Ash generation and distribution from the April-May 2010 eruption of Eyjafjallajökull, Iceland

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The 39-day long eruption at the summit of Eyjafjallajökull volcano in April–May 2010 was of modest size but ash was widely dispersed. By combining data from ground surveys and remote sensing we show that the erupted material was $4.8 \pm 1.2 \times 10^{11}$ kg (benmoreite and trachyte, dense rock equivalent volume $0.18 \pm 0.05$ km$^3$). About 20% was lava and water-transported tephra, 80% was airborne tephra (bulk volume $0.27$ km$^3$) transported by $3–10$ km high plumes. The airborne tephra was mostly fine ash (diameter <1000 µm). At least $7 \times 10^{10}$ kg (70 Tg) was very fine ash (<28 µm), several times more than previously estimated via satellite retrievals. About 50% of the tephra fell in Iceland with the remainder carried towards south and east, detected over ~7 million km$^2$ in Europe and the North Atlantic. Of order $10^{10}$ kg (2%) are considered to have been transported longer than 600–700 km with <10$^8$ kg (<0.02%) reaching mainland Europe.

Tephra dispersal from explosive eruptions is a function of multiple factors, including magma discharge, degree of magma fragmentation, vent geometry, plume height, particle size distribution, and wind velocity. Plinian eruptions, with typical duration of hours, are characterized by magma discharge of $10^7–10^8$ kg s$^{-1}$ supporting 20–45 km high plumes often producing tephra falls covering $\geq 10^7$ km$^2$ (ref. 1–2). Such dispersal has hitherto not been anticipated for smaller explosive events. The explosive summit eruption of Eyjafjallajökull in Iceland in 14 April–22 May 2010 was a moderate size event (VEI = 3 based on maximum plume height and magma discharge) typified by pulsating explosive activity. The magma discharge never came close to that of Plinian eruptions, but it had a sustained character and duration of weeks. Its ash dispersal was also greater than previously reported for an event of this magnitude. Plinian and other explosive eruptions in Iceland have repeatedly produced ash plumes transported over thousands of kilometres and the latest pre-2010 event to do so and produce plumes that reached Europe was the 50-min long initial explosive phase of the 1947 Hekla eruption3. However, since the introduction in the 1960s of jet-engine powered passenger aircraft, which are vulnerable to exposure to volcanic ash, plumes from explosive eruptions in Iceland were not dispersed towards Europe until Eyjafjallajökull erupted in 2010.

The combination of a prolonged and sustained eruption of fine ash and persistent northwesterly winds transporting the ash towards southeast, resulted in dispersal of ash over a large part of Europe, even though concentrations were everywhere quite low$^4–7$. The resulting unprecedented disruption to air traffic and the implications for vulnerability of modern society to even relatively modest eruptions, have made Eyjafjallajökull 2010 a landmark event. The detail of distal detection of minor amounts of ash in the atmosphere through the use of various techniques and observations is without precedent, and provide unique data for various ash and atmospheric studies. Hence, accurate definition of the source parameters of this eruption is important to realise the full potential of these various data. The aims of this paper are therefore to report and analyse:

1. the main phases of the eruption, their characteristics and the in situ measured magnitude and distribution of volcanic ash;
2. the grain sizes observed in the first and most powerful explosive phase of the eruption;
3. in situ measured quantity of fine grained ash and compare with observations of airborne ash concentrations obtained by remote sensing;
4. the lessons learned about the ash dispersal from modest explosive eruptions in Iceland.

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**Results**

**Eruption overview.** When signs of unrest were detected in 1992, the volcano of Eyjafjallajökull had been quiet since the eruption of 1821-23. The pre-2010 unrest was signified by discrete periods of seismic swarms and uplift and culminated in two eruptions in 2010 (ref. 8). Inflation in 1994 and 1999–2000 was interpreted as resulting from intrusions at 6–8 km depth in the crust under the volcano8,9. Intense seismicity and uplift in January-March 2010 (ref. 8) was followed by a small basaltic flank eruption on 20 March–12 April10, forming a 1.3 km² lava flow field with a volume of 0.020 km³ (Fig. 1). The summit eruption started at ca. 1:30 UTC on 14 April and was preceded by an intense three-hour-long earthquake swarm8.

