A pelagic myodocopid ostracod from the Silurian of Arctic Russia

VINCENT PERRIER1*, OLGA K. BOGOLEPOVA2, ALEXANDER P. GUBANOV3, DAVID J. SIVETER1 & MARK WILLIAMS1
1Department of Geology, University of Leicester, University Road, Leicester LE1 7RH, UK
2CASP, Cambridge University, 18a Huntingdon Road, Cambridge CB3 0DH, UK
3Museum of Evolution, Uppsala University, Norbyvägen 16, 752 36, Uppsala, Sweden
*Corresponding author (email: vp110@leicester.ac.uk)

ABSTRACT – The Silurian myodocopid ostracod Richteria migrans is reported from Arctic Russia, from Kotel’ny Island (New Siberian Islands) and the Taimyr Peninsula in strata of Ludfordian (late Ludlow, Late Silurian) age. These occurrences extend the biogeographical range of R. migrans from tropical to mid latitudes of the Early Palaeozoic Rheic Ocean in the palaeo-Southern Hemisphere, into subtropical regions of the palaeo-Northern Hemisphere on, or adjacent to, the Siberia Palaeocontinent. The new records reinforce the idea that R. migrans had wide dispersal capacity and probably possessed a pelagic lifestyle. It also endorses the use of R. migrans as a biostratigraphical marker fossil for the Ludfordian Stage, Ludlow Series, Upper Silurian.

KEYWORDS: myodocopid ostracods, Late Silurian, Arctic Russia, biostratigraphy, palaeobiogeography

INTRODUCTION

Ostracods are a major group of small aquatic crustaceans and are by far the most prolific group of arthropods in the fossil record, ranging over some 500 million years from at least the Ordovician Period onward (Horne et al., 2002; Siveter, 2008; Williams et al., 2008). They are important components of freshwater, marine and brackish-water settings. Among ostracods, myodocopes were present and already diversified by the Silurian (Siveter et al., 2003, 2007, 2010, 2013). Although they are mostly nekto-benthic, Recent myodocope ostracods are also an important component of marine zooplankton, ranging from the ocean ‘Mixed Layer’ to abyssal depths (Angel, 1993). Based on several lines of evidence, Siveter et al. (1991) proposed that myodocopes were pioneer ostracod zooplankton colonists, making the ecological shift from the benthos during the Silurian.

The first myodocopid ostracod groups to colonize pelagic environments were probably the bolbozoids (e.g. Parabolbozoe bohemica (Barrande, 1872) and Bolbozoe anomala (Barrande, 1872)) and the cypridinids (e.g. Silirocypridina calva Perrier et al., 2011), which show a supposedly transcontinental distribution during the late Wenlock–early Ludlow interval (Perrier et al., 2011; Perrier & Siveter, 2013). Richteria migrans (Barrande, 1872), one of the earliest ‘entomozoid’ ostracods (Perrier et al., 2007; Perrier 2012) – a group colloquially known as ‘finger-print’ ostracods because of their characteristic pattern of ribbed ornament – was also probably pelagic. ‘Entomozoaceans’ have particularly wide geographical distribution and biostratigraphical value in the Devonian and the Carboniferous, occurring, for example, in Europe, North Africa, North America and China (Gooday, 1983; Groos-Uffenorde & Wang, 1989; Olempska, 2002a, b). Opinions differ regarding the lifestyle of the ‘entomozoaceans’. Some authors consider that ‘entomozoaceans’ were benthic, possibly living in oxygen-depleted environments (Warshauer & Duffield, 1983; Casier, 1987). However, it is generally thought that most members of the group occupied pelagic, perhaps mid-water niches (Gooday, 1983; Olempska, 1992; and references therein).

The biostratigraphical and palaeoecological significance of R. migrans were tested by Perrier et al. (2007), who showed that the species had wide distribution in the Silurian palaeo-Southern Hemisphere. This range is now extended for the first time into the palaeo-Northern Hemisphere based on new occurrences from the Silurian of Arctic Russia, from Kotel’ny Island and the Taimyr Peninsula, respectively.

