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Variation of the BrO/SO$_2$ Molar Ratio in the Plume of Tungurahua Volcano Between 2007 and 2017 and Its Relationship to Volcanic Activity

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Recent long-term observations of the bromine monoxide (BrO) to sulphur dioxide (SO$_2$) molar ratio in volcanic plumes have suggested a link between changes in the BrO/SO$_2$ ratio and the volcanic activity. Nevertheless, understanding of the mechanisms determining this link is still limited due to the lack of studies on volcanic bromine release from the melt into the atmosphere. We present the results of 10 years (2007–2017) of observations of the BrO/SO$_2$ molar ratio in the volcanic plume of Tungurahua volcano, Ecuador. Following the nearly continuous eruptive activity from 1999 to 2008, Tungurahua showed alternating phases of eruptive activity separated by periods of quiescence between late 2008 and March 2016, after which degassing intensity decreased below detection. By comparing the BrO/SO$_2$ molar ratios collected from 13 eruptive phases to volcanic activity, this study aims to broaden the global observational database investigating their link. For this purpose, we combine three different methods to retrieve the BrO/SO$_2$ molar ratio to analyse variations over different timescales. We identify a cyclic pattern in BrO/SO$_2$ molar ratios for 11 of the 13 eruptive phases. The phases are initialised by low BrO/SO$_2$ molar ratios between 2 and 6 $\times$ 10$^{-5}$ coinciding with vulcanian-type activity followed by a strong increase to ratios ranging between 4 and 17 $\times$ 10$^{-5}$ when eruptive dynamism shifts to strombolian. For five phases, we additionally observe a progressive decrease to the initial values of 2 to 5 $\times$ 10$^{-5}$ toward the end of the phase. This clear pattern indicates a connection between the BrO/SO$_2$ molar ratio and eruptive dynamics. Based on our new data, we propose a conceptual model of the volcanic processes taking place at Tungurahua during the eruptive phases. Our data furthermore indicate that maximal BrO/SO$_2$ molar ratios observed during each phase could be related to the input of volatile-rich magma into the active part of the volcanic system of Tungurahua. This study shows that long-term BrO/SO$_2$ molar ratios can be used as a proxy for the volatile status as well as temporal evolution of the volcanic system.

Keywords: Tungurahua, SO$_2$, BrO, ratio, degassing, volcano, DOAS, NOVAC
INTRODUCTION

The molar ratios of gas species in volcanic plumes offer crucial information about volcanic processes. Sulphur dioxide (SO$_2$) emissions have in some cases proven to be a good forecasting tool for volcanic eruptions (e.g. Malinconico, 1979; Olmos et al., 2007; Hidalgo et al., 2018). Moreover, the halogen to sulphur ratios have been used for detecting changes in volcanic activity due to the different melt-fluid partitioning coefficients of these elements (Noguchi and Kamiya, 1963; Carrol and Holloway, 1994).

Changes in the chlorine to sulphur ratio have been suggested to forecast eruptions (Menyailov, 1975) and to distinguish different types of activity (Pennisi and Cloarec, 1998; Allard et al., 2005; Aiuppa, 2009; Spina et al., 2015). Several studies interpreting these measurements advocated preferred release of chlorine and fluorine at low pressures corresponding to shallow depth beneath the volcanic vent (Noguchi and Kamiya, 1963; Villemant and Boudon, 1999; Aiuppa et al., 2002; Burton et al., 2007; Edmonds et al., 2009). For bromine in contrast, the degassing processes are still not fully understood and lack comparison of measurements with model data. Several studies investigated melt-fluid partitioning of bromine with respect to sulphur using artificial melts (Bureau et al., 2000, 2010; Cochain et al., 2015) and concluded that the partition coefficient highly depends on the melt composition. A recent study using natural melts suggested efficient degassing of bromine at shallow depths beneath the volcanic vent similar to chlorine behaviour (Cadoux et al., 2018).

Since the first detection of bromine monoxide (BrO) inside a volcanic plume (Bobrowski et al., 2003), BrO as well as the BrO/SO$_2$ molar ratio have been retrieved at various volcanoes using ground-based measurements (e.g. Oppenheimer et al., 2006; Bobrowski and Platt, 2007, Kern et al., 2009; Boichu et al., 2011; Kelly et al., 2013) as well as from satellite (Theys et al., 2009; Hörmann et al., 2015). In addition, bromine chemistry in volcanic plumes has also been investigated through plume-chemistry models (e.g. Bobrowski et al., 2007; von Glasow, 2010; Roberts et al., 2014). A review of the current knowledge of bromine chemistry in volcanic plumes as well as a collection of all volcanic bromine measurements since 2003 was reported by Gutmann et al. (2018).

Recent studies of long-term BrO/SO$_2$ datasets have revealed changes in connection with varying volcanic activity at Mount Etna (Bobrowski and Giuffrida, 2012) and Nevado del Ruiz (Lübcke et al., 2014), or due to external forcing such as Earth tides (Dinger et al., 2018). The advantage of the BrO/SO$_2$ molar ratio compared to other halogen/sulphur ratios is that both species can be measured simultaneously via automated remote-sensing using scattered sunlight spectra applying the Differential Optical Absorption Spectroscopy (DOAS) technique (Platt and Stutz, 2008). This allows for continuous long-term data retrieval throughout different states of volcanic activity (Galle et al., 2010).

