Observation of the A Carinid Meteor Shower 2020 Unexpected Outburst

Juan Sebastian Bruzzone1,2, Robert J. Weryk2,3, Diego Janches2, Carsten Baumann4, Gunter Stober5 and Jose Luis Hormaechea6,7

1 Department of Physics, Catholic University of America, 620 Michigan Avenue NE, Washington, DC 20064, USA; juan.s.bruzzone@nasa.gov
2 ITM Physics Laboratory, NASA/Goddard Space Flight Center, 8800 Greenbelt Road, Greenbelt, MD 20771, USA
3 Institute for Astronomy, University of Hawaii, 6860 Woodlawn Drive, Honolulu, HI 96822, USA
4 Deutsches Zentrum für Luft- und Raumfahrt, Institut für Solar-Terrestrielle Physik, Neustrelitz, Germany
5 Institute of Applied Physics & Oeschger Center for Climate Change Research, Microwav Physics, University of Bern, Bern, Switzerland
6 Facultad de Ciencias Astronomicas y Geofisicas, Universidad Nacional de La Plata, Argentina
7 Estación Astronomica Rio Grande, Rio Grande, Tierra del Fuego, Argentina

Received 2021 January 18; revised 2021 February 20; accepted 2021 February 23; published 2021 March 16

Abstract

We present observations of the sudden outburst of the A Carinid meteor shower recorded with the Southern Argentina Agile MEteor Radar Orbital System (SAAMER-OS) near the south toroidal sporadic region. The outburst peaked between 21 UT and 22 UT on 2020 October 14 and lasted 7 days (199° ≤ λ ≤ 205°), with a mean Sun-centered geocentric ecliptic radiant of λ = 271°04, β = −76°4, and a geocentric speed of 33.3 km s−1. Assuming a mass index value of s = 2.0, we compute a peak 24 hr average flux of 0.029 meteoroids km−2 hr−1 to a limit of 9th magnitude, which is equivalent to a zenith hourly rate (ZHR) of 5.7 and comparable to other established showers with similar mass indices. By further estimating the peak fluxes for other typical mass index values, we find that the outburst likely never exceeded a maximum ZHR of ~44, well below the activity of other strong showers. The mean orbital elements resemble those of a short-period object, a = 3.5 ± 0.1 au, q ≈ 1 au, e = 0.72 ± 0.02, i = 55°8 ± 0°3, ω = 1°4 ± 17°5, and Ω = 21°7, and are similar to those derived for two previous shower outbursts observed with SAAMER-OS at high southern ecliptic latitudes. Using the D′ criterion did not reveal a parent object associated with this shower in the known object catalogs.

Unified Astronomy Thesaurus concepts: Meteor showers (1034); Meteors (1041); Radar astronomy (1329)

1. Introduction

The solar system, like other planetary systems (i.e., β Pictoris, Burrows et al. 1995; Fomalhaut, Kalas et al. 2005), contains a circumsolar second-generation dusty disk known as the Zodiacal Dust Cloud (ZDC) populated by debris from asteroid collisions and the breakup and activity of comets and interstellar medium grains (Jenniskens 2006; Nesvorny et al. 2010). Dedicated radar and optical meteor surveys probe the dust content in the inner solar system via detection of meteoroid ablation high in Earth’s atmosphere, providing a reliable way to examine the dissemination of material populating the ZDC (Baggaley 2002; Brown et al. 2010; Jenniskens et al. 2011; Janches et al. 2015). Observations of the influx of material at Earth’s atmosphere reveal two populations clearly distinguishable in the distribution of the ZDC: the sporadic background sources composed of dynamically evolved submillimeter-sized meteoroids and micron-size dust grains largely affected by radiation pressure, and meteor showers, streams of younger and larger meteoroids that in principle could be linked dynamically to parent bodies, namely, asteroids and comets, due to their similarity in orbital elements (Jenniskens 2006).

Surveying, identifying, and studying the meteoroid population is highly relevant, both scientifically and from an operational standpoint. Collisions with 1µg–10g meteoroids moving at average relative speeds in excess of 30 km s−1 could constitute a risk to satellite operations and present a safety concern for manned space missions in low orbit. For example, showers like the Daytime Arietids (ARI), the Geminids (GEM), and the Quadrantids (QUA) attain fluxes for 0.3 cm equivalent particles near the limits for pressure vessel perforation (Moorehead et al. 2019). All-sky meteor surveys capable of conducting uninterrupted observations and delivering timely reports of sudden changes in meteor activity are thus highly desirable. Furthermore, recording long-term seasonal variations of the sporadic background provides constraints to meteoroid stream models (McNamara et al. 2004). Meteor orbit radars currently in operation, like the Canadian Meteor Orbit Radar (CMOR; Webster et al. 2004; Brown et al. 2008) and the Southern Argentina Agile MEteor Radar Orbital System (SAAMER-OS; Janches et al. 2015), aim to address these objectives by continuously monitoring the submillimeter-sized meteoroid population (10−7 to 10−8 kg) in order to collect a large quantity of meteoroid orbit and flux data sets to characterize meteor showers and sporadic sources (Brown & Jones 1995; Brown et al. 2010; Pokorný et al. 2014, 2017; Bruzzone et al. 2015; Campbell-Brown & Brown 2015; Janches et al. 2015).

