A Systematic Look at the Temperature Gradient Contribution to the Dayside Magnetopause Current

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Key Points:
- The magnetopause diamagnetic current is composed of opposing density and temperature gradient generated components.
- The temperature gradient contributes up to 37% of the ion diamagnetic current density along the magnetopause.
- The temperature component typically opposes the classical Chapman-Ferraro current direction.

Abstract
Magnetopause diamagnetic currents arise from density and temperature driven pressure gradients across the boundary layer. While theoretically recognized, the temperature contributions to the magnetopause current system have not yet been systematically studied. To bridge this gap, we used a database of Magnetospheric Multiscale magnetopause crossings to analyze diamagnetic current densities and their contributions across the dayside and flank magnetopause. Our results indicate that the ion temperature gradient component makes up to 37% of the ion diamagnetic current density along the magnetopause and typically opposes the classical Chapman-Ferraro current direction, interfering destructively with the density gradient component, thus lowering the total diamagnetic current density. This effect is most pronounced on the flank magnetopause. The electron diamagnetic current was found to be 5–14 times weaker than the ion diamagnetic current on average.

Plain Language Summary
The solar wind represents a continuous outflow of charged particles from the Sun’s upper atmosphere into the solar system. Upon reaching Earth’s magnetosphere, the solar wind’s dynamic pressure is balanced by the magnetic pressure of Earth’s magnetic field in a boundary layer known as the magnetopause. This boundary layer represents the entry point of the solar wind’s energy into Earth’s magnetosphere and upper atmosphere, playing a crucial role in energy transport throughout the interconnected system. Plasma density and temperature differences across the boundary layer generate an electric current that supports the magnetopause. In this paper, we clarify the physical mechanism of the magnetopause current by using high-resolution data from NASA's Magnetospheric Multiscale mission. We found a significant ion temperature contribution to the magnetopause current not identified in previous studies. Our results also indicated that the plasma electrons’ contribution to the magnetopause current was significantly smaller than the ion contribution.

1. Introduction
The magnetopause is a magnetosphere boundary layer created through the dynamic pressure balance between the solar wind’s kinetic pressure and Earth’s magnetic field. The solar wind causes distortions in the magnetosphere’s magnetic field topology supported by a current sheet first proposed by Chapman and Ferraro (1931), often termed the Chapman-Ferraro (CF) current, which runs in a dawn-to-dusk direction around the magnetopause (Ganushkina et al., 2018). This current structure is believed to be generated through pressure gradients at the magnetopause boundary layer, as explained in Hasegawa (2012), the magnetosheath plasma has a higher plasma density, while the magnetosphere will have a higher ion temperature. The resulting changes in plasma density and temperature across the magnetopause lead to gradients that generate ion and electron diamagnetic currents running perpendicular to the magnetic field (Ganushkina et al., 2018).

Because of the magnetopause's important role in magnetic reconnection and the resulting transfer of plasma and energy into the magnetosphere it has been the focus of numerous studies (Cahill and Amazeen, 1963); Haaland et al. (2014, 2019, 2020); Le and Russell (1994); Phan and Paschmann (1996); Phan et al. (1996); Paschmann et al. (2018); Shuster et al. (2019); etc.) and missions Magnetospheric Multiscale (MMS, THEMIS, and Cluster), which have delved deeper into the current sheet's structure and creation. From Paschmann et al. (2018) and their MMS magnetopause crossing database, the total current density across the dayside magnetopause was studied in detail. The flank magnetopause total current density was then surveyed in Haaland et al. (2019) and Haaland et al. (2020) where the flanks were found to have a weaker current density and a correspondingly thicker boundary layer than the dayside.
While the literature generally recognizes the importance of both the density and temperature in generating diamagnetic currents, a large-scale systematic analysis of their contribution to the magnetopause current system has not yet been accomplished. To help fill this gap in the literature, we considered four years of burst mode MMS mission data over the magnetopause crossing intervals provided by Paschmann et al. (2018)’s MMS database. We measured both the density and temperature diamagnetic current components and created current accumulations of their contributions. In doing so, we found that the temperature diamagnetic current component is a statistically significant factor for both the dayside and flank magnetopause current sheet by acting against the density component and thus reducing the total diamagnetic current density.

