IRON 60 EVIDENCE FOR EARLY INJECTION AND EFFICIENT MIXING
OF STELLAR DEBRIS IN THE PROTOSOLAR NEBULA

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ABSTRACT

Among extinct radioactivities present in meteorites, 60Fe (t1/2 = 1.49 Myr) plays a key role as a high-resolution chronometer, a heat source in planetesimals, and a fingerprint of the astrophysical setting of solar system formation. A critical issue with 60Fe is that it could have been heterogeneously distributed in the protoplanetary disk, calling into question the efficiency of mixing in the solar nebula or the timing of 60Fe injection relative to planetesimal formation. If this were the case, one would expect meteorites that did not incorporate 60Fe (either because of late injection or incomplete mixing) to show 60Ni deficits (from lack of 60Fe decay) and collateral effects on other neutron-rich isotopes of Fe and Ni (coproduced with 60Fe in core-collapse supernovae and AGB stars). Here, we show that measured iron meteorites and chondrites have Fe and Ni isotopic compositions identical to Earth. This demonstrates that 60Fe must have been injected into the protosolar nebula and mixed to less than 10% heterogeneity before formation of planetary bodies.

Subject headings: methods: analytical — nuclear reactions, nucleosynthesis, abundances — solar system: formation — supernovae: general

Online material: color figures

1. INTRODUCTION

Evidence for the presence of live 60Fe has been found in meteoritic materials in the form of its decay product, 60Ni (Birck & Lugmair 1988; Shukolyukov & Lugmair 1993). Recent in situ analyses of minerals from chondrites with high Fe/Ni ratios have improved our knowledge of 60Fe abundance in the early solar system (ESS; Tachibana & Huss 2003; Mostefaoui et al. 2005; Tachibana et al. 2006; Guan et al. 2007). In particular, chondrule pyroxenes show 60Ni excesses corresponding to 60Fe/56Fe ratios of (2.2–3.7) × 10−7 at the time of crystallization (Tachibana et al. 2006). This translates into an initial ratio of (5–10) × 10−7 at the time of condensation of the first solids, calcium-aluminum-rich inclusions (CAIs), in the nebula (assuming a time interval of 1.5–2.0 Myr between formation of chondrules and CAIs; Amelin et al. 2002; Kita et al. 2005). A critical question that remains to be answered is whether or not 60Fe was homogeneously distributed in the ESS. Iron meteorites can help us address this issue because they may sample a different portion of the protoplanetary nebula than that sampled by chondrites (Bottke et al. 2006). In addition, 182Hf (t1/2 = 8.9 Myr)182W chronology indicates that the parent bodies of iron meteorites accreted within ~2 Myr of formation of CAIs (Markowski et al. 2006; Schersten et al. 2006; Qin et al. 2008), when 60Fe would have been extinct if it had been present. Several groups have analyzed Ni in iron meteorites and have found normal (terrestrial) isotopic compositions, within uncertainties (Cook et al. 2006; Quitte´ et al. 2006; Chen et al. 2007; Moyner et al. 2007). However, with improved precision, Bizzarro et al. (2007) reported deficits of 25 ppm (parts per million) in 60Ni in several classes of early-formed differentiated meteorites (including iron meteorites); these deficits were not correlated with Fe/Ni ratios. They attributed the uniform deficits in 60Ni to the accretion of the parent bodies of these differentiated meteorites prior to the injection of 60Fe into the ESS.

In order to address the question of 60Fe distribution in the solar nebula, we have analyzed Fe and Ni isotopic compositions of meteoritic metal. The paper is organized as follows. In §2, we describe the analytical methods used for measuring the isotopic compositions of Fe and Ni (Tables 1, 2, and 3). In §3, we show how inefficient mixing or delayed injection of 60Fe in the solar nebula should be associated with 60Ni deficits (from lack of 60Fe decay) and collateral anomalies in the neutron-rich isotopes 58Fe and 64Ni (coproduced with 60Fe in core-collapse supernovae and AGB stars by neutron-capture reactions). Comparisons between predictions and measurements limit the heterogeneity in the distribution of 60Fe. In §4, we set an upper limit on the abundance of 60Fe at the time of metal-silicate differentiation in planetesimals. Finally, we discuss and summarize the implications of our results in §5.

2. MATERIALS AND METHODS

For Fe analyses, we used aliquots of the samples dissolved by Cook et al. (2006). The protocol used for separation of Fe for the search of isotopic anomalies is described in Dauphas et al. (2004a) and Dauphas & Rouxel (2006). The solutions of purified Fe were analyzed on a Neptune (thermo electron) multicollector inductively coupled plasma mass spectrometer (MC-ICPMS) at the University of Chicago using Al cones. Isotopes 54Fe, 56Fe, 57Fe, and 58Fe were analyzed on Faraday collectors L2, Ax, H1,
and H2. Isobaric interferences from 54Cr and 58Ni were corrected for by monitoring 53Cr and 60Ni on L4 and H4. The corrections from 58Ni on 58Fe ranged from \( \sim 5 \) to 20 \( \varepsilon \). Isotopes 58Ni and 66Zn were analyzed with 1010 and 1012\( \Omega \) amplifiers, respectively. High resolution was used to resolve interferences from 40Ar\( ^{16}O\) and 40Ar\( ^{18}O\). Zoom optics were used to accommodate the mass dispersion (mass 57–66) in the collector array. OPZ intensities from blank solutions were subtracted from all measurements. The reference material used for sample-standard bracketing was SRM-986 (Gramlich et al. 1989). Each sample was analyzed 8–33 times, and the average compositions are reported in \( \epsilon \) units, where

\[
\epsilon = \frac{\frac{N_i}{N_{01}}}{\frac{N_i}{N_{01}}_{\text{SRM-986}}} - 1 \times 10^4
\]

Note.—Uncertainties are 95% confidence intervals.

