{y^2}{\sigma_y(t)^2} \right) \exp \left( -\frac{z^2}{\sigma_z(t)^2} \right)
\]

For the purpose of our demonstration cases, meteorological conditions (wind speed, wind direction, and stability class) are randomly generated for the simulation time-window. Fixed seeds are selected for the generation of random conditions in order to make them reproducible.

### 4.2 Configuration

#### 4.2.1 Modelling set-up

Cases discussed in Sect. 4.3 are based on the Gaussian model computed on a domain of \( 2.5 \times 2 \text{ km}^2 \) with a 100 m horizontal resolution grid of 18 x 12 grid cells. Surface point sources are located on a 100 m regular grid at the center of corresponding grid cells, with flux intensities as represented in Fig. 4. Fifty virtual measurement sites are randomly located in the domain with randomly selected altitudes above ground level as shown in Fig. 4. The inversion time-window spans a period of five days with hourly observations and meteorological forcing conditions. Meteorological conditions are a combination of a wind speed, a wind direction and a stability class applicable to the whole simulation domain for each hour. The three parameters are generated randomly for the period, without respect for realistic relatively smooth transitions in wind speed and direction and stability class.

Truth observations are generated by running the Gaussian model in forward mode with the known true fluxes defined as the sum of prior fluxes \( f \) (used later in the inversions) in Eq. 23, then prior fluxes are generated by perturbing the true fluxes as shown (23) and an arbitrary perturbation as defined in Eq. (22), and illustrated in Fig. 4 (left and right respectively).
\[ f = f_0 \cdot \left\{ \cos \left( \frac{2\pi}{\sigma_x^1} x \right) + \sin \left( \frac{2\pi}{\sigma_y^1} y \right) + \left( \frac{x}{\sigma_x^2} \right)^2 + \left( \frac{y}{\sigma_y^2} \right)^2 \right\} \]  

(21)

\[ \delta f = 0.2 \times f_0 \cdot \left\{ \cos \left( \frac{2\pi}{\sigma_x^3} x \right) + \sin \left( \frac{2\pi}{\sigma_y^3} y \right) \right\} \]  

(22)

with \( f_0 \) an arbitrary reference flux, and scaling lengths \( \sigma_x^1, \sigma_x^2, \sigma_x^3, \sigma_y^1, \sigma_y^2, \sigma_y^3 \) equals 500, 1000, 200, 1000, 1000 and 300 m respectively. Reference fluxes and perturbations are constant over time.

A random noise of 1% of the standard deviation of the forward simulations was added to the truth observations to generate measurements. Prior fluxes and perturbation are generated following the equations below. Please note that the perturbation of the fluxes is generated using an explicit formula and not a random perturbation with a given covariance matrix. For that reason, we discuss results with different possible target vector vectors and covariance matrices. This allows us to assess the sensitivity of an inversion method to the resolution and definition of the target vector and corresponding covariance matrix.

The three sets of inversions differ by the dimension of the target vector and the spatial correlations of errors. The first set uses a target vector based on a grid of \( 3 \times 3 \) pixels-aggregated regions or "bands" independent from each other i.e. with no spatial error correlations. The target vectors of the second and third
sets are defined at the grid cell’s resolution with horizontal isotropic error correlations, following an exponential decay with a horizontal scale of 500 m and 10000–200000 m respectively; the latter case is used to demonstrate that the implementation of correlation lengths is correct as very long correlations are equivalent to having only one spatial scaling factor in the target vector. For all inversion set-ups, the target vectors are defined as constant over time, i.e., only one coefficient per spatial dimension is optimized for the 5 days × 24 hours, computed by the model. In all set-ups, the magnitude of the observation noise used to generate "true" observations is chosen as observation errors in the matrix $R$ for consistency. In all cases, the diagonal terms of the $B$ matrix are set to 100%.

To assess the sensitivity of each set-up to the allocated computational resources, we computed the EnKF EnSRF and the two variational inversions with varying numbers of simulations $N$. In the case of the EnKF EnSRF, $N$ simply depicts the size of the Monte Carlo ensemble. For variational inversions, each step, i.e., each computation of the cost function and its gradient requires one forward simulation and one adjoint simulation. The Gaussian model is a simple auto-adjoint model, which makes the adjoint simulations as long as the forward one. Therefore, $N$ is equal to twice the number of computations of the cost function (one for the forward and one for the adjoint) in the minimization algorithm. Note that in many real application cases, the adjoint of a CTM is more costly than the forward, reducing the number of iterations possible in $N$ times the cost of a forward. Indeed, despite the adjoint being mathematically as expensive as the forward, in practice, the computation of adjoint operations often requires the re-computation of intermediate forward computations, therefore increasing the computational burden of the adjoint model. More precisely, users and developers of adjoint transport models choose the number of forward re-computations to be carried out, based on a space-speed trade off: by saving all forward intermediate states, the adjoint is as costly as the forward, but the disk space burden can be extremely challenging to manage, thus increasing the overall computation time in return.

### 4.3 Results and discussion

Posterior increments for the four inversion methods in the three considered demonstration cases are presented in Fig. 4 (horizontally aggregated target vector), Fig. 5 (target vector) at the pixel resolution with horizontal error correlations a correlation length of 500 m) and Fig. 6 (same with correlations of 10000 m). For the sake of readability, figures for other test cases are grouped in Sect. S2 of the Supplement.

Posterior increments are presented in Fig. 5. Observation locations and heights are reported for information. The color scale of flux increments is the same as in Fig. 4 which represent the true "increments" to be retrieved. In Fig. 7, we present the evolution of the cost function of Eq. 6 depending on the number of simulations used for each inversion method for the three demonstration cases (see details on
the corresponding number of simulations of each inversion methods in Sect. 4.2.2). The x-axis has been cropped at the origin as the EnKF EnSRF value for small sizes of random ensembles diverges to infinity.

In the case with the target vector aggregated on groups of pixels, all inversion methods converge towards a very similar solution. In this case, the posterior increments reproduce the overall structure of the truth-prior difference, with four local maxima surrounding one local minimum in the center of the domain. However, the aggregated control target vector results in too coarse patterns which are not representative of the actual true-prior difference. In the case with the target vector at the grid’s resolution with spatial correlations of 500 m, the fully resolved analytical solution captures all methods capture the true-prior difference structure with four maxima surrounding a minimum. However, posterior increments are rather noisy compared to the truth. This is due to the spatial correlations being inconsistent with the smooth perturbation with fixed length scales in Eq. 22. Correlations help smoothing the posterior fluxes but not perfectly consistently with the truth. For the case with the target vector at the grid’s resolution with spatial correlations of 10000–200000 m, the analytical and variational inversions converge to very similar solutions. The EnKF inversion converges towards a relatively noisy solution, suggesting a possible degeneration of the algorithm; this is due to our very simple implementation of the EnKF, with no localization and a rather small number of samples in the random ensemble. All methods converge towards a very smooth and similar solution, consistently with what is expected with a very long correlation length. However, they do not converge towards the same solution, probably because a larger number of iterations/members would be needed to fully converge.

For the three cases, CONGRAD converges towards a cost function value similar to the analytical solution. It also converges at a faster pace than the two other methods and, after 20 and 50 simulations in the aggregated and full-resolution cases respectively, the convergence rate is close to zero, suggesting additional simulations do not provide significant additional information to CONGRAD. The variational inversion with M1QN3 converges towards the analytical solution only in the case of aggregated target vector. The EnKF inversion converges to the analytical solution in two of the three cases (aggregated and full-resolution with small correlation lengths). For the two full-resolution cases, M1QN3 converges to a local minimum instead of a global minimum, probably due to numerical artifacts in the algorithms which makes it mistake the global minimum with a local critical point with a gradient close to zero. This issue may be overcome by allowing the algorithm to carry on for more iterations. The convergence rate with M1QN3 is similar to the one for the EnKF inversion for full-resolution target vectors, but faster for the aggregated case (although additional iterations bring new constraints on the posterior uncertainty matrix).

Overall, CONGRAD appears to be the most cost-efficient algorithm to estimate the analytical solution in the case of a linear inversion in our very simple demonstration case. Though not as efficient, the EnKF EnSRF method can approximate the analytical solution at reduced cost, but by a reduced cost. By de-
sign, its computation can easily be parallelized, which can allow a faster computation than CONGRAD when computational resources are available in parallel. M1QN3 proves not as efficient as its CONGRAD counterpart, but contrary to CONGRAD, it can accommodate non-linear cases.

The reduction of uncertainties and posterior uncertainty matrices are shown in Fig. 6 and 7, and equivalents in Supplement. Regarding posterior uncertainties, CONGRAD proves relatively efficient to approximate the analytical solution, especially at the pixel resolution. The variational inversion with Monte Carlo and M1QN3 computations and the inversion with EnSRF are much noisier. Approximating posterior matrices requires a large number of Monte Carlo members and proves very challenging in real-world applications.

5 Conclusions

We have introduced here a new generic inversion framework that aims at merging existing inversion systems together, in order to share development and maintenance efforts and to foster collaboration on inversion studies. It has been implemented in a way that is fully independent from the inversion configuration: from the application scales, from the species of interest, from the CTM used, from the assumptions for data assimilation, as well as from the practical operations and transformations applied to the data in pre- and post-processing stages. This framework will prevent redundant developments from participating research groups and will allow for a very diverse range of applications within the same system. New developments will be made in an efficient manner with limited risks of unexpected side effects, and thanks to the generic structure of the code, specific developments for a given application can be directly applied to other applications. For instance, new inversion methods implemented in the CIF can be directly tested with various transport models. With modern inversion methods moving towards an hybrid paradigm of variational and ensemble methods, the flexibility of the CIF will be a valuable asset as abstract methods can be easily mixed with each other.

