A New Approach to the Determination of Mineral and Organic Soil Types Based on Dilatometer Tests (DMT)

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Abstract: In order to identify the soil type in the ground, Marchetti’s nomogram chart is commonly used on the basis of dilatometer tests (DMT). In this chart, the material index values ($I_D$) and the dilatometer modulus ($E_D$) are used to determine the state and type of soils predominant in mineral soils. Unfortunately, this classification is not accurate enough for the identification of organic soils. This article proposes a new classification based on a nomogram chart for both mineral soils and organic soils using ($p_0$, $p_1$) readings and pore water pressure ($u_0$).

Keywords: in situ tests; DMT; mineral and organic soils; chart diagram

1. Introduction

The following tests are commonly used across the world to identify the type of mineral and organic soils in the ground of a designed object: a Standard Static Test (SPT), a Pressuremeter (PMT), a Cone Penetration Test (CPT) or a Marchetti dilatometer (DMT).

In recent years, static probing by “dilatometer penetration tests”, commonly known as DMT (Figure 1), is a widely-used method of in situ investigation of the ground to provide information needed by civil engineers for design, construction, permissions, and operation control. The results of the DMT field tests, complemented by well-established experience, need to be considered to derive the characteristic values ($X_m$) and to later design values ($X_d$) of geotechnical parameters [1–4]:

$$X_d = X_m / \gamma_M$$

(1)

where $\gamma_M$ is the partial factor of a material property.

Figure 1. Marchetti dilatometer: (a)–flat blade, 1—electric wire, 2—pneumatic tubing, 3—steel membrane, (b)—test stages, 4—pushing, 5—contact stress $p_0$, 6—expansion stress $p_1$, 7—pressure $p_2$, (c): view of DMT blade actually used in soil in situ tests.
This paper presents a review of the methods used in the identification of soils from DMT test results. Next, the analysis of the results from field test studies on six sites is presented in [5]: The DMT testing was carried out in the frame of the project of SGGW campus development [6], an experimental embankment (Antoniny site), embankment dams (Nielisz, Koszyce and Mielimaka), “Płocka subway station” and Stegny site in Warsaw. Finally, a new diagram for the identification of mineral and organic soils from DMT tests is offered.

2. Methodology and Interpretation of Dilatometer Tests

Dilatometer tests (DMT) were applied to recognize mineral and organic subsoils distinguished in the above-mentioned test sites. The details of the DMT test operation can be found in [2,7–17]. During the DMT tests, A, B and C readings are carried out as shown in Figure 1 [18,19]. A, B and C readings are adjusted according to the inertia impact resistance of the membrane, which allows to determine pressures $p_0$, $p_1$ and $p_2$ (equ. 2 ÷ 4). The pressures $p_0$, $p_1$ and $p_2$ together with the calculated value of the vertical effective stress component $\sigma'_v$ and the pore water pressure $u_0$ are used to determine the following dilatometer indexes (equ. 5 ÷ 8) [1,13,17]:

- 0.05 mm corrected pressure treading in DMT $p_0$
  
  $$p_0 = 1.05(A - Z_M + \Delta A) - 0.05(B - Z_M - \Delta B)$$  

- 1.10 mm corrected pressure treading in DMT $p_1$
  
  $$p_1 = B - Z_M - \Delta B$$

- corrected third reading in DMT $p_2$
  
  $$p_2 = C - Z_M - \Delta A$$

- Material index $I_D$
  
  $$I_d = \frac{P_1 - P_0}{P_0 - u_0}$$

- Horizontal earth pressure index $K_D$
  
  $$K_d = \frac{P_0 - u_0}{\sigma'_v}$$

- Dilatometer modulus $E_D$
  
  $$E_d = 34.7 \cdot (P_1 - P_0)$$

- Water pressure index $U_D$
  
  $$U_d = \frac{P_2 - u_0}{P_0 - u_0}$$

where:

- $p_0$—A-pressure reading, corrected for $Z_m$, $\Delta A$ membrane stiffness at 0.05 mm expansion, and 0.05 mm expansion itself, to estimate the total soil stress acting normal to the membrane immediately before its expansion into the soil (0.00 mm expansion).
- $p_1$—B-pressure reading corrected for $Z_m$ and $\Delta B$ membrane stiffness at 1.10 mm expansion to give the total soil stress acting normal to the membrane at 1.10 mm membrane expansion.
- $p_2$—C-pressure reading corrected for $Z_m$ and $\Delta A$ membrane stiffness at 0.05 mm expansion and used to estimate pore-water pressure.
- $\sigma'_v$—vertical effective stress at the centre of the membrane before insertion of the DMT blade.


\( u_0 \) — pore-water pressure acting at the centre of the membrane before insertion of the DMT blade (often assumed as hydrostatic below the water table surface).

\( Z_m \) — gage pressure deviation from zero when vented to atmospheric pressure (an offset used to correct pressure readings to the true gage pressure).

In order to reduce the necessity of using various types of equipment, field research methods are being sought that allow interpretation of the results obtained in a wide range. One of the field studies that meets this requirement is the Marchetti dilatometer [1], used more often in the country. The biggest advantage of studying the dilatometer is fast and not very complicated measurement, on the basis of which it is possible to determine the profiles of many soil parameters. The interpretation of ground parameters is based on the use of empirical relationships linking the results of measurements to the values of ground parameters [20,21]. It is standardized in the ASTM and the Eurocode [3]. The DMT has been the object of a detailed monograph by the ISSMGE Technical Committee TC16 [22–24].

