LegacyClimate 1.0: A dataset of pollen-based climate reconstructions from 2594 Northern Hemisphere sites covering the late Quaternary

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Here we describe the LegacyClimate 1.0, a dataset of the reconstruction of mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$), and annual precipitation ($P_{ann}$) from 2594 fossil pollen records from the Northern Hemisphere spanning the entire Holocene with some records reaching back to the Last Glacial. Two reconstruction methods, the Modern Analogue Technique (MAT) and Weighted-Averaging Partial-Least Squares regression (WA-PLS) reveal similar results regarding spatial and temporal patterns. To reduce the impact of precipitation on temperature reconstruction and vice versa, we also provide reconstructions using tailored modern pollen data limiting the range of the corresponding other climate variable. We assess the reliability of the reconstructions using information from the spatial distributions of the root-mean squared error of prediction and reconstruction significance tests. The dataset is beneficial for climate proxy synthesis studies and to evaluate the output of climate models and thus help to improve the models themselves. We provide our compilation of reconstructed $T_{July}$, $T_{ann}$, and $P_{ann}$ as open-access datasets at PANGAEA (https://doi.pangaea.de/10.1594/PANGAEA.930512; Herzschuh et al., 2021). R code for the reconstructions is provided at Zenodo (https://doi.org/10.5281/zenodo.5910989; Herzschuh et al., 2022), including harmonized open-access modern and fossil datasets used for the reconstructions, so that customized reconstructions can be easily established.

1 Introduction

The evaluation of climate model outputs using climate data is essential for model improvements (Eyring et al., 2019). However, the period for which observations are available is only of limited use to validate simulations because it is short and characterized by strong changes in the climate driver. Climate proxy data derived from natural archives are therefore of great value. The extratropical Northern Hemisphere is of particular interest because it is known for complex spatial and temporal temperature and precipitation patterns.

Previous proxy-based climate inferences have contributed to major debates about Holocene climate change. For example, while simulations indicate a gradual warming of the Holocene, temperature proxy data syntheses rather support a mid-Holocene optimum which resulted in the “Holocene conundrum”
debate (Liu et al., 2014). Qualitative proxy-based inferences indicate that the mid-Holocene in the Northern Hemisphere mid-latitudes was rather dry and warm compared with present-day in agreement with modeling outputs (Routson et al., 2019). Also, quantitative precipitation reconstructions from Eastern and Central Asia unveiled the complex monsoon-westerlies interactions (Chen et al., 2019; Herzschuh et al., 2019).

Fossil pollen records are well-established in their use as a palaeoecological and palaeoclimatological proxy and of great value as indicators of past environmental and climatic change for many decades. Considerable efforts have been made to establish regional, continental and even global data repositories like the North American Pollen Database (http://www.ncdc.noaa.gov/paleo/napd.html), the European Pollen Database (http://www.europeanpollendatabase.net) and the Neotoma Paleoecology Database (https://www.neotomadb.org; Williams et al., 2018). Regarding the prevalence of pollen archives across multiple environmental settings such as lakes, wetlands, or marine sediments, fossil pollen records are widely used to quantitatively reconstruct past vegetation and climate variables (Birks, 2019; Chevalier et al., 2020). Pollen data are the only land-derived proxy data that have sufficient temporal and spatial coverage to allow for high-resolution climate model evaluation of the late Quaternary period. A number of methods have been proposed for making pollen-based climate reconstructions (Chevalier et al., 2020): among them, classification approaches like the Modern Analogue Technique (MAT) or regression approaches like Weighted-Averaging Partial-Least Squares regression (WA-PLS) are most commonly used.

For temperature reconstruction time-series, several broad-scale syntheses exist; however, either they originate from different proxies (Kaufman et al., 2020a and 2020b) or are restricted to certain continents or regions (Mauri et al., 2015; Marsicek et al., 2018; Routson et al., 2019). Temperature reconstructions from the large extratropical Asia are mostly lacking. Precipitation syntheses are available from Europe (Mauri et al., 2015), North America (Whitmore et al., 2005) and China and Mongolia (Herzschuh et al., 2019) but, hitherto, no global or hemispheric syntheses of quantitative precipitation changes are available for the Holocene.