Sudden changes in meteor shower activity are not unusual and, in some cases, may be dramatic. For instance, showers like the Leonids (LEO) or the Draconids (DRA) can produce storms with levels of activity thousands of times higher than normal (Kronk 2014). On the other hand, some shower outbursts can be relatively mild enough to only raise the flux slightly above the radar detection threshold, making observations of it possible for the first time. Recent examples of abrupt outbursts include the unexpected Draconid (DRA) meteor storm on 2012 October 8 (Ye et al. 2014) and two shower outbursts detected at austral latitudes: the Volantids shower (VOL) outburst...
The Planetary Science Journal, 2:56 (6pp), 2021 April

by the Advanced Meteor Orbit Radar (AMOR; Baggaley et al. 1994; Galligan & Baggaley 2002), which operated near Christchurch, New Zealand, to probe for clustering of meteor radiants. Since then, it has been applied by other meteor radar surveys (Brown et al. 2008, 2010; Bruzzone et al. 2015; Pokorný et al. 2017; Schult et al. 2018), and more recently, it has been applied to radar and optical meteor observations with SAAMER-OS and CAMS in Bruzzone et al. (2020). We employ the wavelet transform to isolate meteor showers as spatial and temporal enhancements in radiant space of geocentric Sun-centered ecliptic coordinates and geocentric speed: \((\lambda_r - \lambda_0, \beta_r, v_r)\). Meteoroids that belong to a specific shower concentrate in radiant space and time with a characteristic spread in angular coordinates and speed that differs from the sporadic meteor background. For a given radiant distribution, the wavelet transform returns a list of wavelet coefficients \((W_t)\) that serve as a metric for clustering in radiant space enhancing the presence of showers. The wavelet transform can be further optimized to amplify the presence of meteor showers by adjusting the wavelet kernel scale dimensions to resemble the shower’s natural spread in radiant space (Bruzzone et al. 2015). In this way, meteor showers can be effectively separated from the activity of the sparse sporadic meteor background. For SAAMER-OS, we adopt the wavelet kernel scale parameters \(\sigma_r = 2.5\) and \(\sigma_c = 15\%\) derived in Pokorný et al. (2017). For the analysis of SAAMER-OS daily observations, the wavelet-based algorithm is evaluated for \(\lambda_r - \lambda_0 \in [0^\circ, 360^\circ]\) and \(\beta_r \in (-90^\circ, 90^\circ]\) at 0.5° steps and for \(v_r \in [10\, \text{km}\, \text{s}^{-1}, 80\, \text{km}\, \text{s}^{-1}]\) at 5% steps, while advancing at 1° steps in \(\lambda_r\). For each day, this procedure returns a list of \(W_t\) for which each individual entry is compared to its yearly median and standard deviation. Those entries that exceed 3 times the total standard deviation above the yearly median, \(\sigma\), are stored and used in a global maxima search. We proceed to identify a shower core candidate as the radiant returning the maximum in \(W_t\). Each shower core identified is cross-referenced with a compiled list of known meteor showers (Brown et al. 2010; Pokorný et al. 2017). When the location of a core candidate is within 3° and 15% in \(v_r\) of a shower in the reference list, a match is recorded and the candidate is labeled with the shower IAU code in a radiant density map. Radiants in the map are color-coded by the number of adjacent radiants within 2.5° and serve as a proxy for local enhancements in showers. Those shower core candidates that do not match with any known shower are stored and plotted for review. In this study, we revisit the daily reports for the A Carinid outburst and repeat the wavelet-based analysis at 0.1° steps in \((\lambda_r - \lambda_0, \beta_r)\) and 1.5% in \(v_r\) to secure a more precise radiant position and speed. We then follow Bruzzone et al. (2020) to track the outburst progression with time by linking the shower core radiants at each degree in solar longitude.

3. Results and Discussion

3.1. Wavelet-based Activity Profile and Orbits

Figure 1 shows the radiant density plot with meteor detections over a 24 hr period on 2020 October 14 in Sun-centered geocentric ecliptic coordinates. The wavelet-based analysis pipeline labels the position of the Southern Taurids as STA, and the A Carinid outburst as candidate ID-0, as it does not match with any known shower in our reference catalog. The procedure returns the radiant location for the outburst at