Excluding the brief (3–4 hours) subglacial phase when the eruption melted its way through the 200 m thick glacier in the summit caldera, the eruption can be divided into four distinct phases (Fig. 2). The results on Figure 2 are based on combination of fallout data (Figs. 3, 4 and 5) and models of tephra dispersal outside Iceland (Fig. 6). Magma discharges were obtained from combining information from tephra fallout maps (Figs. 3–5) and plume height variations with time (see Methods):

I. **Initial/First explosive phase (14–18 April).** This phase was characterised by magma discharge of 5–10×10⁵ kg s⁻¹ (≈200–400 m³ s⁻¹ DRE) from a water-filled vent, supporting

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**Figure 1 | Setting of the Eyjafjallajökull eruption.** (a) The eruption sites, ice cauldrons and pathways of jökulhlaups (glacial floods). Inset on right: an airborne synthetic aperture radar image from 15 April. Inset on left: Iceland and location of Eyjafjallajökull. (b) A north-south cross-section through the main craters within the summit caldera showing the two main ice cauldrons with the vents (south cauldron active 14–15 April and west cauldron active from 15 April) and pathways of meltwater on 15 April (upper panel) and 5–6 May (lower panel). The location and form of subsurface volcanic conduits is highly schematic.
a 5–10 km high plume that alternated between shades of white (steam-rich) and dark grey (loaded with ash). In the first two days, the eruption formed ice cauldrons over the vents (Fig. 1a and 1b). The main activity on 14 April was in the south ice cauldron, but moved to the west cauldron on 15 April where it remained until the end of the eruption (Fig. 1b). Ice melting caused repeated jökulhlaups (glacial outburst floods) on 14–15 April. Part of the erupted material...
was water-transported pyroclasts, essentially identical to the air transported tephra. This phase is conveniently divided into 14–16 April when the tephra formed a well-defined sector towards the east, and 17- to early 18 April, when northerly winds carried the tephra southwards (Figs. 3–4 and 6, Table 1). The composition of the erupted melt is predominantly evolved benmoreite (Fig. 2e). Airborne ash was detected over 1.5–2 \( \times 10^6 \) km\(^2\) (Fig. 4).

II. Low-discharge effusive phase (18 April–4 May). During this period the melt composition was slightly less silicic, with the SiO\(_2\) content reduced from 60–61 wt% to 58–59 wt% (Fig. 2e) and the magma discharge dropped to 0.5–9 \( \times 10^4 \) kg s\(^{-1}\) (\( \approx \)2–35 m\(^3\) s\(^{-1}\) DRE). The activity was typified by lava effusion as a 3 km long lava flow made its way northwards out of the caldera down the outlet glacier Gigjökull (Fig. 1). Explosive activity was relatively weak but sustained, producing 2–4 km high plumes. Satellite-detected ash spread out over 0.1–0.2 \( \times 10^6 \) km\(^2\) (Fig. 4).

III. Second explosive phase (5–17 May). The onset of this phase is marked by abrupt rejuvenation of explosive activity that coincides with sudden change in melt composition from benmoreite to trachyte (Fig. 2e). The magma output was

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**Figure 3 | Ash concentration and distribution from satellite data.** (a) Ash concentrations in a vertical column for 15 April from SEVIRI retrievals. The inset shows the concentration on a section S-N over the Faroe Islands. On the graph H marks the position of the sampling locality at Hoyvik. (b–d) Distribution patterns of the tephra over Iceland and in its vicinity, based on satellite images from MODIS, MERIS and NOAA AVHRR. One image is used per day but data for some days are missing, usually because of unfavorable cloud cover. The sectors identified for each image denote the area over which ash can be detected on the images and no distinction is made between dense and highly diluted clouds.
Figure 4 | Isopach maps (thickness in cm) of tephra distribution during Phase I (the first explosive phase) 14–18 April and the whole eruption (c). (a) Deposition on land during the 14–16 April, the first three days of the first explosive phase, erupting from a water-filled vent. The nine localities where grain size analysis was carried out and the three zones on land (I–III) used to estimate the grain size distribution for 14–16 April are shown. The most intense fallout occurred from 18:30 on 14 April into the early hours of 15 April. (b) Deposition on land during the second part of the first explosive phase (17 April until early 18 April). A 6–7 km high, two-tiered tephra-laden plume typified the activity on 17 April, causing total darkness for 20 hours in the Eyjafjöll district south the volcano. (c) Total fallout on land in the eruption (14 April–22 May) and estimated fallout thickness (dotted lines) to the south and southeast of Iceland (see Methods and Figs. 5 and 6).
variable, shifting between periods of relatively low (0.3–
0.8 \times 10^5 \text{ kg m}^{-1} \text{ s}^{-1}) and high (2.5–
9 \times 10^5 \text{ kg m}^{-1} \text{ s}^{-1}) discharge (Fig. 2c).

The average ice melting rate in the vent area was
10 m^3 s^{-1} (10^4 \text{ kg s}^{-1}), confirming thermal energy loss from magma
to ice. Variable wind directions coupled with a duration of 13
days resulted in considerably wider dispersal than in the first
explosive phase, with ash being detected over an area of
>5 \times 10^6 \text{ km}^2 (Fig. 4).

