Review of paper: gchron-2020-5:

Resolving the timescales of magmatic and hydrothermal processes associated with porphyry deposit formation using zircon U-Pb petrochronology
Simon J.E. Large, Jörn F. Wotzlaw, Marcel Guillong, Albrecht von Quadt, Christoph A. Heinrich

The paper reports an excellent set of data for zircon from several tonalite phases encompassing the time of formation of a major Cu-Au deposit in Indonesia. The results constrain the period of formation of zircon in the different magma batches, and the time of formation of the mineralization, also providing information of the magma evolution based on trace element data. The analytical results are of very good quality and the interpretation is overall reasonable. The paper is well prepared with just few typos or other problems. I have put a number of questions and suggestions on specific details directly in the file.

Here some comments on aspects of the paper that the authors should consider.

The first concerns the definition of the ‘age of emplacement’. The nice thing with modern ID-TIMS U-Pb data is the high time resolution it achieves, which permits to separate out and date very specific segments of geological processes. This advantage, however, brings new challenges requiring to be more specific on the definitions of the specific parts of the process that are dated. For example, on line 304-305 it states: ‘... zircon dates of 3.736 ± 0.023 Ma, 3.697 ± 0.018 Ma and 3.646 ± 0.022 Ma …[are interpreted] as the time of respective porphyry emplacement …’ which concludes a period of >200 ka of zircon crystallization. This implies that the last zircon in each rock crystallized just as the magma reached its present position. One could wonder why the last zircon couldn’t have crystallized well before the magma reached this final position, or alternatively much later than the emplacement. Some information in merit is provided much later in the Discussion, but clearly, these interpretations and the arguments are quite fundamental in such a paper, and need to be presented before anything else.

A somewhat related problem concerns the question of the validity of the results. I am impressed by the high quality of the data, the superb blanks and the high precision. Nevertheless, a central factor in all data sets is the reproducibility of individual analyses. You are measuring 0.5 pg of Pb, next to nothing, and still achieve a very good precision. But is the precision identical with, or less good than the reproducibility of such measurements? To substantiate the solidity of the work, the authors should present information that backs up the implication at the zircon age of each individual zircon grain is reproducible, the alternative being that the larger spread of the ages in each sample may represent a closer measure of the reproducibility. I have seen some very good such data sets in other papers that support their validity, but the question is central with every new application and needs to be addressed.

My third main point concerns the subsidiary part of the paper, which discussed the comparison of ID-TIMS with ICP and SIMS data and reflects on their applications, correctness, and statistical factors. I find this parts absolutely atrocious, and I highly recommend to cut it out. The ICP analyses of these young zircons achieve intensities of maybe 100 cps for mass 206, for measurements lasting less than a minute, and there is no indication that it even gets to evaluate things like the need to correct for common Pb. So, the results are of very low precision, and it is a wonder that they are even close to the real values. The SIMS data are more substantial, but also they face incredible measurement challenges. So, really, all the arguments on statistics and processes cannot get around these basic limitations. And talking about them to such an extent is like watching children playing in the sand. Boring. Suggest cutting this parts out, keeping them for some contribution to a technical workshop, and not use them to spoil an otherwise interesting paper.

April 2nd, 2020   F. Corfu
Resolving the timescales of magmatic and hydrothermal processes associated with porphyry deposit formation using zircon U-Pb petrochronology

Simon J.E. Large1*,†, Jörn F. Wotzlaw1, Marcel Guillong1, Albrecht von Quadt1, Christoph A. Heinrich1,2

1Department of Earth Sciences, Eidgenössische Technische Hochschule (ETH) Zurich, 8092 Zürich, Switzerland.
2Faculty of Mathematics and Natural Sciences, University of Zurich, 8006 Zürich, Switzerland.
†Current address: Department of Earth Sciences, Natural History Museum, Cromwell Road, London, SW7 5BD, UK

Correspondence to: Simon J.E. Large (s.large@nhm.ac.uk)

Abstract.

Understanding the formation of economically important porphyry-Cu-Au deposits requires the knowledge of the magmatic-to-hydrothermal processes that act within the much larger underlying magmatic system and the timescales on which they occur. We apply high-precision zircon geochronology (CA-ID-TIMS) and spatially resolved zircon geochemistry (LA-ICP-MS) to constrain the magmatic evolution of the magma reservoir at the Pliocene Batu Hijau porphyry-Cu-Au deposit. We then use this extensive dataset to assess the accuracy and precision of different U-Pb dating methods of the same zircon crystals.

Emplacement of the oldest pre- to syn-ore tonalite (3.736 ± 0.023 Ma) and the youngest tonalite porphyry cutting economic Cu-Au mineralisation (3.646 ± 0.022 Ma) is determined by the youngest zircon grain from each sample, which constrains the duration of metal precipitation to less than 90 ± 32 kyr. Overlapping spectra of single zircon crystallisation ages and their trace element distributions from the pre-, syn and post-ore tonalite porphyries reveal protracted zircon crystallisation together with apatite and plagioclase within the same magma reservoir over >300 kyr. The presented petrochronological data constrain a protracted early >200 kyr interval of melt differentiation and cooling within a large heterogeneous magma reservoir leading up to ore formation, followed by magma storage in a highly crystalline state and chemical and thermal stability over several 10s of kyr. Irregular trace element systematics suggest magma recharge or underplating during this final short time interval.

The comparison of high precision CA-ID-TIMS results with in-situ U-Pb geochronology data from the same zircon grains allows a comparison of the applicability of each technique as a tool to constrain dates and rates on different geological timescales. All techniques provide accurate dates with variable precision. Highly precise dates derived by the calculation of the weighted mean and standard error of the mean of zircon dates obtained by in-situ techniques can lead to significantly older suggested emplacement ages than those determined by high-precision CA-ID-TIMS geochronology. This lack in accuracy of the weighted means is due to the protracted nature of zircon crystallisation in upper crustal magma reservoirs, suggesting that standard errors should not be used as a mean to describe the uncertainty in those circumstances. Thus,
geologically rapid events or processes or the tempo of magma evolution are too fast to be reliably resolved by in-situ U-Pb 
geochronology and require ID-TIMS geochronology.

1 Introduction

Zircon geochronology is widely applied to date geological events and constrain timescales of geological processes. Combined with zircon geochemistry it has improved our understanding of crustal magmatic systems, such as those forming economically important magmatic-hydrothermal porphyry Cu-Au deposits. Advances in analytical techniques resulted in a shift from establishing the ages of magma emplacement or crystallisation to resolving the durations of magmatic and associated hydrothermal processes, such as magma accumulation or recharge, fractional crystallisation or hydrothermal ore formation and it has resulted in unprecedented information about the mechanisms and scales of magma ascent and storage in the Earth’s crust (e.g. Vazquez and Reid, 2004; Chamberlain et al., 2014; Barboni et al., 2016; Bucholz et al., 2017).

Porphyry copper deposits provide successively quenched samples of magma extracted from large crustal-scale hydrous magma systems. They are therefore a critical source of information about the processes and rates of magma ascent, magma storage and fluid generation, bridging those of volcanism and pluton formation. The identification of the processes that lead to porphyry deposit formation (e.g. Rohrlach et al., 2005; Audétat et al., 2008; Richards, 2013; Wilkinson, 2013) and the timescales on which they can provide us with valuable information about arc magmatic processes but could also potentially help in discriminating possibly fertile magmatic systems from ubiquitous infertile systems resulting in barren intrusions or volcanic eruptions.

Porphyry Cu-Au deposits commonly display clear field relationships of successive generations of porphyritic stocks or dikes, which were injected into subvolcanic and other upper-crustal rock sequences (Sillitoe, 2010). The injected porphyry magmas thus provide snapshots of the underlying, vertically and laterally extensive, magma reservoirs (e.g. Dilles, 1987; Steinberger et al., 2013). Cross-cutting relationships between veins and intrusive rocks suggest temporal overlap of hydrothermal alteration, ore mineralisation and porphyry emplacement (Proffett, 2003; Seedorff and Einaudi, 2004; Redmond and Einaudi, 2010). Strong hydrothermal alteration of the intrusive rocks associated with ore formation severely disturbs the geochemical information of most minerals and whole-rock compositions. While providing important insights into the hydrothermal history of a deposit (e.g. Roedder, 1971; Dilles and Einaudi, 1992; Landtwing et al., 2005; Cathles and Shannon, 2007; Seedorff et al., 2008; Large et al., 2016) it limits the investigation of the magma evolution, especially for the porphyries that are most intimately associated with ore formation. Zircon is a mineral that is unaffected by nearly all hydrothermal alteration and can thus provide unique information about the evolution of a magmatic system.

Recent advances in high-precision zircon geochronology by chemical abrasion - isotope dilution - thermal ionization mass spectrometry (CA-ID-TIMS: e.g. Mattinson, 2005; Bowring et al., 2011; McLean et al., 2011a; Condon et al., 2015; McLean et al., 2015) now allow dating the porphyritic intrusions associated with ore formation with unprecedented
precision. The dramatically improved precision permits to constrain rapid events, such as individual porphyry emplacement and hydrothermal mineralization phases (<100 kyr: von Quadt et al., 2011; Buret et al., 2016; Tapster et al., 2016) that typically occur at the end of a longer-term period of volcanism and intrusive magma emplacement extending over several million years (e.g. Deino and Keith, 1997; Halter et al., 2004; Maksaev et al., 2004; Rohrlach et al., 2005; Lee et al., 2017).

The integration of the temporal and chemical information gained from zircon is referred to as zircon petrochronology and can yield time-calibrated information about magma chemistry, thermal evolution and crystallinity during zircon crystallisation in magmatic systems (e.g. Schoene et al., 2012; Chelle-Michou et al., 2014; Samperton et al., 2015; Buret et al., 2016; Szymanowski et al., 2017).

Geological events and processes that require highest possible precision to be resolved, essentially rely on the accuracy of the chosen analytical technique. Timescales for magmatic and hydrothermal processes involved in porphyry ore formation have been suggested based on in-situ U-Pb data (e.g. Garwin, 2000; Banik et al., 2017; Lee et al., 2017) and increasingly precise CA-ID-TIMS geochronology (e.g. von Quadt et al., 2011; Chelle-Michou et al., 2014; Buret et al., 2016; Tapster et al., 2016; Gilmer et al., 2017; Large et al., 2018). However, several studies applying multiple techniques on the same sample sets have resulted in differing dates (von Quadt et al., 2011; Chiaradia et al., 2013; Chelle-Michou et al., 2014; Chiaradia et al., 2014; Correa et al., 2016). The discrepancy demands for a more detailed understanding of the precision and accuracy of the techniques and the statistical data treatment that are applied to derive a geological age. This is not only fundamental for resolving dates and rates of geological processes in porphyry research but equally affects magmatic dates and rates obtained by U-Pb geochronology.

For the present paper, we obtained a large dataset of zircon geochemistry and geochronology by laser ablation-inductively coupled plasma-mass spectrometry (LA-ICP-MS) coupled to high-precision geochronology of the same zircon fragments/segments/crystals utilising chemical abrasion-isotope dilution-thermal ionization mass spectrometry (CA-ID-TIMS). The data from the world-class Batu Hijau porphyry Cu-Au deposit allows to resolve the chemical evolution and the changing physical state of the magma reservoir over time as well as the timescales of hydrothermal processes. In addition, previously published data on the same lithologies permit a critical comparison of two in-situ microanalytical methods (SHRIMP data by Garwin (2000), LA-ICPMS presented here) with high-precision U-Pb CA-ID-TIMS geochronology (this study). This allows us to critically compare the effects of variable degrees of precision and of the statistical treatment of data on the resulting interpreted ages and it provides a mean to test the accuracy of the different techniques.

2 Geological Background

The Pliocene, island-arc hosted world-class porphyry deposit of Batu Hijau is located on Sumbawa island, Indonesia (Fig. 1), and it is one of the largest Cu and Au resources in the Southwest Pacific region (7.23 Mt Cu and 572 t Au: Cooke et al.,
It is currently the only mined porphyry deposit in the Banda-Sunda volcanic arc, where Cu-Au porphyries are restricted to a narrow segment of the eastern Sunda-Banda arc from 115°E and 120°E (Fig. 1), where Australian plate is being subducted since the Eocene (Hall, 2002).

The exposed islands of the Sunda-Banda arc are characterised by Late Oligocene to Early Miocene calc-alkaline basaltic to andesitic arc rocks that are overlain or intruded by a Late Miocene to Pleistocene calc-alkaline volcanic and plutonic rock suite ranging from basaltic to rhyolitic compositions (Hamilton, 1979; Hutchison, 1989). The magmatic arc hosts a variety of ore deposit types, including porphyry-Cu-Au deposits, high-, intermediate- and low-sulphidation epithermal deposits and a VMS-type deposit on Wetar (Fig. 1).

![Figure 1: Tectonic map of southeast Asia and the southwest Pacific. The Batu Hijau porphyry Cu-Au deposit (enlarged red square) is located on Sumbawa on the subduction related, magmatic Sunda-Banda arc within a small corridor between 110°E and 120°E that hosts several porphyry deposits. Arrows display plate motion relative to the Eurasian plate (Sundaland shield) but do not indicate velocity. Most deposit locations are from Garwin (2005).](https://doi.org/10.5194/gchron-2020-5)

The geology of Sumbawa Island, hosting the Batu Hijau deposit, is dominated by Early Miocene to Holocene volcanic arc successions deposited on oceanic crust that is 14 – 23 km thick (Hamilton, 1979; Barberi et al., 1987). Thickened continental crust observed in most other porphyry-mineralized magmatic arcs and commonly considered a prerequisite for porphyry-Cu formation (Rohrlach et al., 2005; Lee and Tang, 2020) is lacking beneath Sumbawa (Garwin et al., 2005). The distribution of volcano-sedimentary units, intrusions and the current coastline of Sumbawa are controlled by a major arc-transverse, left-lateral oblique-slip fault zone (Arif and Baker, 2004; Garwin et al., 2005). The fault zone strikes...
SSE-NNE about 30 km east of the Batu Hijau deposit coinciding with the north easterly projection of the Roo Rise oceanic plateau (Fig. 1).