We have presented the first version of this Community Inversion Framework (CIF) alongside with its python-dedicated library pyCIF. As a first step, analytical inversions, variational inversions with M1QN3 and CONGRAD, and EnKF EnSRF have been implemented to demonstrate the general applicability of the CIF. The four inversion techniques were tested here on a test case with a Gaussian Plume Model and with observations generated from known "true" fluxes. The impact of spatial correlations and of spatial aggregation, that drive the shape of the control vectors used in this scientific community, has been illustrated. The analytical inversion is the most accurate approach to retrieve the true fluxes, as expected, followed by variational inversions with the CONGRAD algorithm in our simple linear case. EnKF and M1QN3 are generally less accurate in capturing generally take longer to converge towards the true pattern of the fluxes in our examples, even though EnKF EnSRF inversions...
have the advantage to be fully parallelizable, in contrast to variational inversions, that are sequential by
design and therefore harder to parallelize (e.g., Chevallier, 2013).

The next step of the CIF is the implementation of a large variety of CTMs. The implementation of new
CTMs already interfaced with other inversion systems should not bring particular conceptual challenges as
all interface operations are already written in their original inversion system; in most cases, re-arranging
existing routines is sufficient to interface a model to the CIF. One particular challenge concerns I/O opti-
mizations: the generation of inputs and the processing of outputs can be time consuming and in some very
heavy applications require specific numerical and coding optimizations. The very general formalism of the
CIF may hamper the ability of applying these particular optimizations for some models. Best efforts will
have to be deployed to take full advantage of advanced I/O and data manipulation libraries in python to limit
this weakness.

CHIMERE, LMDz, TM5, FLEXPART, and STILT have already been implemented and a sequel paper
will evaluate and compare their behaviour in similar inversion set-ups. COSMO-GHG and WRF-CHEM are
also planned to be implemented in the near future to widen the developer/user community of the system. The
use of various CTMs in identical inversion configurations (inversion method, observation and target vector,
consistent interface, etc.) will allow extensive assessments of transport errors in inversions. Despite many
past efforts put in inter-comparison exercises, such a level of inter-comparability has never been reached and
will be a natural by-product of the CIF in the future. In addition, comparing posterior uncertainties from
different inversion methods and set-ups will make it possible to fully assess the consistency of different
inversion results.

The flexibility of the CIF allows very complex operations to be included easily. They include the use of
satellite observations, that will be evaluated in a future paper, inversions using observations of isotopic ratios
and optimizing both surface fluxes and source signatures (Thanwerdas et al., 2020) (Thanwerdas et al., 2021)
. In addition, even though GHG greenhouse gas studies have been the main motivation behind the devel-
opment of the CIF, the system will also be tested for multi-species inversions including air pollutants.

*Code and data availability.* The codes, documentation pages (including installation instructions and tutorials) and demon-
stration data used in the present paper are registered under the following DOI: 10.5281/zenodo.4322372 (Berchet et al.,
2020).

*Author contributions.* All authors contributed to the elaboration of the concept of Community Inversion Framework.
Antoine Berchet designed the structure of the CIF. Antoine Berchet, Espen Sollum, Isabelle Pison and Joël Thanwerdas
are the main developers of the python library pyCIF. Antoine Berchet, Espen Sollum and Isabelle Pison maintain the
documentation and websites associated to the Community Inversion Framework.

Acknowledgements. The Community Inversion Framework is currently funded by the project VERIFY (verify.lsce.ipsl.
fr), which received funding from the European Union’s Horizon 2020 research and innovation programme under grant
agreement no. 776810.
Anderson, J., Hoar, T., Raeder, K., Liu, H., Collins, N., Torn, R., and Avellano, A.: The Data Assimilation Research Testbed: A Community Facility, Bull. Amer. Meteor. Soc., 90, 1283–1296, https://doi.org/10.1175/2009BAMS2618.1, https://journals.ametsoc.org/bams/article/90/9/1283/59500/ The-Data-Assimilation-Research-Testbed-A-Community, publisher: American Meteorological Society, 2009.

Ars, S., Broquet, G., Yver Kwok, C., Roustan, Y., Wu, L., Arzoumanian, E., and Bousquet, P.: Statistical atmospheric inversion of local gas emissions by coupling the tracer release technique and local-scale transport modelling: a test case with controlled methane emissions, Atmospheric Measurement Techniques, 10, 5017–5037, https://doi.org/https://doi.org/10.5194/amt-10-5017-2017, https://amt.copernicus.org/articles/10/5017/2017/, publisher: Copernicus GmbH, 2017.

Asefi-Najafabady, S., Rayner, P. J., Gurney, K. R., McRobert, A., Song, Y., Coltin, K., Huang, J., Elvidge, C., and Baugh, K.: A multiyear, global gridded fossil fuel CO$_2$ emission data product: Evaluation and analysis of results, Journal of Geophysical Research: Atmospheres, 119, 10,213–10,231, https://doi.org/10.1002/2013JD021296, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/2013JD021296, _eprint: https://agupubs.onlinelibrary.wiley.com/doi/pdf/10.1002/2013JD021296, 2014.

Babenhauserheide, A., Basu, S., Houweling, S., Peters, W., and Butz, A.: Comparing the CarbonTracker and TM5-4DVar data assimilation systems for CO$_2$ surface flux inversions, Atmospheric Chemistry and Physics, 15, 9747–9763, https://doi.org/10.5194/acp-15-9747-2015, https://acp.copernicus.org/articles/15/9747/2015/acp-15-9747-2015.html, publisher: Copernicus GmbH, 2015.

Bagley, J. E., Jeong, S., Cui, X., Newman, S., Zhang, J., Priest, C., Campos-Pineda, M., Andrews, A. E., Bianco, L., Lloyd, M., Lareau, N., Clements, C., and Fischer, M. L.: Assessment of an atmospheric transport model for annual inverse estimates of California greenhouse gas emissions, J. Geophys. Res. Atmos., p. 2016JD025361, https://doi.org/10.1002/2016JD025361, http://onlinelibrary.wiley.com/doi/10.1002/2016JD025361/abstract, 2017.

Belikov, D. A., Maksyutov, S., Yaremchuk, A., Ganshin, A., Kaminski, T., Blessing, S., Sasakawa, M., Gomez-Pelaez, A. J., and Starchenko, A.: Adjoint of the global Eulerian–Lagrangian coupled atmospheric transport model (A-GELCA v1.0): development and validation, Geoscientific Model Development, 9, 749–764, https://doi.org/https://doi.org/10.5194/gmd-9-749-2016, https://gmd.copernicus.org/articles/9/749/2016/, publisher: Copernicus GmbH, 2016.

Berchet, A., Pison, I., Chevallier, F., Bousquet, P., Bonne, J.-L., and Paris, J.-D.: Objectified quantification of uncertainties in Bayesian atmospheric inversions, Geosci. Model Dev. Discuss., 7, 4777–4827, https://doi.org/10.5194/gmd-7-4777-2014, http://www.geosci-model-dev-discuss.net/7/4777/2014/, 2014.

Berchet, A., Sollum, E., Pison, I., Thompson, R. L., Thanwerdas, J., Fortemps-Cheiney, A., Peet, J. C. A. v., Potier, E., Chevallier, F., and Broquet, G.: The Community Inversion Framework: codes and documentation, https://doi.org/10.5281/zenodo.4322372, https://doi.org/10.5281/zenodo.4322372, 2020.

Bergamaschi, P., Krol, M., Meirink, J. F., Dentener, F., Segers, A., van Aardenne, J., Monni, S., Vermeulen, A. T., Schmidt, M., Ramonet, M., Yver, C., Meinhardt, F., Nisbet, E. G., Fisher, R. E., O’Doherty, S., and Dlugokencky, E. J.: Inverse modeling of European CH$_4$ emissions 2001-2006, J. Geophys.
Bergamaschi, P., Danila, A., Weiss, R. F., Ciais, P., Thompson, R. L., Brunner, D., Levin, I., Meijer, Y., Chevallier, F., Janssens-Maenhout, G., Bovensmann, H., Crisp, D., Basu, S., Dlugokencky, E. J., Engel, R., Gerbig, C., Gv’unther, D., Hammer, S., Henne, S., Houweling, S., Karstens, U., Kort, E. A., Maione, M., Manning, A. J., Miller, J. B., Montzka, S. A., Pandey, S., Peters, W., Peylin, P., Pinty, B., Ramonet, M., Reimann, S., Röckmann, T., Schmidt, M., Stroes, M., Sussams, J., Tarasova, O. A., van Aardenne, J., Vermeulen, A. T., and Vogel, F.: Atmospheric monitoring and inverse modelling for verification of greenhouse gas inventories, EUR 29276 EN JRC111789, Publications Office of the European Union, Luxembourg, https://ec.europa.eu/jrc/en/publication/eur-scientific-and-technical-research-reports/atmospheric-monitoring-and-inverse-modelling-verification-greenhouse-gas-inventories, 2018a.

Bergamaschi, P., Karstens, U., Manning, A. J., Saunois, M., Tsuruta, A., Berchet, A., Vermeulen, A. T., Arnold, T., Janssens-Maenhout, G., Hammer, S., Levin, I., Schmidt, M., Ramonet, M., Lopez, M., Lavric, J., Aalto, T., Chen, H., Feist, D. G., Gerbig, C., Haszpra, L., Hermansen, O., Manca, G., Moncrieff, J., Meinhardt, F., Necki, J., Galkowski, M., O’Doherty, S., Paramonova, N., Scheeren, H. A., Steinbacher, M., and Dlugokencky, E.: Inverse modelling of European CH$_4$ emissions during 2006–2012 using different inverse models and reassessed atmospheric observations, Atmospheric Chemistry and Physics, 18, 901–920, https://doi.org/https://doi.org/10.5194/acp-18-901-2018, https://acp.copernicus.org/articles/18/901/2018/, publisher: Copernicus GmbH, 2018b.

