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The global and regional impacts of climate change under representative concentration pathway forcings and shared socioeconomic pathway socioeconomic scenarios

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Abstract
This paper presents an evaluation of the global and regional consequences of climate change for heat extremes, water resources, river and coastal flooding, droughts, agriculture and energy use. It presents change in hazard and resource base under different rates of climate change (representative concentration pathways (RCP)), and socio-economic impacts are estimated for each combination of RCP and shared socioeconomic pathway. Uncertainty in the regional pattern of climate change is characterised by CMIP5 climate model projections. The analysis adopts a novel approach using relationships between level of warming and impact to rapidly estimate impacts under any climate forcing. The projections provided here can be used to inform assessments of the implications of climate change. At the global scale all the consequences of climate change considered here are adverse, with large increases under the highest rates of warming. Under the highest forcing the global average annual chance of a major heatwave increases from 5% now to 97% in 2100, the average proportion of time in drought increases from 7% to 27%, and the average chance of the current 50 year flood increases from 2% to 7%. The socio-economic impacts of these climate changes are determined by socio-economic scenario. There is variability in impact across regions, reflecting variability in projected changes in precipitation and temperature. The range in the estimated impacts can be large, due to uncertainty in future emissions and future socio-economic conditions and scientific uncertainty in how climate changes in response to future emissions. For the temperature-based indicators, the largest source of scientific uncertainty is in the estimated magnitude of equilibrium climate sensitivity, but for the indicators determined by precipitation the largest source is in the estimated spatial and seasonal pattern of changes in precipitation. By 2100, the range across socioeconomic scenario is often greater than the range across the forcing levels.

1. Introduction
This paper presents a multi-sectoral analysis of the global and regional impacts of climate change through the 21st century, using a set of indicators that are directly relevant to policymakers at national and international scales calculated using consistent climate and socio-economic projections. It uses climate pathways representing seven levels of forcing describing different emissions pathways defined by representative concentration pathways (RCPs) (O’Neill et al 2016), together with five socio-economic scenarios representing future exposure defined by shared socioeconomic pathways (SSPs: O’Neill et al 2017). Scientific
uncertainty in the translation of emissions forcing to change in regional climate is represented by (i) uncertainty in the estimated change in global mean temperature as estimated by the MAGICC energy balance model, and (ii) uncertainty in the spatial variability in relevant climate variables as characterised by 23 climate models.

The study uses around 20 indicators characterising impacts on heat extremes, water resources, river and coastal flooding, agriculture and energy use. It distinguishes explicitly between changes in the physical hazard and resource base—which are dependent on climate change—and socio-economic impacts which are a function of both change in climate and change in exposure. The paper uses a novel approach using relationships between level of warming (or sea level) and impact derived from spatially-explicit impacts models, which are then combined with probabilistic projections of increase in global mean temperature to rapidly estimate impacts under a wide range of alternative climate forcings. Whilst other studies have developed relationships between level of forcing and impact (e.g. Schleussner et al 2016, Seneviratne et al 2016, Arnell et al 2016a, Arnell et al 2019), this is the first time that such relationships have been combined with probabilistic temperature projections to estimate impacts under different forcings.

The analysis develops on earlier work (Arnell et al 2013, 2016b, 2019), and uses more up-to-date climate pathways, climate models and socio-economic scenarios, probabilistic projections of global temperature change, and additional indicators that characterise a wider range of hazards and impacts. This paper presents summary charts of global and continental impacts, whilst more comprehensive regional tables and charts are presented in supplementary material is available online at stacks.iop.org/ERL/14/084046/mmedia. It does not itself explicitly assess or compare impacts across sectors and regions or describe in detail all impacts in all regions: such an assessment requires an explicit judgement of the relative importance of different indicators and thresholds defining ‘significant’ change. However, it provides regional information which will be directly relevant to such assessments, including those made in the forthcoming IPCC 6th Assessment Report. For example, the quantitative indicators can be grouped into ’severity classes’ based on explicit thresholds, as is widely done in national risk assessments. The projected quantitative indicators of impact can also be combined with the qualitative characterisations of drivers of future vulnerability (for example as represented in the SSPs: O’Neill et al 2017) to produce more nuanced narrative storylines suitable for stress testing or strategic evaluations.

In recent years an increasing number of studies have assessed the global scale impacts of climate change. Most of these have concentrated on individual sectors, including heat extremes (e.g. Dosio et al 2018, Harrington et al 2018, Lehner et al 2018, Tebaldi and Wehner 2018), drought (Smirnov et al 2016, Naumann et al 2018), flooding (Arnell and Gosling 2016, Winsemius et al 2016, Alfi e et al 2017), water resources (Arnell and Lloyd-Hughes 2014, Gosling and Arnell 2016) and agriculture (e.g. Ruane et al 2018, Schleussner et al 2018, Tebaldi and Lobell 2018). A small number have considered multiple impacts across sectors (e.g. Schleussner et al 2016, Arnell et al 2016b, Betts et al 2018, Byers et al 2018, O’Neill et al 2018). Studies have used different climate pathways, including RCP forcings and pathways consistent with 1.5 °C and 2 °C climate targets, and have used different sets of climate models to define climate scenarios.

