**Article**

**Spectroscopic, Raman, EMPA, Micro-XRF and Micro-XANES Analyses of Sulphur Concentration and Oxidation State of Natural Apatite Crystals**

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**Abstract:** Sulphur is the third most abundant volatile element in Earth systems. Analytical methods for accurately and efficiently determining the sulphur content and oxidation state in natural minerals are still lacking. Natural apatite is widely distributed in the Earth and incorporates a large amount of sulphur. Therefore, apatite is an ideal mineral for performing sulphur measurements. Here, we used spectroscopic, Raman, X-ray diffraction, laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS), electron microprobe (EMPA) and micro-X-ray fluorescence spectrometry (micro-XRF) analysis techniques and developed a new analytical approach (i.e., micro-X-ray absorption near-edge structure (micro-XANES) analysis of the sulphur K-edge) to investigate the chemical characteristics of natural apatite. These multiple methods were developed to measure in situ sulphur concentration and S oxidation states and to assess a potential natural apatite reference material. Apatite contains chemically homogeneous sulphur, with micro-XANES located at the peak energies corresponding to S^{6+} (sulphate; ~2482 eV), S^{4+} (sulfite; ~2478 eV), and S^{2−} (sulphide; ~2467, 2470 and 2474 eV). The Durango apatite contains total S presented as SO_3 at amount of 0.332 ± 0.012 wt.% (1σ), with a large amount of S^{6+} and a small contribution of S^{4+}. The Kovdor apatite contains 44–100 ppm of S and is dominated by S^{6+}. These results indicate that the Durango apatite crystallised under relative oxidising conditions, and the Kovdor apatite has a higher oxygen fugacity than Durango. In addition, this study indicates the potential use of the natural apatite reference material with its S composition and S oxidation state.

**Keywords:** spectroscopy; Raman and micro-XANES

**1. Introduction**

Sulphur is the third most abundant volatile element in Earth systems [1] and is a significant component in metal complexes in magmatic ore systems, which are the major sources of copper, gold, and molybdenum in the world [2,3]. Methods for accurately and efficiently measuring the sulphur content and oxidation state are essential for understanding volatile evolution and related magmatic mineralisation processes. Natural apatites contain large amounts of sulphur and are widely distributed in the Earth [4]. Sulphur with different oxidation states (e.g., S^{2−}, S^{4+} and S^{6+}) can greatly affect metal solubility in parental melts and partitioning of ore metals between melts and ore-forming fluids [4]. Thus, apatite is an ideal sample for investigating S concentration and S oxidation state. In addition, establishing multiple methods for analysing natural apatite minerals (including major elements and the oxidation state of S in fluorapatite, hydroxylapatite and chlorapatite) is prerequisite for analytical research [5,6]. In this study, we used X-ray diffraction, infrared spectroscopy,
Raman spectroscopy, electron microprobe (EMPA), laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS), micro-X-ray fluorescence spectrometry (micro-XRF) and in situ sulphur X-ray absorption near-edge structure (micro-XANES) analyses to study chemical and S oxidation states in apatite from Durango, Mexico, and Kovdor, Russia. The data were also used to build natural apatite reference materials of major elements and S oxidation states.

2. Standards and Natural Apatite

The apatite group are the most common phosphate minerals, with the general formula of $\text{U}_{10}[\text{VO}_4]_6\text{W}_2$, where U is a divalent cation represented by Ca$^{2+}$ and Ba$^{2+}$, and Sr$^{2+}$, Mg$^{2+}$, Fe$^{2+}$, Ce$^{3+}$ and Mn$^{2+}$ are also incorporated as isomorphs (Figure 1). In addition, alkali ions K$^+$ and Na$^+$ can easily enter the apatite structure by coupling isomorphism. [VO$_4$] is mainly [PO$_4$]$^{3-}$, which can be substituted by [SO$_4$]$^{2-}$, [VO$_4$]$^{3-}$, [CO$_3$]$^{2-}$, [AsO$_4$]$^{3-}$, [SiO$_4$]$^{4-}$ and other complex anions. The anion W in the formula is commonly F$^-$, Cl$^-$, S$^{2-}$, Br$^-$ and OH$^-$. Natural apatites crystallize in the hexagonal system, commonly in the centrosymmetric space group $P6_3/m$, although there are examples of monoclinic $P2_1/b$ chlorapatite and hydroxylapatite end-members [7]. The $P6_3/m$ apatite structure has two non-equivalent cationic sites: Ca1 with nine coordination numbers and Ca2 with seven coordination numbers. The Durango apatite grains from Cerro de Mercado, Mexico [8] are normally homogenous with some carbonatites, and has relatively lower sulphur content ($0.007 \pm 0.002$ wt.%) than the Durango apatite ($0.133 \pm 0.005$ wt.%) [10].

