Article

Microbial Community Composition in Thermal Waters from the Lindian Geothermal Field (Songliao Basin, North-Eastern China)

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Abstract: Geothermal systems represent discrete and relatively homogenous habitats for extremophiles; investigation into the microbial community is key to revealing the geochemical environment and the geochemical evolution of fluids in geothermal reservoirs. The reservoir of the Lindian geothermal field in Northeast China, is highly reducing and rich in methane, but the pathways of methane generation and the related microbial community structure are still unclear. In this research, five thermal water samples were collected and tested, and the microbial community structure and diversity were analyzed. The results show that in the sandstone reservoir belonging to the low-temperature (reservoir temperature < 90 °C) brackish water (total dissolved solids concentration between 1000 and 10,000 mg/L) environment, the richness of the microbial community is relatively high. The microbial community structure is different from other geothermal systems reported but similar to that of oilfields, which may be related to the highly reducing geochemical environment with abundant organic matter. According to the analysis of archaeal function, the biogas production in the Lindian geothermal field is dominated by hydrogen nutrition type methane production, while the H2 reducing methylamine type methane production is secondary, and results of Pearson correlation show that the archaeal communities are more strongly correlated to physicochemical factors than the bacterial communities.

Keywords: thermophilic bacteria and archaea; microbial community; methanogens; Lindian geothermal field

1. Introduction

Microorganisms are the main form of life in subsurface ecosystems. The groundwater environment has nutrients, water, pH, osmotic pressure, temperature and other conditions needed for microbial growth and forms a special living habitat for microorganisms. Meanwhile, microorganisms are also an important driving force in the hydrogeochemical evolution of geofluids [1,2]. Geothermal systems represent discrete and relatively homogenous habitats for extremophiles. These extremophiles can live in the thermoboitic environments that are above the ‘thermophile boundary’ of 55 to 60 °C [3], and participate in many redox reactions, such as sulfate and nitrate reduction, hydrogen oxidation and methanogenesis [4,5], which shapes the water chemistry across time and space. Therefore, an investigation into the composition of the microbial communities in geothermal systems is key to reveal the mechanism of geochemical evolution in geothermal reservoirs, and
moreover, predict the water–rock interactions mediated by microorganisms under human activities, such as reservoir clogging during geothermal reinjection and CO₂ sequestration in deep saline aquifers [6,7].

Recent studies have shown that the structure of microbial communities evolves across time and space, mainly driven by environmental selection and dispersal limitation [8,9]. However, the relative contribution of these ecological factors is still in debate, with different findings reported across terrestrial and marine water bodies [10]. Among terrestrial water bodies, shallow groundwater is the most studied [11,12]. These studies suggest that the microbial community composition is influenced predominantly by substrates (e.g., minerals and organic matter) and water chemistry [13–15]. As the environment of geothermal reservoirs is under relatively high temperature, pressure and salinity, and usually more oligotrophic with respect to shallow groundwater aquifers, the environmental parameters controlling the community structure are different. Most studies have implied that temperature and pH are the primary drivers of bacterial and archaeal communities in these ecosystems, and other geochemical variables, especially substrates, such as hydrogen and nitrogen, may also contribute to community structure [10]. However, among the existing studies on geothermal systems, the microbial communities of relatively high-temperature hot springs, such as Yellowstone National Park in the United States, Taupō Volcanic Zone (TVZ) in New Zealand, Padang Germin hot springs in Indonesia, and Tengchong hot springs in Yunnan in China, are the most studied [16–21], whereas low-temperature geothermal systems are less investigated, a lack of cases has hindered a comprehensive understanding of the microbial biogeography and the influencing environmental parameters.

The Lindian geothermal field located in northern Songliao Basin, Northeast China, is a typical low-temperature geothermal system. Geothermal fluids in the reservoir are abundant in methane. Previous studies have suggested that the methane is mainly biogenic, i.e., from the decomposition of immature organic matter by microorganisms in a reducing environment [22,23], but the characteristics of the microbial communities have not been fully studied. Therefore, the present research collected representative geothermal water samples from the Lindian geothermal field and performed microbial community analysis of the bacterial and archaeal population (16S rRNA gene amplicon sequencing) and gene function prediction based on the FAPROTAX database; then, analyzed the structure and diversity of the microbial communities, and explained the genesis of methane. This study will enrich the understanding of the composition and habitat features of the microbial communities in low-temperature geothermal reservoirs.

