Sensitivity of Enhanced SRTM for Sea-Level Rise Variation on Egyptian Coasts

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Received March 24, 2021; Revised April 19, 2021; Accepted May 28, 2021

Cite This Paper in the following Citation Styles

(a): [1] Alshimaa A. Hussien, Mohamed Gomaa M. Mohamed, Pasent H. A. Yousef, Mohamed R. Hagag, Tarek Abou El Seoud, “Sensitivity of Enhanced SRTM for Sea-Level Rise Variation on Egyptian Coasts,” Civil Engineering and Architecture, Vol. 9, No. 3, pp. 958 - 968, 2021. DOI: 10.13189/cea.2021.900337.

(b): Alshimaa A. Hussien, Mohamed Gomaa M. Mohamed, Pasent H. A. Yousef, Mohamed R. Hagag, Tarek Abou El Seoud (2021). Sensitivity of Enhanced SRTM for Sea-Level Rise Variation on Egyptian Coasts. Civil Engineering and Architecture, 9(3), 958 - 968. DOI: 10.13189/cea.2021.900337.

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Abstract The accuracy of digital elevation models (DEMs) is very critical to planning adaptation strategies for coastal areas under changing scenarios of global climate and sea-level rise. This research assesses the accuracy of the interferometric DEM from the SRTM mission as it is one of the free available DEMs that have been widely used in many applications. The SRTM model is enhanced using the best fit global geoid model for Egypt instead of EGM96 and by applying a mathematical scale factor formula. This research aims to obtain a higher vertical accuracy of the SRTM model that achieves the requirements for coastal inundation studies with minimum field measurement data. The methodology approach consists of three phases. In the first phase, the enhancement scale factor formulas were derived using uniformly distributed ground control points (GCPs) at two different terrains data in Egypt. The second phase contains the evaluation and validation process. It was observed that the enhanced accuracy achieved ranged from 45% to 60% based on the type of terrains. In the last phase, the sensitivity of these DEMs to sea-level projections using the recent available local tide gauge data was analyzed. It is recommended to investigate this method approach to determine the optimal size and distribution of reference ground control points needed to adjust various free open DEMs of low vertical accuracy for further studies and other applications.

Keywords SRTM, Vertical Accuracy, Coastal Inundation, Scale Formula, Enhancement, Tide Gauge

1. Introduction

Nowadays, planning, development, and risk management along the Egyptian coasts need to conduct several studies including topography, water resources, climate change, and sea-level rise [1,2]. Sea level rise (SLR) currently poses a significant threat to coastal areas in Egypt and around the whole world. During the period 1901–2010, the global mean sea level rose by 19 centimetres and it is very likely that the SLR will exceed the observed rate of 2.0 mm/year during 1971–2010, with the rate of 8 to 16 mm/year during 2081–2100[3]. The existence of a reliable and precise digital elevation model (DEM) is greatly required to support the decision-makers for planning the needed adaptation strategies to sea level rise. A Digital Elevation Model (DEM) is referred to as a digital representation of the topographic surface in three dimensions with height (elevation) described above mean sea level. Traditionally, A DEM can be obtained using different sources such as Terrestrial survey, Aerial photo, LIDAR, and Satellite survey. Terrestrial survey uses both the conventional and modern surveying instruments to produce the DEM such as Global Navigation Satellite
System (GNSS), Total Stations, and leveling. Although this method gives reasonably accurate results, the method is time-consuming and less efficient when applied to an area which is not easily accessible. The elevation accuracy of DEM obtained using the sources like LIDAR and aerial survey are good and cover a large area compared to Terrestrial survey. But it is restricted to do aerial surveys in some countries because of security reasons and limited resources of data collection [4,5]. There is now a revolution in the use of satellite survey data to overcome many of the obstacles facing traditional surveying methods, as there is a future promising that higher accuracy DEMs will be freely available for the globe.

