Summer and winter variations of dicarboxylic acids, fatty acids and benzoic acid in PM$_{2.5}$ in Pearl Delta River Region, China

K. F. Ho$^{1,2}$, S. S. H. Ho$^{1,3}$, S. C. Lee$^2$, K. Kawamura$^4$, S. C. Zou$^4$, J. J. Cao$^1$, and H. M. Xu$^1$

$^1$SKLLQG, Institute of Earth Environment, Chinese Academy of Sciences, Xi’an, 710075, China
$^2$Research Center of Urban Environmental Technology and Management, Department of Civil and Structural Engineering, The Hong Kong Polytechnic University, Hung Hom, Kowloon, Hong Kong, China
$^3$Division of Atmospheric Sciences, Desert Research Institute, Reno, NV 89512, USA
$^4$Institute of Low Temperature Science, Hokkaido University, Sapporo, Japan

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Abstract. Ground-based PM$_{2.5}$ samples collected at four different sites in Pearl River Delta region (PRD) during winter and summer (from 14 December 2006 to 28 January 2007 in winter and from 4 July to 9 August 2007 in summer) were analyzed for 30 water-soluble organic species, including dicarboxylic acids, ketocarboxylic acids and dicarbonyls, nine fatty acids, and benzoic acid. Molecular distributions of dicarboxylic acids demonstrated that oxalic acid (C$_2$) was the most abundant species followed by phthalic acid (Ph) in PRD region. The concentrations of total dicarboxylic acids ranged from 99 to 1340 ng m$^{-3}$, with an average of 438 ± 267 ng m$^{-3}$ in PRD. The concentrations of total ketocarboxylic acids ranged from 0.6 to 207 ng m$^{-3}$ (43 ± 48 ng m$^{-3}$ on average) while the concentrations of total α-dicarbonyls, including glyoxal and methylglyoxal, ranged from 0.2 to 89 ng m$^{-3}$, with an average of 11 ± 18 ng m$^{-3}$ in PRD. The total quantified water-soluble compounds (TQWOC) (organic carbon) accounted for 3.4 ± 2.2% of OC and 14.3 ± 10.3% of water-soluble OC (WSOC). Hexadecanoic acid (C$_{16:0}$), octadecanoic acid (C$_{18:0}$) and oleic acid (C$_{18:1}$) were the three most abundant fatty acids in PRD. The distributions of fatty acids were characterized by a strong even carbon number predominance with a maximum (C$_{max}$) at hexadecanoic acid (C$_{16:0}$). Ratio of C$_{18:1}$ to C$_{18:0}$ acts as an indicator for aerosol aging. In PRD, an average of C$_{18:1}$/C$_{18:0}$ ratio was 0.53 ± 0.39, suggesting an enhanced photochemical degradation of unsaturated fatty acid. Moreover, the concentrations of benzoic acid ranged from 84 to 306 ng m$^{-3}$, (165 ± 48 ng m$^{-3}$ on average), which can be emitted as primary pollutant from motor vehicles exhaust, or formed from photochemical degradation of aromatic hydrocarbons. Seasonal variations of the organic specie concentrations were found in the four sampling cities. Higher concentrations of TQWOC were observed in winter (598 ± 321 ng m$^{-3}$) than in summer (372 ± 215 ng m$^{-3}$). However, the abundances of TQWOC in OC mass were higher in summer (0.9–12.4%, 4.5 ± 2.7% on average) than in winter (1.1–5.7, 2.5 ± 1.2% on average), being consistent with enhanced secondary production of dicarboxylic acids in warmer weather. Spatial variations of water-soluble dicarboxylic acids were characterized by higher concentrations in Hong Kong and lower concentrations in Guangzhou (GZ)/Zhaoqing (ZQ) during winter whereas the highest concentrations were observed in GZ/ZQ during summer. These spatial and seasonal distributions are consistent with photochemical production and the subsequent accumulation under different meteorological conditions.

1 Introduction

Organic acids including monocarboxylic acids, dicarboxylic acids and aromatic acids in atmospheric aerosols have been studied owing to their potential roles in affecting the global climate. These acids can reduce the surface tension of particles to form cloud condensation nuclei (Facchini et al., 1999; Kerminen et al., 2000) and are involved in a series of atmospheric chemical reactions occurring in the gas, water and particle phases (Chebbi et al., 1996; Saxena et al., 1996). Dicarboxylic acids and related polar compounds are important components of atmospheric aerosols. Due to the hygroscopic property, they play a role in the global radiative balance (Saxena et al., 1995; Facchini et al., 1999; Kerminen,
Distributions and concentrations of the organic acids in aerosols are important to understand their photochemical reactions and long-range transport. These species are emitted to the atmosphere directly from natural and anthropogenic primary sources (Kawamura and Kaplan, 1987; Rogge et al., 1991, 1993a; Fang et al., 1999; Schauer et al., 1999; Simoneit et al., 2002). They are also produced by secondary atmospheric chemical reactions. Total dicarboxylic acids account for ∼1–3% of the total particulate carbon in urban areas and >10% in remote marine environments (Kawamura and Ikushima, 1993; Kawamura et al., 1996a, b; Kawamura and Sakaguchi, 1999; Kermine et al., 2000).

Fatty acids were one of the most abundant compound classes in the polar organic fraction of aerosols from urban atmospheres (Oliveira et al., 2007). They were found to contribute 6–53% of identified organic compounds from emission sources such as biomass burning (Rogge et al., 1998; Nolte et al., 2001; Schauer et al., 2001; Fine et al., 2002), cooking (Schauer et al., 1999, 2002; Rogge et al., 1991; He et al., 2004), paved road dust (Nolte et al., 2002) and automobiles (Rogge et al., 1993a; He et al., 2006). Benzoic acid is a secondary product from photochemical degradation of aromatic hydrocarbons emitted by automobiles (Suh et al., 2003). This acid has also been measured as primary pollutant in the exhaust of motor vehicles (Kawamura et al., 1985; Rogge et al., 1993b).

The Pearl River Delta (PRD) region in China covers nine prefectures of the province of Guangdong, namely Guangzhou, Shenzhen, Zhuhai, Dongguan, Zhongshan, Foshan, Huizhou, Jiangmen, Zhaoqing, the Hong Kong Special Administrative Region (HKSAR), and the Macau Special Administrative Region. It had a population approximately of 40 million people. The climate of PRD is dominated by the Asian monsoon, with north wind as prevailing wind in winter and south wind in summer. PRD is one of the rapidest economic growth regions in China. As influenced by urbanization and industrialization, air pollution has been becoming more and more severe in PRD, which is one of the four heaviest haze regions in China, together with the Yangtze River Delta, Beijing-Tianjin-Tangshan and Chongqing. Particulate matter with diameter less than 2.5 micrometers (PM$_{2.5}$) has recently received much attention (Yang et al., 2005; Feng et al., 2007; Li et al., 2008). These fine particles can penetrate deeply into human’s lung, and also affect visibility, environment, and radiative forcing (Penner and Novakov, 1996; Menon et al., 2002; Wilkening et al., 2000; Nel, 2005). Adverse health, environment, and climate effects of the fine particles are constitutionally derived from their chemical components and properties.

Previous studies have determined the organic acids in Guangzhou (Feng et al., 2006; Wang et al., 2006). However, the seasonal and spatial variations of organic acids were not studied (e.g., Feng et al., 2006). To better understand the organic compositions in the PRD region, PM$_{2.5}$ samples were acquired at four sampling sites simultaneously during winter and summer. The main objectives of this study are to determine the spatial and seasonal variations of selected organic species and to explore their implications for sources and photochemical reactions.