In a recent effort, pollen records available in the Neotoma Paleoecology Database (Williams et al., 2018) and additional records from China and Siberia (Cao et al., 2013 and 2020) were synthesized and taxonomically harmonized (Herzschuh et al., submitted). Furthermore, all chronologies of these records were recently revised using a Bayesian approach that allows for the inference of temporal uncertainties.
Here we present the pollen-based reconstruction of mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$) and annual precipitation ($P_{ann}$) from these 2594 records from the Northern Hemisphere using WA-PLS and MAT.

2 Methods

2.1 Input data

The objective of this study is to create a dataset of quantitative reconstructions of $T_{July}$, $T_{ann}$ and $P_{ann}$ spanning the Holocene from a set of fossil pollen records. We used fossil data from the Neotoma Paleocoeology Database (Williams et al., 2018; https://www.neotomadb.org; downloaded in July 2020), a dataset from Eastern and Central Asia (Cao et al., 2013; Herzschuh et al., 2019) and a dataset from Northern Asia (Cao et al., 2020). The harmonized dataset is stored on PANGAEA (LegacyPollen 1.0) and presented in Herzschuh et al. (submitted). Ages were taken from age-depth models presented in Li et al. (2022), who recently provided a set of harmonized chronologies under the “LegacyAge 1.0” framework, and applied to our fossil pollen synthesis. A modern pollen training dataset comprised of 15,379 sites includes datasets from Eurasia (EMPD1, Davis et al. 2013; EMPD2, Davis et al. 2020; Herzschuh et al., 2019; Tarasov et al., 2011) and North America (Whitmore et al., 2005). In order to reduce inconsistencies in pollen identification, the modern and fossil pollen datasets were taxonomically harmonized: major tree and shrub pollen were merged to genus level and most of the herbaceous taxa (except the most common ones such as Artemisia, Thalictrum or Rumex) to family level. We excluded aquatic pollen (with the exception of Cyperaceae), spores from ferns and fungi, as well as algae and calculated pollen percentages on the basis of the total number of terrestrial pollen grains. The site specific $T_{ann}$, $T_{July}$, $P_{ann}$ were derived from WorldClim 2 (spatial resolution of 1 km, https://www.worldclim.org, Fick and Hijmans, 2017) by extracting the climate data at the location of the modern sample sites using the raster package in R (version 3.5-11, Hijmans et al., 2021; R Core Team, 2020).

We compiled the fossil data into four sub-continental datasets for Eastern North America (<105°W; Williams et al., 2000), Western North America, Europe and Asia. For consistency with the amount of taxa in the North American training dataset, the fossil datasets were reduced to the 70 most common
taxa on the respective sub-continents, according to Hill’s N² diversity index (i.e., the effective number of occurrences of a species in the dataset; Hill, 1973).

