SILICA IN PROTOPLANETARY DISKS

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ABSTRACT

Mid-infrared spectra of a few T Tauri stars (TTS) taken with the Infrared Spectrograph on board the Spitzer Space Telescope show prominent narrow emission features indicating silica (crystalline silicon dioxide). Silica is not a major constituent of the interstellar medium; therefore, any silica present in the circumstellar protoplanetary disks of TTS must be largely the result of processing of primitive dust material in the disks surrounding these stars. We model the silica emission features in our spectra using the opacities of various polymorphs of silica and their amorphous versions computed from earth-based laboratory measurements. This modeling indicates that the two polymorphs of silica, tridymite and cristobalite, which form at successively higher temperatures and low pressures, are the dominant forms of silica in the TTS of our sample. These high-temperature, low-pressure polymorphs of silica present in protoplanetary disks are consistent with a grain composed mostly of tridymite named Ada found in the cometary dust samples collected from the STARDUST mission to Comet 81P/Wild 2. The silica in these protoplanetary disks may arise from incongruent melting of enstatite or from incongruent melting of amorphous pyroxene, the latter being analogous to the former. The high temperatures of $\sim 1200$–$1300$ K and rapid cooling required to crystallize tridymite or cristobalite set constraints on the mechanisms that could have formed the silica in these protoplanetary disks, suggestive of processing of these grains during the transient heating events hypothesized to create chondrules.

Key words: circumstellar matter -- infrared: stars -- stars: pre-main sequence -- planetary systems: protoplanetary disks

Online-only material: color figures

1. INTRODUCTION

Silica (chemical name for the quartz group of minerals and a synonym for silicon dioxide) is a major ingredient of the Earth’s crust. The basic structural unit of crystalline silica is the tetrahedron, with the oxygen atoms at the corners and the silicon atom in the center. While in crystalline silica, structurally known as a tectosilicate, all oxygen atoms of a given tetrahedron are shared with the adjacent tetrahedra forming a three dimensional network structure, this does not hold for most other silicates. In silicates, $\text{SiO}_4^{4-}$ tetrahedra are coordinated by metal ions such as $\text{Mg}^{2+}$ and $\text{Fe}^{2+}$; examples of silicates are pyroxenes ($\text{Mg}, \text{Fe}_2\text{SiO}_4$) and olivines ($\text{Mg}, \text{Fe}_2\text{SiO}_4$). The spectra of both silica and silicate dust have vibrational spectral features near 10 $\mu$m (due to the $\text{Si}$–$\text{O}$ stretching mode) and $\sim 20$ $\mu$m (due to the $\text{O}$–$\text{Si}$–$\text{O}$ bending mode). For silica, the main features are at $\sim 9$, $\sim 12.6$, and $\sim 20\mu$m, with a minor feature at $\sim 16\mu$m for one of the polymorphs of silica, cristobalite, although Swainson & Dove (1993) note that the $16\mu$m feature of $\beta$-cristobalite is present at room temperature, but grows weaker with increasing temperature until it vanishes at $\sim 525$ K. These are the features that we use to identify the presence of silica grains in our spectra of protoplanetary disks, which are disks of dust and gas thought to be raw material from which larger bodies, including planets, form.

Silica has been used to determine the temperature and pressure conditions at the time of formation in terrestrial rocks (Tuttle & Bowen 1958) and also in enstatite chondrite meteorites (Binns 1967; Dodd 1981). These conditions dictate how silica crystallizes into its various polymorphs (Heaney et al. 1994; Hemley et al. 1994). Generally, for pressures of 1 atm or lower, silica crystallizes as $\alpha$-quartz at temperatures below 846 K; as $\beta$-quartz for temperatures between 846 K and 1140 K; as tridymite between 1140 K and 1743 K; as cristobalite between 1743 K and 2000 K; and as a liquid of silica composition above 2000 K. For a phase diagram illustrating these ranges of stability of silica polymorphs, see Figure 1 in Chapter 1 of Heaney et al. (1994). At higher pressures and a wide range of temperatures, coesite, and then stishovite form. All of the crystalline forms of silica are similar in that their bulk composition is $\text{SiO}_2$ and each of the oxygen atoms of every $\text{SiO}_4^{4-}$ tetrahedron is shared with the adjacent tetrahedron; however, the polymorphs differ in their arrangement of the $\text{SiO}_4^{4-}$ tetrahedra in crystalline networks. If silica is not allowed sufficient time to crystallize into one of its polymorphs, it will form as amorphous $\text{SiO}_2$. In principle, one can estimate the formation conditions (temperature and pressure) of the silica dust in astronomical environments based on the polymorph(s) of silica which can be inferred from their infrared (IR) spectra.

Silica cannot be an abundant interstellar dust species since the expected $\sim 9$ and $\sim 12.6\mu$m bands are not seen in the interstellar medium (ISM; see Li & Draine 2002). The presence of silica dust grains in the protoplanetary disk around the T Tauri stars (TTS) Hen 3–$600$ A was indicated in its 8–13 $\mu$m spectrum (Honda et al. 2003).

In this work, we report the detection of silica dust grains in the disks around five TTS from the prominent emission
features characteristic of submicron silica grains. These emission features are generated in the optically thin uppermost layers of protoplanetary disks. We model the spectra obtained with the Infrared Spectrograph (IRS; Houck et al. 2004) on board the Spitzer Space Telescope (Werner et al. 2004) with silica as a key dust component, with the aim to constrain the polymorph(s) of the silica dust in those disks and, therefore, the formation conditions (temperature and pressure history) of this dust.

2. DATA REDUCTION

2.1. Observations

The five TTS are described in Table 1. All were observed using both orders of the two low spectral resolution modules of the IRS, short-low (SL) and long-low (LL), using the IRS staring mode. For more information on the spectral modules and on the staring mode, we refer the reader to Sargent et al. (2006). These stars were chosen for our study of silica features because all five stars—1RXS J161410.6--230542, IRS 49, ROXs 42C, T51, and ZZ Tau—are Spectral Energy Distribution (SED) Class II young stellar objects; have prominent features at 9, 12.6, and 20 μm wavelengths attributable to silica and often 16 μm features partly attributable to silica, and have high-quality spectra with high signal-to-noise ratios.

2.2. Extraction and Calibration of Spectra

We obtained basic calibrated data (BCD; flat-fielded, straylight-corrected, dark-current-subtracted) products from the Spitzer Science Center (SSC) for each of the targets in our sample from the S15.3.0 IRS data calibration pipeline for IRXS J161410.6--230542 and from the S14.0.0 IRS data calibration pipeline for IRS 49, ROXs 42C, T51, and ZZ Tau. First, we identified bad pixels as described by Watson et al. (2007) in their explanatory supplement. In addition, we also identified as bad pixels a string of pixels with large values of e− s−1 in the BCD data for the first nod of the first order SL observation of ZZ Tau from Campaign 4. Some of the affected pixels were flagged as “radhit” detections (most likely a cosmic ray hit) in the associated bmask; we marked as bad pixels these and additional neighboring pixels that also appeared to be affected by the cosmic ray. As 12.5 μm is about the peak wavelength of a key diagnostic silica feature, fixing these pixels was critical for our purposes of analyzing dust emission features of silica. We correct for bad pixels by interpolation of neighboring pixels in the spectral direction, as described by Watson et al. (2007). First, we take the average of the pixel one pixel above and one pixel below. This is applied to all bad pixels in the BCD data. If a given pixel is not fixed, we then take another iteration; this time, we take the average of the pixel one pixel above and two pixels below the bad pixel. This process is repeated, with subsequent iterations going two pixels above and one pixel below, then two pixels above and two below, then finally one more iteration going one pixel above and one below.

The spectra presented in this paper were extracted from the BCD data using the Spectral Modeling, Analysis, and Reduction Tool (SMART; Higdon et al. 2004) via methods described by Sargent et al. (2006) and Furlan et al. (2006). For all targets in our sample, we subtracted sky for SL and LL by subtracting the pixel-fixed BCD data for the same order but other nod position. We extracted a spectrum each DCE separately, then used the extracted spectrum from each DCE to compute, first, the average spectrum and, second, its uncertainties from the standard deviation of the mean. Wavelength calibration of our data, tapered column extraction for each order of SL and LL, and flux density calibration of spectra using relative spectral response functions (RSRFs) was achieved as described by Sargent et al. (2006). The same high spectral resolution template of α Lac (A1 V; M. Cohen 2004, private communication) used by Sargent et al. (2006) was used to generate RSRFs for SL and LL flux calibration for IRS 49, ROXs 42C, T51, and ZZ Tau, and for SL flux calibration for IRXS J161410.6--230542. For the LL flux calibration of IRXS J161410.6--230542, we applied LL second order, LL bonus order, and LL first order (from its short-wavelength end to 36 μm wavelength) RSRFs generated from data of ξ Dra and the template for ξ Dra from Cohen et al. (2003); past 36 μm, we applied a LL first order RSRF generated from data of Markarian 231 and the template for Markarian 231 (J. Marshall 2006, private communication; Marshall et al. 2007; Armus et al. 2007).

