The Molecular Composition of Shadowed Proto-solar Disk Midplanes Beyond the Water Snowline

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Abstract

The disk midplane temperature is potentially affected by the dust traps/rings. The dust depletion beyond the water snowline will cast a shadow. In this study, we adopt a detailed gas-grain chemical reaction network, and investigate the radial gas and ice abundance distributions of dominant carbon-, oxygen-, and nitrogen-bearing molecules in disks with shadow structures beyond the water snowline around a proto-solar-like star. In shadowed disks, the dust grains at $r \sim 3–8$ au are predicted to have more than $\sim 5–10$ times the amount of ices of organic molecules such as H$_2$CO, CH$_3$OH, and NH$_2$CHO, saturated hydrocarbon ices such as CH$_4$ and C$_2$H$_6$, in addition to H$_2$O, CO, CO$_2$, NH$_3$, N$_2$, and HCN ices, compared with those in non-shadowed disks. In the shadowed regions, we find that hydrogenation (especially of CO ice) is the dominant formation mechanism of complex organic molecules. The gas-phase N/O ratios show much larger spatial variations than the gas-phase C/O ratios; thus, the N/O ratio is predicted to be a useful tracer of the shadowed region. N$_2$H$^+$ line emission is a potential tracer of the shadowed region. We conclude that a shadowed region allows for the recondensation of key volatiles onto dust grains, providing a region of chemical enrichment of ices that is much closer to the star than within a non-shadowed disk, and may explain to some degree the trapping of O$_2$ ice in dust grains that formed comet 67P/Churyumov-Gerasimenko. We discuss that, if formed in a shadowed disk, Jupiter does not need to have migrated vast distances.

Unified Astronomy Thesaurus concepts: Protoplanetary disks (1300); Planet formation (1241); Astrochemistry (75); Interstellar molecules (849); Interstellar abundances (832); Exoplanet atmospheres (487); Planetary atmospheres (1244); Planetary system formation (1257)

1. Introduction

Protoplanetary disks are composed of bare and ice-coated refractory grains, and gas, which are the ingredients of planetesimals and planets (e.g., Williams & Cieza 2011; Öberg & Bergin 2021). Disks are chemically active environments that create simple and complex molecules, including organic material. The physical and chemical conditions of protoplanetary disks determine the properties of forming planets, including mass and chemical composition (e.g., Öberg et al. 2011; Öberg & Bergin 2016, 2021; Madhusudhan et al. 2014; Pontoppidan et al. 2014; Eistrup et al. 2016, 2018; Cridland et al. 2017; Booth & Ilee 2019; Notsu et al. 2020; Schneider & Bitsch 2021; Turrini et al. 2021; Mollière et al. 2022). In addition, molecular abundances in comets and other primitive small bodies in our solar system are determined by the combination of chemical evolution in the proto-solar disk and inheritance from molecular clouds (e.g., Mumma & Charnley 2011; Caselli & Ceccarelli 2012; Walsh et al. 2014; Eistrup et al. 2016, 2018; Altwegg et al. 2019; Drozdovskaya et al. 2019; Öberg & Bergin 2021).

The disk thermal structure plays a predominant role in controlling the disk chemical structure. The midplane temperature $T(r)$ in protoplanetary disks is determined by heating due to viscous dissipation (only in the innermost region) and the stellar irradiation grazing the disk surface (e.g., Oka et al. 2011; Mori et al. 2021). Since the amount of heating per unit volume from both of these sources decreases with increasing disk radius $r$, in a smooth (non-shadowed) disk, the temperature monotonically decreases as the disk radius increases (Kusaka et al. 1970; Kenyon & Hartmann 1987; Chiang & Goldreich 1997; Oka et al. 2011). Inside the water snowline, H$_2$O ice evaporates from the dust grain surfaces into the gas phase, whereas outside it is frozen out onto the dust grain surfaces (Hayashi 1981). In addition, the cold outer disk is needed for the formation of various complex organic molecules, since the sequential hydrogenation of CO on the cold ($T(r) \sim 10–30$ K) dust grain surfaces leads to the formation of H$_2$CO and CH$_3$OH, which are key feedstock molecules that produce more complex organic molecules (e.g., Tielens & Hagen 1982; Watanabe & Kouchi 2002; Cuppen et al. 2009; Fuchs et al. 2009; Drozdovskaya et al. 2014; Furuya & Aikawa 2014; Walsh et al. 2014, 2016; Chiang et al. 2016; Bosman et al. 2018; Aikawa et al. 2020).

Under the assumption of classical monotonically decreasing disk temperature profiles, previous studies have discussed the disk radial locations of formation of planets. Here we introduce the example of Jupiter in the solar system. Recent extensive observations toward Jupiter’s atmosphere (by the Galileo probe, Cassini, and Juno spacecraft) have revealed uniform enrichment patterns of the elemental abundances from protosolar abundances by a factor of two to four, (see, e.g., Atreya...
et al. 2020; Li et al. 2020). The heavy element enrichment has been proposed to originate from planetesimals and/or pebbles dissolved in the atmospheres (e.g., Pollack et al. 1986; Iaroslavitz & Podolak 2007; Hori & Ikoma 2011; Venturini et al. 2016; Shibata & Helled 2022) and/or core erosion (e.g., Moll et al. 2017). It has been noted that the abundances of highly volatile elements, such as N and noble gas elements, are comparable to the other elemental abundances. Owen et al. (1999) suggested that this uniform enrichment originates from planetesimals formed in extremely cold ($T(r) < 30$ K) environments, where nitrogen and noble gases can freeze. Öberg & Wordsworth (2019) and Bosman et al. (2019) suggested that such uniform enrichment could be explained if the Jovian core had formed at $r > 30$ au, where the disk temperature is extremely cold ($T(r) < 30$ K, outside the N$_2$ snowline) under the classical monotonically decreasing disk temperature profile. However, according to theoretical studies on core formation and migration (Bitsch et al. 2019), the occurrence rate of the migration of a core from $r > 30$ au to 5 au is extremely low, and the core migration timescale is of the order of a few megayears even if such a migration occurs. Krijger et al. (2017, 2020) discussed that Jupiter’s core may be formed within 1 Myr to demarcate the inner and outer solar system, on the basis of the isotope analyses of meteorites.

However, the disk temperature profile may not have a monotonically decreasing distribution if a disk has substructures, as found in recent observations of radial gas and dust distributions on many disks (e.g., Isella et al. 2018; Andrews 2020; Öberg et al. 2021). Several previous studies have suggested that a shadowed region where the direct stellar light does not reach is potentially generated, depending on the inner disk density structure. A puffed inner disk rim can block the radiation from the central star and cast a shadow, so-called self-shadowing (e.g., Dullemond et al. 2001; Dullemond & Dominik 2004; Dullemond & Monnier 2010; Flock et al. 2016; Ueda et al. 2017). Ueda et al. (2019) showed that a dust pileup at the inner edge of the magnetorotational instability dead-zone casts a shadow behind it, producing cold regions of $T(r) \approx 50$ K at $r \sim 2$–7 au around Herbig Ae/Be stars. Jang-Condell & Turner (2012), Turner et al. (2012), Isella & Turner (2018), and Okuzumi et al. (2022) described that the outer wall of the gap produced by a giant planet receives extra starlight heating and puffs up, throwing a shadow across the disk beyond. Isella & Turner (2018) showed that the surface brightness contrast between the outer wall and shadow for the 1000 $M_{\odot}$ ($\sim 3.1$ $M_{\text{Jupiter}}$) planet is an order of magnitude greater than a model neglecting the temperature disturbances. In addition, Ueda et al. (2021) and Okuzumi et al. (2022) discussed that a small puffed-up rim and outer shadowed region will result in the formation of ring and gap structures in disks by thermal wave instabilities. Recently, Ohashi et al. (2022) found a steep temperature decrease outside the dust clumps at $r \sim 20$ au in the disk around Class 0/I protostar L1527 IRS. They suggest that the dust clumps create a shadowed region outside, resulting in the sudden drop in temperature.

A dust pileup at the water snowline may also cast a shadow behind it. The dust surface density just inside the water snowline can be enhanced by orders of magnitude. This is because efficient fragmentation slows the radial drift of silicate grains, leading to enhanced production of small grains within the water snowline (e.g., Birnstiel et al. 2010; Banzatti et al. 2015; Cieza et al. 2016; Pinilla et al. 2017; Müller et al. 2021). Such fragmented dust, with higher surface density and scale height, will cast a shadow behind the water snowline and provide cold environments where volatile materials can freeze (Ueda et al. 2019; Ohno & Ueda 2021).

These disk shadow structures may have a significant effect on the atmospheric composition of planets forming within. Ohno & Ueda (2021) computed the temperature structure of a proto-solar (T Tauri) disk that has a shadowed region beyond the water snowline, and investigated the radial volatile distributions. They found that the vicinity of the current orbit of Jupiter ($r \sim 5$ au) could be $T(r) < 30$ K if the small dust surface density decreases by a factor of $\geq 30$ across the water snowline. They discussed that the shadow can cause the condensation of most volatile substances, namely N$_2$ and noble gases, and that the dissolution of shadowed solids can explain the elemental abundance patterns of the Jovian atmosphere even if Jupiter formed near the current orbit. However, Ohno & Ueda (2021) included limited carbon-, nitrogen-, and oxygen-bearing molecules ($\text{H}_2\text{O}, \text{CO}, \text{CO}_2, \text{C}_2\text{H}_6, \text{N}_2$, and NH$_3$ only; see also Öberg & Wordsworth 2019). In addition, they fixed the total (gas+ice) abundances of each volatile and calculated the balance between thermal desorption and freeze-out onto dust grains within each molecular species only. In order to investigate the effects of the shadow beyond the water snowline on disk chemical evolution including $\text{H}_2\text{CO}$, $\text{CH}_3\text{OH}$, and more complex organic molecules (such as, e.g., $\text{NH}_2\text{CHO}$ and HCOOCH$_3$), more detailed gas-grain chemical modeling is needed (e.g., Walsh et al. 2014, 2015; Eistrup et al. 2016, 2018; Bosman et al. 2018; Notsu et al. 2021).

In this study, we calculate the chemical structure of a shadowed disk midplane around a T Tauri star (a proto-solar-like star), using a detailed gas-grain chemical reaction network. We investigate the radial abundance distributions of dominant carbon-, oxygen-, and nitrogen-bearing molecules and the radial distributions of elemental abundance ratios (C/O and N/O ratios) in the gas and ice of disks with shadow structures. We also investigate the dependence of the disk chemical structures on ionization rates and initial abundances. We discuss the effects of disk shadowing on chemical evolution of complex organic molecules and forming planetary atmospheres. We also discuss the implications of our results for the chemical composition of comets and asteroids in the solar system. The outline of our model calculations are explained in Section 2. The results and discussion of our calculations are described in Sections 3 and 4, respectively. The conclusions are presented in Section 5.

2. Methods

2.1. The Physical Model of the Protoplanetary Disk Midplane

We adopt the physical model of a steady, axisymmetric Keplerian disk around a T Tauri star (a proto-solar-like star) with mass $M_* = 1.0$ $M_{\odot}$, radius $R_* = 2.6$ $R_{\odot}$, and effective temperature $T_* = 4300$K. For the disk density structure, we adopted the same parameterized disk model as Ohno & Ueda (2021) in which the dust surface density steeply varies around the water snowline. The radial surface density profiles are described by the following equations:

$$\Sigma_{\text{gas}}(r) = 670 \left( \frac{r}{1 \text{ au}} \right)^{-3/5} \text{ g cm}^{-2},$$  \hspace{1cm} (1)
\[
\Sigma_{\text{dust}}(r) = \begin{cases} 
0.01\Sigma_{\text{gas}}(r) & (r < R_{\text{SL}}(\text{H}_2\text{O})) \\
0.01f_\Sigma_{\text{gas}}(r) & (r \geq R_{\text{SL}}(\text{H}_2\text{O}))
\end{cases},
\]

where \( r \) is the disk radius from the central star, \( \Sigma_{\text{gas}} \) is the gas surface density, \( \Sigma_{\text{dust}} \) is the surface density of the small dust (0.1–100 \( \mu \)m) that contributes to the dust opacity, and \( R_{\text{SL}}(\text{H}_2\text{O}) = 1.3 \) au is the assumed radial position of the water snowline. In protoplanetary disks, the maximum grain size is expected to be larger than that of the typical ISM dust due to significant dust growth. However, large grains (\( \gtrsim 100 \mu \)m) have little impact on the disk temperature structure because they have small opacity and are depleted at the disk surface where stellar irradiation is absorbed. The power index of the radial gas surface density profile (\( \sim 3/5 \)) is calculated under the assumption of a disk with steady accretion and viscous heating (e.g., Nakamoto & Nakagawa 1994; Oka et al. 2011). We adopted various values of the parameter \( f = 1.0, 0.3, 0.03, \) and 0.003) to investigate the effects of the magnitude of shadowing on the disk thermal and chemical structures. The actual values of the parameter \( f \) depend on the efficiency of fragmentation of silicate and icy grains, which is controlled by the turbulence viscosity strength and stickiness of the dust grains (e.g., Birnstiel et al. 2010; Banzatti et al. 2015; Pinilla et al. 2017). Banzatti et al. (2015) reported dust surface density variations of \( f \sim 0.1, 0.001, \) and 0.01 around the water snowline for turbulence viscosity strengths of \( \alpha_t = 10^{-2}, 10^{-3}, \) and \( 10^{-4} \), respectively. The disk midplane temperature profiles were extracted from the results of the 2D thermal model calculation in Ohno & Ueda (2021; see also Section 2 in Ueda et al. 2019). In this calculation, the above radial density profiles (see Equations (1) and (2)) are adopted, and the Monte Carlo radiative-transfer code RADMC-3D (Dullemond et al. 2012) is used. Scattering is assumed to be isotropic because full-scattering treatment is computationally expensive. We expect that the temperature structure is not significantly modified even if we consider the full scattering because the temperature at the shadowed region seems to be insensitive to the presence of scattering (Ueda et al. 2019). The vertical dust density distribution and the temperature structure are mutually dependent; thus, the radiative-transfer calculations were iteratively performed to obtain a self-consistent disk structure (for more details, see Ueda et al. 2019). The calculations also include the internal radiative flux produced by viscous accretion, \( q_{\text{acc}} \), determined by the following equation (Lynden-Bell & Pringle 1974; Pringle 1981; Nomura & Millar 2005):

\[
q_{\text{acc}} = \frac{9}{4} \alpha_{t}\rho_{\text{gas}}c_s^2\Omega_K,
\]

where \( \alpha_t \) is set to be \( 3 \times 10^{-4} \), \( \rho_{\text{gas}} \) is the gas mass density, \( c_s \) is the sound speed, and \( \Omega_K \) is the Keplerian frequency. The surface density coefficient \( \Sigma_{\text{dust}} = 670 \text{ g cm}^{-2} \) and \( \alpha_t \) in Equation (1) are adjusted so that the water snowline position is outside the current Earth orbit (\( = 1.3 \) au) in the calculation of the temperature profiles. We note that we have adopted the same dust surface density within the water snowline in all models so that the position of the water snowline is fixed. This is because in this study, we focus on the effects of shadowing on disk chemical structures beyond the water snowline. Eistrup et al. (2016) adopted a lower surface density model \( (\Sigma_{\text{gas}}(5.2 \text{ au}) = 16 \text{ g cm}^{-2}) \) than our adopted model \( (\Sigma_{\text{gas}}(5.2 \text{ au}) = 249 \text{ g cm}^{-2}) \), and the water snowline position in their model is around \( 0.7 \) au. We assume the same dust and gas temperatures \( (T_{\text{dust}} = T_{\text{gas}} = T) \), which is a valid assumption in the dense molecular regions of the disk. The radial profile of the midplane gas mass density \( \rho_{0,\text{gas}}(r) \) is given by the following equation:

\[
\rho_{0,\text{gas}}(r) = \frac{\Sigma_{\text{gas}}(r)}{\sqrt{2\pi}h_{\text{gas}}},
\]

where \( h_{\text{gas}} = c_s/\Omega_K \) is the gas scale height. We adopted the DSHARP dust opacity model (Birnstiel et al. 2018; see also, e.g., Henning & Stognienko 1996; Draine 2003; Warren & Brandt 2008), and assumed a dust-sized distribution that follows a power law with an index of \( -3.5 \), similar to the Mathis, Rumpl, & Nordsieck (MRN) distribution (Mathis et al. 1977), with minimum and maximum dust radii of 0.1 and 1.0 \( \mu \)m, respectively. We assume that larger grains have been removed from the disk atmosphere by settling and radial drift. The bulk densities of dust grains \( \rho_{\text{internal}} \) are set to be \( 3.0 \text{ g cm}^{-3} \) for the region where the icy component is evaporated (\( T > 160 \text{ K} \)) and 1.4 \( \text{ g cm}^{-3} \) elsewhere. We note that icy grains can efficiently coagulate into larger (\( \gtrsim 1 \text{ mm} \)) dust particles and centimeter-sized pebbles outside the water snowline (e.g., Ros & Johansen 2013; Sato et al. 2016; Drążkowska & Albright 2017; Pinilla et al. 2017), but such large dust particles and pebbles have negligible contribution to the opacity compared with that of smaller dust particles (Miyake & Nakagawa 1993), and they are confined to the disk midplane. Thus, \( \Sigma_{\text{dust}} \) is not necessarily the same as the total surface density of all solid components (dust grains and pebbles), for example, outside the water snowline (Banzatti et al. 2015). We note that in our disk chemical modeling (see Section 2.2), we use the dust density profile in the disk midplane calculated from this small dust (0.1–100 \( \mu \)m) surface density profile \( \Sigma_{\text{dust}} \).

