Revision of the Tanzanian dicynodont *Dicynodon huenei* (Therapsida: Anomodontia) from the Permian Usili Formation

Christian F. Kammerer Corresp. 1

1 North Carolina Museum of Natural Sciences, Raleigh, North Carolina, United States of America

Corresponding Author: Christian F. Kammerer
Email address: christian.kammerer@naturalsciences.org

A single species of the dicynodontid dicynodont *Dicynodon* is currently recognized from the late Permian Usili Formation of Tanzania: *Dicynodon huenei* Haughton, 1932. Restudy of the known Tanzanian materials of *D. huenei* demonstrates that they represent two distinct morphotypes, here considered separate taxa. The holotype of *D. huenei* is not referable to *Dicynodon* and instead is transferred to the genus *Daptocephalus* (but retained as a valid species, *Daptocephalus huenei* comb. nov.) A number of published dicynodontid specimens from the Usili Formation, however, are referable to *Dicynodon*, and are here recognized as a new species (*Dicynodon angielczyki* sp. nov.) *Dicynodon angielczyki* can be distinguished from its South African congener *D. lacerticeps* by the presence of an expansion of the squamosal and jugal beneath the postorbital bar and a curved, posterolateral expansion of the squamosal behind the temporal fenestra. Inclusion of *D. angielczyki* and *D. huenei* in a phylogenetic analysis supports their referral to *Dicynodon* and *Daptocephalus* (respectively). These results indicate higher basinal endemism in large late Permian dicynodonts than previously thought, a sharp contrast to the cosmopolitanism in the group in the earliest Triassic.
Revision of the Tanzanian dicynodont *Dicynodon huenei* (Therapsida: Anomodontia) from the Permian Usili Formation

Christian F. Kammerer

North Carolina Museum of Natural Sciences, Raleigh, USA

ABSTRACT

A single species of the dicynodontoid dicynodont *Dicynodon* is currently recognized from the late Permian Usili Formation of Tanzania: *Dicynodon huenei* Haughton, 1932. Restudy of the known Tanzanian materials of *D. huenei* demonstrates that they represent two distinct morphotypes, here considered separate taxa. The holotype of *D. huenei* is not referable to *Dicynodon* and instead is transferred to the genus *Daptocephalus* (but retained as a valid species, *Daptocephalus huenei* comb. nov.) A number of published dicynodontoid specimens from the Usili Formation, however, are referable to *Dicynodon*, and are here recognized as a new species (*Dicynodon angielczyki* sp. nov.) *Dicynodon angielczyki* can be distinguished from its South African congener *D. lacerticeps* by the presence of an expansion of the squamosal and jugal beneath the postorbital bar and a curved, posterolateral expansion of the squamosal behind the temporal fenestra. Inclusion of *D. angielczyki* and *D. huenei* in a phylogenetic analysis supports their referral to *Dicynodon* and *Daptocephalus* (respectively). These results indicate higher basinal endemism in large late Permian dicynodonts than previously thought, a sharp contrast to the cosmopolitanism in the group in the earliest Triassic.

Subjects: Paleontology, Taxonomy, Zoology

Keywords: Synapsida; Dicynodontia; Africa; Tanzania; Permian; biogeography
INTRODUCTION

The Usili Formation is a sedimentary unit of late Permian age exposed in the Ruhuhu Basin at the southwestern edge of Tanzania (Wopfner, 2002; Sidor & Nesbitt, 2018). The Ruhuhu Basin has been recognized as fossiliferous since the initial geological surveys of G. M. Stockley (Stockley & Oates, 1931; Stockley, 1932), who collected a number of therapsid fossils in what is now considered the Usili Formation and sent them to the South African Museum (Cape Town) for study. Haughton (1932) initially described these fossils, and named four new dicynodont species based on Stockley’s collections: *Dicynodon huenei*, *Dicynodon tealei*, *Megacyclops rugosus*, and *Pachytygos stockleyi*. Of these taxa, only *D. huenei* is considered valid today: *D. tealei* and *M. rugosus* are considered rhachiocephalids of dubious validity (Brink, 1986; Kammerer et al., 2011) and *P. stockleyi* is considered a probable synonym of *Endothiodon bathystoma*, a taxon better known from the Karoo Basin of South Africa (Cox & Angielczyk, 2015).

Haughton (1932) established *Dicynodon huenei* based on a single specimen: SAM-PK-10630, a fragmentary partial skull (Fig. 1) and some associated postcranial elements, most notably the majority of a left scapula, from Stockley’s locality B2. He did not explicitly differentiate the new taxon from the many existing species of *Dicynodon*, and the diagnostic features listed for his new species (e.g., short, wide snout, interorbital width greater than intertemporal) are present in many other dicynodontoids. Additional specimens of Usili Formation dicynodontoids were later collected by Ernst Nowack and sent to Tübingen, Germany for study by Friedrich von Huene (Nowack, 1937; Huene, 1942). Huene (1942) named the new species *Dicynodon bathyrhynchus* (currently *Eutychognathus bathyrhynchus*; Kammerer et al., 2011) for one of these specimens (GPIT/RE/7104), but referred the majority of Nowack’s skulls to *D. huenei*, albeit transferring this species to the genus *Platypodosaurus*. *Platypodosaurus* is a problematic taxon originally established for large dicynodont postcranial remains from South Africa (Owen, 1880), and subsequent authors have not accepted Huene’s referral of *D. huenei* to the genus, instead retaining it as a species of *Dicynodon* (e.g., Haughton & Brink, 1954; King, 1988). Kammerer et al. (2011), in their comprehensive global revision of *Dicynodon*, recognized *D. huenei* as a valid species and considered all dicynodontid specimens from the Usili Formation (with the exception of the aforementioned *E. bathyrhynchus*) to be referable to this taxon. This included Tanzanian specimens previously referred to the typically South African species *Dicynodon lacerticeps* (Wild et al., 1993) and *D. leoniceps* (Gay & Cruickshank, 1999). Kammerer et al. (2011) also referred a number of dicynodontid specimens from the upper Madumabisa Mudstone Formation (Luangwa Basin, Zambia) to *D. huenei*, notably NHMUK PV R37005 (formerly TSK 14), a nearly complete skeleton described by King (1981) as a specimen of *Dicynodon trigonocephalus*. Finally, Angielczyk et al. (2014) provided further details on the distribution and anatomy of the Zambian specimens, and discussed the rationale behind their referral to *D. huenei*.

Recent expeditions (since 2007) to the Ruhuhu and Luangwa basins, led by researchers from the University of Washington and Field Museum of Natural History (USA), have collected a wealth of new therapsid fossils from the Usili and Madumabisa Mudstone Formations (e.g., Angielczyk et al., 2009, 2014; Weide et al., 2009; Sidor et al., 2010; Angielczyk & Cox, 2015; Huttenlocker et al., 2015; Huttenlocker & Sidor, 2016; Sidor & Nesbitt, 2018). Among these specimens are numerous Zambian dicynodontoids, including well-preserved skulls matching NHMUK PV R37005 in general morphology but showing clear differences from all Usili Formation *D. huenei* specimens. These specimens are currently being described (K. Angielczyk, pers. comm., 2019), so will not be discussed in depth here, but suggest that there is at least species-
level distinction between Zambian and Tanzanian “D. huenei”, contra Kammerer et al. (2011) and Angielczyk et al. (2014).

Re-examination of the fragmentary holotype of D. huenei, SAM-PK-10630, provides additional evidence that Kammerer et al. (2011) were overly conservative in treating all Dicynodon materials from Tanzania and Zambia as a single species. Because of the incompleteness of the holotype, Kammerer et al.’s (2011) diagnosis of D. huenei was based primarily on the series of complete Usili Formation dicynodontoid skulls housed in collections in Cambridge (UK) and Tübingen, notably UMZC T1089, GPIT/RE/7175, and GPIT/RE/7177. However, restudy of SAM-PK-10630 shows that, although damaged, this specimen differs in several important regards from those better known skulls. Here, I present a critical review of all published “D. huenei” material from Tanzania, argue that two distinct species are represented in this assemblage, and discuss the phylogenetic and biogeographic implications of this conclusion.

Nomenclatural acts
The electronic version of this article in portable document format will represent a published work according to the International Commission on Zoological Nomenclature (ICZN), and hence the new names contained in the electronic version are effectively published under that Code from the electronic edition alone. This published work and the nomenclatural acts it contains have been registered in ZooBank, the online registration system for the ICZN. The ZooBank Life Science Identifiers (LSIDs) can be resolved and the associated information viewed through any standard web browser by appending the LSID to the prefix http://zoobank.org/. The LSID for this publication is: urn:lsid:zoobank.org:pub:714AFA18-EB4D-4A35-B948-B4E7CB046A77. The online version of this work is archived and available from the following digital repositories: PeerJ, PubMed Central, and CLOCKSS.

COMPARATIVE BACKGROUND

Kammerer et al. (2011) reviewed the original dicynodont genus, Dicynodon Owen, 1845, which had become a notorious wastebasket taxon. As part of this revision, they resurrected the genus Daptocephalus (long considered synonymous with Dicynodon; see Cluver & King, 1983; King, 1988) for the large South African species D. leoniceps, recognizing it as distinct from the type species Dicynodon lacerticeps based on morphometric analysis and discrete-state characters. In their phylogenetic analyses, Kammerer et al. (2011) never recovered D. leoniceps and D. lacerticeps as sister-taxa, supporting recognition of a separate genus for the former. More recent studies have maintained the distinction between Dicynodon lacerticeps and Daptocephalus leoniceps (e.g., Botha-Brink et al., 2014; Jasinoski et al., 2014; Angielczyk & Kammerer, 2017; Olivier et al., 2019), and detailed analysis of the stratigraphic distributions of these two taxa indicates that they had somewhat different ranges (Viglietti et al., 2016, 2018).

Kammerer et al. (2011) considered Dicynodon lacerticeps and Daptocephalus leoniceps to be restricted to the Karoo Basin of South Africa, and considered extra-basinal Permian dicynodontoid records to mostly represent distinct, locally-endemic taxa (e.g., Jimusaria and Turfanodon in China; Peramodon and Vivaxosaurus in Russia; Gordonia in Scotland). They recognized only two Permian dicynodontoid taxa with international ranges: Euptychognathus bathyrynchus (recorded in South Africa and Tanzania) and Dicynodon huenei (recorded in Tanzania and Zambia). Euptychognathus bathyrynchus is a rare taxon, and of the four recorded specimens, only one (the holotype) was found in Tanzania. By contrast, Dicynodon huenei sensu
Kammerer et al. (2011) is relatively common: they referred 14 specimens to the species, and Angielczyk et al. (2014) referred additional specimens from Zambia to *D. huenei*, with further unprepared or juvenile skulls considered possibly referable.