2. Methods

2.1. Overview of approach

Scenario-based climate change impact studies typically estimate impacts from climate scenarios constructed from climate model simulations, and these impacts are therefore conditional on the forcings (such as RCP) used to run the climate models. However, not all climate models are run with all RCP forcings (and forcing scenarios can be updated), and users may be interested in impacts under different rates of forcing. This study uses ‘damage functions’ relating impacts (section 2.2) in a given year to a simple metric of climate change (global mean surface temperature or increase in sea-level rise) together with projections of this metric under a defined level of forcing to estimate impacts over time (figure 1). The different damage functions (left panel) represent uncertainty in the spatial pattern of change in climate variables, and the distribution of change in temperature in a given year (middle panel) represents the effects of uncertainty in the climate system parameters driving response to forcing (section 2.5).

The damage functions are constructed ‘bottom-up’ using spatially-explicit impact models run with scenarios representing specific changes in global mean surface temperature or sea level. The damage functions for the socio-economic indicators are contingent on socio-economic scenario (section 2.3) and therefore time. In this application, the damage functions are used to estimate impacts under the latest set of RCP forcings defined in ScenarioMIP (O’Neill et al 2016), although in principle could be used with any temperature trajectories. There are some similarities with the (independently-developed) method used by Hsiang et al (2017) to estimate economic impacts in the USA, although the details are different: most significantly the damage functions used in the current study take into account the spatial variability in change in climate as represented by the 23 CMIP5 climate model patterns.

It is assumed that each socio-economic scenario can be combined with each level of climate forcing (van Vuuren et al 2014), although in practice high
levels of forcing may not be plausible with some socio-economic scenarios and achieving low levels of forcing will be much more challenging under some than others (Riahi et al. 2017).

The period 1981–2010 is used to represent the reference climate against which the effects of climate change can be compared (the reference period sea level is the average over 1986–2005).

2.2. Indicators of impact

Table 1 summarises the indicators of hazard, resource base and impact, and specific definitions and details of their calculation are given in supplementary material. In each case, there are several plausible alternative indicators, and indeed each global impact study tends to use a different set of indicators for each area of impact. The indicators can be interpreted as proxies for change in hazard and impact. The hazard and resource indicators are expressed in physical terms (such as likelihood of occurrence of a specific magnitude event), and—with two exceptions—represent the regional average value of the indicator at a point (the two exceptions are the indicators calculating the regional area exposed to a significant change in river runoff and the area below the coastal 100 year flood level). The regional average agri-climate hazard indicators are calculated by weighting by cropland areas. The other regional average hazard indicators are calculated by weighting by grid cell area, excluding grid cells with fewer than 1000 people in 2010. This weighting is used because the focus here is on the consequences of climate change relevant to people.

The impact indicators are all expressed as regional aggregations of numbers of people, area of cropland or energy use. The socio-economic indicators depend on socio-economic scenario, and indicators are calculated for each combination of socio-economic scenario and climate forcing.

The two heatwave magnitude indicators (defining heatwaves as either two or more days with temperatures greater than the 98th percentile of reference period warm-season temperatures, or four or more days with temperatures greater than the 99th percentile) represent different size heatwaves, one which currently has a 33% chance (approximately) of occurring in a place, and the other a 5% chance. The indicators based on average annual runoff represent pervasive water resources scarcity, whilst the indicators based on hydrological drought represent short-term availability. Similarly, the agricultural drought indicator represents short-term challenges to crop production in general. The change in crop growth duration is a proxy for crop yield (a shorter duration for crop growth is associated with reductions in yield: Challinor et al. 2016), and the crop heat stress indicator represents the effect of extreme events on yield (Gourdji et al. 2013). River flood hazard characterised by the annual likelihood of the reference period 50 year (2%) flood, and coastal flood hazard is represented by the area of land under the 100 year flood level. River flood impact is defined as the average annual number of people exposed to reference period the 50 year flood, but because the indicator does not incorporate flood defences these people are not necessarily actually flooded. One of the coastal flood impact indicators is the average annual number of people estimated to be actually flooded in events that exceed flood defence standards, and this makes two different assumptions about how flood defence standards improve through time: the difference between the two characterises the effect of adaptation. The second coastal flood impact indicator is the number of people living below the 100 year flood level, ignoring defence levels.

2.3. Socio-economic scenarios

This assessment uses five SSPs (O’Neill et al. 2017), which are based on five different narrative storylines for the future development of societies, economies and governance. They are plausible projections, rather than predictions. National population projections, by age, for each SSP are described by Samir and Lutz (2017), and this analysis uses the population projections downscaled to the 0.5° × 0.5° resolution by Jones and O’Neill (2016). The five SSPs differ in their assumptions about fertility and mortality, rates of urbanisation and international migration. Projections of national Gross Domestic Product are taken from

