Astrochronology of the Paleocene-Eocene Thermal Maximum on the Atlantic Coastal Plain

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The chronology of the Paleocene-Eocene Thermal Maximum (PETM, ~56 Ma) remains disputed, hampering complete understanding of the possible trigger mechanisms of this event. Here we present an astrochronology for the PETM carbon isotope excursion from Howards Tract, Maryland a paleoshelf environment, on the mid-Atlantic Coastal Plain. Statistical evaluation of variations in calcium content and magnetic susceptibility indicates astronomical forcing was involved and the PETM onset lasted about 6 kyr. The astrochronology and Earth system modeling suggest that the PETM onset occurred at an extreme in precession during a maximum in eccentricity, thus favoring high temperatures, indicating that astronomical forcing could have played a role in triggering the event. Ca content data on the paleo-shelf, along with other marine records, support the notion that a carbonate saturation overshoot followed global ocean acidification during the PETM.

The Paleocene-Eocene Thermal Maximum (PETM) was an interval of global warming that occurred ca. 56 million years ago (Ma) and was characterized by a 4–5 °C global mean surface temperature increase1. Estimates of the total amount of carbon released during the PETM range from ~3000 Pg to more than 13,000 Pg2–4, which span the current assessments of remaining fossil fuel reserves5. The PETM is considered to have the highest carbon release rates for the past 66 million years6, although estimates of rate are still limited by the low fidelity of records. Proposed triggers for the PETM include volcanism associated with the North Atlantic Igneous Province7,8, dissociation of methane hydrates (e.g., ref. 8), variations in Earth’s orbit that controlled massive carbon release from permafrost melting or oceanic methane hydrates9–11, and an extraterrestrial impact12,13. To further complicate the matter, estimates for the onset of carbon isotope excursion (CIE) at the PETM range from several years14–16 to thousands and even tens of thousands of years17,18,19,20 in duration.

The PETM CIE onset is defined by a negative shift of δ13C. Over the past few decades, considerable effort has been made to reconstruct the chronology of the CIE using astrochronology17,18,19,21,3He isotope measurements22,23, and modeling experiments6,24. At one extreme, the CIE onset was estimated to have spanned only 13 years based on assumed annual “bedding” couplets at a paleo-shelf section on the mid-Atlantic Coastal Plain15; an assumption contradicted by evidence for coring artefacts produced via biscuiting whereby the formation is fractured during coring and drilling mud is injected in between layers. The 13-year duration is also contradicted by evidence from foraminifer accumulation rates25,26, and carbon cycle/climate modeling6,27. At the other extreme are estimates ranging up to 20 kyr as derived from deep sea cores28,29. These estimates, however, are complicated by slow sedimentation rates coupled with carbonate dissolution and bioturbation12,19,30. Independent astrochronologic studies for the basinal, shallow marine, and terrestrial sites with high sedimentation rates are few18,21,31,32 and can be complicated due to the prevailing

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autogenic sedimentation processes in stratigraphy33. Astronomical age estimates from coastal/shelf records that have high sedimentation rates are still lacking, hindering the evaluation of the timing and the trigger of the PETM. In the coastal/shelf environment, non-orbital, 103–105 year-scale sedimentary noise resulting from storms, tides, bioturbation, variable sedimentation rates, short-term erosion, and diagenesis34, as well as mobile deltaic and continental shelf muds33,35 can be strong, hampering a straightforward interpretation of the astronomical signal in cyclostratigraphy. Moreover, although astronomical cycles have long been recognized in the PETM interval36,37, statistical evaluation of the null hypothesis (H0, no astronomical forcing) is rare, and links between astronomical forcing and proxy oscillations are unclear.

The Aquia Formation and Marlboro Clay from the Howards Tract cores (38.44827°N, 76.14159°W), two vertically offset holes in the coastal/shelf environment that have high temporal resolution paleoclimate proxies for the late Paleocene and early Eocene, e.g., calcium content and magnetic susceptibility (MS). Various studies demonstrate that Ca content and MS are two of the best recorders of astronomical cycles (refs. 39, 40 and references within). Ca content has long been used as a proxy of carbonate productivity in response to astronomically forced climate change41, MS, a measurement of the concentration of magnetic minerals, is a proxy of detrital fluxes from terrestrial sources in the marine environment42.

In this work, time series analysis of the proxy data (i.e., Ca content and MS) enables the recognition of astronomically forced sedimentary cycles at HT, leading to a high-resolution astrochronology for the PETM. The astrochronology is supported by statistical methods of sedimentation rate evaluation and an Earth system model of intermediate complexity. Earth system modeling of the effects of transient astronomical forcing using cGENIE provides a rare chance to elucidate the links between orbital forcing and paleoclimate proxies, e.g., Ca content, as well as the trigger to the PETM.