In the case of dilatometer (DMT) investigations, the diagrams generally developed by Marchetti were commonly used [1,2] (Table 1, Figure 2). Marchetti [1] proposed a soil classification based on the \( I_D \) material index where the value of the material index \( I_D < 0.10 \) points to peat or sensitive clays with no clear discrimination between them. However, it should be pointed out that the diagram was developed based mainly on soil mineral tests. The Marchetti and Crapps diagram shows the relationship between the material index \( I_D \) and the dilatometer modulus \( E_D \) (in a log-log scale). The unit weight of soils (cohesive soils and non-cohesive soils) and their states are also presented in this diagram. Soils are classified as organic soils where the material index \( I_D < 0.6 \) and the dilatometer modulus \( E_D < 1.2 \) MPa.

| Soil Type                  | Material Index \( I_D \) (-) |
|----------------------------|-------------------------------|
| Peat/Sensitive clays       | \(<0.10\)                     |
| Clay                       | \(0.10-0.35\)                 |
| Organic soils and cohesive soils | \(0.35-0.60\)     |
| Silty clay                 | \(0.60-0.90\)                 |
| Clayey silt                | \(0.90-1.20\)                 |
| Silt                       | \(1.20-1.80\)                 |
| Sandy silt                 | \(1.80-3.00\)                 |
| Silty sand                 | \(>3.30\)                     |
| non-cohesive soils         |                               |
| Sand                       |                               |

Based on the analysis of dilatometer test results for pre-consolidated cohesive soils and organic soils, Larsson [7] proposed a revision of the value of the material index \( I_D \) by taking into account the impact of pre-consolidation on the change of its value. The adjusted value of the material index \( I_{D(kor)} \) according to Larsson’s recommendations [7] can be determined from the following relationships (Figure 3):

- for depth \(<2.0 \) m at \( K_D > 2.5 \)

\[
I_{D(kor)} = I_D - 0.075 \cdot (K_D - 2.5)
\]

(9)

- for depth \(\geq2.0 \) m at \( K_D > 2.5 \)

\[
I_{D(kor)} = I_D - 0.035 \cdot (K_D - 2.5)
\]

(10)

If \( K_D < 2.5 \) and/or \( I_D \leq 0.1 \) to \( I_{D(kor)} = I_D \).

The nomogram chart proposed by Larsson [7] to determine the type of soil and its bulk density, based on the adjusted value of the material index \( I_{D(kor)} \) and the dilatometer modulus \( E_D \), is shown in Figure 4. The characterization of soil on the corrected material index \( I_{D(kor)} \leq 0.6 \) (Table 1) is based on undrained shear strength \( \tau_{fu} \), using the division proposed by Leroueil and co-workers [25] (Table 2).
In order to reduce the necessity of using various types of equipment, field research methods are being sought that allow interpretation of the results obtained in a wide range. One of the field studies that meets this requirement is the Marchetti dilatometer [1], used more often in the country. The biggest advantage of studying the dilatometer is fast and not very complicated measurement, on the basis of which it is possible to determine the profiles of many soil parameters. The interpretation of ground parameters is based on the use of empirical relationships linking the results of measurements to the values of ground parameters [20,21]. It is standardized in the ASTM and the Eurocode [3]. The DMT has been the object of a detailed monograph by the ISSMGE Technical Committee TC16 [22–24].

In the case of dilatometer (DMT) investigations, the diagrams generally developed by Marchetti were commonly used [1,2] (Table 1, Figure 2). Marchetti [1] proposed a soil classification based on the ID material index where the value of the material index $ID < 0.10$ points to peat or sensitive clays with no clear discrimination between them. However, it should be pointed out that the diagram was developed based mainly on soil mineral tests. The Marchetti and Crapps diagram shows the relationship between the material index $ID$ and the dilatometer modulus $ED$ (in a log-log scale). The unit weight of soils (cohesive soils and non-cohesive soils) and their states are also presented in this diagram. Soils are classified as organic soils where the material index $ID < 0.6$ and the dilatometer modulus $ED < 1.2$ MPa.

Table 1. Soil classification based on material index $ID$ [1].

| Soil Type          | Material Index $ID$ (-) |
|--------------------|-------------------------|
| Organic soils      |                         |
| Peat/Sensitive clays | $< 0.10$                |
| Clay               | $0.10–0.35$             |
| Silty clay         | $0.35–0.60$             |
| Clayey silt        | $0.60–0.90$             |
| Silt               | $0.90–1.20$             |
| Sandy silt         | $1.20–1.80$             |
| non-cohesive soils |                         |
| Silty sand         | $1.80–3.30$             |
| Sand               | $>3.30$                 |

Figure 2. Chart for estimating soil type and unit weight for soil (normalized to $\gamma_w$) [1,2].

Based on the analysis of dilatometer test results for pre-consolidated cohesive soils and organic soils, Larsson [7] proposed a revision of the value of the material index $ID$ by taking into account the impact of pre-consolidation on the change of its value. The adjusted value of the material index $ID(kor)$ according to Larsson’s recommendations [7] can be determined from the following relationships (Figure 3):

- for depth $< 2.0$ m at $KD > 2.5$
  \[ ID(kor) = ID - 0.075 \times (KD - 2.5) \]  

- for depth $\geq 2.0$ m at $KD > 2.5$
  \[ ID(kor) = ID - 0.035 \times (KD - 2.5) \]  

If $KD < 2.5$ and/or $ID \leq 0.1$ to $ID(kor) = ID$

The nomogram chart proposed by Larsson [7] to determine the type of soil and its bulk density, based on the adjusted value of the material index $ID(kor)$ and the dilatometer modulus $ED$, is shown in Figure 4. The characterization of soil on the corrected material index $ID(kor) \leq 0.6$ (Table 1) is based on undrained shear strength $\tau_{fu}$, using the division proposed by Leroueil and co-workers [25] (Table 2).

