Experimental Investigations of Interactions between Sand Wave Movements, Flow Structure, and Individual Aquatic Plants in Natural Rivers: A Case Study of Potamogeton Pectinatus L.

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Abstract: Long-duration measurements were performed in two sandy bed rivers, and three-dimensional (3D) flow velocity and bottom elevation changes were measured in a vegetated area and in a clear region of a river. Detailed flow velocity profiles downstream and upstream of a single specimen of Potamogeton pectinatus L. were obtained and the bed morphology was assessed. Potamogeton plants gathered from each river were subjected to tensile and bending tests. The results show that the existence of the plants was influenced by both bottom and flow conditions, as the plants were located where water velocity was lower by 12% to 16% in comparison to clear region. The characteristics of the flow and sand forms depended on the cross-sectional arrangement of the river, e.g., dunes were approximately four times higher in the middle of the river than in vegetated regions near the bank. Furthermore, the studied hydrophytes were too sparse to affect water flow and had no discernible impact on the sand forms’ movements. The turbulent kinetic energy downstream of a single plant was reduced by approximately 25%. Additionally, the plants’ biomechanical characteristics and morphology were found to have adjusted to match the river conditions.

Keywords: aquatic plants; flow velocity measurements; river morphology; acoustic Doppler velocimeter; natural sandy bed river; sand waves; turbulent kinetic energy; aquatic plant biomechanics

1. Introduction

Aquatic plants grow in various configurations in rivers, affecting a variety of biological, chemical, and physical processes [1]. They influence both river morphodynamics and flow hydrodynamics by trapping sediments, preventing erosion, affecting turbulent flow fields, and contributing to overall bed roughness, as reported in numerous studies, e.g., [2–8]. One of the most important conclusions of these studies was that the understanding of flow–biota–sediment interactions requires an interdisciplinary approach. O’Hare et al. [9] emphasized that linking plant traits with physical modeling is an example of such multidisciplinary research, where ecological variability is a valid factor. Furthermore, Reid and Thoms [10] highlighted the need to obtain three-dimensional (3D) velocity and turbulence measurements to properly describe river habitats.

In vegetated rivers, sediment transport and channel morphodynamics are affected by vegetation [1,11]. In a vegetated location, the bed load is expected to diffuse transversely into the vegetation and deposit in the lee side of the patches [12], with increased erosion at the lateral sides of the patches, constituting the so-called scouring effect [13]. Suspended sediment is prone to being deposited in the wake of
vegetation [14]. Bouma et al. [15] observed high turbulence and erosion at the leading edges and sides of patches of epibenthic structures such as *Spartina anglica* tussocks. This intertidal vegetation strongly reduced the water velocity, resulting in sedimentation farther downstream. Moreover, Rominger et al. [13] described how vegetative drag diverges and accelerates the flow, potentially causing erosion at the patch edge, which is similar to the scour patterns observed in the field around individual flow obstructions such as bridge piers. In addition, Schnauder and Sukhodolov [16] investigated the seasonal patch effect on sediment transport and concluded that although plant patches affected the transverse flow profiles near the edges, pool scouring did not occur; however, sediment was observed to accumulate in the recirculation zones. Additional data were reported by Cavedon [17], who, based on laboratory experiments, concluded that for a sufficiently high density of stems, the length of bed forms was not influenced by vegetation density, but only by the distance between plants. Velocity profiles in vegetated channels change throughout the year due to plant growth and senescence, therefore shear velocity influencing beds also varies [18]. Our investigation revealed that the height of the forms is reduced first in the vegetated parts of a river. Then it depends on vegetation density and is strongly coupled to the characteristics of the flow field and sediment properties. However, this phenomenon requires further analysis, particularly in natural field conditions.

From a broader perspective, the sedimentation rates in a river are time-dependent, thus natural aquatic habitat characteristics always depend on flow, biota, and morphodynamic feedback over a certain time scale [5,19]. Consequently, the species living in these habitats are called ecosystem engineers because of their control over flow velocity and sediment deposition [20]. To find patterns in the changing and complex ecosystem of a vegetated river, most researchers have focused on flume experiments (e.g., [8,21–23]), and only a few have conducted similar measurements in natural rivers (e.g., [24,25]). Such experiments are typically conducted using an acoustic Doppler velocimeter (e.g., [24,26–29]), which allows measurement of the velocity field at discrete points over bed forms. Although this device was built for laboratory purposes, it is usable in the field (e.g., [30]), albeit with certain limitations in its use and in interpreting the acquired data [31,32]. One can characterize changes in bed structure using either a discrete or continuous approach [33,34]. In the present study, the former approach was used to characterize bed forms in terms of height, length, and celerity, using the echo-sounding ability of current profiler and velocity profiler.