Bocquet, M.: Ensemble Kalman filtering without the intrinsic need for inflation, Nonlin. Processes Geophys., 18, 735–750, https://doi.org/10.5194/npg-18-735-2011, http://www.nonlin-processes-geophys.net/18/735/2011/, 2011.

Bocquet, M., Elber, H., Eskes, H., Žabkar, R., Carmichael, G. R., Flemming, J., Inness, A., Pagowski, M., Pérez Camaño, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J., Baklanov, A., Carnevale, C., Grell, G., and Seigneur, C.: Data assimilation in atmospheric chemistry models: current status and future prospects for coupled chemistry meteorology models, Atmospheric Chemistry and Physics, 15, 5325–5358, https://doi.org/https://doi.org/10.5194/acp-15-5325-2015, https://acp.copernicus.org/articles/15/5325/2015/, publisher: Copernicus GmbH, 2015.

Bousquet, P., Ringeval, B., Pison, I., Dlugokencky, E. J., Brunke, E. G., Carouge, C., Chevallier, F., Fortems-Chiney, A., Frankenberg, C., Hauglustaine, D. A., et al.: Source attribution of the changes in atmospheric methane for 2006-2008, Atmos. Chem. Phys., 11, 3689–3700, http://www.atmos-chem-phys.org/11/3689/2011/acp-11-3689-2011.pdf, 2011.

Bousserez, N. and Henze, D. K.: Optimal and scalable methods to approximate the solutions of large-scale Bayesian problems: theory and application to atmospheric inversion and data assimilation, Quarterly Journal of the Royal Meteorological Society, 144, 365–390, https://doi.org/10.1002/qj.3209, https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/qj.3209, 2018.

Bousserez, N., Henze, D. K., Perkins, A., Bowman, K. W., Lee, M., Liu, J., Deng, F., and Jones, D. B. A.: Improved analysis-error covariance matrix for high-dimensional variational inversions: application to source estimation using a 3D atmospheric transport model, Quarterly Journal of the Royal Meteorological Society, 141, 1906–1921, https://doi.org/10.1002/qj.2495, https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/qj.2495, _eprint: https://rmets.onlinelibrary.wiley.com/doi/pdf/10.1002/qj.2495, 2015.
Brioude, J., Angevine, W. M., Ahmadov, R., Kim, S.-W., Evan, S., McKeen, S. A., Hsie, E.-Y., Frost, G. J., Neuman, J. A., Pollack, I. B., Peischl, J., Ryerson, T. B., Holloway, J., Brown, S. S., Nowak, J. B., Roberts, J. M., Wofsy, S. C., Santoni, G. W., Oda, T., and Trainer, M.: Top-down estimate of surface flux in the Los Angeles Basin using a mesoscale inverse modeling technique: assessing anthropogenic emissions of CO, NO$_x$ and CO$_2$ and their impacts, Atmospheric Chemistry and Physics, 13, 3661–3677, https://doi.org/https://doi.org/10.5194/acp-13-3661-2013, https://acp.copernicus.org/articles/13/3661/2013/, publisher: Copernicus GmbH, 2013.

Broquet, G., Chevallier, F., Rayner, P., Aulagnier, C., Pison, I., Ramonet, M., Schmidt, M., Vermeulen, A. T., and Ciais, P.: A European summertime CO$_2$ biogenic flux inversion at mesoscale from continuous in situ mixing ratio measurements, Journal of Geophysical Research: Atmospheres, 116, n/a–n/a, https://doi.org/10.1029/2011JD016202, http://onlinelibrary.wiley.com/doi/10.1029/2011JD016202/abstract, 2011.

Bruhwiler, L. M. P., Michalak, A. M., Peters, W., Baker, D. F., and Tans, P.: An improved Kalman Smoother for atmospheric inversions, Atmospheric Chemistry and Physics, 5, 2691–2702, https://doi.org/https://doi.org/10.5194/acp-5-2691-2005, https://acp.copernicus.org/articles/5/2691/2005/acp-5-2691-2005.html, publisher: Copernicus GmbH, 2005.

Brunner, D., Henne, S., Keller, C. A., Reimann, S., Vollmer, M. K., O’Doherty, S., and Maione, M.: An extended Kalman-filter for regional scale inverse emission estimation, Atmospheric Chemistry and Physics, 12, 3455–3478, https://doi.org/https://doi.org/10.5194/acp-12-3455-2012, https://acp.copernicus.org/articles/12/3455/2012/, publisher: Copernicus GmbH, 2012.

Brunner, D., Arnold, T., Henne, S., Manning, A., Thompson, R. L., Maione, M., O’Doherty, S., and Reimann, S.: Comparison of four inverse modelling systems applied to the estimation of HFC-125, HFC-134a, and SF$_6$ emissions over Europe, Atmospheric Chemistry and Physics, 17, 10651–10674, https://doi.org/https://doi.org/10.5194/acp-17-10651-2017, https://acp.copernicus.org/articles/17/10651/2017/, publisher: Copernicus GmbH, 2017.

Bréon, F. M., Broquet, G., Puygrenier, V., Chevallier, F., Xueref-Remy, I., Ramonet, M., Dieudonné, E., Lopez, M., Schmidt, M., Perrussel, O., and Ciais, P.: An attempt at estimating Paris area CO2 emissions from atmospheric concentration measurements, Atmos. Chem. Phys., 15, 1707–1724, https://doi.org/10.5194/acp-15-1707-2015, http://www.atmos-chem-phys.net/15/1707/2015/, 2015.

Chatterjee, A., Michalak, A. M., Anderson, J. L., Mueller, K. L., and Yadav, V.: Toward reliable ensemble Kalman filter estimates of CO2 fluxes, Journal of Geophysical Research: Atmospheres, 117, https://doi.org/https://doi.org/10.1029/2012JD018176, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2012JD018176, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/2012JD018176, 2012.

Chevallier, F.: On the parallelization of atmospheric inversions of CO2 surface fluxes within a variational framework, Geosci. Model Dev., 6, 783–790, https://doi.org/10.5194/gmd-6-783-2013, https://www.geosci-model-dev.net/6/783/2013/, 2013.

Chevallier, F., Fisher, M., Peylin, P., Serrar, S., Bousquet, P., Bréon, F. M., Chédin, A., and Ciais, P.: Inferring CO$_2$ sources and sinks from satellite observations: Method and application to TOVS data, J. Geophys. Res., 110, D24309, https://doi.org/10.1029/2005JD006390, http://www.agu.org/pubs/crossref/2005/2005JD006390.shtml, 2005.

Chevallier, F., Ciais, P., Conway, T. J., Aalto, T., Anderson, B. E., Bousquet, P., Brunke, E. G., Ciattaglia, L., Esaki, Y., Fröhlich, M., Gomez, A., Gomez-Pelaez, A. J., Haszpra, L., Krümmel, P. B., Langenfelds, R. L., Leuenberger, M.,
Machida, T., Maignan, F., Matsueda, H., Morguí, J. A., Mukai, H., Nakazawa, T., Peylin, P., Ramonet, M., Rivier, L., Sawa, Y., Schmidt, M., Steele, L. P., Vay, S. A., Vermeulen, A. T., Wofsy, S., and Worthy, D.: CO₂ surface fluxes at grid point scale estimated from a global 21 year reanalysis of atmospheric measurements, J. Geophys. Res., 115, D21 307, https://doi.org/10.1029/2010JD013887, http://onlinelibrary.wiley.com/doi/10.1029/2010JD013887/abstract, 2010.

Chevallier, F., Remaud, M., O’Dell, C. W., Baker, D., Peylin, P., and Cozic, A.: Objective evaluation of surface- and satellite-driven carbon dioxide atmospheric inversions, Atmospheric Chemistry and Physics, 19, 14 233–14 251, https://doi.org/https://doi.org/10.5194/acp-19-14233-2019, https://acp.copernicus.org/articles/19/14233/2019/, publisher: Copernicus GmbH, 2019.

Ciais, P., Dolman, A. J., Bombelli, A., Duren, R., Peregon, A., Rayner, P. J., Miller, C., Gobron, N., Kinderman, G., Marland, G., Gruber, N., Chevallier, F., Andres, R. J., Balsamo, G., Bopp, L., Bréon, F.-M., Broquet, G., Dargaville, R., Battin, T. J., Borges, A., Bovensmann, H., Buchwitz, M., Butler, J., Canadell, J. G., Cook, R. B., DeFries, R., Engelke, R., Gurney, K. R., Heinze, C., Heimann, M., Held, A., Henry, M., Law, B., Luyssaert, S., Miller, J., Moriyama, T., Moulin, C., Myhreni, R. B., Nussli, C., Obersteiner, M., Ojima, D., Pan, Y., Paris, J.-D., Piao, S. L., Poulter, B., Plummer, S., Quegan, S., Raymond, P., Reichstein, M., Rivier, L., Sabine, C., Schimel, D., Tarasova, O., Valentini, R., Wang, R., van der Werf, G., Wickland, D., Williams, M., and Zehner, C.: Current systematic carbon-cycle observations and the need for implementing a policy-relevant carbon observing system, Biogeosciences, 11, 3547–3602, https://doi.org/https://doi.org/10.5194/bg-11-3547-2014, https://bg.copernicus.org/articles/11/3547/2014/, publisher: Copernicus GmbH, 2014.

Crisp, D., Meijer, Y., Munro, R., Bowman, K., Chatterjee, A., Baker, D., Chevallier, F., Nassar, R., Palmer, P. I., Agusti-Panareda, A., and others: A constellation architecture for monitoring carbon dioxide and methane from space, Prepared by the CEOS Atmospheric Constellation Greenhouse Gas Team, Version 1, 2018.