Results and discussion
Paleoclimate proxy records
The studied interval includes from base to top the glauconite-rich quartz sands of the Aquia Formation, the sandy clay to clay of the Marlboro Clay, and the clayey sand of the Nanjemoy Formation. The contact between the Aquia Formation and the Marlboro Clay is gradational with decreasing coarse fraction and CaCO3 content, and a gradual color change from dark greenish gray to brownish gray. In comparison, the highly burrowed interval between the Marlboro Clay and the Nanjemoy Formation indicates a disconformable contact. The high-resolution bulk carbonate δ13C record shows considerable variability at HT. Bulk carbonate δ18O records indicate the PETM CIE onset spans a 60-cm-thick interval (i.e., 200.47 to 199.89 m, pink bars in Figs. 2–3), which is defined by the initial sharp decline in the 8°C series and the changepoint analysis (see Methods and Supplementary Information). The magnitudes of the bulk carbonate δ13C and δ18O shifts at HT (Fig. 3a, b) are far larger than those from most PETM sequences, an artefact of early diagenetic carbonate siderite, common in Marlboro Clay sediments38. In contrast, a lower resolution benthic isotope record shows δ13C and δ18O shifts with magnitudes consistent with other sections along the Atlantic margin (Fig. 2b, c).

We measured Ca content using a Geotek X-ray fluorescence (XRF) scanner and magnetic susceptibility (MS) at 5 mm resolution. The XRF-generated Ca values generally match those measured in the lab (Fig. 2d), confirming the reliability of the Ca content from XRF scanning. The XRF-generated Ca content in the Aquia Formation is low (median 3.3%) with lower amplitude oscillations. The Marlboro Clay interval has a very low Ca content (median 1.9%) with occasionally more elevated values. The Ca content increases abruptly in the basal Nanjemoy Formation (median 6.3%) and remains high throughout the section (median 9.9%). MS values are relatively low in the Aquia Formation and basal Marlboro Clay and then reach their highest values in the middle Marlboro Clay, before gradually decreasing in the Nanjemoy Formation. The exceptionally low values at 182 m coincide with a minor unconformity; values rise again in the upper Nanjemoy sediments (Fig. 2f).

Cyclostratigraphic results
Time series analysis of the proxy data below the unconformity at the base of the Nanjemoy Formation shows astronomical cycle-paced variations of the detrended log10(Ca) and MS series (Fig. 4 and Supplementary Figs. 1–6). The Lomb-Scargle spectra of log10(Ca) and MS show dominant wavelengths of ~12 m and 2.2–3.4 m, respectively. There are also two higher-frequency cycles at 1.0–1.2 m and 0.67–0.73 m wavelength (Fig. 4a, f). The statistical tuning of correlation coefficient (COCO) method38 shows that optimal mean sedimentation rates are 8–15 cm/kyr, and the significance level of the null hypothesis of no orbital forcing is less than 0.05 (Fig. 4). Moreover, the average spectral misfit (ASM) method39, which objectively evaluates potential sedimentation rates, indicates the most likely mean sediment accumulation rate is 10–16 cm/kyr (Fig. 4f, g), at which the significance level of the null hypothesis of no orbital forcing is as low as 0.0014 (Ca).

Abbreviations
BR: Bass River ; CD: Cam-Dor ; HT: Howards Tract ; MCBR: Mattawoman Creek-Billingsley Road ; MV: Millville ; SDB: South Dover Bridge ; WL: Wilson Lake.
Fig. 2 | Time series analyses. a Chronostratigraphy, lithostratigraphy, calcareous nannofossil zones, and lithology for the Howards Tract (HT) cores. The schematic core log shows the grain size and sediment color of the cores. b δ¹³C of the HT cores. c δ¹⁸O of the HT cores. d Calcium content generated from X-ray fluorescence (XRF) scans of the HT cores correlates well with carbonate content from both Howards Tract 1 (HT1) and Howards Tract 2 (HT2). Ca data from 182–185 m is missing due to a coring gap. e Filtered 2.8 m cycles indicative of precession cycles (red, Gaussian filter with a frequency of 0.35 ± 0.15 m⁻¹) of the detrended log₁₀(Ca) content (gray). f Magnetic susceptibility (MS) of the HT cores. g Filtered 2.8 m cycles indicative of precession cycles (red, Gaussian filter with a frequency of 0.35 ± 0.15 m⁻¹) of the detrended magnetic susceptibility. h, i Tuned log₁₀(Ca) (h) and MS (i) and filtered precession (P) cycles (red, Gaussian filter with a frequency of 0.05 ± 0.018 kyr⁻¹). Pink bar: Paleocene-Eocene Thermal Maximum (PETM) carbon isotope excursion (CIE) onset. Shaded region: PETM.
Cibicidoides

The diagenetic setting incorporates low precipitation of abundant siderite. The siderite formed in this early stage is low δ13C, which would be the case during the PETM CIE onset, lying approximately 6 kyr (Fig. 3). The evolutionary fast Fourier transform (FFT), wavelet transform and Spectral Moments of both Ca and MS reveals similar first-order trends in sedimentation (SI): the dominant ~2 m precession cycle at ~205–200 m increases upward gradually to an ~3 m cycle at ~185 m (Supplementary Figs. 7 and 8). This suggests the accumulation rate increased ~1.5 times from the pre-PETM to the recovery phase (Supplementary Fig. 9). Tuning of ~3 m cycles of the Ca and MS series to the ~20 kyr precession cycles enables for the recognition of 5.5 precession cycles from the PETM CIE onset through the hiatus at the top of the Marlboro Clay (Supplementary Tab. 1), suggesting the Howards Tract Bridge section may be an artifact of a more condensed basal Marlboro Clay interval at HT; however, we cannot rule out the possibility that the presence of early diagenetic carbonate has made the CIE onset appear more abrupt than in benthic foraminifera20.