The main goal of this study was to evaluate changes and differences in sand wave movements and flow in vegetated and unvegetated parts of a channel cross-section with regard to sediment composition, flow dynamics, and bed morphology in two natural lowland rivers. This was achieved by recording these changes with long-duration point measurements of 3D velocity and bed elevation, and by scanning bed morphology in chosen parts of rivers with aquatic plants, which gave sufficient quantitative and qualitative information about the investigated subject. Moreover, the biomechanical features of a chosen aquatic macrophyte were studied to investigate how these characteristics differ between plants collected from two distinct river habitats. Due to the challenges inherent in the task of studying multiple aspects of a river ecosystem simultaneously, this case study also represents an attempt to identify the needs and limitations in a multidisciplinary investigation of flow–biota–sediment interactions in field conditions.

### 2. Materials and Methods

#### 2.1. Equipment

Measurement of the instantaneous three-dimensional (3D) velocity field and recording of the bottom elevation were performed using a Vectrino Profiler (VP) (revision 2779/1.32, Nortek AS, Rud, Norway). The following VP setup parameters were chosen: Recording frequency of 25 Hz, minimum ping algorithm with high power level, and velocity range up to 0.1 m s\(^{-1}\) above the maximum observed longitudinal velocity during a trial. VPs were mounted on a steel platform. In addition, the bathymetry and velocity spatial distribution of the rivers were obtained using an acoustic Doppler current profiler.
(ADCP), RiverSurveyor S5 model (SonTek, San Diego, CA, USA). The grain-size distribution was measured using a laboratory shaker with a standard sieve set (10 mm, 2 mm, 1 mm, 0.50 mm, 0.25 mm, 0.10 mm, and 0.071 mm mesh). Before sieving, the sediment sample was dried at 105 °C.

2.2. Study Sites

The first experiment was conducted on the Jeziorka River on 4 July 2017. This river is located south of Warsaw (52°04′55.2″ N, 21°04′02.4″ E). The vegetation in this location covered no more than 20% of the channel cross-section. The water temperature (~15 °C) and other basic characteristics of the river, as well as bed morphology in regard to actual water depth, were measured using the ADCP (Table 1). Bed sediment samples from 3 random locations across the river were taken to conduct granulometry and were found to consist of fine, moderately sorted sand with $D_{16} = 0.11$ mm, $D_{50} = 0.20$ mm, and $D_{84} = 0.39$ mm, and an inclusive graphic standard deviation equal to 0.848 (Figure 1).

The platform in the first part of the experiment, where the velocities in the whole water depth were measured, was placed parallel to the flow, above a single specimen of $P. pectinatus$. Measurements were conducted at a point located 20 cm upstream of the plant and at points 25 and 50 cm downstream, in the wake of the plant, in which flow is supposedly affected by it. Recording of velocity in each point of the velocity profile lasted 3 min. The hydrophyte was then collected for further biomechanical testing.

In the second part of the experiment, the platform was placed perpendicular to the flow. The first VP was placed 20 cm in front of another specimen of $P. pectinatus$ in a 55 cm deep vegetated region in the river adjacent to the right bank, where the majority of the plants grew (Figure 2). The second device was placed near the middle of the river, where the bed was bare. With both probes positioned 15 cm above the bottom level, the velocities and bottom elevation changes were measured for 1.5 h. A survey of the river bottom adjacent to the right bank revealed many small specimens of $P. pectinatus$ and a substantial amount of debris, mostly in the form of dead wood in the vegetated region.

