Genomic Evidence for the Recycling of Complex Organic Carbon by Novel Thermoplasmatota Clades in Deep-Sea Sediments

Peng-Fei Zheng, Zhanfei Wei, Yingli Zhou, Qingmei Li, Zhao Qi, Xiaoping Diao, Yong Wang

Institute of Deep Sea Science and Engineering, Chinese Academy of Sciences, Sanya, Hainan, China
State Key Laboratory of Marine Resources Utilization in the South China Sea, Hainan University, Haikou, China
College of Life Science, Hainan Normal University, Haikou, China
Institute for Ocean Engineering, Shenzhen International Graduate School, Tsinghua University, Shenzhen, China

ABSTRACT Thermoplasmatota have been widely reported in a variety of ecosystems, but their distribution and ecological role in marine sediments are still elusive. Here, we obtained four draft genomes affiliated with the former RBG-16-68-12 clade, which is now considered a new order, “Candidatus Yaplasmales,” of the Thermoplasmatota phylum in sediments from the South China Sea. The phylogenetic trees based on the 16S rRNA genes and draft genomes showed that “Ca. Yaplasmales” archaea are composed of three clades: A, B, and C. Among them, clades A and B are abundantly distributed (up to 10.86%) in the marine anoxic sediment layers (>10-cm depth) of six of eight cores from 1,200- to 3,400-m depths. Metabolic pathway reconstructions indicated that all clades of “Ca. Yaplasmales” have the capacity for alkane degradation by predicted alkyl-succinate synthase. Clade A of “Ca. Yaplasmales” might be mixotrophic microorganisms for the identification of the complete Wood-Ljungdahl pathway and putative genes involved in the degradation of aromatic and halogenated organic compounds. Clades B and C were likely heterotrophic, especially with the potential capacity of the spermidine/putrescine and aromatic compound degradation, as suggested by a significant negative correlation between the concentrations of aromatic compounds and the relative abundances of clade B. The sulfide-quinone oxidoreductase and pyrophosphate-energized membrane proton pump were encoded by all genomes of “Ca. Yaplasmales,” serving as adaptive strategies for energy production. These findings suggest that “Ca. Yaplasmales” might synergistically transform benthic pollutant and detrital organic matter, possibly playing a vital role in the marine and terrestrial sedimentary carbon cycle.

IMPORTANCE Deep oceans receive large amounts of complex organic carbon and anthropogenic pollutants. The deep-sea sediments of the continental slopes serve as the biggest carbon sink on Earth. Particulate organic carbons and detrital proteins accumulate in the sediment. The microbiially mediated recycling of complex organic carbon is still largely unknown, which is an important question for carbon budget in global oceans and maintenance of the deep-sea ecosystem. In this study, we report the prevalence (up to 10.86% of the microbial community) of archaea from a novel order of Thermoplasmatota, “Ca. Yaplasmales,” in six of eight cores from 1,200- to 3,400-m depths in the South China Sea. We provide genomic evidence of “Ca. Yaplasmales” in the anaerobic microbial degradation of alkanes, aliphatic and monoaromatic hydrocarbons, and halogenated organic compounds. Our study identifies the key archaeal players in anoxic marine sediments, which are probably critical in recycling the complex organic carbon in global oceans.

KEYWORDS PAH, archaea, aromatic degradation, anaerobes, metagenome, dehalogenase, South China Sea
Recycling of organic matter in sediments is an important component of biogeochemical cycles because marine sediments are critical for long-term carbon storage (1, 2). Prokaryotes in deep oceanic subsurface account for more than 25% of global cell numbers (3), and it is still controversial whether Archaea or Bacteria are more abundant in this extensive ecosystem formed in marine subsurface sediments (4). Only few bacteria and archaia inhabiting marine sediment have cultivated representatives (5), which poses a major challenge when elucidating the metabolic mechanisms that sustain life, even abundant under the extreme energy-limiting conditions prevailing in the marine deep biosphere (6). Most groups of microorganisms inhabiting deep-sea subsurface sediment were proposed to be fueled by recalcitrant organic carbon (ROC) buried in the marine sediments (7). However, it remains unexplained how these subsurface microorganisms survive under the conditions where the available energy flux appears to be even 1,000-fold lower than the minimum value for maintenance energy estimated from laboratory cultures (6, 8). Considering the vast coverage of marine sediments on earth, our understanding of prokaryotic activities in subsurface marine sediment is still limited.

*Thermoplasmatota* is a globally distributed and ecologically important archaeal phylum that consists of several classes such as *Aciduliprofundales*, *Thermoplasmatales*, *Methanomassiliicoccales*, and “Candidatus Poseidoniales” (formerly called Marine Group II). These organisms were mostly found in surface seawater, while some clades are heterotrophic archaia in deep aphotic waters (9). “Candidatus Pontarchaea” (formerly called Marine Group III) and photic-zone MG-III metagenome-assembled genomes (MAGs) contained numerous photolyase and rhodopsin genes, as well as genes for peptide and lipid uptake and degradation, suggesting a photoheterotrophic lifestyle (10–12); *Thermoproteales* (formerly called Marine Benthic Group D and DHVEG-1) was mixotrophic with detection of extracellular peptidase genes and Wood-Ljungdahl (WL) pathways (13). In a number of uncharacterized lineages, including UBA10834 (now referred to as “Candidatus Gimiplasmatales”), formaldehyde and acetate assimilation might proceed via the WL pathway, which indicates a mixotrophic lifestyle in sediment (14). As a novel clade in *Thermoplasmatota*, RBG-16-68-12 was first reported in terrestrial sediment at a 16-ft depth (15) and was also abundant in deep-sea sediment with water depth over 3,000 m (16). To date, there are few studies on RBG-16-68-12 archaia regarding their distribution depth and ecological function in nutrient-poor marine sediments. Compared to other archaeal phyla, studies on lineages of *Thermoplasmatota* are limited (12); a genome survey found that the *Thermoplasmatota* genomes contain the third highest number of unknown genes (17), suggesting that lineages of *Thermoplasmatota* should be further explored.