2 Experiment

2.1 Sample collection

Four sampling sites were selected in PRD region, including Sun yat-sen University in Guangzhou (GZ), Zhaoqing University in Zhaoqing (ZQ) in Mainland of China, as well as The Hong Kong Polytechnic University (PU) and Hok Tsui (HT) in Hong Kong. Their locations are shown in Fig. 1. These four sites represent different types (urban: GZ; semi-rural: ZQ; urban/roadside: PU; rural: HT). Twenty-four-hour sampling of PM$_{2.5}$ were conducted simultaneously in the four sites from 14 December 2006 to 28 January 2007 in winter and from 4 July to 9 August 2007 in summer. Fifteen samples (i.e. 8 samples in winter, 7 samples in summer) were collected at each site for consequence organic analyses. The PM$_{2.5}$ were acquired on pre-heated (800°C, 3 h) quartz fiber filters (102 mm) by medium volume samplers at a flow rate of 113 L min$^{-1}$. The sampling flow rates were checked before and after sampling with a TSI mass flow meter (model 4040, Shoreview, MN, USA). After sampling, aerosol loaded filters were stored in a refrigerator at 4°C to prevent loss of any volatile components. However, this temperature does not totally prevent the loss of very volatile components and does not avoid some microbial processing. One field blank was collected at each site to subtract positive artifacts that resulted from passive adsorption of gas-phase organic compounds onto the filter during and/or after sampling. Meteorological data (from Hong Kong Observatory and China Meteorological Administration) shows that the southern air mass flow was dominated during summer, therefore GZ and ZQ were the downwind sampling locations in summer. One the contrary, northern air mass flow dominated during winter, thus PU and HT were the downwind sites in winter. Molecular compositions of low molecular weight diacids (C$_2$–C$_{12}$), ketocarboxylic acids (ωC$_2$–ωC$_9$, pyruvic acid), α-dicarboxyls (C$_2$–C$_3$), benzoic acid and fatty acids (C$_{12}$–C$_{25}$) were determined by gas chromatography/mass spectrometry (GC/MS). Samples were also analyzed for organic carbon (OC), elemental carbon (EC) and water-soluble organic carbon (WSOC).

2.2 OC, EC and WSOC analysis

OC and EC were measured on a 0.526 cm$^2$ punch from each filter by thermal optical reflectance (TOR) following the IM-PROVE protocol on a DRI Model 2001 Thermal/Optical Carbon Analyzer (Atmoslytic Inc., Calabasas, CA, USA) (Chow et al., 2004, 2005; Cao et al., 2003a). This produced four OC fractions (OC1, OC2, OC3, and OC4 at 120, 250, 450, 1919–2208, 2011 www.atmos-chem-phys.net/11/2197/2011/
and 550 °C, respectively, in a helium (He) atmosphere), a pyrolyzed carbon fraction (OP, determined when reflected laser light attained its original intensity after oxygen (O₂) was added to the combustion atmosphere), and three EC fractions (EC1, EC2, and EC3 at 550, 700, and 800 °C, respectively, in a 2% O₂/98% He atmosphere). IMPROVE OC is operationally defined as OC1 + OC2 + OC3 + OC4 + OP, whereas EC is defined as EC1 + EC2 + EC3−OP. The minimum detection limit for the carbon analysis is 0.30 and 0.15 µgC m⁻³ for OC and EC, respectively, with a precision better than 10% for total carbon (TC).

For determination of WSOC, five punches (a total area of 2.63 cm²) were taken from each filter and placed into a 15 ml screw-capped vial to which 5 ml of distilled de-ionized water (DDW) was added. Samples were extracted in an ultra-sonic water bath for 1 h. Filter debris and suspending insoluble particles were removed from the water extracts using a syringe filter (0.2 µm PTFE membrane). Each filtered extract was then transferred into a clean auto-sampler fitted vial. The filtered extract was analyzed for total organic carbon (TOC) using a Shimadzu TOC-V CPH high-sensitivity Total Carbon Analyzer (Columbia, MD, USA). The minimum detection limit is 0.01 µgC m⁻³, with a precision of ±5%. Negligible amounts of OC, EC, and WSOC were observed in the field blanks. The data reported here are all corrected for the blanks.

2.3 Inorganic compounds analysis

Inorganic ions were determined using an ion chromatograph (IC) gradient pump (LC40) with a conductivity detector (CD25) (Dionex, Sunnyvale, CA, USA). An analytical column (AS11, 4 mm) with a guard column (AG11, 4 mm) and an anion trap column was used for inorganic anion detection with gradient elution from 0.2 to 5 mM NaOH. A cation analytical column (CS12, 4 mm) with a guard column (CG12, 4 mm) was used to analyze inorganic cations with an eluent of 20 mM methanesulfonic acid.

2.4 Organic acid analysis

The details of sample extraction and derivatization are documented elsewhere (Kawamura 1993; Kawamura and Yasui, 2005; Ho et al., 2010). An aliquot of the sample was extracted with pure water (10 ml × 3) to isolate low molecular weight dicarboxylic acids, ketoacids and α-dicarbonyls as well as free fatty acids. After concentration, the extracts were reacted with 14% BF₃/n-butanol at 100 °C to convert the carboxyl groups to butyl esters and the aldehyde groups to dibutoxy acetals. The derivatized total extracts were analyzed with a Hewlett-Packard 6890 GC installed with HP5 fused silica capillary column (25 m × 0.20 mm i.d., 0.5 µm film thickness) and flame ionization detector. Authentic standards were used for the peak identification based on GC
Retention times. Homologous series of fatty acids were determined as butyl esters (Mochida et al., 2007). Mass spectral confirmation of the compounds was achieved using ThermoQuest Trace GC/MS (Austin, TX, USA) using a similar GC conditions. Recoveries of the dicarboxylic acids, ketocarboxylic acids, α-dicarboxyls, and fatty acids were >70%. The reproducibility errors of the methods for the determination of organic species was <15% (Kawamura and Yasui, 2005; Mochida et al., 2007). Levels of field blanks were within 15% of actual samples, except for phthalic acid (up to 30%). The data reported here were all corrected against the blanks. The term of total quantified water-soluble organic compounds (TQWOC) (carbon) is defined as the sum of dicarboxylic acids, ketoacids and α-dicarboxyls.

3 Results and discussion

3.1 Concentrations of OC, EC and WSOC

Spatial and seasonal distribution of OC, EC and WSOC are shown in Table 1. Average OC concentrations ranged from 1.8 ± 0.8 (HT in summer) to 13.9 ± 4.4 μgC m⁻³ (PU in winter) while average EC concentrations ranged from 0.7 ± 0.2 (HT in summer) to 14.7 ± 4.4 μgC m⁻³ (PU in summer). Among the four sampling sites, the average WSOC at ZQ had the highest concentrations (4.7 ± 2.5 μgC m⁻³ in winter), while the lowest was found at PU and HT (0.4 ± 0.1 and 0.4 ± 0.2 μgC m⁻³, respectively in summer). The OC to EC ratio has been used to infer the origin of carbonaceous particles (Cao et al., 2003b; Novakov et al., 2005). The average OC/EC ratios (0.7 ± 0.4) at PU site (roadside) were significantly lower than those found at the urban/rural sites. The low OC/EC ratio at PU was primarily due to the high EC emissions from automobiles. Higher OC/EC ratios (2.5 ± 0.7) at the HT sites suggest that the transportation of older aerosol as well as secondary organic aerosol (SOA) were significant. The ratio of WSOC to OC ranged from 0.04 to 0.64, with an average of 0.29 ± 0.16. And TQWOC (carbon) accounted for 3.4 ± 2.2% of OC and 14.3 ± 10.3% of WSOC.

