Stratigraphic ranges of selected miospores in coal seams of Upper Coal Measures age in Oxfordshire and S.E. Warwickshire

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ABSTRACT — The extensive programme of boring recently undertaken by the National Coal Board to establish the limits of the concealed coalfield in Oxfordshire and southeast Warwickshire, has provided a unique sequence of high volatile bituminous coals from the upper part of the Upper Coal Measures suitable for spore analysis. The coals are mainly bright and the possible implications for the spore flora are briefly considered. The spores from the coals are characteristic of Westphalian D assemblages but the problem of recognising the base of the Westphalian D is discussed. It is concluded that the seams under investigation are all of Westphalian D age. Two new species, *Vestispora burfordiensis* and *V. witneyensis* are described. The occurrence of these species as well as *Schopfites dimorphus* and *Punctatosporites granifer* in Oxfordshire and in other British coalfields is reviewed. The ranges of these species provide the basis for recognising three new miospore assemblage zones to replace the former *Thymospora obscura* Assemblage Zone XI of Smith & Butterworth (1967). The base of the epible of *Thymospora pseudothiessenii* is also well defined and is a useful biostatigraphic marker. Comparison is made with the zonal sub-division of the Desmoinesian series of the Illinois basin by Peppers (1985).

INTRODUCTION

The presence, near Burford in Oxfordshire, of supposedly barren Upper Coal Measures under Mesozoic rocks has been known for just over 100 years but seams of workable thickness were only discovered when the British Geological Survey borehole at Apley Barn was sunk in 1960-61 near Witney (Poole, 1969). Subsequently, two further B.G.S. boreholes at Steeple Aston (Poole, 1977) and Withycombe Farm near Banbury (Poole, 1978) also proved workable seams. The possible limits of the coalfield were suggested by Dunham & Poole (1974).

The miospores from seams in these boreholes are characteristic of seams of Westphalian D age and tentative correlations between these boreholes, based on the stratigraphic ranges of a limited number of miospores, have been proposed (Smith in Poole, 1977, 1978).

In December 1975, the National Coal Board began a programme of exploratory boring in the area between Oxford and Coventry. The resulting information from a large number of relatively closely spaced boreholes has made possible the detailed correlation of the main coal bearing horizons across the area. There is thus a sound basis from which to examine the stratigraphic ranges of selected miospores in the Oxfordshire Coalfield. The coals are of relatively low rank and therefore represent the fullest succession of seams of Westphalian D age in Britain from which spores can be readily isolated.

As a result of recent work, the species and the ranges proposed earlier have been emended and a scheme for zoning the Upper Coal Measures of Oxfordshire is proposed. It is only concerned with species of proven value for correlation purposes and not with the total assemblage.

GEOLOGY

In Oxfordshire and southeast Warwickshire, up to 1000m of Upper Coal Measures rocks occur, comprising mudstones, sandstones and coals. A variable amount of reddening is developed in the upper portions. No marine horizons are represented in these rocks, which lie unconformably on strata of Cambrian to Devonian age. Strata, in some cases with coal seams, of Middle and possibly Lower Coal Measures age are frequently present at the base of the succession. They occur over a wide area and in many boreholes are separated from the seams of Upper Coal Measures age by igneous rocks varying in thickness from a few centimetres to 165m in the Steeple Aston Borehole. Nowhere in the coalfield is there a conformable development of seams from the Middle to the Upper Coal Measures. The Coal Measures rocks are, in turn, overlain unconformably by Mesozoic rocks.

In the Upper Coal Measures, the thickness of coal bearing strata is greater in Oxfordshire than in southeast Warwickshire due to the influence of the Midland
Barrier on the process of sedimentation and erosion in the two areas. At Apley Barn, the southernmost of the boreholes examined, some 40 coal seams, the majority thin, are developed in the Upper Coal Measures. The thickness of coal bearing strata is 804m. Moving in a northerly direction, both the number of seams and the thickness of coal bearing strata decreases until at Southam the number of seams is reduced to approximately 10 in some 215m of strata. Over most of southeast Warwickshire there are only three seams in 100m of strata and the lower seams in the Upper Coal Measures are missing from the succession. However, the total thickness of Upper Coal Measures (Fig. 2) in the boreholes in the Coventry area, some 15km north of Warwick, exceeds that at Withycombe Farm, near Banbury, due to the accumulation and subsequent preservation of a considerable thickness of barren Upper Coal Measures strata above the coal bearing strata.

The rates of subsidence presumably varied in different parts of the coalfield. Thus at Apley Barn the main coal horizons are represented by many seams and the rate of subsidence must have been rapid compared with the more northerly areas with fewer seams of equivalent age. However, there is no increase in the thickness of individual seams north of Banbury associated with reduced subsidence rates, at least in the boreholes examined. At Apley Barn the total thickness of coal exceeds that of the equivalent horizons farther north. At Southam, for example, many of the seams are thin and dirty.

These facts have a bearing on the development of peat facies (coal lithotypes) and species diversity. Both tend to optimum development in thick coals.

STRATIGRAPHY

At Apley Barn, the B.G.S. has divided the Upper Coal Measures rocks into five lithological formations, which have been named, in descending order:-

- Windrush Formation
- Burford Coal Formation
- Crawley Formation
- Witney Coal Formation
- Arenaceous Coal Formation

These formations were subsequently recognised in the Steeple Aston borehole but, at Withycombe Farm, the strata have been correlated with those parts of the Keele and Halesowen Formations of the Southern English Midlands Coalfields lying to the north and northwest (Poole, 1978, p. 16). The base of the Windrush Formation is roughly equivalent to the base of the Keele Formation in the boreholes of southeast Warwickshire.

With the additional geophysical and chemical information from the relatively closely spaced N.C.B. boreholes, it has been possible not only to recognise and correlate the broad lithological horizons but to demonstrate the lateral persistence of the main seams in the principal coal bearing horizons. Thus, the horizon that corresponds to the base of the Keele and Windrush Formations is better defined by the coal seam which occurs at this level, than by the lithological features, such as reddening of the strata which may not be a contemporary event and is an unreliable basis for chronostratigraphic correlation. The correlation of individual beds within formations, however, is sometimes difficult, particularly where, at a given horizon, there is a discrepancy between the number of leaves in adjacent boreholes. Unfortunately, palynological methods are rarely able to assist correlation in these instances.

The broad correlations which have been established by non-palynological means have been used as a basis for comparing the stratigraphic ranges of selected spores in the different boreholes.