2. Data and Analysis

2.1. MMS Database

For this study, we used four years of data from Paschmann et al. (2018, 2020) and Haaland et al. (2020)’s database of MMS current sheet crossings. MMS is a mission comprised of four separate spacecraft traveling in a tetrahedron pattern through the magnetopause (Burch et al., 2015). This database catalogs MMS current sheet crossings based on Fast Plasma Investigation (FPI) (Pollock et al., 2016), Fluxgate Magnetometer (Russell et al., 2016), and Hot Plasma Composition Analyzer (HPCA) measurements (Young et al., 2016). The magnetopause transit times are captured through an automated minimum variance analysis (MVA) of the magnetic field in boundary normal or LMN coordinates (Paschmann et al., 2018, 2020). The start and end times of the current sheet crossing were then assigned as covering 76% of the primary magnetic field component’s ($B_L$’s) transition across the magnetopause boundary, which represents the thickness of a tanh fit of the $B_L$ profile.

The magnetopause velocities for each crossing are obtained through a de Hoffman-Teller analysis and stored by the database as the de Hoffman-Teller velocity ($V_{HT}$). We then took these velocities and dotted them with the normal direction in the LMN coordinate system to get the normal magnetopause boundary velocity ($V_n$) (Khrabrov & Sonnerup, 1998; Paschmann et al., 2018; Sonnerup & Scheible, 1998). The magnetopause thicknesses are then determined by taking this normal velocity and multiplying it by the time span of the current sheet as determined by the MVA analysis.

The database places identifiers on the individually identified magnetopause crossings, classifying their characteristics and structure. A full description of this process and the current sheet identifiers can be found in Paschmann et al. (2018) and Haaland et al. (2020). An example crossing, with the database defined magnetopause current sheet crossing indicated by the dashed orange lines, is given in Figure 1.

2.2. Magnetopause Current Sheet Selection

We chose database defined crossings from 2015 to 2018 using event identifiers to select for complete and monotonic magnetopause crossings, where monotonic indicated events that had a constant magnetopause velocity so their thicknesses and durations could be computed. Additionally, we included Harris sheet-like events, or simple clean magnetopause crossings that were also complete, monotonic events in our data set. Events with unusually high (above 2,000 nA/m²) current density spikes during the magnetopause crossing time, such as would occur during a reconnection event, or when data flags for the various instruments were non-nominal, were manually removed from our data set.

Alongside the database defined event criteria, we imposed two additional conditions on our events in order to ensure high signal-to-noise ratios and typical magnetopause plasma number densities as described below:

First, we considered the current measured during the magnetopause crossing and selected for events that reported significant peaks in their current densities. Where “significant” in our case was considered to be a current crossing where at least 50% of the crossing duration was within 15% of the maximum current peak during that crossing. This condition enabled us to select events with a strong current and high signal-to-noise ratios, thus ensuring that the temperature and density gradients and their resulting diamagnetic current components were not artificially diminished by the higher noise floor of low current density magnetopause crossings.

Second, we used HPCA measurements to compare the number densities of $H^+$ with $O^+$ for each magnetopause crossing. If $O^+$ exceeded 0.2 cm$^{-3}$ during the transit and $H^+$ remained below 1.5 cm$^{-3}$, then we considered $O^+$ to
dominate the magnetospheric ion mass density by more than a factor of two as described by Fuselier et al. (2019). Events fitting this classification were also removed from our data set as they represented densities not typically found in the magnetopause current sheet.