The particulate organic carbon (POC) export flux reached to 8.2 to 20 mmol of C m⁻² day⁻¹ in the South China Sea (SCS) shelf (18, 19). Recently studies showed that marginal seas, including the continental shelves of SCS, play a key role in the global carbon cycle by linking the terrestrial, oceanic, and atmospheric carbon reservoirs, which are effective in burying POC (20, 21). The POC provides abundant organic matter for microorganisms inhabiting the marine sediment. Among the abundant and prevalent organic matter in the environment, 16 polycyclic aromatic hydrocarbons (PAHs) as the major components of petroleum were declared as priority pollutants (22, 23). However, the roles of different microorganisms in organic matter processing in SCS continental slopes were not clear. In this study, we obtained MAGs of *Thermoplasmatota* dominated by RBG-16-68-12 from sediments of SCS. With analyses of 16S rRNA gene amplicons, we revealed prevalence of RBG-16-68-12 in >10 cm below seafloor (cmbsf) layers of six sediment cores. We could therefore confirm the independent phylogenetic position of RBG-16-68-12 consisting of three clades in *Thermoplasmatota*. Genomics data indicate metabolic potentials of aromatic and halogenated organic compound degradation, alkane utilization, and adaptive schemes to cope with energy limitation.
RESULTS AND DISCUSSION

Environmental parameters of the marine sediment samples. We collected eight sediment cores (SY40, SY152, SY153, SY154, SY155, SY159, SY165, and SY166; Fig. 1) from the SCS continental shelf by the manned submersible Deep-Sea Warrior. Among the determined environmental parameters of the sediment cores, the concentrations of nitrate and nitrite in the surface (0 to 2 cm, nitrate: 2.55 to 16.45 μmol L⁻¹; nitrite: 0.75 to 4.59 μmol L⁻¹) were 2- to 3-fold higher than deeper sediment layers (8 to 24 cm, nitrate: 0.13 to 7.26 μmol L⁻¹; nitrite: 0–0.70 μmol L⁻¹) (see Fig. S1 and Table S3 in the supplemental material). The ammonium concentrations in the different layers (0 to 18 cmbsf) were largely between 13.33 and 35 μmol L⁻¹; the highest ammonium concentration of 81.48 μmol L⁻¹ was measured in the deepest layer of SY159. Therefore, the deep sediment layers showed a higher level of ammonium compared to surface layers where nitrate concentration was highest. Similarly, silicate concentrations decreased with sediment depth, whereas phosphate concentrations did not vary with the depth (see Fig. S1). The phosphate concentrations in SY153 and SY159 were 5.51 to 6.77 μmol L⁻¹, which were significantly higher than the 0.65 to 2.26 μmol L⁻¹ in SY40 (t test; P < 0.01). In contrast, the silicate concentrations in SY40 were significantly higher than those in SY153 and SY159 (Student t test; P < 0.01) (see Table S3). We noticed that the TOC in our study ranged from 1.17 to 3.66%, which was consistent with previous reports (24, 25). The TOC/TN ratio as a frequently used index for organic matter origin was estimated to be 6.85 to 7.39 in SY153. Land plants and algae are diverse in their C/N ratios: marine algae, due to richness in protein and absence of cellulose, typically have C/N ratios between 4 and 10 (26, 27), and new lines of evidence suggested that cyanobacteria may produce more pentadecane (28). Therefore, the TOC/TN ratio of the sediments in this study indicates their origin from plants and algae.

FIG 1 Sampling stations in the South China Sea. The sampling sites of sediment cores are indicated as red (for metagenomic study) and green (for 16S rRNA gene study) dots. All of the sediment cores were obtained by the Deep-Sea Warrior manned submersible with push core samplers. SY40, SY153, SY159, and SY152 were located on a marginal slope at the northwestern South China Sea at an ~1,200-m depth (see Table S1 for more details).
For the PAH measurements in our marine sediment of SY153, the anthracene was the most abundant PAH (460.63 ± 42.34 ng g⁻¹, dry weight), followed by pyrene (436 ± 57.03 ng g⁻¹, dry weight) and fluoranthene (239.89 ± 3.20 ng g⁻¹, dry weight). The other PAHs were <100 ng g⁻¹. The PAH measurement results in SY152 were similar to those in SY153, both much higher than the previous records for the total PAHs (170.7 to 498.2 ng g⁻¹) (24). Among the different layers of SY152 and SY153, the most abundant PAHs were detected in surface layers (4 to 6 cmbsf and 8 to 10 cmbsf) and the deepest layer (24 to 26 cmbsf) (see Fig. S2). The ratio of phenanthrene and anthracene isomers in this study was 0.4 to 0.6, indicating a fossil fuel origin of the PAHs (29, 30).

Identification of MAGs for RBG-16-68-12. We obtained 174 Gb clean metagenomic data for 12 sediment layers between 2 and 26 cmbsf of three push cores (SY40, SY153, and SY159) randomly selected from the eight cores (see Table S1). Metagenome assembly and genome binning resulted in a total of 201 bacterial and archaeal MAGs with a completeness of >50% and contamination at <10%. Five of the MAGs classified as Thermoplasmatota were obtained from the metagenomes for SY40 and SY153; among these, four MAGs were classified as RBG-16-68-12 by GTDB-tk software (31). As shown in Table 1, the genome size of the obtained Thermoplasmatota MAGs ranged from 1.37 to 2.03 Mb with a completeness between 61.84% and 98.93%.

Phylogenomic inference of the Thermoplasmatota genomes showed that the four MAGs of this study were grouped with members of RBG-16-68-12 (15, 32) adjacent to UBA10834 and SG8-5 (Fig. 2a). There are a total of 44 RBG-16-68-12 genomes in the NCBI database. In the tree, three distinguishable clades of RBG-16-68-12 were formed. The MAGs of the clades A and B were both obtained from deep-sea marine sediments with only one previously reported MAG. Therefore, we provide one new MAG in clade A and three new MAGs in clade B in this study. Four additional MAGs were obtained from SCS sediments by us for clade A and had been recently submitted to the NCBI. For the RBG-16-68-12 in NCBI database, 38 of 44 MAGs clustered into an independent clade C consisting of genomes largely from soil (n = 4) and temperate grassland microbiome (n = 31) (see Table S4). In addition, the average nucleotide identity (ANI) values among the RBG-16-68-12 MAGs obtained in this study were between 69 and 93% (see Table S5). The MAG Bin162 for clade A from SY40 was much different from those for clade B (Bin292, Bin295, and Bin296) in SY153 with ANI values of 69 to 70%, which is much lower than the values for a moderate affinity (ANI 80 to 90%) (33) and supports the divergency between the two marine clades. This result confirmed independence of RBG-16-68-12 clade as a novel lineage with more genomes, and we therefore propose “Candidatus Yaplasmales” as the new name of the lineage (34). The word “涯”, “Ya”) in Chinese means limit, and the new clade name here indicates their survival in the deep-sea nutrient-limited sediments.

The ANI values between Methanomassiliicoccales and “Ca. Yaplasmales” range from 67 to 70% (average, 69%), which suggests that they were close relatives. However, the analysis of 16S rRNA gene sequences showed that the similarity between Methanomassiliicoccales and clade A ranged from 87 to 89%, higher than the values between Methanomassiliicoccales and clade B (82 to 88%), indicating that the clades A and B are evidently different from Methanomassiliicoccales. Furthermore, the RED (the relative evolutionary divergence) value of the order Methanomassiliicoccales was 0.554, while the RED value of “Ca. Yaplasmales” lineage (RBG-16-68-12, GTDB-tk database, r202) was 0.528. This suggests that “Ca. Yaplasmales” was a new order of Thermoplasmatota (see Fig. S2).