3.2 Molecular compositions of dicarboxylic acids, ketoacids and α-dicarboxyls

Concentrations of dicarboxylic acids (C₂ – C₁₂) (including straight-chain saturated, unsaturated, branched-chain, and hydroxylated), ketocarboxylic acids (C₂ – C₆), and α-dicarboxyls in the PM₁.₅ samples are listed in Table 1. In general, large variations of these species were observed among the four sampling cities in PRD. The concentrations of total dicarboxylic acids ranged from 99 to 1340 ng m⁻³, with an average of 438 ± 267 ng m⁻³ in PRD. Oxalic acid (C₂) was the most abundant dicarboxylic acid, followed by phthalic acid (Ph). These two species accounted for on average ~60% of TQWOC. The concentrations of oxalic acid ranged from 31 to 1035 ng m⁻³ (260 ± 213 ng m⁻³, on average), which are within the range of the values we reported in the same sampling sites in Hong Kong (Ho et al., 2007). The predominance of oxalic acid was recognized in previous studies also (Ho et al., 2007, 2010). The abundant presence of cis configuration (maleic acid and methylmaleic acid) in the urban atmosphere supports an oxidation of aromatic hydrocarbons (benzene and toluene) as a precursor of oxalic acid. Three phthalic acids including o-, m- and p- isomers were detected. The isomer distribution was characterized by a predominance of phthalic acid followed by terephthalic acid and isophthalic acid, being consistent with those reported in the aerosols in Mt. Tai, China (Fu et al., 2008) and East China Sea (Simoneit et al., 2004). The average concentration of phthalic acid in PRD (81 ± 74 ng m⁻³) is ~2 times higher than that observed in urban area of Tokyo in summer (29 ng m⁻³ on average) (Kawamura and Yasui, 2005), but is close to those reported in the Chinese cities (Ho et al., 2007). Phthalic acid can be formed by photodegradation of naphthalene (NAP) and other polycyclic aromatic hydrocarbons (PAHs) in atmospheric aerosols (Bunce et al., 1997; Jang and McDow, 1997). NAP is a ubiquitous pollutant in the atmosphere. The concentrations of NAP in urban areas such as Hong Kong have been reported to be as high as 3.5 μg m⁻³ (Lee et al., 2001). The products generated in the reaction of gas phase NAP with OH radical have lower vapor pressures than NAP, thus promoting the formation of SOA. Besides C₂ – C₄ dicarboxylic acids, the concentrations of azelaic acid (C₉) were the highest among the straight-chain saturated carboxylic acids in the PRD. Azelaic acid is an oxidation product of unsaturated fatty acids (Kawamura and Gagosian, 1987). An average abundance of azelaic acid was found to be 13.8 ± 9.1 ng m⁻³ in PRD, indicating that aerosols of biological origins are exposed to significant atmospheric processing.

The total concentrations of ketocarboxylic acids ranged from 0.6 to 207 ng m⁻³ (43 ± 48 ng m⁻³ on average). Glyoxylic acid (ωC₂) is the dominant ketocarboxylic acid, followed by pyruvic acid (Pyr) and 4-oxobutanoic acid (ωC₄). Their concentrations are close to those reported in Tokyo, Japan (Kawamura and Yasui, 2005) and other urban sites in China (Ho et al., 2007, 2010). The total concentrations of α-dicarboxyls, including glyoxal and methylglyoxal, ranged from 0.2 to 89 ng m⁻³ with an average of 11 ± 18 ng m⁻³ in PRD. Their concentrations were consistent to those reported in Hong Kong (Li and Yu, 2005) which did not exceed 100 ng m⁻³. α-Dicarboxyls have been demonstrated to be precursors to the formation of SOA via heterogeneous processes (Kroll et al, 2005; Liggio et al., 2005). Photo-oxidation of glyoxal can lead the formation of oxalic acid. The higher concentrations of glyoxal and methylglyoxal may represent the greater potential of subsequence SOA formation processing in PRD.
Table 1. Concentrations of dicarboxylic acids, ketoxylic acids, benzoic acid, and fatty acids in PM$_{2.5}$ samples in PRD during winter and summer in 2006–2007, Guangzhou (GZ), Zhaoping (ZQ), The Hong Kong Polytechnic University (PU), Hok Tsui (HT).

| Compounds                  | GZ (µg m$^{-3}$) | ZQ (µg m$^{-3}$) | PU (µg m$^{-3}$) | HT (µg m$^{-3}$) |
|----------------------------|------------------|------------------|------------------|------------------|
| Winter                     | Summer           | Winter           | Summer           | Winter           |
| Oxalic, C2                 | 182.0 ± 106      | 216.0 ± 97.2     | 209.0 ± 159      | 212.0 ± 96.9     |
| Malonic, C3                | 13.3 ± 955       | 18.5 ± 879       | 20.3 ± 151       | 11.4 ± 720       |
| Succinic, C4               | 18.4 ± 11.2      | 15.2 ± 4.22      | 22.7 ± 9.77      | 11.5 ± 584       |
| Glutaric, C5               | 5.82 ± 3.00      | 4.72 ± 1.84      | 8.05 ± 4.15      | 3.94 ± 1.55      |
| Adipic, C6                 | 5.39 ± 2.92      | 4.85 ± 1.97      | 7.69 ± 3.96      | 6.25 ± 2.26      |
| Pimelic, C7                | 2.68 ± 1.19      | 1.61 ± 1.10      | 4.29 ± 2.38      | 2.34 ± 0.95      |
| Suberic, C6                | 0.29 ± 0.18      | 0.24 ± 0.15      | 0.60 ± 0.33      | 0.47 ± 0.32      |
| Azelaic, C9                | 16.3 ± 6.88      | 14.7 ± 5.52      | 16.2 ± 6.44      | 18.4 ± 9.62      |
| Sebacic, C10               | 1.43 ± 0.87      | 0.94 ± 0.27      | 1.87 ± 1.15      | 1.41 ± 0.83      |
| Undecanedioic, C11         | 0.47 ± 0.49      | 0.00 ± 0.00      | 0.87 ± 0.50      | 0.00 ± 0.00      |
| Dodecanedioic, C12         | 0.20 ± 0.27      | 0.14 ± 0.14      | 0.13 ± 0.23      | 0.00 ± 0.00      |
| Methylomalonic, ic4        | 0.19 ± 0.36      | 0.10 ± 0.10      | 0.06 ± 0.70      | 0.02 ± 0.05      |
| Methylsuccinic, ic5        | 1.96 ± 1.10      | 0.86 ± 0.59      | 3.71 ± 1.93      | 3.09 ± 1.11      |
| 2-Methylglutaric, ic6      | 0.88 ± 0.46      | 0.69 ± 0.40      | 1.65 ± 0.70      | 0.93 ± 0.42      |
| Maleic, M                  | 1.12 ± 0.48      | 1.41 ± 0.81      | 1.69 ± 0.68      | 2.08 ± 2.28      |
| Fumaric, F                 | 0.80 ± 0.32      | 1.75 ± 1.79      | 0.97 ± 0.68      | 1.19 ± 0.75      |
| Methylmaleic, mM           | 1.95 ± 1.28      | 6.21 ± 4.55      | 8.48 ± 0.97      | 28.3 ± 9.56      |
| Pthalic, Ph                | 91.8 ± 38.9      | 215.8 ± 86.1     | 36.0 ± 15.3      | 37.2 ± 21.6      |
| Isothallic, iPh            | 3.13 ± 2.01      | 0.67 ± 0.47      | 4.71 ± 2.21      | 2.94 ± 1.40      |
| Terphthalic, iPh           | 31.8 ± 10.9      | 19.1 ± 8.01      | 22.7 ± 15.4      | 10.7 ± 4.85      |
| Malic, hC4                 | 0.40 ± 0.37      | 0.13 ± 0.14      | 0.23 ± 0.10      | 0.24 ± 0.11      |
| Ketomalonic, kC3           | 1.36 ± 0.58      | 1.28 ± 0.84      | 9.54 ± 8.44      | 4.43 ± 1.79      |
| 4-Ketopimelic, kC7         | 1.73 ± 1.04      | 1.56 ± 0.99      | 3.47 ± 2.78      | 1.55 ± 0.91      |

