An oceanic pathway for Madden–Julian Oscillation influence on Maritime Continent Tropical Cyclones

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While the Madden–Julian Oscillation (MJO) has been shown to affect tropical cyclones (TCs) worldwide through its modulation of large-scale circulation in the atmosphere, little or no role for the ocean has been identified to date in this influence of MJO on TCs. Using observations and numerical model simulations, we demonstrate that MJO events substantially impact TCs over the Maritime Continent (MC) region through an oceanic pathway. While propagating across the MC region, MJO events cause significant sea surface cooling with an area-averaged value of about 0.35 ± 0.12 °C. Hence, TCs over the MC region immediately following the passage of MJO events encounter considerably cooler sea surface temperatures. Consequently, the enthalpy fluxes under the storms are reduced and the intensification rates decrease by more than 50% on average. These results highlight an important role played by the ocean in facilitating MJO-induced sub-seasonal variability in TC activity over the MC region.

INTRODUCTION

The Madden–Julian Oscillation (MJO), a 30–60 days period oscillation, is a dominant mode of variability in the tropical atmosphere. It is one of the most important meteorological phenomena at intraseasonal timescales with wide-ranging impacts on global weather and climate1–3. While the wet or active phase of the MJO is associated with enhanced convection, the dry or suppressed phase is less conducive for convection4,5. Thus, when the active phase of the MJO happens over a region, the large-scale environment becomes more favorable for cyclogenesis, leading to the formation and intensification of tropical cyclones (TCs) in that region6,7,8. It has been observed that TCs tend to form in clusters in different basins coinciding with the favorable or active phase of the MJO9. Reduction in vertical wind shear, enhancement of low-level cyclonic vorticity and an increase in mid-tropospheric humidity are found to be the main factors behind the MJO’s positive influence on TCs in various regions10–13.

While the MJO’s modulation of the atmosphere is well known, its impact on the upper-ocean has been less clear, and consequently, the latter became one of the main focal points of the Dynamics of the Madden–Julian Oscillation (DYNAMO) field campaign14. Besides changes in cloud cover and precipitation, the MJO is also associated with strong surface wind variations. Their joint impact on surface buoyancy and momentum fluxes triggers a significant response from the ocean mixed layer, and consequently the sea surface temperature (SST)21–29. Several studies have shown that a prominent feature of the oceanic response to MJO forcing during the active phase is a SST cooling and a reduction in upper-ocean heat content over a broad area in the eastern equatorial and southeast Indian Ocean regions, which has particular relevance to TCs. Since TCs intensify by extracting heat energy from the ocean, SSTs play a critical role in their development30–33. Despite this, the impact of MJO-induced SST cooling on TCs has not been explored to date. In this study, using a suite of observations and numerical model simulations, we demonstrate that TCs over the Maritime Continent (MC) region are influenced significantly by MJO-induced SST cooling.

RESULTS

Upper-ocean response to MJO

The composite mean upper-ocean response to the active convective phase of the MJO is characterized by a broad region of SST cooling over the MC region, extending approximately between 105°E–150°E and 20°S–Equator (Fig. 1a). This pattern of SST cooling over the MC region stands out as a unique response of the ocean to the MJO in the entire tropics (Supplementary Fig. 1). The maximum SST cooling occurs to the northwest of Australia and exceeds 1 °C. This composite is based on nearly 90 MJO events between 1982 and 2019. MJO-induced SST cooling here is defined as the difference in SST between the end of phase 5 and the beginning of phase 3, which roughly represents the period during the MJO’s lifecycle when it crosses the MC region34,35. The MJO phase and its amplitude correspond to its location and strength, respectively, as defined by the daily real-time multi-variate MJO (RMM) index36. The tendency of the MJO to induce significant cooling over the MC region was also noted in previous studies22,23. This composite is based on the National Oceanic and Atmospheric Administration’s Optimum Interpolation SST (NOAA OI SST), which only assimilates infrared satellite data, and hence could be influenced by cloud-effects. To address this, we also computed the composite using Tropical Rainfall Measuring Mission (TRMM) SST that is based on microwave satellite data (Fig. 1b). The signal, based on 41 MJO events over the period 1998–2014, is broadly consistent with that obtained earlier in terms of the spatial pattern but is stronger in magnitude, especially in certain areas to the south of Java and over the Banda Sea, suggesting that the infrared satellite-based data likely underestimates the SST response to MJO.