We note the Humphreys-α and [Ne II] lines at 12.37 and 12.81 μm, respectively, are located near the central wavelengths of the two narrow features constituting a double peak in the opacity profile of α-quartz (see discussion later in Section 3.1). Sargent et al. (2006) concluded α-quartz was present in the IRS spectrum of TW Hya based upon two narrow emission features located at the wavelengths of the two peaks of the α-quartz opacity profile used in that study. Upon detailed inspection of the Short–High spectrum of TW Hya presented by Uchida et al. (2004), we find unresolved emission lines from Humphreys-α and [Ne II] and no underlying emission from the α-quartz double feature. We have looked at the Short–High spectra from Staring Mode observations of IRXS J161410.6--230542 (AOR ID# 5453824) and ROXs 42C (AOR ID# 6369792) for these lines. The AORs had two DCEs per nod position and two nod positions; in these Short–High data, we found no emission from Humphreys-α or [Ne II].

2.3. Extinction Correction of Spectra

For all objects except the highly extinguished (AV = 10.7) IRS 49, we corrected our RSRF-calibrated spectra in the same manner as described by Sargent et al. (2006), having adopted values of extinction at the V band, AV, as listed in Table 1. For
IRS 49, extinction correction was accomplished via the same method, and also using a lower $\tau_{0.7}$, extinction correction was accomplished via the same method. Due to loss of signal from dispersion-direction mispointing. To account for mispointing along the slits (the cross-dispersion direction) by adjusting the positions of our 3–5 pixel wide extraction boxes. The degree of mispointing is typically much less than 1 pixel (1.8′′ in SL), and we adequately account for this by using the BCD flatfielding. Mispointing in the dispersion direction is harder to detect and causes a loss of signal at every wavelength. We assume that the flux density from one nod being systematically lower than that from the other nod was due to loss of signal from dispersion-direction mispointing. To correct for this signal loss, we multiplied the nod lower in flux density by a scalar to minimize the squared differences in flux at each wavelength in the order. For more about correction for mispointing, see the Appendix.

2.4. Mispointing

We account for mispointing along the slits (the cross-dispersion direction) by adjusting the positions of our 3–5 pixel wide extraction boxes. The degree of mispointing is typically much less than 1 pixel (1.8′′ in SL), and we adequately account for this by using the BCD flatfielding. Mispointing in the dispersion direction is harder to detect and causes a loss of signal at every wavelength. We assume that the flux density from one nod being systematically lower than that from the other nod was due to loss of signal from dispersion-direction mispointing. To correct for this signal loss, we multiplied the nod lower in flux density by a scalar to minimize the squared differences in flux at each wavelength in the order. For more about correction for mispointing, see the Appendix.

2.5. Uncertainties

Uncertainty in flux density was determined by the standard deviation from the mean for the scalar-multiplied spectra from all DCEs. For IRS 49, ROXs 42C, and T51, there were only two DCEs (the two nod positions), so for these the standard deviation from the mean equals half the difference of the spectra at the two nod positions. Also, for the reasons listed by Sargent et al. (2006), we set relative uncertainties lower than 1% to 1%.

2.6. Photosphere Subtraction

Three of the five objects whose spectra are analyzed in this study, 1RXS J161410.6−230542, ROXs 42C, and ZZ Tau, have SEDs indicative of very little mass in small dust grains in the innermost disk regions. The SEDs of these three objects (Silverstone et al. 2006; M. K. McClure et al. 2008, in preparation; Furlan et al. 2006) have optical and $JHK$ photometry that is consistent with stellar photospheric emission, with ZZ Tau photospheric even to the $L$ band (∼3.5 μm). Also in each of these three cases, the excess above stellar photosphere as seen from the IRS spectrum at 5 μm wavelength is very small. Therefore, photosphere subtraction is important for these spectra in order to isolate the emission from circumstellar dust (see Sargent et al. 2006) at the short wavelength end of the IRS spectra. The low extinction to these three objects (see Table 1) gives us confidence that we can reasonably account for stellar photosphere emission in the IRS spectra. Therefore, for each we subtract a Planck function at the stellar effective temperature scaled to equal the flux at the dereddened $J$-band flux of each object. At Spitzer IRS wavelengths (> 5 μm), we are effectively subtracting the Rayleigh–Jeans tail of these Planck functions. The solid angles and temperatures used for the Planck functions representing stellar photosphere emission are listed in Table 1. For the stars T51 and IRS 49, there is significant excess at all IRS wavelengths, so no stellar continuum was subtracted.

3. ANALYSIS

3.1. Silica Opacities

Spitzer & Kleinman (1961) and Wenrich & Christensen (1996) measured the complex dielectric function $\varepsilon(\lambda) = \varepsilon_1(\lambda) + i\varepsilon_2(\lambda)$ for α-quartz at “room temperature” (∼300 K); Gervais & Piriou (1975) obtained $\varepsilon(\lambda)$ for α-quartz at 295 K, 505 K, 620 K, 785 K, and 825 K as well as for β-quartz at 860 K and 975 K.

To compare these dielectric functions, we calculate, for the sake of illustration, the opacities for α-quartz with the same grain shape distributions of either CDE (Bohren & Huffman 1983) or CDE2 (Fabian et al. 2001), using the dielectric functions from Spitzer & Kleinman (1961) for quartz at ∼300 K, Wenrich & Christensen (1996) for quartz at ∼300 K, and Gervais & Piriou (1975) for quartz at 295 K and taking the “$2 \pi$−1−approximation” for the ordinary and extraordinary rays (see Section A in Sargent et al. 2006). We find that the opacities computed from the three sets of dielectric functions are nearly identical. This implies that we can have confidence not only in the optical properties of α-quartz at room temperature, but also in the dielectric functions of α-quartz and β-quartz reported by Gervais & Piriou (1975) at other temperatures.
We show opacity curves for CDE and CDE2 shape distributions using the dielectric function from Gervais & Piriou (1975) for \( \alpha \)-quartz at 295 K in Figure 1. We take the mass density of \( \alpha \)-quartz to be 2.65 g cm\(^{-3} \) (Heaney et al. 1994). In the literature, the CDE and CDE2 grain shape distributions have been widely considered to explore the effects of dust shapes on the opacity profiles. Both CDE and CDE2 assume that the size of the grains under consideration are in the Rayleigh limit; i.e., \( 2\pi a/\lambda \ll 1 \). Thus, the grain sizes for CDE and CDE2 shape distributions used in our models of mid-infrared emission are submicron. The CDE shape distribution (Bohren & Huffman 1983) assumes a distribution of ellipsoidal shapes in which each shape, defined by the axial ratios of the ellipsoids, is statistically weighted equally. The CDE2 shape distribution (Fabian et al. 2001) assumes a distribution of ellipsoidal shapes in which ellipsoids that are almost spherical are weighted more heavily than more extreme (highly elongated along one or two axes or highly flattened along the other axes or both) shapes; the most extreme shapes in CDE2 are given zero weight. CDE is, therefore, more heavily weighted toward very nonspherical ellipsoids than CDE2. Our results confirm what others have shown in the literature (e.g., see the opacities of spheres, CDE, and CDE2 for forsterite by Fabian et al. 2001)—shape distributions of ellipsoids that more heavily weight toward very nonspherical ellipsoids with respect to spheres result in opacity profiles with features centered at increasingly longer wavelengths.

To approximate the effects of quartz of a range of temperatures, in Figure 2 we show the \( \beta \)-quartz opacity profile obtained by averaging the opacity curves of \( \beta \)-quartz of 975 K and 860 K. We take the mass density of \( \beta \)-quartz to be 2.53 g cm\(^{-3} \) (Heaney et al. 1994). Similarly, we obtain the high-temperature \( \alpha \)-quartz profile from averaging the opacity of \( \alpha \)-quartz at 825 K and 785 K, and the lower-temperature \( \alpha \)-quartz profile from averaging the opacity of \( \alpha \)-quartz of 620 K and 505 K, all with a CDE shape distribution. Note the progression of the 12.6 \( \mu \)m feature complex from double-peeked to single-peeked as the quartz temperature increases (see Figure 2). We assume that a given polymorph of silica, once formed, can exist stably at temperatures and pressures within its stability field on the phase diagram or metastably outside the stability field at lower temperatures if the physical conditions change relatively rapidly. However, we assume a polymorph of silica becomes another polymorph if subjected to temperatures higher than allowed by its stability field in the silica phase diagram.