GEOLOGICAL SETTING

R. migrans was found in Upper Silurian strata on Kotel’ny Island during an international expedition in 2011. Kotel’ny Island is the largest island of the Anjou group of the New Siberian Islands located between the Laptev Sea and the East Siberian Sea in the Russian Arctic (Fig. 1a, b). Kotel’ny Island is composed of tectonically deformed platform sequences that include all the systems of the Palaeozoic and Mesozoic eras, except for the Cambrian. The island is a tectonically controlled inlier within Precambrian basement, which includes a significant portion of the marine shelf of the Laptev and East Siberian seas (Natal’in et al., 1999; Kuzmichev & Pease, 2007). Silurian strata are well exposed in the western part of the island (Fig. 1b). The Silurian rocks are up to 1050 m thick, consisting of carbonates and shales of the Urasin (Llandovery Series), Nikola (Wenlock-Ludlow Series) and Eselekh (Pridoli Series) formations (Sobolevskaya, 1976). The Urasin and Nikola formations are composed of predominantly black to dark-grey shales intercalated with thin beds of clayey limestones, whereas the Eselekh Formation is characterized by limestones. R. migrans was recovered from black shales and associated limestone nodules of the Nikola Formation along the Chokurka (75°26′43″N, 137°52′48″E) and Tuor-Yurryakh (75°18′75″N, 139°15′00″E) river sections. Shales yield graptolites of the Saetograptus leintwardinensis Biozone, indicating strata assignable to the early Ludfordian Stage of the Ludlow Series (Sobolevskaya, 1976; Koren’ & Sobolevskaya, 1998; Suyarkova pers. comm. 2013) (Fig. 1d).

The material from the Taimyr Peninsula, Russia (Fig. 1a, c) was recovered from strata of Late Silurian age exposed along the right bank of the Nizhnyaya Taimyra River in the vicinity of Middendorf Cave (Fig. 1c, section P-90218). This outcrop was studied and sampled by AG during fieldwork carried out in 1990.
Fig. 1. Geological setting of localities for *Richteria migrans*. (a) Geographical position of the localities. (b) Geological map of Kotel’ny Island and location of the Tuor-Yuryakh River section. (c) Geological map of the Taimyr Peninsula and location of the Middendorf Cave section (NTD, CTD, STD denote North Taimyr Division, Central Taimyr Division and South Taimyr Division). (d) Sedimentological context of the Tuor-Yuryakh River and Middendorf Cave sections. Map of Kotel’ny Island after Kuzmichev (2009), Taimyr Peninsula after Metelkin et al. (2005). Section for the Taimyr Peninsula after Tesakov (2009); that for Kotel’ny Island after Kuzmichev (2009).
by the Institute of Geology & Geophysics, Novosibirsk. The Taimyr Peninsula lies between the Kara Sea and the Laptev Sea and contains igneous, metamorphic and sedimentary rocks of Proterozoic to Cretaceous ages. Taimyr has generally been divided into three east–west-trending zones, namely North, Central and South Taimyr (Fig. 1c). The Palaeozoic continental margin succession of Central Taimyr represents the stable platform sequence of the northern Siberia passive margin. Silurian strata are well exposed in Central Taimyr, consisting of carbonates and shales up to 500 m in thickness. Zlobin (1962), Tesakov et al. (1995) and Sobolevskaya et al. (1997) established a detailed bio- and lithostratigraphy for this area. The Silurian strata of Central Taimyr comprise the Dvoinaya and Middendorf groups (Tesakov et al., 1995), which are composed of predominantly black to dark-grey shales containing thin lenses and/or interbeds of black limestones. The limestones become thicker and more numerous up the succession.

R. migrans first appears in black mudstones with small micritic nodules in Bed 9 of the Middendorf Group, some 3.5 m above the cephalopod limestone of Bed 8, and is also common in Bed 10 (Bogolepova, 1994; Khiž & Bogolepova, 1995) (Fig. 1d) of section P-90218 (75°20′31.36″N, 99°49′52.83″E). Monograptus fritschi linearis was found in the black shales of Bed 9, indicating an early Ludfordian age (Khiž & Bogolepova, 1995; Tesakov et al., 1995).

**MATERIAL AND METHODS**

The myodocopes consist of partially flattened and 3D-preserved specimens, now housed in the Palaeontological collections, Museum of Evolution, Uppsala University, Uppsala, Sweden, numbers PMU 27241-54. The morphological terminology for the myodocope valve follows that of Siveter et al. (1987). Rock matrix was removed from the specimens mechanically, using fine needles. Casts of the external moulds of ostracods were made using silicone rubber by the technique of Siveter (1982). Photographs were taken using a Leitz Aristophot mounted with a Canon EOS 5D camera, following the methods outlined in Siveter (1990).