Both Kammerer et al. (2011) and Angielczyk et al. (2014) used the large, complete skull UMZC T1089 as their primary exemplar of *D. huenei*, rather than the fragmentary and poorly-preserved holotype (SAM-PK-10630), and referral of additional specimens to the species was largely based on comparisons with the former skull. They considered the main diagnostic feature of the species to be an expanded section of the zygoma beneath the postorbital bar, thickened to form a large plate at the posteroverentral edge of the orbit and making the zygoma appear bowed outwards in dorsal or ventral view, as is very evident in UMZC T1089 and the similar specimens GPIT/RE/7175 (Fig. 2A,C) and GPIT/RE/7177 (Fig. 3A). Kammerer et al. (2011) stated that, although incomplete, this morphology was present in SAM-PK-10630, and Angielczyk et al. (2014) noted its presence in NHMUK PV R37005 and other Zambian specimens.

It is true that the zygomatic arches of SAM-PK-10630 (based on the description of Haughton [1932] and the existing fragments) and the Zambian specimens discussed by Angielczyk et al. (2014) are more bowed than those of the South African *D. lacerticeps* (Fig. 4A) when viewed dorsally. As mentioned above, however, new Zambian dicynodontoid specimens currently under study show several distinctive features differentiating them from UMZC T1089 and similar specimens, including a septomaxillary-lacrimal contact and presence of a deep occipital excavation above the paroccipital process (Kammerer, pers. obs.). Furthermore, although the zygoma in the Zambian specimens is indeed bowed, they do not show the same style of zygomatic expansion as in the Tanzanian material. In GPIT/RE/7175, GPIT/RE/7177, and UMZC T1089, the anterior process of the squamosal is enlarged, forming a thickened edge to the zygoma postorbitaly. However, the jugal in this region is also expanded dorsoventrally, separating the postorbital from the squamosal. In the Zambian specimens, as in most dicynodontoids, there is no such jugal expansion and the postorbital still contacts the dorsal edge of the squamosal. Further commentary on the relationships of the Zambian specimens will have to await their full description, but at present they should not be considered conspecific with the Tanzanian material.

Only part of the zygoma is preserved in SAM-PK-10630. More extensively preserved, however, are the skull roof (including the intertemporal bar) and right snout (Fig. 1). Recent re-examination of these fragments demonstrates that they differ in several important ways from the ‘standard *D. huenei*’ morphotype represented by GPIT/RE/7175, GPIT/RE/7177, and UMZC T1089. In SAM-PK-10630, the intertemporal exposures of the postorbitals are nearly vertical, whereas in the ‘standard *D. huenei*’ morphotype they are more horizontal. A median interorbital ridge is present on the skull roof in SAM-PK-10630, but absent in ‘standard *D. huenei*’. The snout of SAM-PK-10630 is relatively tall and sharply-sloping, but relatively low and gradually-sloping in ‘standard *D. huenei*’. Tusks are ventrally directed in SAM-PK-10630, but anteroventrally angled in ‘standard *D. huenei*’. Intriguingly, this set of characters is not unique to SAM-PK-10630 among Usili Formation dicynodontoids, but is also evident in the more complete skulls GPIT/RE/9316 (Fig. 2B,D) and GPIT/RE/9641 (Fig. 3B, 4B,D). These specimens show additional differences separating them from the ‘standard *D. huenei*’ morphotype, notably a taller, transversely narrower occiput (Fig. 5), proportionally broader interorbital region, proportionally longer, narrower intertemporal bar, and a less constricted median pterygoid plate lacking a distinct crista oesophagea. These other specimens also appear to lack the zygomatic expansion of ‘standard *D. huenei*’, but have a more vertically-oriented, broadly rounded subtemporal arch. Comparison with SAM-PK-10630 indicates that the apparent expansion of the subtemporal arch in that specimen is
just the result of displacement of the more vertical, posterior section of the bar seen in specimens like GPIT/RE/9641, rather than actual dorsoventral and transverse expansion as in UMZC T1089.

The features separating SAM-PK-10630, GPIT/RE/9316, and GPIT/RE/9641 from the ‘standard D. huenei’ morphotype include many of the characters cited by Kammerer et al. (2011) as differentiating Daptocephalus from Dicynodon. The relatively tall, sharply-sloping snout, ventrally-oriented tusks, vertically-oriented, broadly rounded subtemporal arches, narrow, extremely elongate intertemporal bar, and vertical orientation of the postorbitals in the intertemporal bar are all characteristic features of the genus Daptocephalus. The presence of an interorbital ridge, a relatively broad interorbital region, and a weakly constricted median pterygoid plate lacking a crista oesophagea also differentiate Daptocephalus from Dicynodon (Kammerer, pers. obs.), and are newly recognized here as diagnostic features of the former genus (Fig. 3, 4; also see Table 1). (It is worth noting here that the smallest specimens otherwise identifiable as Dicynodon lacerticeps also lack a crista oesophagea, so the presence of this character may be ontogenetically variable in that species. However, all known specimens of Daptocephalus leoniceps, most of which are very large skulls, lack a crista oesophagea, so in mature specimens of the two genera this appears to be a reliable differentiator.) In total, the suite of features listed above indicates that two morphotypes of dicynodontoid are present in the Usili Formation (in addition to the singleton record of Euptychognathus, which can readily be distinguished from both morphotypes by its extremely tall snout and U-shaped naso-frontal ridge). These morphotypes are most similar, among known dicynodonts, to the South African taxa Dicynodon lacerticeps (for the so-called ‘standard D. huenei’ morphotype) and Daptocephalus leoniceps (for the D. huenei holotype and similar specimens) and are here considered congeneric with them (see Phylogenetic Analysis for further rationale behind this). However, each varies somewhat from their Karoo counterparts. The holotype SAM-PK-10630 and GPIT/RE/9316 show dorsoventrally taller lacrimals than in D. leoniceps (Fig. 6; see further discussion below), although otherwise Tanzanian and South African Daptocephalus specimens are very similar. GPIT/RE/7175, GPIT/RE/7177, UMZC T1089 and other representatives of the ‘standard D. huenei’ morphotype are more easily distinguished from their Karoo counterpart D. lacerticeps by their massively expanded suborbital zygoma and bowed subtemporal arches (Fig. 7).

Here, the two morphotypes of Usili Formation dicynodontoid formerly included within Dicynodon huenei are recognized as distinct species, which are closely related to (but distinct from) the well-known South African dicynodontoid taxa Dicynodon lacerticeps and Daptocephalus leoniceps. Confoundingly, the holotype of Dicynodon huenei is not part of the ‘standard D. huenei’ morphotype that has been the basis for most of the recent literature on the species, but instead is referable to the genus Daptocephalus (as Daptocephalus huenei comb. nov.) No preexisting species names are available for what was previously considered ‘standard D. huenei’, requiring the establishment of a new species for this morphotype.

Almost all known dicynodontoid specimens from the Usili Formation can be referred to one of the two taxa recognized here (see full list of referred materials in the Systematic Paleontology section below). However, a few published specimens from the Usili Formation cannot be confidently identified to species at present. Wild et al. (1993) described a dicynodont skull from the Usili Formation that they referred to the South African species Dicynodon lacerticeps. The whereabouts of this specimen are currently unknown; no specimen number or institutional depository were listed in its description. Based on their description, it does appear to be a dicynodontoid, and their Figure 4c shows the skull as having vertically-oriented postorbitals in the intertemporal bar. As such, it may represent a specimen of Daptocephalus huenei. Until this
specimen is relocated and can be examined in detail, however, this record should only be considered Dicynodontoidea indet. UMZC T1280, referred to *Dicynodon huenei* by Kammerer et al. (2011), consists of a partial mandible from Stockley’s Locality B4/7. This jaw has a deeper symphysis than UMZC T1123, suggesting that it could be *Daptocephalus huenei* (the jaw symphysis is relatively deep in *D. leoniceps*), but until associated jaws with definite skulls are known for *D. huenei*, this specimen should also be regarded as Dicynodontoidea indet.

**SYSTEMATIC PALEONTOLOGY**

*Synapsida* Osborn, 1903

*Therapsida* Broom, 1905

*Anomodontia* Owen, 1860a

*Dicynodontia* Owen, 1860a

*Dicynodontoidea* Olson, 1944

**Definition:** All taxa more closely related to *Dicynodon lacerticeps* Owen, 1845 than *Oudenodon bainii* Owen, 1860b or *Emydops arctatus* (Owen, 1876) (Kammerer & Angielczyk, 2009).

**Remarks:** Kammerer & Angielczyk (2009) followed Cluver & King (1983) in treating Dicynodontoidea as a superfamily, and assigned authorship of it to Cope (1871). Cope (1871) was the first to use the orthography Dicynodontidae for the family containing *Dicynodon*, so he was considered to be author of Dicynodontoidea by the Principle of Coordination (Art. 36.1 of The Code; ICZN, 1999). Further research has revealed two problems with this, however. First, although his usage predated standardized familial suffixes in zoological nomenclature, Owen’s (1860a) explicit treatment of his taxon Dicynodontia as a family is considered equivalent to the establishment of Dicynodontidae under Art. 11.7.1.3 of The Code (ICZN, 1999). Thus, Owen (1860a) is the author of both the unranked higher taxon Dicynodontia and the family Dicynodontidae. Second, the earliest usage of the taxon Dicynodontoidea in the literature was not as a superfamily, but explicitly as an infraorder by Olson (1944). Thus, Dicynodontoidea represents a higher taxon outside the strictures of The Code and is not coordinate with Dicynodontidae (despite having the standard suffix for superfamily; this is the case for numerous higher taxa, e.g., Asteroidea, Hyracoidea, etc.)

*Daptocephalus* van Hoepen, 1934

**Type species:** *Daptocephalus leoniceps* (Owen, 1876).

**Included species:** *Daptocephalus huenei* (Haughton, 1932).

**Diagnosis:** Dicynodontoid characterized by the combination of a proportionally tall, steeply sloping snout, ventrally-directed tusks, narrow palatal portion of the premaxilla, median interorbital ridge, long, extremely narrow intertemporal bar, vertical orientation of the postorbitals in the intertemporal bar, zygomatic ramus of squamosal tall and broadly rounded at posterior end, and broad median pterygoid plate lacking a well-developed crista oesophagea. Premaxillary beak tip not sharply ‘hooked’ as in *Dinanomodon*. Weak anterior processes of frontals present, but not elongate, attenuate processes nearing or contacting premaxilla as in *Dinanomodon* and *Vivaxosaurus*.

*Daptocephalus huenei* (Haughton, 1932) comb. nov.
Holotype: SAM-PK-10630, a fragmentary skull, left scapula, and fragments of postcranial elements from Stockley’s locality B2, Kiwohe, Ruhuhu Basin, Tanzania.