This study focuses on Tungurahua volcano, an andesitic subduction volcano located 120 km south of Quito, Ecuador. The average magma supply rates for this volcano have been estimated at $\approx 1.5 \times 10^6$ m$^3$yr$^{-1}$ over the last 2300 years (Hall et al., 1999). Regular eruptions occurred at least once per century (Pennec et al., 2008; Bablon et al., 2018). Hence, Tungurahua is among the most active volcanoes in the northern Andes in recent times (Hall et al., 1999; Wright et al., 2012). The most recent eruptive cycle started in September 1999. From 1999 until late 2008 Tungurahua exhibited long periods of continuous eruptive activity (Arellano et al., 2008; Samaniego et al., 2011; Hidalgo et al., 2015), with typical SO$_2$ emission rates of the order of 1000 t/d derived from ground-based measurements (Arellano et al., 2008; Galle et al., 2010) and from satellite (Carn et al., 2008; McCormick et al., 2014). From 2008 to 2016, however, the activity has been characterised by alternating eruptive phases and relative quiescence. Eruptive phases have displayed strombolian to vulcanian explosions, which were accompanied by persistent ash and gas emissions, and sporadic pyroclastic flows and lava effusion (Hidalgo et al., 2015).

We investigate the variation of the BrO/SO$_2$ molar ratio in the plume of Tungurahua from the spectral data recorded by a network of four autonomous scanning-DOAS stations over the period March 2007 to January 2017. To improve readability, we will refer to all ratios by omitting “molar” in the rest of the manuscript, e.g. BrO/SO$_2$ ratio. The different eruptive phases are decomposed into different stages, in order to identify characteristic patterns in the BrO/SO$_2$ ratio. We identify a pattern which was repeatedly observed throughout phases over the whole period of investigation. We interpret it with respect to magmatic processes and volcanic activity and derive a conceptual model of the eruptive dynamic representative for each eruptive phase. Lastly, we also discuss the behaviour of the BrO/SO$_2$ ratios in relation to different types of explosive activity.
Spectroscopic Retrieval

The amount of SO$_2$ and BrO in the plume was retrieved via a DOAS analysis comparing the spectra measured in the direction of the volcanic plume with spectra in a direction not intercepting the volcanic plume. To improve the signal-to-noise ratio in this work the co-adding retrieval algorithm was used as described in Lübcke et al. (2014): The algorithm identifies 10 plume spectra (having the highest SO$_2$ values) and 10 reference spectra (having the lowest SO$_2$ values) in each scan. The algorithm then co-adds these spectra, respectively from four consecutive scans (i.e. 40 spectra of each type in total). The resulting spectra, which are from now on referred to as “co-added” spectra, are then used as measurement and reference spectrum in the DOAS fitting algorithm, respectively (Platt and Stutz, 2008). This yielded one co-added BrO as well as one SO$_2$ mean slant column density (SCD) roughly every 30 min.

Fit Parameter

In addition to the co-added plume and reference spectra, the DOAS fitting algorithm also included absorption cross-sections for SO$_2$ (Vandaele et al., 2009) and O$_3$ (Burrows et al., 1999) in the SO$_2$ spectral fit region (314.8–326.8 nm) and SO$_2$, BrO (Fleischmann et al., 2004), O$_4$ (Hermans et al., 2003), NO$_2$ (Vandaele et al., 1998) and CH$_3$O (Meller and Moortgat, 2000) in the BrO spectral fit region (330.6–352.75 nm) as well as a third-order polynomial to account for broadband absorption structures, Mie, and Rayleigh scattering. Two spectra were included to account for the Ring effect (Grainger and Ring, 1962) – the standard Ring spectrum as well as a wavelength dependent Ring spectrum (derived from the original Ring spectrum by multiplication with $\lambda^4$, where $\lambda$ represents wavelength; Wagner et al., 2009). The fit settings were chosen as an outcome of the study by Vogel et al. (2013).

The detection limit can be derived as a multiple of the DOAS fit error based on Platt and Stutz (2008). In our study we used three times the fit error – which is usually a quite conservative estimate – yielding here a mean detection limit of $4.9 \times 10^{16}$ molecules/cm$^2$ (SO$_2$) and $4.9 \times 10^{13}$ molecules/cm$^2$ (BrO). The latter fits very well to the random fluctuations of BrO around zero when SO$_2$ SCDs are close to zero (i.e. no gas emissions), which similarly span roughly $\pm 5 \times 10^{13}$ molecules/cm$^2$. For SO$_2$ this fluctuation is hardly visible, because the signal is up to two orders of magnitude larger than the detection limit.

Determining the BrO/SO$_2$ Ratio

The BrO/SO$_2$ ratio for a series of simultaneous measurements of both species inside a volcanic plume is usually obtained by performing a linear regression for BrO as a function of SO$_2$ (e.g. Bobrowski et al., 2003, 2007; Bobrowski and Giuffrida, 2012; Hörmann et al., 2013). The linear regression was performed using the error-weighted orthogonal regression algorithm developed by Cantrell (2008).