\((\beta = -77^\circ.7)\) in late 2015 (Jenniskens et al. 2016; Younger et al. 2016; Pokorný et al. 2017), and the \(\beta\) Tucanid/\(\delta\) Mensid outburst \((\beta = -77^\circ.2)\) in early 2020 (Janches et al. 2020). Jenniskens (2020) reported significant meteor activity from the weak A Carinid shower on the night of 2020 October 13–14, recording 130 orbits with the Cameras for Allsky Meteor Surveillance (CAMS; Jenniskens et al. 2011). The A Carinid is a high-latitude southern shower, with \(\alpha_0 = 103^\circ.2, \delta_0 = -57^\circ, v_g = 30.1\, \text{km}\, \text{s}^{-1}\), and first reported in Jenniskens et al. (2018) based on 121 orbits with CAMS. In this work, we report the unpredicted outburst of the A Carinid shower as detected by SAAMER-OS, peaking on 2020 October 14 (\(\lambda_0 = 201^\circ\)). In Section 2 we describe the radar and the daily shower monitoring methodology. We provide results and characterize the outburst radiants, orbits, and fluxes in Sections 3.1 and 3.2. Lastly, conclusions and final remarks are presented in Section 4.

2. Instrumentation and Data Analysis

SAAMER-OS is a VHF all-sky multistation backscatter meteor orbit radar hosted by the Estación Astronómica Rio Grande (EARG) in Rio Grande, Tierra del Fuego, Argentina (Janches et al. 2015). Here we present a brief overview of the system and refer the reader to Janches et al. (2020) for a more in-depth review of the hardware and data analysis capabilities. SAAMER-OS is a SKYMET radar system (Hocking et al. 1997), currently consisting of a main site (SAAMER-C; 53.7°S, 67.8°W) and four remote stations: SAAMER-S (53.8°S, 67.8°W), located at approximately 7 km south of the SAAMER-C; SAAMER-N (53.6°S, 67.8°W), at 13 km northwest of the central station; SAAMER-W (53.8°S, 67.8°W), at approximately 7 km southwest of SAAMER-C; and SAAMER-E (53.7°S, 67.7°W), at roughly 4 km northeast of the main site. The main site hosts the 64 kW (peak power) transmitter with a single three-element crossed yagi transmitting antenna, and the five three-element crossed yagi receiving antenna interferometer array (Hocking et al. 1997; Jones et al. 1998). At the main site, meteors are detected as backscatter echoes from meteor trails a few kilometers in length (Kaiser & Singer 1956), with average interferometric errors less than 0.5°. SAAMER-OS transmits 32.55 MHz pulses with a repetition frequency of 625 Hz and employs a 7-bit Barker code to achieve a spatial resolution of 1.5 km. Each remote station is equipped with an identical single three-element crossed yagi receiving antenna to detect the slightly forward scattered signals off the meteor trails. The time delays between the detection of meteors at the main site and each remote site allow for the determination of the meteoroid speed and its trajectory. SAAMER-OS currently employs an empirical meteor deceleration correction to better estimate the true out-of-atmosphere meteoroid speed (Bruzzone et al. 2020). SAAMER-OS’s software suite for event detection, correlation, and orbit computation is similar to the one employed in CMOR (Weryk & Brown 2012) and runs in parallel to SKYMET’s standard software routines (SKYCORR; Hocking et al. 2001). Daily detection counts can exceed 10,000 meteoroid orbits (Janches et al. 2020), to a limiting radio magnitude of +9.0, equivalent to meteoroids of mass 10^{-8} \text{kg} (or 300 \mu\text{m} in diameter) at 30\,\text{km}\,\text{s}^{-1} (Verniani 1973).

The SAAMER-OS data reduction pipeline employs a 3D wavelet transform algorithm that is well suited to the daily detection of meteor shower radiants. This method was first used

\[ v_g = 30.1 \, \text{km} \, \text{s}^{-1}, \]
\[ \lambda_e - \lambda_0 = 271^\circ.04 \quad (\alpha_g = 98^\circ.905), \quad \beta_e = -76^\circ.4 \quad (\delta_g = -53^\circ.671), \quad \text{and} \quad v_g = 33.3 \text{ km s}^{-1}, \text{ achieving a maximum} \]

\[ W_c = 1358 \text{ on 2020 October 14 (} \lambda_e = 201^\circ), \text{ with a strong} \]

\[ \text{detection of 16.2} \sigma \text{ based on 1352 meteors. The wavelet} \]

\[ \text{analysis identifies the outburst activity for six consecutive days (} 199^\circ \leq \lambda_e < 205^\circ) \text{ from 2020 October 12 through 17. The} \]

\[ \text{radiant position and speed measured with SAAMER-OS agree} \]

\[ \text{with the mean values from 130 optical meteors reported by} \]

\[ \text{Jenniskens (2020) (} \alpha_g = 98^\circ.7, \delta_g = -54^\circ.3, \quad v_g = 32.4 \text{ km s}^{-1}), \text{ with a difference of} \]

\[ 0.64 \text{ and} \quad 0.9 \text{ km s}^{-1}, \text{ respectively. Such a} \]

\[ \text{difference in radiant position and speed between SAAMER-OS and CAMS video observations is in agreement with mean} \]