2. Materials and Methods

2.1. Study Area

The Lindian geothermal field is located in the Songliao Basin, Northeast China, covering an area of ca. 3500 km² (Figure 1). The area was the sedimentary center of the basin in the Mesozoic and Cenozoic, where a series of fluvial-lacustrine sediments was deposited greater than 3000 m in thickness, including the Craterous, Lower-Middle Eogene and Quaternary. The main geothermal reservoirs consist of part of the Lower and Upper Craterous sandstones; the burial depth of the roof and bottom of the reservoirs are 980–1500 m and 1670–2430 m, respectively. The reservoir temperature falls between 47.4 and 89.0 °C, thus belonging to a low-temperature geothermal system. As some strata constituting the reservoir were deposited under a deep-water environment, the reservoir is abundant in biogas source rocks [24]. Investigation of the geothermal fluids has shown that the main gas component is methane, which accounts for 53.7–95.1% of the dissolved gas in volume [25].
2.2. Sampling and Analysis

The field campaign was carried out in January 2019. Five geothermal wells in operation were chosen for microbial sampling (Figure 1). These wells are located in the center of the geothermal field, and the profile from well L01, L08, L10, L60 to L50 are consistent with the regional flow path of the thermal water. In total 5 water samples (L01, L08, L10, L60 and L50) were collected.

The water samples were analyzed both for water chemistry and microbial communities. Water temperature (T), pH and total dissolved solids (TDS) were measured in situ using a portable multi-parameter (Hach HQ40D, Hach, Loveland, CO, USA), the resolutions are 0.3 °C, 0.01 and 0.1 mg/L, respectively. Water samples for SO$_4^{2-}$ analysis were collected using 50 mL higher-density polyethylene (HDPE) bottles after filtration with a sterile 0.22 µm membrane filter and were measured with Dionex DX-500 ion chromatography (Dionex, Sunnyvale, CA, USA), with a precision better than 5%. The sampling method for rare earth elements (REEs) was similar to that of SO$_4^{2-}$, except for that the water samples were acidified by HNO$_3$ to pH < 2 and measured using NexIONTM 300D ICP-MS (PerkinElmer, Waltham, MA, USA) with a precision better than 3%. Total organic carbon (TOC) samples were collected using 50 mL bottles and measured with Shimadzu TOC-L CPH Analyzer (Shimadzu, Nishinokyo Kuwabara-cho, Kyoto, Japan), with a detection limit of 4 µg/L. δ$^{13}$C$_{DIC}$ ($δ^{13}$C of dissolved inorganic carbon) samples were collected using 500 mL HDPE bottles avoiding any air contamination and were measured with an Isoprime 100 stable isotope ratio mass spectrometer (Isoprime, Cheadle Hulme, Cheadle, UK). The results were reported in ‰ relative to the standard VPDB (Vienna Pee Dee Belemnite) with a precision better than 0.1‰. These measurements were carried out in the Measurement Science and Experiment Center, Jilin University.

The microbial samples were collected using 18 L high-temperature resistant sterile plastic buckets, transported to the laboratory and filtered onto polytetrafluoroethylene membranes with a pore size of 0.22 µm; then the membranes were frozen and sent to Shanghai Sangong Biotech Co., Ltd. for microbial community analysis. The samples were analyzed by Illumina Miseq (Illumina, San Diego, CA, USA) sequencing. Genomic DNA was extracted using an E.Z.N.A® Mag-Bind Soil DNA Kit (Omega Bio-Tec, Doraville, GA, USA).
Two rounds of polymerase chain reaction (PCR) amplification were performed for the bacteria. The first round of amplification with bacterial 16S rDNA genes (V3–V4) from the isolated DNA was performed using the primers 341F (CCTACGGGNGGCWGCAG) and 805R (GACTACHVGGGTATCTAATCC) combined with the Miseq sequencing, while compatible primers of PCR were introduced in the second round of amplification. The PCR amplification of archaea included three rounds. In the first round, the isolated DNA was amplified with M-340F (CCCTAYGGGGYGCASCAG) and GU1ST-1000R (GGCCATGCACYWCTYCTC) primers; the second round of amplification with archaea 16S rDNA genes (V3–V4) was conducted with the primers 349F (GYGCASCAGKCGCMGAAG) and 806R (GGACTACVGGGTATCCTAAT), and the third round of amplification was the same as the second round of bacterial amplification. The DNA after PCR was then purified by Agencourt Ampure XP Magnetic Beads (Beckman Coulter, Indianapolis, IN, USA) and quantified by a Qubit 2.0 TM DNA detection kit (Thermo Fisher Scientific, Waltham, MA, USA). The purified amplicons were sequenced on an Illumina MiSeq platform. The bio-information of raw DNA data was analyzed based on the database of Ribosomal Database Project (RDP), Silva and the National Center for Biotechnology Information (NCBI). The obtained DNA reads were then clustered to operational taxonomical units (OTUs) with 97% similarity.

