PROBLEMATIC PHOSPHATIC PLATES FROM THE SILURIAN–EARLY DEVONIAN OF BOHEMIA, CZECH REPUBLIC

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ABSTRACT—Problematic phosphatic plates are reported for the first time from Bohemia, Czech Republic, and are attributed to Eurytholia bohemiaca n. sp. Similar mineralized elements, interpreted as sclerites, were known only in a very narrow interval from Middle-Late Ordovician beds bordering the Iapetus Ocean. This new report comes from the Silurian and Early Devonian and provides a significant range extension for these Problematica as well as an enlargement of their geographic extent. Comments open new perspectives in the interpretation of these elements.

INTRODUCTION

The principal, the Silurian–Upper Ordovician, biostratigraphic subdivision of the Silurian and Lower–Middle Devonian strata in Bohemia is based upon increasingly elaborate graptolite and conodont biozonation. A joint project is underway for the precise calibration of conodont and graptolite biozonation in specific intervals of the Silurian (responsible, ES and PS). Within this study, detailed resampling of some deﬁnitive levels in famous sections of the Prague Basin (Beroun, Butovice, Hýskov, Kosov Quarry, Amerika Quarry, Lodenice–Černídlava, Mušlovka Quarry, Požáry, U Topolů, Víšeradice), where both conodont- and graptolite-rich layers are present, has been recently performed in several ﬁeld seasons. Samples have been processed with the standard conodont preparation technique using formic or acetic acid and residues concentrated with sodium politungstate. The enigmatic phosphatic elements described below were picked from the heavy fraction of several levels, either from Silurian or Lower Devonian beds. In addition, a collection of these plates, made by one of us (ES) in the Barrandian in 1983, has been included in this study.

Similar small “hat-like” phosphatic plates of unknown origin were reported by Sutton et al. (2001) from the Ordovician of the Iapetus Ocean, specifically from South Wales, United Kingdom (12 specimens), Alabama, USA (nine specimens), Dalarna, Sweden (two specimens), and Estonia (one specimen). All the elements came from a very narrow stratigraphic interval of the Middle-Late Ordovician (Pygodus serra and P. anserinus conodont zones). The plates were attributed to two species of Eurytholia Sutton et al., 2001, and interpreted as dorsal dermal sclerites belonging to an animal of uncertain afﬁnity. A provisional reconstruction of the scleritome was attempted, as a dorsoventrally ﬂattened animal having sclerites arranged in sublongitudinal rows (Sutton et al., 2001, ﬁg. 4).

GEOLOGICAL SETTING

Several-kilometers-thick Cambrian to Middle Devonian deposits, unmetamorphosed and weakly tectonized by the Variscan cycle, unconformably overlie Neoproterozoic basement in the Barrandian area of the Bohemian Massif (Chlupáč et al., 1998). In the Early Ordovician, a northeast-southwest–trending Prague Basin began to subside in axial parts of the Barrandian (Havlíček, 1981; Chlupáč et al., 1998). In the earliest Silurian, the Ordovician shelf, pelitic to psammitic sedimentation was replaced by hemipelagic black graptolitic shales. Depositional settings varied from foreslope to deep shelf/basin. In the middle Wenlock and Ludlow, rarely in the middle Llandovery, several volcanic centers produced basalt lavas, hyaloclastites, and tuffs. Basaltic volcanoes were elevated high above the anoxic basin ﬂoor and were fringed by shallow-water bioclastic limestones (Havlíček and Storch, 1990). The Ludlow Series of the Prague Basin (Fig. 1) is represented by the Kopanina Formation—a sequence of platy, micritic (Butovice bed 10; Fig. 2), and bioclastic, richly fossiliferous limestones (Kosov samples 3, J, L; Amerika Quarry sample 3; Mušlovka Quarry samples A, B, C; Požáry sample 1; Fig. 2) interbedded with shales. Thick-bedded bioclastic brachiopod and cephalopod limestones (Mušlovka Quarry beds 33a–b and 33h–i; Fig. 2) occur particularly in the middle part and, once more, in the uppermost part of the formation (Kříž, 1998). Tuffaceous bioclastic limestones derived from local, brachiopod–coral bioclasts deposited in the vicinity of volcanic highs in the lower part of the Kopanina Formation. The Požáry Formation of Příšov is composed of platy micritic and biomicritic limestones with shaly interbeds and deposited largely under deeper, open-shelf conditions (Chlupáč et al., 1998). Former volcanic highs were capped with shallow-water, crinoidal grainstones. Massive bioclastic limestone with cephalopods and pelagic crinoids is developed in the uppermost part of the unit and in the lower part of the succeeding Lower Devonian Lochkov Formation (U Topolů bed 14; Fig. 2) (Chlupáč et al., 1972). In the Early Devonian, crinoidal limestones, calciturbidites, nodular, and also reeﬁc types gained prevalence.

