Leaf nitrogen and phosphorus of temperate desert plants in response to climate and soil nutrient availability

Mingzhu He1, Feike A. Dijkstra2, Ke Zhang1, Xinrong Li1,3, Huijuan Tan1,3, Yanhong Gao1 & Gang Li1

1Shapotou Desert Research and Experiment Station, Cold and Arid Regions Environmental and Engineering Research Institute, Chinese Academy of Sciences, Lanzhou, 730000, China, 2Department of Environmental Sciences, Centre for Carbon, Water and Food, The University of Sydney, NSW, 2006, Australia, 3Key Laboratory of Stress Physiology and Ecology in Cold and Arid Regions of Gansu Province, Lanzhou, 730000, China.

In desert ecosystems, plant growth and nutrient uptake are restricted by availability of soil nitrogen (N) and phosphorus (P). The effects of both climate and soil nutrient conditions on N and P concentrations among desert plant life forms (annual, perennial and shrub) remain unclear. We assessed leaf N and P levels of 54 desert plants and measured the corresponding soil N and P in shallow (0–10 cm), middle (10–40 cm) and deep soil layers (40–100 cm), at 52 sites in a temperate desert of northwest China. Leaf P and N showed a negative relationship with mean annual temperature (MAT) and no relationship with mean annual precipitation (MAP), but a positive relationship with soil P. Leaf P of shrubs was positively related to soil P in the deep soil. Our study indicated that leaf N and P across the three life forms were influenced by soil P. Deep-rooted plants may enhance the availability of P in the surface soil facilitating growth of shallow-rooted life forms in this N and P limited system, but further research is warranted on this aspect.

Nitrogen (N) and phosphorus (P) availability affect community structure1–2, species diversity3–5, and other ecosystem functions6, such as nutrient cycling and productivity7–8, and the ratio of these two elements in leaf tissue may indicate if a system is limited by N, P or both8. Studies about ecological stoichiometry of N and P have been performed across local10, regional8,11 and global scales12–14. Plant N and P levels can be influenced by various biotic and abiotic factors, such as habitat15, growth stages16, and plant functional groups8.

N- and P-limitation are typically determined by plant nutrient levels and/or soil nutrient availability17. In turn, plant adaptations to soil nutrient levels may exert control over critical N:P values. Studies from terrestrial plant species of China revealed that due to low soil P levels compared to the global average, overall leaf N:P ratios were markedly higher than that of global flora8. However, it was suggested that critical leaf N:P values could not effectively predict nutrient limitation of desert plants. Due to water and nutrient co-limitations and adaptation to low nutrient conditions, desert plants show little plasticity in N:P stoichiometry17 and maintain low tissue nutrient uptake18. This unique pattern highlights the need of considering soil nutrient conditions for plant specific adaptation and plant-soil interactions in desert environments.

In desert ecosystems, low soil moisture coupled with high soil alkalinity acts to decrease both soil N and P availability19. Infrequent and low precipitation limits soil weathering, organic matter production, and mineralization20, leading to slow P release from primary material, low soil organic matter content, and N bound in organic matter13. A study from 224 dryland sites indicated an increased decoupling of carbon (C), N and P with increased aridity resulting in greater P availability compared to N24. Plant N fixation rates in arid regions have long been considered to be low because of low soil moisture and high temperatures21. In contrast, ammonia volatilisation of dryland soils can be high, as volatilisation rates are positively related to soil pH, total salt content and CaCO3, and negatively related to soil organic matter, cation-exchange capacity and clay content22. There is also good evidence that sometimes nutrients, with limited water, can limit plant growth23, although relationships between plant growth and leaf N or P concentration are not always clear16.

What is the intrinsic relationship between soil nutrients and leaf nutrients of desert plants? Soil nutrients are the main driver for leaf nutrient concentrations. Plant available soil P, primarily derived from weathering of primary materials such as apatite, and from dissolution and diffusion of P within the soil solution, is considered lower than that of N21. Available forms of N and P mostly remain in the surface soil because the high temperature...
and low rainfall regimes characteristic of desert ecosystems prevent nutrients to be leached to deeper soil layers\(^{22}\). Desert plants play a vital role in vertically redistributing soil nutrients because nutrients intercepted by roots from soil at depth are recycled to the soil surface layer by throughfall and litterfall\(^{24}\). Jobbágy and Jackson\(^{24}\) found that, as soil nutrients become scarce, that these limiting nutrients are transferred from deeper soil layers. Plant characteristics, such as tissue stoichiometry, above- and below-ground allocation of biomass, nutrient cycling rates, rooting depths and redistribution, may play a crucial role in shaping soil nutrient profiles\(^{24–26}\). However, to our knowledge, limited studies coupled the nutrient distribution in soil profiles to plant life forms and leaf nutrient stoichiometry, and explained the intrinsic relationships between them.

