Deep Soil Layers of Drought-Exposed Forests Harbor Poorly Known Bacterial and Fungal Communities

Beat Frey*, Lorenz Walthert, Carla Perez-Mon, Beat Stierli, Roger Köchli, Alexander Dharmarajah and Ivano Brunner

Forest Soils and Biogeochemistry, Swiss Federal Institute for Forest, Snow and Landscape Research WSL, Birmensdorf, Switzerland

Soil microorganisms such as bacteria and fungi play important roles in the biogeochemical cycling of soil nutrients, because they act as decomposers or are mutualistic or antagonistic symbionts, thereby influencing plant growth and health. In the present study, we investigated the vertical distribution of the soil microbiome to a depth of 2 m in Swiss drought-exposed forests of European beech and oaks on calcareous bedrock. We aimed to disentangle the effects of soil depth, tree (beech, oak), and substrate (soil, roots) on microbial abundance, diversity, and community structure. With increasing soil depth, organic carbon, nitrogen, and clay content decreased significantly. Similarly, fine root biomass, microbial biomass (DNA content, fungal abundance), and microbial alpha-diversity decreased and were consequently significantly related to these physicochemical parameters. In contrast, bacterial abundance tended to increase with soil depth, and the bacteria to fungi ratio increased significantly with greater depth. Tree species was only significantly related to the fungal Shannon index but not to the bacterial Shannon index. Microbial community analyses revealed that bacterial and fungal communities varied significantly across the soil layers, more strongly for bacteria than for fungi. Both communities were also significantly affected by tree species and substrate. In deep soil layers, poorly known bacterial taxa from Nitrospirae, Chloroflexi, Rokubacteria, Gemmatimonadetes, Firmicutes and GAL 15 were overrepresented. Furthermore, archaeal phyla such as Thaumarchaeota and Euryarchaeota were more abundant in subsoils than topsoils. Fungal taxa that were predominantly found in deep soil layers belong to the ectomycorrhizal Boletus luridus and Hydnum vesterholtii. Both taxa are reported for the first time in such deep soil layers. Saprotrophic fungal taxa predominantly recorded in deep soil layers were unknown species of Xylaria. Finally, our results show that the microbial community structure found in fine roots was well represented in the bulk soil. Overall, we recorded poorly known bacterial and archaeal phyla, as well as ectomycorrhizal fungi that were not previously known to colonize deep soil layers. Our study contributes to an integrated perspective on the vertical distribution of the soil microbiome at a fine spatial scale in drought-exposed forests.

Keywords: bacterial and fungal communities, deep forest soil layers, drought, Fagus sylvatica, FUNGuild, Illumina MiSeq sequencing, Quercus spp.
INTRODUCTION

The significance of soil microorganisms is receiving increasing attention, as they are relevant drivers of ecosystem functions (Bardgett and van der Putten, 2014). Soil microorganisms such as bacteria and fungi play important roles in the biogeochemical cycling of soil nutrients because they act as decomposers or are mutualistic or antagonistic symbionts, thereby influencing the growth and health of other organisms (Rousk and Bengtsson, 2014). In addition, they emit greenhouse gases and, thus, affect the climate globally (Oertel et al., 2016; Jansson and Hofmockel, 2020). However, most studies have been focused primarily on the top 20 cm of soil, where microbial biomass, activity and diversity are assumed to be the greatest. In contrast, little is known about the microbiome in the soil subsurface (Jiao et al., 2018), although trees in forest ecosystems have roots down to 2 m or more (Schenk and Jackson, 2002).

In general, living conditions for soil organisms and roots are harsher in deeper soil layers, where the soil density is higher, oxygen concentrations are lower, and carbon and nutrients are less available (Lennon, 2020). Carbon and nutrients typically are more available in the upper soil, mainly due to larger quantities of leaf and fine root litter and root exudates, as well as more biotic activity (Eldridge et al., 2016; Liu et al., 2020). Consequently, microbial biomass is usually one to two orders of magnitude lower in the subsoil than in the topsoil (Eilers et al., 2012; Spohn et al., 2016). Nevertheless, the cumulative biomass of bacteria and fungi inhabiting deeper soil horizons can be as high as in the topsoil, and these microorganisms play similarly important roles in biogeochemical processes, including soil carbon and nitrogen cycling, soil formation, iron redox reactions, and pollutant degradation (Brewer et al., 2019). Some groups of microorganisms appear predominantly in deeper soil horizons, such as the candidate phylum Dormibacteraeota (Brewer et al., 2019), or tree root host-specific mycorrhizal fungal strains, as shown in Eucalyptus forests at 4 m soil depth (Robin et al., 2019). Also, it appears that microbial diversity is not necessarily lower in subsoils than in topsoils, as recently has been found on the Loess Plateau in China in soils down to 20 m depth (Liu et al., 2020).

As climate change is progressing, it is assumed that shifts in vegetation will occur throughout forest biomes (McDowell et al., 2020). In Europe, European beech (Fagus sylvatica L.) forest will move toward cooler sites, most likely toward the north, east, or higher elevations, and oak trees (Quercus spp.) will take over because they are resistant to higher temperatures and lower water availability (Hanewinkel et al., 2013). The soil microbial communities might also change in the long term, not only because climatic conditions have changed, but also because of the shift in dominant tree species. Many oak forests are less dense and shady than beech forests (often little or no understory vegetation, reminding of Gothic Cathedrals – “Buchen-Hallenwald”; Ellenberg, 1986), thus permitting more light and rainfall to reach the forest floor. In addition, the litter quality (i.e., roots, leaves, wood) of oaks might be different than that of beech, affecting soil properties in a different direction, e.g., the quality of the soil organic matter, due to a higher content of tannins and of phenolics in the bark of oaks (Torkaman and Seyam, 2009). Lastly, oak trees mostly host different ectomycorrhizal fungal species than beech trees (van der Linde et al., 2018). Overall, it is assumed that variability in soil moisture, soil temperature, and local conditions such as dominant tree species might affect the phenotypic response of the soil microbiome, with feedbacks to climate change (Jansson and Hofmockel, 2020).

Switzerland hosts some regions where European beech faces its physiological limits due to dry climatic and extreme edaphic conditions such as low soil water holding capacity (Walther et al., 2021). Many of these dry sites are located in the Rhone Valley and the Jura Mountains. Here, European beech forests adjoin forests dominated by more heat- and drought-tolerant pubescent oak or sessile oak, and if drought and heat periods increase due to ongoing climate change, beech will be slowly displaced by oak species. European beech is known to have a suboceanic preference (Meier and Leuschner, 2008), whereas oaks have more of a submediterranean preference (Himrane et al., 2004).

The aim of this study was to investigate the vertical distribution of the soil microbiome to a depth of 2 m in drought-exposed forest ecosystems in Switzerland. We further aimed to disentangle the influence of tree species from soil depth, since the two tree genera, beech and oak, usually host different ectomycorrhizal species, and to verify the general shift from saprotrophic to ectomycorrhizal fungal dominance with increasing soil depth recently observed by Carteron et al. (2020). Finally, we aimed to investigate how the microbial community in bulk soils is related to the fine root microbial community throughout the soil profile, as pioneering studies have shown that root-associated fungi are recruited from the soil fungal community (Goldmann et al., 2016). We hypothesized that:

1. bacterial and fungal abundances decrease with soil depth, with fungal abundance decreasing more strongly than bacterial abundance;
2. microbial diversity and community structure vary strongly with increasing soil depth than ectomycorrhizal abundance;
3. different ectomycorrhizal fungal species than beech trees (van der Linde et al., 2018).
4. relative abundances of saprotrophic fungi decrease more strongly with increasing soil depth than ectomycorrhizal fungi.

Our study contributes to an integrated perspective on the vertical distribution of the soil microbiome in drought-exposed forests at a fine spatial scale.

MATERIALS AND METHODS

Forest Sites

The investigated forest sites are located in the relatively warm and dry Rhone Valley (Chamoson and Saillon) and in the Jura Mountains (Neunkirch) of Switzerland with mean annual precipitation sums (MAP, 1981–2010) of 815–955 mm and mean annual temperatures (MAT, 1981–2010) of 8.5–9.4°C depending on the site (Supplementary Figure 1 and Table 1; Remund et al., 2014). Depending on water availability, the sites are covered mainly by European beech or oak species, with oak sites having slightly warmer and drier soils than European beech sites (Table 2). All forest stands have a near natural tree species composition originating from natural regeneration...
and are unmanaged since several decades. The stands with its admixtures of other tree species had an approximate area of 50–200 m². At the beech sites, there were no admixed tree species, except at Saillon where a few scattered *Abies alba* Mill. tree individuals were recorded. At the oak sites of Chamson and Saillon, a few scattered tree individuals of *Acer opalus* Mill. were detected next to the dominant oak *Quercus pubescens* (Wild.), and at Neunkirch, a few tree individuals of *F. sylvatica* were detected next to the dominant oak *Q. petraea* (Matt.) Liebl. The forest sites are part of the research platforms WSL Forest Soil Database and TreeNet¹, and the forest site Neunkirch is a monitoring plot of the Long-term Forest Ecosystem Research Program LWF (ICP-Forests²). The soil classification followed the system of the Food and Agriculture Organization of the United Nations (IUSS Working Group WRB, 2007) (Table 1).

### Soil and Root Sampling

Soil profiles down to 2 m depth were excavated in the center of a group of dominant trees at each study site in spring 2014. Three soil samples per soil depth were taken with 1 L steel cylinders (10.7 cm height, 11.1 cm diameter) at depths of 0–10 cm (L1), 15–25 cm (L2), 45–55 cm (L3), 75–85 cm (L4), 110–125 cm (L5), 140–155 cm (L6), and 180–200 cm (L7). However, for depths >45 cm at Neunkirch beech and oak sites and for depths >140 cm at the Chamson oak site, the cylinders could not be used for sampling due to the very high content of large stones. Instead, bulk soil samples were excavated and the entire volume was determined by refilling the cavity with quartz sand. For soil chemical and physical analyses, approximately 1 kg of soil was sampled separately from the various depths (L1–L7). All soil samples were then transported in plastic bags to the labs and stored at 4°C until further processing.

In order to estimate fine root biomass and fine-earth densities, the soil samples of the 1 L steel cylinders were sieved with a 4 mm sieve and the roots were sorted out. The roots were then washed, separated according to their diameters (<2 mm: fine roots, 2–5 mm, >5–10 mm, >10 mm) and according to their species affiliation (beech, oak). European beech roots are recognized as being brown reddish with a rigid appearance, and oak roots as being yellowish with a rough appearance. Afterward, the soil was further sieved through a 2 mm mesh sieve.

For soil and root microbial DNA-sampling, aliquots of soils of about 5 g were put into a Ziploc® bag and frozen and stored at −80°C for DNA analyses. Roots were frozen in liquid nitrogen (N₂), freeze-dried, weighed, and stored at −20°C for DNA analyses. In deep soil layers, where steel cylinder samples could not be taken, three soil and root aliquots were sampled from the bulk soil, from which DNA samples were taken independently. Finally, the remaining soil was dried at 105°C and volumetric stone content and fine-earth density were estimated.

### Soil Chemical and Physical Analyses

The soil samples for chemical and particle size analyses were dried a 60°C and sieved through a 2 mm mesh sieve. Soil pH was determined in 0.01 M CaCl₂. For carbon (C) and nitrogen (N) analyses, an aliquot of the dried and sieved samples was ground for 3 min using a vibrating ball mill (MM2000, Retsch GmbH, Germany) with zirconum-grinding tools.

