First Spatial Reconstruction of Past Fires in Temperate Europe Suggests Large Variability of Fire Sizes and an Important Role of Human-Related Ignitions

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The spatial component of past forest fires in temperate Europe has been little studied, despite the value of such data in quantifying human and natural factors driving fire activity and associated forest dynamics. Changes in fire regimes reported across a range of ecosystems call for a better understanding of variability in historic fires and may help define reference points that can be relied upon when discussing climate change effects. We provide the first dendrochronological reconstruction of historical fire sizes in Central Europe and analyze the minimum extent of fires during the last four centuries in a 9.2 km² (920 ha) conifer-dominated section of Białowieża Forest, one of the largest continuous lowland forests of the subcontinent. We recorded 82 fires between 1666 and 1946, using 275 sample trees, while 92% of fires (76 out of 82) spread beyond the studied area. Fires varied considerably in size, from events recorded at only one site (1–200 ha) to fires recorded in more than half of the studied area, thus exceeding 500 ha in size. The fire cycle was 11 years over the whole study period, with three distinct periods revealed by the regime shift analysis. In the years 1670–1750, the fire cycle averaged 12 years. It shortened to 7 years between 1755–1840 and increased to 22 years over the 1845–1955 period. In comparison with present day data, the reconstructed fire density of 3.2 fires per 100 km² (10 000 ha) and year exceeded lightning ignition density by one to two orders of magnitude, suggesting a significant contribution of human-related ignitions. Our results highlight the important role of fire disturbance in Białowieża Forest and provide critical baseline information to design biological conservation strategies for European forests.

Keywords: dendrochronological reconstruction, disturbance regime, fire history, land use history, lowland forests, Pinus sylvestris
INTRODUCTION

Spatial patterns of fire disturbance is a critical feature of fire regimes, controlling their ecological and biochemical impacts at the stand, landscape, and regional scales (Niklasson and Granström, 2000; Beatty and Taylor, 2008). Climatic conditions, topography, and forest structure control the abundance, spatial distribution, and water content of forest fuels. Together with ignition patterns, these factors determine fire occurrence, spread, and severity (Agee, 1993; Taylor and Skinner, 1998; Beatty and Taylor, 2001). Human activities (land use in particular) significantly impact both fuels and ignitions, modifying the natural fire activity (Niklasson and Granström, 2000; Granström and Niklasson, 2008).

Recent changes in fire regimes documented across a range of forest ecosystems (Pausas and Keeley, 2009; Drobyshev et al., 2017) call for a better understanding of factors shaping fire activity within specific regions. In Europe, the number of tree ring studies on such detailed fire regime characteristics like frequency, seasonality, and severity of fires is increasing constantly (Lehtonen et al., 1996; Blanck et al., 2013; Drobyshev et al., 2014; Zin et al., 2015). However, spatially explicit reconstructions are few. They predominantly cover the northerly and southerly sections of European forests (Niklasson and Granström, 2000; Christopoulou et al., 2013; Storaunet et al., 2013; Rolstad et al., 2017; Kitenberga et al., 2019; Ryzhkova et al., 2020; cf. Szymczak et al., 2020). To our knowledge, such studies are missing in the case of temperate Europe.

Spatial tree ring reconstruction of fire occurrence combines information on the fire dates and the exact location of the dated wood samples. The presence of dendrochronologically datable material (dead and living fire-scarred wood) remains one of the main challenges in expanding the network of spatially explicit fire reconstructions. In temperate Europe, the scarcity of old-growth forests and limited availability of old trees and deadwood with fire scars reflects a legacy of long and intensive forest use (Pyne, 1997; Kaplan et al., 2009). In contrast to some of the North American forests, dominated by short-needled conifers existing studies and the knowledge on historical variability of fire regimes in temperate Europe remains limited (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020), especially in regard to fire sizes.

Białowieża Forest is one of the few existing lowland old-growth woodlands in Europe with primary character (Sabatini et al., 2018). It is regarded as a reference landscape for the region (Faliński, 1986; Ellenberg, 1996). Insufficient empirical data on past fire disturbances, a negligible influence of lightning ignition upon modern fire activity (Ubysz and Szczygieł, 2006; Szczygieł et al., 2016; Brackhane et al., 2021), and consideration of European temperate zone as deciduous-dominated and less flammable domain (Ellenberg, 1996; Leuschner and Ellenberg, 2017b) have contributed to the view upon natural fires as very rare events in this region (Leuschner and Ellenberg, 2017a). In fact, a large part of the Central European temperate zone is dominated by flammable conifer forests (Brus et al., 2012; San-Miguel-Ayanz et al., 2021). However, long-term and efficient fire suppression (Granström and Niklasson, 2008; Castellnou et al., 2010; Wallenius, 2011) and high human ignition frequencies (Szczygieł et al., 2016; San-Miguel-Ayanz et al., 2021) resulted in the notion of fire being largely driven by humans. The role of fire as an important driver of natural forest dynamics in this part of Europe has, therefore, been generally disregarded.