For this study, it was employed in two ways:

1. Daily regression: This regression takes all co-added BrO and SO$_2$ data of a particular day to retrieve daily BrO/SO$_2$ ratios in order to observe short-term changes in the BrO/SO$_2$ ratio. To ensure the quality of the retrieved BrO/SO$_2$ ratio, days were considered only if at least one data point exceeded SO$_2$ SCDs of $5 \times 10^{17}$ molecules/cm$^2$ and the regression fit error was below $1.5 \times 10^{-5}$.

2. Period regression: Data acquired over the period of several days/weeks were evaluated using the same methodology as for the daily regression. The beginning and end of these periods were chosen when a change in the daily BrO/SO$_2$ ratios could be detected. For long periods without significant change, the beginning and end of these periods were additionally chosen when changes in the classification of explosive activity from Hidalgo et al. (2015) occurred. Only those periods where at least five co-added SO$_2$ SCDs exceeded $5 \times 10^{17}$ molecules/cm$^2$, the maximum SO$_2$ SCD exceeded $7 \times 10^{17}$ molecules/cm$^2$, and the regression fit error was below $1 \times 10^{-5}$ were taken into account.

We also calculated the BrO/SO$_2$ ratio by taking the BrO/SO$_2$ ratio of every co-added data point with high SO$_2$ signal (SO$_2$ SCD > $7 \times 10^{17}$ molecules/cm$^2$). In a second step, the daily mean BrO/SO$_2$ ratio was calculated as the error-weighted daily mean over these single BrO/SO$_2$ ratios. To ensure robust results we calculated the daily mean BrO/SO$_2$ ratio only if there were at least five data points above the SO$_2$ threshold. An example of the results obtained by this algorithm for 28 December 2010 is given in Figure 1A. This method was used by Lübcke et al. (2014) for NOVAC data and we will refer to it as “daily means” in this study.

In this study, we introduce a scheme including all three methods, the daily means method as well as the two linear regression methods to achieve a sufficient set of BrO/SO$_2$ ratios. They are presented in Figure 1. The daily mean method yielded BrO/SO$_2$ ratios on 277 days, while the daily regression was successful on 425 days. Of the 276 days, in which both methods yielded a BrO/SO$_2$ ratio, we observed a good correlation between the two methods. However, the daily mean method consistently yielded slightly lower ratios (we retrieved a slope of 0.86 ± 0.05; see Figure 2). The results of all daily mean and daily regression BrO/SO$_2$ ratios as a function of time are plotted in Figure 3C.

Volcanic Activity Classification

We used the classification criteria of volcanic activity described by Hidalgo et al. (2015), which includes the eruptive and quiescent phases from 2007 to 2013. We extended this classification in order to cover our study period following the same criteria (i.e. seismo-acoustic records and visual observations).

Three types of activity were distinguished during the study period: (1) Quiescence, (2) Low explosive activity and (3) High explosive activity. Episodes of high explosive activity are those in which discrete volcanic explosions occurred, with the explosions themselves being classified as vulcanian when their mean acoustic excess pressures were greater than 100 Pa at a distance of 1 km from the vent (Johnson, 2003). Episodes of low explosive activity are those in which no vulcanian explosions were present. These episodes were
instead characterised by Strombolian-type activity, during which the explosions had excess pressures smaller than the 100 Pa threshold (Johnson, 2003). In both cases, explosive activity is commonly accompanied by intense gas and ash emissions. During quiescent episodes, the seismic activity was generally weak and merely a few long period and volcano-tectonic earthquakes were detected.

RESULTS

BrO and SO\textsubscript{2} SCDs From Tungurahua
Over the study period between 2007 and 2017, 18 phases of high SO\textsubscript{2} SCDs (of the order of $10^{18}$ molecules/cm\textsuperscript{2}) could be identified from the NOVAC data. They were separated by periods of weak SO\textsubscript{2} degassing, where the SO\textsubscript{2} SCDs generally did not exceed $1.6 \times 10^{17}$ molecules/cm\textsuperscript{2}. BrO SCDs exceeded the detection limit of $4.8 \times 10^{13}$ molecules/cm\textsuperscript{2} only during phases where high SO\textsubscript{2} SCDs were measured. In those phases typical BrO SCDs around $8 \times 10^{13}$ molecules/cm\textsuperscript{2} were observed. During the phases of weak SO\textsubscript{2}, BrO fluctuated around zero and was below detection. The results of BrO and SO\textsubscript{2} from Tungurahua are displayed in Figures 3A,B.

