Hydrometeorological triggering of periglacial debris flows using a Bayesian approach: a case study of the Hailuogou Gully region, China

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Abstract
Many debris flows have occurred in small basins with a small glacier cover or snow cover in periglacial areas, suddenly causing high mortality and great damage. The hydrometeorological conditions that caused such debris flows are complex, making forecasting and early warning difficult. Previous studies for these small-glacial-covered basins have primarily considered rainfall as the only inducing factor of debris flows, and often the effects of temperature are neglected. Thus, we carried out a probabilistic analysis of variables derived from hydrometeorological factors for the Mount Gongga region, Sichuan, China, where debris flows were recorded on 14 days between 1988 and 2019. By analyzing hydrological characteristics when debris flows occurred, three distinct dominant trigger types could be identified. The results show that seven (50%) of the observed debris flow events during the study period, high-intensity rainfall was the dominant trigger, snowmelt by high temperature was identified as the dominant trigger for 2 (14%). Furthermore, five (36%) debris flow events could be attributed to the combined effects of long-lasting (or short-medium) rainfall and sustained higher temperatures. We find that the differences between the trigger types are statistically significant, and a susceptibility prediction differentiating between trigger types can outperform simple rainfall-only situations. This study contributes to an improved understanding of the hydrometeorological impact on debris flow initiation in high elevation watersheds.

Keywords Periglacial debris flows · Meteorological causes · Bayesian probability · Early warning threshold

Abbreviations
HLG Hailuogou
GJTX Guanjingtaixi
HBLG Huangbengliu

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1 Introduction

Debris flows containing large quantities of particles are common phenomena observed in mountainous areas worldwide (Pavlova et al. 2011). It is known that short-term rainfall with a high intensity can trigger debris flows (Berti et al. 1999; Caine 1980; Wilson and Wieczorek 1995). Many studies have analyzed thresholds (antecedent rainfall), rainfall intensity, and duration as triggering factors to understand debris flow initiation (Caine 1980; Glade et al. 2000; Guzzetti et al. 2007, 2008; Osanai et al. 2010). Periglacial debris flows form in glacier and snow-covered areas at a distance from the front of the glacier or near the snow line. The occurrence of debris flows in such regions is triggered by long-term continuous rainfall and short-term heavy rainfall (Chiarle et al. 2007), ice and snow melting (Decaulne et al. 2005; Matsuura et al. 2013), and snow melting and rainfall together (Church and Miles 1987; Malet et al. 2005; Mostbauer et al. 2018). Therefore, studying the association between debris flow events and climatic factors is crucial for predicting and early warning of debris flows. Although many previous studies have been conducted, problems still exist in the effective prediction of debris flows’ initiation in a certain area. Therefore, further research is required to minimize casualties and economic losses.

Most of those methods applied to predict debris flows are derived from the empirical intensity–duration (ID) threshold model used in landslide prediction. This method is used to identify typical rainfall characteristics such as intensity or duration within a certain period that trigger debris flows (Caine 1980; Copeland 2009; Guzzetti et al. 2007; Johnson and Sitar 1990; Zhang et al. 2020). Empirical rainfall thresholds were calculated from historical databases and rely on observed correlations between rainfall amount and debris flows occurrences (Guzzetti et al. 2008; Meyer et al. 2012; Nikolopoulos et al. 2014). However, rainfall is not the only factor that causes debris flows (Pavlova et al. 2014). Undoubtedly, there is a certain degree of uncertainty in determining rainfall thresholds, and as an empirical method, the inconsistency in basic data can often increase uncertainty. Therefore, the researchers have tried to solve these problems by complementing the rainfall threshold information on probability models (Frattini et al. 2010). Walder (1995) first analyzed the impact of temperature on the triggering of debris flows in the Rainier Mountains by a simple conditional probability method. Jomelli et al. (2007) and Pavlova et al. (2014) used the logistic regression (LR) method to study the meteorological conditions of debris flows that occurred in the Alps under different environmental conditions. Many Bayesian probability models have been successfully applied to implement early warning of debris flows in Austria and other Alpine regions (Prenner et al. 2018; van den Heuvel et al. 2016). These models can effectively evaluate the impact of different meteorological parameters on the initiation of debris flows and also use probability thresholds to evaluate the susceptibility to debris flows (Jomelli et al. 2015; Kaitna et al. 2019; Mostbauer et al. 2018; Prenner et al. 2018; Turkington et al. 2016). The Bayesian method has been widely used in Alpine glacier regions, where most of the study areas were watersheds with high latitudes and a large proportion of glaciers, and contrastingly, it has been applied to many basins located in low latitudes with a small proportion of glaciers or no glacier. In the small (or non-) glacier basin around periglacial areas, the melting of ice and snow impacts the initiation of debris flows, but it is often ignored. Currently, research on the characteristics and early warning thresholds of debris flows in this kind of periglacial area are still quite limited.