Paleoecological data from Central Europe suggested fire as one of the factors affecting vegetation development during the Holocene, long before the rise of civilizations based on agriculture (Rösch, 2000). These reconstructions have promoted discussion on the openness of European forests (Vera, 2000; Svenning, 2002; Mitchell, 2005; Bobek et al., 2018; Feurdean et al., 2020) and the human impact on vegetation (Bradshaw and Sykes, 2014; Leuschner and Ellenberg, 2017a). The development gave momentum to the debate of various disturbances as drivers of ecosystem dynamics in the continent. Several paleoecological studies have documented fires in the European temperate forests in the past (Tinner et al., 2005; Novák et al., 2012; Adámek et al., 2015; Latalowa et al., 2015, 2016; Zimny et al., 2017; Feurdean et al., 2020). Although forming very solid evidence of biomass combustion in general, these charcoal-based records provided little insight into specific characteristics of fire regimes, such as fire seasonality, frequency, and fire size. Taphonomy, sediment mixing, diversity of deposition pathways, and landscape morphologies add to the complexity in the reconstruction of the fire regime with paleoecological data (Tolonen, 1986; Clark, 1988; Clark et al., 1998; Ohlson and Tryterud, 2000). Dendrochronological reconstructions may offer detailed data on fire regime characteristics, yet the number of existing studies and the knowledge on historical variability of fire regimes in temperate Europe remains limited (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020), especially in regard to fire sizes.

Nomenclature: Mirek et al. (2002) for vascular plants; Ochyra et al. (2003) for bryophytes.
dendroecological research, which is unique for the Central European region (Niklasson et al., 2010, and literature therein).

Fire-related dendrochronological studies in Białowieża Forest have focused on fire occurrence, stand-level return intervals, fire severity, and the role of fire in shaping tree demography (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020). Paleocological reconstructions using charcoal data from peat bogs have documented the presence of fire over millennia (Dąbrowski, 1959; Mitchell and Cole, 1998; Latalowa et al., 2015, 2016; Zimny et al., 2017). The current study aims to expand the knowledge of fire activity in Białowieża Forest through reconstruction of the minimum sizes of the past fires in the conifer section of this area. We hypothesized that (H1) fires typically exceeded the size of single stands (c. 10^3 ha), (H2) fires were of highly variable spatial extent, and (H3) historical fires occurred at frequencies exceeding frequency of lightning-caused ignitions.

**MATERIALS AND METHODS**

**Study Area**

Białowieża Forest, recognized as one of the best-preserved temperate forests on the European Lowland, extends over approximately 1,500 km² of the borderland between eastern Poland and western Belarus (52°30′–53°N, 23°30′–24°15′E). The area has a warm-summer humid continental climate (Kottek et al., 2006; cf. Jaroszewicz et al., 2019). During the 1955–2001 period, the mean annual precipitation was 633 mm and mean annual temperature was 6.8°C. In the same period, January was the coldest month with a mean temperature of −4.2°C and July was the warmest month with a mean temperature of 17.7°C (Pierzgalski et al., 2002).

The topography of the area is homogeneous and features an old-morainic plateau (135–202 m a.s.l.) which is built from glaciofluvial gravels, clays, and sands. The terrain is locally diversified by small river valleys, melt-out hollows with accumulated peat cover, and an elevated zone of glaciofluvial-kame origin that stretches throughout the central part of the forest and ends at the study area’s highest point – Kozia Góra located in Belarus. The main soil types of Białowieża Forest are oligotrophic podzol and rusty soils, eutrophic and mesotrophic brown and lessive soils, and a range of gley and organic soil types, including peatbog and semi-bog soils (Kwiatkowski, 1994).

The area is characterized by a mosaic of different forest types. The common tree species are pedunculate oak (Quercus robur L.), small-leaved lime (Tilia cordata Mill.), Norway maple (Acer platanoides L.), hornbeam (Carpinus betulus L.), silver birch (Betula pendula Roth.), black alder (Alnus glutinosa (L.) Gaertn.), ash (Fraxinus excelsior L.), and downy birch (Betula pubescens Ehrh.). Among conifer species, Scots pine (Pinus sylvestris L.), and Norway spruce (Picea abies (L.) H. Karst.) dominate. Białowieża Forest is an area with very high habitat diversity, including mesic and mesotrophic habitats typical for deciduous and mixed deciduous forests, oligotrophic habitats of conifer communities, as well as bog alderwoods and streamside ash-alder forests as water-shaped forest types (Faliński, 1986; Jędrzejewska and Jędrzejewski, 1998; Sokolowski, 2004; Jaroszewicz et al., 2019).

