Next-generation ice core technology reveals true minimum natural levels of lead (Pb) in the atmosphere: Insights from the Black Death

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Abstract Contrary to widespread assumptions, next-generation high (annual to multiannual) and ultra-high (subannual) resolution analyses of an Alpine glacier reveal that true historical minimum natural levels of lead in the atmosphere occurred only once in the last ~2000 years. During the Black Death pandemic, demographic and economic collapse interrupted metal production and atmospheric lead dropped to undetectable levels. This finding challenges current government and industry understanding of preindustrial lead pollution and its potential implications for human health of children and adults worldwide. Available technology and geographic location have limited previous ice core investigations. We provide new high- (discrete, inductively coupled plasma mass spectrometry, ICP-MS) and ultra-high resolution (laser ablation inductively coupled plasma mass spectrometry, LA-ICP-MS) records of atmospheric lead deposition extracted from the high Alpine glacier Colle Gnifetti, in the Swiss-Italian Alps. We show that contrary to the conventional wisdom, low levels at or approaching natural background occurred only in a single 4 year period in ~2000 years documented in the new ice core, during the Black Death (~1349–1353 C.E.), the most devastating pandemic in Eurasian history. Ultra-high chronological resolution allows for the first time detailed and decisive comparison of the new glaciochemical data with historical records. Historical evidence shows that mining activity ceased upwind of the core site from ~1349 to 1353, while concurrently on the glacier lead levels have been grossly underestimated, with significant implications for public health and environmental policy, and the history of human exposure to lead. This trans-disciplinary study is a collaboration between the Initiative for the Science of the Human Past at Harvard University and the Climate Change Institute at the University of Maine. It uses next-generation technology and expertise in history, climate science, archaeology, and toxicology, brought to bear in a highly detailed contribution to planetary health, with crucial implications for public health and environmental policy, and the history of human exposure to lead.
1. Introduction

Although scientists and modern historians have documented the devastating effects of lead (Pb) poisoning on humans during the past 2000 years (at the very least) [Hernberg, 2000; Niirjugu, 1983], the extent of population exposure to elevated atmospheric lead levels remains unclear. Despite mitigation and public health measures aimed at reducing human exposure in occupational and residential environments, lead remains a major threat to public health worldwide [Mushak, 2011]. The effects of even minimal human exposure include mental deficiencies [Hernberg, 2000; Lanphear et al., 2005], reduced fertility [Mushak, 2011; Selevan et al., 2003; Chang et al., 2006; De Rosa et al., 2003], and increased aggressive behavior [Mielke and Zahran, 2012; Reyes, 2015]. These symptoms have been observed even at low levels of Pb blood concentration, especially in children [Hernberg, 2000]. Atmospheric lead pollution is both a cause of higher levels of Pb in humans and a proxy for higher concentration of aerosol Pb. Historically, Pb has been mined and smelted (along with silver) and used widely in coinage, water pipes, roofs, and, more recently, as an additive in paint Pb in humans and a proxy for higher concentration of aerosol Pb. Historically, Pb has been mined and smelted (along with silver) and used widely in coinage, water pipes, roofs, and, more recently, as an additive in paint [Hernberg, 2000]. Government and industry standards continue to overestimate the proportion of natural lead (Pb) levels in the environment [United Nations Environment Programme, 2010; Richardson et al., 2001]. Our high- and ultra-high-resolution continuous measurements substantiate and expand upon previously published, pioneering but lower-resolution ice core studies and those from lake sediments and peat cores that suggest a steady increase in Pb levels across western Europe from ~1250 to 900 B.C.E. to the present, with periods of only moderate decline [Hong et al., 1994; Renberg et al., 2001; Shotyk et al., 1998; Le Roux et al., 2004; Martínez Cortizas et al., 2013; Montgomery et al., 2010; Gabrieli and Barbante, 2014; Brännvall et al., 1999].

2. Data: A New High- and Ultra-High Resolution Record of Lead (Pb) Deposition From the Heart of Europe

In this study we provide a new atmospheric Pb deposition record from a ~72 m ice core extracted from the Colle Gnifetti (CG) glacier (4450 m above sea level) in the Swiss-Italian Alps. A discrete, high-resolution inductively coupled plasma mass spectrometry (ICP-MS) Pb record (Figure 1) covers the last ~2000 years; an additional ultra-high-resolution laser ablation (LA)-ICP-MS record provides more detailed evidence of subannual Pb deposition for the years ~1330–1360 C.E.