BrO/SO\textsubscript{2} Ratios
BrO/SO\textsubscript{2} ratios typically ranged between 2 and $8 \times 10^{-5}$, which is similar to ratios observed at other arc volcanoes in the Northern Volcanic Zone of the Andes, such as Nevado del Ruiz (Lübcke et al., 2014) and Cotopaxi (Dinger et al., 2018). In 2009 and the first half of 2010, relatively constant BrO/SO\textsubscript{2} ratios of 2 to $6 \times 10^{-5}$ were observed. This contrasts with periods in 2008 and after 2010, where a larger range of ratios between 2 and $20 \times 10^{-5}$ were detected. During the intermitting periods of weak SO\textsubscript{2} degassing, we could only infer an upper limit for the BrO/SO\textsubscript{2} ratio, calculated as the ratio between the BrO detection limit and the maximal SO\textsubscript{2} SCD during these periods:

$$\left( \frac{\text{BrO}}{\text{SO}_2} \right)_{\text{upper limit}} = \frac{\text{BrO SCD}_{\text{detection limit}}}{\text{SO}_2 \text{SCD}_{\text{max, during low SO}_2 \text{ periods}}} = \frac{4.8 \times 10^{13}}{1.6 \times 10^{17}} = 3 \times 10^{-4}$$

Since this value far exceeds even the largest BrO/SO\textsubscript{2} ratios observed during phases of high SO\textsubscript{2} SCDs, we are not able to draw any conclusion from BrO/SO\textsubscript{2} ratios during these phases of low SO\textsubscript{2} at Tungurahua.

Dynamic of the BrO/SO\textsubscript{2} Ratio at Tungurahua
We identified changing BrO/SO\textsubscript{2} ratios during 11 phases of high SO\textsubscript{2} SCDs, which we number with roman numerals in chronological order (see Figure 3A). We identified a characteristic pattern in the BrO/SO\textsubscript{2} ratio over the course of each phase. For two phases (I and VII) the pattern is observed more than once and we separate this by adding the suffix a and b to the phase numbering, thus adding up to 13 phases, where we observe the BrO/SO\textsubscript{2} pattern. This pattern can be divided into
FIGURE 3 | Time-series of the co-added results for (A) SO$_2$ SCDs and (B) BrO SCDs for the NOVAC stations at Tungurahua volcano from March 2007 until March 2017. (C) the daily mean BrO/SO$_2$ ratios are plotted for both calculation methods. The coloured areas at the bottom of each plot indicates the volcanic activity derived from seismo-acoustic data (Hidalgo et al., 2015): high explosive activity (orange) and low explosive activity (yellow). The SO$_2$ detection limit is $4.9 \times 10^{16}$ molecules/cm$^2$ and the BrO detection limit is $4.9 \times 10^{13}$ molecules/cm$^2$ (indicated by the dashed line). The 11 eruptive phases discussed later in the manuscript are included with roman numerals in (A).
four stages, consisting of two initial stages (0 and 1), followed by an intermediate stage 2 and an ending stage 3. The separation of the initial stages into stage 0 and 1 could only be derived for four phases.

Stage 0 exclusively occurred in phases II, III, IV, and VII and typically lasted for periods of time between 2 and 21 days. We observed BrO and SO$_2$ degassing and BrO/SO$_2$ ratios between 3 and $6 \times 10^{-5}$, with a mean ratio of $4.31 \times 10^{-5}$.

Stage 1 was present in every phase. We observed a constant baseline BrO/SO$_2$ ratio of typically less than $5 \times 10^{-5}$ (only for phase IX was it significantly higher at $6.20 \times 10^{-5}$). Mean ratio: $3.58 \times 10^{-5}$.

Stage 2 generally started with a strong increase in bromine emissions, which resulted in a BrO/SO$_2$ ratio of up to $17 \times 10^{-5}$, often twice as high as the ratio during stage 1. If stage 2 continued for more than a week, the ratio progressively decreased until it reached the former baseline level from stage 1. We will denote the latter part of this stage as stage 2b and the former as stage 2a. Stage 2 was observed in 11 phases (see Table 1). Mean ratio: $9.22 \times 10^{-5}$.

Stage 3 started when the BrO/SO$_2$ ratio decreased to low BrO/SO$_2$ ratios, similar to the ones observed in stage 1 ($2-5 \times 10^{-5}$). It could be seen in 5 phases (Ia, Ib, II, V, and X). Mean ratio: $3.44 \times 10^{-5}$.

The most significant feature is the strong increase in the BrO/SO$_2$ ratios from stage 1 to stage 2. This increase was lowest for phases II and VII and highest for phases Ia, Ib, V, and XI. The evolving pattern of BrO/SO$_2$ ratios through the four stages can be clearly recognised for every phase (see Figure 4 and Table 1).