Figure 1. Illustration of the methodology. Damage functions showing impact in a given year against increase in global mean temperature (1a) are combined with the distribution of the increase in temperature in that year (1b) to produce a distribution of impacts in that year (1c). The individual lines in (1a) are the damage functions constructed from different climate model patterns. The dotted vertical lines in (1c) show the 10th and 90th percentiles, and the solid vertical line shows the median.
| Hazard/ resource indicator | Impact indicator | Additional information |
|----------------------------|------------------|------------------------|
| **Heat extremes**         |                  |                        |
| Heatwave frequency: annual % chance of experiencing at least one at a given location | A heatwave is a period of at least two days with daily maximum temperature greater than the 98th percentile of warm season temperatures, at a given location | Average annual population exposed to at least one heatwave: millions yr$^{-1}$ |
| Major heatwave frequency: annual % chance of experiencing at least one at a given location | A heatwave is a period of at least four days with daily maximum temperature greater than the 99th percentile of warm season temperatures, at a given location | Average annual population exposed to at least one major heatwave: millions yr$^{-1}$ |
| Heatwave duration: average annual number of heatwave-days at given location | A heatwave is a period of at least two days with daily maximum temperature greater than the 98th percentile of warm season temperatures, at a given location | Average annual population exposed to heatwaves: million people-days/year |
| **Water**                 |                  |                        |
| Area with increase or decrease in average annual runoff: % of area | A significant increase or decrease in runoff is more than twice the standard deviation of 30 year mean runoff | Population living in watersheds that become water-stressed, or cease to be water-stressed: millions |
| Duration of hydrological drought: % of time in drought at a given location | Hydrological drought occurs when the 12-month accumulated Standardised Runoff Index (SRI: Shukla and Wood 2008) is less than $-1.5$ | Average annual population exposed to drought: millions yr$^{-1}$ |
| **Floods**                |                  |                        |
| Frequency of reference period 50 year (2%) river flood: annual % likelihood at a given location | Flood frequency is estimated by fitting a Generalised Extreme Value (GEV) distribution to simulated river flows (Arnell and Gosling 2016) | Average annual population exposed to river flooding: millions yr$^{-1}$ |
| Area of coastal land below the 100 year coastal flood level: thousand km$^2$ | 100 year return period coastal flood level estimated using the DIVA model (Vafeidis et al 2008) | Coastal population living below the 100 year flood level: millions |
|                          |                  | 100 year return period coastal flood level estimated using the DIVA model, ignoring coastal flood defences | Average annual number of people flooded in coastal floods: millions yr$^{-1}$ |
|                          |                  | Flood levels estimated using DIVA, with two assumptions about changing levels of coastal defences |
### Table 1. (Continued.)

| Hazard/resource indicator | Impact indicator | Additional information |
|---------------------------|------------------|------------------------|
| **Agriculture**           |                  |                        |
| Duration of agricultural drought: % of time in drought at a given location \(^{(c)}\) | Agricultural drought occurs when the 6 month accumulated Standardised Precipitation Evaporation Index (SPEI: Vicente-Serrano et al 2010) is less than −1.5 | Average annual area of cropland exposed to drought: thousand km\(^2\) yr\(^{-1}\) |
| Change in average annual crop growth duration: days at a given location \(^{(c)}\) | Crop growth duration is based on the time to accumulate reference period thermal degree-days, with thresholds varying between five crops: maize, winter wheat, spring wheat, rice and soybean | Average annual area of cropland with reduction in average crop growth duration of at least 10 days: thousand km\(^2\) yr\(^{-1}\) |
| Frequency of damaging hot spells during crop reproductive season: annual % chance at a given location \(^{(c)}\) | The temperature threshold and timing of reproductive season varies between the five crops: maize, winter wheat, spring wheat, rice and soybean | Average annual area of cropland experiencing a damaging hot spell: thousand km\(^2\) yr\(^{-1}\) |

**Energy**

| Hazard/resource indicator | Impact indicator | Additional information |
|---------------------------|------------------|------------------------|
| Cooling degree days: days/year \(^{(p)}\) | Average annual cooling-degree days relative to 18 °C, in a given location. | Cooling energy demands: PJ |
| Heating degree days: days/year \(^{(p)}\) | Average annual heating-degree days relative to 18 °C, in a given location. | Heating energy demands: PJ |

See supplementary material for more details and references.

\(^{(p)}\) Regional averages weighted by grid cell area, omitting grid cells with fewer than 1000 people in 2010.

\(^{(c)}\) Regional averages weighted by cropland area.
Dellink et al. (2017) and downscaled to the 0.5° × 0.5° resolution by assuming that each grid cell in a country has the same GDP per capita. GDP is used in the energy and coastal flood impact indicators. Figure 2 shows the global total population and GDP through the 21st century under the five SSPs.

For the agricultural indicators, it is assumed that the areas of cropland (total and by crop; supplementary material) remain constant through the 21st century.

2.4. Calculation of damage functions

For the terrestrial indicators (section 2.2), damage functions are constructed using climate scenarios from 23 atmosphere-ocean general circulation model patterns scaled, using ClimGen (Osborn et al. 2016, 2018), to defined increases in global mean temperature and applied to a reference period climatology. These scaled scenarios were constructed using pattern-scaling rather than time-slicing (James et al. 2017) in order to eliminate the effects of year-to-year and decade-to-decade variability on differences between scenarios at different levels of warming. Twenty-three patterns for each climate variable (precipitation, precipitation variability, temperature, humidity, and net radiation) were constructed from 23 CMIP5 climate models (listed in supplementary material), allowing the construction of 23 damage functions for each indicator, time period and socioeconomic scenario. All 23 patterns are assumed to be equally plausible and independent. The changes in monthly climate were applied to the CRU TS4 0.5° × 0.5° 1981–2010 climatology (Harris et al. 2014) using the delta method in ClimGen (Osborn et al. 2016, 2018) to produce perturbed 30 year time series representing specific increases in global mean surface temperature. For precipitation, changes in monthly variability projected by climate models are also diagnosed and used within ClimGen to perturb the observed variability to represent the increased or decreased variability simulated by each climate model (see Osborn et al. 2016, for more details). This is important for those indicators that depend on climate variability as well as the mean climate. The climate scenarios do not incorporate the effects of naturally-forced multi-decadal variability on departures around the climate change trend, which would add to the range in projected impacts. The impacts models are applied at the 0.5° × 0.5° resolution and results averaged or aggregated to the regional and global scales.