Mount Gongga is located in the southeast of the Qinghai–Tibet Plateau in China. Many temperate glaciers are present in the area, and it is the largest modern glacier center in the Hengduan Mountains. The Hailuogou glacier on the eastern slope of Mount Gongga is the
lowest modern temperate glacier in China. Its contraction can be regarded as a response to global warming, and significant changes in debris flow activity are caused by different climatic conditions (Turkington et al. 2016). In most cases, heavy rainfall is the greatest role in triggering debris flows, while in a few cases, it is due to other meteorological combination effects. Based on 32 years of meteorological data in the area, this study classifies the triggering type of debris flows through climatic conditions. The one-dimensional Bayesian method was used to analyze susceptibility to debris flows using different meteorological parameters, and the two-dimensional Bayesian method was used to calculate and obtain the initiation probability of debris flow in different modes. This study aims to analyze the impact of different climatic conditions on the initiation of periglacial debris flows and propose an early warning threshold for this area. Furthermore, these models provide a reference for early warning of periglacial debris flows in other areas.

2 Study area

2.1 Basin overview

Mount Gongga (29.68° N, 101.98° E) is located in the center of the Hengduan Mountain in the southeast of Qinghai–Tibet Plateau. There are 72 temperate glaciers around Mount Gongga, with a total area of 255.1 km², 5 of which glaciers have a length of more than 10 km each. Glacier changes in the region are highly sensitive to temperature changes. Mount Gongga is located at the Longmenshan structural belt—Xianshuihe fault system—Chuanidian S–N tectonic zone junction with active neotectonics and frequent seismic activity. Since the Quaternary, glacier activity has intensified, causing the formation of many moraines on the bottom and sides of the valleys, thereby providing abundant sediment for debris flow development. Geomorphologically, the study area is characterized by the presence of high terrains, high mountains, and deep valleys. Differences in terrain height are large and, thus, provide the potential energy needed to form debris flows. The west slope of Mount Gongga has a subtropical plateau climate and experiences the southwest monsoon with low temperatures and little precipitation. And the east has a subtropical mountainous humid monsoon climate and is warm in winter and wet in summer. In this area, the annual average precipitation and temperature is 1862.5 mm and 4.5 °C, respectively. The cumulative average rainfall from May to September is approximately 1325.74 mm, accounting for approximately 70% of the average annual rainfall. The abundant rainfall, snow, and ice melt in the area provided sufficient water sources of debris flows’ initiation.

The Hailuogou (HLG) basin of the eastern slope of Mount Gongga is between 101° 52’ E–102° 07’ E, and 29° 31’ N–29° 37’ N. The highest altitude point of the basin is the peak of Mount Gongga at 7556 m, the lowest is at 1460 m, and the snow line is approximately 5000 m (a.s.l.) (Liu et al. 2010). There are seven glaciers in the basin, among which glacier no. 3 is the largest, with an area of 24.84 km² and the nadir of its ice tongue at 3000 m (a.s.l.). The smallest glacier (No. 7) is located in Guanjingtaixi (GJTX), with an area of only 0.09 km². There are no glaciers covered in the Huangbengliu (HBLG) Basin, but a large snow-covered area in winter and spring and permanently snow-covered at high elevations. The moraines produced by the ancient and modern glaciers are widely distributed in the two channels: this may provide abundant sediment to initiate debris flows under the action of precipitation or meltwater. The HBLG and the GJTX are located on the left flank of the upper reaches of the main channel of the Hailuogou Gully. The HBLG is
oriented north–southeast, and the GJTX is oriented north–south. The bedrock in the area is composed of various Permian schists intercalated with quartzite, marble, slate, and crystalline limestone. The soil formed by weathering of this type of rock has a high content of clay particles and supports the formation of viscous debris flows.

2.2 Available data

The meteorological data used in this study was collected from the Gongga Alpine Ecosystem Observation Station (GAEOS) of the Chinese Academy of Sciences. The GAEOS meteorological stations are at 3000 m, near the glacier tongue on the Gongga Mountain, which has been operated in 1988. The data collected by this station include precipitation, air temperature, wind speed, humidity, and solar station. Among them, data onto 1998–2002 are missing due to the equipment replacement (we fitted monthly data for missing years). The original time step of data acquisition precipitation and air temperature is hourly. But in the follow-up study, we use rainfall and air temperature variable for research. The precipitation here is almost made up by rainfall since we found nearly no snowfall at 3000 m elevation after June.