Referred material: All referred material is from the Ruhuhu Basin of southwestern Tanzania. GPIT/RE/9316 (=K6), a partial skull (missing the right temporal arch and half of occiput) from Kingori; GPIT/RE/9317 (=K6), a partial skull (missing the snout tip, temporal arches, and edges of the occiput) from Kingori; GPIT/RE/9641 (=K2), a badly distorted skull from Kingori; SAM-PK-10634, a partial skull consisting of the snout and fragmentary skull roof from Stockley’s locality B16; UMZC T799, a skull roof and associated tusk from Stockley’s locality B4/6; UMZC T983, a largely unprepared partial skull (missing the temporal arches) from Stockley’s locality B4/4; UMZC T1282, a small partial skull roof broken into sections (the anteriormost of which appears too large to fit on the rest of the skull and likely represents a different individual) from an uncertain Usili locality.

Diagnosis: Distinguished from Daptocephalus leoniceps by its proportionally taller lacrimal and possibly by a longer premaxillary beak tip.

Remarks: The rationale for taxonomically differentiating these specimens from Dicynodon and referring them to Daptocephalus is provided above. Known material of Daptocephalus huenei is extremely similar to that of the South African type species D. leoniceps. However, there is at least one consistent difference: Tanzanian Daptocephalus specimens have an enlarged lacrimal compared to their South African congeners, with a taller facial portion between the prefrontal and maxilla (Fig. 6). A relatively tall lacrimal is present in the holotype, SAM-PK-10630, meaning that despite the extreme inadequacy of that specimen, it is diagnosable, and the name D. huenei can be retained for this morphotype. Although less certain of an autapomorphic feature due to frequent damage (it is missing in the holotype), in specimens of D. huenei that preserve the tip of the premaxillary beak, this structure is proportionally longer than in comparably intact, well-preserved D. leoniceps specimens. This is especially evident in GPIT/RE/9316 (Fig. 2D, 6A).

Most of the published material referable to Daptocephalus huenei is rather poor, and attempting a thorough redescription of this species here would be premature. The best-preserved of the described specimens, GPIT/RE/9316 (Fig. 2B,D, 5A,D), is missing most of the right side of the skull and its braincase is largely unprepared; other specimens are either highly distorted, fragmentary, or unprepared. However, recent expeditions to the Ruhuhu Basin have recovered several new specimens of D. huenei (K. Angielczyk, pers. comm., 2019), including a complete, relatively well-preserved skull (NMT RB43; C. Kammerer, pers. obs., see also Supplementary Data 4 of Angielczyk et al., 2018). The description of these new specimens should greatly improve our understanding of the anatomy of D. huenei, and may reveal additional features distinguishing it from D. leoniceps.

Dicynodon Owen, 1845

Type species: Dicynodon lacerticeps Owen, 1845.

Included species: Dicynodon angielczyki sp. nov.

Diagnosis: Dicynodontoid characterized by the combination of a relatively low, weakly-sloping snout, anteroventrally-directed tusks, broad, usually squared-off palatal portion of the premaxilla, median pterygoid plate distinctly constricted relative to rest of pterygoid in adults, bearing well-developed crista oesophagea, horizontal orientation of the postorbitals in the intertemporal bar,
relatively short intertemporal bar, and zygomatic and quadrate rami of the squamosal forming an
acute angle.

**Dicynodon angielczyki** sp. nov. (Fig. 2A,C, 3A, 5C,E, 7A,C, 8, 9, 10, 11, 12)

**Holotype:** UMZC T1089, a complete skull from Stockley’s locality B19, Kingori, Ruhuhu Basin, Tanzania.

**Referred material:** All referred material is from the Ruhuhu Basin of southwestern Tanzania. GPIT/RE/7175 (=K110), a complete skull from Kingori; GPIT/RE/7177 (=K101), a complete skull from Kingori; UMZC T979, a fragmented but nearly complete skull (missing parts of the snout roof and postorbital bars) from Stockley’s locality B4/6; UMZC T982, a fragmented but nearly complete skull (missing parts of the palate and skull roof) from Stockley’s locality B4/7; UMZC T1123, a complete skull, lower jaws, and possible skeletal elements preserved in association with two gorgonopsians from Stockley’s locality B4/3; UMZC T1126, a partial skull roof from Stockley’s locality B4/5.

**Diagnosis:** Distinguished from *Dicynodon lacerticeps* by expansion of the zygomatic ramus of the squamosal anteriorly, becoming dorsoventrally and mediolaterally swollen below the level of the postorbital bar; dorsoventral expansion of the jugal below the postorbital bar, such that the postorbital bone is widely separated from the squamosal; absence of a distinct postfrontal; and extension of the squamosal at the posterolateral corner of the temporal fenestra, forming a flange that curves slightly medially (Fig. 7).

**Etymology:** Named after the preeminent dicynodont researcher Kenneth Angielczyk, in particular recognition of his work on the Tanzanian and Zambian dicynodont faunas.

**Description:** Three intact, nearly-complete, thoroughly-prepared crania are known for *Dicynodon angielczyki*: GPIT/RE/7175 (Fig. 2A,C, 5C,E), GPIT/RE/7177 (Fig. 3A, 7A), and UMZC T1089 (Fig. 7C, 8, 9, 10). The following description is based primarily on UMZC T1089, which best illustrates the cranial sutures in this taxon (the skull roofs of the Tübingen specimens are somewhat overprepared). No associated mandibles are preserved with the aforementioned specimens, however, so the mandibular description is based on UMZC T1123. UMZC T1122–T1123 (Fig. 11, 12) is a large block of jumbled fossils from Stockley’s locality B4 (Katumbi viwili, also variously written as Katumbi vawili and Katumbi mwili), one of the most productive Permian fossil sites in the Ruhuhu Basin (Gay & Cruickshank, 1999; Angielczyk, 2007). The remains of two gorgonopsians and one dicynodont can be identified in this block. The two gorgonopsians appear to represent the same species, which can be recognized as a rubidgeine on the basis of the greatly expanded height of the zygomatic arch and extreme transverse width of the temporal region (Kammerer, 2016). Definite species level identification of these gorgonopsians is not possible at present due to incomplete exposure of their skulls (notably, the palates are not visible), but based on general proportions they are probably referable to the common Usili Formation rubidgeine *Sycosaurus nowaki*. The dicynodont is represented by a complete skull exposed in ventral view (33.0 cm basal skull length, making it slightly larger than the 31.5 cm holotype), some possible postcrania (although most of the postcranial material appears gorgonopsian), and a lower jaw disarticulated into its component rami. Only the posterior tip of the right mandibular ramus is exposed; the rest of the jaw descends into the block (Fig. 11B). However, the left mandibular ramus has been separated from the main block and fully prepared (Fig. 12). This specimen can be identified as a dicynodontid on the basis of an enlarged labial fossa and the uniformly rugose
palatine pad flush with the surrounding palate (as opposed to the condition in geikiids, where the
palatine pad is smoother and flush with the rest of the palate anteriorly but raised and extremely
rugose posteriorly), and can be recognized as *Dicynodon angielczyki* rather than *Daptocephalus huenei* based on the relatively narrow median pterygoid plate bearing a well-developed crista
oesophagea. It can further be distinguished from *Euptychognathus bathyrhynchus* by the greater
transverse width of the premaxilla and occiput and relatively low snout. Because of limited
exposure of the skull, this specimen provides little data on the cranium of this species not already
known in the specimens discussed above. However, it is important in providing the only available
information on mandibular morphology for *D. angielczyki*, and is the basis for the mandibular
description presented below.

The premaxilla of *D. angielczyki* is a fused median element forming a beak, as in all known
dicynodonts. It makes a triangular contribution to the dorsal surface of the snout (Fig. 8),
terminating above the nares (Fig. 9). A broad, low median ridge is present on the anterodorsal face
of the premaxilla. Lateral to this ridge the premaxillary surface is noticeably sculptured (Fig. 10),
probably associated with a keratinous beak covering in life. Ventrally, the premaxilla makes up a
broad secondary palatal plate extending posteriorly to contact the palatines and vomer. Well-
developed, parallel anterior palatal ridges are present at the tip of the beak. These do not converge
posteriorly, but are confluent with a weak median shelf anterior to the median palatal ridge. This
shelf is bounded laterally by elongate grooves, which continue posteriorly along the edges of the
tall, blade-like median palatal ridge.

The septomaxilla is a plate-like bone confined to the naris. The naris in *D. angielczyki* is
relatively large, and in addition to the actual opening into the nasal cavity incorporates a wide,
rounded embayment of the lateral snout surface composed of premaxilla, septomaxilla, and
maxilla (Fig. 9, 10A,B). The posterior edge of this embayment is bounded by a sharp, near-
vertically-oriented ridge running posterodorsal to anteroveltral on the maxilla. The ventral edge
of the embayment features a shallow groove in the maxillary surface, immediately lateral to its
contact with the septomaxilla. Ventrally, the maxilla forms a pointed caniniform process (Fig. 9)
housing the tusks. The tusks of UMZC T1089 are broken off (Fig. 8C,D), but other specimens of
*D. angielczyki* in which they are preserved (Fig. 2C) show that they were directed anteroveltrally,
as in *D. lacerticeps*. The posterior face of the caniniform process is depressed, such that its margin
is distinctly concave in cross-section (Fig. 8C,D). The bone surface of this depression is smooth,
unlike the rugose lateral surface of the maxilla. The posterior process of the maxilla that extends
into the zygomatic arch is relatively short in *D. angielczyki*, as it is ‘crowded out’ by the enlarged
anterior contributions of the squamosal and jugal. It has only a thin, attenuating ventral extension
below the posterior half of the orbit (Fig. 9).

The nasal is a broad bone making up much of the dorsal surface of the snout (Fig. 8A,B).
The dorsal surface of the nasals is rugose and densely foraminated. A lengthy mid-nasal suture
separates the premaxilla from the frontals. No anterior frontal process is present between the nasals
posteriorly (like *Dicynodon lacerticeps* but unlike many other dicynodontoids; e.g., frontal
processes are present in *Peramodon* and, albeit more weakly developed, in *Daptocephalus*, and
nearly contact the premaxilla in *Dinonomodon* and *Vivaxosaurus*; Kammerer et al., 2011). Discrete
nasal bosses are not present, but rather take the form of a single, raised, shelf-like area extending
across much of the nasal surface in the internarial region. The edges of this shelf extend as pointed
projections into the naris (Fig. 10A,B), giving the dorsal narial margin a ‘notched’ appearance.
Ventrally, the nasal contacts a dorsal extension of the maxilla, separating the septomaxilla and
lacrimal. The lacrimal is a small bone largely restricted to the anterior orbital margin, but has a
small anterior contribution between the nasal and maxilla (Fig. 9). A single, large lacrimal foramen
is present on the posterior face of the bone within the orbit, behind a knob-like lacrimal process
(Fig. 10A,B).