Exchangeable cations (Na, K, Mg, Ca, Mn, Al, Fe) were extracted (in triplicate) in an unbuffered solution of 1 M NH₄Cl. The element concentrations in the extracts were determined by ICP-OES (Optima 3000, Perkin–Elmer, United States). In samples with a pH (CaCl₂) < 6.5, the exchangeable proton content was calculated as the difference between the total and the Al-induced exchangeable acidity, as determined by the KCl method (Thomas, 1982). In samples with a higher pH, the contents of exchangeable protons were assumed to be negligible. The effective cation-exchange capacity (CEC) was obtained by summing the charge equivalents of exchangeable Na, K, Mg, Ca, Mn, Al, Fe, and H. The base saturation (BS) was calculated as the ratio of the sum of exchangeable Na, K, Mg, and Ca to the CEC. Organic C (Corg) and total N (Ntot) were determined with a CN elemental analyzer (NC2500; CE Instruments, Italy). The organic C (Corg) content of samples with a pH (CaCl₂) < 6.0 was assumed to be equal to the total carbon (Ctot) content, whereas samples with a pH > 6.0 were assumed to potentially contain carbonates and, thus, were fumigated with HCl vapor prior to the CN analysis (Walthert et al., 2010). Soil particles were fractionated into sand, silt and clay using the pipette method according to Gee and Bauder (1986). Plant-available water storage capacity (AWC), i.e., the storable amount of water between field capacity and permanent wilting point, was calculated with a pedotransfer function according to Teepe et al. (2003) for each sampled soil layer for a thickness of 10 cm considering fine-earth density and soil texture, i.e., particle size distribution.

### Measurement of Soil Temperature and Soil Water Potential

MPS-2 sensors and EM-50 data loggers (Decagon Devices, United States) for continuously recording soil temperature and soil water potential (Ψsoil) were installed at each study site in spring 2014. Two to three MPS-2 sensors per soil depth were embedded in the front wall of the soil pit at 20, 80, and 140 cm depth (at 180–200 cm depth only in Chamson and Saillon). Measurements of Ψsoil were temperature corrected to 22°C according to Walthert and Schleppi (2018). Soil temperature and soil water potential (Ψsoil) have been recorded from spring 2014 onward in hourly intervals. In order to characterize the forest sites, daily data from year 2015 were calculated and divided into four annual quarters (January–March, April–June, July–September, October–December).

### Soil and Root DNA Extraction, MiSeq Sequencing, and Bioinformatic Analyses

For DNA extraction, either frozen soil material or lyophilized fine root material, which was chopped with sterilized scissors and afterward grounded in 2 ml Eppendorf tubes containing eight sterile steel balls (2 mm diameter) on a FastPrep-24 bead beating grinder (MP Biomedicals, United States) three times for 30 s, was used. DNA was extracted using the PowerSoil DNA Isolation Kit (Qiagen, Germany). For the bead beating procedure, either 0.5 g soil or 0.02–0.1 g root material was used according

¹www.treenet.info
²www.icp-forests.net
to Herzog et al. (2019). DNA was quantified using the high-sensitivity Qubit assay (Thermo Fisher Scientific, Switzerland). The V3–V4 region of the bacterial small-subunit (16S) rRNA gene and the internal transcribed spacer region 2 (ITS2) of the eukaryotic (fungal groups, some groups of protists, and green algae) ribosomal operon were PCR amplified using primers previously described by Frey et al. (2016). PCR amplification was performed with 40 ng soil DNA and the GoTaq® G2 HotStart Polymerase (Promega, Switzerland) in a final volume of 50 µL per samples (16S: 5 min at 95°C/36 cycles: 40 s at 94°C, 40 s at 58°C, 1 min at 72°C/10 min at 72°C; ITS2: 5 min at 95°C/38 cycles: 40 s at 94°C, 40 s at 58°C, 1 min at 72°C/10 min at 72°C). A mixture of peptide nucleotide acid (PNA) blocker oligos (PNA Bio Inc., United States) targeted at plant mitochondrial and plastidic genomes was added in the PCR reaction mixture of root DNA to increase the fraction of bacterial sequences and reduce the PCR-bias, resulting in more accurate sequencing (Lundberg et al., 2013). Thus, the PCR mixture contained 1.25 mPNA blocker (0.25 µM) and 1.25 pPNA blocker (0.25 µM). A PNA clamping at 75°C for 10 s after denaturation at 94°C was included in the PCR conditions. Both positive and negative PCR controls were included to exclude false positives and negatives. Bacterial and fungal amplicons were sent to the Génome Québec Innovation Centre at McGill University (Montreal, QC, Canada) for barcoding using the Fluidigm Access Array technology and paired-end sequencing on the Illumina MiSeq v3 platform (Illumina Inc., United States). Raw sequences have been deposited in the NCBI Sequence Read Archive under the BioProject accession number PRJNA704912.

Quality filtering, clustering into operational taxonomic units (OTUs), and taxonomic assignment were performed as previously described in Frey et al. (2016). In brief, a customized pipeline largely based on UPARSE (Edgar, 2013; Edgar and Flyvbjerg, 2015) and implemented in USEARCH (v9;2; Edgar, 2010) was used. Filtered reads were de-replicated and singleton reads were removed prior to clustering. Sequences were clustered into OTUs at 97% sequence identity (Edgar, 2013). For taxonomic classification of the OTUs, corresponding centroid sequences were queried against selected reference databases using the naïve Bayesian classifier (Wang et al., 2007) and a minimum bootstrap support of 80%. Prokaryotic sequences were queried against the SILVA database (v132; Quast et al., 2013). Eukaryotic ITS2 sequences were first queried against a custom-made ITS2 reference database retrieved from NCBI GenBank, and sequences assigned to fungi were subsequently queried against the fungal ITS database UNITE (v8.0; Abarenkov et al., 2010). Prokaryotic centroid sequences identified as originating from organelles (chloroplast, mitochondria), as well as eukaryotic centroid sequences identified as originating from soil animals (Metazoa), plants (Viridiplantae, except green algae), or of unknown eukaryotic origin, were removed prior to data analysis.

### Quantitative PCR of Bacterial and Fungal Ribosomal Markers

Relative abundances of the bacterial 16S rRNA genes and the fungal ITS genes of soils from 0–10 cm, 45–55 cm and 140–155 cm depth were determined by quantitative PCR using an ABI7500 Fast Real-Time PCR system (Applied Biosystems, United States) and using 27f-512r (16S V1–V3) and ITS2 primers (without barcodes) and cycling conditions as for the sequencing approach (see also Frey et al., 2020). For qPCR analyses, 20 ng DNA in a total volume of 6.6 µL and 8.4 µL GoTaq® qPCR Master Mix (Promega, Switzerland) containing 1.8 µM of each primer and 0.2 mg mL⁻¹ of BSA were used. PNA blockers were also added in the reaction mix of the root DNA as described above. The specificity of the amplification products was confirmed by melting-curve analysis. Three standard curves per target region (correlations ≥ 0.997) were obtained using tenfold serial dilutions (10⁻¹ to 10⁻⁹ copies) of plasmids generated from cloned targets (Frey et al., 2011). Data were converted to represent the average copy number of targets per µg DNA and per g soil.

### Data Analysis

All statistical analyses were performed using R (v4.0.2; R Core Team, 2020). We used analysis of variance (ANOVA) to test for effects of soil depth (L1: 0–10 cm, L2: 15–25 cm, L3: 45–55 cm, L4: 75–85 cm, L5: 110–125 cm, L6: 140–155 cm and L7: 180–200 cm)

---

**TABLE 1 | Forest site, tree and soil characteristics.**

| Lat. | Long. | Elev. m a.s.l. | Exp. | Slope % | MAT °C | MAP mm | Age Year | Height m | DBH cm | Soil type |
|------|-------|---------------|------|---------|-------|--------|----------|----------|--------|-----------|
| **Beech sites:** | | | | | | | | | | |
| Chamoson | 46°12' | 07°12' | 880 | NE | 65 | 8.5 | 870 | 74 | 29 | 39 | Eutric Cambisol |
| Neunkirch | 47°41' | 08°32' | 560 | N | 58 | 8.8 | 953 | 166 | 24 | 66 | Rendzic Leptosol |
| Salton | 46°10' | 07°09' | 890 | SE | 55 | 9.4 | 829 | 100 | 24 | 38 | Calcaric Regosol |
| **Oak sites:** | | | | | | | | | | |
| Chamoson | 46°12' | 07°12' | 870 | S | 90 | 8.6 | 815 | 66 | 9 | 19 | Calcaric Regosol |
| Neunkirch | 47°40' | 08°32' | 640 | S | 35 | 9.1 | 949 | 134 | 17 | 38 | Rendzic Leptosol |
| Salton | 46°10' | 07°09' | 790 | SE | 50 | 9.4 | 828 | 97 | 18 | 28 | Calcaric Regosol |

Lat., latitude; Long., longitude; Elev., elevation; Exp., exposition; MAT, mean annual temperature; MAP, mean annual precipitation sum; DBH, stem diameter at breast height (mean of the three sampled trees); FAO class., classification according to IUSS Working Group WRB (2007). World Reference Base for Soil Resources 2006, first update 2007. World Soil Resources Reports No. 103, FAO, Rome.
TABLE 2 | Mean values of soil temperature (°C) and soil water potential (kPa) in 2015 (L2, L4, L6: n = 3 forest sites; L7: n = 2 forest sites).

| Soil layers | Beech sites | Oak sites | p-ANOVA† |
|-------------|-------------|-----------|----------|
|              | L2 | L4 | L6 | L7 | L2 | L4 | L6 | L7 | Tree | Depth | Interact. |
| Soil temperature: | | | | | | | | | | | | |
| January–March | 3.8 | 4.9 | 5.9 | 7.2 | 4.4 | 5.4 | 6.7 | 8.2 | 0.018 | <0.001 | 0.93 |
| April–June | 10.7 | 9.6 | 8.8 | 8.6 | 12.5 | 11.7 | 10.7 | 10.6 | <0.001 | 0.002 | 0.99 |
| July–September | 15.5 | 14.5 | 13.3 | 12.5 | 17.0 | 16.3 | 14.8 | 14.1 | 0.001 | <0.001 | 0.99 |
| October–December | 8.4 | 9.6 | 10.2 | 10.9 | 9.1 | 10.3 | 11.1 | 12.0 | 0.011 | <0.001 | 0.94 |
| Year | 9.6 | 9.7 | 9.5 | 9.8 | 10.8 | 10.9 | 10.8 | 11.2 | 0.11 | 0.99 | 1.00 |
| Soil water potential: | | | | | | | | | | | | |
| January–March | −10 | −57 | −288 | −431 | −45 | −387 | −572 | −833 | 0.10 | 0.07 | 0.83 |
| April–June | −30 | −21 | −77 | −203 | −60 | −93 | −231 | −640 | 0.008 | 0.002 | 0.13 |
| July–September | −614 | −497 | −509 | −418 | −814 | −778 | −689 | −850 | 0.07 | 0.94 | 0.93 |
| October–December | −322 | −579 | −623 | −612 | −531 | −874 | −831 | −984 | 0.09 | 0.31 | 0.98 |
| Year | −244 | −289 | −374 | −416 | −362 | −533 | −581 | −827 | 0.002 | 0.05 | 0.54 |

† Effects of tree species, soil depth and their interaction were assessed by analysis of variance (ANOVA); significant values (p < 0.05) are in bold.

‡ Soil layers: L2: 15–25 cm, L4: 75–85 cm, L6: 140–155 cm, L7: 180–200 cm.

Soil temperature: for different years (April–June, September–November, December–March, and January–April) and for different tree species (beech and oak).

Soil water potential: for different months (January–March, April–June, July–September, October–December, and November–January).

The differences in soil temperature and water potential between the different years and tree species were significant (p < 0.05).

Soil layers: L2, L4, L6, L7.

To evaluate the differences in microbial β-diversity, Bray–Curtis dissimilarities were calculated based on square root transformed relative abundances of OTUs. The differences were compared using the R package phyloseq (v1.32; McMurdie and Holmes, 2013).

The differences in microbial β-diversity between soils were visualized with canonical analysis of principal coordinates (CAP) ordinations for the tested factors of soil depth, tree species and substrate. Microbial β-diversities were also visualized using unconstrained non-metric multidimensional scaling (NMDS) ordinations.