\[ \text{values found for a selection of 20 established meteor showers in Bruzzone et al. (2020). Figure 2 shows the annual activity} \]

\[ \text{profiles by computing} \quad W_c \quad \text{at} \quad \lambda_e - \lambda_0 = 271^\circ.04, \quad \beta_e = -76^\circ.4, \quad \text{and} \quad v_g = 33.3 \text{ km s}^{-1} \text{ while advancing at} \]

\[ 1^\circ \text{ steps in} \lambda_e \text{ for} \]

\[ 2017 \text{ through 2020. A horizontal dashed line indicates the} \]

\[ 3\sigma \text{ level above the median} \quad W_c \quad \text{of 51.4 for 2020. The profiles} \]

\[ \text{confirm the absence of this shower in past years with} \]

\[ \text{SAAMER-OS data, as well as its sudden appearance in 2020. The insert in Figure 2 displays the hourly flux from 0 UT 2020 October 12} \]

\[ \text{through 0 UT 2020 October 18 for a mass index value of} \quad 2; \text{ times of very low and zero meteor flux indicate reduced radar detectability of meteors when the shower} \]

\[ \text{radiant is close to the local zenith.} \]

---

**Figure 1.** SAAMER-OS daily radiant density plot in Sun-centered geocentric ecliptic coordinates on 2020 October 14. The Southern Taurids (STA) is labeled on the map, while the A Carinid outburst is labeled as shower candidate ID-0.

**Figure 2.** Wavelet coefficient (Wc) profiles at 1° steps in \( \lambda_e \) for 2017 through 2020. Wc values are estimated at the radiant position and speed of the outburst during the peak at \( \lambda_e = 201^\circ \). The horizontal dashed line marks the 3\( \sigma \) level above the annual median \( W_c \) for 2020. The inset shows the hourly flux from 0 UTC 2020 October 12 through 0 UTC 2020 October 18 for a mass index value of 2; times of very low and zero meteor flux indicate reduced radar detectability of meteors when the shower radiant is close to the local zenith.
elaborate on SAAMER-OS’s collecting area and meteoroid flux estimates in Section 3.2. The geocentric ecliptic radiant positions, geocentric speed, and orbital elements derived with the wavelet analysis are listed in Table 1. We employ 10,000 iterations in a Monte Carlo procedure to draw orbital element uncertainties from the errors in radiant position and speed. We follow Bruzzone et al. (2020) to estimate the error in the outburst radiant position as the angular separation of the wavelet radiant position and the position of peak radiant density at $\lambda_0 = 201^\circ$. On the same date, we adopt the error in the outburst speed as the standard error of the outburst mean geocentric speed. The errors in the outburst radiant position and speed are $0^\circ.7$ and $0.09$ km s$^{-1}$, respectively. The orbital elements derived with SAAMER-OS closely resemble those obtained by Jenniskens (2020) with video observations. Employing the $D'$ criterion (Drummond 1981) to look for potential parents for this outburst from the Minor Planet Center Orbit Database, results in no clear parent object of the A Carinid meteor shower.

In addition to the outburst reported here, two more sudden outbursts have been detected with SAAMER-OS south of the south toroidal region: the Volantids outburst on 2015 December 31 (Bruzzone et al. 2020), with $\lambda_0 = 280^\circ$, $\lambda_g - \lambda_0 = 304^\circ.1$, $\beta_g = -77.7$, and $v_g = 30.2$ km s$^{-1}$, and the $\beta$ Tucanid/$\delta$ Mensid outburst on 2020 March 12 (Janches et al. 2020), with $\lambda_0 = 352^\circ$, $\lambda_g - \lambda_0 = 305.73^\circ.1$, $\beta_g = -77.2$, and $v_g = 30.7$ km s$^{-1}$. Radiant and orbital elements are listed in Table 1. All three outburst orbits resemble those of a short-period comet but display similarly higher inclinations than expected for Jupiter-family comets. The orbits share similar shape, size, and inclination; however, their orientations differ, and both $\omega$ and $\Omega$ display more scatter. The orbits are close to several important mean motion resonances (MMRs) with Jupiter, especially to the 2:1, 5:3, 8:5, and 7:3 at $e = 0.7$ and $i = 55^\circ$ (Gallardo 2020), and differ from those at the south and north toroidal regions: $a \sim 1$ au, $e \sim 0.2$, $i \sim 60^\circ\text{-}70^\circ$ (Campbell-Brown 2008; Janches et al. 2015). Furthermore, the duration of these outbursts may indicate that they are part of relatively younger streams as opposed to the older sporadic meteoroids composing the toroidal ring that likely evolved from long-period comet-type objects (Pokorny et al. 2014). We note, however, that the duration for these outbursts (between 2 and 7 days) suggests fairly evolved streams as opposed to very young ones such as the Camelopardalids. Janches et al. (2020) report asteroid (248590) 2006 CS ($a = 2.91$ au, $e = 0.7$, $i = 52^\circ.3$, $\Omega = 172^\circ.4$, $\omega = 346^\circ.4$) as a promising parent candidate ($D' = 0.055$) of the $\beta$ Tucanid/$\delta$ Mensid shower outburst. The latter may suggest that the three detected outbursts could have originated from a short-period object. However, further analysis including dynamical simulations will be needed to properly address the origin of these outbursts.