Today, approximately 44% of the total forest area is covered by old-growth naturally regenerated stands (Jędrzejewska and Jędrzejewski, 1998; Jaroszewicz et al., 2019). Białowieża Forest encompasses forest stands that are both strictly protected (c. 53%) and managed (c. 47%). Various nature conservation, forest management, and land use regimes are in place, some of which include commercial timber production (Sokołowski, 2004; Krzyściak-Kosińska et al., 2012; Jaroszewicz et al., 2019). The Polish part of Białowieża Forest (approximately 635 km²) consists of the Białowieża National Park (c. 105 km²) and the adjacent managed forest (c. 500 km²), including several smaller nature reserves encompassing approximately 120 km² in total. The whole Belarusian section of Białowieża Forest is approximately 875 km², forming Belovezhskaya Pushcha National Park (Sokołowski, 2004; Jaroszewicz et al., 2019).

Twelve sites were used to study the spatial extent of fires in the conifer section of Białowieża Forest during the last four centuries (Table 1). Five of them (P1–P5) were the subject of the stand-scale fire history reconstructions (Zin et al., 2015;...
Zin et al. (2016). They were all mixed coniferous (P. sylvestris – P. abies) stands with tree density of approximately 700–800 stems·ha⁻¹. Additional seven sites (S1–S7) were situated in neighboring conifer forest on flat terrain and lacked potential fire breaks (streams, bogs, swamps). Only one site (P3) was separated from the other by a small river valley (Figure 1).

Ground vegetation at all twelve study sites was composed of Vaccinium myrtillus L., V. vitis-idaea L., Trisetum europaeum L., Dryopteris carthusiana (Vill.) H.P. Fuchs and Calamagrostis arundinacea (L.) Roth, and the ground layer by mosses: Pleurozium schreberi (Willd. ex Brid.) Mitt., Ptilium crista-castrensis (Hedw.) De Not., Hylocomium splendens (Hedw.) Schimp., and Dicranum undulatum Schrad. ex Brid. The surrounding stands were mostly conifer forests of analogous type, or humid conifer forests (with some share of Sphagnum spp. in the ground layer) and, occasionally, mixed forests with a higher share of deciduous tree species, like oak, birch, and hornbeam.

Study Design and Field Sampling

There is no previous knowledge from Białowieża Forest nor from temperate Europe that could indicate an appropriate size of the area to be sampled for the study of fire sizes. Earlier stand-scale fire history studies in the area covered 13 ha (Niklasson et al., 2010), 8.5 ha (Zin et al., 2015), and 43 ha (Spînu et al., 2020). It was not possible to delineate any fires in these areas with certainty, suggesting that fires had operated on a scale exceeding the size of those study areas.

Our earlier work in Białowieża Forest has shown a highly uneven distribution of datable wood material under the c. 1 km² (100 ha) scale (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020). We, therefore, designed a grid of nine study sites aiming at a resolution of c. 1 km² (100 ha). Collectively, the grid covered the area of 9.2 km² (920 ha) and represented an area of a size one–two magnitudes higher than the previous stand-scale studies. In an effort to find some indication of the potential spread of very large fires (reaching c. 10⁴ ha in size), we included in this study three additional satellite study sites (P3–P5) located outside the studied area of 9.2 km² (920 ha) at the distance of 3 and 6.5 km. The resulting network of twelve study sites (Table 1) was expected to represent four magnitudes of the spatial fire occurrence: 10¹, 10², 10³, and 10⁴ ha (Figure 1).

In the study sites, we collected partial or full cross-sections from all available Scots pine stumps, logs and snags with fire scars with a chain saw, following the procedure of Arno and Sneck (1977) and McBride (1983). We also sampled all old living pines with fire scars (if such trees were present) and individuals with lowest DBH (diameter at breast height, 1.3 m above the ground) considering them as probable representatives of the youngest populations of that species (Zin et al., 2015). Living trees were sampled with an increment borer. During the field sampling, we did not note fire scars on any tree taxa other than Scots pine.

On each of the seven study sites (S1–S7), we sampled between four and 26 trees, on average 15 trees, with 104 trees in total. Along with the two previously sampled sites (P1 and P2), the
total number of trees that were used for fire reconstructions in the studied area of 9.2 km² (920 ha) reached 275. The comparison satellite sites (P3, P4, and P5) contained an additional 169 sample trees altogether (Table 1).

**Laboratory Procedure**

All wood samples were sanded to enable wood cell visibility under a dissecting microscope with 6–40 × magnification. The tree ring material was visually cross-dated by identifying local pointer years according to standard dendrochronological methodology (Stokes and Smiley, 1968; Yamaguchi, 1991). This allowed us to identify the exact years of fire occurrence. A fire was recorded in a given year and site, if we found at least one fully developed fire scar among all trees sampled in that site (Zin et al., 2015).

**Spatial Fire Reconstruction**

We assessed the extent of past fires by developing GIS-assisted reconstruction of the minimum fire areas by analyzing the spatial distribution of study sites, which were either affected (i.e., burned) or not affected (i.e., unburned) by fire in a given year (cf. Niklasson and Granström, 2000) among the population of recording sites. A recording site was a site with tree ring data covering the year in question.