Currently, numerous freely open-access global DEMs with large global coverage are available, for example, the Advanced Space borne Thermal Emission and Reflection Radiometer (ASTER), the Shuttle Radar Topography Mission (SRTM), the Global Multi-resolution Terrain Elevation Data 2010 (GMTED2010), the Global 30 Arc-Second Elevation (GTOPO30) and EarthEnv-DEM90 digital elevation models [6]. ASTER and SRTM are the two most commonly Global DEMs that are addressed in several studies. ASTER is a cooperative effort between Japan’s Ministry of Economy, Trade, and Industry (METI) and U.S. National Aeronautics and Space Administration (NASA) that was released in October 2011. The characteristics of ASTER are a spatial resolution of 1 arc sec (~30 m) and a height accuracy of about 17m [7,8,9]. On the other hand, SRTM is a joint effort of NASA, the National Geospatial-Intelligence Agency (NGA), the German Aerospace Center (DLR), and the Italian Space Agency (ASI). The recent version (4.1) of SRTM with a spatial resolution of 1 arc sec (30 m) and around 16 m of vertical height accuracy was released in September 2014 [10,11,12]. EarthEnv-DEM90 with a spatial resolution of 3 arc sec (~90m) is a compilation dataset that is merged from ASTER GDEM v2 and SRTM v4.1[13]. GMTED2010 is developed by a collaborative effort of the U.S. Geological Survey (USGS) and the NGA. GMTED2010 has been produced at three separate spatial resolutions: 7.5 arc sec (~250 m), 15 arc sec (~500 m), and 30 arc sec (~ 1 km) [14,15]. Also, GTOPO30 is produced by the staff at the USGS Center for Earth Resources Observation and Science with a horizontal cell size of 30 arc sec (~1 km) and was obtained from several rasters and vector sources of topographic information [16]. These freely DEMs differ in the data resolution and accuracy, according to the technology of data acquisition and the methodology of processing concerning a particular terrain and land cover type. The positional and attributive accuracy of these DEMs is often unknown and non-uniform within each dataset [17,18].

On the other hand, the height supplied by open-source data is related to the reference ellipsoid which is called ellipsoidal height (h). However, the height-related to an equipotential surface of the terrestrial gravity field is the most important for DEM generation. This height is called the orthometric height (H). The relation between these two heights is the geoid undulation (N), i.e., for obtaining orthometric height, based on the ellipsoidal height. It is necessary to understand the geoid undulation. The geoid undulation varies from one site to another. The Earth Gravitational Models (EGM96 and EGM2008) are some of the global geo-potential models that are used to have geoid undulation the orthometric height [19,20]. The conversion of ellipsoidal height (h) of these DEMs into orthometric height (H) is fundamental in most geoscience and engineering applications. This conversion requires the geoid undulation (N) related to World Geodetic System (WGS84) ellipsoid by using the simple mathematical relation $H = h - N$ [20,21].

In fact, no official precise local DEM for Egypt is published yet. So, using the free source data requires accordingly to examine the qualities and assesses the accuracy before attempting any data extraction from these freely DEM and the Global Geopotential Model (GGM) products. Rabah et al. [22] have checked the accuracy of three types of global DEMs (namely SRTM 1, SRTM 3, and ASTER) using precise GPS/leveling control points. The results showed that the most accurate one was the SRTM 1 that produced a mean height difference and standard deviations equal 2.89 and ±8.65 m respectively. El-Quilish et al.[23] have compared eight types of global DEMs in the Nile delta region including (EarthEnv-D90, SRTM 1, SRTM 3, ASTER, GMTED2010, and GTOPO30) using 416 GPS/leveling control points. They have introduced a statistical measure for DEMs accuracy evaluation called the reliability Index, which is based on the weighted average mean concept. The results showed that in the Nile delta region, EarthEnv-D90 and SRTM models have a high-reliability Index. Furthermore, Al-Krargy et al.[24] investigated three types of global DEMs (namely ASTER, SRTM 3 arc-second, and GTOPO30) and compared seven GGM including (EGM2008, and EGM96) over precise local gravity and GPS/Leveling data. The results showed that the SRTM 3 DEM produces a mean standard deviation of ±4.3 m, using 1227 observed orthometric heights control points in Egypt. As well, it has been indicated that the EGM2008 is the most precise global model, as it produces a mean standard deviation of geoid undulation differences which equals ±0.23 m using observed 1074 GPS/Leveling stations.