Ultra-high resolution sampling of this ice core (~120 μm, allowing ~550 measurements within the year dated ~1300 C.E.) was produced using the Climate Change Institute’s (CCI at the University of Maine) W. M. Keck Laser Ice Facility laser ablation inductively coupled plasma mass spectrometer (LA-ICP-MS) [Sneed et al., 2015]. This new method allowed us to count highly thinned annual layers previously not detectable by conventional cm-resolution analyses. The ultra-high resolution time series permitted us to apply the layer-counting procedure down to the beginning of the first millennium of the Common Era. Back to 1900 C.E., known time markers such as documented Saharan dust events were used to constrain the chronology of the ice core, as already demonstrated in similar Alpine cores [Gabrieli and Barbante, 2014; Bohleber et al., 2013; Jenk et al., 2009; Schwikowski et al., 2004; Eisen et al., 2003; Wagenbach and Geis, 1989]. For the most recent ~800 years, the resulting time scale was further corroborated by direct time series comparisons with a neighboring CG ice core dated with conventional cm-resolution analysis [Bohleber et al., 2013]. The time scale for the layers dated for years before this period is currently under development using our ultra-high resolution technique. Figure 2 shows an example of annual layer counting for the period ~1310 C.E. to 1317 C.E., illustrating seasonal variability in dust-source Ca. Figure 2 (as well as Figure S5) presents the raw data (red) and a smoothed line as a visual aid (black). In the time range between ~500 C.E. and 1500 C.E. the ultra-high resolution annual layer counting data are also backed up by 14C ages, retrieved from the analysis of the particulate organic carbon fraction [Hoffmann, 2016]. These 14C data were developed completely independently from the layer counting. Comparison reveals very good agreement with the annual layer counting within a 1σ error range (Figure S1 in the supporting information).

It is important to note that all annual layers counting for the CG ice core were completed prior to comparison with written historical evidence collected and analyzed by the Initiative for the Science of the Human Past at Harvard. The independently developed subannual resolution record derived from the CG ice core thus allowed testing against subannually resolved historical records. Sources in Latin, Middle English, English, French, German, Dutch and Italian provided subannual dates (months, year) for the arrival and spread of the Black Death throughout Europe, as well as decadal to subannual trends in mining and smelting activity.
from the last two millennia. Extensive archaeological and historiographical evidence corroborated our conclusions and time scale (Tables S2–S5 in the supporting information).

Prior to ultra-high-resolution LA-ICP-MS analysis, high-resolution ICP-MS discrete analysis (~4.27 cm average sample resolution over the ~2000 year record), also conducted by CCI, independently revealed a dramatic drop in atmospheric Pb levels falling exactly within the period of the Black Death (1349–1353 C.E.), the greatest pandemic to ravage Eurasia in recorded history. Previous studies of atmospheric lead in low-resolution ice core records available for the last two millennia did not document this same, sharp, multiyear decline to undetectable levels [Hong et al., 1994; Gabrieli and Barbante, 2014] (Figure 3). Potential uncertainty in our layer-counted depth-age time scale was initially estimated to be less than 35 years at this interval, based on the lag generated by comparing our CG time series and previous annually dated CG time series [Bohleber et al., 2013]. Further, we found that our Pb deposition record was in good agreement with shorter, multiyear resolution CG ice core records for the period ~1650–2000 C.E. [Schwikowski et al., 2004; Gabrieli and Barbante, 2014].

3. Results: Consilience of Highly Detailed Historical Evidence and the Glaciochemical Record

Remarkably, the drop in Pb concentration, captured by both the discrete ICP-MS and the continuous LA-ICP-MS methods, coincides with written historical evidence of the effects of the Black Death pandemic on European populations and metal production and parallels data on similar downward trends in atmospheric CO₂ levels in the same period, due to population decline [van Hoof et al., 2006]. The coincidence of the two
independently derived time series (ice core and written record) and in particular the unique nature of the Pb drop in the ice core record at this confluence confirms the ice core dating of this event. The discrete Pb levels corresponding to the layers counted as years 1349–53 C.E. are the lowest in our record and are much lower than levels documented in even the deepest CG layers, indicating that for at least the past two millennia human mining and smelting activities have been the originator of detectable lead pollution in the European continent.