Here, we tested the response of leaf \(N\) and \(P\) stoichiometry to climate, soil \(N\) and \(P\) pools, for different life forms at 52 sites in the Alxa desert, northwest China. Our objectives were to understand the effects of potential drivers (climate, soil \(N\) and \(P\)) on leaf \(N\) and \(P\) stoichiometry of life forms in this temperate desert. We address the following questions: Under \(N\) - and \(P\)-limited conditions, how do leaf \(N\) and \(P\) concentrations and stoichiometry vary among different life forms (annuals, perennials, shrubs) of the temperate desert flora? How do leaf \(N\) and \(P\) vary with mean annual temperature (MAT) and mean annual precipitation (MAP) and with \(N\) and \(P\) levels in the soil? How are plant nutrient concentrations among life forms affected by the vertical distribution of soil nutrients?

**Results**

The concentrations of leaf \(N\), \(P\) and \(N: P\) mass ratio of the desert plants varied markedly, ranging between 1.20 and 37.4 mg g\(^{-2}\) for \(N\); 0.11 and 4.51 mg g\(^{-1}\) for \(P\), and 2.55 and 36.9 for \(N: P\). Leaf \(P\) was higher in annuals and perennials than in shrubs, while leaf \(N: P\) in shrubs was higher than in annuals and perennials (Table 1). Among these traits, greater variation was found in leaf \(N\) and \(P\) (overall mean CV of 0.62 and 0.67 respectively) than in leaf \(N: P\) (CV = 0.37). Among all species, leaf \(N\) was positively correlated to leaf \(P\) (Table 1).

Across all sites, total soil \(N\) and \(P\) were strongly correlated (Figure 1). Total soil \(N\) and \(P\) concentrations were highest at the surface and decreased with depth (Figure 2). On the other hand, total soil \(N: P\) did not show a clear pattern with soil depth and there were no significant differences among the three layers.

Life form effects on leaf \(N\), \(P\) and \(N: P\) did not depend on MAT or MAP at the site (i.e., no significant interactions with MAT or MAP in the ANCOVA), but the covariate MAT had a significant effect on leaf \(N\) mass (Table 1). Leaf \(P\) and \(N\) (averaged across species for each site) significantly decreased with increased MAT (Figure 3). Leaf \(N\) and \(P\) showed no relation with MAP.

Life form effects on leaf \(N\), \(P\) and \(N: P\) did not depend on total soil \(N\) concentration at 0–10, 10–40, or 40–100 cm soil depth (no significant life form*soil \(N\) interactive effects in the ANCOVA). Leaf \(N\), \(P\), and \(N: P\) were also not related to total soil \(N\) at any of the three soil depths (no significant covariate effect, data not shown). On the other hand, leaf \(N\) showed positive relationships with soil \(P\) at 0–10 and 10–40 cm soil depth (Table 3). Life form effects on leaf \(P\) depended on soil \(P\) at 0–10 and 10–40 cm soil depth (significant life form*soil \(P\) interactive effects in the ANCOVA) (Table 3). Leaf \(P\) was positively related to soil \(P\) in 0–10, 10–40 and 40–100 cm soil depth. Life form effects on leaf \(N: P\) did not depend on total soil \(P\) concentration at 0–10, 10–40, or 40–100 cm soil depth (no significant life form*soil \(P\) interactive effects in the ANCOVA) (Table 3). We conducted linear regressions between total soil \(P\) concentration at 0–10, 10–40 and 40–100 cm soil depth and leaf \(P\) for each life form separately to examine which life forms caused significant relationships with total soil \(P\). Only regressions between total soil \(P\) and leaf \(P\) were significant, and only at specific soil depths for each life form (Figure 4).