The damage functions for the coastal indicators (section 2.2) are constructed differently because impact in a given year is a function of sea-level rise rather than temperature increase, and the relationship between temperature increase in a given year and sea level rise depends on the rate of change in temperature. The coastal damage functions relating impact to global average sea level rise were therefore developed by running the Dynamic Interactive Vulnerability Assessment (DIVA) model (Vafeidis et al. 2008, Hinkel et al. 2014) with a range of sea-level rise scenarios describing change through the 21st century and plotting estimated impacts in a given year against sea-level rise in that year. The coastal indicators are calculated for coastal segments (which vary in length) and are aggregated to the regional and global levels.

2.5. Climate forcings and increases in temperature and sea level

The ScenarioMIP initiative (O’Neill et al. 2016) has defined a series of climate forcings for use by the climate modelling community to drive global climate model experiments, representing different trajectories of future greenhouse gas emissions. These include revisions to earlier RCPs (RCP2.6, RCP4.5, RCP6.0 and RCP8.5) and three new pathways with radiative forcing at 2100 of 1.9, 3.4 and 7.0 W m⁻² (another overshoot pathway is not considered here). The new pathways are designed to fill gaps in the original range, and the revised pathways differ from the previous versions primarily through (i) the socio-economic scenarios used to define emissions of different greenhouse gases and aerosols, and (ii) the use of more...
up-to-date historical emissions and harmonisation between observations and scenarios. Impacts were calculated under all seven of the pathways (see supplementary material for full results), but the plots here focus for clarity on RCPs 2.6, 4.5 and 8.5. ScenarioMIP does not define a ‘business-as-usual’ emissions scenario, because ‘business-as-usual’ depends on socio-economic assumptions. The IPCC Fifth Assessment Report (Clarke et al 2014) concluded that emissions scenarios with no specific assumptions about emissions reductions produced forcings between 6.0 and 8.5 W m$^{-2}$ by 2100. RCP8.5 is here used in the plots as an illustrative upper limit even though it may only arise under a relatively narrow set of circumstances (Riahi et al 2017), for two primary reasons. First it continues to be widely used for climate model simulations, and second it is consistent with calls for the presentation of the consequences of high-impact and ‘worst-case’ scenarios (King et al 2015, Sutton 2018).

Projections of the increase in global mean surface temperature for each of the RCP forcings (figure 3(a)) were made using a probabilistic implementation of version 4.2 of the MAGICC energy balance model (Lowe et al 2009, Lowe and Bernie 2018). This probabilistic implementation samples across 1863 combinations of feasible values of equilibrium climate sensitivity (ECS), ocean diffusivity and carbon cycle feedback strength to produce 1863 projections of global mean temperature over time (each with a relative probability derived from the probability distributions of the ocean diffusivity and carbon cycle feedback strength parameters). The ECS values used are taken from the CMIP5 climate models assessed in the IPCC Fifth Assessment Report (Flato et al 2013, Forster 2013), and assumed to be equally plausible.

Impacts are estimated in each year (for a given socio-economic scenario) by combining the projected changes in temperature with the 23 equally-plausible damage functions (in practice percentiles from the distribution of temperature change in each year were used rather than the individual projections). The range in estimates of impact in each year under a specific climate and socio-economic pathway therefore represents uncertainty in (i) the increase in temperature in that year, which depends on ECS, ocean diffusivity and carbon cycle feedback strength, and (ii) the spatial and seasonal distribution of change in relevant climate variables. The magnitude of impacts is characterised by the median, and the range is represented by the 10th and 90th percentiles of the distribution of impacts in each year (but should be interpreted as ‘low’ and ‘high’ rather than specific percentiles).

Sea level rise scenarios corresponding to the temperature forcings (figure 3(b)) were constructed from the projected temperature changes using an empirical relationship between accumulated global temperature increase since 1985 and sea level rise relative to the 1986–2005 average level (supplementary material). This empirical relationship emulates the increase in sea level as presented in the IPCC AR5 (Church et al 2013). The scenarios are globally uniform. For each emissions scenario, a central estimate of sea level rise is calculated from the time series of the median increase in temperature, and low and high sea level scenarios are calculated from the time series of the 10th and 90th percentile temperature changes. The range in estimated impacts in a year under a specific climate and socio-economic pathway just represents uncertainty in the sea-level rise by that year. The median sea level rise assumed here for RCP8.5 is similar to that projected by Kopp et al (2014) and slightly lower than the estimate produced by Vousdoukas et al (2018). The range here is between Kopp et al’s (2014) ‘likely’ and ‘very likely’ ranges, but smaller than the range in Vousdoukas et al (2018). Several projections of sea level rise made since AR5 have suggested larger high-end increases than presented in AR5 (Bamber et al 2019), but there is considerable uncertainty over the effect of ice sheet melt. The high-end increases in sea level used in this study should therefore be regarded as

Figure 3. Increase in global mean surface temperature and global mean sea level, under the seven forcing scenarios. For temperature, the solid line represents the median estimate and the shaded area the 10th to 90th percentiles. For sea level, the solid line represents a central estimate and the shaded area the range between a ‘low’ and ‘high’ estimate.
3. Results

3.1. Changes in hazard

Figure 4 shows the global-scale hazard indicators through the 21st century (tables in supplementary material), under the seven RCP forcings. At the global aggregate scale, heatwave frequency increases (but at a different rate for the two heatwave definitions), drought frequency increases, flood frequency increases, the crop growth duration decreases (implying reduced yields), the likelihood of hot spells damaging to crops increases, and cooling degree-days increase: all these are adverse consequences of climate change. Heating degree-days decrease, which is a potential benefit of climate change. Figure 4 shows the wide range in estimated future hazard, particularly for the indicators based on precipitation, but the range shown here does not suggest climate change could result in a reduction in hazard at the global scale for any of the indicators. The figure also demonstrates the large difference in future hazard occurrence between the different climate forcings, with the difference conservative: coastal impacts could be considerably higher.