### Table 1
Nickel Isotopic Compositions of Meteoritic Metal (\( \epsilon^{58}\text{Ni}/^{60}\text{Ni} \) normalized to \( ^{65}\text{Ni}/^{64}\text{Ni} = 0.0533886 \))

| Sample (Type) | Number | \( \epsilon^{58}\text{Ni} \) | \( \epsilon^{56}\text{Ni} \) | \( \epsilon^{60}\text{Ni} \) | \( \epsilon^{58}\text{Ni} \) | \( \epsilon^{60}\text{Ni} \) | \( \epsilon^{64}\text{Ni} \) |
|---------------|--------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| Coahuila (IIAB) | 15     | 0.006 ± 0.031 | 0.011 ± 0.093 | 0.011 ± 0.093 | 0.0317 ± 0.200 |
| Santa Luzia (IIAB) | 8      | -0.047 ± 0.050 | -0.046 ± 0.211 | -0.046 ± 0.211 | 0.3017 ± 0.200 |
| Casas Grandes (IIIAB) | 30     | -0.030 ± 0.022 | 0.053 ± 0.081 | 0.053 ± 0.081 | 0.023 ± 0.085 |
| Henbury (IIIB) | 17     | 0.003 ± 0.023 | 0.029 ± 0.053 | 0.029 ± 0.053 | 0.053 ± 0.131 |
| Molong (PMG) | 33     | 0.002 ± 0.031 | 0.048 ± 0.076 | 0.048 ± 0.076 | 0.019 ± 0.134 |
| Bishunpur (LL3.1) | 14     | 0.022 ± 0.031 | 0.215 ± 0.143 | 0.215 ± 0.143 | 0.095 ± 0.233 |

### Table 2
Nickel Isotopic Compositions of Meteoritic Metal (\( \epsilon^{58}\text{Ni}/^{60}\text{Ni} \) normalized to \( ^{61}\text{Ni}/^{64}\text{Ni} = 0.0167442 \))

| Sample (Type) | Number | \( \epsilon^{58}\text{Ni} \) | \( \epsilon^{56}\text{Ni} \) | \( \epsilon^{60}\text{Ni} \) | \( \epsilon^{61}\text{Ni} \) | \( \epsilon^{62}\text{Ni} \) | \( \epsilon^{64}\text{Ni} \) |
|---------------|--------|-----------------|-----------------|-----------------|-----------------|-----------------|-----------------|
| Coahuila (IIAB) | 15     | -0.001 ± 0.061 | 0.014 ± 0.123 | 0.014 ± 0.123 | -0.124 ± 0.269 |
| Santa Luzia (IIAB) | 8      | -0.016 ± 0.123 | 0.061 ± 0.279 | 0.061 ± 0.279 | 0.407 ± 0.325 |
| Casas Grandes (IIIAB) | 30     | -0.057 ± 0.050 | 0.070 ± 0.107 | 0.070 ± 0.107 | -0.125 ± 0.141 |
| Henbury (IIIB) | 17     | -0.023 ± 0.032 | 0.038 ± 0.070 | 0.038 ± 0.070 | -0.003 ± 0.151 |
| Molong (PMG) | 33     | -0.019 ± 0.057 | 0.063 ± 0.100 | 0.063 ± 0.100 | -0.126 ± 0.194 |
| Bishunpur (LL3.1) | 14     | -0.121 ± 0.092 | -0.284 ± 0.189 | -0.284 ± 0.189 | -0.324 ± 0.287 |

Note.—Uncertainties are 95% confidence intervals.
(cc-SN) or asymptotic giant branch star (AGB) shortly before formation of the solar system (Cameron & Truran 1977; Wasserburg et al. 1998, 2006; Meyer & Clayton 2000; McKeegan & Davis 2007). Interestingly, a cc-SN explosion or passing AGB star could have triggered the collapse of the molecular cloud core that made our solar system (Cameron & Truran 1977; Cameron 1993; Foster & Boss 1996; Vanhala & Boss 2000). The probability of encounters between AGB stars and molecular clouds is small, whereas cc-SNs are associated with star-forming regions. Thus, cc-SNs are more likely candidates for explaining the presence of $^{60}$Fe in meteorites than AGBs (Kastner & Myers 1994). An alternative to the triggered molecular cloud core collapse hypothesis is direct injection of $^{60}$Fe into the protoplanetary disk (Chevalier 2000; Hester & Desch 2005; Ouellette et al. 2007). The likelihood of such an injection is <1% (Williams & Gaidos 2006), then one would expect to find $^{60}$Ni deficits in a hypothetical reservoir free of $^{60}$Fe relative to the chondritic uniform reservoir (CHUR, see Appendix A for details),

$$
\varepsilon^{60}\text{Ni}_0 = -Q\left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{CHUR}},
$$

(1)

### 3.1. $^{60}$Ni Deficits?

All Ni measurements have been corrected for mass-dependent fractionation by internal normalization to the terrestrial $^{62}$Ni/$^{58}$Ni ratio. Accordingly, terrestrial rocks related by the laws of mass-dependent fractionation should have $\varepsilon = 0$ for the $^{60}$Ni/$^{58}$Ni, $^{61}$Ni/$^{58}$Ni, and $^{64}$Ni/$^{58}$Ni ratios. As shown in Figure 2, we found normal (terrestrial) $^{60}$Ni isotopic compositions in all analyzed meteorites (weighted average $\varepsilon^{60}$Ni = 0.006 ± 0.011). This conclusion agrees with previous studies (Cook et al. 2006; Quitte et al. 2006; Chen et al. 2007; Moynier et al. 2007), but with significantly improved precision. The discrepancy with Bizzarro et al. (2007) is not due to differences in sampling (meteorites of the same type were analyzed in Bizzarro et al. [2007] and this study) or the choice of the ratio used for correcting mass-dependent fractionation (Fig. 1). If $^{60}$Fe were heterogeneously distributed in the solar nebula, one would expect to see $^{60}$Ni deficits in a hypothetical reservoir free of $^{60}$Fe relative to the chondritic uniform reservoir (CHUR, see Appendix A for details),

![Fig. 1 – Comparison between Ni data from this study (Table 2) and Bizzarro et al. (2007) for differentiated meteorites. Identical notation (Birck & Lugmair 1988; Bizzarro et al. 2007; $\varepsilon^{58}$Ni/$^{58}$Ni) and correction for mass-dependent fractionation (internal normalization to a fixed $^{61}$Ni/$^{58}$Ni ratio of 0.0167442 using the exponential law; Gramlich et al. 1989; Marechal et al. 1999) were used. Because of its low abundance (0.91 atomic percent) and the presence of a major $^{64}$Zn interference (48.63 atomic percent), $^{64}$Ni data were not reported in Bizzarro et al. (2007). [See the electronic edition of the Journal for a color version of this figure].](image)