Crowell, S., Baker, D., Schuh, A., Basu, S., Jacobson, A. R., Chevallier, F., Liu, J., Deng, F., Feng, L., McKain, K., Chatterjee, A., Miller, J. B., Stephens, B. B., Eldering, A., Crisp, D., Schimel, D., Nassar, R., O’Dell, C. W., Oda, T., Sweeney, C., Palmer, P. I., and Jones, D. B. A.: The 2015–2016 carbon cycle as seen from OCO-2 and the global in situ network, Atmospheric Chemistry and Physics, 19, 9797–9831, https://doi.org/https://doi.org/10.5194/acp-19-9797-2019, https://acp.copernicus.org/articles/19/9797/2019/, publisher: Copernicus GmbH, 2019.

Curier, R. L., Timmermans, R., Calabretta-Jongen, S., Eskes, H., Segers, A., Swart, D., and Schaap, M.: Improving ozone forecasts over Europe by synergistic use of the LOTOS-EUROS chemical transport model and in-situ measurements, Atmospheric Environment, 60, 217–226, https://doi.org/10.1016/j.atmosenv.2012.06.017, http://www.sciencedirect.com/science/article/pii/S1352231012005687, 2012.

Enting, I. G.: Inverse problems in atmospheric constituent transport, Cambridge University Press, http://books.google.fr/books?hl=fr&lr=&id=b61nA7D1XXUC&oi=fnd&pg=PP1&dq=enting+inverse+problem+atmospheric+constituent+transport+&ots=vc8hwiQ_YK&sig=u5npCEt4WdBNLyrDn9V2EfD1SQ, 2002.

Errico, R. M.: What Is an Adjoint Model?, Bull. Amer. Meteor. Soc., 78, 2577–2592, https://doi.org/10.1175/1520-0477(1997)078<2577:WIAAM>2.0.CO;2, https://journals.ametsoc.org/bams/article/78/11/2577/55799/What-Is-an-Adjoint-Model, publisher: American Meteorological Society, 1997.

Feng, L., Palmer, P. I., Bösch, H., and Dance, S.: Estimating surface CO₂ fluxes from space-borne CO₂ dry air mole fraction observations using an ensemble Kalman Filter, Atmospheric Chemistry and Physics, 9, 2619–2633,
Feng, L., Palmer, P. I., Bösch, H., Parker, R. J., Webb, A. J., Correia, C. S. C., Deutscher, N. M., Domingues, L. G., Feist, D. G., Gatti, L. V., Gloor, E., Hase, F., Kivi, R., Liu, Y., Miller, J. B., Morino, I., Sussmann, R., Strong, K., Uchino, O., Wang, J., and Zahn, A.: Consistent regional fluxes of CH\(_4\) and CO\(_2\) inferred from GOSAT proxy XCH\(_4\) : XCO\(_2\) retrievals, 2010–2014, Atmospheric Chemistry and Physics, 17, 4781–4797, https://doi.org/https://doi.org/10.5194/acp-17-4781-2017, https://acp.copernicus.org/articles/17/4781/2017/, publisher: Copernicus GmbH, 2017.

Filges, A., Gerbig, C., Chen, H., Franke, H., Klaus, C., and Jordan, A.: The IAGOS-core greenhouse gas package: a measurement system for continuous airborne observations of CO\(_2\), CH\(_4\), H\(_2\)O and CO, Tellus B: Chemical and Physical Meteorology, 68, 27989, https://doi.org/10.3402/tellusb.v67.27989, https://www.tandfonline.com/doi/abs/10.3402/tellusb.v67.27989%40zelb20.2016.68.issue-s1, publisher: Taylor & Francis eprint: https://www.tandfonline.com/doi/pdf/10.3402/tellusb.v67.27989%40zelb20.2016.68.issue-s1, 2016.

Fisher, M.: Minimization algorithms for variational data assimilation, in: Seminar on Recent Developments in Numerical Methods for Atmospheric Modelling, 7-11 September 1998, pp. 364–385, ECMWF, Shinfield Park, Reading, https://www.ecmwf.int/node/9400, backup Publisher: ECMWF, 1998.

Fortems-Cheiney, A., Pison, I., Broquet, G., Dufour, G., Berchet, A., Potier, E., Coman, A., Siour, G., and Costantino, L.: Variational regional inverse modeling of reactive species emissions with PYVAR-CHIMERE-v2019, Geoscientific Model Development, 14, 2939–2957, https://doi.org/10.5194/gmd-14-2939-2021, https://gmd.copernicus.org/articles/14/2939/2021/, publisher: Copernicus GmbH, 2021.

Foster-Wittig, T. A., Thoma, E. D., and Albertson, J. D.: Estimation of point source fugitive emission rates from a single sensor time series: A conditionally-sampled Gaussian plume reconstruction, Atmospheric Environment, 115, 101–109, https://doi.org/10.1016/j.atmosenv.2015.05.042, http://www.sciencedirect.com/science/article/pii/S135223101530114X, 2015.

Ganesan, A. L., Rigby, M., Zammit-Mangion, A., Manning, A. J., Prinn, R. G., Fraser, P. J., Harth, C. M., Kim, K.-R., Krummel, P. B., Li, S., Mühle, J., O’Doherty, S. J., Park, S., Salameh, P. K., Steele, L. P., and Weiss, R. F.: Characterization of uncertainties in atmospheric trace gas inversions using hierarchical Bayesian methods, Atmos. Chem. Phys., 14, 3855–3864, https://doi.org/10.5194/acp-14-3855-2014, http://www.atmos-chem-phys.net/14/3855/2014/, 2014.

Gaubert, B., Emmons, L. K., Raeder, K., Tilmes, S., Miyazaki, K., Arellano Jr., A. F., Elguindi, N., Granier, C., Tang, W., Barré, J., Worden, H. M., Buchholz, R. R., Edwards, D. P., Franke, P., Anderson, J. L., Saunois, M., Schroeder, J., Woo, J.-H., Simpson, I. J., Blake, D. R., Meinardi, S., Wennberg, P. O., Crounse, J., Teng, A., Kim, M., Dickerson, R. R., He, H., Ren, X., Pusede, S. E., and Diskin, G. S.: Correcting model biases of CO in East Asia: impact on oxidant distributions during KORUS-AQ, Atmospheric Chemistry and Physics, 20, 14 617–14 647, https://doi.org/10.5194/acp-20-14617-2020, https://acp.copernicus.org/articles/20/14617/2020/, publisher: Copernicus GmbH, 2020.

Gilbert, J. C. and Lemaréchal, C.: Some numerical experiments with variable-storage quasi-Newton algorithms, Mathematical Programming, 45, 407–435, https://doi.org/10.1007/BF01589113, https://link.springer.com/article/10.1007/BF01589113, 1989.
Gurney, K. R., Law, R. M., Denning, A. S., Rayner, P. J., Baker, D., Bousquet, P., Bruhwiler, L., Chen, Y.-H., Ciais, P., Fan, S., Fung, I. Y., Gloor, M., Heimann, M., Higuchi, K., John, J., Kowalczyk, E., Maki, T., Maksyutov, S., Peylin, P., Prather, M., Pak, B. C., Sarmiento, J., Taguchi, S., Takahashi, T., and Yuen, C.-W.: TransCom 3 CO2 inversion intercomparison: 1. Annual mean control results and sensitivity to transport and prior flux information, Tellus B, 55, 555–579, https://doi.org/10.1034/j.1600-0889.2003.00049.x, http://www.tellusb.net/index.php/tellusb/article/view/16728, 2003.

Hanna, S. R., Briggs, G. A., and Hosker, J.: Handbook on atmospheric diffusion, Tech. Rep. DOE/TIC-11223, National Oceanic and Atmospheric Administration, Oak Ridge, TN (USA). Atmospheric Turbulence and Diffusion Lab., https://doi.org/10.2172/5591108, https://www.osti.gov/biblio/5591108, 1982.

Henne, S., Brunner, D., Oney, B., Leuenberger, M., Eugster, W., Bamberger, I., Meinhardt, F., Steinbacher, M., and Emmenegger, L.: Validation of the Swiss methane emission inventory by atmospheric observations and inverse modelling, Atmos. Chem. Phys., 16, 3683–3710, https://doi.org/10.5194/acp-16-3683-2016, http://www.atmos-chem-phys.net/16/3683/2016/, 2016.

Houweling, S., Badawy, B., Baker, D. F., Basu, S., Belikov, D., Bergamaschi, P., Bousquet, P., Broquet, G., Butler, T., Canadell, J. G., Chen, J., Chevallier, F., Ciais, P., Collatz, G. J., Denning, S., Engelen, R., Enting, I. G., Fischer, M. L., Fraser, A., Gerbig, C., Gloor, M., Jacobson, A. R., Jones, D. B. A., Heimann, M., Khalil, A., Kaminski, T., Kasibhatla, P. S., Krakauer, N. Y., Krol, M., Maki, T., Maksyutov, S., Manning, A., Meesters, A., Miller, J. B., Palmer, P. I., Patra, P., Peters, W., Peylin, P., Pouzi, Z., Prather, M. J., Randerson, J. T., Röckmann, T., Rödenbeck, C., Sarmiento, J. L., Schimel, D. S., Scholze, M., Schuh, A., Suntharalingam, P., Takahashi, T., Turnbull, J., Yurganov, L., and Vermeulen, A.: Iconic CO2 Time Series at Risk, Science, 337, 1038–1040, https://doi.org/10.1126/science.337.6098.1038-b, https://science.sciencemag.org/content/337/6098/1038.2, publisher: American Association for the Advancement of Science Section: Letters, 2012.

