Shallow Subsurface Stratigraphy Inferred from the Use of Vertical Electrical Soundings (VES) Survey in Central Chihuahua, Mexico

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Abstract

Vertical Electrical Soundings (VES) provide fast and economical measurements used in geophysical exploration. VES were carried out in El Sauz-Encinillas (ESE) aquifer, in northern Mexico, to determine apparent resistivity and geoelectrical units’ thickness. Despite it being one of the three main aquifers feeding Chihuahua city, a lack of available geophysical data prevails in its northern portion. The main goal of this study was the determination of the geoelectrical units in the subsurface stratigraphy via electrical-resistivity soundings. The ESE’ aquifer is located within alluvial Quaternary sediments, with varying granulometry and reaching from a few meters to more than 600 meters of thickness at the center of the valley. Forty-five vertical electrical resistivity soundings (Schlumberger array, maximum AB/2 distance of 500 m) were performed throughout ESE aquifer’s northern portion. Field data were analyzed using software. Results illustrate a wide variability in resistivity values throughout the study area. Five geoelectrical units were identified: 1) a hardpan topsoil, with resistivity values ranging from 200 - 800 Ω-m; 2) an alluvial material mixture (sand/silt) with resistivity values ranging from 25 to 100 Ω-m; 3) playa lake-type material (clay/evaporites mixture) with resistivity values ranging from 0.2 to 15 Ω-m; 4) a gravel/sand mixture with resistivity values from 100 to 300 Ω-m; and 5) a partly fractured rock or conglomeratic material with resistivity values ranging from 400 to 3500 Ω-m. The electrical resistivity data, therefore gives reasonably accurate results that can be used to understand the subsurface stratigraphy and basement configuration in groundwater exploration.

Keywords
Geophysics, Vertical Electric Soundings, Geoelectrical Sections
1. Introduction

Vertical Electrical Soundings (VES) provide fast and economical measurements used in geophysical exploration and groundwater studies. Rapid expansion of urbanized zones within arid climate areas, plus growing populations and an ever-increasing need for water has led to a growing consumption and an overexploitation of water resources. In desertic areas such as Chihuahua, in northern Mexico, surficial water sources are limited and generally with poor quality, thus groundwater resources are the most important source of water for both human consumption and economic activities. Groundwater depths have increased as well as its quality degraded with time. The El Sauz-Encinillas’ aquifer (ESE), located 90 kilometers north of Chihuahua City, is an important source of water to its population. A spike in groundwater extraction from water wells within the area has led to degradation of water quality, rapid drawdowns, and change of groundwaters’ flow direction [1]. Previous work in this aquifer was related to geology [2] [3] [4] [5], geophysics [6] [7] [8] [9] [10], hydrogeology [1] [11]-[16] and groundwater flow models [17] [18] [19]. However, this work has focused mostly on the south-central portion of the ESE’s aquifer, where most of wells are located. The fine-grained sediments of the Laguna, product of the weathering of igneous rocks in surrounding ranges, hinder surface water’s infiltration, favoring evaporitic deposits and providing a chemical signature to northern ESE’s groundwater [13] [15] [16].

The main purpose of the present study was to provide information regarding the shallow subsurface stratigraphy, to determine the distribution and thickness of the sedimentary materials within ESE’s north-central aquifer’s portion.

2. Location and Climate

The study zone is located within the central portion of the Chihuahua state in Mexico. El Sauz-Encinillas’ valley, within an endorreic basin, is surrounded by the Sierra de la Tinaja Lisa (NW), the Del Nido block (W) and several (Peña Blanca, San Martín and Los Carneros) ranges to the East, covering a surface of approximately 1200 Km². The area is geographically located between parallels 28˚85’N and 29˚39’N, and meridians 106˚03’W and 106˚44’W, approximately at a straight distance of 92 kilometers North from Chihuahua city (Figure 1).

Climate in the area is commonly described as extremely arid and dry. According to the Köppen climate classification, modified specifically for Mexico [20], it can be classified as template-dry (BSk). Mean annual temperature is 15.64˚C, the highest monthly mean is in June (24.5˚C) while the coldest months are December and January (mean of 8˚C) [2]. Mean annual precipitation is 356.02 mm, while it rains the most in July and August (84.5 mm mean precipitation), and the driest months are March and April (mean monthly precipitation of 3.2 mm), thus summer is the monsoonal season [14]. All the above, plus the nature of precipitation, make though conditions for the aquifer’s recharge, because the months with higher precipitation are also the ones with higher...
temperatures, enhancing evapotranspiration and evaporite precipitation (playa lake conditions).