Prevalence of “Ca. Yaplasmales” in deep sediment layers. The relative abundance of each “Ca. Yaplasmales” MAG in the three sediment cores was calculated as the percentage of the metagenomic reads that mapped to the MAGs. Our result showed that MAG Bin162 of clade A was only detected in 25 to 26 cmbsf of SY40 with 0.09% in relative abundance, and could not be found in all layers of the SY153 and SY159. The MAGs in clade B (Bin292, Bin296, and Bin 295) recruited 0.02 to 0.22% of metagenomic reads in 8 to 10 cmbsf and 16 to 18 cmbsf of SY153, respectively, which indicates a higher relative abundance of clade B in deep, possibly anoxic sediment layers (>10 cmbsf) compared to surface layers.
| Category                                | Bin162 | Bin292 | Bin296 | Bin295 | Bin344 |
|-----------------------------------------|--------|--------|--------|--------|--------|
| Classification                          | "Ca. Yaplasmales" clade A | "Ca. Yaplasmales" clade B | "Ca. Yaplasmales" clade B | "Ca. Yaplasmales" clade B | Thermoprofundales |
| No. of scaffolds                        | 136    | 286    | 109    | 241    | 268    |
| Genome size (Mb)                        | 2.03   | 1.46   | 0.81   | 1.37   | 0.65   |
| N₅₀ value (kbp)                         | 34.38  | 6.44   | 10.64  | 9.65   | 2.68   |
| No. of protein coding genes             | 1,974  | 1,744  | 947    | 1,732  | 937    |
| Completeness (%)                       | 98.93  | 94.35  | 64.33  | 61.84  | 54.32  |
| Contamination (%)                      | 1.6    | 4.4    | 0      | 4.46   | 1.87   |
| GC (%)                                  | 57.1   | 57.6   | 64.5   | 62.9   | 35.9   |
| RED value                              | 0.57   | 0.58   | 0.59   | 0.58   | 0.73   |
| No. (%) of KEGG genes                   | 1,049 (53.14) | 1,030 (59.06) | 593 (62.62) | 792 (45.73) | 523 (55.82) |
| No. (%) of COGs                        | 1,395 (70.67) | 1,267 (72.65) | 723 (76.35) | 1,051 (60.68) | 638 (68.09) |
| Prevalence depth (cm)                   | 25–26  | 16–18  | 16–18  | 16–18  | 22–24  |

*The five genomes were annotated in eggNOG database (v2020.08). Genome completeness and contamination were estimated using CheckM (v1.1.2). RED, relative evolutionary divergence.
FIG 2 Phylogenetic trees of “Ca. Yaplasmales” (RBG-16-68-12). (a) An ML phylogenetic tree was constructed using alignment of concatenated 43 conserved proteins obtained from MAGs of Thermoplasmataota. The (Continued on next page)
To confirm the distribution of the “Ca. Yaplasmales” clades, we sequenced 16S rRNA gene amplicons in different layers of the eight sediment cores. Overall, 3,397,112 raw sequencing reads of 16S rRNA gene amplicons were obtained (see Table S2); of these, 2,914,782 qualified 16S rRNA gene sequences were used for operational taxonomic unit (OTU) sorting. The communities were dominated by Proteobacteria, Chloroflexi, and Thaumarchaeota (Fig. 3). Archaea accounted for 19.26% ± 5.96% of the whole microbial communities in the eight cores for this study (Fig. 2). For the archaeal community structure, Thaumarchaeota was the only dominant phylum in all sediment layers with an average relative abundance 56.27% of archaeal communities in 39 samples, and always occupied >90% of the archaeal communities in the top surface sedi-

FIG 3 Distribution of “Ca. Yaplasmales” in eight sediment cores. (a) Prokaryotic community structures in eight sediment cores were revealed by sequencing and taxonomic sorting of 16S rRNA gene amplicons at the phylum level using SILVA 132 as a reference database. (b and c) Archaeal communities were separated to demonstrate their relative abundance at phylum (b) and class (c) levels. Ca. Ya. clade A, “Ca. Yaplasmales” clade A; Ca. Ya. clade B, “Ca. Yaplasmales” clade B.

FIG 2 Legend (Continued)
reference genomes of clades A and C are listed in Table S4. (b) A phylogenetic tree was constructed using 16S rRNA genes extracted from three “Ca. Yaplasmales” MAGs, representative sequences of OTUs from 16S rRNA amplicons, and the reference sequences within Thermoplasmato from SILVA SSU database version 138 and NCBI. Bootstrap values were based on 1,000 replicates, and red solid nodes refer to the bootstrap values of >50%. The tree was drawn to scale, with branch lengths representing the number of substitutions per site.
The relative abundance of *Thaumarchaeota* in all samples also decreased along with the sediment depth. *Nitrosoarchaea*, known as ammonia-oxidizing archaea (AOA), dominated the *Thaumarchaeota* in all the samples. *Nanoarchaeota*, *Thermoplasmatota*, and *Bathyarchaeota* were abundantly distributed in the deeper layers. The relative abundances of *Thermoplasmatota* were between 7.54% (SY159; 10 to 12 cmbsf) and 55.89% (SY165; 10 to 12 cmbsf) of the archaeal communities in the sediment samples over 10 cmbsf. In samples of SY152 and SY153, “Ca. Yaplasmales” clade B was much more abundant in >10 cmbsf than upper sample layers, since its relative abundance ranged from 15.29% (10 to 12 cmbsf of SY153) to 43.00% (12 to 14 cmbsf of SY152) of the archaeal communities (Fig. 2). Most (≥50%) of *Thermoplasmatota* were assigned to “Ca. Yaplasmales,” and the rest were largely unclassified.

A phylogenetic tree was built with the 16S rRNA genes extracted from the MAGs, the 16S rRNA gene amplicons, and references from public databases. The phylogeny revealed more species in the “Ca. Yaplasmales” clades (Fig. 2b), some of which exhibit affinity to the sequences obtained from marine sediments and cold seeps located in SCS and the Gulf of Mexico. Clade C was exclusively composed of the sequences isolated from land soils.

The global distribution of “Ca. Yaplasmales” was examined by searching their 16S rRNA gene fragments in published data sets for a study of prokaryotic communities of 299 deep-sea sediment samples from 41 different sites (35). Within the published 16S rRNA amplicons, “Ca. Yaplasmales” were identified in 25 sediment samples from 14 sites with their relative abundance in prokaryotes almost all <0.1% (see Fig. S4). As expected, clade C was not detected in the marine sediments. Furthermore, “Ca. Yaplasmales” occupied 9.25% of the whole archaeal community (0.075% of the detected prokaryotes) in depth of 37.2m below seafloor of the Peru Basin (ODP leg 201, site 1231), which was the highest relative abundance among the 41 sites and was comparable to the highest abundance in this study. A sediment sample from the South Pacific Gyre (329U1367D1H-2) harbored >2% of the archaeal community. In brief, the distribution of “Ca. Yaplasmales” was wide in marine sediments, despite the low relative abundances. However, the usage of the different universal primers (U515F and U806R), compared to this study, in the published data sets might be causal to the low relative abundance.