The differences in microbial β-diversity, soil environmental variables were projected as vectors onto the NMDS ordinations using the envfit function in the R package ‘vegan.’ The direction of the vectors reflects the gradient of variation of the environmental variables, from the smallest (origin) to the largest (arrowhead) values. Permutation tests (10⁵ permutations) implemented with the envfit function were used to evaluate the statistical significance of the correlation (R²) between the projected variables and the NMDS axes.

To identify bacterial and fungal phyla and genera that were differentially distributed between topsoil (L1 and L2; 0–25 cm) and subsoil (L5 and L6; 110–155 cm) we performed DESeq2 pairwise analyses (DESeq2 package version 1.30.0; Love et al., 2014, 2020). The DESeq2 comparisons were performed at the OTU level, and the differentially abundant OTUs identified in the analyses were grouped a posteriori into the taxa displayed in the graphical representations. To avoid the influence of rarely occurring (and potentially erroneous) OTUs in the analyses, the DESeq2 comparisons only included OTUs for which the sum of the sequences over all soil samples belonging to the same tree species at each forest site was ≥ 10, for at least two independent combinations of forest site and tree species. For each DESeq2 comparison of topsoil vs. subsoil, the log₂ fold change (LFC) of each OTU was calculated as the log₂ (genera abundance of topsoil/genera abundance of subsoil). OTU abundances were normalized with the median of ratios method to account for differences in sequencing depth between samples. The statistical significance of the LFC was evaluated with Wald tests (Rüthi et al., 2020). P values were adjusted for multiple testing using the Benjamini–Hochberg method with a false discovery rate threshold of 5%.

The topsoil vs. subsoil DESeq2 comparisons were performed independently for (1) the subset of fine roots and (2) the subset of bulk soils, collected across all beech and oak sites at the distinct soil depths. Soil layer (topsoil or subsoil) was implemented as the main factor in the DESeq2 tests, controlling for the additional factors of site and tree species (~ site + tree species + soil layer). DESeq2 analyses were also applied to identify differentially abundant taxa between beech and oak trees (~ site + tree species), and soils and fine roots (~ site + tree species + substrate). In these later cases, the DESeq2 tests included all fine roots and soil samples.
TABLE 3 | Mean values of soil biological, chemical and physical properties (L1–L6: n = 3 forest sites; L7: n = 2 beech sites, n = 1 oak site).

| Soil layers* | Beech sites | Oak sites | p-ANOVA+ |
|--------------|-------------|-----------|----------|
|              | L1 | L2 | L3 | L4 | L5 | L6 | L7 | Tree | Depth | Interaction |
| **Soil biological properties:** | | | | | | | | | | |
| Microbial biomass (µg DNA g⁻¹ soil)* | 18.8 | 14.5 | 9.66 | 2.40 | 2.06 | 1.40 | 0.07 | 11.3 | 9.43 | 3.01 | 1.86 | 1.49 | 1.40 | 2.32 | 0.050 | <0.001 | 0.37 |
| Fine root biomass (g dm⁻³ soil) | 3.29 | 1.78 | 0.69 | 0.37 | 0.21 | 0.37 | 0.37 | 3.90 | 2.65 | 1.03 | 0.50 | 0.20 | 0.41 | 0.44 | 0.017 | <0.001 | 0.27 |
| **Soil chemical properties:** | | | | | | | | | | |
| pH (CaCl₂) | 6.56 | 6.97 | 7.28 | 7.81 | 7.71 | 7.60 | 7.75 | 6.32 | 7.32 | 7.66 | 7.68 | 7.70 | 7.71 | 7.75 | 0.71 | 0.002 | 0.93 |
| Base saturation (%) | 98.6 | 99.3 | 99.5 | 99.9 | 100 | 100 | 100 | 99.5 | 99.9 | 99.9 | 99.9 | 99.9 | 99.9 | 99.9 | 0.33 | 0.26 | 0.87 |
| Corg (%) | 6.89 | 3.42 | 1.46 | 0.91 | 0.81 | 0.68 | 0.56 | 6.48 | 2.29 | 1.05 | 0.95 | 0.48 | 0.48 | 0.68 | 0.55 | <0.001 | 1.00 |
| Ntot (%) | 0.45 | 0.27 | 0.13 | 0.07 | 0.09 | 0.08 | 0.09 | 0.39 | 0.20 | 0.12 | 0.11 | 0.08 | 0.08 | 0.05 | 0.56 | 0.002 | 0.99 |
| C/N ratio | 15.6 | 11.9 | 10.6 | 14.3 | 14.2 | 13.1 | 11.6 | 16.4 | 11.9 | 8.8 | 9.9 | 8.0 | 8.4 | 14.5 | 0.42 | 0.80 | 0.95 |
| **Soil physical properties:** | | | | | | | | | | |
| Fine-earth density (g cm⁻³) | 0.64 | 0.81 | 0.98 | 1.06 | 1.20 | 1.04 | 1.22 | 0.73 | 0.88 | 0.99 | 1.01 | 1.10 | 1.12 | 1.21 | 0.75 | <0.001 | 0.79 |
| Stones (%) | 15.8 | 15.8 | 28.3 | 54.2 | 43.3 | 33.8 | 46.3 | 9.7 | 24.2 | 39.2 | 47.5 | 47.5 | 47.5 | 5.0 | 0.81 | 0.32 | 0.92 |
| Sand (%) | 17.6 | 16.7 | 18.5 | 31.4 | 32.3 | 32.3 | 45.6 | 45.4 | 22.6 | 36.9 | 34.3 | 40.6 | 42.8 | 44.0 | 36.8 | 0.09 | 0.029 | 0.63 |
| Silt (%) | 45.3 | 48.0 | 53.1 | 47.4 | 49.8 | 43.4 | 47.1 | 41.7 | 35.7 | 45.1 | 44.4 | 44.1 | 43.4 | 57.6 | 0.58 | 0.97 | 0.98 |
| Clay (%) | 37.2 | 35.3 | 28.4 | 21.3 | 18.0 | 14.0 | 7.7 | 35.8 | 27.5 | 20.6 | 15.0 | 13.2 | 12.7 | 5.7 | 0.34 | 0.032 | 1.00 |
| AWC3 (mm)‡ | 21.6 | 23.1 | 17.3 | 11.2 | 13.3 | 15.5 | 12.9 | 27.0 | 40.0 | 15.1 | 12.6 | 11.9 | 11.8 | 23.2 | 0.40 | 0.14 | 0.83 |

*Effects of tree genus, soil depth and their interaction were assessed by analysis of variance (ANOVA); significant values (p < 0.05) are in bold.

†Soil layers: L1: 0–10 cm, L2: 15–25 cm, L3: 45–55 cm, L4: 75–85 cm, L5: 110–125 cm, L6: 140–155 cm, L7: 180–200 cm.

‡Plant-available water storage capacity (AWC) per 10 cm soil thickness.

Fungal functional guilds were assigned within the six most abundant guilds, namely ectomycorrhizal fungi, arbuscular mycorrhizal fungi, endophyte, undefined saprotrophs, animal pathogens, and plant pathogens, using an open annotation tool (FUNGuild) according to Nguyen et al. (2016). Only guild assignments with ‘highly probable’ confidence rankings were accepted (Perez-Mon et al., 2020).

RESULTS

Forest Site Characteristics

Analysis of soil chemical and physical parameters at the six forest sites showed that they were not dependent on the tree genus, as none of the variables revealed significances. In contrast, some of the soil parameters showed significant changes with depth (Table 3). Soil pH showed a significant increase with soil depth, whereas soil base saturation remained at a high level around 100%. Both, Corg and Ntot decreased significantly with soil depth by about tenfold, as did clay content by about fivefold. In contrast, fine-earth density, sand content, and stone content increased with soil depth, whereas C:N ratio, silt content, and the plant-available water storage capacity (AWC) remained more or less constant (Table 3).

Analysis of fine root variables showed, that oak sites had a slightly but significant greater fine root biomass compared with beech sites, and mainly at depths down to 1 m (Table 3). However, a strong significant decrease of the fine root biomass with increasing soil depth occurred down to 2 m for both tree species, with the strongest decrease within the first 50 cm, from around 3.5 to 0.6 g dm⁻³ soil.

At the oak sites, soil temperatures of each of the yearly quarters were significantly higher than at the beech sites, by 1–2°C, whereas the yearly means within each of the soil layers were not significantly different (Table 2). The soil water potentials varied strongly within each site and within each of the yearly quarters, resulting in significantly lower water potentials for the oak sites only for the first quarter. Averaged over the whole year, however, the oak sites had significantly (1.5–2 times) lower water potentials compared with the beech sites (Table 2).

Soil Bacteria and Fungi Abundance and Diversity

Bacteria are here defined as being bacterial and archaeal taxa. Bacterial and fungal abundances in bulk soil, as indicated by quantitative PCRs, showed no significant differences between the beech and oak sites, with a bacteria to fungi ratio of about 75–110 in the upper 50 cm of soil, and a ratio of about 400 at a depth of 150 cm (Table 4). Soil depth, in contrast, had a significant effect on fungal abundances in soils, with values decreasing up to 500 times with greater soil depth. In contrast, bacterial abundance in bulk soils tended to be higher in the subsoil. The bacteria to fungi ratio in bulk soils of both beech and oak sites increased by about five times, from around 80 to 400, from the top layer (L1) to the lowest layer (L6) (Table 4). Soil depth did not affect the bacterial or fungal abundance or their ratios in the fine root samples, whereas fungal abundance was significantly higher in beech compared with oak roots.
After quality filtering we identified among the 234 soil samples a total of 9,256,287 bacterial sequences with an average of 39,556 ± 7,668 per sample. These sequences were distributed among 12,543 OTUs (without singletons), 95.2% of which could be classified at the phylum level (32 phyla) and 15.3% at genus level (382 genera). The predominant bacterial phyla were Proteobacteria (22.3%; 2385 OTUs), Planctomycetes (20.1%; 2235 OTUs), Acidobacteria (12.8%; 935 OTUs), Verrucomicrobia (11.2%; 684 OTUs) and Actinobacteria (11.1%; 957 OTUs). The most abundant genera were Bradyrhizobium (2.0% of relative abundance), Streptomyces (1.1%) and Rhodomicrobium (0.8%). We further identified 89 archaeal OTUs with a total of 85,662 sequences.

For fungi, sequence analysis yielded a total of 10,960,463 sequences, averaging 46,840 ± 4,711 per sample. These sequences represented a total of 3,133 fungal OTUs (without singletons), 95% of which could be classified at the phylum level (8 phyla) and 55% at genus level (467 genera). The predominant fungal phyla were Boletus (29.7% of relative abundance), Mortierellomycota (20.1%; 2385 OTUs), Planctomycetes (20.1%; 2235 OTUs), Acidobacteria (12.8%; 935 OTUs), Verrucomicrobia (11.2%; 684 OTUs) and Actinobacteria (11.1%; 957 OTUs). The most abundant genera were Inocybe (29.7% of relative abundance), Boletus (5.3%) and Conocybe (2.9%). A complete list of all bacterial and fungal OTUs, including their taxonomic assignment, the number of sequences and abundance information can be found in the Supplementary Data Sheets 3, 4.

The α-diversity of bacteria showed highly significant effects of site, substrate (bulk soil vs. fine roots), tree species, and soil depth, with the one exception that the Shannon index was not significantly affected by tree species (Table 5). A similar pattern was observed for the fungal α-diversity, with the one exception that the sites had no significant effect on the Shannon index (Table 5).

Both, richness and the Shannon index showed a decreasing trend with increasing soil depth for both bulk soil and fine roots (Figure 1). However, the decrease in richness was much more pronounced than the decline in the Shannon index. For both bacteria and fungi, the Shannon index remained more or less constant across the soil layers L1–L6 (0–155 cm depth) and only decreased in the deepest layer L7 (180–200 cm; Figure 1). Comparing the microbial communities in the bulk soil with those of the fine roots, it is clear that both richness and the Shannon index were higher in the soil. The richness of bacteria in the soil was about 2 × higher and of fungi about 4 × higher compared with values for fine roots. Similarly, the Shannon index of bacteria in the bulk soil was about 1.2 × and of fungi about 1.5–3 × higher compared with values for fine roots (Figure 1). It is noteworthy that fungi associated with the fine roots of oaks had a distinctly higher Shannon index (about 2) than those associated with the fine roots of European beech (about 1), while fungi of the bulk soils of both forest types had about the same Shannon index (about 3) (Figure 1).