LOCATION OF BOREHOLES AND SEAMS EXAMINED

Fig. 1 shows the location of the nine boreholes which were the sources of the coals used in this investigation. The boreholes were selected to provide a roughly N-S section through the coalfield from Apley Barn in the north to Southam in the north. The Kineton and Southam boreholes are located in the South Warwickshire Prospect. The distance between the extreme boreholes is approximately 50km.

The main seams have been named by R. F. Goossens to assist in correlation and their usage is well established within the N.C.B. The representation of these seams in the boreholes examined is shown in Fig. 2. Named seams may comprise several beds of coal separated by dirts. Unless otherwise indicated, at least one member of each of the important groups of seams in each borehole has been examined for spores. This diagrammatic presentation of the succession conceals the northward thinning of coal bearing strata and the considerable reduction in the number of seams referred to above.

PETROGRAPHY OF THE COALS

Few palynologists consider the evidence which a coal provides on the environment in which it was formed and the possible effect of that environment on the plants growing on the deposit and on the derived miospore assemblages. The petrographic differences between seams may be due to one or more of the following influences: climate, the depositional environment and the seral changes in the vegetation taking place during the growth of the deposit. They may also reflect differences in the diagenetic and catagenetic stages of coalification resulting from geological processes. Petrographic differences attributable to the latter
cause will not be reflected in the floral record. To assist in the interpretation of the petrographic evidence Styan & Bustin (1983 a & b) have investigated the microstructure of a modern delta peat deposit in an attempt to trace the origin of the petrographic components of Hard coals.

Harvey & Dillon (1985) were able to demonstrate both long term changes of maceral composition which they related to climatic conditions, and short term variations in maceral composition within individual seams which they attributed to localised palaeoenvironmental influences. They record an enrichment of inertinite in coals belonging to the Missourian Series in Illinois when compared with the older Desmoinesian coals. The latter are believed to be equivalent to European coal of Westphalian D age. This change was considered to mark the onset of a drier climatic regime, which according to Phillips et al. (1974) and Phillips & Peppers (1984) was responsible for the change of the dominant peat forming vegetation from arborescent lycopods to tree ferns and no doubt accounts for the increasing proportion of monolet spores in the younger seams of the Upper Coal Measures in Britain.

The broad relationships between petrography and spore assemblages have been established for British coals of Middle and Lower Coal Measures age (Smith, 1962, 1964). No study has been made of coals of Upper Coal Measures age in Britain although coals of this age from the Saar-Lorraine basin have been studied by Navale (1964) and Alpern et al. (1964). Such studies depend on a fine subdivision of the seam profile for petrographic and spore analyses. Such investigations have not been made on the seams in the Oxfordshire coalfield but, nevertheless, some observations can be made by comparing the petrography of these seams with those from older horizons on a whole seam basis.

The most meaningful technique for characterising the composition of the coal-peat is to determine its microlithotype content (I.C.C.P., International Handbook of Coal Petrography, 1963, 1971, 1976). Microlithotypes are the natural associations of macerals (the basic organic components of coal) broadly corresponding to the different coal lithotypes which are recognisable by eye without the aid of a microscope.

Table 1 compares the average microlithotype composition of the seams in a southeast Warwickshire borehole and two Oxfordshire boreholes (located at the extremes and in the middle of the N/S transect across the coalfield) with the composition of seams from the Upper, Middle and Lower Coal Measures of Yorkshire. It can be seen that the Oxfordshire seams are enriched in the bright coal microlithotypes (vitrite and clarite) and deficient in durite (the main component of dull coal) compared with the older seams in Yorkshire, but do not differ significantly in composition from the seam in the Upper Coal Measures of Yorkshire. The mean of the combined totals of vitrite and clarite in the individual seams from the three boreholes is 74% ranging from 42-96%. However, four seams from Withycombe Farm have combined totals of these microlithotypes of less than 60%. These seams all contained between 35-43% of trimacerite which is significantly above the average value of seams in the Oxfordshire coalfield. These seams tend to be among the thinner seams, which is unexpected. In general, thicker seams tend to have a greater diversity of microlithotypes. The seams at Southam, where the succession is condensed, are also thin, and several contain significant amounts of pyrite, carbonaceous shale and shale.

The scarcity of Densospores and the absence of crassidurite from the seams in the Oxfordshire Basin is unlikely to be due to climatic influence since the genus occurs in contemporary strata in the Bristol and Somerset Coalfields. It suggests that the hydrological conditions associated with the occurrence of Densospores were lacking in the Oxfordshire Basin in Upper...
Coal Measures times. In the contemporary Bristol and Somerset strata and in the Lower and Middle Coal Measures in other paralic basins in Britain, the environmental conditions associated with the formation of many of the thicker seams intermittently favoured the growth of plants producing the larger forms of *Densosporites*, apparently to the exclusion of most other plants. At these Densospore rich horizons the processes of decay and preservation of organic matter were also different, in so far as it is preserved as macrinite rather than vitrinite.

Tenuidurite is also scarce in the Oxfordshire seams, the greatest amount being 7% in the upper bed of the Apley Seam at Apley Barn. This seam also contained 14% of cannel coal. The amounts of fusite are not significantly different from the amounts in the seams of Middle Coal Measures age.

It is reasonable to assume that the mainly bright coals of the Oxfordshire Coalfield represent the formation of autochthonous or subautochthonous peat in a more or less rapidly subsiding basin where the organic matter was relatively well preserved under waterlogged anoxic conditions. Peat formed under these conditions might be expected to support relatively few types of vegetation and overall to have a relatively poor diversity of miospores. In contrast are the petrographically variable seams of the Lower and Middle Coal Measures. In the case of these seams, subsidence was probably interrupted or peat growth may have exceeded the amount of subsidence. Ground water levels probably showed greater fluctuations with the result that part of the peat was periodically exposed to relatively aerobic conditions. Peat formed under these varying conditions would provide a greater diversity of habitat type and vegetation.