The surface of the prefrontal is also sculptured, but with finer ornamentation than on the
nasal (Fig. 9A, 10A). No prefrontal boss is present, although the dorsal orbital margin (extending
across the prefrontal, frontal, and postorbital) is uniformly weakly swollen. The frontal is a broad,
roughly rectangular bone making up most of the dorsal orbital margin. No interorbital ridge is
present on the mid-frontal suture; instead, there is an interorbital depression. Several large,
irregular pits are present on the frontal surface in UMZC T1089 (Fig. 8A). These appear to be
natural features of the bone (although they may have been exaggerated by acid preparation), and
are visible in other specimens of D. angielczyki where the skull roof has not been overprepared
(e.g., UMZC T979). Ventrally, the frontal curves to form the roof of the orbit, then extends
ventromedially to contact the orbital plate (Fig. 9). This element, also known as the anterior plate,
represents a fusion of the orbitosphenoid and mesethmoid (Cluver, 1971) and forms a median wall
separating the orbits.

The jugal typically has limited lateral exposure in dicynodontoids, mostly forming a thin
strip below the anterior half of the orbit. In D. angielczyki, the jugal is greatly expanded in size
relative to D. lacerticeps and other Permian dicynodontoids, with a tall contribution forming the
base of the postorbital bar and separating the postorbital bone from the squamosal (Fig. 7C, 9).
Ventrally, it forms much of the medial surface of the subtemporal arch, curving anteromedially to
cover the posterior face of the maxilla (Fig. 8C,D). In this region, the jugal, maxilla, and palatine
bound a large labial fossa, as is typical of dicynodontoids (Angielczyk & Kurkin, 2003; Kammerer
& Angielczyk, 2009).

The zygomatic ramus of the squamosal is greatly expanded at its anterior end, terminating
in a broadly-rounded tip that covers most of the suborbital portion of the maxilla in lateral view
(Fig. 9). Its greatest expansion is immediately posterior to the postorbital bar (Fig. 2C, 7C, 9),
resulting in the temporal arch being sharply bowed in this area (Fig. 2A, 7A, 8A, 10A). The
subtemporal zygoma is horizontally-oriented, as in D. lacerticeps, in contrast to the condition in
Daptocephalus where it is more vertical. Also as in D. lacerticeps, the posterior contact between
the zygomatic and quadrate squamosal rami forms an acute angle (Fig. 2C, 7C, 9) rather than a
broadly-rounded arc (as in Daptocephalus). The anterolateral surface of the quadrate squamosal
ramus is strongly depressed, producing a fossa for attachment of the M. adductor mandibulae
externus lateralis (Angielczyk et al., 2018). Posteriorly, the squamosal has a large contribution to
the lateral edge of the occipital plate (Fig. 10C,D). The dorsal and lateral edges of the squamosal
are attenuate occipitally, and extend somewhat posterior to the main portion of the occipital plate.
Medially on the occiput, the squamosal surface is depressed, particularly where it forms the lateral
margin of the post-temporal fenestra.

The parietal surrounds the anterior half of the oval pineal foramen (Fig. 2A, 7A, 8A,B). Its posterior portion is bounded laterally by thin anterior processes of the parietals. In all specimens
the parietal expands in transverse width anteriorly, and in specimens with well-exposed sutures
(e.g., UMZC T1089; Fig. 8A) the anterior margin is ragged with three distinct tips. The parietal
surface is depressed, and this depression is not contiguous with the parallel depressions on the
posterior frontal processes and postorbitals.

The postorbital contribution to the postorbital bar is gently curved and expands in width
ventrally (Fig. 8A,B), where it overlies the jugal. The anterodorsal margin of the postorbital is
somewhat rugose and forms part of the generally swollen dorsal rim of the orbit. The raised
anterior edge of the postorbital contribution to the skull roof bounds a deep, falciform depression that would have served as an attachment site for jaw adductor musculature. A distinct postfrontal is not present in *D. angielczyki*, but it is possible that this anterior portion of the postorbital incorporates the postfrontal, as a postfrontal is present in this region in *D. lacerticeps* (Cluver & King, 1983; Kammerer et al., 2011) and fusion between the postfrontal and postorbital over the course of ontogeny occurs in other dicynodonts (Kammerer & Smith, 2017; Angielczyk et al., 2018). The postorbital contribution to the intertemporal bar is relatively broad and more horizontally-oriented than vertical. It makes up the entire medial margin of the temporal fenestra and terminates in a curved process along the posterior edge of the temporal fenestra, overlaying the occiput. At its medial border the postorbital forms a thin crest around the parietal. The parietals are barely visible dorsally, with only very narrow exposure between the postorbitals and short anterior processes around the prepirietal. Laterally, the parietal is exposed below the postorbitals as part of the dorsal portion of the braincase, and is visible where it contacts the ascending process of the epipterygoid (Fig. 9).

The anteriormost visible portion of the vomer is immediately behind the median palatal ridge of the premaxilla, and continues as a similar blade-like structure posteriorly (Fig. 8C,D). Because the palate has been fully acid-prepared in UMZC T1089, the dorsal portion of the vomer is also visible in ventral view, a rarity among dicynodont specimens. It forms paired, vaulted laminae curving ventrolaterally to contact the palatines laterally and pterygoids posterolaterally. Medially, it surrounds an elongate, ‘teardrop’-shaped interpterygoid vacuity (narrow end anterior), bearing thin ridges along the margin of the vacuity.

A distinct ectopterygoid is not clearly present in any of the known specimens of *D. angielczyki*, but sutural edges of what may be the ectopterygoid are visible in UMZC T982 and UMZC T1123. The ectopterygoid is definitely absent as a separate ossification (either not ossifying or fused with one of the surrounding bones, probably the maxilla) in many Triassic dicynodonts, such as *Lystrosaurus* and some kannemeyeriforms (Cluver, 1971; Maisch, 2002; Angielczyk et al., 2018), but is usually present in Permian, “*Dicynodon*”-grade dicynodontoids (Kammerer et al., 2011; there are exceptions, however—see Angielczyk & Kurkin, 2003). Here, a probable outline for the ectopterygoid is shown (Fig. 8D) based on the morphology in the aforementioned specimens, but this is tentative.

The palatine of *D. angielczyki* is typical for dicynodontoids, with a raised, rugose palatine pad anteriorly and a smooth, laminar section forming part of the lateral wall of the choana posteriorly (Fig. 8C,D). A small, rounded lateral palatal foramen is present between the maxilla, (possibly) ectopterygoid, and palatine, at around the midpoint of the latter. The pterygoids form a roughly X-shaped unit composed of anterior and posterior (or quadrate) rami united by a median pterygoid plate, as is the case in all dicynodonts (King, 1988). The anterior pterygoid rami are bowed laterally, surrounding a broad choana. Narrow ridges are present on the posterior halves of the anterior pterygoid rami, which unite at the median pterygoid plate to form a well-developed crista oesophagea. The median pterygoid plate is strongly constricted relative to the anterior and posterior rami, as is usual for *Dicynodon* (Fig. 3A,C, 8C,D) but not *Daptocephalus* (Fig. 3B,D,E).

The posterior or quadrate rami are thin, ribbon-like structures extending from the median pterygoid plate towards the quadrates at a 30–45° angle relative to the long axis of the skull. In all three specimens where these fragile structures are preserved, the posterior pterygoid rami are slightly twisted through their length (Fig. 3A, 8C, 11B). Twisting of the quadrate ramus of the pterygoid is naturally present in some therapsid taxa (e.g., gorgonopsians; Kammerer, 2016), but can probably be attributed to taphonomic distortion in *D. angielczyki*—this ramus is usually straight.
in dicynodonts (e.g., Fig. 3C) and its shape is asymmetrical in all the studied specimens of D. angielczyki (note that the left posterior ramus in UMZC T1123 is straight for most of its length then broken at tip, whereas twisting in the right ramus appears to be the result of the anterior face of the ramus being displaced ventrally through crushing; Fig. 11B). Dorsally, the pterygoid bears a median, laminar cultriform process extending anteriorly and terminating below the orbits, where it is underlain by the vomer. Resting on top of the pterygoids posteriorly are paired epipterygoids, which consist of an anteroposteriorly elongate footplate and a thin ascending process, which expands dorsally where it contacts the parietal (Fig. 9).

A clear sutural boundary between the pterygoid and parabasisphenoid is not visible in any of the studied specimens, but the anterior extent of the latter can be recognized by the presence of paired, ventrally-directed internal carotid canals (Fig. 8C,D). Posterior to these openings, the ventral surface of the parabasisphenoid bears paired ridges (with a marked depression between them) that curve posterolaterally and expand to join the basal tubera. The tubera are semi-oval in shape and are angled somewhat ventrolaterally. A deep intertuberal depression is present medially between the parabasisphenoid and basioccipital. Only the anterior edges of the tubera are composed of parabasisphenoid, the rest is basioccipital. The basioccipital does not show distinct sutures with the opisthotic (Fig. 8B,C), exoccipitals, or supraoccipital (Fig. 10C,D) and it appears that they form a single fused element, also incorporating the prootic and thus representing a periotic. Fusion of some or all of these occipital and basiocranial elements is common in dicynodonts (e.g., Surkov & Benton, 2004; Boos et al., 2016; Angielczyk & Kammerer, 2017). The contributions of the basioccipital and exoccipitals can still be discerned in the tripartite occipital condyle, as they each make up a distinct, knob-like process, but these processes are fused at the base (as can be seen due to damage to the condyle in UMZC T1089, revealing uniform bone internally; Fig. 10C). The paroccipital processes are large, wing-like structures in D. angielczyki, with a curved, protruding ventral edge and a broad depression of their dorsal posterior surface, below their contribution to the margin of the post-temporal fenestra. Dorsal to the post-temporal fenestra, what is presumably the supraoccipital bears prominent, dorsolateral-to-ventromedially angled ridges. The dorsal portion of the supraoccipital is weakly depressed on its posterior face, and its dorsal margin is split by a ventral process of the postparietal.

The postparietal (or interparietal) is a roughly trapezoidal, plate-like bone situated at the dorsal midpoint of the occiput (Fig. 10C,D). It does not make a noticeable contribution to the skull roof. It is bounded laterally by the tabulars, which are small, flat bones made up of a narrow dorsal process and a broader ventral plate.

The quadrate and quadratojugal are fused to form a large element anteroventral to the quadrate ramus of the squamosal (Fig. 9) and exhibit the usual morphology for dicynodonts (King, 1988). The quadratojugal forms a broad but thin plate separated from the dorsal portion of the quadrate by a large quadratojugal foramen or channel. It is mostly occluded in posterior view by a ventral process of the squamosal (Fig. 10D). The articular surface of the quadrate ventrally is made up of lateral and medial condyles of roughly equal size separated by a trochlea (Fig. 8C,D). The quadrate is angled such that the medial condyle is situated somewhat posterior to the lateral one.