IV. Final phase (18–22 May), the period when the eruption
decreased and eventually continuous activity stopped.
Explosions occurred on 4–8 June, but these involved minute
amounts of magma, and caused limited dispersion of tephra,
confined to an area within 1 km of the vents. A steam plume
from the summit crater persisted for months after the eruption.

On the basis of the magma petrology and its evolution with time^{12}
and the deformation and seismicity before and during the eruption^{9},
the course of events has been explained by injection of basaltic
magma from the mantle and its mixing and mingling with pre-
existing silicic magma residing in the crust under the volcano^{9}. Repeated peaks in activity observed during the second explosive
phase (Phase III, 5–17 May, see Fig. 2c) were preceded by deep earth-
quakes explained by injection of new basaltic magma^{6,12}. The lack of
injection of new basalt in late May is considered to have stopped the
eruption, not the exhaustion of silicic magma in the crust^{12}.

Tephra fallout and distribution. Although the tephra fallout was
most prominent in the sectors to the south and east, the prolonged explosive activity resulted in full 360° dispersal around the volcano
forming a continuous tephra blanket extending from 30 up to 80 km
distance from the source vents (Fig. 4). On 14–16 April ash was
carried eastward by westerly upper-tropospheric winds, while on
17 April northerly winds directed the plume to the south (Fig. 4a and 4b). Ash clouds were reported in Northern and Western Europe
in the period 15–21 April and again between 6 and 17 May^{6,7}. The
strong northwesterly flow persisted until 24 April, when it was
replaced by weaker westerly flow (Fig. 2b). In early May, when the
vigour of the eruption again intensified, the atmospheric flow turned
northwesterly and remained through the latter half of May although
the strength of the air flow gradually weakened. The eruption plume
was directed east- or southwards for 28 out of 39 days of sustained
activity. The weather situation during the spring of 2010 was indeed
rather unusual with frequent northwesterly winds and a large
deviation from the climatological mean atmospheric circulation^{13}.
The persistent northwesterly winds contributed strongly to the
dispersal of ash clouds to Europe.

The maximum thickness of the tephra fallout exceeds 30 meters in
the cone ramparts lining the rims of the ice cauldrons, falling
abruptly to 1 m over a distance of 2 km from the source vents. With
the exception of the northwest peninsula (the region of Vestfirðir), dusting of ash was reported in most parts of Iceland. The
total amount of airborne tephra produced in the eruption is
270 \pm 70 \times 10^6 \text{ m}^3 (bulk volume, density 1400 \text{ kg m}^{-3}) of which
140 \pm 20 \times 10^6 \text{ m}^3 fell in Iceland. This value is obtained by combining
isopach map integration on land with integrating a piecewise expo-
nential model of decline of thickness with distance for the area south
of Iceland (see Methods). The volume of 130 \pm 50 \times 10^6 \text{ m}^3 outside
Iceland is based on the form of the exponential segments using the
estimate for the fallout south of the Faroe Islands (see Methods) as
one point of data and the reported minor fallout in various parts of

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Figure 5 | Airborne ash detected outside Iceland within 40°–70°N and 40°W–30°E. The map is constructed by dividing SEVIRI retrievals into three separate time intervals: 14–17 April, 18 April–4 May and 5–18 May. Daily ash mass loadings that exceed 0.2 g m^{-2} within each 0.25°×0.25° grid cell are summed up. No attempt is made to indicate the amount of ash retrieved or the concentration, rather the plot shows the spatial distribution of the ash over the domain considered.
part for the whole eruption on Fig. 6a yields about $10^7$ m$^3$ to the total estimated volume beyond the coast of Iceland, or about 8%. A further $30\times10^9$ m$^3$ of tephra were transported out of the craters with meltwater and $23\times10^9$ m$^3$ (bulk density 2400 kg m$^{-3}$) were emplaced as lava. When taking into account the water transported tephra and the lava flow, the total mass erupted is $4.8\times10^{14}$ kg, or $0.18\pm0.05$ km$^3$ DRE (using $\rho_{DRE} = 2600$ kg m$^{-3}$). The partitioning of the erupted material with total volumes and mass are given in Table 1. Estimates of production rates of tephra for each day of the eruption, main direction of dispersal (estimated from ground observations and satellite images) and grain size characteristics, when available, are given in Table 2.

Out of the $130\times10^6$ m$^3$ of ash dispersed and deposited beyond the shores of Iceland $35\times10^6$ m$^3$ were emitted on 14–16 April, $18\times10^6$ m$^3$ on 17–18 April, $10\times10^6$ m$^3$ on 18 April–4 May, $65\times25\times10^6$ m$^3$ on 5–17 May and finally $2\times10^7$ m$^3$ in the final declining phase on 18–22 May (Table 1).