The left panel of Figure 1 shows \( \Sigma_{\text{gas}}(r) \) and \( \Sigma_{\text{dust}}(r) \) for different values of the parameter \( f = 1.0, 0.3, 0.03, \) and 0.003). The right panel of Figure 1 shows the radial profiles of the gas number density \( n_{\text{gas}}(r) \) and temperature \( T(r) \) in the disk midplane for different values of the parameter \( f \). \( n_{\text{gas}}(r) \) is calculated by dividing \( \rho_{0,\text{gas}}(r) \) by the mean particle mass \( (= 2.33 \text{ amu}) \). The values of \( n_{\text{gas}} \) are largest in the innermost region \( (\gtrsim 10^{14} \text{ cm}^{-2} \text{ at } r \lesssim 1.4 \text{ au}) \), and decrease as the value of \( r \) increases \( (\lesssim 10^{12} \text{ cm}^{-2} \text{ at } r \gtrsim 17 \text{ au}) \). The value of \( T \) within the water snowline is independent of \( f \) and is primarily controlled by viscous heating, yielding the radial dependence of \( \propto r^{-9/10} \) (Oka et al. 2011). As demonstrated in Ohno & Ueda (2021), the depletion of small dust grains outside the water snowline casts a shadow, changing the disk midplane temperature significantly. In the cases of smaller values of \( f \), the decrease of dust surface density outside the water snowline results in reduced dust opacity, which reduces the amount of stellar radiation received per unit volume and cools the outer region compared with the case of the non-shadowed disk \( (f = 1.0) \). The values of \( T \) at \( r \sim 1.3–17 \text{ au} \) decrease with decreasing \( f \), as the shadow

\footnote{\( \alpha_t \) is a dimensionless parameter to quantify the disk viscosity (Shakura & Sunyaev 1973).}
extends farther out. Even in the case of a small variation in the dust density profile with \( f = 0.3 \), \( T \) is around 20 K lower than that in the non-shadowed disk with \( f = 1.0 \). For \( f > 0.03 \), \( T \) is <30 K at \( r \sim 3–8 \) au, and <25 K at the current orbit of Jupiter (~5 au). In contrast, the values of \( T \) at \( r > 20 \) au increase with decreasing \( f \), although the degree of increase is small and \( T \) is <35 K even for \( f = 0.003 \). We interpret that reduced dust surface density beyond the water snowline allows the reprocessed stellar radiation to enter the outer disk midplane from the upper disk surface (Cleeves 2016). At the outer disk midplane (\( r > 20 \) au), the heating effect by such radiation overcomes the cooling effect, which reduces as \( r \) increases. The disk temperature is expected to be higher in the outer disk (beyond \( r > 20 \) au) if we assume lower surface densities than we initially chose. The values of disk masses are determined by the density profile in the outer region of the disk. We note that we have adopted the surface densities following a single power-law profile for the sake of simplicity (see Equations (1) and 2). The disk gas mass within 20 au and 40 au is \( \sim 2.2 \times 10^{-3} M_{\odot} \) and \( \sim 5.8 \times 10^{-3} M_{\odot} \), respectively. Eistrup et al. (2016) adopted a disk surface density profile with an outer exponential cutoff radius of \( r_c = 20 \) au. If we include such an outer exponential cutoff with \( r_c \sim 20 \) au in the disk surface density profile, the total disk mass within 40 au will be around half of the value without cutoff. Recently, Trapani et al. (2022) estimated from HD and C\(^{18}\)O line observations that the disk masses for some typical Class II T Tauri disks (such as TW Hya) are around a few \( \times 10^{-2} M_{\odot} \).

We assumed that the water snowline position as 1.3 au, but it might be outside the current orbit of Mars (1.52 au). Hansen (2009) discussed that Mars forms from a protoplanet that has moved by scattering from the current orbit of Earth.

In our calculations, the far-UV (FUV) and X-ray radiation from both the central T Tauri star and interstellar radiation fields are neglected, since the disk has a high surface density (\( \Sigma_{\text{gas}}(r) \sim 10^4 - 10^5 \) g cm\(^{-2}\)) and the disk midplane is effectively shielded (e.g., Nomura & Millar 2005; Nomura et al. 2007; Walsh et al. 2015). Cleeves et al. (2014b), Eistrup et al. (2016, 2018), and Schwarz et al. (2018, 2019) discussed that disk chemical evolution (such as chemical processing of CO) is affected by the degree of (cosmic-ray) ionization. In order to investigate the effects of ionization on chemistry in the shadowed disk, we adopt two values of radially constant ionization rates \( \xi_{\text{CR}}(r) = 1.0 \times 10^{-17} \text{ s}^{-1} \) and \( 1.0 \times 10^{-18} \text{ s}^{-1} \) (e.g., Cleeves et al. 2014b; Eistrup et al. 2016, 2018). The former high ionization value includes contributions from the decay products of short-lived radionuclides (SLRs; Cleeves et al. 2013a, 2013b) and from cosmic-rays, originating externally to the disk (e.g., Umebayashi & Nakano 2009; Padovani et al. 2018), whereas the latter low ionization value considers contribution from the decay products of SLRs only. We note that, in reality, the ionization rate has radial dependence, since the contribution from SLRs decreases as \( r \) increases (Cleeves et al. 2013b) and that from cosmic-rays increases as \( r \) increases (e.g., Umebayashi & Nakano 2009; Padovani et al. 2018; Aikawa et al. 2021; Seifert et al. 2021; Fujii & Kimura 2022).

### 2.2. Calculation of the Disk Chemical Structure

We calculated the chemical evolution of the shadowed disk midplane around a T Tauri star (a proto-solar-like star) using a detailed gas-grain chemical reaction network. The chemical reaction network adopted in this work is basically similar to that in Notsu et al. (2021). The detailed background theories and procedures are also discussed in our previous works (e.g., Walsh et al. 2010, 2012, 2014, 2015, Heinzeller et al. 2011; Eistrup et al. 2016, 2018; Eistrup & Walsh 2019, Notsu et al. 2016, 2017, 2018, 2019), although there are some differences between the models in these studies and our adopted model, which we explain in this Section. Our model calculates the time-dependent chemical evolution at the disk midplane of each radial distance and does not include physical mass transport in the radial direction by viscous accretion and in the vertical direction by diffusive turbulent mixing and disk winds (Heinzeller et al. 2011).
2.2.1. Gas-phase Reactions

Our gas-phase chemistry includes the complete network from the release of the UMIST Database for Astrochemistry, called RATE12, and is publicly available8 (McElroy et al. 2013). RATE12 includes gas-phase two-body reactions, photoionization and photodissociation, direct cosmic-ray ionization, and cosmic-ray-induced photodissociation and photodissociation. Since the FUV radiation fields from the central T Tauri star and interstellar FUV fields are neglected in our calculations (see also Section 2.1), the photodissociation and photoionization by FUV radiation are effectively zero. In the cosmic-ray-induced photoreactions, UV photons are generated internally via the interaction of secondary electrons produced by cosmic-rays with H2 molecules (Greder et al. 1987, 1989).

As in Walsh et al. (2015), we also added a set of three-body reactions and “hot” H2 chemistry, although they are not expected to be important around the water snowline (∼100–200 K). Moreover, the gas-phase chemical network is supplemented with reactions for important species, for example the CH3O radical, which are not included in RATE12. The gas-phase formation and destruction reactions for these species are from the Ohio State University (OSU) network (Garrod et al. 2008).

2.2.2. Gas-grain Interactions

As in Notsu et al. (2021), we include the freeze-out of gas-phase molecules on dust grains, and the thermal and nonthermal desorption of molecules from dust grains (Hasegawa et al. 1992; Walsh et al. 2010, 2012; Notsu et al. 2016). As nonthermal desorption mechanisms, cosmic-ray-induced photodissociation, reactive desorption (see Section 2.2.3), and direct cosmic-ray-induced (thermal) desorption (Leger et al. 1985; Hasegawa & Herbst 1993; Hollenbach et al. 2009) are adopted. We note that the direct cosmic-ray-induced (thermal) desorption has no significant impact on chemistry, since its reaction timescale is typically much longer (>10⁷ yr) than the age of a Class II disk (Hollenbach et al. 2009). We adopt the value for the integrated cosmic-ray-induced UV photon flux as 10⁹ photons cm⁻² s⁻¹ (Prasad & Tarafdar 1983; Walsh et al. 2014). We scale the internal UV photon flux by the cosmic-ray ionization rate. On the basis of the adopted dust-sized distribution (see Section 2.1), we assume compact spherical grains with an average radius <0.1 μm. We note that the total dust surface area is mainly dominated by the smallest dust grains, under the MRN distribution. The values of photodesorption yields adopted in this work, Y_{des}(j), are the same as those adopted in Notsu et al. (2021). We use experimentally determined photodesorption yields, where available (e.g., Öberg et al. 2007, 2009a, 2009b; Fillion et al. 2014; Bertin et al. 2016; Cuppen et al. 2017, see Table 1 of Notsu et al. 2021). For all species without experimentally determined photodesorption yields, a value of 10⁻³ molecules photon⁻¹ is used. As in Notsu et al. (2021), we include the fragmentation pathways for photodesorption of water and methanol molecules (e.g., Öberg et al. 2009b; Arasa et al. 2010, 2015; Bertin et al. 2016; Cruz-Díaz et al. 2016).

The sticking coefficient is assumed to be 1 for all species, except for H, which leads to H₂ formation (for more details, see Appendix B.2 of Bosman et al. 2018). We adopt the same values of the binding (desorption) energies E_{des}(j) for all molecules as used in Notsu et al. (2021). Table 1, E_{des}(j) for several dominant molecules and their estimated snowline positions R_{SL}(j) in the non-shadowed disk (f = 1.0) are listed (see also Section 3). We defined the snowline positions as the radii where the gas and ice abundances of each molecule are the same because of the balance between thermal desorption and freeze-out. We note that the binding energies of molecules depend on the chemical compositions and physical structures (e.g., crystal or amorphous) of the ice mantles on the dust grains (e.g., Cuppen et al. 2017; Penteado et al. 2017; Kouchi et al. 2021, see also Section 4.2). For CO and N₂, we assume the values of pure CO and N₂ ices, respectively (Öberg et al. 2005).

2.2.3. Grain-surface Reactions

For the grain-surface reactions, we adopt the same reaction network as Notsu et al. (2021) used. Those reactions are mostly based on the OSU network (Garrod et al. 2008), with including some extended network for some molecules (such as CH₃OH and NH₂CHO; see Section 2.2.3 of Notsu et al. 2021 and, e.g., Noble et al. 2015; Chuang et al. 2016). In addition to grain-surface two-body reactions (Hasegawa et al. 1992) and reactive desorption, grain-surface cosmic-ray-induced photodissociation is also included in our calculations (Garrod et al. 2008; Walsh et al. 2014, 2015). Only the top two monolayers of the ice mantle are chemically active. We assume that the size of the barrier to surface diffusion is 0.3 × E_{des}(j) (Walsh et al. 2015). For the lightest reactants, H and H₂, we adopt either the classical diffusion rate or the quantum tunneling rate depending on which is fastest (Hasegawa et al. 1992; Bosman et al. 2018). For the quantum tunneling rates, we adopt a rectangular barrier of width 1.0 Å (Hasegawa et al. 1992; Bosman et al. 2018).

2.2.4. Initial Abundances

As Eistrup et al. (2016, 2018) conducted, the disk chemical evolution is calculated with two different sets of initial abundances: molecular species and atomic species. All abundances in our calculations of this paper are with respect to the total number density of H nuclei. For both sets of initial abundances (molecular and atomic), the elemental ratios are consistent. The choice of these initial abundances is motivated

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8 http://udfa.ajmarkwick.net
by the following two extreme scenarios about the history of the disk midplane material (Öberg & Bergin 2021; van Dishoeck et al. 2021).

The use of molecular initial abundances assumes that the disk material is wholly inherited from the molecular cloud and the pre-stellar core from which the central star formed, implying a more quiescent mode of disk formation. The enhanced deuterium fractionation of water and organic molecules in disks is considered to be a probe of pre-stellar inheritance, since efficient deuteration of water and organic molecular ices is only possible in cold and embedded pre-stellar cloud cores where UV radiation is almost completely attenuated and CO is frozen out onto dust grains (e.g., Clevees et al. 2014a; Furuya et al. 2016, 2017; Drozdovskaya et al. 2019; Jensen et al. 2021). Drozdovskaya et al. (2019) described that the relative abundances of volatile organic molecules (with respect to, e.g., methanol) correlate, with some scatter, between Atacama Large Millimeter/submillimeter Array (ALMA) data toward the protostar IRAS 16293-2422 B and Rosetta’s in situ monitoring data of comet 67P/Churyumov-Gerasimenko. This result implies that the volatile composition of solar system comets may be partially inherited from the pre-stellar and protostellar phases. In contrast, the use of atomic initial abundances implies that full chemical reset has occurred during disk formation. Throughout disk formation, infalling material is subject to radiation from the central star and accretion shocks, which may alter the chemical composition of material (e.g., Visser et al. 2009; Drozdovskaya et al. 2016; Miura et al. 2017; Notsu et al. 2021). Öberg & Bergin (2021) and van Dishoeck et al. (2021) discussed that neither a complete inheritance nor a complete reset scenario can explain the results of recent disk observations (by, e.g., ALMA) and the full solar system record. They described that both must have played an important role during disk formation.

As Walsh et al. (2015) and Notsu et al. (2021) adopted, the values of volatile elemental abundances for He, O, C, N, and S are, respectively, $9.75 \times 10^{-2}$, $3.20 \times 10^{-4}$, $1.40 \times 10^{-4}$, $7.50 \times 10^{-5}$, and $8.00 \times 10^{-8}$ relative to total hydrogen nuclei density. For other elements, we use the low-metallicity elemental abundances from Graedel et al. (1982). In the reset scenario, we use the above elemental abundances as initial abundances of our disk chemical calculations, except hydrogen. We assume that the abundances for H$_2$ and H are, respectively, $5.0 \times 10^{-1}$ and $5.0 \times 10^{-5}$.

Table 2 shows the gas and ice fractional molecular abundances with respect to total hydrogen nuclei density, which are used as molecular initial abundances in the inheritance scenario. In this Table 2, we include dominant species and molecules that are important in disk chemical evolution. We adopt the same molecular abundances as Notsu et al. (2021) used for the initial abundances in their chemical modeling. Previous chemical calculations (e.g., Walsh et al. 2015; Eistrup et al. 2016, 2018; Drozdovskaya et al. 2016) adopted a similar water-rich molecular abundances as initial conditions.

### 3. Results

In Sections 3.1–3.2, we investigate the results for our standard model (assuming molecular initial abundances and $\zeta_{\text{CH}\_4}\_r = 1.0 \times 10^{-17}$ [s$^{-1}$]), and describe the effects of a disk shadow on the radial molecular abundance distributions. In Appendix A, we also study the disk chemical structures for different disk ionization rates and for atomic initial abundances (see also Section 3.3). In addition, in Sections 3.1–3.2 we show the chemical abundance distributions at $t = 10^6$ yr, which is the typical age of Class II disks and consistent with previous studies (e.g., Eistrup et al. 2016). In Appendix B, we also investigate the time evolution of radial molecular abundance distributions (see also Section 3.4).

#### 3.1. Dominant Carbon, Oxygen, and Nitrogen Carriers around the Current Orbit of Jupiter

In this Section, we show the composition of the dominant carbon, oxygen, and nitrogen carriers around the current orbit of Jupiter. Figure 2 shows pie charts of the percentage contributions of the dominant carbon-, oxygen-, and nitrogen-bearing molecules to the total elemental abundances ($=1.4 \times 10^{-4}$ for carbon, $=3.2 \times 10^{-4}$ for oxygen, $=7.5 \times 10^{-5}$ for nitrogen) at...
Figure 2. Pie charts of the percentage contributions of the dominant carbon-bearing molecules to the total elemental carbon abundance ($=1.4 \times 10^{-4}$, top panels), the dominant oxygen-bearing molecules to the total elemental oxygen abundance ($=3.2 \times 10^{-4}$, middle panels), and the dominant nitrogen-bearing molecules to the total elemental nitrogen abundance ($=7.5 \times 10^{-5}$, bottom panels), at $r = 5.3$ au (around the current orbit of Jupiter) and $t = 10^6$ yr. The left panels show the contributions for the disk midplane with the monotonically decreasing density and temperature profile ($f = 1.0$), whereas the right panel shows the contributions for the shadowed disk midplane ($f = 0.03$). These panels show the results for molecular initial abundances (the “inheritance” scenario) and $\xi_{CR}(r) = 1.0 \times 10^{-17}$ [s$^{-1}$]. The filled and hatched slices are, respectively, the contributions of gaseous and icy molecules. (Top panels): the dark blue, red, black, purple, green, orange, yellow, and light blue slices are, respectively, the contributions of CO$_2$ ice, CO ice, CO gas, H$_2$CO ice, CH$_4$ ice, HCN ice, C$_2$H$_6$ ice, and other molecules (such as NH$_2$CHO ice and CH$_3$OH ice for $f = 0.03$, and CH$_3$CCH ice, CH$_3$CH$_2$ ice, and CH$_3$CH$_3$H$_2$ ice for $f = 1.0$). (Middle panels): the dark blue, red, black, purple, green, and light blue slices are, respectively, the contributions of CO$_2$ ice, CO ice, CO gas, H$_2$CO ice, H$_2$O ice, and other molecules (such as NH$_2$CHO ice and CH$_3$OH ice). (Bottom panels): the dark blue, red, black, purple, green, orange, and light blue slices are, respectively, the contributions of N$_2$ gas, N$_2$ ice, NH$_3$ ice, NH$_2$CHO ice, NH$_2$OH ice, HCN ice, and other molecules (such as HC$_3$N ice).
$r = 5.3$ au (around the current orbit of Jupiter) and $t = 10^6$ yr, in the non-shadowed disk ($f = 1.0$) and a shadowed disk ($f = 0.03$). These results are from the model in which we assume molecular initial abundances and a high ionization rate (see Figures 3–5 in Section 3.2).

In the non-shadowed disk, the dominant carbon carriers at $r = 5.3$ au are CO$_2$ ice ($\sim 36\%$) and CO gas ($\sim 40\%$). In the shadowed disk, CO freezes out onto the dust grain surface, and the amounts of CO$_2$ ice ($\sim 30\%$), CO ice ($\sim 25\%$), and ices of unsaturated hydrocarbon molecules (such as C$_3$H$_4$) decrease, whereas those of H$_2$CO ice ($\sim 19\%$), CH$_4$ ice ($\sim 6\%$), C$_2$H$_6$ ice ($\sim 5\%$), HCN ice ($\sim 4\%$), NH$_2$CHO ice ($\sim 4\%$), and CH$_3$OH ice ($\sim 3\%$) increase. The enhanced abundances of these organic molecules in the shadowed disks are owing to the reaction pathways in which the sequential hydrogenation of CO on the dust grain surfaces and cosmic-ray-induced photodissociation of CH$_3$OH are the key reactions, as we explain in Sections 3.2.2 and 3.2.3.

In the non-shadowed disk, the dominant oxygen carriers at $r = 5.3$ au are H$_2$O ice ($\sim 47\%$), CO$_2$ ice ($\sim 31\%$) and CO gas ($\sim 18\%$). In the shadowed disk, the amounts of CO$_2$ ice ($\sim 26\%$) and CO ice ($\sim 11\%$) decrease, whereas that of H$_2$O ice is similar ($\sim 50\%$), and that of H$_2$CO ice ($\sim 8\%$) increases.

