FORMING CHONDRITES IN A SOLAR NEBULA WITH MAGNETICALLY INDUCED TURBULENCE

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ABSTRACT

Chondritic meteorites provide valuable opportunities to investigate the origins of the solar system. We explore impact jetting as a mechanism of chondrule formation and subsequent pebble accretion as a mechanism of accreting chondrules onto parent bodies of chondrites, and investigate how these two processes can account for the currently available meteoritic data. We find that when the solar nebula is \( \approx 5 \) times more massive than the minimum-mass solar nebula at \( a \approx 2-3 \) au and parent bodies of chondrites are \( \leq 10^{25} \) g (\( \leq 500 \) km in radius) in the solar nebula, impact jetting and subsequent pebble accretion can reproduce a number of properties of the meteoritic data. The properties include the present asteroid belt mass, the formation timescale of chondrules, and the magnetic field strength of the nebula derived from chondrules in Semarkona. Since this scenario requires a first generation of planetesimals that trigger impact jetting and serve as parent bodies to accrete chondrules, the upper limit of parent bodies’ masses leads to the following implications: primordial asteroids that were originally \( \geq 10^{24} \) g in mass were unlikely to contain chondrules, while less massive primordial asteroids likely had a chondrule-rich surface layer. The scenario developed from impact jetting and pebble accretion can therefore provide new insights into the origins of the solar system.

Key words: magnetic fields – meteorites, meteors, meteoroids – minor planets, asteroids: general – planets and satellites: formation – protoplanetary disks – turbulence

1. INTRODUCTION

Chondritic meteorites can be regarded as a fossil record of how the solar system formed (e.g., Davis et al. 2014, p. 809). The most invaluable records are contained in chondrules, which are the most abundant ingredient in chondrites. Chondrules are millimeter-sized spherical particles that are formed as the consequence of transient heating events occurring in the solar nebula (e.g., Hewins et al. 2005). Since such events may be relevant to important processes regulating nebula evolution (e.g., Desch et al. 2012), investigating chondrule formation can shed light on the origins of the solar system. Chondrites themselves also possess profound information. It is widely accepted that chondrites are fragments of planetesimals originating from the asteroid belt (e.g., DeMeo et al. 2015). Taking into account the consideration that planetesimals are the building blocks of planets, understanding how chondrites are generated from their parent bodies can provide fundamental insight into the time evolution of the asteroid belt, planetesimal formation, and even (perhaps) the formation of (exo)planetary systems (e.g., Vernazza et al. 2014; Johansen et al. 2015).

Significant progress has recently been made on chondrule formation and accretion. For instance, Johnson et al. (2015) have produced a pioneering work in which planetesimal collisions and the resultant impact jetting are proposed as a mechanism of chondrule formation. They have found that ejected materials from the collisional surface can reproduce the thermal history of chondrules, and the collisional frequency during planetary accretion is high enough to account for the present asteroid belt mass (also see Hasegawa et al. 2016). In addition, an experimental study revealed that primitive chondrites and the chondrules embedded in them can retain the information of the magnetic field strength of the solar nebula (Fu et al. 2014). Furthermore, Lambrechts & Johansen (2012) have performed numerical simulations and shown that gas drag acting on pebbles can substantially accelerate the growth of massive bodies due to the efficient accretion of pebbles onto the bodies (also see Ormel & Klahr 2010; Guillot et al. 2014; Kretke & Levison 2014). This process can be viewed as a mechanism of chondrule accretion onto parent bodies of chondrites (Johansen et al. 2015).

Here we synthesize these recent steps on chondrule formation and accretion, and explore how impact jetting and subsequent pebble accretion can be consistent with the meteorite conditions. We show below that the coupling of these processes with the meteorite data can provide tight constraints on the mass of the solar nebula at \( a \approx 2-3 \) au and the size of parent bodies of chondrites in the solar nebula. We will therefore conclude that both impact jetting and pebble accretion are intriguing processes for obtaining some valuable clues about the origins of the solar system.