**Discussion**

Leaf \(N\) and \(P\) stoichiometry of typical desert plants, across the Alxa desert of China, were analysed in this study. The average leaf \(N\) and \(P\) concentration of all 54 species was 10.9 and 1.13 mg g\(^{-1}\), respectively, lower than reported in other studies (Table 4). The average leaf \(N\) : \(P\) ratio of plants in the Alxa desert was 10.8, also lower than that reported elsewhere (Table 4). The leaf \(N\) : \(P\) mass ratio has been used to indicate plant \(N\) or \(P\) limitation conditions (i.e., \(N\) limitation when \(N: P < 14\), \(P\) limitation when \(N: P > 16\)). The low \(N\) : \(P\) ratios in our study suggests that the desert flora in the Alxa desert were more \(N\)-limited than \(P\)-limited (Table 4). In general, desert ecosystems tend to have a high \(N\)-efficiency due to limited plant \(N\) fixation\(^{27,28}\) and limited \(N\) losses as dissolved organic \(N\)^\(^{29}\) or as gaseous \(N\)^\(^{22,30}\). Nevertheless, the extremely low plant \(N\) and \(P\) concentrations suggest that both nutrients were in low supply in this temperate desert.

Plant \(P\) concentrations are often positively related to soil \(P\) concentrations at the ecosystem scale\(^{27,31}\). Recent studies focusing on \(N\) and \(P\) stoichiometry of Chinese flora\(^{10,11}\) and soil\(^{32,33}\) have indicated that an underlying cause for the low \(P\) concentrations of Chinese plants is the relatively low soil \(P\) content\(^{32–33}\). Soil \(P\) concentrations of Chinese soils are lower compared to global levels\(^{31}\); however average total soil \(P\) content of the Alxa desert (0.70 and 0.61 mg g\(^{-1}\) at 0–10 and 10–40 cm soil depth) was above the average of that in other parts of China\(^{4}\). A large fraction of the total soil \(P\) of the Alxa desert was in inorganic form that is mostly unavailable to plants (e.g., insoluble calcium phosphates\(^{34}\)).

Some significant differences in leaf \(P\) and \(N: P\) were observed among plant life forms of the desert flora, but not in leaf \(N\) (Table 1). Leaf \(P\) concentrations were significantly higher (and leaf \(N: P\) lower) in annuals and perennials than in shrubs, consistent with other studies where higher \(P\) concentrations were observed in

| Functional group | \(N\) (mg g\(^{-1}\)) Mean CV | \(P\) (mg g\(^{-1}\)) Mean CV | \(N: P\) mass ratio | \(rs\) |
|------------------|-----------------|-----------------|-----------------|-------|
| Shrub            | 51              | 11.16*          | 0.63            | 51    | 0.91*          | 0.63          | 51    | 13.18*          | 0.25          | 0.86*          |
| Annual           | 32              | 10.39*          | 0.62            | 32    | 1.26*          | 0.61          | 32    | 9.33*          | 0.39          | 0.85*          |
| Perennial        | 31              | 10.88*          | 0.64            | 31    | 1.37*          | 0.67          | 31    | 8.57*          | 0.41          | 0.83*          |
| ANOVA P-value    | 0.93            | 0.68            | 0.02            | <0.0001 | 0.02          | | 0.01          | | 0.86          | |

1) Different superscript letters (a, b) indicate significant differences (Tukey’s HSD test, \(P < 0.05\)) in the mean value. \(N\), sample size; CV, coefficient of variation.
2) Spearman’s rank correlation coefficient \(rs\) between species-specific leaf \(N\) and \(P\) was calculated for each life form. Correlations with *, \(P < 0.001\).
Figure 1 | Relationships between soil total nitrogen (TN) and total phosphorus (TP) concentrations at three soil layers. Spearman’s correlation efficient between TN and TP were 0.548 at 0–10 cm (P < 0.0001), 0.513 at 10–40 cm (P < 0.0001) and 0.526 at 40–100 cm (P < 0.0001).

Figure 2 | Soil N, P and N : P with soil depth. The values of (a) total soil nitrogen (TN); (b) total soil phosphorus (TP) and (c) soil N : P ratio averaged by all sample sites (n = 52). Letters indicate significant differences of the mean ± SE at P ≤ 0.05 using Tukey’s HSD test.
fast-growing, short-lived species than in slow growing, long-lived species3,35. Fast growing plants often require more P-rich RNA to support synthesis of proteins necessary for rapid growth36. Possibly, P uptake from deeper soil layers and redistribution to shallow soil layers by slow-growing shrubs may have contributed to meet the relatively high leaf P requirement in fast-growing annuals and perennials in this P-poor system.