As a burned site for a year, we regarded a study site with at least one sample tree with a fire scar in that year. This might potentially inflate our estimates of burned area as portions of the sites might have remained unburned. However, this bias is believed to be less of a concern for fires recorded by multiple trees at the same site and by multiple sites. The situation with 1- and 2-year fire intervals recorded in the same tree in Białowieża Forest (cf. Zin et al., 2015) might be an evidence of a very quick fuel build-up. This suggests that fuel amounts in multiple trees at the same site and by multiple sites. The situation with 1- and 2-year fire intervals recorded in the same tree in Białowieża Forest (cf. Zin et al., 2015) might be an evidence of a very quick fuel build-up. This suggests that fuel amounts in multiple sites, varied from one site up to eight, with the majority (42

\[
FC = \frac{TSA \times PS}{TBA}
\]

where: FC is the fire cycle (years), TSA is the total studied area (ha), PS is the time period studied (years) and TBA is the total burned area over this time period (ha). To obtain 10 and 90% confidence limits for the calculated fire cycle values, we resampled with replacement the pool of fire years 1000 times (Ryzhkova et al., 2020).

To study temporal changes in the fire cycle, we applied a regime shift detection algorithm (Rodionov, 2004). It uses sequential t-tests, which diagnose regime change when the cumulative sum of normalized deviations from the mean of a new regime differs from the mean value of the current regime, calculated on a predefined moving time frame. The algorithm is fed with a threshold value, below which the capability of the algorithm to determine a regime change declines (Rodionov, 2004). To increase the ability of the algorithm to detect the short-term changes in fire activity, we set this threshold to 10 years. We also designated the significance level of t-tests to 0.05 and the Huber’s weight parameter, which controlled for weights assigned to the outliers, to 1.

**RESULTS**

During the 1666–1946 period, we recorded 82 individual fires that included nine cases of two spatially separated fires dated in the same year within the studied area of 9.2 km² (920 ha). The earliest fire in our analyses (1666) marked the onset of the period when the number of recording sites (i.e., study sites with tree ring chronology) exceeded 50% of the total number of study sites. We documented nine fire years (1702, 1718, 1731, 1732, 1742, 1748, 1828, 1838, 1882) when burned sites were clearly dated to the year in question by the number of recording sites for that year. We then used the following formula:

As an unburned site, we considered a recording site without fire scars dated to the year in question. We drew fire borders halfway between burned and unburned study sites. In areas lacking recording sites, the border was drawn to circumscribe the burned site at the same distance as the shortest of the halfway distances to the existing unburned sites. To develop spatial reconstruction of the minimum fire areas we used QGIS software (v. 2.8.1-Wien; QGIS Development Team).

We calculated fire cycle as the time required to burn an area equivalent to the study area (Reed et al., 1998). To obtain area estimates of burned areas, we calculated the proportion of study sites burned in each fire and reconstructed minimum fire areas as described previously. To this end, we divided the number of study sites which actually recorded a fire (i.e., were burned) in the year in question by the number of recording sites for that year. We then used the following formula:

\[
FC = \frac{TSA \times PS}{TBA}
\]

where: FC is the fire cycle (years), TSA is the total studied area (ha), PS is the time period studied (years) and TBA is the total burned area over this time period (ha). To obtain 10 and 90% confidence limits for the calculated fire cycle values, we resampled with replacement the pool of fire years 1000 times (Ryzhkova et al., 2020).

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The spatial extent of fires, expressed by the number of burned sites, varied from one site up to eight, with the majority (42

![FIGURE 2](https://www.frontiersin.org/articles/10.3389/fec.2015.00170/data-supplementary/1.32863/27586001-sup-fig1.png)

**FIGURE 2** | Distribution of fire sizes approximated by the number of sites burned during a particular fire within the study area.
the grid, thus indicating a fire size of up to 200 ha. 

**Figure 3**

An example of a one-site fire in 1755, with the reconstructed area of 188.3 ha (black dashed line) within the 9.2 km² (920 ha) study area (black solid line). White circles: burned trees, i.e., sample trees which recorded a fire (i.e., had a fire scar) in the year in question, black dots: unburned trees, i.e., sample trees with tree ring chronology covering the year in question, gray solid line: watercourse, dark gray dashed line: forest compartments.

of the 82 analyzed) dated at one site (**Figure 2**). Only six of the fires, all one-site fires, were reasonably contained inside the grid, thus indicating a fire size of up to 200 ha. **Figure 3** gives an example of such a fire in 1755. The remaining 76 fires had one or more open borders, their calculated sizes to be regarded as minimum areas burned. By using the GIS-assisted reconstruction, we obtained minimum fire area values that varied from 33 up to c. 800 ha (**Figure 4**). The three largest fires within the studied area occurred in 1795, 1809, and 1825 (**Figure 5**), with estimated minimum areas of 700–900 ha. Out of those, the 1795 fire was also documented at the P3 site and the 1809 fire at all of the satellite sites (P3–P5) outside the studied area. In total, eleven of the dated fire years were recorded at the adjacent P3 site c. 3 km from the SE border of the studied area. Twelve of the dated fire years were recorded at the P4 and P5 sites located 6.5 km to the W (**Supplementary Figure 1**).