Our findings are in sharp contrast with a consensus among policy makers and industry experts that ascribes a significant portion of preindustrial atmospheric lead levels to natural, e.g., crustal or volcanic sources [United Nations Environment Programme, 2010; Richardson et al., 2001]. The new measurements indicate that this consensus overestimates the contribution of such natural sources to current lead levels in the atmosphere. The location of the CG ice core in the heart of Europe provides a geographically specific signal. Whereas, for example, the first polar ice core detections of historic metal pollution were unable to distinguish clearly Roman and Chinese Empires’ production areas, the new CG ice core’s location is relatively close to the mining and smelting centers of western Europe from the historical beginnings of smelting activities to the present. The long-range transport necessary for heavy lead particles to reach and be trapped in polar ice is more difficult to interpret than the shorter distances between source and the Alpine core. Therefore, while long-range transport is necessary for heavy lead particles to be trapped in polar ice, the proximity of potential Pb sources offers in the Alpine core a more precise, definitive, continuous, and regionally specific signal.

Historical records show that massive mortalities in the spring and summer of 1349 C.E. halted metal production in all the major Pb-producing regions of Western Europe (Figure 4 and Table 1). During the pandemic, 30–50% of the European population died [DeWitte and Wood, 2008]. Extensive archaeological
investigation has recently a 45% mortality rate in Eastern England [Lewis, 2016], principally due to bubonic plague (*Yersinia pestis*), now definitively identified by genome sequencing [Bos et al., 2011]. Throughout its ~2000 year record, the CG ice core shows levels of Pb significantly higher than those recorded during the Black Death.

The lowest Pb levels recorded in our study occurred during the Black Death (0.4 ng/L at 1353 C.E. in the high-resolution discrete ICP-MS and below the limit of detection at 1351 C.E. in the ultra-high resolution LA-ICP-MS) and likely represent dispersal of Pb from the Earth’s crust, that is, as close to natural background Pb levels as were achieved in the full ~2000 year record. The new measurements significantly alter our understanding of atmospheric Pb pollution hitherto labeled as natural background and therefore assumed to be safe. Thus, they challenge the assumption that preindustrial atmospheric Pb levels had no discernible effect on human physiology. These new data show that human activity has polluted European air almost uninterruptedly for the last ~2000 years. Only a devastating collapse in population and economic activity caused by pandemic disease reduced atmospheric pollution to what can now more accurately be termed “background” or natural

**Figure 5.** Lead concentration in CG ice core, from ultra-high-resolution LA-ICP-MS, 1330–1360 C.E. (with an average of 279 measurements per year in 1349–1353). The grey histogram represents declining number of active major mining regions as they were progressively hit by the plague and ceased operations; the red histogram represents the number of mining regions resuming metal production, based on written sources. At present, there are no estimates of volume of aggregate metal production, and thus, the histograms reflect only regions that were active, not volume of Pb produced. The values below 1ng/L here are calculated using semiquantitative calibration data. Smoothing (black line) is provided only as a visual aid, while the red plot presents the raw data. As shown in the methods section in the supporting information, the LA-ICP-MS technique [Sneed et al., 2015] measures total element concentration; the spikes can thus be related to individual particles and/or storm event concentrations.

| City/Region | Black Death Arrives | Year Pb/Ag Mining Ceases |
|-------------|---------------------|--------------------------|
| Britain     |                     |                          |
| Mendip      | 1348/1349           | 1340s                    |
| Devon       | March 1349          | 1349                     |
| Flintshire  | April–June 1349     | 1349–1350                |
| Derbyshire (Peak) | May 1349   | 1349–1352                |
| York        | May 1349            |                          |
| Harz (Goslar) |             | 1350                     |
| Harz (Halberstadt) | May 1350 |                          |
| Magdeburg   | May 1350            |                          |

For details, see Tables S2–S5. Dates adjusted to modern calendar, whenever appropriate (see supporting information for details).
levels. Pb crustal enrichment factors (EF_c, Figure 3, see also Figure S4 for potential volcanic EF influences and further discussion) evaluating the extent of anthropogenic soil contamination for the years corresponding to the Black Death corroborate our interpretation. They show a marked decline, reaching a value of 2.82 in the year 1352, the second lowest in the entire record (the past ~2000 years) with the lowest EF_c occurring in 1366, with a corresponding value of 2.36. The latter date corresponds very closely to the date range of a further plague pandemic between 1367 and 1369, the impact of which is dramatically documented in the Halesowen manorial court rolls in the West Midlands of England [Razi, 1980].