2.2 Reconstruction methods

Our reconstruction approach included MAT (Overpeck et al., 1985) and WA-PLS (ter Braak and Juggins, 1993) by applying the `MAT` and `WAPLS` functions from the `rioja` package (version 0.9-21, Juggins, 2019) for R (R Core Team, 2020) on our Northern Hemispheric fossil pollen synthesis. For each fossil location, we calculated the geographic distance between each modern sampling site and the fossil pollen record using the `rdist.earth` function from the `fields` R-package (version 10.3, Nychka et al., 2020) and selected a unique calibration set from modern sites within a 2000 km radius. For the reconstruction with MAT, we used the original pollen percentages of the selected fossil pollen taxa, looking for 7 analogues between the pollen data and the selected calibration dataset. The dissimilarity between the fossil samples and the modern pollen assemblages was determined by squared-chord distance metrics (Simpson, 2012; Cao et al., 2014). For the reconstruction with WA-PLS, we used the square-root transformed pollen percentages in a leave-one-out cross-validation approach (Cao et al., 2014). In addition to the classic WA-PLS reconstruction, we provide WA-PLS_tailed. To reconstruct $T_{July}$ we “tailored” our modern training dataset with respect to the $P_{ann}$ range. For this purpose, we identified the range of the reconstructed $P_{ann}$ and extended it by 25% to both ends of the modern $P_{ann}$ range in order to reduce the influence of $P_{ann}$ on $T_{ann}$ and $T_{July}$ reconstruction due to co-variation. We applied the same method to the reconstruction of $P_{ann}$; $T_{ann}$ and $T_{July}$ were tailored by $P_{ann}$; $P_{ann}$ was tailored by $T_{July}$ and, additionally, by $T_{ann}$ (illustrated for an example in Appendix Fig. 1). A statistical significance test (Telford and Birks, 2011) was performed for the reconstruction by using the `randomTF` function in the `palaeoSig` R-package (version 2.0-3, Telford, 2019). The reconstructed climate parameters were tested as single variables, as well as with partialling out the respective other variable. We applied a Canonical Correlation Analysis (CCA) to the modern training dataset in order to infer the explained variance in the modern dataset by using the `cca` function in the `vegan` R-package (version 2.5-7, Oksanen et al., 2020). The ratio between constrained ($\lambda_1$) and unconstrained ($\lambda_2$) explained variance was determined for all modern training datasets.
Dataset description LegacyClimate 1.0: input data, reconstructions and reconstruction model statistics

LegacyClimate 1.0 provides pollen-based reconstructions and sample-specific reconstruction errors of $T_{ann}$, $T_{July}$ and $P_{ann}$ for 2594 fossil pollen records (i.e., a total of 146,067 single pollen samples) from three reconstruction methods (WA-PLS, WA-PLS_tailored, MAT). Furthermore, we provide the method-specific model metadata and statistics for each record and each climate variable (Table 1). To ease data handling, the dataset files are separated into Western North America, Eastern North America, Europe and Asia.

**Table 1.** Structure and content of the LegacyClimate 1.0 data with details about the information contained in the input, datasets, in the reconstructions and the reconstruction model statistics.

| Datasets | Content |
|----------|---------|
| **Input datasets** | Modern pollen dataset of 15,379 sites |
| | Modern dataset of $T_{ann}$, $T_{July}$, $P_{ann}$ |
| | Fossil pollen data (LegacyPollen 1.0) for 2594 sites with a total of 146,067 samples |
| **LegacyClimate 1.0: Climate reconstructions** | Reconstructions and sample-specific reconstruction errors of $T_{ann}$, $T_{July}$ and $P_{ann}$ for 2594 sites using MAT, WA-PLS and WA-PLS_tailored |
| **LegacyClimate 1.0: Reconstruction model statistics** | **Site information** (Event label, Source, ID, Site name, Longitude, Latitude) |
| | **Modern pollen dataset information** (number of modern analogues, range of climate variables) |
| | **Model statistics for each site for MAT, WA-PLS, WA-PLS_tailored** (including $r^2$ observed vs. predicted, RMSEP, no. of WA-PLS components) |
4 Dataset assessment

4.1 Spatial and temporal coverage of LegacyClimate 1.0

In total, we provide reconstructions for 2594 fossil pollen records, among them 670 records from Eastern North America, 361 records from Western North America, 1075 records from Europe and 488 Asian records (Fig. 1). The temporal coverage of the records is rather uneven: 119 and 289 records cover the periods before 30,000 years (Fig. 2) and the Last Glacial Maximum, respectively. A total of 1229, 1845, 2052 records are available for 12-11 ka, 6-5 ka BP and 2-1 ka BP, respectively.

Figure 1. left: map indicating the spatial distribution and record lengths covered by the LegacyPollen 1.0 dataset (Herzschuh et al., submitted) for which climate reconstructions are provided in LegacyClimate 1.0 with a total of 2594 records; right: spatial distribution of modern pollen dataset used for reconstruction with a total of 15,379 sites.