Infrared transmission spectroscopy by Plyusnina et al. (1970) and absorption measurements by Williams et al. (1993) for coesite and stishovite rule out these higher pressure polymorphs of silica in our sample of TTS disks. For the intermediate pressure polymorph, coesite, both Plyusnina et al. (1970) and Williams et al. (1993) show four features with widths similar to that of the 12.6 \( \mu \)m feature between 14 and 23 \( \mu \)m, but all with strengths greater than that of the 12.6 \( \mu \)m feature—one feature near 14 \( \mu \)m, a pair of slightly longer features at \( \sim 16-17 \mu \)m, and the strongest of the features at \( \sim 20-23 \mu \)m. The spectra with prominent silica features do often have a feature at \( \sim 16-17 \mu \)m, but it is single, and, in addition, such spectra do not have a narrow feature at 14 \( \mu \)m of greater prominence than the 12.6 \( \mu \)m feature. This rules out coesite. Stishovite can be more easily ruled out, as both Plyusnina et al. (1970) and Williams et al. (1993) show stishovite to lack a 9 \( \mu \)m feature, while the spectra of the five silica exemplars analyzed in this study have prominent 9 \( \mu \)m features.

To represent the two higher temperature, low pressure polymorphs of silica, tridymite and cristobalite, we use the opacity obtained from the transmission measurements of submicron annealed silica grains embedded in Potassium Bromide, KBr, by Fabian et al. (2000). This silica was formed by heating amorphous SiO\(_2\) grains at 1220 K for 5 hr, and was found to be mostly cristobalite but also partially tridymite. This is supported by the fact that the cristobalite presented by Plyusnina et al. (1970) and Rokita et al. (1998) have a 16 \( \mu \)m feature of prominence similar to but slightly less than that of the 12.6 \( \mu \)m feature, whereas for tridymite the 16 \( \mu \)m feature is almost nonexistent (see the tridymite absorbance shown by Hofmeister et al. 1992). The annealed silica presented by Fabian et al. (2000) has a 16 \( \mu \)m feature of prominence closer to that of “low cristobalite” as measured by Rokita et al. (1998). We also note that the opacity profile of the silica annealed by Fabian et al. (2000) at 1220 K for 5 hr bears a close resemblance to the IR spectrum of the products of annealing for 120 hr at 1300 K of silica formed from iron-bearing silicates by Hallenbeck et al. (1998). We present in Figure 3 the opacity of the annealed silica used in our models.

For amorphous SiO\(_2\), the dielectric functions at mid-IR wavelengths were measured by Henning & Mutschke (1997), Steyer et al. (1974), and Koike et al. (1989). In our dust modeling we adopt the dielectric functions of amorphous SiO\(_2\) at 300 K of Henning & Mutschke (1997). In addition, we consider obsidian, as measured by Koike et al. (1989). To compute the opacities of obsidian and amorphous SiO\(_2\), we take their mass densities to be 2.384 g cm\(^{-3} \) (Koike & Hasegawa 1987) and 2.21 g cm\(^{-3} \) (Fabian et al. 2000), respectively. We plot in Figure 3 also the opacities of obsidian and amorphous SiO\(_2\) in the CDE shape distribution. The shape distribution affects the opacities of obsidian and amorphous SiO\(_2\) less than it does for \( \alpha \)- and \( \beta \)-quartz; for this reason, we only use the opacities of obsidian and amorphous SiO\(_2\) with a CDE shape distribution, and we do not plot the opacities of these grains in the CDE2 shape distributions.
3.2. The KBr Effect

Embedding small grains in KBr is commonly used to measure the grain opacity in the IR. This technique has limitations—optical constants are not directly measured, so it is not possible to explore grain shapes, porosity, or size. The embedded grains may clump together, reducing the contrast in sharp features. Moreover, the index of refraction of KBr (∼1.5) can distort strong resonance features in the spectrum from how they would appear in a spectrum of the same grains in a vacuum (i.e., in a protoplanetary disk).

Annealed silica and crystalline pyroxene are the only grain components in our models which depend on a KBr-embedded measurement. We model the KBr effect on β-quartz (see Figure 4) to illustrate the probable effects of KBr on our annealed silica opacities. We compute the opacities of solid, small β-quartz grains in a KBr medium and in vacuum for both CDE and CDE2 shape distributions. We use the optical constants of β-quartz at 975 K computed by Gervais & Piriou (1975) as the optical properties of silica grains that are embedded in a KBr medium whose 6–28 μm optical constants we take from Heaton (1971). As shown in Figure 4, for the CDE shape distribution, the peak wavelengths of the 9, 12.6, and 20 μm features are little altered by embedding in KBr. The peak wavelength of the 9 μm feature for β-quartz in the CDE2 shape distribution embedded in KBr is at a slightly longer wavelength than that for β-quartz grains in the CDE2 distribution in vacuum. This also holds for the 20 μm feature. The peak wavelength of the 12.6 μm and 16 μm features are much less affected by either being in KBr or with a different shape distribution.

Note in Figure 4 that if the shape distribution for the dust in KBr is already at an extreme (e.g., CDE), models with the same shape distribution but in vacuum will have spectral features not much different (in width and central wavelength) from those in KBr. We note that all opacities of β-quartz are somewhat higher for dust in a KBr medium, also found by Jaeger et al. (1998).

3.3. Other Dust Opacities

For submicron amorphous silicates, we compute opacities for the CDE2 shape distribution (Fabian et al. 2001).

We take the dielectric functions of MgFeSiO₄ of Dorschner et al. (1995) for amorphous olivine. For amorphous pyroxene we use the dielectric functions of Mg₀.₅Fe₀.₅SiO₃ from Dorschner et al. (1995). We take the mass densities of both olivine and pyroxene to be 3.3 g cm⁻³.

To account for the effects of grain growth, we consider porous dust. Porous material is simulated by applying the Bruggeman effective medium theory (Bohren & Huffman 1983), assuming the porous grains are 60% vacuum by volume. The effective dielectric functions for porous amorphous pyroxene and olivine are then computed. Mie theory (Bohren & Huffman 1983) is then used to compute the opacities of spherical porous grains of radii of 5 μm. As noted by Sargent et al. (2006), the opacity of 5 μm radius, 60% porous amorphous silicate grains is similar to that for 2 μm radius solid dust, the size of dust used by Bouwman et al. (2001) to investigate grain growth in Herbig Ae/Be disks.

We also include pyroxene (crystalline) and forsterite (of which the density is also 3.3 g cm⁻³) in our analysis. For pyroxene, we use the opacity of “En90” (Chihara et al. 2002). With a stoichiometric composition of Mg₀.₅Fe₀.₅SiO₃, “En90” is quite effective in fitting the pyroxene features at 9.3, 10.5, 11.2, and 11.6 μm within the 10 μm complex of FN Tau (see Sargent et al. 2006). For forsterite, we use the optical constants of Sogawa et al. (2006) which are similar to that of Servoin & Piriou (1973), to compute the opacity for a distribution of ellipsoidal shapes.

Using the formalism of Bohren & Huffman (1983), we have constructed an opacity profile for a distribution of forsterite grains in which all ellipsoidal axial ratios are given an equal weight, as in CDE. However, we exclude the most extreme ellipsoidal axial ratios in the following manner.

Our shape distribution, which we call “CDE” (truncated Continuous Distribution of Ellipsoids), is that we confine the L₁ and L₂ (which are the L_j parameters described by Bohren &
Figure 5. Model fit to ROXs 42C using annealed silica, minimizing chi-squared from 6 to 37 μm, comparing fit including forsterite in the tCDE (black solid line; see text for description of tCDE) and CDE (orange dashed line) shape distributions. The model was constructed as described in Section 3 using parameters listed in Table 2. All non-silica model components for the model using forsterite grains in the tCDE shape distribution are plotted as solid lines at bottom of graph. Model components from hot dust are plotted in red, and those from cold dust are plotted in blue, and those from cold dust are plotted in red. Silica components are plotted with a dashed line for the hot silica and a dotted line for the cold silica. Note the lack of the 6.2 μm PAH feature.