The temporal evolution of MJO-induced SST cooling based on NOAA OI SST, averaged over the region 110°E–130°E and 15°S–5°S, demonstrates the persistence of the oceanic response to MJO forcing (Fig. 1c). Beginning on day 0, which represents the first day of MJO phase 3, SST cools rapidly until around day 20 with the magnitude of area-averaged cooling reaching a maximum value of about 0.35 ± 0.12 °C. The time taken by the MJO event to progress from the beginning of phase 3 to the end of phase 5 is
15 ± 7 days, which is in good agreement with the timing of the peak SST cooling. Beyond day 20, the recovery of SST is a rather slow process. On day 50, nearly a month after the maximum cooling occurs, the average cooling is still nearly 0.2 °C below the pre-active phase MJO SST. The time evolution of MJO-induced SST cooling based on TRMM microwave SST (Fig. 1d) is in good agreement with that obtained using NOAA OI SST. Considering the composite mean, the maximum SST cooling occurs on day 22 albeit with a stronger magnitude of about 0.54 ± 0.23 °C. Beyond this, the SST anomaly recovers slowly and the cooling reaches a value of about 0.30 ± 0.21 °C on day 50.

To identify the role of various upper-ocean processes in the MJO-induced SST cooling over the MC region, a mixed layer heat budget analysis was performed using an eddy-permitting ocean reanalysis that assimilates satellite SSTs and all available hydrographic measurements, including data from Argo floats. The analysis, broadly consistent with previous studies, reveals that both surface fluxes at the air–sea interface and vertical processes in the sub-surface ocean are responsible for the SST change under the MJO, and that the shallow seas and climatological mean mixed layers in the MC region play an important role in its larger SST response compared to other regions (see Supplementary Note 1).

**Impact on TCs**

The region to the north of Australia over the MC region is home to nearly 15% of global TC activity. Further, the region to the northwest of Australia, between 110°E−130°E and 20°S−10°S, happens to be the most TC-prone region in the entire Southern Hemisphere with 75% of the most intense Australian TC landfalls occurring along the coast here. Although this region is sparsely populated, it has much commercial significance, particularly oil and gas production. Earlier, we have seen how MJOs induce substantial sea surface cooling over this region while crossing the MC region. Further, the TC season over this region extends from December to April, coinciding with the seasonal peak of MJO activity over the MC region. Thus, the natural question it leads us to is whether the MJO-induced SST cooling can influence TCs to the northwest of Australia. To answer this, we performed a Lagrangian along-track composite analysis of the large-scale ocean-atmosphere environment affecting TCs for the 34-year period 1985–2018 (see "Methods").

Between days 5 and 24 after the beginning of the MJO event, the atmosphere is highly favorable for TCs. As the sea surface begins to cool, the atmosphere is very conducive for TC development over this period. On average, the vertical shear of zonal wind is more easterly, the 700 hPa relative humidity increases by 4.5%, the low-level relative vorticity is more cyclonic, the outgoing longwave radiation is reduced by 15.65 W m⁻² and the sea-level pressure drops by nearly 200 Pa (Supplementary Table 1). Consequently, the mean TC intensification rate increases by about 0.6 m s⁻¹ 6-hr⁻¹, and more than 40% of all western Australian 6-hourly TC track locations are found during this period. These results are in line with those from previous studies that have suggested the MJO’s favorable influence on TCs in this region during phases 3–5.

Subsequently, between days 24 and 41, which roughly represents the period after the MJO-active phase has left the MC region, the positive atmospheric influence of the MJO subsides to a large extent and significant sea surface cooling prevails over this region (Fig. 1c, d). The number of 6-hourly TC occurrences along the coast here, particularly oil and gas production. Earlier, we have seen how MJOs induce substantial sea surface cooling over this region while crossing the MC region. Further, the TC season over this region extends from December to April, coinciding with the seasonal peak of MJO activity over the MC region. Thus, the natural question it leads us to is whether the MJO-induced SST cooling can influence TCs to the northwest of Australia. To answer this, we performed a Lagrangian along-track composite analysis of the large-scale ocean-atmosphere environment affecting TCs for the 34-year period 1985–2018 (see "Methods").