### Table 1

| Phase | Period (dd.mm.yy) | Stages observed (duration in d) | BrO/SO$_2$ stage 0 | BrO/SO$_2$ stage 1 | BrO/SO$_2$ stage 2 | BrO/SO$_2$ stage 3 |
|-------|------------------|---------------------------------|---------------------|---------------------|---------------------|---------------------|
| Ia    | 19.01.08–25.03.08| 1-2-3                          | Below detection     | $3.88 \times 10^{-5}$ | $1.06 \times 10^{-4}$ | $5.18 \times 10^{-5}$ |
|       |                  | (20-18-29)                      |                     |                     |                     |                     |
| Ib    | 28.04.08–04.08.08| 1-2-3                          | Below detection     | $2.86 \times 10^{-5}$ | $1.22 \times 10^{-4}$ | $3.26 \times 10^{-5}$ |
|       |                  | (30-6-26)                       |                     |                     |                     |                     |
| II    | 16.12.08–01.03.09| 0-1-2                          | $3.07 \times 10^{-5}$ | $3.56 \times 10^{-5}$ | $5.57 \times 10^{-5}$ | $2.12 \times 10^{-5}$ |
|       |                  | (2-30-8-5)                      |                     |                     |                     |                     |
| III   | 28.03.09–03.07.09| 0-1-2                          | $4.84 \times 10^{-5}$ | $3.19 \times 10^{-5}$ | $4.17 \times 10^{-5}$ | Below detection     |
|       |                  | (21-69-8)                       |                     |                     |                     |                     |
| IV    | 30.12.09–04.03.10| 0-1                            | $3.04 \times 10^{-5}$ | $2.28 \times 10^{-5}$ | Below detection     | Below detection     |
|       |                  | (12-45)                         |                     |                     |                     |                     |
| V     | 26.05.10–04.08.10| 1-2                            | Below detection     | $2.80 \times 10^{-5}$ | $9.49 \times 10^{-5}$ | $3.16 \times 10^{-5}$ |
|       |                  | (24-22-25)                      |                     |                     |                     |                     |
| VI    | 22.11.10–25.12.10| 1-2                            | Below detection     | $5.10 \times 10^{-5}$ | $9.00 \times 10^{-5}$ | Below detection     |
|       |                  | (34-9)                          |                     |                     |                     |                     |
| VIIa  | 20.04.11–08.05.11| 0-1                            | $6.31 \times 10^{-5}$ | $4.32 \times 10^{-5}$ | $6.57 \times 10^{-5}$ | below detection     |
|       |                  | (2-12-5)                        |                     |                     |                     |                     |
| VIIb  | 16.05.11–26.05.11| 1-1                            | Below detection     | $4.36 \times 10^{-5}$ | Below detection     | Below detection     |
|       |                  | (11)                            |                     |                     |                     |                     |
| VIII  | 28.04.13–16.05.13| 1-2                            | Below detection     | $2.24 \times 10^{-5}$ | $8.29 \times 10^{-5}$ | Below detection     |
|       |                  | (8-11)                          |                     |                     |                     |                     |
| IX    | 14.07.13–05.08.13| 1-2                            | Below detection     | $6.20 \times 10^{-5}$ | $9.19 \times 10^{-5}$ | Below detection     |
|       |                  | (6-17)                          |                     |                     |                     |                     |
| X     | 02.04.14–24.04.14| 1-2                            | Below detection     | $2.63 \times 10^{-5}$ | $8.54 \times 10^{-5}$ | $3.27 \times 10^{-5}$ |
|       |                  | (13-6-4)                        |                     |                     |                     |                     |
| XI    | 02.08.14–25.10.14| 1-2                            | Below detection     | $3.11 \times 10^{-5}$ | $1.78 \times 10^{-4}$ | Below detection     |
|       |                  | (41-44)                         |                     |                     |                     |                     |

### Evolution of the BrO/SO$_2$ Ratio Over the Course of Each Phase

Here, we discuss in detail the evolution of the BrO/SO$_2$ ratio through every phase in chronological order. Phase V – serving as an example – will be discussed in more detail.

Phase I: During phase I (from 19 January to 22 July 2008), we observed the pattern twice. We will refer to them as phase Ia and phase Ib. Phase Ia started with a BrO/SO$_2$ ratio of $4.54 \times 10^{-5}$ during the initialisation from 19 January 2008 (stage 1), lasting until an increase to $1.7 \times 10^{-4}$ was observed on 8 February (stage 2a, lasting for 5 days). The BrO/SO$_2$ ratio decreased over the following weeks (stage 2b) until baseline levels were reached in March/April 2008 (stage 3, see Figure 5). Phase Ib started on 28 April 2008 with a ratio of $2.86 \times 10^{-5}$, reaching a short intermediate stage 2 on 28 May 2008, lasting only 6 days. It then decreased to $3.26 \times 10^{-5}$ until 27 July 2008.

Phases II and III: The pattern is less clear for these phases, where only a small enhancement in the BrO/SO$_2$ ratio could...
FIGURE 4 | The evolution of the BrO/SO\textsubscript{2} ratio during the stages of the different phases. Depicted are the mean BrO/SO\textsubscript{2} ratios (period regression) for each of the four stages: stage 0, stage 1, stage 2, and stage 3. The colour at the bottom indicates the most predominant activity classification during each stage (Hidalgo et al., 2015). Detailed discussion in the text.

be observed on 19 January 2009 (phase II) and on 26 June 2009 (phase III).

Phase IV: This phase exclusively consisted of the initial stages 0 and 1 with low BrO/SO\textsubscript{2} ratio (2–3 × 10\textsuperscript{-5}). After stage 1, the gas signal decreased below the detection limit and no BrO/SO\textsubscript{2} ratio could be retrieved for stage 2.