In the Alpine region of Europe, high-level regional delivery and lower-level atmospheric circulation transport pollutants [Schwikowski et al., 2004; Gabrieli and Barbante, 2014]. Modern atmospheric circulation patterns associated with the Azores High (Figures 4 and S2) point to potential British, French, and German sources of pollution transported to CG. Our record of the multiyear Black Death period is not associated with any anomalous atmospheric circulation patterns, based on Ca and Fe as crustal air mass proxies (Figure S3). Comparison with historical evidence from SoHP’s geodatabase of climate events also presented no substantial change in observed climate patterns in the region at the time (Data Set S1 in the supporting information). This leaves a dramatic decline, if not complete interruption of anthropogenic emissions of Pb at the time of the Black Death as the most likely dominant control, especially in light of documented Pb residence times in the troposphere, averaging a week to ten days [Papastefanou, 2006].

Historical documentary evidence—fiscal, legal, and chronicle sources—shows that while Pb production in the Harz Mountains was already in severe decline in the 1330s, Britain dominated western European Pb production until the plague reached its regions of most labor-intensive mining between January and September 1349 C.E. (Tables 1 and S2–S5). We argue that British mines and smelting sites were the likely dominant source of Pb captured in the CG ice core at the time of the 1349–1353 C.E. collapse, since they were by far the principal producers in this period. Extensive and large-scale mining and smelting were largely constrained within the principal British lead producing regions by 1348, such as the High Peak District region of Derbyshire [Blanchard, 2005] (Table S5), the Bere Ferrers mine in Devon [Claufton, 2010], and, to a lesser extent at that time, in the Yorkshire Dales and the hills of Shropshire and Flintshire [Claufton et al., 2016]. Coincident locations of both galena ore sources and woodland for fuel were the key factors governing the largest regional concentrations of these activities in Britain. The movement of the raw ore of metals, such as iron by water, is attested archaeologically when coastal waters and shipping were immediately available, indicated from the mid-13th century Magor Pill ship from the Welsh shore of the Bristol Channel [Claufton et al., 2016; Nayling, 1998]. Dressed galena ore is recorded as having been paid by miners in the Peak District in the form of renders to local landowners for smelting, usually to the king or major aristocrats, from the 12th century onward, but evidence of the movement of galena for smelting outside the Peak or other principal mining regions is currently lacking. Lead is only attested textually and archaeologically as having been moved inter-regionally and over long distances in its smelted form between the 9th and 14th centuries, over land and by water, as ingots or sheet [Rieuwerts, 1988; Allen, 2011; Kelly and Brooks, 2013]. The constraint of lead production to paramount mining and smelting regions in England is further demonstrated by specific traits within their regional economies, for example, the payment of rents and tithes to local landowners in dressed ore or smelted lead and the use of the metal as a medium of barter exchange [Rieuwerts, 1988; Barnatt and Smith, 2004; Blanchard, 2005] (Table S5). Other potential non-British sources of pollution, such as the French mines and woodland smelting sites west of CG, at Mont-Lozère, had already ceased activities by 1280 C.E. [Baron et al., 2006]. Sardinia, the most significant Mediterranean Pb producer, was in deep decline already in the 1330s; moreover, the island lies outside the dominant atmospheric transportation pattern and had already been ravaged by plague in 1348 C.E. (Figure 4 and Tables S2–S5).