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track locations during this period decreases to 20%, likely due to the reduced MJO’s positive influence on cyclogenesis through the atmospheric pathway. We denote this period as the MJO’s Window of Oceanic Influence (MWOI). To clearly isolate the role of MJO-induced oceanic cooling on TC intensification during this period, we sub-sample the data so that in addition to distributions of initial storm characteristics, distributions of pre-storm atmospheric parameters are also statistically similar between sample sets used for comparison (see Supplementary Table 2). This is necessary because a period of suppressed MJO convection typically follows the convectively active phase\cite{22}, thus raising a possibility for some potential atmospheric influence to act on top of that from the ocean. The pre-storm SST anomaly for TCs outside the MWOI is 0.86 °C, which reduces significantly to 0.33 °C inside the MWOI representing a decrease of 0.53 °C. In response to this cooling, the anomalous dry stability in the lower troposphere (1000–850 hPa) increases by 0.0024 K Pa$^{-1}$ and the anomalous enthalpy flux out of the ocean reduces from about 47 W m$^{-2}$ outside the MWOI to about 9 W m$^{-2}$ inside the MWOI, indicating a reduction of about 38 W m$^{-2}$. Consequently, the TC intensification rates are significantly reduced inside the window. While the mean intensification rate outside the MWOI is 1.29 m s$^{-1}$ 6-hr$^{-1}$, it reduces by more than 50% to 0.45 m s$^{-1}$ 6-hr$^{-1}$ within the MWOI.

The probability distribution functions (PDFs) of pre-storm SSTs, enthalpy fluxes under TCs and the 24 hr TC intensification rates further reinforce these results (Fig. 2). Compared to the PDFs for TC locations outside the MWOI, the PDFs for TC track locations inside the MWOI are skew to the left indicating relatively lower pre-storm SSTs (Fig. 2a) and air–sea enthalpy fluxes (Fig. 2b), and a reduction in TC intensification rates (Fig. 2c) for the latter. These differences are statistically significant at the 95% level based on the Student’s t-test for difference of the means. Although we attribute differences in enthalpy fluxes under TCs to those in pre-storm SST differences, minor contributions are also possible from differences in surface winds (see Supplementary Note 2). Some previous studies suggested that the potential intensity for a TC, and consequently its intensification, is more closely related to the relative SST or the absolute SST minus the tropical mean SST\cite{45}. A sensitivity test indicates that similar results are obtained when absolute SST is replaced with relative SST in our analysis. Although these results strongly support the role of MJO-induced oceanic cooling in suppressing TC intensification inside the MWOI, we acknowledge that the ocean’s negative impact on TCs may be modulated by some lingering atmospheric effects of the MJO.

Model

While results presented thus far are based on observations, numerical simulations of Cyclone Olga using a high-resolution regional atmospheric model (see “Methods”) provide further support. Olga, a tropical storm from the 1999–2000 Australian
Fig. 3 Simulating MJO’s impact on TC Olga (2000). a Track of TC Olga from observations (black) and model simulations (WAKE (green) and NO-WAKE (orange)). While the “WAKE” case represents the set of simulations where the model is forced by observed SST, the “NO-WAKE” case represents the set of simulations where the model is forced by SSTS from which the influence of MJO is removed. Both “WAKE” and “NO-WAKE” each consist of 8 simulations and the ensemble mean tracks are shown. The color in the background represents the SST change induced by prior MJO, just before Olga’s formation on 15 March 2000. Dates are shown near tracks to indicate the direction of Olga’s movement. b Time evolution of Olga’s 10 m maximum winds (kt) from observations and model simulations. Time evolution of difference between NO-WAKE and WAKE simulations for c 10 m maximum wind (kt, green) and minimum sea-level pressure (hPa, red), and d along-track pre-storm SST (°C, brown) and enthalpy flux at the air-sea interface under the storm (W m⁻², purple). As in panel (a), ensemble mean values are shown in panels (b–d). The shading in panel (b), and the vertical bars in panels (c, d), indicate the ensemble spread as represented by the s.d.