The cumulative number of fires showed some distinct periods of abrupt increases and relaxations in fire activity (around 1718, 1775, 1803, and 1838) and exhibited a decrease in fire occurrence which started in the mid-1800s (**Figure 6**).

The historical fire cycle varied in time. It was 11 years during the 1666–1946 period, the 90% confidence envelope being 8 to 50 years. During 1670–1750, the fire cycle was 12 years (with 90% confidence envelope of 10 to 16 years). During 1755–1840, it shortened to 7 years (with the confidence envelope of 6 to 8 years). During 1845–1955, fire activity declined with the fire cycle being 22 years (with 90% confidence envelope of 15 to 40 years) (**Figure 7**).

Over the studied area and the whole period analyzed (1666–1946), the average fire return interval was 3.4 years. This corresponds to 3.2 ignitions per 100 km² (10 000 ha) and year. The only fire broadly known from historical sources – that of 1811 – was confirmed in our analyzes, burning over half of the studied sites (**Supplementary Figure 1**).

**DISCUSSION**

We present the first spatial reconstruction of the forest fire regime in temperate Europe. It adds a new dimension to the understanding of the long-term forest fire dynamics in this subcontinent. Fires have undoubtedly been inferred over most of the Holocene in the region and in Bialowieza Forest from the analyses of charcoal remains (Dąbrowski, 1959; Mitchell and Cole, 1998; Rösch, 2000; Tinner et al., 2005; Novák et al., 2012; Adámek et al., 2015; Latalowa et al., 2015, 2016; Zimny et al., 2017), with several paleoecological studies particularly discussing their role in shaping Scots pine-dominated forests (Novák et al., 2012; Adámek et al., 2015). However, the knowledge on spatial characteristics of fires and the link up to the present day high-resolution fire monitoring has been lacking. Existing tree ring data from Bialowieza Forest produced firm evidence of the key role of fire over the last ~400 years both in conifer and deciduous communities (Niklasson et al., 2010; Zin et al., 2015; Spinu et al., 2020). It suggested that fires were possibly a landscape-scale phenomena, affecting a wide spectrum of habitats, from xeric to mesic ones, as has been documented for the temperate region of North America (Flatley et al., 2015). However, these data gave little information on fire sizes, spatial extent, and perhaps most importantly, the ignition densities.

**Fire Extent**

Our reconstruction supported both (H1) and (H2): fire size typically exceeded stand scale (c. 101 ha) and fires varied strongly in size. Small fires (up to c. 200 ha) dominated the reconstructed period, with a few events attaining large sizes – a pattern that has earlier been revealed in other forest fire reconstructions (Niklasson and Granström, 2000; Heyerdahl et al., 2001; Storaunet et al., 2013; Rolstad et al., 2017). A number of features of the studied landscape apparently favoured fire spread: the area is dominated by generally flat or very gently undulating topography, with few depressions and swampy areas, and lacks lakes (Kwiatkowski, 1994). Such landscape evidently has the potential for large fires to develop, as evidenced by the fires in 1795, 1809, and 1825, recorded at most of the study sites (**Figure 5**), including sites well outside the studied area. The fire in 1809 seems to be an extreme case, since it was documented at all of the satellite sites (P3–P5), located at the distance of c. 3 km to the E (P3) and c. 6.5 km to the W (P4–P5) (**Supplementary Figure 1**).

Given that most of the conifer-dominated landscapes of Bialowieza Forest have similar topographical features with few wetlands, rivers, and lakes (Kwiatkowski, 1994), we suspect that large fires have been common all over the Bialowieza Forest area.

For establishing more firmly maximum fire sizes, we estimate that the sampling effort and aerial extent needs to be approximately 5–10 times larger than this study. Such sampling should preferably also include more mesic, deciduous-dominated parts of the landscape, where past fire occurrence has been already confirmed (Spinu et al., 2020).
The 1811 fire dated on half of the study sites was actually the only fire broadly known from historical sources (Ronke, 1830; Genko, 1902–1903; Faliński, 1986). We documented other fires, like those of 1803 or 1809, that seemed to be similarly widespread (Figure 5; Supplementary Figure 1). Clearly, there is a very strong discrepancy between the fire record derived from written sources and our tree ring reconstructed fires (cf. Niklasson et al., 2010). The reasons behind this remain unclear. One of the explanations could be possible gaps in the historical source material (Samojlik et al., 2016), even though the amount of archival data on the environmental history of Bialowieża Forest is growing constantly (Samojlik, 2010; Samojlik et al., 2020). An alternative explanation is that fires were so common and not worth mentioning, unless humans were negatively affected in some way, as has been demonstrated in a recent study of anthropogenic fire regimes in North America (McClain et al., 2021).