The mandible (based on UMZC T1123) is edentulous and primarily composed of a large, robust dentary (Fig. 12). As usual for dicynodonts, the dentary is a single fused structure; separation of the left mandibular ramus in this specimen was managed by cutting it through the symphysis, not through natural disarticulation of the two hemimandibles. The anterior and lateral surface of the dentary is rugose; this is at least partially the result of acid preparation but some of the rugosity on the anterior and dorsal edges of the symphysis appears natural and probably
corresponds to coverage by the keratinous beak. Anterodorsally the dentary extends to a curved, pointed tip terminating well above the dorsal edge of the rest of the mandible. The border between the anterior and lateral faces of the dentary are weakly demarcated, without a sharp ridge between them. Posterolaterally, the dentary terminates in two processes above and below the mandibular fenestra. Below the mandibular fenestra, the posterior process of the dentary is triangular and fits into a deep facet on the angular. Above the fenestra, the dentary has a longer posterior process, which attenuates somewhat posteriorly but does not come to a discrete point. Rather, the posterior edge of the dentary in this region remains tall where it overlies the surangular. The posterior margin of this process has a distinct concavity, as figured by Cluver & King (1983) for South African *Dicynodon*. A lateral dentary shelf is present at the anterodorsal edge of the mandibular fenestra. It is weakly developed and thin, as in many other Permian bidentalians, but unlike *Aulacephalodon* where it extends anteriorly to join an enlarged, rounded boss (Kammerer et al., 2011).

The mandibular fenestra is elongate and bounded dorsally by the dentary and surangular and ventrally by the angular. The splenial, like the dentary, is a single fused element. It forms a large portion of the ventral margin of the symphysis (as is typical of dicynodontoids; see Kammerer, 2018) and continues posteriorly along the medial face of the jaw, terminating in an attenuate process overlying the angular (Fig. 12E). A thin process of the angular extends forwards between the splenial and dentary to contribute to the symphysis. Posteriorly, the angular consists of a flat, ribbon-like element covered by dentary laterally and splenial medi ally. It is well-exposed below the mandibular fenestra and bears a posteroventrally-directed reflected lamina beneath the posterior edge of the fenestra. The reflected lamina is somewhat damaged but appears typical for dicynodontoids: it is a free-standing structure (i.e., the posterior edge is not bound to the main jaw ramus) and bears one major ridge surrounded by two surficial concavities (with weaker ridges along the anteroventral and posterodorsal edges of the lamina). A flat portion of angular forms most of the lateral surface of the jaw posterior to the reflected lamina.

The surangular and prearticular are clearly separate medially (Fig. 12E), but the distinction between these bones and the articular is unclear and at least partial fusion between these three elements is likely. Laterally, the surangular is exposed as a narrow strip above the angular along the dorsal edge of the post-dentary jaw ramus. Due to taphonomic distortion, it is somewhat displaced in UMZC T1123 and would originally have had a greater degree of lateral exposure. The prearticular is also damaged, with its anterior tip broken off. When complete it would have extended anteriorly to overlie the splenial. Medially, both the surangular and prearticular are thin, ribbon-like elements bracing the angular wall. The articular morphology is similar to that of other dicynodontoids, consisting of lateral and medial condyles around a median trochlea, where it would articulate with the quadrate (Fig. 12C). No retroarticular process is evident, but this is probably attributable to damage, as this structure is present in other *Dicynodon* (Cluver & King 1983).

**PHYLOGENETIC ANALYSIS**

*Dicynodon huenei* has been a problematic taxon in recent analyses of dicynodont phylogeny. In part, this is due to the general instability of the Permian dicynodontoid portion of the tree (see discussion in Angielczyk & Kammerer, 2017), but here, its problematic status is also recognized as the result of chimaerical codings, including data from what are probably three distinct species (*Daptocephalus huenei*, *Dicynodon angielczyki*, and so-called “*D. huenei””/“*D. trigonocephalus” from Zambia). Kammerer et al. (2011) were the first to include *D. huenei* in a cladistic analysis of
anomodont phylogeny, and recovered it as the sister-taxon of *D. lacerticeps* in their primary tree. However, support for this relationship was extremely low, and in variant analyses the genus *Dicynodon* sensu Kammerer *et al.* (i.e., containing *D. lacerticeps* and *D. huenei*) was not found to be monophyletic. Kammerer *et al.*'s (2011) discrete state codings for *D. huenei* were based on cranial data mostly from UMZC T1089, mandibular data mostly from UMZC T1123, and postcranial data from the Zambian specimen NHMUK PV R37005 (=TSK 14), and continuous codings were based on GPIT/RE/7175 (=K110), GPIT/RE/7177 (=K101), GPIT/RE/9316 (=K2), NMT RB43, NHMUK PV R37005, NHMUK PV R37374 (=TSK 37), UMZC T979, UMZC T982, UMZC T987, UMZC T1089, and UMZC T1123.

In the current analysis, the previous "*Dicynodon huenei*" operational taxonomic unit (OTU) has been deleted and replaced with separate OTUs for *Dicynodon angielczyki* (coded based on GPIT/RE/7175, GPIT/RE/7177, UMZC T979, UMZC T982, UMZC T1089, and UMZC T1123) and *Daptocephalus huenei* (coded based on GPIT/RE/9316, GPIT/RE/9317, GPIT/RE/9641, NMT RB43, and SAM-PK-10630). These OTUs were added (see Supplemental Information) to the most recent iteration of the anomodont character matrix originally published by Kammerer *et al.* (2011), namely that of Kammerer *et al.* (2019). The emydopoid *Thliptosaurus imperforatus* (Kammerer, 2019) was also added to that analysis, but the recently-described Laotian dicynodontoids *Couvillonia superoculis* and *Repelinosaurus robustus* (Olivier *et al.*, 2019) were not included, as their holotypes were not available for study during the production of this paper.

The data were analyzed using parsimony in TNT v1.5 (Goloboff *et al.*, 2008) using New Technology search parameters (tree drifting, parsimony ratchet, and tree fusing), starting at level 65 and forced to find the shortest tree at least 20 times. Discrete-state characters 58, 61, 79, 140, 150, 151, and 166 were treated as ordered. Symmetric resampling values were calculated based on 10000 replicates.

Two most parsimonious trees of length 1158.54 were recovered (consistency index=0.238, retention index=0.720), differing only in the positions of genera within Placeriinae (consensus topology for bidentalian dicynodonts shown in Fig. 13). The overall tree topology is very similar to that of Kammerer (2019) and Kammerer *et al.* (2019), and the relationships of non-bidentalian anomodonts in the most parsimonious trees are identical to those of Kammerer (2019). Relationships among bidentalians remain, with few exceptions, poorly-supported and labile. Unlike the analysis of Kammerer *et al.* (2019), but as in Kammerer (2019; see also Boos *et al.*, 2016; Angielczyk & Kammerer, 2017; Olroyd *et al.*, 2018), Cryptodontia is not recovered as monophyletic in its traditional sense (i.e., rhachiocephalids+geikiids are more closely related to "*Dicynodon*-grade dicynodontoids thanoudenodontids). However, despite continued uncertainty surrounding the intergeneric relationships of bidentalians, the current analysis does recover *Dicynodon* (*D. lacerticeps*+*D. angielczyki*) and *Daptocephalus* (*D. leoniceps*+*D. huenei*) as monophyletic, with *Daptocephalus* being one of the few strongly supported clades in the resampling analysis (Fig. 13). The *Dicynodon* clade is supported by two characters (continuous character 14, shape of mandibular fenestra, and discrete state character 51, postorbital contribution to intertemporal bar relatively flat) and the *Daptocephalus* clade is supported by four characters (continuous characters 8, median pterygoid plate width, and 11, relative area of internal nares, and discrete state characters 92, ventral surface of median pterygoid plate smooth and flat, and 111, absence of central circular depression on occipital condyle) (refer to Supplementary Information for lists of all character states).

**DISCUSSION**
The discovery that multiple taxa of large dicynodontoids are present in the Usili Formation should not be surprising—this clade is common in the late Permian, and sympatric dicynodontoid taxa are known within more poorly-sampled, probably coeval Laurasian basins (Li et al., 2008; Kurkin, 2012; Olivier et al., 2019) in addition to the heavily sampled Karoo Basin of South Africa (Kammerer et al., 2011). What is notable about the Usili species, however, is their close similarity to (and recovery as sister-taxa of) particular Karoo species (Dicynodon lacerticeps and Daptocephalus leoniceps), to the extent that they are here considered congeneric. Although genera are an arbitrary taxonomic construct, in the current case the differences between the Tanzanian and South African species are so few (especially in the case of Daptocephalus huenei vs. D. leoniceps) that generic separation seems unwarranted.

The present work on “Dicynodon huenei” has benefited from a detailed view of dicynodont variation only possible after decades of taxonomic revision of the group (see reviews in Kammerer & Angielczyk, 2009; Kammerer et al., 2011). Research on the expansive sample of fossils from the Karoo Basin in particular has provided substantial insight into what is most parsimoniously interpreted as intraspecific (mostly ontogenetic, sexual, and taphonomic) variation in dicynodont species, which was often historically interpreted as representing distinct species (e.g., Broom, 1932). Given the unwieldy and egregiously oversplit nature of historical taxonomic schemes for dicynodonts (recognized as such even at the time; see review in Kammerer et al., 2011), latter-day revisions of dicynodont taxonomy have understandably focused on synonymizing the many nominal taxa. However, another result of the more nuanced view of dicynodont taxonomy currently available is the ability to tease apart instances of lumping: cases where previously-synonymized taxa are shown to be distinct (e.g., Kammerer et al., 2015) or where some specimens previously referred to common taxa represent unrecognized species (e.g., Kammerer et al., 2016; Kammerer & Smith, 2017). Such cases are likely to increase in number as previously poorly-sampled basins become better known, and taxa previously known from limited material once difficult to distinguish from Karoo species become more robustly diagnosed. For example, Keyser (1975) recognized only a single valid rhachiocephalid species, Rhachiocephalus magnus, including material from the Usili Formation, but more recent study (Maisch, 2005) has shown that the Tanzanian specimens are specifically distinct from South African Rhachiocephalus. Kammerer et al. (2011) could not find consistent differences separating Zambian specimens of Oudenodon (previously known as O. luangwanensis) and the South African type species O. bainii, and considered only O. bainii to be valid. However, the type materials of Zambian Oudenodon they examined were generally poor (the holotype of O. luangwanensis, SAM-PK-11310, for example, is almost entirely unprepared, with only the skull roof exposed), and would not necessarily have shown the ‘species-level’ variation recognized for, e.g., Daptocephalus here. Angielczyk et al. (2014) figured additional, complete and well-prepared Zambian Oudenodon specimens, and although they also referred these specimens to O. bainii (following Kammerer et al., 2011), the new specimens do show some consistent proportional differences from typical South African O. bainii. Even more recently-discovered, well-preserved Zambian Oudenodon specimens are now known (K. Angielczyk, pers. comm., 2019). Detailed study of these specimens is needed to determine whether they form a discrete morphotype from Karoo specimens, but preliminary information is suggestive of their distinctness. As a final example, Angielczyk (2019) recently described the first specimen of the rare dicynodont Digalodon (previously known only from South Africa) from the Luangwa Basin of Zambia. Although clearly referable to Digalodon, the Zambian specimen (NHCC LB830) differs from South African skulls in several notable regards (non-
diverging anterior median palatal ridges, rounded postcaniniform keel, horizontal zygoma, continuous rim of the pineal foramen, proportionally broader intertemporal bar). Angielczyk (2019) considered this specimen taxonomically uncertain and classified it as *Digalodon cf. D. rubidgei*, a reasonable approach given its singleton nature and our poor knowledge of variation in South African *Digalodon* (known from a few, mostly poorly-preserved and prepared specimens; Kammerer et al., 2015). However, the unique features of NHCC LB830 are not known to vary intraspecifically in other emydopoids, and it is likely to represent a distinct species.