The second experiment was conducted on the ´Swider River (52°07′59.2″ N, 21°15′41.5″ E) on 1 October 2017. The riverbed in this location was covered with sand and was almost clear of plants, which grew only along the left bank. The water temperature (~11.6 °C) and other basic characteristics of the river, as well as bed morphology regarding actual water depth, were measured using ADCP (Table 1). In addition to $P. pectinatus$, two other species ($Myriophyllum spicatum$ L. and $Potamogeton crispus$ L.) were also found growing downstream at the left bank. The platform was oriented perpendicular to the flow, 1 m from the left bank (Figure 3). The first velocimeter was positioned 10 cm upstream along a 0.6 m long and 1 m wide strip where $P. pectinatus$ specimens were located. The second velocimeter was placed in a clear region 3 m from the first one. The bed elevation and velocity recording lasted 3 h.

Granulometry revealed that the bed sediment in the ´Swider River consisted of medium poorly sorted sand with $D_{16} = 0.14$ mm, $D_{50} = 0.24$ mm, and $D_{84} = 0.47$ mm, and an inclusive graphic standard deviation equal to 1.02 (Figure 1).

Table 1. Basic hydraulic statistics of the studied rivers in the experimental locations.

|                | Mean $U$ (m·s$^{-1}$) | Discharge $Q$ (m$^3$·s$^{-1}$) | Mean $H$ (m) | Width (m) | Reynolds Number $U·H·ν^{-1}$ | Froude Number $U·(H·g)^{-0.5}$ |
|----------------|------------------------|--------------------------------|--------------|-----------|-----------------------------|-------------------------------|
| Jeziorka R.    | 0.30                   | 2.01                           | 0.6          | 11        | $1.58 \times 10^5$         | 0.123                         |
| ´Swider R.     | 0.51                   | 4.26                           | 0.46         | 18        | $1.88 \times 10^5$         | 0.24                          |
Figure 1. Granulometric distribution curves of bed sediment collected from the Jeziorka and Świder Rivers.

Figure 2. Arrangement of the platform with Vectrino Profilers and individuals of *P. pectinatus* during the experiment in the Jeziorka River (second part).

2.3. Plant Characterization

Hydrodynamic measurements were conducted near the same aquatic plant species, *P. pectinatus*, which is a slender, submerged macrophyte [35]. Each field experiment was followed by laboratory measurements of biomechanical traits.
In the Jeziorka River, the chosen individual plant had a length of 1 m and floated almost horizontally approximately 10 to 20 cm above the bottom. Other macrophytes in the vegetated region of the channel grew in a dense patch 5 m in length, and their stems floated higher above the bottom than the tested hydrophyte. In the Świder River, individuals of *P. pectinatus* were much shorter, with an average length of 35 cm. The investigated plants grew in a different pattern than the macrophytes in the Jeziorka River; instead of one slender and long patch, the plants were spread across a 1.5 m wide and 0.6 m long area (Figure 3), with stems floating at a maximum of 5 cm above the riverbed.

**Figure 3.** Arrangement of the platform with Vectrino Profilers and individuals of *P. pectinatus* during the experiment in the Świder River.

### 2.4. Measurements of Biomechanical Traits

Biomechanical tests (3-point bending and tension tests) were conducted using a Tinius Olsen Bench Top Testing Machine, 5ST Model (Tinius Olsen, Redhill, UK), and the data were recorded using Horizon software. The plants were transported to the laboratory and kept in a 112 L tank (more details are available in [36]). After a plant was removed from the aquarium, the stem was cut into 7 cm pieces. Next, the diameter of each sample was measured using a microscope. In the tension tests, short strips of sandpaper were glued to the ends of the samples to prevent them from slipping out of the machine clamps [36]. The measurements were conducted in submerged conditions [36]. In addition, to minimize the time interval between removing a plant from the aquarium and conducting the test, no more than 10 samples were prepared at the same time.

The investigated parameters in the 3-point bending test included maximum force, maximum stress, maximum deflection, flexural strain, flexural rigidity, and flexural modulus, and the tension parameters included breaking force, stress and strain, and Young’s modulus. All parameters were calculated following formulas from Łoboda et al. [36], based on Niklas [37] and ASTM [38].
2.5. Data Processing

The VP was used to record velocities in a profile placed 11 cm above the bottom at the start of the measurement in the Jeziorka River and 15 cm in the Świder River. The depth was chosen to perform measurements in a water volume as close to the bottom as possible, but without danger of burying the probe. In case of profiles upstream and downstream of *P. pectinatus* in the Jeziorka River, each point in the profile represented a 3-min recording, and the vertical distance between the following points was 3 cm. The collected datasets were then processed and filtered before further analysis. Parts of code included in the MITT open source MATLAB algorithms [39] were used for velocity data processing. The time-dependent analysis of the recorded velocity data revealed significant drops in data quality caused by noise that affected certain VP beams. Thus, a significant part of velocity data records was discarded during the second experiment.