**Metabolic potentials of “Ca. Yaplasmales.”** The metabolic potentials of “Ca. Yaplasmales” clades were predicted based on annotation results of available MAGs affiliated with clades A and B, including nine reference MAGs from the NCBI and five from this study (Fig. 4). The six MAGs of clade C with a high completeness (2 MAGs > 97% and 4 MAGs >70%) were used for prediction of metabolic pathways. Based on the phylogenetic position of new class “Ca. Yaplasmales” (Fig. 3b), we compared their metabolic potentials to the closely related *Thermoplasmatota* in marine sediment inhabitants, such as *Methanomassiliicoccales*, SG8-5 and *Thermoprofundales* (Fig. 5). *Methanomassiliicoccales* was known for its methane metabolism (36), and *Thermoprofundales* was probably composed of mixotroph in mangrove and intertidal mudflat sediments (13). SG8-5 was inferred to have capability for extracellular protein degradation in sediment (37).

**(i) Carbon metabolism.** Alkane is ubiquitous carbon source for microbial inhabitants in marine sediment (38). 1-Methyl-alkyl-succinate synthase gene (*assA*) as a marker gene of anaerobic alkane degradation to acyl coenzyme A (acyl-CoA) (39, 40) was identified in all “Ca. Yaplasmales” MAGs, indicating their capacity of alkane utilization. The genes encoding acetyl-CoA acyltransferase (FadA) and long-chain acyl-CoA synthetase (ACSL and FadD) involved in further oxidation of acyl-CoA via the beta-oxidation were encoded in the MAGs. Therefore, alkanes are probably an important carbon source for all the clades of “Ca. Yaplasmales.” Genes encoding the bifunctional enzyme 3-hexulose-6-phosphate synthase/6-phospho-3-hexulonoserase (HPS-PHI) were identified in all clades, suggesting the potential capacity in pentoses biosynthesis (41). The genes coding for acetate-CoA ligase (ADP-forming, AcdAB) and acetyl-CoA synthetase (Acs) were found in almost all “Ca. Yaplasmales” MAGs, which indicates that “Ca. Yaplasmales” might benefit from acetate assimilation or secretion. Acetate is an important energy and carbon source in marine sediment (42) and is also the central
intermediate regulating community interactions and biogeochemical cycling in these deep-sea sediments (16). These inferences are in line with mounting evidence that acetogens are important community members in energy-limited seafloor ecosystems (42).

Regarding autotrophy, using a range of inorganic and organic substrates, the WL pathway produces acetate by acetogenic CO₂ reduction (43) and was thought to be the largest carbon fixation pathway under anaerobic conditions (43, 44). Almost all of the MAGs from clade A harbor a complete set of WL genes, including the CO dehydrogenase (CooS and CooF) (45, 46) and acetyl-CoA synthase (CdhDE/AcsBE) genes (47) (Fig. 4; see also Table S6). The absence of WL pathway in Bin25 from clade A was likely due to its low genome completeness (60.67%), while all the MAGs from the clades B and C did not have a WL pathway (Fig. 4; see also Tables S4 and S8), as well as other known carbon fixation pathways. These results indicate that clade A has the genetic potential for inorganic carbon assimilation via the WL pathway for autotrophy, and clades B and C were likely heterotrophic archaea (Fig. 4; see also Table S6). The acetate metabolism of clade A might be associated with the WL pathway, which indicates the acetogenic capacity of this clade. Moreover, the WL pathway can probably be reversed and is associated with alkane degradation for CO₂ production in marine sediment (48). Therefore, the WL pathway in this study is possibly a part of alkane degradation, rather than an autotrophic approach.

(ii) Sulfur metabolism. The sulfide quinone oxidoreductase gene (sqr) responsible for sulfide oxidation was present in all the MAGs for clade A (Fig. 5 and 6c), and most of MAGs for clades B (2 of 3 MAGs) and C (4 of 6 MAGs) of “Ca. Yaplasmales,” which indicates that “Ca. Yaplasmales” might utilize sulfide at micro- to millimolar concentrations as an electron donor (49). Sqr is a membrane-associated protein that oxidizes sulfide to zero-valent sulfur and transfers electrons to flavin adenine dinucleotide (FAD). In particular, hydrogen sulfide biotransformation might also provide electrons, reduced FAD, and hydrogen to fix CO₂ via WL pathway using H₄-folate as C1 carriers for clade A. It is interesting that this enzyme in clade B is a potentially new type of Sqr with respect to the phylogenetics position (Fig. 6c) (50, 51).

Degradation of halogenated organic compounds. Halogenated compounds are widely reported in marine sediments as anthropogenic contaminants and natural products (52). From the Pfam annotation (53), we identified rdh gene encoding putative reductive dehalogenase (RDase) in clade A but not in clades B and C (Fig. 6d). The
domain organizations of RDases are variable and fall into at least two categories: a C-terminal 4Fe-4S dicluster binding domain (Pfam13484) and a 4Fe-4S dicluster domain (Pfam12838) (54). In the present study, the RDases contain Pfam13484 or Pfam12838, which indicates that "Ca. Yaplasmal" clade A could metabolize halogenated organic compounds.

Halogenated organic compounds (organohalides) are a large class of natural and synthetic chemicals that contain one or more halogens (fluorine, chlorine, bromine, or iodine) combined with carbon and other elements, which are globally prevalent, recalcitrant toxic, and carcinogenic environmental pollutants (55). They are produced through abiotic and biotic processes in marine ecosystems and may serve as electron acceptors in deep-sea sediments (16). Reductive dehalogenation is a process that...
FIG 6 ML phylogenetic trees of the alkyl-succinate synthase (Ass), benzyl-succinate synthases (Bss), sulfide-quinone oxidoreductase (Sqr), and reductive dehalogenase (RDase) proteins in “Ca. Yaplasmales.” The phylogenetic trees of Ass (a), Bss (b), Sqr (c), and RDase (d) are shown.
removes a halogen substituent from a molecule with concurrent reduction or adds two electrons to it, which is one of six enzyme-catalyzed dehalogenation reactions in microorganisms growing on halogenated organic compounds (56). It has been reported that some organisms can use halogenated substance as an energy source by transferring electrons to halogenated aromatic and aliphatic substances, a process termed organohalide respiration (57). The organohalide respiration pathway was one of the possible electron-acceptor systems in anoxic marine sediments (58, 59). Reductive dehalogenases are responsible for biological dehalogenation in organohalide respiring bacteria with a substrate source, not only from natural biological processes (59, 60) such as marine sponges and chlorinated biodegradation products of lignin but also some with an anthropogenic origin, including polychlorinated biphenyls or dioxins (60), and the results of this study indicate that “Ca. Yaplasmales” clade A plays an important role in recycling marine chlorinated pollutants (61).

