Modelling maximum discharge of the catastrophic flood at the Iya River (Irkutsk region, Russia) in 2019

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Abstract. The aim of the study is to estimate the maximum discharge of the catastrophic flood in June 2019 at the Iya River (Irkutsk Region, Russia). The main cause of this flood was extreme precipitation (170 mm for 3 days). The distributed deterministic hydrological model Hydrograph was applied. The schematization of the Iya river basin, parametrization and verification of the Hydrograph model were performed. The median value of the Nash-Sutcliff criteria was 0.69 for the period 1970-1996 for three catchments of the Iya River basin. Based on the data of weather stations and global weather forecast model ICON, maximum daily discharge values of the flood were estimated as 6570 and 4780 m3s⁻¹ respectively with the possible value range assessed by the dependence of Q(H) 6250-7500 m3s⁻¹. The flood hydrograph estimated from weather station data coincides in magnitude of flood peak, but its formation is delayed by 1 day. ICON data underestimates maximum value but provides proper timing of the flood peak. The ensemble of input meteorological data from various sources could potentially be used to satisfactorily predict the magnitude and duration of the catastrophic flood and minimize the consequences of the flood.

1 Introduction

Irkutsk Region, located in the South-Eastern part of Siberia (Russia), regularly experiences flooding caused by extreme precipitations. Earlier, such floods in this area occurred in 1984, 1996, 2001 [1].

The most hazardous flood was observed at the rivers originating from the Eastern Sayan mountains in June 2019. As a result of the flood, 107 settlements were affected, 25 people died and 8 were missing [2]. According to preliminary assessment, economic damage from the flood in 2019 amounted up to half a billion Euro in Irkutsk region [3]. The main reasons of hazardous flood discussed in the media are melting of snow and glaciers in the mountains, deforestation and forest fires, heavy, long rains as a result of climate change [4].

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The assessment of design flood characteristics in Russia is based on a historical observed data. But the issues related to the questionable applicability of historical observed data to assess hydrological characteristics at present are widely discussed [5, 6]. The necessity for development of new methods for assessment, including the methods based on mathematical modeling of flow formation processes is manifested [7, 8]. Due to the increase in the frequency and amplitude of extreme hydrological events in the future [9], it is necessary to develop new methods for calculating and predicting hydrological characteristics that can take into account the recent and future changes.

The aim of the study is to estimate the maximum discharge of the catastrophic flood in June 2019 at the Iya River at the Irkutsk Region (Russia) based on hydrological modelling approach and assess other possible reasons which could cause the flood.

2 Study area

The Iya River basin (14500 m²) (the South-Eastern part of Siberia, Russia) belongs to the zone of dark coniferous taiga (Fig. 1). The Iya River originates from the northern slopes of the Eastern Sayan up to 2789 m of altitude (the Holba peak). The study area is divided into mountain and lowland. There are significant areas occupied by goltsy in the mountain part. The area has a continental climate. Affected by both climate and terrain, the rainfall distribution in space and time is uneven. The annual precipitation in different parts of the basin changes from 350 mm yr⁻¹ to 760 mm yr⁻¹. The main amount of precipitation falls in the summer period.

![Fig. 1. The scheme of the Iya River basin at Tulun.](image-url)
The hydrological regime of the Iya River is characterized by spring-summer flood, high summer rainfall floods and a low winter flow. The main part of runoff is formed in mountain (56% of the total area) and reaches about 280 mm yr\(^{-1}\). On the lowland, about 50 mm yr\(^{-1}\) is formed. There are three hydrological gauges where streamflow measurements are conducted in the Iya River basin (Table).

### 3 Verification of hydrological model

Distributed process-based hydrological model Hydrograph is used in this study. It describes all components of the land hydrological cycle, including precipitation and its interception; snow accumulation and melting; evaporation from snow, soil, and vegetation cover; surface flow and infiltration; soil water dynamics and flow; heat dynamics and phase change in soil layers; underground flow formation, slope and channel flow transformation [10, 11]. The model is a workable tool for studying hydrological processes and assessing the characteristics of runoff under various physical and geographical conditions and on objects of any size from the soil column or elementary slope to large river basins without changing the structure and algorithms. The description of the model and methods for its parameterization are presented in detail in earlier studies [10-15].

The input to the model is a limited set of meteorological forcing data (air temperature and humidity, precipitation) that can be obtained for most of the weather stations. The main output includes streamflow, water balance components and variable states (snow depth, soil moisture, temperature and other).

The basin of the Iya River was divided into three runoff formation complexes (RFC) based on satellite imagery Landsat-8 and the landscape map of the USSR: goltsy, dark coniferous forest and river valleys (Fig.1). RFC – are the areas which can be described by one set of the model parameters presenting some kind of hydrological landscapes where runoff formation processes are taken as uniform. Parameters of this RFC (vegetation and soil cover, surface) were determined based on reference materials (observed data under identical climatic conditions). The catchment was also covered with a regular grid of representative points (RP). The hydrological processes are simulated at RPs that possess the unique topography characteristics and the set of the model parameters related to one of the RFCs.

We used daily meteorological data, such as air temperature, air moisture, and precipitation from 7 weather stations (Fig. 1) to simulate streamflow for verification of the hydrological model. The station Beloziminskaya (755 m), which is the most representative for the mountainous part of the catchment, was ceased after 1996. The modelling period (1970-1996) was determined by the availability of data for all weather stations. The interpolation of precipitation was carried out taking into account the distribution of snow and rainfall in the mountains.