Since so few fires were clearly contained within our studied area, we suggest that the fire cycle is a robust measure to characterize the past fire activity.

**Fire Activity**

The dynamics of reconstructed fire activity corresponds broadly to changes in land use in Bialowieża Forest.

During the 1670–1750 period, the fire cycle was 12 years (Figure 7), being one of the shortest cycles reported from conifer-dominated European forests (Storaumnet et al., 2013; Rolstad et al., 2017; Kitenberga et al., 2019; Pinto et al., 2020;
Ryzhkova et al., 2020). During the second half of the 1600s, a series of uprisings and wars were responsible for a decline in human population of the Białowieża Forest area and an increased demand for additional income to the royal treasury. Tar and potash production were introduced (Hedemann, 1939; Samojlik, 2010), activities that involved use of fire (Brincken, 1826; Samojlik, 2016) and likely promoted fire occurrence (Niklasson et al., 2010). Traditional forest beekeeping, cattle pasturing, and collecting of the resinous Scots pine wood for kindling (that was chopped off from the bottom of the scorched tree trunks), all activities associated with the use of fire (Brincken, 1826; Genko, 1902–1903; Samojlik et al., 2020), were common in the area at least since the 1500s (Hedemann, 1939; Samojlik et al., 2016, 2019).

The period of 1755–1840 marked an even shorter fire cycle of 7 years (Figure 7). Along with the fire-related activities observed in the previous period, charcoal burning was introduced during the second half of the 1700s (Hedemann, 1939; Samojlik, 2010), which may have further increased fire occurrence (Niklasson et al., 2010). This period of high fire activity is corroborated with other data. Stand-scale fire return intervals of 5 years were documented in Białowieża Forest during the 1700s (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020) and large influxes of charcoal and pollen of fire-associated plants were recorded in sediments during the 17th and 18th centuries (Latałowa et al., 2015, 2016; Zimny et al., 2017).

During the 1845–1955 period, the fire cycle increased to 22 years (Figure 7). This period encompassed the 1946 fire, the last fire on the studied sites (Figure 6). We propose that changes in the land management following incorporation of the area into the Russian Empire in 1795 were responsible for changes in the fire dynamics. Traditional forest beekeeping, tar, potash and charcoal production were progressively eliminated starting in the late 1700s (Brincken, 1826; Samojlik et al., 2020). In the
case of forest beekeeping, the process of removal was gradual, supported by the government-ordered transition to apiaries and finalized after 1888, when Białowieża Forest became part of the Tsar’s private properties (Samojlik et al., 2019). Acquisition of the resinous wood chips from pine trees was prohibited in the 1840s and several measures were undertaken to eliminate that practice. However, they were unsuccessful and this activity lasted until the 20th century (Samojlik et al., 2019). Fire prevention in Białowieża Forest started already in the late 18th century (Hedemann, 1939), but it became rigorous during the 1800s (Genko, 1902–1903), as evidenced by the archival documents from the 19th century which mention several regulations concerning fire control (Samojlik et al., 2020, p. 92–93). Stand-scale fire histories in the area show decline in fire activity during the 1800s (Niklasson et al., 2010; Zin et al., 2015; Spînu et al., 2020). Paleoecological records confirm a substantial decrease in fire disturbance in the 19th century (Latalowa et al., 2015, 2016; Zimny et al., 2017). Like in many areas worldwide (Gränström and Niklasson, 2008; Wallenius, 2011; Rolstad et al., 2017; Ryzhkova et al., 2020), active fire suppression in Białowieża Forest was strongly linked to the increasing interest in modern forest management, focusing on timber production (Genko, 1902–1903; Wińcko, 1984; cf. McGrath et al., 2015; Samojlik et al., 2020).

The decline in fire activity in Białowieża Forest occurred approximately 50 to 150 years later than in Northern Europe, where fires disappeared in the second half of the 1700s and in the early 1800s (Niklasson and Granström, 2000; Niklasson and Drakenberg, 2001; Wallenius, 2011; Drobyshev et al., 2014; Pinto et al., 2020). However, in the eastern section of the northern boreal forests fire decline occurred only in the late 19th century (Ryzhkova et al., 2020). Some areas escaped effective fire suppression policies, even until the middle of the 20th century (Kitenberga et al., 2019). At present, fire suppression is efficient and the fire cycle of Białowieża Forest equals c. 26 000 years, with an average fire size of c. 0.5 ha (Szczygieł et al., 2016). Data for the broader region reflect a similar picture with current fire cycles of several thousand years (Szczygieł et al., 2009; Drobyshev et al., 2021b; San-Miguel-Ayanz et al., 2021).