Environmental changes associated with the PETM, the global carbon isotope excursion of ~4–5 ‰ is amplified to ~13 ‰ at HT. Moreover, because the onset of the CIE at HT coincides closely with the base of the Marlboro Clay, the possibility exists that the timing of the isotope excursion reflects both the depositional and early diagenetic environment of the mobile mud belt as well as the input of isotopically light carbon that fueled the PETM warming recorded at sites globally. To attempt to deconvolve these two factors, we compare the timing of the CIE onset at HT with other sections in Maryland and New Jersey where siderite is also present, yet the bulk carbonate δ13C still captures the global carbon isotope signal as represented in high resolution planktonic and benthic foraminifera records from the same sections. We compare the onset with the timing of the base of the Marlboro Clay as well as three nannoplankton datums to determine whether the initial stage of the CIE at HT was more abrupt than in the other sections; such an abrupt onset could be a result of a relationship with the deposition of the mobile mud belt or early diagenetic conditions within it (Fig. 5).

Identification of datums used in this analysis can be subjective, including the base of the Marlboro Clay36, change points in the carbon isotope excursion, and biostratigraphic datums, and we attempt to be as consistent as possible with the definition of all three types of datums (see Supplementary Information for more discussion). The analysis shows that the base of the onset of the CIE lies in an identical position to the base of the Marlboro Clay in HT as in the other two sections (Fig. 5). Moreover, the onset does not appear to be more abrupt at HT compared to Wilson Lake, New Jersey, as determined by its position relative to the three nannofossil datums, but it does appear to be two times more abrupt at HT than at South Dover Bridge, Maryland. The more abrupt onsets at HT relative to the relatively close by South Dover Bridge section may be an artifact of a more condensed basal Marlboro Clay interval at HT; however, we cannot rule out the possibility that the presence of early diagenetic carbonate has made the CIE onset appear more abrupt than the original global signal. Indeed this looks to be the case at the Mattawoman Creek-Billingsley Road section in Maryland, where siderite is abundant, and the CIE onset in bulk carbonate δ13C is more abrupt than in benthic foraminifera20.

Sedimentation rate at Howards Tract

Our estimation for the PETM CIE onset duration assumes a constant sedimentation rate within the first precession cycle (P1, in Figs. 2–3), a cycle that includes the transition between the Aquia Formation and the Marlboro Clay. The onset of the deposition of the mobile mud belt in the Marlboro Clay may have involved a significant increase in sedimentation rate that would weaken the constant sedimentation rate

Mattawoman Creek-Billingsley Road (MCBR) cores in stratigraphic depth20. δ13C values of the Wilson Lake core in stratigraphic depth. The data source for Wilson Lake: Cibicidoides spp. and bulk carbonate (206 denotes ref. 89); Anomalinoides acutus and Cibicidoides spp. (S12 denotes ref. 53); organic matter of dinocysts90.

and 0.0012 (MS). In other words, confidence levels of astronomically forced variations in Ca and MS are higher than 98.6%. Therefore, the ~12 m, 2.2–3.4 m, 1.0 m, and 0.63–0.75 m cycles represent ~100 kyr short eccentricity, ~20 kyr precession, and sub-Milankovitch cycles (~10 kyr and ~7 kyr), respectively. The δ13C values of the HT cores.

δ13C values of the Howards Tract (HT) cores. b Tuned δ18O values of the HT cores. c Tuned calcium content generated from X-ray fluorescence (XRF) scans (gray line) and filtered precession (P) cycles (red line, detailed in Fig. 2b) of the HT cores. d δ13C values of the HT cores.

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assumption, thus assuming a constant sedimentation rate would overestimate the duration of the CIE onset, which lies almost entirely in the Marlboro Clay. Nonetheless, spectral moments of both log$_{10}$(Ca) and MS series indicate the mean sedimentation rate within each 4 m sliding window increases only slightly between the Aquia Formation and the Marlboro Clay (Supplementary Figs. 7–9); we consider the ca. 6 kyr duration of the CIE onset determined at HT given the uncertainty related to the definition of the CIE onset and the impact of diagenesis as discussed above.

High detrital accumulation rates are thought to enhance organic carbon sequestration during the PETM, which along with the silicate weathering feedback, drove the recovery of the Earth system from
PETM CO₂ emissions as indicated in part by the termination of the CIE\(^\text{28,29}\). Indeed, many studies demonstrate that the hydrological cycle intensified during the PETM\(^\text{49,50}\), which is reflected in large part by the dramatic increase in sedimentation rates during the PETM in continental margin settings including Belluno Basin, Italy\(^\text{51}\), Tunisia\(^\text{52}\), Paleotethys\(^\text{53}\), Atlantic Coastal Plain sections\(^\text{47,53}\), west coast of North America\(^\text{47}\), Lomonosov Ridge in Arctic Ocean\(^\text{54}\), and North Sea Basin\(^\text{55}\).