Total: 384 ± 171

Ketocarboxylic acids

| Compounds                  | Winter (µg m$^{-3}$) | Summer (µg m$^{-3}$) |
|----------------------------|----------------------|----------------------|
| Pyruvic                    | 4.68 ± 2.12          | 2.64 ± 1.63          |
| Glyoxylic, oc2             | 13.0 ± 6.70          | 11.0 ± 5.46          |
| 3-Oxopropanoic, oc3        | 0.51 ± 0.31          | 0.19 ± 0.07          |
| 4-Oxobutanonic, oc4        | 4.63 ± 3.61          | 2.31 ± 1.54          |
| 9-Oxononanoic, oc9         | 1.91 ± 1.52          | 3.34 ± 1.64          |
| Total                      | 24.7 ± 12.0          | 19.5 ± 9.59          |
| 3,86 ± 1.78                | 1.96 ± 1.07          | 6.00 ± 2.19          |
| Methylglyoxal, MeGly       | 1.37 ± 1.70          | 0.74 ± 0.51          |
| Total                      | 5.23 ± 3.27          | 5.11 ± 2.13          |
| TQWOC                      | 414.0 ± 184          | 547.0 ± 144          |

Benzoic acid

| Compounds                  | Winter (µg m$^{-3}$) | Summer (µg m$^{-3}$) |
|----------------------------|----------------------|----------------------|
| Benzoic acid               | 159.0 ± 45.6         | 215.0 ± 34.5         |

Fatty acids

| Compounds                  | Winter (µg m$^{-3}$) | Summer (µg m$^{-3}$) |
|----------------------------|----------------------|----------------------|
| Dodecanoic acid, C_{12}    | 2.60 ± 1.28          | 1.95 ± 1.59          |
| Tetraconanoic acid, C_{14} | 7.21 ± 2.97          | 4.80 ± 0.74          |
| Hexadecanoic acid, C_{16}  | 24.0 ± 14.55         | 28.2 ± 7.96          |
| Heptadecanoic acid, C_{17} | 0.48 ± 0.34          | 0.50 ± 0.22          |
| Octadecanoic acid, C_{18}  | 7.95 ± 3.91          | 9.07 ± 3.51          |
| Octadecanoic acid, C_{19}  | 5.89 ± 8.32          | 3.34 ± 1.92          |
| Eicosanoic acid, C_{20}     | 1.20 ± 0.41          | 1.00 ± 0.24          |
| Docosanoic acid, C_{22}     | 2.05 ± 0.74          | 1.47 ± 0.27          |
| Tetracosanoic acid, C_{24}  | 2.05 ± 0.76          | 1.25 ± 0.22          |
| Sum of fatty acids          | 53.4 ± 28.9          | 51.6 ± 13.8          |

OC (µg m$^{-3}$) | 9.29 ± 5.18          | 5.40 ± 2.12          |
| EC (µg m$^{-3}$) | 5.12 ± 4.45          | 2.86 ± 1.66          |
| WSOC (µg m$^{-3}$) | 2.68 ± 1.43          | 1.47 ± 0.51          |
| OC/EC              | 2.08 ± 0.50          | 2.23 ± 0.74          |
| WSOC/OC%            | 2.02 ± 0.73          | 1.47 ± 0.27          |

TQWOC/OC% = 2.88

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3.3 Molecular compositions of fatty acids and benzoic acid

Concentrations of a homologous series of straight chain saturated fatty acids (C_{12,0}-C_{25,0}), unsaturated acid (oleic C_{18:1}) and benzoic acid are shown in Table 1 as well. The total quantified fatty acids concentrations ranged from below MDL (spell out) to 103 ng m\(^{-3}\), with an average of 43.4 \pm 27.3 ng m\(^{-3}\). Hexadecanoic acid (C_{16:0}), octadecanoic acid (C_{18:0}) and oleic acid (C_{18:1}) are the three most abundant fatty acids in PRD. This was consistent with the data reported by Zheng et al. (2000). The odd number fatty acids with C\(_{>19}\) were below MDL in the sites, demonstrating a strong even to odd predominance for fatty acids. Both biogenic and anthropogenic sources are the essential inputs of fatty acids. Microbial activity is one of important biogenic sources (Simonet and Mazurek, 1982). For anthropogenic sources, C\(_{16:0}\), C\(_{18:0}\) and C\(_{18:1}\) are predominantly emitted from the meat cooking while C\(_{16:0}\) can also be directly formed in fossil fuel combustion (Rogge et al., 1991; Schauer et al., 1999, 2002; Zhao et al., 2007a,b). The high concentration of total fatty acids suggests that both cooking and vehicular emissions are important pollution sources in PRD as well as vegetations.

Molecular distributions of fatty acids are characterized by a strong even carbon number predominance with a maximum (C\(_{max}\)) at hexadecanoic acid (C\(_{16:0}\)). Similar distribution patterns of fatty acids have been reported in other urban and rural areas in Hong Kong and China (Zheng et al; 2000; Hou et al., 2006; Fu et al., 2008).

The dominance of even carbon number fatty acid to odd carbon number fatty acid isomers is quantified by Carbon Preference Index (CPI) and is calculated as:

\[
\text{CPI_{fatty acid}} = \frac{\sum \text{Even carbon number fatty acids}}{\sum \text{Odd carbon number fatty acids}}
\]

The predominance of even carbon numbered fatty acids emphasizes that a significant influence from biological sources of aerosols such as microbial activities and epicuticular waxes of vascular plant (Simonet and Mazurek, 1982; Simont, 1984). Here the CPI was calculated with homologous series of fatty acids (C\(_{12,0}-0\)to C\(_{25,0}\)). The CPI values of the fatty acids are 19.4, 13.8, 40.4 and 3.26, in GZ, ZQ, PU and HT, respectively. The CPI values are higher in summer than in winter indicate that biogenic source has a larger contribution in hot weather.

Octadecenoic acid (oleic acid, C\(_{18:1}\)) was detected in most of the urban samples which had concentrations ranging from below MDL to 26 ng m\(^{-3}\) (4.1 \pm 4.7 ng m\(^{-3}\) on average) in PRD. Automobile engine exhaust is one of the pollution sources for C\(_{18:1}\) (Rogge et al., 1993b). Ratio of C\(_{18:1}\) to C\(_{18:0}\) can be used as an indicator for aerosol aging. A lower ratio was observed in aged aerosol as unsaturated fatty acids can be photo-chemically degraded, while saturated fatty acids are more stable in the atmosphere (Kawamura and Gagosian, 1987; Wang et al., 2006). In PRD, an average of C\(_{18:1}/C_{18:0}\) ratio was 0.53 \pm 0.39, suggesting an enhanced photochemical degradation of unsaturated fatty acid.

Benzoic acid was detected in most of the samples in PRD, with an average concentration of 165 \pm 48 ng m\(^{-3}\). Benzoic acid was proposed to be a primary pollutant in the motor vehicles exhaust (Kawamura et al., 1985; Rogge et al., 1993b), and a secondary product from photochemical degradation of aromatic hydrocarbons such as toluene emitted by automobiles (Suh et al., 2003). Guo et al. (2004) found that high daily concentration of toluene, with a maximum of 53 \(\mu\)g m\(^{-3}\), was determined in Hong Kong. This suggests that major portion of benzoic acid in the PRD aerosols is probably produced by the oxidation of toluene in the atmosphere.