Figure 2: Geological map (a) and north to south cross-sections with lithological information and grade contours (b-e) of the open pit at Batu Hijau. The Intermediate and Young tonalite intruded into a Volcanic lithic breccia and the equigranular quartz diorite (a, b). Note that the extent of the Old tonalite is not displayed but is included in the Intermediate tonalite. Dashed line in a) is the N-S section displayed in b)-e). Thin grey lines in a) indicate mine-benches. Cu- and Au-grades (c + d) are enveloped around the tonalites and a deep, central barren core. High-grade Cu- and Au-mineralisation is cut by the Young tonalite. The ratios of Cu to Au (e) illustrate strong Au-enrichment proximal to the Intermediate tonalite and Cu-dominated distal mineralization. Map, section and grades are based on company information from May 2016.

The hypabyssal stocks in the Batu Hijau district are intruded into an Early to Middle Miocene volcano-sedimentary rock sequence (< 21 Ma based on biostratigraphy: Adams, 1984; Berggren et al., 1995) that reaches thicknesses of up to 1500 m in southwestern Sumbawa. The low K$_2$O, calc-alkaline, sub-volcanic intrusive rocks in the Batu Hijau district have andesitic to quartz-dioritic and tonalitic compositions (Foden and Varne, 1980; Garwin, 2000) and were emplaced in several pulses during the Late Miocene and Pliocene (Garwin, 2000). Over this multi-million year magmatic history, a continuous geochemical evolution towards more fractionated lithologies is indicated by whole-rock chemistry and Fe-isotopic evidence.
of the magmatic rock suite in the Batu Hijau district (Garwin, 2000; Wawryk and Foden, 2017). Within the Batu Hijau deposit, andesite porphyries and different quartz-diorite bodies are the earliest recognized stocks, whereas three tonalite porphyries are the youngest exposed intrusions (Clode, 1999). These tonalite porphyries, which are associated with economically important Cu-Au mineralization and pervasive hydrothermal alteration at Batu Hijau, were emplaced as narrow semi-cylindrical stocks into a broad ENE trending structural dome between ~3.9 – 3.7 Ma (Fig. 2: Garwin, 2000). Based on petrography and crosscutting relationships they were termed Old Tonalite, Intermediate Tonalite and Young Tonalite (Fig. 3: Meldrum et al., 1994; Clode, 1999).

All three tonalite intrusions are petrographically similar and are geochemically described as low-K calc-alkaline tonalites (Idrus et al., 2007). Least altered specimens contain phenocrysts of plagioclase, hornblende, quartz, biotite, magnetite ± ilmenite hosted in an aplitic groundmass of plagioclase and quartz (Fig. 3: Mitchell et al., 1998; Clode, 1999; Garwin, 2000; Idrus et al., 2007). Notably, all three porphyry intrusions lack potassium feldspar. Identified accessory minerals include apatite, zircon and rare titanites. Relicts of clinopyroxene can be identified within the tonalites. Vein density, ore grade and alteration intensity decrease from the Old to Young Tonalite. The Old Tonalite is the volumetrically smallest occurring mostly at the edges of the composite stock. It can clearly be identified in drill-core where its veins are truncated by later intrusions (Fig. 3) but it is currently not separated from the Intermediate Tonalite by the mine geology department at Batu Hijau, because their phenocryst proportion is almost indistinguishable. Thus, it is not displayed as a separate unit in Figure 2 but mapped together with the Intermediate Tonalite. It locally contains the highest ore-grades (>1 % Cu and >1 g/t Au) and its matrix is characteristically coarsest of the three tonalite intrusions. The Intermediate Tonalite is the volumetrically largest of the three porphyry intrusions and strongly mineralized (Fig. 2). The Intermediate Tonalite is porphyritic with phenocrysts, including characteristic euhedral quartz phenocrysts, <8 mm in diameter (Fig. 3b). The Young Tonalite is the youngest intrusive rock in the district cutting most vein generations, ore mineralization and alteration (Fig. 3c, e). It is strongly porphyritic with largest observed phenocrysts, including euhedral quartz phenocrysts, and contains elevated but sub-economic metal grades (<0.3 % Cu and <0.5 g/t Au).
Figure 3: Rock specimens of the different tonalite porphyries and zircon CL images at Batu Hijau. Mineral assemblage in all tonalites is dominated by plagioclase, quartz and biotite. a) Phenocrysts in the slightly propylitically altered, pre- to syn-Cu-Au-mineralisation, equigranular Old tonalite are <3 mm. b) Phenocrysts of the syn-Cu-Au-mineralisation, porphyritic Intermediate tonalite are < 5 mm. c) post-mineralisation porphyritic Young tonalite contains largest phenocrysts <8 mm and is characterized by the higher abundance in ‘quartz eyes’. d) Abundant veins in the equigranular Old tonalite are truncated by the later porphyritic Intermediate tonalite. e) Strongly veined Intermediate tonalite is truncated by the barren and little altered Young tonalite. Dashed-green lines indicate intrusive contacts. f) Representative zircons that display dominant oscillatory zoning and areas with little zoning. Circles indicate domains selected for LA-ICP-MS analyses (30 μm in diameter).

Copper and gold are not distributed uniformly within the deposit. High Au-zones are tightly enveloped around the tonalite stocks whereas high copper grades extend further out into the volcanic lithic breccia and the equigranular quartz diorite (Fig. 2). Lowest Cu/Au ratios occur towards and below the current pit floor and higher Cu/Au ratios are recorded peripheral to the central porphyry stock and towards the upper, already mined part of the ore body. A positive correlation between vein density and Cu and Au contents was described at Batu Hijau (Mitchell et al., 1998; Clode, 1999; Arif and Baker, 2004). A-veins were suggested to comprise ~80 % of all quartz veins and contain a similar fraction of the Cu (Mitchell et al., 1998). Most authors suggested that the bulk of the Cu and Au were precipitated as bornite during early A-vein formation and converted to later chalcopyrite and gold associated with AB and B vein formation (Clode, 1999; Arif and Baker, 2004; Proffett, 2009). Other studies on vein relationships and mineralogy using SEM-CP petrography combined with fluid inclusion analyses suggests that Cu-Au ore mineralization including bornite, chalcopyrite and gold all precipitated with a late quartz generation postdating high-temperature A and AB vein quartz, at lower temperature together with the formation of C-veins (Zwyer, 2011; Schirra et al., 2019). Irrespective of the relative timing of stockwork quartz veins and economic ore mineral deposition in the Old and Intermediate tonalites, the Young Tonalite cuts through all high-grade Cu and Au zones demonstrating its late, post-mineralisation emplacement (Fig. 2, 3e). Therefore, the maximum duration of economic mineralisation is bracketed by the emplacement ages of the Cu-Au-rich Old Tonalite pre-dating it and the Young Tonalite post-dating it.

3 Materials and Methods

Based on detailed core logging and outcrop mapping with company geologists in May 2016, one sample each from the Old Tonalite, Intermediate Tonalite, the Young Tonalite and the equigranular quartz diorite were selected from locations where the lithologies were in unequivocal time relationship (See Supplementary Material for sample locations). Rocks were crushed and zircons separated with conventional techniques, including Selfrag™ disintegration, panning and heavy liquid mineral separation (methylene iodide; 3.3 g/cm³). Selected zircons were annealed for 48 hours at 900°C, mounted in epoxy.
resin and polished to reveal their crystal interior. Polished zircons were carbon coated and imaged using scanning electron microscopy cathodoluminescence (SEM-CL; Tescan Eoscan VEGA XLSeries 4 Scanning Electron Microscope) prior to in situ LA-ICP-MS analysis for trace elements and U-Pb isotopes employing a 193 nm ASI Resolution (S155) ArF excimer laser with a 30 µm spot diameter, 5Hz repetition rate and 2 J cm\(^{-2}\) energy density coupled to an Element SF-ICP-MS. A detailed description of the method including data reduction can be found in Guillong et al. (2014) and the supplementary material, including results on secondary reference materials. Generally, at least one spot was chosen in the interior (core) and one in the exterior (rim) part of the zircon but up to four individual spots were analysed per zircon (Fig. 3f) to obtain in-situ geochemical information and U-Pb dates. All \(^{206}\text{Pb}/^{238}\text{U}\) dates were corrected for initial \(^{230}\text{Th}-^{238}\text{U}\) disequilibrium in the \(^{238}\text{U}-^{206}\text{Pb}\) decay chain (e.g. Schärer, 1984). Ratios of Th/U recorded by zircons cluster around 0.3 – 0.6 and the dates were therefore corrected assuming a constant Th/U\(_{\text{melt}}\) of 2 based on partition coefficients (0.25) by Rubatto and Hermann (2007). Variation of the assumed Th/U\(_{\text{melt}}\) by ±0.5 would result in changes of individual \(^{238}\text{U}-^{206}\text{Pb}\) dates of <10 kyr, far below analytical uncertainty.

Titanium concentrations in zircon have been calibrated as a proxy for the crystallization temperature of zircons (Watson and Harrison, 2005;Watson et al., 2006;Ferry and Watson, 2007) and have been widely used in igneous and ore deposit petrology (e.g. Claiborne et al., 2010b;Reid et al., 2011;Chelle-Michou et al., 2014;Dilles et al., 2015;Buret et al., 2016;Lee et al., 2017). The determination of accurate zircon crystallisation temperatures by Ti-in-zircon thermometry (Ferry and Watson, 2007) requires reliable estimates for the activity of SiO\(_2\) and TiO\(_2\) (aSiO\(_2\) and aTiO\(_2\)) during zircon crystallization. Based on previous studies on porphyry deposits we utilize an aSiO\(_2\) of 1 and an aTiO\(_2\) of 0.7 (Chelle-Michou et al., 2014;Buret et al., 2016;Tapster et al., 2016;Lee et al., 2017;Large et al., 2018) reflecting quartz and titanite saturation (Claiborne et al., 2006;Ferry and Watson, 2007). Titanite saturation during zircon crystallization is ambiguous at Batu Hijau (see discussion) but changes in the assumed aTiO\(_2\) result in systematic changes of all zircon crystallization temperatures and will therefore not affect the interpretation of relative temperature changes: a change of the aTiO\(_2\) by ±0.2 would result in a variation of about ±30°C.

Imaging by CL and low-precision but spatially resolved LA-ICP-MS U-Pb dates and geochemical data were used to evaluate inherited zircon populations and to select inheritance-free zircons for subsequent dissolution and analysis by high-precision U-Pb geochronology by CA-ID-TIMS. Selected crystals were removed from the epoxy mount chemically abraded (CA) for 12-15 hours at 180°C using techniques modified from Mattinson (2005). Zircons were spiked with 6-8 µg of the EARTHTIME \(^{202}\text{Pb}-^{205}\text{Pb}-^{233}\text{U}-^{235}\text{U}\) tracer solution (ET2535; Condon et al., 2015;McLean et al., 2015) and dissolved in high-pressure Parr bombs at 210°C for >60 hours. Dissolved samples were dried down and redissolved in 6N HCl at 180°C for 12 hours. Sample dissolution, ion exchange chromatography modified from Krogh (1973) and loading onto zone-refined Re filaments were conducted at ETH Zürich and are described in detail by Large et al. (2018). High-precision U-Pb isotopic data were obtained employing thermal ionization mass spectrometry at ETH Zürich (Thermo Scientific TRITON Plus). Pb was measured sequentially on a dynamic MassCom secondary electron multiplier and U was measured in static mode as U-oxide using Faraday cups fitted with 10\(^{13}\) Ω resistor amplifiers (von Quadt et al., 2016;Wotzlaw et al., 2017). Data reduction
and age calculation were performed using the algorithms and software described in McLean et al. (2011) and Bowring et al. (2011). All \(^{206}\text{Pb}/^{238}\text{U}\) dates were corrected for initial \(^{230}\text{Th}/^{238}\text{U}\) disequilibrium in the \(^{238}\text{U}/^{206}\text{Pb}\) decay chain (e.g. Schärer, 1984) using a constant \(\text{Th}/\text{U}\) partition coefficient ratio of 0.25 (Rubatto and Hermann, 2007) assuming that variations in \(\text{Th}/\text{U}\) of the zircons result from different \(\text{Th}/\text{U}\) of the crystallising melt and not from variations in relative zircon-melt partitioning of Th and U. High-precision U-Pb dates were obtained from 45 zircons, all of which were previously analysed by LA-ICPMS.

4 Results

4.1 Optical zircon appearance and SEM-CL petrography

Zircons were extracted from the three tonalites (Old, Intermediate and Young Tonalite) and the equigranular quartz diorite. Zircon crystals from all three tonalite samples are colourless, euhedral to subhedral and variable in size with c-axis lengths of 100 – 500 µm and aspect ratios between 1:2 and 1:4 (Fig. 3f). Thin section observations reveal zircons that are enclosed by phenocrysts and those that occur within the fine-grained groundmass suggesting protracted zircon crystallisation within the magma until emplacement of the tonalite porphyries. Investigation of mineral separates and mounts with a binocular microscope reveals that many zircons contain small (<<20 µm) mineral or melt inclusions. SEM-CL imaging reveals few unzoned and sector zoned zircon domains, but most zircons exhibit oscillatory zoning (Fig. 3f).