However, both the compilation of hydrologic records and Earth system modeling suggests the climate response had significant regional variability – some areas are characterized by increased precipitation-evaporation, whilst others are associated with a decrease\(^\text{47,53}\). Our high-resolution astrochronology indicates that the mean sedimentation rate during the PETM at HT was \(-10\) cm/kyr, which is consistent with the estimates from other sites on the Atlantic Coastal Plain\(^\text{\textbar}\). The evolutionary FFT, wavelet, and Spectral Moments analyses indicate a generally smooth increase in the mean sedimentation rate in the HT cores (See SI), rather than a dramatic 2.8- to 220-fold increase (i.e., from 0.1-1.0 cm/kyr for the pre-CIE to 2.8-22 cm/kyr during the CIE) in regional sedimentation rates\(^\text{\textbar}\). The previous sedimentation rates were estimated via the division of the stratigraphic thickness by the corresponding duration\(^\text{\textbar}\), which was determined by stratigraphic correlation using biozones and the CIE shape, both of which are low resolution and can be affected by sporadic deposition and erosion in the Atlantic Coastal Plain. In comparison, the astrochronology as applied here and elsewhere is high resolution and relies on the net sediment accumulation rate.

Astrochronology of the Paleocene-Eocene thermal maximum

Our analysis of the Howards Tract cores is generally consistent with and more resolved than published astrochronologies and \(^{1}\text{He}\) chronological models for the PETM. Cyclostratigraphy of deep-sea cores at ODP Sites 1051 (western North Atlantic) and 690 (Weddell Sea, Southern Ocean) suggested the PETM spanned 11 precession cycles yielding a duration of 210-220 kyr, and the PETM CIE onset of initial decrease in \(8^{\text{13}}\text{C}\) took over 20 kyr, while 52 kyr elapsed between the onset and the nadir of the \(8^{\text{13}}\text{C}\) excursion\(^\text{28,29}\). About two-thirds of the duration, assuming a constant sedimentation rate within each precession cycle\(^\text{28}\), is consistent with those from Earth system modeling experiments that suggest the PETM CIE onset spanned at least 4 kyr\(^\text{24}\) or less than 5 kyr\(^\text{24}\). The initial release of carbon at a rate of \(0.6\) Pg C/yr during the PETM, assuming an \(-20\) kyr duration of the onset, could be doubled when a 5 kyr duration is considered\(^\text{\textbar}\), but anthropogenic carbon release rates at \(-10\) Pg C/yr\(^\text{56}\), which is one order of magnitude higher than that of the PETM. This study provides direct constraints on the carbon cycle and paleoclimate changes in the shelf environment, supporting the emerging consensus view of a few millennia for the onset interval.

Precession forced \(^{15}\text{C}\) excursions

This study can improve our understanding of the linkage between orbital forcing and changes in paleoclimate proxies such as \(^{15}\text{C}\) content. Based on previous work using the cGENIE Earth-system model with transient orbital forcing (cf. ref. 61), we simulate the influence of transient astronomical forcing on paleoclimate to compare to our Howards Tract record. Modeling of the \(8^{\text{13}}\text{C}\) excursion using cGENIE has already been undertaken\(^\text{57}\), which forced the model to conform to observed isotope excursions, providing insightful constraints of the rate of carbon release and isotope fingerprint of the carbon source. Alternatively, we focus on astronomically forced climate change without simulating the effect of carbon release. In cGENIE model, variations of insolation are controlled by astronomical forcing\(^\text{\textbar}\) (Fig. 6a-c). The upper envelope of mean daily insolation at HT was paced by 20 kyr precession cycles and modulated by eccentricity cycles (Fig. 6b, d). The same is true for sea surface temperature (SST, Fig. 6e) and \(^{15}\text{C}\) ion concentration (Fig. 6f). The upper envelope of mean daily \(^{15}\text{C}\) export fluxes of biological production (Fig. 6g) or annual \(^{15}\text{C}\) fluxes (Fig. 6h) at Howards Tract is dominated by precession cycles. The annual \(^{15}\text{C}\) export fluxes compare well with the filtered precession cycles of the \(^{15}\text{C}\) content in the Aquia Formation and the Marlboro Clay (Fig. 6i). For example, the modeled \(^{15}\text{C}\) fluxes at \(-110\) kyr (i.e., 70 kyr after the PETM CIE onset) capture the minimal \(^{15}\text{C}\) content at 194 m (Fig. 6i).
The above results can be explained by the fact that insolation forcing controls surface temperature and thus determines the rate of carbonate weathering and silicate weathering in cGENIE and cGENIE66,67. Moreover, summer insolation is dominated by precession and eccentricity forcing (Supplementary Fig. 10), while the annual insolation intensity is controlled by obliquity forcing at HT (Supplementary Fig. 11). Therefore, CaCO3 flux in HT cores could be paced by astronomically forced maximum insolation, in other words, the intensity of temperature in summer season or summer half-year. A positive swing in Ca time series at HT occurs when the northern summer occurs during perihelion, and vice versa. The amplitude of Ca concentration variations is higher at eccentricity maxima because of the modulation of eccentricity, during which Earth can be either particularly close to or away from the Sun in northern hemisphere summer on the 100–405 kyr time-scale (Fig. 6f, g). Another scenario is that siliciclastic fluxes that control CaCO3 content are modulated by astronomically forced changes in weathering, precipitation, runoff, and sediment discharge. In this scenario, siliciclastic dilution of CaCO3 is driven partly by precipitation on the mid-Atlantic Coastal Plain which would impact sediment discharge and nutrient fluxes. In addition, MS is considered to be a record of terrigenous material supplied to the depositional basin by runoff from the continent, which suggests MS should be out of phase with the Ca concentration time series. Figure 2 shows more Ca generally corresponds to relatively less terrigenous material (thus drier summers), and vice versa. Therefore, the climate processes influencing the character of local sedimentation are not mutually exclusive and might enhance the lithologic cycle pattern.