Many studies focused on the broad biogeographic patterns with climatic indices (e.g., MAT or MAP). At the global scale, Reich and Oleksyn34 identified that leaf P, and to a lesser degree leaf N, decreased as MAT increased and that leaf N:P increased as MAT increased41. These global patterns differed among taxonomic groups and were independent of changes in species composition35 and were confirmed by studies in Chinese terrestrial flora and grasslands36,37.

In contrast, leaf N, P and N:P of the desert flora in this study were unrelated to MAP, while relationships with MAT were weak (Figure 3). Possibly, the range in MAT (ranging from 5.8 to 8.8 °C) and MAP (from 40 to 200 mm) in our area was too small to detect strong patterns with leaf N and P. Leaf N and P stoichiometry of different desert plants can fluctuate remarkably in time because of seasonal climate variation, obscuring spatial relationships between leaf N and P stoichiometry and mean annual climate parameters38.

Leaf N and P were significantly related to total soil P (Table 3, Figure 4), but not to total soil N, suggesting that total soil P availability influences plant uptake of both N and P more than total soil N in this system. Total soil N in desert ecosystems may not be a good indicator of N availability to plants. In desert ecosystems, important N sources include ammonium (NH₄⁺) and nitrate (NO₃⁻) derived from rainfall, aeolian deposition with nitrate dust, and biological assimilation by N-fixing soil organisms39. Nitrogen uptake by plants is affected by temperature, soil water condition, microbial community and N-fixation. All these factors may result in total soil N being a poor indicator of N availability. However, P can be an equal or a more important limiting factor in many desert ecosystems compared to N, especially in calcareous soils with high pH40,41. Several studies from regional to global scales found that leaf P, for various plant groups and in different geographical regions, can be more variable and more strongly related to climate conditions and soil nutrients compared to leaf N or N:P42,43. Our results suggest that, leaf N and P are at least partially controlled by total soil P levels in the Alxa desert.

Because most N and P sources, derived from dust, rock weathering, and the decomposition of litter, often occur in the topsoil44, higher N and P concentration are found in surface soils than that of deep soil in most environments. This phenomenon is pronounced in desert ecosystems because insufficient rainfall limits N and P leaching downward into soil45, and is supported by decreasing soil N and P with depth in this study (Figure 2). This decrease in soil N and P with depth can be enhanced by the desert vegetation in this system. The relationship between leaf P and soil P at 40–100 cm suggests that shrubs may have taken up P (and N to a lesser degree) from deeper soil layers and returned these nutrients to the surface soil via litterfall. By enhancing P availability in the surface soil, deep-rooted shrubs could then relieve P limitation of shallow rooted and fast growing annuals and perennials with relatively high N and P requirements44,46. Indeed, leaf P of annuals and perennials were more responsive to soil P at shallower depths. This upwards movement of N and P could have important consequences for the vegetation structure in this N- and P-limited desert ecosystem, possibly facilitating co-existence of different plant life forms. We did not measure the rooting depths of the different life forms that could support this notion. However, a global analysis on rooting depths showed that shrubs had relatively more roots at greater soil depth than grasses47. Most studies in desert ecosystems have focused on horizontal nutrient patterns, including the development of “islands of fertility”48. However, desert plants may also play an important role in structuring vertical distributions of soil nutrients49. Soil nutrient distribution has been associated with plant characteristics including nutrient stoichiometry, allocation of above- and belowground biomass, biomass cycling rate and rooting depth50, which could potentially have important consequences for plant community structure.

| Table 2 | Summary of ANCOVA results (P values) for the effects of MAT, MAP (covariates) and life form on leaf N and P stoichiometry |
|-----------------|-----------------|-----------------|-----------------|
| Effect          | Log N           | Log P           | Log N:P         |
| MAT             | 0.04            | 0.01            | -0.30           |
| Life form       | 0.85            | 0.01            | -0.90001        |
| MAT × life form | 0.78            | 0.70            | 0.73            |
| MAP             | 0.54            | 0.91            | 0.25            |
| Life form       | 0.88            | 0.02            | -0.90001        |
| MAP × life form | 0.83            | 0.91            | 0.94            |

P-values are in bold when P < 0.05.
In conclusion, we document low concentrations of N and P in desert plants that have adapted to low nutrient conditions. There was no strong relationship between climate (MAT or MAP) and leaf nutrients, but soil P positively related to leaf N and P. In addition, we found that shrub N and P showed the strongest relationship with P in the deep soil (40–100 cm), while N and P in annuals and perennials were only significantly related to N and P in the shallow soil. Our results suggest that deep-rooted plants may enhance the availability of P in the surface soil facilitating growth of shallow-rooted life forms in this N and P limited system. This mechanism needs to be tested in more comprehensive arid/desert regions where plant cycling dominates the vertical distribution of soil nutrients.