The identification of alkyl-succinate synthase (Ass) and benzyl-succinate synthase (Bss; Fig. 6b) coding genes in clades A and B also suggests metabolic potentials of the degradation of alkanes and aromatic hydrocarbons. Bss catalyzes a highly unusual reaction: the addition of toluene to fumarate to form (R)-benzylsuccinic acid (62, 63), which drives anaerobic toluene degradation. Bss is now regarded as the prototype enzyme of a much larger family of fumarate-adding enzymes, which play important roles in the anaerobic metabolism of further aromatic and even aliphatic hydrocarbons (64). For the anaerobic oxidation of alkanes, an essential mechanism was also involved in the addition to fumarate, producing alkyl-substituted succinates (65), which was analogous to the anaerobic activation of aromatic hydrocarbons (66).

Furthermore, the bcrABCD genes in the class I ATP-dependent benzoyl-CoA reductase pathway were found in all clade A MAGs and in some MAGs in clades B (2/3 MAGs) and C (2/6 MAGs), which also indicates that “Ca. Yaplasmales” might utilize aromatic compounds (Fig. 4; see also Table S6) via a different approach. Benzoyl-CoA can be anaerobically degraded to yield a nonaromatic dienoyl-CoA intermediate by the bcr genes in anaerobes, including Dehalococcoidia, Anaerolineae, Deltaproteobacteria, Batharchaeota, and Thermoplasmata (67).

The correlation analysis was performed between 16 PAH concentrations and “Ca. Yaplasmales” relative abundances (based on 16S rRNA gene amplicons), which showed a significantly negative correlation between clade B and anthracene (Pearson correlation; \( R = -0.69; P < 0.05 \)) and between clade B and phenanthrene (Pearson correlation; \( R = -0.78; P < 0.01 \)). This indicates that “Ca. Yaplasmales” clade B probably benefits from PAH degradation. The negative correlation between microorganism abundance and the phenanthrene concentration has been reported recently in cultivation (68). Hence, the distribution of “Ca. Yaplasmales” may be influenced by the anthracene and phenanthrene accumulation (see Fig. S3). The physical properties of PAHs, such as low aqueous solubility and high solid/water distribution ratios, affect their distribution in marine sediment, which may explain the low abundance of “Ca. Yaplasmales” in some of our cores (Fig. 3).

Archaea have been reported to mediate oxidation of methane and other short-chain alkanes in sediments (16, 69), whereas anaerobic degradation of larger and complex hydrocarbons (for example organohalides) is rarely reported. Recently, anaerobic alkane and aromatic compound degradation with alkyl-coenzyme A and benzoyl-coenzyme A oxidation were reported in a new archaeal phylum Hermodarchaeota (70). Similarly, the possible ability of “Ca. Yaplasmales” in degrading aromatic compounds in marine sediment may also benefit them with fitness. This is the first report of potential anaerobic aromatic degradation for Thermoplasmata.

**Energy conservation mechanisms.** In the present study, hydrogenase genes were
widely detected in “Ca. Ypsilonales” clades A and C (Fig. 5; see also Table S6). The encoded hydrogenases might supply intracellular reducing equivalents needed for various redox reactions and redox homeostasis (71). The MvhADG/HdrABC complex catalyzes an iron-sulfur cluster-assisted disulfide reduction reaction, which can be integrated into a flavin-based electron bifurcation (FBE) process, a mode of energy coupling that optimizes the energy yield of the cell (72). The complex was frequently found in anaerobic bacteria and archaea (73), as well as in “Ca. Ypsilonales” MAGs (Fig. 4; see also Table S6), which indicated that the sediment layers dominated by “Ca. Ypsilonales” were probably oxygen limiting or anoxic. Here, the MAGs affiliated with clades A and C possessed the hppA gene that encodes K⁺-stimulated pyrophosphate-energized proton pump. HppA utilizes the energy of pyrophosphate hydrolysis as the driving force for proton movement across the membrane (74). The MAGs of “Ca. Ypsilonales” clade B harbor the mnh and mrg genes coding for multicomponent Na⁺:H⁺ antiporters as the proton or Na⁺ pumpers (Fig. 5), suggesting that clade B probably formed a functional complex involved in Na⁺ extrusion to maintain osmotic balance (75).

Environmental adaptation of “Ca. Ypsilonales.” As mentioned above, “Ca. Ypsilonales” archaea are probably anaerobes. The thioredoxin 1 (TrxA) that is able to protect anaerobes against oxygen (76, 77) was encoded by all “Ca. Ypsilonales” clades with up to five copies. TrxA plays a crucial role in maintaining the redox balance in various cells (78), with two thiol moieties that accept reducing elements from NADPH in reaction catalyzed by thioredoxin reductase (TR). The reduced TrxA can reduce disulfide bonds in proteins and helps them to maintain native structure or change conformation in response to various stimuli (79). The lack of the genes coding for a cytochrome bc complex and cytochrome bd terminal oxidase (Fig. 4; see also Tables S4 and S8) provides further evidence for anaerobic lifestyle of “Ca. Ypsilonales.” MAGs in clade A encode nitrate reductase/nitrite oxidoreductase complex (NarBGHI) and nitrite reductase (NirBD), which indicates capacity of nitrate utilization as electron acceptor for anaerobic respiration (Fig. 4; see also Table S6). Nitrite reductase coding genes nirBD and nirKS were also present in clades B and C MAGs, suggesting the two clades might reduce nitrite for respiration and detoxification.

Biotic arsenic oxides accumulate in marine sediment, which may impose environmental stress to microbial inhabitants (80). We found that almost all “Ca. Ypsilonales” MAGs (except Bin295 of clade B) harbored the arsenite methyltransferase gene (arsM) that is responsible to transformation of biotoxic arsenite to methylarsenite (81). The gene coding for arsenite transporter (ACR3) was also identified a few MAGs of different clades. Arsenate reductase (ArsC) was widely found in clade C (5/6 MAGs), but was rarely detected in clades A and B (Fig. 4; see also Table S6). This result indicates that “Ca. Ypsilonales” had the genetic potential to conduct biotransformation of arsenic oxides in the marine sedimentary environment (81).

**Conclusion.** In this study, we describe “Ca. Ypsilonales,” a novel order of *Thermoplasmatota* from deep-sea subsurface sediments collected from SCS. The metabolic potentials of all “Ca. Ypsilonales” clades indicated the utilization of sulfide oxidation in anoxic layers for energy and alkane and butanoate for a carbon source. The ecological importance of “Ca. Ypsilonales” is highlighted by its potential in the degradation of organohalides and aromatic compounds, which might significantly contribute to carbon and energy budgets in these deep-sea settings. Variations in the distribution of “Ca. Ypsilonales” in the deep-sea sites are probably accounted for by different carbon and sulfide sources in the sediments. Overall, this study expands our understanding of *Thermoplasmatota* by shedding light on its role in global element cycles and the biotransformation of marine pollutants.