Houweling, S., Krol, M., Bergamaschi, P., Frankenberg, C., Dlugokencky, E. J., Morino, I., Notholt, J., Sherlock, V., Wunch, D., Beck, V., Gerbig, C., Chen, H., Kort, E. A., Röckmann, T., and Aben, I.: A multi-year methane inversion using SCIAMACHY, accounting for systematic errors using TCCON measurements, Atmos. Chem. Phys., 14, 3991–4012, https://doi.org/10.5194/acp-14-3991-2014, http://www.atmos-chem-phys.net/14/3991/2014/, 2014.

Houweling, S., Bergamaschi, P., Chevallier, F., Heimann, M., Kaminski, T., Krol, M., Michalak, A. M., and Patra, P.: Global inverse modeling of CH4 sources and sinks: An overview of methods, Atmospheric Chemistry and Physics Discussions, pp. 1–30, https://doi.org/10.5194/acp-2016-572, http://www.atmos-chem-phys-discuss.net/acp-2016-572/, 2016.

Houweling, S., Bergamaschi, P., Chevallier, F., Heimann, M., Kaminski, T., Krol, M., Michalak, A. M., and Patra, P.: Global inverse modeling of CO2 sources and sinks: an overview of methods, Atmospheric Chemistry and Physics, 17, 235–256, https://doi.org/10.5194/acp-17-235-2017, https://acp.copernicus.org/articles/17/235/2017/, publisher: Copernicus GmbH, 2017.

Ide, K., Courtier, P., Ghil, M., and Lorenc, A. C.: Unified notation for data assimilation: operational, sequential and variational, J. Meteorol. Soc. Japan, pp. 181–189, http://citeseerx.ist.psu.edu/viewdoc/download?doi=10.1.1.47.2607&rep=rep1&type=pdf, 1997.
Janssens-Maenhout, G., Pinty, B., Dowell, M., Zunker, H., Andersson, E., Balsamo, G., Bézy, J.-L., Brunhes, T., Bösch, H., Bojkov, B., Brunner, D., Buchwitz, M., Crisp, D., Ciais, P., Coumet, P., Dee, D., Denier van der Gon, H., Dolman, H., Drinkwater, M. R., Dubovik, O., Engelen, R., Fehr, T., Fernandez, V., Heimann, M., Holmlund, K., Houweling, S., Husband, R., Juvyns, O., Kentarchos, A., Landgraf, J., Lang, R., Löscher, A., Marshall, J., Meijer, Y., Nakajima, M., Palmer, P. I., Peylin, P., Rayner, P., Schulze, M., Sierk, B., Tamminen, J., and Veeck, P.: Toward an Operational Anthropogenic CO2 Emissions Monitoring and Verification Support Capacity, Bull. Amer. Meteor. Soc., 101, E1439–E1451, https://doi.org/10.1175/BAMS-D-19-0017.1, https://journals.ametsoc.org/bams/article/101/8/E1439/345576/Toward-an-Operational-Anthropogenic-CO2-Emissions, publisher: American Meteorological Society, 2020.

Kang, J.-S., Kalnay, E., Liu, J., Fung, I., Miyoshi, T., and Ide, K.: “Variable localization” in an ensemble Kalman filter: Application to the carbon cycle data assimilation, Journal of Geophysical Research: Atmospheres, 116, https://doi.org/10.1029/2010JD014673, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2010JD014673, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/2010JD014673, 2011.

Kang, J.-S., Kalnay, E., Miyoshi, T., Liu, J., and Fung, I.: Estimation of surface carbon fluxes with an advanced data assimilation methodology, Journal of Geophysical Research: Atmospheres, 117, n/a–n/a, https://doi.org/10.1029/2012JD018259, http://onlinelibrary.wiley.com/doi/10.1029/2012JD018259/abstract, 2012.

Karion, A., Sweeney, C., Tans, P., and Newberger, T.: AirCore: An Innovative Atmospheric Sampling System, J. Atmos. Oceanic Technol., 27, 1839–1853, https://doi.org/10.1175/2010JTECHA1448.1, https://journals.ametsoc.org/jtech/article/27/11/1839/3375/AirCore-An-Innovative-Atmospheric-Sampling-System, publisher: American Meteorological Society, 2010.

Kopacz, M., Jacob, D. J., Henze, D. K., Heald, C. L., Streets, D. G., and Zhang, Q.: Comparison of adjoint and analytical Bayesian inversion methods for constraining Asian sources of carbon monoxide using satellite (MO-PITT) measurements of CO columns, J. Geophys. Res., 114, D04 305, https://doi.org/10.1029/2007JD009264, http://onlinelibrary.wiley.com/doi/10.1029/2007JD009264/abstract, 2009.

Krol, M., Houweling, S., Bregman, B., van den Broek, M., Segers, A., van Velthoven, P., Peters, W., Dentener, F., and Bergamaschi, P.: The two-way nested global chemistry-transport zoom model TM5: algorithm and applications, Atmospheric Chemistry and Physics, 5, 417–432, https://doi.org/https://doi.org/10.5194/acp-5-417-2005, https://acp.copernicus.org/articles/5/417/2005/acp-5-417-2005.html, publisher: Copernicus GmbH, 2005.

Kuhlmann, G., Broquet, G., Marshall, J., Clément, V., Löscher, A., Meijer, Y., and Brunner, D.: Detectability of CO2 emission plumes of cities and power plants with the Copernicus Anthropogenic CO2 Monitoring (CO2M) mission, Atmospheric Measurement Techniques, 12, 6695–6719, https://doi.org/10.5194/amt-12-6695-2019, https://amt.copernicus.org/articles/12/6695/2019/, publisher: Copernicus GmbH, 2019.

Kumar, P., Broquet, G., Yver-Kwok, C., Laurent, O., Gichuki, S., Caldow, C., Croke, F., Lauvaux, T., Ramonet, M., Berthe, G., Martin, F., Duclaux, O., Juery, C., Bouchet, C., and Ciais, P.: Mobile atmospheric measurements and local-scale inverse estimation of the location and rates of brief CH4 and CO2 releases from point sources, Atmospheric Measurement Techniques Discussions, pp. 1–27, https://doi.org/https://doi.org/10.5194/amt-2020-226, https://amt.copernicus.org/preprints/amt-2020-226/, publisher: Copernicus GmbH, 2020.

Liu, J., Bowman, K. W., and Henze, D. K.: Source-receptor relationships of column-average CO2 and implications for the impact of observations on flux inversions, Journal of Geophysical Research: Atmospheres,
Liu, Y., Haussaire, J.-M., Bocquet, M., Roustan, Y., Saunier, O., and Mathieu, A.: Uncertainty quantification of pollutant source retrieval: comparison of Bayesian methods with application to the Chernobyl and Fukushima Daiichi accidental releases of radionuclides, Quarterly Journal of the Royal Meteorological Society, 143, 2866–2901, https://doi.org/10.1002/qj.3138, https://rmets.onlinelibrary.wiley.com/doi/abs/10.1002/qj.3138, _eprint: https://rmets.onlinelibrary.wiley.com/doi/pdf/10.1002/qj.3138, 2017.

Locatelli, R., Bousquet, P., Chevallier, F., Fortems-Cheney, A., Szopa, S., Saunois, M., Agusti-Panareda, A., Bergmann, D., Bian, H., Cameron-Smith, P., Chipperfield, M. P., Gloor, E., Houweling, S., Kawa, S. R., Krol, M., Patra, P. K., Prinn, R. G., Rigby, M., Saito, R., and Wilson, C.: Impact of transport model errors on the global and regional methane emissions estimated by inverse modelling, Atmospheric Chemistry and Physics Discussions, 13, 10961–11021, https://doi.org/10.5194/acpd-13-10961-2013, http://www.atmos-chem-phys-discuss.net/13/10961/2013/acpd-13-10961-2013.html, 2013.

Lunt, M. F., Rigby, M., Ganesan, A. L., and Manning, A. J.: Estimation of trace gas fluxes with objectively determined basis functions using reversible-jump Markov chain Monte Carlo, Geosci. Model Dev., 9, 3213–3229, https://doi.org/10.5194/gmd-9-3213-2016, http://www.geosci-model-dev.net/9/3213/2016/, 2016.

Masarie, K. A., Pétron, G., Andrews, A., Bruhwiler, L., Conway, T. J., Jacobson, A. R., Miller, J. B., Tans, P. P., Worthy, D. E., and Peters, W.: Impact of CO2 measurement bias on CarbonTracker surface flux estimates, Journal of Geophysical Research: Atmospheres, 116, https://doi.org/10.1029/2011JD016270, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2011JD016270, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/2011JD016270, 2011.

Masarie, K. A., Peters, W., Jacobson, A. R., and Tans, P. P.: ObsPack: a framework for the preparation, delivery, and attribution of atmospheric greenhouse gas measurements, Earth System Science Data, 6, 375–384, https://doi.org/10.5194/essd-6-375-2014, https://www.earth-syst-sci-data.net/6/375/2014/essd-6-375-2014.html, 2014.

Meirink, J. F., Bergamaschi, P., and Krol, M. C.: Four-dimensional variational data assimilation for inverse modelling of atmospheric methane emissions: method and comparison with synthesis inversion, Atmospheric Chemistry and Physics, 8, 6341–6353, https://doi.org/10.5194/acp-8-6341-2008, https://acp.copernicus.org/articles/8/6341/2008/acp-8-6341-2008.html, publisher: Copernicus GmbH, 2008.

Michalak, A. M.: Technical Note: Adapting a fixed-lag Kalman smoother to a geostatistical atmospheric inversion framework, Atmospheric Chemistry and Physics, 8, 6789–6799, https://doi.org/10.5194/acp-8-6789-2008, https://acp.copernicus.org/articles/8/6789/2008/acp-8-6789-2008.html, publisher: Copernicus GmbH, 2008.