3. Results and Discussion

3.1. Characteristics of Geothermal Reservoir Environment

The hydrochemistry type of the geothermal waters is HCO$_3$·Cl-Na. The depth of the sampling wells is 1800–2718 m, the reservoir temperature predicted by the non-steam loss non-mixing quartz geothermometer proposed by Fournier [26] is 72.9–89.0 °C [27], while the pH is 8.14–8.93, TDS is 1440–1699 mg/L, TOC is 323.4–477.1 mg/L and $^{13}$C$_{DIC}$ ranges from $-12.11$ to $-8.07$‰VPDB (Table 1).

| Sample ID | Depth (m) | T* (°C) | pH | TDS (mg/L) | TOC (mg/L) | $^{13}$C$_{DIC}$ (%VPDB) | Ce/Ce* | Eu/Eu* | SO$_4^{2-}$ (mg/L) |
|-----------|-----------|---------|----|------------|------------|-------------------------|--------|--------|------------------|
| L01       | 2718      | 89.0    | 8.17 | 1563       | 323.4      | $-9.60$                 | ND     | 10.95972 | 0.101            |
| L08       | 1802      | 77.6    | 8.40 | 1440       | 477.1      | $-9.28$                 | 0.539  | 6.467552 | 0.053            |
| L10       | 2002      | 72.9    | 8.93 | 1544       | 367.4      | $-10.36$                | ND     | 6.818943 | 0.189            |
| L50       | 2250      | 78.9    | 8.14 | 1699       | 417.6      | $-8.07$                 | ND     | 7.392759 | 0.189            |
| L60       | 2000      | 84.3    | 8.20 | 1670       | 441.2      | $-12.11$                | 0.295  | 5.778223 | 0.466            |

Rare earth elements Ce and Eu in groundwater are sensitive to the redox environment. Groundwater environments different in redox state often show positive or negative anomalies of Eu and Ce normalized against North American Shale Composite (NASC) [28–30], i.e., the concentration of Eu or Ce in the sample is higher or lower than that in NASC (Eu* or Ce*), respectively. Generally, Eu/Eu* $< 1$ or Ce/Ce* $> 1$ indicates an oxidizing environment, while Eu/Eu* $> 1$ or Ce/Ce* $< 1$ represents a reducing environment [31]. As all the thermal water samples collected show Ce/Ce* $< 1$ and Eu/Eu* much higher than 1 (Table 1), indicating that the geochemical environment of the geothermal reservoir is highly reducing.

3.2. Microbial Community Structure

3.2.1. Bacterial Community Structure

In the thermal water samples, Proteobacteria (average abundance 94.97%) and Bacteroidetes (1.82%) are the dominant phyla, among which Gammaproteobacteria (78.75%) dominates, followed by Alphaproteobacteria (11.41%) and Betaproteobacteria (4.86%) (see Figure S1). In previous studies, Proteobacteria usually do not have absolute advantage over the other phyla in most low-temperature geothermal fields. For example, the Firmicutes
and Proteobacteria in the deep karst-fissure geothermal reservoir in the Xiongan New Area account for 38.70% and 37.61% of the microorganism, respectively [32], while Bacteroidetes are dominant (ca. 80%) in the Unkeshwar hot springs of India (water temperature 42–60 °C) [33]. Actinobacteria and Aquificae are the dominant bacteria in the hot springs in Coffee Pot of Yellowstone National Park, USA (water temperature 39–74 °C, pH 5.8–6.9), while Proteobacteria and Aquificae accounts for 34.2% and 31% of total average relative abundance in TVZ, respectively [10]. Thus, the absolute domination of Proteobacteria in the thermal waters from Lindian is not very common, which probably relates to the high organic matter content (143 g/kg on average) in the reservoir rock and thermal water [27].