DISCUSSION

Past research has documented the negative effects of lingering cold wakes of TCs on subsequent TCs that encounter them. Similarly, it was found that strong MJO events weaken subsequent events that immediately follow them through their SST footprints. The results from this study, where we demonstrate the cross-phenomena interaction between MJOs and TCs through the ocean, represent a shift in the prevailing understanding that MJOs can primarily influence TCs through a modulation of the large-scale atmospheric circulation. Also, they point towards an improvement in prediction skill through realistic representation of air-sea coupled processes in models.

Furthermore, the oceanic influence of the MJO on Australian TCs presented here can also have longer-term consequences. For instance, stronger MJOs tend to induce more sea surface cooling as they pass over the MC region (Supplementary Fig. 2a). Also, the strength of MJOs while over the MC region appears to be on the decline (Supplementary Fig. 2b) as a result of gradual eastward expansion of the background climatological warm pool. Thus, changes in MJO characteristics over time can induce corresponding changes in the MJO’s influence on TCs over this region. Finally, the MJO-induced SST cooling over the MC is tightly related to the upper-ocean stratification over that region (Supplementary Fig. 3). With the latter projected to change under global warming, the “braking” effect of MJOs on TC intensification over the MC region could be altered in future.

METHODS

Data

We obtain daily NOAA OI SST from https://psl.noaa.gov/data/gridded/data.noaa.oisst.v2.highres.html for the period 1983–2019 to estimate MJO-induced SST change. We also obtain TRMM Microwave Imager daily SST for the period 1998–2014 from http://www.remss.com/missions/tmi/ to verify the signal based on NOAA OI SST. Five-day mean vertical ocean temperature and salinity profiles, and surface ocean fluxes from SODA3 (version 3.3.1) ocean reanalysis, obtained from https://www.soda.umd.edu for the period 1980–2015, are used to understand the role of various oceanic processes.
upper-ocean processes in MJO-induced SST change. To identify MJOs, we use time series of daily RMM index values, obtained for the period 1980–2018 from the Australian Bureau of Meteorology (http://www.bom.gov.au/climate/mjo/graphics/rmm.74toRealtimetext.txt).

TC track data for the period 1985–2018 are obtained from https://emanuel.mit.edu/products and used to estimate the values of various atmospheric parameters along TC tracks. WHOI OAFlux daily latent and sensible heat fluxes, obtained from http://oaflux.whoi.edu are used to estimate the enthalpy flux transfer at the air–sea interface under TCs. Daily NCEP-DOE 2 atmospheric reanalysis data39, obtained from (https://psl.noaa.gov/data/gridded/data.ncep.reanalysis2.html), and daily NOAA outgoing longwave radiation (OLR)40 obtained from https://www. psl.noaa.gov/data/gridded/data.interp_OLR.html, are used to estimate various TC-relevant atmospheric parameters and outgoing longwave radiation anomalies along TC tracks.

Calculations
MJO events are identified using the following procedure. First, a date is identified in the RMM index time series on which the phase 3 amplitude is larger than 1. The next 30 days are then examined to see if phases 4, 5, and 6 occur at least once irrespective of the amplitude. If these two conditions are met, the MJO event is said to have propagated over the MC region and selected for our analysis35.

The following parameters are estimated along TC tracks to understand the pre-storm large-scale environment: vertical shear of zonal wind estimated between 200 and 850 hPa, relative humidity at 700 hPa, relative vorticity at 850 hPa, vertical velocity at 300 hPa, sea-level pressure, outgoing longwave radiation, and SST5. Specifically, we follow each MJO event and look for TC track locations that fall within a period. For each identified track location, we compute various pre-storm atmospheric parameters, pre-storm SST, the enthalpy flux under the storm and the intensification rate. Next, we follow a similar procedure to compute the same environmental and storm parameters for those TC track locations that fall outside that period. Finally, we examine the composite mean differences between the two sets of data and test for statistical significance in various parameters.