### Table 1

| Object | $\lambda_g - \lambda_0$ (deg) | $\beta_g$ (deg) | $v_g$ (km s$^{-1}$) | $a$ (au) | $q$ (au) | $e$ | $i$ (deg) | $\omega$ (deg) | $\Omega$ (deg) |
|--------|-------------------------------|-----------------|-----------------|--------|--------|----|----------|-------------|-------------|
| A Carinid | 271.04 | $-76.4$ | 33.3 | 3.5 $\pm$ 0.3 | 0.9972 $\pm$ 0.0001 | 0.72 $\pm$ 0.02 | 55.8 $\pm$ 0.3 | 1 $\pm$ 173 | 21.7 |
| $\beta$ Tucanid/$\delta$ Mensid | 305.7 | $-77.2$ | 30.7 | 3.2 $\pm$ 0.1 | 0.976 $\pm$ 0.001 | 0.69 $\pm$ 0.01 | 50.8 $\pm$ 0.2 | 345.3 $\pm$ 0.5 | 172.0 |
| VOL | 304.1 | $-77.7$ | 30.2 | 3.1 $\pm$ 0.7 | 0.970 $\pm$ 0.001 | 0.69 $\pm$ 0.04 | 49.7 $\pm$ 0.8 | 166.2 $\pm$ 0.7 | 280.0 |

3.2. Estimating SAAMER-OS’s Collecting Area and Meteoroid Fluxes

In order to estimate meteoroid fluxes with SAAMER-OS, a measure of the radar collecting area and a correction for observational biases is needed. The meteoroid flux $\Phi$ can be estimated by dividing the debiased meteor rate $\Sigma$ by the collecting area $A$ (Campbell-Brown 2004; Campbell-Brown & Brown 2015; Bruzzone et al. 2015). Biases affecting radar observations are numerous and pertain to the specific radar system parameters, the interaction of the scattered waves within the atmosphere, and the inherent scattering mechanism of radio waves by free electrons (see Cephecha et al. 1998; Galligan & Baggaley 2004, 2005). Such biases include the initial trail radius effect, in which the observability of meteors occurring higher in the atmosphere is reduced owing to the increase of the mean free path with height, resulting in the attenuation of the echo amplitude for large trail widths due to destructive interference. Other effects include Faraday rotation, the change of the polarization plane of the radar wave as it passes through the ionosphere, the diffusion of meteor trails during formation, and the decay time of established meteor trails. The combination of these effects results in a decrease in the meteor rates for any radar system. To correct for observational biases, we make use of the derivation of SAAMER-OS’s RRF in Janches et al. (2015) and refer the reader to that study for an in-depth description and derivation of correction factors for this system. We model the transmitting and receiving antennas with an NEC-2D code to determine the beam patterns. We find that the total gain and beam patterns for the antennas are the same and well described by a smooth function on elevation alone and proceed to fit it with a nine-order polynomial. The peak gain is 8.8 dB at the zenith ($z = 0^\circ$), with the $-3$ dB point at $z = 79^\circ$. Correction for Faraday rotation is not necessary since SAAMER-OS receiving antennas are cross-yagis, and thus the system receives both linear polarizations and is not sensitive to this effect (Janches et al. 2015). To estimate the meteor collecting area, we use the methodology in Kaiser (1960), Brown & Jones (1995), and Brown et al. (1998) as a guide. The collecting area is a strip of space that is perpendicular to the meteor radiant and has a width given by the mean vertical trail length, which describes the altitude range in which ablation occurs. The length of the strip is the length of the echo line that extends from horizon to horizon. In practice, the echo line is truncated out to a limiting range for the radar. For the vertical trail height, we use the empirical relation with mass index $s$ reported in Brown et al. (1998) from TV observations of faint meteors by Flemming et al. (1993). Values of $s$ below 2.0 indicate that there is more mass in larger particles, where the opposite holds for larger values. In general, values for showers are in the 1.6–2.0 range, whereas values for sporadics are larger than 2.0 (Blaauw et al. 2011a, 2011b). We adopt $s = 2.0$ for the flux estimates in this work. However, we also include fluxes for a list of $s$ values and leave the development of a method to

---

8. [http://www.minorplanetcenter.net/iau/MPCORB.html](http://www.minorplanetcenter.net/iau/MPCORB.html), retrieved on 2020 October 20.
determine specific shower mass indices with SAAMER-OS for a future study.