The 14–16 April tephra fall sector is well constrained. By integrating the model of piecewise exponential decline in tephra thickness with distance (Fig. 4) from $A^{1/2} = 260$ km (corresponding to the Faroe Islands) out to infinity, and converting the volume to mass using bulk density of 1400 kg m$^{-3}$ the result is that the mass of ash transported beyond the Faroe Islands on 14–16 April amounted to $m = 2\times6\times10^7$ kg ($2\times6$ Tg). The satellite data (Fig. 3a) indicate that the cloud southwest of the Faroe Islands on 15 April was 80–100 km wide and about 1 km thick. The upper tropospheric wind speed ($u$) at 5–7 km elevation for 15 April was 20 m s$^{-1}$. Combining these data, assuming that the transport took place over time $\Delta t \approx 24$ hours, with plume width being $w$ and $z$ the plume thickness, we estimate the average mass concentration ($c$) of ash in the cloud from

$$c = \frac{m}{wz\Delta t}$$

The result is an average concentration of ash in the cloud of 10–35 mg m$^{-3}$. By the time this same cloud reached Norway (1300 km from vent, $A^{1/2} \approx 600$ km), it had split into two, with each part having width of about 200 km (Fig. 5). Ash transported beyond that distance was of order 0.5–10$^7$ kg ($0.5\sim1$ Tg) (Fig. 6c). Using wind speed of 18 m s$^{-1}$ the mean ash concentration obtained is of order 1 mg m$^{-3}$.

During 5–17 May, ash transport in any single phase probably never reached the intensity of 14–16 April. An estimate of the mass transported beyond $\sim650$ km from source ($A^{1/2} \approx 250$ km) over the whole eruption is obtained by integrating the exponential thickness curve (Fig. 6a), as done above for 14–16 April, and converting to mass. The result is $4\times40\times10^7$ kg ($4\times40$ Tg). Finally, integration beyond $A^{1/2} \approx1000$ km (Fig. 5c) and converting to mass as before suggests that the amount of ash reaching mainland Europe was most likely less than $1\times10^8$ kg (<0.1 Tg).

About 85% of the total airborne tephra was produced during the two explosive phases (Table 1, Fig. 2). On the basis of analysis from eight sampling localities (Fig. 4a, and ref. 14) the main characteristics of the deposit from 14–16 April have been studied (Fig. 7). By partitioning the sector into four zones on the basis of distance from vent (Zone I: 0–6 km, Zone II: 6–30 km, Zone III: 30–80 km, with Zone IV including all fallout outside the coast) grain sizes for each sector have been estimated and combined into a whole deposit estimate. No grain size analyses exist for the offshore part (Zone IV) so we use the Zone III data there also. The results indicate that 94% of tephra erupted on 14–16 April was finer than 1000 µm, and 48–50% was ash finer than 63 µm. Furthermore, 35% of the total was finer than 30 µm, which is probably an underestimate since it is expected that the finest ash was carried further and preferably deposited beyond the coast. This is supported by modal and maximum grain sizes observed in the Faroe Islands (mean = 40 µm; max ~100 µm), Norway (mean = 48 µm; max ~90 µm), the British Isles

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Figure 6 | Estimates of distribution outside Iceland and the volume of tephra erupted. The tephra from 14–16 April is estimated separately since it forms a well defined sector and was responsible for the first ash transported to Europe. (a) Tephra thickness as a function of square root of area within each isopach, using fallout in the Faroe Islands, satellite data and the established first order behaviour of tephra thicknesses declining exponentially with distance from source$^{33,34}$. (b) Estimated cumulated bulk volume of fallout volume and its distribution with distance. The volume within Iceland is found by direct integration of the onland part of the maps in Fig. 4. The volume outside Iceland is found by integrating the exponential curves (e.g. ref. 37). (c) Estimate of volume (left axis) and mass (right axis) that was transported beyond a specified area. The volume is found by integrating the exponential curves for 14–16 April and the whole eruption from the point on the square root of area axis on (a) out to infinity, using eqs. 13 and 14 in ref. (35). Values outside $A^{1/2} = 300$ km are rough order of magnitude estimates.
become dominant. Dellino et al. suggest that stress-induced reaction of magmatic fragmentation were found during the second weak phase, while in the third, explosive phase, evidence of both magmatic and phreatomagmatic activity was influential in the initial phase of the eruption.