![Fig. 2 — Fe and Ni isotope data for meteoritic metal (Tables 1 and 3). The compositions are in $\varepsilon$ units corrected for mass-dependent fractionation by fixing the ratio $^{56}$Fe/$^{56}$Fe = 0.3625663 (Taylor et al. 1992) and $^{58}$Ni/$^{58}$Ni = 0.0533863 (Graham et al. 1989) using the exponential law (Marechal et al. 1999). (a) Fe has normal (terrestrial) isotopic composition. (b) If iron meteorites did not incorporate $^{60}$Fe and the initial $^{56}$Fe/$^{56}$Fe$_{\text{CHUR}}$ ratio was $\sim 5 \times 10^{-7}$ (Tachibana et al. 2006), then one would expect to find $\varepsilon^{60}$Ni$_0 = -10^{6} (^{56}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}} \times (^{56}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}} \approx -0.31$ in iron meteorites (large cross) relative to chondrites and Earth, which is not observed. This conclusion does not depend on the pair of isotopes ($^{58}$Ni/$^{58}$Ni or $^{60}$Ni/$^{58}$Ni) that is used for correcting the measurements for mass-dependent fractionation (Tables 1 and 2, Fig. 1). [See the electronic edition of the Journal for a color version of this figure].](image)
where $Q = (^{56}\text{Fe}/^{60}\text{Ni})_{\text{CHUR}} \times 10^4 = 61.35 \times 10^4$ (Lodders 2003) and $t$ is the time of CAI formation. Here it is assumed that the terrestrial reference material has the same Ni isotopic composition as CHUR ($^{60}\text{Fe}$ was present in the two reservoirs at a similar level). For $(^{60}\text{Fe}/^{56}\text{Ni})_{t,\text{CHUR}} = 5 \times 10^{-7}$ (Tachibana et al. 2006), the predicted effect is $\epsilon^{60}\text{Ni} = -0.31$, nearly the same as the value reported in differentiated meteorites by Bizzarro et al. (2007).

Our high-precision measurements rule out the presence of such a deficit (Figs. 1 and 2). The isotopic composition of a hypothetical reservoir containing $^{60}\text{Fe}$ at a lower or higher level than CHUR (denoted with a subscript) is given by $Q(^{60}\text{Fe}/^{56}\text{Ni})_{t,\text{CHUR}}$ (see Appendix A, eq. [A8]). Using $(^{60}\text{Fe}/^{56}\text{Ni})_{t,\text{CHUR}} = 5 \times 10^{-7}$, a reservoir with an initial ratio of $>5.8 \times 10^{-7}$, would have $\epsilon^{60}\text{Ni} > +0.05$. Similarly, a reservoir with an initial ratio of $<4.2 \times 10^{-7}$ would have $\epsilon^{60}\text{Ni} < -0.05$. Thus, the $\pm 0.05$ dispersion in $^{60}\text{Fe}$ around zero only allows for a heterogeneity of less than $\sim 15\%$ in the initial $(^{60}\text{Fe}/^{56}\text{Ni})$ ratio (if the ratio is $5 \times 10^{-7}$; a higher initial value would allow even less heterogeneity). This observation is consistent with the degree of homogeneity inferred for $^{26}\text{Al}$ in the accretion region of meteorite parent bodies (McKeegan & Davis 2007), as well as results of dynamical modeling, which show that passive tracers are mixed in the nebula at the 10% level on timescales of a few thousand years (Boss 2007).

3.2. Collateral Nucleosynthetic Effects on $^{58}\text{Fe}$ and $^{64}\text{Ni}$?

A powerful means of investigating whether $^{60}\text{Fe}$ was heterogeneously distributed in the ESS is to search for collateral anomalies in other isotopes of Fe (Nichols et al. 1999; Sahijpal & Soni 2006; Gounelle & Meibom 2007). Solar system Fe was produced by different nucleosynthetic processes (nuclear statistical equilibrium and neutron capture) in different stars ($\sim 1/2$ comes from cc-SN and the rest from SN Ia; Iwamoto et al. 1999). Its isotopic composition represents the integrated contribution over $~9$ Gyr of SN Ia and cc-SN of different metallicities and initial masses. Thus, Fe isotopes are unlikely to be present in exact solar proportions in the cc-SN or AGB star that injected $^{60}\text{Fe}$ into the nascent solar system. Addition of even minute amounts of stellar ejecta containing $^{60}\text{Fe}$ should leave a diagnostic fingerprint on nonradioactive Fe isotopes if it is missing from other parts of the solar system (as is suggested by Sugiyama et al. 2006; Bizzarro et al. 2007; Quitte et al. 2007). The virtue of this approach is that no chemical fractionation can decouple $^{60}\text{Fe}$ from stable Fe isotopes (whether $^{60}\text{Fe}$ is injected in the form of dust or gas, Ouellette et al. 2007) is inconsequential.

The bulk ejecta of a cc-SN could not have been injected into the ESS because such a scenario would overproduce $^{41}\text{Ca}$, $^{53}\text{Mn}$, and $^{60}\text{Fe}$ relative to $^{26}\text{Al}$ (Meyer & Clayton 2000). A possible solution to this problem is to invoke an injection mass cut different from the mass cut of the remnant (Cameron et al. 1995; Meyer & Clayton 2000). The meteoritic abundances of $^{26}\text{Al}$, $^{41}\text{Ca}$, and $^{60}\text{Fe}$ can be explained with a $25\%$ (solar mass) cc-SN using a free decay interval of $~1$ Myr and an injection mass cut of $~7 M_\odot$ that corresponds to the edge of the He-exhausted core (Meyer & Clayton 2000). In a cc-SN, Fe isotopes are produced in different layers by different nuclear reactions. Iron 54, 56, and 57 are produced as radioactive progenitors (e.g., $^{56}\text{Ni}$) in the inner region of the cc-SN by nuclear statistical equilibrium associated with explosive Si burning. This process results in synthesis of isotopes with approximately equal numbers of neutrons and protons. For this reason, $^{58}\text{Fe}$ and $^{60}\text{Fe}$, which have large excesses of neutrons and are shielded by stable $^{58}\text{Ni}$ and $^{60}\text{Ni}$, are not produced in appreciable quantities. These two isotopes are synthesized in more external regions by neutron capture reactions on preexisting Fe isotopes (e.g., Fig. 3). Using an injection mass cut similar to that used in Meyer & Clayton (2000), one would predict that $^{58}\text{Fe}/^{54}\text{Fe}$ and $^{57}\text{Fe}/^{54}\text{Fe}$ ratios should be similar to solar values because the inventory of these isotopes in the ejecta is dominated by unprocessed material from the envelope. However, one would also predict that the $^{58}\text{Fe}/^{54}\text{Fe}$ ratio should be higher than solar because $^{60}\text{Fe}$ is accompanied by $^{57}\text{Fe}$. Therefore, any portion of the solar system deficient in $^{60}\text{Fe}$ should also exhibit negative anomalies in $^{58}\text{Fe}$. The $^{58}\text{Fe}$ isotopic composition of a reservoir missing the cc-SN or AGB component relative to CHUR is (in $t$ with internal normalization, although in this case the isotopic variations are nucleosynthetic in origin and are not due to mass fractionation)

