Emiliania huxleyi coccolith calcite mass modulation by morphological changes and ecology in the Mediterranean Sea

Barbara D’Amario¹*, Patrizia Ziveri¹,²*, Michaël Grelaud¹, Angela Oviedo¹

1 Institute of Environmental Science and Technology (ICTA), Universitat Autònoma de Barcelona (UAB), Bellaterra, Spain, 2 Institució Catalana de Recerca i Estudis Avançats (ICREA), Barcelona, Spain

* 3damario@gmail.com (BD); Patrizia.Ziveri@uab.cat (PZ)

Abstract

To understand the response of marine calcifying organisms under high CO₂ scenarios, it is critical to study their calcification patterns in the natural environment. This paper focuses on a major calcifying phytoplankton group, the coccolithophores, through the analysis of water samples collected along a W-E Mediterranean transect during two research cruises, in April 2011 (Meteor cruise M84/3) and May 2013 (MedSeA cruise 2013). The Mediterranean Sea is a marginal sea characterized by large biogeochemical gradients. Currently, it is undergoing both warming and ocean acidification, processes which are rapidly modifying species distribution and calcification. The species Emiliania huxleyi largely dominates the total coccolithophore production in present day oceans and marine basins, including the Mediterranean Sea. A series of morphometric measurements were performed on the coccoliths of this species to estimate their mass, length and calculate a calcification index (proxy for the size-normalized calcification degree). The most abundant morphotype of E. huxleyi in the Mediterranean Sea is Type A. Coccoliths of this morphotype were additionally analyzed based on scanning electron microscopy images: four calcification varieties were quantified, according to the relationship between slit length-tube width, and the state of the central area (open or closed). The average Emiliania huxleyi coccolith mass along the Mediterranean oceanographic transect depended strongly on both the average coccolith length and calcification index. The variability in average coccolith length and calcification index across samples reflected oscillations in the relative abundance of the calcification varieties. We also demonstrated that the distribution of the calcification varieties followed the main environmental gradients (carbonate chemistry, salinity, temperature, nutrient concentrations). Hence, shifts in the distribution of the calcification varieties and of the average E. huxleyi coccolith mass are to be expected in the Mediterranean Sea under climate change. These physiological and ecological responses will modulate the net coccolithophore calcification and, ultimately, the regional carbonate export to the seafloor.
1 Introduction

The accumulation of human-induced atmospheric CO$_2$ is altering the global climate and driving rapid changes in the carbonate chemistry of surface seawaters. For example, it has been estimated that since the industrial revolution, the global ocean surface pH has decreased by 0.1 units and a supplementary decrease of 0.06 to 0.32 pH units is expected by the end of the 21st century [1,2]. This process, termed ocean acidification, is thought to impact the calcification process of many marine organisms [3,4]. However, it is still unknown how large such potential reduction in calcification would be, or what could be the net effect on coccolithophores, which represent the most prominent calcifying phytoplankton group on Earth: their calcite builds up 20 to 40% of the total open ocean carbonate sedimentation, from equatorial to sub-polar regions [5–7].

Several studies have shown the sensitivity to ocean acidification of the family Noelaerhabdaceae, the most abundant in present-day coccolithophore communities [8–12]. On a global scale, coccolith mass within this family seems correlated to the seawater carbonate system (carbonate ion concentration, calcite saturation state, pH, pCO$_2$): carbonate chemistry could control the distribution of differently calcified taxa, primarily species and morphotypes of the genera Gephyrocapsa and Emiliania [12–14]. Most laboratory experimental results on selected strains of Emiliania huxleyi, as reviewed by [8], showed a tendency for decreasing PIC production and PIC: POC ratio in high pCO$_2$ conditions. Nonetheless, variable and sometimes contradictory responses have been found with respect to ocean acidification for $E$. huxleyi, e.g. [12,15,16]. This is probably due to the large number of genotypes included in the species considered [17–20]. Also, in the natural environment, the regional distribution of differently calcified specimens of $E$. huxleyi is likely affected by additional parameters [21], such as temperature [22], nutrient concentrations [23] and salinity [24].

Overall, the responses of $E$. huxleyi to changing environmental conditions have been examined through two main approaches: field studies and laboratory experiments. Field studies describe natural coccolithophore communities and their geographical distributions: they can highlight statistically significant correlations between the distribution of morphotypes, calcification varieties and the environmental variables, but cannot be used to establish cause-effect relationships. Conversely, laboratory experiments can prove the direct effects of different environmental pressures on $E$. huxleyi, but only for a few strains and environmental parameters at a time: they cannot replicate the complexity of the natural environment, and as such they cannot account for genotype sorting.