Figure 4. Global hazard indicators to 2100, under RCP2.6, RCP4.5 and RCP8.5. The bars on the right show impacts in 2100 under seven RCPs. The dotted line shows the value of the indicator under the 1981–2010 climate (1986–2005 sea level). The solid line represents the median and the shaded area the range between the 10th and 90th percentiles (‘low’ and ‘high’ for the coastal indicator).
increasing after around 2040. Some of the hazard indicators level off at high levels of forcing. For the heatwave frequency indicator this is simply because everywhere experiences a heatwave every year. The runoff change indicator is—unlike the other indicators—just defined on the basis of whether or not a threshold is crossed, and this levels off with high levels of forcing as fewer and fewer additional areas exceed the threshold.

The variation in aggregated hazard indicator by continent in 2100 is shown in figure 5 (variation across regions is shown in supplementary material, which also includes plots showing each hazard indicator by region). The height of the individual bars represents scientific uncertainty (section 3.2). For the indicators dependent on temperature, the bars typically do not overlap, indicating that uncertainty in the rate of forcing is large relative to the scientific uncertainty; for

Figure 5. Hazard indicators across continents in 2100: RCP2.6, RCP4.5 and RCP8.5. The solid black lines show indicators under the 1981–2010 climate (1986–2005 sea level). The solid line represents the median and the shaded area the range between the 10th and 90th percentiles ('low' and 'high' for the coastal indicator).
the precipitation-based indicators, the bars overlap indicating that scientific uncertainty is large relative to emissions uncertainty.

There is greater variability between continents with the major heatwave indicator than for the more moderate heatwave indicator (because this saturates at 100% with high emissions), and even more variability with the heatwave duration indicator. This indicator shows the greatest increase in heatwave duration in Africa, South America, Asia (especially south east Asia) and Australasia.

Hydrological drought frequency increases in each continent, but with low climate forcing there could be very little change in Asia and North America. With the highest forcing, the greatest increase in hydrological drought frequency is in Europe, South America (especially central America and Brazil), the Middle East and North Africa, and Australasia, but with wide uncertainty between climate model patterns. Agricultural drought frequency increases in every continent under all scenarios, but the difference between regions is less marked.

The greatest increase in river flood frequency is in Asia (especially south and south east Asia) and Africa, with relatively small change with low forcings in Europe and North America (and frequency could decrease).

Europe and North America see the greatest reductions in crop growth duration—and therefore potentially yield—for all crops (to the least extent for winter wheat). Maize and winter wheat hot spell frequencies increase most in Europe and North America (especially the USA), but the frequency of hot spells for soybean, spring wheat and rice increases most in Africa and Asia: Europe and North America are relatively unaffected.

The absolute changes in cooling degree days are in Asia and Africa, but the greatest relative increases are in Europe and North America. The greatest absolute decreases in heating degree days are in Asia, Europe and North America, but the greatest relative decreases are in Africa.

3.2. Sources of uncertainty
The plots in figures 4 and 5 show the median estimate, plus the range between `low’ and `high’ hazard indicator for each climate forcing. For a given level of climate forcing, the range for an indicator represents uncertainty in (i) the change in global mean temperature (a function of ECS, ocean diffusivity and carbon cycle feedback) and (ii) the spatial pattern of change in temperature and precipitation. Figure 6 shows the relative importance of each of these sources of uncertainty for each hazard indicator (except for the coastal indicator) under the highest level of forcing (RCP8.5) and at the global scale. The maximum uncertainty for each indicator is shown when the uncertainty contributions sum to 100% in figure 6.

The uncertainty range of most indicators increases over time and reaches its maximum in 2100. The uncertainty ranges of the two heatwave and the rice heat stress indicators peak earlier in the century and then decline because at large changes in climate everywhere is impacted under all parameter combinations so uncertainty reduces. The plots show that uncertainty in the ECS is much more important than uncertainty in the strength of the carbon cycle feedback or ocean diffusivity, but that for the indicators determined by precipitation change the scientific uncertainty is dominated by the uncertainty in the spatial distribution of change in rainfall as represented by different climate models. The selection of models used to estimate impacts therefore has a greater effect on the estimated uncertainty range.

3.3. Socio-economic impact
Figure 7 shows the global socio-economic impact of changes in hazard indicators in 2050 and 2100 (tables in supplementary material). The figure focuses for clarity on three climate forcings (RCP2.6, RCP4.5 and RCP8.5) and for the population indicators shows impact under all five socio-economic scenarios (the agricultural indicators are all expressed in terms of area of cropland affected, which is assumed to remain constant over time). In 2050 there is relatively little difference in impact between the three levels of forcing, but by 2100 the difference is much greater. Again, the height of the individual bars represents scientific uncertainty.

Even by 2050, however, there is a difference in impact between the five socio-economic scenarios, and this difference increases further by 2100. Scenario SSP3 has the highest total population, so has the highest impact on heatwaves, floods, droughts and the population living in watersheds which cease to be water-stressed. More people live in watersheds that become water-stressed under SSP2 than the other scenarios, and this is because of the geographical distribution of the increase in population.