Phase V: This phase took place from 29 May until 4 August 2010 and comprised stages 1 to 3. Stage 1, spanning from 29 May until 18 June, was characterised by high SO\textsubscript{2} emissions with SO\textsubscript{2} SCDs of up to 2 × 10\textsuperscript{18} molecules/cm\textsuperscript{2} while BrO SCDs did not exceed 1 × 10\textsuperscript{14} molecules/cm\textsuperscript{2} (see Figure 6, bottom left). During this stage, a mean BrO/SO\textsubscript{2} ratio of 2.80 × 10\textsuperscript{-5} was retrieved. Stage 2 started on 19 June with a strong increase in the BrO/SO\textsubscript{2} ratio by more than 1 × 10\textsuperscript{-4} to 1.7 × 10\textsuperscript{-4} as a result of a sharp rise in BrO SCDs up to 2 × 10\textsuperscript{14} molecules/cm\textsuperscript{2} (see Figure 6, bottom middle). As stage 2 progressed, the BrO/SO\textsubscript{2} ratio slowly decreased (stage 2b), reaching previous baseline levels of roughly 3 × 10\textsuperscript{-5} by 11 July (after 22 days), where it remained throughout stage 3 until the gas emission dropped below the detection limit on 4 August. Both the BrO/SO\textsubscript{2} ratio as well as the linear regression results for the three stages of phase V are plotted in Figure 6. The period regression of each of these three stages reveals that the gas composition was similar during stages 1 and 3. While the SO\textsubscript{2} SCDs were similar for all three stages, BrO SCDs were higher during stage 2, if compared to stages 1 and 3 (Figure 6, bottom).

Phases VI to IX and XI exhibited only the initial stages and the intermediate stage 2, with a varying increase in strength from initialisation to stage 2 of 2.3 × 10\textsuperscript{-5} (during phase VII) to 1.4 × 10\textsuperscript{-4} (during phase XI).

Phase X: This phase behaved similar to phases VI-IX, but also had a short stage 3.

The detailed depiction of the co-added SO\textsubscript{2} and BrO SCDs as well as the BrO/SO\textsubscript{2} ratios for every phase can be found in the Supplementary Materials.

In 2012, in the second half of 2015, and for six phases between 2013 and 2016, only short periods with few BrO SCDs above the detection limit could be observed and longer period of BrO/SO\textsubscript{2} ratios could not be identified. For the six phases between 2013 and 2016 constant BrO/SO\textsubscript{2} ratio were calculated, with values ranging from 1.8 to 5.2 × 10\textsuperscript{-5}, similar to the range during stages 0 and 1.

Relation of the BrO/SO\textsubscript{2} Ratio With Explosive Activity

The observed changes in the BrO/SO\textsubscript{2} ratio between the different stages in the cycle observed during each phase at Tungurahua can be linked to changes in the explosive activity classification. The classification is closely related to the pattern of the BrO/SO\textsubscript{2} ratio in the following way: Stage 0 was typically related to low explosive activity. During Stage 1 – characterised by low BrO/SO\textsubscript{2} ratios – explosive activity was always defined as high. The subsequent increase in BrO/SO\textsubscript{2} ratios between stage 1 and 2 coincides with a change from high to low explosive activity. Only phases VIII, IX, and X (all after 2012), which are the shortest phases, did not exhibit this change in the classification. For these phases the activity classification remained high throughout stage 2. Stage 3 was generally governed by low explosive activity.
FIGURE 5 | BrO/SO$_2$ ratio over the course of each phase. Mean BrO/SO$_2$ ratio of the different periods (blue area, separated by black lines), the corresponding stage number is marked below the blue area. Furthermore, daily BrO/SO$_2$ from daily regression (red crosses) as well as from daily mean method (yellow crosses) are included. Explosivity classification is marked by yellow (low explosive) and orange area (high explosive).
The fact that for most of the phases the change in activity from high to low explosive activity was linked to an increase in the BrO/\( \text{SO}_2 \) ratio suggests that this might be inherent to the volcanic system of Tungurahua during the study period. In order to compare the explosive activity classification with the BrO/\( \text{SO}_2 \) ratio throughout the whole dataset, we performed period regressions for every period of high and low explosive activity. The results are shown in Figure 7. While the mean of all BrO/\( \text{SO}_2 \) ratio during high explosive activity \( (3.87 \times 10^{-5}) \) is slightly lower than the mean of all BrO/\( \text{SO}_2 \) ratios during low explosive activity \( (5.19 \times 10^{-5}) \), there is large variance caused by other parameters and thus this difference seems to be not statistically significant.

**DISCUSSION**

In Figure 4, a recurring temporal pattern in the eruptive phases of Tungurahua can be observed. We can summarise our observations as follows:

1. Initiation of degassing with low BrO/\( \text{SO}_2 \) ratios is observed during the initial stages. These are associated with high \( \text{SO}_2 \) emissions and strong seismicity, i.e. vulcanian explosions, occurring at the beginning of the eruptive phases.
2. After the initial stages, increased BrO values and nearly unchanged \( \text{SO}_2 \) emissions are observed in the plume, leading to an increased BrO/\( \text{SO}_2 \) ratio (observed for 11 of 13 eruptive phases). This typically coincided with the termination of high explosive activity.
3. For five phases, BrO/\( \text{SO}_2 \) decreased afterward and remained low until \( \text{SO}_2 \) SCDs decreased below detection. For the other eight phases, degassing had decreased already to a point that \( \text{SO}_2 \) and BrO were below the limit to calculate meaningful BrO/\( \text{SO}_2 \) ratios.