Michalak, A. M. and Kitanidis, P. K.: A method for the interpolation of nonnegative functions with an application to contaminant load estimation, Stoch Environ Res Ris Assess, 19, 8–23, https://doi.org/10.1007/s00477-004-0189-1, http://link.springer.com/article/10.1007/s00477-004-0189-1, 2005.

Michalak, A. M., Hirsch, A., Bruhwiler, L., Gurney, K. R., Peters, W., and Tans, P. P.: Maximum likelihood estimation of covariance parameters for Bayesian atmospheric trace gas surface flux inversions, Journal of Geophysical
Miller, S. M., Michalak, A. M., and Levi, P. J.: Atmospheric inverse modeling with known physical bounds: an example from trace gas emissions, Geosci. Model Dev., 7, 303–315, https://doi.org/10.5194/gmd-7-303-2014, http://www.geosci-model-dev.net/7/303/2014/, 2014.

Miller, S. M., Saibaba, A. K., Trudeau, M. E., Mountain, M. E., and Andrews, A. E.: Geostatistical inverse modeling with very large datasets: an example from the OCO-2 satellite, Geoscientific Model Development Discussions, pp. 1–25, https://doi.org/https://doi.org/10.5194/gmd-2019-185, https://www.geosci-model-dev-discuss.net/gmd-2019-185/, 2019.

Mizzi, A. P., Arellano Jr., A. F., Edwards, D. P., Anderson, J. L., and Pfister, G. G.: Assimilating compact phase space retrievals of atmospheric composition with WRF-Chem/DART: a regional chemical transport/ensemble Kalman filter data assimilation system, Geoscientific Model Development, 9, 965–978, https://doi.org/https://doi.org/10.5194/gmd-9-965-2016, https://gmd.copernicus.org/articles/9/965/2016/, publisher: Copernicus GmbH, 2016.

Monteil, G. and Scholze, M.: Regional CO\textsubscript{2} inversions with LUMIA, the Lund University Modular Inversion Algorithm, v1.0, Geoscientific Model Development Discussions, pp. 1–35, https://doi.org/https://doi.org/10.5194/gmd-2019-227, https://www.geosci-model-dev-discuss.net/gmd-2019-227/, 2019.

Monteil, G., Broquet, G., Scholze, M., Lang, M., Karstens, U., Gerbig, C., Koch, F.-T., Smith, N. E., Thompson, R. L., van der Laan-Luijkkx, I. T., White, E., Meesters, A., Ciais, P., Ganesan, A. L., Manning, A., Mischurow, M., Peters, W., Peylin, P., Tarniewicz, J., Rigby, M., Rödenbeck, C., Vermeulen, A., and Walton, E. M.: The regional EUROpean atmospheric transport inversion COMparison, EUROCOM: first results on European wide terrestrial carbon fluxes for the period 2006\&ndash;2015, Atmospheric Chemistry and Physics Discussions, pp. 1–40, https://doi.org/https://doi.org/10.5194/acp-2019-1008, https://acp.copernicus.org/preprints/acp-2019-1008/, publisher: Copernicus GmbH, 2019.

Morzfeld, M., Hodyss, D., and Snyder, C.: What the collapse of the ensemble Kalman filter tells us about particle filters, Tellus A: Dynamic Meteorology and Oceanography, 69, 1283809, https://doi.org/10.1080/16000870.2017.1283809, https://doi.org/10.1080/16000870.2017.1283809, publisher: Taylor & Francis _eprint: https://doi.org/10.1080/16000870.2017.1283809, 2017.

Nassar, R., Hill, T. G., McLinden, C. A., Wunch, D., Jones, D. B. A., and Crisp, D.: Quantifying CO2 Emissions From Individual Power Plants From Space, Geophysical Research Letters, 44, 10,045–10,053, https://doi.org/10.1002/2017GL074702, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1002/2017GL074702, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1002/2017GL074702, 2017.

Niwa, Y., Tomita, H., Satoh, M., Imasu, R., Sawa, Y., Tsuboi, K., Matsueda, H., Machida, T., Sasakiwa, M., Belean, B., and Saigusa, N.: A 4D-Var inversion system based on the icosahedral grid model (NICAM-TM 4D-Var v1.0) – Part 1: Offline forward and adjoint transport models, Geoscientific Model Development, 10, 1157–1174, https://doi.org/https://doi.org/10.5194/gmd-10-1157-2017, https://gmd.copernicus.org/articles/10/1157/2017/gmd-10-1157-2017.html, publisher: Copernicus GmbH, 2017.

Palmer, P. I., Suntharalingam, P., Jones, D. B. A., Jacob, D. J., Streets, D. G., Fu, Q., Vay, S. A., and Sachse, G. W.: Using CO2:CO correlations to improve inverse analyses of carbon fluxes, Journal of Geophysical Research: Atmo-
spheres, 111, https://doi.org/10.1029/2005JD006697, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2005JD006697, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/2005JD006697, 2006.

Palmer, P. I., Feng, L., Baker, D., Chevallier, F., Bösch, H., and Somkuti, P.: Net carbon emissions from African biosphere dominate pan-tropical atmospheric CO$_2$ signal, Nature Communications, 10, 3344, https://doi.org/10.1038/s41467-019-11097-w, https://www.nature.com/articles/s41467-019-11097-w, number: 1 Publisher: Nature Publishing Group, 2019.

Peters, W., Miller, J. B., Whitaker, J., Denning, A. S., Hirsch, A., Krol, M. C., Zupanski, D., Bruhwiler, L., and Tans, P. P.: An ensemble data assimilation system to estimate CO$_2$ surface fluxes from atmospheric trace gas observations, Journal of Geophysical Research: Atmospheres, 110, https://doi.org/10.1029/2005JD006157, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2005JD006157, 2005.

Peters, W., van der Velde, I. R., van Schaik, E., Miller, J. B., Ciais, P., Duarte, H. F., van der Laan-Luijkx, I. T., van der Molen, M. K., Scholze, M., Schaefer, K., Vidale, P. L., Verhoef, A., Wårlind, D., Zhu, D., Tans, P. P., Vaughn, B., and White, J. W. C.: Increased water-use efficiency and reduced CO$_2$ uptake by plants during droughts at a continental scale, Nature Geoscience, 11, 744–748, https://doi.org/10.1038/s41561-018-0212-7, https://www.nature.com/articles/s41561-018-0212-7, number: 10 Publisher: Nature Publishing Group, 2018.

Peylin, P., Law, R. M., Gurney, K. R., Chevallier, F., Jacobson, A. R., Maki, T., Niwa, Y., Patra, P. K., Peters, W., Rayner, P. J., Rödenbeck, C., van der Laan-Luijkx, I. T., and Zhang, X.: Global atmospheric carbon budget: results from an ensemble of atmospheric CO$_2$ inversions, Biogeosciences, 10, 6699–6720, https://doi.org/10.5194/bg-10-6699-2013, http://www.biogeosciences.net/10/6699/2013/, 2013.

Pisso, I., Sollum, E., Grythe, H., Kristiansen, N. I., Cassiani, M., Eckhardt, S., Arnold, D., Morton, D., Thompson, R. L., Groot Zwaaitink, C. D., Evangelou, N., Sodemann, H., Haimberger, L., Henne, S., Brunner, D., Burkhart, J. F., Fouilloux, A., Brioude, J., Philipp, A., Seibert, P., and Stohl, A.: The Lagrangian particle dispersion model FLEX-PART version 10.4, Geoscientific Model Development, 12, 4955–4997, https://doi.org/https://doi.org/10.5194/gmd-12-4955-2019, https://gmd.copernicus.org/articles/12/4955/2019/, publisher: Copernicus GmbH, 2019.

Ramonet, M., Ciais, P., Apadula, F., Bartyzel, J., Bastos, A., Bergamaschi, P., Blanc, P. E., Brunner, D., Caracciolo di Torchioarlo, L., Calzolari, F., Chen, H., Chmura, L., Colomb, A., Conil, S., Cristofanelli, P., Cuevas, E., Curcoll, R., Delmotte, M., di Sarra, A., Emmenegger, L., Forster, G., Frumau, A., Gerbig, C., Gheusi, F., Hammer, S., Haszpra, L., Hatakka, J., Hazan, L., Heliasz, M., Henne, S., Hensen, A., Hermansen, O., Keronen, P., Kivi, R., Komínková, K., Kubistin, D., Laurent, O., Laurila, T., Lavric, J. V., Lehner, I., Lehtinen, K. E. J., Leskinen, A., Leuenberger, M., Levin, I., Lindauer, M., Lopez, M., Myhre, C. L., Mammarrella, I., Manca, G., Manning, A., Marek, M. V., Marklund, P., Martin, D., Meinhardt, F., Mihalopoulos, N., Mölder, M., Morgui, J. A., Necki, J., O’Doherty, S., O’Dowd, C., Ottosson, M., Philippon, C., Piacentino, S., Pichon, J. M., Plass-Duelmer, C., Resovsky, A., Rivier, L., Rodó, X., Sha, M. K., Scheeren, H. A., Sferlazzo, D., Spain, T. G., Stanley, K. M., Steinbacher, M., Trisolino, P., Vermeulen, A., Vítková, G., Weyrauch, D., Xueref-Remy, I., Yala, K., and Yver Kwok, C.: The fingerprint of the summer 2018 drought in Europe on ground-based atmospheric CO$_2$ measurements, Philosophical Transactions of the Royal Society B: Biological Sciences, 375, 20190513, https://doi.org/10.1098/rstb.2019.0513, https://royalsocietypublishing.org/doi/abs/10.1098/rstb.2019.0513, publisher: Royal Society, 2020.
Rayner, P. J., Raupach, M. R., Paget, M., Peylin, P., and Koffi, E.: A new global gridded data set of CO2 emissions from fossil fuel combustion: Methodology and evaluation, Journal of Geophysical Research: Atmospheres, 115, https://doi.org/10.1029/2009JD013439, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2009JD013439, _eprint: https://agupubs.onlinelibrary.wiley.com/doi/pdf/10.1029/2009JD013439, 2010.