Figure 2. Number of records that cover certain millennia of the last 30 ka.
4.2 Prediction errors of LegacyClimate 1.0

The mean RMSEPs and their standard deviations for $T_{ann}$ are 1.98±0.52°C (MAT), 2.61±0.53°C (WA-PLS) and 2.24±0.61°C (WA-PLS_tailored) and mean RMSEPs as a percentage of modern $T_{ann}$ range are 7.68±1.93% (MAT), 10.09±2.05% (WA-PLS) and 10.26±2.79% (WA-PLS_tailored). The largest mean RMSEP values are located in Central Asia in Kazakhstan, Mongolia and the north-western parts of the Tibetan Plateau and are consistent across all three reconstruction methods. Other areas with large mean RMSEP values are located in Western North America, Southern and Central Europe and south-east Asia. The smallest RMSEPs can be found along the east coast of North America. Relative to the modern temperature range, the RMSEP from this region also reveals the lowest fraction. In general, MAT has the lowest mean error fraction relative to the modern temperature range of all three methods.

The mean RMSEPs of $T_{July}$ are 1.90±0.63°C (MAT), 2.50±0.73°C (WA-PLS) and 2.21±0.75°C (WA-PLS_tailored) and mean percentages of $T_{July}$ range are 8.11±1.64% (MAT), 10.71±1.94% (WA-PLS) and 10.70±2.60% (WA-PLS_tailored). Thus, they are slightly smaller than those of $T_{ann}$ but slightly larger as a percentage of the range. The spatial patterns, however, are largely similar to those of $T_{ann}$.

The mean RMSEPs of $P_{ann}$ are 176.38±51.40 mm (MAT), 244.48±75.84 mm (WA-PLS) and 232.71±98.57 mm (WA-PLS_tailored) and mean percentages of $P_{ann}$ range are 6.78±1.48% (MAT), 9.27±1.70% (WA-PLS) and 10.26±2.67% (WA-PLS_tailored). High RMSEPs are found for Western North America, Europe and along the coastline of south-east Asia, while the lowest RMSEP values are found for Central Asia. A clear division in RMSEPs are found on the North American continent: while the western part of North America (with the exception of Alaska) has a rather high RMSEP, the eastern part of North America has a smaller RMSEP. This pattern is found for all three methods (Fig. 3).
Figure 3. Spatial distribution of root mean squared error of prediction (RMSEP) as inferred from leave-one out cross-validation presented as absolute values and as a percentage of the range of mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$), mean annual precipitation ($P_{ann}$) in the modern pollen data used for reconstruction for the three methods applied (weighted-averaging partial least
squares (WA-PLS), WA-PLS using a training set from within a limited climate range (WA-PLS_tailored) and modern analogue technique (MAT)).

A significance test ($p < 0.1$) according to Telford and Birks (2011) for the whole reconstructed time period was run for each record and for the reconstructions with WA-PLS and WA-PLS_tailored (Fig. 4; Table 2). The $T_{\text{July}}$ reconstruction is significant for 30.9% (WA-PLS) and 35.2% (WA-PLS_tailored) when included as a single variable in the significance test. Partialling out precipitation as a conditional variable causes an increase in the amount of significant records to 35.5% for WA-PLS, but a decrease for WA-PLS_tailored to 33.6% of all records. For $T_{\text{ann}}$, 32.8% (WA-PLS) and 36.1% (WA-PLS_tailored) of all records pass the significance test when tested as a single variable. When partialling out precipitation, the amount of significant records decreases for both WA-PLS and WA-PLS_tailored. 32.1% (WA-PLS) and 33.4% (WA-PLS_tailored) of all records pass the significance test when testing $P_{\text{ann}}$ as a single variable. In contrast to the significance tests for $T_{\text{ann}}$, partialling out the mean July temperature as a conditional variable increases the number of significant records for both WA-PLS and WA-PLS_tailored.
Figure 4. Maps showing mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$), mean annual precipitation ($P_{ann}$) records that passed the reconstruction significance test ($p<0.1$). Color indicates the significance level.