(A color version of this figure is available in the online journal.)

Huffman (1983) within the triangle in the \( L_1-L_2 \) space specified by vertices \( (L_1 = 0.1, L_2 = 0.05) \), \( (0.1, 0.895) \), and \( (0.99, 0.005) \). For CDE, all shapes are allowed: the triangle specified by the vertices \( (0, 0) \), \( (0, 1) \), and \( (1, 0) \). The opacity profile of forsterite for CDE has a feature peaking at 11.3 μm which is close to the feature seen in astronomical data (but often peaking at a slightly longer wavelength). The opacity profile for tCDE “rounds” this peak somewhat, pushing the peak toward slightly shorter wavelengths, and this results in marginally better fits to astronomical spectra.

We will discuss “tCDE” at a greater length in B. Sargent et al. (2008, in preparation), but here we justify this choice by showing in Figure 5 our model fits to ROXs 42C with forsterite grains having either a CDE shape distribution (“CDE” model) or a CDE distribution (“tCDE” model). The former (CDE) has \( \chi^2 \) per degree of freedom (d.o.f.) of \( \sim 5.1 \), while for the latter (tCDE) \( \chi^2 / \text{d.o.f.} \approx 4.3 \) which is somewhat better; the tCDE model does a slightly better job in fitting the 11 and 23 μm complexes of the spectrum of ROXs 42C but is overall very similar to CDE. Sargent et al. (2006) found great similarity between the opacity of forsterite and silica grains with a CDE shape distribution and that of porous dust with a porosity of 60%. Thus the effects of shape distribution and porosity cannot be clearly distinguished in the spectra.

No hint of polycyclic aromatic hydrocarbon (PAH) emission is seen in our spectra, except possibly ZZ Tau at 6.2 μm. Especially telling is the lack of the ubiquitous 6.2, 7.7 μm PAH features, which are clearly distinguishable from silicate or silica emission; in ZZ Tau, no 11.3 μm PAH emission is evident.

We note that the opacity profiles of some solids are similar to that of silica, but nevertheless can be ruled out. Kimura & Nuth (2007) measured the transmittance of \( \text{H}_2\text{Si}_2\text{O}_4 \) and \( \text{Si}_2\text{O}_3 \). They found in some of their spectra the features peak at wavelengths very close to those of silica (\( \sim 9, 12.6, \) and 21.5 μm); however, many of their spectra also show a narrow feature at 11.36 μm with a width similar to their 12.6 μm feature but of greater prominence. This feature is not seen in the spectrum of ZZ Tau. In the spectra of 1RXS J161410.6–230542 and ROXs 42C there is a feature near 11.1 μm, but as will be shown in our model fits to their spectra, this peak can be attributed to forsterite. Also, the opacity curves computed using the optical constants of Na-bearing amorphous aluminosilicates determined by Mutschke et al. (1998) exhibit features near 9 μm and 21 μm similar to those of silica of various versions; however, the Na-bearing aluminosilicates lack the distinctive 12.6 μm feature which is seen in the emission spectra of the 5 T Tauri stars and in the opacity spectra of all versions of silica.

3.4. Models

We model the spectra of each of five T Tauri stars as a sum of featureless continuum emission (from optically thick disk midplane, iron or amorphous carbon grains, and very large >10 μm silicate grains) and optically thin emission from dust grains with strong infrared resonances in the disk atmosphere (see Sargent et al. 2006). It is not our goal to model self-consistently the entire spectral energy distribution from the near-IR to submm; such would require radiative transfer models such as those of Calvet et al. (1992) and D’Alessio et al. (2001).

We reject silicate self-absorption, a mix of emission and absorption by similar kinds of dust but with cooler dust in front of warmer dust, as an explanation of the 9 and 12.6 μm features we see in our spectra. Self-absorption in the 10 μm silicate complex would give apparent emission features in the 8 and 12 μm wings of the 10 μm complex; however, such would require the optical depth of the cool amorphous silicate dust to exceed 1. Since amorphous silicate dust in the ISM gives \( A_V / T_{9.7} \) of 18 or 25 (see Section 2.3), this would mean the intervening dust has \( A_V \) of 18 to 25. Such dust would not be at the disk surface; otherwise it would heat up to the same temperature as surrounding dust and not rise to an absorption feature. Such dust would have to be further away from the disk and star, and therefore would also be in front of the star in the line-of-sight to the star, and we would, therefore, find \( A_V \) of 18 to 25 to all stars with what appear to be prominent 9 and 12.6 μm silica features. All of our objects have \( A_V \) less than 18, the minimum \( A_V \) required to result in self-absorption.

We approximate the IR emission from dust with a large range of temperatures by two temperatures. A similar approach was employed by Kastner et al. (2006) to model the IRS emission spectra of the dust around the hypergiant R66 in the Large Magellanic Cloud and by Chen et al. (2006) to model that of five debris disks (see especially the two dust temperatures used in the model for \( \eta \) Tel). We assume that, for a given protoplanetary disk, all dust grains, regardless of composition, within the same arbitrarily small volume of the part of the disk atmosphere giving rise to optically thin silicate and silica emission are at the same temperature. Kamp & Dullemond (2004) find in their disk models that, in a layer above the layer giving rise to dust emission features, the gas and dust temperatures are within 10% of each other. As gas density increases closer towards the disk midplane, the gas and dust temperatures in the layer giving rise to dust emission features should be even closer to each other. The dense gas effectively should equalize the temperatures of dust grains of different optical properties (resulting from, e.g., different compositions, different sizes). Our two-temperature models, therefore, model dust composition for two disk regions: the warmer inner disk regions and the cooler outer disk regions.
Our model is a sum of blackbodies at two temperatures plus optically thin emission from dust at those temperatures. Mathematically, these models are given by

\[ F_\nu(\lambda)^{\text{mod}} = B_\nu(T_c) \left[ \Omega_c + \sum_i a_{c,i} \kappa_i(\lambda) \right] + B_\nu(T_w) \left[ \Omega_w + \sum_j a_{w,j} \kappa_j(\lambda) \right], \]

(1)

where \( F_\nu(\lambda)^{\text{mod}} \) is the model flux density, \( B_\nu(T) \) is the Planck function at temperature \( T \) and wavelength \( \lambda \), \( T_c \) is the temperature of “cool” dust and \( T_w \) is the temperature of “warm” dust, \( \Omega_c \) (\( \Omega_w \)) is the solid angle of a blackbody representing the continuum emission from cool (warm) dust, \( a_{c,i} \) (\( a_{w,j} \)), the mass weight for cool (warm) dust, equals \( m_{c,i}/d^2 \) (\( m_{w,j}/d^2 \)), and \( d \) is the distance to the T Tauri star (see Table 1 for the stellar distances). We specify a grid of temperature pairs, \( T_c \) and \( T_w \), not allowing either of the temperatures to exceed 1401 K, as this temperature is intermediate between the stability limits of forsterite and enstatite of 1430 K and 1370 K, respectively (Posch et al. 2007). For a given temperature pair, our model equation is linear in solid angles and optically thin dust mass weights. We minimize

\[ \chi^2 = \sum_k \left[ \frac{F_\nu(\lambda_k)^{\text{irs}} - F_\nu(\lambda_k)^{\text{mod}}}{\Delta F_\nu(\lambda_k)^{\text{irs}}} \right]^2, \]

(2)

with respect to each of the solid angles and mass weights, where \( F_\nu(\lambda_k)^{\text{irs}} \) is the observed flux density at the \( k \)th wavelength band and \( \Delta F_\nu(\lambda_k)^{\text{irs}} \) is the uncertainty in the observed flux density at the \( k \)th wavelength band. If any of the mass weights computed from an iteration of \( \chi^2 \) minimization are negative, the component with the most negative integrated flux from 7.7 to 23 \( \mu \)m is set to zero, and we then minimize \( \chi^2 \) again. This process is iterated until there are no components with negative mass weights. The solution found in the iteration for which no components are negative is the best fit for a given temperature pair. The process is carried out for another temperature pair. The global best fit for a spectrum is at the temperature pair that minimizes \( \chi^2 \) per degree of freedom. In all of our best fits, the silica component at both low and high temperatures was positive. The number of degrees of freedom is equal to the number of data points minus four (since we consider two dust temperatures and two solid angles) minus the number of mass weights at both temperatures altogether.