The above results do not include the upper (dry) 2-m thick soil layer
- ● - Swedish clay,
- ▼ - Norwegian clay,
- ♦ - London clay

Figure 3. Difference between material index ($ID(Mar)$) and material index $ID(kor)$ for different clay content depending on the horizontal stress index $KD$ [7].
### Zone Classification*

| Zone | Classification* | Material index $I_D$ (-) |
|------|------------------|-------------------------|
| 1    | clay with thin silt layers/silty clay/(organic clay) | 1.45 |
| 2    | clay/(organic clay) | 1.45 |
| 3    | clay/(gyttja/peat) | 1.4 |
| 4    | Silty clay/(gyttja/peat) | 1.4 |
| 5    | organic clay/(clay with thin silt layers/silty clay) | 1.3 |
| 6    | organic clay/(clay) | 1.3 |
| 7    | gyttja/peat/(clay) | 1.25 |
| 8    | gyttja/peat/(silty clay) | 1.25 |
| 9    | organic clay/gyttja/peat | 1.15 |

*For clayey and organic soils, the notations are very loose. Loose, medium stiff, stiff and very stiff are given on the basis of the estimated shear strength.

**Figure 4.** Chart for estimating mineral and organic soils [7].

**Table 2.** Soil states based on undrained shear strength [25].

| Description of Soil State | Undrained Shear Strength $\tau_{fu}$ [kPa] |
|--------------------------|------------------------------------------|
| Very soft                | <12.5                                    |
| Soft                     | 12.5–25                                  |
| firm                     | 25–50                                    |
| Stiff                    | 50–100                                   |
| Very stiff               | 100–200                                  |
| Hard                     | >200                                     |
Geotechnical Conditions of Test Sites

This paper presents the test results of mineral and organic subsoil obtained from Antoniny and Koszyce sites located in the Noteć river valley in the Wielkopolska province, Nielisz site located in the Wieprz river valley in Lublin province, the SGGW Campus with the Department of Geotechnical Engineering SGGW, and the Stegny site located in Warsaw, where a laboratory and field testing programme has been carried out under and outside of the main dam embankment [26–28]. The location of all analyzed objects is shown in Figure 5. The grain size distribution curve obtained from laboratory tests for mineral soil from the described sites is presented in Figure 6.

Figure 5. Location of test sites in the region of Poland.

Figure 6. Grain size distribution curve obtained from laboratory tests for the mineral soils from the described sites.
The Antoniny test embankment was designed and performed in the frame of cooperation between the Department of Geotechnical Engineering SGGW and the Swedish Geotechnical Institute (SGI). The physical properties of the soil at the Antoniny site were determined during previous WULS-SGGW tests. The peat and gyttja layers have a thickness of 7.5 m; the ground is preconsolidated with the overconsolidated ratio OCR at 3–5 for peat, and 1.5–2.5 for gyttja [29–31]. The embankment was located in the Noteć river valley on organic sediments, which contain two layers: peat with a thickness of 4.1 m and gyttja with a thickness of 3.7 m. Generally, the organic subsoil is composed of amorphous peat with varying carbonate gyttja, and variable content of organic matter and calcium carbonate. In the amorphous peat, the content of organic parts I\text{om} ranges from 65% to 75%, the content of calcium carbonate CaCO\text{3} is equal to 10 \div 15% at moisture w\text{n} between 310 and 340%, and the determined liquidity limit W_{L} is 305 \div 450%. The unit density ρ is 1.05 \div 1.10 g/cm\text{3}, when the specific density is ρ\text{s} = 1.45 \div 1.50 g/cm\text{3}. In the calcareous gyttja, the content of organic parts I\text{om} ranges from 5% to 20%, the content of calcium carbonate CaCO\text{3} is equal to 65 \div 90% at moisture w\text{n} between 105 and 140%, and the determined liquidity limit W_{L} is 80 \div 110%. The unit density ρ is 1.25 \div 1.40 g/cm\text{3}, whereas the specific density ρ\text{s} = 2.2 \div 2.30 g/cm\text{3} [26–28] (Table 3).

The Koszyce test dam was located in the Ruda river valley. A subsoil layer of soft organic soils was discovered in the central part of the dam. The organic soils are Quaternary deposits of an oxbow lake. The thickness of organic soils in this region generally exceeds 10 m and locally even 20 m. Dense sand occurs under the organic soils. The upper organic soils in the test area consist of a 2.5 m thick peat layer on the top of a 10.5 m thick gyttja layer underlain by a sand layer. In the amorphous peat, the content of organic parts I\text{om} ranges from 70% to 85% and the content of calcium carbonate CaCO\text{3} is equal to 5 \div 15% at moisture w\text{n} between 400% and 550%, and the determined liquidity limit W_{L} is 450%. The unit density ρ is 1.05 \div 1.10 g/cm\text{3}, at the specific density ρ\text{s} = 1.45 \div 1.50 g/cm\text{3}. Based on the index properties, the gyttja layer was sub-divided into three layers, the first one with a thickness of 2.5 to 6.3 m. In the calcareous gyttja (G\text{y}), the content of organic parts I\text{om} ranges from 10% to 20% and the content of calcium carbonate CaCO\text{3} is equal to 65 \div 80% at moisture w\text{n} in the range of 120% and 160% and the determined liquidity limit W_{L} is 80 \div 110%. The unit density ρ is 1.20 \div 1.35 g/cm\text{3}, at the specific density ρ\text{s} = 2.1 \div 2.25 g/cm\text{3}. The second layer has a thickness from 6.3 to 10.5 m and represents calcareous gyttja (G\text{cy}). Its content of organic parts I\text{om} ranges from 10% to 20% and the content of calcium carbonate CaCO\text{3} is equal to 65 \div 75% at moisture w\text{n} between 180% and 220%, and the determined liquidity limit W_{L} is 100 \div 110%. The unit density ρ is 1.25 \div 1.30 g/cm\text{3}, whereas the specific density ρ\text{s} is 2.20 g/cm\text{3}. The third gyttja (calcareous-organic) (G\text{cy}) layer has the following properties: the content of organic parts I\text{om} ranges from 10% to 15% and the content of calcium carbonate CaCO\text{3} is equal to 70 \div 75% at moisture w\text{n} between 135% and 140%, and the determined liquidity limit W_{L} is 105%. The unit density ρ is 1.30 \div 1.35 g/cm\text{3}, whereas the specific density ρ\text{s} = 2.2 g/cm\text{3} [26–28]. The static ground water level is present in the peat layer at the depth of 0.5 m below the ground surface. The preconsolidation pressure obtained from oedometer tests is higher than the initial values of effective vertical stresses, which shows that organic soils are overconsolidated with an overconsolidation ratio, OCR, in the range of 1.5 \div 4 [26–28] (Table 3).