To exclude spikes in the velocity records from the VP, a 3-stage method was used. First, from each measurement point containing 30 cells within it, one cell within the “sweet spot” was identified with the best signal-to-noise ratio (SNR) value [32,40,41] and subjected to further processing. Second, points in the recording with correlation below 70% and SNR values below 15 dB were discarded and replaced using linear interpolation. The last step included using the modified 3D phase-space thresholding method filter to remove the spikes [42–44].

After the filtering, the following flow characteristics were calculated: Mean velocity, standard deviation of velocity, turbulent kinetic energy, turbulence intensity, and Reynolds stress (Tables 3 and 4). The turbulent kinetic energy (TKE) was calculated as:

\[
TKE = \frac{1}{2} \left( (U')^2 + (V')^2 + (W')^2 \right)
\]

where \(U', V',\) and \(W'\) are velocity fluctuations in the longitudinal, lateral, and vertical directions, respectively, and the overbars denote mean values. Normal Reynolds stresses, turbulence intensities, and TKE were calculated as noise free, which was possible due to redundant data of 2 independent vertical velocities recorded by VP [45,46]. As the first vertical velocity was obtained from the same pair of beams as longitudinal velocity, it was used for all relevant calculations [41].

The power spectral densities were calculated using scripts implemented in MATLAB, i.e., the Welch [47] method with discrete Fourier transform points equal to 512, with 50% overlap and a Hamming window function [41].

The characteristics of the bed forms were calculated from the data acquired by the current profiler, where the lengths of forms were given as the height from the local maximum crest to the local minimum trough [48]. The mean celerity of the bed forms was estimated from the current profiler and VP records and was calculated using the equation proposed by Nikora et al. [33]:

\[
C = \frac{0.666 U Fr^{2.9}}{\lambda^2}
\]

where \(\overline{U}\) is mean velocity, \(Fr\) is the Froude number, \(g\) is gravitational acceleration, and \(\lambda\) is the length of a bed form.

3. Results

3.1. Bed Elevation Long-Duration Behavior and Bed Morphology

The bed morphology revealed by a river cross-section (Figure 4) showed uneven elevations near the left bank of the river and a relatively flat bottom through the middle part up to the deepest section near the right bank, where the majority of the *P. pectinatus* specimens grew. The results from the
experiment conducted in the Jeziorka River showed fluctuations of the bed elevation of less than 0.5 cm in the vegetated area and up to 2 cm in the clear area (Figure 5).

The experiment conducted in the Świder River showed ripples 1.0 to 2.5 cm high in the vegetated area (Figure 6). In the clear region, no distinct ripples were observed. In the clear region, a sand dune 10 cm high was observed (Figure 6). The occurrence of smaller ripples was then observed on the upstream side of that sand form, which were clearly distinguishable as a second type of wave form [33,49]. The streamwise depth elevation profiles revealed easily distinguishable deep and shallow areas close to the bank (Figure 4, profile I in Figure 7). Moving away from the bank, profile II, which was at the same distance from the bank as the long-duration measurement point in the vegetated area, showed continuous changes in elevation approximately 3 cm high. Profile IV, which was at the same distance from bank as the long-duration measurement point in the clear area, showed a series of dunes 10 cm high. Both profiles II and IV were thus consistent with time-dependent changes in the vegetated and clear regions (Figure 6). The measured properties of the sand forms showed that height and length increased with increasing distance from the channel’s left bank, while dune celerity decreased (Table 2, Figure 7).

Table 2. Statistics of sand waves in the considered profiles in the Świder River. L denotes length of a dune; H denotes height of the dune; SD denotes standard deviation.

| Location | Mean L (cm) | Mean H (cm) | SD L (cm) | SD H (cm) | Mean Time τ (s) | Celerity (ms$^{-1} \times 10^3$) | Theoretical Celerity (ms$^{-1} \times 10^3$) |
|----------|-------------|-------------|-----------|-----------|-----------------|-------------------------------|------------------------------------------|
| Profile II | 127.1       | 3.2         | 36.0      | 2.2       | 1338            | 0.949                         | 0.112                                    |
| Profile III | 257.5       | 12.8        | 67.9      | 5.2       | –               | –                             | –                                        |
| Profile IV | 345.8       | 13.2        | 138.5     | 4.0       | 10.400          | 0.332                         | 0.041                                    |

Figure 4. Bed morphology of studied sites in the (A) Świder River and (B) Jeziorka River.
Figure 5. Smoothed longitudinal velocity and depth changes in the Jeziorka River at one point in clear, and one in vegetated region.