While the atmospheric parameters are averaged over a 5° × 5° box centered over the storm, pre-storm SST and enthalpy fluxes are averaged over a 2° × 2° box centered over the storm. The atmospheric parameters and pre-storm SST are estimated 2 days before the arrival of the storm. The enthalpy flux, on the other hand, is estimated on the day of the storm. The TC intensification rate at a location is estimated as the slope of the linear regression of the maximum wind speed over five successive 6-hourly track locations beginning with the current location. The dry stability of the atmosphere is estimated as $\delta$ where $\delta$ is the potential temperature and $p$ is the pressure. Throughout the study for various analysis, the seasonal cycle is removed from different parameters. We use the Monte Carlo method of repeated random sampling to generate error bars for the PDFs. From a given distribution, we randomly select half of the samples to generate a PDF and repeat this process a hundred times. Following this, the mean and standard deviation computed across the PDFs yield the mean PDF and error bars, respectively.

Numerical experiments
To illustrate the MJO’s oceanic influence on TCs, we performed numerical simulations of tropical storm “Olga” (15–19 March 2000), which occurred 5 days after the end of a MJO’s phase 5, using the non-hydrostatic WRF-ARW version 4.053. The model domain covers the region between 95°E–125°E and 30°S–Equator, with 4 km horizontal resolution and 35 terrain following vertical levels up to 50 hPa. The initial and boundary conditions, including SST, are from ERAS reanalysis34 (https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/eras5) with 31 km horizontal resolution, available from https://www2.mmm.ucar.edu/wrf/users/download/free_data.html. Boundary conditions are updated every 6 h. The physical parameterizations used are detailed in Supplementary Table 3. Eight ensemble members are created using combinations of four cloud microphysical schemes and two planetary boundary layer schemes. The integration of the model is carried out from 00 UTC 15 March 2000 to 12 UTC 19 March 2000. Using aforementioned model design, two sets of eight-member ensemble experiments are performed.

In the first set of experiments (WAKE), which represents the control case, TC Olga is forced with observed SST. In the second set of experiments (NO-WAKE), TC Olga is forced with a SST from which the cooling induced by the prior MJO is removed. More specifically, the magnitude of negative SST anomaly on 00 UTC 15 March 2000, calculated with respect to the 10-day mean SST prior to phase 4 of the MJO, is added to the initial SST as perturbation in the NO-WAKE experiment. The initial perturbation is reduced every 6 h using the composite mean evolution of SST cooling averaged over the region 110°E–130°E and 15°S–5°S (Fig. 1c).

DATA AVAILABILITY
Model output are available on request from KB. All other datasets used in this study are freely available for download from the various links provided in “Methods”.

CODE AVAILABILITY
The WRF-ARW model code is freely available for download from the link provided in “Methods”.

Received: 24 May 2021; Accepted: 29 September 2021; Published online: 27 October 2021

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ACKNOWLEDGEMENTS

K.B., L.R.L., S.M.H. and S.K. are supported by the Office of Science (BER) of the U.S. Department of Energy as part of the Regional & Global Model Analysis (RGMA) program area. The Pacific Northwest National Laboratory is operated for DOE by Battelle Memorial Institute under contract DE-AC05-76RL01830. The computations were mainly carried out using the computing resources at the National Energy Research Scientific Computing Center (NERSC), a U.S. Department of Energy Office of Science User Facility located at Lawrence Berkeley National Laboratory, operated under Contract No. DE-AC02-05CH11231.

AUTHOR CONTRIBUTIONS

K.B., L.R.L. and S.M.H. developed the main idea. K.B. performed the analysis of observations. K.B., L.R.L. and S.M.H. wrote the paper. S.K. conducted model simulations and analyzed the output.

COMPETING INTERESTS

The authors declare no competing interests.

ADDITIONAL INFORMATION

Supplementary Information The online version contains supplementary material available at https://doi.org/10.1038/s41612-021-00208-4.

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