Figure 5 is a sample fit of a model with annealed silica to the 7.7–37 \( \mu \)m IRS spectrum of ROXs 42C (all other models, including the other models of ROXs 42C, discussed or shown in this paper are fit over 7.7–23 \( \mu \)m). It is seen that our two-temperature model fits the spectrum of ROXs 42C over almost the entire wavelength span of the IRS. What interests us most in this study are the spectral features diagnostic of silica, the 9, 12.6, 16, and 20 \( \mu \)m features. For this reason, we minimize \( \chi^2 \) over the smallest range of wavelengths sufficient to adequately cover these features, 7.7–23 \( \mu \)m. We determine a very conservative estimate of the uncertainty for each nonzero dust component and solid angle in the model as the amount that resulted from increasing \( \chi^2 \) per degree of freedom by \( \sim 1 \) from the best-fit value. The models shown in all figures after Figure 5 and whose parameters are summarized in Table 2 are constructed by minimizing \( \chi^2 \) over 7.7–23 \( \mu \)m.

We obtained models of quartz in each of the CDE and CDE2 shape distributions (where the same shape distribution is used for quartz at the low temperature, \( T_w \), as for quartz at the high temperature, \( T_w \)). We used the opacity of the 295 K \( \alpha \)-quartz for the “cool” silica component. For the opacity of the “warm” silica component, we use the \( \beta \)-quartz, high-temperature \( \alpha \)-quartz, and lower-temperature \( \alpha \)-quartz opacity curves described in Section 3.1 in reference to Figure 2. The opacities computed from the dielectric constants of \( \alpha \)-quartz at 505 K and 620 K were averaged to give an “average 562.5 K \( \alpha \)-quartz opacity,” 785 K and 825 K \( \alpha \)-quartz opacities were averaged to give an “average 805 K \( \alpha \)-quartz opacity,” and 860 K and 975 K \( \beta \)-quartz opacities were averaged to give an “average 917.5 K \( \beta \)-quartz opacity.” The model with the lowest difference between \( T_w \) and “average \( \alpha \)–or \( \beta \)-quartz opacity temperature” (562.5 K, 860 K, or 917.5 K) is the best fit model for quartz grains in the shape distribution under consideration (only CDE and CDE2 were considered). A model is rejected if \( T_w \) is greater than the highest temperature allowed according to the stability field for a given silica polymorph (for silica polymorph stability fields, see Section 1).

4. RESULTS

4.1. ROXs 42C

ROXs 42C was identified by Bouvier & Appenzeller (1992) as the eastemmost stellar source within the positional error of the X-ray source ROX 42 found by the Einstein satellite (Montmerle et al. 1983). Mathieu et al. (1989) found it to have a double-lined spectroscopic binary with a 36 day period, while Ghez et al. (1993) found it also to be a visual binary separated by \( \sim 0.16 \); Jensen & Mathieu (1997) conclude that ROXs 42C is a hierarchical triple, a spectroscopic binary orbited by single star further out. Ratzka et al. (2005) agree with the assessment that ROXs 42C is a hierarchical triple, but report that the visual binary is now separated by \( 0.277 \). Assuming a distance to Ophiuchus of 140 pc (see Table 1), this implies a separation of \( \sim 39 \) AU between the double-lined spectroscopic binary and the single star. Jensen et al. (1996) calculate the separation of the double-lined spectroscopic binary to be 0.27 AU. We note the similarity between ROXs 42C and Hen 3–600 itself, a spectroscopic binary wit another star at 70 AU orbiting the binary. We also note that the disk in the Hen 3–600 system also has abundant silica as indicated by its IR spectrum (Honda et al. 2003; Uchida et al. 2004). The state of accretion in the disk(s) in this system is unknown, as Bouvier & Appenzeller (1992) report the equivalent width of \( \text{H-}\alpha \) to be neither in absorption or emission, but such that the photospheric line was filled in by 1.6 Å of emission. However, because of the multiple stars in this system, and because it is not known where the disk(s) is (are) located, we do not know whether active accretion could be occurring currently onto one star and its spectral signature is obscured, or whether there is no active accretion in this system.

The IRS spectrum of ROXs 42C has prominent narrow features at 9 and 12.6 \( \mu \)m from silica, in addition to features around 11, 23, 28, and 33.5 \( \mu \)m identified with forsterite and pyroxene. We show the model fits to ROXs 42C using each of annealed silica, obsidian, amorphous SiO\(_2\), and quartz grains in the CDE and CDE2 shape distributions in Figure 6.

We rule out the models using the amorphous forms of silica—obsidian and amorphous SiO\(_2\)—because the 12.3–12.4 \( \mu \)m
features of these two silicas are of much less strength with respect to their 9 \( \mu \)m features than for the crystalline forms of silica. This translates into being able to fit the 9 and 20 \( \mu \)m features somewhat reasonably but not having enough strength in the 12.3–12.4 \( \mu \)m feature to fit the data. In addition, the 12.3–12.4 \( \mu \)m feature from the amorphous versions of silica is centered shortward of the feature in the data by \( -0.2 \) \( \mu \)m and is noticeably wider, with a full width at half-maximum (FWHM) of ~1 \( \mu \)m, than that in the data (with a FWHM \( \sim 0.5 \) \( \mu \)m). The \( \chi^2 \) per degree of freedom for the model using obsidian is 5.4; for the model using amorphous SiO\(_2\), it is 5.2.

Next, consider \( \alpha \)-quartz, \( \beta \)-quartz, tridymite and cristobalite (tridymite and cristobalite being represented together by annealed silica). For quartz in the CDE2 shape distribution, the lowest difference between model \( T_{\alpha} \) and “average opacity temperature” (temperature at which quartz optical constants were measured) happened for \( \beta \)-quartz. A \( \chi^2 \) per degree of freedom of 4.5 was obtained for this model. The fit to the 9 \( \mu \)m feature is good, coming from a combination of \( \beta \)-quartz and pyroxene. Also, the model is within the error bars of a little over half the data points between 20.5 and 22 \( \mu \)m. However, the spectrum of ROXS 42C lacks the double feature at 12.6 \( \mu \)m in the model, and the model does not have a strong 16 \( \mu \)m feature like ROXS 42C. Forsterite gives a 16 \( \mu \)m feature, but forsterite is insufficient by itself to fit the 16 \( \mu \)m feature in ROXS 42C.

A \( \chi^2 \) per degree of freedom of 5.0 was obtained for the model using quartz in the CDE shape distribution. The fit to the 12.6 \( \mu \)m feature is improved due to the feature being single-peaked in higher-temperature \( \alpha \)-quartz. However, this model provides a poor fit to 20.5–22.0 \( \mu \)m. The model feature at 9 \( \mu \)m peaks slightly longward of the feature in the data. A shape distribution intermediate between CDE and CDE2 could fit the 9 \( \mu \)m feature, but would still provide a 20 \( \mu \)m feature that would overshoot the ROXS 42C spectrum between 20.5 and 22.0 \( \mu \)m. Furthermore, the model does not have a 16 \( \mu \)m feature sharp enough to fit the 16 \( \mu \)m feature in the data.
believe this definitely rules out $\beta$-quartz as a candidate to explain the silica features of ROXs 42C; $\alpha$-quartz at temperatures of 505 K–620 K is a possibility to explain the features, but we believe it to be of lower probability.

This leaves annealed silica. We show the model using annealed silica at low and high model temperatures in Figure 6. A $\chi^2$ per degree of freedom of 3.2 was obtained for this model, the lowest for any of the models of ROXs 42C. The fit to the 9, 12.6, 16, and 20 $\mu$m features and the 20.5–22.0 $\mu$m region is the best of all models fit to ROXs 42C, and this is reflected in the $\chi^2$ per degree of freedom (Table 2 and Figure 7). We note the opacity of annealed silica that we use comes from KBr pellet transmission measurements, but, as discussed in Section 3.2, we expect that the opacity curve for the annealed silica grains in KBr would be very similar in terms of spectral feature central wavelength and width to the opacity in vacuum. We conclude that annealed silica best matches the spectrum of ROXs 42C.

4.2. ZZ Tau

Simon et al. (1996) found ZZ Tau to be a binary system with a separation of stars that increased from 0′0338 in 1994 to 0′0420 in 1996. Subsequent measurements of the positions of the stars in the binary let Schaefer et al. (2006) solve for the orbital parameters of this system. Schaefer et al. (2006) compute the semi-major axis expressed as an angle of 0′061, implying a separation of 8.5 AU. Prato & Simon (1997) report EWH from this system of 0.0222 separated from the primary and is 0.21 magnitudes fainter at the K band (Metchev 2006). At an assumed distance to this object of 145 pc (Table 1), this implies a separation of 32 AU between the components. Preibisch et al. (1998) report EWH from IRXS J161410.6–230542 of 0.96 Å, which, when combined with the K0 spectral type assigned to this system by Pascucci et al. (2007), suggests little or no accretion is occurring from the disk to the star in the IRXS J161410.6–230542 system.

