ZETA POTENTIAL CHANGING IN COMPRESSED CLAYS

Natalia A. Medvedeva, Karine A. Alvanyan, Yulia O. Malgina, Valery V. Seredin
Perm State National Research University (15, Bukireva st., Perm, 614068, Russian Federation)

ИЗМЕНЕНИЕ ДЗЕТА-ПОТЕНЦИАЛА ГЛИН, ПОДВЕРЖЕННЫХ СЖАТИЮ

N.A. Medvedeva, K.A. Alvanian, Ю.О. Мальгина, В.В. Середин
Пермский государственный национальный исследовательский университет (614068, Россия, г. Пермь, ул. Букireva, 15)

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Experimental results concerning the effect of loading pressure on the electrokinetic potential of kaolinite and montmorillonite particles in aqueous suspension are presented. It has been determined that the investigated suspensions of kaolinite and montmorillonite are aggregatively and sedimentally stable. The aggregative stability is evidenced by the calculated energy of the molecular motion (10^{-8} \text{ J/m}^3). The monitoring of particle size indicates sedimentation stability. It was found that the change in the electrokinetic potential of clay particles depends on the loading pressure in different ways. Three classes were distinguished: 0–125 MPa, 150–750 MPa, 800–1200 MPa. In each class, characteristic changes in the electrokinetic potential are observed, due to the nature of the clay and the fractional composition. The fractional composition of the clay is also changed while compression. Pressure load has a different effect on the change in electrokinetic potential for kaolinite and montmorillonite. To establish the influence of the particle size distribution on the conditions of formation of the zeta potential on the surface of clay particles, a statistical correlation analysis was performed. It was established that for montmorillonite in the 1st and 2nd classes, the particle size distribution and zeta-potential have statistical correlations, and in 3rd class - not, whereas for kaolinite, statistical correlations between the particle size distribution and zeta-potential are observed in the 1st and 3rd classes, and in the 2nd class - not. The resulting changes in the zeta potential are associated with the processes of dispersion and aggregation, which are implemented more intensively at low pressures (p < 150 MPa).

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Natalia A. Medvedeva (Author ID in Scopus: 55164437300) – PhD in Chemistry, Associate Professor at the Department of Physical Chemistry (tel.: +007 902 478 43 11, e-mail: nata-kladova@yandex.ru).
Karine A. Alvanyan – Senior Lecturer at the Department of Engineering Geology (tel.: +007 342 239 64 39, e-mail: karishuta@yandex.ru).
Yulia O. Malgina – Master's student at the Department of Engineering Geology (tel.: +007 342 239 64 39, e-mail: karishuta@yandex.ru).
Valery V. Seredin (Author ID in Scopus: 56974440000) – Doctor of Geology and Mineralogy, Professor, Head of the Department of Engineering Geology (tel.: +007 342 239 64 39, e-mail: sendin@nedra.perm.ru). The contact person for correspondence.

Медведева Наталья Александровна – кандидат химических наук, доцент кафедры физической химии (тел.: +007 902 478 43 11, e-mail: nata-kladova@yandex.ru).
Алвания Карине Антоновна – старший преподаватель кафедры инженерной геологии (тел.: +007 342 239 64 39, e-mail: karishuta@yandex.ru).
Мальгина Юлия Олеговна – магистр кафедры инженерной геологии (тел.: +007 342 239 64 39, e-mail: karishuta@yandex.ru).
Середин Валерий Викторович – профессор, доцент геолого-минералогических наук, заведующий кафедрой инженерной геологии (тел.: +007 342 239 64 39, e-mail: seredin@nedra.perm.ru). Контактное лицо для переписки.
Introduction

The clays properties are largely determined by the energy potential formed on the clay particles surface [1, 2]. One of the components of the energy potential is the electrokinetic potential (zeta potential), which is formed at the sliding boundary between the adsorption and diffusion layers [3]. Zeta potential is determined both by the charge of the particle itself and by external environmental conditions. There are several factors that determine the electrical charge on the clay particles surface: defects in the crystal lattice [4–13], pH of the solution, type and concentration of cations in pore solution, effect of indifferent electrolytes [9, 14, 15], effect of non-indifferent electrolytes, effect of temperature [15, 16], influence of the dispersion medium nature [15–17]. The magnitude of the particle charge significantly affects on formation of strength [18] and adsorption properties of clays [19–22], as well as transformation of fluid in the soil mass [23]. So, in the article V.I. Osipova [24] process of water molecules adsorption by silicate surfaces of clay minerals, leading to the formation of nanoscale films of adsorbed water is described.