Table 2. Percentage of records that pass the reconstruction significance test ($p<0.1$) sensu Telford and Birks (2011).

|                | WA-PLS  | WA-PLS_tailored |
|----------------|---------|-----------------|
| $T_{July}$     | 30.9%   | 35.2%           |
| $T_{July}$ partialling out $P_{ann}$ | 35.5%   | 33.6%           |
| $T_{ann}$      | 32.8%   | 36.1%           |
| $T_{ann}$ partialling out $P_{ann}$ | 32.6%   | 34.1%           |
| $P_{ann}$      | 32.1%   | 33.4%           |
| $P_{ann}$ partialling out $T_{July}$ | 34.3%   | 36.5%           |

4.3 Assessment of major temporal patterns of LegacyClimate 1.0

For analyzing the temporal variation, we calculated means of all three climate variables for the time periods between 6.5 and 5.5 ka BP and between 1.5 and 0.5 ka BP and subtracted those means from every record in order to evaluate the changes between the reconstructed mid-Holocene conditions and those of modern times. Differences between these time periods reveal warmer and drier conditions especially in Eastern North America but also in Central and Northern Europe. The overall patterns are in good agreement for all three methods but show differences on a regional scale, especially when comparing the reconstructions with WA-PLS and MAT. For $T_{July}$, the reconstruction with MAT shows greater temperature differences in Western North America and south-east Asia. Compared to the reconstruction with WA-PLS, there is a reduced cooling in Eastern Europe and a warming in the Western Mediterranean region and along the south-eastern Asian coastline. Comparing the reconstructions of $T_{ann}$, more gradual patterns are seen in the reconstruction with WA-PLS: Western North America reveals a mid-Holocene warming, while Eastern North America shows a cooling. In Europe records that report a cooling are more concentrated in the northern and western parts of the continent. In the reconstruction with MAT, Eastern North America is divided into a reported cooling in the northern part and a warming...
in the southern part. In Western North America, there is a mixture of locations with a warming and a cooling since the mid-Holocene. In Europe, only France and Southern Scandinavia show a cooling; in Central and parts of Southern Europe, a warming can be found in the reconstructions. For large areas in North America and Europe, the reconstructions with WA-PLS suggest an increase in precipitation since 6 ka BP. A shift to drier conditions can be found along the south-eastern coastline in North America, in the Mediterranean Region and especially in south-east Asia. The reconstruction with MAT reveals a gradient from increasing precipitation in south-western Europe to decreasing precipitation in north-eastern Europe. In contrast to the reconstructions with WA-PLS, records along the south-eastern Asian coastline suggest an increase in precipitation with MAT rather than a decrease (Fig. 5).

**Figure 5.** Difference from 6 ka to 1 ka for mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$), mean annual precipitation ($P_{ann}$) as reconstructed from weighted-averaging partial least squares (WA-PLS), WA-PLS using a training set from within a limited climate range (WA-PLS_tailored) and modern analogue technique (MAT).

Time-series of absolute $T_{ann}$ reconstructions reveal temporal as well as latitudinal spatial variation on the single continents. Eastern North America and Asia show the most variation in the low latitudes. It is also Eastern North America which shows the most pronounced latitudinal gradient. In Western North America, the most variation takes place in the high latitudes, while the variation is concentrated to the
mid-latitudes in Europe. Especially in North America, the warming since the last deglaciation and the beginning of the Holocene is well shown in the temporal variation of the time-series (Fig. 6).

**Figure 6.** Time-series of absolute mean annual temperature ($T_{\text{ann}}$) reconstruction for each (sub-)continent. Colors denote the latitude of record origin. Note logarithmic x-axis.