$$
\epsilon^{58}\text{Fe}_0 = -\frac{\rho_{58} - \mu_{58}/\mu_{54}}{\rho_{58} + \mu_{58}/\mu_{54}} \left(\frac{^{56}\text{Fe}/^{58}\text{Fe}}{^{58}\text{Fe}/^{56}\text{Fe}}\right)_{\text{ccSN}/\text{AGB}} \times 10^4,$$

where $\rho_{58} = (^{58}\text{Fe}/^{60}\text{Ni})_{\text{ccSN}/\text{AGB}}/(^{58}\text{Fe}/^{56}\text{Fe})_{\text{ccSN}/\text{AGB}} - 1$, $\mu_{58} = (58 - 54)/(57 - 54)$, $\Delta t$ is the free decay interval between production in the cc-SN or AGB and injection into the solar system (the derivation of the general equation is given in Appendix B, also see Dauphas et al. 2004b).

If one allows for $\Delta t > 0$, the sizes of the collateral effects increase because $^{60}\text{Fe}$ has time to decay before being injected and the contribution of material from the cc-SN or AGB must increase in order to explain $(^{60}\text{Fe}/^{58}\text{Fe})_{t,\text{CHUR}}$ which is fixed. In this study, we use $\Delta t = 0$ and present collateral isotopic effects that must therefore be treated as lower limits. In Figure 4a, we have computed $^{58}\text{Fe}$ deficits using yields from 15, 19, and 25 $M_\odot$ cc-SN models as a function of injection mass cut (Rauscher et al. 2002). Note that the $^{60}\text{Fe}$ yields in these models are thought to be overestimated (Limongi & Chieffi 2006; Woosley & Heger 2007) and the collateral effects on $^{58}\text{Fe}$ may be larger than what is calculated here (the $^{60}\text{Fe}/^{58}\text{Fe}$ production ratio appears in the denominator of eq. [2]). However, the combined uncertainties in the neutron capture reactions governing the synthesis of $^{60}\text{Fe}$ are not expected to reduce the production of $^{60}\text{Fe}$ by more than a factor of 2 (Woosley & Heger 2007). The predicted deficits on
make the assumption that Fe and Ni from the cc-SN are not decoupled during injection into the ESS. In the 15, 19, and 25 $M_\odot$ cc-SN models, $^{60}$Fe is produced in layers where C/O $< 1$ and Fe would presumably condense as metal (Lattimer et al. 1978). Under those circumstances, Ni is unlikely to be decoupled from Fe. The predicted collateral $^{64}$Ni isotope effects are given by

$$
e^{64}{Ni} = -c \frac{\rho_{Ni}^{64} - \rho_{Ni}^{56}}{1 + \rho_{Fe}^{60}} \frac{(^{60}Fe/^{56}Fe)_{CHUR}}{(^{60}Fe/^{56}Fe)_{ccSN/AGB}} e^{\Delta t} \times 10^4,$$

(3)

where $\rho_{Ni}^{64} = (^{64}Ni/^{58}Ni)_{CHUR} / (^{64}Ni/^{58}Ni)_{CCSN/AGB} - 1$, $\rho_{Ni}^{64} = (64 - 58)/(62 - 58)$, and $c = (^{58}Ni/^{56}Ni)_{CCSN/AGB} / (^{58}Ni/^{56}Ni)_{CHUR}$ (the derivation of the general equation is given in Appendix C; also see Dauphas et al. 2004b). If some parts of the solar system did not incorporate $^{60}$Fe from a cc-SN, then one would predict $^{64}$Ni of $< 2$ or lower (Fig. 4b). This contradicts measurements, which show normal $^{64}$Ni within 0.2. This limits the possible heterogeneity of the initial ($^{60}Fe/^{56}Fe)_{ccSN}$ ratio to within 10% (if the value is $5 \times 10^{-7}$; see Appendix C, eq. [C5]).

Although less likely than cc-SN (Kastner & Myers 1994; Williams & Gaidos 2007), AGB stars represent an alternative source for $^{26}$Al and $^{60}$Fe in the ESS (Cameron 1993; Wasserburg et al. 1998, 2006). In AGBs, neutrons produced by the $^{22}$Ne($\alpha$, n)$^{25}$Mg reaction can be captured on $^{54}$Fe, $^{56}$Fe, and $^{58}$Fe to produce $^{58}$Fe and $^{60}$Fe. The predicted $^{60}$Ni deficit (eq. [2]) for a 3 $M_\odot$ initial mass star with standard $^{13}$C pocket (Wasserburg et al. 2006) is $\sim 133$ (for an initial $^{60}$Fe/$^{56}$Fe = $5 \times 10^{-7}$). Again, no such deficits were found ($^{58}$Fe = $0 \pm 0.3$). Similarly, the predicted $^{60}$Ni deficit of $-125$ is not observed in meteorites ($^{60}$Ni = $0 \pm 0.2$). Thus, if the source of $^{60}$Fe in the ESS was a nearby AGB star, limits on the possible heterogeneous distribution of $^{60}$Fe in the protosolar nebula are even more stringent ($\pm 0.2$%) than in the case of a cc-SN.