Much attention is paid to the study of the electric charge of dispersed soils particles and the laws of its formation [4–6, 25–27]. The factor that determines the clays adsorption activity and, as a consequence, the particles charge, is the loading pressure. As a result of high pressure and pressure with a shift in certain areas of a solid, a stress field is created, the relaxation of which occurs in different ways: heat generation, formation of new surfaces and short-lived active centers, formation of plastic waves, the interaction of which leads to the formation of various types of defects in the structure and metastable states [28–31]. Experimental study of kaolinite, halloysite, montmorillonite, sepiolite, palygorskite and illite structures behavior under the influence of different types of pressure made it possible to establish that the degree of structures perfection and the interpacket material distribution depend both on the interlayer bonds nature and the type of pressure [32].

The effect of pressure (technogenic load) on the zeta potential of clay minerals is rather poorly studied; therefore the purpose of this study is to establish the pattern of zeta potential changing in clays prone to shear compression. The objects of study are Lobanov’s montmorillonite and Chelyabinsk’s kaolinite. According to the results of X-ray analysis [33], montmorillonite clay consists of: montmorillonite (75 %), kaolinite (3.6 %), quartz (11.4 %), albite (6.7 %), calcite (3.3 %). Kaolinite clay contains: kaolinite (76.7 %), montmorillonite (15.6 %), quartz (7.7 %).

The clay was subjected to compression, which was produced on a specially designed and manufactured device according to the method described in [34]. Before experiments, clays were cleaned by sedimentation method of 2 % aqueous clay suspensions. For pressing, clay samples with fractions greater than 0.01 mm were used.

In research suspensions content 0.5 % of clay.

**Patterns of changes in zeta potential in clays subjected to mechanical pressure treatment**

Determination of the clays electrokinetic characteristics (electrokinetic potential) in water, as well as structural parameters (particle size) was carried out using an automated system of electrokinetic properties studying of colloidal systems and surfaces DelsaNanoC of Beckman Coulter (USA). The particle size was measured in a cell with a quartz cuvette, and the zeta potential of clay particles was measured in a flow cell. Measurements of particle size and electrokinetic characteristics are based on photon correlation spectroscopy (PCS), which use principle of dynamic light scattering and electrophoretic light scattering [35].
Simultaneously with the measurement of particle size and zeta potential of clay suspensions, the structural parameters were monitored over time. The criterion for these measurements was the accumulation time that characterizes the accumulation of data and is expressed in arbitrary units [1]. The determination of polydispersity (P.I.) was carried out on the basis of the intensity distribution depending on the particle size.

One of the important factors in the electrokinetic characteristics measurement is the stability of the systems. In this case, we are talking about both aggregative and sedimentation stability. The attainment of sedimentation stability of the clay suspensions was confirmed by the results of over time particle size monitoring (Fig. 1).

![Fig. 1. Monitoring the size of montmorillonite clay particles in water](image)

In order to assess the aggregative stability of particles of clay minerals, the energy of thermal motion of particles was calculated

$$\sigma_m = \gamma \frac{kT}{d^2},$$

where $\gamma \approx 10$ – dimensionless coefficient; $k$ – the Boltzmann constant (1,38·10$^{-23}$J/K); $T$ – temperature, K; $d$ – the size of the structural unit, m [37].

This characteristic is used as a criterion for aggregative stability, to calculate which the average diameter of clay particles in the aqueous suspension was determined. For some pressures, the results are shown in Table 1.

| $P$, MPa | $d_{av}$, 109, m | P.I. | $\sigma_m$, 108, J/m$^2$ | $d_{av}$, 109, m | P.I. | $\sigma_m$, 108, J/m$^2$ |
|---------|----------------|------|----------------|----------------|------|----------------|
| 0       | 720 ± 15       | 0.30 | 7.93           | 800 ± 59       | 0.34 | 6.43           |
| 60      | 883 ± 85       | 0.36 | 5.26           | 680 ± 40       | 0.28 | 8.88           |
| 105     | 776 ± 90       | 0.32 | 6.81           | 877 ± 23       | 0.24 | 5.34           |
| 200     | 592 ± 63       | 0.30 | 11.71          | 1023 ± 85      | 0.35 | 3.92           |
| 550     | 1450 ± 100     | 0.34 | 1.96           | 714 ± 47       | 0.20 | 8.06           |