Most of the Egyptian coastal lands are becoming vulnerable to inundation as their elevation will likely be below the projected sea-level rise under the impact of global warming and climate change [25]. From the previous studies, most of these recent studies investigated the assessment analysis and comparisons between different DEMs, but did not put forwards methods of accuracy enhancement for specific applications, especially in Egypt. The accuracy of DEM data are critical regarding the results of coastal flood risk assessments, because topographic data
are the most important factor in determining the extent of flooding and therefore the accuracy of flood maps. Especially in low-lying coastal zones, the simulation of flood extents was found to be very sensitive to the terrain representation [26,27]. Therefore, this research aims to obtain the highest possible vertical accuracy of SRTM that achieve the requirements for coastal inundation risk studies with minimum volume of field measurements data. In this paper, an enhanced SRTM model is introduced by using the best fit global geoid model for Egypt instead of EGM96. The enhancement was completed by applying a scale factor formula that were derived mathematically by using actual field measurements of uniformly distributed ground control points along the study area. Then, the improved accuracy was evaluated at the two study areas of different types of terrain surfaces e.g., flat terrain and hilly one. This method approach was validated by using high precision local DEMs and the Sensitivity of these DEMs to sea level projections was statistically analysed.

2. Study Areas and Available Data

Egypt has coastlines about 1550 km on the Mediterranean Sea on the north side, including the River Nile in the middle. However, on the east side, Egypt's Red Sea coastline is about 1705 km [28]. The Nile Delta is the delta formed in Lower Egypt where the Nile River spreads out and drains into the Mediterranean Sea. From west to east, it covers around 240 km of coastline. Two case studies were selected to achieve the purpose of this research. The study area (1) of around 200 km long and 500 m wide was selected at the Nile Delta coast. This area has the advantage of being a flat terrain. The mainland used in this area is agriculture. Fig.1 represents the location of the Delta study area. The green-colored points represent the reference (GCPs) that were measured and processed using the GNSS instrument and Leveling. The Red-colored points along the whole study area are the grid of points measured using the GNSS post-processing kinematic (PPK) technique with an observation interval of 5 seconds. These points will be used for local DEM creation and validation purposes.

The second study area is along the Red Sea coast. It is about 194 km long and 500 m wide as well. This area was selected as the case study of hilly terrain and it is considered as a worldwide known touristic zone. The location of the study area (2) and the used land surveying data is shown in Fig.2.

Generally, the topographic and tidal Data are the two main types of data that were used in this research as follows:

- Leveling measurements of 50 GCPs on the Nile Delta coast and 46 GCPs on the Red Sea coast were used. They were distributed uniformly about 5 km spacing parallel to the coastlines and observed by using the precise level instrument. These high-precision 96 GCPs were obtained by using the international standards and specifications by the survey research institute in Egypt.

- Two local DEMs surfaces were generated with 30 m cell size using GIS application. They were created based on land surveying data which contains around 25778 Grid points distributed from Rashid at 30.38°E to Brullus at 30.96°E along the Nile delta coast (refer to Fig.1) and about 19887 Grid points with around 60 km between 26.2° N and 25.8° N along the red sea coast (refer to Fig.2). These points were collected and processed by GNSS instruments, Dual Frequency GNSS Receivers, using the kinematic GNSS technique and adjusted to the local vertical datum by the survey research institute in Egypt.