**MATERIALS AND METHODS**

**Sample collection.** Eight push cores with lengths of 24 to 26 cm were collected from eight different sites in the SCS (Fig. 1; see also Table S1) by the manned submersible *Deep-Sea Warrior* in dives 40, 152, 153, 154, 155, 159, 165, and 166 during R/V *Tansuiyihao* research cruises TS7 (March 2018) and TS12 (June and July 2019). Sampling sites were located in the Northwestern SCS slope at depths ranging from 1,200 to 3,400 m (Fig. 1; see also Table S1). Subsequently, each sediment core was segmented into...
sequential 2-cm layers in the shipboard lab at room temperature. The sediment subsamples obtained were immediately preserved in sterile tubes and frozen at −80°C for analytical measurements.

**Environmental parameter and PAH measurements.** The sediment samples were centrifuged at 5,000 rpm for 10 min (Thermo Fisher, USA) to retrieve pore water (filtered through rinsed 0.45-μm-pore-size cellulose acetate membranes; Millipore, Germany) for the immediate analysis of environmental parameters. The concentrations of ammonium (NH₄⁺), nitrate (NO₃⁻), nitrite (NO₂⁻), and soluble reactive phosphate (PO₄³⁻) in the sediment pore waters were determined with a nutrient AutoAnalyzer (Seal, Norderstedt, Germany). The total organic carbon (TOC) and total nitrogen (TN) contents in the sediments were measured with a Vario Micro Cube elemental analyzer (Elementar, Langenselbold, Germany).

In order to measure the 16 priority PAHs in sediment, the wet sediment samples were processed by the freeze drying for 72 h to eliminate pore water, and then the grinded sediment was filtered with dichloromethane via a preconditioned solid-phase extraction C₁₈ cartridge. After drying by nitrogen blowing, the resulting residue was redissolved with 1 mL of dichloromethane (82, 83). The analyses were performed on an Agilent 7890A GC (Agilent, USA) in the Analytical and Testing Center of Hainan University.

**DNA extraction and metagenome sequencing.** The total DNA of the sediment layer was extracted using a PowerSoil DNA isolation kit (MOBIO, Germany) according to the manufacturer's instructions. The quality and quantity of genomic DNA and cDNA were checked by gel electrophoresis. DNA concentration was determined using a Qubit dsDNA HS assay kit with a Qubit 2.0 fluorometer (Invitrogen, Carlsbad, CA). DNA samples with a low concentration of <2 ng/μL were concentrated using AMPure XP 546 beads (Beckman Coulter, CA) prior to library preparation. A total of 100 ng of DNA was randomly fragmented to ~350 bp by Covaris M220 Focused ultrasonicator (Covaris, MA) using an Illumina TruSeq Nano DNA sample prep kit (Illumina, San Diego, CA). Libraries were sequenced on a Novaseq6000 platform to produce paired-end reads (ILLUMINA) (v1.1) (85) to remove duplicated reads (see Table S2).

The qualified reads for each sediment layer were merged for assembly using Megahit (v1.2.8) (86) with a kmer range of 21 to 141 and k-step of 10 (−min-contig-len 300 -m 0.9 −k-min 21 −k-max 141 −k-step 10) to achieve the best assembly result. MetaWRAP v1.2.1 (87) was used for genome binning and subsequent refinement. The MAGs were checked by CheckM v1.0.11 (88) to filter those with low completeness (<50%) and high contamination (>10%). Taxonomic annotation of the MAGs was carried out using GTDB-tk software v0.22.3 (31). Then, the MAGs belonging to Thermoplasmatota archaea were retrieved for downstream analyses. The relative evolutionary divergence (RED) was calculated when a query genome could not be classified based on the ANI via GTDB-tk software.

**Genome annotation.** Reference Thermoplasmatota genomes were downloaded from the NCBI. Open reading frames of our MAGs and reference genomes were predicted by Prodigal (v 2.6.2) with option “-p meta” (89). The predicted genes were annotated by GhostKOALA online service (https://www.xekgp.jip/ghostkoala, v2.2) and KofamScan (90) (version 1.1.0) against KEGG databases (91), by HMM search against protein families downloaded from Pfam v34.0 (53), and by eggNOG-mapper v1.0.3 (92).

**Revealing microbial communities using 16S rRNA gene amplicons.** The V3-V4 region of 16S rRNA genes was amplified using a pair of universal primers: forward primer 341F (5' - CCTACGGGGBGCASCAG-3') and reverse primer 802R (5' - ATTACCCTTGTTACGACTT-3'). The PCR products of partial 16S rRNA genes were purified using a Cycle-Pure kit (Omega, Norcross, GA) to degrade excess primers and nucleotides. An Illumina library for the equally pooled 16S rRNA gene amplicons was prepared by using a TruSeq Nano DNA LT kit (Illumina) and sequenced on an Illumina MiSeq platform.

The raw sequencing data of 16S rRNA gene amplicons were trimmed using an NGS QC Toolkit (v2.3.3) (95), and those with a low-quality score of <20 and shorter than 100 bp were discarded. The paired-end reads were assembled using multiple_join_paired_ends.py from QIIME (v1.9.1) with the default parameters. Afterward, the assembled reads were analyzed by using QIIME (v1.9.1) pipeline (96).

The 16S rRNA gene amplicons were clustered to the OTUs using UCLUST (v1.0.0) with a similarity threshold of 97% (97). The representative reads of the OTUs were used for further taxonomic classification with the SILVA 132 (98) database using UCLUST (97) at a similarity threshold of 97%. The OTUs assigned to mitochondria, chloroplasts, and eukaryotes were removed, and the singletons were also ignored. The RBG-16-68-12 OTUs were further confirmed by the phylogenetic relationships with the reference 16S rRNA genes and those extracted from the RBG-16-68-12 MAGs using a threshold similarity of 97% in BLASTN searching result.

**Phylogenetic analyses.** Phylogenetic analysis of 16S rRNA genes was performed with the sequences from the MAGs and NCBI-nt database, with 16S rRNA sequences of MG II and III groups from SILVA SSU database (version 132) serving as the outgroup. All 16S rRNA gene sequences were aligned with MAFFT v7.407 (99) with default setting, and sequences were trimmed via trimAl (v1.4) (100). Finally, a phylogenetic tree of 16S rRNA genes was constructed by the maximum-likelihood (ML) algorithm (GTR/Gamma model) using IQ-TREE 2 (-m MFP -B 1000 -alrt 1000 -T AUTO -bnni (101).

Phylogenomic analysis of Thermoplasmatota genomes was conducted with 43 ribosomal proteins identified by CheckM program with default setting (88). Alignment of the obtained conserved protein...
sequences from the MAGs and reference sequences from NCBI SRA database was performed using MAFFT v7.407 (99) with the default settings, and poorly aligned regions were removed by trimAI (v1.4) (100). An ML phylogenomic tree was constructed using the concatenated aligned protein sequences with IQ-TREE tool (101) and the best-fit substitution model (LG + R10 model) for 1,000 replicates. In addition, all phylogenetic trees were visualized using the interactive Tree Of Life (ITOL) tool v6 (102).