MATERIAL

About 400 problematic phosphatic plates were recovered from six Bohemian sections (Butovice, Kosov Quarry, Amerika Quarry, Mušlovka Quarry, Požáry, and U Topolů; Fig. 1). All samples were precisely constrained biostratigraphically using conodonts. Five distinct Silurian conodont zones (Ludlow) and one Devonian (Lochkovian) conodont zone were documented (Fig. 2). All stratigraphic intervals sampled for conodonts within the project are in Ludlow and Lochkovian sediments and produced phosphatic plates. It is therefore probable that other Silurian and Devonian beds might be discovered in Bohemia or elsewhere, thus giving a continuous stratigraphic record of these elements.

Phosphatic plates recovered from Bohemia occur in variable quantities, apparently not strictly related to the weight of processed material. The number of plates recovered from the same section in the Kosov Quarry (Storch, 1995) ranges from a sole specimen (sample L: 5.6 kg of processed material) to three elements (sample J: 6.7 kg) or 14 plates (sample 3: 4.9 kg) and reaches up to 284 elements in the sample Mušlovka Quarry A (5.4 kg) (Fig. 2).

Plate dimensions range between 0.3 mm and 2 mm and their color varies from amber to black in massive specimens. If compared with conodonts, plates share the same C.A.I. 5 (Color Alteration Index) except of Kosov, where conodont elements have a slightly lower C.A.I. (3–4). Preservation of Silurian material is
good. Early Devonian elements are, on the contrary, poorly preserved. Detailed SEM observation of the outer surface revealed no significant ornamentation patterns or growth lines. Peripheral striae were observed along the margin of a single specimen. Rare angular pores (2 μm in diameter) are present distally and apically (Fig. 3.8c).

In addition, the electron microscope investigation confirmed the remarks given by Sutton et al. (2001) on the shell structure. Phosphatic plates are composed of two clearly separated distinct layers (Fig. 4.10). The outer (dorsal) layer, 100 μm thick, is apparently homogeneous and massive in unetched specimens (Fig. 4.9) and appears thicker than the corresponding one observed in Eurytholia prattensis Sutton, Holmer, and Cherns, 2001 (20–40 μm). An inner (ventral) reticulated layer (Fig. 4.10) develops along all the inner surfaces (Fig. 4.12b), with a 2 μm hole diameter of the “net-like” pattern. A similar ornamentation was described in the “laminated ventral secondary layer of variable thickness” of a single specimen of E. prattensis Sutton et al. (2001, p. 1), having a hole diameter of 5 μm. Some shell margins expose a shell lamination which might be confirmed only with a detailed study of the shell ultrastructure.

A combination of X-rays, Gandolfi camera, and EDAX analysis revealed a shell composition of calcium phosphate. It appears to be primary phosphatization since no other phosphatized fossils have been found with the fauna. Analysis of the outer layer also showed a minor presence of iron, while a fluorine enrichment was detected in the inner reticulated layer. It is noteworthy that any calcareous layer that may have been present would have been destroyed by lab processing.

SYSTEMATIC PALEONTOLOGY

The material described in this paper has been deposited at the Paleontological Museum of the “Dipartimento del Museo di Paleobiologia e dell’Orto Botanico” of the University of Modena and Reggio Emilia (IPUM).

We follow the orientation proposed by Sutton et al. (2001) for the description of the phosphatic elements.

Genus Eurytholia Sutton, Holmer, and Cherns, 2001

Type species.—Eurytholia prattensis Sutton, Holmer, and Cherns, 2001

Diagnosis (slightly emended from Sutton et al., 2001).—Plates phosphatic, small, and transversely ovoid, symmetrical or sub-symmetrical about transverse and longitudinal axes. Transverse profile subtriangular, formed by high longitudinal ridge positioned medially or submedially; lateral slopes of ridge concave. Longitudinal profile subsemicircular with sublinear anterior and posterior slopes. Ventral surface of concave plate thickened somewhat at anterior and posterior margins. Shell structure consisting of homogenous dorsal primary layer and laminated ventral secondary layer of variable thickness. Dorsal surface lacking ornamentation and growth lines.

Eurytholia bohemica new species

Figures 3, 4

Diagnosis.—Elliptical plates having a linear or slightly curved median, submedian, or marginal ridge. A well-defined girdle with an inner furrow runs all along the margins. Wide basal cavity expands to the apex of the ridge.

Description.—Great variability in shape, size, and other main morphological features typifies our material (even within the same sample). Plate margins linear to curved. General outline transversally and longitudinally symmetrical to asymmetrical. The most common plate type (Fig. 3.8a, 3.8b) elliptical, longitudinally asymmetrical, with a ridge in eccentric position in relation to lateral margins. Long, linear anterior and posterior margins. Plate maximum length in a variable position, in some specimens corresponding to the plate ridge axis. Plate ridge perpendicular to oblique with respect to long margins. Longitudinal profile of ridge frequently asymmetrical, with highest point in a marginal position
(Fig. 4.1, 4.6). Height and width of median ridge strongly variable and independent from general dimensions of the plate. The ridge may be, in fact, large (Fig. 3.4) or slim (Fig. 3.2), rounded or sharp, or having a small apical depression (Fig. 4.11b). The ridge is sometimes apically worn (Fig. 3.11b) but no true abrasions, scratches, or cuts have been observed along the ridge that might be convincingly attributed to in vivo damage. Lateral slopes of ridge symmetrical to asymmetrical in transverse view. Plate halves often dorsally bent (Fig. 4.2).