From the application of the database identifiers as well as our conditions, we were able to identify 561 events. The locations and corresponding years of the selected crossings are denoted in Figure 2.
2.3. Current Calculations

MMS’s four separate spacecraft allows the total current to be calculated by the curlometer method (Dunlop et al., 1988) that uses all four spacecraft to perform the curl of the observed magnetic field in order to approximate Ampere's law in the MHD approximation (e.g., Ganushkina et al. (2018)). This current we call the curlometer current:

$$\mathbf{J}_{\text{curl}} = \frac{\mathbf{v} \times \mathbf{B}}{\mu_0}. \quad (1)$$

Using the curlometer method, we also calculated the gradient of the ion density, and the divergence of the temperature tensor to get the total ion diamagnetic current and its current components ($\mathbf{J}_{\text{dia} \mathbf{v} \mathbf{N}}$ and $\mathbf{J}_{\text{dia} \mathbf{v} \mathbf{T}}$). We found the electron diamagnetic current to be at least one order of magnitude smaller, on average, than the ion diamagnetic current. This is in agreement with the results of Dong et al. (2018) in their case study where they found that the ion diamagnetic current was the primary source of perpendicular current in the magnetopause. Thus, we are presenting results for the ion components and will drop the signifier “i” going forward. The components of the total diamagnetic current are found by inverting the ideal gas law, $P = N k_b T$, into the expression for the total diamagnetic current $J_L = \frac{eN k_b T}{\mu_0} \frac{\partial v}{\partial x}$, see for example, Ganushkina et al. (2018) and the references therein, resulting in the following expressions:

$$\mathbf{J}_{\text{dia} \mathbf{v} \mathbf{N}} = \frac{\mathbf{B} \times \left( K_s \mathbf{T} \cdot \nabla N \right)}{|\mathbf{B}|^2}, \quad \mathbf{J}_{\text{dia} \mathbf{v} \mathbf{T}} = \frac{\mathbf{B} \times \left( K_s N \mathbf{v} \cdot \mathbf{T} \right)}{|\mathbf{B}|^2}. \quad (2)$$

Where, by definition, $\mathbf{J}_{\text{dia} \mathbf{v} \mathbf{N} \text{ Total}} = \mathbf{J}_{\text{dia} \mathbf{v} \mathbf{N}} + \mathbf{J}_{\text{dia} \mathbf{v} \mathbf{T}}$.

These currents were then averaged over our selected magnetopause crossings for each event, leading to single averaged current values for each of our 561 crossings. See Figure 1 for an example crossing on the dayside magnetopause. Note the two vertical orange lines represent the database defined current sheet crossings where the resulting currents would be averaged over.

All of the data taken from MMS, as well as the current calculations, was interpolated to the 30 ms FPI electron time resolution. As our main results involve ion diamagnetic currents and the total current as computed from the curlometer method, any sub 150 ms variations in the ion measurements will not impact our results. For non-curlometer calculations, we averaged over all four spacecraft to create a single data stream where, on average, the MMS separation during 2015–2018 was 10–60 km while the magnetopause current crossings analyzed typically had thicknesses greater than several hundred km, sufficiently larger than the max 60 km tetrahedron separation. These calculations were completed in GSE coordinates and then converted to a local spherical coordinate system built off of the Cartesian GSE coordinates. See Figure 2 for a depiction of our spherical coordinate system and the definition of the dusk and dawn sectors.

Note, one limitation with applying the curlometer method to a magnetopause current sheet crossing is that the curlometer method requires simultaneous measurements from all four spacecraft to calculate a gradient, curl, or divergence of a quantity. Thus, errors occur when one or more of the MMS constellation is outside of the magnetopause boundary as the spacecraft are no longer all measuring the same medium. For our study, this caveat is mitigated by the fact that the magnetopause boundary is, on average, at least one order of magnitude larger than the average MMS separation, with median magnetopause thicknesses often reported as approximating 1,000 km (e.g., Haaland et al. (2019, 2020)). This makes times where the curlometer method results in erroneous measurements brief and limited to the outskirts of a current sheet crossing.
2.4. Current Accumulations

The averaged currents from the 561 magnetopause crossings were put into bins corresponding to MMS’s physical location in our local spherical coordinate system. The angle \( \phi \) was used to create 1-dimensional bins from 90° to −90° in 20° increments. This was done for each component in the spherical coordinate system as described in Figure 2 and shown in Figure 3. Error bars for each figure were computed using the standard error or \( \sigma/\sqrt{N} \), with \( \sigma \) the standard deviation of the values in each bin and \( N \) the number of events that fell inside that bin, with the bins visually depicted in Figures (3a–3d) by the gray-dashed lines.