![Optimised crystal structure of apatite.](image)

Figure 1. Optimised crystal structure of apatite.

In addition, we use other natural minerals as reference materials during XANES measurements. Natural alabandite, scapolite, lazurite, pyrite, gypsum and sodium sulfite are used as standards during analysis. The chemical formula of scapolite is Na$_4$[AlSi$_3$O$_8$](Cl, OH)-Ca$_4$[Al$_2$Si$_2$O$_8$](CO$_3$, SO$_4$), and the main component of gypsum is CaSO$_4$. These minerals contain S$^{6+}$. The chemical formula of lazurite is generally (Na, Ca)$_{7-8}$(Al, Si)$_{12}$(O, S)$_{24}$[SO$_4$, Cl$_2$(OH)$_2$], and lazurite contains S$^{6+}$ and S$^{2-}$ [11,12]. Pyrite (FeS) and alabandite (MnS) are used as the peak of S$^{2-}$ [13]. The chemical formula of sodium sulfite is Na$_2$SO$_3$, containing S$^{4+}$. 
3. Experiment

3.1. X-ray Diffraction (XRD)

The powder crystal X-ray diffraction (XRD) of the samples was acquired using a D8 Advance X-ray diffractometer (Bruker Scientific Instruments, China) at the China University of Geosciences, Beijing. The testing voltage was 40 KV, the current was 100 mA, the scanning pace was 8°/min. Furthermore, the test used continuous scanning mode with 2θ scanning of 10–70°. The results are listed in Table 1.

Table 1. Unit-cell parameters of Durango and Kovdor apatites.

| Samples                     | Unit-Cell Parameters |
|-----------------------------|----------------------|
|                             | a(Å)     | c(Å)     |
| DUR-1                       | 9.3657 (77) | 6.8685 (16) |
| DUR-2                       | 9.3659 (74) | 6.8984 (1)  |
| KOV                         | 9.3892 (25) | 6.8891 (1)  |
| Ca₅(PO₄)₃(OH) PDF#09-0432  | 9.4180    | 6.8840    |
| Ca₅(PO₄)₃F PDF#15-0876     | 9.3690    | 6.8840    |

3.2. Infrared Spectra

The infrared spectra of analysed samples were performed at the China University of Geosciences, Beijing by using the Tensor 27 Infrared Spectrometer produced by Germany. During the test, the temperature in the laboratory was controlled at 16–18 °C and the relative humidity was 36–38%. The voltage is 220–240 V, the frequency is 56–60 Hz, and the power is 250 W.

3.3. Raman Spectra

The Raman spectra of the Durango apatite were acquired using a Raman spectrometer (HR-Evolution, HORIBA) at the China University of Geosciences, Beijing, and the spectrometer was equipped with a 50×objective. The spectrometer used a 600-lines-per-millimetre diffraction grating and a Peltier-cooled, Si-based CCD detector, and a solid-state 100-mW 532-nm laser was used as the excitation source. The spectra were collected in the range of 200–2100 cm⁻¹ (10-s accumulation time, 3 scans). The instrument was calibrated by a silicon and synthetic standard.

3.4. EMPA

The major elements of apatite, presented in Table 2, were determined using an electron probe (EMPA) at the Institute of Geology, Chinese Academy of Geological Sciences, with a JEOL JXA-8800R microprobe. The representative points with different grains were analysed with a 2–5-microm beam diameter at 15-kV voltage and 20-nA beam current. Detection limits of oxides are below 0.01 wt.%.