Only few broken zircon fragments could be identified from heavy mineral separates of the equigranular quartz diorite, but these indicate originally euhedral to subhedral shapes. Five grains, typically <200 µm long with aspect ratios of ~1:4, could be identified and were mounted. Four zircons were unzoned and one was oscillatory zoned.

4.2 Spatially resolved zircon trace element composition

At Batu Hijau, zircon geochemical analyses from the three tonalites display largely overlapping arrays and ranges for all described trace element concentrations and ratios (Fig. 4). Most zircons from the tonalites display systematically higher HREE (e.g. Yb) over MREE (e.g. Dy) and LREE (e.g. Nd) contents in their rims relative to their cores (Fig. 4a). This strongly correlates with core-rim systematics of other differentiation proxies, like increasing Hf or decreasing Th/U (Fig. 4: Hoskin and Ireland, 2000; Claiborne et al., 2006; Schaltegger et al., 2009; Claiborne et al., 2010b; Samperton et al., 2015). However, some core-rim trends, especially, from the Young Tonalite display increasing Th/U and decreasing Yb/Dy ratios (Fig. 4b).
Figure 4: Covariation diagrams (a-e) and probability density plots (f-h) of in-situ geochemical data obtained by LA-ICP-MS. a) - c) are plotted against Th/U as an indicator for fractionation, whereas d) + e) are plotted against Hf as the fractionation proxy. Arrows labelled ‘fractional crystallisation’ indicate the approximate predicted direction zircon geochemistry would migrate given fractional crystallization of zircon±apatite ± titanite ± amphibole. Arrows labelled ‘plag’ and ‘tit’ points into the predicted direction of zircon geochemistry evolution during co-crystallisation with plagioclase or titanite. Zircons from the three tonalite porphyries are considered to have crystallised from the same magma reservoir, whereas zircons from the equigranular quartz diorite (purple) are unrelated (see text for discussion). Temperature lines in d) are calculated with an assumed aSiO2 = 1 and aTiO2=0.7 based on Ferry and Watson (2007: see text for discussion). Cross in top right corners illustrates average analytical 2σ uncertainties. Probability density plots (after Vermeesch et al., 2013) illustrate differences between different samples and core and rim analyses within each sample. Axes of probability density plots in f) + g) are aligned with axes of d) + e).

In most zircons, Ti-concentrations decrease from core to rim (Fig. 4d, f). This decrease correlates well with increasing Hf and decreasing Th/U. Maximum and minimum values for all intrusions are ~10 ppm and ~2 ppm resulting in model crystallization temperatures of 770°C to 650°C (see methods for details). The majority of zircons from the Batu Hijau...
deposit contain lower U concentrations (<75 ppm) compared to zircons from most other porphyry deposits (several 100 ppm) but individual zircons can contain up to 300 ppm (Fig. 4c). The zircons with high U-concentration do not correspond to the lower Th/U zircons but also contain high Th-concentrations and cover the whole spectra of Th/U ratios observed at Batu Hijau (Fig. 4c). The Eu-anomaly (Eu/Eu*, which is a mean to quantify the negative inflexure of the normalised REE diagram) increases (Eu/Eu* decrease) with increasing Hf concentrations (Fig. 4e). Zircon analyses from the equigranular quartz diorite plot towards the lowest Hf, Yb/Dy, Yb/Nd, Eu/Eu* highest Th/U and Ti end of the trends displayed by the tonalite zircons (Fig. 4).

Probability density functions (Vermeesch, 2012) are used to test for statistically significant differences between the overlapping zircon populations of the different tonalites and between core and rim analyses from the same tonalite porphyries (Fig. 4f, g, h). The Hf and Ti concentrations as well as the europium anomaly of zircons display overlapping distributions for the Intermediate and Young tonalites. The Old Tonalite zircon population peaks at higher Ti concentrations and Eu/Eu* as well as lower Hf concentrations than the younger tonalites. Core and rim analyses from zircons of the Old Tonalite document decreasing Ti and Eu/Eu* together with increasing Hf concentrations from cores to rims. Hafnium contents of the rim analyses peak at higher concentrations than the core analyses within the Intermediate and Young Tonalite with the Eu/Eu* displaying the opposite effect. Populations illustrating titanium concentrations of the two younger tonalites however, display no systematic changes between core and rim.

4.3 CA-ID-TIMS geochronology

We dated 16 zircons each of the Old and Intermediate Tonalite and 13 zircons of the Young Tonalite by high-precision CA-ID-TIMS geochronology. The youngest zircons of the Old, Intermediate and Young Tonalite yield $^{230}$Th/$^{238}$U disequilibrium corrected $^{206}$Pb/$^{238}$U zircon dates of 3.736 ± 0.023 Ma, 3.697 ± 0.018 Ma and 3.646 ± 0.022 Ma (Fig. 5). We interpret these dates as the time of respective porphyry emplacement (c.f. Oberli et al., 2004; von Quadt et al., 2011; Samperton et al., 2015; Large et al., 2018)) consistent with field observations (Fig. 3). The time intervals between emplacement of the Old and Intermediate Tonalite and between the Intermediate and Young Tonalite can therefore be constrained to 39 ± 29 ka and 51 ± 28 ka, respectively. Recorded duration of zircon crystallization, as defined by the oldest and youngest zircon of each sample, spreads over 246 ± 28 kyr, 212 ± 32 kyr and 171 ± 26 kyr for the Old, Intermediate and Young Tonalite (Fig. 5). The overall duration of recorded zircon crystallization is 336 ± 27 ka. Using the youngest zircon population, rather than the youngest individual zircon as the best approximation for porphyry emplacement (c.f. Samperton et al., 2015; Buret et al., 2016; Tapster et al., 2016) would result in slightly older emplacement ages (~20 kyr) but nearly identical durations of zircon crystallisation and time intervals between porphyry emplacement events (see Supplementary Material). Our high-precision CA-ID-TIMS dates precisely constrain protracted zircon crystallization over several 100 ka and successive emplacement of the three porphyritic tonalite bodies at Batu Hijau within 90 ± 32 ka.
Ratios of Th/U obtained by CA-ID-TIMS analyses on the same sample volume illustrate no systematic variation with time. Values vary inconsistently between 0.4 and 0.6 over the whole recorded time interval (Fig. 6).

Figure 5: High-precision U-Pb CA-ID-TIMS zircon dates from the three tonalite porphyries and comparison with weighted mean averages by in-situ U-Pb geochronology. Vertical bars are individual analyses including analytical uncertainty (2σ). The youngest crystallization age is used as the best approximation for porphyry emplacement. The extended range in zircon crystallization ages in each sample indicates protracted crystallization. Yellow box indicates maximum duration of ore formation as constrained by the emplacement age of the Old tonalite and the Young tonalite. Grey box illustrates the duration of zircon crystallization recorded by CA-ID-TIMS geochronology. Vertical bars in the grey box are emplacement ages of the tonalites, demonstrating >200 ka of zircon crystallization before emplacement of the first porphyry intrusion and start of Cu-Au mineralisation.

4.4 In-situ U-Pb geochronology

Trace element and U-Pb isotopic data were obtained for each LA-ICP-MS spot (Fig. 7) prior to CA-IC-TIMS dating. Low Uranium concentrations and the young ages of the analysed zircons resulted in high individual uncertainties for individual in-situ U-Pb dates (Mean: 10%; Minimum: 3%; Maximum: 41%). All individual spot analyses of the three tonalites that were not discarded due to common Pb or strong discordance yield Pliocene dates (2.98 ± 1.06 – 4.95 ± 0.54 Ma: Fig. 7) with no apparently inherited zircons. All in-situ dates of individual samples illustrate continuous arrays and do not indicate more than one population of zircons per sample (Fig. 7). Weighted means of all zircon analyses from each tonalite are 3.879 ± 0.027, 0.065, 0.32 (n = 207, MSWD = 2.1), 3.778 ± 0.023, 0.061, 0.62 (n = 189, MSWD = 2.5) and 3.751 ± 0.023, 0.060,
0.29 Ma (n = 158, MSWD = 2.6) from oldest to youngest (Fig. 7), where the stated uncertainties are the standard error of the weighted mean, the standard error including an external uncertainty of 1.5 % as suggested by Horstwood et al. (2016) to incorporate of excess variance, and the standard deviation of zircons dates from each sample. These weighted averages are not overlapping within uncertainty with the emplacement ages constrained by CA-ID-TIMS but overlap with the mean of the respective population. The few in-situ analyses (n = 8) on zircons from the diorite result in overlapping Late Miocene dates. The weighted mean of all LA-ICPMS analyses of the equigranular diorite results in an apparent age of the equigranular diorite of 6.37 ± 0.40, 0.41, 0.37 Ma (n = 8, MSWD = 0.46).

Figure 6: Th/U ratios plotted against time. Both values obtained from CA-ID-TIMS analyses of the same sample volume.

Garwin (2000) presented the first SHRIMP U-Pb data on zircons from the Batu Hijau tonalites. Similar to the LA-ICP-MS analyses in this study, individual uncertainties of the dates were elevated (0.12 – 0.30 Ma: ~5 – 10 % uncertainty) due to low U concentrations and the young zircon crystallisation ages. As all dates of each sample appear to represent the same populations (Supplementary Material) weighted means were calculated from of all zircons of a sample, these were interpreted as the intrusion ages of the tonalites by Garwin (2000). The reported zircon dates were not corrected for 230Th-
$^{238}$U disequilibrium. For comparability we will only consider zircon dates that are corrected for initial Th/U disequilibrium (Schärer, 1984: for details consult the Supplementary Material). Correction increases individual zircon dates by ~60 – 100 kyr and recalculation of the weighted means averages and standard errors results in dates of 3.74 ± 0.14 Ma (MSWD = 1.2, n = 8), 3.843 ± 0.094 (MSWD = 1.2, n = 18) Ma and 3.81 ± 0.2 Ma (MSWD = 2.35, n = 7) for the Old, Intermediate and Young tonalite, respectively.

Figure 7: In-situ U-Pb geochronology by LA-ICP-MS of zircons from the equigranular quartz diorite a), the Old tonalite b), the Intermediate tonalite c) and the Young tonalite d). Vertical lines illustrate individual U-Pb dates including analytical uncertainty ($2\sigma$). As no zircon populations can be separated the weighted mean average of all analyses is calculated. Standard error (2SE, lightest grey), standard error + 1.5% (light grey) to incorporate of excess variance (Horstwood et al., 2016) and standard deviation of all individual dates (dark grey) are calculated and plotted. Duration of zircon crystallisation as obtained by CA-ID-TIMS is illustrated as a box for comparison. Lower
boundary of the coloured box indicates youngest CA-ID-TIMS date – or the emplacement age. Note the different vertical scale in a). Zircon dates from the equigranular quartz diorite are ~2 Ma older than zircons from the tonalites.
5 Discussion

5.1 Timing and duration of magmatic and hydrothermal processes leading to porphyry Cu formation

The three tonalite intrusions each record protracted zircon crystallisation over ~200 kyr, as resolved by high-precision ID-TIMS geochronology. The older zircon dates from the Young and Intermediate tonalites overlap with the younger zircons of the older intrusion/s (Fig. 5). This overlap together with the consistent trace element systematics of the three samples (Fig. 4) strongly suggests crystallisation of all zircons within the same magma reservoir. High-precision geochronology records a total duration of zircon crystallisation of 336 ± 27 kyr, which is also a minimum estimate for the lifetime of the deeper reservoir underlying Batu Hijau. The first exposed and highly mineralised tonalite intrusion (Old Tonalite) was injected into the upper crust 246 ± 28 kyr after the onset of zircon crystallization. Emplacement of the three tonalites occurred within 90 ± 32 kyr. Emplacement of the Old Tonalite was followed by the emplacement of the Intermediate Tonalite after 39 ± 29 kyr and the Young Tonalite was emplaced after a further 51 ± 28 kyr.

The maximum duration of ore formation is defined by the timespan between the emplacement of the pre- to syn-mineralisation Old Tonalite and the post-mineralisation Young Tonalite (Fig. 3d, e) and can be therefore constrained to less than 122 kyr. This maximum duration is in good agreement with previous geochronological studies indicating timescales of ore formation from <100 kyr to <29 kyr (Fig.8, 9: von Quadt et al., 2011; Buret et al., 2016; Tapster et al., 2016). It is also coherent with results from thermal modelling studies (Cathles, 1977; Weis et al., 2012) and modelling of diffusive fluid-rock equilibration (Cathles and Shannon, 2007; Mercer et al., 2015; Cernuschi et al., 2018) suggesting timescales of ore formation between a few ka and 100 kyr. Strongly elevated Cu- and Au-grades in the Old Tonalite and somewhat lower, but still economic, grades within the Intermediate Tonalite (Clode, 1999; Garwin, 2000; Arif and Baker, 2004) together with cross-cutting relationships (Fig. 3d, e) indicate that mineralisation occurred within at least two but possibly more pulses: (i) one strong mineralisation pulse associated with or slightly postdating the emplacement of the Old Tonalite but predating the injection of the Intermediate Tonalite (Fig. 3d); (ii) a second pulse is bracketed by the intrusion of the Intermediate and the Young Tonalite (Fig. 3e).