In the non-shadowed disk, the dominant nitrogen carriers at $r = 5.3$ au are N$_2$ gas ($\sim 76\%$) and NH$_3$ ice ($\sim 17\%$). In the shadowed disk, N$_2$ mostly freezes out onto the dust grain surface, and the amount of N$_2$ gas ($\sim 5\%$) decreases, whereas that of NH$_3$ ice is almost similar ($\sim 20\%$), and those of N$_2$ ice ($\sim 51\%$) and other nitrogen-bearing molecules such as NH$_2$OH, NH$_2$CHO, and HCN ices ($\sim 7\%$–$\sim 8\%$ each) increase. The enhanced abundances of NH$_2$OH, NH$_2$CHO, and HCN in the shadowed disks are due to hydrogenation reactions on the dust grain surfaces (see Section 3.2.3).

According to these results, in shadowed disks the dust grains at around the current orbit of Jupiter ($r \sim 5.3$ au) are expected to have significant amounts of saturated hydrocarbon ices such as CH$_4$ and C$_2$H$_6$, ices of organic molecules such as H$_2$CO, NH$_2$CHO, and CH$_3$OH, in addition to H$_2$O, CO, CO$_2$, NH$_3$, N$_2$, HCN, and NH$_2$OH ices, compared with those in the non-shadowed disks (mostly ices of H$_2$O, CO$_2$, NH$_3$, and unsaturated hydrocarbon molecules). These abundant saturated hydrocarbons and various organic molecules such as H$_2$CO, NH$_2$CHO, and CH$_3$OH were not reported by Ohno & Ueda (2021), since they approximated various organic molecules by considering C$_2$H$_6$ alone. In addition, our results provide intriguing implications for the compositions of small bodies (such as comets and asteroids) formed at around the current orbit of Jupiter, as we discuss in Section 4.2.
Figure 4. Same as Figure 3, but for CH$_4$ ($n_{\text{CH}_4}/n_{\text{H}}$, top-left panel), C$_2$H$_6$ ($n_{\text{C}_2\text{H}_6}/n_{\text{H}}$, top-right panel), H$_2$CO ($n_{\text{H}_2\text{CO}}/n_{\text{H}}$, bottom-left panel), and CH$_3$OH ($n_{\text{CH}_3\text{OH}}/n_{\text{H}}$, bottom-right panel).

Figure 5. Same as Figure 3, but for N$_2$ ($n_{\text{N}_2}/n_{\text{H}}$, top-left panel), NH$_3$ ($n_{\text{NH}_3}/n_{\text{H}}$, top-right panel), HCN ($n_{\text{HCN}}/n_{\text{H}}$, bottom-left panel), and NH$_2$CHO ($n_{\text{NH}_2\text{CHO}}/n_{\text{H}}$, bottom-right panel).
3.2. Radial Molecular Abundance Distributions for Our Standard Model

In this Section, we describe the effects of disk shadowing on the radial molecular abundance distributions. The radial profiles of fractional abundances with respect to total hydrogen nuclei densities \( n_X/n_H \) at \( t = 10^6 \) yr for dominant carbon-, oxygen-, and nitrogen-bearing molecules are shown in Figures 3–5 in Sections 3.2.1–3.2.3, with separate panels for each molecule, for different values of the parameter \( f (= 1.0, 0.3, 0.03, \text{ and } 0.003) \), respectively. In Figures 6–8 in Section 3.2.4, the radial fractional abundance distributions at \( t = 10^6 \) yr of other major carbon-, nitrogen-, and oxygen-bearing molecules, mainly complex organic molecules, are shown. In Figure 9 in Section 3.2.5, the radial profiles of fractional abundances with respect to total hydrogen nuclei densities at \( t = 10^6 \) yr for sums of larger (complex) organic molecules are shown. We show this to demonstrate the efficiency of conversion from more simple to more complex...
molecules for different dust depletion factors. The panels in these Figures 3–9 show the results for the radially constant cosmic-ray ionization rate of $\zeta_{\text{CR}}(r) = 1.0 \times 10^{-17}$ [s$^{-1}$] with assuming molecular initial abundances (the “inheritance” scenario).

Since previous studies (e.g., Eistrup et al. 2016, 2018) have discussed the chemical structure of the non-shadowed disk ($f = 1.0$) in detail, in this paper we mainly focus on the effects of shadowing on disk chemical structures. We note that the results of our calculations for the non-shadowed disk are largely consistent with those in Eistrup et al. (2016, 2018) with some exceptions, such as HCN (see Section 3.2.3 of this paper).

### 3.2.1. H$_2$O, CO, CO$_2$, and O$_2$

Figure 3 shows the radial profiles of fractional abundances for dominant oxygen-bearing molecules: H$_2$O (water), CO, CO$_2$, and O$_2$. The Astrophysical Journal, 936:188 (39pp), 2022 September 10 Notsu et al.
CO (carbon monoxide), CO$_2$ (carbon dioxide), and O$_2$ (molecular oxygen). The water snowline position (∼1.3 au, $T(r) \sim 140$ K) and the H$_2$O gas abundances (∼2 × 10$^{-8}$) within the water snowline do not change for various values of $f$, since the disk midplane temperature is significantly changed only beyond the water snowline. We note that the gas abundances of other molecules are also unchanged within the water snowline.

Outside the water snowline, H atoms are supplied onto the dust grain surface from the gas or produced in situ within the grain mantles. We note that in the gas phase, H atoms are mainly produced by cosmic-ray-induced photodissociation of H$_2$. Since the fractional abundances of H$_2$ with respect to total H nuclei (∼5 × 10$^{-1}$) are much larger than those of other molecules, the gas-phase production is dominant. Thus, if we conduct disk chemical modeling with atomic initial abundances

---

**Figure 8.** Same as Figure 3, but for CH$_3$CHO ($n_{\text{CH}_3\text{CHO}}/n_{\text{H}}$, top-left panel), CH$_3$OCH$_3$ ($n_{\text{CH}_3\text{OCH}_3}/n_{\text{H}}$, top-right panel), HCOOH ($n_{\text{HCOOH}}/n_{\text{H}}$, middle-left panel), HCOOCH$_3$ ($n_{\text{HCOOCH}_3}/n_{\text{H}}$, middle-right panel), CH$_3$COOH ($n_{\text{CH}_3\text{COOH}}/n_{\text{H}}$, bottom-left panel), and C$_2$H$_5$OH ($n_{\text{C}_2\text{H}_5\text{OH}}/n_{\text{H}}$, bottom-right panel).
unsaturated hydrocarbon molecules for which we show the results in Figure 6.

\[
\text{OH}_{\text{ice}} + \text{H}_{\text{ice}} \rightarrow \text{H}_2\text{O}_{\text{ice}}.
\]

(5)

However, because we have adopted abundant initial water ice abundances (=1.984 × 10^−4, see Table 1), further water ice formation in the outer disk midplane does not proceed efficiently. Thus, the H_2O ice abundances beyond the water snowline are relatively constant (~(1−2) × 10^−4) for various values of f. We note that the H_2O ice abundances increase from ~1.5 × 10^−4 to ~2 × 10^−4 in the coldest regions with T(r) < 20 K (at r ≈ 3−8 au for f = 0.003 and at r ≥ 25 au for f = 1.0), where the formation of CO_2 ice does not proceed efficiently (see below).

In the non-shadowed disk (f = 1.0), the CO snowline position is r ~ 22 au (T(r) ~ 25 K), and CO gas abundances inside the water snowline are 10^−6. In addition, in the non-shadowed disk, CO gas abundances between the H_2O and CO snowlines are ~(4−8) × 10^−5.

As introduced later, the chemical conversion of CO plays vital roles in producing other carbon-bearing molecules. CO molecules accreting onto the dust grain surfaces can react with OH radicals produced on the ice by cosmic-ray-induced photodissociation of H_2O ice (e.g., Drozdovskaya et al. 2016; Eistrup et al. 2016, 2018; Schwarz et al. 2018, 2019). This produces CO_2 ice via the following grain-surface reaction:

\[
\text{CO}_{\text{ice}} + \text{OH}_{\text{ice}} \rightarrow \text{CO}_2\text{ice} + \text{H}_{\text{ice}}.
\]

(6)

In addition, for the ISM level ionization rate (ζ_cr(r) = 10^−17 [s^−1]), the following destruction pathway of gas-phase CO by He^+ is also efficient:

\[
\text{CO} + \text{He}^+ \rightarrow \text{C}^+ + \text{O} + \text{He},
\]

(7)

which leads to the formation of CH_4, C_2H_6, and other hydrocarbons (e.g., Aikawa et al. 1999; Furuya & Aikawa 2014; Eistrup et al. 2016, 2018; Yu et al. 2016; Bosman et al. 2018, see also Sections 3.2.2 and 3.2.4). We note that He^+ is produced by the direct cosmic-ray ionization of He. Moreover, the sequential hydrogenation of CO on the dust grain surfaces leads to the formation of H_2CO and CH_3OH ices (e.g., Watanabe & Kouchi 2002; Drozdovskaya et al. 2014; Bosman et al. 2018, see also Section 3.2.2 and Appendix A.2). Several studies have suggested that the above chemical processes (partly) explain the depletion of CO in the Class II disks reported by recent observations with ALMA (e.g., Nomura et al. 2016, 2021; Schwarz et al. 2016; Krijt et al. 2018, 2020; Bergner et al. 2020; Zhang et al. 2020, 2021).

In the shadowed disk (f ≤ 0.03), CO freezes out onto dust grains at around the current orbit of Jupiter (r ~ 3−8 au, T(r) ≤ 25 K), and CO ice abundances are ~ (3−6) × 10^−5. In addition, CO returns to the gas phase at r > 8 au. At r > 20 au, CO (re-)freezes out onto dust grains, and the adsorption front moves outward with decreasing f, since the values of T(r) at r > 20 au increase (≤35 K at the maximum) with decreasing f (see also Section 2.1). This CO abundance profile is consistent with the results reported in Ohno & Ueda (2021). We suggest that reducing the dust surface density beyond the water snowline allows the reprocessed stellar radiation to enter the outer disk midplane from the disk upper layers, and that the penetration of such radiation causes the temperature to rise in the outer disk (r > 20 au).

In the non-shadowed disk (f = 1.0), the CO_2 snowline position is r ~ 3.3 au (T(r) ~ 75 K). CO_2 ice abundances are ~5 × 10^−5 between the CO_2 and the CO snowlines, and decrease (<10^−6) beyond the CO snowline. The CO_2 snowline position moves inward with decreasing f. In the shadowed disk (f ≤ 0.03), CO_2 ice abundances decrease at r ~ 3 − 8 au (from ~5 × 10^−5 to <10^−6) with decreasing f, whereas H_2O ice abundances at such radii slightly increase (from ~1.5 × 10^−4 to ~2 × 10^−4). Eistrup et al. (2016) discussed that at the coldest conditions (T(r) ≤ 20 K), formation of H_2O ice (Equation 5) is faster than that of CO_2 ice (Equation 6). This is because at the coldest conditions (T(r) ≤ 20 K), the mobility of H_ is higher than those of CO and OH, and because the adsorption rate of atomic hydrogen onto dust grains at T(r) ≤ 20 K becomes larger than that at T(r) > 20 K (E_{des}(H) = 650 K).

In our standard disk model, O_2 abundances are much smaller (≤10^−6 for gas and ≤10^−9 for ice) than those of H_2O, CO, and CO_2. O_2 is formed in the gas phase via the following reaction (Walsh et al. 2015; Eistrup et al. 2016; Notsu et al. 2021):

\[
\text{O} + \text{OH} \rightarrow \text{O}_2 + \text{H}.
\]

(8)
Since O$_2$ is very volatile (the binding energy $E_{\text{des}}$(O$_2$) = 898 K; Noble et al. 2012), it remains in the gas phase at $r \lesssim 20$ au in the non-shadowed disk ($f = 1.0$). The O$_2$ gas abundances between the water and O$_2$ snowlines are $\sim 10^{-7} - 10^{-6}$.

In the shadowed disk ($f \lesssim 0.03$), O$_2$ freezes out onto dust grains at $r \sim 2-10$ au, and returns to the gas phase at $r > 10$ au. Outside the O$_2$ snowline, O$_2$ ice is formed via the following grain-surface reaction (Taquet et al. 2016; Eistrup & Walsh 2019):

$$O_{\text{gas}} + O_{\text{ice}} \rightarrow O_{\text{ice}}$$  \hspace{1cm} (9)

In our standard disk model, the initial atomic oxygen is zero (see Table 2). In addition, the gas-phase production route of atomic oxygen from CO by the reaction of Equation (7) is not efficient outside the O$_2$ snowline since the binding energies of CO and O$_2$ are almost similar (see Table 1). Thus, in our standard disk model, O$_2$ ice abundances are low ($\ll 10^{-8}$) both in the shadowed and non-shadowed disks.

We note that our calculations also yield higher O$_2$ ice abundances ($\sim 10^{-7} - 10^{-6}$) for atomic initial abundances and low ionization rates. Such larger abundances are consistent with the measured cometary abundances (Bieler et al. 2015; Rubin et al. 2015b, see also Section 4.2 of this paper). Moreover, we find that in the shadowed disk, such an O$_2$ abundant region is located in the inner region ($r \sim 2-10$ au) compared with that in the non-shadowed disk ($r \geq 20$ au), as detailed in Section 4.2 and Appendix A.1.

3.2.2. Other Dominant Carbon-bearing Molecules

Figure 4 shows the radial profiles of fractional abundances for other dominant carbon-bearing molecules (CH$_4$, C$_2$H$_6$, H$_2$CO, and CH$_3$OH). We note that CH$_4$ (methane) and C$_2$H$_6$ (ethane) are dominant acyclic saturated hydrocarbon molecules (alkanes). In the non-shadowed disk ($f = 1.0$), the CH$_4$ snowline position is $r \sim 15$ au ($T(r) \sim 38$ K), and the CH$_4$ gas abundances between the CO$_2$ and CH$_4$ snowlines are $\sim 10^{-7}$. These values are much smaller than those in the disk with a lower ionization rate ($10^{-18}$ [s$^{-1}$]; see Appendix A) and that for the initial CH$_4$ gas abundance ($1.120 \times 10^{-6}$). Outside the CH$_4$ snowline, CH$_4$ ice abundances are slightly enhanced ($\sim 10^{-5}$) compared with the initial CH$_4$ ice abundances ($7.384 \times 10^{-6}$).

Here we describe the dominant formation and destruction routes of CH$_4$ both in the gas phase and icy phase. The destruction of gas-phase CO by He$^+$ (see Section 3.2.1) produces C$^+$, and a sequence of reactions with, e.g., H$_2$ transform C$^+$ to, e.g., CH$_2^+/CH_3^+/CH_4^+$, which lead to CH$_4$ gas and ice (e.g., Aikawa et al. 1999; Furuya & Aikawa 2014; Bosman et al. 2018, see Section 3.2.1). In addition, Bosman et al. (2018) and Eistrup et al. (2018) discussed that CH$_4$ ice is also produced both inside/outside the CO snowline through the hydrogenation of CH$_3$ ice, which is formed by cosmic-ray-induced photodissociation of CH$_3$OH (see also below). Inside the CH$_4$ snowline, CH$_4$ gas is destroyed by cosmic-ray-induced photodissociation and ion-molecule reactions (such as CH$_2$ + C$^+$), and the carbon is thus converted from CH$_4$ gas to CO$_2$, H$_2$CO, and hydrocarbons such as, e.g., C$_2$H$_2$, C$_2$H$_6$, and C$_3$H$_4$ (Aikawa et al. 1999; Eistrup et al. 2016, 2018; Yu et al. 2016, see also Section 3.2.4).

The CH$_4$ snowline positions moves inward with decreasing $f$. In the shadowed disk ($f \lesssim 0.03$), CH$_4$ snowline positions are $r \lesssim 2$ au, and CH$_4$ ice abundances at $r \sim 2-10$ au are around $\sim 10^{-5}$. This result indicates that the CH$_4$ ice abundances are comparable in both CO sublimation and frozen regions. Thus, we interpret that the above reaction pathway starting from cosmic-ray-induced photodissociation of CH$_3$OH is efficient for CH$_4$ ice formation in this region.

In the non-shadowed disk ($f = 1.0$), the C$_2$H$_6$ snowline position is $r \sim 3.2$ au, which is similar to that for CO$_2$ ($E_{\text{des}}$(CO$_2$) = 2267 K, $E_{\text{des}}$(C$_2$H$_6$) = 2320 K). The C$_2$H$_6$ ice abundances outside its snowline increase with increasing $r$. The ice abundances are $\sim (3-6) \times 10^{-6}$ outside the CH$_4$ snowline, which is larger than the initial ice abundance ($= 2.417 \times 10^{-6}$).

Previous modeling studies suggested that unsaturated hydrocarbon molecules (such as C$_2$H$_2$, C$_3$H$_2$, and C$_3$H$_4$) are efficiently produced with the chemical reaction pathways starting from CH$_4$ + C$^+$, where C$^+$ is formed by CO + He$^+$ (e.g., Aikawa et al. 1999; Eistrup & Aikawa 2014; Eistrup et al. 2016, 2018; Yu et al. 2016; Bosman et al. 2018, see Section 3.2.1). In addition, such unsaturated hydrocarbon molecules are also supplied by the cosmic-ray-induced photodissociation of CH$_4$ and subsequent gas-phase combination reactions of CH$_3$ radicals (e.g., Aikawa et al. 1999; Eistrup et al. 2016; Yu et al. 2016). These reaction pathways are efficient in the inner warm disk, at least within the CH$_4$ snowline (see also Section 3.2.4). The saturated hydrocarbons including C$_2$H$_6$ are formed on the grain surfaces by the hydrogenation of such unsaturated hydrocarbon ices (e.g., Aikawa et al. 1999; Bosman et al. 2018). In addition, Bosman et al. (2018) and Eistrup et al. (2018) discussed that C$_2$H$_6$ ice is also produced both inside/outside the CO snowline by CH$_3$ice + CH$_3$ice, where CH$_3$ ice is formed by cosmic-ray-induced photodissociation of CH$_3$OH ice. Thus, we conclude that grain-surface reactions are needed to form icy-phase saturated hydrocarbon molecules, including C$_2$H$_6$.

In the shadowed disk ($f \lesssim 0.03$), C$_2$H$_6$ snowline positions are $r < 2$ au, and C$_2$H$_6$ ice abundances at $r \sim 2-10$ au significantly increase ($\sim 5 \times 10^{-6}$) compared with the values in the non-shadowed disk. We interpret that the reaction pathway starting from CH$_3$OH ice destruction becomes more important at these radii of the shadowed disk, since the abundances of unsaturated hydrocarbon ices (such as C$_2$H$_2$, C$_3$H$_2$, and C$_3$H$_4$) decrease in the shadowed disk (see Section 3.2.4 for details).