The ultra-high-resolution CG data (Figure 5) show a steep progressive decline in Pb deposition, from layers dated 1349 C.E. to 1352 C.E., corresponding to the progression of the pandemic through different lead-producing areas. The arrival of the plague in the most productive British mining regions in the second half of 1349 C.E. corresponds to subannual LA-ICP-MS data points showing the beginning of the most severe drop in Pb concentration in the ice core. Table 1 summarizes the dates when the plague reached the British, German, and Italian mining regions and when mining and smelting operations were interrupted. Mining resumed sporadically and progressively from 1352 C.E. (Table S4) when some of the pre-plague mining sites reopened in Britain (High Peak District in Derbyshire) along with new mines (North Yorkshire), but production
levels fluctuated for a century due to the more limited demands of a population reduced by ~50%. There is no evidence of new mining or smelting in Sardinia until ~1420 C.E. nor in the Harz until the 1460s C.E. [Blanchard, 2005; Dyer, 2000].

In the high-resolution discrete CG Pb data (Figure 1), a second severe drop corresponds to the period 1460–65 C.E. Historical records show that British mining activities declined drastically at this time due to market oversupply, probably linked to another series of epidemics that affected Britain, as well as lower demand due to an economic downturn (Tables S2–S4) [Blanchard, 2005; Dyer, 2000; Nightingale, 2005; Gottfried, 1977; Hatcher, 2003; Creighton, 1891]. Resurgence of Harz mining activities in the 1460s [Bartels, 2010] is not detected at CG, suggesting that German mines were either not a major contributor to Pb deposition at CG at that time or that their emissions from smelting were relatively low. Pb crustal enrichment factors also reflect this second decline (Figure 3).

The third lowest level of Pb deposition in the discrete ice core record corresponds to the year 1885. Mining activities slumped in that year due to the long-term economic collapse that affected Western countries in 1882–1885 [Brayshaw, 1980]. A similar trend is observed in the United States in 1885, the year in which Pb production levels declined most severely in extant historical records dating back to the late 18th century [Mushak, 2011; Brayshaw, 1980; U.S. Geological Service, 2013]. The most recent decrease in atmospheric Pb levels in Europe began in 1974. This decline reflects legislative efforts to phase out leaded fuel in Western countries, which resulted in decreased blood levels of Pb throughout Europe and the United States [Schwikowski et al., 2004; Strömberg et al., 2008; Gabrielli and Barbante, 2014].

4. Conclusions

Ultra-high-resolution measurements from the heart of Europe, combined with a densely documented historical and archaeological archive, usher in a new era in the detailed reconstruction of human interaction with the environment. Anticipating the forthcoming reduction of dating uncertainty in the deepest ice core sections, the examination of pre-Black Death Pb deposition levels (Figure 1) points to intriguing areas of future research such as Europe’s shift from gold to silver coinage with the opening of new Ag/Pb mines in France (Melle), between 640 and 680 C.E. Similarly, our new measurements of Pb deposition suggest that Europe’s booming metal production ~1180–1220 C.E. (the highest preindustrial Pb peak in our record) may have generated pollution levels rivaling those ~1650 C.E. Since previous research has correlated deposition levels to volume of emissions of sulfur, copper, uranium, arsenic, and lead in earlier ice cores, for example [Schwikowski et al., 2004; Mayewski et al., 1986], we expect that future research will elucidate whether deposition levels captured by the new ultra-high-resolution method can be correlated more precisely and quantitatively with historical volume of emissions and, potentially, of production levels. Our study also points to the need to explore possible connections between historic atmospheric Pb pollution and ecosystem health, including human fertility, intelligence, and behavior. Such trans-disciplinary research will represent a significant contribution to the field of planetary health, in line with the aims outlined in Almada et al. [2017].

In this paper we have mobilized more than a million new environmental data points using ICP-MS and LA-ICP-MS in conjunction with highly detailed historical records to show the devastating impact of the Black Death on European metal production, an insight into the pandemic’s effect on human activity, demographics, and population health. In the last ~2000 years, only two other instances (in the 1460s C.E. and in 1885 C.E.) even remotely approached Black Death Pb deposition levels, either due to economic decline or epidemic disease, or both. Our findings imply that what were once believed to be background Pb levels represent, in fact, a significant anthropogenic component of the atmosphere over the last ~2000 years. The sole exception was a 4-year period at the time of the Black Death when atmospheric Pb pollution dropped to levels analytically undetectable by LA-ICP-MS. The geographic proximity to pollution sources and ultra-high resolution of the data presented here provide the most detailed, updated, regional record of European Pb pollution for the past two millennia and indicate that man-made pollution has been, and continues to be, a major contributor to lead levels in the atmosphere. Current policies and industry consensus, based on the assumption that current Pb atmospheric levels contain a significant “natural” Pb contribution, are thus clearly misleading. The health implications of such anthropogenically elevated levels of Pb in the atmosphere need further investigation in light of these new data.
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4.1. Author Contributions