Analyses of Fe and Ni stable isotopic compositions of meteorites limit the heterogeneity in the distribution of $^{60}$Fe to less than $\sim 10\%$. This is consistent with contemporaneous injection of $^{60}$Fe and $^{26}$Al followed by efficient mixing in the protosolar disk (Boss 2007). It is worth noting that the scale relevant to this study is that of planetesimals. Stable isotopic anomalies produced by neutron-rich nuclear statistical equilibrium (Hartmann et al. 1985) have been detected in CAIs for Fe and Ni (Birck & Lugmair 1988; Völkening & Papanastassiou 1989) and one cannot exclude that $^{60}$Fe was heterogeneously distributed at such scale (subcentimeter objects).

4. UPPER LIMIT ON $^{60}$Fe ABUNDANCE

Having established that $^{60}$Fe was homogeneously distributed in the ESS, we can now turn to estimating its abundance. Evidence from $^{18}$F-W systematics indicates that separation between metal and silicate occurred within 2 Myr of condensation of the first solids in the nebula (Markwick et al. 2006; Schersten et al. 2006; Qin et al. 2008). In Figure 5, we plot $^{60}$Ni (using $^{62}$Ni/$^{58}$Ni normalization) as a function of the Fe/Ni fractionation factor for several types of samples. For Bishunpur metal, we use the Fe/Ni ratio measured by Cook et al. (2006); $f_{Fe/Ni} = (Fe/Ni)(Fe/Ni)_{CHUR} - 1 = -0.221$. Part of the fractionation between Fe and Ni in iron meteorites was produced during fractional crystallization of the metallic core, which most likely took place after $^{60}$Fe had decayed. Under those circumstances, the
relevant Fe/Ni fractionation factor is that of the parental melt, established when metallic cores segregated from silicate mantles. We therefore plot the average \( \varepsilon^{60}\text{Ni} \) for each magmatic iron meteorite group (Table 1) as a function of the fractionation factor of the parent melt \( f_{\text{Fe/Ni}} \) (Tachibana et al. 2006). If \( \varepsilon^{60}\text{Ni} \) was present when metal segregated from silicate (taken to represent a single event precisely defined in time), there should be a linear correlation between \( \varepsilon^{60}\text{Ni} \) and the Fe/Ni fractionation factor relative to chondrites (Jacobsen & Wasserburg 1984; see Appendix D for details),

\[
\varepsilon^{60}\text{Ni} = \mathcal{Q} \left( \frac{^{60}\text{Fe}}{^{56}\text{Fe}} \right) t_{\text{Fe/Ni}},
\]

where \( \mathcal{Q} \) has the same definition as in equation (1) and \( t_{\text{Fe/Ni}} \) is the time of Fe/Ni fractionation. Bishunpur metal and IIIAB irons provide the most stringent constraints on the \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio at the time of metal-silicate differentiation, which is conservatively estimated to be \( <2 \times 10^{-7} \) in both cases. This value is in agreement with the \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio measured in Bishunpur by in situ techniques (Tachibana & Huss 2003; Tachibana et al. 2007; 1.08 ± 0.23 \( \times 10^{-7} \) in sulfides and 1.9 \( \pm 1.1 \times 10^{-7} \) in chondrules) and the value inferred from pallasite and magmatic iron meteorite measurements (Cook et al. 2006; marginally different from 0, \( \sim 1 \) to \( 8 \times 10^{-7} \)).

The \( ^{182}\text{Hf}/^{184}\text{Hf} \) date of core formation in the parent body of IIIAB iron meteorites is \( \sim 1.6 \) Myr after CAI formation (Markowski et al. 2006a; Schersten et al. 2006; Qin et al. 2008). No precise \( ^{182}\text{Hf}/^{184}\text{Hf} \) age of metal in LL chondrites is available (Yin et al. 2002), but Kleine et al. (2008) showed that in primitive H chondrites, metal-silicate differentiation occurred \( 1.7 \pm 0.7 \) Myr after CAI and was coeval with chondrule formation (Amelin et al. 2002; Kita et al. 2005). If we tentatively correct for a possible delay of up to 2 Myr between CAI formation and metal-silicate differentiation, the \( \varepsilon^{60}\text{Ni} \) values measured in meteoritic metal translate into an upper limit of \( <6 \times 10^{-7} \) for the initial \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio at the time of condensation of the first solids in the solar nebula. This result is consistent with the initial value of \( (5-10) \times 10^{-7} \) derived from in situ isotopic analysis of chondrule pyroxenes (Tachibana et al. 2006).

5. SUMMARY AND CONCLUSION

The presence of \( ^{60}\text{Fe} \) in meteorites can only be explained by injection into the solar nebula of debris from a nearby AGB or cc-SN. Homogeneity of short-lived isotopes in the solar nebula is one of the key unsettled questions at the present time. In order to address this question, we have measured the isotopic compositions of Fe and Ni in meteoritic metal. The data have clear implications.

1. If \( ^{60}\text{Fe} \) was not well mixed or if it was injected late into the solar nebula, one would expect the meteorites that did not incorporate \( ^{60}\text{Fe} \) to show \( ^{60}\text{Ni} \) deficits (from lack of \( ^{60}\text{Fe} \) decay). Our measurements show that within uncertainties, the isotopic abundance of \( ^{60}\text{Ni} \) in metal from IIIAB, IIAB, PMG, and LL meteorites is constant, consistent with homogeneous distribution of \( ^{60}\text{Fe} \) in the solar nebula (less than 15% heterogeneity in the \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio).

2. Iron \( 60 \) is synthesized in cc-SN and AGB stars by neutron-capture reactions. If some parts of the solar system did not incorporate \( ^{60}\text{Fe} \), they should show deficits in the abundances of other isotopes also produced by neutron-capture, \( ^{58}\text{Fe} \) and \( ^{64}\text{Ni} \). These isotopes have very low isotopic abundances and are difficult to analyze. Our study reveals that the abundances of \( ^{58}\text{Fe} \) and \( ^{64}\text{Ni} \) in meteorites are identical within uncertainties to abundances measured in terrestrial materials. Comparisons with model predictions limit the possible heterogeneous distribution of \( ^{60}\text{Fe} \) to less than 10%.