**Methods**

**Site description.** This study was conducted at 52 vegetated sites of the Alxa Desert in northwestern China (Supporting information, Table S1), which is situated in the arid zone of the East Asian continent and spans from 37°24’N, 97°10’E to 42°47’N, 106°53’E. The sampling sites experienced harsh climatic conditions, with MAT...
Table 4 | Comparison of leaf N and P stoichiometry between Alxa desert and other regions

| Data source | N (mg g⁻¹) | P (mg g⁻¹) | N:P (mg g⁻¹) | References |
|-------------|------------|------------|--------------|------------|
| Alxa desert | 10.87 ± 7.91 | 1.13 ± 0.81 | 9.50 ± 6.06 | This study |
| China       | 20.20 ± 8.41 | 1.46 ± 0.99 | 14.00 ± 6.92 | 10 |
| Grassland of China | 27.60 ± 8.60 | 1.90 ± 0.84 | 14.62 ± 6.92 | 14 |
| Global Flora | 20.10 ± 8.71 | 1.77 ± 1.12 | 11.63 ± 6.82 | 13 |
| Alxa desert | 20.60 ± 12.23 | 1.99 ± 1.49 | 10.82 ± 6.32 | 11 |

1. This study

8. Han, W., Fang, J., Guo, D., & Zhang, Y. Leaf nitrogen and phosphorus stoichiometry across 753 terrestrial plant species in China. *New Phytol.* 168, 377–385 (2005).
9. Koerselman, W. & Meuleman, A. F. M. The vegetation N:P ratio: A new tool to detect the nature of nutrient limitation. *J. Appl. Ecol.* 33, 1441–1450 (1996).
10. Zheng, S. X. & Shangee, Z. Spatial patterns of leaf nutrient traits of the plants in the Loess Plateau of China. *Trees-Struct. Funct.* 21, 357–370 (2007).
11. He, J. S. et al. Leaf nitrogen: phosphorus stoichiometry across Chinese grassland biomes. *Oecologia* 155, 301–310 (2008).
12. McGrady, M. E., Daufreune, T. & Hedlin, L. O. Scaling of C:N:P stoichiometry in forests worldwide: Implications of terrestrial redfield-type ratios. *Ecology* 85, 2390–2401 (2004).
13. Elser, J. J. et al. Nutritional constraints in terrestrial and freshwater food webs. *Nature* 408, 578–580 (2000).
14. Reich, P. B. & Oleksyn, J. Global patterns of plant leaf N and P in relation to temperature and latitude. *Proc. Natl. Acad. Sci. U. S. A.* 101, 11001–11006 (2004).
15. Schlesinger, W. H. & Pimentel, A. M. Plant-soil interactions in deserts. *Biogeochemistry* 46, 169–198 (1998).
16. James, J. I., Tiller, R. L. & Richards, J. H. Multiple resource limits plant growth and function in a saline-alkaline desert community. *J. Ecol.* 93, 113–126 (2005).
17. Drenovsky, R. E. & Richards, J. H. Critical N:P values: Predicting nutrient deficiencies in desert shrublands. *Plant Soil* 259, 59–69 (2004).
18. Delgado-Baquetizo, M. et al. Decoupling of soil nutrient cycles as a function of aridity in global drylands. *Nature* 502, 672–676 (2013).
19. Noy-Meir, I. Desert ecosystems: environment and producers. *Annu. Rev. Ecol. Evol. S.* 4, 25–51 (1973).
20. Schlesinger, W. H. *Biogeochemistry: an analysis of global change* (Academic Press, San Diego, 1997).
21. Hartley, A., Barger, N., Belnap, J. & Okin, G. [Dryland ecosystems] *Nutrient cycling in terrestrial ecosystems* [Marschner, P. & Rengel, Z. (eds.)] [271–307] (Springer, Heidelberg, 2007).
22. McCalley, C. K. & Sparks, J. P. Abiotic Gas Formation Drives Nitrogen Loss from a Desert Ecosystem. *Science* 326, 837–840 (2009).
23. Walvoord, M. A. et al. A Reservoir of Nitrate Beneath Desert Soils. *Science* 302, 1021–1024 (2003).
24. Jobbagy, E. G. & Jackson, R. B. The uplift of soil nutrients by plants: Biogeochemical consequences across scales. *Ecology* 85, 2380–2389 (2004).