N.S., S.S., P.B., and M.H. conducted the sampling, analysis, and annual layer counting. E.K. calculated enrichment factors. H.H. conducted radiocarbon analysis. P.A.M. and A.K. contributed climatological, glaciological, and atmospheric circulation analysis and expertise. A.F.M., C.L., and M.MCc. researched historical and current health aspects of Pb poisoning and mining, as well as historical epidemiology, archaeological, and historical data. A.F.M. wrote the initial paper draft, and all authors met to produce the final draft. All authors discussed the results and commented on the manuscript.

References

Allen, M. (2011). Silver production and the money supply in England and Wales, 1086 to c. 1500, Econ. Hist. Rev., 64, 114–131.

Almada, A. A. C. D., Golden, S. A. Osofsky, and S. S. Myers (2017), A case for planetary health/GeoHealth, GeoHealth, 1, 75–78, doi:10.1002/2017GH000084.

Andersen, K. K., et al. (2006), The Greenland ice core chronology 2005, 15–42 ka. Part 1: Constructing the time scale, Quat. Sci. Rev., 25, 3246–3257, doi:10.1016/j.quascirev.2006.08.002.

Barnatt, J., and K. Smith (2004), The Peak District Landscapes Through Time, pp. 111–113, Windgather Press, Macclesfield.

Baron, S., J. Carignan, S. Laurent, and A. Ploquin (2006), Medieval lead making at Mont-Lozère Massif (Cévennes-France): Tracing ore sources using Pb isotopes, Appl. Geochem., 21, 241–252, doi:10.1016/j.applgeochem.2005.09.005.

Barret, C. (2010), The production of silver, copper, and lead in the Harz mountains from late medieval times to the onset of industrialization, in Materials and Expertise in Early Modern Europe, edited by U. Klein and E. C. Spary, pp. 71–100, Univ. of Chicago Press, Chicago.

Blanchard, I. (2005), Mining, Metallurgy and Minging in the Middle Ages, vol. 3, Steiner, Stuttgart, Germany.

Bohleber, P., D. Wagenbach, W. Schöner, and R. Böhm (2013), To what extent do water isotope time series from low accumulation Alpine ice cores reproduce instrumental temperature series?, Tellus B, 65, 2014, doi:10.3402/tellusb.v65i2014.14887.

Brock Ramsey, C. (1995), Radiocarbon calibration and analysis of stratigraphy: The OxCal program, Radiocarbon, 37, 425–430, doi:10.1017/0033822220030903.

Bos, K. L., et al. (2011), A draft genome of Yersinia pestis from victims of the Black Death, Nature, 478, 506–510, doi:10.1038/nature10549.

Brännvall, M.-L., et al. (1999), The medieval metal industry was the cradle of modern large-scale atmospheric lead pollution in northern Europe. Environ. Sci. Technol., 33, 4391–4395.

Brayshay, M. (1980), Depopulation and changing household structure in the mining communities of West Cornwall, Local Popul. Stud., 25, 26–47.

Chang, S. H., et al. (2006), Low blood lead concentration in association with infertility in women, Environ. Res., 101, 380–386, doi:10.1016/j.envres.2005.10.004.

Claupe, P., et al. (2010), The crown silver mines and the historic landscape in Devon, England, ArchéSciences, 34, 299–306.

Claupe, P., et al. (2016), Iron and ironstone, in The Archaeology of Mining and Quarrying in England. A Research Framework for the Archaeology of the Extractive Industries in England, edited by P. Newman, pp. 115–126, National Association of Mining History Organisations, Matlock Bath.

Creighton, C. (1989), A History of Epidemics in Britain, vol. 1, Cambridge Univ. Press, Cambridge, U. K.

De Rosa, M., et al. (2003), Traf

DeWitte, S. N., and J. Wood (2008), Selectivity of Black Death mortality with respect to pre-existing health, J. Environ. Monit., 10, 502–510, doi:10.1039/B401500B.