Annealed silica provided the best fit to the Spitzer IRS spectrum of IRXS J161410.6–230542. As seen at the bottom of Figure 8, the fits to the 9, 12.6, 16, and 20 $\mu$m features are all adequate with our model. The $\chi^2$ per degree of freedom is 6.6 for this model. Of all the versions of silica discussed in this paper, only annealed silica has a single, isolated peak at 16 $\mu$m. The 16 $\mu$m feature in our model is coming equally from forsterite and cold annealed silica, and the model fits the narrow (∼1 $\mu$m FWHM) peaks in the spectrum fairly well at 11 and 19 $\mu$m, which are features associated with forsterite, meaning the fit to the 16 $\mu$m feature would be worse without the contribution from annealed silica.

We note that the curious trapezoidal shape of the 12.6 $\mu$m feature does not perfectly match the rounded shape of the 12.6 $\mu$m annealed silica feature in our model, but we also note the size of the errorbars in the data at the top of the 12.6 $\mu$m feature do not preclude a more rounded shape like the feature belonging to annealed silica. The 12.6 $\mu$m complex of $\alpha$-quartz does evolve from double-peaked at 295 K to somewhat trapezoidal and eventually single-peaked at progressively higher temperatures of 505 K, 620 K, 785 K, and 825 K; the feature is single-peaked and centered at 12.7 $\mu$m for $\beta$-quartz. As opposed to annealed silica, $\alpha$- and $\beta$-quartz in the CDE shape distribution do not fit the 20 $\mu$m feature, and both quartzes in the CDE2 shape distribution do not fit the 9 $\mu$m feature (see discussions in Sections 4.1 and 4.2 on ROXs 42C and ZZ Tau, respectively). Further, models with annealed silica fit the 16 $\mu$m features of IRXS J161410.6–230542 and ROXs 42C noticeably better than does quartz in either the CDE or CDE2 shape distributions.

4.4. IRS 49

IRS 49, also known as WLY 2–49, has been searched for multiplicity but has not been found to be multiple (Barsony et al. 2003; Ratzka et al. 2005). Ratzka et al. (2005) established upper limits on the relative brightness at the K band of any unseen companion around IRS 49 of 0.04 for 0.15 separation and 0.02 for 0.50 separation. Gatti et al. (2006) report significant accretion of $M = 1.3 \times 10^{-8} M_{\odot}$ yr$^{-1}$ from Paschen-β and Brackett-γ emission.

IRS 49 is in the Ophiuchus star-forming region; it has a sizeable extinction of $A_V$ of 10.7 (Table 1). How extinction at visible wavelengths translates to extinction at mid-infrared
Figure 7. Comparison of chi-squared per degree of freedom for dust emission models using different types of silica for the spectra of the five objects modeled, including the dereddened spectra obtain by two different computations of optical depth at 9.7 μm of the material in front of IRS 49. The type of quartz—β-quartz and an average of α-quartz at 785 K and 825 K—is denoted by β and αH, respectively. (A color version of this figure is available in the online journal.)

Figure 8. Best fits to ZZ Tau (top) and 1RXS J161410.6−230542 (bottom) using annealed silica. Same color and linestyle scheme for model components as for Figure 5. The dominant non-continuum dust component in the 8–12 μm region of ZZ Tau is large amorphous pyroxene. (A color version of this figure is available in the online journal.)

wavelengths is still uncertain (see discussion by Draine 2003 on values of AV/τ9.7 determined from various studies); for this reason, we present fits to the spectrum of IRS 49 dereddened assuming AV/τ9.7 = 18 (Figure 9, top, brown points and solid black line; χ² per degree of freedom of 5.5) and assuming AV/τ9.7 = 25 (Figure 9, top, orange points and solid violet line; χ² per degree of freedom of 6.3). As discussed by Sargent et al. (2006) with regard to the dust model of V410 Anon 13, greater extinction correction of at mid-infrared wavelengths results in spectral profiles that look more like the interstellar medium profile; namely, more like amorphous silicates. The effect of extinction correction on the silica features is most important for the 9 and 20 μm silica features, but, as can be seen from Table 2 and Figure 7, χ² per degree of freedom is lowest for both extinction corrections using models with annealed silica.

4.5. T51

Lastly, we consider T51. Also known as Sz 41, this object was found to be double by Reipurth & Zinnecker (1993). Another star, named Sz 41C, had also been thought to be associated with Sz 41, but was shown to be a background star by Walter (1992). This is noted by Correia et al. (2006), whose observations show the Sz 41 binary to be separated by 1"974. Guenther et al. (2007) assign EWHα of −2.0 Å, i.e., absorption. Spectroastrometry conducted by Takami et al. (2003) determine that the Hα profiles from the two components of the Sz 41 binary are roughly equivalent, suggesting there is negligible accretion in the disk(s) of the Sz 41 system.

We fit models with different silicas to T51. The spectrum of T51 includes a very prominent peak at 9 μm, a triangular 20 μm feature, a modest 12.6 μm feature, and a slight 16 μm feature. Once again, the model with annealed silica fit the dust excess spectrum the best (see Figure 9, bottom; χ² per degree of freedom of 3.6). However, we note that of all objects whose spectra were analyzed in this study, T51 provides the weakest tests discriminating between the various versions of silica (see Figure 7). The models using obsidian and amorphous SiO₂ are not much worse than the model using annealed silica.
5. DISCUSSION

5.1. Cristobalite Dominance

All the models (see Table 2) indicated the presence of both amorphous and crystalline silicates and grains grown to larger sizes as well as silica, suggesting production of silica accompanies both dust processing and grain growth. Models using the opacity of annealed silica provided the best fit of all models for all five objects whose spectra we analyzed. From our discussion of the KBr effect in Section 3.2, we believe the models for all five objects whose spectra we analyzed. From using the opacity of annealed silica provided the best fit of all.

5.2. Silica in the Solar System

Dodd (1981) reports that tridymite and cristobalite are found in chondritic meteorites (see also Binns 1967); more specifically, they were found in enstatite chondritic meteorites, which are characterized by calcium-aluminum inclusions (CAIs; Guan et al. 2000). Dodd (1981) also summarizes findings on euclidian meteorites. Euclides are achondrites, which means they lack chondrules. Chondrules are spheroidal inclusions in chondritic meteorites, are submillimeter to centimeter in diameter, and have the appearance of having formed as molten drops (Norton 2002). The findings show the volume percentage of silica (as quartz, tridymite, and cristobalite polymorphs) in these eucrites to be between 0% and 4%. In the review by Brearley & Jones (1998), it is mentioned that less than 2% of chondrules are silica-bearing chondrules, chondrules that are up to 40% by volume silica, and they have been observed in a number of ordinary chondrites.

Silica has not been widely reported as detected in the infrared spectra of comet comae (Hanner 2003; Hanner & Bradley 2004). Neither Comet Kohoutek (Merrill 1974) nor Comet Hale-Bopp (Harker et al. 2002) showed any hint of the 9 and 12.6 \( \mu \text{m} \) silica features. Spitzer IRS spectra of the comet nucleus ejecta from the Deep Impact mission to Comet Tempel 1 also show no hint of silica features, and models of emission from this dusty eject do not require silica for a good fit (Lisse et al. 2007). Bouwman et al. (2001) reported silica present based on the spectrum of Comet Halley; however, the data are not of high quality, being coarse and noisy, so the detection is tentative at best. As discussed by Harker et al. (2002), magnesium-rich crystalline silicates absorb stellar radiation poorly over visible and near-infrared wavelengths, where the Sun emits a large fraction of its radiation. Harker et al. (2002) argue that Mg-rich pyroxene should have optical constants very similar to those of amorphous Mg-rich pyroxene given by Dorschner et al. (1995). The imaginary part of the index of refraction is as low as \( \sim 0.0003 \) at 0.5 \( \mu \text{m} \) wavelength for amorphous Mg\(_{0.95}\) Fe\(_{0.05}\)SiO\(_3\), so we follow Harker et al. (2002) and take this value of \( k \) for Mg-rich crystalline pyroxene, which is found in comets (see Harker et al. 2002). Optical properties for the various silica polymorphs at visible and near-infrared wavelengths are not readily available in the literature (macroscopic rocks of quartz are often transparent or translucent). We estimate an upper limit on \( k \) for both ordinary and extraordinary rays of \( \alpha \)-quartz of \( 10^{-3} \) (see relevant discussion by Palik 1985) and for amorphous SiO\(_2\) of \( 10^{-5} \) (again see Palik 1985). This suggests that, at wavelengths over which the Sun emits most of its power, any silica in comets is either as poorly absorptive as Mg-rich pyroxenes, which are seen in comets, or is even less absorptive. Thus, silica may be present in comets but may be at temperatures too low to emit enough to be seen in mid-infrared spectra of comet comae.