At the genus level, Acinetobacter (average abundance 73.80%), Novosphingobium (6.21%), Pseudomonas (2.91%), Alphaproteobacteria, Rhizobium (1.46%), Aquabacterium (1.40%), Brevundimonas (1.37%) and Hydrogenophaga (1.03%) dominate, among which Acinetobacter is the dominant species, and Proteobacteria are the only bacterial species with an abundance greater than 1% at the genus level (Figure 2).

Acinetobacter and Pseudomonas are found abundant in oil reservoirs and are currently known as the genera with hydrocarbon degradation function [34], which are aerobic mesophilic bacteria [35,36], e.g., Acinetobacter are the dominant taxa (19–85%) in the Daqing Oilfield adjacent to the Lindian geothermal field [1], indicating that the reservoir environment is to some extent similar to oil reservoirs, such as similar temperature range and organic matter content in reservoir rocks.

Figure 2. Relationship between samples and taxa of bacterial communities at the genus level.
3.2.2. Archaeal Community Structure

The archaeal community (abundance > 1%) is mainly composed of Euryarchaeota (average abundance 79.12%), Thaumarchaeota (8.85%) and Woesearchaeota (2.83%). Among them, Euryarchaeota dominates. Euryarchaeota, traditionally known as methanogens, are typically involved in sulfur, nitrogen and iron cycling in addition to hydrocarbon conversion [37]. The archaeal community in the deep karst geothermal reservoir in the Xiongan New Area is dominated by Euryarchaeota and Crenarchaeota (average abundance 93%) [38]. The archaeal taxa with abundance greater than 1% mainly include Methanosarcinales (48.66%), Methanomicrobiales (16.43%), Nitrososphaerales (8.78%), Methanobacteriales (7.60%), Methanomassiliicoccales (4.85%) and Woesearchaeota (2.83%) (Figure S2).

At the genus level, the archaeal community (abundance >1%) mainly includes Methanothrix (average abundance 35.62%), Methanosarcina (10.97%), Methanocystococcus (10.32%), Methanothrix (35.62%), Methanosarcina (10.97%), Methanoculleus (10.32%), Nitrososphaera (8.78%), Methanomassiliicoccus (4.85%), Methanobacterium (3.67%) and Methanospirillum (3.29%), Methanothermobacter (2.17%), Woesearchaeota Incertae Sedis Ar16 (2.13%), Methanomethylovorans (1.95%), Methanosphaerula (1.28%) and Methanobrevibacter (1.07%). Except for Woesearchaeota Incertae Sedis AR16 and Nitrososphaera, which are part of a branch of ammonia-oxidizing archaea (AOA) generated during microbial nitrification [39], all the archaea with an abundance greater than 1% belong to Euryarchaeota (Figure 3).

Figure 3. Relationship between samples and taxa of archaeal communities at the genus level.
3.3. Mechanism of Methanogenesis

The results show that methanogens dominate the archaeal community in the Lindian geothermal field. The existence of methanogens indicates that the groundwater environment is extremely reducing. Methanogens combine H\(_2\) and CO\(_2\) to drive energy metabolism and produce CH\(_4\) as a metabolite. Methanogens use an ancient carbon dioxide fixation pathway called the reductive acetyl-coA pathway \[40\]. Methanation is the final step in the anaerobic fermentation process. In this process, acetic acid and methylamine, H\(_2\) and CO\(_2\), and methylamine and methanol are transformed into CH\(_4\) and CO\(_2\) through acetoclastic methanogens, hydrogenotrophic methanogenesis and methylotrophic methanogenesis, respectively \[41\]. Methane production is a strictly anaerobic process occurring in sulfate-deficient environments; methanogenic activity occurs only in the case when sulfate is depleted or at extremely low concentrations. As the concentration of SO\(_4^{2−}\) in the reservoir of the thermal water is very low in the Lindian geothermal field (Table 1), the reservoir environment is highly sulfate deficient, thus is favorable for methanogenesis. Moreover, the methanogen community structure is another important factor affecting the intensity of methane production. According to the analysis of available substrates at methanogen level in the study area (Table 2), the most available substrates of methanogens with high abundance in the reservoir are H\(_2\) and CO\(_2\), which also confirms that hydrogenotrophic methanogenesis is the primary pathway.