$Emiliania$ huxleyi was initially classified into four genotypically-controlled morphotypes: A, B, C and corona [25]. The combined approach of microscopic and molecular techniques has allowed the distinction of other morphotypes, highlighting their genetic diversity [17,18,26–31] and biogeographical/seasonal distributions [18,32–34]. Up to present, five morphotypes of $E$. huxleyi (A, B, B/C, C and R) are proven to remain consistent when reproducing [25,35,36].

The existence of different degrees of calcification within $E$. huxleyi Type A has been documented in the past [35], mainly from the North Atlantic Ocean [37–39], the Pacific Ocean [40] and the Mediterranean Sea [41–47]. The region of focus for this study is the Mediterranean Sea: a ‘small-scale ocean’ with steep W–E biogeochemical and physical gradients and a fast overturning circulation (80 to 100 years) [48].

$E$. huxleyi calcification (coccolith mass) in the Mediterranean Sea is controlled by the average coccolith size and calcification degree. The present work hypothesizes that such morphological parameters are linked to different proportions of Type A calcification varieties vs. Type B/C. Four Type A calcification varieties are described here and their spatial distribution is evaluated to identify plausible environmental controls.
2 Material and methods

2.1 Environmental conditions

The oceanographic data presented and discussed in this paper were retrieved along a W-E transect extending from the Atlantic Ocean (−6.64° E, 36.03° N) to the Levantine Basin of the Mediterranean Sea (31.00° E, 33.70° N; Fig 1), during the Meteor M84/3 (6 – 28 April 2011) and the MedSeA (1 – 31 May 2013) oceanographic cruises, following similar longitudinal transects (S1 Fig) [49,50]. Three main water masses were identified within the upper 200 m during the MedSeA and M84/3 cruises, as shown in the temperature-salinity diagram (S2 Fig). Detailed descriptions of the oceanographic settings and the methods of analysis used have been published previously for both the Meteor M84/3 cruise [49,51,52] and the MedSeA cruise [50]. Permission for navigation and research operations in exclusive economic zones of the Mediterranean Sea was granted from the Governments of Spain, Greece, France, Italy and Cyprus. Sampling did not involve endangered or protected species.

2.2 Phytoplankton samples

A total of 55 water samples from the two cruises were analyzed through morphometric measurements (mass, length, width) and/or detailed morphology recognition. They were collected along the transect (Fig 1) between 5 and 150 m depth (Table 1), following the eastward deepening of the photic zone. The depth of the photic zone was estimated from the available fluorescence values (see fluorescence profile in the supplementary material of [50]). Using Niskin bottles, 2.5–5 liters of water per sample were collected and filtered through cellulose acetate-nitrate filters (Millipore, Ø 47 mm, 0.45 μm). A hydraulic vacuum pump system (Eyela, A-1000S) was used at low pressure, to obtain an even distribution of particles on the filter. Each filter was then rinsed with distilled water, buffered with ammonia (63 ml NH₃ + 500 ml of distilled water), in order to remove salt residues, and oven-dried at 40°C for 8–12 hours.

2.2.1 *Emiliania huxleyi* coccolith mass and length. For each sample, a portion of filter was mounted on a glass slide, soaked in immersion oil and secured applying a coverslip and tape on the borders of the slide. The sample slides were then analyzed with a Leica DM6000B cross polarized light microscope fitted to a SPOT Insight Camera, at x 1000 magnification. The luminosity level of the microscope was set prior to analysis, following the guidelines of [21]. At least 40 pictures were taken per sample, to count a minimum of 300 intact and isolated coccoliths (100 in a few low abundance samples). The resulting images were analyzed with an automated system of coccolith recognition (SYRACO, [53,54]). The mass (pg) and length...
(μm) of 23255 individual isolated coccoliths were measured following the guidelines of [21,55]. The input of reworked calcareous nannofossils from the continent is a common process in the Mediterranean Sea, as shown by previous studies on Holocene sediments (i.e. [56–60]). However, scanning electron microscopy (SEM) observations on the same samples did not reveal any evident sign of diagenesis on the specimens. Because of this, we considered the analyzed coccoliths to be a representation of those originally interlocked in *Emiliania huxleyi* coccospheres.