In the non-shadowed disk ($f = 1.0$), the H$_2$CO (formaldehyde) and CH$_3$OH (methanol) snowline positions are $r \sim 1.8$ au ($T(r) \sim 100$ K) and $\sim 1.5$ au ($T(r) \sim 120$ K), respectively, and they are just outside the water snowline ($r \sim 1.3$ au). The H$_2$CO and CH$_3$OH ice abundances outside their snowlines increase with increasing $r$. The H$_2$CO ice abundances between the H$_2$CO and CO snowlines are $\sim 8 \times 10^{-6}$ at most, which is similar to the initial H$_2$CO ice abundance ($= 8.437 \times 10^{-6}$). The CH$_3$OH ice abundances between the CH$_3$OH and CO snowlines are $\sim 2 \times 10^{-7}$, which is smaller than the initial CH$_3$OH ice abundance ($= 6.027 \times 10^{-7}$). We suggest that the differences in the CH$_3$OH and H$_2$CO abundances with respect to the initial abundances inside the CO snowline are also related to whether or not gas-phase formation pathways are present (e.g., Fockenberg & Preses 2002; Atkinson et al. 2006; Loomis et al. 2015; Walsh et al. 2016; Pegues et al. 2020). Both H$_2$CO and CH$_3$OH ice abundances increase outside the CO snowline ($\sim (2-5) \times 10^{-5}$ for H$_2$CO ice and $\sim (2-9) \times 10^{-6}$ for CH$_3$OH ice).
The freeze-out of CO onto dust grains plays a central role in producing organic molecules, such as H2CO and CH3OH. According to previous studies (e.g., Tielens & Hagen 1982; Watanabe & Kouchi 2002; Cuppen et al. 2009; Fuchs et al. 2009; Drozdovskaya et al. 2014; Furuya & Aikawa 2014; Walsh et al. 2014, 2016; Chuang et al. 2016; Bosman et al. 2018; Aikawa et al. 2020; see also Appendix A.2), the following sequential hydrogenation of CO on the dust grain surfaces leads to the formation of H2CO and CH3OH ices:

\[
\text{CO}_{\text{ice}} + \text{H} \rightarrow \text{HCO} + \text{H} \rightarrow \text{H}_2\text{CO}_{\text{ice}},
\]

and

\[
\text{H}_2\text{CO}_{\text{ice}} + \text{H} \rightarrow \text{CH}_3\text{OH}_{\text{ice}} + \text{H} \rightarrow \text{CH}_3\text{OH}_{\text{ice}}.
\]

This grain-surface reaction pathway produces a large amount of H2CO and CH3OH ices in the cold regions where CO freezes out onto dust grains. The cosmic-ray-induced photodissociation of CH3OH ice produces many radicals such as H2CO ice and of H2CO and CH3OH ices in the cold regions where CO abundances are relatively constant (the disk for various values of r, the H2CO and CH3OH ice abundances are relatively constant with the values in the non-shadowed disk, since the temperature is low (T(r) < 30 K) and the sequential hydrogenation of CO ice (Equations (10) and (11)) proceeds efficiently. In contrast, their ice abundances at r > 20 au become a bit smaller (∼10−5 for H2CO ice and ∼(1−3) × 10−6 for CH3OH ice), since the CO adsorption front moves outward with decreasing f.

3.2.3. Dominant Nitrogen-bearing Molecules

Figure 5 shows the radial profiles of fractional abundances for dominant nitrogen-bearing molecules (N2, NH3, HCN, and NH2CHO). In the non-shadowed disk (f =1.0), the N2 (molecular nitrogen) snowline position is r ∼ 24 au (T(r) ∼ 23 K), which is slightly outside the CO snowline. For the atomic initial abundances, atomic N quickly forms N2 gas and it freezes out onto dust grains outside its snowline (Schwarz & Bergin 2014; Eisner et al. 2016). In addition, for the molecular initial abundances, the initial atomic N abundance (=2.1 × 10−5; see Table 1) is similar to the initial N2 gas + ice abundance (=1.5 × 10−7; see Table 1), and this atomic N is also quickly converted to N2 in the disk. Thus, the N2 gas + ice abundances are relatively constant (∼(2−3) × 10−5) throughout the disk for various values of f and various initial conditions (see also Appendix A.3).

In the shadowed disk (f ≤ 0.03), like CO, N2 freezes out onto dust grains at around the current orbit of Jupiter (r ∼ 3–8 au), and N2 ice abundances are ∼ 2 × 10−5. In addition, N2 returns to the gas phase at r > 8 au. At r > 20 au, N2 (re-)freezes out onto dust grains, and the adsorption front moves outward with decreasing f, since T(r) at r > 20 au increases (<35 K at the maximum) with decreasing f (see also Section 2.1). This N2 abundance profile is consistent with the results reported in Ohno & Ueda (2021).

In the non-shadowed disk (f =1.0), the NH3 (ammonia) snowline position is r ∼ 2.3 au, which is slightly outside the water snowline (r = 1.3 au). NH3 ice abundances just outside the NH3 snowline (r ∼ 2–4 au) are <10−6, and those at r > 5 au are ∼1 × 10−5, which is similar to the initial NH3 ice abundance (∼1.327 × 10−5). We suggest that on the warm grain surfaces at r ∼ 2–4 au, NH3 ice is converted to NH3CHO ice via the following grain-surface reaction route (Jones et al. 2011; Walsh et al. 2014; López-Sepulcre et al. 2015):

\[
\text{NH}_2\text{CHO}_{\text{ice}} + \text{HCO}_{\text{ice}} \rightarrow \text{NH}_3\text{CHO}_{\text{ice}},
\]

where NH3 is formed by the cosmic-ray-induced photodissociation of NH3 ice.

In the shadowed disk (f ≤ 0.03), the NH3 snowline positions are r < 2 au. NH3 ice abundances at r ∼ 2–5 au are significantly enhanced, since the disk temperature decreases, and the above destruction route of NH3 ice (see Equation (12)) becomes inefficient.

Formamide (NH2CHO) has been proposed as a key precursor of various (pre)metabolic and (pre)genetic molecules such as pyruvic acid and adenine (Saladin et al. 2012). Gas-phase NH2CHO has been detected in pre-stellar cores and protostar envelopes (e.g., Kahane et al. 2013; López-Sepulcre et al. 2015; Okada et al. 2021) and comets (e.g., Bockelée-Morvan et al. 2000), but has not yet been detected in Class II disks. The NH2CHO snowline position (r ∼ 1.2 au, T(r) ∼ 150 K) is slightly inside the water snowline (Edes(NH2CHO) = 5560 K, Edes(H2O) = 4880 K); thus, the NH2CHO snowline position does not move in our chemical modeling.

In the non-shadowed disk (f =1.0), the NH2CHO ice abundances are ∼ 10−6 in the outer cold disk (around the CO/N2 snowlines, r ∼ 14–33 au), and <10−7 at r < 10 au. In addition, NH2CHO is efficiently formed (∼ 10−6) just outside its snowline (r ∼ 1.2–4 au). We interpret that just outside its snowline, the above radical–radical formation route on the warm dust grains is efficient (see Equation (12)).

In the shadowed disk (f ≤ 0.03), NH2CHO ice abundances at around the current orbit of Jupiter (r ∼ 3–8 au) are enhanced (∼ 10−6). We propose that the following hydrogenation routes become efficient in this region. Garrod et al. (2008), Walsh et al. (2014), and López-Sepulcre et al. (2015) discussed that grain-surface NH2CHO can form via the atom addition reactions in the cold region, such as the following reaction routes:

\[
\text{OCN}_{\text{ice}} + \text{H} \rightarrow \text{HOCN}_{\text{ice}} + \text{H} \rightarrow \text{HCN}_{\text{ice}}/\text{NHCHO}_{\text{ice}}
\]

and

\[
\text{NH}_2\text{CO}_{\text{ice}}/\text{NHCHO}_{\text{ice}} + \text{H} \rightarrow \text{NH}_2\text{CHO}_{\text{ice}},
\]

where OCNice is formed via the grain-surface reactions (CNice + Oice and ONice + Cice) or gas-phase radical–radical reactions. We note that our chemical reaction network includes a competing hydrogenation abstraction pathway during hydrogenation from HNCO to NH2CHO, on the basis of the experimental results in Noble et al. (2015).

HCN (hydrogen cyanide) and CN (cyanide) are the simplest molecules containing both C and N atoms, and reactions including HCN lead to the formation of more complex cyanides (such as CH2CN and aminoacetonitrile H2NCH2CN;
Öberg et al. 2011; Noble et al. 2013). In the non-shadowed disk ($f = 1.0$), the HCN snowline position is $r \sim 1.5$ au ($T(r) \sim 120$ K), which is just outside the water snowline ($r = 1.3$ au). HCN ice abundances are $\sim (1-3) \times 10^{-6}$ at $r \sim 3$–$20$ au and $\sim 4 \times 10^{-6}$ at $r \geq 20$ au, which are slightly smaller/larger than the initial HCN ice abundance ($=2.772 \times 10^{-5}$). Eisrtp et al. (2016) showed much lower HCN ice abundances ($\ll 10^{-7}$ outside the HCN snowline) for their model calculations with molecular initial abundances, since they set the initial HCN ice abundance to zero.

In the shadowed disk ($f \leq 0.03$), the HCN ice abundances at around the current orbit of Jupiter ($r \sim 3$–$8$ au) are slightly enhanced ($\sim 5 \times 10^{-6}$). According to Aikawa et al. (1999) and Eisrtp et al. (2016), HCN is formed through the gas-phase reaction of HCO with an N atom, with subsequent freeze-out onto dust grains. HCO gas is efficiently formed by hydrogenation of CO on the cold dust grain surface and cosmic-ray-induced desorption of HCO gas (see also Sections 3.2.1 and 3.2.2). Schwarz & Bergin (2014) and Eisrtp et al. (2018) showed that HCN ice becomes abundant in the outer cold region of the disk where CO freezes out onto dust grains, although it is efficiently produced within a few megayears. Our calculations for both non-shadowed and shadowed disks also show that HCN ice is efficiently formed at the cold regions where CO freezes out onto dust grains. Thus we confirm that the above HCN formation route, which starts from hydrogenation of CO, is dominant in the disks.

### 3.2.4. Other Organic Molecules

In this Section 3.2.4, we show the radial abundance distributions of other major carbon-, nitrogen-, and oxygen-bearing molecules, mainly larger organic molecules.9 Figure 6 shows the radial profiles of fractional abundances for $\text{C}_2\text{H}_2$ (acetylene), $\text{C}_3\text{H}_4$ (ethylene), $\text{C}_5\text{H}_2$ (cyclop propane), $\text{C}_9\text{H}_4$ (propyne $\text{CH}_2\text{CCH}$ and allene $\text{CH}_2\text{CCH}_2$), and $\text{C}_1\text{H}_2$ (propylene, $\text{CH}_3\text{CHCH}_3$). These are dominant unsaturated hydrocarbon molecules. In the non-shadowed disk ($f = 1.0$), $\text{C}_2\text{H}_2$ and $\text{C}_2\text{H}_4$ ice abundances are largest ($\sim (1-2) \times 10^{-6}$ and $\sim (3-8) \times 10^{-6}$, respectively) at $r \sim 10$–$15$ au, which is just inside the CH$_4$ snowline. In the shadowed disk ($f \leq 0.03$), these ice abundances significantly decrease ($\ll 10^{-7}$) at $r \sim 3$–$15$ au, and increase ($\sim 10^{-6}$) at $r \sim 2$ au. Moreover, in the non-shadowed disk, the ice abundances of other molecules ($\text{C}_2\text{H}_2$, $\text{C}_3\text{H}_4$, and $\text{C}_5\text{H}_2$) are enhanced at $r \sim 2$–$15$ au, whereas in the shadowed disk, they decrease at $r \sim 2$–$15$ au and increase at $r \sim 1$–$2$ au.

As we describe in Section 3.2.2 (see also, e.g., Aikawa et al. 1999; Furuya & Aikawa 2014; Eisrtp et al. 2016, 2018; Yu et al. 2016; Bosman et al. 2018), the formation pathways of these unsaturated hydrocarbon molecules are efficient within the CH$_4$ snowline, since the gas-phase chemical reaction pathways starting from CH$_3$ and C$^+$, and the cosmic-ray-induced photodissociation of CH$_3$ and subsequent gas-phase combination reactions of CH$_3$ radicals are active. In the shadowed disks ($f \leq 0.03$), the CH$_4$ snowline positions are $r \leq 2$ au, and such gas-phase formation reactions do not proceed efficiently at $r \geq 2$ au. Thus, the unsaturated hydrocarbon ices are deficient at $r \sim 3$–$15$ au (including the current orbit of Jupiter), in contrast to saturated hydrocarbons such as CH$_3$ and C$^+$.

Figure 7 shows the radial profiles of fractional abundances for CH$_3$NH$_2$ (methyl amine), CH$_2$NH (methylene imine), CH$_3$CN (acetoc nitrile), HC$_3$N (cyanoc acetylene), NH$_2$OH (hydroxylamine), and HNCO (isocyanic acid). Figure 8 shows the radial profiles of fractional abundances for CH$_3$CHO (acetaldehyde), CH$_3$OCH$_3$ (dimethyl ether), HCOOH (formic acid), HCOOH$_3$ (methyl formate), CH$_3$COOH (acetic acid), and C$_2$H$_5$OH (ethanol). The ice abundances of CH$_3$NH$_2$ and NH$_2$OH at $r \sim 3$–$8$ au (around the Jupiter orbit) are smaller in the non-shadowed disk ($\sim 10^{-8}$ and $\sim 2 \times 10^{-6}$ for $f = 1.0$, respectively) than those in the shadowed disk ($\sim (2-7) \times 10^{-6}$ and $(3-8) \times 10^{-6}$ for $f \leq 0.03$). These are saturated molecules. In contrast, the ice abundances of CH$_3$NH$_2$, HC$_3$N, HCOOH, HCOOH$_3$, CH$_3$COOH, and HNCO at $r \sim 3$–$8$ au are larger in the non-shadowed disk (e.g., $\sim 10^{-7}$ and $\sim 6 \times 10^{-6}$ for HCOOH and $f = 0.03$) than those in the shadowed disk (e.g., $\lesssim 10^{-8}$ for HCOOH and $f \leq 0.03$).

Moreover, the ice abundances of CH$_3$CN, CH$_3$CHO, CH$_3$OCH$_3$, and C$_2$H$_5$OH at $r \sim 3$–$8$ au are larger for $f = 0.03$ (e.g., $\sim 10^{-6}$ for CH$_3$CN) than those for $f = 1.0$, 0.3, and 0.003 (e.g., $< 10^{-7}$ for CH$_3$CN).

Walsh et al. (2014) discussed that CH$_3$NH$_2$ is formed via the sequential hydrogenations of CH$_3$NH$_2$ ice, where CH$_3$NH$_2$ ice is originated from atom addition to small hydrocarbon radicals (CH$_3$, CH$_2$) on the dust grain surfaces. In addition, Garrod et al. (2008) and Walsh et al. (2014) described that on the warm dust grain surfaces, the association of the methyl and amine radicals (CH$_3$ice + NH$_2$ice) is also dominant. Such small hydrocarbon radicals are efficiently formed both in the warm and cold grain surfaces by cosmic-ray-induced photodissociation of CH$_3$OH ice, which is efficiently produced by Equations (10) and (11) (see Section 3.2.2 and Eisrtp et al. 2018). Thus, we find that the former atom addition reaction route is efficient in the shadowed region ($T(r) < 30$ K).

Garrod et al. (2008) explained that NH$_2$OH is formed initially by NH + OH addition on grains, followed by hydrogenation. Also, the addition reaction of OH + NH$_2$ becomes dominant on the warm dust grains (see also, e.g., Molyarova et al. 2018). We propose that the former reaction route is efficient in the shadowed region.

We find that the dependence of CH$_3$CN and HC$_3$N ice abundances on the values of $f$ are determined by the combination of the following reaction routes. Walsh et al. (2014) described that similar to H$_2$CO, both CH$_3$CN and HC$_3$N can form in the gas phase through multiple pathways. On the cold dust grain surfaces ($T(r) \sim 10$–$30$ K), CH$_3$CN ice is formed via the sequential hydrogenations of C$_2$N ice, and C$_2$N can form via C$_{ice}$ + CN$_{ice}$ and N$_{ice}$ + C$_2$$_{ice}$, or freeze out from the gas phase (Walsh et al. 2014). In warmer regions, CH$_3$CN ice can also form via the following radical–radical reaction with no reaction barrier, CN$_{ice}$ + CH$_3$$_{ice}$. In addition, on the cold dust grain surface ($T(r) \sim 10$–$30$ K), HC$_3$N ice is formed via the hydrogenation of C$_2$N$_{ice}$.

Garrod et al. (2008), Walsh et al. (2014), and López-Sepulcre et al. (2015) described that HNCO is formed on the cold dust grain surfaces by hydrogenation of OCN, but is further hydrogenated to NH$_2$CHO (see Equations (13) and (14) in Section 3.2.3).

For other organic molecules, grain-surface association of large radical–radical reactions in the warm regions ($T(r) \sim 50$ K) is needed for their formation (e.g., Garrod & Herbst 2006; Garrod et al. 2008; Herbst & van Dishoeck 2009; Vasyunin &

---

9 We note that NH$_2$OH is technically an inorganic molecule since it does not include carbon.
Herbst 2013; Walsh et al. 2014). In addition, the dependence of ice abundances on the values of \( f \) is determined by the combination of the following radical–radical reactions and radical formation reactions from CO and/or CH\(_3\)OH on the dust grain surfaces. CH\(_3\)CHO, CH\(_3\)OCH\(_3\), HCOOCH\(_3\), and C\(_2\)H\(_5\)OH ices can form via the following radical–radical reactions with no reaction barriers on the dust grain surfaces, respectively (Garrod et al. 2008; Walsh et al. 2014):

\[
\begin{align*}
\text{CH}_3\text{ice} + \text{HCOice} & \rightarrow \text{CH}_3\text{CHOice}, \\
\text{CH}_3\text{ice} + \text{H}_2\text{Oice} & \rightarrow \text{CH}_3\text{OCH}_3\text{ice}, \\
\text{HCOice} + \text{CH}_3\text{Oice} & \rightarrow \text{HCOOCH}_3\text{ice},
\end{align*}
\]

and

\[
\text{CH}_3\text{ice} + \text{CH}_2\text{OHice} \rightarrow \text{C}_2\text{H}_5\text{OHice}.
\]

The radicals of HCO, CH\(_3\), CH\(_2\)OH, and CH\(_3\)OH are produced during the sequential hydrogenation of CO to form CH\(_3\)OH and/or the cosmic-ray-induced photodissociation of CH\(_2\)OH ices. The abundance of CH\(_3\)OH ice at \( r \sim 3–8 \) au is larger in the shadowed disk than that in the non-shadowed disk (see Section 3.2.2), whereas the shadowed disks with lower values of \( f \) (=0.003) exhibit lower ice abundances for CH\(_3\)CHO, CH\(_3\)OCH\(_3\), HCOOCH\(_3\), and C\(_2\)H\(_5\)OH. This trend is owing to the temperature dependence of the radical–radical reaction efficiency: the mobility of such radicals is lower than that of hydrogen and light atoms, especially on the cold dust grain surfaces (\( T(r) \sim 10–30 \) K).