The physical properties of the soil at the Nielisz site were determined during previous WULS-SGGW tests. The layer of soft subsoil has a thickness of 3 m to 5 m; the ground is slightly preconsolidated [29–31]. Two layers of organic subsoil were distinguished at the Nielisz site. In the first layer, the content of organic parts ranges from 20% to 30% at moisture between 120% and 150% and the determined liquidity limit W_{L} is 130 \div 150%. The bulk density is 1.25 \div 1.30 g/cm\text{3}, whereas the specific density is 2.25 \div 2.30 g/cm\text{3}. In the second layer lying below, the content of the organic part is 10 \div 20% at a moisture of 105 \div 120% and a liquidity limit of 110 \div 130%; the bulk density of the layer is 1.30 \div 1.45 g/cm\text{3}, and the specific density is 2.30 \div 2.40 g/cm\text{3}. These layers are separated by sandy silt. Beyond the existing embankment under the downstream berm and the upstream slope, the soft soils are overconsolidated with an overconsolidation ratio, OCR, decreasing from 3 to 2 with depth [28] (Table 3).
Taking into account the physical and mechanical properties of the soils, five geotechnical layers were distinguished in the grounds of the WULS-SGGW Campus (Figure 6). Layer I consists of fluvioglacial layers of the Warta Glaciation (⁶QₚW)—medium and fine sands, with relative density $D_r = 0.35 \div 0.55$, and clay sands, sandy clays and silt with $I_L = 0.15 \div 0.20$. Layer II represents the meltout sediments of the Warta Glaciation (bQₚW)—medium and fine sands with $I_D = 0.3 \div 0.5$, and sandy clay and clay sands with $I_L = 0.0 \div 0.20$ and $I_L = 0.25 \div 0.54$. Layer III is brown glacial clay from the Warta Glaciation (gQₚW)—sandy clays with $I_L = 0.0 \div 0.11$. Layer IV is grey glacial clay from the Odra Glaciation (⁶QₚO)—sandy boulder clays with $I_L = 0.0 \div 0.12$. Layers III and IV are similar in terms of plasticity, but clearly differ in the sand fraction content. The sandy clays layer III contains a few percent more of the sand fraction, which together with the analysis of the results of DMT sounding were the basis for the separation of these layers into sublayers. Layer V comprises river sediments of the Mazovian Interglacial (⁶QₚM)—fine and medium sands, in the roof very compact layers with a relative density $D_r = 0.8 \div 0.9$ (Figure 6). Boulder clays with the OCR = 3 \div 7 are similar in terms of plasticity, but clearly differ in the content of the sand fraction [6] (Table 3).

The Stegny site is located in southern Warsaw, where a few sedimentation cycles, from sands to clays, were observed in vertical succession. The entire complex of Pliocene clays comprises of clays, silty clays (60–70%), silts (10–25%), and sands (10–20%). The CaCO₃ and organic matter contents do not exceed 5% and 1%, respectively. The basic properties of the Mio-Pliocene clays are presented in Table 3.

### Table 3. Index properties of organic soils at the Antoniny, Koszyce, Nielisz, Stegny and SGGW Campus test sites.

| Site         | Type of Soil          | Organic Content $I_{om}$ [%] | CaCO₃ Content [%] | Water Content $w_s$ [%] | Liquid Limit $w_L$ [%] | Unit Weight of Soil $\rho$ [t/m³] | Specific Weight of Soil $\rho_s$ [t/m³] |
|--------------|-----------------------|-----------------------------|-------------------|-------------------------|------------------------|-----------------------------------|----------------------------------|
| Antoniny     | amorphous Peat        | 65–75                       | 10–15             | 310–340                 | 305–450                | 1.05–1.10                        | 1.45–1.50                        |
|              | calcareous Gyttja     | 5–20                        | 65–90             | 105–140                 | 80–110                 | 1.25–1.40                        | 2.2–2.30                         |
| Koszyce      | amorphous Peat (Gₚₓ) | 70–85                       | 5–15              | 400–550                 | 450                    | 1.05–1.1                         | 1.45–1.50                        |
|              | calcareous Gyttja (Gₚₓ) | 10–20                  | 65–80             | 120–160                 | 80–110                 | 1.20–1.35                        | 2.1–2.25                         |
|              |                        | 15–20                       | 65–75             | 180–220                 | 100–110                | 1.25–1.30                        | 2.2                               |
| Nielisz      | Organic mud (Mₓ)      | 20–30                       | -                 | 120–150                 | 130–150                | 1.25–1.30                        | 2.25–2.3                         |
|              | Organic mud (Mₓ)      | 10–20                       | -                 | 105–120                 | 110–130                | 1.30–1.45                        | 2.30–2.40                        |
| Stegny       | Pliocene clays        | -                           | 19.20–28.50       | 67.6–88.0               | 2.1–2.2                | 2.68–2.73                        |                                  |
| SGGW Campus  | Boulder clay          | -                           | 5.20–20.10        | 21.9–26.6               | 2.0–2.2                | 2.68–2.73                        |                                  |

Based on laboratory tests, Figure 4 shows the grain distribution curve for all analyzed objects.