An example from Oxfordshire of a petrographic difference between seams unrelated to their miospore assemblages is the occurrence of significant amounts of structureless vitrinite including pseudovitrinite in seams from the Apley Barn Borehole and the occurrence of vitrinite with a patchy texture in the coals from the other two boreholes. Jones *et al.* (1984) have suggested that such differences may be the result of different coalification gradients.
Miospores of Upper Coal Measures age

| Source                        | Number of Seams | Samples | Vitrite | Clarite | Microlithotypes (volume percent) | Carbominerite |
|-------------------------------|-----------------|---------|---------|---------|---------------------------------|---------------|
|                               |                 |         |         |         | Trimacerite | Durite | Fusite | Vitrinertite | Carbopyrite |
| Oxfordshire/ Warwickshire     |                 |         |         |         |                   |               |        |              |            |
| Upper CM                      |                 |         |         |         |                   |               |        |              |            |
| Apley Barn BH                 | 15              | 47      | 25      | 19      | 1                  | 4          | 1      | 3            |            |
|                               |                 | (28-63) | (13-43) | (7-34)  | (0-7)              | (1-9)      |        |              |            |
| Withycombe Fm BH              | 10              | 43      | 27      | 23      | 1                  | 4          | 1      | 1            |            |
|                               |                 | (32-52) | (10-48) | (3-43)  | (0-5)              | (1-10)     |        |              |            |
| Southam BH                    | 6               | 54      | 31      | 8       | 1                  | 3          | <0.5   | 3            |            |
|                               |                 | (48-58) | (23-35) | (5-13)  | (0-2)              | (2-6)      |        |              |            |
| Yorkshire (Section worked excl. gross dirt) |     |         |         |         |                   |               |        |              |            |
| Upper CM                      | 1               | 7       | 48      | 28      | 17                 | 4          | n.d.   | 2            |            |
|                               |                 | (43-59) | (20-36) | (6-25)  | (0-2)              | (2-6)      |        |              |            |
| Middle CM                     | 5               | 45      | 42      | 20      | 21                 | 10         | 6      | n.d.         | 1           |
|                               |                 | (22-58) | (8-35)  | (8-34)  | (2-27)             | (2-14)     |        |              |             |
| Lower CM                      | 6               | 51      | 42      | 25      | 20                 | 6          | 6      | n.d.         | 1           |
|                               |                 | (24-57) | (16-39) | (6-36)  | (1-22)             | (2-12)     |        |              |             |
| n.d. = not determined         |                 |         |         |         |                   |               |        |              |            |

Table 1. Microlithotype composition of seams from the Oxfordshire Coalfield (Upper Coal Measures) and Yorkshire Coalfield (Upper, Middle and Lower Coal Measures).

**SYSTEMATIC DESCRIPTIONS**

*Turma* Triletes (Reinsch) Dettmann, 1963

*Suprasubturma* Acavatitriletes Dettmann, 1963

*Subturma* Azonotritiletes (Luber) Dettmann, 1963

*Infraturma* Apiculati (Bennie and Kidston) Potonie, 1956

*Subinfraturma* Verrucati, Dybova & Jachowicz, 1957

**Genus Schopfites** Kosanke, 1950

Type species *Schopfites dimorphus* Kosanke, 1950 (Pl. 2, figs. 22-24)

**Remarks.** Peppers (1985), in his comparative studies of the ranges of miospores in Illinois and in Western Europe, recognises two species, *S. dimorphus* and *S. colchesterensis* Kosanke, 1950 and used the name of the latter species to designate an assemblage zone. In the present work all forms of *Schopfites* have been assigned to *S. dimorphus* for the reasons given in Smith & Butterworth (1967). A smaller, less coarsely ornamented form referred to *Schopfites* sp. by Peppers (1985) and to *S. colchesterensis* by Loboziak (1971 pl. 4, fig. 25) was not seen in the present work.

**Occurrence.** Smith & Butterworth (1967) assumed the species to occur throughout the Westphalian D of Great Britain. However in Oxfordshire it is now known to have a more restricted stratigraphic range occurring in seams between the base of the Witney Coal Formation and the base of the Windrush Coal Formation. Re-examination of its occurrence elsewhere suggests that its range may be similar to that in Oxfordshire. It may be found in most seams as a scarce component of the miospore flora, seldom exceeding 0.4% of the total seam assemblage.

*Turma* Monoletes Ibrahim, 1933

*Suprasubturma* Acavatomonoletes Dettmann, 1963

*Subturma* Azonomonoletes (Luber) Dettmann, 1963

*Infraturma* Sculptatomonoletes Dybova & Jachowicz, 1957

**Genus Punctutosporites** (Ibrahim 1933. Krutzsch 1959) Alpern & Doubinger, 1973

Type species *Punctutosporites grunifer* (Potonie & Kremp 1956) Alpern & Doubinger, 1973 (Pl. 2, figs. 1-8)

**Remarks.** The figures show the range of size, shape and ornament of forms assigned to *P. grunifer* by the author. Some of these forms resemble specimens assigned to *P. minutus* and *P. rotundus* Bharadwaj,
1957 by Alpern & Doubinger (1973) in their revision of Palaeozoic monolete genera. However, the author has assigned to *P. granifer* all monolete spores without a pseudocingulum, within the emended size range 14-42 μm (Alpern & Doubinger, 1973) and with a distinctly granulate ornament clearly visible at low powers of magnification (×10 objective). The grana clearly project at the amb. Monolete spores with scabrate ornament and with a smooth amb. are assigned to *P. punctatus* (Kosanke) Alpern and Doubinger or *P. minutus*. It is significant that Van Wijhe & Bless, 1974 refer to the distinctive appearance of forms of *C. granifer* occurring in the upper Westphalian C and Westphalian D of the Netherlands. These forms are relatively large, frequently less regularly oval and possess a thicker exine than specimens from the lower part of the Westphalian C.

**Occurrence.** Smith & Butterworth (1967) gave the stratigraphic range of the species as the Westphalian C and D of Great Britain without the knowledge of the precise age of the coals examined. In Oxfordshire the species occurs throughout the Arenaceous and Witney Coal Formations. It occurs in most seams, with an abundance generally between 1% and 3% of the seam assemblage; maximum recorded abundance 8%.

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**Explanation of Plate 1**

All figures are ×500

Figs. 1-7. *Vestispora burfordiensis* sp. nov. (figs. 1-6, from type locality).

Fig. 1. Holotype, oblique polar compression; preparation T/129/1.

Fig. 2. Paratype, equatorial compression; operculum partly detached, several relatively narrow encircling costae running parallel to the equator; preparation T/129/2.

Fig. 3. Oblique polar compression. Note smooth amb and thickness of exoexine as seen at margin; laesura and operculum faintly visible, poor costal development.

Fig. 4. Partial spore with distinct foveo-retticate ornament and unusually well developed costae projecting from amb.

Fig. 5. Equatorial compression; split between operculum and remainder of exoexine, costae indistinct.

Fig. 6. Torn exine, one narrow encircling costa and several incompletely preserved developed costae.