### 3. Geological Setting

The Quaternary sediments present in El Sauz-Encinillas’ valley have a wide ranged granulometry, which vary from boulder-sized sediments close to the mountains, coarse to medium sediments in the alluvial fans, all the way to sand in the fluvial deposits of the intermittent streams, to clays and evaporites in the Laguna de Encinillas. Del Nido Block extends almost 140 km with an N-S orientation. Limited by faults due to the Basin and Range province (Miocene), this mountain range is composed mostly of thick Tertiary alkaline igneous rocks [3].

Regarding the valley’s eastern limit, the Peña Blanca range contains a similar volcanic sequence; but slightly variable in lithology, as well as in age and thickness. The volcanic sequence rests discordantly overlying Cretaceous limestones [21].

Physiographically speaking, the study area is located within the Basin and Range Province [22]. Ranges are low and abrupt, generally trending NNW-SSE, next to bajadas and alluvial-filled basins. River patterns in ESE due to the meteorological circumstances already explained, and the valley’s sedimentology do not allow the presence of important currents, being mainly ephemeral streams (arroyos) only carrying water during rainy seasons. Several of these arroyos feed the Laguna de Encinillas, previously considered as a perennial water body, but due to the continuous drought and excessive well water drawdowns, are nowa-
The ESE’s aquifer is emplaced in Quaternary alluvial deposits, with granulometry that varies from boulders and gravels close to the ranges to clays and evaporitic deposits reaching thicknesses of up to 600 meters in the center of the valley. These sediments have medium to high permeabilities in most of the aquifer and with transmissibility values ranging from $1.5 \times 10^{-3}$ to $12 \times 10^{-3}$ m$^2$/s [14]. The alluvial material interstratifies with fractured volcanic rocks, behaving as a single geohydrological unit. In addition, as mentioned before, there are some outcrops of Cretaceous limestones mostly towards East of the area. In the ESE’s aquifer, most of the wells were drilled in its southern part, near El Sauz town, where an important area of agricultural irrigation exists, causing big drawdowns in this area.

4. Methodology

Among all the electrical resistivity methods, Vertical Electrical Soundings (VES) have been applied most widely in groundwater exploration studies [24]. They have a wide range of applications, such as determination of alluvium thickness and unconsolidated water-bearing aquifers, detection of faults, fractures and groundwater movement [25] and/or weathered/fractured zones with reasonable accuracy [7]. The initial application of electrical resistivity methods to geophysical prospecting started with the original work [26] [27] both authors proposed similar techniques with four-point electrode configuration for field measurements but with different electrode displacements. Since then, these electrode arrays (Wenner & Schlumberger) became very popular and are widely used in hydrogeological prospecting. The Schlumberger array offers good depth of penetration, as well as good conductive surface layer penetration, and excellent resolution of horizontal layers [28]. This array was used throughout this study because of its simpler and straightforward field logistics, and relatively quick/economic analysis using available free software (IPI2WIN).

In the Schlumberger method (Figure 2), the electric current (DC or low frequency AC) is injected into the ground, between two points through 2-feet stainless-steel current electrodes (AB) driven into the ground surface, and the resulting potential difference between two other points (MN, also 2-feet stainless-steel potential electrodes) is measured. The four electrodes are arranged in such a way that the distance between the two inner potential electrodes (MN) is kept small compared to the distance between the outer current electrodes (AB), while making sure that the four electrodes (AMNB) are placed along a line. Distance between the potential electrodes should never exceed more than one fifth of that of the current electrodes at any stage during the operation of the probing [29]. As a rule of thumb, under ideal conditions, penetration depth is about 1/3 of the total AB distance. The injected current, measured potential and geometry of arrangement give the apparent resistivity of the strata at the center of the electrode array.
Field Work and Data Interpretation

In this study, electrical resistivity sounding data were obtained using two resistivity instruments: 1) StingR1 and 2) IRIS-Syscal resistivity meter while performing the 45 Vertical Electric Soundings (VES). The maximum separation of AB ranged from 250 to 500 m depending upon the availability of free space along the profile line. In most cases, field data were generally of high quality, with some exemptions, when the instrument’s battery had operating problems. Fieldwork required at least a three (but ideally a four) men brigade. Two men laid the cable, moved and stood by the two current electrodes A & B. A third man, remained at the center point, and was responsible for taking the measurements in the instrument; and a fourth man oversaw the movement of the M and N electrodes.