**Orbital trigger for the Paleocene-Eocene thermal maximum**

Forcing the cGENIE model to conform to the boron isotope pH proxy and carbon isotope data indicates a mixed source of carbon release, i.e., volcanic outgassing plus methane hydrates and/or permafrost, during the PETM onset. Both the astrochronology in the HT cores and model results demonstrate the PETM CIE onset occurred at an extreme in precession, favoring high temperature in summers (Fig. 6b, d, e) and a maximum in the eccentricity cycles (Fig. 6a, b), indicating an astronomical trigger. The possibility that volcanism pulsed at the maximum in the eccentricity cycles cannot be precluded; nonetheless, increased ocean temperatures could have triggered the release of methane hydrates and/or carbon ejection from permafrost (cf. refs. 9, 11). This mechanism implies that hyperthermal warming events could have occurred at other times with similar orbital configurations. Time series analysis of the deep-sea records demonstrates both the PETM and Eocene Thermal Maximum 2 (ETM-2) occurred during the eccentricity maxima that post-date the very long, Myr-scale eccentricity minima that occurred at other times with similar orbital configurations. Forcing the cGENIE model to conform to the boron isotope pH proxy and carbon isotope data indicates a mixed source of carbon release, i.e., volcanic outgassing plus methane hydrates and/or permafrost, during the PETM onset. Both the astrochronology in the HT cores and model results demonstrate the PETM CIE onset occurred at an extreme in precession, favoring high temperature in summers (Fig. 6b, d, e) and a maximum in the eccentricity cycles (Fig. 6a, b), indicating an astronomical trigger. The possibility that volcanism pulsed at the maximum in the eccentricity cycles cannot be precluded; nonetheless, increased ocean temperatures could have triggered the release of methane hydrates and/or carbon ejection from permafrost (cf. refs. 9, 11). This mechanism implies that hyperthermal warming events could have occurred at other times with similar orbital configurations. Time series analysis of the deep-sea records demonstrates both the PETM and Eocene Thermal Maximum 2 (ETM-2) occurred during the eccentricity maxima that post-date the very long, Myr-scale eccentricity minima that occurred at other times with similar orbital configurations.
Eocene global warmth was punctuated by recurrent, rapid hyperthermal events, which are mainly paced by cyclicities in Earth’s orbit eccentricity. Moreover, coupled climate model simulations indicate that eccentricity-forced changes in ocean circulation and seawater temperature (through variations in seasonality) caused the destabilization of methane hydrates, which could explain the increasing frequency and decreasing amplitude of hyperthermal warming events in the early Eocene. Therefore, the conjunction of 100 kyr, 405 kyr, and very long, Myr-scale eccentricity cycles may have facilitated the build-up of a major mobile reservoir of reduced carbon such as methane hydrates, marine dissolved organic carbon, and/or organic-rich peat before its release during the hyperthermal events.

Unlike those deep ocean records that are complicated by a major dissolution interval and bioturbation in the PETM interval, high sedimentation rates and almost non-existent bioturbation in the HT cores allow for an unprecedented resolution for astrochronology of the PETM CIE onset that is supported by Earth system modeling, pointing to a possible orbital trigger for the PETM.