2. CHONDRULE FORMATION AND ACCRETION

2.1. Disk Model

We adopt the minimum-mass solar nebular (MMSN) model (Hayashi 1981), which can be given as (see Table 1 for the definition of variables),

\[
\Sigma_d = 1.7 \times 10^3 f_d \left( \frac{a}{1 \text{ au}} \right)^{-3/2} \text{ g cm}^{-2},
\]

\[
\Sigma_s = 7 f_d \left( \frac{a}{1 \text{ au}} \right)^{-3/2} \text{ g cm}^{-2},
\]

and

\[
T_d = 280 \left( \frac{a}{1 \text{ au}} \right)^{-1/2} \text{ K},
\]

where \( f_d \) is a factor to examine the effect of disk mass on chondrule formation,\(^3\) and \( T_d \) is derived under the optically thin conditions.

\(^3\) Note that both the gas and the dust densities in Hasegawa et al. (2016) are increased by a factor of 1.5 for the standard model, following Kokubo & Ida (2000).
can reach $M_{\text{iso}}$, within a given disk lifetime (Hasegawa et al. 2016). This formula is obtained because chondrule-forming collisions can be realized when protoplanets undergo oligarchic growth and acquire most of their masses (e.g., Kokubo & Ida 1998).

The total abundance of chondrules for $2 \text{ au} \lesssim a \lesssim 3 \text{ au}$ can be readily computed as (see Hasegawa et al. 2016 for a complete discussion and references)

$$m_{\text{ch,ij}}(\tau_{\text{ch}}) \approx \sum_a f_{\text{ij}} M_{\text{iso}}(a),$$  

(6)

where

$$M_{\text{iso}} \simeq 9.4 \times 10^{22} M_{\odot} \left(\frac{\Sigma_{\text{ch}}}{7 \text{ g cm}^{-2}}\right)^{3/2} \left(\frac{a}{1 \text{ au}}\right)^3.$$

(7)

Equation (6) assumes that protoplanets are separated by their feeding zones ($\approx 10_{\text{M}}$), with each other, and that chondrules were formed over a time interval $\tau_{\text{ch}}$. Since meteorite records indicate that chondrule formation began around the time of CAI formation and continued for $3-5 \times 10^6$ years (e.g., Connelly et al. 2012; Bollard et al. 2015), two characteristic values of $\tau_{\text{ch}}$ are considered ($\tau_{\text{ch}} = 3 \times 10^6$ years and $\tau_{\text{ch}} = 5 \times 10^6$ years; see Figure 1). In the summation, terms contribute only if $\tau_{\text{ch}}$ is longer than the time ($\tau_{\text{ij}}$) for protoplanets to grow up to $M_{\text{iso}}$, which is given as (Ida & Makino 1993)

$$\tau_{\text{ij}} \equiv f_{\text{ij}} \frac{M_{\text{p}}}{dM_{\text{p}}/dt} \bigg|_{M_{\text{p}}=M_{\text{iso}}},$$  

(8)

where

$$\frac{dM_{\text{p}}}{dt} \simeq 4\pi \Sigma_{\text{ch}} GM_{\text{p}}R_{\text{pl}} \left(\frac{a_{\text{eq}}}{a_{\text{pl}}}\right)^2 \Omega,$$

(9)

and

$$\langle e_{\text{eq}}^2 \rangle^{1/2} \simeq 2.8 \times 10^{-2} \left(\frac{m_{\text{pl}}}{10^{23} \text{ g}}\right)^{1/15} \left(\frac{\rho_{\text{pl}}}{2 \text{ g cm}^{-3}}\right)^{2/15} \times \left(\frac{1.4 \times 10^{-9} \text{ g cm}^{-3}}{\rho_e}\right)^{-1/2} \times \left(\frac{a}{1 \text{ au}}\right)^{-1/5} \left(\frac{M_{\text{p}}}{0.1 M_{\odot}}\right)^{1/3}.$$

(10)

As found by Hasegawa et al. (2016), $f_{\text{ij}} = 3$ is needed to reproduce the results of more detailed calculations. Also, the oligarchic phase is assumed in computing planetesimal eccentricity ($\langle e_{\text{eq}}^2 \rangle$), wherein the balance between the pump up of ($e_{\text{eq}}^2$) by a protoplanet and its damping by the disk gas is considered. Thus, $m_{\text{ch,ij}}$ is calculated based on the semi-analytical formulation developed in Hasegawa et al. (2016), under the assumption that protoplanet–planetesimal collisions play a dominant role for $m_{\text{ch,ij}}$.