- On the other hand, a Free source of DEM data with global 1-arcsecond (30-m) SRTM elevation data for the study areas was downloaded from Earth Explorer (http://earthexplorer.usgs.gov) website of the United States Geological Survey. This data was processed and projected to the Universal Transverse Mercator (UTM) coordinate system of zone 36.

- The available nearest tide gauge stations for the two different study areas are at Alexandria and Safaga respectively. These stations were established by the Survey Research Institute (SRI) of the National Water Research Center (NWRC) coop with the Egyptian Navy Hydrographic Department (ENHD). Table 1 presents the locations and characteristics of the collected tide gauge datasets at these stations.

| Station | Period (2008-2018) | Location-coordinates |
|---------|-------------------|----------------------|
| Alex.   | pressure-type devices | 31°11′55.91″ N, 29°52′2.23″ E |
| Safaga  |                    | 26°44′56.93″ N, 33°57′04.58″ E |
3. Methodology

The methodology described to achieve the objective of this paper is based on three main phases:

The first phase: SRTM data enhancement

The Horizontal datum of SRTM data is World Geodetic System (WGS 84). However, the vertical datum is Earth Gravitational Model (EGM96). To assess the quality of the SRTM data, several procedures were conducted. The 96 used reference ground control points are considered to be uniformly distributed parallel to the coastline at both study areas. The corresponding values of these points from SRTM DEM were extracted under GIS Environment by using bilinear interpolation. Then, the enhancement introduced approach was applied using these GCPs based on two steps. In the first step, the Earth Gravitational Models EGM96 which is included in SRTM was replaced to EGM2008. EGM2008 was selected as it was proven to be the most precise global model for Egypt referring to the previously mentioned studies.

The geoid heights from EGM96 and EGM2008 were computed at each location using NIMA EGM96 calculator (ver1.0) and Altrans. EGM2008 calculator, respectively. These software applications compute the accurate geoid undulation values which are used in the conversion of SRTM orthometric to ellipsoidal heights and vice versa. In the second step, a mathematical scale factor formula for each study area using the GCPs and the extracted corresponding elevation points from SRTM is derived. At
both different terrains study areas, the simple scale factor formulas using a linear transformation method were used to achieve better agreement with the reference ground control points.

\[ Z_0 = a \cdot (Z_m + N_{\text{EGM96}} - N_{\text{EGM2008}}) + b \]  

(1)

Where \( Z_0 \) is the observed elevation at reference GCPs, \( Z_m \) is the Model’s elevation from SRTM, \( N_{\text{EGM96}} \) and \( N_{\text{EGM2008}} \) are the geoid undulation values from EGM96 and EGM2008 respectively. The values of \( a \) and \( b \) are the estimated transformation parameters.

The second phase: Enhancement evaluation and validation by statistical analysis

The values of the actual levels of the 96 used reference points were compared with the data extracted from SRTM DEMs and after applying the scaling formulas at the same exact control point locations. The evaluation of the vertical accuracy for these points from DEMs is performed by computing the vertical Root-Mean-Square-Error (RMSE) and mean error (ME). RMSE measures the difference between the values of the DEM elevations and the values of reference GCPs elevations. These individual point differences are also called residuals. ME lets us know whether the set of measurements higher or lower than the true values. RMSE and ME may be estimated by the following equations [29]:

\[
\text{ME} = \frac{1}{n} \sum_{i=1}^{n} (E_o - E_m) \\
\text{RMSE} = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (E_o - E_m)^2} 
\]

(2)

(3)

Where \( E_o \) is the observed elevation, \( E_m \) is the model’s elevation and \( n \) is the number of tested elevation points.

The validation of the accuracy enhancement approach for the free available SRTM DEM was done by using the same scaling mathematical formulas derived from the reference GCPs in phase 1. These simple scaling formulas were applied to the whole raster surface of the SRTM DEM at both study areas after the conversion to the global model of EGM2008. The raster of the gravity models of EGM2008 and EGM96 were downloaded from International Centre for Global Earth Models (ICGEM) to be used in the conversions process of SRTM [30]. The resulted scaled SRTM model surfaces have been compared with LDEMs at both study areas. The RMSE was computed by subtracting SRTM DEMs from the accurate LDEMs and producing new raster DEMs showing all the elevation differences (DOD). All these calculations related to the DEMs raster were done using the calculate statistics and raster math tool in GIS application.