A better approach to searching for silica from comets is to analyze physical samples of comet dust. In situ measurements of comet dust to date do not determine mineralogy, only the abundances of elements constituting comet dust. Such measurements determined the rocky material in the coma of Comet Halley to have solar, or chondritic, abundances of the major rock-forming elements Mg, Si, Ca, and Fe (Hanner 2003).

The STARDUST mission to Comet 81P/Wild 2 returned one several-micron sized grain (“Ada”) composed of tridymite and fayalite (Zolensky et al. 2006). Mikouchi et al. (2007) report a grain of crystalline silica that is either tridymite or cristobalite or both from Comet 81P/Wild 2. It remains to be seen how much silica will be inferred for this comet based upon analysis of the returned samples of its dust.

As Bradley (2003) summarizes, some Interplanetary Dust Particles (IDPs) are believed to originate from comets. Rietmeijer & McKay (1986) reported a 1.5 \( \mu \text{m} \) size grain of SiO\(_2\) in the phenocryst of the IDP W7010*A2 constituting less than 5% of the abundance of all minerals in the IDP. This SiO\(_2\) grain is of unknown polymorph (Rietmeijer & McKay 1985). Rietmeijer & McKay (1986) also reported SiO\(_2\) fragments of <10% abundance in the same IDP. These crystals are subhedral and \( \sim 0.1 \mu \text{m} \) in diameter (MacKinnon & Rietmeijer 1987). However, Bradley (2003) does not mention silica as a...
significant mineral phase in IDPs. Grün et al. (2001), in the discussion of micrometeorite mineralogy and petrography, do not list silica as a major phase of micrometeorites, which are IDPs that have landed on Earth near the poles.

In a study of a large number of Kuiper Belt Objects (KBOs) using near-infrared spectra, Barkume et al. (2008) found the spectra were well-fit by models assuming water ice and featureless spectral continuum from unknown material. However, other compositions are inferred from near-infrared spectra of KBOs, such as methane and ethane ice, pyroxenes and olivines, and possible organic material (Emery et al. 2007). It is unknown whether KBOs contain silica.

Cristobalite and tridymite commonly occur in terrestrial siliceous volcanic rocks, such as in rapidly-cooled obsidian and rhyolite lava flows of the San Juan district, Colorado, USA, both as the lining of cavities and in the fine-grained groundmass (e.g., Klein & Hurlbut 1977).

5.3. Origin of Silica in Protoplanetary Disks

Less than 5% of the Si in the diffuse ISM is in crystalline silicates (Li & Draine 2001), and Kemper et al. (2004) did not detect silica in the interstellar medium 8–12.7 μm absorption profile observed towards the Galactic Center. The profile has no 9 or 12.3–12.6 μm features to indicate the presence of amorphous or crystalline silica, and they very nicely fit the optical depth profile without silica. This optical depth profile was shown by Sargent et al. (2006) to be very similar to the 8–12.7 μm emissivities derived for CoKu Tau/4, DM Tau, and GM Aur, whose IRS spectra were shown by D’Alessio et al. (2005) and Calvet et al. (2005) to be consistent with protoplanetary disks with many-AU-sized almost empty inner holes (CoKu Tau/4 and DM Tau) or optically thin inner regions and an almost empty gap (GM Aur). The 10 μm silicate emission from these systems, therefore, originates from their outer disk regions, where the dust is cooler and where dust processing is less likely to occur. Their 10 μm silicate emission profiles were very well fit with amorphous olivine and amorphous pyroxene and negligible silica, similar to the ISM absorption. We infer the average starting mixture of grains in protoplanetary disks is like those of CoKu Tau/4, DM Tau, and GM Aur and has negligible amounts of silica. Therefore, the silica we see must have formed in the disk.

In the creation of smoke from a laser-ablated enstatite grain, Fabian et al. (2000) found that the smallest smoke grains, 10–50 nm in size, were close to SiO₂ in composition. The smallest particles in a smoke created from a forsterite grain were not pure SiO₂, but had Mg/Si ratios between 0.5 and 1. The largest grains in both smokes were close to the composition of their parent enstatite and forsterite grains. This could be an equilibrium process related to one in which enstatite melts incongruently (Bowen & Anderson 1914) to forsterite plus a high temperature polymorph of silica, either cristobalite or tridymite:

\[ 2\text{MgSiO}_3 \rightleftharpoons \text{Mg}_2\text{SiO}_4 + \text{SiO}_2 \]

(Enstatite) + (Cristobalite or Tridymite),

except that in this case amorphous pyroxene, \([\text{Mg}_{1-x}, \text{Fe}_x]_3\text{SiO}_5\), replaces MgSiO₃, and amorphous olivine, \([\text{Mg}_{1-x}, \text{Fe}_x]_2\text{SiO}_4\), replaces Mg₂SiO₄ (Bowen & Schairer 1935). Here, x is probably around 0.5—we will call this reaction “incongruent melting of amorphous pyroxene.” We invoke amorphous pyroxene as the starting material in this reaction because amorphous pyroxene along with amorphous olivine are hypothesized to be the main ingredients of the interstellar medium (ISM; see Kemper et al. 2004) and the starting mixture of dust for protoplanetary disks (see discussion of unprocessed dust of transitional disks CoKu Tau/4, DM Tau, and GM Aur by Sargent et al. 2006). We also allow that incongruent melting of enstatite could be happening, where the parent enstatite could itself be the daughter product of prior processing of “grandparent” amorphous silicate.

Also, shocked enstatite grains in meteorites are known to have formed cristobalite inclusions (Benzzerara et al. 2002). Shock heating within the protoplanetary disk (Harker & Desch 2002), heating during disk-processing, flares, and lightning (Plipp et al. 1998), heating by intense light during the grains’ removal from the vicinity of the central star(s) by the T Tauri phase solar wind (de Pater & Lissauer 2001) or by an X-wind (Shu et al. 1996) all may cause melting of enstatite. Experiments by Rietmeijer et al. (1986), Hallenbeck et al. (1998), Hallenbeck & Nuth (1998), Rietmeijer et al. (1999), and Rietmeijer et al. (2002) also found silica grains (amorphous SiO₂ and tridymite) condensed from vapors of silicate compositions. This suggests silica can form in protoplanetary disks as a smoke or condensate from parent silicate dust grains, preferentially of low Mg/Si ratio like pyroxene. However, silica is not expected to condense from vapor of solar composition in chemical equilibrium (Gail 2004) in inner disk regions, suggesting the silica grains in the protoplanetary disks in our study did not arise from chemical equilibrium condensation in the solar nebula close to the Sun.

5.4. Silica as a Tracer of Dust Processing

After condensation of grains of the various silicates and silica, the experiments by Rietmeijer et al. (1986), Hallenbeck et al. (1998), Hallenbeck & Nuth (1998), Rietmeijer et al. (1999), Fabian et al. (2000), and Rietmeijer et al. (2002) subjected the condensates to high temperatures to anneal them. The two experiments of annealing at the highest temperatures for the longest durations, the heating of silica from iron-silicate smokers at 1300 K for 120 hr by Hallenbeck et al. (1998) and of pure silica grains at 1220 K for 5 hr by Fabian et al. (2000), gave very similar opacity profiles. Fabian et al. (2000) determined that their annealed silica was mostly cristobalite with some tridymite. The similarity of the opacity spectra indicates a similar composition for the Hallenbeck et al. (1998) annealed silica. The silica grains remaining after annealing at lower temperature for lesser durations in the other annealing experiments of Fabian et al. (2000) remained tridymite and amorphous SiO₂. Tridymite forms at lower temperatures and does not have an appreciable 16 μm feature. Therefore the 16 μm features of 1RXS J161410.6–230542 and ROXs 42C, which were better fit with annealed silica (which has a 16 μm feature) than with other versions of silica, suggests the presence of at least some cristobalite in our silica exemplars. Thus, we expect the silica grains in these protoplanetary disks were subjected to temperatures for durations of time like the most extreme annealing experiments of Hallenbeck et al. (1998) and Fabian et al. (2000).