Plate girdle is delimited by two well-defined thickened margins (Figs. 3.3, 3.5, 4.8). A linear inner furrow runs in between all along the plate. The upper margin of the girdle may protrude outside the basal cavity margin (Fig. 3.9).

Rare pores (2 μm in diameter) with a geometric outline (Fig. 3.8c) may occur with no preferential arrangement either in the apical or distal part of the plates.

The wide basal cavity expands deeply below the ridge, and in plates having a marginal ridge the cavity is edged on one side by the ridge wall and on the other by a wide platform (Fig. 3.10). A more rounded aspect characterizes some subcircular plates (Fig. 3.7), longitudinally and transversally symmetrical, having nonlinear anterior and posterior margins and lateral margins not easily detectable from the former. Others have, however, a sub-rectangular profile (Fig. 3.1, 3.2).

One specimen reveals a hole with circular outline on one side of the plate, and with perpendicular walls (Fig. 4.4). Excavations and predatory perforations are reported in various phosphatic groups such as Cambrian shells (Conway Morris and Bengtson, 1994), Ordovician brachiopods (Holmer, 1989), and conodonts (Müller and Nogami, 1972). Two plates (Fig. 4.12a) were found fused together. Similar clusters are common in conodonts and represent elements fused together by diagenetic minerals which may (e.g., Nicoll, 1985; Nicoll and Rexroad, 1987) or may not (e.g., Nowlan, 1979; Dzik and Drygant, 1986) retain the functional orientation they had in the living animal (Purnell and Donoghue, 1998). The rocks from which these plates come are mainly bioclastic packstones, rich in trilobites and echinoderm fragments, suggesting energy for transport which would have dissociated unfused elements of an assemblage.

**Etymology.**—From Bohemia.

**Type.**—Holotype, IPUM 27845, Figure 3.8a, 3.8b. Sample Kosov 3, Kopanina Formation, Kosov Quarry, Ludlow (Polygnathoides siluricus Zone).

**Other material examined.**—Four hundred fifteen elements.

**Occurrence.**—Silurian (Ludlow) and Devonian (early Lochkovian) of Bohemia, Czech Republic.

**Discussion.**—The Bohemian specimens appear bigger than the older reports, having a maximum width of 2 mm (compared to 1.2 mm in E. prattensis and E. elibata Sutton, Holmer, and Cherns, 2001). The Ordovician material bears anterior and posterior margins often thickened in “rolls,” extending possibly also to lateral margins. The Bohemian material reveals, on the contrary, a well-defined girdle, bordered by two margins, running all along the entire plate border with a more or less regular thickness.

![Figure 2](image)

**Figure 2**—Age and abundance of phosphatic plates in the Bohemian samples. Conodont biozones mostly refer to Corradini and Serpagli (1999).
E. bohemica differs from E. pratensis in lacking the median indentation ("waist") in anterior and posterior margins, having median length often representing maximum plate length, sometimes even expanding outwards. E. bohemica differs from E. pratensis in being more rounded and less angular in general aspect. E. bohemica differs from E. elibata in having no median arching of posterior and anterior margins above the resting plane.

CONCLUSION

Sutton et al. (2001) carefully analyzed possible affinities of these Problematica and regarded the phosphatic plates as disarticulated skeletal elements of a sclerite possessing only one sclerite morphotype. Phosphatic elements were possibly arranged in soft tissue (presumably chitinous) and exerted a protective function.

Our observations of the Bohemian material confirmed many of the morphological and structural features already pointed out by previous authors. In addition, small pores have been observed. Moreover, compared to the simple thickening of the anterior and posterior margins of earlier reports, Eurytholita bohemica n. sp. bears a peculiar girdle along all margins which appears to be a more sophisticated site for attachment or insertion into the animal's soft tissue.

The possible evidence of predation in the Bohemian material, even if from a sole specimen, could support a protective role for these plates. In spite of that, a predator could have more easily attacked in unprotected parts of the body (such as plate intervals).

The Bohemian material considerably enlarges the geographic extent and significantly extends the range of these Problematica, which appear to be a common constituent of the late Silurian–Early Devonian fauna in Bohemia and possibly elsewhere. A recent finding of E. bohemica from the late Silurian of the Austrian Carnic Alps (AF and ES personal obs.) corroborates the idea. The stratigraphical range of Eurytholita was a period when the earliest vertebrates were evolving. The possible affinity with the vertebrates needs to be fully explored, mainly by histological analysis.

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