More than one episode of mineralisation is also inferred based on detailed mineralogy and vein petrography (see Geology section: Arif and Baker, 2004; Zwyer, 2011). This further strengthens the hypothesis that individual ore-forming hydrothermal pulses are relatively short events, possibly on the millennial or sub-millennial scale (Cathles, 1977; Weis et al., 2012; Mercer et al., 2015), but that the formation of large economic Cu-Au deposits occurs in several pulses occurring over a few 10s of kyr but ≤100 kyr (von Quadt et al., 2011; Weis et al., 2012; Cernuschi et al., 2018).
5.2 Reconstructing the chemical and physical evolution of a porphyry-forming magma reservoir

Trace element systematics of zircons are powerful geochemical proxies, if applied correctly, as they record the magma evolution and characterise the magmatic system, that they crystallised from (e.g. Hoskin and Schaltegger, 2003; Reid et al., 2011; Schoene et al., 2012; Wotzlaw et al., 2012; Chamberlain et al., 2014; Samperton et al., 2015). The largely overlapping trace element systematics recorded by zircons together with the protracted nature of zircon crystallisation are here used to infer zircon crystallisation within the same mid- to upper crustal magma reservoir that sourced magmas forming the three tonalitic porphyry stocks but most likely also volatiles and metals to form the porphyry Cu-Au deposit. At Batu Hijau we are able to reconstruct the magmatic evolution over 236 ± 27 kyr of recorded zircon crystallization.

Th/U ratios and Hf concentrations are commonly used as proxies for the degree of crystal fractionation within a magma reservoir (e.g. Claiborne et al., 2006; Claiborne et al., 2010b; Samperton et al., 2015). The systematically decreasing Th/U ratios and increasing Hf concentrations between samples and from cores to rims (Fig. 4) are indicative of progressive melt differentiation during zircon crystallisation. The good correlation of these melt evolution proxies with decreasing Ti-contents (Fig. 4) further suggests progressive cooling during differentiation. Ratios of HREE over MREE or LREE (e.g. Yb/Dy, Yb/Nd) can be utilised to make inferences about the co-crystallising mineral assemblage. Titanite for example preferentially depletes the melt in MREE resulting in distinct trace element patterns recorded by co-crystallizing zircon (e.g. Reid et al., 2011; Wotzlaw et al., 2013; Samperton et al., 2015; Loader et al., 2017). The systematically higher HREE (e.g. Yb) over MREE (e.g. Dy) and LREE (e.g. Nd) contents in the rims of most zircons relative to their cores (Fig. 4a) thus indicate zircon crystallisation from a fractionally crystallising magma with co-crystallisation of minerals that preferentially incorporate MREE and LREE (e.g. apatite, titanite, amphibole). At Batu Hijau apatites were petrographically identified, whereas magmatic titanite occurs very subordinately. The apparent lack of magmatic titanite is unusual as it is reported as a common accessory phase in many other porphyry-Cu deposits (e.g. Bajo de la Alumbrera, El Salvador, Ok Tedi, Oyu Tolgoi). The absence of euhedral titanite within the mineral separate could be the result of dissolution during intense hydrothermal alteration (van Dongen et al., 2010). The decrease of Eu/Eu* correlating with proxies of increased fractionation (Hf, Fig. 4e) and during zircon growth (Fig. 4f) suggests co-crystallisation of plagioclase and could indicate a lack of titanite crystallisation, or subordinate crystallisation, as already minor titanite crystallisation strongly increases the Eu/Eu* recorded by zircon (Loader et al., 2017). This apparent lack of titanite crystallisation identify apatite as the main REE fractionating mineral during zircon crystallization.

Trace element compositions of zircons from the equigranular quartz diorite suggest crystallisation within a hotter and less evolved magma than the zircons from the tonalites (Fig. 4). In principle, this might indicate that all zircons analysed in this study have crystallised from the same reservoir. However, the >2 Ma time gap is longer than the thermal lifetime of any recognised upper-crustal magmatic body (e.g. Schoene et al., 2012; Wotzlaw et al., 2013; Caricchi et al., 2014; Samperton et al., 2015; Eddy et al., 2016; Karakas et al., 2017) and longer than considered possible based on thermal modelling (Jaeger,
We therefore consider the zircons within the equigranular Diorite to be part of a separate crustal magmatic system not directly related to the ore-forming system that sourced the three tonalitic intrusions. Trace element populations of zircons from the three tonalites demonstrate that the crystallising magma at the time of emplacement of the Old Tonalite was hotter and less fractionated (Fig. 4) than at the time of emplacement of the younger Intermediate and Young Tonalite (i.e. $39 \pm 29$ ka and $90 \pm 32$ ka after emplacement of the Old Tonalite, respectively). The good correlation of proxies indicating progressive differentiation (Th/U and Hf) with decreasing Ti concentrations (Fig. 4d) indicates that the magma reservoir cooled during concurrent crystallisation and melt evolution. In-situ analyses of cores and rims are evidence for an evolving magma reservoir over the course of individual zircon crystallisation (decreasing Hf: Fig. 4h). Core-rim systematics of zircons from the Old Tonalite further demonstrate cooling during protracted zircon growth (Fig. 4f). Rarely recorded coherent zircon trace element systematics recording melt differentiation over time are commonly inferred to result from zircon crystallisation within a homogeneous magma that best resembles near-closed-system behaviour (e.g. Wotzlaw et al., 2013;Large et al., 2018). The lack of such systematic temporal changes in the chemistry of the zircons (Fig. 6) indicates that the magma reservoir at Batu Hijau was not evolving homogenously. This could be explained by incremental recharge or assembly of the magma reservoir. However, this would imply at least partial resetting of the intra-grain systematics recorded in zircons from the Old Tonalite (Buret et al., 2016;Large et al., 2018). To explain the intra-grain and inter-sample systematics but absence of temporal trends (Fig. 4, 6), we favour different degrees of crystallinity in the magma reservoir. Overall the reservoir is generally hotter and less evolved at the time of emplacement of the Old Tonalite than thereafter (Fig. 4). We therefore suggest that the magma reservoir underlying Batu Hijau progressively but heterogeneously cooled and crystallised over at least $246 \pm 28$ ka with potential incremental recharges until emplacement of the Old Tonalite.

A change from a differentiating, crystallising and cooling magma reservoir to a state of chemical and thermal stability is recorded between emplacement of the Old and Young Tonalite (separated by $90 \pm 32$ kyr) as demonstrated by the trace element systematics of the Intermediate and Young Tonalite porphyries. The indistinguishable highly fractionated and low temperature zircon characteristics (Fig. 4) indicate that the magma reservoir remained in near steady-state conditions between emplacement of the Old and Young Tonalite as coherent intra-grain systematics are not pronounced (Hf) or absent (Ti) in zircons from the younger tonalites (Fig. 4f, h).

Irregular zircon trace element systematics in other intrusive magmatic settings have been associated with crystallisation in non-homogenised and small melt batches sometimes with contemporaneous incremental magma addition to the mushy magma reservoir (e.g. Schoene et al., 2012;Buret et al., 2016;Tapster et al., 2016). Geochemically similar zircon chemistries of the Intermediate and Young Tonalite could also result from chemical stability as the magma reservoir reached the ‘petrological trap’ at a crystallinity of $\sim 55 – 65\%$ (Caricchi and Blundy, 2015) where the crystal fraction does not change over a broad temperature interval. Rim analyses that plot outside the mineral co-crystallisation trends than the respective core analyses (Fig. 4) could suggest late-stage crystallisation within a nearly solidified magma that can be characterized by unsystematically variable trace element systematics (Buret et al., 2016;Lee et al., 2017). Alternatively, they could indicate
thermal and possibly chemical rejuvenation of the magma (Buret et al., 2016). The latter would help explaining the recorded thermal stability over several 10s of kyr. It is not possible to unambiguously identify one of the two mechanisms as dominant and a concurrence of both is feasible. We therefore propose that in between emplacement of the Old and Young Tonalite the underlying magma reservoir was in a thermally and chemically stable and crystal-rich state and was most likely affected by incremental magma recharge or underplating.

Our data of a porphyry-Cu fertile magmatic system constrain a heterogeneous magma reservoir that was initially dominated by cooling and melt differentiation and evolved into a thermally and chemically stable, crystal-rich magma that possibly experienced incremental recharge. The likely transitional change of reservoir behaviour can be temporally constrained to have occurred between emplacement of the Old and Young tonalites and coincides with the formation of a world-class Cu-Au reserve. This suggests that porphyry Cu-Au deposits form after a few 100 ka of cooling and crystallisation, potentially within an originally melt-rich magma reservoir.

5.3 Different timescales of processes related to porphyry Cu-Au ore-formation

To date no clear relationship between the duration of magmatic-hydrothermal activity and the size of porphyry deposits can be identified from studies applying high-precision CA-ID-TIMS geochronology. Comparison of published datasets (Buret et al., 2016;Tapster et al., 2016;Large et al., 2018) reveals maximum durations of metal forming events between a few $10^4$ to $10^5$ yr (Fig. 9). Although these studies are so far constrained to deposits of <10 Mt contained Cu they range over at least one order of magnitude in size (Koloula vs. Batu Hijau). A correlation between the duration of the mineralizing/magmatic event and the total mass of deposited copper had been previously suggested based on compilations of different geochronological data-sets (Chelle-Michou et al., 2017;Chiaradia and Caricchi, 2017;Chelle-Michou and Schaltegger, 2018;Chiaradia, 2020). High durations of ore formation (>1 Ma) were suggested based on Re-Os geochronology on Mo-pybdenite at the giant porphyry deposits and deposit clusters in Chile (>50 Gt Cu: El Teniente, Cannell et al. (2005) and Maksaev et al. (2004); Rio Blanco, Deckart et al. (2012); and Chuquicamata, Barra et al. (2013)). Copper (-gold) mineralising timescales were calculated by subtracting the youngest from the oldest Re-Os date. However, recent Re-Os dates from El Teniente (Spencer et al., 2015) indicate that the spread in dates is more consistent with several short (≤200 kyr) hydrothermal events separated by hiatuses of ~500 kyr. Thus, the large tonnage of these deposits could be the result of the superimposition of several ore forming mid- to upper crustal magmatic systems. As the correlation of deposit size and timescales of shallow magmatic-hydrothermal systems is currently ambiguous we would argue that other variables could be the dominant factors controlling the deposit size, such as magma reservoir size, magma or fluid chemistry, fluid release and focussing mechanisms or the metal precipitation efficiency.

Zircon crystallisation over ~200 kyr before the onset of porphyry-ore formation recorded at Batu Hijau is consistent with other high-precision geochronological studies on porphyry deposits (Fig. 8, 9: Buret et al., 2016;Tapster et al., 2016;Large et al., 2018). The lack of variation observed in these deposits suggests the necessity of a long-lived and continuously crystallising magma reservoir preceding economic ore formation. The recorded ~200 kyr of protracted zircon
crystallisation could indicate a period of volatile enrichment as a result of fractional crystallisation and cooling of the magma reservoir before porphyry emplacement.

Figure 8: Compilation of high-precision data-sets on several pre-, syn- and post-ore intrusions at magmatic-hydrothermal Cu-Au deposits. Data for Ok Tedi, Bajo de la Alumbrera and Koloula are from Large et al. (2018) Buret et al. (2016) and Tapster et al. (2016), respectively. Coloured vertical bars are individual analyses including analytical uncertainty (2σ). Intrusions are categorised in pre-ore, pre-/syn-ore and post-ore intrusion. Decreasing deposit size from left to right (tonnages from Cooke et al., 2005).

The geochronological data from the Hijau district are further evidence that rapid porphyry emplacement and ore formation (<100 ka) are the product of a longer-term evolution (a few 100 ka) of a large magma reservoir underlying the porphyry deposit that is the main driver of ore formation (von Quadt et al., 2011; Chelle-Michou et al., 2014; Buret et al., 2016; Tapster et al., 2016; Buret et al., 2017; Large et al., 2018). Magma reservoirs capable of forming porphyry deposits are in turn part of a longer-term (several Myr) evolution of lithosphere-scale magma systems (Sasso, 1998; Rohrlach et al., 2005; Longo et al., 2010; Rezeau et al., 2016), which is consistent with the >>2 Myr record of intrusive rocks preceding porphyry emplacement and ore formation recorded in the Batu Hijau district (Garwin, 2000; Wawryk and Foden, 2017).
5.4 Resolving lower crustal magmatic processes from Zircon petrochronology

The lack of inheritance within the zircon record at Batu Hijau suggests that the crustal magmas experienced very minor crustal assimilation. Typically, magmas that are associated with porphyry ore formation contain diverse suites of inherited zircons (e.g. Tapster et al., 2016; Lee et al., 2017; Large et al., 2018), which have been interpreted to represent extended interaction with arc lithologies (Miller et al., 2007). This apparent lack of crustal contamination is consistent with the juvenile isotopic signatures (Pb-Pb, Sm-Nd, Rb-Sr) of intrusions in the Batu Hijau district (Garwin, 2000; Fiorentini and Garwin, 2010). The juvenile and “porphyry-fertile” magmas at Batu Hijau have been explained by asthenospheric mantle upwelling through a tear in the subducting slab that resulted from the collision with the Roo rise (Garwin, 2000; Fiorentini and Garwin, 2010). This would also explain why the only mined porphyry-deposit in the Sunda-Banda arc (Batu Hijau) and the most promising prospects (Elang and Tumpangpitu) are located above the inferred margin of the subducting Roo rise (Fig. 1).