Figure 6. Smoothed longitudinal velocity and depth changes in the Świder River at one point in clear, and one in vegetated region. Parts of the velocity recording were removed due to bad signal characteristics.
3.2. Velocity Long-Duration Behavior

The long-duration velocity measurements in two distinct river regions were quasi-stationary, given various phases of sand waves moving beneath, and represent only quantitative differences between those regions and the rivers themselves (Tables 3 and 4). The measured mean longitudinal velocity was higher in the clear flow path in the Jeziorka and Świder rivers by 12% and 16%, respectively, with mean squared error two times higher in the latter (Table 3). The lateral and vertical velocity fluctuations were also higher in the clear area in the Świder River, while in the Jeziorka River there was no difference between regions (Table 3). The normalized turbulence intensities in the Świder showed increased mixing in longitudinal and vertical directions (higher in clear region than in vegetated region), in contrast to the Jeziorka River, where lateral mixing increased with no differences between regions (Table 3). The turbulent kinetic energy was the highest in the clear region in the Świder River, three times higher than in the other three regions. The normal stresses in the clear region of the Świder River were larger than those in the vegetated area (Table 4), particularly the shear stress (U'W'), which showed typical values as reported by Sukhodolov et al. [27] above the bed with a sandbar of similar height (albeit immobile). The ratio of normal transverse and vertical stresses was close to 1.65 for the Jeziorka River, which is consistent with the coefficient for open channel flows found by Nezu and Nakagawa [50]. The fluctuating longitudinal velocity power spectral density (PSD) plot (Figure 8) showed the occurrence of a $-5/3$ Kolmogorov scaling range at higher frequencies, and a $-1$ scaling range at lower frequencies, which are attributed to the inertial and production subranges of turbulence, respectively [51]. In the Jeziorka River, there were no apparent differences in the PSD between the clear and vegetated regions, in contrast to the Świder River, where the PSD was flatter for the region with plants.
3.3. Velocity Profiles in the Proximity of P. pectinatus

The mean longitudinal velocities upstream of *P. pectinatus* were distributed in the expected logarithmic profile, with a maximum value of 0.58 m·s⁻¹ (Figure 9). Two downstream profiles were similar in each velocity direction, but with increased deviations below \(Z/H = 0.3\), especially for the profile 0.25 m behind the plant. The profile upstream had increased mean lateral velocity and vertical velocity oscillating near 0 m·s⁻¹, while downstream of the plant the velocity tended to be directed toward the bottom. The turbulent kinetic energy was elevated at a height of \(Z/H = 0.1\) upstream of the plant, but downstream it was lower at that height (Figure 9).

![Figure 8. Power spectrum of longitudinal velocities calculated from the filtered signal of long-duration measurements.](image-url)
1.14 mm for hydrophytes from the Jeziorka and the Świder rivers, respectively (Table 5), whereas in the measurements the values ranged from 0.98 mm for species from the Jeziorka to 1.33 mm for individuals from the Świder (Table 6). The maximum force, maximum stress, and flexural strain were similar for both locations (Table 5). However, *P. pectinatus* from the Jeziorka River displayed a mean Young’s modulus of 104.50 MPa and 175.20 MPa, respectively (Table 5). The same species from the Świder had a stiffness that was approximately twice as high: The flexural rigidity was 16.09 N·mm² (Table 5). The flexural modulus was also much higher, reaching 280.39 MPa.

The tension test results show that the breaking force of *P. pectinatus* from the Jeziorka River was higher, 3.32 N compared to the value associated with the Świder River, 2.34 N (Table 6). The differences in strength were more prominent; the plant from the first experiment had a breaking stress of 4.56 MPa, whereas the value was lower for the second plant, at 1.91 MPa (Table 6). The strain of the studied species from the Jeziorka River was 12.33%, whereas for hydrophytes from the Świder River, the strain was 9.08%.