3.4 Summer/winter variations and spatial distribution

Summer/winter variations of the organic species were found in the four sampling sites. TQWOC concentrations ranged from 145 to 1340 ng m\(^{-3}\) (544 ng m\(^{-3}\) on average) in winter and from 99 to 665 ng m\(^{-3}\) (318 ng m\(^{-3}\) on average) in summer. These values are similar to those (90–1370 ng m\(^{-3}\), 480 ng m\(^{-3}\) on average) reported in urban Tokyo (Kawamura and Ikushima, 1993), but are lower to those reported in other urban cities in China (Ho et al., 2007). Total ketocarboxylic acid concentrations ranged from 4.5 to 178 ng m\(^{-3}\) (43.9 ng m\(^{-3}\) on average) in winter and from 0.59 to 207 ng m\(^{-3}\) (42.3 ng m\(^{-3}\) on average) in summer, while total dicarboxylic acid concentrations ranged from 1.3 to 88.6 ng m\(^{-3}\) (11.0 ng m\(^{-3}\) on average) in winter and from 0.15 to 68.2 ng m\(^{-3}\) (11.6 ng m\(^{-3}\) on average) in summer. These concentrations are similar to those reported (ketocarboxylic acids = 53 ng m\(^{-3}\); dicarboxyls = 12 ng m\(^{-3}\)) at the Gosal site on Jeju Island, South Korea (Kawamura et al., 2004). Total quantified fatty acids ranged from 2.9 to 103 ng m\(^{-3}\) (45.3 ng m\(^{-3}\) on average) in winter and from below MDL to 96.1 ng m\(^{-3}\) (41.3 ng m\(^{-3}\) on average) in summer, while benzoic acid concentrations ranged from 101 to 256 ng m\(^{-3}\) (157 ng m\(^{-3}\) on average) in winter and from 83.9 to 306 ng m\(^{-3}\) (175 ng m\(^{-3}\) on average) in summer.

In order to investigate the transport and the source region of air pollutants, 2-day air mass back trajectory analyses were conducted using NOAA HY-SPLIT model (HYbrid Single-Particle Lagrangian Integrated Trajectory, NOAA/ARL) with a starting elevation of 100 m. In winter, prevailing northeasterly wind travels across South China before reaching Hong Kong (Fig. 2a). Total dicarboxylic acids were the most abundant in downwind location of Hong Kong (i.e., PU and HT sites). Poor air quality in Hong Kong in winter is due to the influence of local sources and polluted air mass transported from South China. In contrast, the highest dicarboxylic concentration was found in the downwind locations in GZ and ZQ during summer, when prevailing southwesterly wind brings warm and damp air masses from the South China Sea through Hong Kong to PRD region.
Comparisons of winter and summer concentrations for both individual quantified compounds are shown in Table 1 while comparisons of winter and summer concentrations for TQWOC and sum of fatty acids are shown in Fig. 3. The highest average concentration of the TQWOC was found at PU in winter, which is attributable to the mixed contribution of local and regional sources. The concentrations of the organic species in winter were statistically higher in PU than in GZ. The high abundances of organic aerosols in downwind urban locations (PU) are due to the emissions from urban local sources and regional long-range transport from PRD when the air mass came from the north. Reversibly, statistically lower concentrations of the TQWOC were found at PU and HT in summer because of the upwind locations. The local emission sources were diluted by marine air masses transported from the South China Sea. In contrast, the concentrations of the TQWOC at downwind locations (GZ and QZ) were 2–3 times higher than those in Hong Kong during summer. Concentrations of total quantified fatty acids at urban sites (GZ, ZQ and PU) were found to be 3–23 times higher than those at background site (HT).

TQWOC were normalized by OC and WSOC to better discuss summer/winter variations (Fig. 3). The relative abundances of TQWOC in WSOC were higher in summer (5.9–50.7%, 20.2 ± 10.3% on average) than in winter (winter: 2.2–40.8, 9.6 ± 7.6% on average) while the relative abundances of TQWOC in OC were also higher in summer (0.9–12.4%, 4.5 ± 2.7% on average) than in winter (winter: 1.1–5.7, 2.5 ± 1.2% on average) except at PU site, being consistent with enhanced secondary production of dicarboxylic acids under warmer weather conditions.

### 3.5 Correlation analysis and the ratios of selected species

Low molecular weight dicarboxylic acids can be primarily produced from anthropogenic emissions. Photochemical reactions in the atmosphere also play an important role in the formation of dicarboxylic acids. The dicarboxylic acids are secondarily generated in the atmosphere by photochemical chain reactions of unsaturated hydrocarbons or fatty acids as well as their oxidation products (Kawamura and Sakaguchi, 1999; Kawamura et al., 1996b), even though their formation mechanisms are poorly understood. The correlation coefficients of selected species were examined in different sites in both seasons. Table 2 shows the correlation coefficients of selected dicarboxylic acids, ketocarboxylic acids and α-dicarbonyls. Strong correlations were observed for C$_2$, C$_3$, C$_4$, C$_9$, ωC$_2$ and ωC$_9$ in downwind locations, that is, GZ and ZQ in summer and PU and HT in winter, respectively. Other than direct vehicular emission, photochemical processes control the atmospheric concentrations of these species. For instance, ωC$_2$, the most abundant ketocarboxylic acid, can be further oxidized to C$_2$ dicarboxylic acid; thus a good correlation was found between ωC$_2$ and C$_2$ (r = 0.93, P < 0.01 in GZ; r = 0.98, P < 0.01 in ZQ in summer; and r = 0.42, P < 0.5 in HK; r = 0.89, P < 0.01 in HT in winter).

Furthermore, positive correlations of ωC$_2$ with Gly (r = 0.85 – 0.99, P < 0.01) are observed in the downwind sites. This is consistent with the atmospheric oxidation process proposed for Gly to ωC$_2$ (Kawamura et al., 1996c). Malonic (C$_3$) and succinic (C$_4$) acids can be oxidized to C$_2$ via the breakdown of intermediates such as ketomalonic acid (kC$_3$) (Kawamura and Ikushima, 1993), thus strong...
Table 2. Correlation coefficients of selected dicarboxylic acids, ketocarboxylic acids and dicarbonyls at four sampling sites in PRD in winter and summer seasons.

|       | GZ winter | C2 | C3 | C4 | C9 | kC3 | GZ summer | C2 | C3 | C4 | C9 | kC3 |
|-------|-----------|----|----|----|----|-----|-----------|----|----|----|----|-----|
| C2    | 1.00      | 0.74* | 0.77* | 0.78* | 0.46 | 0.82* | 0.25 | 0.62 |     |    |     |
| C3    | 1.00      | 0.95** | 0.74* | 0.81* | 0.84** | 0.00 | 0.63 |     |    |    |    |
| C4    | 1.00      | 0.83* | -0.08 | 0.77* |       |     |     |     |    |    |    |
| C9    | 1.00      | 0.28 | 0.98** | -0.06 | 0.92** |     |     |     |    |    |    |
| kC3   | 1.00      | 0.44 | 0.05 | 0.13 |       |     |     |     |    |    |    |
| ωC2   | 1.00      | 0.72 | 0.85** | -0.14 |       |     |     |     |    |    |    |
| ωC9   | 1.00      | 0.76* |       |       |     |     |     |     |    |    |    |
| Gly   | 1.00      |     |     |       |     |     |     |     |    |    |    |

|       | ZQ winter | C2 | C3 | C4 | C9 | kC3 | ZQ summer | C2 | C3 | C4 | C9 | kC3 |
|-------|-----------|----|----|----|----|-----|-----------|----|----|----|----|-----|
| C2    | 1.00      | 0.67 | 0.84** | 0.87** | 0.67 | 0.97** | 0.68 | 0.94** |     |    |    |
| C3    | 1.00      | 0.47 | 0.69 | 0.32 | 0.65 | 0.96** | 0.57 |     |    |    |    |
| C4    | 1.00      | 0.49 | 0.94** | 0.85** | 0.43 | 0.88** |     |     |    |    |    |
| C9    | 1.00      | 0.24 | 0.79* | 0.72* | 0.71* |     |     |     |    |    |    |
| kC3   | 1.00      | 0.75* | 0.31 | 0.83* |     |     |     |     |    |    |    |
| ωC2   | 1.00      | 0.68 | 0.99** |       |     |     |     |     |    |    |    |
| ωC9   | 1.00      | 0.62 |       |       |     |     |     |     |    |    |    |
| Gly   | 1.00      |     |     |       |     |     |     |     |    |    |    |

|       | PU winter | C2 | C3 | C4 | C9 | kC3 | PU summer | C2 | C3 | C4 | C9 | kC3 |
|-------|-----------|----|----|----|----|-----|-----------|----|----|----|----|-----|
| C2    | 1.00      | 0.94** | 0.92** | 0.80* | 0.54 | 0.42 | 0.56 | 0.42 |     |    |    |
| C3    | 1.00      | 0.98** | 0.81* | 0.67 | 0.41 | 0.44 | 0.33 |     |    |    |    |
| C4    | 1.00      | 0.74* | 0.50 | 0.45 | 0.37 | 0.34 |     |     |    |    |    |
| C9    | 1.00      | 0.73* | 0.69 | 0.85** | 0.61 |     |     |     |    |    |    |
| kC3   | 1.00      | 0.18 | 0.53 | 0.29 |     |     |     |     |    |    |    |