Runoff formation processes were simulated for the subbasins of the Iya River basin (Table) with a daily time step for the model verification. The median value of the Nash-Sutcliffe criteria was from 0.66 to 0.72 for the period 1970-1996 for three catchments (Table 1). The difference in the mean annual value of the calculated and observed flow does not exceed 7%. In addition, we verified the model by maximum water discharges. The curves at Fig. 2 show the exceedance probability of calculated and observed annual maximum water discharge for the simulation period 1970-1996 at the Iya River basin at Tulun. The difference in values does not exceed 300 m\(^3\)s\(^{-1}\) (8%). The results have been considered satisfactory, and the Hydrograph model being suitable for modelling the processes of runoff formation in the territory under consideration.
Table. Characteristics of runoff gauge stations, water balance and efficient criteria, 1970-1996.

| Index | River – gauge | Period         | S (km²) | H (m)  | Yo  | Ys  | P   | E   | NS (m/av) |
|-------|--------------|----------------|---------|--------|-----|-----|-----|-----|-----------|
| 8230  | Iya – Arshan | 1963-2017      | 5140    | 1483   | 540 | 528 | 771 | 243 | 0,69/0,62 |
| 8233  | Iya – Tulun  | 1941-2017      | 14500   | 979    | 326 | 338 | 586 | 247 | 0,72/0,67 |
| 8241  | Kirej – Ujgat| 1959-2017      | 2950    | 873    | 374 | 402 | 688 | 286 | 0,66/0,57 |

Note: S – basin area, km²; H – average catchment elevation (m); Yo and Ys – observed and calculated mean annual streamflow, mm; P, E – calculated precipitation and evaporation, mm; m and av – median and average value of Nash-Sutcliffe (NS) criteria.

Fig. 2. The curves of frequency of maximum water discharge according to modeling results (red points) and the observations (black points), the Iya river – Tulun, 1970-1996

4 Assessment of the hazardous flood causes

We have analyzed satellite image Landsat-8 for June 23 which shows that even in the mountainous part of the basin less than 10% of the area was covered with snow at that time. This could not cause flooding of such magnitude. Deforestation also could not be the reason of this flood because the deforested area in the basin occupies less than 4% of the total catchment area (Fig. 1) [16].

The analysis of changes in hydrometeorological characteristics by the Mann-Kendall method with a significance level p<0.05 and the Tain-Sen method to assess the magnitude of trends for the period 1966-2019 were carried out. The air temperature over the indicated period increased by 1.2 - 2.1 °C; there is an increase of rainfall in June by 36-61% or 31-46 mm. Increase of streamflow is observed only in the cold period, from November to April. The series of maximum daily streamflow for the gauge the Iya River – Tulun is characterized by a negative trend (on average 37% in May and July at p<0.05 and 24% in June and August at p<0.08).
5 Simulations of maximum discharge based on different sources of meteorological data

Two experiments with different types of input meteorological data due to the lack of representative observational data were conducted to simulate the hazardous flood in June 2019. We used 1) the daily meteorological data from two stations (Ikey and Tulun) and 2) the data from the global weather forecast model ICON [17] as the input. Both types of data were available only for the period of 1-30 June, therefore the year 1989 was adopted as the initial conditions for running the hydrological model. In 1989 the spring flood conditions by magnitude and timing were similar to the conditions of 2019.

Using the weather stations data the following distribution of water balance was obtained based on simulations: basin precipitation during 25-27 June – 239 mm, including 79 mm – on June 25, 132 mm – on June 26, 28 mm – on June 27. The estimated maximum daily discharge reaches 6570 m$^3$s$^{-1}$ (Fig. 3).

For the 2nd types of input (ICON model data) we used daily data for 1989 until May 31 and 3-hour ICON data from June 1 to 30. Total basin precipitation amounted to 170 mm during 25-27 June. As a result, the maximum calculated 3-hour discharge reached 5260 m$^3$s$^{-1}$, maximum daily value was 4780 m$^3$s$^{-1}$ (Fig. 3).

Using water level (H) and discharge (Q) historical data we estimated possible discharge based on the curve Q(H). The possible value is within the range from 6250 to 7500 m$^3$s$^{-1}$ (blue area in Fig. 3).

The maximum discharge based on ICON data is 1400 m$^3$s$^{-1}$ lower than the observed, however, its formation coincides by timing. According to weather station data, the maximum discharge coincides in magnitude, but its formation is delayed by 1 day.

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Fig. 3. The results of flood modeling at the Iya River – Tulun in June 2019: 1, 2 – the amount of precipitation for the catchment – 3-hour precipitation according to the ICON weather model and daily precipitation based on data from weather stations; 3 – the observed flow hydrograph (based on extrapolation of the dependence of water flow on the level); 4, 5 – calculated 3-hour and averaged daily flow hydrograph according to the ICON weather model; 6 – calculated daily runoff hydrograph based on data from weather stations.
6 Conclusions

In this study, precipitation as the main reason for the catastrophic flood at the Iya River in 2019 is revealed. Based on the weather station and weather model data, maximum daily discharge values were estimated as 6570 and 4780 m³s⁻¹ respectively with the possible value range assessed by the dependence Q(H) 6250 – 7500 m³s⁻¹.

By the study we also attempted to show the need to expand the meteorological and hydrological network in the region. We also demonstrate the capabilities of the modern calculation methods and forecasts in case of insufficient observed data. In our opinion, the ensemble of input meteorological data from various sources could potentially be used to satisfactorily predict the magnitude and duration of the catastrophic flood. So, the results of the study confirm the fundamental possibility of a short-term flood forecast, using the method of deterministic hydrological modeling.

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