The emerging pattern of ‘low-level’ or ‘species-level’ endemism among dicynodonts in the African late Permian basins is interesting in the context of recent proposals concerning changing biogeographic patterns across the Permo-Triassic boundary. Based on network analyses of Permo-Triassic vertebrate assemblages, Sidor et al. (2013) argued that there is a sharp increase in provincialism among tetrapods between the Permian and Middle Triassic. It is true that in the late Permian, dicynodont genera such as *Pristerodon*, *Oudenodon*, *Dicynodontoides*, and *Endothiodon* (cited by Sidor et al. (2013) as cosmopolitan taxa) have a broad, interbasinal distribution, and that Middle Triassic faunas show greater taxonomic/phylogenetic separation (at least between the African basins, though some evidence suggests greater similarity between Zambian/Tanzanian and coeval South American assemblages; see Peecook et al., 2018). The recognition that a number (potentially many) of these wide-ranging dicynodont genera contain multiple, locally endemic species adds a new wrinkle to this proposal, however. I would suggest that rather than a simple transition from cosmopolitan Permian tetrapod faunas to provincialized Triassic ones following the Permo-Triassic mass extinction, there is a shift between ‘weakly provincialized’ faunas (substantial phylogenetic propinquity between basins, but frequent species-level endemism in each) in the late Permian, to true cosmopolitanism associated with the spread of ‘disaster taxa’ in the wake of the Permo-Triassic extinction (as in individual species like *Lystrosaurus murrayi* with well-supported circum-Gondwanan distributions; Colbert, 1974; Ray, 2005), and finally ‘strongly provincialized’ faunas in the Middle Triassic (distinct at higher taxonomic levels; probably, as Sidor et al., 2013 argued, resulting from heterogeneous re-occupation of empty ecospace during ecosystem recovery). Continued research on late Permian faunas, particularly from basins outside of the well-sampled Karoo, is needed to test this proposal. Additional paleoecological data is also needed—given the apparent importance of local climate in driving Triassic tetrapod distributions (Whiteside et al., 2011), it needs to be determined whether the heterogeneous repopulation of faunas in the Middle Triassic is stochastic, or whether taxon composition was driven by local environments.

**CONCLUSIONS**

*Dicynodon huenei*, a supposed taxon of “*Dicynodon*-grade dicynodontoid from upper Permian strata of Tanzania and Zambia, is here recognized as being made up of several distinct dicynodont species. The Tanzanian specimens of “*D. huenei*” constitute two species, here reclassified as *Daptocephalus huenei* comb. nov. and *Dicynodon angielczyki* sp. nov. Permian dicynodontoids have proven to be one of the most troublesome regions of dicynodont phylogeny, exhibiting substantial instability between recent phylogenetic analyses. Separation of “*Dicynodon huenei*” into multiple OTUs resolves some problems in recent analyses (i.e., occasional polyphyly of the genus *Dicynodon*), but support for the current topology remains low. Re-evaluation and expansion of the character data for dicynodontoids is required for better resolution in this part of the tree.
ACKNOWLEDGEMENTS

Thanks to Ken Angielczyk for invaluable assistance during this project, including permitting me to examine undescribed dicynodont specimens collected in the recent TZAM expeditions. Thanks also to Christian Sidor and Roger Smith for various discussions on the fauna and stratigraphy of the Usili Formation. I thank Paul Barrett, Sifelani Jirah, Tom Kemp, Mathew Lowe, Robert and Marion Rubidge, Zaituna Skosan, and Ingmar Werneburg for access to specimens in their care, and Michael Day for catalogue information on recently-accessioned NHMUK specimens. Reviews by Jun Liu, Savannah Olroyd, and Brandon Peecook improved the manuscript, as did the efforts of editor Andrew Farke.

REFERENCES

Angielczyk KD. 2007. New specimens of the Tanzanian dicynodont “Cryptocynodon parringtoni” von Huene, 1942 (Therapsida, Anomodontia), with an expanded analysis of Permian dicynodont phylogeny. Journal of Vertebrate Paleontology 27:116–131.

Angielczyk KD. 2019. First occurrence of the dicynodont *Digalodon* (Therapsida, Anomodontia) from the Lopingian upper Madumabisa Mudstone Formation, Luangwa Basin, Zambia. Palaeontologia africana 53:219–225.

Angielczyk KD, Cox CB. 2015. Distinctive emydopoid (Therapsida, Anomodontia) mandibles from the Permian Ruhuhu and Usili formations (Songea Group), Ruhuhu Basin, Tanzania. Journal of Vertebrate Paleontology e1008699.

Angielczyk KD, Kammerer CF. 2017. The cranial morphology, phylogenetic position and biogeography of the upper Permian dicynodont *Compsodon helmoedi* van Hoepen (Therapsida, Anomodontia). Papers in Palaeontology 3:513–545.

Angielczyk KD, Kurkin AA. 2003. Phylogenetic analysis of Russian Permian dicynodonts (Therapsida: Anomodontia): implications for Permian biostratigraphy and Pangaean biogeography. Zoological Journal of the Linnean Society 139:157–212.
Angielczyk KD, Hancox PJ, Nabavizadeh A. 2018 (for 2017). A redescription of the Triassic kannemeyeriiform dicynodont *Sangusaurus* (Therapsida, Anomodontia), with an analysis of its feeding system. *Society of Vertebrate Paleontology Memoir* 17:189–227.

Angielczyk KD, Sidor CA, Nesbitt SJ, Smith RMH, Tsuji L.A. 2009. Taxonomic revision and new observations on the postcranial skeleton, biogeography, and biostratigraphy of the dicynodont genus *Dicynodontoides*, the senior synonym of *Kingoria* (Therapsida, Anomodontia). *Journal of Vertebrate Paleontology* 29:1174–1187.

Angielczyk KD, Steyer J-S, Sidor CA, Smith RMH, Whatley RL, Tolan S. 2014. Permian and Triassic dicynodont (Therapsida: Anomodontia) faunas of the Luangwa Basin, Zambia: Taxonomic update and implications for dicynodont biogeography and biostratigraphy. In: Kammerer CF, Angielczyk KD, Fröbisch J, eds. *Early Evolutionary History of the Synapsida*. Dordrecht: Springer, 93–138, 337.

Boos ADS, Kammerer CF, Schultz CL, Soares MB, Ilha ALR. 2016. A new dicynodont (Therapsida: Anomodontia) from the Permian of southern Brazil and its implications for bidentalian origins. *PLoS ONE* 11(5):e0155000.

Botha-Brink J, Huttenlocker AK, Modesto SP. 2014. Vertebrate paleontology of Nooitgedacht 68: a *Lystrosaurus maccaigi*-rich Permo-Triassic boundary locality in South Africa. In: Kammerer CF, Angielczyk KD, Fröbisch J, eds. *Early Evolutionary History of the Synapsida*. Dordrecht: Springer, 289–304, 337.

Brink AS. 1986 (for 1982). Illustrated Bibliographic Catalogue of the Synapsida. *Handbook of the South African Geological Survey* 10.

Broom R. 1905. On the use of the term Anomodontia. *Records of the Albany Museum* 1:266–269.

Broom R. 1932. *The Mammal-Like Reptiles of South Africa and the Origin of Mammals*. London: H. F. & G. Witherby, 376.

Cluver MA. 1971. The cranial morphology of the dicynodont genus *Lystrosaurus*. *Annals of the South African Museum* 56:155–274.

Cluver MA, King GM. 1983. A reassessment of the relationships of Permian Dicynodontia (Reptilia, Therapsida) and a new classification of dicynodonts. *Annals of the South African Museum* 91:195–273.

Colbert EH. 1974. *Lystrosaurus* from Antarctica. *American Museum Novitates* 2535:1–44.

Cope ED. 1871. On the homologies of some of the cranial bones of the Reptilia, and on the systematic arrangement of the class. *Proceedings of the American Association for the Advancement of Science* 19:194–247.
Cox CB, Angielczyk KD. 2015. A new endothiodont dicynodont (Therapsida, Anomodontia) from the Permian Ruhuhu Formation (Songea Group) of Tanzania and its feeding system. *Journal of Vertebrate Paleontology* e935388.

Gay SA, Cruickshank ARI. 1999. Biostratigraphy of the Permian tetrapod faunas from the Ruhuhu Valley, Tanzania. *Journal of African Earth Sciences* 29:195–210.

Goloboff PA, Farris JS, Nixon KC. 2008. TNT, a free program for phylogenetic analysis. *Cladistics* 24:774–786.

Haughton SH. 1932. On a collection of Karroo vertebrates from Tanganyika Territory. *Quarterly Journal of the Geological Society of London* 88:634–668.

Haughton SH, Brink AS. 1954. A bibliographical list of Reptilia from the Karroo beds of Africa. *Palaeontology africana* 2:1–187.

Huene F von. 1942. Die Anomodontier des Ruhuhu-Gebietes in der Tübinger Sammlung. *Palaeontographica Abteilung A: Paläozoologie-Stratigraphie* 94:154–184.

Huttenlocker AK, Sidor CA. 2016. The first karenitid (Therapsida, Therocephalia) from the upper Permian of Gondwana and the biogeography of Permo-Triassic therocephalians. *Journal of Vertebrate Paleontology* e1111897.

Huttenlocker AK, Sidor CA, Angielczyk KD. 2015. A new eutherocephalian (Therapsida, Therocephalia) from the upper Permian Madumabisa Mudstone Formation (Luangwa Basin) of Zambia. *Journal of Vertebrate Paleontology* e969400.

ICZN. 1999. *International Code of Zoological Nomenclature*. Fourth Edition. London: The International Trust for Zoological Nomenclature, 306.

Jasinoski SC, Cluver MA, Chinsamy A, Reddy BD. 2014. Anatomical plasticity in the snout of *Lystrosaurus*. In: Kammerer CF, Angielczyk KD, Fröbisch J, eds. *Early Evolutionary History of the Synapsida*. Dordrecht: Springer, 139–149, 337.

Kammerer CF. 2016. Systematics of the Rubidgeinae (Therapsida: Gorgonopsia). *PeerJ* 4:e1608.