3. Metal from IIAB, IIIAB, and LL meteorites have fractionated Fe/Ni ratios relative to CHUR. If Fe/Ni fractionation occurred in the presence of \( ^{60}\text{Fe} \), one would expect to find isotopic variations in the abundance of \( ^{60}\text{Ni} \), possibly correlated with Fe/Ni ratios. No correlation was found, providing an upper limit on the initial \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio in the ESS of around \( <6 \times 10^{-7} \). This upper limit agrees with recent in situ isotopic analyses of chondrule pyroxenes, which give an initial ratio of \( (5-10) \times 10^{-7} \) (Tachibana et al. 2006).

The degree of mixing of \( ^{60}\text{Fe} \) in the ESS is quantified for the first time in this study. We conclude that \( ^{60}\text{Fe} \) must have been homogeneously distributed (less than 10% heterogeneity in the \( ^{60}\text{Fe}/^{56}\text{Fe} \) ratio). This agrees with results of dynamical modeling (Boss 2007), showing that passive tracers are mixed in a turbulent nebula at the 10% level within several thousand years. A second important conclusion is that \( ^{60}\text{Fe} \) was injected before formation of planetary bodies. This supports the idea that the presolar molecular cloud core was triggered into collapse by interaction with a nearby star (Cameron & Truran 1977; Cameron 1993). Iron \( 60 \) was homogeneously distributed in the solar nebula and can therefore be used as a reliable short-lived chronometer for ESS events. This is particularly important given the converging lines of evidence indicating that the key events that shaped our solar system took place shortly after condensation of the first solids in the solar nebula.
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APPENDIX A

DERIVATION OF EQUATION (1)

Over time, all initial $^{60}\text{Fe}$ present in CHUR was converted into $^{60}\text{Ni}$,

$$^{60}\text{Ni}_{\text{CHUR}} = ^{60}\text{Ni}_0 + ^{60}\text{Fe}_{0,\text{CHUR}}. \quad (A1)$$

Introducing isotopic ratios into this equation,

$$\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_{\text{CHUR}} = \left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_0 + \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{CHUR}} \left(\frac{^{56}\text{Fe}}{^{58}\text{Ni}}\right)_{\text{CHUR}}, \quad (A2)$$

where $\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_0$ is also the Ni isotopic composition of any reservoir missing $^{60}\text{Fe}$ (either because of delayed injection of incomplete mixing). Dividing both sides by $\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_{\text{CHUR}}$ gives

$$\frac{\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_0}{\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_{\text{CHUR}}} = 1 - \left(\frac{^{56}\text{Fe}}{^{60}\text{Ni}}\right)_{\text{CHUR}} \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{CHUR}}. \quad (A3)$$

Expressed in $\epsilon$ notation relative to CHUR, this translates into

$$\epsilon^{60}\text{Ni}_0 = -10^4 \left(\frac{^{56}\text{Fe}}{^{60}\text{Ni}}\right)_{\text{CHUR}} \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{CHUR}}. \quad (A4)$$

Introducing $Q = \left(\frac{^{56}\text{Fe}}{^{60}\text{Ni}}\right)_{\text{CHUR}} \times 10^4$, this takes the form (eq. [1]),

$$\epsilon^{60}\text{Ni}_0 = -Q \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{CHUR}}. \quad (A5)$$

Let us now consider a hypothetical reservoir $r$ with an initial $\left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{r}}$ ratio different from CHUR but with chondritic Fe/Ni ratio. An equation similar to equation (A2) can be posed,

$$\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_r = \left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_0 + \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,r} \left(\frac{^{56}\text{Fe}}{^{58}\text{Ni}}\right)_{\text{CHUR}}. \quad (A6)$$

Dividing the two sides by $\left(\frac{^{60}\text{Ni}}{^{58}\text{Ni}}\right)_{\text{CHUR}}$, it is easy to show that

$$\epsilon^{60}\text{Ni}_r = \epsilon^{60}\text{Ni}_0 + Q \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,r}, \quad (A7)$$

which can also be written in the form

$$\epsilon^{60}\text{Ni}_r = Q \left[\left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,r} - \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{0,\text{CHUR}}\right]. \quad (A8)$$

This equation can be used to quantify the degree of heterogeneity in $^{60}\text{Fe}/^{56}\text{Fe}$ ratios based on $^{60}\text{Ni}$ measurements.

APPENDIX B

DERIVATION OF EQUATION (2)

Let us denote $^{60}\text{Fe}_{\text{inj}}$ the amount of $^{60}\text{Fe}$ injected into the ESS. If $\Delta t$ is the time that separates nucleosynthesis in AGB or cc-SN and injection into the ESS,

$$^{60}\text{Fe}_{\text{inj}} = ^{60}\text{Fe}_{\text{ccSN/AGB}} e^{-i\omega \Delta t}. \quad (B1)$$
The amount of $^{60}\text{Fe}$ injected is equal to the amount of $^{60}\text{Fe}$ in chondrites (back-calculated at the time of condensation of the first solids in the protosolar nebula),

$$
\left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{inj}} = \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{CHUR}}^{^{56}\text{Fe}}.
$$

If we define $x = ^{56}\text{Fe}_{\text{inj}}/^{56}\text{Fe}_{\text{CHUR}} = ^{56}\text{Fe}_{\text{ccSN/AGB}}/^{56}\text{Fe}_{\text{CHUR}}$, it follows from the previous equations that

$$
x = \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{CHUR}} \cdot e^{\Delta t}.
$$

Let us now define $y = ^{54}\text{Fe}_{\text{ccSN/AGB}}/^{54}\text{Fe}_{\text{CHUR}}$,

$$
y = \left(\frac{^{54}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{ccSN/AGB}} \cdot \frac{^{56}\text{Fe}_{\text{ccSN/AGB}}}{^{56}\text{Fe}_{\text{CHUR}}},
$$