The formation of porphyry Cu(-Au) deposits has been commonly associated with the fractionation of amphibole ± garnet in thickened crust (e.g. Rohrlach et al., 2005; Lee and Tang, 2020) within lower crustal magma reservoirs that are
active over several Myr (Rohrlach et al., 2005). Zircons have been suggested to directly track this extended lower crustal history (Rohrlach et al., 2005). At Batu Hijau no zircon was identified that crystallised resolvably before the main crystallisation period, which we consider to have occurred in the mid- to upper crust (Fig. 5, and discussion above). Unzoned cores surrounded by oscillatory zoned rims (Fig. 3f) could be interpreted to reflect a two-stage crystallisation process however, the depth of these two processes cannot be resolved and they would have occurred within the few 100 kyr of recorded zircon crystallisation (Fig. 5). As most crystals within a mount are not polished exactly to their centre, the unzoned cores could equally likely represent a polishing effect where the surface of one zone appears as an unzoned core. Therefore, it is highly speculative to directly relate zircon textures to a locus or style of zircon crystallisation.

In the case of Batu Hijau, petrochronology data was used to reconstruct the mid- to upper crustal magma evolution but the data can only provide indirect information about the lower crustal processes involved in the formation of the deposit. For example, the overall elevated Eu/Eu* of the investigated zircons (0.4 – 0.7; cf. Loader et al., 2017) could be the result of amphibole fractionation in the lower crust, which would have, relatively, enriched the residual melt Eu compared to the other REE. This would be analogous to elevated whole-rock Sr/Y ratios in exposed rocks being indicative of the lower crust fractionating assemblage (Rohrlach et al., 2005; Chiaradia, 2015). The intra-crystal and intra-sample trends of decreasing Eu/Eu* discussed above describe the evolution within the mid- to upper crustal magma reservoir that was dominated by plagioclase crystallisation and do not reflect any lower crustal process. Zircon can thus directly record the mid- to upper crustal magma evolution but the information about lower crustal processes is limited to potentially identifying the chemistry of melt and magma that was injected from below into the mid- to upper crust, where zircon started crystallising.

5.5 An assessment of the accuracy and precision of CA-ID-TIMS and in-situ U-Pb zircon geochronology

The obtained U-Pb dataset from Batu Hijau, allows a critical comparison of the two zircon U-Pb geochronology techniques (LA-ICP-MS, CA-ID-TIMS) that have different analytical precision and can analyse samples on varying spatial scales. Previous investigation of the same lithologies by SHRIMP (Garwin, 2000) allows further comparison. The spatially resolved and fast in-situ U-Pb geochronology techniques (LA-ICP-MS or SIMS/SHRIMP) allow the investigation of different crystal domains, whereas the much more time-consuming CA-ID-TIMS analysis of zircons or zircon fragments provides the highest analytical precision. The in-situ techniques can discriminate between different zircon populations within single crystals (e.g., inheritance), whereas CA-ID-TIMS geochronology allows for an >10-fold analytical precision for individual grains that is required to resolve rapid geochronological events. To increase precision of the in-situ techniques large numbers of individual dates that are considered to represent the same geological event are commonly used to calculate a weighted mean date and standard error of the mean (Wendt and Carl, 1991). On the other hand, the CA-ID-TIMS community has started to measure only small zircon fragments to increase spatial resolution (e.g. Samperton et al., 2015; Smith et al., 2019). Here, the comparison of the different U-Pb zircon techniques applied to the same rock suite allows an assessment of the accuracy of the techniques and of the effect of statistical treatment on the accuracy and precision of the different techniques.
At Batu Hijau, the youngest individual CA-ID-TIMS U-Pb date of each sample is used as the best approximation for the emplacement age of the respective porphyry. This is based on the assumption that the magma cooled rapidly upon injection into the subvolcanic environment (cf. Schaltegger et al., 2009; von Quadt et al., 2011; Samperton et al., 2015; Large et al., 2018). The resulting porphyry emplacement ages are 3.736 ± 0.023 Ma, 3.697 ± 0.018 Ma, 3.646 ± 0.022 Ma for the Old, Intermediate and Young Tonalite, respectively (Fig. 5). The extended range of concordant zircon dates obtained by CA-ID-TIMS does not allow to distinguish between different stages of zircon crystallisation within each sample (e.g., inherited vs. autocrystic) but the common geochemical trends indicate crystallisation within the same magma reservoir (see above). Thus, the range in zircon dates preceding this emplacement age is interpreted to represent zircon crystallisation within the underlying source magma reservoir over parts or, depending on the timing of onset of zircon saturation, the entirety of its lifetime. The recorded duration of zircon crystallisation is 336 ± 27 kyr.

Similar to the CA-ID-TIMS dates, in-situ analyses by LA-ICP-MS illustrate an extended range of zircon dates that cannot be separated into different stages of zircon crystallisation. However, the span in zircon dates is about a magnitude higher for the in-situ analyses (1.41 ± 0.5 – 2.1 ± 1.1 Myr) than obtained by CA-ID-TIMS (0.171 ± 0.026 – 0.246 ± 0.028 Myr). This could indicate that the LA-ICP-MS data records an extended period of zircon crystallisation not covered by CA-ID-TIMS data, potentially due to sampling bias, that one data-set is inaccurate or that the span within the in-situ data is the result of analytical scatter.

Sampling bias in the selection of the zircons for CA-ID-TIMS geochronology can be excluded as the analyses were conducted on chemically abraded zircons (Mattinson, 2005) that cover the oldest and youngest dates obtained by LA-ICP-MS (Fig. 10). High accuracy of both, the CA-ID-TIMS and LA-ICP-MS, data-sets are suggested by routine measurements of secondary standards during the LA-ICP-MS analytical run (See Supplementary Material) and regular measurements of zircon standards by CA-ID-TIMS over the period of data acquisition (von Quadt et al., 2016; Wotzlaw et al., 2017). The distributions of the zircon dates of each sample, as illustrated by probability density plots (Fig. 11), illustrate that the peak of the LA-ICP-MS and SHRIMP dates falls within the mean of zircon crystallisation as defined by the CA-ID-TIMS data-set. This suggests that all datasets are accurate but that the in-situ data displays more scatter and lower precision. LA-ICP-MS analyses record younger zircon dates for core analyses than rim analyses in 17 of 49 cases, however the dates are always overlapping within uncertainty. Direct comparison of U-Pb dates from the same zircon crystals by the two techniques (Fig. 10) reveals that suggested dates from the two techniques do not overlap within uncertainty in some cases (6/49 for rim analyses). This could indicate that uncertainties associated with the LA-ICP-MS data have been underestimated in relation to the achieved precision of the technique. However, due to the high number of analyses it is more likely that it is purely an effect of analytical scatter where 5% of the data do not fall within the 95% confidence interval. This is corroborated by ~7% (39/554) of LA-ICP-MS dates not overlapping with the minimum overall duration of zircon crystallisation identified by CA-ID-TIMS dates from all porphyries (336 ± 27 kyr). It is therefore concluded that all three techniques are accurate and represent the ~300 – 350 kyr of zircon crystallisation. The high number of analyses obtained by LA-ICP-MS together with...
the lower precision results in extreme outliers that extend the apparent duration of zircon crystallisation but can be regarded purely as an analytical artefact.

Figure 10: Comparison of in-situ LA-ICP-MS dates and CA-ID-TIMS dates on the same zircons. In upper panel each CA-ID-TIMS date is aligned with the rim (filled) and core (empty) LA-ICP-MS date of the same zircon. Coloured bars indicate individual CA-ID-TIMS analysis including analytical uncertainty (2σ). Downward pointing black arrow indicates that core analyses are older than rim analyses of the respective zircon, whereas red upward pointing arrow indicates the opposite. Note that CA-ID-TIMS dates can be plotted several times, with core and rim analyses of the same zircon. Lower panel compares CA-ID-TIMS date with respective rim analysis.
5.6 Determining geological ages, uncertainties and rates from in-situ U-Pb data

Understanding the timing of magma emplacement, crystallisation or eruption is essential for determining dates and rates of magmatic processes and those directly related or bracketed by them. Where high-precision CA-ID-TIMS data is not available porphyry emplacement ages are commonly inferred by calculating a weighted mean and standard error from the youngest overlapping population of in-situ U-Pb dates (e.g. Correa et al., 2016; Rezeau et al., 2016; Lee et al., 2017). In the case of Batu Hijau the calculation would include all LA-ICP-MS zircon dates for each sample as there is no apparent inheritance within the datasets (Fig. 7). The resulting weighted mean dates for the Old, Intermediate and Young Tonalite are 3.879 ± 0.027/0.064 (MSWD = 2.1, n = 207), 3.783 ± 0.023/0.061 (MSWD = 2.5, n = 189), 3.751 ± 0.023/0.060 (MSWD = 2.6, n = 158) where the first stated uncertainty is the standard error including internal uncertainties and those associated with tracer calibration (Schoene, 2014) and the second includes the added 1.5% external uncertainty as suggested by Horstwood et al. (2016) to account for excess variance. The MSWD for each data-set (2.1 – 2.6) is elevated in respect to the sample size (n=150-200; Wendt and Carl, 1991) suggesting an underestimation of the individual uncertainties or that the data do not represent a normal distribution, e.g. by prolonged zircon crystallisation. However, there is no obvious treatment of the data to obtain more appropriate MSWDs. Under these conditions weighted means and standard errors should not be calculated (Wendt and Carl, 1991) but we will ignore this here, as is commonly done in the scientific literature, and will use these numbers to illustrate a few points below. Analogous, the weighted mean and standard error of all zircons analysed by SHRIMP from each sample results in weighted means of 3.74 ± 0.14 Ma (MSWD = 1.2, n = 8), 3.843 ± 0.094 (MSWD = 1.2, n = 18) Ma and 3.81 ± 0.2 Ma (MSWD = 2.35, n = 7) for the Old, Intermediate and Young tonalite, respectively (Fig. 11). The weighted means of the different tonalites obtained by LA-ICP-MS would be in accordance with cross-cutting relationships, whereas the SHRIMP dates overlap within uncertainty. The calculated standard errors for the LA-ICP-MS dates are significantly smaller than for the SHRIMP data. The decrease in the standard errors is directly correlated with the increasing sample size (Wendt and Carl, 1991; McLean et al., 2011b) implying that a comparably high number of SHRIMP analyses would result in similarly low standard errors. Irrespective of the different standard errors the calculated weighted means by SHRIMP and LA-ICP-MS are overlapping within uncertainty, thus suggesting that both are accurate or similarly inaccurate.

At Batu Hijau, emplacement ages determined by CA-ID-TIMS geochronology are systematically younger than the weighted mean dates calculated from in-situ data (100 – 150 kyr: except CA-ID-TIMS and SHRIMP for the Old Tonalite) and the ages determined by CA-ID-TIMS do not overlap with the LA-ICP-MS values within the attributed uncertainties (Fig. 11). Indeed, disparities between different U-Pb data-sets on the same porphyry samples have been noted in several studies comparing high-precision CA-ID-TIMS data with in-situ data (von Quadt et al., 2011; Chiaradia et al., 2013; Chelle-Michou et al., 2014; Chiaradia et al., 2014; Correa et al., 2016). As discussed above all presented datasets are considered accurate and thus the discrepancy in dates is most likely the result of differences in the statistical handling and geological interpretation.
Figure 11: Probability density plots for the geochronology data for each analytical technique. All dates of each technique are combined in a). Plots in b), c), and d) are constructed from the data of the Old, Intermediate and Young tonalite. Dashed lines indicate the youngest and oldest zircon crystallization age as determined by CA-ID-TIMS for the respective investigated data-set. Weighted means, standard error, standard error + 1.5% (Horstwood et al., 2016) and standard deviation are identical to those in Fig. 5.

The protracted zircon crystallisation identified at Batu Hijau has profound implications for the determination of magma emplacement, crystallisation or eruption ages. Extended magma reservoir lifetimes are not unique to Batu Hijau but a commonly described feature (e.g. Miller et al., 2007; Claiborne et al., 2010a; Reid et al., 2011; Buret et al., 2016). A weighted mean is a measure to quantify the mean of a population whereby emphasising the importance of values with low uncertainties over those with high uncertainties (Reiners et al., 2017) and is only allowed to be used in cases where the data is normally distributed around the expected value (Wendt and Carl, 1991). The presented in-situ data-sets record protracted zircon crystallisation (>300 kyr) in the magma reservoir that results in zircon population distributions that cannot be easily defined statistically (Fig. 5, 7: cf. Keller et al., 2018), negating a normal distribution around a single geological event – the emplacement age. The calculated weighted mean rather represents the mean of the duration of zircon crystallisation, corroborated by the weighted means of the LA-ICP-MS and SHRIMP dates approximately describing the mean of the zircon populations defined by CA-ID-TIMS (Fig. 11). Therefore, the calculated weighted mean does not describe any specific geological event, especially as the uncertainties indicated by the standard error for the LA-ICP-MS data are too small to even cover the entire recorded duration of zircon crystallisation. It has to be noted that more complex settings where xenocrystic zircons overlap within uncertainty of the in-situ techniques with the auto- and antecrystic zircon population (e.g. Chelle-Michou et al., 2014) would result in even less reliable geological dates estimated by in-situ techniques. Therefore, using the weighted mean of a zircon population to determine emplacement or intrusion ages can be a broad oversimplification if dates...
do not represent a normal distribution around the dated event. Similar problems can occur when calculating eruption ages for non-homogeneous zircon populations from tuffs or other volcanic rocks (Schoene, 2014).