### 2.2.2 *Emiliania huxleyi* morphotypes and Type A calcification varieties.

A portion of filter was radially cut from each sample, attached to a stub and coated with a gold/palladium (Au/Pd) alloy, to be observed at 30000 X magnification through SEM (Zeiss EVO MA 10). A longitudinal transect of 5 mm or more was scanned (∼4.1 ml of water), until a minimum of 80 – 100 coccospheres were counted: *Emiliania huxleyi* morphotypes were identified and their coccospheres counted separately. Individual *E. huxleyi* coccoliths from Type A coccospheres (a mean of 377 per sample, for a total of 13192) were identified by SEM observations and assigned to one of the following categories: low–calcified (A1), medium–calcified (A2) and high–calcified (A3a, A3b). The main distinguishing feature between A1, A2 and A3b was the ratio between slit length (SL) and tube width (TW) on the distal shield of the coccolith (Fig 2): in A1, SL > TW; in A2, SL ≈ TW; in A3b, SL < TW. The main feature characterizing A3a coccoliths was instead the nearly, or completely closed, central area; in this case, the ratio between SL and TW was very variable and not taken into account. Each coccolith was assigned to a single calcification variety and the relative abundance of each variety to the total *E. huxleyi* Type A coccoliths was calculated. Our observations suggest that the norm is for a Type A coccosphere to be entirely formed of coccoliths of one calcification variety. However, a mix of calcification varieties was clearly identified on two coccosphere specimens from the SE Mediterranean (S3 Fig).

### 2.2.3 Calcification index.

The coccolith mass measured through SYRACO depends on both coccolith length and its degree of calcification. We obtained an indicator of the calcification degree, the calcification index $C_i$, as an alternative to the “relative tube width” [61].

$$C_i = \frac{M_i}{M_n}$$ (1)

### Table 1. List of samples analyzed using SYRACO (a) and/or SEM (b).

| Station | Date      | Coordinates | Depth (m) |
|---------|-----------|-------------|-----------|
|         | dd/mm/yy  | E E N       | 5 10 25 40 50 75 80 100 110 125 150 |
| 1       | 02/05/13  | -6.64 36.03 | a,b a a,b |
| 2       | 03/05/13  | -5.56 35.95 | a,b a,b a a,b |
| 4       | 07/05/13  | 1.45 37.49  | a,b a a,b |
| 5       | 08/05/13  | 5.55 38.52  | a a a,b a |
| 7       | 11/05/13  | 12.68 37.12 | a a a,b a |
| 9       | 12/05/13  | 18.29 35.11 | a a,b a |
| 11      | 15/05/13  | 28.00 33.50 | a a,b a a,b a,b a |
| 294     | 10/04/11  | 31.00 33.70 | a,b a,b a,b |
| 302     | 13/04/11  | 20.35 35.07 | a,b a,b a,b |
| 324     | 21/04/11  | 5.60 38.65  | a,b a,b a,b |
| 329     | 23/04/11  | 2.00 37.90  | a,b a,b a,b |
| 337     | 25/04/11  | -5.36 36.00 | a,b a,b a,b |

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In Eq (1), $M_s$ is the average coccolith mass measured with SYRACO for the sample under consideration, while $M_n$ is the “normalized mass”. $M_n$ was calculated based on [62], as in Eq (2):

$$M_n = V \times \rho_c$$

in which $\rho_c$ is the density of calcite ($= 2.7 \text{ pg } \mu m^{-3}$) and $V$ corresponds to the coccolith volume calculated as in Eq (3):

$$V = k_s \times L_c^3$$

where $L_c$ corresponds to the average coccolith ‘corrected length’ obtained from SYRACO measurements (see next paragraph) and $k_s$ is an estimated shape dependent constant ($k_s = 0.02$) for *E. huxleyi* type A normally calcified coccoliths [62].

We expected the coccolith distal shield length to be systematically underestimated by SYRACO. In fact, when the coccolith is observed in cross-polarized light, the calcitic coccolith tube appears extremely bright in comparison to the peripheral and thinner area of the shield [63]. To quantify this underestimation, we selected six samples from those already analyzed with SYRACO, so as to include a wide range of coccolith lengths. Fifty micrographs of flat-lying *E. huxleyi* Type A coccoliths were captured per sample, using the SEM at 30000 X magnification. Those images were analyzed with the open source software Fiji [64], a distribution of ImageJ [65]. The Coccobiom2 macro (http://ina.tmsoc.org/nannos/coccobiom/Usernotes.html; [61]), developed specifically for coccoliths, was chosen to facilitate the measurements. The image analysis done through this method allowed accurate measurements of the central area + tube length, and of the total distal shield length for each coccolith (S4 Fig). As expected, the coccolith lengths obtained from SYRACO were systematically greater than the SEM-derived sums of coccolith central area + tube length. Still, both of these quantities were smaller than the total
coccolith lengths measured from SEM micrographs. The equation of the linear regression in the last plot of S4 Fig \( y = 0.585 x + 0.4537 \) was used to correct the SYRACO-derived coccolith lengths of all remaining samples. As mentioned in the previous paragraph, the corrected length values were used in Eq (3).