### 4.4 Assessment of consistency among reconstruction methods

Reconstructions with MAT are, in general, in good agreement with those derived from the WA-PLS. Comparing MAT with WA-PLS, 37.3% ($T_{\text{July}}$), 38.9% ($T_{\text{ann}}$) and 30.4% ($P_{\text{ann}}$) of all records have a positive correlation of $r \geq 0.6$. Strong positive correlations ($r \geq 0.9$) can mainly be identified in Eastern North America, while weak correlation can be found for large areas in central North America and most of Europe (Fig. 7).
Figure 7. Correlation between time-series of the 3 different reconstruction methods used – weighted-averaging partial least squares using a global training set (WA-PLS), WA-PLS using a training set with a limited modern climate range (WA-PLS_tailored) and the modern analogue technique (MAT) for the three climate variables of mean July temperature ($T_{July}$), mean annual temperature ($T_{ann}$) and mean annual precipitation ($P_{ann}$).

WA-PLS_tailored used a reduced modern training dataset (illustrated for an example in Appendix Fig. 1). The tailoring successfully reduced the co-variation of temperature and precipitation in the modern dataset as indicated by the distribution of the correlation coefficient in Fig. 8. Nevertheless, the obtained reconstructions are largely consistent between WA-PLS and WA-PLS-tailored: a correlation of $r >= 0.9$ is found for 59.2% of all records for $T_{July}$, 60.7% for $T_{ann}$ and 56.5% for $P_{ann}$.

Figure 8. Violin plot of the correlation coefficients between $T_{July}$ and $P_{ann}$ in the 15,379 training datasets used for the reconstructions. Left: used for WA-PLS reconstructions; middle: WA-PLS $T_{July}$-tailored (used for the reconstruction of $P_{ann}$); WA-PLS pann-tailored (used for the reconstruction of $T_{July}$).

A CCA was performed to infer the ratio between constrained and unconstrained explained variance for all modern training datasets ($\lambda_1/\lambda_2$) for the modern datasets used for WA-PLS and WA-PLS_tailored. Modern datasets used for WA-PLS constrained by $T_{July}$ reveal a concentration of high ratios in Eastern North America while low ratios can be found in Central Asia. While the spatial pattern of $\lambda_1/\lambda_2$ constrained...
by $T_{ann}$ is similar, the ratios are slightly higher for $T_{ann}$ than for $T_{July}$. Reconstructions for $P_{ann}$ show low
ratios in Europe and Eastern North America. Areas with high ratios are concentrated in Alaska and East
Asia.

**Figure 9.** Maps showing $\lambda_{1}/\lambda_{2}$, representing the ratio of explained variance of first axis (constrained) vs.
second (unconstrained) axis as revealed by applying a CCA to all modern training datasets that were
used for the reconstructions. Constraining variable as well as tailoring of the dataset (see methods) is
indicated in the map captions.

### 5 Discussion

#### 5.1 Impact of the fossil pollen data source on LegacyClimate 1.0 quality

LegacyClimate 1.0 contains reconstructions of climate variables from fossil pollen data derived from
open-access data repositories. The fossil records were derived from multiple natural archives, most
commonly, assemblages from continuous lacustrine and peat accumulations (Herzschuh et al.,
submitted). Different sizes of lakes and peat areas result in varying sizes of pollen source areas and
thus the spatial representativeness of a record, as small lakes and peatlands are considered to provide
information about the (extra-)local scale, while pollen assemblages from large lakes are considered as
a regional signal (Jackson, 1990; Sugita, 1993). However, such signals might be impacted by
taphonomy of the record, for example pollen from azonal riverine vegetation might be over-represented
in fluvially impacted pollen records.

Our dataset is based on taxonomically harmonized modern and fossil pollen datasets using a restricted
number of taxa (i.e., the most common 70 taxa on each (sub-)continent). Such an approach guarantees
that all records are handled consistently. Although losing taxonomic information when merging taxa together into a higher taxonomic level, it also increases the possibility of matching climate analogues in the modern and the fossil datasets. However, one needs to keep in mind that species with different ecological requirements may be merged together into one genus or family, for example, *Pinus* species that are restricted to tropical or subtropical areas in China or ones that grow in boreal forests (Cao et al., 2013; Tian et al., 2017).