**Genome abundance and prevalence.** The relative abundance of MAGs in the metagenomes was estimated by the recruitment rates of the qualified reads by BWA (103) and CoverM (with the settings -m relative_abundance -min-read-aligned-length 50 –min-read-percent-identity 0.99 –min-covered-fraction 0.1 –proper-pairs-only in genome mode) v0.6.1 (https://github.com/wwood/CoverM).

**Data availability.** The MAGs obtained from the marine sediment samples in this study have been submitted to the NCBI database under BioProject ID PRJNA757197.

**SUPPLEMENTAL MATERIAL**

Supplemental material is available online only.

**FIG S1**, TIF file, 0.6 MB.

**FIG S2**, TIF file, 0.3 MB.

**FIG S3**, TIF file, 0.5 MB.

**FIG S4**, TIF file, 0.5 MB.

**TABLE S1**, DOCX file, 0.01 MB.

**TABLE S2**, DOCX file, 0.01 MB.

**TABLE S3**, DOCX file, 0.01 MB.

**TABLE S4**, DOCX file, 0.01 MB.

**TABLE S5**, DOCX file, 0.01 MB.

**TABLE S6**, XLSX file, 0.1 MB.

**ACKNOWLEDGMENTS**

We thank the Supercomputing Center of University of Sanya for providing the computation assistance. We also thank the Hainan Province Postdoctoral Science Foundation and the CAS Special Research Assistant Program.

This study was financially supported by the Natural Science Foundation of Hainan Province (420MS077).

**REFERENCES**

1. Orsi WD, Richards TA, Francis WR. 2018. Predicted microbial secretomes and their target substrates in marine sediment. Nat Microbiol 3:32–37. https://doi.org/10.1038/s41564-017-0047-9.

2. Arndt S, Jørgensen BB, LaRowe DE, Middelburg JJ, Pancost RD, Regnier P. 2013. Quantifying the degradation of organic matter in marine sediments: a review and synthesis. Earth Sci Rev 123:33–86. https://doi.org/10.1016/j.earscirev.2013.02.008.

3. Flemming H-C, Wurzt S. 2019. Bacteria and archaea on Earth and their abundance in biosilts. Nat Microbiol 17:247–260. https://doi.org/10.1038/s41556-019-0158-9.

4. Lipp JS, Morono Y, Inagaki F, Hinrichs K-U. 2008. Significant contribution of Archaea to extant biomass in marine subsurface sediments. Nature 454:991–994. https://doi.org/10.1038/nature07174.

5. Steen AD, Crits-Christoph A, Carini P, DeAngelis KM, Fierer N, Lloyd KG, A, House CH, Hinrichs K-U. 2006. Heterotrophic archaea dominate sediment. Nat Rev Microbiol 11:83–94. https://doi.org/10.1038/nrmicro2939.

6. Hoehler TM, Jørgensen BB. 2013. Microbial life under extreme energy limitation. Nature 499:1–10. https://doi.org/10.1038/nature12527.

7. D’Hont D, Jørgensen BB, Miller DJ, Batzke A, Blake R, Cragg BA, Cypionka H, Dickens GR, Fiedelken T, Hinrichs K-U, Holm NG, Mertzerer R, Spivack A, Wang G, Bekins B, Engelen B, Ford K, Gettemy G, Humphrey F, Skelby SC, Säss H, Skillebeck CG, Alhio MW, Güerin G, House CH, Inagaki F, Meister P, Naehr T, Nishiyama S, Parkes RJ, Schippers A, Smith DC, Teske A, Wiegel J, Padilla CN, Acosta JLS. 2004. Distributions of microbial activities in marine sediments. Science 306:2216–2221. https://doi.org/10.1126/science.1011155.

8. Biddle JF, Lipp JS, Lever MA, Lloyd KG, Sørensen KB, Andersen R, Fredricks HF, Elvert M, Kelly TJ, Schrag DP, Sogin ML, Brenchley JE, Teske A, House CH, Hinrichs K-U. 2006. Heterotrophic archaea dominate deep subsurface ecosystems off Peru. Proc Natl Acad Sci U S A 103:3846–3851. https://doi.org/10.1073/pnas.0600035103.
55. Judger B-E, Ertan H, Lee M, Manefield M, Marquis CP. 2015. Reductive dehalogenases come of age in biological destruction of organohalides. Trends Biotechnol 33:595–610. https://doi.org/10.1016/j.tibtech.2015.07.004.

56. Jassend DB, Pries F, van der Ploeg JR. 1994. Genetics and biochemistry of dehalogenating enzymes. Annu Rev Microbiol 48:163–191. https://doi.org/10.1146/annurev.mi.48.070194.001115.

57. Greenhalgh ED, Kunze C, Schubert T, Diekert G, Brunold TC. 2021. A spec-troscopically validated computational investigation of viable reaction intermediates in the catalytic cycle of the reductive dehalogenase PceA. Biochemistry 60:202–2032. https://doi.org/10.1021/acs.biochem.9b02717.

58. Bossert ID, Häggblom MM, Young LY. 2003. Microbial ecology of dehalo-genation, p 33–52. In Häggblom MM, Bossert ID (ed), Dehalogenation: microbial processes and environmental applications. Springer, Boston, MA. https://doi.org/10.1007/0-306-48011-5_2.

59. Kawai M, Futagami T, Toyoda A, Takaki Y, Nishi S, Hori S, Arai W, Tsushouichi T, Morono Y, Uchiyama I, Ito T, Fujyuma I, Inagaki F, Takami H. 2014. High frequency of phylogenetically diverse reductive dehalogena-nase-homologous genes in deep subsurface sedimentary metage-nomes. Front Microbiol 5:80. https://doi.org/10.3389/fmicb.2014.00800.

60. Payne KAP, Quezada CP, Fisher K, Dunstan MS, Collins FA, Sjuts H, Levy T, Hay J, Hovmöller RU. 2016. Reconstructing metabolic pathways of a member of the genus Pelotomaculum ruw254. ISME J 15:1055–1066. https://doi.org/10.1038/ismej.2015.215.

61. Li L, Patterson DP, Fox CC, Lin B, Coschigano PW, Marsh ENG. 2009. Subunit structure of benzenesulfinate synthase and benzoyl-CoA pathway. ISME J 15:47–58. https://doi.org/10.1038/ismej.2008.176.

62. Mai Z, Wang L, Li Q, Sun Y, Zhang S. 2020. Biodegradation and metabolic pathway of phenanthrene by a newly isolated bacterium Gordonia sp. isolates, single cells, and metagenomes. Microbiome 6:158. https://doi.org/10.1186/s40128-018-0599-3.