* Correlation is significant at the 0.05 level (2-tailed)
** Correlation is significant at the 0.01 level (2-tailed)
correlations were observed among C₂, C₃ and C₄ in this study. Other acids such as fumaric (F), maleic (M) and methylmaleic (mM) acids are fairly correlated each other \((r = 0.61 - 0.70, P < 0.1)\). These three dicarboxylic acids are known to be the photooxidation products of toluene, benzene, and xylene. Maleic acid (M) can isomerize to trans-fumaric acid (F) by photochemical transformations (Kawamura and Ikushima, 1993). C₂, sometimes regarded as secondary organic aerosol tracer, has fair correlation with sulfate \((r = 0.66, P < 0.1)\), which is consistent with previous studies. Yu et al. (2005) argue that in-cloud processing has been established to be the dominant formation pathway for oxalate.

Kawamura and Ikushima (1993) suggested that ratio of C₃ to C₄ can be used as an indicator of enhanced photochemical production of dicarboxylic acids. It is known that C₄ can serve as a precursor of C₃. In this study, C₃/C₄ ratios ranged between 0.24 and 5.42 with an average of 1.29, which are higher than those reported from vehicular emissions (0.3–0.5) (Kawamura and Kaplan, 1987), and for aerosols in Northern China in summer (0.61) and winter (1.12) (Ho et al., 2007). Our findings also suggest that in addition to primary exhaust, secondary formation of particulate dicarboxylic acids by photo-oxidation reaction is also important in PRD. The (F+M+mM)/EC ratios at the downwind sampling locations were much higher than those at the upwind sampling sites. The elevated abundance of M, F and mM in aged aerosols indicates that the photooxidation of aromatic compounds to F, M, and mM is important during long-range transport. Good correlations \((r = 0.69, P < 0.1\) in downwind sites; \(r = 0.85, P < 0.01\) in upwind sites) were observed between the TQWOC and WSO (Fig. 4). These results suggest that dicarboxylic acids, ketocarboxylic acids and dicarbonyls are the major water-soluble organic species in PRD, which are linked to the photochemical chain reactions. The TQWOC contributed more than 15\% of WSO in downwind sites (except HT), suggesting that the water-soluble organic species are one of the major contributors of WSO in PRD.

Fig. 3. Box plots of TQWOC and sum of fatty acid concentrations (ng m⁻³) as well as TQWOC/WSOC and TQWOC/OC ratios in PM₂.₅ samples during winter and summer in four cities in PRDR.
It is reasonable because there is a sufficient duration for the precursors to form SOC during the long-distance transport to the downwind sampling locations.

4 Summary and conclusions

Molecular compositions of low molecular weight dicarboxylic acids (C$_2$–C$_{12}$), ketocarboxylic acids (oC$_2$–oC$_9$, pyruvic acid), α-dicarboxyls (C$_2$–C$_3$), fatty acids (C$_{12}$–C$_{25}$) and benzoic acid were studied in PM$_{2.5}$ samples collected from four sampling locations in PRD during the winter and summer to better understand their spatial and seasonal variations of water-soluble organic species. Oxalic (C$_2$) acid was found as the most abundant dicacid, followed by phthalic acid (Ph) which are similar to other urban cites in China (Ho et al., 2007). The TQWOC contributed a significant fraction in WSOC (14.3 ± 10.3%). The fatty acids had an average total concentration of 43.4 ± 27.3 ng m$^{-3}$ and are derived from both biogenic and anthropogenic sources. The strong even carbon number predominance in fatty acid distributions represents significant influences from biological sources such as microbial activities and epicuticular waxes of vascular plant in PRD region. Octadecanoic acid (oleic acid, C$_{18:1}$) was detected in most of the urban samples with concentrations ranging from below MDL to 26 ng m$^{-3}$ (4.1 ± 4.7 ng m$^{-3}$ on average) in PRD. Automobile engine exhaust may be one of the pollution sources for C$_{18:1}$.

The concentrations of the organic species in winter were generally higher in PU and HT than in GZ and ZQ. The high abundances of organic aerosols in downwind locations (PU and HT) are due to the emissions from urban local sources and long-range transport from PRD when the air mass came from the north. In contrast, lower concentrations of the TQWOC were found at the upwind locations of PU and HT in summer. The local emission sources were diluted by marine air masses transported from South China Sea. However, the relative abundances of TQWOC in OC were higher in summer (0.9–12.4%, 4.5 ± 2.7% on average) than in winter (1.1–5.7, 2.5 ± 1.2% on average) except at PU site, being consistent with enhanced secondary production of dicarboxylic acids under warmer weather conditions. These spatial and seasonal variations are consistent with photochemical production and the subsequent accumulation under different meteorological conditions.

Relatively high C$_3$/C$_4$ ratios (0.24–5.42 with an average of 1.29) were found for the molecular distributions of dicarboxylic acids in this study, further suggesting that in addition to primary emissions from vehicular emissions, secondary formation of particulate dicarboxylic acids via photooxidation reaction is important in PRD. Good correlations ($r = 0.69$ in downwind sites; $r = 0.85$ in upwind sites) were observed between the TQWOC and WSOC, suggesting that dicarboxylic acids, ketocarboxylic acids and dicarbonyls are the major water-soluble organic species in PRD. The TQWOC contributed more than 15% of WSOC in downwind sites (except HT), suggesting that the water-soluble organic species are one of the major contributors of WSOC in PRD. In light of WSOC being the most abundant component of PM$_2.5$, future work is suggested to further speciate and quantify this fraction (e.g., humic-like substances (HULIS)) in PRD and other megacities in China.

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References

Bunce, N. J., Liu, L., Zhu, J., and Lane, D. A.: Reaction of naphthalene and its derivatives with hydroxyl radicals in the gas phase, Environ. Sci. Technol., 31, 2252–2259, 1997.

Cao, J. J., Lee, S. C., Ho, K. F., Zhang, X. Y., Zou, S. C., Fung, K., Chow, J. C., and Watson, J. G.: Characteristics of carbonaceous aerosol in Pearl River Delta Region, China during 2001 winter period, Atmos. Environ., 37, 1451–1460, 2003a.

Cao, J. J., Lee, S. C., Ho, K. F., Zou, S. C., Zhang, X. Y., and Pan, J. G.: Spatial and seasonal distributions of atmospheric carbonaceous aerosols in Pearl River Delta Region, China, China Particulol., 1, 33–37, 2003b.

Chebbi, A. and Carlier, P.: Carboxylic acids in the troposphere occurrence, sources, and sinks: A review, Atmos. Environ., 30, 4233–4249, 1996.

Chow, J. C., Watson, J. G., Chen, L.-W. A., Arnott, W. P., Moosmüller, H., and Fung, K. K.: Equivalence of elemental carbon by thermal/optical reflectance and transmittance with different temperature protocols, Environ. Sci. Technol., 38, 4414–4422, 2004.

Chow, J. C., Watson, J. G., Chen, L.-W. A., Paredes-Miranda, G., Chang, M.-C. O., Trimble, D., Fung, K. K., Zhang, H., and Zhen Yu, J.: Refining temperature measures in thermal/optical carbon analysis, Atmos. Chem. Phys., 5, 2961–2972, doi:10.5194/acp-5-2961-2005, 2005.