One possibility is that tridymite or cristobalite forms close to the central star, where the stellar irradiation heats grains to the high temperatures required for formation of tridymite or cristobalite, and then are transported to cooler disk regions (Bockelée-Morvan et al. 2002). Our dust models have annealed silica (cristobalite/tridymite admixture) at temperatures ranging from 153 K to 1401 K, corresponding to distances from the star in the disk of a few AU down to the dust sublimation radius (sub-AU range). To explain the existence of tridymite or
cristobalite in the cool outer disk regions, however, the tridymite or cristobalite, once formed, would have to be cooled sufficiently quickly so that their crystalline structure does not revert to that of either of the lower-temperature polymorphs of silica, β- and α-quartz, and then transported to the outer disk. It seems unlikely that any of the proposed transport mechanisms could move these silica grains to a lower-temperature region of the disk rapidly enough to quench the cristobalite or tridymite structure in place.

If the tridymite or cristobalite do not form in the inner disk, this would suggest transient heating events, like shocks (Harker & Desch 2002) or lightning (Plipp et al. 1998), occurring in situ in the outer disk. The shock model of Desch & Connolly (2002) provides cooling similar to that required to explain chondrules—rates of 10–1000 K hr\(^{-1}\)—which may be quite sufficient to allow silica grains to retain their high-temperature polymorph crystalline structure. Furthermore, some of the Desch & Connolly (2002) shock models reached temperatures high enough that chondrule precursors would be evaporated; this may allow formation of silica via condensation. The shock model of production of crystalline silicate grains has been invoked to explain the presence of crystalline silicates in outflows from asymptotic giant branch (AGB) stars by Edgar et al. (2008). The lightning theory by Plipp et al. (1998) also attempts to explain chondrule formation and, therefore, meets similar requirements for cooling rates. The heating of pyroxene in some interplanetary dust particles (IDPs) to temperatures above 1258 K and their subsequent rapid cooling of ~1000 K per hour (Alexander et al. 2007) also suggest heating and rapid cooling in protoplanetary disks, as the chondrules do.

We note again the chondrules bearing silica by up to 40% by weight indicates that silica could have formed in the disks. The principal results of this paper are the following:

1. The spectra rule out the highest temperature polymorph of silica, stishovite, and we also fairly confidently rule out the intermediate pressure polymorph of silica, coesite.

2. Amorphous versions of silica like amorphous SiO\(_2\) (Henning & Mutschke 1997) and silica-rich glass, obsidian (Koike et al. 1989) are quite firmly ruled out to explain the silica features in our five silica exemplars due to their inability to fit the 12.6 \(\mu\)m features seen in these spectra. The amorphous SiO\(_2\) features peak at too short a wavelength for this feature, are too wide, and are too weak with respect to the 9 \(\mu\)m feature.

3. The opacity of annealed silica obtained by heating amorphous SiO\(_2\) for 5 hr at 1220 K (Fabian et al. 2000) provides the best fit to the spectra of all five of the strong silica exemplars in this study. Fabian et al. (2000) report this annealed silica to be mostly cristobalite but partly tridymite, so these silica polymorphs are the best candidates to account for the silica features seen in the spectra of our five silica exemplars. This opacity is very similar to that reported by Hallenbeck et al. (1998) for silica annealed at 1300 K for 120 hr.

4. The opacity of α-quartz at intermediate temperatures (505 K, 620 K) provided a good fit to the 12.6 \(\mu\)m feature of ROXs 42C, but the fits to features at other wavelengths characteristic of silica (9, 16, 20 \(\mu\)m) were not as good as the fits obtained from using annealed silica, indicating α-quartz is not the dominant silica in these disks. β-quartz, however, may be fairly firmly ruled out, as its feature at 12.7 \(\mu\)m cannot fit the 12.6 \(\mu\)m feature we observe in our dust excess spectra, regardless of the assumed shape distribution of the β-quartz grains.

5. Neither amorphous nor crystalline silica are seen in interstellar medium dust, so the silica present in protoplanetary disks of the five silica exemplars analyzed in this study must have formed in the disks.

6. The silica grains giving rise to emission from our five silica exemplars may have formed in similar conditions to those that gave rise to Ada, a grain composed mostly of tridymite surrounded by fayalite, found in the coma of Comet Wild 2 and the silica polymorphs found in primitive meteorites.

7. Tridymite and cristobalite are common products of annealing of amorphous dust smokes and condensates in laboratory experiments (e.g., Rietmeijer et al. 1986, 1999, 2002; Fabian et al. 2000), designed to mimic protoplanetary disk conditions thought to produce crystalline silicates by annealing amorphous silicate smokes. Cristobalite, the major component of the annealed silica we use, is present in silica grains annealed at temperatures above 1220 K for durations.
more than 5 hr and has a 16 \mu m feature in its opacity profile that is needed to fit the spectra of our silica exemplars. Silica, as a tridymite or cristobalite polymorph, might form in protoplanetary disks in manners similar to these laboratory experiments. Protoplanetary disk temperature and heating duration conditions giving rise to the silica we see in our sample of protoplanetary disks must be more like the most extreme conditions from the annealing experiments giving rise to cristobalite.

8. The silica grains may form via the incongruent melting reaction of amorphous pyroxene or the incongruent melting reaction of enstatite.

9. The tridymite or cristobalite grains, once formed, must be cooled quickly enough to retain their crystalline structure and not revert to the lower temperature polymorphs of silica, \( \beta \)- and \( \alpha \)-quartz. Mechanisms like spiral shocks and lightning devised to explain the rapidity with which chondrules cool may also satisfy the rapidity with which \( \beta \)-silica, as a tridymite or cristobalite polymorph, might form in the formation of the annealed silica.

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**APPENDIX**

For ZZ Tau, average flux densities for each DCE of each order of each nod were computed, averaging over 5.18–6.76 \mu m wavelength for SL second order, 7.57–8.37 \mu m wavelength for SL bonus order, 7.57–11.81 \mu m wavelength for SL first order, 14.15–17.97 \mu m wavelength for LL second order, 19.82–21.19 \mu m wavelength for LL bonus order, and 22.02–29.99 \mu m wavelength for LL first order. Each DCE of ZZ Tau was then scaled so that its average flux was the same average flux as the other DCEs of the same order. For SL second order, all DCEs were scaled so that their average flux was the same as that of the second nod from the Campaign 4 AOR; for SL bonus order, the average flux to match was that of the first nod of the Campaign 4 AOR; for LL second order, the second DCE of the second nod of the Campaign 29 AOR; for LL bonus order, the first DCE of the first nod of the Campaign 29 AOR; and for LL first order, the second DCE of the first nod of the Campaign 29 AOR. In almost all cases, the average fluxes of these named DCEs were the highest of their given order; in the case of LL second order, the average flux of the named DCE was only 1% discrepant from the highest average flux of all LL second order DCE spectra. The range of scalars applied to each DCE of ZZ Tau is provided also in Table 3.

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**Table 3**

| Object     | Correlation to SL2nod1 | Correlation to SL2nodd2 | Correlation to SL1nodd1 | Correlation to SL1nodd2 | Correlation to LL2nodd1 | Correlation to LL2nodd2 | Correlation to LL1nodd1 | Correlation to LL1nodd2 |
|------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|------------------------|
| RXS J161410.6–230542 | 1.03                     | 1.03                    | 1.03                    | 1.03                    | 1.00                    | 1.00                    | 1.00                    | 1.00                       |
| IRS 49     | 1.00                     | 1.04                    | 1.00                    | 1.04                    | 1.00                    | 1.04                    | 1.00                    | 1.04                       |
| ROXs 42C   | 1.00                     | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                       |
| TS1        | 1.00                     | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                    | 1.00                       |
| ZZ Tau     | 1.00–1.10                | 1.00–1.10               | 1.00–1.14               | 1.02–1.07               | 0.99–1.09               | 1.00–1.09               | 1.00–1.08               | 1.02–1.08                  |

**Notes.** Column 1: object name. Columns 2–9: multiplicative scalars applied to the spectrum of one order of one nod for a given object to match the flux density of the other nod as described in the text. SL2 means short-low order 2, SL1 is short-low order 1, LL2 is long-low order 2, LL1 is long-low order 1. For ZZ Tau, the range of scalars used for each order of each nod, derived as described in the text, is provided.