The eruption plume never exceeded 10 km in elevation (Fig. 2a) and the clast populations (Fig. 5) in the first and second explosive phases (Phases I and III) are predominantly angular, indicating brittle fragmentation of the magma. The grains are typically poorly to non-vesicular (<30% vesicles) with moderate to high abundances (30–90 modal %) of microlites and their groundmass glass is outgassed. Analyses of grain morphology for the different phases indicate that phreatomagmatic activity was influential in the first explosive phase, evidence of both magmatic and phreatomagmatic fragmentation were found during the second weak phase, while in the third, explosive phase, magmatic fragmentation had become dominant. Dellino et al. suggest that stress-induced reaction of microvesicular, fragile melt was responsible for the fragmentation once water access to the vents was exhausted.

**Discussion**

The eruption plume never exceeded 10 km in elevation (Fig. 2a) and classifies as weak, having a 0–1 km high gas thrust region with the wind diverting the buoyant plume above this elevation. The prolonged duration (Fig. 2) contrasts with Plinian eruptions that usually last for some hours rather than days. Moreover, the 2010 eruption demonstrates that despite its moderate size, atmospheric loading at upper tropospheric levels (4–9 km) can be substantial and widespread exceeding areas well over 10° km². It is likely that the high proportion of fine-grained ash was of major importance, allowing a fraction of the finest ash to disperse widely. The settling time in the atmosphere of <60 μm particles, initially at 5–10 km elevation, is of the order of a few days, or significantly longer than the travel time between Iceland and Northern Europe. However, the abundance of very fine ash beds (often loaded with ash aggregates) on land in Iceland testifies to the importance of aggregation in explaining the heavy fallout observed in the proximity of source vents. At larger distances, sedimentation of the very fine ash was much slower within the widespread diluted ash clouds over Europe (Fig. 5) as was observed in the days following the eastwards directed plume of 14–16 April. However, aggregation is also important during sedimentation in the far-field regions as clearly indicated by the observations in the UK. It should be noted that the synoptic meteorology also influences the ash sedimentation rate, as indicated by lidar observations in the UK and Germany showing the ash cloud to be a sloping structure and parallel to a weather front.

Application of satellite data to infer mass loading of the atmosphere during volcanic eruptions holds great potential in enhancing aviation safety and successful routing of aircraft in the vicinity of ash clouds. Stohl et al. estimated the total amount of ash finer than 28 μm in diameter in the Eyjafjallajökull eruption as 8.7 ± 4.2 Tg. Our grain size data are confined to land and at present only allow a full grain size distribution estimate for 14–16 April, indicating 33 ± 12 Tg of <28 μm for that phase alone. This leaves out substantial fallout of <28 μm particles on 17 April and in the second explosive phase on 5–17 May (Table 2), during which it may have been 10–20% of the erupted tephra. Thus, the 14–16 April estimate is in all likelihood accounting for less than half of the total of the <28 μm fraction generated in the eruption. Thus, the true value of <28 μm ash erupted in all phases combined is estimated to have been at least 70 Tg, one order of magnitude higher than the published estimates based on satellite retrievals. A likely reason for such a discrepancy might be that satellite retrievals for mass concentration do not work in dense, optically opaque eruptions clouds, and may therefore not provide reliable estimates close to a volcano where aggregation-enhanced settling can be a major factor.

Although only partial information on grain sizes exists at present, it is instructive to compare the Eyjafjallajökull eruption with other eruptions. It holds for the first explosive phase that the proportion of fine ash (<1000 μm) is comparable to that of the VEI 4–5 eruptions of Mount St. Helens in 1980 (dacite), El Chichon in 1982 (trachyandesite) and the phreatoplinian C phase of the Askja eruption in 1875 (rhyolite), and considerably larger than in e.g. the VEI 3 anidesitic eruption of Mount Spurr in Alaska in 1992 (ref. 23).

**Table 1 | Results on tephra and lava erupted at Eyjafjallajökull, April 14–May 22, 2010. In converting volumes to mass the density of tephra is taken as 1400 kg m⁻³, and that of lava 2400 kg m⁻³ [See Methods for details]**