### 3. Results

#### 3.1. Dilatometer Tests Results

The test results obtained for selected sites are presented in Figures 7 and 8. They were taken into account in the construction of a new classification system for organic soils presented in the following chapter.

#### 3.2. Proposed Classification Chart

The diagram chart developed in this paper is based on the diagram proposed by Marchetti and Crapps [2] (Figure 9). In this paper, based on the analysis of the dilatometer test results for pre-consolidated mineral and organic soils, it is proposed to introduce direct values $p_1$—B-pressure reading corrected for $Z_m$ and $\Delta B$ membrane stiffness at 1.10 mm expansion to give the total soil stress acting normal to the membrane at 1.10 mm membrane expansion on the vertical axis, $p_0$ and $u_0$. Figure 10 shows the classification chart proposed in this paper based on $p_1$ and soil type index $I_{SDMT}$.
values to determine the soil type, its bulk weight, and undrained shear strength. Based on the $u_0$, $p_0$ and $p_1$ values, the soil states were separated (Table 1); they were established on the undrained shear strength $\tau_{fu}$, using the division proposed by Leroueil together with co-workers [24,32].

Based on laboratory tests, Figure 4 shows the grain distribution curve for all analyzed objects.

### Table 3.

**Index properties of organic soils at the Antoniny, Koszyce, Nielisz, Stegny and SGGW Campus test sites.**

| Site       | Type of Soil        | Organic Content | Iom [%] | CaCO3 Content | [%] | Water Content | wn [%] | Liquid Limit | wL [%] | Density Unit Weight of Soil $\rho$ [t/m³] | Specific Weight of Soil $\rho_s$ [t/m³] |
|------------|---------------------|-----------------|---------|---------------|-----|---------------|--------|--------------|--------|------------------------------------------|------------------------------------------|
| Antoniny   | amorphous Peat      | 65–75           | 10–15   | 310–340       | 305–450 | 1.05–1.10     | 1.07   |                          |        | 1.45–1.50                                |                                          |
|            | calcareous Gyttja   | 5–20            | 65–90   | 105–140       | 80–110  | 1.25–1.40     | 2.2    |                          |        | 2.2–2.30                                 |                                          |
| Koszyce    | amorphous Peat      | 70–85           | 5–15    | 400–550       | 450    | 1.05–1.1      | 1.05   |                          |        | 1.45–1.50                                |                                          |
|            | calcareous Gyttja (Gy) | 10–20          | 65–80   | 120–160       | 80–110  | 1.20–1.35     | 2.1    |                          |        | 2.1–2.25                                 |                                          |
|            | calcareous Gyttja (Gy) | 15–20          | 65–75   | 180–220       | 100–110 | 1.25–1.30     | 2.2    |                          |        | 2.2                                     |                                          |
| Nielisz    | Organic mud (Mor)   | 20–30           | -       | 120–150       | 130–150 | 1.25–1.30     | 2.25   |                          |        | 2.25–2.3                                  |                                          |
|            | Organic mud (Mor)   | 10–20           | -       | 105–120       | 110–130 | 1.30–1.45     | 2.3    |                          |        | 2.30–2.40                                |                                          |
| Stegny     | Pliocene clays      | -               | 19.20–28.5 | 67.6–88.0   | 2.1–2.2 | 2.68–2.73   |        |                          |        |                                          |                                          |
| SGGW Campus| Boulder clay        | -               | 5.20–20.1 | 21.9–26.6   | 2.0–2.2 | 2.68–2.73   |        |                          |        |                                          |                                          |

Analysis of the traced points on Marchetti’s nomogram by inserting the dilatometer modulus $E_D$ (MPa) on the vertical axis and the index $I_D$ (-) value on the horizontal axis shows that only this part of the non-cohesive soil area ($I_D > 1.8$) gives clear discrimination of this group. However, the remaining soils of these divisions are not visible. Therefore, an action was undertaken to determine the boundaries of the division of a particular group of soils (residual mineral soils and organic soils); presented in Figure 10.

In order to create a mechanism for sub-dividing the area for each soil, a new interpretation of dilatometer results was proposed as follows by introducing the values $p_1$ (MPa) and $S_{DMT} = (p_0 - u_0)/p_1$. Namely, it was necessary to enter $p_1$ (MPa) values on the vertical axis and $S_{DMT} = (p_0 - u_0)/p_1$ values on the horizontal axis. On the basis of this technique, it can be noticed that in the case of non-cohesive soils, results similar as in Marchetti’s nomogram will be obtained. However, for both cohesive soils and organic soils, this new approach gives a clear subdivision of the area for a specific...
soil. As shown in Figure 10, seven areas may currently be distinguished. The areas are depicted by variously coloured dashed lines: the line in black represents non-cohesive soils, the line in blue stands for silt soils, the line in brown represents clay soils, the line in violet is the transition area, the line in red is gyttja, the line in green is mud and organic mud, and the line in grey represents peat.