HCOOH and CH\(_3\)COOH ices can also form via the radical–radical reaction routes on the warm dust grain surfaces. However, CH\(_2\)CO, the precursor of CH\(_3\)COOH and CH\(_3\)CHO (the hydrogenation route), cannot form on the cold dust grain surfaces, since the reaction barrier of CH\(_3\)ice + CO\(_2\)ice is large (\( \sim 3460 \) K; Walsh et al. 2014). Moreover, COOH, the precursor of CH\(_3\)COOH and HCOOH, cannot form on the cold dust grain surfaces, since the reaction barrier of OH\(_2\)ice + CO\(_2\)ice is large (\( \sim 3000 \) K; Walsh et al. 2014). We suggest that HCOOH ice can be formed via OH\(_2\)ice + H\(_2\)CO, although the mobility of both ices is low on the cold dust grain surfaces (Garrod et al. 2008).

### 3.2.5. The Total Complex Organic Reservoir

Figure 9 shows the radial profiles of sums of fractional abundances with respect to total hydrogen nuclei densities at \( t = 10^6 \) yr for larger (complex) organic molecules. In the left (“All”) panel of Figure 9, we show the sums of all larger (complex) organic molecules introduced in Section 3.2 (Figures 4–8); C\(_6\)H\(_6\), H\(_2\)CO, CH\(_3\)OH, NH\(_2\)CHO, C\(_2\)H\(_2\), C\(_3\)H\(_4\), C\(_4\)H\(_4\), CH\(_2\)CCH, CH\(_2\)CCH\(_2\), CH\(_3\)CHCH\(_2\), HC\(_3\)N, CH\(_2\)N\(_2\), CH\(_3\)CHO, CH\(_3\)OCH\(_3\), CH\(_3\)NH\(_2\), HCOOH, HCOOCH\(_3\), CH\(_3\)COOH, CH\(_2\)NH, HNCO, and C\(_2\)H\(_5\)OH. We note that CO, H\(_2\)O, O\(_2\), CH\(_4\), N\(_2\), NH\(_3\), HCN, and NH\(_2\)OH are not included in this Figure, since they are smaller molecules and/or inorganic molecules.

In the non-shadowed disks, the total icy fractional abundances of the listed organic molecules is \( \sim 3 \times 10^{-6} \) just outside the water snowline (\( r \sim 1.5 \) au), and gradually increases with increasing \( r \). It is \( (1-2) \times 10^{-5} \) at \( r \sim 3-8 \) au, around the current orbit of Jupiter, and is \( (4-6) \times 10^{-5} \) outside the CO snowline. In the shadowed disk, it is \( (3-5) \times 10^{-5} \) at \( r \sim 3-8 \) au, and \( 2 \times 10^{-5} \) at \( r \gtrsim 10 \) au. Thus in the shadowed disk, our model predicts that the dust grains at \( r \sim 3–8 \) au contain \( \sim 2–5 \) times the amount of larger (saturated+unsaturated) organic molecules in total compared with those in the non-shadowed disks.

In the right (“Reduced”) panel of Figure 9, we show the sums of all larger (complex) organic molecules but exclude the dominant unsaturated hydrocarbon molecules shown in Figure 6: C\(_2\)H\(_2\), C\(_2\)H\(_4\), C\(_3\)H\(_2\), CH\(_2\)CCH, CH\(_2\)CH\(_2\), and CH\(_3\)CCH\(_2\). In the non-shadowed disks, at \( r \sim 3–8 \) au, the icy abundances of sums of these “Reduced” organic molecules are smaller (\( \sim (3-9) \times 10^{-6} \)) than those of “All” organic molecules (\( (1-2) \times 10^{-5} \)). In contrast, in the shadowed disks, the icy abundances of sums of organic molecules are similar (\( (3-5) \times 10^{-5} \)) between these two cases. This is because such dominant unsaturated hydrocarbon molecules are frozen onto the dust grains at \( r \sim 3–8 \) au, and their ice abundances significantly decrease with decreasing \( f \) (see Figure 6). This result demonstrates that the shadowed region promotes the synthesis of saturated organic molecules rather than unsaturated hydrocarbons. Thus in the shadowed disk, the dust grains at \( r \sim 3–8 \) au are predicted to have around 5–10 times the amount of larger saturated organic molecules in total compared with those in the non-shadowed disks.

### 3.3. The Dependence of the Disk Chemical Structure on Ionization Rate and Initial Abundances

In this subsection, we briefly summarize key results of disk chemical evolution for atomic initial compositions and different ionization rates, which are detailed in Appendix A. Figures 14–16 in Appendix A show the radial profiles of fractional abundances at \( t = 10^6 \) yr for dominant oxygen-, carbon-, and nitrogen-bearing molecules (H\(_2\)O, CO, CO\(_2\), O\(_2\), CH\(_4\), C\(_2\)H\(_6\), H\(_2\)CO, CH\(_3\)OH, N\(_2\), NH\(_3\), HCN, and NH\(_2\)CHO) in the shadowed and non-shadowed disk midplane \( (f = 1.0 \) and \( f = 0.03 \), respectively). In these Figures, we assume either molecular or atomic initial abundances and either low or high ionization rates \( (\xi_{\text{CR}}(r) = 10^{-18}, 10^{-17} \text{ [s}^{-1}\text{]}\)\)). We note that Figures 3–5 in Section 3.2 show the radial abundance profiles for the same molecules with molecular initial abundances and \( \xi_{\text{CR}}(r) = 10^{-17} \text{ [s}^{-1}\text{]}\)\).

In Appendix A, we show these results and explain the dependence of the disk chemical evolution on disk ionization rates and initial abundances in detail. According to our calculations, CO/CO\(_2\) abundances become smaller/larger with increasing ionization rates, respectively. CH\(_4\) and C\(_2\)H\(_6\) gas abundances within their snowline become smaller as the ionization rates become larger. In addition, abundances of H\(_2\)O and carbon molecules are larger for molecular initial abundances than those for atomic initial abundances. It is worth noting that O\(_2\) ice abundances are \( \sim 10^{-5} – 10^{-4} \) (consistent with the measured cometary abundances; see also Section 4.2) only for atomic initial abundances and the low ionization rates.

### 3.4. Time Evolution of Molecular Abundances

In this subsection, we briefly summarize some key results for the time evolution of molecular abundances, which are detailed in Appendix B. Figures 17–19 in Appendix B show the time evolution of the radial profiles of fractional abundances for dominant oxygen-, carbon-, and nitrogen-bearing molecules (H\(_2\)O, CO, CO\(_2\), O\(_2\), CH\(_2\), C\(_2\)H\(_6\), H\(_2\)CO, CH\(_3\)OH, N\(_2\), NH\(_3\), HCN, and NH\(_2\)CHO) in the shadowed and non-shadowed disk.
midplane \((f=1.0\) and 0.03\), when assuming molecular initial abundances and high ionization rates \(\xi_{\text{CR}}(r)=10^{-17} \text{ [s}^{-1}]). These initial conditions are similar to those in Sections 3.1 and 3.2.

In Appendix B, we describe these results and explain the time evolution of molecular abundances in detail. According to our calculations, in the shadowed region \((r \sim 3-8 \text{ au})\), the icy abundances of CO2 and organic molecules such as H2CO, CH3OH, and NH2CHO become larger with time. Thus, we find that if the shadowed region is maintained for a relatively long time \((t \sim 10^6 \text{ yr})\), chemical evolution may produce dust grains and solid objects with large amounts of CO2 and organic molecular ices (see also Section 4.2). In addition, the ice abundances of these molecules can be a clue in constraining the formation age of solid bodies in the shadowed region.

4. Discussion

4.1. The C/O and N/O Ratios and Implication for Planetary Atmospheres

The carbon-to-oxygen (C/O) ratios of exoplanetary atmospheres have been proposed to be a possible tool to link giant exoplanets to their formation sites in the natal protoplanetary disk (e.g., Öberg et al. 2011; Öberg & Bergin 2016; Madhusudhan et al. 2014; Pontoppidan et al. 2014; Booth & Ilee 2019; Cridland et al. 2020b; Notsu et al. 2020; Ohno & Ueda 2021; Schneider & Bitsch 2021; Turrini et al. 2021; Dash et al. 2022; Mollière et al. 2022). This is because the radial-dependent positions of snowlines of abundant oxygen- and carbon-bearing molecules result in systematic radial variations in the C/O ratios in the gas and ice. However, disk chemistry can affect the C/O ratios in the gas and ice, thus potentially erasing the chemical fingerprint of snowlines in atmospheres (Eistrup et al. 2016, 2018; Notsu et al. 2020). Notsu et al. (2020) discussed that hot Jupiters with C/O > 1 can only form between the CO2 and CH4 snowlines in the non-shadowed disk, which has fully inherited interstellar abundances, and where negligible chemistry has occurred because of a low ionization rate. They also discussed that carbon-rich planets are likely rare unless efficient transport of hydrocarbon-rich ices via pebble drift to within the CH4 snowline (Booth et al. 2017; Booth & Ilee 2019) is a common phenomenon. We note that disk chemistry significantly affects the C/O ratios, through the change of abundances of hydrocarbons and O2 gas abundances (e.g., Helling et al. 2014; Eistrup et al. 2016, 2018; Notsu et al. 2020; Ohno & Ueda 2021).

The left panels of Figures 10 and 11 show the radial profiles of C/O ratios at \(t = 10^6 \text{ yr}. Different color lines show the profiles for different values of the parameter \(f = 1.0, 0.3, 0.03, \) and 0.003, respectively. As shown in our previous studies (Eistrup et al. 2016, 2018; Notsu et al. 2020), the inclusion of chemistry has a significant impact on the disk elemental abundance ratios of both gas and ice. The ices remain, on the whole, dominated by oxygen (i.e., C/O < 0.5). In the non-shadowed disk \((f = 1.0)\), between the H2O and CH4 snowlines, the gas is carbon rich relative to the initial elemental value \((=0.44)\) for molecular initial abundances, whereas gas-phase C/O ratios are \(<0.44\) for atomic initial abundances. For molecular initial abundances, the gas-phase C/O ratios in these regions are \(\sim 1.0-1.1\) for the low ionization rate and \(\sim 0.7-1.0\) for the high ionization rate.

In the shadowed disk \((f \leq 0.03)\), however, the gas-phase C/O ratios in the shadowed region \((r \sim 2-10 \text{ au})\) are almost unity for the high ionization rate and/or molecular initial abundances (as suggested by Ohno & Ueda 2021) and 0.8 for the low ionization rate and atomic initial abundances. This is because CO carries most of the gas-phase C and O there. In addition, the icy-phase C/O ratios reach the initial elemental value \((=0.44 \text{ in our model})\) since almost all C and O reservoirs freeze out onto dust grains in the shadowed region \((r \sim 3-8 \text{ au})\). Thus, if planets acquire their atmospheres from the gas in the disk shadowed region, the atmospheres will have substellar metallicities and larger C/O ratios than the initial elemental value \((=0.44 \text{ in our model})\) regardless of disk initial abundances and ionization rates. In addition, the C/O ratios are unity unless the disk has fully atomized initial abundances and negligible chemistry has occurred because of the low ionization rate. We note that the atmospheres can be polluted by dissolution of accreting oxygen-rich icy planetesimals/pebbles (e.g., Hori & Ikoma 2011; Mordasini et al. 2016) and erosion of cores (e.g., Moll et al. 2017), potentially lowering C/O ratios in planetary atmospheres.

In the bottom-left panel of Figure 11 (the case for low ionization and atomic initial abundances), the icy-phase C/O ratios just outside the water snowline are larger than the initial elemental value, and they exceed 1.0 for \(f \leq 0.3\). The peak position shifts inside with decreasing \(f\). In this region, H2O ice abundances are much smaller \((\sim 10^{-6})\) than those in other models \((>10^{-5})\). In addition, this region is between the HCN and CO2 snowline, and the HCN ice abundance is larger than those of other molecules (such as H2O and H2CO) in this region (see Appendix A). We note that the presence or otherwise of this large peak depends on the relative binding energies of HCN and CO2 assumed in the model (see also Eistrup et al. 2016).

Since previous studies suggested that atmospheric N/O ratio is also a useful tracer of planet formation locations (Piso et al. 2016; Cridland et al. 2020b; Turrini et al. 2021; Ohno & Ueda 2021), we also study the radial distributions of N/O ratios in our disk models. The right panels of Figures 10 and 11 show the radial profiles of N/O ratios at \(t = 10^6 \text{ yr}\). The ices remain, on the whole, dominated by oxygen (i.e., N/O < 0.23). In the non-shadowed disk \((f = 1.0)\) and for molecular initial abundances, the gas-phase N/O ratios increase \((>0.4)\) outside the water snowline; they gradually increase \((0.4-0.5\) and \(\sim 1.0\) outside the H2O and CO2 snowlines, respectively) for the high ionization rate, whereas they are \(<0.5\) for the low ionization rate. This is because of the significant enhancement of CO2 abundances in the disk with a high ionization rate (see Appendix A.1). In the non-shadowed disk \((f = 1.0)\) and for atomic initial abundances, the gas-phase N/O ratios do not increase outside the water snowline, since gas-phase abundances of NH3, HCN, and NH2CHO are at least around an order of magnitude smaller than those for molecular initial abundances (see Appendix A.3). For the high ionization rate, they reach around unity outside the CO2 snowline. This is because CO/CO2 abundance ratios are smaller in the high ionization rate case than those in the low ionization rate case (see Appendix A.1). In addition, the gas-phase N/O ratios exceed unity and reach around \(10\) just outside the CO snowline in all models, because the binding energy of N2 is lower than that of CO (see Table 1), leading to the higher N2 gas.
abundances around these radii (see also, e.g., Turrini et al. 2021).

In the shadowed disk, the radial profiles of the gas-phase N/O ratios show spatial variations that are much larger than those of the C/O ratios, and thus the N/O ratio would be a useful tracer of the shadowed region in the disks. For molecular initial abundances, the gas-phase N/O ratios are much larger than unity in the shadowed region (r ≈ 3–8 au), and they are ≈ 6–9 for the low ionization rate and ≈ 10–20 for the high ionization rate. This is because of the difference in the binding energies between N$_2$ and CO, similar to the situation outside the snowline of the non-shadowed disk. In addition, the icy-phase N/O ratios reach the initial elemental value (∼ 0.23 in our model) since almost all N and O reservoirs freeze out onto dust grains in the shadowed region (r ≈ 3–8 au). For atomic initial abundances, the overall pictures are similar, and around the shadowed region, the gas-phase N/O ratios are >10 and icy-phase N/O ratios reach 0.23. We note that in our results, the icy-phase elemental carbon, oxygen, and nitrogen abundances at r = 5.3 au are ∼ 1.4 × 10$^{-5}$, ∼ 3.2 × 10$^{-4}$, and ∼ (6.5–7.1) × 10$^{-5}$, respectively, which are similar to the initial elemental abundances.

The results for molecular initial abundances and the low ionization rate are consistent with those in Ohno & Ueda (2021), although the gas-phase N/O ratios in the shadowed region are much larger (>100) than that in our disk model (~10). This is likely because they used larger literature values of binding energies of CO and N$_2$ ($E_{\text{des}}$(CO) = 1180 K and $E_{\text{des}}$(N$_2$) = 1051 K) than ours, as the gas-phase N/O ratio is sensitive to those values when both CO and N$_2$ are frozen. In addition, Ohno & Ueda (2021) included several dominant molecules only (i.e., N$_2$ and NH$_3$ only) and did not include the effects of disk chemical evolution. Moreover, they assumed slightly larger N$_2$ abundances (∼ 3.50 × 10$^{-5}$) and smaller NH$_3$ abundances (∼ 7.78 × 10$^{-6}$) than those in our calculations (see Figure 16). Our calculations also investigate the dependence on initial disk conditions.

Therefore, if planets acquire their atmospheres from the gas in the shadowed region (r ≈ 3–8 au in our disk model), they are expected to have the super-stellar N/O ratios of >1 and substellar metallicities. On the other hand, they are expected to have the stellar N/O ratios and super-solar metallicities if the planetary atmospheres are efficiently polluted by solid components (including the case of Jupiter; see Ohno & Ueda 2021).

As discussed in Ohno & Ueda (2021), Saturn, in contrast to Jupiter, may not have the uniform enrichment in the elemental abundances of their atmospheres, if formed in their vicinity of
the current orbits. This is because the current Saturn orbit is outside the shadowed region \((r \sim 3\text{–}8\text{ au})\), and \(\text{N}_2\) has sublimated into the gas phase. In the shadowed disk, the icy-phase \(\text{N}/\text{O}\) ratios at around the current orbit of Saturn \((r \sim 10\text{ au})\) are \(\lesssim 0.1\) for all models, whereas they reach the initial elemental value \((= 0.23\) in our model\) at around the current orbit of Jupiter. In addition, regardless of the initial abundances and ionization rates, the icy-phase \(\text{N}/\text{O}\) ratios at \(r \sim 10\text{–}40\text{ au}\) are \(\lesssim 0.1\) in the shadowed disk, whereas in the non-shadowed disk, they reach the initial elemental value at \(r > 24\text{ au}\) (outside the \(\text{N}_2\) snowline). Future observations and entry probe missions on Saturn and also outer icy planets \((\text{Uranus and Neptune})\) would help to distinguish the shadow formation scenario proposed in Ohno & Ueda \((2021)\) from other scenarios, since the elemental abundances such as \(\text{N}\) in such planets are still uncertain \((\text{see, e.g., Atreya et al.} 2018; \text{Mandt et al.} 2020)\).

Nitrogen abundances within the atmospheres of exoplanets \((\text{such as hot Jupiters})\) can be constrained by the observations of HCN and \(\text{NH}_3\) \((\text{e.g., MacDonald} \text{ & Madhusudhan} 2017; \text{Hawker et al.} 2018; \text{Giacobbe et al.} 2021)\), which are the major nitrogen-bearing species in hot gas-giant atmospheres along with \(\text{N}_2\) \((\text{e.g., Moses et al.} 2011, 2013)\). Through upcoming observations with the James Webb Space Telescope (JWST) and observations with next-generation facilities \((\text{such as ARIEL and ground-based telescopes})\), it is anticipated that the elemental composition of carbon, oxygen, and nitrogen, and the ratios between them will be determined with much higher precision than currently possible for the atmospheres of many exoplanets such as hot Jupiters \((\text{see, e.g., Tinetti et al.} 2018; \text{Madhusudhan} 2019; \text{Changeat et al.} 2020; \text{Turrini et al.} 2022)\).