3.5. LA-ICP-MS

The Durango and Kovdor apatite were examined to determine their Sr element content by using an Agilent 7900 ICP-MS with a NWR193¹²C Elemental Scientific Lasers at the Chinese Academy of Geological Sciences. The spot size was 25 μm, the laser energy density 2.5 J/cm², and the repetition rate 10 Hz. ⁴³Ca was used as the internal standard to calculate the trace-element compositions of unknown samples. The raw data were processed using the online software program GLITTER 4.4. The data are presented in Table 3.
Table 2. Representative EMPA major-element composition of Durango and Kovdor apatite.

| Sample | Na$_2$O | SiO$_2$ | CaO | P$_2$O$_5$ | SO$_3$ | FeO | F | Cl | Total |
|--------|---------|---------|-----|------------|--------|-----|---|----|-------|
| DUR-1  | 0.23    | 0.5     | 53.6| 39.97      | 0.32   | 0.07| 2.97| 0.43| 99.04 |
| DUR-2  | 0.27    | 0.53    | 54.2| 41.03      | 0.31   | 0.02| 3.15| 0.42| 100.81|
| DUR-3  | 0.28    | 0.51    | 53.82| 40.26      | 0.32   | 0.01| 3.05| 0.44| 99.73 |
| DUR-4  | 0.26    | 0.57    | 53.95| 41         | 0.33   | 0.05| 3.01| 0.39| 100.48|
| DUR-5  | 0.28    | 0.41    | 54.15| 40.99      | 0.32   | 0.07| 3.17| 0.43| 100.6 |
| DUR-6  | 0.27    | 0.34    | 54.25| 40.86      | 0.34   | 0.02| 3.01| 0.42| 100.4 |
| DUR-7  | 0.25    | 0.49    | 54.12| 40.66      | 0.32   | 0.02| 3.03| 0.4  | 100   |
| DUR-8  | 0.21    | 0.44    | 54.25| 40.86      | 0.34   | 0.07| 3.06| 0.41| 100.34|
| DUR-9  | 0.25    | 0.49    | 54.07| 40.62      | 0.34   | 0.03| 3.08| 0.42| 100.01|
| DUR-11 | 0.29    | 0.37    | 54.34| 40.98      | 0.34   | 0.03| 3.05| 0.43| 100.52|
| DUR-12 | 0.28    | 0.34    | 54.29| 41         | 0.33   | 0.04| 3.03| 0.42| 100.43|
| DUR-13 | 0.28    | 0.38    | 54.55| 41.18      | 0.34   | 0.04| 2.96| 0.41| 100.92|
| DUR-14 | 0.26    | 0.43    | 54.44| 40.67      | 0.35   | 0.03| 3.01| 0.41| 100.34|
| DUR-15 | 0.28    | 0.42    | 54.41| 40.68      | 0.34   | 0.04| 3 | 0.41| 100.37|
| DUR-16 | 0.3     | 0.36    | 54.45| 40.94      | 0.35   | 0.03| 3.11| 0.41| 100.69|
| KOV1-1 | 0.269   | 0.294   | 52.208| 40.151     | 0.015  | 0.007| 1.496| 0.014| 94.621|
| KOV1-2 | 0.301   | 0.259   | 52.375| 39.979     | 0.025  | -   | 1.595| 0.013| 94.694|
| KOV1-3 | 0.227   | 0.476   | 52.266| 39.83      | 0.011  | 0.009| 1.557| 0.013| 94.676|
| KOV2-1 | 0.272   | 0.61    | 52.007| 39.265     | 0.019  | 0.002| 1.63 | 0.018| 94.065|
| KOV2-2 | 0.251   | 0.554   | 52.111| 39.738     | 0.016  | 0.03 | 1.638| 0.027| 94.738|
| KOV2-3 | 0.231   | 0.579   | 52.211| 39.733     | 0.025  | 0.009| 1.572| 0.017| 94.731|
| KOV3-1 | 0.218   | 0.447   | 52.448| 39.941     | 0.017  | -   | 1.625| 0.02 | 94.92 |
| KOV3-2 | 0.234   | 0.527   | 52.259| 40.227     | 0.013  | 0.021| 1.877| 0.017| 95.252|
| KOV3-3 | 0.258   | 0.495   | 52.112| 39.968     | 0.013  | 0.028| 1.894| 0.017| 94.923|