Along with the pollen assemblages, data repositories also provide chronological information for fossil records. The quality of such chronologies varies strongly with respect to dating methods, calibration and numerical algorithms for determining an age-depth relationship (Blois et al., 2011; Trachsel and Telford, 2017). Having accurate and precise chronologies is thus of pivotal importance for reconstructing past climate in order to identify temporo-spatial patterns and therefore in helping to evaluate climate model outputs. The advantage of the fossil pollen dataset used for the reconstruction presented here (i.e., LegacyPollen 1.0; Herzschuh et al., submitted) is that it has harmonized chronologies (LegacyAge 1.0) along with information about uncertainties as well as related metadata and scripts that allow a customized re-establishment of the chronologies (Li et al., 2022). This, for example, allows the calculation of the temporal uncertainties when presenting reconstruction uncertainties of a specific time-slice.

5.2 Modern pollen and climate data sources and LegacyClimate 1.0 quality

Palaeoclimate reconstruction methods such as MAT and WA-PLS rely on extensive collections of modern training data. Designing a robust calibration dataset from modern pollen assemblages is a crucial part of the reconstruction process. A suitable calibration dataset should cover a wide range of climatic and environmental gradients in order to represent an empirical relationship between pollen assemblages and climate (Birks et al., 2010; Chevalier et al., 2020). Like with fossil pollen records, data syntheses and repositories also exist for modern surface pollen data. Most of the records in our modern dataset were compiled from well-established pollen assemblages from North America (Whitmore et al., 2005), Eurasia (Davis et al., 2013 and 2020) and China (Cao et al., 2013; Herzschuh et al., 2019). For fossil pollen records in areas with an insufficient coverage of modern surface pollen samples (e.g., Central Asia or Western Siberia), it might be difficult to create a calibration dataset that maps the required variety of environmental and climatic gradients and therefore find enough modern analogues.
for reconstructions with a classification approach such as MAT. Our routine uses the modern pollen data from within a radius of 2000 km around the site of the fossil record. The information provided in the reconstruction metadata including number of modern pollen samples and ranges of reconstructed variables, allows an assessment of the modern dataset used for reconstruction.

5.3 Reconstruction method and LegacyClimate 1.0 quality

Climate reconstruction methods all have different strengths and weaknesses. MAT and WA-PLS used in this study heavily rely on extensive collections of modern assemblage data covering diverse climatic and environmental gradients and are applicable on a broad spatial scale. However, both methods may struggle with complex species responses, are sensitive to spatial autocorrelation, can only deal with a certain extent of non-analogous situations and may produce poor results in so-called “quantification deserts” (Chevalier, 2019), where fossil pollen is hardly preserved or nearby modern surface pollen samples are missing (Chevalier et al., 2020). Nonetheless, for reconstructions on a local or regional scale, MAT and WA-PLS are most commonly used in climate reconstructions. The format of the modern and fossil datasets as well as the provided scripts could also be easily adapted to apply to other reconstruction methods such as CREST, a Bayesian approach that combines presence-only occurrence data and modern climatologies to estimate the conditional response of a given taxon to a climate variable (Chevalier et al., 2014 and 2020).

Through numerous physical processes that vary with both location and time, temperature and precipitation are interconnected, especially within the extratropical regions (Adler et al., 2008; Trenberth, 2011) and thus temperature and precipitation may not be treated as independent variables. Due to the numerical mechanisms in the transfer function, the correlation between both climate variables may reduce the reliability of the reconstructions. This is especially true for regions with a temperature-moisture driven circulation system such as the East Asian Summer Monsoon (EASM) that can heavily affect precipitation patterns in certain regions (Herzschuh et al., 2019). With our tailoring approach we are able to reduce the influence of co-variation of these two climate variables for the reconstruction and increase the number of records that pass a significance level of $p < 0.1$ (Telford and Birks, 2011).
5.4 Potential use of LegacyClimate 1.0