| Phase no. and dates | All airborne | Tephra fallen in Iceland | Tephra outside Iceland | Tephra water transp | Lava (21.4.–3.5.) | Mass total |
|---------------------|-------------|--------------------------|------------------------|---------------------|------------------|-----------|
| Phase I - south: 17–18.4 | 70±15 | 10±3 | 10±3 | 10±3 | 10±3 | 0.3±0.3 |
| Phase I - east: 14–16.4 | 25±7 | 7±1 | 18±6 | 25±5 | 0.4±0.1 |
| Phase I - total: 14–18.4 | 95±22 | 42±6 | 53±16 | 25±5 | 1.7±0.4 |
| Phase II: 18.4–4.5 | 30±10 | 20±3 | 10±7 | 5±1 | 23±5 | 1.1±0.3 |
| Phase III: 5–17.5 | 135±35 | 70±9 | 65±26 | 0 | 1.9±0.5 |
| Phase IV: 18–22.5 | 10±3 | 8±2 | 2±1 | 0 | 0.1±0.03 |
| **Total** | **270±70** | **140±20** | **130±50** | **30±6** | **23±5** | **4.8±1.2** |
The 2010 eruption at Eyjafjallajökull acutely demonstrated that a moderate size explosive eruption can produce widely dispersed volcanic plumes and cause major disruptions to air travel. In Europe, a number of Italian volcanoes have and will in the future produce intermediate to silicic explosive eruptions with large amounts of fine ash. However, in terms of eruption frequency and potential for disruption, Icelandic volcanoes are the most potent. Upper tropospheric winds are usually most influential in determining the direction of dispersal of tephra from Icelandic eruptions, tending to have a component of eastbound flow, with northwesterly winds slightly less common than southwesterlies. Quantitative analysis of the probabilities for magnitude and duration of explosive eruptions in Iceland is beyond the scope of this study. However, the existing records show that about 10–15 eruptions with a volcanic explosivity index VEI = 3 (ejecta volume 0.01–0.1 km³) or larger occur in Iceland every century, and between 1970 and 2010 seven such eruptions occurred with tephra generation lasting from about an hour to 1.5 days. Southwesterly upper tropospheric winds were dominant during these events carrying tephra towards northeast. Hence, only minor disruption to air traffic occurred and the busy air routes in central and western Europe were not affected. Finally, basaltic phreatomagmatic eruptions in Iceland such as some of those from Katla, the larger and neighbouring volcano to Eyjafjallajökull, typically last several days or even weeks. Such eruptions may have higher ash discharge than observed at Eyjafjallajökull. The most recent explosive eruption in Iceland, the one at Grimsvötn in May 21–28, 2011, was basaltic and of greater intensity than the 2010 Eyjafjallajökull eruption, as indicated by its 15–20 km high eruption.

Table 2 | Approximate estimate of airborne tephra produced per day 14 April–22 May and main tephra dispersion directions. $\Delta M/\Delta t$ is estimated tephra mass erupted per day, based on partitioning of tephra erupted using the scaled version of the Mastin et al. equation (see Methods). The three columns on the right give grain size data where it is available: distance from vent where grain size samples were taken and the proportion of particles $<63 \mu m$ ($\phi > 4$) and $<31 \mu m$ ($\phi > 5$). Same information is given at other dates for representative samples. The first nine grain size samples are from the tephra deposited during the period 14 April to 16 April, including data from Gislason et al.

| Date   | $\Delta M/\Delta t$ | Direction of dispersal | Grain size sample distance | Grain size % | Grain size % |
|--------|---------------------|------------------------|----------------------------|--------------|--------------|
|        | 10$^6$ kg/day       |                       | km                         | $<63 \mu m$ | $<31 \mu m$ |
| 14 Apr | 40–50               | E                      | 1                          | 21           | 11           |
|        |                     |                        | 2                          | 24           | 15           |
|        |                     |                        | 10                         | 13           | 11           |
|        |                     |                        | 10                         | 17           | 11           |
| 15 Apr | 5–10                | E                      | 58                         | 44           | 33           |
|        |                     |                        | 60                         | 56           | 45           |
|        |                     |                        | 58                         | 70           | 51           |
|        |                     |                        | 56                         | 65           | 45           |
| 16 Apr | 30–40               | E                      | 21                         | 26           | 17           |
| 17 Apr | 30–40               | S                      | 11                         | 47           | 33           |
| 18 Apr | 5–10                | SE                     | -                          | -            | -            |
| 19 Apr | 1–3                 | SE                     | -                          | -            | -            |
| 20 Apr | <1                  | SE                     | -                          | -            | -            |
| 21 Apr | <1                  | SE                     | -                          | -            | -            |
| 22 Apr | <1                  | SE                     | 4                         | 7            | 4            |
| 23 Apr | 1–2                 | E                      | -                          | -            | -            |
| 24 Apr | 1–2                 | E                      | -                          | -            | -            |
| 25 Apr | 2–3                 | NE                     | -                          | -            | -            |
| 26 Apr | 3–5                 | E                      | -                          | -            | -            |
| 27 Apr | <1                  | N                      | -                          | -            | -            |
| 28 Apr | 10–15               | W                      | -                          | -            | -            |
| 29 Apr | 5–10                | SE                     | -                          | -            | -            |
| 30 Apr | <1                  | SE                     | -                          | -            | -            |
| 31 May | <1                  | SE                     | -                          | -            | -            |
| 3 May  | 1–2                 | E                      | -                          | -            | -            |
| 4 May  | 5–10                | E                      | -                          | -            | -            |
| 5 May  | 30–40               | SE                     | 30                         | 46           | 29           |
| 6 May  | 5–10                | E                      | -                          | -            | -            |
| 7 May  | 5–10                | SE                     | -                          | -            | -            |
| 8 May  | 3–5                 | SE                     | 13                         | 12           | 10           |
| 9 May  | 2–5                 | SE                     | -                          | -            | -            |
| 10 May | 2–5                 | SE                     | 13                         | 12           | 10           |
| 11 May | 10–20               | SE                     | -                          | -            | -            |
| 12 May | 3–5                 | E                      | -                          | -            | -            |
| 13 May | 20–30               | SE                     | 10,13                      | 38           | 23           |
| 14 May | 20–30               | W                      | 14                         | 10           | 8            |
| 15 May | 15–25               | SW                     | -                          | -            | -            |
| 16 May | 30–40               | SE                     | -                          | -            | -            |
| 17 May | 20–30               | E                      | -                          | -            | -            |
| 18 May | 3–5                 | NE                     | -                          | -            | -            |
| 19 May | 3–5                 | NE                     | 8                          | 42           | 27           |
| 20 May | 1–3                 | NE                     | -                          | -            | -            |
| 21 May | <1                  | E                      | -                          | -            | -            |
| 22 May | <1                  | W                      | -                          | -            | -            |
Methods