**CaCO₃ preservation**

The HT carbonate record exhibits signs of a carbonate saturation “overshoot” in the later recovery stage of the PETM. Theory, supported by recent observations, indicates that a large and rapid release of carbon into the Earth’s surface system induces a two-phase response in ocean carbonate saturation. The first phase of carbon ejection will cause short-term ocean acidification lowering seawater ocean saturation (Ω), while the second phase could be characterized by carbonate oversaturation caused by elevated rates of silicate weathering and elevated carbonate deposition. This phenomenon, known as carbonate saturation overshoot, could have led to an over-deepening of the calcite compensation depth (CCD) relative to its pre-event depth. Globally distributed PETM sites ranging from deep ocean to shelf support ocean acidification and the shoaling of the CCD possibly to even shallow shelf depths. For example, multiple cores on the Atlantic paleo-shelf in Maryland and New Jersey record an interval devoid of carbonate during the onset of the PETM and the disappearance of nannofossils and planktic foraminifera. The dissolution of calcareous material was considered to be syndepositional possibly due to the significant shoaling of the CCD, although there are other possible explanations involving local influences, including dilution coupled with euxinia. In the second phase, the recovery and overshoot in carbonate saturation is best captured in hemi-pelagic and pelagic records. The Forada section in particular, with a distinct clay layer indicates resumption of carbonate deposition roughly 20 kyrs after the acidification. In the Atlantic, the CCD gradually deepened over several tens of thousand of years before a state of oversaturation was reached, resulting in carbonate deposition at depths previously below the CCD. Collectively, these observations support carbon cycle models that include a silicate weathering feedback. At HT, the Nanjemoy Formation preserves a high CaCO₃ content, i.e., up to 18% in the PETM late recovery phase versus 3.3% in the pre-PETM and 1.9% during the PETM body interval (Fig. 2), demonstrating the occurrence of the overshoot in carbonate saturation. This overshoot could explain the enhanced nannofossil preservation right above the dissolution interval in cores on the mid-Atlantic paleo-shelf. The trend in the Atlantic paleoshelf demonstrates the carbon saturation overshoot impacted even the shallow ocean.

To review, a statistically significant astronomical signal in the Marlboro Clay has been detected, in this case from the Howards Tract cores in Maryland. The astrochronology suggests that the Marlboro Clay at this site preserves a 110-kyr record of the PETM and that the onset of the event lasted -6 ky. A combination of astrochronology and Earth system modeling suggests that the PETM CIE onset occurred at an extreme in precession favoring high temperature and at the maxima of 405 kyr and 100 kyr eccentricity cycles, indicating a possible orbital trigger. Astronomically paced silicilastic and nutrient fluxes, along with precession-forced temperature-dependent changes in global weathering rates of carbonate and silicate rocks, could have contributed to oscillations of Ca content as exemplified at Howards Tract. Carbonate content data on the Atlantic paleo-shelf, along with other deep-sea records, suggest that carbonate saturation overshoot occurred not just in the deep sea but also in coastal regions during the PETM recovery.

**Methods**

**Lithology**

Two cores were drilled at Howards Tract (HT1 and HT2, 5 m apart with offset coring intervals) to minimize loss due to coring gaps. Spliced data from HT1 and HT2 resulted in relatively complete coverage for the Aquia Formation, Marlboro Clay, and Nanjemoy Formation. The Aquia Formation-Marlboro Clay contact is located at 200.43 m (657.6 ft), and the Marlboro Clay-Nanjemoy Formation contact is at 187.5 m (615.2 ft). Observation of the HT cores suggests the contact between the Aquia Formation and the underlying unit has been reported at other sites, including the Medford Auger Project (MAP) cores, Millville, CamDor, and Wilson Lake on the mid-Atlantic Coastal Plain. Unlike cores at Millville and Wilson Lake, Howards Tract cores show no evidence of couplets. Core photos are shown in Supplementary Figs. 12–15.

**Proxy data measurement**

The elements in the Howards Tract cores (both HT1 and HT2) were measured using the XRF scanner of Geotek’s Multi-Sensor Core Logger at Pennsylvania State University. The measurement time for calcium is 30 seconds and the spatial resolution is 5 mm. To test the reliability of the XRF-scan calcium, we measured carbonate content on a UIC Inc. coulometers Coulometer at the University of California Santa Cruz with a precision of ± 0.05%. δ¹³C and δ¹⁸O of bulk carbonate and benthic foraminifera (3-5 specimens from the 180-212 cm depth) were measured using the XRF scanner of Geotek at Pennsylvania State University. The measurement time for calcium is 30 seconds and the spatial resolution is 5 mm. To test the reliability of the XRF-scan calcium, we measured carbonate content on a UIC Inc. coulometers Coulometer at the University of California Santa Cruz with a precision of ± 0.05%. δ¹³C and δ¹⁸O of bulk carbonate and benthic foraminifera (3-5 specimens from the 190-212 cm depth) is based on the replicate analyses of standards (i.e., Carrara Marble). All data are reported relative to Vienna Pee Dee Belemnite. The sampling rate for bulk samples is ca. 0.1 m for the top of the Aquia Formation and increases to 0.03-0.05 m for the lower part of the Marlboro Clay. Nonetheless, the base of the Marlboro Clay is characterized by a prominent interval in which CaCO₃ content decreases to close to zero during the CIE onset, i.e., the low carbonate interval (LCI) on the New Jersey and Maryland paleoshelf. This LCI can be further supported by a gap in the foraminifera and very poor coccolith shield preservation, due to a lack of calcareous material during the CIE onset at many mid-Atlantic paleoshelf sites, such as South Dover Bridge, MCBR, and HT. The LCI and missing cores prevent a uniformly sampling strategy and are responsible for the sampling rate over 0.3 m in the bottom of the Marlboro Clay.