Figure 1 shows the resultant value of $m_{\text{ch,ij}}$ as a function of $f_{\text{ij}}$. The value of $m_{\text{pl}}$ varies from $10^{22}$ to $2 \times 10^{23}$ g in each

### Table 1

**List of Quantities**

| Symbol | Meanings | Value |
|--------|----------|-------|
| $a$    | Semimajor axis | ... |
| $\Omega$ | Angular frequency | ... |
| $c_s$ | Sound speed | ... |
| $\Sigma$ | Gas surface density | ... |
| $\rho_g$ | Gas volume density at the disk midplane | ... |
| $f_d$ | Increment factor | ... |
| $T_d$ | Disk temperature | ... |
| $h_g$ | Gas pressure scale height | ... |
| $\langle B \rangle$ | Magnetic field strength estimated from a chondrite | 50–540 mG |
| $\langle B \rangle_{\text{min}}$ | Minimum value of $\langle B \rangle$ required for pebble accretion | ... |
| $\alpha_{\text{eff}}$ | Effective value of $\alpha$ computed from $\langle B \rangle$ | ... |
| $\gamma_{\text{hub}}$ | Quantity related to the nature of turbulence | 2 |
| $M_{\text{p}}$ | Mass of protoplanets | ... |
| $R_{\text{ch}}$ | Radii of protoplanets | ... |
| $\rho_{\text{ch}}$ | Mean density of protoplanets | 5 g cm$^{-3}$ |
| $M_{\text{ch}}$ | Isolation mass of protoplanets | ... |
| $r_{\text{ch}}$ | Hill radius of protoplanets | ... |
| $\rho_{\text{pl}}$ | Root mean square equilibrium eccentricity of field planetesimals in oligarchic growth | ... |
| $h_{\text{pl}}$ | Vertical height of planetesimals | ... |
| $\rho_{\text{pl}}$ | Mean density of field planetesimals | 2 g cm$^{-3}$ |
| $\tau_{\text{ch}}$ | Duration of the chondrule-forming epoch (since CAI formation) | 3–5 $\times 10^6$ yr |
| $m_{\text{ch,ij}}$ | Total mass of chondrules formed via impact jetting for $\tau_{\text{ch}}$ | ... |
| $f_{\text{ij}}$ | Mass fraction of planetesimals that can eventually generate chondrules via impact jetting | $10^{-2}$ |
| $\tau_{\text{ij}}$ | Timescale for forming chondrules via impact jetting | ... |

2.2. Impact Jetting

We consider impact jetting as one of the primary processes for generating chondrules, which can occur during planetary accretion. For this case, the resultant abundance of chondrules can be characterized by $f_{\text{ij}}M_{\text{iso}}$ as long as formed protoplanets
2.3. Magnetic Records of Chondrules

Magnetic fields can play an important role in the evolution of circumstellar disks (e.g., Turner et al. 2014, p. 411). The existence of these fields in the solar nebula is supported by measurements of the Semarkona ordinary chondrite (Fu et al. 2014). In Fu et al.’s experiment, the thermoremanent magnetization of chondrules in the chondrite and the direction of magnetic fields are measured simultaneously. These measurements suggest that some chondrules in this sample very likely recorded the magnetic field strength that they experienced well before being accreted by their parent body. In other words, the data trace the magnetic field strength of the solar nebula around the chondrule-forming region, which is \( B \approx 50-540 \text{ mG} \). This information is of fundamental importance for characterizing the strength of turbulence in the nebula, which can be computed as (e.g., Wardle 2007; Okuzumi & Hirose 2011)

\[
\alpha_{\text{eff}} = \left( \frac{B_r B_\phi}{\sum q_h \Omega^2} \right) \leq \frac{B_r^2}{\sum q_h \Omega^2}, \tag{11}
\]

where \( B_r \) and \( B_\phi \) are the radial and the azimuthal components of magnetic fields, respectively, and the famous \( \alpha \)-prescription is adopted to label the turbulent strength (Shakura & Sunyaev 1973). We will use the resultant value of \( \alpha_{\text{eff}} \) (and \( B \)) below.