Using Figure 3 we can make several observations. The first is that the \( J_{\text{out}} \), \( J_{\text{dia Total}} \), and \( J_{\text{dia VN}} \) \( \phi \)-component currents are all in the +\( \phi \) direction across the magnetopause, or in the classical CF, dawn-to-dusk direction. However, the \( J_{\text{dia V T}} \) \( \phi \) component is consistently in the −\( \phi \) direction, or from dusk-to-dawn across the magnetopause. Therefore the two components of the ion diamagnetic current, \( J_{\text{dia VN}} \) and \( J_{\text{dia V T}} \), are oppositely directed across the magnetopause as can be seen in Figure 3b. The directions of these components are as expected when using the magnetospheric quantities evaluated by Hasegawa (2012) where the density component should run in the traditional CF current direction as the plasma density is higher in the magnetosheath and lower in the magnetosphere. At the same time, it is expected that the ion temperature is lower in the magnetosheath and higher in the magnetosphere, leading to the ion temperature component typically running counter to the CF current direction. On average, however, \( J_{\text{dia VN}} \) is stronger than \( J_{\text{dia V T}} \) which allows the total diamagnetic current, \( J_{\text{dia Total}} \), to still flow in the classical CF direction.

2.5. Current Measurement Results

We then used our data to create a table of results over the dusk, dayside, and dawn magnetopause including the mean, median, and standard errors for our 561 magnetopause crossings as seen in Table 1.

From this table, \( J_{\text{out}} \) is strongest on the dayside, with a dusk-dawn asymmetry as the dawn curlometer current is stronger than the dusk. Both \( J_{\text{dia Total}} \) and \( J_{\text{dia VN}} \) show similar distributions with the dayside again being the strongest sector, but the dusk and dawn results are now in agreement within their standard errors. The \( J_{\text{dia V T}} \) component shows a dusk-dawn asymmetry with a significantly stronger dusk current density than the dawn. The total electron diamagnetic current, \( J_{\text{e dia Total}} \) is the weakest current component studied and shows a dusk-dawn asymmetry with the dusk being significantly weaker than both the dusk and the dayside.

Comparing these results with past studies, our magnetopause thickness are, in general, higher than those found by Haaland et al. (2020); however, this is to be expected as we are considering a specific subset of database events as described in the previous sections. For similar reasons, the current densities found by Haaland et al. (2020) show differences, with our dawn current densities matching closely, but the dusk and dayside values showing deviations.

3. Temperature Gradient’s Impact on the Magnetopause Current System

Using our results from Figure 3 and Table 1, we can posit three primary ways the ion temperature gradient impacts the magnetopause current density and, in doing so, create a 2D diagram to summarize our findings as shown in Figure 3e.

1. The divergence of the ion temperature tensor generates up to one third of the total ion diamagnetic current density in the \( \phi \) direction. Specifically, in the \( \phi \) direction, the ratio of the temperature-driven component to the density-driven component’s contribution to the total ion diamagnetic current gives the following break down with \( J_{\text{dia V T}} \) making up 37% of the diamagnetic current density on the dusk, 20% on the dayside, and 30% on the dawn.

2. \( J_{\text{dia V T}} \) goes in the opposite direction of the classical CF Current. \( J_{\text{dia V T}} \) is clearly in the −\( \phi \) direction across the magnetopause when considering Figure 3b and Table 1. This results in \( J_{\text{dia V T}} \) lowering the contribution of \( J_{\text{dia VN}} \) making the \( J_{\text{dia Total}} \) less than \( J_{\text{out}} \) on average as seen in Table 1.