For the inversion and interpretation of the field-data, the IPI2WIN software was used [30]. This automatic curve-matching computer program results in a geoelectric model, where the calculated apparent resistivity matches the field curve almost exactly. However, a cautionary note should be stated: even if the interpretation as determined by the program is mathematically correct, it may not necessarily correspond to reality. In other words, due to the ambiguity inherent to geophysics, results may, on some occasions, tend to exceed the limitations of the VES methods. Nonetheless, the geoelectric model determined by the program generally helps in estimating the parameters of a four- or five-layer geoelectrical model [31]. The result of a VES via curve interpretation is a layered model of the shallow-subsurface with thicknesses (H) and apparent electrical resistivities (ρ) for each individual geoelectrical layer. These models can provide valuable information on the lithology of the shallow subsoil and/or salinity of the groundwater [32] [33] [34].

To improve interpretations, available geological and hydrogeological data were incorporated into the modeling and inversion results to make a sturdier and more realistic geoelectrical model.

5. Results and Discussion

Fifty-five Vertical Electrical Soundings, randomly distributed, were conducted in the north-central portion of El Sauz-Encinillas’ valley, their locations are shown in Figure 3. Data were collected and later fed to the IPI2WIN software. For the best results, the better fit was selected. The quadratic mean error (rms) in the one-dimensional inverse model of all the VES ranged between 0.5% and 25%.
Results depict a wide variability in resistivity values throughout the study area. After correlating the different individual analyses, five overall geoelectrical units were identified (Table 1).

Interpretation of resistivity data as discussed before, provided information about the apparent resistivity and thickness of the different geoelectrical units (Table 2).

**Geoelectric Cross Sections**

Five geoelectrical cross-sections (Figure 4) were constructed to reveal the lateral and vertical geoelectrical-lithological variations in the study area. Of these, two geoelectrical sections are oriented with a W-E trend (Sections 1 & 2), approximately perpendicular to the federal highway 45 (Chihuahua-Juárez city), while two other geoelectrical sections (Sections 4 & 5) are oriented with a N-S trend, parallel to highway 45. The remaining section (Section 3) has a N-SW direction.

The five sections provided insight into the subsurface sequence in the study area. All the geoelectrical sections were built with at least 5 and a maximum of 10 VES, from them the geoelectrical stratigraphy can be inferred (Figure 5).

Please note that not all the VES have the same depth penetration, so once the sections are made, they may not have the same depth throughout its entire length, as clearly depicted in geoelectrical Sections 2 and 4, in Figure 5.

Geoelectrical Section 1 has a W-E alignment, cutting the whole valley. It was built with 10 VES (1, 2, 4, 6, 19, 15, 9, 7, 12 and 32) crossing the alluvial fan’s sediments all the way through the Laguna de Encinillas’ sediments. It starts showing colluvial material, possibly related to mass wasting processes near Del Nido block, these coarse-grain materials show high resistivities. As you go through the whole section, fine grained material (sands and silt) starts to show and dominate sedimentation. At its easternmost part, an interdigitation of gravel/sandy material can be inferred from this section, as the sections gets away from the playa lake material and closer to colluvial materials from the northern ranges.
Table 1. Summary of geoelectrical units found within the El Sauz Encinillas valley.

| Geoelectrical Unit | Apparent Resistivity (in Ω-m) | Thickness (in m) | Inferred Geologic Material |
|--------------------|-------------------------------|------------------|-----------------------------|
| 1                  | 200 - 800                     | 0.43 - 15.6      | Hardpan topsoil             |
| 2                  | 25 - 100                      | 0.37 - 341.62    | Alluvial material (sand/silt), |
| 3                  | 0.2 - 15                      | 0.43 - 161.53    | Playa lake material (clay/evaporites), |
| 4                  | 100 - 300                     | 0.54 - 39.8      | Gravel/sand mixture         |
| 5                  | 400 - 3500                    | 1.03 - 26.3      | Partly fractured rock/conglomeratic material |

Table 2. Resistivity and thickness results for each VES using the IPI2WIN program.