PERMANOVA analyses revealed significant differences in bacterial and fungal communities (β-diversity) with site, soil

![Table 4](https://example.com/table4.png)

**Table 4**: Mean values of bacterial and fungal abundances (16S and ITS copy numbers) and the bacteria to fungal ratio in bulk soils and fine roots (n = 3 forest sites).

| Soil layers | Beech sites | Oak sites | p-ANOVA + |
|-------------|-------------|-----------|-----------|
|             | L1 | L3 | L6 | L1 | L3 | L6 | Tree | Depth | Interaction |
| Abundance in soils: | | | | | | | | | |
| Bacterial 16S (× 10⁶)* | 3.58 | 2.91 | 4.94 | 3.83 | 3.67 | 4.77 | 0.66 | 0.14 | 0.83 |
| Fungal ITS (× 10⁶) | 6.47 | 4.03 | 1.89 | 5.80 | 3.64 | 1.27 | 0.60 | 0.012 | 0.99 |
| 16S:ITS ratio | 75 | 87 | 384 | 77 | 109 | 402 | 0.82 | 0.002 | 0.99 |
| Abundance in roots: | | | | | | | | | |
| Bacterial 16S (× 10⁶) | 3.17 | 2.35 | 2.62 | 3.59 | 3.70 | 2.60 | 0.10 | 0.20 | 0.26 |
| Fungal ITS (× 10⁶) | 8.18 | 6.61 | 9.74 | 1.14 | 1.33 | 1.31 | <0.001 | 0.66 | 0.65 |
| 16S:ITS ratio | 0.47 | 0.45 | 0.38 | 4.45 | 8.56 | 6.31 | 0.06 | 0.79 | 0.79 |

*Effects of tree genus, soil depth and their interaction were assessed by analysis of variance (ANOVA); significant values (p < 0.05) are in bold.

†Soil layers: L1: 0–10 cm, L3: 45–55 cm, L6: 140–155 cm.

‡16S and ITS copy numbers are given in µg−1 DNA.

![Table 5](https://example.com/table5.png)

**Table 5**: Main effects on bacterial and fungal α-diversity (Richness and Shannon index) and β-diversity (Bray-Curtis dissimilarities) of forest sites (Chamounx, Neurkirch, Salion), soil depths (soil layers: L1: 0–10 cm, L2: 15–25 cm, L3: 45–55 cm, L4: 75–85 cm, L5: 110–125 cm, L6: 140–155 cm, L7: 180–200 cm), tree genera (beech, oaks), and substrates (bulk soil, fine roots).

| α-diversity* | β-diversity* |
|--------------|--------------|
|               | Bacteria:     | Fungi:      |
|               | Site | Depth | Tree | Substrate | Site | Depth | Tree | Substrate | Site | Depth | Tree | Substrate |
|               | 2    | 6     | 1    | 1        | 2    | 6     | 1    | 1        | 2    | 6     | 1    | 1        |
|               | 51.3 | 16.8  | 8.4  | 459      | 16.9 | 38.5  | 8.8  | 517      | <0.001 | <0.001 | <0.001 | <0.001 |
|               | <0.001 | <0.001 | <0.004 | <0.001 | <0.001 | <0.001 | <0.003 | <0.001 | <0.001 | <0.001 | <0.001 |
|               | 17.5 | 4.9   | 0.6   | 288      | 1.7  | 4.8   | 0.43  | 444      | 12.8  | 5.7   | 7.2   | 60.8     |
|               | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 | <0.001 |

*Effects of site, soil depth, tree genus, and substrate were assessed by analysis of variance (ANOVA).

†Effects of site, soil depth, tree species and substrate were assessed by permutational multivariate analysis of variance (PERMANOVA).

‡Values represent degrees of freedom (DF), the F-ratio (F), and the level of significance (p); significant values (p < 0.05) are in bold.
depth, tree species, and substrate (Table 5). The canonical analysis of principal coordinates (CAP) based on Bray–Curtis dissimilarities for microbial communities showed that bacterial communities were significantly separated across the soil layers (L1–L7; $F = 5.7$, $p < 0.001$). However, they were less strongly dependent on tree species (beech vs. oak; $F = 7.2$, $p < 0.001$), in contrast to the substrate (bulk soil vs. fine roots; $F = 77.7$, $p < 0.001$) (Table 5 and Figures 2A–C).

The non-metric multidimensional scaling (NMDS) ordination based on Bray–Curtis dissimilarities for the microbial communities in the various soil layers (L1–L6) of beech and oak sites indicated a strong separation of the soil bacterial communities (stress = 0.1; Figure 3A). We further identified the individual soil environmental variables explaining the bacterial and fungal community structures. Correlations with the soil environmental variables revealed that fine root biomass, clay content, $C_{\text{org}}$, and $N_{\text{tot}}$ were highly significantly related ($R^2 = 0.39–0.78$, $p < 0.001$) to the bacterial communities in the upper soil layers, whereas communities in the lower soil layers corresponded closely to the stone content, fine-earth density, pH, and sand content ($R^2 = 0.28–0.38$, $p < 0.001$; Table 6). The fungal communities were not or only weakly related to the soil parameters silt content, C:N ratio, base saturation, and AWC (Figure 3B).

**Differential Abundance of Bacterial and Fungal Taxa in Response to Soil Depth, Tree Genus and Substrate**

Overall, bacterial and archael taxa belonging to the phyla Bacteroidetes, Verrucomicrobia, Patescibacteria, Proteobacteria, Planctomycetes and Acidobacteria were significantly more abundant in topsoils than in subsoils, independently from the substrate (Figure 4). In contrast, poorly known bacterial phyla, such as Nitrospirae, Chloroflexi, Rokubacteria, Gemmatimonadetes and GAL 15, were overrepresented in the subsoils. Furthermore, archael phyla, such as Thaumarchaeota and Euryarchaeota, were also more abundant in subsoils.

---

**FIGURE 1** | Richness and Shannon index of bacteria and fungi of soils and fine roots in the various soil layers (L1–L7) at beech and oak sites. Soil layers: L1: 0–10 cm, L2: 15–25 cm, L3: 45–55 cm, L4: 75–85 cm, L5: 110–125 cm, L6: 140–155 cm, L7: 180–200 cm.
Interestingly, we found many poorly characterized taxa (i.e., abundant in bulk soils than in root samples (data not shown). That were overrepresented in the subsoil were also more abundant in bulk soils than in root samples (data not shown). In contrast, the genera Edaphobacter, Granulicella, and Bryobacter (both Acidobacteria) were significantly increased in topsoils relative to subsoils. Similarly, the genera Singulisphera (Planctomycetes) and Rhodanobacter (Proteobacteria) were significantly more abundant in topsoils (Figure 5).

The abundance of Anaerolineaceae taxa (belonging to Chloroflexi) was significantly higher in bulk soil samples from beech compared with oak sites (Supplementary Figure 2). We also found that several poorly characterized bacterial taxa from Chloroflexi (Anaerolineaceae, Ktedonobacteraceae; candidate classes Gitt-GS-136, P2-11E and TK10; candidate orders SBR1031 and S085) were more abundant in bulk soil samples than in fine root samples (Supplementary Figure 3). In contrast, several root-associated bacterial taxa, such as Pseudomonas, Rhodomicrobium and Sterptomycetes exhibited a positive log$_2$-fold change in root samples relative to bulk soil samples.

Within the fungal kingdom, the majority of genera that exhibited either positive or negative log$_2$-fold changes in abundance with soil depth (topsoil L1 and L2 vs. subsoil L5 and L6) were members of the phyla Ascomycota and Basidiomycota. The ectomycorrhizal genera Amanita, Melanogaster, Odontia, Sistotrema, and Telephora were among the fungal taxa that showed the highest positive log$_2$-fold change in the topsoils. In contrast, a different group of genera of ectomycorrhizal fungi were found in the subsoils such as Boletus, Clavulina, Hydnum, Mallochybe, Membranomyces, and Pseudomomentella, and (Figure 6). The majority of fungi with the largest log$_2$-fold changes in abundance with soil depth, however, were saprotrophs. Abundant saprotrophs in the topsoils were Oidiodendron, Saitozyma, and Trichoderma, which are all litter decomposers. In addition, Sagenomella was recorded, which is in contrast to the saprotrophs an animal pathogen. In the subsoils, abundant saprotrophs were members of Xylaria, a well-known wood decomposer (Figure 6).

Abundant ectomycorrhizal genera predominantly found in beech sites were Hebeloma and Hymenogaster, and in the oak sites Cenococcum and Russula (Supplementary Figure 4). One exception is Glomus, which belongs to Glomeromycota and forms arbuscular mycorrhizas with roots of herbs, grasses, and shrubs. Further abundant genera in beech sites were the ascomycetous yeast Lipomyces and the plant pathogen Idriella. Cistella, Leohumicola, and Xenasma were the most dominant saprotrophs in oak sites (Supplementary Figure 4).

Ectomycorrhizal fungi predominantly and abundantly found in the bulk soils but not in the fine roots were, e.g., Clavulina and Melanogaster (Supplementary Figure 5). In contrast, no ectomycorrhizal fungi were found exclusively in the fine roots. Differentially abundant saprotrophic fungi recorded in soils were Geminibasidium, Saitozyma, and Solicoccymenia.
all basidiomycetous yeasts; *Mortierella*, a mucoromycetous saprotroph; *Oidiodendron* and *Trichoderma*, ubiquitous decomposing fungi; and *Sagenomella*, an animal pathogen ([Supplementary Figure 5](#)). When comparing the fungi from bulk soils with those from fine roots, we observed that only a few genera predominantly occurred in fine roots, and these were known decomposers of litter and wood (*Lachnum*, *Mycena*, *Mycenella*, *Pezicula*, *Xylaria*) ([Supplementary Figure 5](#)).

### Fungal Guilds

An analysis of the fungi in soils and fine roots with FUNGUILD showed that in both forest types “saprotrophs” and “ectomycorrhizal fungi” dominate over “plant pathogens,” “animal pathogens,” “ericoid mycorrhizal,” and “endophyte fungi.” The latter four guilds all had a relative abundance of <10% with a “highly probable” classification (data not shown). In contrast, saprotrophs and ectomycorrhizal fungi were...
Our understanding of plant ecophysiology, community ecology and geochemical cycles” (Maeght et al., 2013), Germon et al. (2020) pointed out “the critical need to include depth and timescale of ecological processes below the soil surface in a modern conceptualization of forest ecosystems.” Only recently, it was found that particular microbial taxa are consistently more abundant in deep soils and are preferentially adapted to low-nutrient conditions, with the ability to synthesize and store carbohydrates, the potential to use carbon monoxide as a supplemental energy source, and the ability to form spores (Brewer et al., 2019). Interestingly, the few existing studies on forest ecosystems showed that the richness and diversity of soil microorganisms are not necessarily lower in deeper soil layers compared to topsoils. Moreover, specific groups of microorganisms, such as Chloroflexi, Nitrospirae, or Euryarchaeota emerge in the subsurface and are clearly specialized to these conditions (Brewer et al., 2019).

Extreme conditions in soils affect microbial life and biogeochemical processes (Schimel, 2018). Under conditions of periodic or permanent low water availability, which are unquestionably challenging to life, it might be expected that microbial populations would be either scarce or absent. Gremer et al. (2015) estimated that conditions become critical for soil microorganisms below a water potential of −3 MPa. To exist in dry habitats, microorganisms must be capable of surviving extended periods of low water potential. While the production of resistant spores is one option under such conditions (Manzanera, 2020), this capacity is clearly restricted to a limited range of taxa. For organisms lacking this capacity, one of the alternatives is to enter a state of anhydrobiosis (Crowe, 2014), where much of cellular metabolism is shut down and metabolic activity is restricted to processes related to cellular maintenance and repair (Manzioni and Katul, 2014; Lebre et al., 2017).