Using the notation

$$
\rho^i_{\text{Fe}} = \left(\frac{^{i}\text{Fe}}{^{54}\text{Fe}}\right)_{\text{ccSN/AGB}} / \left(\frac{^{i}\text{Fe}}{^{54}\text{Fe}}\right)_{\text{CHUR}} - 1,
$$

equation (B4) can be rewritten in the form

$$
y = x / (1 + \rho^{56}_{\text{Fe}}).
$$

For any isotopes $^i\text{Fe}$, we can pose a conservation equation:

$$
\left(\frac{^{i}\text{Fe}}{^{54}\text{Fe}}\right)_{\text{CHUR}}^{^{54}\text{Fe}} = \left(\frac{^{i}\text{Fe}}{^{54}\text{Fe}}\right)^{^{0}}_{\text{CHUR}} + \left(\frac{^{i}\text{Fe}}{^{54}\text{Fe}}\right)_{\text{ccSN/AGB}}^{^{54}\text{Fe}}.
$$

where the 0 subscript denotes the composition of CHUR prior to $^{60}\text{Fe}$ injection. It is also the composition of any planetary or meteoritic component that did not receive $^{60}\text{Fe}$, either because it formed before $^{60}\text{Fe}$ was injected into the ESS or because it was accreted in a region of the nebula that did not incorporate $^{60}\text{Fe}$ due to incomplete mixing. Introducing $y$ and $\rho_{\text{Fe}}$ in this equation, it is straightforward to show that the isotopic composition of a reservoir missing the cc-SN or AGB component relative to CHUR is

$$
\frac{\left(\frac{^i\text{Fe}}{^{54}\text{Fe}}\right)^{^{0}}_{\text{CHUR}}}{\left(\frac{^i\text{Fe}}{^{54}\text{Fe}}\right)_{\text{CHUR}}} = \frac{1 - (1 + \rho_{\text{Fe}}) y}{1 - y}.
$$

In $\epsilon_i = (R/R_{\text{CHUR}} - 1) \times 10^4$ notation, this takes the form

$$
\epsilon^i_{\text{Fe}} = -\rho_{\text{Fe}} y \frac{\rho_{\text{Fe}}}{1 - y} \times 10^4,
$$

where the asterisk superscript indicates that the ratios have not been internally normalized. Equation (B9) can be applied to $^{57}\text{Fe}$,

$$
\epsilon^{57}_{\text{Fe}} = -\rho_{\text{Fe}} y \frac{\rho_{\text{Fe}}}{1 - y} \times 10^4.
$$

Isotopic analyses are corrected for mass fractionation by internal normalization to a constant $^{57}\text{Fe}^{54}\text{Fe}$ ratio. Internally normalized $\epsilon$ values are computed using the following formula:

$$
\epsilon^i_{\text{Fe}} = \epsilon^i_{\text{Fe}} - \mu^i_{\text{Fe}} \epsilon^{57}_{\text{Fe}},
$$

where $\mu^i_{\text{Fe}} = (i - 54)/(57 - 54)$. We therefore have

$$
\epsilon^i_{\text{Fe}} = -\frac{y}{1 - y} \left(\rho_{\text{Fe}} - \mu^i_{\text{Fe}} \rho_{\text{Fe}}\epsilon^{57}_{\text{Fe}}\right) \times 10^4.
$$

From equation (B6), we have $y = x / (1 + \rho_{\text{Fe}}^{56})$, and thus $y(1 - y) = x / (1 + \rho_{\text{Fe}}^{56} - x)$. The fraction of the solar system inventory of Fe contributed by nucleosynthesis in a nearby star must have been small ($x \ll 1 + \rho_{\text{Fe}}^{56}$). Using $y(1 - y) \approx x / (1 + \rho_{\text{Fe}}^{56})$ and the expression of $x$ given in equation (B3), equation (B12) takes the form (eq. [2]),

$$
\epsilon^i_{\text{Fe}} \approx -\frac{\rho_{\text{Fe}}^i - \rho_{\text{Fe}}^i \rho_{\text{Fe}}^{57}}{1 + \rho_{\text{Fe}}^{56}} \left(\frac{^{60}\text{Fe}}{^{56}\text{Fe}}\right)_{\text{CHUR}} \cdot \epsilon^{\Delta \eta t} \times 10^4.
$$
Similarly to equation (A8), the isotopic composition of a reservoir formed with an initial \((^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}}\) ratio different from CHUR is

\[
e^{\prime}\text{Fe}_{r} \approx \frac{\rho_{\text{Fe}}^{60} - \rho_{\text{Fe}}^{56} (^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}} - (^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}}}{1 + \rho_{\text{Fe}}^{56}} (^{60}\text{Fe}/^{56}\text{Fe})_{\text{ccSN/AGB}} e^{\Delta \text{t}} \times 10^{4},
\]

where \(\rho_{\text{Fe}}^{i}\) = \((i - 58)/(62 - 58)\) and \(\rho_{\text{Ni}}^{i} = (^{58}\text{Ni}/^{56}\text{Ni})_{\text{ccSN/AGB}}/(^{58}\text{Ni}/^{56}\text{Ni})_{\text{CHUR}} - 1\). As discussed before, \(cx \approx 1 + \rho_{\text{Fe}}^{56}\). For that reason, the expression \(z(1 - z) = cx(1 + \rho_{\text{Ni}}^{62} - cx)\) can be approximated by \(cx(1 + \rho_{\text{Ni}}^{62})\). We therefore have (eq. [3])

\[
e^{\prime}\text{Ni}_{0} \approx - \frac{z}{1 - z} (\rho_{\text{Ni}}^{i} - \rho_{\text{Ni}}^{62}) \times 10^{4},
\]

Derivation of Equation (3)

The derivation of equation (3) is in many respects similar to that presented for equation (2). An important difference lies in the fact that Fe and Ni have different chemical behaviors and can therefore be decoupled during or after injection. In the following, we shall assume that Fe and Ni are not fractionated. Let us define \(z = \frac{^{58}\text{Ni}_{\text{ccSN/AGB}}}{^{58}\text{Ni}_{\text{CHUR}}}\). This can be rewritten in the form

\[
z = \frac{(^{58}\text{Ni}/^{54}\text{Fe})_{\text{ccSN/AGB}}}{(^{58}\text{Ni}/^{54}\text{Fe})_{\text{CHUR}}}.
\]

If we now introduce \(c = (^{58}\text{Ni}/^{54}\text{Fe})_{\text{ccSN/AGB}}/(^{58}\text{Ni}/^{54}\text{Fe})_{\text{CHUR}}\), it follows

\[
z = \frac{cx}{1 + \rho_{\text{Ni}}^{62}}.
\]