| VES | Long. | Lat. | ρ1 | H1 | ρ2 | H2 | ρ3 | H3 | ρ4 | H4 | ρ5 | H5 | ρ6 | H6 |
|-----|-------|------|----|----|----|----|----|----|----|----|----|----|----|----|
| 1   | 358,056 | 3,248,552 | 937 | 1.52 | 355 | 6.2 | 91.4 | 6.32 | 524 | 17.7 | 38.9 | 35.2 | 684 | ∞   |
| 2   | 359,349 | 3,248,568 | 471 | 4.79 | 3572 | 4.67 | 107 | 14.6 | 4112 | 33.3 | 6.4 | ∞ | ∞ |
| 3   | 360,353 | 3,248,577 | 2377 | 3.7 | 306 | 39.8 | 1600 | ∞ |
| 4   | 361,225 | 3,248,652 | 132 | 4.22 | 475 | 5.4 | 28.7 | 13.7 | 169 | 45.2 | 0.37 | ∞ |
| 5   | 362,048 | 3,248,756 | 902 | 2.59 | 66.1 | 3.44 | 3626 | 9.61 | 11.3 | 29 | 5966 | ∞ |
| 6   | 367,098 | 3,249,341 | 518.1 | 2.87 | 167.5 | 8.52 | 67.2 | 77.45 | 49.83 | 147.8 | 380.4 | ∞ |
| 7   | 366,333 | 3,249,263 | 465 | 7.26 | 81 | 29.4 | 49.2 | 61.9 | 119 | ∞ |
| 8   | 365,483 | 3,249,186 | 463 | 3.26 | 178 | 4.69 | 63.8 | 337 | 1127 | ∞ |
| 9   | 368,586 | 3,249,985 | 448 | 4.6 | 42.4 | 8.73 | 601 | 18.3 | 4.5 | 31.8 | 2605 | ∞ |
| 10  | 368,646 | 3,249,488 | 59.7 | 2.97 | 304 | 2.35 | 65.4 | 39.8 | 55.3 | ∞ |
| 11  | 368,481 | 3,250,241 | 127.2 | 2.445 | 191.8 | 3.954 | 64.68 | 47.72 | 43.1 | ∞ |
| 12  | 367,942 | 3,250,748 | 851 | 4.07 | 74.3 | 38.1 | 16 | 48.7 | 5065 | ∞ |
| 13  | 363,074 | 3,248,001 | 235 | 7.94 | 54.4 | 38.4 | 1084 | ∞ |
| 14  | 363,575 | 3,249,026 | 190.5 | 2.10 | 354.6 | 3.93 | 75.57 | 14.9 | 51.77 | 126.3 | 120.1 | ∞ |
| 15  | 362,690 | 3,251,518 | 442 | 8.21 | 72.9 | 31.9 | 46.8 | 189 | 3667 | ∞ |
| 16  | 368,897 | 3,251,518 | 500 | 3.11 | 71.9 | 77.3 | 350 | ∞ |
| 17  | 364,591 | 3,249,159 | 398 | 7.9 | 58.1 | 150 | 11,399 | ∞ |
| 18  | 362,865 | 3,248,817 | 20.8 | 0.58 | 6.26 | 15.5 | 13.1 | 40.6 | 2.52 | 20 | 12 | ∞ |
| 19  | 376,324 | 3,257,229 | 5.19 | 1.23 | 39.6 | 0.80 | 6.17 | 10.2 | 10.2 | 30 | 1.33 | 28.8 | 23.4 | ∞ |
| 20  | 374,656 | 3,257,104 | 41.4 | 16.2 | 3.12 | 11 | 192 | 28.8 | 0.11 | ∞ |
| 21  | 372,653 | 3,260,369 | 9.1 | 0.43 | 69.3 | 0.58 | 4.83 | 12.2 | 13.9 | 57.5 | 8.4 | ∞ |
| 22  | 372,299 | 3,260,731 | 3.14 | 0.99 | 41.2 | 1.04 | 1.54 | 2.37 | 25.8 | 9.42 | 1.96 | 25.4 | 1996 | ∞ |
| 23  | 372,666 | 3,260,128 | 82.5 | 1.3 | 1.4 | 32.6 | 0.38 | 25.5 | 12.4 | ∞ |
| 24  | 372,430 | 3,257,052 | 47.2 | 1.12 | 1341 | 1.67 | 36 | 8.38 | 383 | 12 | 52.6 | ∞ |