The investigated beech and oak sites, however, had maximum water potentials around −1 MPa far above values considered critical for microorganisms. However, they were critical for the trees. In the summer months of 2015, most parts of Central Europe suffered from an exceptional heat wave and dry period. The amount of precipitation in that season was among the lowest since 1901 and soil moisture was even lower than during the centennial heat wave of 2003 (Orth et al., 2016). This was also manifested specifically at our oak sites. In particular in the deep soil layers, the soil water potential reached values of −850 to −984 kPa. For young European beech trees in the oak stands, such low water potentials could be critical, as their leaves begin to drop at about −800 kPa (Walthert et al., 2021). Furthermore, such critical soil water potentials could also lead to the death of fine roots, which can then stimulate microbial activity and thereby increase soil organic matter (SOM) degradation (referred to as the rhizosphere priming effect; Adamczyk et al., 2019a).

### DISCUSSION

#### Life in Deep and Dry Soils

The subsurface of soils might be one of the last frontiers of undiscovered sites, also because “an increasing number of studies clearly indicate that ‘looking deeper’ is essential to increase

| Soil biological variables: | Bacteria | Fungi |
|----------------------------|----------|-------|
| Tree genus                 | 0.09     | 0.34  |
| Fine root biomass          | 0.78     | 0.55  |

| Soil chemical variables:  | Bacteria | Fungi |
|---------------------------|----------|-------|
| pH (CaCl₂)                | 0.55     | 0.29  |
| Base saturation           | 0.30     | 0.13  |
| Corg                      | 0.47     | 0.38  |
| Ntot                      | 0.39     | 0.38  |
| C:N ratio                 | 0.14     | 0.14  |

| Soil physical variables:  | Bacteria | Fungi |
|---------------------------|----------|-------|
| Fine-earth density        | 0.55     | 0.38  |
| Stones                    | 0.58     | 0.38  |
| Sand                      | 0.39     | 0.28  |
| Silt                      | 0.36     | 0.14  |
| Clay                      | 0.65     | 0.60  |
| AWC10                      | 0.31     | 0.27  |

Compare also Figure 3 for the arrow vectors and abbreviations, and Table 3 for the units.

*Significant values (p < 0.05) are in bold.

*Plant-available water storage capacity (AWC) per 10 cm soil thickness.

Soil Physicochemical Variables Affect Microbial Community Structures

Environmental conditions, such as oxygen levels, soil moisture, temperature, pH, fine-earth density, soil organic carbon content (SOC), and N content, can influence the structure of the soil
FIGURE 4 | Differentially abundant bacteria phyla of topsoil (soil layers L1 and L2) vs. subsoil (soil layers L5 and L6). Shown are only significant log$_2$-fold changes (LFC) of phyla based on a significance level of $p < 0.01$ after false discovery rate correction. Error bars represent standard deviation. Positive LFC values indicate a higher occurrence of phyla in the topsoil, whereas negative LFC indicate a higher occurrence in the subsoil. The number of differentially abundant OTUs comprised in each phyla are indicated in brackets. Poorly known phyla are highlighted in red. Relative abundance (%): sum of the abundance of the differentially abundant OTUs belonging to the same phyla as a percentage of the total number of sequences across the compared soils. Soil layers: L1: 0–10 cm, L2: 15–25 cm, L5: 110–125 cm, L6: 140–155 cm.

microbial community (Fierer, 2017; Schimel, 2018). Therefore, changes in soil physicochemical characteristics with depth can have significant positive (e.g., SOC and N) or significant negative correlations (e.g., pH and fine-earth density) with the diversity of the microbial community. At the continental scale, soil pH has been identified as the dominant factor shaping microbial communities in surface soils (Rousk et al., 2010). In our study, however, we assumed that pH was not an important factor causing variation in microbial communities throughout the soil profiles. Our forest soil samples contained a very narrow pH range (6.6–7.8), so this variable most likely was not important in structuring the microbial communities. We expect that changes in the quality of SOC (e.g., labile SOC and resistant SOC) with soil depth have stronger effects on the community structure than the total quantity of SOC (Seuradge et al., 2017). Shifts in microbial communities between surface and subsurface soils could partially result from differences in the availability of labile SOC, leading to physiologically altered (adapted) organisms capable of utilizing more recalcitrant sources of organic carbon (Fierer et al., 2003).

A non-metric multidimensional scaling (NMDS) analysis revealed differences in microbial community structures within the soil profiles for both bacteria and fungi, although more
distinctly for bacteria. The major factor that was related to microbial community structure was soil depth together with fine root biomass and C\textsubscript{org}, N\textsubscript{tot}, and clay content. This result is consistent with the observations of Fierer et al. (2007) that the availability and quality of nutrient resources appear to have a major influence on microbial community composition in soil profiles. Our topsoils contained significantly larger amounts of fine roots, C\textsubscript{org}, N\textsubscript{tot}, and clay than subsoils, both at beech and oak forest sites. This result also indicates that fine roots strongly influence the amounts of C\textsubscript{org} and N\textsubscript{tot} in soils (see also Rasse et al., 2005). Clay minerals, on the other hand, are associated with C\textsubscript{org} because they bind strongly together and can persist for thousands of years (Rümpel and Kögel-Knabner, 2011). The decrease in C\textsubscript{org} and N\textsubscript{tot} content in soils as the depth increased was consistent with findings from previous studies (Du et al., 2017, 2020; Jiao et al., 2018; Feng et al., 2019). Organic C content typically declines with soil depth because of reduced nutrient input from plant litter and root exudates (Fierer et al., 2003). The decrease in microbial biomass, measured in our study as DNA content, with increasing soil depth was also consistent with results from previous studies (Eilers et al., 2012; Hu et al., 2019; Zhang et al., 2019). At our forest sites, DNA content decreased

---

**FIGURE 5** | Differentially abundant bacteria genera of topsoil (soil layers L1 and L2) vs. subsoil (soil layers L5 and L6). Shown are only significant log\textsubscript{2}-fold changes (LFC) of genera based on a significance level of \( p < 0.01 \) after false discovery rate correction. Error bars represent standard deviation. Positive LFC values indicate a higher occurrence of the taxa in the topsoil, whereas negative LFC indicate a higher occurrence in the subsoil. Only 30 genera with the highest (positive) and lowest (negative) mean LFC values were represented. The number of differentially abundant OTUs comprised in each genus, and the phyla to which they belong are indicated in brackets. Poorly known taxa are highlighted in red. Relative abundance (%): sum of the abundance of the differentially abundant OTUs belonging to the same genera as a percentage of the total number of sequences across the compared soils. Relative abundance depicted in red exceeds the abundance scale. Ac, Acidobacteria; At, Actinobacteria; Cf, Chloroflexi; F, Firmicutes; Ni, Nitrospirae; Pi, Planctomycetes; Pr, Proteobacteria; R, Rokubacteria; T, Thaumarchaeota. Soil layers: L1: 0–10 cm, L2: 15–25 cm, L5: 110–125 cm, L6: 140–155 cm.
drastically from the surface soil to a depth of about 0.5 m, and then the value kept relatively steady at greater depths.

**Microbial Abundance and Diversity in Deep Soils**

Our understanding of soil microbial ecology remains particularly limited with respect to changes in microbial communities with depth in particular in forests (Pickles and Pither, 2014). Nonetheless, many important biogeochemical processes, such as C$_{\text{org}}$ and N cycling, are carried out differently throughout the soil profile by distinct microbial assemblages (Eilers et al., 2012). Here, soil microbiomes followed a vertical distribution in terms of their abundance, diversity and community structure. We found consistent decreases in bacterial alpha-diversity with increasing depth, corroborating findings from previous studies on forests (Fierer et al., 2003; Seuradge et al., 2017; Deng et al., 2018; Jiao et al., 2018; Brewer et al., 2019; Feng et al., 2019). Our results go further, showing that fungal diversity in both beech and oak soil profiles also decreased with increasing depth. This trend of decreasing fungal richness was much stronger than the pattern observed for bacteria. Bacterial and fungal community structure also showed a clear vertical stratification with depth, in agreement with previous studies (Deng et al., 2018;
Fungal functional guilds were analyzed by FUNGuild. Dashed lines have been drawn for better visualization. Soil layers: L1: 0–10 cm, L2: 15–25 cm, L3: 45–55 cm, L4: 75–85 cm, L5: 110–125 cm, L6: 140–155 cm.

Tang et al., 2018; Feng et al., 2019; Hao et al., 2021). In particular, bacterial communities in the topsoils (0–25 cm soil depth) were clearly separated from those in the deeper layers. The vertical differences in these microbiomes can be attributed in large part to the decrease in availability of various resources with increasing soil depth (Wu et al., 2020). If oxygen levels also decrease with soil depth, then this could also have contributed to the observed microbiome differences, since different microorganisms prefer different oxygen conditions – archaeal taxa are mainly anaerobic, while bacteria and fungi are mainly aerobic (Fontaine et al., 2007; Haroon et al., 2013).

**Shifts in Bacterial Taxa With Soil Depth**

Our results demonstrate that the composition of the soil microbiome shows stratum specificity across the 2 m o soil profile. At a coarse taxonomic resolution we identified several bacterial and archaeal phyla that consistently increased in abundance with increasing soil depth, as measured by DESeq2 analyses across the entire data set: *Nitrospirae, Chloroflexi, Firmicutes, Rokubacteria, Thaumarchaeota, Gemmatimonadetes, Euryarchaeota,* and GAL15. Their presence in subsoils may indicate an adaptation of these members to the low-nutrient environments characteristic of bulk soils.

Members of *Chloroflexi* and *Nitrospirae* have previously been found to increase in abundance with increasing soil depth in individual soil profiles (Will et al., 2010; Brewer et al., 2019; Feng et al., 2019), while candidate phylum GAL15 has been shown to be abundant in sedimentary subsoils (Lin et al., 2012; Gocke et al., 2017; Brewer et al., 2019; Yan et al., 2019; Liu et al., 2020; Hao et al., 2021), floodplain soils (Steger et al., 2019), and forest subsoils (Feng et al., 2019). Members of these phyla are likely oligotrophic taxa adapted to surviving under the resource-limited conditions found in deeper layers. In the study of Yan et al. (2019) the candidate phylum GAL15 was sensitive to variation in most of the soil properties and was significantly negatively correlated with SOM, N$_{tot}$, available K, and available P in soil profiles down to 3 m in agricultural land planted with legumes. In addition to GAL15, the subsoil also contained a greater relative abundance of *Rokubacteria* than in topsoils, which is in accordance with our findings.

Several members of *Chloroflexi* were shown to be enriched in the subsoil. In particular, we found several uncultivable bacteria from the candidate classes Gitt-GS-136 and SBR1031 within *Chloroflexi*. Studying uncultivable bacteria provides an opportunity to explore poorly characterized species of bacteria, thus providing insight into their diversity and also helping to understand their role in deep soil layers. *Chloroflexi* is a diverse
Our study sheds light on the vertical spatial variation of the forest community is supported by chemolithotrophy, with limited Archaea GAL15, and may prefer a micro-aerobic environment rather than aerobic Verrucomicrobia, studies (Will et al., 2010; Feng et al., 2019). significantly higher in topsoils (0–25 cm) than in subsoils (e.g., 0–10 cm; Egidi et al., 2019). One exception is the study by Jiao et al. (2018) investigating bacterial, archaeal, and fungal communities in soils down to 3 m on reforestation sites with Robinia pseudoacacia L. on the Loess Plateau in China. The authors observed only a slight decrease of the Shannon index from around 5.0 to about 4.4 in the deepest soil layers. As R. pseudoacacia is not considered an ectomycorrhizal forming tree species, dominant fungal taxa at the deep layers were ascomycetous fungi, such as Cladosporium, Chaetomium and Trichoderma, all three of which are saprotrophic taxa, Gibberella and Stagonospora, both plant pathogens, or Metarhizium, an animal pathogen (Jiao et al., 2018). The additional occurring and frequent genus Lysurus was an exception, as it belongs to the saprotrophic Basidiomycota and forms stinkhorns as fruiting bodies. Recent studies exploring soils below 1 m depth have been dedicated to searching in particular for ectomycorrhizal fungi associated with deep-rooting trees. Robin et al. (2019) recorded Pisolithus and Scleroderma species in association with roots of Eucalyptus grandis W. Hill ex Maiden throughout the soil profiles down to 4 m in Brazil.