Rayner, P. J., Michalak, A. M., and Chevallier, F.: Fundamentals of data assimilation applied to biogeochemistry, Atmospheric Chemistry and Physics, 19, 13 911–13 932, https://doi.org/https://doi.org/10.5194/acp-19-13911-2019, https://acp.copernicus.org/articles/19/13911/2019/, publisher: Copernicus GmbH, 2019.

Rödenbeck, C., Gerbig, C., Trusilova, K., and Heimann, M.: A two-step scheme for high-resolution regional atmospheric trace gas inversions based on independent models, Atmos. Chem. Phys, 9, 5331–5342, http://atmos-chem-phys.net/9/5331/2009/acp-9-5331-2009.pdf, 2009.

Saunois, M., Stavert, A. R., Poulter, B., Bousquet, P., Canadell, J. G., Jackson, R. B., Raymond, P. A., Dlugokencky, E. J., Houweling, S., Patra, P. K., Ciais, P., Arora, V. K., Bastviken, D., Bergamaschi, P., Blake, D. R., Brailsford, G., Bruhwiler, L., Carlson, K. M., Carrol, M., Castaldi, S., Chandra, N., Crevoisier, C., Crill, P. M., Covey, K., Curry, C. L., Etope, G., Frankenberg, C., Gedney, N., Hegglin, M. I., Höglund-Isaksson, L., Hugelius, G., Ishizawa, M., Ito, A., Janssens-Maenhout, G., Jensen, K. M., Joos, F., Kleinert, T., Krummel, P. B., Langenfelds, R. L., Laruelle, G. G., Liu, L., Machida, T., Maksyutov, S., McDonald, K. C., McNorton, J., Miller, P. A., Melton, J. R., Morino, I., Müller, J., Murguia-Flores, F., Naik, V., Niwa, Y., Noce, S., O’Doherty, S., Parker, R. J., Peng, C., Peng, S., Peters, G. P., Prigent, C., Prinn, R., Ramonet, M., Regnier, P., Riley, W. J., Rosentreter, J. A., Segers, A., Simpson, I. J., Shi, H., Smith, S. J., Steele, L. P., Thornton, B. F., Tian, H., Tohjima, Y., Tubiello, F. N., Tsuruta, A., Viovy, N., Voulgarakis, A., Weber, T. S., van Weele, M., van der Werf, G. R., Weiss, R. F., Worthy, D., Wunch, D., Yin, Y., Yoshida, Y., Zhang, W., Zhang, Z., Zhao, Y., Zheng, B., Zhu, Q., Zhu, Q., and Zhuang, Q.: The Global Methane Budget 2000–2017, Earth System Science Data, 12, 1561–1623, https://doi.org/https://doi.org/10.5194/essd-12-1561-2020, https://essd.copernicus.org/articles/12/1561/2020/, publisher: Copernicus GmbH, 2020.

Say, D., Manning, A. J., O’Doherty, S., Rigby, M., Young, D., and Grant, A.: Re-Evaluation of the UK’s HFC-134a Emissions Inventory Based on Atmospheric Observations, Environ. Sci. Technol., 50, 11 129–11 136, https://doi.org/10.1021/acs.est.6b03630, https://doi.org/10.1021/acs.est.6b03630, publisher: American Chemical Society, 2016.

Schuh, A. E., Jacobson, A. R., Basu, S., Weir, B., Baker, D., Bowman, K., Chevallier, F., Crowell, S., Davis, K. J., Deng, F., Denning, S., Feng, L., Jones, D., Liu, J., and Palmer, P. I.: Quantifying the Impact of Atmospheric Transport Uncertainty on CO2 Surface Flux Estimates, Global Biogeochemical Cycles, 33, 484–500, https://doi.org/https://doi.org/10.1029/2018GB006086, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2018GB006086, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/2018GB006086, 2019.

Seibert, P. and Frank, A.: Source-receptor matrix calculation with a Lagrangian particle dispersion model in backward mode, Atmospheric Chemistry and Physics, 4, 51–63, https://doi.org/10.5194/acp-4-51-2004, http://www.atmos-chem-phys.org/4/51/2004/acp-4-51-2004.html, 2004.

Stauffer, J., Broquet, G., Bréon, F.-M., Puygrenier, V., Chevallier, F., Xueref-Rémy, I., Dieudonné, E., Lopez, M., Schmidt, M., Ramonet, M., Perrussel, O., Lac, C., Wu, L., and Ciais, P.: The first 1-year-long estimate of the
Paris region fossil fuel CO2 emissions based on atmospheric inversion, Atmos. Chem. Phys., 16, 14 703–14 726, https://doi.org/10.5194/acp-16-14703-2016, http://www.atmos-chem-phys.net/16/14703/2016/, 2016.

Stohl, A., Seibert, P., Arduini, J., Eckhardt, S., Fraser, P., Greally, B. R., Lunder, C., Maione, M., Mühle, J., O’Doherty, S., Prinn, R. G., Reimann, S., Saito, T., Schmidbauer, N., Simmonds, P. G., Vollmer, M. K., Weiss, R. F., and Yokouchi, Y.: An analytical inversion method for determining regional and global emissions of greenhouse gases: Sensitivity studies and application to halocarbons, Atmos. Chem. Phys., 9, 1597–1620, https://doi.org/10.5194/acp-9-1597-2009, http://www.atmos-chem-phys.net/9/1597/2009/, 2009.

Tarantola, A. and Valette, B.: Generalized nonlinear inverse problems solved using the least squares criterion, Reviews of Geophysics, 20, 219–232, https://doi.org/https://doi.org/10.1029/RG020i002p00219, http://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/RG020i002p00219, _eprint: https://onlinelibrary.wiley.com/doi/pdf/10.1029/RG020i002p00219, 1982.

Thanwerdas, J., Saunois, M., Berchet, A., Pison, I., Vaughn, B. H., Michel, S. E., and Bousquet, P.: Variational inverse modelling within the Community Inversion Framework to assimilate $\delta^{13}$C(CH$_4$) and CH$_4$: a case study with model LMDz-SACS, Geoscientific Model Development Discussions, pp. 1–29, https://doi.org/10.5194/gmd-2021-106, https://gmd.copernicus.org/preprints/gmd-2021-106/#discussion, publisher: Copernicus GmbH, 2021.

Thompson, R. L. and Stohl, A.: FLEXINVERT: an atmospheric Bayesian inversion framework for determining surface fluxes of trace species using an optimized grid, Geosci. Model Dev. Discuss., 7, 3751–3801, https://doi.org/10.5194/gmd-7-3751-2014, http://www.geosci-model-dev-discuss.net/7/3751/2014/, 2014.

Tippett, M. K., Anderson, J. L., Bishop, C. H., Hamill, T. M., and Whitaker, J. S.: Ensemble Square Root Filters, Monthly Weather Review, 131, 1485–1490, https://doi.org/10.1175/1520-0493(2003)131<1485:ESRF>2.0.CO;2, https://journals.ametsoc.org/view/journals/mwre/131/7/1520-0493_2003_131_1485_esrf_2.0.co_2.xml, publisher: American Meteorological Society Section: Monthly Weather Review, 2003.

Tolk, L. F., Dolman, A. J., Meesters, A. G. C. A., and Peters, W.: A comparison of different inverse carbon flux estimation approaches for application on a regional domain, Atmos. Chem. Phys., 11, 10 349–10 365, https://doi.org/10.5194/acp-11-10349-2011, http://www.atmos-chem-phys.net/11/10349/2011/, 2011.

Trusilova, K., Rödenbeck, C., Gerbig, C., and Heimann, M.: Technical Note: A new coupled system for global-to-regional downscaling of CO2 concentration estimation, Atmos. Chem. Phys., 10, 3205–3213, https://doi.org/10.5194/acp-10-3205-2010, https://www.atmos-chem-phys.net/10/3205/2010/, 2010.

Turner, A. J. and Jacob, D. J.: Balancing aggregation and smoothing errors in inverse models, Atmos. Chem. Phys., 15, 7039–7048, https://doi.org/10.5194/acp-15-7039-2015, https://www.atmos-chem-phys.net/15/7039/2015/, 2015.

van der Laan-Luijkx, I. T., van der Velde, I. R., van der Veen, E., Tsuruta, A., Stanislawskas, K., Babenhauserheide, A., Zhang, H. F., Liu, Y., He, W., Chen, H., Masarie, K. A., Krol, M. C., and Peters, W.: The CarbonTracker Data Assimilation Shell (CTDAS) v1.0: implementation and global carbon balance 2001–2015, Geosci. Model Dev., 10, 2785–2800, https://doi.org/10.5194/gmd-10-2785-2017, https://www.geosci-model-dev.net/10/2785/2017/, 2017.

van der Velde, I. R., Miller, J. B., van der Molen, M. K., Tans, P. P., Vaughn, B. H., White, J. W. C., Schaefer, K., and Peters, W.: The CarbonTracker Data Assimilation System for CO$_2$ and $\delta^{13}$C (CTDAS-C13 v1.0): retrieving information on land–atmosphere exchange processes, Geoscientific Model Development, 11, 283–304, https://doi.org/10.5194/gmd-11-283-2018, https://www.geosci-model-dev.net/11/283/2018/, 2018.
Wang, F., Maksyutov, S., Tsuruta, A., Janardanan, R., Ito, A., Sasakawa, M., Machida, T., Morino, I., Yoshida, Y., Kaiser, J. W., Janssens-Maenhout, G., Dlugokencky, E. J., Mammarella, I., Lavric, J. V., and Matsunaga, T.: Methane Emission Estimates by the Global High-Resolution Inverse Model Using National Inventories, Remote Sensing, 11, 2489, https://doi.org/10.3390/rs11212489, https://www.mdpi.com/2072-4292/11/21/2489, 2019.