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|   |   |   |   |   |   |   |   |   |   |
|---|---|---|---|---|---|---|---|---|---|
| 25 | 378,193 | 3,257,463 | 40.7 | 0.70 | 1311 | 1.01 | 81.8 | ∞ |   |
| 26 | 377,281 | 3,257,088 | 99   | 1.27 | 3.16 | 0.82 | 1.65 | ∞ |   |
| 27 | 374,205 | 3,257,609 | 10.9 | 0.92 | 2.38 | 6.55 | 10.9 | ∞ |   |
| 28 | 373,732 | 3,258,438 | 6.94 | 0.73 | 1.85 | 4.04 | 0.38 | 3.26 | 1.13 | 43.84 | 3.61 | ∞ |   |
| 29 | 372,841 | 3,257,128 | 15.23 | 0.91 | 688 | 3.51 | 98.2 | 15.2 | 28.5 | 22.9 | 137 | 46.9 | 0.42 | ∞ |   |
| 30 | 371,070 | 3,256,640 | 118  | 1.3  | 955 | 14.3 | 48.1 | ∞ |   |   |   |   |   |   |   |   |
| 31 | 377,700 | 3,254,295 | 710  | 1.02 | 1786 | 1.2  | 372 | 9.54 | 71.2 | 195 | 3.75 | ∞ |   |   |   |   |
| 32 | 364,665 | 3,252,962 | 162  | 0.91 | 688 | 3.51 | 98.2 | 15.2 | 28.5 | 22.9 | 137 | 46.9 | 0.42 | ∞ |   |
| 33 | 365,538 | 3,253,900 | 20.4 | 1.01 | 6.41 | 0.93 | 531 | 5.99 | 40.5 | ∞ |   |   |   |   |   |   |
| 34 | 367,089 | 3,255,009 | 51.7 | 1.05 | 7.18 | 1.29 | 269 | 3.13 | 28.3 | ∞ |   |   |   |   |   |   |
| 35 | 368,763 | 3,255,826 | 19.42 | 0.96 | 55.84 | 1.09 | 6.11 | 2.331 | 26.06 | 4.98 | 7.24 | ∞ |   |   |   |
| 36 | 370,933 | 3,256,597 | 1246 | 0.43 | 34.85 | 0.37 | 188.6 | 3.50 | 25.52 | 26.92 | 11.95 | ∞ |   |   |   |
| 37 | 375,545 | 3,257,360 | 233  | 1.71 | 62.5 | 16.1 | 8.71 | 32.1 | 8588 | ∞ |   |   |   |   |   |   |
| 38 | 370,231 | 3,270,164 | 5.54 | 0.54 | 0.49 | 0.59 | 1.98 | 12.4 | 41.2 | 13.4 | 1.75 | ∞ |   |   |   |
| 39 | 368,031 | 3,268,939 | 98.1 | 0.47 | 727 | 1.03 | 35.9 | 9.48 | 2.2 | 10.8 | 52.6 | ∞ |   |   |   |
| 40 | 365,349 | 3,264,618 | 139  | 0.52 | 884 | 1.11 | 48.3 | 6.32 | 13.3 | 62.5 | 30.7 | ∞ |   |   |   |
| 41 | 362,129 | 3,264,247 | 215  | 2.85 | 5.53 | 2.59 | 46,798 | ∞ |   |   |   |   |   |   |   |   |
| 42 | 362,132 | 3,256,991 | 53.3 | 0.99 | 359 | 1.13 | 47.9 | 11.9 | 11.2 | 14.9 | 417 | 26.3 | 0.82 | ∞ |   |
| 43 | 363,732 | 3,259,333 | 8317 | 2.5  | 99.9 | 2.19 | 8530 | 7.83 | 1.84 | ∞ |   |   |   |   |   |   |
| 44 | 362,091 | 3,257,336 | 146  | 0.55 | 2143 | 0.53 | 319 | 8.57 | 55.7 | ∞ |   |   |   |   |   |   |
| 45 | 362,610 | 3,252,361 | 128  | 0.61 | 2185 | 0.53 | 247 | 7.41 | 58.1 | 36 | 13.5 | 19.3 | 75.6 | ∞ |   |   |

**Figure 4.** Location of all the geoelectrical sections built with VES data. Source: “Laguna de Encinillas” 3258213.90 N and 376327.93 W. Google Earth. January 26, 2017.
Figure 5. Geoelectrical sections built from the VES data using the IPI2WIN software.
Geoelectrical Section 2 also has a W-E trend. Made up by 10 VES (2, 16, 33, 35, 36, 37, 30, 28, 20 and 26), it is the longest section, since it starts at near Del Nido Block and ends when almost reaching the San Martin range. It starts at its western margin material with high resistivity material possibly related to colluvial material at the apex of the alluvial fans. As the section goes toward the East, resistivity values slightly decrease due to the gravelly/sandy material in the middle part of the fan. In this part sandy material is detected. As the section gets closer to the Laguna the finer playa lake deposits’ material is detected, but again at its easternmost section again colluvial material from the San Martin range dominates at its upper part seemingly interdigitating with the finer (less resistive material). The rock basement was not found since resistivity values throughout the section indicate a dominance of granular material.