The third phase: Sea Level rise Analysis

Sea level projections were based on analysing hourly, monthly, and annual tide records at Alexandria (city located on the Mediterranean Sea) and Safaga (city located on the Red Sea). Linear regression has been carried out based on this data to compute the rising rates of sea level at the two tide gauge sites. Then, the Sensitivity analysis of DEMs to sea level projections was performed using GIS application to determine the areas that are topographically lower than the elevated water levels. The inundation extent and water depth on each DEM dataset for different SLR scenarios were studied taking in consideration to check hydrologic connectivity.

4. Results and Discussion

The statistics of the variation of elevations data from SRTM DEMs with the corresponding actual values from GCPs have appeared in Table 2 for the two study regions.

As appeared in the above table, the height variations range from 0.90 m to 6.53m, with an average of 2.58 m and standard deviations equalling ± 1.19 m referring to the 50 GCPs used at the Nile Delta. Moreover, the height variations at the red sea area referring to the 46 GCPs used are in the order of 31m with an average of 6.69 m and standard deviations equalling ± 6.23m (see Table 2). It can be realized that compared with the SRTM, the DEM after enhancement and scaling (Scale SRTM) produces the smallest differences with a standard deviation equalling ± 0.18 m and ± 5.56m at the Nile delta coast and the red sea coast study areas. Also, the error statistics at different terrains shows RMS error of 1.16m and 2.76m for the introduced enhanced SRTM. The figures showed that the RMS error was significantly decreased about 60% and 45% after enhancement and scaling of SRTM at the Nile Delta and the red sea coast study areas respectively (shown in Fig. 3). Generally, the RMS error of SRTM’s were clearly lower in plain regions like the Nile Delta coast compared to hilly regions like the red sea coast study area.

| Table 2. Height variation in SRTM DEM’s and GCPs for study areas along Egyptian coasts |
|-----------------------------------------------|----------|----------------|----------------|----------------|----------------|----------------|
| Height | Nile Delta coast | Red sea coast | Nile Delta coast | Red sea coast | Nile Delta coast | Red sea coast |
|--------|--------------------|---------------|------------------|---------------|-----------------|----------------|
|        | GCPs | SRTM | scale-SRTM | GCPs | SRTM | scale-SRTM |
| Min    | 0.90  | -5.00 | 2.13   | 1.13  | 0.00  | -1.56  |
| Max    | 6.53  | 9.00  | 3.07   | 32.21 | 35.00 | 24.45  |
| Average| 2.58  | 1.78  | 2.58   | 6.69  | 10.54 | 6.69   |
| St.Dev.| 1.19  | 2.72  | 0.18   | 6.23  | 7.50  | 5.56   |
In this section, the introduced methodology was validated by comparing DEMs raster with GIS application. The local DEMs (LDEMs) at two different terrains are considered as a reference to assess the vertical accuracy of SRTM after the enhancement by the same derived scaling formulas from GCPs in phase 1. The two references local DEMs properties were described in section 2. Figures (4-7) highlighted the selected area of interest (AOI) and contained layout views in different scales for each DEM raster to show the surface of elevations for the (AOI) at each study area.

The local DEM in Fig. 4(a) for the Nile Delta study area
showed variations in elevations from +0.08 m to +5.45 m, while in Fig. 4(b) the variations in elevations were from -20 m to +20 m for the SRTM. Fig. 4(c) showed the enhanced elevation of SRTM from +1.12 m to +3.80 m by applying EGM2008 and the scale factor formula derived using the 50 GCPs in the Nile Delta coast.