It should be noted that, along with nanoscale particles, the presence of their agglomerates, as well as micron-range grains, is characteristic of dispersed systems. These large particles cannot be determined, since they do not possess sedimentation stability (they quickly settle) and are not suspended in the measurement process [36, 37]. The calculated values of $\sigma_m$ are in the range from 10$^{-7}$ to 10$^{-8}$ J/m$^2$, which indicate the aggregative stability of the system, but the kinetically – studied systems are weakly stable.

The homogeneity of the system is indicated by the polydispersity index, whose values do not exceed 0.4.

After carrying out a series of tests, graphs of changes in the zeta-potential of kaolinite and montmorillonite clay particles were plotted depending on the loading pressure (Fig. 2).

![Figure 2](image)

Figure 2, a shows the dependence of the zeta potential of montmorillonite clay from pressure. From the graph it can be seen that at a pressure from 0 to 125 MPa, the zeta potential decreases in absolute value from –36 to –16 mV. At a pressure from 150 to 750 MPa, the zeta potential increases in the absolute value from –18 to –38 mV, and at pressure from 800 to 1200 MPa, the zeta potential varies chaotically from –30 to –40 mV. Figure 2, b shows the dependence of the zeta potential of
kaolin clay from the pressure; three areas can be distinguished in the graph: the first, at pressures from 0 to 125 MPa, the zeta potential varies from –55 to –45 mV, the second, from 150 to 750 MPa, where zeta potential increases from –52 to –30 mV. The third is in the pressure range from 800 to 1200 MPa, in which the zeta potential varies chaotically.

It can be seen that the load affects kaolinite and montmorillonite differently. This circumstance may be due to several reasons. First, isomorphic substitutions in the crystal lattice of clay minerals: for example, for clay minerals of type 2: 1 (minerals that consist of two tetrahedral and one octahedral sheet, which include montmorillonite), substitutions of Si⁴⁺ by Al³⁺ in the tetrahedral sheet and substitutions of Al³⁺by Mg²⁺ and Fe²⁺ in the octahedral sheet [6, 9]. It is believed that isomorphic substitutions are not characteristic of kaolinite [6], however, a number of authors note that isomorphic substitutions of Al³⁺ by Si⁴⁺ occur in tetrahedral sheets in kaolinite [7, 10]. Isomorphic substitutions are known to affect the charge of clay minerals basal surfaces (siloxane and hydroxyl) [12]. The structural sheet of kaolinite has a siloxane and hydroxyl surfaces, while the structural sheet of montmorillonite has only siloxane surfaces.

Secondly, the formation of a constant structural charge is also due to the electronegativity of the chemical elements atoms that the clay mineral is consisted of [38–41].

Third, the fractional composition. An increase in the content of the fine fraction due to the clays mechanical activation leads to an increase in the zeta potential absolute value [29, 30]. In addition, it is necessary to take into account the electric charge localization heterogeneity in different parts of the particles of a clay mineral.

Thus, the classes allocated for Lobanov’s montmorillonite according to the graph of dependence of R on P are the following: 1st class – 0–125 MPa, 2nd class – 150–600 MPa, 3rd class – 650–1200 MPa. For the Chelyabinsk kaolinite according to the dependence of R on P, the following: 1st class – 0–125 MPa, 2nd class – 150–400 MPa, 3rd class – 500–1200 MPa.
The influence of particle size distribution on the zeta-potential formation conditions on the clay particles surface

To study the effect of the particle size distribution on the zeta potential value change three graphs of the change in the zeta potential of clays were plotted depending on the percentage of fractions with a particle size: less than 0.1; 0.1–0.2; 0.2–0.5; 0.5–1.0; 1–2; 2–5; 5–50 microns. The method of the granulometric composition determining of the clays under pressure, as well as the rationale of the effect of pressure on the formation of the granulometric composition of the clays, were previously studied and described in [33, 34]. Experimentally it was found that with increasing pressure on the clay, there is a general tendency of the clay fractions content decreasing and the silt fraction increasing. Such changes occur more intensively in kaolinite clay than in argillite-like montmorillonite. It was revealed that in the pressure range of 0–125 MPa, the change in the clays fractional composition proceeds more intensively than at higher pressures.