\[\text{Figure 1. Mass of chondrules (} m_{ch,pl} \text{) produced by impact jetting via planetary accretion at 2–3 au as a function of } f_d \text{ (see Table 1). The value of } m_{pl} \text{ varies from } 10^{22} \text{ g (top)} \text{ to } 2 \times 10^{24} \text{ g (bottom). For comparison purposes, both the primordial and the present asteroid belt masses are plotted as the horizontal dotted lines. The minimum value of } f_d \text{ above which } m_{ch,pl} \text{ is large enough to reproduce the required abundance of chondrules increases with increasing } m_{pl} \text{ (see the vertical dotted line). This occurs because a disk becomes massive enough at that time, so a growing protoplanet at } a = 2 \text{ au can reach } M_{\text{iso}} \text{ within given values of } \tau_{\text{ch}}. \text{ Once } f_d \text{ exceeds these critical values, the increment of } m_{ch,pl} \text{ becomes gradual until other sudden jumps in } m_{ch,pl} \text{ appear. These jumps originate from an additional contribution arising from another protoplanet that can now reach } M_{\text{iso}} \text{ within } \tau_{\text{ch}} \text{ by increasing } f_d. \text{ Note that this second formed protoplanet is separated from the first formed protoplanet by } \Delta a \approx 0.1 \text{ au.}
\]

The results also show that the first sudden rises of \( m_{ch,pl} \) require a higher value of \( f_d \) as \( m_{pl} \) increases (see the vertical dotted line on each panel in Figure 1). This is a simple reflection of \( \tau_f \), which is an increasing function of \( m_{pl} \) (see Equations (8)–(10)); when massive planetesimals are involved in planetary accretion, the value of \( \epsilon_* \) becomes higher, since the eccentricity damping by the disk gas becomes less effective. Then, the growth rate of a protoplanet slows down. As a result, a higher value of \( f_d \) is needed to accelerate the formation of protoplanets, in order to satisfy the condition that protoplanets obtain \( M_{\text{iso}} \) within \( \tau_{\text{ch}} \).

Thus, the abundance of chondrules formed by impact jetting through planetesimal collisions is very likely significant enough to reproduce the present asteroid belt mass; for given values of \( \tau_{\text{ch}} \), the formation of protoplanets at \( a \approx 2 \text{ au} \) generates a large amount of chondrules for reasonable values of \( f_d \) and \( m_{pl} \). This finding is consistent with the previous studies (Johnson et al. 2015; Hasegawa et al. 2016). Nonetheless, it is obvious that the results degenerate for certain ranges of \( f_d \) and \( m_{pl} \). We will determine the ranges in these two quantities below, using further constraints from meteorite measurements.
2.4. Pebble Accretion

Pebble accretion is a promising candidate process for forming the chondrite parent bodies (e.g., Lambrechts & Johansen 2012; Johansen et al. 2015). For this case, the vertical scale heights of chondrules ($h_{ch}$) and the planetesimals that accrete them ($h_{pl}$) are important for determining the accretion efficiency. Specifically, pebble accretion is most efficient for planetesimals with a scale height $h_{pl} < h_{ch}$, namely, when they lie among the pebbles (Levison et al. 2013). We assume this holds in the following. Note that $\tau_{ch}$ (~3–5 $\times$ 10$^6$ years) is longer than the timescale at which a steady state of $h_{pl}$ is achieved, which is less than 10$^6$ years (Kokubo & Ida 2000).

Chondrules are typically about 1 mm in size, and large enough to settle toward the midplane due to stellar gravity. In addition, the solar nebula is very likely turbulent, which diffuses chondrules vertically. Based on Equation (11), $\alpha_{eff} \approx 10^{-4} - 10^{-2}$ for $f_d \gtrsim 4$ at $a \simeq 2-3$ au. This is a manifestation that shows that magnetic field can play an important role in the solar nebula. As a result, $h_{ch}$ can be determined by the balance between stirring and settling. Then, $h_{ch}$ can be written as (Dubrulle et al. 1995)

$$\frac{h_{ch}(r_{ch})}{h_g} = \frac{H}{\sqrt{1 + H^2}},$$

where

$$H = \left( \frac{1}{1 + \gamma_{turb}} \right)^{1/4} \alpha_{eff} \Sigma_g \sqrt{2\pi \rho_{ch} r_{ch}},$$

This formula is consistent with other studies (e.g., Youdin & Lithwick 2007), and MHD simulations show that this approach works well for turbulent regions (e.g., Fromang & Papaloizou 2006).