An analysis and comparisons have been made to assess the enhancement of SRTM vertical accuracy by subtracting DEM raster in Fig. 4(b) and Fig. 4(c) from the reference local DEM in Fig. 4(a) using the raster math tool in GIS application and producing a new raster map which is a DEM showing all the elevations differences (DOD). The results referring to Fig. 5(a) showed that the differences in elevations between LDEM and the SRTM are varying from +20.61 m to -17.87 m above MSL, while the differences in elevations between LDEM and the scaled SRTM are varying from +2.99 m to -3.09 m as shown in Fig. 5(b). The RMSEs were computed by using elevations differences (DOD) in Figs. 5(a,b), the results show that RMSE for SRTM was 2.63 m. However, after the enhancement, it decreases to 1.35 m. The achieved enhancement of SRTM vertical accuracy was about 48% for study region 1 at the Nile delta coast.

Figure 5. DEM of elevations differences (DOD) (a) Between Local DEM & SRTM; (b) Between Local DEM & Scale SRTM DEM for Nile Delta study area

Figure 6. (a) Local DEM; (b) SRTM DEM; (c) Scale SRTM DEM for Red sea coast study area
The local DEM in Fig. 6 (a) for the Red sea coast study area showed variations in elevations from -0.04 m to +29.37 m, while in Fig. 6 (b) the variations in elevations were from -15 m to +32 m for the SRTM. Fig. 6(c) showed the enhanced elevation of SRTM from -3.1 m to +15 m by applying the same formula derived using 46 GCPs in the Red sea coast.

The same-mentioned analysis for the first study region has been used. The results referring to Fig. 7 (a) show that the differences in elevations between LDEM and the SRTM vary from +24.49 m to -25.54 m above MSL, while the differences in elevations between LDEM and the scaled SRTM vary from +25.50 m to -9.70 m as shown in Fig. 7 (b). As well, the RMSEs using elevations differences (DOD) in Fig. 7 (a,b) were 5.01 m for SRTM and decreased after the enhancement to 2.29 m. The SRTM vertical accuracy for the scaled SRTM was improved by about 54%.

The sensitivity of these DEMs to sea-level projections using the recent available local tide gauge data was analysed. Tide records spanned from 2008 to 2018 are used for measuring rates of present-day relative sea-level rise at both tide stations. The annual averages of sea level have been analysed taking with considering the zero-value for mean sea level (MSL) as a datum, it has been found that the MSL at Alexandria, during the period 2008-2018, varies from 41.96 cm to 45.89 cm above that datum, with a mean value of 44.76 cm. At Safaga, MSL changes between 47.99 cm to 53.15 cm, with an average of 51.33 cm. Fig. 8 presents the annual MSL values for both tide gauges. Furthermore, linear regression has been carried out based on this data to compute the rising rates of sea level at the two tide gauge sites. Table 3 presents the obtained trend formulas at both tide gauge stations with their corresponding coefficients of determination.
Table 3. Relative Sea Level Rise at Alexandria and Safaga tide gauges

| Tide Station | Sea Rise Trend (mm) | R² |
|--------------|---------------------|----|
| Alexandria   | MSL = 3.3955 YD + 427.18 | 0.75 |
| Safaga       | MSL = 5.188 YD + 482.22   | 0.89 |

MSL is the value of MSL height in millimeter, and YD is the year difference or number of years since 2008.

The analyses show that the two locations have positive sea rise trends with significant quality of fitting. These trends indicate that the sea level rising rate of the Mediterranean Sea at Alexandria is 3.4 mm/year and in Safaga at the Red Sea is 5.1 mm/year. Moreover, the data records collected from both stations showed that in the period 2008-2018, the Red Sea level at Safaga is higher than the Mediterranean Sea level at Alexandria. These estimates are close to results of similar previous studies [31,32]. The difference in annual mean sea level value between the Alexandria and Safaga is 7.25 cm in 2018.