We parameterize the echo line as a function of its elevation \( \phi \), weighted by the antenna gain, and obtain the length through numeric integration. Since the vertical trail length is only dependent on \( s \), the collecting area \( A(z, h, s) \) can be found by multiplying the echo line length by the vertical trail length. Following Campbell-Brown (2004), Campbell-Brown & Brown (2015), and Bruzzone et al. (2015), we set \( h = 100 \text{ km} \). After some algebra (see Baumann 2012), the collecting area \( A(z, h, s) \) can be expressed as

\[
A(z, h, s) = 2 \int_0^{\pi/2} \left( G(\phi) \cos z \right)^{s-1} R_E \cos \phi \times \left( \frac{R_E}{R_e + h} \sin \phi \right)^2 \frac{d\phi}{\sin z} \left( 1.15 + 14.6e^{-hr} \right),
\]

where \( G(\phi) = \sqrt{G_{Ts}(\phi)G_{Rs}(\phi)} \) is the radar gain pattern, \( z \) is the radiant zenith distance, \( R_e \) is Earth’s radius, \( h \) is the meteor height, \( s \) is the mass index, and \( \phi \) is the echo line elevation. Here \( G_{Ts} \) and \( G_{Rs} \) are the radar transmit and receive antenna gain powers, respectively, which in this case are identical.

Outburst fluxes are estimated at 1 hr intervals within a 6° \( \times 6° \) window centered on the position of wavelet-based outburst radian for each day from 2020 October 12 through 17. Windows are partitioned in 0'1 steps and the debiased meteor rates and collecting area computed. Hourly fluxes are then determined by finding the ratio \( \Sigma A \). Hourly A Carinid fluxes in Figure 2 are determined as the total sum of fluxes in the window within 3° from the position of the wavelet radian. The shower daily flux is computed from the averaged hourly fluxes, where we subtract the equivalent sporadic background averaged over 15 consecutive days before 2020 October 12 and 15 days after 2020 October 17. To help the comparison with results from visual observers, we follow Koschack & Rendtel (1990) and adjust the flux to a +6.5 limiting magnitude, \( \Phi_{+6.5} \), using

\[
\Phi_{+6.5} = \Phi_{+9.0} \times 10^{(6.5-9.0)(r-1)/2.5},
\]

and estimate the zenithal hourly rate (ZHR) with

\[
\text{ZHR}(r) = 37200 \text{ km}^2 \times \Phi_{+6.5} \times \left( \frac{r - 1.3}{13.1r - 16.45} \right),
\]

where \( r = 10^{(s-1)/2.5} \) is the population index. Figure 3 displays the average meteoroid fluxes for the outburst, before sporadic background subtraction, and sporadic background fluxes. The dashed line indicates the 30-day median sporadic flux. After background subtraction, we record a peak average flux, \( \Phi_{+9.0} = 0.029 \text{ meteoroids km}^{-2} \text{ hr}^{-1} \), down to a +9 limiting magnitude, on 2020 October 14. The peak flux corresponds to a ZHR\(_{\text{max}}\) of approximately 5.7 at \( r = 2.5 \). The peak ZHR is comparable to values for other established showers from optical observations at similar \( r \) values: Southern and Northern Taurids (ZHR\(_{\text{max}} = 5, \ r = 2.3 \)), September \( \epsilon \) Perseids (ZHR\(_{\text{max}} = 5, \ r = 2.5 \)), Aurigids (ZHR\(_{\text{max}} = 6, \ r = 2.5 \)), \( \alpha \) Capricornids (ZHR\(_{\text{max}} = 4, \ r = 2.5 \)), and \( \gamma \) Normids (ZHR\(_{\text{max}} = 4, \ r = 2.4 \)) (Rendtel 2014). Instead of displaying

![Figure 3. A Carinid 24 hr average fluxes with SAAMER-OS with \( s = 2.0 \).](image)

Table 2

| \( s \)     | \( \Phi_{+9.0} \) | \( \Phi_{+6.5} \) | ZHR   |
|-----------|-----------------|-----------------|------|
| 1.65      | 2.24e-2         | 5.33e-3         | 43.8 |
| 1.70      | 2.25e-2         | 4.89e-3         | 31.1 |
| 1.75      | 2.52e-2         | 4.50e-3         | 22.7 |
| 1.80      | 2.60e-2         | 4.12e-3         | 16.8 |
| 1.85      | 2.68e-2         | 3.78e-3         | 12.6 |
| 1.90      | 2.75e-2         | 3.47e-3         | 9.6  |
| 1.95      | 2.83e-2         | 3.17e-3         | 7.4  |
| 2.00      | 2.90e-2         | 2.90e-3         | 5.7  |
| 2.05      | 3.00e-2         | 2.66e-3         | 4.4  |
| 2.10      | 3.04e-2         | 2.43e-3         | 3.5  |
| 2.15      | 3.11e-2         | 2.23e-3         | 2.7  |
| 2.20      | 3.18e-2         | 2.04e-3         | 2.2  |

in general the shower mass index drops as Earth intersects the core of the stream where larger particles are located (Blauw et al. 2011b). For this reason, we include peak flux estimates for a list of mass index values typical for showers in Table 2. The wide range in ZHR values reflects the sensitivity of shower fluxes with mass index. Our flux estimates suggest that the outburst flux likely never rises above the level seen in showers like the \( \eta \) Aquariids, ZHR\(_{\text{max}}\) = 50–80, adopting \( s = 1.9 \) (Campbell-Brown & Brown 2015), or the Daytime Arietids, ZHR\(_{\text{max}}\) = 189 at \( s = 1.75 \) (Bruzzone et al. 2015).