Kammerer CF. 2019. A new dicynodont (Anomodontia: Emydopoidea) from the terminal Permian of KwaZulu-Natal, South Africa. *Palaeontology africana* 53:179–191.

Kammerer CF, Angielczyk KD. 2009. A proposed higher taxonomy of anomodont therapsids. *Zootaxa* 2018:1–24.

Kammerer CF, Smith RMH. 2017. An early geikiid dicynodont from the *Tropidostoma* Assemblage Zone (late Permian) of South Africa. *PeerJ* 5:e2913.
Kammerer CF, Bandyopadhyay S, Ray S. 2016. A new taxon of cistecephalid dicynodont from the upper Permian Kundaram Formation of India. *Papers in Palaeontology* **2**:569–584.

Kammerer CF, Angielczyk KD, Fröbisch J. 2011. A comprehensive taxonomic revision of *Dicynodon* (Therapsida, Anomodontia) and its implications for dicynodont phylogeny, biogeography, and biostratigraphy. *Society of Vertebrate Paleontology Memoir* **11**:1–158.

Kammerer CF, Angielczyk KD, Fröbisch J. 2015. Redescription of *Digalodon rubidgei*, an emydopoid dicynodont (Therapsida, Anomodontia) from the Late Permian of South Africa. *Fossil Record* **18**:43–55.

Kammerer CF, Viglietti PA, Hancox PJ, Butler RJ, Choiniere JN. 2019. A new kannemeyeriiform dicynodont (*Ufudocyclops mukanelai* gen. et sp. nov.) from Subzone C of the *Cynognathus* Assemblage Zone (Triassic of South Africa) with implications for biostratigraphic correlation with other African Triassic faunas. *Journal of Vertebrate Paleontology* e1596921.

Keyser AW. 1975. A re-evaluation of the cranial morphology and systematics of some tuskless Anomodontia. *Memoirs of the Geological Survey of South Africa* **67**:1–110.

King GM. 1981. The functional anatomy of a Permian dicynodont. *Philosophical Transactions of the Royal Society of London Series B* **291**:243–322.

King GM. 1988. *Anomodontia. Handbuch der Paläoherpetologie* 17 C. Stuttgart: Gustav Fischer Verlag, 174.

Kurkin AA. 2012. Dicynodontids of Eastern Europe. *Paleontological Journal* **46**:187–198.

Li J, Wu X, Zhang F. 2008. *The Chinese Fossil Reptiles and Their Kin*. Beijing: Science Press, 474.

Maisch MW. 2002. A new basal lystrosaurid dicynodont from the Upper Permian of South Africa. *Palaeontology* **45**:343–359.

Maisch MW. 2005. Observations on Karoo and Gondwana vertebrates. Part 6. A new rhachiocephalid dicynodont (Therapsida) from the Upper Permian of the Ruhuhu Basin (Tanzania). *Neues Jahrbuch für Geologie und Paläontologie Abhandlungen* **237**:313–355.

Nowack E. 1937. Zur kenntniss der Karruformation in Ruhuhu-Graben (D.O.A.). *Neues Jahrbuch für Mineralogie, Geologie und Paläontologie Beilage-Band Abteilung B* **78**:380–412.

Olivier C, Battail B, Bourquin S, Rossignol C, Steyer J.-S., Jali N.-E. 2019. New dicynodonts (Therapsida, Anomodontia) from near the Permo-Triassic boundary of Laos: implications for dicynodont survivorship across the Permo-Triassic mass extinction and the paleobiogeography of southeast Asian blocks. *Journal of Vertebrate Paleontology* e1584745.
Olroyd SL, Sidor CA, Angielczyk KD. 2018. New materials of the enigmatic dicynodont Abajudon kaayai (Therapsida, Anomodontia) from the lower Madumabisa Mudstone Formation, middle Permian of Zambia. *Journal of Vertebrate Paleontology* e1403442.

Olson EC. 1944. Origin of the mammals based upon cranial morphology of the therapsid suborders. *Geological Society of America Special Paper* 55:1–136.

Osborn HF. 1903. On the primary division of the Reptilia into two sub-classes, Synapsida and Diapsida. *Science* 17:275–276.

Owen R. 1845. Report on the reptilian fossils of South Africa. Part I.—Description of certain fossil crania, discovered by A. G. Bain, Esq., in sandstone rocks at the south-eastern extremity of Africa, referable to different species of an extinct genus of Reptilia (*Dicynodon*), and indicative of a new tribe or sub-order of Sauria. *Transactions of the Geological Society of London, Second Series* 7:59–84.

Owen R. 1860a. On the orders of fossil and recent Reptilia, and their distribution in time. *Report of the Twenty-Ninth Meeting of the British Association for the Advancement of Science* 1859:153–166.

Owen R. 1860b. On some reptilian fossils from South Africa. *Quarterly Journal of the Geological Society of London* 16:49–63.

Owen R. 1876. *Descriptive and Illustrated Catalogue of the Fossil Reptilia of South Africa in the Collection of the British Museum*. London: Taylor and Francis, 88.

Owen R. 1880. Description of parts of the skeleton of an anomodont reptile (*Platypodosaurus robustus*, Ow.) from the Trias of Graaff Reinet, S. Africa. *Quarterly Journal of the Geological Society* 36:414–425.

Peecook BR, Steyer JS, Tabor NJ, Smith RMH. 2018 (for 2017). Updated geology and vertebrate paleontology of the Triassic Ntawere Formation of northeastern Zambia, with special emphasis on the archosauromorphs. *Society of Vertebrate Paleontology Memoir* 17:8–38.

Ray S. 2005. *Lystrosaurus* (Therapsida, Dicynodontia) from India: taxonomy, relative growth and cranial dimorphism. *Journal of Systematic Palaeontology* 3:203–221.

Sidor CA, Nesbitt SJ. 2018 (for 2017). Introduction to vertebrate and climatic evolution in the Triassic rift basins of Tanzania and Zambia. *Society of Vertebrate Paleontology Memoir* 17:1–7.

Sidor CA, Angielczyk KD, Weide DM, Smith RMH, Nesbitt SJ, Tsuji LA. 2010. Tetrapod fauna of the lowermost Usili Formation (Songea Group, Ruhuhu Basin) of southern Tanzania, with a new burnetiid record. *Journal of Vertebrate Paleontology* 30:696–703.

Sidor CA, Vilhena DA, Angielczyk KD, Huttenlocker AK, Nesbitt SJ, Peecook BR, Steyer JS, Smith RMH, Tsuji LA. 2013. Provincialization of terrestrial faunas following the end-
Permian mass extinction. *Proceedings of the National Academy of Sciences of the United States of America* **110**:8129–8133.

Stockley GM. 1932. The geology of the Ruhuhu coalfields, Tanganyika Territory. *Quarterly Journal of the Geological Society of London* **88**:610–622.

Stockley GM, Oates F. 1931. Report on geology of the Ruhuhu Coalfields. *Bulletin of the Geological Survey of Tanganyika* **2**:1–68.

Surkov MV, Benton MJ. 2004. The basicranium of dicynodonts (Synapsida) and its use in phylogenetic analysis. *Palaeontology* **47**:619–638.

van Hoepen ECN. 1934. Oor die indeling van die Dicynodontidae na aanleiding van nuew vorme. *Paleontologiese Navorsing van die Nasionale Museum* **2**:67–101.

Viglietti PA, Smith RMH, Angielczyk KD, Kammerer CF, Fröbisch J, Rubidge BS. 2016. The *Daptocephalus* Assemblage Zone (Lopingian), South Africa: A proposed biostratigraphy based on a new compilation of stratigraphic ranges. *Journal of African Earth Sciences* **113**:153–164.

Viglietti PA, Smith RMH, Rubidge BS. 2018. Changing palaeoenvironments and tetrapod populations in the *Daptocephalus* Assemblage Zone (Karoo Basin, South Africa) indicate early onset of the Permo-Triassic mass extinction. *Journal of African Earth Sciences* **138**:102–111.

Weide DM, Sidor CA, Angielczyk KD, Smith RMH. 2009. A new record of *Procynosuchus delaharpeae* (Therapsida: Cynodontia) from the Upper Permian Usili Formation, Tanzania. *Palaeontologia africana* **44**:21–26.

Whiteside JH, Grogan DS, Olsen PE, Kent DV. 2011. Climatically driven biogeographic provinces of Late Triassic tropical Pangea. *Proceedings of the National Academy of Sciences of the United States of America* **108**:8972–8977.

Wild R, Kaaya CZ, Kreuser T, Markwort S, Semkiwa PZ. 1993. Discovery of a skull of *Dicynodon lacerticeps* in the Uppermost Permian (Tatarian) of Tanzania. *Sonderveröffentlichungen, Geologisches Institut der Universität zu Köln* **70**:231–242.

Wopfner H. 2002. Tectonic and climatic events controlling deposition in Tanzanian Karoo basins. *Journal of African Earth Sciences* **34**:167–177.
Figure 1

Main cranial fragments of SAM-PK-10630, holotype of *Dicynodon huenei* (=*Daptocephalus huenei* comb. nov.)

Specimen in (A) dorsal and (b) left semi-lateral views. Note narrow intertemporal bar made up of vertically-oriented postorbital bones. Abbreviations: cp, caniniform process; ib, intertemporal bar; pb, postorbital bar. Scale bar equals 5 cm. Photos: Christian Kammerer.
Figure 2

Comparisons between the two morphotypes of Usili Formation dicynodontoids previously included in “Dicynodon huenei.”

GPIT/RE/7175, referred specimen of Dicynodon angielczyki sp. nov., in (A) dorsal and (C) right lateral views. GPIT/RE/9316, referred specimen of Daptocephalus huenei, in (B) dorsal and (D) left lateral views. Distinguishing features of the two morphotypes labeled on the figure: (1) width of the intertemporal bar and orientation of the postorbitals (relatively broad with horizontal postorbitals in D. angielczyki vs. relatively narrow with vertical postorbitals in D.huenei), (2) thickness of the zygoma below the postorbital bar (squamosal and jugal expanded in height, with distinct lateral bowing visible in dorsal view, in D. angielczyki vs. narrow base with no bowing in D. huenei), (3) angulation of squamosal at junction between zygomatic and quadrate rami (highly acute with relatively narrow flange at posterior end of zygomatic ramus in D. angielczyki vs. less acute with broad flange in D. huenei). Note also the generally taller skull and especially deeper snout in D. huenei, as well as the proportionally broader interorbital region and longer intertemporal region of that species. The features noted above are generally typical of Dicynodon and Daptocephalus at the generic level (Kammerer et al., 2011), with the exception of the squamosal expansion, which is a species level autapomorphy of D. angielczyki. Scale bars equal 5 cm. Photos: Christian Kammerer.
Manuscript to be reviewed

PeerJ reviewing PDF | (2019:05:37306:1:2:NEW 1 Jul 2019)
Figure 3

Palatal comparisons between *Dicynodon* and *Daptocephalus*.