Dyer, C. (2000), Everyday Life in Medieval England, 14 pp., Cambridge Univ. Press, Cambridge, U. K.

Elsen, O., et al. (2003), Alpine ice cores and ground penetrating radar: combined investigations for glaciological and climatic interpretations of a cold Alpine ice body, Tellus, 55B, 1007–1017, doi:10.1034/j.1600-0889.2003.tb00849.x.

Gabrieli, J., and C. Barbante (2014), The Alps in the age of the Anthropocene: The impact of human activities on the cryosphere recorded in the Colle Gnifetti glacier, Rend. Fis. Acc. Lincei, 25, 71–83, doi:10.1007/s10933-013-9705-y.

Gallagher, S. (1985), The Archaeology of the Extractive Industries in England. A Research Framework for the Archaeology of the Extractive Industries in England, Oxford Univ. Press, Oxford.

Hatcher, J. (2003), Understanding the population history of England, Past Present, 208, 83–130.

Henningsberg, S. (2000), Lead poisoning in a historical perspective, Am. J. Ind. Med., 38, 244–254.

Hinkley, T. K., P. J. Lamothe, S. Wilson, D. L. Finnegan, and T. M. Gerlach (1999), Metal emissions from Kilauea, and a suggested revision of the estimated worldwide metal output by quiescent degassing of volcanoes, Earth Planet. Sci. Lett., 170, 315–325, doi:10.1016/S0012-821X(99)00103-X.

Hoffmann, H. M., (2016) Micro radiocarbon dating of the particulate organic carbon fraction in Alpine glacier ice: Method refinement, critical evaluation and dating applications, PhD dissertation, Univ. of Heidelberg, doi:10.1126/sciadv.1400712.

Hong, S. J., P. Candelone, C. C. Patterson, and C. F. Bouton (1994), Greenland ice evidence of hemispheric lead pollution two millennia ago by Greek and Roman civilizations, Science, 265, 1841–1843, doi:10.1126/science.265.5180.1841.

Hutchinson, D., et al. (2007), Radiocarbon dating technique applied to an ice core from the Alps indicating late Pleistocene ages, J. Geophys. Res., 111, D14405, doi:10.1029/2005JD009180.

Kauffman, P. R., et al. (2008), An improved continuous flow analysis system for high-resolution field measurements on ice cores, Environ. Sci. Technol., 42, 8044–8050, doi:10.1021/es8007722.

Kelly, S. E., and N. P. Brooks (2013), Charters of Christ Church Canterbury, pp. 636–638, Oxford Univ. Press, Oxford.

Le Roux, G., et al. (2004), Identifying the sources and timing of ancient and medieval atmospheric lead pollution, in England using a paint profile from Lindog bow, Manchester, J. Environ. Monit., 6, 502–510, doi:10.1039/B401500B.

Lewis, C. (2016), Disaster recovery, new archaeological evidence for the long-term impact of the ‘calamitous’ fourteenth century, Antiquity, 90, 777–797, doi:10.15184/aqy.2016.69.

Martinez Cortizas, A., et al. (2013), Atmospheric Pb pollution in N Iberia during the late Iron Age/Roman times reconstructed using the high-resolution record of La Molina mine (Asturias, Spain), J. Paleol., 50, 71–86, doi:10.1002/10933-013-9705-y.
Mayewski, P. A., et al. (1986), Sulfate and nitrate concentrations from a South Greenland ice core, Science, 232, 975–977, doi:10.1126/science.232.4753.975.

Mielke, H. W., and S. Zahrani (2012), The urban rise and fall of air lead (Pb) and the latent surge and retreat of societal violence, Environ. Int., 43, 48–55, doi:10.1016/j.envint.2012.03.005.

Montgomery, J., et al. (2010), Gleaming, white and deadly, J. Roman. Archaeol. Suppl., 78, 199–226.

Mushak, P. (2011), Lead and Public Health: Science, Risk and Regulation, pp. 401–816, Elsevier, Boston.

Nayling, N. (1998), The Major Pill Medieval Wreck (York, Council for British Archaeology), pp. 105–115.

Nightingale, P. (2005), Some new evidence of crises and trends in mortality in late medieval England, Past Present, 187, 33–68, doi:10.1093/past/igt009.