We revisit archival observations of the previous \( \beta \) Tucanid shower outburst detected with SAAMER-OS on 2020 March 12 (Janches et al. 2020) and apply the procedure developed here to estimate the average peak flux. We find a maximum 9 mag flux for the \( \beta \) Tucanid shower of 0.01 meteoroids km\(^{-2}\) hr\(^{-1}\) at \( s = 2.0 \), which corresponds to a ZHR\(_{\text{max}}\) slightly above 2.

### 4. Conclusions

We reported radar observations of an unexpected outburst of the A Carinid meteor shower recorded with SAAMER-OS. Our wavelet-based analysis returned the shower radiant location south of the south toroidal ring at \( \lambda_\odot - \lambda_\odot = 271°.04 \).
\( \beta_k = -76^\circ, \text{ with } v_k = 33.3 \text{ km s}^{-1} \) during the peak at \( \lambda_0 = 201^\circ \) on 2020 October 14. The wavelet-based technique unequivocally confirms the sudden appearance of the outburst in 2020, rising 16 times the total standard deviation above the annual median. The outburst lasted for approximately 6 days from 2020 October 12 through 17. The radiant location, speed, and period of observation agree with those reported with video observations by Jenniskens (2020). We measured a 9 mag peak hourly flux of 0.09 meteoroids km\(^{-2}\) hr\(^{-1}\), assuming \( s = 2 \), down to a limiting mass of \( 1.9 \times 10^{-8} \) kg between 21 UTC and 22 UTC on October 14. To compute fluxes, we debiased the observed meteor rates and estimated the radar collecting area at 1 hr intervals. The 6 mag equivalent peak average daily flux corresponds to a ZHR of approximately 6, comparable to other known meteor showers at similar \( s \) values. We computed peak average fluxes for several mass index values to derive probable ZHR estimates returning limits between 2 and 44 approximately. The latter suggest that A Carinid fluxes remain well below those recorded for strong showers like the Daytime Arietids or Geminids. Based on 1352 events during the peak, the orbital elements resemble those of a short-period object: \( \alpha = 3.5 \pm 0.3 \text{ au}, q \approx 1 \text{ au}, e = 0.72 \pm 0.02, i = 55^\circ 8 \pm 0^\circ 73, \omega = 1^\circ 7 \pm 17^\circ 9, \text{ and } \Omega = 21^\circ 7 \). Comparably, two other auroral shower outburst previously recorded with SAAMER-OS, the \( \beta \) Tucanid/\( \delta \) Mensid and the Volantids (VOL), have orbits with shape, size, and inclination similar to the A Carinid. Our search for a parent object using the \( D^* \) criterion (Drummond 1981) did not reveal any clear candidate. While the duration suggests that the shower is not as old as the Arietids or Taurid streams (thousands to tens of thousands of years), the significant spread in nodal crossing may indicate a fairly evolved stream.

The work of J.S.B. and D.J. is supported by the NASA SSO and ISFM Programs and the NASA NESC. R.J.W. was supported through NASA grant 80NSSC18K0656. SAAMER’s operation is supported by NASA SSO, NESC assessment TI-17-01204, and NSF grant AGS-1647354. The authors appreciate the invaluable support of Carlos Ferrer, Gerardo Conlan, Luis Barbero, and Leandro Maslov with the operation of SAAMER.

**ORCID iDs**

Juan Sebastian Bruzzone @ https://orcid.org/0000-0002-2731-0397
Robert J. Weryk @ https://orcid.org/0000-0002-0439-9341
Diego Janches @ https://orcid.org/0000-0001-8615-5166
Carsten Baumann @ https://orcid.org/0000-0001-7104-5992
Gunter Stober @ https://orcid.org/0000-0002-7909-6345
Jose Luis Hormaechea @ https://orcid.org/0000-0003-4533-3282

**References**

Baggaley, W. J. 2002, in Meteor's in theEarth's Atmosphere, ed. E. Murad & I. P. Williams (Cambridge : Cambridge Univ. Press), 123
Baggaley, W. J., Bennett, R. G. T., Steel, D. I., & Taylor, A. D. 1994, QJRAS, 35, 293