Paleoecological records support the notion of fire as a human-mediated phenomenon over most of temperate Europe. Charcoal influx in sediments from forests in the region has been shown to correlate with anthropogenic indicators, such as pollen of cultivated plants or spores of coprophilous fungi (Bradshaw and Lindbladh, 2005; Tinner et al., 2005; Pędziszewska and Latalowa, 2016; Czerwiński et al., 2021). Well-preserved old-growth woodlands with primary character such as Białowieża Forest also show signs of long-term human impact (Samojlik et al., 2013; Latalowa et al., 2015, 2016; Leuschner and Ellenberg, 2017a; Zimny et al., 2017; Jaroszewicz et al., 2019).

Climate and human land use might have jointly affected past fire activity in Białowieża Forest (cf. Bradshaw and Lindbladh, 2005; Novák et al., 2012; Pędziszewska and Latalowa, 2016). However, land use policies and increasingly efficient fire suppression were likely more important than the impact of climate variability on the fire regime in the most recent period. The warmer climate in Białowieża Forest that the region recorded since the second half of the 1800s (Jędrzejewska et al., 1997; Pierzgalski et al., 2002; Büntgen et al., 2011; Boczoń et al., 2018) was instead associated with a decline in fire activity, as revealed by our results. Similarly, increased early season fire hazard over most of Belarus and Poland, as indicated by the dynamics of the monthly drought code over the 20th century (Drobyshev et al., 2021b), has not been reflected in fire activity until now.

Disentangling the climatic and anthropogenic impacts on ecosystem dynamics is complex and probably not always possible (Bradshaw and Lindbladh, 2005; Pędziszewska and Latalowa, 2016). We propose that to explore causes and mechanisms behind the recorded changes in fire activity in Białowieża Forest in even greater detail, a larger sampling area should be considered and further, parallel research in historical and paleoecological archives should be continued (cf. Latalowa et al., 2015; Czerwiński et al., 2021).

Ignition Density: Its Relation to Fire Causes and Fuel Dynamics

Historical fires in Białowieża Forest occurred at frequencies exceeding frequency of lightning-caused ignitions (H3), since the number of fires we recorded exceeded the present day natural background level of ignition by a factor of 10–100, when compared with data on lightning ignition frequencies from Poland, Germany, and Latvia (Sannikov and Goldammer, 1996; Ubyss and Szczygiel, 2006; Szczygiel et al., 2016; Brackhane et al., 2021). The sizes of the largest fires, given that they originate from one or few ignitions, have implications on calculations of the ignition density on the landscape and hence, on the interpretation of the possible fire causes – whether they are caused by lightning or by people (Niklasson and Granström, 2000). Therefore, we think that our results must be interpreted with caution, since most fire areas were only partly reconstructed, and especially for the larger ones, the values we obtained may have been greatly underestimated. Obviously, our reconstruction lacks direct information on fire cause, which could potentially be corroborated with more extensive analyzes of independent historical data (court protocols, eye-witness accounts, diaries, etc.). However, such high ignition density level on the landscape points toward a strong anthropogenic influence on the fire regime for the studied area and period (Niklasson and Granström, 2000).

The multi-site approach with a high degree of sample replication at each study site also yielded a spatial dimension of the occurrence of short-interval fires (1–2 years) recorded both at the scale of a given tree stand, i.e., study site, and at the scale of an individual sample tree, rarely recorded in *P. sylvestris* trees (Ivanova et al., 2016; Zin et al., 2015). We documented several cases in which two multi-site fires were recorded in consecutive years at the same study site, although not in the same tree – as in the years: 1757 and 1758, 1769 and 1770, 1789 and 1790 (Supplementary Figure 1). However, only a 1-year interval recorded in the very same tree is definite evidence of fire passing the same point in space, as was actually documented by one of our sample trees (Figure 8). Nevertheless, the intermixed positions of scars in two consecutive years suggest that fire actually spread over the same stand in the succeeding year (Supplementary Figure 1), but likely with highly varying severity. We believe that a 1-year scarring interval is simply very rare because of
feedback mechanisms between fire severity and post-fire stand regeneration (Zin et al., 2015). Similar sudden growth depressions recorded in two surviving pines, and we had a clear evidence of its high severity, as was proven by (Zin et al., 2015). In case of the 1718 fire (in the site P1), Granström, 1996, 1997). In the cumulative fire curve for our site, the fires in Białowieża Forest seems to be enough to sustain 1-year intervals for fires on a larger scale of a tree stand (Niklasson et al., 2010; Zin et al., 2015; Spanu et al., 2020), at the scale of an individual tree the fuel accumulation seems to be usually not sufficient to produce enough heat for a renewed scarring. The reason for this is likely a variation in fire behavior and heterogeneous fire severity within the same fire event, which impact several features of the post-fire tree stand (e.g., Keyser et al., 2008; Iniguez et al., 2009; Parro et al., 2009; Marozas et al., 2007; Parro et al., 2009). The strong grass dominance in the fuel bed may also relate to the open, park-like stand structure resulting from repeated low- and mixed-severity fire disturbances (Brown and Wu, 2005; Zin, 2016). However, while fuel build-up in Białowieża Forest seems to be enough to sustain 1-year intervals for fires on a larger scale of a tree stand (Niklasson et al., 2010; Zin et al., 2015; Spanu et al., 2020), at the scale of an individual tree the fuel accumulation seems to be usually not sufficient to produce enough heat for a renewed scarring. The reason for this is likely a variation in fire behavior and heterogeneous fire severity within the same fire event, which impact several features of the post-fire tree stand (e.g., Keyser et al., 2008; Iniguez et al., 2009; Parro et al., 2009; Marozas et al., 2007; Parro et al., 2009; Adámek et al., 2016), including the following fuel succession (Skre et al., 1998; Adámek et al., 2016),