Nriagu, J. O. (1983), Lead and Lead Poisoning in Antiquity, Wiley, New York.

Osterberg, E. C., M. J. Handley, S. B. Sneed, P. A. Mayewsky, and K. J. Kreutz (2006), Continuous ice core melter system with discrete sampling for major ion, trace element, and stable isotope analyses, Environ. Sci. Technol., 40, 3355–3361, doi:10.1021/es052536w.

Papastefanou, C. (2006), Residence time of tropospheric aerosols in association with radioactive nuclides, Appl. Radiat. Isot., 64, 93–100, doi:10.1016/j.apradiso.2005.07.006.

Riazi, Z. (1980), Life, Marriage and Death in a Medieval Parish. Economy, Society and Demography in Halesowen 1270–1400, pp. 124–128, Cambridge Univ. Press, Cambridge, U. K.

Renberg, I., R. Bindler, and M. L. Brannvall (2001), Using the historical atmospheric lead-deposition as chronological marker in sediment deposits in Europe, Holocene, 11, 511–516, doi:10.1191/09596830168023468.

Reyes, J. (2015), Lead exposure and behavior: Effects on aggression and risky behavior among children and adolescents, Econ. Inq., 53, 3, doi:10.1111/ecin.12202.

Richardson, G. M., Garrett R. Mitchell I, Mah-Poulson M, Hackbarth T (2001), Critical review of natural global and regional emissions of six trace metals to the atmosphere. (International Lead Zinc Research Organisation, International Copper Association, Nickel Producers. Environmental Research Association).

Rieuwerts, J. H. (1988), A History of the Laws and Customs of the Derbyshire Lead Mines, pp. 15–16, Derbyshire Barmote Court, Sheffield.

Schwikowski, M., et al. (2004), Post-17th-century changes of European lead emissions recorded in high-altitude alpine snow and ice, Environ. Sci. Technol., 38, 957–964, doi:10.1021/es034715o.

Selevan, S. G., D. C. Rice, K. A. Hogan, S. Y. Euling, A. Pfahles-Hutchens, and J. Bethel (2003), Blood lead concentration and delayed puberty in girls, N. Engl. J. Med., 243, 1527–1536, doi:10.1056/NEJMoa020880.

Shotyk, W., D. Weiss, P. G. Appleby, A. K. Cheburkin, R. Frei, M. Gloor, J. D. Kramers, S. Reese, and W. O. van der Knaap (1998), History of atmospheric lead deposition since 12,370 14C yr from a peat bog, Jura mountains, Switzerland, Science, 281, 1635–1640, doi:10.1126/science.281.5383.1635.

Strömbärg, U., T. Lundth, and S. Skerfving (2008), Yearly measurements of blood lead in Swedish children since 1978: The declining trend continues in the petrol-lead-free period 1995–2007, Environ. Res., 107, 322–335, doi:10.1016/j.envres.2008.03.007.

Snee, S. B., et al. (2015), New LA-ICP-MS cryocell and calibration technique for sub-millimeter analysis of ice cores, J. Glaciol., 61, 233–242, doi:10.3189/2015JoG141139.

United Nations Environment Programme (2010), Final review of scientific information on lead: 73–76.

U.S. Geological Service (2013), Salient lead statistics. [Available at http://minerals.usgs.gov/minerals/pubs/commodity/lead/stat/tbl1.txt.]

van Hoof, T. B., F. P. M. Bunnik, J. G. M. Waucomont, W. M. Kürschner, and H. Visscher (2006), Forest re-growth on medieval farmland after the Black Death pandemic—Implications for atmospheric CO2 levels, Palaeogeogr. Palaeoclimatol. Palaeoecol., 237, 396–411, doi:10.1016/j.palaeo.2005.12.013.

Wagenbach, D., and K. Geis (1989), The mineral dust record in a high altitude Alpine glacier (Colle Gnifetti, Swiss Alps), in Palaeoclimatology and Paleometeorology: Modern and Past Patterns of Global Atmospheric Transport, NATO ASI series, 282, edited by M. Leinen and M. Sarnthein, pp. 543–564.

Wedepohl, K. H. (1995), The composition of the continental crust, Geochim. Cosmochim. Acta, 59, 1217–1232, doi:10.1016/0016-7037(95)00038-2.