Changes of the BrO/\( \text{SO}_2 \) ratio in the plume of a single volcano have previously been linked to changes in volcanic processes (Bobrowski and Giuffrida, 2012; Lübcke et al., 2014; Dinger et al., 2018), when observing 3 to 20 min old volcanic plumes (e.g. Platt and Bobrowski, 2015; Gutmann et al., 2018). This was assumed to be caused by different melt-fluid partitioning of bromine and sulphur in magmatic melts. Moreover, differences in the BrO/\( \text{SO}_2 \) ratio among different volcanoes have been attributed to differences in magmatic composition (Platt and Bobrowski, 2015).

The melt-fluid partitioning of halogens (especially of chlorine and fluorine) with respect to sulphur has been the subject of multiple studies (e.g. Spilliaert et al., 2006; Burton et al., 2007). They found lower melt-fluid partitioning of chlorine compared to sulphur and thus concluded that (1) sulphur is usually released from magma earlier than chlorine (Carrol and Holloway, 1994; Spilliaert et al., 2006; Aiuppa, 2009; Webster et al., 2009) and (2) that the Cl/S ratio decreases with increasing degassing source depth (e.g. Burton et al., 2007). Measurements of chlorine and sulphur suggest that chlorine is more significantly released only
Based on our observations and data from previous studies (Hidalgo et al., 2015), we therefore suggest that each eruptive phase is driven by an intrusion of gas-rich magma, which subsequently degasses under different conditions. However, the amount of magma driving the phase might vary between the different eruptive phases.

**Comparison of the Conceptual Model to Petrological Studies**

Petrological studies showed that major eruptive phases at Tungurahua, like August 2006 (VEI 3) and May 2010 (phase V in this study), have been triggered by the input of new, hot, volatile-rich magma into a long-lived magmatic reservoir (Samaniego et al., 2011; Myers et al., 2014). Both studies propose an andesitic magma reservoir located at 8–10 km depth, which is coupled to a deeper mafic reservoir, probably located at depth of 15–16 km (Andújar et al., 2017). Myers et al. (2014) suggest that phase V was driven by the ascent of volatile-rich magma from the deeper basaltic andesite magma reservoir resulting in mixing with the upper andesitic, degassed, reservoir. This mechanism can account for the major eruptive phases documented at Tungurahua, and is in agreement with our conceptual model, which proposes the intrusion of volatile-rich magma (at least sulphur- and bromine-rich) as the main triggering process.

The lowest BrO/SO$_2$ ratios are observed for phases III (June 2009) and IV (February 2010), where the BrO/SO$_2$ ratios varied only between 2 and $4 \times 10^{-5}$. Thus, the high ratio recorded for phase V ($18 \times 10^{-5}$ at the beginning of stage 2) could indicate a comparatively high input of mafic magma leading toward this eruption (as proposed by Myers et al., 2014) and reciprocally the low BrO/SO$_2$ ratios during phase III and IV indicate low...
FIGURE 8 | Conceptual model for each of the 13 eruptive phases observed at Tungurahua. (A,B) Low BrO/SO$_2$ ratios during initialisation indicate degassing at depth, interpreted as the intrusion of gas-rich magma and deep degassing signature. BrO/SO$_2$ ratios typically between 3 and 6 $\times 10^{-5}$ (stage 0, mean ratio: 4.31 $\times 10^{-5}$), and 2 and 6 $\times 10^{-5}$ (stage 1, mean ratio: 3.58 $\times 10^{-5}$) were observed. (C) Strong increase in the BrO SCDs indicates degassing at shallow depth. BrO/SO$_2$ ratio increased to 4 to 17 $\times 10^{-5}$ (stage 2, mean ratio: 9.22 $\times 10^{-5}$). The subsequent decrease in the BrO/SO$_2$ ratio indicates depletion of first Br in the magmatic melt. (D) Degassing from magma depleted in gas during stage 3 with low BrO/SO$_2$ ratios between 2 and 5 $\times 10^{-5}$ (stage 3: mean ratio: 3.44 $\times 10^{-5}$). Detailed description given in the text.

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or no input of volatile-rich mafic magma. Therefore, we can potentially link the BrO/SO$_2$ ratios observed during efficient degassing (stage 2) of the eruptive phases to the respective volatile status of the degassing magma.

Furthermore, it is important to note that during phase V, the energy of the vulcanian explosions was extremely high (Hidalgo et al., 2015) and larger than for the other eruptive phases between 2007 and 2016.

Comparison to Other Studies With Respect to Volcanic Activity

There are other studies similarly discussing long-term trends of BrO/SO$_2$ ratios with respect to changes in volcanic activity. Similar to our study, Bobrowski and Giuffrida (2012) found lower BrO/SO$_2$ ratios during periods of higher volcanic activity at Mount Etna during an observation period of more than 3 years including two long lasting periods of effusive eruptive activity and several more violent events (e.g. lava fountaining). Lübcke et al. (2014) obtained a continuous 3-year time-series of BrO/SO$_2$ ratios at the andesitic arc volcano Nevado del Ruiz. They observed low BrO/SO$_2$ ratios of $\approx 2\times 10^{-5}$ during a period of high seismic activity and an increase of the BrO/SO$_2$ ratios to $\approx 5\times 10^{-5}$ when the seismicity later decreased. This is similar to the behaviour which we observed during the eruptive phases of Tungurahua between the high explosive stage 1 and the low explosive stage 2.