For Ni, we can derive an equation very similar to equation (B12) with \(z\) in place of \(y\),

\[
e^{\prime}\text{Ni}_{0} \approx - \frac{z}{1 - z} (\rho_{\text{Ni}}^{i} - \rho_{\text{Ni}}^{62}) \times 10^{4},
\]

Note that for \(i = 60\), the nucleosynthetic anomalies would superimpose on \(^{60}\text{Ni}\) deficits from lack of \(^{60}\text{Fe}\) decay (eq. [1]). Similarly to equation (A8), the isotopic composition of a reservoir formed with an initial \((^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}}\) different from CHUR is

\[
e^{\prime}\text{Ni}_{r} \approx c \frac{\rho_{\text{Ni}}^{i} - \rho_{\text{Ni}}^{62} (^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}} - (^{60}\text{Fe}/^{56}\text{Fe})_{\text{CHUR}}}{1 + \rho_{\text{Fe}}^{56}} (^{60}\text{Fe}/^{56}\text{Fe})_{\text{ccSN/AGB}} e^{\Delta \text{t}} \times 10^{4}.
\]

This equation can be used to quantify the degree of heterogeneity in \(^{60}\text{Fe}/^{56}\text{Fe}\) ratios based on \(e^{64}\text{Ni}\) measurements.

Appendix D

Derivation of Equation (4)

This equation was derived in another context by Jacobsen & Wasserburg (1984). Let us consider a reservoir \(r\) that evolved with chondritic composition until time \(t_{r}\), at which point Fe/Ni fractionation occurred. The present inventory of \(^{60}\text{Ni}\) is

\[
^{60}\text{Ni}_{r} = ^{60}\text{Ni}_{t_{r}} + ^{60}\text{Fe}_{t_{r}}.
\]

This can be rewritten as

\[
(^{60}\text{Ni}/^{58}\text{Ni})_{r} = (^{60}\text{Ni}/^{58}\text{Ni})_{t_{r}} + (^{60}\text{Fe}/^{56}\text{Fe})_{t} (^{56}\text{Fe}/^{58}\text{Ni})_{r}.
\]

We note that because the system evolved with chondritic composition until \(t_{r}\), we have \((^{60}\text{Ni}/^{58}\text{Ni})_{t_{r}} = (^{60}\text{Ni}/^{58}\text{Ni})_{t_{r}}\),

\[
(^{60}\text{Ni}/^{58}\text{Ni})_{r} = (^{60}\text{Ni}/^{58}\text{Ni})_{t_{r}} + (^{60}\text{Fe}/^{56}\text{Fe})_{t} (^{56}\text{Fe}/^{58}\text{Ni})_{r}.
\]

The same equation can be posed for CHUR,

\[
(^{60}\text{Ni}/^{58}\text{Ni})_{\text{CHUR}} = (^{60}\text{Ni}/^{58}\text{Ni})_{t_{r}} + (^{60}\text{Fe}/^{56}\text{Fe})_{t} (^{56}\text{Fe}/^{58}\text{Ni})_{\text{CHUR}}.
\]
Forming the difference between the last two equation,

\[
(60\text{Ni}/58\text{Ni})_r - (60\text{Ni}/58\text{Ni})_{\text{CHUR}} = (60\text{Fe}/56\text{Fe})_r \left[ (56\text{Fe}/58\text{Ni})_r - (56\text{Fe}/58\text{Ni})_{\text{CHUR}} \right].
\]

(D5)

Dividing the two sides by \((60\text{Ni}/58\text{Ni})_{\text{CHUR}}\), it is easy to rearrange the equation and show that (eq. [4])

\[
\epsilon^{60\text{Ni}} = Q \left( \frac{56\text{Fe}}{56\text{Ni}} \right)_r \frac{f_{\text{Fe}}}{f_{\text{Ni}}},
\]

(D6)

where \(Q = (56\text{Fe}/60\text{Ni})_{\text{CHUR}} \times 10^4\) and \(f_{\text{Fe}}/f_{\text{Ni}} = (\text{Fe}/\text{Ni})_r/(\text{Fe}/\text{Ni})_{\text{CHUR}} - 1\).
ERRATUM: “IRON-60 EVIDENCE FOR EARLY INJECTION AND EFFICIENT MIXING OF STELLAR DEBRIS IN THE PROTOSOLAR NEBULA” (2008, ApJ, 686, 560)

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Due to an error in processing the data, some values of $\epsilon_{60}^{\text{Ni}}$ corrected for instrumental mass fractionation using $^{61}\text{Ni}/^{58}\text{Ni}$ ratios are incorrect. The corrected Table 2 is appended below. Nevertheless, the shifts that result from this correction are all within error bars and do not affect the conclusions of the paper. The figures are also unaffected.

| Sample (type) | Number | $^{58}\text{Ni}$ | $^{60}\text{Ni}$ | $^{61}\text{Ni}$ | $^{62}\text{Ni}$ | $^{64}\text{Ni}$ |
|---------------|--------|------------------|------------------|------------------|------------------|------------------|
| Coahuila (IIAB) | 15 | 0.001 ± 0.061 | 0.124 ± 0.269 | 0.001 ± 0.061 | 0.124 ± 0.269 | 0.001 ± 0.061 |
| Santa Luzia (IIAB) | 8 | 0.016 ± 0.123 | 0.129 ± 0.279 | 0.016 ± 0.123 | 0.129 ± 0.279 | 0.016 ± 0.123 |
| Casas Grandes (IIIAB) | 30 | 0.062 ± 0.052 | 0.125 ± 0.141 | 0.062 ± 0.052 | 0.125 ± 0.141 | 0.062 ± 0.052 |
| Henbury (IIIAB) | 17 | 0.017 ± 0.031 | 0.003 ± 0.151 | 0.017 ± 0.031 | 0.003 ± 0.151 | 0.017 ± 0.031 |
| Molong (PMG) | 33 | 0.048 ± 0.055 | 0.121 ± 0.194 | 0.048 ± 0.055 | 0.121 ± 0.194 | 0.048 ± 0.055 |
| Bishunpur (LL3.1) | 14 | 0.122 ± 0.092 | 0.189 ± 0.287 | 0.122 ± 0.092 | 0.189 ± 0.287 | 0.122 ± 0.092 |

\textbf{Note.} Uncertainties are 95\% confidence intervals.