Fig. 7. Pseudo reticulate pattern formed by crossing costae on opposing surfaces of exoexine, seam at 676.68m Apley Barn Borehole, Crawley Formation.

Figs. 8-13. *Vestispora witneyensis* sp. nov.

Fig. 8. Holotype, oblique polar compression, aperture on proximal face left by partially detached operculum, single well-defined encircling costa; preparation T/130/1.

Fig. 9. Paratype, oblique polar compression, operculum in situ but incipient suture clearly visible, costae thin and indistinct; preparation T/130/2.

Fig. 10. Laesura showing through aperture left by detached operculum, seam at 676.68m Apley Barn Borehole, Crawley Formation.

Fig. 11. Equatorial compression with operculum partly detached but still in contact with part of exoexine; compare ornamentation of operculum with that in fig. 2; specimen from type locality.

Fig. 12. A large specimen with well-developed evenly spaced costae to one side, passing over the amb but hardly projecting, seam at 676.68m Apley Barn Borehole, Crawley Formation.

Fig. 13. Specimen with detached operculum and poorly developed costae; note the crenulate appearance of the amb due to the foveolate ornament, specimen from type locality.

Figs. 14-17. *Vestispora magna* (Butterworth & Williams) Wilson & Venkatachala, 1963, from type locality.

Fig. 14. Equatorial compression, well-defined broad costae encircling spore parallel to the quator and projecting at amb.

Fig. 15. Small specimen showing fine reticulate ornament but lacking distinct costae.

Figs. 16, 17. Two specimens showing fine and coarse reticulation, the latter produced by branching costae.
Miospores of Upper Coal Measures age
Genus *Thymospora* (Wilson & Venkatachala, 1963)
Alpern & Doubinger, 1973
Type species *V. thiessenii* (Kosanke) Wilson & Venkatachala, 1963

*Thymospora pseudothiessenii* (Kosanke, 1950) Alpern & Doubinger, 1973

(Pl. 2, figs. 9-21)

**Remarks.** The figures show the range of size and considerable variation of ornament of spores assigned to *T. pseudothiessenii*. In earlier work on the Oxfordshire Coalfield, Smith in Poole (1977, 1978) assigned the forms with extremely coarse ornament to *T. pervalvosa* (Alpern, 1959) Wilson & Venkatachala, 1963 and considered the species to have stratigraphic significance. The subsequent detailed study of all forms of *Thymospora* shows that spores with ornament intermediate between *T. pervalvosa* and *T. pseudothiessenii* occur, thus making the recognition of these species a subjective matter. This accords with the emended diagnosis of *T. pseudothiessenii* by Alpern & Doubinger (1973). Thus, whilst the coarse forms are more likely to be found in seams above the base of the epibole of the genus, the species can no longer be regarded as having stratigraphic significance within the Westphalian D.

**Occurrence.** Oxfordshire Coalfield, ranges throughout the Westphalian D. It is uncommon in the seams of the Arenaceous and Witney Coal Formations and in the Crawley Formation, seldom exceeding 0.4% of the total assemblage. In the Burford Coal and Windrush Coal Formations it generally exceeds 1% and may exceed 10% of the assemblage. The base of its epibole occurs at the horizon of the Low Broughton Seam. The species may show considerable variation in its abundance within the same seam.

*Turma Hilates* Dettmann, 1963
*Suprasubturma Cavatihilates* Smith & Butterworth, 1967
*Subturma Azonocavatihilates* Smith & Butterworth, 1967
*Infraturma Epitygmati* (Spode) Smith & Butterworth, 1967

**Genus Vestispora** (Wilson & Hoffmeister) Wilson & Venkatachala, 1963
Type species *V. profunda* Wilson & Hoffmeister, 1956

*Vestispora burfordiensis* sp. nov.

(Pl. 1, figs. 1-7)

**Derivation of name.** The specific name refers to the type stratum.

**Diagnosis.** Miospores, radial, trilete, amb circular to subcircular. Relatively thick operculate exoexine with fine foveo – reticulate ornament overlain by open system of relatively coarse unbranched encircling costae. Operculum more coarsely ornamented. Costae may project slightly beyond amb. Inner body and laesura seldom seen.

**Holotype.** Pl. 1, fig. 1. Preparation T/129/1 in collection of N.C.B. Yorkshire Regional Laboratory.

**Paratype.** Pl. 1, fig. 2. Preparation T/129/2.

**Locality and horizon.** Apley Barn Borehole, Oxfordshire, England. Top Broughton Seam at 492.46m. Burford Coal Formation, Upper Coal Measures.

**Description.** In lateral compression the proximal surface bearing the operculum often appears flattened. Eoexine thickness 2.5-3.7μm, generally recognisable at margin of spore as clear rim. Foveo-reticulate ornament weakly to strongly developed. Muri about 1.5-2.5μm in width, lumina 1.5-3.5μm, shape variable but generally polygonal. Costae usually prominent and unbranched 3μm in width. May project at amb up to 5μm but often hardly recognisable. Generally two to four costae encircle the spore. Mean size of the operculum 33μm measurable in about half the specimens. Muri on operculum tend to form a concentric pattern with lumina somewhat larger and more irregular in shape than on the remainder of exine.

**Dimensions.** 55 (70) 84μm (23 specimens macerated in fuming nitric acid). Holotype 78μm.

**Remarks.** The lumina are similar to those of *V. luminata* Ravn in size and shape but the latter species lacks costae.

**Comparison.** Two other species *V. pseudoreticulata* (Spode) in Smith & Butterworth, 1967 and *V. magna* (Butterworth & Williams) Wilson & Venkatachala, 1963 have a dual component exoexine ornament. *V. pseudoreticulata* may be distinguished by its less prominent, narrower carinate costae and *V. magna* by the reticulate pattern of its fine and coarse ornamentation. The muri comprising the latter clearly project at the amb.

**Occurrence.** Oxfordshire Coalfield, range in the coals extends from the upper part of the Witney Coal Formation (Upper Milton Seam) to the base of the Windrush Coal Formation (Newington Seam). Except for the Upper Milton Seam in which it is sporadic, it is generally present but with an abundance of <0.1% of the seam assemblage.

*Vestispora witneyensis* sp. nov.