Figure 8. Profiles of material indexes $I_D$ from the DMT tests carried out for the subsoil of the Nielisz, Stegny and the SGGW Campus sites: (a)—before loading, (b)—after loading for Nielisz site; 1—measured values, 2—average values ± one standard deviation.

To recognize the residual mineral soils and organic soils in a more detailed manner compared to Marchetti and Crapps [2], the proposed diagram shows the relationship between the second reading $p_1$ and the soil type index $I_{SDMT}$ (Equation (11)). The $I_{SDMT}$ soil type index values can be calculated using Equation (11). Subsoils are classified as organic soils when the soil type index $0.40 < I_{SDMT} \leq 1.0$ and the second pressure dilatometer reading $p_1$ is in the range of $0.01\text{ MPa} < p_1 \leq 1.0\text{ MPa}$ (Tables 4 and 5). The new diagram contains 10 areas: 1 ÷ 4—non-cohesive soils, 5—silt, 6—clay, 7—gyttja, 8—mud/organic mud, and 9 ÷ 10—peat (Figure 11a–c). For a more precise subdivision of organic soils, an additional nomogram chart was developed for areas 7, 8, 9 and 10. The detailed form and the use of this nomogram chart is shown in Figure 11b,c.

Soil type index:

$$I_{SDMT} = \frac{p_0 - u_o}{p_1}$$ (11)
Figure 9. Chart for estimating soil type and unit weight (normalized to $\gamma_w$) [2].

Figure 10. Variation of $p_1$ with soil type index $S_{LDMT} = (p_0 - u_0)/p_1$. 

Equation of Lines A-D

$E_D (\text{MPa}) = 10^m \log m + \text{log ID}$

| Line   | $m$   | $n$   |
|--------|-------|-------|
| A      | 0.585 | 0.737 |
| B      | 0.621 | 1.013 |
| C      | 0.657 | 1.289 |
| D      | 0.694 | 1.564 |

$Dilatometer\ modulus\ E_D (\text{MPa})$

$SILT$ | Very $SAND$ | (2.15) | Line D

$CLAY$ | Hard | (2.05)

$Stiff$ | (1.90) |

$Medium\ Stiff$ | (1.80) | Compress

$Soft$ | (1.70) |

$Very$ | (1.60) |

$MUCK/PEAT$

$G_{B} = \text{Approximate\ Bulk\ Specific}\ Gravity\ shown\ in\ parentheses$ 

$\text{Reduce}\ G_B\ by\ 0.10\ if\ Plasticity\ Index > 50\%$
Table 4. Proposed soil classification based on $I_{SDMT}$ and $p_1$ from in situ (DMT) tests.

| Zone | Description          | $\rho$ (t/m$^3$) | $\rho_d$ (t/m$^3$) | Void Ratio $e$ (-) | $w_n$ (%) | $p_1$ (MPa) | $I_{SDMT}$ (-) | $\tau_{fu}$ (MPa) from FVT |
|------|----------------------|------------------|--------------------|-------------------|-----------|------------|---------------|--------------------------|
| Residual mineral soils | | | | | | | | |
| 1 | Coarse sand (CSa) | 1.70 $\div$ 2.15 | 1.65 $\div$ 2.0 | 0.20 $\div$ 0.48 | 4.0 $\div$ 25 | 0.2 $< p_1 \leq$ 10 | 0.0 $\div$ 0.1 | 0.10 $\leq \tau_{fu} < 0.30$ |
| 2 | Medium sand (MSa) | 1.70 $\div$ 2.15 | 1.65 $\div$ 2.0 | 0.20 $\div$ 0.48 | 4.0 $\div$ 25 | 0.2 $< p_1 \leq$ 10 | 0.1 $\div$ 0.2 | 0.10 $\leq \tau_{fu} < 0.30$ |
| 3 | Fine sand (Fsa) | 1.70 $\div$ 2.15 | 1.60 $\div$ 2.0 | 0.20 $\div$ 0.48 | 5.0 $\div$ 28 | 0.2 $< p_1 \leq$ 10 | 0.2 $\div$ 0.3 | 0.10 $\leq \tau_{fu} < 0.30$ |
| 4 | Silty sand(siSa) | 1.70 $\div$ 2.15 | 1.59 $\div$ 2.0 | 0.20 $\div$ 0.48 | 5.0 $\div$ 28 | 0.2 $< p_1 \leq$ 10 | 0.3 $\div$ 0.4 | 0.10 $\leq \tau_{fu} < 0.30$ |
| 5 | SILT(Si) | 1.60 $\div$ 2.10 | 1.4 $\div$ 1.90 | 0.18 $\div$ 0.70 | 10 $\div$ 30 | 2.0 $< p_1 \leq$ 10 | 0.4 $\div$ 0.75 | 0.02 $\leq \tau_{fu} < 0.5$ |
| 6 | CLAY(Cl) | 1.60 $\div$ 2.10 | 1.50 $\div$ 1.70 | 0.18 $\div$ 0.40 | 20 $\div$ 40 | 0.5 $< p_1 \leq$ 2 | 0.4 $\div$ 0.75 | 0.02 $\leq \tau_{fu} < 0.6$ |
| Organic soils | | | | | | | | |
| 7 | Gyttja (Gy) | 1.20 $\div$ 1.40 | 0.50 $\div$ 0.60 | 2.5 $\div$ 3.2 | 110% $\div$ 150% | 0.09 $< p_1 < 0.2$ | | 0.0125 $\leq \tau_{fu} < 0.0255$ |
| 8 | Mud (M) or Organic mud (Mor) | 1.25 $\div$ 1.70 | 0.54 $\div$ 0.67 | 2.6 $\div$ 3.2 | 110% $\div$ 140% | 0.2 $< p_1 \leq$ 0.5 | | 0.0255 $< \tau_{fu} < 0.0505$ |
| 9 | peat | 1.05 $\div$ 1.10 | 0.17 $\div$ 0.244 | 4.5 $\div$ 7.3 | 350% $\div$ 500% | $p_1 < 0.09$ | | $\tau_{fu} < 0.0125$ |
| 10 | | | | | | | | $\tau_{fu} > 0.0125$ |