### 2.3 Statistics

A Canonical Correspondence Analysis (CCA) and a series of Spearman’s rank correlations were performed on the dataset. The CCA is a multivariate analysis [66], used here (Fig 3) to relate the observed *Emiliania huxleyi* morphotypes and Type A calcification varieties distributions to the environmental gradients.

The Spearman's analyses (Tables 2–4) tested the correlations between the morphological parameters \( M_s, L_c, C_i \), the environmental variables, and the relative abundances of the *Emiliania huxleyi* morphotypes and Type A calcification varieties (A1, A2, A3a, A3b); coefficients were regarded significant for \( p \leq 0.05 \). The software PAST 3.14 [67] was used for these statistical analyses.

#### Table 2. Correlations among morphological parameters: Coccolith mass \( (M_s) \), calcification index \( (C_i) \) and corrected length \( (L_c) \).

|          | \( M_s \) | \( C_i \) | \( L_c \) |
|----------|-----------|-----------|-----------|
| \( C_i \) | 0.86**    |           |           |
| \( L_c \) | 0.91**    | 0.61**    |

Significant Spearman coefficients are in bold \( (N = 54; ^* p \leq 0.01) \).
tests. Regression analyses and correspondent coefficients ($R^2$) were obtained plotting data in Grapher™ 12 (Golden Software, LLC).

### 3 Results

Data from the Meteor M84/3 (April 2011) and MedSeA cruises (May 2013) have been combined to obtain the average values for all morphological variables, the relative abundances of the *Emiliania huxleyi* morphotypes, and those of the Type A calcification varieties.

The CCA (Fig 3) revealed that 81.87% of the variance in the relative abundance of *Emiliania huxleyi* Type A and of its four calcification varieties was explained by axis 1 (52.23%) and 2 (29.64%). Axis scores showed that, among the environmental parameters, the three major contributors to axis 1 were $pCO_2$ (-0.48), pH (0.47) and $[CO_3^{2-}]$ (0.38); the three major contributors to axis 2 were instead salinity (-0.33), total alkalinity (-0.33) and temperature (-0.30).

The average coccolith mass ($M_s$) increased eastward between the Atlantic-Gibraltar Strait and the SE Mediterranean: from 3.77 to 4.76 pg. The average values of $L_c$ and $C_i$ increased in the same direction, respectively from 3.25 to 3.45 μm and from 2.01 to 2.14 (Fig 4). The lowest coccolith mass values were found in the upper photic zone of the Atlantic-Gibraltar Strait (minimum in station 1, at 40 m, 2.24 pg; Fig 5), while the highest values were found in the middle-lower photic zone of the SW and SE Mediterranean (maximum in station 11, at 100 m, 5.85 pg).

The largely dominant *Emiliania huxleyi* morphotype in our samples was Type A (Fig 4). Type B/C, the only other morphotype of *Emiliania huxleyi* observed, was poorly represented: on average, it corresponded to 5% of the total *Emiliania huxleyi* coccospheres, except for St. 337 (the westernmost station sampled during the Meteor M84/3 cruise), where it reached 73%.

### Table 3. Correlations between morphological parameters, Type A coccospheres % (with respect to the total *Emiliania huxleyi* coccospheres; N = 52) and calcification varieties % (with respect to the total Type A coccoliths; N = 35).

| Parameter | Type A% | A1% | A2% | A3a% | A3b% |
|-----------|---------|-----|-----|------|------|
| $M_s$     | 0.22    | -0.33* | 0.29 | -0.03 | 0.75** |
| $C_i$     | 0.03    | -0.11 | 0.05 | 0.30  | 0.45** |
| $L_c$     | 0.29*   | -0.40* | 0.39* | 0.21  | 0.81** |

Significant Spearman coefficients are in bold (*$p$ ≤ 0.01, *$p$ ≤ 0.05).