(Pl. 1, figs. 8-13)

**Derivation of name.** The specific name refers to the type stratum.

**Diagnosis.** Miospores, radial, trilete, amb circular. Relatively thick operculate exoexine with fine foveolate ornament overlain by a poorly developed open system of relatively narrow unbranched costae.
Miospores of Upper Coal Measures age

Explanation of Plate 2

All figures are ×500

Figs. 1-8. *Punctatosporites granifer* (Potonić & Kremp) Alpern & Doubinger, 1973 (figs. 1, 2, 4, 8, seam at 467.40m (Lower Milton), Steeple Aston Borehole, Witney Formation; figs. 3, 5, 6, 7, seam at 836.44m (?) Aston), Withycombe Farm Borehole, Arenaceous Coal Formation).

Figs. 1, 2. Small and large specimens with fine grana.

Figs. 3, 4, 5. Oval and round specimens with medium sized grana.

Figs. 6, 7, 8. Specimens with coarse grana; fig. 8, specimen with thickened exine giving appearance of pseudo-cingulum.

Figs. 9-21. *Thymospora pseudothiessenii* (Kosanke) Alpern & Doubinger, 1973 (figs. 9-13, 15-19, seam at 586.93m (Overthorpe) Withycombe Farm Borehole, Windrush Formation; fig. 14, seam at 637.30m (bottom leaf of Newington) Hollies Barn Borehole, Windrush Formation; figs. 20, 21, seam at 714.65m (Top Broughton) Hollies Barn Borehole, Burford Coal Formation).

Figs. 9, 10. Specimens with fine ornament.

Figs. 11-15. Specimens with medium sized ornament; fig. 14, specimen with pseudo-cingulum.

Figs. 16-19. Specimens with coarse ornament, formerly assigned to *T. perverrucosa* (Alpern) Wilson and Venkatachala 1963.

Fig. 17. Specimen showing approach to *Torispora* type crassitude.

Figs. 20, 21. Large specimens with fine grana interspersed among coarse ornament.

Figs. 22-24. *Schopfites dimorphus* Kosanke, 1950 (figs. 22, 23, seam at 450.84m (Upper Milton) Steeple Aston Borehole, Witney Coal Formation; fig. 24, seam at 492.46m (Top Broughton) Apley Barn Borehole, Burford Coal Formation). Specimens to show variation in size and density of distal ornament.
Costae hardly project beyond amb. Inner body and laesura seldom seen.

**Holotype.** Pl. 1, fig. 8. Preparation T/130/1 in collection of NCB Yorkshire Regional Laboratory.

**Paratype.** Pl. 1, fig. 9. Preparation T/130/2.

**Locality and horizon.** Steeple Aston borehole, Oxfordshire, England. Upper Milton Seam at 450.84m, Witney Coal Formation, Upper Coal Measures.

**Description.** Distinctly foveolate ornament, fovea circular to oval, mean diameter 2μm and approximately the same distance apart. Operculum is similarly ornamented. Costae relatively weak, few in number, narrow 1.5-3.0μm in width, often fragmentary. Only the most strongly developed costae project beyond the amb. but not by more than 2.5μm. Operculum often partly detached or missing.

**Dimensions.** 63 (73.5) 87μm (19 specimens, macerated in fuming nitric acid). Holotype 75μm.

**Remarks.** The fovea closely resemble those of *V. fenestrata* (Kosanke & Brokaw) Wilson & Venkatachala, 1963. However *V. fenestrata* lacks costate sculpture.

**Comparison.** Differs from *V. burfordiensis* sp. nov. in having distinctly foveolate ornament and poorly developed costae.

**Occurrence.** Oxfordshire Coalfield, range in the coals extends from the middle of the Arenaceous Coal Formation to the upper part of the Burford Coal Formation. Below the Upper Milton Seam it is only sporadic but at other horizons within its range it occurs regularly but with an abundance of <0.1% of the seam assemblage.

**Vestispora magna** (Butterworth & Williams, 1954)

Wilson & Venkatachala, 1963

(Pl. 1, figs. 14-17)

**Remarks.** The figured specimens are from the type material and have been included to show some of the variation in exoexine ornament for comparison with *V. burfordiensis* sp. nov. and *V. witneyensis* sp. nov.

**RANGES AND ABUNDANCE OF SELECTED SPORES IN COAL SEAMS OF UPPER COAL MEASURES OF OXFORDSHIRE AND S.E. WARWICKSHIRE**

The species listed by Smith & Butterworth (1967), as being characteristic of their Assemblage Zone XI, were assumed by these authors to range throughout their Zone. The absence of a species from a particular horizon was not regarded as significant because of the somewhat fragmentary nature of the seam successions examined and the often sporadic occurrence of species within these successions. Also, although many samples were examined, they were often representative of thin seams whose precise correlation and stratigraphic position is uncertain.

With the exploration in Oxfordshire the opportunity became available to examine the ranges of species through a continuous succession of seams. The work on the three B.G.S. boreholes (Smith in Poole, loc. cit.) suggested that certain species had restricted ranges which could serve a useful stratigraphic purpose. The subsequent work on coals from the N.C.B. boreholes, involving the systematic study of the morphographic variations within the populations of the species in question, has made it necessary to modify the earlier conclusions. In particular, the coarsely verrucate forms of *Thymospora* referred to *T. perverrucosa* are now considered to represent an extreme development of the ornamentation of *T. pseudothiessenii* and the form referred to *Vestispora* of *V. magna* has been resolved into two new species of *Vestispora* with overlapping ranges.

Table 2 shows the frequency of occurrence of five species of miospores of proven stratigraphical value in the correlation of the Upper Coal Measures of Oxfordshire. The numerical abundance of these species is summarised below.

*T. pseudothiessenii* only occurs sporadically in the assemblages from seams of the Arenaceous and Witney Coal Formations and the Crawley Formation with a frequency generally <0.1% and only once exceeding 0.4% of the total assemblage. It is present in most seams of the Burford Coal Formation and Windrush Formation often exceeding 2% of the assemblage and reaching 14% in the Overthorpe Seam at Withycombe Farm. The base of the epibole occurs at the horizon of the Low Broughton Seam. It is, however, less abundant in seams at these horizons in the Kineton and Southam boreholes and is apparently absent from the assemblages from the Witney Coal Formation in these boreholes. However, in the Kineton borehole the species is well represented (3%) in the Marston (Napton) Seam, the youngest seam in the Windrush Formation.

*Vestispora burfordiensis* is a relatively uncommon (<0.1%) but persistent member of the assemblages from coal seams in the Burford Coal and Crawley Formations. Outside this range it was rarely recorded from the Newington and Upper Milton seams.