**Changepoint analysis and the definition of the carbon isotope excursion onset**

The changepoint analysis of δ¹³C data is able to provide the objective detection of changepoints at HT. Detailed search methods and test statistics of the changepoint analysis can be found in ref. 80. We use the cpt.meanvar function of the changepoint R package because the carbon isotope data show changes in both the mean and variance. Four
change points are detected at depths of 186.61, 199.12, 200.47, and 202.92 m (Supplementary Text 1 and Supplementary Fig. 16). Among these, 200.47 m is used as the base of the CIE onset and coincides with the Aquia Formation-Marlboro Clay contact.

The change point analysis doesn’t provide direct constraints for the top of the CIE onset. We choose 199.39 m (5.8 ky after the CIE onset) as the top of the onset because this position records the largest negative 5°C C13 change, which is constrained by over one data point. The position at 199.34 m (11 ky after the CIE onset) has the most negative 5°C C13 value, however, this position is only constrained by one datapoint, which is thus not used in the main paper. Even if it is used as the top of the CIE onset, the comparison of the onset with the positions for the Marlboro Clay and the three nanoplanckton datums shows the onset is no more abrupt at HT compared to Wilson Lake (Supplementary Fig. 17). The results do not contradict our conclusion on the sedimentation rate variation during the CIE onset.

**Time series methods**

The identification of astronomical cycles takes advantage of Acycle v2.4.1 software and follows typical procedures39. The Ca and magnetic susceptibility (MS) series carry a long-term trend that can be high amplitude and non-periodic, leading to power leakage from low-frequency components into the frequency band of interest40. Therefore, both series were detrended after subtracting a 20-m “loess” (local regression using weighted linear least squares and a 2nd degree log10(Ca) series. Because regularly spaced time series is required for MS series were interpolated using a 2nd degree polynomial model trend for the MS series and a linear trend for the log10(Ca) series. Because regularly spaced time series is required for many powerful techniques in this study, the detrended log10(Ca) and MS series were interpolated using a “linear” method. To reveal the dominant wavelength of the proxy series and search for potential astronomical cycles, the Lomb-Scargle spectrum is calculated and shown with confidence levels test against robust AR(1) red noise models fitting to 30% median-smoothed power spectrum using the “Spectral Analysis” function in Acycle. Gauss and Taner bandpass filters were applied to isolate potential astronomical parameters41.

Astronomical tuning is conducted using 20 ky precession cycles and the “Age Scale” function in Acycle. In order to identify the sediment accumulation rate and test the null hypothesis that no astronomical forcing drove oscillations of the proxy series derived from the HT cores, we calculated the average spectral misfit (ASM)42 using Astrochron package2 and the correlation coefficient (COCO) spectra of the detrended log10(Ca) and detrended MS series and six astronomical target periodicities (i.e., 125, 95, 39.8, 23.3, 22.0, and 18.7 ky), which is based on the power spectrum of astronomical target series (La2004 solution from 55 Ma to 57 Ma). Details of the parameters for the ASM calculation can be found in the Supplementary Information. In the COCO calculation, classic red noise models of both spectra were removed to suppress the very high amplitude for the low frequencies. Test sedimentation rates range from 0.13 cm/kyr to 30 cm/kyr with a step of 0.1 cm/kyr. The number of Monte Carlo simulations is 2000.

**TimeOpt** is a statistical method for the estimation of optimal sedimentation rate for a given paleoclimate proxy series43. At each test sedimentation rate, the proxy series was converted from depth domain to time domain. Then TimeOpt used the Taner filter and the Hilbert transform to isolate the potential precession cycles and the corresponding amplitude envelope. The envelope was linearly regressed on a synthetic time series that was generated using eccentricity frequencies retrieved from astronomical models (e.g., La2004 solution) for a given age. The correlation coefficient of the regression at the test sedimentation rate was recorded as $r_{\text{envelope}}^2$. Meanwhile, the data were linearly regressed to another synthetic dataset using frequencies of eccentricity and precession. And the regression was reported with the correlation coefficient $r_{\text{power}}$. The product $r_{\text{opt}}^2 = r_{\text{envelope}}^2 \times r_{\text{power}}^2$ was used to evaluate the most likely sedimentation rate. The test sedimentation rate with the highest $r_{\text{opt}}^2$ might be the optimal sedimentation rate. The null hypothesis of no orbital forcing and the confidence of the optimal sedimentation rate can be evaluated using Monte Carlo simulations. The lag-1 correlation coefficient of the proxy series was calculated and used for the generation of many (e.g., 2000) red noise series. Then $r_{\text{opt}}^2$ at each test sedimentation rate of these noise series was calculated and recorded. Consequently, the percentile of the $r_{\text{opt}}^2$ using real proxy series indicates the chance of this $r_{\text{opt}}^2$ that can occur randomly. The optimal sedimentation rate can be considered to be significant when the null hypothesis of $r_{\text{opt}}^2$ is lower than 0.05. These calculations can be done either using astrochron package in R44 or using Acycle software58.