Baumann, C. 2012, Master's thesis, Univ. Rostock https://www.iap-kborn.de/fileadmin/user_upload/MAIN-abteilung/optik/Forschung/Dipl_Baumann-Dipl-2012.pdf
Blauw, R. C., Campbell-Brown, M. D., & Weryk, R. J. 2011a, MNRAS, 412, 2033
Blauw, R. C., Campbell-Brown, M. D., & Weryk, R. J. 2011b, MNRAS, 414, 3322
Brown, P., Hocking, W. K., Jones, J., & Rendtel, J. 1998, MNRAS, 295, 847
Brown, P., & Jones, J. 1995, EM&P, 68, 223
Brown, P., Weryk, R. J., Wong, D. K., & Jones, J. 2008, Icar, 195, 317
Brown, P., Wong, D. K., Weryk, R. J., & Wiegert, P. 2010, Icar, 207, 66
Bruzzone, J. S., Brown, P., Weryk, R. J., & Campbell-Brown, M. D. 2015, MNRAS, 446, 1625
Bruzzone, J. S., Janches, D., Jenniskens, P., Weryk, R., & Hormaechea, J. L. 2020, PASJ, 188, 104936
Burrows, C. J., Krist, J. E., Stapelfeldt, K. R. & WFPC2 Investigation Definition Team 1995, BAAS, 27, 1329
Campbell-Brown, M., & Brown, P. G. 2015, MNRAS, 446, 3669
Campbell-Brown, M. D. 2004, MNRAS, 352, 1421
Campbell-Brown, M. D. 2008, Icar, 196, 144
Cephelea, Z., Borovička, J., Elfford, W. G., et al. 1998, SSRv, 84, 327
Drummond, J. D. 1981, Icar, 45, 545
Flemming, D. B., Hawkes, R. L., & Jones, J. 1993, in Meteoroids and their Parent Bodies, ed. J. Stoh & I. P. Williams (Bratislava: Slovak Academy of Sciences), 261
Gallardo, T. 2020, CeMDA, 132, 9
Galligan, D. P., & Baggaley, W. J. 2002, in IAU Colloq. 181, Dust in the Solar System and Other Planetary Systems, Vol. 15, ed. S. F. Green et al. (Oxford: Pergamon).
Galligan, D. P., & Baggaley, W. J. 2004, MNRAS, 353, 422
Galligan, D. P., & Baggaley, W. J. 2005, MNRAS, 359, 551
Hocking, W. K., Fuller, B., & Vandepeer, B. 2001, JASTP, 63, 155
Hocking, W. K., Thayaparan, T., & Jones, J. 1997, grl, 24, 2977
Janches, D., Bruzzone, J. S., Weryk, R. J., et al. 2020, ApJL, 895, L25
Janches, D., Close, S., Hormaechea, J. L., et al. 2015, ApJ, 809, 36
Jenniskens, P. 2006, Meteor Showers and Their Parent Comets (Cambridge: Cambridge Univ. Press)
Jenniskens, P. 2020, eMeNet, 5, 369
Jenniskens, P., Baggaley, J., Crompton, I., et al. 2018, PlanSS, 154, 21
Jenniskens, P., Gural, P. S., Dynneson, L., et al. 2011, Icar, 216, 40
Jenniskens, P., Nénon, Q., Albers, J., et al. 2016, Icar, 266, 331
Jones, J., Webster, A. R., & Hocking, W. K. 1998, RaSc, 33, 55
Kaiser, T. R. 1960, MNRAS, 121, 284
Kaiser, T. R., & Singer, S. F. 1956, P&ST, 9, 31
Kalas, P., Graham, J. R., & Clampin, M. 2005, Nature, 435, 1067
Koschack, R., & Rendtel, J. 1990, JMO, 18, 119
Kronk, G. W. 2014, Meteor Showers (New York: Springer)
Mcnamara, H., Jones, J., Kauffman, B., et al. 2004, EM&P, 95, 123
Moorehead, A. V., Egal, A., Brown, P. G., Moser, D. E., & Cooke, J. W. 2019, JSRRo, 56, 1531
Nesvorny, D., Jenniskens, P., Levison, H., et al. 2010, ApJ, 713, 816
Pokorný, P., Janches, D., Brown, P. G., & Hormaechea, J. L. 2017, Icar, 290, 162
Pokorný, P., Vokrouhlický, D., Nesvorny, D., Campbell-Brown, M., & Brown, P. 2014, ApJ, 789, 25
Rendtel, J. 2014, Meteor Shower Workbook 2014 (Hove: Int. Meteor Organization)
Schult, C., Brown, P., Pokorný, P., Stober, G., & Chau, J. L. 2018, Icar, 309, 177
Verniani, F. 1973, JGR, 78, 8429
Webster, A. R., Brown, P. G., Jones, J., Ellis, K. J., & Campbell-Brown, M. 2004, ACP, 4, 679
Weryk, R. J., & Brown, P. G. 2012, PlanSS, 62, 132
Ye, Q., Wiegert, P. A., Brown, P. G., Campbell-Brown, M. D., & Weryk, R. J. 2014, MNRAS, 437, 3812
Younger, J., Reid, I., & Murphy, D. 2016, in Proc. of the Int. Meteor Conf., ed. A. Roggemans & P. Roggemans (Hove: Int. Meteor Organization ), 352