In our study, we also identified a Pisolithus species, Pisolithus arrhizus (Scop.: Pers.) Rauschert, at depths of 0.5–1.5 m in the soil samples of an oak site. However, this species was only rarely recorded, and thus not recognized by our differential abundance analysis. More commonly, we recorded in deep soil layers the ectomycorrhizal genera Boletus and Hydnum, and to a lesser extent Clavulina, Malloxybe, Membranomyces, and Pseudotomentella. Boletus luridus Schaeff. (Syn. Suillus luridus (Schaeff.) Murrill) was the only member of Boletus, but was commonly found in deep soil layers in beech and oak sites, and in both, soil and fine root samples. In addition three Hydnum taxa were recorded, but only Hydnum vesterholtii Olariaga, Grebenc, Salcedo, Martin, occurred on an abundant scale. Hydnum vesterholtii was recently newly described in the French Pyrenees under Fagus sylvatica and Abies alba on calcareous soils (Olariaga et al., 2017). More records are available from the Toscana region in Italy, the Huesca region in Spain, and Andorra under Fagus, Quercus, Castanea, Corylus, Abies, and Pinus (Olariaga et al., 2017). To our knowledge, our observations with B. luridus and H. vesterholtii are the first to suggest that these fungal taxa most likely colonize tree roots in deep soil layers.

Shifts in Fungal Taxa With Soil Depth
Previous studies on microbial communities in deep soil layers have been focused predominantly on bacteria (e.g., Li et al., 2014; Gocke et al., 2017; Liu et al., 2020). In studies where fungal taxa were considered, soil samples were mainly taken from topsoils (0–10 cm; Egidi et al., 2019). Although archaeal taxa were relatively rare in all soil layers in our study, some individual members of Archaea were abundant. Thaumarchaeota was the dominant taxon of archaeal organisms in our forest soils. Interestingly, Thaumarchaeota was more abundant in subsoils, which is in contrast findings from other studies (Mikkonen et al., 2014) but supports a study by Liu et al. (2020). Thaumarchaeota are generally negatively correlated with Corg (Liu et al., 2020) and are able to compete with various groups of bacteria in soils (Hansel et al., 2008). Thaumarchaeota have been shown to have ammonia oxidation (Treusch et al., 2005) and C sequestration activities (Koenneke et al., 2014). The relative abundance of Euryarchaeota increased significantly with increasing depth in our forest soils, which could be attributed to their preference for anaerobic environment. Notably, members of Euryarchaeota produce methane (Bodelier and Steenbergh, 2014), but also oxidize methane, fix N, and reduce nitrates (Cabello et al., 2004). With respect to archaeal phyla, which were overrepresented in the subsoils in our study, we conclude that they play a major role in soil nutrient cycling in deep layers.

Overall, bacteria are known to be involved in various soil processes and global biogeochemical cycling, such as organic matter degradation, and N cycling. The presence of chemolithotrophic prokaryotes (i.e., Chloroflexi, Nitrospirae, GAL15, and Archaea) in deep soils suggests that the microbial community is supported by chemolithotrophy, with limited utilization of organic carbon. Besides chemolithotrophy in the deep-soil community, metabolic cooperation via syntrophy between archaeal and bacterial groups plays a critical role in the survival of the whole community under oligotrophic conditions. Our study sheds light on the vertical spatial variation of the forest soil microbiomes at a fine scale.

In contrast, the relative abundance of Verrucomicrobia, Bacteroidetes, Planctomycetes and Proteobacteria was significantly higher in topsoils (0–25 cm) than in subsoils (110–155 cm) in our study, which is consistent with previous studies (Will et al., 2010; Feng et al., 2019). Verrucomicrobia may prefer a micro-aerobic environment rather than aerobic and extreme anoxic environments (Eilers et al., 2012). The decreased abundances of Proteobacteria and Bacteroidetes with increasing soil depth could be attributed to the general copiotrophic properties and responses to labile sources of Corg at shallow depths (Fierer et al., 2007). Soil organic carbon may affect the relative microbial abundance of Planctomycetes across the soil profile. Liu et al. (2020) reported that the abundance of Planctomycetes was significantly correlated with soil organic carbon, indicating that copiotrophic conditions benefit the growth of this phylum.

Phylum including autotrophic, heterotrophic and mixotrophic taxa (Hanada et al., 2002; Bennett et al., 2020). Thus, different physiological strategies might be responsible for them coping with the harsh environments in subsoils. Chloroflexi have also been found to be highly abundant in C-poor alpine environments at high elevations (Rime et al., 2016; Adameczk et al., 2019b). Two other phyla, Gemmatimonadetes and Nitrospirae, were also found in greater relative abundance in the subsoil. Previous research has highlighted that Gemmatimonadetes members are strongly adapted to low soil moisture conditions (DeBruyn et al., 2002; Steger et al., 2019). Furthermore, members from the Nitrospirae phylum have been shown to prefer non-rhizosphere environments (Dunbar et al., 2002; Steger et al., 2019). Will et al. (2010) hypothesized that heterotrophic microorganisms that are associated with roots suppress the growth of autotrophic Nitrospirae. Considering the general decrease in Corg and N tot with depth observed in our study, the greater abundance of members of the Nitrospirae phylum in deeper soil may be due to a selection advantage that dark-adapted chemolithoautotrophic members have in subsoil systems (Will et al., 2010).

Although archaeal taxa were relatively rare in all soil layers in our study, some individual members of Archaea were abundant. Thaumarchaeota was the dominant taxon of archaeal organisms in our forest soils. Interestingly, Thaumarchaeota was more abundant in subsoils, which is in contrast findings from other studies (Mikkonen et al., 2014) but supports a study by Liu et al. (2020). Thaumarchaeota are generally negatively correlated with Corg (Liu et al., 2020) and are able to compete with various groups of bacteria in soils (Hansel et al., 2008). Thaumarchaeota have been shown to have ammonia oxidation (Treusch et al., 2005) and C sequestration activities (Koenneke et al., 2014). The relative abundance of Euryarchaeota increased significantly with increasing depth in our forest soils, which could be attributed to their preference for anaerobic environment. Notably, members of Euryarchaeota produce methane (Bodelier and Steenbergh, 2014), but also oxidize methane, fix N, and reduce nitrates (Cabello et al., 2004). With respect to archaeal phyla, which were overrepresented in the subsoils in our study, we conclude that they play a major role in soil nutrient cycling in deep layers.

Overall, bacteria are known to be involved in various soil processes and global biogeochemical cycling, such as organic matter degradation, and N cycling. The presence of chemolithotrophic prokaryotes (i.e., Chloroflexi, Nitrospirae, GAL15, and Archaea) in deep soils suggests that the microbial community is supported by chemolithotrophy, with limited utilization of organic carbon. Besides chemolithotrophy in the deep-soil community, metabolic cooperation via syntrophy between archaeal and bacterial groups plays a critical role in the survival of the whole community under oligotrophic conditions. Our study sheds light on the vertical spatial variation of the forest soil microbiomes at a fine scale.

In contrast, the relative abundance of Verrucomicrobia, Bacteroidetes, Planctomycetes and Proteobacteria was significantly higher in topsoils (0–25 cm) than in subsoils (110–155 cm) in our study, which is consistent with previous studies (Will et al., 2010; Feng et al., 2019). Verrucomicrobia may prefer a micro-aerobic environment rather than aerobic and extreme anoxic environments (Eilers et al., 2012). The decreased abundances of Proteobacteria and Bacteroidetes with increasing soil depth could be attributed to the general copiotrophic properties and responses to labile sources of Corg at shallow depths (Fierer et al., 2007). Soil organic carbon may affect the relative microbial abundance of Planctomycetes across the soil profile. Liu et al. (2020) reported that the abundance of Planctomycetes was significantly correlated with soil organic carbon, indicating that copiotrophic conditions benefit the growth of this phylum.
The most abundant group of fungi with the largest log₂-fold changes in deep soil layers, however, were the saprotrophic litter and wood decomposers Chaetomium, Pezicula, and Xylaria. Chaetomium taxa are typically ascomycetous saprotrophs and have the ability to degrade cellulose (Dreyfuss, 1976). More than 400 species have been described, some of which can cause human infection (Wang et al., 2014). In our subsols, an unknown Chaetomium taxa dominates mainly in the oak sites with a few records in the beech sites. Pezicula, another ascomycetous fungal genera, is predominantly reported from temperate regions, where they occur as saprotrophs on recently dead branches and twigs. However, they have also been frequently isolated from living twigs and roots without showing symptoms of disease, thus exhibiting an endophytic lifestyle (Chen et al., 2016). The dominant taxa in our study was Pezicula brunnea (Sigler) P. R. Johnst., which occurred throughout the soil profiles in beech and oak sites and in bulk soil and fine roots, but with a strong preference for deep soil layers. This fungus was originally isolated from roots of an ericaceous plant from Canada and described as Cryptosporiopsis brunnea Sigler (Sigler et al., 2005). The genus Xylaria is a well-known group of fungi that decompose wood. As a special feature, these fungi produce the black pigment melanin that makes their hyphae more resistant against stressors such as drought (Tudor et al., 2013). Of the several Xylaria taxa found, two unknown taxa were most abundant in deep soil layers, one predominantly in beech sites, the other in oak sites.

Noteworthy in deeper soil layers is the record of an additional basidomycetous genus, Naganishia. The most abundant species was Naganishia albida (Saito) X.Z. Liu (Syn. Cryptococcus albidus), an extremophilic black-melanized yeast, which can be found frequently in soils of extreme environments, such as Antarctica or the Atacama Desert in Chile (Schmidt et al., 2017). Naganishia albida is also known as a human pathogen, but it is not harmful to plants. In addition, we recorded in deep soil layers an unknown tree pathogen of the genus Neonectria. Many species of Neonectria can induce bark cankers, such as N. major (Castlebury et al., 2006), which we also recorded in our forest sites but mainly in the topsoils.

The ectomycorrhizal genera Amanita, Melanogaster, and Telephora were among the fungal taxa that showed the largest positive log₂-fold change to colonize the topsoils. But overall, the most common ectomycorrhizal genera of the two forest types were Inocybe with >30 taxa, Cortinarius with >25 taxa, and Russula with >15 taxa. Among them, Amanita pantherina and Cortinarius subtilis occurred exclusively in the topsoils of both forest types.

Effects of Tree Genus and Substrate Type on Microbial Communities

Our findings demonstrate that tree genera affected both the bacterial and fungal community structure. However, the fungal community structure (F = 28.9; p < 0.001) showed a much stronger dependence on tree species than the bacterial community structure (F = 7.2; p < 0.001). Thus, our results confirm the hypothesis that fungal communities are more affected by the predominant forest vegetation than the bacterial communities, which supports the findings of other studies (e.g., Urbanová et al., 2015). These authors explained the lower dependence of bacterial communities from tree roots as resulting from the unicellularity and size of most bacteria. Bacteria are expected to be exposed to the conditions of their immediate microhabitats, which are often the size of a single soil pore and which have very specific conditions that differ from the average properties of their environmental matrix (Vos et al., 2013). Due to their size, most of these microniches are likely separated from the direct influence of plant roots and may be the reason why bacterial communities in soils under different trees are affected to lesser extent than fungi (Urbanová et al., 2015).