*Schopfites dimorphus* is a distinctive spore which is not easily overlooked. It ranges from the base of the Witney Coal Formation to the base of the Windrush Formation but is only a persistent member of the assemblages in the seams of the Burford Coal Formation where it may account for 1% of the assemblage. In general, however, it represents no more than 0.2% of the assemblage.

*Vestispora witneyensis* first appears as a rare and sporadic component of the assemblages in seams of the
### Ranges of Selected Spores in Some Other British Coalfields

**Thymospora pseudothiessenii**

In north Staffordshire thin coals from the Etruria Marl Formation, generally considered to be of Upper Westphalian C age, were sampled at an exposure and from two boreholes (Smith & Butterworth, 1967). Spores diagnostic of the *T. obscura* Zone including *Schopfites dimorphus* were present at all three localities and *T. pseudothiessenii* at two. It is difficult to reconcile the occurrence of these species in strata of presumed Upper Westphalian C age with the stratigraphic ranges established in Oxfordshire. However, Besley (1983) has shown that the base of the Etruria Marl Formation in north Staffordshire lacks chronostratigraphic significance, the base of the continuous ‘red bed’ facies being strongly diachronous. There is the possibility that locally the upper strata of the Formation may be of Westphalian D age. On the other hand, the spore evidence from one of the boreholes may be unreliable. Besley has pointed out that the material used for spore analysis was in the form of chippings from an uncased borehole thus allowing the possibility of contamination by spores from the overlying Newcastle Formation of Westphalian D age.

In Britain, seams of suitable rank are poorly rep-
sent in that part of the succession which includes the Westphalian C/D boundary. This is due in part to the hiatus of the Symon Unconformity. In the Forest of Dean, *T. pseudothiessenii* was not recorded from either the Trenchard or Coleford High Delf seams near the top of the Trenchard and the bottom of the Pennant Formations respectively. In South Wales, the species was not recorded from the Tillery (Brithdir) and Tillery Rider (Brithdir Rider) seams in the lower part of the Pennant Formation.

In coalfields situated to the north of the Wales-Brabant Massif, *T. pseudothiessenii* is common in certain seams in the Halesowen, or equivalent formations in strata which may be contemporary with the Burford Coal Formation in Oxfordshire. Unfortunately, the precise lithostratigraphic correlation of these beds is uncertain.

South of the Wales-Brabant Massif, *T. pseudothiessenii* is apparently absent from seams in the Farrington Formation of the Bristol and Somerset Coalfields and from the two seams in the middle of the Pennant Formation of the Forest of Dean. These horizons may therefore be the equivalent of strata from below the Burford Coal Formation in Oxfordshire in which *T. pseudothiessenii* is uncommon. However, this does not explain its absence from the Woor Green seam in the Supra Pennant Formation of the Forest of Dean in which *Schopfites dimorphus* occurs. No samples have been examined from the crucial parts of the species range in South Wales.

*T. pseudothiessenii* has been recorded, sometimes as an abundant element of the assemblages, from seams of the Keele Formation, or strata of equivalent age, in all the coalfields subject to palynological investigation in Britain.

Thus, with the exception of north Staffordshire, the range and abundance of *T. pseudothiessenii* in the British coalfields is generally in agreement.

**Schopfites dimorphus**

Apart from its apparently anomalous occurrence in seams of the Etruria Marl in north Staffordshire, *Schopfites dimorphus* has a similar range in all the British coalfields in which it occurs. It is confined to a relatively restricted portion of the Upper Coal Measures, namely to seams in the upper part of the Halesowen Formation and at the base of the overlying Keele Formation. Its absence from the Mynyddllswyn seam and from the Tillery and Tillery Rider seams in South Wales conforms to this pattern. The apparent absence of the species from seams in the Radstock Formation of the Bristol and Somerset coalfields is evidence that these beds are younger than the Burford Coal Formation of Oxfordshire, and may correlate with the Keele Formation elsewhere. However, its presence in the Woor Green seam from the Forest of Dean is interesting in view of the possible Cantabrian age assigned to the highest beds above the Household Coals on the basis of the fossil plants (Cleal in Ramsbottom, 1978).

**Punctatosporites granifer**

Small monolete spores with scabrate to granulate ornament were regarded as a single morphographic species by Smith & Butterworth (1967) and referred to *P. granifer*. In later work in Oxfordshire, Smith (in Poole, 1977, 1978) separated the more finely ornamented forms which he referred to *P. punctatus*. This was supported by the systematic study of monolete spores by Alpern & Doubinger (1973). There is thus no evidence in Britain relating to the range and abundance of *P. granifer* having the form illustrated in this paper.

**Vestispora spp.**

The two new species of *Vestispora* have only been recorded from seams in the small coalfield at Newent and from the Oxfordshire and southeast Warwickshire basins. It seems unlikely that these spores are confined to these coalfields.

**MIOspore Zones in the Upper Coal Measures of Oxfordshire**

In their study of miospore assemblages in British coalfields, Smith & Butterworth (1967) placed all assemblages containing *Thymospora* into a single zone, the base of which was roughly equivalent to the Westphalian C/D boundary. In his work on the Oxfordshire Coalfield, Smith, in Poole (1977, 1978) laid the foundation for subdivision of this zone. The present work proposes the replacement of Assemblage XI of Smith & Butterworth (1967) by a tripartite division, which is designated numerically XI-XIII to continue the zones recognised by Smith & Butterworth (1967). The relationship between the new zones and the stratigraphic ranges of the selected species used for their recognition is shown in Fig. 3.

The lithological horizons and the species associated with these zones are summarised in descending order below:

- **Zone XIII** – Seams of the Windrush Formation above the Newington Seam. *T. pseudothiessenii* is characteristic (coarse forms common). Other species listed below are absent.

- **Zone XII** – Seams of the Witney and Burford Coal Formations and the intervening Crawley Formation. The top of the Zone is placed immediately above the Newington Seam at the base of the Windrush Formation. *Schopfites dimorphus*, *Vestispora burfordiensis* and *V. witneyensis* all occur as characteristic species.
Fig. 3. Zonation of seams in the Upper Coal Measures of the Oxfordshire Coalfield based on the stratigraphic ranges of selected spores.
The base of the epibole of *T. pseudothiessenii* occurs at the base of the Burford Coal Formation in the upper part of the Zone. *P. granifer* is confined to the coals of the Milton Group.

**Zone XI** – Seams of the Arenaceous Coal Formation. *P. granifer* usually common. *Thymospora* represented by *T. obscura* and *T. pseudothiessenii* in small numbers. *V. witneyensis* is scarce.