### Table 2. Characteristics of the methanogens at order level \[42\].

| Methanogens | Methanobacteriales | Methanomicrobiales | Methanosarcinales | Methanomassiliicoccales |
|-------------|--------------------|--------------------|-------------------|------------------------|
| Representatives | *Methanobacterium bryanlii*, *Methanobrevibacter wolfeii*, *Methanosphaera stadtmanae* | *Methanoculleus bourgensis*, *Methanogenium cariaci* | *Methanosarcina barkeri*, *Methanaosa concilii* | *Methanomassiliicoccus luminyensis* |
| Substrate | H\(_2\) + CO\(_2\) (formic acid, CO, ethyl alcohol, methyl alcohol) | H\(_2\) + CO\(_2\), formic acid (ethyl alcohol) | Methanol, methylamine, acetic acid (H\(_2\) + CO\(_2\)) | H\(_2\) + methylamine (methyl alcohol) |
| Temperature range (°C) | 15–97 | 0–60 | 2–70 | 37 |

Based on the FAPROTAX database, the bacterial community function is predicted. The results show that the bacterial community functions with the highest abundance are chemoheterotrophy, aerobic_chemoheterotrophy, animal_parasites_or_symbionts, human_pathogens and aromatic_compound_degradation (Figure 4). The main function is to degrade macromolecular organic matters, i.e., organic matter is broken down by fermentation bacteria into substrates available to methanogens (e.g., hydrogen, carbon dioxide, Acetic acid, methanol and dimethyl sulfide), which is also identified by the negative correlation between archaeal richness and \(^{13}\)δ\(_{\text{DIC}}\), as \(^{13}\)δ\(_{\text{DIC}}\) is depleted due to fractionation during the decomposition of organic matters \[43\].

The predicted functions of the archaeal community with the highest abundance are methanogenesis, hydrogenotrophic_methanogenesis and methanogenesis by CO\(_2\) reduction with H\(_2\) (Figure 5). It can be seen that the predicted abundance of methane production function of CO\(_2\) reduction by H\(_2\) is higher, while the methanogenesis by the reduction of methyl compounds with H\(_2\) is secondary (abundance 1.9%). This is consistent with the reservoir lithology, which is mainly sandstone and mudstone rich in organic matters, thus facilitating H\(_2\) and CO\(_2\) production and hydrogenotrophic methanogenesis \[44\].
3.4. Microbial Diversity

Alpha diversity analysis is conducted to characterize the abundance and diversity of microbial communities, including a series of statistical indexes, such as Shannon, Simpson, Chaol1, ACE indexes and Coverage, among which Shannon, Simpson indexes and Coverage are used to reveal the distribution diversity of the microbial communities. A larger Shannon index indicates a higher community diversity, while a larger Simpson index indicates a lower community diversity; Coverage refers to the coverage of each sample library. The higher the value is, the lower the probability that the sequence in the sample is not detected. This index actually reflects to what extent the sequencing result represents the real composition of the sample; Chaol1 and ACE indexes are used to estimate the number of OTUs in the community, which can reflect the abundance of the microbial community. Moreover, Chaol1 is often used to estimate the total number of species in the microbial community. Higher Chaol1 and ACE indices indicate higher community richness.
A total of 677,437 original sequences were obtained by high-throughput sequencing of five samples in this study, and 21,377 OTUs were generated, among which 13,079 OTUs were generated by the bacterial community, and 8298 OTUs by the archaeal community, accounting for 68% and 32% of the whole microbial community, respectively. Generally, OTUs of geothermal waters are low, e.g., the OTUs of the thermal waters are 284–564, 49–2997, 224 and 2398 from the Xiongan New Area [45], TVZ [10], the Western Plain of Romania [46] and Sri Lanka [47], respectively. Thus, the total number of microbial taxa in thermal waters from the Lindian geothermal field is remarkably abundant, indicating abundant substrates, i.e., the high organic matter contents in the reservoir (Table 1).