Our results are based on measurements of tephra thickness on land in Iceland, records of tephra fallout outside Iceland, reconstruction of plume dispersal from satellite images, monitoring and mapping of ice melting, eruption plume height, grain size analysis of tephra, tephra grain morphology and chemistry of the erupted products.

Tephra volume in Iceland. The tephra on land consisted of airfall (dominant), and water transported (minor). For the airfall tephra thicknesses were measured in about 400 locations and used to construct the isopach maps. Tephra volume within the mapped area was found by directly integrating the isopach maps. Volume outside the 0.01 cm contour on land (approximately 2·10⁻⁹ m⁻³) was estimated using exponential decline in thickness with distance. In addition, the material left within the two main ice cauldrons at the end of the eruption was estimated from the difference between the depth of the partly tephra-filled cauldrons and the depth to pre-eruption bedrock (~ 33·10⁻⁹ m). Error estimates for isopach maps on land are considered to be 15%. Separate isopach maps were constructed for the east-directed sector formed on 14–16 April, the south-directed sector formed on 17 April and for the whole deposit (Fig. 4). The volume of tephra deposited on land for the period 18 April–22 May was obtained by subtracting the volume erupted on 14–18 April from the whole-deposit map. The partitioning between 18 April–4 May, 5–17 May and 18–22 May is estimated from (1) the integrated plume height over the two periods (eq. 2, see below), and (2) relative thickness of deposits for the two periods on Eyjafjallajökull based on our measurements in July 2010 and August 2011.

Tephra volume outside Iceland. We assume piecewise exponential decline with distance. The magnitude is constrained from distribution of ash as observed from satellite images showing eruption cloud distribution pattern analysis of MODIS, MERIS and NOAA AVHRR satellite images (Figs. 3 and 4). Very fine ash fell in Hoyvik in the Faroe Islands (about 1 g/m², ref. 40) at the northern margin of the main plume of April 15–16. We assume that fallout was proportional to estimates of the satellite retrievals of volcanic ash mass loadings (see below) and in that way derive a rough estimate of maximum fallout south of the islands as 40 g/m², corresponding a thickness of 0.03 cm. The form of the dotted contours outside the coast on Fig. 4c is based on the plume locations documented on Figs. 3b–c, while the area covered by each contour is defined in such a way that direct integration would yield the same result as the piecewise exponential model. Error estimates for material deposited outside Iceland are based on a minimum extrapolation that still allows for some material to be transported past the Faroe Islands and a maximum estimate that still limits fallout equivalent to 0.1 g m⁻² e.g. 0.5–1% of a surface would contain a particle if all particles were 10 μm in diameter) to the northern coast of mainland Europe. These estimates place the relative error at about 40%.

Conversion from tephra volume to mass is obtained by using bulk tephra density of 1400–240 kg m⁻³, found by averaging the results from 14 sampling locations obtained at various times between 15 April and 20 May (range 1195–1689 kg m⁻³ and standard deviation of 130 kg m⁻³). The same value is used for water-transported tephra. Our results are essentially identical to the mean value of 1443 kg m⁻³ obtained by Bonadonna et al. for fallout from 4–8 May. Representative samples yielding a reliable value for the lava flow are lacking but a value of 2400 kg m⁻³ is used, a reasonable value for lava that has experienced considerable water interaction (e.g. ref. 42). Grain size analysis was done by hand sieving down to 125 or 63 μm and a Sympatec laser diffraction analyser down to 2 μm. Total grain size partitioning for the 14–16 April east-directed sector was done by defining the three zones on land (Fig. 4a) as Voromoi Teraaz (nearest neighbour) cells using average values for the grain size samples within each zone (two in Zone I, two at the same locality in Zone II and one on the southern part of the sector, and four in Zone III, all at 58–60 km distance from vent in Zone III). The values from Zone III are used for Zone IV (offshore).