3.4. Biomechanical Measurements

*P. pectinatus* samples from the Jeziorka River were characterized by thinner stems than those collected from the Świder River. The mean diameters in the three-point bending tests were 0.99 and 1.14 mm for hydrophytes from the Jeziorka and the Świder rivers, respectively (Table 5), whereas in the tension tests, these values ranged from 0.98 mm for species from the Jeziorka to 1.33 mm for individuals from the Świder (Table 6). The maximum force, maximum stress, and flexural strain were similar for both locations (Table 5). However, *P. pectinatus* from the Jeziorka River was characterized by lower flexural rigidity and flexural modulus of 7.82 N·mm² and 159.80 MPa, respectively (Table 5). The same species from the Świder River had a stiffness that was approximately twice as high: The flexural rigidity was 16.09 N·mm² (Table 5). The flexural modulus was also much higher, reaching 280.39 MPa.

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### Table 5. Diameter and three-point bending test results for *P. pectinatus*.

| Parameter                  | Jeziorka River | Świder River |
|----------------------------|----------------|--------------|
| Diameter (mm)              | 0.99 ± 0.17    | 0.99         |
| Maximum force (N)          | 0.022 ± 0.007  | 0.020        |
| Maximum stress (MPa)       | 0.030 ± 0.013  | 0.025        |
| Flexural strain (%)        | 4.38 ± 1.02    | 4.27         |
| Max. deflection (mm)       | 16.14 ± 1.55   | 16.13        |
| Sec. m. of area (mm⁴)      | 0.057 ± 0.049  | 0.046        |
| Flexural rigidity (N·mm²)  | 7.82 ± 4.22    | 6.22         |
| Flexural modulus (MPa)     | 175.20 ± 91.49 | 159.80       |

Figure 9. Profiles from the Jeziorka River upstream and downstream of *P. pectinatus*.
Table 6. Diameter and tension test results for *P. pectinatus*.

| Parameter       | Jeziorka River | Świder River |
|-----------------|----------------|--------------|
|                 | 5 July         | 1 October    |
| Diameter (mm)   | 0.98 ± 0.12    | 1.00         |
| Breaking force (N) | 3.32 ± 0.63    | 3.45         |
| Breaking stress (MPa) | 4.56 ± 1.30    | 4.63         |
| Breaking strain (%) | 12.33 ± 3.42   | 13.18        |
| Young’s modulus (MPa) | 104.50 ± 35.19 | 104.00       |

4. Discussion

The typical configuration of aquatic plants found in both studied rivers consisted of a strip of plants growing along the river bank, similar to the configuration observed in the study by Naden et al. [24]. We expect aquatic macrophytes to grow closer to river banks due to the favorable conditions for seeds to be trapped [52]. In the case of *P. pectinatus*, Chambers et al. [53] showed that it thrives in rivers with slower currents; therefore, in the vegetated region, the water current was expected to be lower than in the main stream due to the plant’s growth requirements. This theory was borne out by the obtained results, with measured velocity being 12% and 16% lower than that in the clear regions in the Świder and Jeziorka rivers, respectively (Table 3). However, the study of Kemp et al. [54] showed that submerged macrophytes with fine leaves thrive in habitats with a wide range of Froude numbers.

The investigated rivers exhibited similar granulometry (Figure 1), but the Świder River had a Froude number two times greater (Table 1), and it produced different sand wave patterns. In the Jeziorka River, approximately 20% of the width of the studied cross-section was considered the vegetated region, while in the Świder River, the living plants spread to no more than 10% of the channel width. Note that the experiments were conducted in meandering rivers, which is reflected in the cross-sectional structure (Figure 4), and which in turn had a major influence on bed scour and deposition [55,56]. Three regions were distinguished in both rivers: The relatively flat, deepest region with vegetation, a middle region without plants but with sand forms, and a shallow region opposite the deep one (Figure 4), which in the Jeziorka River was also vegetated. This pattern of deep and shallow fragments of a channel, a sequence of pools and riffles, is typical for natural meandering rivers [57–59]. However, the pattern of plant patches in a channel varies not only yearly but also seasonally, changing the bed shear stress over time [18,24]. Furthermore, the occurrence of plants and flattening of sand waves may be connected to the appearance of dead wood debris, which was found in the vegetated regions in both rivers, as this type of debris can trap seeds and organic matter [60]. Submerged vegetation in this case performs a similar role as emergent or riparian vegetation: Trapping fine sediment, building landforms, and ultimately stabilizing the river banks [19,61,62]. This hypothesis is consistent with the occurrence of sandbars, visible in vegetated regions near banks in both rivers (Figure 4). These bars could serve as plant-rooting locations, where vegetation decreases the flow velocity within the patch and increases it around the main body of the plant [63]. The present case study revealed that sparsely rooted plant specimens did not create blockage factor big enough to deflect flow and scouring effect was not observed, which otherwise would protect a plant from being buried by passing sand forms.