Traditionally, problems associated with the oversimplification associated with calculating weighted means and their standard errors in geochronology were hidden by the higher uncertainties resulting from higher analytical uncertainties and smaller sample sizes. The standard error of the mean is a measure for the reproducibility of an experiment (i.e. how likely is it to obtain the same weighted mean if the same amount of zircons from the same sample are analysed again) but one of the main assumptions for using the standard error as the uncertainty of a weighted mean is that the data is normally distributed around the expected value. Due to rapid data acquisition by in-situ techniques calculated standard errors can result in uncertainty envelopes of <0.1% for a sample. In the case of the LA-ICP-MS dates from Batu Hijau the standard error of the weighted mean (~1%; ~40 ka) is on the same order of magnitude as an individual CA-ID-TIMS date and therefore smaller than the geological spread of zircon crystallisation dates, which negates a normal distribution of the zircon data. The combination of using a weighted mean to describe a non-gaussian sample distribution with the very small attributed uncertainties results in highly precise dates that have no relation to a specific geological event.

The MSWD (A reduced chi-square statistic) of a data-set provides a first measure to indicate whether your dates are normally distributed around an expected value and thus whether the calculated weighted mean and standard error are of significance (\(\sqrt{\frac{2}{n-1}}\) rule by Wendt and Carl (1991)). As discussed before, the MSWDs for the LA-ICP-MS data are elevated, suggesting an underestimation of the individual uncertainties or, in this case, that the data do not represent a normal distribution, and thus implying that weighted means and standard errors should not be calculated to characterise a geological event. However, the MSWDs for the SHRIMP zircon analyses of the Intermediate and Old Tonalite are acceptable, mainly due to the higher individual uncertainties. Still, they are similarly affected by protracted zircon crystallisation, which biases the weighted mean to higher values (Fig. 11). Furthermore, overestimated individual uncertainties can result in acceptable MSWDs but similarly inaccurate dates and low standard errors. For example, increasing individual uncertainties for the \(^{206}\text{Pb}\)-\(^{238}\text{U}\) dates obtained for the Old Tonalite by LA-ICP-MS by a factor of 1.5 would result in an acceptable MSWD (0.95) but the weighted mean and standard error would be nearly identically precise but inaccurate (3.880 ± 0.041, 2SE) to those calculated with the actual uncertainties. Based on the presented data it is advised not to characterise a geological event by a weighted mean with an associated standard error if the MSWD is elevated (Wendt and Carl, 1991) and if presented it should be referred to as date and not an age. Even if the MSWD of a dataset is acceptable this is no absolute confirmation that the presented date is accurate within the presented uncertainty, unless there is evidence that the data are uniformly distributed around the dated event or uncertainties are sufficiently high.

An attempt to obtain reliable porphyry emplacement ages from convoluted datasets could be to apply the weighted mean on the youngest or geochemically most evolved population of zircons. Differentiating zircon populations based on geochemical affinity could potentially work in situations where there are clear temporally resolved chemical trends (e.g. Wotzlaw et al., 2013; Samperton et al., 2015; Large et al., 2018). In most scenarios however, there are general geochemical
trends but they are strongly convoluted on a temporal scale (e.g. Schoene et al., 2012; Rivera et al., 2014; Buret et al., 2016). For example, at Batu Hijau early crystallised zircons can have the same chemical signature (e.g. Th/U: Fig. 6) as some of the youngest zircons. In summary, identifying the youngest zircon population (e.g. by its geochemical signature) and applying a weighted mean to it could significantly increase the accuracy of the calculated emplacement age but it requires a detailed understanding of the geochemical evolution of the crystallising magma reservoir.

Based on the presented data we would recommend to use an uncertainty attributed to the weighted mean that is more representative of the uncertainty of the individual analyses, so that it will most likely cover the actually dated event. Here, we tested the standard deviation of zircons dates from each sample as a measure for the uncertainty of the weighted mean. This approach would give a more realistic estimation of the uncertainty associated with calculating a weighted mean of a data-set as it describes the variability in the measurements, (0.29 – 0.62 Ma// Fig. 7, 11) and, importantly, it would be independent of the number of analyses. The resulting values at least for the Pliocene Batu Hijau deposit results in appropriate uncertainties for the weighted mean, as it would cover an appreciable part of the range of in-situ dates and thus the >300 kyr of zircon crystallisation and the emplacement age. Another approach would be the calculation of the dispersion of a data-set (Vermeesch, 2010, 2018) where not all data is treated as part of a single population but where the possibility of data dispersion of the analysed sample set is considered. For the presented LA-ICP-MS data-sets this would result in apparent dispersions of 212 +43/-39 kyr, 229 +43/-39 kyr and 191 +41/-36 kyr for the data of the Old, Intermediate and Young Tonalite, similar to the actually recorded durations of zircon crystallisation. However, this approach requires a precise estimate of the associated individual uncertainties. Similar to calculating the MSWD, over- or under-estimated uncertainties would significantly modify the result.

The presented data highlights the importance of CA-ID-TIMS zircon U-Pb geochronology to resolve complicated zircon crystallisation patterns, which in turn allow to adjust in-situ techniques. While the individual LA-ICP-MS data appear to be accurate (Fig. 11), weighted means and standard errors of high-n datasets (e.g. 3.879 ± 0.039 Ma for Old Tonalite) may provide precise mean zircon crystallisation ages but are likely to be inaccurate in determining emplacement or eruption ages if there is only a minor degree of protracted zircon crystallisation. Taking into account the dispersion of the dataset (e.g. 212 +43/-39 kyr for the Old Tonalite) or using the standard deviation (320 kyr), results in estimates of the emplacement age that overlap with the emplacement age suggested by ID-TIMS and appear to be a more honest way of treating the data. The resulting emplacement ages may not be precise enough to resolve short timescales but can provide a timeline of broader scale magmatic events. In the case of porphyry research high-precision ID-TIMS dates are required to resolve the durations of porphyry emplacement and hydrothermal processes but in-situ data can reliably reconstruct a timeline of magma emplacement events within porphyry districts over by Myr timescales (e.g. Rezeau et al., 2016). More generally, it is understandable that the highest possible precision is strived for from a single dataset. However, it should be refrained from increasing the precision purely by statistical measure without valid assumptions or knowledge that the boundary conditions are met as this can result in hugely precise but inaccurate dates. Furthermore, combination with in-situ petrochronology...
techniques (i.e. U-Pb isotope and geochemical data from the same analyte) allows to screen zircons for inheritance and more importantly provides spatially resolved geochemical information that can be integrated with high-precision dates.

6 Conclusions

High-precision zircon geochronology by CA-ID-TIMS combined with in-situ zircon geochemistry provides valuable datasets that allow the reconstruction of geological processes with the highest temporal resolution. At Batu Hijau zircons record the magmatic to hydrothermal evolution of the world-class Batu Hijau porphyry Cu-Au deposit from the onset of zircon crystallisation to emplacement of the post-ore Young Tonalite. The magma reservoir that sourced the tonalites and the Cu-Au mineralising fluids records zircon crystallisation over 336 ± 27 kyr. Emplacement of the first exposed tonalite at the Batu Hijau deposit (Old Tonalite) occurred after 246 ± 28 kyr of uninterrupted zircon crystallisation in this subjacent reservoir. Zircon trace element signatures support a dominantly crystallising and cooling magma reservoir over 285 ± 24 kyr until emplacement of the Intermediate Tonalite. After emplacement of the Intermediate Tonalite the chemistry of the reservoir remained in rather steady conditions for 51 ± 28 kyr during which it could have been disturbed by magmatic recharge or underplating until final emplacement of the Young Tonalite. Ore formation is most probably associated with the last stages of the chemically and thermally evolving magma reservoir. The maximum duration of ore formation can be constrained to <122 kyr by the emplacement ages of pre-to syn-ore Old Tonalite and the post-ore Young Tonalite. This maximum duration of ore formation covers different pulses of mineralisation that could have lasted only a few kyr. We record a magmatic system that was active over ~250 kyr before emplacement of the first porphyry intrusion and onset of several pulses of hydrothermal activity forming the world-class ore reserve in less than 100 kyr.

Comparison between in-situ LA-ICP-MS and SHRIMP as well as CA-ID-TIMS U-Pb geochronology reveals that all techniques provide accurate individual dates (within the stated confidence interval). However, statistical treatment of in-situ data by calculating a weighted mean and standard error can result in highly precise but inaccurate and therefore geologically meaningless ages. The tempo of magma evolution and hydrothermal processes associated with magmatic-hydrothermal systems, such as porphyry deposits is too fast to be reliably resolved by in-situ U-Pb geochronology and requires ID-TIMS geochronology. Combination of high-precision geochronology with in-situ or TIMS-TEA geochemistry is currently the most powerful tool in deciphering these geologically rapid processes.

7 Data availability

All data used in this manuscript is available from the supplementary files.
8 Supplement link

9 Author contribution

Simon Large conducted the ID-TIMS measurements. Simon Large together with Marcel Guillong conducted the LA-ICP-MS measurements. The study was designed by Simon Large, Christoph Heinrich, Albrecht von Quadt and Jörn-Frederik Wotzlave. Discussion of the data involved all authors and Simon Large wrote the manuscript and drafted the figures with input from all co-authors.

10 Acknowledgement

This work was supported through the Swiss National Science Foundation project 200026-166151. Wotzlave acknowledges funding through the ETH Zurich Postdoctoral Fellowship Program. Extensive logistical support by the Geology Department of the Batu Hijau mine, especially Eddy Priowasono, and the technical staff at the mine site was hugely appreciated.

References

Adams, C. G.: Neogene larger foraminifera, evolutionary and geological events in the context of datum planes, Pacific Neogene datum planes, 47-67, 1984.

Annen, C.: From plutons to magma chambers: Thermal constraints on the accumulation of eruptible silicic magma in the upper crust, Earth and Planetary Science Letters, 284, 409-416, 10.1016/j.epsl.2009.05.006, 2009.

Arif, J., and Baker, T.: Gold paragenesis and chemistry at Batu Hijau, Indonesia: implications for gold-rich porphyry copper deposits, Mineralium Deposita, 39, 523-535, 2004.

Audétat, A., Pettke, T., Heinrich, C. A., and Bodnar, R. J.: Special paper: the composition of magmatic-hydrothermal fluids in barren and mineralized intrusions, Economic Geology, 103, 877-908, 2008.

Banik, T. J., Coble, M. A., and Miller, C. F.: Porphyry Cu formation in the middle Jurassic Yerington batholith, Nevada, USA: Constraints from laser Raman, trace element, U-Pb age, and oxygen isotope analyses of zircon, Geosphere, 10.1130/ges01351.1, 2017.

Barberi, S., Bigioggero, B., Boriani, A., Cattaneo, M., Cavallin, A., Eva, C., Cioni, R., Gelmini, R., Giorgetti, F., and Iaccarino, S.: The Island of Sumbawa; a major structural discontinuity in the Indonesian Arc, Bollettino della Societa Geologica Italiana, 106, 547-620, 1987.

Barboni, M., Annen, C., and Schoene, B.: Evaluating the construction and evolution of upper crustal magma reservoirs with coupled U/Pb zircon geochronology and thermal modeling: A case study from the Mt. Capanne pluton (Elba, Italy), Earth and Planetary Science Letters, 432, 436-448, 10.1016/j.epsl.2015.09.043, 2015.

Barboni, M., Boehnke, P., Schmitt, A. K., Harrison, T. M., Shane, P., Bouvier, A.-S., and Baumgartner, L.: Warm storage for arc magmas, Proceedings of the National Academy of Sciences, 113, 13959-13964, 10.1073/pnas.1616129113, 2016.
Barra, F., Alcota, H., Rivera, S., Valencia, V., Munizaga, F., and Maksaev, V.: Timing and formation of porphyry Cu–Mo mineralization in the Chuquicamata district, northern Chile: new constraints from the Toki cluster, Mineralium Deposita, 48, 629-651, 10.1007/s00126-012-0452-1, 2013.

Berggren, W. A., Kent, D. V., Swisher III, C. C., and Aubry, M.-P.: A revised Cenozoic geochronology and chronostratigraphy, 1995.

Bowring, J. F., McLean, N. M., and Bowring, S.: Engineering cyber infrastructure for U-Pb geochronology: Tripoli and U-Pb_Redux, Geochemistry, Geophysics, Geosystems, 12, 2011.

Bucholz, C. E., Eddy, M. P., Jagoutz, O., Bowring, S. A., Schmidt, M. W., and Sambuu, O.: Constraining the time scales of magmatic differentiation with U-Pb zircon geochronology, Geology, 45, 11-14, 10.1130/g38505.1, 2017.

Buret, Y., von Quadt, A., Heinrich, C., Selby, D., Wälle, M., and Peytcheva, I.: From a long-lived upper-crustal magma chamber to rapid porphyry copper emplacement: Reading the geochemistry of zircon crystals at Bajo de la Alumbrera (NW Argentina), Earth and Planetary Science Letters, 450, 120-131, 10.1016/j.epsl.2016.06.017, 2016.

Buret, Y., Wotzlaw, J.-F., Roozen, S., Guillong, M., von Quadt, A., and Heinrich, C. A.: Zircon petrochronological evidence for a plutonic-volcanic connection in porphyry copper deposits, Geology, 45, 623-626, 10.1130/g38994.1, 2017.

Cannell, J., Cooke, D. R., Walshe, J. L., and Stein, H.: Geology, mineralization, alteration, and structural evolution of the El Teniente porphyry Cu-Mo deposit, Economic Geology, 100, 979-1003, 2005.

Caricchi, L., Simpson, G., and Schaltegger, U.: Zircons reveal magma fluxes in the Earth's crust, Nature, 511, 457-461, 10.1038/nature13532, 2014.

Caricchi, L., and Blundy, J.: Experimental petrology of monotonous intermediate magmas, Geological Society, London, Special Publications, 422, 105-130, 10.1144/sp422.9, 2015.