3. \( J_{\text{dia V T}} \)’s contribution to the magnetopause current density becomes more important toward the flank magnetopause.
Figure 3. (a) Through (d) depict current density accumulations for $J_{\text{curl}}$, $J_{\text{dia Total}}$, $J_{\text{dia } \nabla N}$, and $J_{\text{dia } \nabla \cdot \mathbf{T}}$ across the dayside and flank magnetopause sectors, represented by the orange dashed lines. The gray dashed lines represent the edges of the 20° accumulation bins, centered at 80°, 60°, 40°, etc. Moving from top to bottom: (a). Represents the R-component of the current in our local spherical coordinate system (described in Figure 2). (b). $\phi$ - component, (c). $\theta$ - component, (d). Magnitude of the current components. (e) diagram of the contributions and directions of $J_{\text{dia Total}}$, $J_{\text{dia } \nabla N}$, and $J_{\text{dia } \nabla \cdot \mathbf{T}}$ across the dusk, dayside, and dawn magnetopause. Note. The size of the arrows in each sector indicates the relative magnitude of their current densities and the direction indicates the current component's flow around the magnetopause.
Table 1
Magnetopause Parameters, Magnitudes, and \( \phi \) Components of Current Densities Across Dusk (+50° to +90°), Dawn (−50° to −90°), and Dayside (>−50° and <+50°) Crossings With the Following Format: Mean (Median) ± Standard Error of the Values Measured Over Each Event’s Magnetopause Crossing

| Parameter            | Dusk        | Dayside     | Dawn        |
|----------------------|-------------|-------------|-------------|
| Number of Crossings  | 215         | 225         | 121         |
| Thickness (km)       | 1961 (1429) ± 132 | 2032 (1379) ± 135 | 1950 (1393) ± 180 |
| Thickness (\( d_l \))| 21.4 (14.3) ± 1.5  | 26.7 (16.4) ± 1.8  | 28.0 (15.9) ± 3.0  |
| Thickness (\( R_p \))| 35.6 (18.7) ± 4.3  | 85.9 (33.2) ± 12  | 33.2 (17.9) ± 4.0  |
| \( V_a \) (km/s)    | 131.5 (109.8) ± 7.0 | 99.8 (71.1) ± 7.5  | 94.3 (86.0) ± 6.3  |
| Duration (s)         | 16.0 (11.2) ± 0.8  | 25.0 (18.7) ± 1.3  | 20.6 (14.7) ± 1.3  |
| \( J_{\text{curl}} \) (nA/m²) | 30.0 (17.6) ± 3.2  | 83.2 (67.7) ± 5.6  | 52.1 (35.0) ± 6.3  |
| \( J_{\text{dua total}} \) (nA/m²) | 19.4 (8.4) ± 5.8  | 59.0 (47.5) ± 6.9  | 25.1 (17.5) ± 5.0  |
| \( J_{\text{dua V||}} \) (nA/m²) | 50.0 (28.3) ± 8.5  | 78.2 (64.7) ± 7.2  | 43.6 (32.7) ± 6.6  |
| \( J_{\text{dua V⊥}} \) (nA/m²) | 31.5 (16.6) ± 7.0  | 20.0 (13.5) ± 5.8  | 19.6 (10.8) ± 2.9  |
| \( J_{\text{dia V||}} \) (nA/m²) | 21.1 (10.9) ± 2.7  | 83.0 (67.4) ± 5.3  | 45.3 (30.3) ± 5.9  |
| \( J_{\text{dia V⊥}} \) (nA/m²) | 18.2 (8.3) ± 5.2  | 58.7 (47.3) ± 5.2  | 22.6 (16.6) ± 5.0  |
| \( J_{\text{dia V T}} \) (nA/m²) | 43.2 (26.1) ± 6.7  | 78.0 (64.7) ± 6.7  | 39.4 (30.4) ± 6.6  |
| \( J_{\text{dia V T}} \) (nA/m²) | −25.0 (−16.1) ± 6.2 | −19.3 (−13.3) ± 4.1 | −16.8 (−10.7) ± 2.7 |
| \( J_{\text{dia total}} \) (nA/m²) | 1.7 (0.5) ± 0.7  | 5.9 (4.4) ± 1.0  | 5.4 (3.6) ± 0.7  |

Note. \( V_a \) values are unsigned averages and medians of the normal magnetopause velocity.

\( J_{\text{dua total}} \) and \( J_{\text{dia V||}} \) are strongest on the dayside and grow steadily weaker on the dusk and the dawn flanks, with both flanks showing similar results for the current densities. This is in contrast to \( J_{\text{dia V T}} \) which increases in strength, particularly on the dusk flank, resulting in the total diamagnetic current being decreased even further by \( J_{\text{dia V T}} \)’s impact on the flank magnetopause.