The collected records of bed elevation changes show differences between various regions in river cross-section morphology in terms of the discrete characterization method. The Jeziorka River featured completely disturbed sand waves in the vegetated region (Figure 5), while in the Świder River, these waves occurred in the vicinity of *P. pectinatus*; however, these dunes were three times lower than the fully developed wave forms in the clear region (Figure 6). The bed form celerity shows that larger dunes were migrating at slower rate than ripples close to the river bank (Table 2). Large bed forms contain more bed load, therefore a significant part of the sediment must be transported in the clear
region of the river through dune movements [64]. The flow in that region was characterized by higher Reynolds stresses due to higher shear caused by larger dunes and higher mean longitudinal velocity, however reduced in the dune trough (Figure 6), which is in line with those of, e.g., Venditti and Bauer [65] and Kostaschuk [66]. In comparison to the study of Sukhodolov et al. [27], mean Reynolds shear stress was 10 times higher in spite of the same mean velocity, however it can be explained by the fact that measured point in the present case study in the Świder River was relatively close to the bottom and a lot of negative velocities were recorded, which was not the case in study of Sukhodolov [27].

A spike visible in the power spectral density plot in the clear region of the Świder River at 0.4–0.5 Hz (Figure 8) can be related to eddy-like flow structures that originated from the flow separation zone in the lee side of the dune [50]. However, Venditti and Bauer [65] indicated that eddies could not be attributed to a certain frequency, and what is more, analysis of a long-duration velocity power spectrum can be misleading due to the occurrence of local nonstationarity in the time series (particularly prominent in Figure 6). Furthermore, Singh et al. [67] reported that eddies with a vertical size smaller than the distance from the bottom to the measurement point could not be found in the spectra at all. Nevertheless, the spectrum of the velocity in the vegetated region in the Świder River was flatter in the inertial subrange than in the clear region. Consequently, more energy was conducted into higher frequencies in the clear region, as would be expected in the dissipation wake of a bed form [65]. The elevated Reynolds stress in that clear region (Table 4) was also in agreement with the observation of energy distribution. Obtained results indicate that there was a connection between the height of sand dunes and Reynolds shear stress in the water layer above. However, in the pool region, Reynolds stress distribution might be not correlated with sediment size or flow characteristics [59].

Plants, even of the same species, that grow in similar rivers can be characterized by changes in biomechanical properties that depend on the habitat conditions and the phase of plant growth [68]. *P. pectinatus* from two investigated rivers differed in terms of length, thickness, and biomechanical traits. Hydrophytes from the Świder River were shorter, thicker, and stiffer (Tables 5 and 6), and were more resistant to bending forces than specimens from the Jeziorka River. Nevertheless, they were more durable in the tension tests (Table 6). In other words, the specimens growing in the region where the bed was flat grew longer and had altered characteristics such that they were more prone to bend to withstand higher forces resulting from their length. This behavior may be an example of an avoidance strategy, as described by Puijalon et al. [69]. The occurrence of drag on a single specimen was visible in the measured velocity profiles (Figure 9), in which a region of disturbed velocities in the plant wake was noted. In the Świder River, the increased resistance to bending by specimens growing out of the riverbed, where the ripples were as high as the plants themselves, may be a natural strategy to withstand unfavorable conditions. Then, the most likely reason for the above-mentioned differences in biomechanical properties of hydrophytes may be connected to their diversity of growth phases [68].