By 2050, the range in impact across the climate forcings is greater than the range across the socio-economic scenarios for major heatwaves, heatwave duration, and the population exposed to river flooding (supplementary material). The range across socio-economic scenarios is greater than the range across the climate forcings for heatwave impacts (possibly here because of the saturation at high levels of forcing noted above), energy demand, exposure to water resources scarcity and coastal flood exposure. By 2100, the range across climate forcings is greater than the range across socio-economic scenarios only for major heatwaves and heatwave duration impacts, and the greater range across socio-economic scenarios is more apparent than in 2050.

The socio-economic impacts of climate change are—at the global scale—adverse for most, but not all the socio-economic indicators. Residential heating energy requirements decline with higher temperatures, and
by 2050 the global total residential heating and cooling demands are lower than they would be without climate change under all the climate forcings and socio-economic scenarios. By 2100, however, the increases in cooling energy demand outstrip the reductions in heating demand in three of the socio-economic scenarios (SSP1, SSP2 and SSP5), so total demands are increased. At the global scale, more people live in watersheds that ceased to be water-stressed than live in watersheds that become water-stressed, particularly under the high population SSP3 and increasingly through the 21st century. This is different to the distribution of change in runoff (figure 4), because the areas with an increase in runoff are more populous than areas with a decrease. Note that the drought indicator represents a different dimension to water resources stress, as droughts can occur in regions that are not water-stressed.

The variation in impact between continents and regions can be expressed in two ways. Expressing

Figure 6. Relative importance of different sources of scientific uncertainty in the projected hazard indicators: RCP8.5, global scale. The plots show the contribution of uncertainties in equilibrium climate sensitivity (ECS), carbon cycle feedback, ocean diffusivity and climate model pattern to the total uncertainty in the projected impacts in each year. Uncertainty is characterised by the average variance in the distribution of estimated impacts in a year across each source of uncertainty. The contributions of each source are expressed as a % of the maximum total sum of the variances across all years: this is usually 2100.
impacts in \textit{absolute} terms—people, area of cropland and energy use—gives an indication of the absolute magnitude of impact in a region and the contribution of a region to the global total impact. Expressing impacts in \textit{relative} terms—as a proportion of population or cropland, for example—gives an indication of the relative significance of an impact in a region. A third approach would be to show impacts as a proportional change from a reference period. Each is policy relevant. The distribution of \textit{absolute} socio-economic impact in 2100 across continents by indicator is shown in figure \ref{fig:7a} (just for SSP2 for the population indicators: the variation between continents is broadly similar with the other SSPs. Supplementary material also shows distribution by region and allows the calculation of impacts in relative terms). The plots show the indicators in 2100 with the value for 2100 assuming no climate change, and give an indication of the effect of climate change relative to the reference climate. The river flooding and cropland indicators also show the
total floodplain population and cropland area, and therefore give an indication of the relative importance of the impact in each region.

In absolute terms, the greatest impact of climate change on people exposed to heatwaves is in Asia (especially south and east), followed by Africa. The greatest number of people exposed to an increase in water scarcity are in west Africa, and most people exposed to river and coastal flooding are in Asia, especially south Asia. Asia also sees the largest number of people living in watersheds that cease to be defined as water-scarce by 2100 (mostly in south Asia). Africa and Asia have the greatest number of people exposed to hydrological drought. The greatest absolute cooling energy requirements by 2100 are in (south) Asia and (west) Africa, whilst heating energy requirements are concentrated in (east) Asia, Europe and North America but are projected to decrease sharply with greater climate warming. The greatest absolute areas of cropland exposed to drought are in Asia and Europe. Large areas of maize and winter wheat across all continents see a reduction in crop growth duration of at least 10 d, but extensive areas with reductions in duration for soybean and spring wheat are concentrated in North America and Asia. Large proportions of the rice-growing
areas in Asia see reductions in duration of at least 10 d. The greatest proportion of maize-growing areas affected by damaging hot spells is in North America (USA), and an increasing frequency of damaging hot spells for soybean is most widespread in south Asia. Winter wheat growing areas are more affected by damaging hot spells than spring wheat areas (except for south Asia). The rice growing areas in west Africa, south and southeast Asia are most affected by hot spells; east Asia is little affected.

3.4. A summary: multiple indicators across a region

The previous figures have shown each indicator separately, across each continent (figures 5 and 8) and region (supplementary material). Figures 9 and 10 show all the hazard and impact indicators together for each continent (by region in supplementary material). This form of presentation allows an evaluation of how a region is affected by change in each indicator.
4. Discussion

There are, of course, several caveats with this study. The temperature scenarios incorporate current best estimates of uncertainty in climate sensitivity, carbon cycle feedbacks and ocean diffusivity, and these estimates change as more evidence becomes available. The climate scenarios were constructed using pattern-scaling and the delta method and making specific assumptions about disaggregating monthly climate data to the daily scale: other methods are available. It was assumed that each of the CMIP5 climate model...
patterns of change was equally plausible and independent, and can be matched with any increase in global mean temperature. A different ensemble of model runs—for example using higher resolution models—could give a different spread of results. The sea level rise scenarios use an empirical relationship between