There is also a runoff station in Huangbengliu (HBLG), and the runoff data were available from 1990 to 2008. However, many runoff data are missing due to the impact on debris flow; we only use runoff data for auxiliary analysis.

3 Data and methodology

3.1 Debris flow history

Guanjingtaixi (GJTX) and Huangbengliu (HBLG) in the Hailuogou periglacial area were chosen for the case study. The GJTX has a small glacier, and the HBLG has a part of firn (Fig. 1). During 1988–2019, debris flows were recorded in these areas on 14 days, and six of them were recorded in the two channels simultaneously. The first documented debris flow occurred in GJTX on July 8, 1989. On the afternoon of July 26 of the same year, debris flows occurred in HBLG and GJTX at the same time, with a total accumulation of more than 500,000 m$^3$. On July 12, 1990, heavy rainfall caused a debris flow in HBLG. In the early morning of July 28, 1995, debris flow occurred in both HBLG and GJTX, and it went on for a long time. On August 11 of the same year, a new debris flow developed in HBLG due to moraine debris in the basin’s upper reaches collapsed, feeding debris flow along the channel. On July 27, 1996, HBLG recorded a high-speed debris flow during a nearly full day of clear weather. Heavy rains on July 1 to 3, 1997, led to widespread flash floods, and on July 4, debris flows occurred in the Hailuogou Basin, including in HBLG and GJTX. On August 1 and 15 of the same year, debris flows occurred in both gullies. Among them, the rainfall recorded on August 15 is the highest 91 mm in the history of the station. Intense rainfall on July 26, 2003, triggered debris flow in HBLG, which caused severe damage to the cableway in the scenic spot (Ruren et al. 2001; Ni et al. 2010). Due to continuous rainfall, large-scale debris flows occurred in Gongga Mountain on August 11, 2005, including HBLG. In 2010, the temperature in the study area reached the highest during 1988–2019, debris flow, respectively, triggered on July 17 and July 31 in 2010 at HBLG. The latest debris flow of HBLG occurred on July 29, 2019, with the total amount of moraine soil exceeding 20,000 cubic meters.
Debris flows occurred at the same time in adjacent channel are usually triggered by the same rainfall event or an extreme temperature event in the area. Therefore, it is better to count several debris flows that occurred simultaneously as a single event while historically accounting for debris flow events. For example, debris flows occurred in both channels on July 26, 1989. If we calculate this for each channel separately, it will be counted as two events, but we only recorded one debris flow event because they were triggered by the same set of climatic conditions, ensuring that the number of debris flow events does not exceed the number of extreme rainfall events. Therefore, the debris flow events mentioned in this work do not refer to the event in a particular gully but to the event days (14 days) with recorded debris flow events.

For the purposes of this study, we adopted the following methods: (1) classifying debris flows according to the influence degree of meteorological events with different exceedance probabilities; (2) analysis of the sensitivity of debris flows to different climatic conditions by using the one-dimensional Bayesian method; and (3) analysis of debris flow susceptibility caused by the effect of multiple climatic conditions through the two-dimensional Bayesian method and proposing an early warning threshold.

### 3.2 Methods

#### 3.2.1 Debris flow trigger analysis

Due to all the debris flow days were only recorded in summer (June–August), we restricted the definition of probabilities of the years which recorded debris flows, i.e., from June 1 to August 31, 1988–2019. Therefore, all the probabilities after that in this paper are depending on this duration.

As debris flows often occur on days of extreme meteorology, we categorize the starting pattern of the periglacial debris flows based on an extreme meteorological event on
To make out potentially trigger mechanism for debris flows initiation, we discussed the hydrometeorological variables at different time scales, while debris flows occurred. There were some observed variables, like daily rainfall \( R \) (mm d\(^{-1}\)), daily mean air temperature \( T \) (°C), daily runoff \( Q \) (mm\(^3\) d\(^{-1}\)), as well as calculated variables as antecedent rainfall (\( AR_3, AR_7, AR_{15}, AR_{30} \) (d\(^{-1}\))), and mean temperature (\( MT_3, MT_7, MT_{15}, MT_{30} \) (°C)). The antecedent rainfall here is considered to be a function of degressive index and time, which can be calculated by the following equation (Bruce and Clark 1966):

\[
R_a = \sum_{i=1}^{n} R_i(K)^i
\]

where \( R_a \) = the effective antecedent rainfall fell during \( i \) days before debris flow occurred (mm), \( K \) = the degressive index which is 0.84 adopted widely in the Southwest China (Cong et al. 2006; Ni and Song 2020; Xu and Liu 2019) at present, \( R_i \) = the total rainfall during \( i \) days before debris flow occurred (mm).