Here, TimeOpt analysis for the detrended log10(Ca) series of the Aquia Formation and the Marlboro Clay using Acycle shows the most likely sedimentation rate is ca. 16.0 cm/kyr (Supplementary Fig. 3) at which the null hypothesis significance level of no orbital forcing is 0.001, that is, the confidence level of orbital forcing is 99.9% (Supplementary Fig. 4). The wide range of sedimentation rates at 10–17 cm/kyr demonstrates the sedimentation was variable at HT (Supplementary Fig. 4). This sedimentation rate is slightly higher than the COCO- and ASM-generated sedimentation of 8–16 cm/kyr (Fig. 4). In comparison, TimeOpt analysis for the MS series shows the most likely sedimentation rate is 8.1 cm/kyr (range of 6–10 cm/kyr; Supplementary Fig. 5) at which the null hypothesis significance level of no orbital forcing is 0.001 (Supplementary Fig. 5). This range is slightly lower than the COCO- and ASM-generated sedimentation rate of 8–16 cm/kyr. Nonetheless, all results point to the conclusion that variable 2–3 m wavelengths in our proxy records represent 20 ky precession cycles. This is within expectation because multiple proxies and approaches can usually lead to different, but comparable within error, results. Taken together, ASM, COCO and TimeOpt analyses indicate the PETM interval of the Marlboro Clay was paced by precession cycles, which were modulated by eccentricity cycles.

**Variable sedimentation rate**

In order to reveal the secular trend of dominant frequencies, the evolutionary fast Fourier transform (FFT)45 were calculated with Acycle “Evolutionary Spectral Analysis” function45 using a sliding window of 7 m and a step of 0.01 m. Because the window size is smaller than the reported eccentricity cycles (8–10 m), precession-related cycles (2–4 m) are the strongest signal in the evolutionary FFT result. The evolutionary FFT of both Ca and MS reveals a similar trend: the dominant ~2 m precession cycles at ~205–220 m increases upward smoothly to ~3 m cycles at ~185 m (Supplementary Fig. 7), indicating an increasing upward sedimentation rate from ca. 10 cm/kyr to ~15 cm/kyr. Wavelet analysis shows cyclicities of the data series are generally stable (Supplementary Fig. 8). The ca. 11 m wavelengths that are interpreted as 100 ky short eccentricity cycles are mostly unchanged. The 2.3 m cycles (~20 ky precession cycles) show a similar increase upward trend.

The Spectral Moments methods evaluate the first order change in sedimentation rate via investigating the analyzed series using a periodogram with two spectral moments: mean frequency ($\mu$) and bandwidth (B) using a sliding window approach46. We detect shifts and changes in sedimentation rate using the Spectral Moments method in Acycle 2.4.16. The edge of the data series is filled using the zero-padding method. The bandwidth and mean frequency are calculated using a sliding window of 4 m with a running step of 0.1 m. The mean sedimentation rate, required by the Spectral Moments algorithm, is set to 10 cm/kyr based on the ASM and COCO analysis. We estimate the trend in sedimentation rate by taking the LOESS trend of the bandwidth.

Spectral moments of both series are shown in Supplementary Fig. 9. The first-order changes in sedimentation rate using both Ca and MS series show an increasing upward trend. The minor discrepancies...
between the two estimated sedimentation rate maps indicate complex climate responses of different proxies\(^{29}\), and/or the ability of the spectral moments method in the estimation of fine-scale changes in sedimentation rate.

**Earth system modeling**

We used the cGENIE Earth system model to simulate the variability of the Ca content in the HT cores. The model consists of a 3D ocean circulation model\(^{66}\) coupled to a 2D energy-moisture balance model (EMBM) of the atmosphere and a dynamic-thermodynamic sea-ice model\(^{67}\). It also includes a 3D module of marine biogeochemical cycling of major nutrients, trace elements, and isotopes in the ocean\(^{68}\), a 2D atmospheric chemistry module, and a module for interactions between sediments and ocean\(^{69}\) and terrestrial weathering\(^{70}\). The model used a Paleogene bathymetry and continental configuration and was initialized with a value of alkalinity (1975 umol eq kg\(^{-1}\)) to produce a mean global CaCO\(_3\) content of 47%\(^{85}\). Two-stage spin-up weathering\(^{66}\). An initial spin-up phase, the ocean-atmosphere carbon cycle is set to CaCO\(_3\) at all times and no bioturbation is allowed in the sediments\(^{84}\).

The simulation results should not be largely affected by the choice of the onset age because both options occurred with similar cycles\(^{88}\). The model was set to reproduce the astronomically forced Ca variations at HT core, the release rate unprecedented during the past 66 million years. The proxy series of calcium content, magnetic susceptibility, and sedimentation rates and astronomical forcing in Phanerozoic paleoclimate proxy series with evolutionary correlation coefficients and hypothesis testing. Earth Planet. Sci. Lett. 501, 165–179 (2018).

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Author contributions

M.L. and T.J.B. designed the study, T.J.B., J.M.S, and M.M.R. collected and prepared cores, M.L. conducted XRF scanning, W.D.R. collected and identified the benthic foraminifera and carried out the isotope analyses, M.L., T.J.B., L.R.K., J.M.S., and J.C.Z. interpreted data. M.L. wrote the paper, and all authors contributed to editing the paper.

Competing interests

The authors declare no competing interests.