**Figure 10.** Overview of continental impact indicators: 2100: RCP2.6, RCP4.5 and RCP8.5. SSP2 socio-economic scenario for the population indicators. The solid line represents the median and the shaded area the range between the 10th and 90th percentiles ('low' and 'high' for the coastal indicator). The axis limits for each indicator are shown at the top of each column. The axis limits vary across continents. For the cropland indicators the limits are the total regional continental cropland area, and for the river flood indicator the limits are the total regional river floodplain population.
accumulated temperature and sea level rise tuned to results from the IPCC AR5, and assume that sea level rise is globally-uniform. The study uses a series of indicators that represent the consequences of climate change for future hazards and socio-economic impacts, but do not describe actual impacts: these will depend on current and future adaptation decisions. Different global-scale studies have used different indicators of similar dimensions of hazard and impact, hindering comparisons of results between the different studies. The impacts as defined by the different socio-economic indicators are not directly comparable, because they are expressed—for practical reasons—in different units (numbers of people, numbers of people per year etc). Even where the units are the same, it is not necessarily straightforward to compare the magnitude of impacts. Is one person exposed to flooding equal to one person exposed to drought equal to one person exposed to heatwave, for example?

Nevertheless, the study has determined in a consistent way multiple indicators of hazards, resource base and impacts across regions, sectors, climate forcings and socio-economic scenarios. The supplementary material includes information presented in different formats. The study provides the foundation for more nuanced assessments of implications for the distribution of impact across regions and for the development of narrative storylines describing implications of climate change for, for example, resilience, supply chains and security. The study shows the wide uncertainty range in estimated changes in hazard and impact. For the water-related indicators, this is primarily due to uncertainty in the projected change in precipitation across space. This wide range, together with the observation that the estimated shape of the distribution of potential consequences in a year is not necessarily uni-modal, demonstrates that the selection of climate models for an assessment can have a major effect on the estimated range—and potentially even direction—in change.

5. Conclusions

This paper has used a consistent set of climate and socio-economic scenarios to present changes in a wide range of indicators of hazard and socio-economic impact at regional and global scales through the 21st century. It incorporates uncertainty in future emissions and socio-economic scenarios, alongside the effects of scientific uncertainty. It provides a quantitative foundation (in the paper and in supplementary material) both for risk assessments based on explicit categories of impact and for more nuanced qualitative assessments of the consequences of climate change using narrative characterisations of changes in key drivers such as governance and policy.

At the global scale, all the aggregated consequences of climate change considered here are adverse, with the exception of requirements for heating energy. However, the uncertainty range is large, primarily due to uncertainty in the projected regional change in precipitation, and the ‘high end’ consequences under each of the climate forcings considered can be much greater than the median estimate. For example, by 2100 and under RCP8.5 the median estimate of the global average proportion of time in hydrological drought is 27%, but the high-end estimate is 36%; the median estimate of the global average return period of the current 50 year flood is 14 years (7% likelihood), but the high-end estimate is nine years (11% likelihood).

The uncertainty range for a few of the precipitation-related indicators (floods and droughts) in some regions includes both adverse and beneficial consequences, reflecting regional variability in the direction of change in rainfall. There is a clear difference in hazard and impact under the different levels of climate forcing, with the difference varying across the indicators. Even under the lowest forcing there are substantial changes in hazard and impact by 2100. The assumed future socio-economic scenario has a very large effect on the estimated impacts of a given level of forcing, and for most socio-economic indicators the range in impact is greater across the socio-economic scenarios than the climate forcings. This highlights the strong dependence of future impacts of climate change on socio-economic change.

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Data availability statement

Any data that support the findings of this study are included within the article and in online supplementary material.
References