**BIOSTRATIGRAPHICAL SIGNIFICANCE**

Perrier et al. (2007) proposed that R. migrans is a marker fossil for rocks of the lower and middle part of the Ludfordian Stage, Ludlow Series across Europe; this can now be extended into Arctic Russia, in the Nikola Formation of Koteln’y Island and the Middendorf Group of the Taimyr Peninsula. R. migrans was previously recorded in the Kopanina Formation, Bohemia, Czech Republic (Perrier et al., 2007); the Long Mountain Silstone Formation, Welsh basin, UK (Siveter, 2009); the La Lande-Murée and La Tavelle formations of the Armorican Massif and La Rouquette Formation of the Montagne Noire, France (Perrier et al., 2007); the Fluminimaggiore Formation, Sardinia (Perrier et al., 2007; Gnoli et al., 2009); the Niewachów (Greywacke) Beds, Poland (Gürich, 1896); the Kurgan Horizon, north Nura-Tau Range, Uzbekistan (Mikhailova & Siveter unpublished data, in Perrier et al., 2007) and possibly in the Ludlow Series of Germany (Perrier & Siveter, 2013) and the Hemse Formation of Gotland (Hede, 1921) (Fig. 2a).

In the Czech Republic R. migrans ranges from the Monograptus fritschi linearis to Bohemograptus bohemicus s. 1. Interzone graptolite biozones (Perrier et al., 2007). In Britain it ranges from the Saetograptus leintwardinensis graptolite Biozone to the Bohemograptus proliferation interval (Siveter, 2009). The graptolites associated with some of the French and Sardinian occurrences of R. migrans are insufficiently studied or too ill-preserved to provide precise independent stratigraphic control; however, based on the available geological information and its consistent stratigraphic range in the Czech Republic and Britain, Perrier et al. (2007) suggest a mid to late Ludlow Epoch age. Furthermore, although the Polish, central Asian, German and Swedish material remains unrevise, R. migrans was always recorded in these areas from the mid Ludfordian Series. The proposition of Perrier et al. (2007) that R. migrans can be used as a marker fossil for rocks assignable to the Ludfordian Stage of the Ludlow Series (Fig. 2a), is strengthened by the new records in the Kotel’ny Island (Ludfordian, Saetograptus leintwardinensis Biozone) and the Taimyr Peninsula (Ludfordian, Monograptus fritschi linearis Biozone).

**PALAEOGEOGRAPHICAL SIGNIFICANCE**

The widespread palaeogeographical distribution of R. migrans, in southerly tropical to mid latitudes on both sides of the Early Palaeozoic Rheic Ocean, was noted by Perrier et al. (2007). The new occurrences expand this distribution to the subtropical region of the palaeo-Northern Hemisphere on, or next to, the Siberia palaeocontinent (Fig. 2b). These new records therefore reinforce the idea that R. migrans had wide dispersal capacity.

The New Siberian Islands along with the western part of the East Siberian Sea belong to the New Siberian Islands microcontinent (Bogdanov et al., 1998; Kośko, 2007; Drachev, 2011; Cocks & Torsvik, 2011; Vernikovsky et al., 2013). During the Late Silurian, the New Siberian Islands Terrane was located in northerly subtropical, 20–25° latitudes. Cocks & Torsvik (2011) placed the New Siberian Islands microcontinent between the Siberia palaeocontinent in the NE, the Kolyma–Omolon Terrane in the NW, and Bactica further south (Fig. 2b).

In contrast, based on palaeomagnetic data, Vernikovsky et al. (2013) and D.V. Metelkin (pers. comm.) place the New Siberian Islands north of the Chukotka Terrane, at about 45°N (Fig. 2b). However, this position seems unlikely because of the Siberian/Baltic affinities of the New Siberian Islands trilobite fauna in Early Palaeozoic strata (Holm & Westergaard, 1930; Danukalova et al., 2012), contrasting with the mix of Laurentian and Siberian faunas found in the Lower Palaeozoic of the Arctic Alaska–Chukotka microcontinent (Cocks & Torsvik, 2011).

The Taimyr Peninsula samples belong to the Central Taimyr Domain, a microplate that was accreted to the Siberian margin in the Late Neoproterozoic (c. 600Ma), and covered by Late Neoproterozoic (Vendian) and Palaeozoic successions (Vernikovsky, 1996; Vernikovsky et al., 2004). During the Late Silurian the Central Taimyr Domain was located in subtropical latitudes, at about 20°N, on the southern part of the Siberia palaeocontinent (Metelkin et al., 2005; Cocks & Torsvik, 2011) (Fig. 2b).