Satellite retrievals of volcanic ash mass loadings are made using measurements from infrared channels of the geosynchronous Meteosat Second Generation (MSG) Spin-stabilized Enhanced Visible and Infrared Imager (SEVIRI), with a temporal sampling time of 15 min. SEVIRI observes the earth disk over a total field of view of 70° in 12 channels from the visible to the infrared with a temporal sampling time of 15 min, and a spatial resolution of ~10 km at the sub-satellite point, increasing to ~100 km near the edges of the scan.

Eruption plume height (Fig. 2) was monitored with weather radar at Keflavik, 150 km to the west of the volcano, sampling every 5 minutes, but six hourly values of mean and maximum height are presented. From the isopach maps on land and piecewise exponential decline outside Iceland we obtain total volume for different periods. In order to partition the erupted material into shorter periods of six hours (Fig. 2), we use the record of plume height and then apply a scaled version of the empirical formula of Mastin et al. to

\[
Q = 0.0564k \left[ H_{\text{max}} + H_{\text{med}} \right]^{1.15} / 2
\]

Here \(Q\) is average magma discharge for the 6 hour interval (in m³ s⁻¹ of magma). We convert the plume density to mass discharge using DRE density 2600 kg m⁻³, applicable to magmas of intermediate composition. \(H_{\text{med}}\) and \(H_{\text{max}}\) are respectively the median and maximum values of plume height over vent for the interval. \(k\) is the scaling factor found by comparing the value obtained by integrating eq. (1) (with \(k = 1\)) with the actual mapped fallout over three different periods of the eruption.

Figure 7 | Grain size distributions for 14-16 April. Results for samples from four different localities within the on-land part of the 14–16 April eastwards directed tephra sector of Phase I at distances of 2, 10, 21 and 60 km (see Fig. 4a for location).

Vent development and ice melting. Monitoring of vents and ice cauldron formation (Figs. 1, 2) was done using the Icelandic Coast Guard Dash 9 aircraft, equipped with a Synthetic Aperture Radar (SAR)³¹. Ice melting rates were estimated from (1) growth of ice cauldrons as measured from the SAR images, (2) evolution of crevasse patterns and other subsidence structures around the ice cauldrons from photographs and the SAR images, and (3) mapping of ice surface with GPS in July 2010.
scoring factor $k$, is found to take the values 2.15 for 14–16 April, 1.58 for 17 April and 1.59 for 18 April–22 May. Note that these scoring factors are dependent on our choice of using the mean of median and maximum height for each interval; slightly different scoring values are found if 5 minute plume height data is integrated directly$^{46,57}$. We are not attempting to refine the Mastin formula in any general way, only scale it in such a way that when integrated it yields values that fit our total volume estimates. The plume was pulsating over long periods, displaying considerable variations in height over time scales of hours. The maximum height might also be applied or the use of shorter intervals, leading to slightly different values of $k$. However, we find that the results shown in Table 1 are not significantly affected by the interval length (5 min., 3 or 6 hours) or whether maximum or mean plume height is used. The six hourly means,$p_{6h}$, required minimal editing or interpolation of gaps in the plume height record$^{46,57}$ and are therefore used here.

As noted above, we also use the scaled plume discharge values from equation (2) to estimate partitioning of the volume obtained for 18 April–22 May into the individual main phases (Table 1).

Wind speed and direction at approximately 5 and 7 km altitude is retrieved from the operational analysis of the European Centre for Medium Weather Forecasts (ECMWF).

Composition of the magma was obtained by Cameca SX100 electron microprobe at the University of Edinburgh$^{46}$. Scanning electron microscope (SEM) images were obtained with a Leo Supra 25 instrument.

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data for distant distribution while IJ and TH defined fallout patterns south and southeast of Iceland from satellite images. HB and GNP analysed data on wind patterns and, together with MTG and BO, weather radar data. Data from airborne monitoring were analysed by MTG, EM, BO and TH. CLH and TT performed chemical analyses and JAS analysed ash from the UK, Norway and the Faeroe Islands. MTG and TT wrote the manuscript with input from all team members.

Additional information
Competing financial interests: The authors declare no competing financial interests.
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