Table 3. The Sr and Y content of apatites.

| Samples | Sr (ppm) | Y (ppm) |
|---------|----------|---------|
| DUR-01  | 477.86   | 555.13  |
| DUR-02  | 478.34   | 556.79  |
| DUR-03  | 477.73   | 574.91  |
| DUR-04  | 480.6    | 577.63  |
| DUR-05  | 480.55   | 577.54  |
| DUR-06  | 474.93   | 558.05  |
| DUR-07  | 477.54   | 560.46  |
| DUR-08  | 477.57   | 560.44  |
| DUR-09  | 482.11   | 550.49  |
| DUR-10  | 480.93   | 549.23  |
| DUR-11  | 480.76   | 549.08  |
| KOV-01  | 2767.97  | 222.11  |
| KOV-02  | 2771.66  | 222.77  |
| KOV-03  | 2726.27  | 261.3   |
| KOV-04  | 2739.79  | 262.33  |
| KOV-05  | 2740.4   | 262.35  |
| KOV-06  | 2771.72  | 229     |
| KOV-07  | 2784.01  | 229.8   |
| KOV-08  | 2785.06  | 229.85  |
| KOV-09  | 2612.97  | 194.45  |
| KOV-10  | 2604.21  | 193.44  |
| KOV-11  | 2601.36  | 193.28  |
| KOV-12  | 2801.04  | 230.65  |
| KOV-13  | 2785.84  | 228.78  |
| KOV-14  | 2781.77  | 228.53  |
| KOV-15  | 2688.25  | 207.2   |
3.6. Micro-XRF Elemental Mapping

The elemental mapping of the Durango apatite was performed using a micro-XRF spectrometer (BRUKER) at Macquarie University, Australia. The spectrometer was equipped with an X-ray tube and a vacuum pump and was operated at 50-kV voltage and 600-μA current. The data were processed using the mineral analyser software from BRUKER.

3.7. Sulphur X-ray Absorption Near-Edge Structure Spectroscopy (S XANES)

In situ sulphur XANES measurements at the K-edge were performed on the I-13 beamline at the Diamond Light Source (UK national synchrotron). XANES measurements were performed in a He atmosphere using a 6 × 6-μm, 3-keV beam, and the fluorescence spectra were acquired using a Vortex silicon drift detector and quantified using the PyMca software. For each sample, we also collected XRF maps at 3 keV, i.e., above the Cl K-edge but below the Ar K-edge, and P was used as an internal reference for S/Cl quantification. These conditions were optimised for S/Cl quantification because Ar is always present in a He atmosphere but in variable quantities; therefore, it affects the quantification [14].

4. Results and Discussion

4.1. X-ray Diffraction (XRD)

The unit-cell parameters of Durango apatite $a = 9.3657(77)$–$9.3659(74)$ Å, $c = 6.8984(1)$–$6.8685(16)$ Å. The unit-cell parameters of Kovdor apatite $a = 9.3892(25)$ Å, $c = 6.8891(1)$ Å (Table 1). Compared with PDF cards: 09-0432 (Ca$_5$(PO$_4$)$_3$(OH)) and 15-0876 (Ca$_5$(PO$_4$)$_3$F), the unit-cell parameters of two Durango apatite are consistent with Ca$_5$(PO$_4$)$_3$F, while Kovdor apatite is more inclined to Ca$_5$(PO$_4$)$_3$(OH) (Table 1, Figure 2).

4.2. Infrared Spectra

Typical infrared reflectance spectra for apatite are shown in Figure 3. DUR and KOV apatites exhibit peaks at 431, 465, 575, 607, 962, 1058, and 1105 cm$^{-1}$. The strong bands near 600 and 575 cm$^{-1}$, which are present in the infrared reflectance spectra of all apatite samples, are due to the bending vibration of PO$_4^{3-}$. The weak bands at 950–960 cm$^{-1}$ are due to the symmetrical stretching vibration of PO$_4^{3-}$. The shoulder peaks near 1050–1060 and 1100 cm$^{-1}$ are due to the anti-symmetric stretching vibration of PO$_4^{3-}$. There was such small amount of SO$_4^{2-}$ vibration bands at 609 cm$^{-1}$, as observed (Figure 3).