Geoelectrical Section 3 has a S-N orientation, it consists of 9 VES (14, 15, 16, 45, 44, 43, 41, 40 and 39). It starts parallel to Del Nido block and later stats to bend towards the NE direction and ends surrounding the Laguna de Encinillas. It displays a drastic change from the upper coarse sediments (hard& compacted top soil) and colluvial material possibly related to the alluvial fans parallel to the Del Nido Block, to a finer sandy material at depth. More resistive material possibly related to coarse materials is found at its central portion. This behavior is interesting and should be studied with more detail in further work. As you get closer to its northernmost section a dominance of finer material is found, as the SEVS are closer the Laguna’s sediments.

Geoelectrical Section 4, has a S-N trend in the western portion of the El Sauz Encinillas’ valley. It was built with data from 7 VES (10, 11, 12, 13, 35, 43 & 42) in the valley’s central portion. In it, sandy material predominates throughout the section. Coarser materials are present at its south-central tip, possibly correlated to colluvial materials near the alluvial fans parallel to the mountains. As you go to the north the finer-grained material near the Laguna de Encinillas is clearly seen, but as you get to the northern part of the section sandy and maybe even silty material again appears.

Finally, geoelectrical Section 5, also with a S-N direction it was made up with 7 VES (32, 21, 28, 29, 22, 23 and 40) in the eastern part of the valley, focusing in the Laguna sediments. In its southern part some coarse material was detected (high resistivity values) but as you go north the fine-grained material (low resistivity values) dominate as expected both in extension and at depth throughout the whole section. In both its southern and northernmost portions, high resistivity units are detected in the upper part of the section. But again, no material possibly related to the basement was detected.

Something to notice in all sections in Figure 5, are the interdigitations of coarse and sandy materials close to the alluvial fans in the range-valley transitional zones. The playa lake deposits and their low resistivities associated with the evaporites and clays were detected in the Laguna. Also relevant is how the basement could not be found, implying that the normal faults that made up this typical basin and range depocenter must have a very steep angle and/or the
basement is deeper and out of reach with the equipment and AMNB array used.

6. Conclusions and Recommendations

The shallow sub-surface stratigraphy was delineated through investigations conducted via electrical resistivity surveys. After correlating the different individual analyses of each VES, five overall geoelectrical units were identified: 1) hardpan topsoil, with resistivity values ranging from 200 - 800 Ω-m; 2) alluvial material mixture (sand/silt) with resistivity values ranging from 25 to 100 Ω-m; 3) playa lake type material (clay/evaporites) with resistivity values ranging from 0.2 to 15 Ω-m; 4) gravel/sand mixture with resistivity values from 100 to 300 Ω-m; and 5) partly fractured rock/conglomeratic material with resistivity values ranging from 400 to 3500 Ω-m.

From the geoelectrical sections, a decrease in resistivity data values can be seen, possibly related to a decrease in its granulometry (from colluvial material to sandy sediments related to alluvial fans), both frontally and laterally in the areas flanking the Laguna de Encinillas. Resistivity values (much lower) as well as both permeability and the water quality in its distal portion, seem to be affected by the playa lake deposits, due to a raising ratio of clay-size sediments (and evaporites) in the center of the valley, near the Laguna de Encinillas. The basement was not detected in the SEVs, either implying that normal faults that created this depocenter must have a very steep angle and/or the basement is deeper and out of reach with the equipment and AMNB array used.

The electrical resistivity data, therefore gives reasonably accurate results that can be used to understand the subsurface stratigraphy and when possible, structures and the basement configuration, so valuable in groundwater exploration/management. These results must to be complemented by using another geophysical technique with greater depth capabilities to be able to locate the rock basement in the valley. This would allow to build a more robust conceptual model of the aquifer, since data from its southern section is available.

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Conflicts of Interest

The authors declare no conflicts of interest regarding the publication of this paper.

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