It is considered premature to designate the zones by spore names before the miospore flora has been thoroughly documented. This will avoid the use of inappropriate names and confusion with the names already introduced by Peppers (1985) for a similar suite of biozones in strata of the Illinois Basin in North America considered to be equivalent in age to the Westphalian D of Europe. A comparison of the two zonal schemes shows some differences.

Peppers designated his three Assemblage Zones M1, CP and GD in ascending order. The base of his M1 Zone corresponds to the first appearance of *Cadiospora magna*, *Mooreisporites inusitatus* and the sporadic appearance of *T. pseudothiessenii*. These spores were recorded from the lowest seams from certain of the boreholes in the Oxfordshire Basin but because the lower part of the Upper Coal Measures is not present in the basin, the horizon at which these species first appear is not known. The lower and upper limits of Peppers' CP Zone and the limits of Assemblage Zone XII of the present work correspond to the first and last appearance of *Schopfites colchesterensis* and *S. dimorphus*. However, in Illinois the lower limit of the CP Zone is coincident with the base of the *T. pseudothiessenii* epibole, whereas in Oxfordshire the base of the epibole occurs in the upper part of Zone XII. In Illinois *T. pseudothiessenii* declines in abundance in seams above the range of *S. colchesterensis* whereas in Oxfordshire no such decline is apparent. The reduction in the numbers of *T. pseudothiessenii* is characteristic of the assemblages in Peppers' GD Zone, the uppermost of the three zones under discussion. The two species which designate this zone are *Lycospora granulata* and *Cappasporites distortus*. It is difficult to compare the range and abundance of *L. granulata* (Kosanke) since in this work the species has been included along with other Lycospores having a narrow flange and cingulum under the collective name *L. pusilla*. *C. distortus* (Urban) may be conspecific with *Apiculatisporites irregularis* (Alpern) Smith & Butterworth. The latter species is present throughout the Oxfordshire sequence but varies in abundance both within and between seams. It is unlikely that the features which mark the top of Peppers' GD Zone will be observed in Oxfordshire where strata of the appropriate age are probably lacking.

**BIOSTRATIGRAPHY**

Miospore assemblages from seams at the base of the Upper Coal Measures in Oxfordshire are characterised by species such as *T. obscura*, *Cadiospora magna* and *Mooreisporites inusitatus* which are diagnostic of the *T. obscura* Assemblage Zone XI of Smith & Butterworth (1967) and the lower part of the *T. obscura-T. thiessenii* (OT) Zone of Clayton et al. (1977). The base of the *T. obscura* Zone is determined by the first appearance of the monolete verrucate spores of the genus *Thymospora*. This genus is widely known from Carboniferous strata in Canada (Hacquebard, 1961), North America, Europe and Russia (Butterworth & Smith, 1976) and is obviously an important horizon for correlation (Laveine, 1976).

It has long been known that there is an approximate coincidence between the appearance of *Thymospora* and the Westphalian C/D boundary. However, attempts to use spores to define this boundary have been hindered partly by the lack of agreement concerning the palaeontological criteria used to define the boundary and partly from the lack of faunal and floral controls on the sections examined for spores. However, some progress has been made to resolve the problem. Since fossil plants provide a more useful basis for the biostratigraphic subdivision of the uppermost division of the Westphalian stage than the mussels, it is appropriate to examine the first appearance of *Thymospora* in relation to the record of the leaves of fossil plants. Laveine (1976) has shown that in those basins of N.W. Europe from which he examined the fossil floras, *Thymospora*, *Linopteris obtusa* var. *bunburi* and *Neuropteris ovata* appear in sequence over a very short stratigraphic interval. Recently Cleal (1984) has compared the plant distribution in the Saar with that in the Pennant Formation of South Wales and has proposed a tripartite division of the Westphalian D. He equates the base of the Westphalian D with the appearance of *L. obtusa* var. *bunburi* which in South Wales occurs near the top of the Rhondda Beds. This horizon is lower than that shown by Ramsbottom et al. (1978) who placed the base of the Westphalian D at the Hughes seam which marks the arbitrary boundary between the Phillipsi and Tenuis Chronozones and the base of Cleal's middle zone.

In South Wales only two seams have been examined from the appropriate horizon, namely the Tillery (Brithdir) and Tillery Rider (Brithdir Rider). According to the boundary established by Cleal, these seams are just above the base of the Westphalian D stage. *Thymospora obscura* was present in the assemblage from the Tillery Rider. Thus, on the basis of this admittedly somewhat slender evidence, *Thymospora* in South Wales only occurs in seams of Westphalian D age.
Unfortunately, Cleal has not established the base of Westphalian D in other coalfields from which *Thymospora* has been recorded and correlations in these coalfields are based on other evidence. It is worth detailing the stratigraphical evidence in order to highlight the problem of correlating the first occurrence of *Thymospora* with the Westphalian stage boundaries in Britain. Only coalfields with seams at the appropriate horizons are considered.

In Oxfordshire, *T. obscura* has been found in the lowest seams of the Arenaceous Coal Formation whose stratigraphy, excluding the evidence of the spores, is based on lithological considerations.

In the Forest of Dean, *T. obscura* has been recorded from the Trenchard Seam which is believed to be of late Westphalian C age (Phillipsii chronozone) though fossil evidence is lacking (Ramsbottom et al., 1978). In the Bristol and Somerset coalfields three seams from the Farrington Formation have been examined and all contained *T. obscura*. This formation contains typical Westphalian D floras (Ramsbottom et al., 1978). No seams from the underlying Mangotsfield Formation of Westphalian C, or possibly of lowest D, age (Ramsbottom et al., 1978) have been examined.

There remains the anomalous spore association previously mentioned from the thin coals at the base of the Etruria Marl facies in north Staffordshire. These red beds occur between the grey measures of the Blackband Formation containing faunas of the Phillipsii chronozone and the overlying Newcastle Formation belonging to the Tenuis chronozone (Ramsbottom et al., 1978).

If the stratigraphically lower base of the Westphalian D proposed by Cleal is accepted, then there is at present no unequivocal evidence for the occurrence of *Thymospora* spp in strata of Westphalian C age in Britain.