Cathles, L.: An analysis of the cooling of intrusives by ground-water convection which includes boiling, Economic Geology, 72, 804-826, 1977.

Cathles, L., and Shannon, R.: How potassium silicate alteration suggests the formation of porphyry ore deposits begins with the nearly explosive but barren expulsion of large volumes of magmatic water, Earth and Planetary Science Letters, 262, 92-108, 2007.

Cernuschi, F., Dilles, J. H., Grocke, S. B., Valley, J. W., Kitajima, K., and Tepley, F. J.: Rapid formation of porphyry copper deposits evidenced by diffusion of oxygen and titanium in quartz, Geology, 46, 611-614, 10.1130/g40262.1, 2018.

Chamberlain, K. J., Wilson, C. J. N., Wooden, J. L., Charlier, B. L. A., and Ireland, T. R.: New Perspectives on the Bishop Tuff from Zircon Textures, Ages and Trace Elements, Journal of Petrology, 55, 395-426, 10.1093/petrology/egt072, 2014.

Chelle-Michou, C., Chiaradia, M., Ovtcharova, M., Ulianov, A., and Wotzlaw, J.-F.: Zircon petrochronology reveals the temporal link between porphyry systems and the magmatic evolution of their hidden plutonic roots (the Eocene Corocohuayco deposit, Peru), Lithos, 198-199, 129-140, 10.1016/j.lithos.2014.03.017, 2014.

Chelle-Michou, C., Rottier, B., Caricchi, L., and Simpson, G.: Tempo of magma degassing and the genesis of porphyry copper deposits, Sci Rep, 7, 40566, 10.1038/srep40566, 2017.
Chelle-Michou, C., and Schaltegger, U.: U–Pb dating of mineral deposits: From age constraints to ore-forming processes, 2018.

Chiaradia, M., Schaltegger, U., Spikings, R., Wotzlaw, J.-F., and Ovtcharova, M.: How accurately can we date the duration of magmatic-hydrothermal events in porphyry systems?—an invited paper, Economic Geology, 108, 565-584, 2013.

Chiaradia, M., Schaltegger, U., and Spikings, R. A.: Time scales of mineral systems—advances in understanding over the past decade, 2014.

Chiaradia, M.: Crustal thickness control on Sr/Y signatures of recent arc magmas: an Earth scale perspective, Sci Rep, 5, 8115, 10.1038/srep08115, 2015.

Chiaradia, M., and Caricchi, L.: Stochastic modelling of deep magmatic controls on porphyry copper deposit endowment, Sci Rep, 7, 44523, 10.1038/srep44523, 2017.

Chiaradia, M.: Gold endowments of porphyry deposits controlled by precipitation efficiency, Nat Commun, 11, 248, 10.1038/s41467-019-14113-1, 2020.

Claiborne, L. L., Miller, C., Walker, B., Wooden, J., Mazdab, F., and Bea, F.: Tracking magmatic processes through Zr/Hf ratios in rocks and Hf and Ti zoning in zircons: an example from the Spirit Mountain batholith, Nevada, Mineralogical Magazine, 70, 517-543, 2006.

Claiborne, L. L., Miller, C. F., Flanagan, D. M., Clynne, M. A., and Wooden, J. L.: Zircon reveals protracted magma storage and recycling beneath Mount St. Helens, Geology, 38, 1011-1014, 10.1130/g31285.1, 2010a.

Claiborne, L. L., Miller, C. F., and Wooden, J. L.: Trace element composition of igneous zircon: a thermal and compositional record of the accumulation and evolution of a large silicic batholith, Spirit Mountain, Nevada, Contributions to Mineralogy and Petrology, 160, 511-531, 10.1007/s00410-010-0491-5, 2010b.

Clode, C.: Relationships of intrusion, wall-rock alteration and mineralization in the Batu Hijau copper-gold porphyry deposit, Proceedings, PACRIM'99 Congress, Bali, Indonesia, 1999, 485-498.

Condon, D., Schoene, B., McLean, N., Bowring, S., and Parrish, R.: Metrology and traceability of U–Pb isotope dilution geochronology (EARTHTIME Tracer Calibration Part I), Geochimica et Cosmochimica Acta, 164, 464-480, 2015.

Cooke, D. R., Hollings, P., and Walshe, J. L.: Giant porphyry deposits: characteristics, distribution, and tectonic controls, Economic Geology, 100, 801-818, 2005.

Correa, K. J., Rabbia, O. M., Hernández, L. B., Selby, D., and Astengo, M.: The timing of magmatism and ore formation in the El Abra porphyry copper deposit, northern Chile: Implications for long-lived multiple-event magmatic-hydrothermal porphyry systems, Economic Geology, 111, 1-28, 2016.

Deckart, K., Clark, A. H., Cuadra, P., and Fanning, M.: Refinement of the time-space evolution of the giant Mio-Pliocene Rio Blanco-Los Bronces porphyry Cu–Mo cluster, Central Chile: new U–Pb (SHRIMP II) and Re–Os geochronology and 40Ar/39Ar thermochronology data, Mineralium Deposita, 48, 57-79, 10.1007/s00126-012-0412-9, 2012.

Deino, A., and Keith, J. D.: Ages of volcanic and intrusive rocks in the Bingham mining district, Utah, Geology and ore deposits of the Oquirrh and Wasatch mountains, Utah: Society of Economic Geologists, Guidebook Series, 29, 91-95, 1997.
Dilles, J. H.: Petrology of the Yerington Batholith, Nevada; evidence for evolution of porphyry copper ore fluids, Economic Geology, 82, 1750-1789, 1987.

Dilles, J. H., and Einaudi, M. T.: Wall-rock alteration and hydrothermal flow paths about the Ann-Mason porphyry copper deposit, Nevada; a 6-km vertical reconstruction, Economic Geology, 87, 1963-2001, 1992.

Dilles, J. H., Kent, A. J., Wooden, J. L., Tosdal, R. M., Koleszar, A., Lee, R. G., and Farmer, L. P.: Zircon compositional evidence for sulfur-degassing from ore-forming arc magmas, Economic Geology, 110, 241-251, 2015.

Eddy, M. P., Bowring, S. A., Miller, R. B., and Tepper, J. H.: Rapid assembly and crystallization of a fossil large-volume silicic magma chamber, Geology, 44, 331-334, 10.1130/g37631.1, 2016.

Ferry, J. M., and Watson, E. B.: New thermodynamic models and revised calibrations for the Ti-in-zircon and Zr-in-rutile thermometers, Contributions to Mineralogy and Petrology, 154, 429-437, 10.1007/s00410-007-0201-0, 2007.

Fiorentini, M. L., and Garwin, S. L.: Evidence of a mantle contribution in the genesis of magmatic rocks from the Neogene Batu Hijau district in the Sunda Arc, South Western Sumbawa, Indonesia, Contributions to Mineralogy and Petrology, 159, 819-837, 10.1007/s00410-009-0457-7, 2010.

Foden, J., and Varne, R.: The petrology and tectonic setting of Quaternary—Recent volcanic centres of Lombok and Sumbawa, Sunda arc, Chemical Geology, 30, 201-226, 1980.

Garwin, S., Hall, R., and Watanabe, Y.: Tectonic setting, geology, and gold and copper mineralization in Cenozoic magmatic arcs of Southeast Asia and the West Pacific, Economic Geology 100th anniversary volume, 891-930, 2005.

Garwin, S. L.: The Setting, Geometry and Timing of Intrusion-related Hydrothermal System in the Vicinity of the Batu Hijau porphyry Copper-gold deposit, Sumbawa, Indonesia, PhD Thesis, University of Western Australia, 2000.

Gilmer, A. K., Sparks, R. S. J., Rust, A. C., Tapster, S., Webb, A. D., and Barfod, D. N.: Geology of the Don Manuel igneous complex, central Chile: Implications for igneous processes in porphyry copper systems, Geological Society of America Bulletin, 129, 920-946, 10.1130/b31524.1, 2017.

Guillong, M., von, von Quadt, A., Sakata, S., Peytcheva, I., and Bachmann, O.: LA-ICP-MS Pb–U dating of young zircons from the Kos–Nisyros volcanic centre, SE Aegean arc, Journal of Analytical Atomic Spectrometry, 29, 963-970, 2014.

Hall, R.: Cenozoic geological and plate tectonic evolution of SE Asia and the SW Pacific: computer-based reconstructions, model and animations, Journal of Asian Earth Sciences, 20, 353-431, 2002.

Halter, W. E., Bain, N., Becker, K., Heinrich, C. A., Landtwing, M., VonQuadt, A., Clark, A. H., Sasso, A. M., Bissig, T., and Tosdal, R. M.: From andesitic volcanism to the formation of a porphyry Cu-Au mineralizing magma chamber: the Farallón Negro Volcanic Complex, northwestern Argentina, Journal of Volcanology and Geothermal Research, 136, 1-30, 10.1016/j.jvolgeores.2004.03.007, 2004.

Hamilton, W. B.: Tectonics of the Indonesian region, 1078, US Govt. Print. Off., 1979.

Horstwood, M. S. A., Kößler, J., Gehrels, G., Jackson, S. E., McLean, N. M., Paton, C., Pearson, N. J., Sircombe, K., Sylvester, P., Vermeesch, P., Bowring, J. F., Condon, D. J., and Schoene, B.: Community-Derived Standards for LA-ICP-MS Pb–U dating of young zircons from the Kos–Nisyros volcanic centre, SE Aegean arc, Journal of Analytical Atomic Spectrometry, 29, 963-970, 2014.
MS U-(Th-)Pb Geochronology - Uncertainty Propagation, Age Interpretation and Data Reporting, Geostandards and Geoanalytical Research, 40, 311-332, 10.1111/j.1751-908X.2016.00379.x, 2016.

Hoskin, P. W., and Ireland, T. R.: Rare earth element chemistry of zircon and its use as a provenance indicator, Geology, 28, 627-630, 2000.

Hoskin, P. W., and Schaltegger, U.: The composition of zircon and igneous and metamorphic petrogenesis, Reviews in mineralogy and geochemistry, 53, 27-62, 2003.

Hutchison, C. S.: Geological evolution of South-east Asia, Clarendon Press Oxford, 1989.

Idrus, A., Kolb, J., and Meyer, F. M.: Chemical Composition of Rock-Forming Minerals in Copper-Gold-Bearing Tonalite Porphyries at the Batu Hijau Deposit, Sumbawa Island, Indonesia: Implications for Crystallization Conditions and Fluorine-Chlorine Fugacity, Resource Geology, 57, 102-113, 2007.

Jaeger, J.: The temperature in the neighborhood of a cooling intrusive sheet, American Journal of Science, 255, 306-318, 1957.

Karakas, O., Degruyter, W., Bachmann, O., and Dufek, J.: Lifetime and size of shallow magma bodies controlled by crustal-scale magmatism, Nature Geoscience, 10, 446-450, 10.1038/ngeo2959, 2017.

Keller, C. B., Schoene, B., and Samperton, K. M.: A stochastic sampling approach to zircon eruption age interpretation, Geochemical Perspectives Letters, 31-35, 10.7185/geochemlet.1826, 2018.

Krogh, T.: A low-contamination method for hydrothermal decomposition of zircon and extraction of U and Pb for isotopic age determinations, Geochimica et Cosmochimica Acta, 37, 485-494, 1973.

Landtwing, M., Pettke, T., Halter, W., Heinrich, C., Redmond, P., Einaudi, M., and Kunze, K.: Copper deposition during quartz dissolution by cooling magmatic–hydrothermal fluids: The Bingham porphyry, Earth and Planetary Science Letters, 235, 229-243, 10.1016/j.epsl.2005.02.046, 2005.

Large, S. J. E., Bakker, E. Y. N., Weis, P., Wälle, M., Ressel, M., and Heinrich, C. A.: Trace elements in fluid inclusions of sediment-hosted gold deposits indicate a magmatic-hydrothermal origin of the Carlin ore trend, Geology, 44, 1015-1018, 10.1130/g38351.1, 2016.

Large, S. J. E., von Quadt, A., Wotzlaw, J.-F., Guillong, M., and Heinrich, C. A.: Magma Evolution Leading to Porphyry Au-Cu Mineralization at the Ok Tedi Deposit, Papua New Guinea: Trace Element Geochemistry and High-Precision Geochronology of Igneous Zircon, Economic Geology, 113, 39-61, 10.5382/econgeo.2018.4543, 2018.

Lee, C.-T. A., and Tang, M.: How to make porphyry copper deposits, Earth and Planetary Science Letters, 529, 10.1016/j.epsl.2019.115868, 2020.

Lee, R. G., Dilles, J. H., Tosdal, R. M., Wooden, J. L., and Mazdab, F. K.: Magmatic Evolution of Granodiorite Intrusions at the El Salvador Porphyry Copper Deposit, Chile, Based on Trace Element Composition and U/Pb Age of Zircons, Economic Geology, 112, 245-273, 10.2113/econgeo.112.2.245, 2017.
Loader, M. A., Wilkinson, J. J., and Armstrong, R. N.: The effect of titanite crystallisation on Eu and Ce anomalies in zircon and its implications for the assessment of porphyry Cu deposit fertility, Earth and Planetary Science Letters, 472, 107-119, 10.1016/j.epsl.2017.05.010, 2017.

Longo, A. A., Dilles, J. H., Grunder, A. L., and Duncan, R.: Evolution of calc-alkaline volcanism and associated hydrothermal gold deposits at Yanacocha, Peru, Economic Geology, 105, 1191-1241, 2010.