4.3. Raman Spectra

The Raman spectra of two apatite samples are shown in Figure 4. There are four major vibration peaks at 429–446, 588–579, 962–965, and 1048–1058 cm$^{-1}$ for the Durango apatite. The peak at 962–965 cm$^{-1}$ is a symmetrical stretching vibration of [PO$_4$]$^{3-}$, and it is the strongest characteristic peak of apatite [15]. The peaks at 1048–1058 and 588 cm$^{-1}$ are the asymmetric stretching vibration and anti-symmetric bending vibration of [PO$_4$]$^{3-}$. A vibration of [SO$_4$]$^{2-}$ should show peaks at 1005 cm$^{-1}$ and 1123 cm$^{-1}$, which did not appear in the DUR and KOV apatites as expected.
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4.4. Elemental Maps by Micro-XRF Mapping and Major-, Trace-Element Compositions of Apatite

The elemental maps of apatite made by micro-XRF are shown in Figure 5. Iron, calcium, chlorine, phosphorus, sulphur and strontium are uniformly distributed in the samples, and all of the above-mentioned elements show the same distribution in the DUR apatite. The major elements compositions of the analysed Durango and Kovdor apatite are shown in Table 2. In general, the Durango apatite grains contain 39.97–41.18 wt.% of P$_2$O$_5$ and 53.60–54.55 wt.% of CaO. The SO$_3$ concentration in Durango varies from 0.31 to 0.35 wt.% The Kovdor apatite grains contain 39.27–40.22 wt.% of P$_2$O$_5$ and 52.01–52.45 wt.% of CaO. The concentration of SO$_3$ in the Kovdor apatite ranges from 0.01 to 0.03 wt%. The Sr and Y content of the Durango apatites range from 474.93–482.11 ppm (average: 478.99 ppm) and 549.08–577.63 ppm (average: 560.88 ppm). Meanwhile, the Kovdor apatite grains contain higher Sr (2601.36–2801.04 ppm, average: 2730.82 ppm) and lower Y (193.28–262.35 ppm, average: 226.39 ppm) than the Durango apatite (Table 3). The high S and low Y indicate...
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![Figure 4. Raman spectra from selected grains of Durango apatite.](image)

**Figure 4.** Raman spectra from selected grains of Durango apatite.

![Figure 5. Durango apatites and elemental maps obtained by micro-XRF mapping: (A) The gem-like Durango apatite crystals; (B) the distribution of Ca, Cl, Fe, P, S and Sr is shown.](image)

**Figure 5.** Durango apatites and elemental maps obtained by micro-XRF mapping: (A) The gem-like Durango apatite crystals; (B) the distribution of Ca, Cl, Fe, P, S and Sr is shown.
In this study, all measurements were conducted in the parallel c-axis to avoid the effects of orientation. The Durango apatite samples with different points selected for this study were consistent (as determined by infrared reflectance and Raman spectra). The micro-XRF mapping of calcium, phosphorus and sulphur indicate the chemical homogeneity (Figure 5). Electron microprobe analyses (Table 2; Figure 5) were performed on selected single crystals and generally showed low-to-moderate variability for major and minor compositions within each mineral grain. The analyses of the Durango apatite grains (where the c-axis was parallel) yielded SO$_3$ composition of 0.332 ± 0.012 wt.% (1σ), which suggested homogeneity in terms of sulphur distribution within a single grain. The Kovdor have SO$_3$ of 0.017 ± 0.005 wt.% (1σ). These results indicate that Durango apatite have higher SO$_3$ content, and may be more suitable than for those from Kovdor for sulphur reference materials.