The base of the *Thymospora* epibole is a widely reported event of biostratigraphic significance occurring in mid Westphalian D times. In Oxfordshire it can be recognised in seams at the base of the Burford Coal Formation and in other British coalfields in seams which may correlate with this horizon in Oxfordshire. According to Laveine (1976) using the data of Alpern (1969) the base of the *Thymospora* epibole equates with the horizon of Tonstein 60 in Lorraine. Using the macrofloral ranges given by Donsonini (1981), Cleal (1984) considers Tonstein 60 corresponds to the base of his *Lobatopteris micromiltoni* Zone (middle zone of three into which the Westphalian D is subdivided) which he equates with the base of the Hughes Beds in South Wales. Laveine (1976), on the other hand, had earlier equated Tonstein 60 with the Five Feet seam in the Swansea Beds which overlie the Hughes Beds. It is worth noting that Cleal broadly equates his *L. micromiltoni* Zone with the lower part of Dix's Floral Zone I, whereas the Hughes Beds were placed in the older floral zone H by Dix (1934). In the present state of knowledge, it would be unwise to equate the base of the Burford Coal Formation in Oxfordshire with either the base of the *L. micromiltoni* Zone or with the Hughes Seam in South Wales.

In North America, Peppers (1985) comparing the ranges of miospores in the Pennsylvanian of Illinois with those in Western Europe, recognises an epibole in the numbers of *T. pseudothiessenii* whose base lies near the upper limit of the Spoon Formation (Desmoinesian Series).

Other species whose ranges in the Lower and Middle Westphalian D of Oxfordshire makes them locally useful for correlation purposes, and which may have a wider biostratigraphical significance, are *Schopfites* spp. and *Punctatosporites granifer*.

In northern France, Loboziak (1971) showed *S. colchesterensis* appearing somewhat later in the succession than *T. obscura*. However, according to Peppers (1985) the spore identified and illustrated by Loboziak (1971, pl. 4, fig. 25) as *S. colchesterensis* is closer to his *Schopfites* sp., which in Illinois appears at the same time as *Thymospora*. This form is less coarsely ornamented than *S. colchesterensis* and has not been recognised in Oxfordshire. In Illinois, Peppers (loc. cit.) shows that the ranges of *S. dimorphus* and *S. colchesterensis* (Smith & Butterworth, 1967) do not distinguish between these species) are confined to parts of the Spoon and Carbondale Formations and terminate well before the disappearance of *T. pseudothiessenii*. He equates these ranges with strata of mid Westphalian D age in Europe. Peppers cites the range in Britain, on the authority of Smith & Butterworth (1967), as in coals throughout the Westphalian D and in the Etruria Marl. However, the range now established in Oxfordshire shows much closer agreement with the range span in Illinois although in North America *S. colchesterensis* first appears at the base of the *T. pseudothiessenii* epibole, whereas in Oxfordshire *S. dimorphus* occurs somewhat earlier.

The abundance of *P. granifer* in seams of uppermost Westphalian C and Westphalian D age in boreholes outside the mining district of Limburg in the Netherlands led Van Wijhe & Bless (1974) to name their youngest miospore zone after this species. Unfortunately, due to the presence of a red bed sequence in the upper part of the succession in these Dutch boreholes, resulting in the elimination of the spores, no comparison is possible with the younger strata in Oxfordshire.

The range in Northern France (Loboziak, 1971) and in the Saar-Lorraine coalfields (Alpern et al., 1969) is given as Westphalian B-D; in the Campine basin of Belgium (Somers, 1971) as Westphalian C-D and in the Ruhr coalfield (Grebe, 1972) as Westphalian C although it is rare in the lower part. However, the
possible confusion between morphologically similar forms (Van Wijhe & Biess, 1974) combined with the limited development of seams in the basal Westphalian D in certain of these coalfields makes comparison difficult. Similarly, Peppers (1979), in his comparative studies, considers *P. granifer* to be a synonym of *Laevigatosporites punctatus*. There is thus no unequivocal evidence from N.W. Europe or from North America to corroborate the range in Oxfordshire where it is restricted to the presumed lower Westphalian D strata.

The Westphalian-Stephanian boundary has been traditionally placed at the unconformity at the base of the Holz Conglomerate. According to Cleal (1984) the fossil plants provide no unequivocal evidence of Cantabrian (basal stage of Stephanian) from strata below the Holz Conglomerate in the Saar-Lorraine Basins.

Above the Holz Conglomerate in these basins and at similar horizons in the Decize Basin and the basins of the Central Massif, significant changes take place in the miospore assemblages with an increase in the abundance of such species as *Spinosporeites spinosus* and *Polyomorphopsis spp.*, and the appearance of *Potonioisporites novicus* and *Aumancisporites striatus* accompanied by a reduction in the numbers of *Thymospora* spp. and the disappearance of *Lycospora* spp.

The disappearance of *Lycospora* spp., *Thymospora pseudothiessenii* and *Cappasporites distortus* (possibly conspecific with *Apiculatisporis irregularis* (Alpern, Smith & Butterworth) has been recorded in part by Kosanke (1950) and by Peppers (1964, 1970) in Illinois within the lower part of the Modesto Formation at the boundary of the Desmoinesian and Missourian Series. This event has been used by Peppers (1985) to define the boundary between his *L. granulata C. distortus* (GD) and *P. minitus P. obliquus* (MO) Zones which he equates with the Westphalian D/Stephanian boundary in Europe. It should, however, be noted that *Lycospora* spp. have been recorded from younger strata of the Appalachian area (Clendening, 1974). According to Clayton et al. (1977) in Western Europe *Lycospora* spp. remain abundant throughout the Stephanian until the Autunian.

In Oxfordshire both *Lycospora* spp. and *A. irregularis* are well represented in the younger seams and there is no evidence of miospore assemblages of post-Westphalian character.

It is hoped that the recognition of the proposed miospore assemblage zones may assist the inter-coalfield correlation of strata of Upper Coal Measures age in Britain. It may prove more difficult to relate the miospore zone boundaries to horizons established from macroplant remains since the classic collections of the latter from the South Wales Coalfield are from areas where the rank of the coals of the appropriate age is too high to obtain recognisable miospores.

ACKNOWLEDGEMENTS

The author wishes to thank the National Coal Board for the opportunity to carry out this work and for permission to publish the results. Special thanks are due to Ron Goossens, Chief Geologist of the N.C.B., and to the two members of his staff, Phil Eaton and Eddie McNeirnie, with special responsibilities for the Oxfordshire and South Warwickshire Prospects, for helpful information on seam correlation and nomenclature. The author is also indebted to Pamela Spriggs for practical help and for information based on her own observations.

Manuscript received February 1986
Revised manuscript accepted June 1986

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