Based on the obtained results of particle size analysis, a statistical analysis was carried out and the correlation coefficient for each fraction by classes was derived. The results for montmorillonite and kaolinite are presented in Tables 2 and 3 respectively.

Table 2 shows that for montmorillonite in classes 1 and 2, size distribution and zeta potential have a statistical correlation, since the calculated correlation coefficients are larger than table values. In class 3, the correlation is weak, due to the lack of influence of particle size distribution on the formation of the electrokinetic potential.

It should be noted that in class 1 (0–125 MPa), the zeta potential decreases with an increase in the percentage of the clay fraction (less than 0.1–5.0 µm), whereas with an increase in the content of the silt fraction (5–50 µm), the potential increases. In class 2 (150–750 MPa), an increase in the zeta potential is observed with an increase in the percentage of fractions 0.1–5.0 µm. And with an increase of content of the fraction less than 0.1 and 5–50 µm, the zeta potential value decreases. It is assumed that an increase in the content of the clay fraction particles, which remain in suspension, leads to an increase in zeta potential.

Table 3 presents the results of statistical analysis for kaolinite, it is shown that in classes 1 and 3, the particle size distribution and the zeta potential do not have statistical relationships.

Table 2

| Fraction  | Class 1 $r(n=21)$ | Class 2 $r(n=30)$ | Class 3 $r(n=15)$ |
|-----------|------------------|------------------|------------------|
| Less than 0.1 | -0.52 | -0.43 | -0.36 |
| 0.1–0.2   | -0.50 | -0.43 | -0.36 |
| 0.2–0.5   | -0.59 | 0.70 | 0.36 |
| 0.5–1.0   | -0.66 | 0.66 | 0.36 |
| 1–2       | -0.48 | 0.69 | 0.36 |
| 2–5       | -0.54 | 0.67 | 0.36 |
| 5–50      | 0.69 | 0.43 | -0.60 |

Note: here and in Table 3 red indicates insignificant correlation coefficients.

Table 3

| Fraction  | Class 1 $r(n=18)$ | Class 2 $r(n=30)$ | Class 3 $r(n=15)$ |
|-----------|------------------|------------------|------------------|
| Less than 0.1 | -0.16 | -0.47 | 0.58 |
| 0.1–0.2   | -0.20 | -0.47 | -0.15 |
| 0.2–0.5   | -0.21 | -0.47 | -0.66 |
| 0.5–1.0   | -0.24 | -0.47 | -0.64 |
| 1–2       | -0.22 | -0.47 | -0.72 |
| 2–5       | -0.10 | -0.47 | -0.75 |
| 5–50      | 0.20 | 0.47 | 0.61 |

Values of correlation coefficients between the zeta potential and granulometric composition of kaolinite clay

In class 3 (800–1200 MPa), with an increase in the percentage of fractions 0.1–5.0 µm, the zeta potential decreases. With an increase in the percentage of fractions less than 0.1 and 5–50 µm, the value of the zeta potential increases.

Table 3 presents the results of statistical analysis for kaolinite, it is shown that in classes 1 and 3, the particle size distribution and the zeta potential do not have statistical relationships.
In class 2, excepting the fraction 0.1–0.2, statistical relationships are observed between the particle size distribution and the zeta potential.

Thus, the described changes in the zeta potential can be associated with the processes of dispersion and aggregation, which are implemented more intensively at low pressures ($P < 150$ MPa). It is assumed that the particles of the silt fraction precipitate, and the content of sedimentation-resistant particles of the clay fraction, which remain suspended, decreases, which leads to a drop in the zeta potential.

Conclusion

Physico-chemical properties of soils are largely determined by the size of their structural elements. The magnitude of the electrokinetic potential is a criterion, the use of which will make it possible to predict the structural changes of clay minerals in various environments.

With increasing pressure on clay, the electrokinetic potential of the particles changes in different directions. Zeta-potential formation is associated with both the size of the particles and their mineral composition.

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