Maksaev, V., Munizaga, F., McWilliams, M., Fanning, M., Mathur, R., Ruiz, J., and Zentilli, M.: New chronology for El Teniente, Chilean Andes, from U-Pb, 40Ar/39Ar, Re-Os, and fission-track dating: Implications for the evolution of a supergiant porphyry Cu-Mo deposit, Society of Economic Geologists, 11, 15-54, 2004.

Mattinson, J. M.: Zircon U–Pb chemical abrasion (“CA-TIMS”) method: combined annealing and multi-step partial dissolution analysis for improved precision and accuracy of zircon ages, Chemical Geology, 220, 47-66, 2005.

McLean, N. M., Bowring, J. F., and Bowring, S.: An algorithm for U-Pb isotope dilution data reduction and uncertainty propagation, Geochemistry, Geophysics, Geosystems, 12, 2011a.

McLean, N. M., Bowring, J. F., and Bowring, S. A.: An algorithm for U-Pb isotope dilution data reduction and uncertainty propagation, Geochemistry, Geophysics, Geosystems, 12, n/a-n/a, 10.1029/2010gc003478, 2011b.

McLean, N. M., Condon, D. J., Schoene, B., and Bowring, S. A.: Evaluating uncertainties in the calibration of isotopic reference materials and multi-element isotopic tracers (EARTHTIME Tracer Calibration Part II), Geochimica et Cosmochimica Acta, 164, 481-501, 2015.

Meldrum, S. J., Aquino, R. S., Gonzales, R. I., Burke, R. J., Suyadi, A., Irianto, B., and Clarke, D. S.: The Batu Hijau porphyry copper-gold deposit, Sumbawa Island, Indonesia, Journal of Geochemical Exploration, 50, 203-220, 10.1016/0375-6742(94)90025-6, 1994.

Mercer, C. N., Reed, M. H., and Mercer, C. M.: Time scales of porphyry Cu deposit formation: insights from titanium diffusion in quartz, Economic Geology, 110, 587-602, 2015.

Miller, J. S., Matzel, J. E. P., Miller, C. F., Burgess, S. D., and Miller, R. B.: Zircon growth and recycling during the assembly of large, composite arc plutons, Journal of Volcanology and Geothermal Research, 167, 282-299, 10.1016/j.jvolgeores.2007.04.019, 2007.

Mitchell, P., Proffett, J., and Dilles, J.: Geological review of the Batu Hijau porphyry copper-gold deposit, Sumbawa Island, Indonesia, Newmont Nusa Tenggara Company Final report. PT Newmont Nusa Tenggara, Batu Hijau project, Sumbawa, Indonesia, 164p, 1998.

Oberli, F., Meier, M., Berger, A., Rosenberg, C. L., and GierÉ, R.: U-Th-Pb and 230Th/238U disequilibrium isotope systematics: Precise accessory mineral chronology and melt evolution tracing in the Alpine Bergell intrusion, Geochemica et Cosmochimica Acta, 68, 2543-2560, 10.1016/j.gca.2003.10.017, 2004.

Proffett, J. M.: Geology of the Bajo de la Alumbrera porphyry copper-gold deposit, Argentina, Economic Geology, 98, 1535-1574, 2003.
Proffett, J. M.: High Cu grades in porphyry Cu deposits and their relationship to emplacement depth of magmatic sources, Geology, 37, 675-678, 10.1130/g30072a.1, 2009.

Redmond, P. B., and Einaudi, M. T.: The Bingham Canyon porphyry Cu-Mo-Au deposit. I. Sequence of intrusions, vein formation, and sulfide deposition, Economic Geology, 105, 43-68, 2010.

Reid, M. R., Vazquez, J. A., and Schmitt, A. K.: Zircon-scale insights into the history of a Supervolcano, Bishop Tuff, Long Valley, California, with implications for the Ti-in-zircon geothermometer, Contributions to Mineralogy and Petrology, 161, 293-311, 10.1007/s00410-010-0532-0, 2011.

Reiners, P. W., Carlson, R. W., Renne, P. R., Cooper, K. M., Granger, D. E., McLean, N. M., and Schoene, B.: Interpretational approaches: making sense of data, in: Geochronology and thermochronology, John Wiley & Sons, 2017.

Rezeau, H., Moritz, R., Wotzlaw, J.-F., Tayan, R., Melkonyan, R., Ulianov, A., Selby, D., d’Abzac, F.-X., and Stern, R. A.: Temporal and genetic link between incremental pluton assembly and pulsed porphyry Cu-Mo formation in accretionary orogens, Geology, 44, 627-630, 10.1130/g38088.1, 2016.

Richards, J. P.: Giant ore deposits formed by optimal alignments and combinations of geological processes, Nature Geoscience, 6, 911-916, 10.1038/engeo1920, 2013.

Rivera, T. A., Schmitz, M. D., Crowley, J. L., and Storey, M.: Rapid magma evolution constrained by zircon petrochronology and 40Ar/39Ar sanidine ages for the Huckleberry Ridge Tuff, Yellowstone, USA, Geology, 42, 643-646, 10.1130/g35808.1, 2014.

Roedder, E.: Fluid inclusion studies on the porphyry-type ore deposits at Bingham, Utah, Butte, Montana, and Climax, Colorado, Economic Geology, 66, 98-118, 1971.

Rohrlach, B. D., Loucks, R. R., and Porter, T.: Multi-million-year cyclic ramp-up of volatiles in a lower crustal magma reservoir trapped below the Tampakan copper-gold deposit by Mio-Pliocene crustal compression in the southern Philippines, 2005.

Rubatto, D., and Hermann, J.: Experimental zircon/melt and zircon/garnet trace element partitioning and implications for the geochronology of crustal rocks, Chemical Geology, 241, 38-61, 10.1016/j.chemgeo.2007.01.027, 2007.

Samperton, K. M., Schoene, B., Cottle, J. M., Brenhin Keller, C., Crowley, J. L., and Schmitz, M. D.: Magma emplacement, differentiation and cooling in the middle crust: Integrated zircon geochronological–geochemical constraints from the Bergell Intrusion, Central Alps, Chemical Geology, 417, 322-340, 10.1016/j.chemgeo.2015.10.024, 2015.

Sasso, A.: The Farallón Negro Group, northwest Argentina: magmatic, hydrothermal and tectonic evolution and implications for Cu-Au metallogeny in the Andean back-arc, Soc. Econ. Geol. Newsletter, 34, 8-18, 1998.

Schaltegger, U., Brack, P., Ovtcharova, M., Peytcheva, I., Schoene, B., Stracke, A., Marocchi, M., and Bargossi, G. M.: Zircon and titanite recording 1.5million years of magma accretion, crystallization and initial cooling in a composite pluton (southern Adamello batholith, northern Italy), Earth and Planetary Science Letters, 286, 208-218, 10.1016/j.epsl.2009.06.028, 2009.
Schärer, U.: The effect of initial $^{230}$Th disequilibrium on young U Pb ages: The Makalu case, Himalaya, Earth and Planetary Science Letters, 67, 191-204, 1984.

Schirra, M., Driesner, T., and Heinrich, C. A.: Relative timing of sulfide precipitation at the Batu Hijau porphyry Cu-Au deposit, Sumbawa, Indonesia, Life with Ore Deposits on Earth, Proceedings of the 15th SGA Biennial Meeting, Glasgow, 2019, 2019, 1009-1012.

Schoene, B., Schaltegger, U., Brack, P., Latkoczy, C., Stracke, A., and Günther, D.: Rates of magma differentiation and emplacement in a ballooning pluton recorded by U–Pb TIMS-TEA, Adamello batholith, Italy, Earth and Planetary Science Letters, 355-356, 162-173, 10.1016/j.epsl.2012.08.019, 2012.

Schoene, B.: U–Th–Pb Geochronology, in: Treatise on Geochemistry, 341-378, 2014.

Seedorff, E., and Einaudi, M. T.: Henderson porphyry molybdenum system, Colorado: I. Sequence and abundance of hydrothermal mineral assemblages, flow paths of evolving fluids, and evolutionary style, Economic Geology, 99, 3-37, 2004.

Seedorff, E., Barton, M. D., Stavast, W. J., and Maher, D. J.: Root zones of porphyry systems: Extending the porphyry model to depth, Economic Geology, 103, 939-956, 2008.

Sillitoe, R. H.: Porphyry copper systems, Economic Geology, 105, 3-41, 2010.

Smith, W. D., Darling, J. R., Bullen, D. S., Lasalle, S., Pereira, I., Moreira, H., Allen, C. J., and Tapster, S.: Zircon perspectives on the age and origin of evolved S-type granites from the Cornubian Batholith, Southwest England, Lithos, 336-337, 14-26, 10.1016/j.lithos.2019.03.025, 2019.

Spencer, E. T., Wilkinson, J. J., Creaser, R. A., and Seguel, J.: The distribution and timing of molybdenite mineralization at the El Teniente Cu-Mo porphyry deposit, Chile, Economic Geology, 110, 387-421, 2015.

Steinberger, I., Hinks, D., Driesner, T., and Heinrich, C. A.: Source plutons driving porphyry copper ore formation: combining geomagnetic data, thermal constraints, and chemical mass balance to quantify the magma chamber beneath the Bingham Canyon deposit, Economic Geology, 108, 605-624, 2013.

Szymanowski, D., Wotzlaw, J.-F., Ellis, B. S., Bachmann, O., Guillogl, M., and von Quadt, A.: Protracted near-solidus storage and pre-eruptive rejuvenation of large magma reservoirs, Nature Geoscience, 10, 777-782, 10.1038/ngeo3020, 2017.

Tapster, S., Condon, D. J., Naden, J., Noble, S. R., Petterson, M. G., Roberts, N. M. W., Saunders, A. D., and Smith, D. J.: Rapid thermal rejuvenation of high-crystallinity magma linked to porphyry copper deposit formation; evidence from the Koloula Porphyry Prospect, Solomon Islands, Earth and Planetary Science Letters, 442, 206-217, 10.1016/j.epsl.2016.02.046, 2016.

van Dongen, M., Weinberg, R. F., and Tomkins, A. G.: REE-Y, Ti, and P remobilization in magmatic rocks by hydrothermal alteration during Cu-Au deposit formation, Economic Geology, 105, 763-776, 2010.

Vazquez, J. A., and Reid, M. R.: Probing the accumulation history of the voluminous Toba magma, Science, 305, 991-994, 10.1126/science.1096994, 2004.
Vermeesch, P.: HelioPlot, and the treatment of overdispersed (U–Th–Sm)/He data, Chemical Geology, 271, 108-111, 10.1016/j.chemgeo.2010.01.002, 2010.

Vermeesch, P.: On the visualisation of detrital age distributions, Chemical Geology, 312-313, 190-194, 10.1016/j.chemgeo.2012.04.021, 2012.

Vermeesch, P.: IsoplotR: A free and open toolbox for geochronology, Geoscience Frontiers, 9, 1479-1493, 10.1016/j.gsf.2018.04.001, 2018.

von Quadt, A., Erni, M., Martinek, K., Moll, M., Peytcheva, I., and Heinrich, C. A.: Zircon crystallization and the lifetimes of ore-forming magmatic-hydrothermal systems, Geology, 39, 731-734, 10.1130/g31966.1, 2011.

von Quadt, A., Wotzlaw, J.-F., Buret, Y., Large, S. J., Peytcheva, I., and Trinquier, A.: High-precision zircon U/Pb geochronology by ID-TIMS using new 10 13 ohm resistors, Journal of Analytical Atomic Spectrometry, 31, 658-665, 2016.

Watson, E., and Harrison, T.: Zircon thermometer reveals minimum melting conditions on earliest Earth, Science, 308, 841-844, 2005.

Watson, E., Wark, D., and Thomas, J.: Crystallization thermometers for zircon and rutile, Contributions to Mineralogy and Petrology, 151, 413, 2006.

Wawryk, C. M., and Foden, J. D.: Iron-isotope systematics from the Batu Hijau Cu-Au deposit, Sumbawa, Indonesia, Chemical Geology, 466, 159-172, 10.1016/j.chemgeo.2017.06.004, 2017.

Weis, P., Driesner, T., and Heinrich, C. A.: Porphyry-copper ore shells form at stable pressure-temperature fronts within dynamic fluid plumes, Science, 338, 1613-1616, 10.1126/science.1225009, 2012.

Wendt, I., and Carl, C.: The statistical distribution of the mean squared weighted deviation, Chemical Geology: Isotope Geoscience Section, 86, 275-285, 1991.

Wilkinson, J. J.: Triggers for the formation of porphyry ore deposits in magmatic arcs, Nature Geoscience, 6, 917-925, 10.1038/ngeo1940, 2013.

Wotzlaw, J.-F., Schaltegger, U., Frick, D. A., Dungan, M. A., Gerdes, A., and Günther, D.: Tracking the evolution of large-volume silicic magma reservoirs from assembly to supereruption, Geology, 41, 867-870, 10.1130/g34366.1, 2013.

Wotzlaw, J.-F., Buret, Y., Large, S. J., Szymanowski, D., and von Quadt, A.: ID-TIMS U–Pb geochronology at the 0.1% level using 10¹³ Ω resistors and simultaneous U and ¹⁸O/¹⁶O isotope ratio determination for accurate UO₂ interference correction, Journal of Analytical Atomic Spectrometry, 32, 579-586, 2017.

Zwyer, T.: Temporal and Spatial Evolution of Hydrothermal, Ore-Related Fluids in the Batu Hijau Porphyry Copper-Gold Deposit, Sumbawa (Indonesia), M.Sc. ETH, Institute of Petrology and Geochemistry, ETH Zurich, 52 pp., 2011.