However, it is worth mentioning that the data utilized in the research are the most recent available local tide records which are consistent with the growing climate change. Hence, it could be considered reliable for long-term MSL rise determination. SLR, planning scenarios for climate adaptation are mostly limited to a time horizon of 100 years. Consequently, the Inundation height at 2100 for the two different terrains using the formulas in Table 3 will be 28 and 43 cm above the MSL at Alexandria and Safaga, respectively. Most of the low-lying delta coastlines are subjected to natural subsidence including the Nile delta. However, the red sea coast does not subside, since the subsurface sediments that are made of rocky layers are incompressible. Land subsidence can be estimated using different techniques including the differential interferometric synthetic aperture radar (InSAR) [33]. Based on several recent studies, the average of subsidence rates was calculated as 4 mm/y for the study area region 1 [34,35]. Consequently, the Inundation height values will be approximated to 50 cm above MSL at the two tide gauge stations by 2100. There are other certain aspects based on local and global factors that affect coastal flooding and SLR. So, the sensitivity of SRTM DEMs will be examined under another SLR scenario using 1 m inundation height. For each scenario specified, the total area of inundation in (km²) was calculated with each DEM dataset and the results are given in Table 4. Regarding LDEMS at the flat terrain type of the Nile delta, the percentage of the area inundated at 0.5 m SLR was increased from 1.5% to 34.6% at 1 m SLR. The inundated areas from SRTM DEM were significantly higher than the local DEM for both scenarios. Also, the scale SRTM didn’t show values below 1 m as detailed and described in Fig.(4,5). The results were so different at the hilly terrain type of the red sea coast. As the percentage of the area inundated for LDEMS were increased from 0.25 % to 1.24% in both scenarios. Also, the inundated area of the scaled SRTM was enhanced to be 25% only higher than LDEM in comparison with the SRTM inundated area that was 61% higher for the 1 m inundation height.

5. Conclusions

The present study was an attempt to obtain a practical framework for enhancing the vertical accuracy of freely available DEM such as SRTM model at two different terrains along the Egyptian coasts. The proposed methodology was using uniformly distributed GCPs to extract mathematical formulas that can be applied for the enhancement of the SRTM raster DEMs. The RMSE of SRTM data extracted at the reference GCPs concerning plain and hilly regions were 2.88 m and 5.07 m respectively. These results match the specified accuracy of SRTM, which is about ±16 m at a 95% confidence level [36]. The evaluation of the enhancement results reached 60% and 45% at the two terrains of the Nile delta and red sea coasts. The validation process by applying the derived scaled formulas on 60 km strip of SRTM raster DEMs ensured that RMSEs were reduced by 48% to 54% in comparison with local accurate DEMs at the flat and hilly regions. These results are very well-fitting the obtained ones from the used reference ground control points. With respect to the sensitivity of these DEMs to sea-level projections using the recent local tide gauge data, the scaled SRTM DEM at the red sea region can be reliable as low-cost data with medium accuracy fulfil the requirement of monitoring inundation height response to sea-level rise. However, the accuracy achieved for the enhanced SRTM DEM didn’t fit the requirement for coastal inundation studies at the plain area of the Nile delta.

Table 4. Calculated inundated areas in (km²) due to sea level rise

| Inundation height (m) | Nile Delta coast Area(1) =26.72 km² | Red sea coast Area(2)=24.26 km² |
|-----------------------|---------------------------------------|-------------------------------|
|                       | Local DEM | Scale DEM | SRTM DEM | Local DEM | Scale DEM | SRTM DEM |
| 0.5                   | 0.396     | 0.000     | 10.719   | 0.061     | 0.445     | 0.744    |
| 1                     | 9.247     | 0.000     | 17.977   | 0.310     | 0.468     | 0.789    |
Therefore, it is concluded that this method approach should be investigated under different mathematical models and parameters to determine the optimal size and distribution of reference ground control points needed to adjust free open DEMs depending on the vertical accuracy demand.

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