From Table 1 and Figure 3 we may also notice the difference between the curlometer and the total diamagnetic current densities across the magnetopause. This difference is generally expected as the curlometer current represents the total current density as measured by MMS across the magnetopause current layer, which includes contributions from both ion and electron currents and their perpendicular and parallel components. The total ion diamagnetic current is thus one component of the curlometer current. This being said, the total diamagnetic current density still represents the main contributor to the curlometer current, accounting for almost 70% of the current density on the dayside.

4. Discussion

4.1. Gradient Formation

These observations are generally consistent with previous literature regarding ion populations in the magnetosphere. From Chappell et al. (2008), the warm plasma cloak is defined as a population of 10 eV to 3 keV ions energized in the polar cap and magnetotail, which circulates in a dawn-to-dusk circulation pattern throughout the inner magnetosphere out to the magnetopause. As Chappell et al. (2008) notes, the warm plasma cloak ions can occupy the same space as the much warmer and more energetic ring current ions, which circulate in the opposite direction across the magnetosphere, from dusk-to-dawn. Thus it is possible, on a simplified level of magnetospheric circulation, that the colder warm plasma cloak ions provide generating pressure for the density gradient component across the magnetopause in its dawn-to-dusk, CF current like direction, while the warmer ring current ions provide generating pressure for the temperature gradient component in its dusk-to-dawn direction. The
density component's dayside-flank asymmetry could also be explained by additional density gradients generated by the plasmasphere drainage plume (Borovsky & Denton, 2008), which exhausts through the dayside magnetopause during storm conditions, enhancing the dayside with more cold ions, thus leading to an enhanced dayside $J_{\phi V N}$ while leaving the dusk and dawn components reliant solely on the warm plasma cloak ion population. The presence of magnetosphere particle populations in the magnetopause and their effects on the magnetopause current sheet, specifically regarding magnetic reconnection, is explored in the following studies: for example, Borovsky and Denton (2008), Fuselier et al. (2017), and Walsh et al. (2013). However, additional study regarding the effects of the warm plasma cloak and plasmasphere particle populations on the diamagnetic current and its component generation is needed.

4.2. Large-Scale Current System Closure

From Figure 3 and Table 1, we can see the total diamagnetic current density is lower on the flanks, helped by an increasingly prominent temperature component. This indicates current closure with the larger 3D current system where the magnetopause current steadily curves toward the parallel as it goes around the flanks. As the total current through the magnetopause current system must remain constant, as charge is conserved, the current density of the total current's components must then fluctuate appropriately along the dayside and flanks to “transfer” the current density from the perpendicular dominated dayside to the increasingly parallel dominated flanks. This indicates that the total diamagnetic (perpendicular) current decreases in order for more of the total current density to be diverted toward parallel currents such as the field aligned currents. Additionally, the magnetopause is generally thinner on the dayside and thicker on the flanks while the magnetopause current must remain the same as it flows through the magnetopause (e.g., Haaland et al. (2019, 2020)). Thus, the total perpendicular current density must change to compensate for either a thinner or thicker magnetopause. From Figure 3, we can see this is, indeed, the case as the total diamagnetic current density decreases on the flanks and is strongest toward the dayside.

4.3. Ion Versus Electron Diamagnetic Current Densities

Even though the electron diamagnetic current was found to be significant on electron scale current sheets by Shuster et al. (2019, 2021), we found the electron current density to be less significant over our ion scale magnetopause current sheets. Specifically, the electron diamagnetic current density is 7% of the ion diamagnetic current density in the $\phi$ direction on the dusk, 10% on the dayside, and 20% on the dawn. This presents an interesting asymmetry for the electron current as it is noticeably weaker on average on the dusk than it is on the dawn; however, in both cases, the electron diamagnetic current is also weaker than the contribution made by the ion current. The weaker electron current density may be explained based on the fact that we are averaging over many electron scale current sheets when considering our ion scale magnetopause crossing, thus lowering the resulting current density.