Alfieri L et al 2017 Global projections of river flood risk in a warmer world Earths Future 5 171–82
Arnell N W et al 2013 A global assessment of the effects of climate policy on the impacts of climate change Nat. Clim. Change 3 512–9
Arnell N W and Gosling S N 2016 The impacts of climate change on river flood risk at the global scale Clim. Change 134 387–401
Arnell N W and Lloyd-Hughes B 2014 The global-scale impacts of climate change on water resources and flooding under new climate and socio-economic scenarios Clim. Change 122 127–40
Arnell N W et al 2016a Global-scale climate impact functions: the relationship between climate forcing and impact Clim. Change 134 475–87
Arnell N W et al 2016b The impacts of climate change across the globe: a multi-sectoral assessment Clim. Change 134 457–74
Arnell N W et al 2019 Global and regional impacts of climate change at different levels of global temperature increase Clim. Change (https://doi.org/10.1007/s10584-019-02464-z)
Bamber J L et al 2019 Ice sheet contributions to future sea-level rise from structured expert judgment Proc. Natl Acad. Sci. 116 11195–200
Bets R A et al 2018 Changes in climate extremes, fresh water availability and vulnerability to food insecurity projected at 1.5 °C and 2 °C global warming with a higher-resolution global climate model Phil. Trans. R. Soc. A 376 20160452
Byers E et al 2018 Global exposure and vulnerability to multi-sector development and climate change hotspots Environ. Res. Lett. 13 035012
Challinor A J et al 2016 Current warming will reduce yields unless maize breeding and seed systems adapt immediately Nat. Clim. Change 6 954–8
Church J A et al 2013 Sea level change Climate Change 2013: The Physical Science Basis ed T F Stocker et al (Cambridge: Cambridge University Press)
Clarke L et al 2014 Assessing transformation pathways Climate Change 2014: Mitigation of Climate Change. Contribution of Working Group III to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change ed O Edenhofer (Cambridge: Cambridge University Press)
Dellink R et al 2017 Long-term economic growth projections in the shared socioeconomic pathways Glob. Environ. Change 42 200–14
Dosio A et al 2018 Extreme heat waves under 1.5 °C and 2 °C global warming Environ. Res. Lett. 13 055006
Flato G et al 2013 Evaluation of climate models Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change ed T F Stocker (Cambridge: Cambridge University Press)
Forster P M et al 2013 Evaluating adjusted forcing and model spread for historical and future scenarios in the CMIP5 generation of climate models J. Geophys. Res. Atmos. 118 1139–50
Gosling S N and Arnell N W 2016 A global assessment of the impact of climate change on water scarcity Clim. Change 134 371–85
Goudjii S M, Silbey A M and Lobell D B 2013 Global crop exposure to critical high temperatures in the reproductive period: historic trends and future projections Environ. Res. Lett. 8 24041
Harrington L J et al 2018 How uneven are changes to impact-relevant climate hazards in a 1.5 °C world and beyond? Geophys. Res. Lett. 45 6672–80
Harris I et al 2014 Updated high-resolution grids of monthly climatic observations—the CRU TS3.10 dataset Int. J. Climatol. 34 623–42
Hinkel J et al 2014 Coastal flood damage and adaptation costs under 21st century sea-level rise Proc. Natl Acad. Sci. USA 111 3292–7
Hsiang S et al 2017 Estimating economic damage from climate change in the United States Science 356 1362–9
James R et al 2017 Characterizing half-a-degree difference: a review of methods for identifying regional climate responses to global warming WIREs Clim. Change 8 e457
Jones B and O’Neill B C 2016 Spatially explicit global population scenarios consistent with the shared socioeconomic pathways Environ. Res. Lett. 11 084003
Samir K C and Lutz W 2017 The human core of the shared socioeconomic pathways: population scenarios by age, sex and level of education for all countries to 2100 Glob. Environ. Change 42 181–92
King D et al 2015 Climate Change: A Risk Assessment (Cambridge: Centre for Science and Policy)
Kopp R E et al 2014 Probabilistic 21st and 22nd century sea-level projections at a global network of tide-gauge sites Earths Future 2 383–406
Lehner P et al 2018 Future risk of record-breaking summer temperatures and its mitigation Clim. Change 146 363–75
Lowe J A et al 2009 How difficult is it to recover from dangerous levels of global warming? Environ. Res. Lett. 4 014012
Lowe J A and Bernie D 2018 he impact of Earth system feedbacks on different levels of warming in the United States Proc. Natl Acad. Sci. USA 115 171
Naumann G et al 2018 Global changes in drought conditions under different levels of warming Geophys. Res. Lett. 45 3283–96
O’Neill B C et al 2016 The scenario model intercomparison project (ScenarioMIP) for CMIP6 Geosci. Model Dev. 9 3461–82
O’Neill B C et al 2017 The roads ahead: narratives for shared socioeconomic pathways describing world futures in the 21st century Glob. Environ. Change 42 169–80
O’Neill B C et al 2018 The benefits of reduced anthropogenic climate change (BRACE): a synthesis Clim. Change 146 287–301
Osborn T J et al 2016 Pattern-scaling using ClimGen: monthly-resolution future climate scenarios including changes in the variability of precipitation Clim. Change 134 353–69
Osborn T J et al 2018 Performance of pattern-scaled climate projections under high-end warming: I. Surface air temperature over land J. Clim. 31 5667–80
Rahi K et al 2017 The shared socioeconomic pathways and their energy, land use, and greenhouse gas emissions implications: an overview Glob. Environ. Change 42 153–68
Ruan A C, Phillips M M and Rosenzweig C 2018 Climate shifts within major agricultural seasons for +1.5 and +2 °C worlds: HAPPI projections and AgMIP modeling scenarios Agric. For. Meteorol. 259 329–44
Schleussner C F et al 2016 Differential climate impacts for policy-relevant limits to global warming: the case of 1.5 °C and 2 °C Earth Syst. Dyn. 7 327–51
Schleussner C F et al 2018 Crop productivity changes in 1.5 °C and 2 °C worlds under climate sensitivity uncertainty Environ. Res. Lett. 13 064007
Seneviratne S et al 2016 Allowable CO2 emissions based on regional and impact-related climate targets Nature 529 477–83
Shukla S and Wood A W 2008 Use of a standardized runoff index for characterizing hydrologic drought Geophys. Res. Lett. 35 L02405
Smirnov O et al 2016 The relative importance of climate change and population growth for exposure to future extreme droughts Clim. Change 138 41–53
Sutton R 2018 ESD Ideas: a simple proposal to improve the contribution of IPCC WG1 to the assessment and communication of climate change risks Earth Syst. Dyn. 9 1155–8

Tebaldi C and Lobell D 2018 Estimated impacts of emission reductions on wheat and maize crops Clim. Change 146 533–45

Tebaldi C and Wehner M F 2018 Benefits of mitigation for future heat extremes under RCP4.5 compared to RCP8.5 Clim. Change 146 349–61

Vafeidis A T et al 2008 A new global coastal database for impact and vulnerability analysis to sea-level rise J. Coast. Res. 24 917–24

van Vuuren D P et al 2014 A new scenario framework for climate change research: scenario matrix architecture Clim. Change 122 373–86

Vicente-Serrano S M et al 2010 A multi-scalar drought index sensitive to global warming: the standardised precipitation evapotranspiration index—SPEI J. Clim. 23 1696–711

Vousdoukas M I et al 2018 Global probabilistic projections of extreme sea levels show intensification of coastal flood hazard Nat. Commun. 9 2360

Winsemius H C et al 2016 Global drivers of future river flood risk Nat. Clim. Change 6 381–5