For the planetesimals ($h_{pl}$), an eccentricity pump-up by a protoplanet is very likely to dominate over the effect of the random torque caused by turbulence in the solar nebula. This is because $\langle e_{eq}^2 \rangle^{1/2} \sim 10^{-1} - 10^{-2}$ for the former, while $\langle e_{eq}^2 \rangle^{1/2} \sim 10^{-2} - 10^{-3}$ for the latter, given that $\alpha_{eff} \approx 10^{-2} - 10^{-4}$ (e.g., Ida et al. 2008; Gressel et al. 2012). If this is the case,

$$h_{pl} \simeq a \langle e_{eq}^2 \rangle^{1/2}/2,$$

under the assumption that the planetesimals’ random motions reach equipartition between eccentricity and inclination. Note that $h_{pl}$ depends both on $f_d$ and on $m_{pl}$ (see Equation (10)). Then, we can find out the minimum value of magnetic fields ($B_{min}$) above which pebbles are stirred enough to make their accretion efficient by equating $h_{ch}$ with $h_{pl}$ (using Equation (11)):

$$\langle B \rangle \gtrsim (2\pi (1 + \gamma_{turb}))^{1/4} \sqrt{\rho_{ch} h_{g} \theta^2 (h_{pl}/h_g)^2} / \left( 1 - (h_{pl}/h_g)^2 \right)^{1/4},$$

$$\equiv \langle B \rangle_{min}.$$

Thus, we can compare $\langle B \rangle_{min}$ with that derived from the chondrite (see Section 2.5).

Figure 2 shows the results of $\langle B \rangle_{min}$ as a function of both $f_d$ and $m_{pl}$. The calculations are done at $a = 2$ au, because this is where the main contribution to form chondrules occurs (see Section 2.2). We find that $\langle B \rangle_{min}$ increases with increasing $f_d$.

This trend can be understood as the following. As a disk becomes more massive (a higher value of $f_d$), the value of $M_{\text{iso}}$ also becomes higher. This in turn pumps up the planetesimals’
eccentricities ($\langle e^2 \rangle$)) and their scale heights ($h_{pl}$). Consequently, a higher value of $\langle B \rangle_{\text{min}}$ is needed to fulfill the condition that $h_{ch} = h_{pl}$. For the dependence of $m_{pl}$ on $\langle B \rangle_{\text{min}}$ (see each panel), the results show that when massive planetesimals are considered, the required value of $\langle B \rangle_{\text{min}}$ becomes higher for a given value of $f_d$. This is simply because $h_{pl}$ is an increasing function of $m_{pl}$; the eccentricity damping by the disk gas becomes weaker for massive planetesimals (see Equation (10)). As a result, $\langle B \rangle_{\text{min}}$ should be stronger in order to expand the pebble sea for the vertical direction to catch up with the increment of $h_{pl}$ by increasing $m_{pl}$.

Thus, the coupling of pebble accretion with the dynamics of planetesimals that is regulated by growing protoplanets enables an estimate of magnetic fields that are needed to achieve efficient pebble accretion in the system.

2.5. New Constraints on $f_d$ and $m_{pl}$

We are now in a position to derive new constraints on the solar nebula and parent bodies of chondrites, under the assumption that impact jetting and subsequent pebble accretion play a dominant role in forming the bodies.

Based on the discussion in Section 2.2, a lower limit of $f_d$ is obtained. Above this limit both the abundance of chondrules and their formation timescales are consistent with the meteoritic data (see the vertical dotted line in Figure 1). An upper limit of $f_d$ comes from comparing $\langle B \rangle_{\text{min}}$ with the magnetic fields inferred from Semarkona (see the vertical dashed line in Figure 2; also see Section 2.4). In the end, the most likely values of $f_d$ can be specified as a function of $m_{pl}$ (see the hatched region in Figure 2). It is important that the hatched region shrinks as $m_{pl}$ increases, and it eventually disappears for the case in which $m_{pl} > 10^{24}$ g.