5. Conclusions

In this paper, we have quantified the diamagnetic current's composite nature. Based on our systematic analysis of four years of MMS magnetopause crossings, the diamagnetic current is composed of two competing components: one generated by density gradients and one by temperature gradients.

We have found that the temperature generated component acts against the density component, weakening the total diamagnetic current's net strength, particularly on the flanks where the temperature component's contribution can reach up to 37% of the diamagnetic current density along the magnetopause. We also found that ions contribute the majority of the current density to the diamagnetic current, with electrons accounting for only 7%–20% of the ion's contribution.

Taking these findings into account, we can posit that enhancements of the temperature gradient along the magnetopause boundary may lead to a corresponding weakening of both the diamagnetic current and, by extension, the total current in the magnetopause. This implied weakening of the magnetopause current by the temperature gradient leads to a more complicated picture of the interaction between the solar wind and Earth's magnetosphere, especially on small scales where situations can arise where the two components of the diamagnetic current
become equal yet opposite, leading to the net cancellation of the diamagnetic current in that region. Studying the small-scale consequences of this interaction is the basis of our future work.

Data Availability Statement
The Magnetospheric Multiscale (MMS) data used in this study is publicly available at https://asp.colorado.edu/mms/sdc/public/datasets/ from the FPI, FIELDS, and HPCA data sets. The averaged MMS data over the 561 MP crossings used for this study are available through a Harvard Dataverse public database: https://doi.org/10.7910/DVN/SRB1CR.