Wang, J. S., Kawa, S. R., Collatz, G. J., Sasakawa, M., Gatti, L. V., Machida, T., Liu, Y., and Manyin, M. E.: A global synthesis inversion analysis of recent variability in CO₂ fluxes using GOSAT and in situ observations, Atmospheric Chemistry and Physics, 18, 11 097–11 124, https://doi.org/https://doi.org/10.5194/acp-18-11097-2018, https://acp.copernicus.org/articles/18/11097/2018/, publisher: Copernicus GmbH, 2018.

Wang, Y., Broquet, G., Bréon, F.-M., Lespinas, F., Buchwitz, M., Reuter, M., Meijer, Y., Loescher, A., Janssens-Maenhout, G., Zheng, B., and Ciais, P.: PMIF v1.0: an inversion system to estimate the potential of satellite observations to monitor fossil fuel CO₂ emissions over the globe, Geoscientific Model Development Discussions, pp. 1–27, https://doi.org/https://doi.org/10.5194/gmd-2019-326, https://gmd.copernicus.org/preprints/gmd-2019-326/, publisher: Copernicus GmbH, 2020.

Whitaker, J. S. and Hamill, T. M.: Ensemble Data Assimilation without Perturbed Observations, Monthly Weather Review, 130, 1913–1924, https://doi.org/10.1175/1520-0493(2002)130<1913:EDAWPO>2.0.CO;2, https://journals.ametsoc.org/view/journals/mwre/130/7/1520-0493_2002_130_1913_edawpo_2.0.co_2.xml, publisher: American Meteorological Society Section: Monthly Weather Review, 2002.

Yadav, V. and Michalak, A. M.: Improving computational efficiency in large linear inverse problems: an example from carbon dioxide flux estimation, Geoscientific Model Development, 6, 583–590, https://doi.org/https://doi.org/10.5194/gmd-6-583-2013, https://gmd.copernicus.org/articles/6/583/2013/gmd-6-583-2013.html, publisher: Copernicus GmbH, 2013.

Zammit-Mangion, A., Cressie, N., Ganesan, A. L., O’Doherty, S., and Manning, A. J.: Spatio-temporal bivariate statistical models for atmospheric trace-gas inversion, Chemometrics and Intelligent Laboratory Systems, 149, Part B, 227–241, https://doi.org/10.1016/j.chemolab.2015.09.006, http://www.sciencedirect.com/science/article/pii/S0169743915002221, 2015.

Zammit-Mangion, A., Bertolacci, M., Fisher, J., Stavert, A., Rigby, M. L., Cao, Y., and Cressie, N.: WOMBAT: A fully Bayesian global flux-inversion framework, arXiv:2102.04004 [astro-ph, stat], http://arxiv.org/abs/2102.04004, arXiv: 2102.04004, 2021.

Zheng, T., French, N. H. F., and Baxter, M.: Development of the WRF-CO2 4D-Var assimilation system v1.0, Geoscientific Model Development, 11, 1725–1752, https://doi.org/https://doi.org/10.5194/gmd-11-1725-2018, https://gmd.copernicus.org/articles/11/1725/2018/, publisher: Copernicus GmbH, 2018.

Zupanski, D., Denning, A. S., Uliasz, M., Zupanski, M., Schuh, A. E., Rayner, P. J., Peters, W., and Corbin, K. D.: Carbon flux bias estimation employing Maximum Likelihood Ensemble Filter (MLEF), Journal of Geophysical Research: Atmospheres, 112, https://doi.org/10.1029/2006JD008371, https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2006JD008371, _eprint: https://agupubs.onlinelibrary.wiley.com/doi/pdf/10.1029/2006JD008371, 2007.

Zupanski, M.: Maximum Likelihood Ensemble Filter: Theoretical Aspects, Mon. Wea. Rev., 133, 1710–1726, https://doi.org/10.1175/MWR2946.1, https://journals.ametsoc.org/mwr/article/133/6/1710/67374/Maximum-Likelihood-Ensemble-Filter-Theoretical, publisher: American Meteorological Society, 2005.
**Table 1.** Elementary operations required for each data assimilation method. An = Analytical inversion; EnKF = Ensemble Kalman filter; Var = Variational; Fwd = forward simulation; AdTest = Test of the adjoint. We note \( \mathcal{X} \) and \( \mathcal{Y} \) the target and observation spaces respectively, \( \mathfrak{A} \) the regularization space in the minimization algorithm of variational inversions; the (·)* symbol depicts the adjoint of corresponding spaces.

| Elementary operation                          | Mathematical formulation | Data assimilation |
|-----------------------------------------------|--------------------------|-------------------|
| Forward observation operator                   | \( \mathcal{X} \rightarrow \mathcal{Y} \) | An X Var Fwd      |
|                                               | \( \mathbf{x} \rightarrow \mathcal{H} (\mathbf{x}) \) or \( \mathbf{Hx} \) |                 |
| Adjoint observation operator                   | \( \mathcal{Y}^* \rightarrow \mathfrak{A}^* \) | An X          |
|                                               | \( \delta \mathbf{y} \rightarrow \mathcal{H}^* (\delta \mathbf{y}) \) or \( \mathbf{HT} \delta \mathbf{y} \) | X               |
| Normalisation of the observation increments    | \( \mathcal{Y}^* \rightarrow \mathcal{Y}^* \) | X               |
|                                               | \( \delta \mathbf{y} \rightarrow \mathbf{R}^{-1} \delta \mathbf{y} \) |                 |
| Regularization of the control target space    | \( \mathfrak{A} \rightarrow \mathcal{X} \) | X               |
|                                               | \( \mathbf{x} \rightarrow \mathbf{x} = \mathbf{B}^{1/2} \mathbf{\chi} + \mathbf{x}^b \) |                 |
| Adjoint of the control target space regularization | \( \mathfrak{A}^* \rightarrow \mathcal{X}^* \) | X               |
|                                               | \( \delta \mathbf{x} \rightarrow \delta \mathbf{\chi} \equiv \mathbf{B}^{1/2} \delta \mathbf{x} \) |                 |
| Control Target space sampling                  | \( \mathcal{X} \times \mathcal{X}^2 \rightarrow \mathcal{Y} \) | X               |
|                                               | \( (\mathbf{x}, \mathbf{P}) \rightarrow (\mathbf{x}_1, \mathbf{x}_2, \ldots, \mathbf{x}_N) \) |                 |
|                                               | \( \mathcal{X} \times \mathcal{X}^2 \rightarrow \mathcal{X}^N \) |                 |
|                                               | \( (\mathbf{x}, \mathbf{B}) \rightarrow (\mathbf{x}_1, \mathbf{x}_2, \ldots, \mathbf{x}_N) \) |                 |

43
Figure 1. Call structure of the CIF.

Figure 2. Observation operator structure. Emissions are processed from the target vector to generate model inputs, as well as other inputs, no optimized by the inversion; in this example, some background for the simulations is also optimized by the inversion and is added to simulations at the end of the pipeline just before stacking outputs to the observation vector format.
**Community Inversion Framework**

- **Central gateway**
  - community-inversion.eu/
  - installation instructions
  - tutorials
  - set-up examples

- **GitLab server**
  - git.nilu.no/VERIFY/CIF

**pyCIF**
- python library
- driver for CTMs
- carries out all processing operations

**CTMs**
- FORTRAN sources for executing CTMs

**Tests**
- Test and demonstration cases

**Figure 3.** Organisation of the Community Inversion Framework.

---

**Figure 4.** (left) True Prior fluxes and observation sites. (right) Difference between prior fluxes and Perturbation from the truth prior used to generate "true" observations. Observation sites are shown as circles coloured according to their height in meters above ground level (m a.g.l.). Fluxes are reported in arbitrary units (a.u.)
Figure 5. Posterior increments for analytical, EnKF, EnSRF, variational with CONGRAD and variational with M1QN3 (from top to bottom, left to right) for an inversion set-up at the pixel resolution with aggregated bands horizontal correlation length of 500 m.

Figure 6. Same Uncertainty reduction for analytical, EnSRF, variational with CONGRAD and variational with M1QN3 (from top to bottom, left to right) as in Fig. ?? with an inversion set-up at the pixel resolution with horizontal correlation length of 500 m.
Figure 7. Same Prior (left) and posterior (middle and right) uncertainty matrices for analytical, EnSRF, variational with CONGRAD and variational with M1QN3 (from top to bottom, middle and right columns), with the same set-up as in Fig. ?? 5. All matrices are reported with an inversion set-up at unitless values, i.e., a "1" on the pixel resolution with horizontal correlation length of 10000 m diagonal corresponds to a 100% uncertainty.
Figure 8. Sensitivity of the inversions to the number of computed simulations for analytical (red), EnSRF (green), variational with CONGRAD (orange) and variational with M1QN3 (blue) methods. (top) inversion set-up with aggregated regions of 3 pixels × 3 pixels; (middle) inversion set-up at the pixel resolution with horizontal correlation length of 500 m; (bottom) inversion set-up at the pixel resolution with horizontal correlation length of 10000 m. For each sub-panel: (top) Distance between the true and posterior fluxes for a given number of simulations; (bottom) Evolution of the cost function depending on the number of simulations.