Even though data is lacking, nothing precludes the idea that some fires in Białowieża Forest did consume more fuel, went deeper into the humus layer and/or were so severe that a subsequent build-up of fuel was slower (Schimmel and Granström, 1997). In the cumulative fire curve for our studied area (Figure 6), there are at least three cases in which fires were followed by much longer fire-free periods than usual. Perhaps these gaps in the fire record were a product of high-severity fires which occasionally occurred in Białowieża Forest (Zin et al., 2015). In case of the 1718 fire (in the site P1), we had a clear evidence of its high severity, as was proven by sudden growth depressions recorded in two surviving pines, and the subsequent cohort regeneration (Zin et al., 2015). Similar feedback mechanisms between fire severity and post-fire stand dynamics have been described, for example, in North American ponderosa pine (Pinus ponderosa Douglas ex Lawson and C. Lawson) ecosystems (Iniguez et al., 2009). However, further analysis of past fire behavior and subsequent fire intervals is needed to convincingly show the effect of high-severity fires in Białowieża Forest.

Management Implications

Fires have been present in Białowieża Forest throughout the Holocene (Dąbrowski, 1959; Latalowa et al., 2016) with the proportion of human ignitions likely to be increasing over time. However, it is still unknown whether human-influenced fire regimes differ fundamentally from natural, lightning-driven ones in terms of the resulting ecosystem structure and species composition. In a natural fire regime, fires may grow very large when unsuppressed, partially outweighing the effect of relatively few ignitions (Niklasson and Granström, 2000; Rolstad et al., 2017). Both natural- and human-driven fire activity has been effectively suppressed in the modern times and the lack of fire puts at risk the regeneration of fire-adapted canopy dominants, such as pine and oak (Spanu et al., 2020; Drobyshiev et al., 2021a). This is especially true for Scots pine (Niklasson et al., 2010; Zin et al., 2015), which natural regeneration is proven to decline in Białowieża Forest over the last decades (Kuijper et al., 2010; Drozdowski et al., 2012; Paluch, 2015; Brzeziecki et al., 2020) – a pattern now regarded a region-wide phenomenon (Matuszkiewicz, 2007; Heinken, 2008). Thus, our results contribute directly to the ongoing discussion on management and conservation of Central European pinewoods (Walentowski et al., 2007; Heinken, 2008; Adámek et al., 2016). Moreover, a number of fire- and light-demanding species in Białowieża Forest have declined due to the long period of fire suppression and have been shown to respond positively to recent fires (Gutowski et al., 2020). Finally, vegetation (and fuel) changes currently experienced by the area due to the effective fire control (Szczygieł et al., 2016) may result in strong and unpredictable feedbacks upon future fire regimes (Pifol et al., 2005; Stephens et al., 2013; Steel et al., 2015). To sum up, future conservation and management strategies for Białowieża Forest and for the region need to seriously re-consider the fundamental role of fire in temperate forest ecosystems.

CONCLUSION

Although limited in the area covered, the study provided valuable quantitative information on the past forest fire activity, that has not been available for temperate European forests. In addition, the spatial study design brought the first quantification of the ignition density, pointing toward a strong human impact on ignitions. Fires in Białowieża Forest varied considerably in size, with the largest events probably exceeding 1 000 ha. Future studies could elucidate the role of wet forest types and depressions as fire breaks, as well as further document the size distribution of past fires (including both very large and very small events), which would further strengthen the discussion on possible fire causes. This type of study would also be important for the deeper understanding of
the role of climate in driving the past fire regime in this region and potential future fire activity under climate change (cf. e.g., Drobotysh et al., 2014). Our results highlight the importance of fire disturbance in Bialowieża Forest, show the potential for further spatial reconstructions for studies of past fire regimes in temperate Europe, and provide critical baseline information to design biological conservation and management strategies in European forests.

DATA AVAILABILITY STATEMENT

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

AUTHOR CONTRIBUTIONS

EZ and MN developed study ideas and design, collected the data, and wrote the first draft. ID developed parts of statistical setup. EZ legally organized and realized field work. EZ, ŁK, ID, and MN analyzed the data. All authors contributed equally to the writing of the text and have given their approval for publication.

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SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fevo.2022.768464/full#supplementary-material

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