63. Li L, Diouf S, Green-Diniz R, Sullivan O, Greene-Diniz R, Cole JR, Ross RP, Pusztai A, Bierne N, Janssens S, De Vos P, Fournier D, Mercado-Blanco J. 2018. Distribution of arsenic resistance genes in prokaryotes. Front Microbiol 9:2473. https://doi.org/10.3389/fmicb.2018.02473.

64. Li P, Cao J, Diao X, Wang B, Zhou H, Han Q, Zhang P, Li Y. 2015. Spatial distribution, sources and ecological risk assessment of polycyclic aromatic hydrocarbons in surface seawater from Yangpu Bay, China. Mar Pollut Bull 93:53–60. https://doi.org/10.1016/j.marpolbul.2015.02.015.

65. Liu L, Zhen X, Wang B, Li Y, Sun X, Tang J. 2020. Legacy and novel halogenated flame retardants in seawater and atmosphere of the Bohai Sea: spatial trends, seasonal variations, and influencing factors. Water Res 184:11617. https://doi.org/10.1016/j.watres.2020.11617.

66. Chen S, Zhou Y, Chen Y, Gu J. 2018. fastp: an ultra-fast all-in-one FASTQ preprocessor. Bioinformatics 34:4884–4890. https://doi.org/10.1093/bioinformatics/bty560.

67. Xu H, Luo X, Qian J, Pang X, Song J, Qian G, Chen J, Chen S. 2012. FastUMI: a fast de novo duplicates removal tool for paired short reads. PLoS One 7:e32249. https://doi.org/10.1371/journal.pone.0032249.

68. Li D, Liu C-M, Luo R, Watanabe K, Tan H-T. 2015. MEGAHIT: an ultra-fast single-node solution for large and complex metagenomes assembly via succinct de Bruijn graph. Bioinformatics 31:1674–1676. https://doi.org/10.1093/bioinformatics/btv333.

69. Uritsky GV, DiRuggiero J, Taylor J. 2018. MetaWRAP: a flexible pipeline for genome-resolved metagenomic data analysis. Microbiome 6:158. https://doi.org/10.1186/s40168-018-0541-1.

70. Facklam DP, Imelfort M, Skennerton CT, Hugenholtz P, Tyson GW. 2015. CheckM: assessing the quality of microbial genomes recovered from isolates, single cells, and metagenomes. Genome Res 25:1043–1055. https://doi.org/10.1101/gr.190714.117.

71. Hyatt D, Chen G-L, LoCascio PF, Land ML, Larimer FW, Hauser LJ. 2010. Prodigal: prokaryotic gene recognition and translation initiation site identification. BMC Bioinformatics 11:119. https://doi.org/10.1186/1471-2105-11-119.

72. Arakami T, Blanc-Mathieu R, Endo H, Ohkubo K, Kanehisa M, Goto S, Ogata H. 2020. KofamKOALA: KEGG Ortholog assignment based on profile HMM and adaptive score threshold. Bioinformatics 36:2251–2252. https://doi.org/10.1093/bioinformatics/btz989.

73. Kanehisa M, Sato Y, Morishima K. 2016. BlastKOALA and GhostKOALA: KEGG tools for functional characterization of genome and metage-nome sequences. J Mol Biol 428:726–731. https://doi.org/10.1016/j.jmb.2018.01.006.

74. Huerta-Cepas J, Forslund K, Coelho LP, Szlaczky D, Jensen LJ, von Mering C, Bork P. 2017. Fast genome-wide functional annotation through orthology assignment by eggNOG-Mapper. Mol Biol Evol 34:2112–2122. https://doi.org/10.1093/molbev/msx148.

75. Clausen MS, Wang Q, O’Sullivan O, Greene-Diniz R, Cole JR, Ross RP, O’Toole PW. 2010. Comparison of two next-generation sequencing technologies for resolving highly complex microbiota composition using tandem variable 16S rDNA gene regions. Nucleic Acids Res 38:e200. https://doi.org/10.1093/nar/gkq873.

76. Klimoworth A, Pruesse E, Schweer T, Peples J, Quast C, Horn M, Glöckner FO. 2013. Evaluation of general 16S ribosomal RNA gene PCR primers for
classical and next-generation sequencing-based diversity studies. Nucleic Acids Res 41:e1. https://doi.org/10.1093/nar/gks808.

95. Patel RK, Jain M. 2012. NGS QC Toolkit: a toolkit for quality control of next generation sequencing data. PLoS One 7:e30619. https://doi.org/10.1371/journal.pone.0030619.

96. Caporaso JG, Kuczynski J, Stombaugh J, Bittinger K, Bushman FD, Costello EK, Fierer N, Peña AG, Goodrich JK, Gordon JJ, Huttley GA, Kelley ST, Knights D, Koenig JE, Ley RE, Lozupone CA, McDonald D, Muegge BD, Pirrung M, Reeder J, Sevinsky JR, Turnbaugh PJ, Walters WA, Widmann J, Yatsunenko T, Zaneveld J, Knight R. 2010. QIIME allows analysis of high-throughput community sequencing data. Nat Methods 7:335–336. https://doi.org/10.1038/nmeth.f.303.

97. Edgar RC. 2010. Search and clustering orders of magnitude faster than BLAST. Bioinformatics 26:2460–2461. https://doi.org/10.1093/bioinformatics/btp461.

98. Quast C, Pruesse E, Yilmaz P, Gerken J, Schweer T, Yarza P, Peplies J, Glöckner FO. 2013. The SILVA ribosomal RNA gene database project: improved data processing and web-based tools. Nucleic Acids Res 41:DS90–DS96. https://doi.org/10.1093/nar/gks1219.

99. Rozewicki J, Li S, Amada KM, Standley DM, Katoh K. 2019. MAFFT-DASH: integrated protein sequence and structural alignment. Nucleic Acids Res 47:W5–W10. https://doi.org/10.1093/nar/gkz342.

100. Capella-Gutiérrez S, Silla-Martínez JM, Gabaldón T. 2009. trimAl: a tool for automated alignment trimming in large-scale phylogenetic analyses. Bioinformatics 25:1972–1973. https://doi.org/10.1093/bioinformatics/btp348.

101. Minh BQ, Schmidt HA, Chernomor O, Schrempf D, Woodhams MD, von Haeseler A, Lanfear R. 2020. IQ-TREE 2: new models and efficient methods for phylogenetic inference in the genomic era. Mol Biol Evol 37:1530–1534. https://doi.org/10.1093/molbev/msaa015.

102. Letunic I, Bork P. 2019. Interactive Tree Of Life (iTOL) v4: recent updates and new developments. Nucleic Acids Res 47:W256–W259. https://doi.org/10.1093/nar/gkz239.

103. Li H, Durbin R. 2009. Fast and accurate short read alignment with Burrows-Wheeler transform. Bioinformatics 25:1754–1760. https://doi.org/10.1093/bioinformatics/btp324.