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### Table 4. Environmental correlations of morphological parameters (N = 54), Type A coccospheres % (N = 52) and calcification varieties % (N = 35).

| Parameter | $M_s$ | $C_i$ | $L_c$ | Type A% | A1% | A2% | A3a% | A3b% |
|-----------|-------|-------|-------|---------|-----|-----|------|------|
| Temperature | -0.28* | -0.47** | -0.11 | 0.32* | 0.04 | 0.15 | -0.73** | 0.03 |
| Salinity   | 0.61** | 0.37** | 0.65** | 0.54*  | -0.42** | 0.19 | -0.20 | 0.67** |
| Total alkalinity | 0.57** | 0.37** | 0.60** | 0.54*  | -0.34*| 0.10 | -0.17 | 0.63** |
| pH         | 0.41** | 0.29*  | 0.39*  | 0.53*  | 0.36* | -0.39* | -0.36* | 0.16 |
| $pCO_2$    | -0.42** | -0.25 | -0.45** | -0.51** | -0.20 | 0.22 | 0.46** | -0.29 |
| $CO_3^{2-}$ | 0.37** | 0.13  | 0.47** | 0.69** | 0.10 | -0.17 | -0.55** | 0.37* |
| $NO_2+NO_3$ | -0.01 | 0.17  | -0.08 | -0.43** | 0.16 | -0.16 | 0.51** | -0.20 |
| $PO_4^{3-}$ | -0.12 | 0.10  | -0.20 | -0.46** | 0.25 | -0.15 | 0.44** | -0.33 |
| PAR%       | -0.54** | -0.44** | -0.47** | 0.12  | 0.41  | -0.21 | -0.13 | -0.57* |

$P$ values were available only for the MedSeA cruise (N = 35 in relation to morphological parameters; N = 33 in relation to Type A coccospheres %; N = 16 in relation to calcification varieties %). Significant Spearman coefficients are in bold (*$p$ ≤ 0.01, *$p$ ≤ 0.05).
Coccospheres were almost never retrieved in the SE Mediterranean (0–1% of the total Emiliania huxleyi).

A1 and A2 were the most abundant calcification varieties along the transect. The average percentage of A1 decreased from the Atlantic-Gibraltar Strait (47% of the total E. huxleyi Type A) towards the SE Mediterranean (30%), while A2 increased (from 47% to 54%). The high-calciﬁed forms included A3a and A3b: A3a was present almost exclusively in the Atlantic-Gibraltar Strait province, at low relative abundance (2% of the E. huxleyi Type A coccoliths), and in the SW Mediterranean, where it reached up to 15%. Contrasting with the patchy distribution of A3a, A3b was found along the whole transect and increased in relative abundance eastwards, 4% to 16% (Fig 4).

Emiliania huxleyi coccolith mass ($M_s$) was positively correlated with $C_i$ and $L_c$ (S5 Fig, Table 2), and with the relative abundance of A3b (Fig 6, Table 3); while it was negatively correlated with the relative abundance of A1 (Fig 6, Table 3). Coccolith length ($L_c$) had similar coefﬁcients with respect to A1 and A3b, but was also signiﬁcantly correlated with A2; $C_i$ was correlated only with A3b. Other signiﬁcant correlations were found between morphological and environmental parameters (Table 4; S6 Fig): $M_s$, $C_i$ and $L_c$ were all strongly correlated with the salinity, total alkalinity, pH and Photosynthetically Active Radiation (PAR%). Weaker but statistically signiﬁcant relationships with $pCO_2$, [$CO_3^{2-}$], temperature and dissolved oxygen were also found. The relative abundance of Type A cccospheres showed the strongest correlations with the carbonate system, salinity and nutrients, followed by temperature (Table 4; S7 Fig). All of the Type A calcification varieties displayed signiﬁcant correlations with at least one carbonate system parameter. Furthermore, the calcification variety A1 was signiﬁcantly correlated with salinity, A3a with temperature and nutrient concentrations, A3b with salinity and PAR% (Table 4).

4 Discussion

4.1 Morphological controls over E. huxleyi coccolith mass

Along the studied transect, the average E. huxleyi coccolith mass per sample oscillated between 2.24 and 5.85 pg. This range of values is similar to that registered in sediment trap material from the NW Mediterranean (2.81 – 5.61 pg; [68]) and is overall higher than that measured in surface sediments from the South Atlantic-southwestern Indian Ocean (1.73–4.85 pg; [21]). The largest variations in average E. huxleyi coccolith mass should mainly reﬂect the geographical distribution of its morphotypes [12]. Still, in oceanic regions where one morphotype is
largely predominant, as in the Mediterranean Sea [46,47,49], the calcification varieties of that particular morphotype can be expected to exercise a major control.