Our fossil pollen synthesis contains records from all over the Northern Hemisphere extratropics and thus can be used to infer spatio-temporal patterns in climate reconstructions that are not only limited to a local or regional scale. Although several hemispheric or global reconstruction studies exist, they have been largely restricted to temperature or have included relatively few records (Marcott et al., 2013; Marsicek et al., 2018; Routson et al., 2019; Kaufman et al., 2020a and 2020b). Our dataset is therefore a valuable addition. It may be used in a multi-proxy approach, synthesizing marine and terrestrial records in order to assess temperature development during the Holocene and can help to highlight possible interdependencies between oceans and land masses and such contribute to the “Holocene conundrum” debate (Liu et al., 2014). Temperature reconstructions from proxy data indicate peak temperatures during the Holocene Thermal Maximum around 6000 years BP followed by a pronounced cooling trend toward the late Holocene (Liu et al., 2014; Bova et al., 2021), which is also visible in our pollen-based reconstructions (Fig. 6). In contrast, climate models simulate a monotonic warming throughout the Holocene, which resulted in the “Holocene conundrum” debate (Liu et al., 2014). Temperature reconstructions are often derived from sea-surface temperatures as either mean annual temperatures (Birks, 2019; Bova et al., 2021) or global mean surface temperature (Marcott et al., 2013; Marsicek et al., 2018; Kaufman et al., 2020a and 2020b). However, it is argued that proxy-based climate reconstructions are seasonally biased and therefore might be the reason for the observed proxy-model divergence (Liu et al., 2014; Rehfeld et al., 2016; Bova et al., 2021). In this respect, it might help that we provide $T_{\text{July}}$ along with $T_{\text{ann}}$ reconstructions, which provides the opportunity to assess seasonal impacts on the reconstruction.

So far, reconstructions of precipitation have not been implemented on a hemispheric scale. The interconnection between temperature and precipitation (Trenberth, 2011) and its spatio-temporal variation across the Northern Hemisphere is therefore an important aspect of evaluating climate models (Wu et al., 2013; Hao et al., 2019; Herzschuh et al., submitted). A broad-scale quantitative reconstruction of temperature and precipitation would therefore be of great value for evaluating transient model runs performed by climate models such as TraCE 21k (He, 2010).

6 Data and code availability

The compilation of reconstructed $T_{\text{July}}$, $T_{\text{ann}}$, and $P_{\text{ann}}$, is open access and available at PANGAEA (https://doi.pangaea.de/10.1594/PANGAEA.930512; in the “Other version” section; Herzschuh et al.,...
2021). The dataset files are stored in machine-readable data format (.CSV), which are already separated into Western North America, Eastern North America, Europe, and Asia for easy access and use.

The R code to run the reconstructions for single sites is available at Zenodo (https://doi.org/10.5281/zenodo.5910989; Herzschuh et al., 2022) including harmonized open-access modern and fossil pollen datasets so that customized reconstructions can be easily established.

**Author contributions.** UH designed the study design and reconstruction dataset. CL and TB compiled the metadata and the harmonized pollen dataset. TB wrote the R scripts and ran the analyses under the supervision of UH. UH, TB and MC wrote the first draft of the manuscript. All authors discussed the results and contributed to the final manuscript.

**Competing interests.** The authors declare that they have no conflict of interest.

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**Appendix Figures**

![Mean July Temperature Reconstruction](image1)

![Annual Precipitation Reconstruction](image2)

![Temperature-Tailoring](image3)

![Precipitation-Tailoring](image4)
Appendix Figure 1. Example to illustrate the effect of tailoring the modern dataset for the location “Yellow Dog Pond” in Eastern North America. Upper part: reconstruction of $T_{July}$ and $P_{ann}$ with WA-PLS (red) and WA-PLS_tailored (blue); lower part: correlation of $T_{July}$ and $P_{ann}$ in the modern dataset and the effect of tailoring the modern dataset (indicated with the red box). Correlations are given for non-tailored (red) and tailored (blue) data.