Fierer et al. (2003) suggested that the influence of plants on soil microbial community composition is typically limited to the 20-cm topsoil layer. In contrast, we observed that the effect of tree species on fungal community could reach a depth of 2 m. At least part of the effect of tree species on soil fungi could be due to indirect effects mediated through specific understory vegetation and litter chemistry (Urbanová et al., 2015). Tree species induce a specific signature in terms of litter quality, and litter is the primary source of soil Corg and N, meaning that it determines the substrate quality for the growth of soil microorganisms (Baldrian, 2019). For example, oak bark has a higher tannin and total phenolic content compared to beech bark, with values of about 22% for oak and 15% for beech (Torkaman and Seyam, 2009). For the ectomycorrhizal communities we also observed a distinct difference between the two tree species, which is not surprising, as a recent study throughout Europe also showed tree specificity (van der Linde et al., 2018). Urbanová et al. (2015) concluded, that ectomycorrhizal fungi are more strongly associated with tree composition than with soil properties, in contrast to saprotrophic fungi, which are more strongly correlated with soil properties. Schappe et al. (2020) found similar relationships in tropical rainforests in Panama.

In our study, we also found that the ratio of saprotrophs to ectomycorrhizal fungi remained more or less constant at the beech sites with increasing soil depth, at values of about 1, whereas at the oak sites the ratio decreased from about 2 to 1. Therefore, our fourth hypothesis, that the relative abundances of saprotrophic fungi decrease more strongly with soil depth than ectomycorrhizal fungi, was confirmed for oak but not for beech sites. This observation at the oak sites supports the so-called Gadgil effect, which describes the increase of saprotrophic fungi when ectomycorrhizal fungi are excluded by trenching (Gadgil and Gadgil, 1971, 1975). In other words, "competition between these two fungal guilds is hypothesized to lead to suppression of decomposition rates" (Fernandez and Kennedy, 2016). At our oak sites, the ectomycorrhizal fungi occurred in a relatively lower abundance, so that competition between these two guilds would be reduced, which would promote the saprotrophs and, consequently, the decomposition of litter. The need for more saprotrophs could be explained by the fact that oak litter has a higher recalcitrance than beech litter, due to a high phenolic content. For the beech sites, however, a suppression of decomposition due to the high relative abundance of ectomycorrhizal fungi was not observed. Nevertheless, the soil Corg content of the two forest site types was similar.
It is likely that the microbial communities of fine roots are similar to those in the surrounding soil, as the soil represents the inoculum pool that can potentially colonize the roots (Goldmann et al., 2016). However, only a portion of this soil inoculum pool is actively associated with the roots. In German beech-dominated forests, Goldmann et al. (2016) found that root-associated fungi are recruited from the soil fungal community. Similarly, in Scots pine forests in France and Spain, Rincón et al. (2015) found that the total fungal richness in soils largely doubled that of root tips, and that more than 93% of the taxa in root tips were also found in bulk soil. These authors pointed out that bulk soil is a good proxy for estimating ectomycorrhizal fungal richness. Our study leads to a similar conclusion, that the community structure in the fine roots is, for both bacteria and fungi, well represented in the soils.

CONCLUSION

Microorganisms in deeper soil layers are considered to play important roles in soil development and biogeochemical cycling, given their proximity to parent material, their potential for nutrient and carbon storage, and their lower turnover rates than surface soil microbiota. Our results from 2 m soil profiles demonstrate that the richness, abundance, and composition of the soil microbiome display layer specificity and distinct differences across strata. Soil depth is an important factor affecting microbial diversity and structure, and the gradient in the soil environment caused by soil depth, reflecting shifts in availability of resources and local chemical conditions, may be the most important factor in the vertical distribution characteristics of the microbial taxa. In particular, for bacteria we found several poorly characterized taxa in the subsoils with unknown function. Given the presence of chemolithotrophic bacteria at depth, we conclude that the microbial community is supported by chemolithotrophy with limited utilization of organic carbon. We also observed, to our knowledge for the first time, that Boletus luridus and Hydnum versterholtii colonize tree roots, most likely in deep soil layers. Further research is needed to evaluate their functional roles in these often-dry subsoils. Our results provide important baseline data for how microorganisms are distributed within soils, demonstrating the potential for increased phylogenetic novelty in relation to soil depth. Additional investigations surveying deeper soils may reveal novel organisms that are responsible for key functions in subsurface soils (Fierer, 2017; Baldrian, 2019). By focusing on deep soil layers in particular, future studies could provide insight into unique and potentially significant processes affecting drought-prone forest ecosystems.

DATA AVAILABILITY STATEMENT

The datasets presented in this study can be found in online repositories. The names of the repository/repositories and accession number(s) can be found in the article/Supplementary Material.

AUTHOR CONTRIBUTIONS

BF and IB designed the microbial study. LW established and maintained the six forest sites within the TreeNet network. LW, RK, and IB provided soil, fine root, and climatic data. BS performed genetic analyses in the lab. BF, CP-M, AD, and IB performed statistical analyses. BF and IB wrote the main parts of the manuscript. All authors contributed to the final version of the manuscript.

FUNDING

We thank the Swiss Federal Office for the Environment (FOEN) for partly funding the logging infrastructure and maintenance costs for TreeNet (Grants 00.0365.PZI0427-0562 and 09.0064.PJ/R301-0223).

ACKNOWLEDGMENTS

We acknowledge Rafael Boduryan for processing soils and roots, the WSL Central Laboratory for soil chemical analyses, and Stephan Zimmermann for providing data from the WSL soil database. We also acknowledge the Genetic Diversity Centre (GDC) of the ETH Zurich and scientists at the McGill University and Génome Québec Innovation Centre in Montréal, Canada, for performing Illumina MiSeq sequencing. The six forest sites are part of TreeNet (https://treenet.info/), a biological drought and growth indicator network in Switzerland. Evaluations from the European beech site Neunkirch were based on data from the Swiss Long-term Forest Ecosystem Research Program LWF (www.lwf.ch), which is part of the UNECE Co-operative Program on Assessment and Monitoring of Air Pollution Effects on Forests ICP Forests (www.icp-forests.net). We thank Melissa Dawes for help in editing the manuscript and improving the English language.

SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: https://www.frontiersin.org/articles/10.3389/fmicb.2021.674160/full#supplementary-material.

Supplementary Data Sheet 1 | Supplementary Tables 1–4.
Supplementary Data Sheet 2 | Supplementary Figures 1–6.
Supplementary Data Sheet 3 | Bacterial OTU, Tax table.
Supplementary Data Sheet 4 | Fungal OTU, Tax table.
Supplementary Data Sheet 5 | Fungal guild classification, Table.
References

Abarenkov, K., Henrik Nilsson, R., Larsson, K., Alexander, I. J., Eberhardt, U., Erland, S., et al. (2010). The UNITE database for molecular identification of fungi—recent updates and future perspectives. New Phytol. 186, 281–285. doi: 10.1111/j.1469-8137.2009.03160

Adamczyk, B., Sietio, O.-M., Straková, P., Prommer, J., Wild, B., Hagner, M., et al. (2019a). Plant roots increase both decomposition and stable organic matter formation in boreal forest soil. Nat. Comm. 10:3982. doi: 10.1038/s41467-019-11993-1

Adamczyk, M., Hagedorn, F., Wipf, S., Donhauser, J., Víttoz, P., Rixen, C., et al. (2019b). The soil microbiome of Gloria mountain summits in the Swiss Alps. Front. Microbiol. 10:1080. doi: 10.3389/fmicb.2019.01080

Bardgett, R. D., and van der Putten, W. H. (2014). Belowground biodiversity and ecosystem functioning. Nature 515, 505–511. doi: 10.1038/nature13855

Bennett, A. C., Murugappiran, S. K., and Hamilton, T. L. (2020). Temperature impacts community structure and function of phototrophic Chloroflexi and Cyanobacteria in two alkaline hot springs in Yellowstone. Environ. Microbiol. Rep. 12, 503–513. doi: 10.1111/1758-2229.12863

Bodelier, P. L. E., and Steenbergh, A. K. (2014). Interactions between methane and the nitrogen cycle in light of climate change. Curr. Opin. Environ. Sustain. 9–10, 26–36. doi: 10.1016/j.cosust.2014.07.004

Brewer, T. E., Aronson, E. L., Arogyaswamy, K., Billings, S. A., Botthoff, J. K., Campbell, A. N., et al. (2019). Ecological and genomic attributes of novel bacterial taxa that thrive in subsurface soil horizons. mBio 10, e01318-e01319. doi: 10.1128/mBio.01318-19

Cabello, P., Roldan, M. D., and Moreno-Vivian, C. (2004). Nitrate reduction and the nitrogen cycle in archaea. Microbiology 150, 3527–3546. doi: 10.1099/mic.0.27303-0

Carteron, A., Beigas, M., Joly, S., Turner, B. L., and Laliberté, E. (2020). Temperate forests dominated by arbuscular or ectomycorrhizal fungi are characterized by strong shifts from saprotrophic to mycorrhizal fungi with increasing soil depth. Microb. Ecol. doi: 10.1007/s00248-020-01540-7

Castlebury, L. A., Rossman, A. Y., and Hyten, A. S. (2006). Phylogenetic relationships of Chaetomium and related genera. Fungal Biol. 120, 1291–1322. doi: 10.1016/j.funbio.2015.09.013

Crowe, J. H. (2014). Anhydrobiosis: an unsolved problem. Plant Cell Environ. 37, 1491–1493. doi: 10.1111/pce.12304

DeBruyn, J. M., Nixon, L. T., Fawaz, M. N., Johnson, A. M., and Radosevich, M. (2011). Global biogeography and quantitative seasonal dynamics of mycorrhizal and ecosystem functioning. New Phytol. 189, 629–630. doi: 10.1111/j.1469-8137.2010.03168

Deng, Y. E., Ning, D., Qin, Y., Xue, K., Wu, L., He, Z., et al. (2018). Spatial scaling in microbial community structure are associated with increased litter input rather than temperature in High Arctic soils. Soil Biol. Biochem. 151:108054. doi: 10.1016/j.soilbio.2020.108054

Dreyfuss, M. (1976). Taxonomische Untersuchungen innerhalb der Gattung Chaetomium. Sydowia 28, 50–133.

Du, C., Geng, Z., Wang, Q., Zhang, T., He, W., Hou, L., et al. (2017). Variations in bacterial and fungal communities through soil depth profiles in a Betula albosinensis forest. J. Microbiol. 55, 684–693. doi: 10.1007/s12725-017-6466-8

Du, X., Deng, Y., Li, S., Esclais, A., Feng, K., He, Q., et al. (2020). Steeper spatial scaling patterns of soil microbial biota are shaped by deterministic assembly processes. Mol. Ecol. 30, 1072–1085. doi: 10.1111/mec.17577

Dunbar, J., Barns, S. M., Ticknor, L. O., and Kuske, C. R. (2002). Empirical and theoretical bacterial diversity in four Arizona soils. Appl. Environ. Microbiol. 68, 3035–3045. doi: 10.1128/AEM.68.6.3035-3045.2002

Edgar, R. C. (2010). Search and clustering orders of magnitude faster than BLAST. Bioinformatics 26, 2460–2461.

Edgar, R. C. (2013). UPARSE: highly accurate OTU sequences from microbial amplicon reads. Nat. Methods 10, 996–998.

Edgar, R. C., and Flyvbjerg, H. (2015). Error filtering, pair assembly and error correction for next-generation sequencing reads. Bioinformatics 31, 3476–3482. doi: 10.1093/bioinformatics/btv401

Frey, B., Niklaus, P. A., Kremer, J., Lüscher, P., and Zimmermann, S. (2011). Heavy-rotten soils: microbial community structure in oxic forest soils. Appl. Environ. Microbiol. 77, 6060–6068. doi: 10.1128/AEM.05206-11

Frey, B., Rimé, T., Phillips, M., Sterli, B., Haidas, I., Widmer, F., et al. (2016). Microbial diversity in European alpine permafrost and active layers. FEMS Microbiol. Ecol. 92:fw018. doi: 10.1093/femsdc/fw018

Gadgil, R. L., and Gadgil, P. D. (1971). Mycorrhiza and litter decomposition. Nature 233:133. doi: 10.1038/233133a0

Gadgil, R. L., and Gadgil, P. D. (1975). Suppression of litter decomposition by mycorrhizal roots of Pinus radiata. N. Z. J. For. Sci. 5, 35–41.