We verified that Type A was by far the dominant morphotype along the studied Mediterranean transect (Fig 4), as in [49,60,69]. The concomitance of high Type B/C percentages and of low average coccolith mass values in the Atlantic-Gibraltar province is likely not a coincidence: Type B/C is typically less calcified than Type A. Still, the Spearman analysis did not reveal any significant correlation between coccolith mass and the percentage of Type A coccospheres with respect to the total *Emiliania huxleyi* population (Table 3). Our results suggest instead that most of coccolith mass variability in the Mediterranean Sea is actually controlled by the relative abundance of the Type A calcification varieties. The significance of the correlations between A1, A3b and the average coccolith mass, indicates that our morphological subdivision of Type A into calcification varieties is of relevance to addressing coccolith calcite content. Moreover, we can infer that the high-calcified A3b coccoliths were usually larger than those of other
calcification varieties, especially if compared to the low-calcified A1 coccoliths. Similar positive trends between size and calcification degree were noticed in oceanic communities of the North Atlantic [61]. Overall, the comparison between morphological variables suggests that *E. huxleyi* coccolith mass is controlled by the distribution of the Type A calcification varieties, characterized by different size ranges and calcification degrees. The Eastern Mediterranean was richer in large, high-calcified specimens (A3b), explaining the high coccolith mass values obtained from this province (Figs 4 and 5). Coccoliths of A3a were present almost exclusively in the Western Mediterranean, where they occasionally became more abundant than A3b (in stations 2, 4, 7, 324 and 329). The two high-calcified varieties A3a and A3b should have both contributed positively to the average calcification degree. However, increases in the relative abundance of A3a did not always correspond to increases in the average coccolith length. For example, in station 329, A3a increased from 22 to 58%, while the average *E. huxleyi* coccolith length decreased from 3.30 to 3.27 μm: A3a might actually contain less calcite than A3b, due to its smaller average size, resulting in a lower mass. We conclude that the contribution of A3a to the average coccolith calcification degree, length, and ultimately mass, was negligible at the time of sampling, both at basin (whole W-E transect) and province (SW Mediterranean) scale. The calcification varieties A1 and A3b, instead, controlled most of the *E. huxleyi* coccolith mass variability along the transect.

### 4.2 Ecology of *E. huxleyi* morphotypes and Type A calcification varieties

*Emiliania huxleyi* morphotypes are genetically-controlled and have distinct environmental preferences [14,49,70,71]. For example, Type B/C commonly lives in cold waters, rich in nutrients and with low calcite saturation state. These preferences explain the limited distribution of Type B/C in the Mediterranean Sea compared to Type A, especially in the Eastern province (Fig 4); the high percentages of Type B/C in the Atlantic–Gibraltar province have also been interpreted as a biological signal of the Atlantic water inflow into the Mediterranean Sea [49].

The distribution of the Type A calcification varieties seems to be controlled by an array of environmental parameters: carbonate chemistry, salinity, temperature and nutrient concentrations (Fig 3, Table 4). However, It is not clear if this distribution reflects their ecological affinities or their phenotypic plasticity. With the currently available data, it is difficult to determine if the Type A calcification varieties are characterized by distinct genotypes or not.
Previous field observations have demonstrated the existence of a connection between multiple environmental parameters and the calcification degree of *E. huxleyi* populations [12,14,38,46,47,58]. Besides, distinct sets of environmental parameters appear to control the average coccolith mass, depending on the oceanic region [21].

Our results suggest that carbonate chemistry could have an important role in regulating the distribution of the Type A calcification varieties (Table 4, S7 Fig). However, there is not a straightforward relationship between the calcification degree of Type A coccoliths and the carbonate parameters: the distributions of the two high-calcified varieties (A3a and A3b) were related differently to total alkalinity, pH, pCO$_2$ and especially in respect to [CO$_3$$^{2-}$]. This is in accord with previous field studies [12,38], showing that low-calcified and high-calcified specimens can both occur in waters having low pH, low calcite saturation state and low [CO$_3$$^{2-}$]. Hence, the significant correlation we observed between the calcification degree (C$_i$) and water pH (Table 4; S6 Fig) should be evaluated with caution.