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References
Borovsky, J. E., & Denton, M. H. (2008). A statistical look at plasmaspheric drainage plumes. Journal of Geophysical Research: Space Physics, 113(A9). https://doi.org/10.1029/2007JA012994
Burch, J. L., Moore, T. E., Torbert, R. B., & Giles, B. L. (2015). Magnetospheric multiscale overview and science objectives. Space Science Review, 199, 5–21. https://doi.org/10.1007/s11214-015-0164-9
Cahill, L. J., & Amazeen, P. G. (1963). The boundary of the geomagnetic field. Journal of Geophysical Research, 68(7), 1835–1843. https://doi.org/10.1029/JD068i007p01835
Chapman, S., & Ferraro, V. C. A. (1931). A new theory of magnetic storms. Terrestrial Magnetism and Atmospheric Electricity, 36, 77–97. https://doi.org/10.1029/et036i002p00077
Chappell, C. R., Huddleston, M. M., Moore, T. E., Giles, B. L., & Delcourt, D. C. (2008). Observations of the warm plasma cloak and an explanation of its formation in the magnetosphere. Journal of Geophysical Research, 113(A9). https://doi.org/10.1029/2007JA012945
Dong, X. C., Dunlop, M. W., Wang, T. Y., Cao, J. B., Trattner, K. J., Bamford, R., et al. (2018). Carriers and sources of magnetopause current: MMS case study. Journal of Geophysical Research: Space Physics, 123, 5464–5475. https://doi.org/10.1002/2018JA025292
Dunlop, M. W., Southwood, D. J., Glassmeier, K.-H., & Neubauer, F. M. (1988). Analysis of multipoint magnetometer data. Advanced Space Research, 8, 273–277. https://doi.org/10.1016/0273-1177(88)90141-x
Fuselier, S. A., Mukherjee, J., Denton, M. H., Petrinec, S. M., Trattner, K. J., Toledo-Redondo, S., et al. (2019). High-density $\alpha^+$ in earth’s outer magnetosheath and its effect on dayside magnetopause magnetopause reconnection. Journal of Geophysical Research: Space Physics, 124(12), 10,257–10,269. https://doi.org/10.1029/2019JA027396
Fuselier, S. A., Vines, S. K., Burch, J. L., Petrinec, S. M., Trattner, K. J., Casak, P., et al. (2017). Large-scale characteristics of reconnection diffusion regions and associated magnetopause crossings observed by MMS. Journal of Geophysical Research: Space Weather, 122, 5466–5486. https://doi.org/10.1002/2017JA024024
Ganushkina, N. Y., Liemohn, M. W., & Dubyagin, S. (2018). Current systems in the earth’s magnetosphere. Reviews of Geophysics, 56, 309–332. https://doi.org/10.1002/2017RG000950
Haaland, S., Paschmann, G., Øhlerose, M., Phan, T., Hasegawa, H., Fuselier, S. A., et al. (2020). Characteristics of the flank magnetopause: MMS results. Journal of Geophysical Research: Space Physics, 125. https://doi.org/10.1029/2019JA027623
Haaland, S., Reistad, J., Tenfjord, P., Gjerloev, J., Maes, L., DeKeyser, J., et al. (2014). Characteristics of the flank magnetopause: Cluster observations. Journal of Geophysical Research: Space Physics, 119, 9019–9037. https://doi.org/10.1002/2014JA020539
Hasegawa, H. (12). Structure and dynamics of the magnetopause and its boundary layers. Monogr. Environ. Earth Planets, 1(2), 71–119. https://doi.org/10.5047/meep.2012.00102.0071
Khrokhov, A. V., & Sonnerup, B. U. (1998). Dehoffmann-teller analysis. In Analysis methods for multi-spacecraft data (p. 221). ESA Publication Division.
Le, G., & Russell, C. T. (1994). The thickness and structure of high beta magnetopause current layer. Geophysical Research Letters, 21(23), 2451–2454. https://doi.org/10.1029/94GL02292
Paschmann, G., Haaland, S. E., Phan, T., Sonnerup, B., Burch, J., Torbert, R., et al. (2018). Large-scale survey of the structure of the dayside magnetopause by MMS. Journal of Geophysical Research: Space Physics, 123, 2018–2033. https://doi.org/10.1002/2017JA025121
Paschmann, G., Sonnerup, B., Haaland, S. E., Phan, T., & Denton, R. E. (2020). Comparison of quality measures for wavenumber relation. Journal of Geophysical Research: Space Physics, 125, e2020A026044. https://doi.org/10.1029/2020JA026044
Phan, T. D., Larson, D. E., Lin, R. P., McDalden, J. F., Anderson, K. A., Carlson, C. W., et al. (1996). The subsolar magnetosheath and magnetopause for solar wind ram pressure: Wind observations. Geophysical Research Letters, 23(10), 1289–1292. https://doi.org/10.1029/96GL00945
Phan, T. D., & Paschmann, G. (1996). Low-latitude dayside magnetopause and boundary layer for high magnetic shear: I. Structure and motion. Journal of Geophysical Research: Space Physics, 101(A4), 7801–7815. https://doi.org/10.1029/95JA03752
Pollock, C., Moore, T., Jacques, A., Burch, J., Gliese, D., Saito, Y., et al. (2016). Fast plasma investigation for magnetospheric multiscale. Space Science Reviews, 199, 331–406. https://doi.org/10.1007/s11214-016-0245-4
Russell, C. T., Anderson, B. J., Baumjohann, W., Bromund, K. R., Dearborn, D., Fischer, D., et al. (2016). The magnetospheric multiscale magnetometers. Space Science Reviews, 199, 189–256. https://doi.org/10.1007/s11214-014-0057-5
Shuster, J. R., Gershman, D. J., Chen, L.-J., Wang, S., Bessho, N., Dorelli, J. C., & Vilas, A. F. (2019). MMS measurements of the vlasov equation: Probing the electron pressure divergence within thin current sheets. Geophysical Research Letters, 46, 7862–7872. https://doi.org/10.1029/2019GL083549
Sonnerup, B. U., & Scheible, M. (1998). Minimum and maximum variance analysis. In Analysis methods for multi-spacecraft data (pp. 1850). ESA Publication Division.
Walsh, B. M., Sibeck, D. G., Nishimura, Y., & Angelopoulos, V. (2013). Statistical analysis of the plasmaspheric plume at the magnetopause. *Journal of Geophysical Research: Space Physics, 118*(8), 4844–4851. https://doi.org/10.1002/jgra.50458

Young, D. T., Burch, J. L., Gomez, R. G., Santos, A. D. L., Miller, G. P., IV, P. W., et al. (2016). Hot plasma composition analyzer for the magnetospheric multiscale mission. *Space Science Reviews, 199*, 407–470. https://doi.org/10.1007/s11214-014-0119-6