We note that in addition to \(\text{C}/\text{O}\) ratios, the \(\text{C}/\text{H}\) ratios \((\text{carbon elemental abundances})\) of exoplanetary atmospheres have also been proposed to be a possible tool to link gas-giant exoplanets to their formation sites in the protoplanetary disk \((\text{e.g., Öberg et al.} 2011; \text{Madhusudhan et al.} 2014; \text{Pontoppidan et al.} 2014; \text{Line et al.} 2021; \text{Pelletier et al.} 2021)\). Recently, super-solar and Jupiter-like \(\text{C}/\text{H}\) ratios \((\gtrsim 10^{-3})\) have been confirmed for some hot Jupiters \((\text{e.g., Brogi \& Line} 2019; \text{Gandhi et al.} 2019; \text{Pelletier et al.} 2021)\). Pelletier et al. \((2021)\) analyzed thermal emission spectra of a nontransiting hot Jupiter \(\tau\) Boo b with high spectral resolutions \((R = \lambda/\Delta \lambda = 70000)\) and reported that the planet’s atmosphere has a super-solar \(\text{C}/\text{H}\) ratio and possibly super-solar \(\text{C}/\text{O}\) ratio. We suggest that the super-solar \(\text{C}/\text{H}\) and \(\text{C}/\text{O}\) ratios might be explained if the atmosphere was formed near the inner edge of the shadowed region and polluted by solid components, since the solid \(\text{C}/\text{O}\) ratio exceeds unity there when disk chemistry starts from the atomized initial abundances and the cosmic-ray ionization rate is low \((\text{see Figure} 11; \text{see also Eistrup et al.} 2016)\).

Figures 20, 21, and 22 in Appendix C show, respectively, the radial profiles of \(\text{C}/\text{H}, \text{O}/\text{H}, \text{and N}/\text{H}\) ratios at \(t = 10^6\) yr. In the shadowed disk \((f \lesssim 0.03)\), at \(r \sim 3\text{–}8\text{ au}\) (around the current orbit of Jupiter), the gas-phase \(\text{N}/\text{H}\) ratios \((\sim 10^{-7} \text{–} 10^{-5})\) are
larger than the gas-phase C/H and O/H ratios ($\sim 10^{-8} - 10^{-6}$) in each model, which produce super-stellar N/O ratios of $\gg 1$ (see Figures 10 and 11).

4.2. Implications for the Small Bodies in the Solar System

On the basis of our calculations, in the shadowed region, the snowline positions of molecules with smaller $E_{\text{des}}$ than that of H$_2$O move inward, and even the most volatile species CO and N$_2$ freeze out onto dust grains at around the current orbit of Jupiter ($r \sim 3–8$ au), as found in Ohno & Ueda (2021). In the shadowed region, the dust grains at $r \sim 3–8$ au are expected to have significant (more than $\sim 5–10$ times) amounts of saturated hydrocarbon ices such as CH$_4$ and C$_2$H$_6$, ices of organic molecules (which are mostly saturated) such as H$_2$CO, NH$_2$CHO, CH$_3$OH, and CH$_2$NH$_2$, in addition to H$_2$O, CO, CO$_2$, NH$_3$, N$_2$, HCN, and NH$_2$OH ices, compared with those in the non-shadowed disks (mostly ices of H$_2$O, CO$_2$, NH$_3$, and unsaturated hydrocarbon molecules). In addition, the icy abundances of unsaturated hydrocarbons and other organic molecules in the shadowed disks are much smaller than those in the non-shadowed disks.

Here it is worth discussing whether the presence of the shadowed region influences the chemical composition of small objects, such as primitive comets in the solar system. If primitive comets formed from the icy dust grains in the shadowed region of the disks (such as $r \sim 3–8$ au for $f \leq 0.03$), they are expected to include more saturated hydrocarbons and complex organic molecules such as H$_2$CO, NH$_2$CHO, and CH$_3$OH, and less unsaturated organic molecules than those formed in the non-shadowed disks.

Reactions of radicals, which are mainly formed by cosmic-ray-induced photodissociation of CH$_3$OH ice, are needed to form complex organic molecules in disks (see Sections 3.2.2 and 3.2.4), assuming that they are not already formed in the molecular cloud phases. In addition, Garrod et al. (2008) discussed that (CH$_2$OH)$_2$ (ethylene glycol) is formed on the dust grains around the water snowline from its precursor radical CH$_2$OH, which becomes mobile just as water and other species are beginning to desorb. Moreover, icy grains can efficiently coagulate into larger ($\gg 1$ mm) dust particles and centimeter-sized pebbles outside the water snowline (see Section 2.1 and, e.g., Ros & Johansen 2013; Sato et al. 2016; Drążkowska & Alibert 2017; Pinilla et al. 2017). If CH$_3$OH-rich large dust grains and/or pebbles migrate into the shadowed region (such as around the water snowline), and/or if the shadowed region disappears due to the disk evolution, the ice abundances of various complex organic molecules in such dust grains and pebbles are also expected to increase (see also Section 4.4). This is because the temperature of such dust grains and pebbles increases ($> 30$ K) and radical–radical reactions can proceed efficiently in such warm conditions. Thus in the shadowed disk, both efficient CH$_3$OH ice formation and the formation of ices of further complex organic molecules such as (CH$_2$OH)$_2$ may be realized due to the dissipation of shadowed structures and/or migration inside the shadowed region, and they can be realized without dust grains and/or pebbles migrating vast distances, compared with the non-shadowed disk. We suggest that in shadowed disks, complex organic molecules can be formed in situ rather than being fully inherited from molecular clouds.

To date, various complex organic molecules have been detected in comets (such as Hale-Bopp and 67P/Churyumov-Gerasimenko). Several studies discussed that the molecular abundances in comets are determined by the combination of chemical evolution in the proto-solar disk and inheritance from molecular clouds (e.g., Mumma & Charnley 2011; Caselli & Ceccarelli 2012; Walsh et al. 2014; Eistrup et al. 2016, 2018; Altwegg et al. 2017, 2019; Drozdovskaya et al. 2019; Öberg & Bergin 2021). Walsh et al. (2014) discussed that grain-surface fractional abundances (relative to water ice) for the outer region of the non-shadowed disk ($T < 50$ K) are consistent with

![Figure 12](image_url)
abundances derived for comets, suggesting a grain-surface route to the formation of complex organic molecules observed in cometary comae.

Recently, the abundances of various volatiles and organic molecules (including hydrocarbons) toward comet 67P/Churyumov-Gerasimenko were reported by, e.g., Le Roy et al. (2015), Rubin et al. (2015a, 2015b, 2019, 2020), Altwegg et al. (2017, 2019), and Schuhmann et al. (2019). They reported that the relative abundances of CO, CO₂, and C₂H₆ with respect to H₂O are larger than those of other Jupiter-family comets, which might suggest the formation of cometary grains at lower-temperature regions (such as below 30 K). In addition, Rubin et al. (2015a) reported the first cometary detection of N₂ toward comet 67P, and discussed that the lower N₂/CO abundance ratio (∼3 × 10⁻⁵ based on Rubin et al. 2019, 2020) compared with the protostar value (∼10⁻³) suggested the cometary grains of comet 67P are formed in cold regions with ∼24–30 K. Schuhmann et al. (2019) also reported the existence of unsaturated hydrocarbons in the coma of comet 67P. We note that gas-phase reactions within the CH₄ snowline are needed to efficiently form unsaturated hydrocarbons (see Sections 3.2.2 and 3.2.4).

In Figure 12, we compare the molecular ice abundances (with respect to water ice) of the coma of comet 67P (Rubin et al. 2019, 2020) and our standard disk model calculations (see Sections 3.1 and 3.2). On the basis of Figure 12, the ice abundances of CO, CH₄, and N₂ in the coma of comet 67P cannot be explained by the icy grains within the CH₄ snowline in non-shadowed disks. In addition, as for results in the cold regions where both CO and N₂ are frozen onto dust grains (both in the shadowed and non-shadowed disks), the CO ice abundances are closer to the value of comet 67P than the N₂ ice abundances. These results also suggest that the icy grains of comet 67P were formed between the CO and N₂ snowlines (see above and Rubin et al. 2015a, 2019, 2020). Such cold regions, where the icy grains of comet 67P are formed, can be located at r ∼ a few astronomical unit for the shadowed disks, whereas they are located at r > 20 au for the non-shadowed disk.

In the results of our standard disk model calculations, the CO₂ ice abundances at r ∼ 5.3 au in the shadowed disk are around two orders of magnitude larger than those at r ∼ 30 au in the non-shadowed disk, although the differences in other molecular ice abundances between the two cases are within around one order of magnitude. This is because CO₂ ice formation on the dust grains is not efficient in the coldest regions with T(r) ≤ 20 K (see Section 3.2.1). Thus, the CO₂ ice abundances can be used to distinguish these two regions (the inner cold region (r ∼ 3–8 au, T(r) ∼ 20–30 K) in the shadowed disk and outermost coldest region (r ∼ 30 au, T(r) ≤ 20 K) in the non-shadowed disk) as formation sites of icy dust grains. The CO₂ ice abundance at r ∼ 30 au in the non-shadowed disk is around one order of magnitude smaller than that in comet 67P, whereas that at r ∼ 5.3 au in the shadowed disk is four to seven times larger than that in comet 67P. Thus, for explaining the CO₂ ice abundance in comet 67P, the formation of CO₂ ice at r ∼ 5.3 au in the shadowed disk may be more suitable than that at r ∼ 30 au in the non-shadowed disk. In addition, if we also include the above discussion about the CO and N₂ abundances, we suggest that the icy grains of comet 67P may have been formed at the region with T(r) ∼ 25 K, such as the shadowed region.

For the model results in the shadowed region (see Figure 12), the abundances of ices such as H₂CO and CH₃OH are around one to two orders of magnitude larger than those of comet 67P, whereas those of C₂H₆OH and CH₃COOH are more than one and three orders of magnitude smaller, respectively. We suggest that if the cometary grains of comet 67P are originally from the shadowed region, a temperature rise (due to the inward migration and/or dissipation of the shadowed region) after CH₃OH ice formation in the shadowed region may be needed to trigger radical–radical reactions (see also above) to produce more complex molecules such as C₂H₆OH and CH₃COOH. Future detailed chemical modeling including such physical evolutions may be need needed to construct the formation scenario that explains the ice abundances of all molecules in comet 67P simultaneously.

Bieler et al. (2015) and Rubin et al. (2015b) reported that O₂ ice is abundant in comet 67P and the ratio of O₂/H₂O is a few percent (see also Luspay-Kuti et al. 2022). On the basis of Figure 12, our standard disk model is unable to explain the cometary O₂ ice abundances (see Sections 3.1 and 3.2). According to previous studies (e.g., Eistrup et al. 2016, 2018; Taquet et al. 2016) and our calculations, such higher O₂ ice abundances are possible in disks only if chemical starting conditions were purely atomized and the disk ionization level was low (see Appendix A.1). In addition, O₂ is very volatile and freezes out onto dust grains at r ∼ 2 – 10 au only in the shadowed disk (f ≤ 0.03, see Appendix A.1). We note that Taquet et al. (2016) suggested that O₂ trapping in H₂O ice at earlier evolutionary stages (such as in the molecular clouds) may also be an explanation.

Other small objects such as asteroids will have similar abundances (such as CO and CH₃OH rich), if they are formed from the dust grains in the shadowed region. Yada et al. (2021) reported the results of preliminary analyses of the Hayabusa2 samples returned from C-type asteroid Ryugu, and showed the infrared spectral profile with weak absorptions at 2.7 and 3.4 μm that imply a carbonaceous composition with indigenous aqueous alteration. Kurokawa et al. (2022) compared the infrared spectra (including 3.1 μm absorption features of ammoniated phyllosilicate) of main belt asteroids collected by the AKARI space telescope and their models of water-rock reactions, and suggested that multiple large main belt asteroids formed beyond the NH₃ and CO₂ snowlines and have been transported to their current locations. Fujiya et al. (2019) inferred from the measurements of CO₂ and H₂O abundances and carbon isotope ratios of carbonate minerals in the Tagish Lake meteorite (an carbonaceous chondrite) that at least some D-type asteroids were formed beyond the CO₂ snowline. We note that D-type asteroids are discovered mainly at the outer edge of the main asteroid belt and in the Jupiter Trojan regions (DeMeo & Carry 2014). Tsuchiyama et al. (2021) reported the discovery of primitive CO₂-bearing fluids in an aqueously altered carbonaceous chondrite (one of the primitive meteorites), and discussed that its parent body was formed outside the CO₂ snowline and later transported to the inner solar system.

Since the CO₂ snowline position moves inward in the shadowed disk, the supply of dust grains and small objects with significant amounts of CO₂ and complex organic molecules to the inner region may be relatively easier than that in the non-shadowed disk. We note that if the shadowed region is maintained for a relatively long time (∼10⁶ yr), chemical evolution may produce dust grains and solid objects.
with large amounts of CO$_2$ ice and ices of complex organic molecules such as H$_2$CO and CH$_3$OH.

The solid bodies formed at the Kuiper Belt may be CO$_2$ rich if we consider the disk shadowing effect. This is because at $r > 20$ au, the CO$_2$ ice abundance in the shadowed disk is around one order of magnitude larger than that in the non-shadowed disk (see Section 3.2.1 and Figure 3).

We note that the CO$_2$ snowline position will be important both for chemistry and dust grain growth in the disk. Recent laboratory experiments showed that CO$_2$ ice is less sticky compared to H$_2$O ice (Musiolik et al. 2016a, 2016b; Fritscher & Teiser 2021). Thus, efficient grain growth is only expected between the H$_2$O and CO$_2$ snowlines (Okuzumi & Tazaki 2019; Arakawa & Krijt 2021). If the disk has a shadowed region beyond the water snowline, the region where efficient grain growth can take place might become narrower.

Dartois et al. (2013, 2018) found that ultracarbonaceous Antarctic micrometeorites (UCAMMs) have higher C/Si abundance ratios ($\gtrsim 10^2$) and N/C abundance ratios ($\sim 0.05-0.12$) in organic compounds than those in other primitive meteorites and the local interplanetary dust particles. They discussed that UCAMMs might be formed in a cold nitrogen-rich environment, such as outside the N$_2$ snowline. We suggest that if the disk has a shadowed region beyond the water snowline, the formation sites of UCAMMs may be located in the inner region of the disk, such as $r \sim 3-8$ au.

4.3. Implication for the Observations of Protoplanetary Disks

Many of the complex (organic) molecules described in Sections 3.2.2–3.2.4 have been observed by previous observations with, e.g., ALMA toward hot cores/corinos in high-/low-mass star-forming regions, respectively (see, e.g., Herbst & van Dishoeck 2009; Sakai & Yamamoto 2013; Jørgensen et al. 2020; Yang et al. 2021), and have been detected in the comae of multiple comets (see Section 4.2 and, e.g., Mumma & Charnley 2011; Walsh et al. 2014; Altwegg et al. 2019; Drozdovskaya et al. 2019; Rubin et al. 2020). Lee et al. (2019) reported the detections of CH$_3$CHO and CH$_3$CN, in addition to CH$_3$OH, toward Class I disks around FU Ori–type young star V883 Ori with ALMA. In addition, some of these molecules have been observed toward Class II disks by previous infrared observations (C$_2$H$_2$: e.g., Pontoppidan et al. 2010) and ALMA observations, such as H$_2$CO and CH$_3$OH (e.g., Loomis et al. 2015; Walsh et al. 2016, 2018; Booth et al. 2021a; van der Marel et al. 2021; Guzmán et al. 2021), and C$_2$H$_2$, CH$_3$CN, HC$_5$N, and HCOOH (e.g., see, Q. Li et al. 2013; Öberg et al. 2015; Bergner et al. 2018; Favre et al. 2018; Loomis et al. 2018, 2020; Ilee et al. 2021).

Booth et al. (2021a) reported the first detection of CH$_3$OH in the disk around a Herbig Ae star, HD 100546. They reported that the CH$_3$OH to H$_2$CO abundance ratio is higher ($\gtrsim 10$) at the inner edge of the dust ring in the disk ($\sim 10$–30 au) than those ($\sim 1-2$) in the outer disk ($> 100$ au; Carney et al. 2016) and in the TW Hya disk (Walsh et al. 2016). They discussed that at the inner edge, CH$_3$OH is proposed to originate from thermal desorption. van der Marel et al. (2021) and Brunken et al. (2022) reported the detections of CH$_3$OH, H$_2$CO, and CH$_3$OCH$_3$ line emission in the vicinity of the asymmetric dust trap in the disk around a Herbig Ae star Oph IRS 48 (see also Booth et al. 2021b). They discussed that these molecules are thermally desorbed from icy dust grains and that such dust traps provide huge icy grain reservoirs in the disk midplane.

Recent dust continuum survey observations of protoplanetary disks (with, e.g., ALMA) have shown that disk substructures, most prominently gas-depleted gaps and dust-rich rings, are common (e.g., Andrews et al. 2016, 2018; Andrews 2020; Dullemond et al. 2018; Huang et al. 2018; Isella et al. 2016, 2018; Tsukagoshi et al. 2016). In addition, the recent Molecules with ALMA at Planet-Forming Scales survey obtained spatial distributions (with $\Delta r \sim 10-50$ au) of line emission for CO (Zhang et al. 2021) and some other organic molecules (H$_2$CO, HCN, CH$_3$CN, HC$_3$N, and C$_2$H$_2$: see Guzmán et al. 2021; Öberg et al. 2021; Ilee et al. 2021) toward three T Tauri disks (IM Lup, GM Aur, and AS 209) and two Herbig Ae disks (HD 163296 and MWC 480). They found that dust substructures (such as rings and gaps) are also common in molecular gas emission.

Since a shadowed region can be formed in the midplanes beyond such dust rings/traps of these disks by blocking the radiation from the central star, we propose that the effects of shadows need to be taken into account when we discuss chemical evolution of complex organic molecules in disks with such dust rings and asymmetric traps. Alarcón et al. (2020) calculated the temperature and chemical structures of disks with gas-depleted gaps. They showed that the disk midplane temperature in the gap increases, producing local sublimation of key volatiles, while it decreases in the ring, causing a higher volatile deposition onto the dust grain surfaces. Isella et al. (2018) and Okuzumi et al. (2022) described that the outer wall of the gap produced by a giant planet receives extra starlight heating and puffs up, throwing a shadow across the disk beyond. In addition, Ohashi et al. (2022) suggested from their ALMA observations that the dust clumps in the disk around a Class 0/I protostar create the shadowed region outside, resulting in the sudden drop in temperature. We predict that, as well as the shadowed region beyond the water snowline (discussed in this paper), the freeze-out of molecules onto dust surfaces and efficient formation of organic molecules are also expected to occur in the shadowed region formed by the presence of a giant planet and/or dust clumps.