Table 5. Proposed soil classification based on $I_{SDMT}$ and $p_1$ from in situ (DMT) tests.

| Soil Types \ Parameters | Residual Mineral Soils | Organic Soils |
|-------------------------|-----------------------|---------------|
|                         | Non-Cohesive Soils | Cohesive Soils | Gyttja (Gy) | Mud (M) and Organic Mud (Mor) | Peat |
| $I_{SDMT}$ | 0.0 $< I_{SDMT} \leq$ 0.4 | 0.4 $< I_{SDMT} \leq$ 0.75 | | 0.4 $< I_{SDMT} \leq$ 1.0 |
| $p_1$ (MPa) | 0.1 $< p_1 \leq$ 10 | 1.0 $< p_1 \leq$ 10 | 0.09 $< p_1 < 0.2$ | 0.2 $< p_1 \leq$ 0.5 | $p_1 < 0.09$ |
Figure 11. Chart estimating the type and state of residual mineral soils and organic soils.
4. Conclusions

Peat and gyttja, as well as organic mud, are located in the lower boundary zone of mineral soils, and in some cases, slightly above the limit set by Marchetti for organic soils. A very good and effective complement to this system are the methods developed in this article. The new system facilitates a more accurate separation of soils in the group into gyttja, peat, organic mud, clay, silt, and sands. This possibility is also confirmed in this study, in which gyttjas are relatively well-discriminated from organic mud, and peat is clearly separated from organic mud. In addition, distinct subdivisions can also be seen between silt, clay, and sand.

An analysis of the subsoil of mineral and organic soils that have been carried out in the DMT study allow for the conclusion that the identification of the zone and the range of non-cohesive, cohesive, and organic soils in the subsoil is possible by means of DMT classification systems using $p_0$, $p_1$ and $u_o$ parameters. The effectiveness of the systems depends on the complex number of factors affecting the parameters measured in DMT ($p_0$, $p_1$).

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References
1. Marchetti, S. In Situ Tests by Flat Dilatometer. J. Geotech. Geoenviron. Eng. 1980, 106, 299–321.
2. Marchetti, S.; Crapps, D.K. Flat Dilatometer Manual; Internal Report of GPE; GPE Inc.: Gainesville, FL, USA, 1981.
3. EN 1997-1. Eurocode 7: Geotechnical Design—Part 1; General Rules; CEN: Brussels, Belgium, 2001.
4. Frank, R.; Bauduin, C.; Driscoll, R.; Kavvadas, M.; Krebs Ovesen, N.; Orr, T.; Schuppener, B. Designers’ Guide to EN 1997-1 Eurocode 7: Geotechnical Design—General Rules; Thomas Telford Service Ltd.: London, UK, 2004; 216p.
5. Lechowicz, Z. Profiling of stress history in organic soils by dilatometer test. Ann. Wars. Agric. Univ. SGGW Land Reclam. 1997, 28, 97–105.
6. Interim Reports—Geotechnical Documentations for Design SGGW Campus Building (2000–2005); Department of Geotechnical Engineering, SGGW: Warsaw, Poland, 2000–2005. (In Polish)
7. Larsson, R. Dilatometr Försök för Bedömning av Jordläggerföljd och Egenskaper i Jord. Information No. 10 [Dilatometer Test. An In-Situ Method for the Determination of Layer Sequences and Soil Properties. Information No. 10]; Swedish Geotechnical Institute: Linköping, Sweden, 1989. (In Swedish)
8. Mayne, P.W. Evaluating effective stress parameters and undrained shear strength of soft-firm clays from CPT and DMT. In Pursuit of Best Practices—Proceedings of the 5th International Conference on Geotechnical & Geophysical Site Characterization (ISC-5), Queensland, Australia, 5-9 September; Australian Geomechanics Society: Sydney, Australia, 2016; Volume 1, pp. 19–40. ISBN 978-0-9946261-1-0.
9. Long, M.; Boylan, N.; Powell, J.; O’Connor, S.; Donohue, S. Characterisation of the soils beneath the flood banks along the River Thames estuary. In Proceedings of the 4th International Workshop—Soil Parameters from In Situ Tests, Poznań, Poland, 27 September 2010; pp. 395–411.
10. Schmertmann, J.H. Suggested method for performing the flat dilatometer test. ASTM Geotech. Test. J. 1988, 9, 93–101.
11. Briand, J.; Miran, J. The Flat Dilatometer Test; Report No. FHWA-SA-91-044; US DOT Federal Highway Administration: Springfield, VA, USA, February 1992; pp. 1–102.
12. Młynarek, Z.; Niedzielski, A.; Tschuschke, W. Variability of shear strength and physical parameters of peat. In Proceedings of the 7th Danube European Conference on Soil Mechanic and Foundation Engineering, Kishinev, Moldawi, 2–5 October 1983; pp. 93–98.
13. Lechowicz, Z.; Rabarijoely, S. Dilatometer C-reading to help determine stratigraphy of organic subsoil. Ann. Wars. Agric. Univ. Land Recl. 2000, 29, 71–78.
14. Młynarek, Z.; Tschuschke, W.; Wierzbicki, J.; Marchetti, S. Interrelationship between shear and deformation parameters for gyttja and peat from CPT and DMT tests. In Proceedings of the 13th Danube European Conference on Geotechnical Engineering, Ljubljana, Slovenia, 29–31 May 2006.
15. Młynarek, Z.; Wierzbicki, J.; Long, M. Factors affecting CPTU and DMT characteristics in organic soils. In Proceedings of the 11th Baltic Sea Geotechnical Conference: Geotechnical in Maritime Engineering, Gdańsk, Poland, 15–18 September 2008; Młynarek, Z., Sikora, Z., Dembicki, E., Eds.; Polish Committee on Geotechnics & Gdańsk University of Technology: Gdańsk, Poland, 2008; Volume 1, pp. 407–417.