Ge, G. W., and Bauder, J. W. (1986). “Particle Size Analysis,” in Methods of Soil Analysis, Part I, Chemical and Physical Methods, Vol. 9, ed. A. Klute (Madison WI: American Society of Agronomy), 383–411.

Gersum, A., Lacaun, J.-P., Robin, A., and Jouyard, C. (2020). Tamm review: deep fine roots in forest ecosystems: why dig deeper? For. Ecol. Manage. 466:118135. doi: 10.1016/j.foreco.2020.118135

Gocke, M. I., Huguet, A., Derenne, S., Kolb, S., Dippold, M. A., and Wiesenberg, G. L. (2017). Disentangling interactions between microbial communities and roots in deep subsoil. Sci. Total Environ. 575, 135–145. doi: 10.1016/j.scitotenv.2016.09.184

Goldmann, K., Schröter, K., Pena, R., Schöning, I., Schrumpf, M., Buscot, F., et al. (2016). Divergent habitat filtering of root and soil fungal communities in temperate beech forests. Sci. Rep. 6:31439. doi: 10.1038/srep31439

Gremer, J. R., Bradford, J. B., Munson, S. M., and Duniway, M. C. (2015). Desert grassland responses to climate and soil moisture suggest divergent vulnerabilities across the southwestern United States. Glob. Chang. Biol. 21, 4049–4062. doi: 10.1111/gcb.12043
Hanada, S., Takaichi, S., Matsuura, K., and Nakamura, K. (2002). *Roseiflexus castenholzii* gen. nov., sp. nov., a thermophilic, filamentous, photosynthetic bacterium that lacks chlorosomes. *Int. J. Syst. Evol. Microbiol.* 52, 187–193. doi: 10.1099/00227773-52-1-187

Hanewinkel, M., Cullmann, D. A., Schelhas, M. J., Nabnurs, G. J., and Zimmermann, N. E. (2013). Climate change may cause severe loss in the economic value of European forest land. *Nat. Clim. Change* 3, 203–207. doi: 10.1038/nclimate1687

Hansel, C. M., Fendosh, S., Jardine, P. M., and Francis, C. A. (2008). Changes in bacterial and archaeal community structure and functional diversity along a geochemically variable soil profile. *Appl. Environ. Microbiol.* 74, 1620–1633. doi: 10.1128/AEM.01787-07

Hao, J., Chai, Y. N., Shi, Y., Imelfort, M., Keller, J., Hugenholtz, P., et al. (2013). Hao, J., Chai, Y. N., Lopes, L. D., Ordóñez, R. A., Wright, E. E., Archontoulis, S., Hanewinkel, M., Cullmann, D. A., Schelhaas, M. J., Nabuurs, G. J., and Manzoni, S., and Katul, G. (2014). Invariant soil water potential at zero microbial respiration explained by hydrological discontinuity in dry soils. *Geophys. Res. Lett.* 41, 7151–7158. doi: 10.1002/2014GL061467

McDowell, N. G., Allen, C. D., Anderson-Texeira, K., Aukema, B. H., Bond-Lamberty, B., Chini, L., et al. (2020). Pervasive shifts in forest dynamics in a changing world. *Science* 366, eaaz9463. doi: 10.1126/science.aaz9463

McMurtrie, P. J., and Holmes, S. (2013). Phyloseq: an R package for reproducible interactive analysis and graphics of microbiome census data. *PloS One* 8:e61217. doi: 10.1371/journal.pone.0061217

Meier, I. C., and Leuschner, C. (2008). Belowground drought response of European beech: fine root biomass and carbon partitioning in 14 mature stands across a precipitation gradient. *Glob. Change Biol.* 14, 2081–2095. doi: 10.1111/j.1365-2486.2008.01634.x

Mikkonen, A., Santalahi, M., Lappi, K., Pulkkinen, A.-M., Montonen, L., and Suominen, L. (2014). Bacterial and archaeal communities in long-term contaminated surface and subsurface soil evaluated through coexpressed RNA and DNA. *FEMS Microbiol. Ecol.* 90, 103–114. doi: 10.1111/1574-6941.12376

Nguyen, N. H., Song, Z. W., Bates, S. T., Branco, S., Tedersoo, L., Menke, J., et al. (2016). FUNGuild: an open annotation tool for parsing fungal community datasets by ecological guild. *Fungal Ecol.* 20, 241–248.

Oertel, C., Matschullat, J., Zurbia, K., Zimmermann, F., and Erasmis, S. (2016). Microbial communities: the soil’s memory for the past. *Prog. Soil Sci.* 185, 566–575. doi: 10.1007/00468-004-0340-0

Orth, R., Zscheischler, J., and Seneviratne, S. I. (2016). Record dry summer in 2015 challenges precipitation projections in Central Europe. *Sci. Rep.* 6:28334.

Perez-Mon, C., Frey, B., and Frossard, A. (2020). Functional and structural responses of Arctic and alpine soil prokaryotic and fungal communities under freeze-thaw cycles of different frequencies. *Front. Microbiol.* 11:982. doi: 10.3389/fmicb.2020.00982

Pickles, B. J., and Pither, J. (2014). Still scratching the surface: how much of the ‘black box’ of soil ectomycorrhizal communities remains in the dark? *New Phytol.* 201, 1101–1105. doi: 10.1111/nph.12616

Quast, C., Pruesse, E., Yilmaz, P., Gerken, J., Schweer, T., Yarza, P., et al. (2013). The SILVA ribosomal RNA gene database project: improved data processing and web-based tools. *Nucleic Acids Res.* 45, D590–D596. doi: 10.1093/nar/gks1219

R Core Team (2020). R: A Language and Environment for Statistical Computing. Vienna: R Foundation for Statistical Computing.

Rasse, D. P., Rumpel, C., and Dignac, H.-F. (2005). Is soil carbon mostly root carbon? Mechanisms for a specific stabilisation. *Plant Soil* 269, 341–356. doi: 10.1007/s11104-004-0907-y

Remund, J., Rihm, B., and Huguenin-Landi, B. (2014). Klimadaten für die Waldmodellierung für das 20. und 21. Jahrhundert. Birmensdorf: Eidg. Forschungsanstalt für Wald, Schnee und Landschaft WSL. doi: 10.3929/ethz-a-019395731

Rime, T., Hartmann, M., Brunner, I., Widmer, F., Zeyer, J., and Frey, B. (2015). Vertical stratification of subsurface microbial community composition across geological formations at the Hanford Site. *Environ. Microbiol.* 16, 414–425. doi: 10.1111/1462-2920.12659.x

Rime, T., Hartmann, M., Brunner, I., Widmer, F., Zeyer, J., and Frey, B. (2015). Vertical distribution of the soil microbiota along a successional gradient in a glacier forefield. *Mol. Ecol.* 24, 1091–1108. doi: 10.1111/mec.13051

Rime, T., Hartmann, M., and Frey, B. (2016). Potential sources of microbial colonizers in an initial soil ecosystem after retreat of an alpine glacier. *ISME J.* 10, 1625–1641. doi: 10.1038/ismej.2015.238

Robin, A., Pradier, C., Sanguin, H., Salhi, F., Lambais, G. R., de Araujo Pereira, A. P., et al. (2019). How deep can ectomycorrhizas go? A case study on *Pisolithus*...
down to 4 meters in a Brazilian eucalypt plantation. *Mycorrhiza* 29, 637–648. doi: 10.1007/s00572-019-00917-y

Rousk, J., Bäath, E., Brookes, P. C., Lauber, C. L., Lozupone, C., Caporaso, J. G., et al. (2010). Soil bacterial and fungal communities across a pH gradient in an arable soil. ISME J. 4, 1340–1351.

Rousk, J., and Bengtsson, P. (2014). Microbial regulation of global biogeochemical cycles. *Front. Microbiol*. 5:103. doi: 10.3389/fmicb.2014.00103

Rümpl, C., and Kogel-Knabner, I. (2011). Deep soil organic matter—a key but poorly understood component of terrestrial C cycle. *Plant Soil* 338, 143–158. doi: 10.1007/s11104-010-0391-5

Rüthi, J., Bölsterli, D., Pardi-Comensoli, L., Brunner, I., and Frey, B. (2020). The “plastisphere” of biodegradable plastics is characterized by specific microbial taxa of Alpine and Arctic soils. *Front. Environ. Sci*. 8:562263. doi: 10.3389/fenvs.2020.562263

Schappe, T., Albornoz, F. E., Turner, B. L., and Jones, F. A. (2020). Co-occurring fungal functional groups respond differently to tree neighborhoods and soil properties across three tropical rainforests in Panama. *Microb. Ecol*. 79, 675–685. doi: 10.1007/s00248-019-01446-z

Schimel, J. P. (2018). Life in dry soils: effects of drought on soil microbial communities and processes. *Annu. Rev. Ecol. Evol. Syst.* 49, 409–432. doi: 10.1146/annurev-ecolsys-110617-062614

Schindler, S., Will, C., Thurmer, A., Wollherr, A., Nacke, H., Herold, N., Schrumpf, M., Will, C., Thurmer, A., Wollherr, A., Nacke, H., Herold, N., Schrumpf, M., et al. (2017). From the comfort zone to crown dieback: sequence of physiological stress thresholds in mature European beech trees across progressive drought. *Sci. Total Environ*. 753:141792. doi: 10.1016/j.scitotenv.2020.141792

Walgren, H., Graf, U., Kammer, A., Luster, J., Pezzotta, D., Zimmermann, S., et al. (2010). Determination of organic and inorganic carbon, 813C, and nitrogen in soils containing carbonates after acid fumigation with HCl. *J. Soil Sci. Plant Nutr*. 173, 207–216. doi: 10.1002/jpln.200900158

Walthert, L., and Schleppi, P. (2018). Equations to compensate for the temperature effect on readings from dielectric Decagon MPS-2 and MPS-6 water potential sensors in soils. *J. Soil Sci. Plant Nutr*. 181, 619–759. doi: 10.1002/jpln.201700620

Wang, Q., Garrity, G. M., Tiedje, J. M., and Cole, J. R. (2007). Naive Bayesian classifier for rapid assignment of RNA sequences into the new bacterial taxonomy. *Appl. Environ. Microbiol*. 73, 5261–5267. doi: 10.1128/AEM.00620-07

Wang, X.-W., Wang, X.-L., Liu, F.-J., Zhao, X.-M., Li, J., and Cai, L. (2014). Phylogenetic assessment of *Chaetomium indicum* and allied species, with the introduction of three new species and epitypification of *C. funicola* and *C. indicum*. *Mycol. Proc*. 13, 657–713. doi: 10.1017/s00572-013-0955-x

Wild, C., Thurmer, A., Wollherr, A., Naeke, H., Herold, N., Schrumpf, M., et al. (2010). Horizon-specific bacterial community composition of German grassland soils, as revealed by pyrosequencing-based analysis of 16S rRNA genes. *Appl. Environ. Microbiol*. 76, 6751–6759. doi: 10.1128/AEM.01063-10

Wu, H., Adams, J. M., Shi, Y. U., Li, Y., Song, X., Zhao, X., et al. (2020). Depth-dependent patterns of bacterial communities and assembly processes in a typical red soil critical zone. *Geomicrobiol. J.* 37, 201–212. doi: 10.1080/01490451.2019.1688432

Yang, H., Yang, F., Gao, Z., Peng, Z., and Chen, W. (2019). Subsoil microbial community responses to air exposure and legume growth depend on soil properties across different depths. *Sci. Rep.* 9:18536. doi: 10.1038/s41598-019-55089-8

Zhang, C., Li, J., Wang, J., Liu, G., Wang, G., Guo, L., et al. (2019). Decreased pigment formation by lignicolous fungi. *Environ. Microbiol*. 21, 5985–6002. doi: 10.1111/1462-2920.1490451.2019.1688432

Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Copyright © 2021 Frey, Walthert, Perez-Mon, Sterl, Köchli, Dharmarajah and Brunner. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner(s) are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.