In this study, we adopt a 1D physical model of the protosolar disk (T Tauri disk) in order to investigate the effects of shadow structures on the chemical evolution of the disk midplane and compare the results with the molecular composition of small objects in the solar system (see also Section 4.2). The influence of disk shadowing is expected to be larger in the disk midplane than in the disk surface, since it is easier for direct stellar light to reach the surface than the midplane (Ueda et al. 2019; Okuzumi et al. 2022). We note that the lines from more rare isotopologues trace deeper regions in the disks, but the disk midplane is obscured if the dust emission is optically thick. In addition to future observations of molecular lines with much higher spatial resolutions toward the disks using, e.g., ALMA and the next-generation Very Large Array (ngVLA), further calculations of disk chemical evolution with 2D physical models of both T Tauri disks and Herbig Ae disks are also important to investigate whether or not the effects of shadows are needed to explain the observed distributions of molecular emission. Since the dust opacities at the frequencies of ngVLA are lower than those of ALMA, observations with ngVLA will be useful to trace the inner molecular gas abundances, such as within 10 au. In addition,
ngVLA expects to resolve the dust emission at centimeter wavelengths with $\Delta r \lesssim 1$ au, and thus will be able to confirm the variation in the dust density profile at the water snowline and presence of the shadowed region beyond the water snowline even in disks around T Tauri stars.\(^{11}\) We note that future calculations of disk chemical evolution with 2D physical models would be also useful to investigate the effects of shadowing on the composition of planetary atmospheres. This is because some recent theoretical studies (e.g., Tamigawa et al. 2012; Morbidelli et al. 2014) and observational studies (e.g., Teague et al. 2019) suggested that the flow of gas into the gap produced by a growing planet is dominated by gas falling vertically from a height of at least one scale height (see, e.g., Crildland et al. 2020a).

$\text{NH}_2^+$ is considered to be a useful probe of the CO and N$_2$ snowlines in disks (e.g., Qi et al. 2013, 2019; Aikawa et al. 2015; van’t Hoff et al. 2017; Murillo et al. 2022), because it is destroyed by proton transfer to CO,

$$\text{N}_2\text{H}^+ + \text{CO} \rightarrow \text{HCO}^+ + \text{N}_2.$$  (19)

Aikawa et al. (2015) suggested from their disk modeling that the $\text{N}_2\text{H}^+$ abundance can have a peak at the temperature slightly below that of CO sublimation, even if the binding energies of CO and N$_2$ are nearly the same. Figure 13 shows the radial profiles of the $\text{N}_2\text{H}^+$ abundances for various values of $f$. The CO freezes out onto dust grains in the shadowed region, and thus $\text{N}_2\text{H}^+$ abundances at $r \sim 3$–8 au become larger in the shadowed disk ($\sim 10^{-15}$ for $f = 0.03$) than those in the non-shadowed disk ($< 10^{-17}$ for $f = 1.0$) by several orders of magnitudes. Thus, $\text{N}_2\text{H}^+$ line emission may potentially trace the shadowed region of the disk beyond the water snowline, although previous observations of $\text{N}_2\text{H}^+$ lines do not resolve the spatial scales around the water snowline ($r \sim 1$–10 au). We note that the $\text{N}_2\text{H}^+$ abundances in the disk midplane are lower than those within the disk surface (Aikawa et al. 2015; van’t Hoff et al. 2017), because of the high densities and low ionization rates. Further chemical and radiative-transfer modeling in the 2D disk structures with shadow structures are needed to confirm whether $\text{N}_2\text{H}^+$ line emission can really be used as an observational tracer of such shadowed regions, since the line emission also traces the vertical distributions of CO and N$_2$ (Qi et al. 2019), and observations of the disk midplane directly are not easy because of the high optical depths at ALMA wavelengths.

4.4. Other Model Caveats

In our disk chemical modeling, we implicitly assumed that ices on grains are formed by homogeneous layers regardless of their composition or crystallinity, as most astrochemical models of disks assume. However, recently, Kouchi et al. (2021) suggested from their transmission electron microscopy studies that the macroscopic morphology of icy dust grains are as follows: amorphous H$_2$O covered the refractory grain uniformly, CO$_2$ nano-crystals were embedded in the amorphous H$_2$O, and a polyhedral CO crystal is attached to the amorphous H$_2$O. Such morphology of dust grains would affect the chemical evolution in the disks, since the binding energies and nonthermal desorption rates depend on the chemical composition of the ice mantles on the dust grains (e.g., Bertin et al. 2016; Cuppen et al. 2017; Penteado et al. 2017).

Heinzeller et al. (2011) investigated the effects of physical mass transport phenomena in the radial direction by viscous accretion and in the vertical direction by diffusive turbulent mixing and disk winds. They showed that the gas-phase molecular abundances of such as H$_2$O and CH$_3$OH are enhanced in the warm surface layer due to the effects of vertical mixing. van der Marel et al. (2021) discussed that this vertical transport may be important to explain the observed abundance of CH$_3$OH in the disk around IRS 48, in addition to icy dust concentrations at the dust trap. In addition, we assume that the disk physical structure is steady and that the shadow structure is maintained for $10^6$ yr. We note that the inward migration of solids and/or destruction of the shadow structure after the efficient formation of CH$_3$OH ice in the shadowed region may increase the abundances of various complex organic molecules in the inner disks around the water snowline (see also Section 4.2). Future observations of organic molecular lines with, e.g., ngVLA and infrared telescopes (such as JWST and GREX-PLUS) are expected to constrain the abundances in such inner warm regions.

5. Conclusions

In this study, we investigated the radial abundance distributions of dominant carbon-, oxygen-, and nitrogen-bearing molecules and the radial distributions of elemental abundance ratios (C/O and N/O ratios) in the gas and ice of disks with shadow structures. We used a detailed gas-grain chemical reaction network and calculated chemical structures in the shadowed disk midplane around a T Tauri star (a protosolar-like star). Gas-phase reactions, thermal and nonthermal gas-grain interactions, and grain-surface reactions were included in our adopted reaction network. We investigated the dependence of the disk chemical structures on ionization rates and initial abundances. We discussed the effects of disk shadowing on chemical evolution of complex organic molecules and forming planetary atmospheres. We also compared the results of our calculations with the molecular composition of small bodies in the solar system (such as

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\(^{11}\) See ngVLA-J memo series, https://ngvla.nao.ac.jp/researcher/memo/.
comets and asteroids) and recent observational results of protoplanetary disks. Our findings can be summarized as follows:

1. In the shadowed disks \((f \leq 0.03)\), the snowline positions of molecules with smaller \(E_{\text{les}}\) than that of \(H_2O\) move inward, and even the most volatile species, \(CO\) and \(N_2\), freeze out onto dust grains at around the current orbit of Jupiter \((r \sim 3–8 \text{ au})\). Our detailed calculations confirm the results of Ohno & Ueda (2021) who showed the freeze-out of \(CO\) and \(N_2\) onto dust grains in the shadowed region with using more simplified calculations. Meanwhile, we newly find that the dust grains within the shadowed regions have significant (more than 5–10 times) amounts of saturated hydrocarbon ices such as \(CH_4\) and \(C_2H_6\), ices of organic molecules such as \(H_2CO\), \(NH_2CHO\), and \(CH_3OH\), in addition to \(H_2O\), \(CO\), \(CO_2\), \(NH_3\), \(N_2\), \(HCN\), and \(NH_2OH\) ices, compared with those in the non-shadowed disks (mostly ices of \(H_2O\), \(CO_2\), \(NH_3\), and unsaturated hydrocarbon molecules). We note that these abundant saturated hydrocarbons and various organic molecules such as \(H_2CO\), \(NH_2CHO\), and \(CH_3OH\) were not reported by Ohno & Ueda (2021), since they approximated various organic molecules by considering \(C_2H_6\) alone.

2. The icy abundances of \(C_2H_2\), \(HCN\), \(NH_2OH\), and complex organic molecules (which are mostly saturated) such as \(H_2CO\), \(CH_3OH\), \(NH_2OH\), \(NH_2CHO\), and \(CH_3NH_2\) at \(r \sim 3–8 \text{ au}\) in the shadowed disks are enhanced compared with the values in the non-shadowed disk, although the snowline positions of these molecules are \(r \lesssim 3 \text{ au}\) regardless of the values of \(f\). We concluded that the sequential hydrogenation reactions of (especially) \(CO\) on the cold dust grains play a vital role in efficiently forming these molecules, in addition to reaction pathways starting from the cosmic-ray-induced photodissociation of \(CH_3OH\). We found that if the shadowed region is maintained for a relatively long time \((t \sim 10^6 \text{ yr})\), chemical evolution may produce dust grains and solid objects with large amounts of \(CO_2\) and organic molecular ices (see also Appendix B).

3. The icy abundances of unsaturated hydrocarbons such as \(C_2H_2\), \(C_2H_4\), \(C_3H_2\), \(C_3H_4\), \(C_2H_6\), and \(HC_3N\) at \(r \sim 3–8 \text{ au}\) (just outside their snowlines) are much smaller than those of a non-shadowed disk, since gas-phase chemical reactions especially within the \(CH_4\) snowline \((\sim 15 \text{ au}\) for \(f = 1.0\) and \(\lesssim 2 \text{ au}\) for \(f \leq 0.03\)) mainly drive the formation of these molecules. Moreover, the icy abundances of unsaturated complex organic molecules such as \(HCOOCH_3\), \(HCOOH\), and \(CH_3COOH\) outside the snowlines are much smaller than those of a non-shadowed disk, since grain-surface association of large radical–radical reactions in the warm regions \((T(r) \sim 50 \text{ K})\) is needed for the formation of these molecules. In addition, the icy abundances of other molecules such as \(CH_3CN\), \(CH_3CHO\), \(CH_3OCH_3\), and \(C_2H_5OH\) are larger for \(f = 0.03\) than those for \(f = 1.0\), 0.3, and 0.003, since they are determined by the combinations of radical–radical reactions and gas-phase reactions/radical formation reactions from \(CO\) and/or \(CH_3OH\) on the dust grain surfaces.

4. We also studied the impacts of different ionization rates and initial chemical abundances on disk chemical structures in Appendix A. We found that \(CO/CO_2\) abundances become smaller/larger with increasing ionization rates, respectively. In addition, \(CH_4\) and \(C_2H_6\) gas abundances within their snowline become smaller as the ionization rates become larger. Abundances of \(H_2O\) and organic molecules are larger for molecular initial abundances than those for atomic initial abundances. Moreover, \(O_2\) ice abundances are \(\sim 10^{-3} – 10^{-4}\) (consistent with the measured cometary abundances) only for atomic initial abundances and the low ionization rates. \(O_2\) is very volatile and freezes out onto dust grains at \(r \sim 2–10 \text{ au}\) only in the shadowed disk.

5. In the shadowed region, the gas-phase \(C/O\) ratios are almost unity for the high ionization rate and/or molecular initial abundances and 0.8 for the low ionization rate and atomic initial abundances. In addition, the radial profiles of the gas-phase \(N/O\) ratios show spatial variations that are much larger than those of the \(C/O\) ratios, and thus the \(N/O\) ratio would be a useful tracer of the shadowed regions of disks. For molecular initial abundances, the gas-phase \(N/O\) ratios are much larger than unity in the shadowed region, and they are \(\sim 6–9\) for the low ionization rate and \(\sim 10–20\) for the high ionization rate. The icy-phase \(C/O\) and \(N/O\) ratios reach the initial elemental values in the shadowed region. Therefore, if the planets acquire their atmospheres from the gas in the shadowed region \((r \sim 3–8 \text{ au}\) in our disk model), they are expected to have super-stellar \(N/O\) ratios of \(\gg 1\), super-stellar \(C/O\) ratios of around unity in most cases, and substellar metallicities. In contrast, they are expected to have stellar \(N/O\) and \(C/O\) ratios and super-solar metallicities if the planetary atmospheres are efficiently polluted by solid components (including the case of Jupiter; see Ohno & Ueda 2021). Upcoming and future observations by JWST and ARIEL will constrain such elemental ratios precisely for atmospheres of many exoplanets such as hot Jupiters.

6. We discussed whether the presence of a shadowed region influences the chemical composition of small objects, such as primitive comets and asteroids in the solar system. Recently the abundances of various volatiles and organic molecules toward comet 67P/Churyumov-Gerasimenko were reported, and some of the results (including \(CO\), \(CO_2\), \(C_2H_6\), \(N_2\), and \(O_2\) abundances) would suggest the formation of cometary grains in lower-temperature regions with \(T \sim 25 \text{ K}\), which can be found in the shadowed region.

7. We propose that \(CO_2\) ice abundances can be used to distinguish the inner cold region \((r \sim 3–8 \text{ au}\), \(T(r) \sim 20–30 \text{ K}\)) in the shadowed disk and the outermost coldest region \((r \sim 30 \text{ au}\), \(T(r) \lesssim 20 \text{ K}\)) in the non-shadowed disk as formation sites of icy dust grains. Moreover, if \(CH_3OH\)-rich large dust grains and/or pebbles migrate inside the shadowed region (such as around the water snowline), and/or if the shadowed region disappears due to the disk evolution, the ice abundances of various complex organic molecules are also expected to increase. This is because radical–radical reactions may be able to proceed efficiently due to the rapid heating.

8. \(N_2H^+\) line emission could potentially trace the shadowed region of the protoplanetary disks beyond the water
snowline, although further modeling and observations with much higher spatial resolution than currently conducted are needed.

On the basis of our calculations, we conclude that a shadowed region allows the recondensation of key volatiles onto dust grains, and may explain to some degree the trapping of icy molecules such as CO$_2$, CO, N$_2$, and O$_2$ in the dust grains that formed comet 67P/Churyumov-Gerasimenko. In addition, in a shadowed disk, Jupiter need not have migrated vast distances to explain its atmospheric composition. In the formation of ices of further complex organic molecules such as CH$_3$OH driven by rapid heating may be realized without dust grains and/or pebbles migrating vast distances, compared with the non-shadowed disk. Thus, we propose that in the shadowed disks, various complex organic molecules can be formed in situ rather than being fully inherited from molecular clouds. Further chemical modeling (such as using 2D physical models of both T Tauri disks and Herbig Ae disks) and comparison with observations of disks (using, e.g., ALMA and ngVLA), planetary atmospheres, and small objects (comets and asteroids) in the solar system will further constrain the effects of disk shadowing on chemical evolution.

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Software: RADMC-3D (Dullemond et al. 2012).

Appendix A

The Dependence of the Disk Chemical Structures on Ionization Rates and Initial Abundances

In this Appendix A, we explain the dependence of the disk chemical evolution on disk ionization rates and initial abundances. Figures 14–16 show the radial profiles of fractional abundances at $t=10^6$ yr for dominant oxygen-, carbon-, and nitrogen-bearing molecules (H$_2$O, CO, CO$_2$, O$_2$, CH$_4$, C$_2$H$_6$, H$_2$CO, CH$_3$OH, N$_2$, NH$_3$, HCN, and NH$_2$CHO) in the shadowed and non-shadowed disk midplane ($f=1.0$ and $f=0.03$, respectively). In these Figures, we assume either molecular or atomic initial abundances and either low or high ionization rates ($\xi_{CR}(r) = 10^{-18}$, $10^{-17}$ [s$^{-1}$]). We note that Figures 3–5 in Section 3.2 show the radial abundance profiles for the same molecules with molecular initial abundances and $\xi_{CR}(r) = 10^{-17}$ [s$^{-1}$].

A.1. H$_2$O, CO, CO$_2$, and O$_2$

Figure 14 shows the radial profiles of fractional abundances for dominant oxygen-bearing molecules: H$_2$O, CO, CO$_2$, and O$_2$. For molecular initial abundances, the H$_2$O ice abundances outside the water snowline are $\sim (1-2) \times 10^{-4}$ (see also Section 3.2.1), which are similar to the assumed value of the initial H$_2$O ice abundance (1.984 $\times 10^{-4}$) for atomic initial abundances, they decrease outside the water snowline, and at $r>2-10$ au they are $\sim 10^{-6}$ $\times 10^{-5}$ for $f=1.0$ and $\sim (2-5) \times 10^{-5}$ for $f<0.03$. In addition, just outside the H$_2$O snowline, the H$_2$O ice abundances are about an order of magnitude larger in the high ionization case ($\sim 10^{-5}$) than those in the low ionization case ($\sim 10^{-6}$). Eistrup et al. (2016) described that in the case of atomic initial abundances, the H$_2$O ice is not efficiently produced in the outer disk, and the ion-molecule reactions in the gas phase (Hollenbach et al. 2009) are contributing to the formation of water. Previous studies (e.g., Schmalzl et al. 2014; Notsu et al. 2021; van Dishoeck et al. 2021) discussed that it takes more than 1 Myr of the pre-stellar phase to produce water ice with an abundance of $\gtrsim 10^{-4}$.

For molecular initial abundances, the total CO (gas+ice) abundances outside the water snowline are $\sim 10^{-4}$ for the low ionization rate and $\sim (3-6) \times 10^{-5}$ for the high ionization rate. This decrease of CO coincides with the overall enhancement of CO$_2$ ice outside the CO$_2$ snowline. Under the high ionization rate, CO destruction pathways of Equations (6) (icy phase) and (7) (gas phase) are efficient (see Section 3.2.1). In the shadowed disk ($f=0.03$), CO freezes out onto dust grains at around the current orbit of Jupiter ($r<3-8$ au), and for atomic initial abundances and the high ionization rate, a larger decrease of total CO abundance ($\lesssim 10^{-5}$) beyond the CO$_2$ snowline is shown.

CO$_2$ gas abundances between the H$_2$O and CO$_2$ snowlines and CO$_2$ ice abundances outside CO$_2$ snowline are larger for the high ionization rate than those for the low ionization rate. In addition, they are also larger for atomic initial abundances than those for molecular initial abundances. For atomic initial abundances and the high ionization rate, CO$_2$ ice abundances are larger ($\gtrsim 10^{-4}$) outside the CO$_2$ snowline, with the decreases of CH$_4$ gas and total (gas+ice) CO abundances.

We note that in the non-shadowed disk ($f=1.0$), CO$_2$ ice abundances at $r>30$ au are significantly smaller ($<10^{-6}$). This is because this region has the coldest conditions ($T(r) \lesssim 20$ K), and the formation of H$_2$O ice (Equation (5)) is faster than that of CO$_2$ ice (Equation (6); Eistrup et al. 2016).