25. Belnap, J. [Biological phosphorus cycling in dryland regions] *Phosphorus in Action* [Bünemann, E. K., Oberson, A. & Frossard, E. (eds.)] [371–406] (Springer-Verlag Berlin Heidelberg, 2011).
26. Jobbagy, E. G. & Jackson, R. B. The distribution of soil nutrients with depth: global patterns and the imprint of plants. *Biogeochemistry* 53, 51–77 (2001).
27. Aerts, R. & Chapin, F. S. The mineral nutrition of wild plants revisited: A re-evaluation of processes and patterns. *Adv. Ecol. Res.* 30, 1–67 (2000).
28. Rundel, P. W. et al. Seasonal dynamics of nitrogen cycling for a prosopis woodland in the Sonoran Desert. *Plant Soil* 67, 343–353 (1982).
29. Persakis, S. S. & Hedlin, L. O. Nitrogen loss from unpolluted South American forests mainly via dissolved organic compounds. *Nature* 415, 416–419 (2002).
30. Peterjohn, W. T. & Schlesinger, W. H. Nitrogen Loss from Deserts in the Southwestern United States. *Biogeochemistry* 10, 67–79 (1990).
31. Hedlin, L. O. Global organization of terrestrial plant-nutrient interactions. *Proc. Natl. Acad. Sci. U. S. A.* 101, 10849–10850 (2004).
32. Tian, H. Q., Chen, G. S., Zhang, C., Melillo, J. M. & Hall, C. A. S. Pattern and variation of C:N:P ratios in China’s soils: a synthesis of observational data. *Biogeochemistry* 98, 139–151 (2010).
33. National Soil Survey Office of China. Survey of soils in China (Chinese Agriculture Press, 1997).
34. Ma, B., Zhou, Z. Y., Zhang, C. P., Zhang, G. & Hu, Y. J. Inorganic phosphorus fractions in the rhizosphere of xerophytic shrubs in the Alxa Desert. *J. Arid Environ.* 73, 55–61 (2009).
35. Thompson, K., Parkinson, J. A., Band, S. R. & Spencer, R. E. A comparative study of leaf nutrient concentrations in a regional herbaceous flora. *New Phytol.* 136, 679–689 (1997).
36. Matzek, V. & Vitousek, P. M. N : P stoichiometry and protein:RNA ratios in vascular plants: an evaluation of the growth-rate hypothesis. *Ecol. Lett.* 12, 765–771 (2009).
37. Chen, Y. H., Han, W. X., Tang, L. Y., Tang, Z. Y. & Fang, J. Y. Leaf nitrogen and phosphorus concentrations of woody plants differ in responses to climate, soil and plant growth form. *Ecography* 36, 178–184 (2013).

ranging between 5.8 and 8.8 °C, MAP ranging between 40 and 200 mm. There are four main vegetation subtypes, namely, psammophytic vegetation, typical desert vegetation, desert steppe vegetation and steppe/desert vegetation **4**.

**Sampling and measurement.** Collection and measurement of samples were conducted in late August of 2012. The distribution and locations of dominant plant communities have been initially identified based on the Vegetation Map of Inner Mongolia **4**. During the field surveys and plant samplings, the research sites were located away from grazing activity and anthropogenic disturbances. Quantitative surveys of the vegetation was carried out at each site and fully expanded sun-exposed leaves were collected from 5 to 10 individuals for each dominant plant of three life forms (i.e., annual, perennial and shrub). In total, we collected 276 plant samples belonging to 12 families, 41 genera and 54 dominant species (See Table S2) from 52 research sites. The leaf samples were rinsed with deionized water at least two times to reduce the influences of dust or soil. After oven-drying at 60 °C for 72 h, plant samples were ground and then measured for N and P concentrations. Leaf N was measured with a CHNS/O Elemental Analyzer (Pekin-Elmer, USA). Leaf P was measured colorimetrically after H₂SO₄-H₂O₂-HF digestion using the molybdate/stannous chloride method **46**.