Facchini, M. C., Mircea, M., Fuzzi, S. and Charlson, R. J.: Cloud albedo enhancement by surface-active organic solutes in growing droplets, Nature, 410, 257–259, 1999.

Fang, M., Zheng, M., Wang, F., To, K. L., Jaafar, A. B., and Tong, S. L.: The solvent-extractable organic compounds in the Indonesia biomass burning aerosols-characterization studies, Atmos. Environ., 33, 83–795, 1999.

Fine, P. M., Cass, G. R., and Simonet, B. R. T.: Chemical characterization of fine particle emissions from the fireplace combustion.
of woods grown in the southern United States, Environ. Sci. Technol., 36, 1442–1451, 2002.

Feng, J. L., Hu, M., Chan, C. K., Lau, P. S., Fang, M., He, L. Y., and Tang, X. Y.: A comparative study of the organic matter in PM2.5 from three Chinese megacities in three different climatic zones, Atmos. Environ., 40, 3983–3994, 2006.

Feng, J. L., Guo, Z. G., Chan, C. K., and Fang, M.: Properties of organic matter in PM2.5 at Changdao Island, China – A rural site in the transport path of the Asian continental outflow, Atmos. Environ., 41, 1924–1935, 2007.

Fu, P., Kawamura, K., Okuzawa, K., Aggarwal, S. G., Wang, G. H., Kanaya, Y., and Wang, Z.: Organic molecular compositions and temporal variations of summertime mountain aerosols over Mt. Tai, North China Plain, J. Geophys. Res., 113, D19107, doi:10.1029/2008JD009900, 2008.

Guo, H., Lee, S. C., Louie, P. K. K., Ho, K. F.: Characterization of hydrocarbons, halocarbons and carbonyls in the atmosphere of Hong Kong, Chemosphere 57, 1363–1372, 2004.

He, L.-Y., Hu, M., Huang, X.-F., Yu, B.-D., Zhang, Y.-H., and Liu, D.-Q.: Measurement of emissions of fine particulate organic matter from Chinese cooking, Atmos. Environ., 38, 6557–6564, 2004.

He, L.-Y., Hu, M., Huang, X.-F., Zhang, Y.-H., and Tang, X.-Y.: Seasonal pollution characteristics of organic compounds in atmospheric fine particles in Beijing, Sci. Total Environ., 359, 167–176, 2006.

Ho, K. F., Cao, J. J., Lee, S. C., Kawamura, K., Zhang, R. J., Chow, J. C., and Watson, J. G.: Dicarboxylic acids, ketocarboxylic acids and dicarbonyls in 14 cities of China, J. Geophys. Res., 112, D22S27, doi:10.1029/2009JD009900, 2010.

Ho, K. F., Lee, S. C., Ho, S. S. H., Kawamura, K., Tachibana, E., Cheng, Y., and Zhu, T.: Dicarboxylic Acids, Ketocarboxylic acids, α-dicarbonyls, Fatty Acids and Benzoic Acid in urban aerosols collected during 2006 Campaign of Air Quality Research in Beijing (CAREBeijing-2006), J. Geophys. Res., 115, D19312, doi:10.1029/2010JD001304, 2010.

Hou, X. M., Zhuang, G. S., Sun, Y. L., and An, Z. S.: Characteristics and sources of polycyclic aromatic hydrocarbons and fatty acids in PM2.5 aerosols in dust season in China, Atmos. Environ., 40, 3251–3262, 2006.

Jang, M. and McDow, S. R.: Products of benz[a]anthracene photodegradation in the presence of known organic constituents of atmospheric aerosols, Environ. Sci. Technol., 31, 1046–1053, 1997.

Kawamura, K.: Identification of C2−C10 α-oxocarboxylic acids, pyruvic acid, and C2−C3 α-dicarbonyls in wet precipitation and aerosol samples by capillary GC and GC/MS, Anal. Chem., 65, 3505–3511, 1993.

Kawamura, K., Ng, L.-L., and Kaplan, I. R.: Determination of organic acids (C1−C10) in the atmosphere, motor exhausts, and engine oils, Environ. Sci. Technol., 19, 1082–1086, doi:10.1021/es00141a010, 1985.

Kawamura, K. and Kaplan, I. R.: Motor exhaust emissions as a primary source for dicarboxylic acids in Los Angeles ambient air, Environ. Sci. Technol., 21, 105–110, 1987.

Kawamura, K. and Gagosian, R. B.: Implications of α-oxocarboxylic acids in the remote marine atmosphere for photo-oxidation of unsaturated fatty acids, Nature, 325, 330–332, 1987.

Kawamura, K. and Ikushima, K.: Seasonal changes in the distribution of dicarboxylic acids in the urban atmosphere, Environ. Sci. Technol., 27, 2227–2235, 1993.

Kawamura, K., Sempéré, R., Imai, Y., Hayashi, M., and Fujii, Y.: Water soluble dicarboxylic acids and related compounds in Antarctic aerosols, J. Geophys. Res., 101, 18721–18728, 1996a.

Kawamura, K., Steinberg, S., and Kaplan, I. R.: Concentrations of monocarboxylic and dicarboxylic acids and aldehydes in Southern California wet precipitations: comparison of urban and non-urban samples and compositional changes during scavenging, Atmos. Environ., 30, 1035–1052, 1996b.

Kawamura, K., Kasukabe, H., and Barrie, L. A.: Source and reaction pathways of dicarboxylic acids, ketoacids and dicarbonyls in Arctic aerosols: one year of observations, Atmos. Environ., 30, 1709–1722, 1996c.

Kawamura, K. and Sakaguchi, F.: Molecular distribution of water soluble dicarboxylic acids in marine aerosols over the Pacific Ocean including tropics, J. Geophys. Res., 104, 3501–3509, 1999.

Kawamura, K., Kobayashi, M., Tsuennonuma, N., Mochida, M., Watanabe, T., and Lee, M.: Organic and inorganic compositions of marine aerosols from East Asia: Seasonal variations of watersoluble dicarboxylic acids, major ions, total carbon and nitrogen, and stable C and N isotopic composition. In: Geochemical Investigation in Earth and Space Science: A Tribute to Isaac R. Kaplan, edited by: Hill, R. J., Leventhal, J., AizenstaiJ. C., Seinfeld, J. H.: Chamber studies of secondary organic aerosol growth by reactive uptake of simple carbonyl compounds, J. Geophys. Res., 110, D23207, doi:10.1029/2005JD006004, 2005.

Lee, S. C., Ho, K. F., Chan, L. Y., Zielinska, B., and Chow, J. C.: PAHs and carbonyl compounds in the urban Tokyo atmosphere, Atmos. Environ., 39, 1945–1960, 2005.

Kerminen, V.-M.: Relative roles of secondary sulfate and organic aerosols in atmospheric cloud condensation nuclei production, J. Geophys. Res., 106, 17321–17333, 2001.

Kerminen, V.-M., Ojanen, C., Pakkanen, T., Hillamo, R., Aurela, M., and Merilainen, J.: Low-molecular-weight dicarboxylic acids in an urban and rural atmosphere, J. Aerosol Sci., 31, 349–362, 2000.

Kroll, J. H., Ng, N. L., Murphy, S. M., Varutbangkul, V., Flagan, R. C., Seinfeld, J. H.: Chamber studies of secondary organic aerosol growth by reactive uptake of simple carbonyl compounds, J. Geophys. Res., 110, D23207, doi:10.1029/2005JD006004, 2005.

Kerminen, V.-M., Ojanen, C., Pakkanen, T., Hillamo, R., Aurela, M., and Merilainen, J.: Low-molecular-weight dicarboxylic acids in an urban and rural atmosphere, J. Aerosol Sci., 31, 349–362, 2000.