The distributions of the A1 and A3b varieties followed, in addition to carbonate chemistry, the salinity gradients; in the case of A3b, we also observed an inverse relationship with PAR% (Fig 3, Table 4). Overall, A1 was predominantly encountered in the upper photic zone of the Atlantic–Gibraltar Strait and SW Mediterranean, characterized by low salinity waters; while A3b was mainly found in the middle-lower photic zone of the SE Mediterranean, characterized by high salinity waters. A similar vertical distribution of the calcification varieties was observed, at smaller scale, in the NE Aegean Sea [47]: low-calcified *E. huxleyi* Type A coccoliths were particularly abundant in the upper 20 m of the water column, within the low salinity-low density layer of inflowing Black Sea water. The calcification varieties’ distribution observed in our samples could thus depend on the salinity preferences of the strains and/or could be regulated by physical factors (density and vertical stratification of water masses).

Temperature and nutrient concentration seem to exercise a major control on the distribution of A3a coccoliths (Fig 3, Table 4). Our correlations can help explain a relative increase in high-calcified coccoliths during winter-spring in the Aegean Sea, when temperatures were lower and nutrient concentrations higher [46,47].

As anticipated, we provide two interpretations for the existence of the *E. huxleyi* Type A calcification varieties: each calcification variety could represent a population characterized by (a) one genotype, or (b) multiple genotypes.

The first interpretation contemplates the potential production of several calcification varieties during the lifespan of a single *E. huxleyi* cell, triggered by changing environmental conditions. In support of this hypothesis, we report the occurrence of seasonal unimodal patterns of calcification in the Aegean Sea [46] and of occasional ‘mixed’ coccospheres, composed of multiple calcification varieties, retrieved along our Mediterranean transect (S3 Fig) and in the Atlantic Ocean [38]. Furthermore, laboratory experiments have demonstrated that the calcification process in *E. huxleyi* can be strongly influenced by carbonate chemistry [8], salinity [24,72–74], temperature [22,75] and nutrient concentrations [23]: the environmental conditions regulate the cellular energetic consumption dedicated to the calcification process. A complication is that, in laboratory and field studies, two different concepts of ‘calcification’ are used: calcification rate and calcification degree. During cultures, PIC production and PIC:POC ratio are obtained, while in filtered marine water samples and sediments it is necessary to refer to coccolith calcite mass, degree of calcification and/or size. Recent work [22] however, indicates that high coccolith mass for *E. huxleyi* might be a good indicator of high PIC production.

The second interpretation envisages the idea that each calcification variety represents a group of strains adapted to a specific ecological niche. In this case, the correlations observed in our samples (Table 4) would indicate the ecological preferences of each calcification variety.
Emiliania huxleyi has a high genetic diversity in today’s oceans, as proved by both microsatellite and laboratory data [12,18,76]. Although the relationship between E. huxleyi genome and its morphotypes is not straightforward [18], it has been observed that shifts in the relative abundance of morphotypes generally correspond to shifts in the relative abundance of distinct genetic codes [12]. Such genetic diversity translates into strain-specific responses to environmental change [23,77]: at least two clades of E. huxleyi Type A, characterized by different preferential habitats, are present in the Mediterranean Sea; temperature and phosphate are likely the most important environmental constrains on their distribution [17,18]. Another interesting point raised by a recent study is that inter-strain genetic variability can potentially induce larger phenotypic differences than phenotypic plasticity itself [78].

Although the occurrence of ‘mixed’ coccospheres would induce a quick conclusion (i.e. each Type A cell can produce all four calcification varieties), we cannot exclude the possibility that those rare specimens are artefacts: a xenosphere formed by a mix E. huxleyi coccolith morphologies, similarly to the specimens observed in our samples and in the Atlantic Ocean [38], was identified by [79] as a probable artefact.

The four calcification varieties described in this work should be ideally isolated and subjected to controlled environmental perturbations. That would be a direct way to assess their nature and their specific responses to environmental change. Among the four calcification varieties, A3a has a characteristic distribution in the SW Mediterranean and a peculiar affinity for low temperatures-high nutrients, making it a very interesting candidate to test in the future.