It is generally believed that in the geothermal environment of low temperatures (50–90 °C), microbial communities are dominated by bacteria, while archaea may only be dominant when the temperature is higher than 90 °C [48]. In this research, the reservoir temperatures of the sampling wells are 72.9–89.0 °C, and the sequencing result shows that archaea are not dominant, which is consistent with common findings, but the proportion of archaea in the microbial communities is much higher than those in most cases reported in previous studies. For example, in the deep groundwater of Yinchuan (sampling depth 230–600 m), the proportion of archaea is 0.84% [49], while in the Xiongan New Area, the proportion is 2.5% (sampling depth 1500–3853 m, temperature 60–83 °C, slightly alkaline, brackish thermal water) [38], and ca. 0.4% in TVZ (temperature range of hot springs 13.9–100.6 °C) [10]. This is unusual, which is probably related to the specific geochemical environment of the reservoir, such as high organic matter content (TOC up to 477.1 mg/L) and highly reducing (dissolved oxygen < 2 mg/L) [27].

Biostatistical analysis shows that the community coverage of the samples is all greater than 95%, indicating that the sequencing amount can represent the composition of the microbial community. The Chao1 index of the samples ranges from 17,465.49 to 80,566.93, while the Shannon and Simpson indexes ranges are 0.79–3.49 and 0.13–0.81, respectively (see Table S1); the Chao1 index of the archaea communities ranges from 25,917.06 to 67,299.77, while the Shannon and Simpson indexes are 2.45–4.53 and 0.03–0.30, respectively (see Table S1), indicating a relatively large variation in microbial diversity. The Mann–Whitney U test was performed to compare the diversity indexes, as the sample is small (<30), and $p < 0.1$ was interpreted to be statistically significant.

According to the ACE and Shannon indexes, the total number of the archaeal taxa of samples L10 and L50 are greater than those of the bacterial taxa, respectively, and the diversity of the archaea taxa is greater than that of the bacteria for the other four samples, except for L10 (see Table S1). Samples L08 and L10 are richer in archaea taxa, while the other three samples are more abundant in bacterial taxa (see Table S1). With respect to bacterial communities, sample L08 has the highest richness but the lowest diversity, while sample L10 has the lowest richness but the highest diversity. In the archaea communities, sample L10 has the highest richness, while sample L08 has the lowest richness, sample L50 has the highest diversity, and sample L08 has the lowest diversity. However, no significant difference exists in the ACE index between the archaeal and bacterial communities ($p = 0.310$, both for ACE), whereas the difference in the Shannon index is significant ($p = 0.095$). The variation in microbial richness and diversity reflects the spatial heterogeneity of physicochemical conditions typical in geothermal systems.

4. Conclusions

Both the bacterial and archaeal community compositions of the thermal waters from the Lindian geothermal field were investigated and accounted for 68% and 32%, on average, of the microbial taxa, respectively. The proportion of archaeal taxa is far higher than other thermal waters reported, and the microbial richness is remarkably high, which can be related to the high organic matter content and methanogenesis in the geothermal reservoir.

The existence of methanogens indicates a strong reducing groundwater environment in the geothermal reservoir. The prediction of microbial community function shows that the main function of bacteria is to decompose macromolecular organic matter into substrates
facilitating methanogens, while the main function of the archaea is to produce methane through CO₂ reduction by H₂, indicating that hydrogenotrophic methanogenesis is the primary pathway of methane production in the Lindian geothermal field.

**Supplementary Materials:** The following supporting information can be downloaded at: https://www.mdpi.com/article/10.3390/w14040632/s1, Table S1: OTU and taxonomy of archaea; Table S2: OTU and taxonomy of bacterial; Table S3: Diversity of archaeal and bacterial communities of the thermal waters; Figure S1: Phylum level bar chart for bacterial taxa; Figure S2: Phylum level bar chart for archaeal taxa.

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**Data Availability Statement:** All data are reported here and the raw sequencing data are available from the authors upon request.

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