The question arises as to whether the studied plants affect the flow and sediment movements or vice versa. The measurements in the Jeziorka River revealed lower mean longitudinal and lateral velocities in the wake of the single plant, but most outstanding was the lower TKE relative to that of the point upstream (Figure 9). The TKE results were not consistent with those obtained by Biggs et al. [70] with another hydrophyte of similar morphology, *Ranunculus penicillatus* (Dumort.) Bab., for which TKE was elevated. However, that study focused on a dense patch and not on a thin individual plant. In the Jeziorka River, the studied macrophyte was streamlining, which is a natural behavior for flexible plants in high-flow conditions [71,72]. Moreover, Siniscalchi and Nikora [72] showed that single hydrophytes exhibited flapping-like motions to minimize drag force, indicating their susceptibility to turbulent eddies [63]. With the frontal area of *P. pectinatus* being approximately four times smaller than that of the patch in Biggs’s experiment, the effect on flow in the plant’s wake was barely discernible (Figure 9). The studied individuals were too small to show any effect on the bottom morphology; however, there were denser patches downstream (Figure 2), which were followed by a shallower area (Figure 4). In the Świder River, plants grew too small and sparsely to have a significant impact on the velocity profile or bed shear stress (canopies of a/h << 0.1, where a is diameter and h is water height; see [71])
or even on the sand forms; therefore, plants were periodically covered by migrating dunes, as we observed in our measurements. This result is consistent with a theory by Green [63], who suggested that certain hydrophytes in regions close to the bank and the riverbed do not contribute significantly to resistance due to drag of the boundary layer itself. The present case study investigated such species with even lower shoot density and more sharp-ended stems than *R. penicillatus*, which have practically no blockage effect. This can be a hint, which species are likely to positively contribute to biodiversity of river but not significantly affect flow resistance. Furthermore, Chen and Kao [73] showed that flow through sparse vegetation was similar to flow that occurs in a solid-boundary channel.

Studies performed in situ by Naden et al. [24] and Sukhodolova and Sukhodolov [74] represent a goal for how studies of flow–biota–sediment interactions should be conducted. The present investigation was smaller in scale and details, as the requirement of maintaining stationarity of the measurement and revision of local conditions forced us to skip some of the observations. In spite of this, the findings of the presented study show agreement with general theories on this subject and add scarcely available data to it. Therefore, the next step would be to perform subsequent measurements in the next summer season in the same spots to observe changes in bed morphology and vegetation cover. Additionally, the creation of a denser grid of velocity profiles around *P. pectinatus* patches, including big canopies present in the Jeziorka River, would be preferable in order to find how these plants’ assembles affect flow field.

5. Conclusions

We present the following observations with regard to the plant occurrence, sediment movements, plant traits, and effects of hydrophytes on flow and sand forms:

- The aquatic plant *P. pectinatus* grew in a larger quantity in the river with the lower Reynolds number, tending to be located in a pool section of a channel, where smaller ripples occurred. This finding suggests an emerging pattern in which flow conditions and bed morphology are connected with the occurrence of the studied hydrophyte.

- The bed forms measured in the Świder River tended to exhibit increased height and length with increased distance from the vegetated region close to the bank. These forms were also more developed than those in the Jeziorka River, which is characterized by finer sediment and faster but steadier flow, with lower turbulent kinetic energy and a lower Froude number.

- Major differences were identified in the plants’ morphology and biomechanics: Older individuals from the faster-flowing Świder River were thicker, shorter, and stiffer than their younger counterparts from the Jeziorka River, which were more prone to bending. These traits may have resulted from the aquatic macrophytes’ adjustment to the habitat conditions.

- Strips of short individual *P. pectinatus* plants within a mobile bed in the Świder River did not seem to detectably affect passing sand ripples. Turbulence statistics suggest a much steadier flow than that in the clear area.

- The single hydrophyte from the Jeziorka River did not affect the velocities measured downstream; only elevated TKE was visible in the wake of the plant.

- The mean velocity in front of the plants was approximately 12% and 16% lower than in the clear regions in the Świder River and Jeziorka River, respectively. Therefore, bed conditions such as small sandbars and wood debris trapping organic material could have made major contributions to the creation of a habitat suitable for plants to grow.
Author Contributions: Ł.P., A.M.L., and R.J.B. designed the study and carried out the field experiments. Ł.P. conducted the majority of the data analysis and wrote the majority of the paper. A.M.L. and Ł.P. carried out the biomechanical tests and granulometry. A.M.L. described and analyzed the biomechanical test results and prepared bed morphology maps. R.J.B. prepared and depicted bed elevation data. All authors contributed to the final version of the manuscript.

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