Under climate change, we can expect shifts within the E. huxleyi populations. The Mediterranean Sea could experience not only a shift in the proportion of E. huxleyi morphotypes (Type A vs. Type B/C), but also in the proportion of the Type A calcification varieties (A1, A2, A3a, A3b). The average E. huxleyi coccolith mass will depend mainly on the response of the Type A calcification varieties. Based on their present distributions, we foresee that A3a will likely become even rarer under warming and enhanced water column stratification. The fate of A1 and A3b could be instead mainly linked to carbonate chemistry and salinity, but the relative importance of these parameters cannot be discerned with certainty based only on available data. Overall, the proportion of A1 and A3b calcification varieties is expected to retain the greatest control over the average E. huxleyi coccolith mass in the Mediterranean Sea.

5 Conclusions

Emiliania huxleyi coccolith mass variability in the Mediterranean Sea is primarily modulated by the relative abundance of Type A calcification varieties, being Type A the largely dominant E. huxleyi morphotype in this oceanographic region. The large, high-calcified specimens (A3b) are more common in the SE basin, while the small, low-calcified specimens (A1) are more common in the Atlantic-Gibraltar Strait and in the SW basin. Our results indicate that seawater carbonate chemistry, water salinity, temperature and nutrient concentrations might all contribute in regulating the distribution of different E. huxleyi Type A calcification varieties. The nature of the calcification varieties observed is possibly linked to the cellular energetic consumption in relation to the calcification process in different environmental settings. Overall, they might represent a mix of genotypes. The available data suggest that E. huxleyi Type B/C and the Type A calcification variety A3a will become rarer in the Mediterranean Sea under warming and, especially, under enhanced water column stratification. The average E. huxleyi coccolith mass in the Mediterranean Sea is basically controlled by the proportion of the Type A calcification varieties A1 and A3b. The fate of A1 and A3b is probably closely linked to changes in seawater carbonate chemistry and needs to be assessed through laboratory perturbation experiments.
Supporting information

S1 Fig. Total alkalinity along the studied transect, during the Meteor M84/3 cruise and the MedSea cruise.
(TIF)

S2 Fig. Main water masses within the upper 200 m during the MedSeA and M84/3 cruises.
A-G = Atlantic Water from the eastern Atlantic Ocean and the Gibraltar Strait; SW Med. = Modified Atlantic Water from the South Western Mediterranean; SE Med. = Modified Atlantic Water from the South Eastern Mediterranean.
(TIF)

S3 Fig. *Emiliania huxleyi* Type A ccccospheres composed by a mix of coccolith calcification varieties. Blue = A1, green = A3b, orange = A3a. Specimens are from St. 11, collected at 75 m and 100 m depth.
(TIF)

S4 Fig. Coccolith morphometry and correspondent standard deviations for six samples: Comparisons between SYRACO results and those obtained from the analysis of SEM micrographs. The linear regression function in the last plot was used to calculate the average corrected length ($L_c$) for all samples.
(TIF)

S5 Fig. Linear regressions between coccolith mass ($M_s$), corrected length ($L_c$) and calcification index ($C_i$).
(TIF)

S6 Fig. Plots combining the morphological and carbonate system parameters. Linear regressions are shown only for combinations of parameters which have significant Spearman correlation coefficients ($p \leq 0.05$).
(TIF)

S7 Fig. Plots combining the percentage of low-calcified / high-calcified Type A varieties and the carbonate system parameters. Linear regressions are shown only for combinations of parameters which have significant Spearman correlation coefficients ($p \leq 0.05$).
(TIF)

S1 Table. Comparison between the *E. huxleyi* Type A calcification varieties defined in this work (A1, A3a, A3b) and those appeared in the literature concerning the North Atlantic Ocean and the Mediterranean Sea (Med). The variety A2 is not included being very common (showing no special patterns in our samples) and medium-calcified.
(DOCX)

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Author Contributions
Conceptualization: Barbara D’Amario, Patrizia Ziveri, Michaël Grelaud, Angela Oviedo.
Data curation: Michaël Grelaud.
Formal analysis: Barbara D’Amario.
Funding acquisition: Patrizia Ziveri.
Investigation: Barbara D’Amario, Angela Oviedo.
Methodology: Barbara D’Amario, Patrizia Ziveri, Michaël Grelaud, Angela Oviedo.
Project administration: Barbara D’Amario, Patrizia Ziveri.
Resources: Patrizia Ziveri.
Supervision: Patrizia Ziveri, Michaël Grelaud.
Visualization: Barbara D’Amario.
Writing – original draft: Barbara D’Amario.
Writing – review & editing: Barbara D’Amario, Patrizia Ziveri, Michaël Grelaud, Angela Oviedo.

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