At each sampling site, soil samples were randomly collected at three soil depths with three replicates: shallow layer (0–10 cm), middle layer (10–40 cm), and deep layer (40–100 cm). Soil samples in each soil layer were mixed evenly. Air-dried and ground soil samples were sieved using a 100-mesh sieve. Soil total nitrogen (TN) was analysed with a Kjeltec System 2300 Analyzer Unit (Tecator, Höganas, Sweden). Soil total phosphorus (TP) content was determined with the molybdate/ascorbic acid blue method **46** after digestion with HClO₄ and H₂SO₄ acid. The MAP and MAT used in this study were obtained through linear interpolation models based on variables of latitude, longitude, and altitude, using the climate database from the Inner Mongolia Weather Bureau.
38. Kerkhoff, A. J., Enquist, B. J., Ehler, J. J. & Fagan, W. F. Plant allometry, stoichiometry and the temperature-dependence of primary productivity. *Global Ecol. Biogeogr.* 14, 585–598 (2005).
39. Lovelock, C. E., Feller, I. C., Ball, M. C., Ellis, J. & Sorrell, B. Testing the growth rate vs. geochemical hypothesis for latitudinal variation in plant nutrients. *Ecol. Lett.* 10, 1154–1163 (2007).
40. Ordon˜ez, J. C. et al. A global study of relationships between leaf traits, climate and soil measures of nutrient fertility. *Global Ecol. Biogeogr.* 18, 137–149 (2009).
41. Blaser, W. J., Sitters, J., Hart, S. P., Edwards, P. J. & Venterink, H. O. Facilitative or competitive effects of woody plants on understory vegetation depend on N-fixation, canopy shape and rainfall. *J. Ecol.* 101, 1598–1603 (2013).
42. Sitters, J., Edwards, P. J. & Venterink, H. O. Increases of soil C, N, and pools along an *Acacia* tree density gradient and their effects on trees and grasses. *Ecosystems* 16, 347–357 (2013).
43. Jackson, R. B. et al. A global analysis of root distributions for terrestrial biomes. *Ecology* 108, 389–411 (1996).
44. Li, X. R., Tan, H. J., He, M. Z., Wang, X. P. & Li, X. J. Patterns of shrub species richness and abundance in relation to environmental factors on the Alxa Plateau: prerequisites for conserving shrub diversity in extreme arid desert regions. *Sci. China Ser. D* 52, 669–680 (2009).
45. Integrative Investigation Team for Vegetation of Inner Mongolia. *Vegetation of Inner Mongolia* (Science Press, Beijing, 1985).
46. Kuo, S. [Phosphorus] Methods of soil analysis. Part 3. Chemical methods. *Soil Science Society of America, American Society of Agronomy* [Bigham, J. M. (eds.)] [869–919] (Madison, 1996).
47. John, M. K. Colorimetric determination of phosphorus in soil and plant materials with ascorbic acid. *Soil Sci.* 109, 214–220 (1970).

**Acknowledgments**

This work was supported by the National Natural Science Foundation of China (No. 41101054). We gratefully acknowledge the help of students, technicians and field workers in the field investigation.

**Author contributions**

M.Z.H. conceived and designed the experiment. M.Z.H. and F.A.D. analysed the data and wrote the paper. M.Z.H., K.Z., X.R.L., Y.H.G. and G.L. carried out the field investigation and sample analyses, and contributed to the draft manuscript.

**Additional information**

Supplementary information accompanies this paper at http://www.nature.com/scientificreports

**Competing financial interests:** The authors declare no competing financial interests.

**How to cite this article:** He, M. et al. Leaf nitrogen and phosphorus of temperate desert plants in response to climate and soil nutrient availability. *Sci. Rep.* 4, 6932; DOI:10.1038/srep06932 (2014).

This work is licensed under a Creative Commons Attribution-NonCommercial-NoDerivs 4.0 International License. The images or other third party material in this article are included in the article’s Creative Commons license, unless indicated otherwise in the credit line; if the material is not included under the Creative Commons license, users will need to obtain permission from the license holder in order to reproduce the material. To view a copy of this license, visit http://creativecommons.org/licenses/by-nc-nd/4.0/