(A) GPIT/RE/7177, referred specimen of *Dicynodon angielczyki* sp. nov. (B) GPIT/RE/9641, referred specimen of *Daptocephalus huenei*. (C) RC 38, referred specimen of *Dicynodon lacerticeps* (holotype of *D. aetorhamphus*). (D) NHMUK PV OR 47047, holotype of *Daptocephalus leoniceps*. (E) BP/1/2784, referred specimen of *Daptocephalus leoniceps*. All specimens in ventral view, with anterior right. GPIT/RE/9641 is badly anteroposteriorly distorted, whereas NHMUK PV OR 47047 and BP/1/2784 are slightly laterally compressed and GPIT/RE/7177 is slightly dorsoventrally compressed (RC 38 is largely undistorted). However, note narrower span between tusks in *Daptocephalus* specimens regardless of style of deformation, and their narrower pre-caniniform region of the premaxilla relative to *Dicynodon*. Note also the relatively narrow, distinctly constricted median pterygoid plate bearing a sharp median ridge (the crista oesophagea) in *Dicynodon*. In *Daptocephalus*, the median pterygoid plate is comparatively broad, less sharply constricted from the anterior and posterior pterygoid rami, and lacks a distinct crista oesophagea. Abbreviations: co, crista oesophagea; mpp, median pterygoid plate; tu, tusk. Scale bars equal 5 cm. Photos: Christian Kammerer.
Figure 4

Comparisons between similarly-distorted skulls of *Dicynodon* and *Daptocephalus*.

RC 38, a referred specimen of *Dicynodon lacerticeps* (holotype of *D. trigonocephalus*), in (A) dorsal and (C) right lateral views. GPIT/RE/9641, a referred specimen of *Daptocephalus huenei*, in (B) dorsal and (D) left lateral views. Specimens scaled to equal anteroposterior length. White vertical bars between (A) and (B) illustrate the least interorbital width of RC 38 (left) and GPIT/RE/9641 (right), showing the greater interorbital width of *Daptocephalus* relative to *Dicynodon* even when their skulls are otherwise (and atypically) similar in shape due to anteroposterior compression. Note also the greater depth of the snout, more vertically-oriented postorbital contributions to the intertemporal bar, and broader, less sharply angled junction between the zygomatic and quadrate rami of the squamosal in *Daptocephalus*. Scale bars (black horizontal) equal 5 cm. Photos: Christian Kammerer.
Figure 5

Occipital and anterior comparisons between *Dicynodon* and *Daptocephalus*.

GPIT/RE/9316, referred specimen of *Daptocephalus huenei*, in (A) occipital and (D) anterior views. MB.R.992, referred specimen of *Daptocephalus leoniceps*, in (B) occipital view. GPIT/RE/7175, referred specimen of *Dicynodon angielczyki* sp. nov., in (C) occipital and (E) anterior views. RC 23, referred specimen of *Dicynodon lacerticeps* (holotype of *D. cadlei*), in (F) occipital view. Note the broader, lower occiput and snout of *Dicynodon* specimens, and (1) the autapomorphic lateral expansions of the zygomatic arch in *D. angielczyki*. Scale bars equal 5 cm. Photos: Christian Kammerer.
Figure 6

Comparisons between *Daptocephalus huenei* comb. nov. and *Daptocephalus leoniceps*.

(A) GPIT/RE/9316, referred specimen of *Daptocephalus huenei*, in left lateral view. (B) UCMP 33431, referred specimen of *Daptocephalus leoniceps*, in right lateral view (mirrored for comparative purposes). (C) MB.R. 992, referred specimen of *D. leoniceps*, in left lateral view.

Major facial bones colored to show arrangement in various specimens and highlight comparatively large size of the lacrimal in *D. huenei*, here interpreted as autapomorphic for the species. Abbreviations: la, lacrimal; mx, maxilla; na, nasal; prf, prefrontal; smx, septomaxilla. Scale bars equal 5 cm. Photos: Christian Kammerer.
Figure 7

Comparisons between *Dicynodon angielczyki* sp. nov. and *Dicynodon lacerticeps*.

GPIT/RE/7177, referred specimen of *Dicynodon angielczyki*, in (A) dorsal view. RC 23, referred specimen of *Dicynodon lacerticeps* (holotype of *D. cadlei*), in (B) dorsal and (D) left lateral views. UMZC T1089, holotype of *D. angielczyki*, in (C) left lateral view. Autapomorphies of *D. angielczyki* labeled on the figure: (1) medially-curved squamosal flange at posterolateral edge of temporal fenestra (extensively restored in plaster in this specimen, but sharp curvature at rear edge of squamosal shows it was present; this morphology is clearly preserved in other *D. angielczyki* specimens, see Fig. 2A, 7A); (2) zygomatic arched bowed laterally (shown in A) and dorsoventrally expanded (shown in C) below and immediately behind the postorbital bar. Scale bars equal 5 cm. Photos: Christian Kammerer.
Figure 8

UMZC T1089, holotype of *Dicynodon angielczyki* sp. nov. in dorsal and ventral views.

(A) photograph and (B) interpretive drawing of the specimen in dorsal view. (C) photograph and (D) interpretive drawing of the specimen in ventral view. Abbreviations: af, adductor fossa of postorbital; apr, anterior palatal ridge; apt, anterior pterygoid ramus; bt, basal tuber; co, crista oesophagae; ect, ectopterygoid; fr, frontal; ic, internal carotid canal; ipv, interpterygoid vacuity; j, jugal; la, lacrimal; lf, labial fossa; lp, lateral palatal foramen; mx, maxilla; mpp, median pterygoid plate; mpr, median palatal ridge; na, nasal; oc, occipital condyle; pa, parietal; pe, periotic; pf, pineal foramen; pl, palatine; pmx, premaxilla; po, postorbital; pp, preparietal; prf, prefrontal; q, quadrate; qj, quadratojugal; qpt, quadrate ramus of pterygoid; sq, squamosal; tu, tusk; v, vomer. Scale bar equals 5 cm.

Photos/drawings: Christian Kammerer.
Figure 9

UMZC T1089, holotype of *Dicynodon angielczyki* sp. nov. in right lateral view.

(A) photograph and (B) interpretive drawing. Abbreviations: af, adductor fossa of squamosal; cp, caniniform process of maxilla; ept, epipterygoid; fr, frontal; j, jugal; la, lacrimal; mx, maxilla; na, nasal; op, orbital plate; pa, parietal; pmx, premaxilla; po, postorbital; prf, prefrontal; q, quadrate; qf, quadratojugal foramen; qj, quadratojugal; smx, septomaxilla; sq, squamosal. Scale bar equals 5 cm. Photo/drawing: Christian Kammerer.
Figure 10

UMZC T1089, holotype of *Dicynodon angielczyki* sp. nov. in anterior and posterior views.

(A) photograph and (B) interpretive drawing of specimen in anterior view. (C) photograph and (D) interpretive drawing of specimen in posterior view. Abbreviations: bt, basal tuber; fm, foramen magnum; fr, frontal; j, jugal; la, lacrimal; mx, maxilla; na, nasal; oc, occipital condyle; par, paroccipital process of periotic; pe, periotic; pmx, premaxilla; po, postorbital; ppa, postparietal; prf, prefrontal; ptf, post-temporal fenestra; q, quadrate; qj, quadratojugal; smx, septomaxilla; sq, squamosal; ta, tabular; tu, tusk; vr, vertical ridge on maxilla. Scale bar equals 5 cm. Photos/drawings: Christian Kammerer.
Figure 11

UMZC T1122–T1123, a fossil block containing the jumbled remains of two gorgonopsians and a specimen of *Dicynodon angielczyki* sp nov.

(A) entire block. (B) close-up of *Dicynodon angielczyki* skull in ventral view, with gorgonopsian scapula removed and additional gorgonopsian material lightened to highlight the dicynodont. Both jaw rami of this dicynodont specimen are preserved: the right ramus is descending into the block and visible in (B), whereas the left ramus has been prepared out and is shown in Figure 12. Abbreviations: ar, articular; co, crista oesophagea; tu, tusk. Scale bars equal 5 cm. Photos: Christian Kammerer.
Figure 12

Left mandibular ramus of UMZC T1123, referred specimen of *Dicynodon angielczyki* sp. nov.

Jaw in (A) left lateral, (B) right medial, and (C) posterior views with (D, E) interpretive drawings. Abbreviations: an, angular; ar, articular; d, dentary; ds, sagittal section through dentary; lds, lateral dentary shelf; mf, mandibular fenestra; pra, prearticular; rla, reflected lamina of angular; sa, surangular; sp, splenial; sps, sagittal section through splenial. Scale bar equals 5 cm. Photos/drawings: Christian Kammerer.
Figure 13

Phylogeny of Bidentalia.

Usili Formation dicynodontoids in bold. Numbers at nodes represent symmetric resampling values >50.
Table 1 (on next page)

Cranial measurements (in cm) of the more complete specimens of *Daptocephalus huenei* and *Dicynodon angielczyki*.

Anterior intertemporal width taken at the junction between the intertemporal and postorbital bars, posterior intertemporal width taken at the junction between the intertemporal bar and occiput, following Kammerer *et al.* (2011).
|                         | *Daptocephalus huenei* | *Dicynodon angielczyki* |
|-------------------------|------------------------|-------------------------|
|                         | GPIT/RE/9316     | GPIT/RE/9317   | GPIT/RE/9641 | GPIT/RE/7175 | GPIT/RE/7177 | UMZC T1089 |
| Dorsal skull length     | 23.6           | ~25          | 23.0        | 27.1        | 21.1        | 27.3        |
| Basal skull length      | 28.2           | ~28          | 29.8        | 29.2        | 22.2        | 31.5        |
| Snout length            | 9.4            | NA           | 5.1         | 11.3        | 6.3         | 10.2        |
| Interorbital width      | 8.1            | 8.5          | 10.2        | 7.8         | 5.5         | 7.9         |
| (minimum)               |                |              |             |             |             |             |
| Anterior intertemporal  | 5.6            | 5.0          | 6.7         | 6.9         | 5.7         | 8.0         |
| width                   |                |              |             |             |             |             |
| Posterior intertemporal | 3.1            | 3.0          | 3.0         | 5.8         | 3.5         | 6.8         |
| width                   |                |              |             |             |             |             |
| Temporal fenestra length| 17.9           | NA           | 17.0        | 16.7        | 13.7        | 19.4        |
| (left maximum)          |                |              |             |             |             |             |
| Temporal fenestra length| NA             | NA           | 16.8        | 17.2        | NA          | 17.3        |
| (right maximum)         |                |              |             |             |             |             |
| Median pterygoid plate   | 3.7            | 4.2          | 4.7         | 3.1         | 2.3         | 2.9         |
| width (minimum)         |                |              |             |             |             |             |