Our new observations confirm that, as noted by Perrier et al. (2007), the occurrences of R. migrans are always associated with either black micritic limestones (e.g. Kote’ny Island, Taimyr, Sardinia, Montagne Noire) or shales (e.g. Bohemia, Armorican Massif, central Asia). These lithologies contain cephalopod-, graptolite-, phyllocarid- and bivalve-dominated assemblages. The myodocope-bearing facies typify possible deep-shelf environments or topographic lows on the shelf (Siveter et al., 1991; Perrier et al., 2011) and are characterized by the lack of bioturbation, the
presence of lamination and by a low diversity, mostly pelagic, fauna. These facies data, added to the cosmopolitan distribution, suggest that *R. migrans* probably possessed a pelagic lifestyle, and this is consistent with the timing of a proposed ecological shift in pioneer pelagic (myodocope) ostracods from benthic to pelagic modes of life during the late Early or early Late Silurian (see Siveter, 1984; Siveter *et al.*, 1987, 1991; Siveter & Vannier, 1990; Perrier *et al.*, 2007, 2011).

**SYSTEMATIC PALAEONTOLOGY**

Order *Myodocopa* Sars, 1866 (*nom. correct.* Pokorný, 1953)

Family *Entomozoidae* Přibyl, 1950 (*= Entomidae Jones, 1873*)

**Remarks.** For the complex nomenclatural status of the Entomozoacea, see Siveter & Vannier (1990).
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Explanation of Plate 1. *Richteria migrans* (Barrande, 1872), Ludlow Series, Silurian: figs A–G, Kotel’ny Island, New Siberian Islands, Russia; figs H–N, Taimyr Peninsula, Russia. A, right valve (PMU 27241); B, right valve (PMU 27242); C, left valve (PMU 27243); D, left valve (PMU 27244); E, right valve (PMU 27245); F, right valve, mould of the inner part of the carapace (PMU 27246); G, left valve (PMU 27247); H, left valve (PMU 27248); I, right valve (PMU 27249); J, left valve (PMU 27250); K, left valve (PMU 27251); L, right valve (PMU 27252); M, right valve (PMU 27253); N, left valve (PMU 27254). All pictures are lateral stereo-pairs. Figs A, D, M and N are silicone casts.
Genus Richteria Jones, 1874
Richteria migrans (Barrande, 1872)
(Pl. 1, figs A–N; Fig. 3)

1995 ‘Entomis’ aff. lamarmorai Canavari, 1900; Kříž & Bogolepova: 580–581, pl. 70, fig. 12.
2003 Richteria migrans taimyrica Abushik, subsp. nov.; Abushik et al.: 122–123, pl. 27, figs 1–2.
2007 Richteria migrans (Barrande, 1872); Perrier et al.: 156–160, figs 6–10. (q.v. for full synonymy).

Material. More than 20 valves from the New Siberian Islands and 8 valves from the Taimyr Peninsula.

Dimensions. Maximum valve length – valve height of largest and smallest well-preserved specimens (Fig. 3): Kotel’ny Island: 3.55–2.05 mm (PMU 27241; Pl. 1, fig. A), 1.15–0.85 mm (PMU 27245; Pl. 1, fig. E); Taimyr: 2.15–1.40 mm (PMU 27254; Pl. 1, fig. N), 0.95–0.75 mm (PMU 27249; Pl. 1, fig. I).

Distribution. Known from strata of Late Silurian age (Ludlow Epoch) of the Czech Republic, France, Sardinia, UK, Poland, Central Asia, Arctic Russia, and possibly Germany and Sweden.

Remarks. Based on what seems to be a more delicate ornament, Kříž & Bogolepova (1995) assigned the Taimyr material to ‘Entomis’ aff. lamarmorai Canavari, 1900, and Abushik et al. (2003) created a new sub-species Richteria migrans taimyrica. Based on our evaluation of the global dataset of R. migrans, we consider this ornamental variation to be intraspecific (see Perrier et al., 2007). The largest Kotel’ny Island specimens are larger (3.5 mm vs. 2 mm) than those described by Perrier et al. (2007) from France and the Czech Republic, but of comparable size with the Sardinian specimens (Fig. 3).

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