4.5. Sulphur X-ray Absorption Near-Edge Structure Spectroscopy (S XANES)

The measured S$^{6+}$/ΣS ratios were consistent for five Durango apatite grains and four Kovdor apatite grains in mounts (Figure 6). Five standards with distinct peak energy are used to identify different valences of sulphur. The average XANES spectra for scapolite and gypsum contained dominant S$^{6+}$ (~2482 eV) peaks. The average XANES spectra for sodium sulfite contained dominant S$^{4+}$ (~2478 eV) peaks. The average XANES spectra for alabandite and pyrite contained dominant S$^{2-}$ (~2467 eV, 2470 eV and 2474 eV peak). The dominant peaks in lazurite were attributed to S$^{6+}$ (~2483 eV) and S$^{2-}$ (~2467 eV, 2470 eV and 2474 eV peak). The Durango apatite contained peaks at ~2482 and ~2478 eV, which represented the characteristic spectral features of S$^{6+}$ and spectral evidence for the presence of S$^{4+}$, respectively, whilst the Kovdor apatite contained a peak at ~2482 eV, which was the characteristic spectral feature of S$^{6+}$. In the studied Durango apatite, the intensity of the ~2482 eV peak of S$^{6+}$ is significantly higher than that of the ~2478 eV peak. In this study, to evaluate the effect of beam damage during sulphur XANES measurements, we also analysed all sample line transects and time series. Time series measurements on studied samples using analytical durations ranging from 10 to 30 min per analytical spot produced an insignificant variation between spectra. Therefore, irradiation damage resulting in possible oxidation or reduction of sulphur in apatite during the analyses can be ruled out.

![Figure 6. Sulphur XANES spectra of selected apatite samples from Durango, Kovdor and standard samples.](image-url)
4.6. Homogeneous Chemistry and S Oxidation States, Implications for Potential Natural Apatite Reference Materials

An in situ apatite chemical study is a powerful tool for studying magmatic evolution [16–18] and is a potential method for mineral exploration [19–21]. However, natural apatite standards with multiple chemical and S oxidation states are the precondition for analyses. In this study, all measurements were performed at the parallel apatite c-axis to avoid the effect of orientation. From the chemical measurement, the Durango apatite samples have homogeneous sulphur data. It has been documented that apatite contains S\(^{2-}\) but is dominated by S\(^{6+}\) [22]. Our study confirms that S\(^{6+}\) is dominant in the Durango and Kovdor apatites, but the Durango apatite still contains a minor amount of S\(^{4+}\) (Figure 6). The consistent nine XANES spectra with different area and time series suggest that the Durango and Kovdor apatites have homogeneous S oxidation states. This study is also consistent with the results that the higher sulphur content can be obtained from oxidised melts for which parental melts contain only S\(^{2-}\) owing to the ability of S\(^{2-}\) to partition into the apatite mineral structure [22,23]. Therefore, apatite sulphur oxidation states can be developed and applied to igneous rocks from various mineral systems. Sulphur is the most important metal complex in magmatic-hydrothermal ore deposits such as porphyry copper systems [24]. In 2017, Konecke [24] demonstrated that apatite contains different oxidation states of sulphur. Thus, sulphur XANES can be used as a function of the redox conditions of the magma. This study also indicates the substitution of S\(^{4+}\) and S\(^{6+}\) for P\(^{5+}\) in apatite. In this study, the co-variation of S content and S\(^{4+}\) and S\(^{6+}\) observed in the Durango and Kovdor apatites suggest that both of them are derived from an oxidised environment and metasomatised by SO\(^{2-}\) fluids [25–28]. Thus, apatite coexisting with a given melt in the f\(O_2\) range is anticipated to contain both S\(^{6+}\) and S\(^{4+}\), and thus, it can be used to quantify the f\(O_2\) of the system at the time of apatite crystallisation during mineralization [29]. Therefore, our study presents the improvement of the Durango and Kovdor natural apatite standards based on the S\(^{6+}/S_{\text{total}}\) of apatite and provides a new suit of data for additional analytical work.

5. Conclusions

A combination of spectroscopic, chemical and micro-XANES techniques allows an integrated chemical and S oxidation state characterisation of natural Durango and Kovdor apatite. The Durango apatite crystals show a variation of S of less than 5%. Therefore, Durango apatite can be attributed to reduced matrix effects during the analytical work. Synchrotron-based XANES reveals that Durango apatite is dominated by S\(^{6+}\) and contains a small amount of S\(^{4+}\) and Kovdor apatite is dominated by S\(^{6+}\), which indicates that it was crystallised from an oxidised environment. In addition, this study suggests that the micro-XANES technique can be used for the quantitative research of S oxidation states.

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