The gas-phase O$_2$ abundances are approximately two orders of magnitude larger for atomic initial abundances than those for molecular initial abundances. The O$_2$ gas abundances between the water and O$_2$ snowlines are $\sim 10^{-5}$ $\sim 10^{-4}$ for atomic initial abundances and $<10^{-6}$ for molecular initial abundances. This is because in the gas-phase O$_2$ is formed from atomic oxygen via Equation (8) (Walsh et al. 2015; Taquet et al. 2016; Eistrup...
Figure 14. The radial profiles of fractional abundances with respect to total hydrogen nuclei densities $n_X/n_H$ at $t = 10^6$ yr for H$_2$O (top panels), CO (second-row panels), CO$_2$ (third-row panels), and O$_2$ (bottom panels). The left panels show the results for the disk midplane with the monotonically decreasing density and temperature profile ($f = 1.0$), and the right panels show the results for the shadowed disk midplane ($f = 0.03$). In each panel, the dashed and solid lines show the profiles for gaseous and icy molecules, respectively. The green and blue lines show the results for molecular initial abundances (the “inheritance” scenario), and the red and black lines show the results for atomic initial abundances (the “reset” scenario). The green and red lines show the results for the higher cosmic-ray ionization rate $\xi_{\text{CR}}(r) = 1.0 \times 10^{-17}$ [s$^{-1}$], whereas the blue and black lines show the results for the lower cosmic-ray ionization rate $\xi_{\text{CR}}(r) = 1.0 \times 10^{-18}$ [s$^{-1}$].
Figure 15. Same as Figure 14, but for CH₄ (top panels), C₂H₆ (second-row panels), H₂CO (third-row panels), and CH₃OH (bottom panels).
Figure 16. Same as Figure 14, but for N$_2$ (top panels), NH$_3$ (second-row panels), HCN (third-row panels), and NH$_2$CHO (bottom panels).
& Walsh 2019), and the initial atomic oxygen abundance is $3.2 \times 10^{-4}$ for atomic initial abundances whereas it is zero for molecular initial abundances. In the shadowed disk ($f \leq 0.03$), O$_2$ freezes out onto dust grains at $r \approx 2-10$ au, and returns to the gas phase at $r > 10$ au. O$_2$ ice is mainly formed from atomic oxygen via Equation (10) (Taquet et al. 2016; Eistrup & Walsh 2019). For the high ionization rate, O$_2$ ice photodissociation becomes efficient and the released oxygen is contained in other major oxygen-bearing molecules (Eistrup et al. 2016; Taquet et al. 2016). Thus, O$_2$ ice abundances are low ($\leq 10^{-6}$) for molecular initial abundances and/or the high ionization rates, and they are $\sim 10^{-5} - 10^{-4}$ only for atomic initial abundances and the low ionization rates, which is consistent with the results of previous studies (e.g., Eistrup et al. 2016; Eistrup & Walsh 2019; Taquet et al. 2016). Such larger abundances are consistent with the measured cometary abundances (Bieler et al. 2015; Rubin et al. 2015b; see also Section 4.2 of this paper).

According to the right-hand panels in Figure 14, both O$_2$ and H$_2$O ice abundances are similar ($\sim$$(3-5) \times 10^{-5}$) in the shadowed region ($r \sim 2-10$ au) in the disk with atomic initial abundances and low ionization rates. As we explain in Section 4.2, the ratio of O$_2$/H$_2$O in comet 67P is a few percent (Bieler et al. 2015; Rubin et al. 2015b). Thus in reality, the models between our assumed parameters (such as partially inherited initial abundances and/or the ionization rates of $\sim 5 \times 10^{-15}$) might well reproduce the observed O$_2$ and H$_2$O ice abundance ratios.

A.2. Other Dominant Carbon-bearing Molecules

For molecular initial abundances, the CH$_4$ and C$_2$H$_6$ gas abundances within their snowline become smaller as the ionization rate becomes larger. This is because CH$_2$ and C$_2$H$_6$ gases are not efficiently formed in these regions, and they are destroyed by cosmic-ray-induced photodissociation and ion-molecule reactions (such as CH$_4$+C$^\bullet$). Thus, carbon is converted from CH$_4$ gas to CO$_2$, H$_2$CO, and hydrocarbons such as, e.g., C$_2$H$_2$, C$_2$H$_4$, C$_3$H$_2$, and C$_3$H$_4$ (Aikawa et al. 1999; Eistrup et al. 2016, 2018; Yu et al. 2016; see also Sections 3.2.2 and 3.2.4). In addition, for atomic initial abundances, the CH$_4$ and C$_2$H$_6$ gas abundances within the CH$_4$ snowline are much smaller ($\ll 10^{-5}$) than those for molecular initial abundances. For atomic initial abundances, the CH$_4$ and C$_2$H$_6$ ice abundances outside their snowlines ($<10^{-6}$ for CH$_4$ ice and $<10^{-6}$ for C$_2$H$_6$ ice) are much smaller than those for molecular initial abundances ($\sim 10^{-3}$ for CH$_4$ ice and $\geq 10^{-6}$ for C$_2$H$_6$ ice).

For molecular initial abundances, H$_2$CO and CH$_3$OH ice abundances between their snowlines and the CO snowline are larger for the low ionization rate ($\sim 2 \times 10^{-7}$ for CH$_3$OH ice) than those for the high ionization rate ($\sim 5 \times 10^{-7}$ for CH$_3$OH ice). We note that cosmic-ray photodissociation of these molecules is efficient in these regions and produces many radicals, which creates more complex organic molecules (see Section 3.2.2). In addition, H$_2$CO and CH$_3$OH ice abundances outside the CO snowline (including the shadowed region) are smaller for the low ionization rate ($\geq 10^{-6}$ for CH$_3$OH ice) than those for the high ionization rate ($\geq 3 \times 10^{-6}$ for CH$_3$OH ice). This is because H$_2$CO and CH$_3$OH ice are mainly formed by the sequential hydrogenation of CO on the grain surfaces (see Equations (10) and (11)), and the atomic hydrogen needed to hydrogenate CO is an end product of H$_2$ ionization (e.g., Schwarz et al. 2018).

For atomic initial abundances, both H$_2$CO and CH$_3$OH ice abundances at $r \sim 2-10$ au significantly decrease. The H$_2$CO ice abundances at such radii are $\ll 10^{-15}$ in the non-shadowed disk ($f = 1.0$) and $\sim (1-6) \times 10^{-6}$ in the shadowed disk ($f \leq 0.03$). The CH$_3$OH ice abundances are $\ll 10^{-8}$ in the non-shadowed disk ($f = 1.0$) and $\sim (1-3) \times 10^{-7}$ in the shadowed disk ($f \leq 0.03$). Schwarz et al. (2018) discussed that the timescales needed to build a reservoir of CH$_3$OH ice are longer than those needed for CO$_2$, because CO$_2$ can form directly from CO ice (see Equation (6)) while CH$_3$OH ice is formed via a series of hydrogenation reactions.

A.3. Dominant Nitrogen-bearing Molecules

Unlike the case for CO, N$_2$ gas and ice abundance profiles barely change for different initial abundances and ionization rates.

For molecular initial abundances, NH$_3$ gas and ice abundances around the NH$_3$ snowline decrease with increasing the ionization rate. This is because NH$_3$ is converted into N$_2$ both through ion-molecule reactions as well as through cosmic-ray-induced photoreactions (Eistrup et al. 2016). In addition, for molecular initial abundances, NH$_3$ ice abundances at $r > 5$ au are $\sim 1 \times 10^{-5}$. In contrast, for atomic initial abundances, NH$_3$ ice abundances at $r > 5$ au decrease ($\sim 10^{-6} - 10^{-5}$ for the high ionization rate and $\sim 10^{-7} - 10^{-6}$ for the low ionization rate). Schwarz & Bergin (2014) and Eistrup et al. (2016) noted that under atomic initial abundances, atomic N quickly forms N$_2$ gas.

Between the HCN and CO snowlines, HCN ice abundances are larger for molecular initial abundances ($\sim 10^{-6} - 10^{-5}$) than those for atomic initial abundances ($\sim 10^{-7} - 10^{-6}$). In contrast, HCN ice abundances just outside the CO snowline (including the shadowed region) are similar ($\sim (3-5) \times 10^{-6}$) both for the atomic and molecular initial abundances. This is because HCN is formed through the gas-phase reaction of HCO with an N atom, with subsequent freeze-out onto dust grains, where HCO is formed by hydrogenation of CO on the dust grain surface (see Section 3.2.3 and, e.g., Aikawa et al. 1999; Eistrup et al. 2016).

Between the NH$_2$CHO and CO snowlines, NH$_2$CHO ice abundances are larger for molecular initial abundances ($\sim 10^{-6} - 10^{-5}$ at $r \sim 1-5$ au) than those for atomic initial abundances ($\sim 10^{-9} - 10^{-8}$ at $r \sim 1-5$ au). Just outside the CO snowline (including the shadowed region), NH$_2$CHO ice abundances for molecular initial abundances are around $10^{-6} - 10^{-5}$. In addition, NH$_2$CHO ice abundances for atomic initial abundances significantly increase with increasing the ionization rate. We find that NH$_2$CHO ice can form via the subsequent hydrogenation of OCN in the cold region (see Equation (13) in Section 3.2.3, and Garrod et al. 2008; Walsh et al. 2014; López-Sepulcre et al. 2015), and the atomic hydrogen is an end product of H$_2$ ionozation.

Appendix B

Time Evolution of Molecular Abundances

Figures 17–19 show the time evolution of the radial profiles of fractional abundances for dominant oxygen-, carbon-, and nitrogen-bearing molecules (H$_2$O, CO, CO$_2$, O$_2$, CH$_4$, C$_2$H$_6$, H$_2$CO, CH$_3$OH, N$_2$, NH$_3$, HCN, and NH$_3$CHO) in the
Figure 17. The time evolution of the radial profiles of fractional abundances with respect to total hydrogen nuclei densities for H$_2$O (top panels), CO (second-row panels), CO$_2$ (third-row panels), and O$_2$ (bottom panels). The dashed and solid lines show the profiles for gaseous and icy molecules, respectively. The green, blue, red, and black lines show the profiles for different evolutional stages ($t = 10^4, 10^5, 5 \times 10^5$, and $10^6$ yr), respectively. These panels show the results when assuming molecular initial abundances (the “inheritance” scenario) and for $\zeta_{\text{CR}}(r) = 1.0 \times 10^{-17}$ [s$^{-1}$]. The left panels show those for the disk midplane with the monotonically decreasing temperature and density structure ($f = 1.0$), and the right panels show the results for the shadowed disk midplane ($f = 0.03$).
Figure 18. Same as Figure 17, but for CH$_4$ (top panels), C$_2$H$_6$ (second-row panels), H$_2$CO (third-row panels), and CH$_3$OH (bottom panels).
Figure 19. Same as Figure 17, but for N$_2$ (top panels), NH$_3$ (second-row panels), HCN (third-row panels), and NH$_2$CHO (bottom panels).
shadowed and non-shadowed disk midplane (f =1.0 and f =0.03, respectively), for molecular initial abundances and a high ionization rate (ξCR(r)=10^{-17}[s^{-1}]). These initial conditions are the same as those in Sections 3.1 and 3.2.

According to these Figures, H2O and N2 abundances do not change with time (see Table 2). O2 gas abundances inside its snowline increase with time (from ≲10^{-9} at t ≤ 10^{5} yr to ∼10^{-6} at t ∼ 10^{6} yr).

Both CO gas and ice abundances (inside and outside the CO snowline, respectively) decrease with time (from ∼10^{-4} at t ≤ 10^{5} yr to ∼(3-5) × 10^{-5} at t ∼ 10^{6} yr). In contrast, both CO2 gas and ice abundances (inside and outside the CO2 snowline, respectively) increase with time (from ≲10^{-5} at t < 10^{5} yr to ∼5 × 10^{-5} at t ∼ 10^{6} yr). Eisipur et al. (2018) discussed that the decreasing abundance of CO gas and increasing abundance of CO2 ice with time between CO2 and CO snowlines can be explained by CO2 gas collisions with the grains, followed by fast reactions with OH (faster than CO can desorb) that produce CO2 on dust grains. Bosman et al. (2018) discussed that for ξCR(r)=1.0 × 10^{-17}[s^{-1}], the CO abundance can be reduced by two orders of magnitude in t ∼(2-3) × 10^{6} yr.

As previous studies have shown (see, e.g., Eisipur et al. 2018), CH4 gas abundances within its snowline decrease with time (from ≳5 × 10^{-6} at t ≲ 10^{5} yr to ≲10^{-7} at t ≳ 5 × 10^{5} yr) and convert to CO2 (see Section 3.2.2). In addition, we show that abundances of other molecules such as C2H6, H2CO, CH3OH, and NH3 decrease with time around their own snowlines. We predict that in these regions, gas-phase destruction processes (ion-molecule reactions and cosmic-ray-induced photodissociation) and cosmic-ray-induced photodesorption/photodissociation overcome the production processes in gas and on the dust grain surfaces, respectively.

In the region where CO is frozen out onto dust grains (including the shadowed region for f ≲ 0.03), the abundances of, e.g., H2CO, CH3OH, and NH2CHO increase with time because of the grain-surface reactions (from ≲10^{-6} at t ≤ 10^{5} yr to ≥5 × 10^{-6} at t ∼ 10^{6} yr for CH3OH). On the basis of these results, the timescales of gas-phase reactions (such as ion-molecule reactions and cosmic-ray-induced photodissociation) and cosmic-ray-induced photodesorption/photodissociation are generally shorter than those of the grain-surface formation reactions, such as sequential hydrogenation reactions.

The HCN ice abundances are slightly enhanced (from ∼3 × 10^{-6} at 10^{4} yr to ∼5 × 10^{-6} at 10^{5} yr) in the region where CO is frozen out onto dust grains (including the shadowed region for f ≲ 0.03). As we describe in Section 3.2.3, HCN ice is efficiently formed within the cold regions where CO freezes out onto dust grains. Schwarz & Bergin (2014) and Eisipur et al. (2018) described that HCN ice is efficiently produced by a few megayears.

On the basis of these results, in the shadowed region (r ∼3−8 au) the ice abundances of CO2 and organic molecules such as H2CO, CH3OH, and NH2CHO become larger with time. Thus, we find that if the shadowed region is maintained for a relatively long time (t ∼ 10^{7} yr), chemical evolution may produce dust grains and solid objects with large amounts of CO2 and organic molecular ices (see also Section 4.2). In addition, the ice abundances of these molecules can be a clue in constraining the formation age of solid bodies in the shadowed region.

We note that in our modeling, we assume a constant central stellar luminosity and a constant viscous accretion heating (see Section 2.1). Kunimoto et al. (2021) calculated that in the case of a T Tauri star with mass M* =1.0M_{⊙}, the central star luminosity L_{*} gradually decreases from ∼5L_{⊙} at t ∼10^{7} yr to L_{*}c =10^{3} at t ≥3 × 10^{6} yr. In addition, the viscous accretion heating determines the temperature in the inner disk (see Section 2.1). Thus, the disk temperature in the inner disk (around the water snowline) does not change very much within the timescale of 10^{6} yr by evolving stellar luminosity.

In addition, in reality the efficiency of viscous accretion heating also changes with time. Oka et al. (2011) described the relation between the midplane temperature and mass accretion rate: T ∝ M^{1/4}. Thus, based on Oka et al. (2011), the water snowline position expects to move inward (from ∼5 au to ∼1.3 au) as the mass accretion rate becomes around 10 times smaller with time (from t ∼ 10^{3} yr (Class 0-I disk) to t ∼ 10^{6} yr (Class II disk)).

Appendix C

The C/H, O/H, and N/H Ratios

Figures 20, 21, and 22 show, respectively, the radial profiles of C/H, O/H, and N/H ratios at t ∼10^{6} yr. Differently colored lines show the profiles for different values of the parameter f (=1.0, 0.3, 0.03, and 0.003), respectively. As described in our previous studies for the case of non-shadowed disks (see Section 2.3 and Figure 2 of Notsu et al. 2020), the inclusion of chemistry has a significant impact on the disk elemental abundances of both gas and ice.

In the non-shadowed disk (f =1.0), for ξCR(r)=1.0 × 10^{-18}[s^{-1}], the gas-phase C/H and O/H ratios are, respectively, ∼10^{-4} and ∼(1−3) × 10^{-4} at r < 22 au (within the CO snowline). In addition, for ξCR(r)=1.0 × 10^{-17}[s^{-1}], the gas-phase C/H and O/H ratios are, respectively, ∼10^{-4} and ∼10^{-4} within the CO2 snowline and, respectively, ∼10^{-6} × 5 × 10^{-5} and ∼10^{-5} − 10^{-4} between the CO2 and CO snowlines. This is because under the high ionization rate (ξCR(r)=1.0 × 10^{-17}[s^{-1}]), CO and CH4 gas are converted to less volatile molecules, such as CO2, H2CO, and hydrocarbons (see Sections 3.2.1 and 3.2.2). Outside the CO snowline (r > 22 au), both gas-phase C/H and O/H abundances are ≲10^{-7}. The icy-phase C/H and O/H ratios increase with increasing r, and reach a similar value to the initial elemental abundance outside the CO snowline.

In the non-shadowed disk (f =1.0), the gas-phase N/H ratios are ∼(5−7) × 10^{-5} at r < 24 au (within the N2 snowline). Outside the N2 snowline (r > 24 au), the gas-phase N/H abundances are ≲10^{-7}. The icy-phase N/H ratios increase with increasing r, and reach a similar value to the initial elemental abundance outside the N2 snowline.

In the shadowed disk (f ≲ 0.03), at r ∼3−8 au (around the current orbit of Jupiter), the gas-phase C/H and O/H ratios are both ∼10^{-8} − 10^{-6}. In addition, at r ∼3−8 au (around the current orbit of Jupiter), the gas-phase N/H ratios (10^{-7} − 10^{-5}) are larger than the gas-phase C/H and O/H ratios in each model, which produce super-stellar N/O ratios of ≫1 (see Figures 10 and 11). The icy-phase C/H, O/H, and N/H ratios reach similar values to the initial elemental abundances.
Figure 20. The radial profiles of C/H ratios at $t=10^6$ yr. The dashed and solid lines show the profiles for gaseous and icy molecules, respectively. The black, red, blue, and green lines show the profiles for different values of the parameter $f$ ($=1.0, 0.3, 0.03, \text{ and } 0.003$), respectively. The top panels show the results when assuming molecular initial abundances (the "inheritance" scenario), whereas the bottom panels show those when assuming atomic initial abundances ("reset" scenario). The left panels show the results for $\xi_{\text{CR}}(r)=1.0 \times 10^{-17}$ [s$^{-1}$], whereas right panels show the results for $\xi_{\text{CR}}(r)=1.0 \times 10^{-18}$ [s$^{-1}$]. The initial elemental C/H ratio is $1.40 \times 10^{-4}$. 

\[ \text{Initial elemental C/H ratio is } 1.40 \times 10^{-4}. \]
Figure 21. Same as Figure 20, but for the radial profiles of O/H ratios at $t=10^6$ yr. The initial elemental O/H ratio is $3.20 \times 10^{-4}$. 
Figure 22. Same as Figure 20, but for the radial profiles of N/H ratios at $t = 10^6$ yr. The initial elemental N/H ratio is $7.50 \times 10^{-5}$. 