Our results therefore demonstrate that the combination of impact jetting with pebble accretion can satisfy all the currently available meteoritic data when the solar nebula is $\lesssim 5$ times more massive than the MMSN at $a \simeq 2$–$3$ au and parent bodies of chondrite are $\lesssim 10^{25}$ g in the solar nebula.

3. DISCUSSION

So far we have assumed that chondrules generated from forming protoplanets via impact jetting are accreted by the surrounding field planetesimals as long as $h_{pl} < h_{ch}$. What happens if the condition $h_{pl} < h_{ch}$ is not satisfied? For this case, it is expected that the majority of chondrules will be accreted by the protoplanets and/or the central star. In the current asteroid belt, chondrules are apparently ubiquitous. In addition, protoplanets are too large to be disrupted to form asteroids. Furthermore, it is obvious that smaller planetesimals collectively have a larger surface area than more massive bodies (Johnson et al., 2015). Thus, it is necessary that (at least) some fractions of chondrules should be accreted by small planetesimals.

Our results can be compatible with several previous findings. Morbidelli et al. (2009) suggest that the planetesimals originally at the asteroid belt’s location were at least $50$ km in radius. This planetesimal radius or even a higher value (up to $500$ km in radius) is needed to reproduce a bump observed at the planetesimal radius of $\sim 50$ km in the size distribution of the current asteroid belt. In addition, Johansen et al. (2015) have recently shown that planetesimals with a characteristic radius of $\sim 100$ km can experience efficient pebble accretion within $\sim 3 \times 10^6$ years. Since planetesimals that are smaller than $50$–$100$ km in radius can be generated as fragments of larger bodies (e.g., Morbidelli et al. 2009; DeMeo et al. 2015), their results may provide a lower limit of parent bodies’ mass. If this is the case, coupling our results with theirs enables the identification of a certain range of $m_{pl}$ for primordial asteroids. Furthermore, a higher value of $f_d(>1)$ is consistent with the finding derived from the standard core accretion picture applied to the solar system (e.g., Pollack et al. 1996; Hubickyj et al. 2005), as well as the populations of exoplanets (e.g., Thommes et al. 2008; Hasegawa & Pudritz 2013); more massive disks than the MMSN are required to form Jupiter and Saturn.

Another important implication derived from our results may be the mass segregation of parent bodies in terms of chondrule inclusion. In our scenario, the presence of a first generation of planetesimals is needed both for impact jetting and for pebble accretion. Then, the resultant constraint on parent bodies’ size may imply that chondrules might have had less of a chance to be accreted by primordial asteroids that were $> 500$ km in radius, while smaller primordial asteroids likely had a surface layer in which chondrules were abundant. It is also interesting that the new constraint on $m_{pl}(\leq 10^{24}$ g) is roughly comparable to the mass of Vesta/Ceres. A careful examination of Vesta/Ceres may lead to further clues about our scenario.

It is clear that a more detailed investigation is needed to verify the picture discussed here. For instance, our scenario requires a first generation of planetesimals as discussed above. Possible mechanisms to form such planetesimals (i.e., streaming instabilities, Johansen et al. 2007) should be further explored in this context. Also, detailed studies of how the resultant chondrules can be accreted onto planetesimals under the presence of a protoplanet should be undertaken. One of the key steps may be to perform numerical simulations that treat the random torque on protoplanets from turbulent disks, which may affect how and where chondrules form and accrete.

Finally, our current setup is local, that is, it is isolated entirely from its surrounding. It is important to examine how the picture developed here can be affected by the presence of nearby planets such as Jupiter. These kind of global simulations may show how our scenario could favor or disfavor the Grand-tack scenario (Walsh et al., 2011), which proposes that a mixture of volatile-rich and volatile-poor asteroids in the current asteroid belt may have originated from gas-induced migration of Jupiter and Saturn.

In the next few years, potential sample-return missions and laboratory studies should yield a broader and deeper view of the history recorded in primitive asteroids. This will enable tests of the ideas that chondrules formed through impact jetting, cooled in the asteroid belt’s ambient magnetic field, and were quickly swept up by planetesimals with the aid of gas drag forces.

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