In contrast to the laboratory measurements from Cadoux et al. (2018), Bobrowski and Giuffrida (2012) assumed an earlier bromine melt-fluid partitioning in comparison to sulphur as an interpretation of the spectral measurements, i.e. higher BrO/SO$_2$ ratios during the intrusion of gas-rich magma into the system. A similar hypothesis was proposed by Lübcke (2014), who associated the low BrO/SO$_2$ ratios to shallow degassing during
a period of increased seismicity. The increase in BrO/SO$_2$ ratios during the activity of Cotopaxi in 2015 (Dinger et al., 2018) were, like our study, similarly interpreted in terms of efficient degassing of rising magma in the conduit following the opening of the volcanic system by four phreato-magmatic explosions (Hidalgo et al., 2018).

In contrast to our measurements, the data of Bobrowski and Giuffrida (2012) and Lübcke et al. (2014) additionally enabled to determine BrO/SO$_2$ ratios during non-eruptive periods and to include those in their interpretation, which could be a reason for the divergence in the interpretation to our study. In order to resolve this divergence in the interpretation of the BrO/SO$_2$ ratio with respect to magmatic processes and partitioning of gas constituents from the melt, further field studies as well as laboratory studies should be undertaken. For such field studies, we therefore recommend even more comprehensive data collection, including seismic tomography to better determine source regions from a geophysical standpoint.

Multiple studies have investigated the behaviour of other halogens with respect to volcanic activity, predominantly through Cl/S ratio or HCl/SO$_2$ ratios. Measurements at Mount Etna revealed a decrease in the Cl/S ratio when Strombolian activity followed a period of violent lava fountaining, and even lower Cl/S during a phase of passive degassing (Aiuppa et al., 2004). Furthermore, Spina et al. (2015) suggested a correlation with eruption intensity derived from FTIR measurements of HCl/SO$_2$ ratios at Mount Etna. This contrasts to what we found for the BrO/SO$_2$ ratio, which could potentially be an indication that bromine and chlorine do not behave similarly as has been assumed here and in several previous studies (Seo and Zajacz, 2016; Cadoux et al., 2018). However, other measurements at Mount Etna did display an increase in the HCl/SO$_2$ ratio, when explosive activity subsided (Allard et al., 2005).

To our knowledge there is only one study with simultaneous measurements of BrO/SO$_2$ and Cl/S (Bobrowski et al., 2017). In this work the authors find a positive correlation of both ratios during a one-week period at Nyiragongo volcano in June 2011. Both ratios change concurrently with a significant decrease (several tens of metre) in the height of the lava lake. Another study by Aiuppa et al. (2005), investigating Br/S and Cl/S ratios during the transition from quiescence to eruption at Mount Etna starting in October 2004, does not demonstrate a clear change of either ratio, both staying within the variance of the measurements. Further studies will be essential to confirm or disprove the assumptions made regarding the similarities of fluid-melt partitioning of chlorine and bromine.

**CONCLUSION**

We successfully retrieved BrO/SO$_2$ ratios for 13 eruptive phases at Tungurahua volcano between 2007 and 2017. The data show typical values of BrO/SO$_2$ ratios between 2 and $8 \times 10^{-5}$, similar to other andesitic arc volcanoes.

For 11 of the 13 eruptive phases, we identified an evolutionary pattern in the BrO/SO$_2$ ratio, which we interpret as characteristic of the volcano’s eruptive phases. We propose a conceptual eruptive model for Tungurahua linking this pattern in the BrO/SO$_2$ ratio to the eruptive dynamics of the volcanic system. The appearance of low BrO/SO$_2$ ratios in the plume at the beginning of the eruptive phases is directly linked to a volatile-rich magma intrusion, which acts as a trigger for the eruptive activity and is initially characterised by high explosivity. A following increase in molar ratios is associated with a shift to shallower depth at which efficient bromine degassing occurs, and is typically associated with a sustained period of low explosive activity. The subsequent depletion in volatiles in the degassing magma body finally leads to a decrease in BrO/SO$_2$ ratios altogether, and an end to the eruptive phase. The evolutionary pattern of the BrO/SO$_2$ ratios is thus linked to the explosive activity of the volcano, with low BrO/SO$_2$ ratios associated to high explosive activity later evolving to high ratios during a period of low explosive activity.

High BrO/SO$_2$ ratios observed after a reported high input of volatile-rich magma into the upper volcanic system could infer a relationship between bromine degassing to the amount of volatiles in the active system. The BrO/SO$_2$ can thus potentially be used as a proxy for the degree of depletion of volatiles. This study therefore demonstrates the increasing potential of long-term measurements of the BrO/SO$_2$ molar ratio to reveal important information about status as well as temporal evolution of the volcanic system.

**AUTHOR CONTRIBUTIONS**

NB, SW, PL, and UP contributed conception and design of the study. SW performed the data evaluation. NB, HS, and SW performed the statistical analysis. SW wrote the draft of the manuscript. NB, SH, FD, SA, UP, TW, and SW contributed to the interpretation of the data. All authors contributed to manuscript revision, read and approved the submitted version.

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SUPPLEMENTARY MATERIAL
The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/feart.2019.00132/full#supplementary-material

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