Li, Y. C. and Yu, J. Z.: Simultaneous Determination of Mono- and Di-carboxylic Acids, α-Oxo-carboxylic Acids, Mid-chain Ketocarboxylic acids, and Aldehydes in Atmospheric Aerosol Samples, Environ. Sci. Technol. 39, 7616–7624, 2005.

Li, J., Zhuang, G. S., Huang, K., Lin, Y. F., Xu, C., and Yu, S. L.: Characteristics and sources of air-borne particulate in Urumqi, China, the upstream area of Asia dust, Atmos. Environ., 42, 776–787, 2008.

Liggio, J., Li, S.-M., and McLaren, R.: Reactive uptake of glycolal by particulate matter, J. Geophys. Res., 110, D10304, doi:10.1029/2004JD005113, 2005.

Menon, S., Hansen, J., Nazarenko, L. and Luo, Y.: Climate effects of black carbon aerosols in China, Science, 297, 2250–225, 2002.
Mochida, M., Umemoto, N., Kawamura, K., Lim, H., and Turpin, B. J.: Bimodal size distributions of various organic acids and fatty acids in the marine atmosphere: Influence of anthropogenic aerosols, Asian dusts, and sea spray off the coast of East Asia, J. Geophys. Res., 112, D15209, doi:10.1029/2006JD007773, 2007.

Nel, A.: Air pollution-related illness: Effects of particles, Science, 308, 804–806, 2005.

Nolte, C. G., Schauer, J. J., Cass, G. R., Simoneit, B. R. T.: Highly polar organic compounds present in wood smoke and in the ambient atmosphere, Environ. Sci. Technol., 35, 11912–11919, 2001.

Nolte, C. G., Schauer, J. J., Cass, G. R., Simoneit, B. R. T.: Trimethylsilyl derivatives of organic compounds in source samples and in atmospheric fine particulate matter, Environ. Sci. Technol., 36, 4273–4281, 2002.

Novakov, T., Menon, S., Kirchstetter, T. W., Koch, D., and Hansen, J. E.: Aerosol organic carbon to black carbon ratios: analysis of published data and implications for climate forcing, J. Geophys. Res., 110, D21205, 2005.

Olivera, C., Pio, C., Alves, C., Evtyugina, M., Santos, P., Gonçalves, V., Nunes, T., Silvestre, A. J. D., Palme, F., Wällin, P., and Harrad, S.: Seasonal distribution of polar organic compounds in the urban atmosphere of two large cities from the North and South of Europe, Atmos. Environ., 41, 5555–5570, 2007.

Penner, J. E. and Novakov, T.: Carbonaceous particles in the atmosphere: a historical perspective to the fifth international conference on carbonaceous particles in the atmosphere, J. Geophys. Res., 101, 19373–19378, 1996.

Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R. T.: Sources of fine organic aerosol. 1. Charbroilers and meat cooking operations, Environ. Sci. Technol., 25, 1112–1125, 1991.

Rogge, W. F., Mazurek, M. A., Hildemann, L. M., Cass, G. R., and Simoneit, B. R. T.: Quantification of urban organic aerosols at a molecular level: identification, abundance and seasonal variation, Atmos. Environ., 27A, 1309–1330, 1993a.

Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., Simoneit, B. R. T.: Sources of fine organic aerosol. 2. Noncatalyst and catalyst-equipped automobiles and heavy duty diesel trucks, Environ. Sci. Technol., 27, 636–651, 1993b.

Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R. T.: Sources of fine organic aerosol 9. Pine, oak, and synthetic log combustion in residential fireplaces, Environ. Sci. Technol., 32, 13–22, 1998.

Rogge, W. F., Medeiros, P. M., and Simoneit, B. R. T.: Organic marker compounds for surface soil and fugitive dust from open lot dairies and cattle feedlots, Atmos. Environ., 40, 27–49, 2006.

Saxena, P. and Hildemann, L.: Water-soluble organics in atmospheric particles: a critical review of the literature and application of thermodynamics to identify candidates compounds, J. Atmos. Chem. 24, 57–109, 1996.

Saxena, P., Hildemann, L. M., Mcmurry, P. H., and Seinfeld, J. H.: Organics after hygroscopic behavior of atmospheric particles, J. Geophys. Res., 100, 18755–18770, 1995.

Schauer, J. J., Kleeman, M. J., Cass, G. R., Simoneit, B. R. T.: Measurement of emissions from air pollution sources, 1. C1 through C29 organic compounds from meat charbroiling, Environ. Sci. Technol., 33, 1566–1577, 1999.

Schauer, J. J., Kleeman, M. J., Cass, G. R., and Simoneit, B. R. T.: Measurement of emissions from air pollution sources, 3. C1 – C29 organic compounds from fireplace combustion of wood, Environ. Sci. Technol., 35, 1716–1728, 2001.

Schauer, J. J., Kleeman, M. R., and Simoneit, B. R. T.: Measurement of emissions from air pollution sources, 4. C1 – C27 organic compounds from cooling with seed oils, Environ. Sci. Technol., 36, 567–575, 2002.

Simoneit, B. R. T.: Organic matter of the troposphere-III. Characterization and sources of petroleum and pyrogenic residues in aerosols over the western United States, Atmos. Environ., 18, 51–67, 1984.

Simoneit, B. R. T.: Biomass burning – a review of organic tracers for smoke from incomplete combustion, Appl. Geochem, 17, 129–162, 2002.

Simoneit, B. R. T. and Mazurek, M. A.: Organic matter of the troposphere II. Natural background of biogenic lipid matter in aerosols over the rural western US, Atmos. Environ., 16, 2139–2159, 1982.

Simoneit, B. R. T., Kobayashi, M., Mochida, M., Kawamura, K., and Huebert, B. J.: Aerosol particles collected on aircraft flights over the northwestern Pacific region during the ACE-Asia campaign: Composition and major sources of the organic compounds, J. Geophys. Res., 109, D19S09, doi:10.1029/2004JD004565, 2004.

Sin, D. W. M., Fung, W. H., Choi, Y. Y., Lam, C. H., Louie, P. K. K., Chow, J. C., and Watson, J. G.: Seasonal and spatial variation of solvent extractable organic compounds in fine suspended particulate matter in Hong Kong, JAWMA 55, 291–301, 2005.

Suh, I., Zhang, R., Molina, L. T., and Molina, M. J.: Oxidation mechanism of aromatic peroxo and bicyclic radicals from OH-toluene reactions, J. Am. Chem. Soc., 125, 12655–12665, doi:10.1021/ja0350280, 2003.

Wang, G. H., Kawamura, K., Lee, S. C., Ho, K. F., and Cao, J. J.: Molecular, seasonal and spatial distribution of organic aerosols from fourteen Chinese cities, Environ. Sci. Technol., 40, 4619–4625, 2006.

Wilkening, K. E., Barrie, L. A., and Engle, M.: Trans-Pacific air pollution, Science, 290, 65–66, 2000.

Yang, H., Yu, J. Z., Ho, S. S. H., Xu, J., Wu, W. S., Wán, C. H., Wang, X., Wang, X., and Wang, L.: The chemical composition of inorganic and carbonaceous materials in PM2.5 in Nanjing, China, Atmos. Environ., 39, 3735–3749, 2005.

Yu, J. Z., Huang, X. F., Xu, J. H., and Hu, M.: When aerosol sulfate goes up, so does oxalate: Implication for the formation mechanism of oxalate. Environ. Sci. Technol., 39, 128–133, 2005.

Zhao, Y. L., Hu, M., Slanina, S., and Zhang, Y. H.: Chemical compositions of fine particulate organic matter emitted from Chinese cooking, Environ. Sci. Technol., 41, 99–105, 2007.

Zehng, M., Fang, M., Wang, F., and To, K. L.: Characterization of the solvent extractable organic compounds in PM2.5 aerosols in Hong Kong, Atmos. Environ., 34, 2691–2702, 2000.