16. Młynarek, Z.; Wierzbicki, J.; Stefaniak, K. CPTU, DMT, SDMT results for organic and fluvial soils. In Proceedings of the 2nd International Symposium on Cone Penetration Testing (CPT ’10), Huntington Beach, CA, USA, 9–11 May 2010; pp. 455–462.

17. Bihs, A.; Long, M.; Marchetti, D.; Ward, D. Interpretation of CPTU and SDMT in organic, Irish soils. In Proceedings of the 2nd International Symposium on Cone Penetration Testing, CPT10, Huntington Beach, CA, USA, 9–11 May 2010; Volume 2, pp. 257–264.

18. Lutenegger, A.J.; Kabir, M.G. Dilatometer C-reading to help determine stratigraphy. Proc. Int. Sym. Penetr. Test. 1988, 1, 549–553.

19. Totani, G.; Calabrese, M.; Marchetti, D.; Ward, D. Interpretation of CPTU and SDMT in organic, Irish soils. In Proceedings of the 2nd International Symposium on Cone Penetration Testing, CPT10, Huntington Beach, CA, USA, 9–11 May 2010; Volume 2, pp. 257–264.

20. Jamiolkowski, M.; Leroueil, S.; LO Presti, D.C.F. Design parameters from theory to practice. Theme lecture. In Proceedings of the International Conference on Geotechnical Engineering for Coastal Development, Yokohama, Japan, 3–6 September 1991.

21. Młynarek, Z. Badania laboratoryjne oraz badania podłoża in situ. Referat generalny. Mat. na XI Kraj. Konf. Mech. Grunt. i Fund 1997, 3, 113–126.

22. Marchetti, S.; Monaco, P.; Totani, G.; Marchetti, D. In situ tests by Seismic Dilatometer (SDMT). In From Research to Practice in Geotechnical Engineering; ASCE Geotechnical Special Publication N°. 180 Honouring; Schmertmann, J.H., Ed.; American Society Of Civil Engineers: New Orleans, LA, USA, March 2008; pp. 292–311.

23. Amoroso, S. Use of seismic dilatometer for design infrastructures. In Proceedings of the Conference on 18th Southeast Asian Geotechnical & Inaugural AGSSEA Conference, Singapore, 29–31 May 2013; pp. 663–669.

24. Ouyang, Z.; Mayne, P. Effective Stress Strength Parameters of Clays from DMT. Geotech. Test. J. 2018, 41, 851–867. [CrossRef]

25. Lerouel, S.; Magnan, J.; Tavenas, F. Embankments on Soft Clays; English Edition; Wood, D.M., Translator; Ellis Horwood: New York, NY, USA, 1990; 360p.

26. Wolski, W.; Szymański, A.; Milecki, J.; Lechowicz, Z.; Larsson, R.; Hartlen, J.; Garbulewski, K.; Bergdahl, U. Two Stage-Constructed Embankments on Organic Soils; Report 32; Swedish Geotechnical Institute: Linköping, Sweden, 1988.

27. Wolski, W.; Szymański, A.; Lechowicz, Z.; Larsson, R.; Hartlen, J.; Bergdahl, U. Full-Scale Failure Test on Stage-Constructed Test Fill on Organic Soil; Report 36; Swedish Geotechnical Institute: Linköping, Sweden, 1989.

28. Lechowicz, Z.; Rabarijoely, S. Wykorzystanie badań in situ w ocenie wzmocnienia słabonośnego podłoża zapory budowanej etapowo. The use in situ tests in the evaluation of subsoil strengthening at stages in dam construction. Mat. na VII Konf. Tech. Kon. Zap. 1996, 231–240. (In Polish)

29. Rabarijoely, S. The Use of Dilatometer Test for Evaluation of Organic Soil Parameters. Ph.D. Thesis, Warszaw Agricultural University—SGGW, Land Reclamation, Warszawa, Poland, 2000. (In Polish)

30. Lechowicz, Z.; Rabarijoely, S.; Szczypiński, P. Wykorzystanie badań dylatometrycznych do określania rodzaju gruntów organicznych. The use of dilatometer test to determine the type and state of organic soils. In Przegląd Naukowy Wydziału Melioracji i Inżynierii Środowiska; SGGW: Warszawa, Poland, 2004; pp. 191–201. (In Polish)

31. Lechowicz, Z.; Fukue, M.; Rabarijoely, S.; Sulewska, M.J. Evaluation of the Undrained Shear Strength of Organic Soils from a Dilatometer Test Using Artificial Neural Networks. Appl. Sci. 2018, 8, 1395. [CrossRef]

32. Smagin, A.V.; Sadovnikova, N.B.; Vasenev, V.I.; Smagina, M.V. Biodegradation of Some Organic Materials in Soils Band Soil Constructions: Experiments, Modeling and Prevention. Materials 2018, 11, 1889. [CrossRef] [PubMed]