To evaluate the influence of different variables when debris flow triggers, we compared the magnitude of each variable with its marginal distribution, and assign an “exceedance probability” to each value. The exceedance probabilities \( P_e \) refer to the probability of meteorological events of a certain magnitude, whether debris flow occurs or not, within the time range studied in this paper. It can be simply calculated by dividing the number of meteorological events exceeding a certain magnitude by the total number of meteorological events (i.e., \( P_e = N_A/N \), \( N \) is the total number of rainfall events recorded; \( N_A \) is the total number of rainfall events over \( i \) mm d\(^{-1}\) occurred).

To promote a more intuitive comparison of the variables, we defined the classes of exceedance probabilities \( P_e \) for the independent variables as follows: \( 1 \geq P_e > 0.5 \) equaled to high, \( 0.5 \geq P_e > 0.1 \) was moderate, \( 0.1 \geq P_e > 0.01 \) was low, and \( P_e \leq 0.01 \) as very low corresponds to extreme events of rainfall and temperature. These exceedance probabilities of the observed and calculated data were used to efficient analyses whether there are different dominant trigger mechanisms of debris flow. The \( P_e \leq 0.01 \) in this paper equaled to \( R \geq 43.2 \) mm d\(^{-1}\) and \( T \geq 16.4 \) °C, respectively.

Using the \( P_e \) of the recorded variables daily rainfall \( R \) and daily air temperature \( T \), together with eight calculated variables at the event days, then relative assessments of which variable played the greatest role in triggering a debris flow and how the relative impact of those variables changes as time goes on depend on the meteorological conditions at that time. While a certain variable reaches the value with \( 1 \geq P_e > 0.5 \), it can be considered that this variable is hard to trigger debris flow. By comparison, it was accordingly to have an extremely high relevance for triggering debris flow on days with extremely low exceedance probabilities (\( P_e \leq 0.01 \)) (Mostbauer et al. 2018).

### 3.2.2 One-dimensional analysis

Conditional probability refers to the probability of occurrence of event \( A \) (hereinafter referred to as debris flows) under event \( B \) (a certain scale of rainfall, temperature, or other variables) (Casella and Berger 2002; Stoffel et al. 2005; Walder and Driedger 1995). Conditional probability can be expressed as \( P(A|B) \), which can be understood as “the probability of occurrence of debris flows (A) given rainfall (B)” (Bayes 1763). Conditional probability can be directly applied to the Bayes’ theorem (Bayes 1763) expressed by the following formula:
where \( P(B|A) \) = the probability of occurrence of \( B \) after a given \( A \) (also called likelihood), that is, the probability of observing a \( B \)-level rainfall event when a debris flow occurs; \( P(A) \) = the prior probability of \( A \), that is, the probability that a debris flow can occur regardless of whether a \( B \)-level rainfall event occurs; \( P(B) \) = the marginal probability of \( B \), that is, the probability of observing a \( B \)-level rainfall regardless of whether debris flows occur; \( P(A|B) \) = the probability of occurring \( A \) after given \( B \) (also known as the posterior probability), i.e., the probability of debris flows observed when a \( B \)-level rainfall event occurred.

One-dimensional Bayesian analysis can be effectively used to assess the importance of variable \( B \) in explaining an event (Berti et al. 2012). The analysis was performed using the data set presented in Sect. 3.1. The marginal rainfall probability \( P(B) \) was calculated from the selected 2300 rainfall and temperature events, and \( P(A|B) \) was calculated using the rainfall conditional probability of the 14 debris flow events. The prior probability of debris flows was \( P(A) = 14/2300 = 0.006 \). Ten explanatory variables were tested: event rainfall (\( R \)), antecedent rainfall in the previous \( n \) days (\( AR_{3}, AR_{7}, AR_{14}, AR_{30} \)), and event mean temperature (\( T \)), \( n \) days mean temperature (including \( MT_{3}, MT_{7}, MT_{14}, MT_{30} \)) (the last day of the total and the average is considered the day when the debris flows occurred).

### 3.2.3 Two-dimensional analysis

The formula of two-dimensional Bayes’ theorem can be presented by:

\[
P(A|B, C) = \frac{P(B, C|A) \cdot P(A)}{P(B, C)}
\]

where the symbols \( B \) and \( C \) indicate specific values with two variables. For instance, if \( B \) is daily rainfall (\( R \)), \( C \) is the mean daily temperature (\( T \)). Equation (3) represents the probability of debris flows occurring at a given rainfall and temperature.

Any pair of variables (such as daily rainfall and daily average temperature) can be considered in the two-dimensional Bayesian analysis, and the posterior debris flows probability \( P(A|B) \) is compared with the prior debris flows probability \( P(A) \) to assess their importance. Although the Bayesian method can be used to deal with the multi-dimensional analysis of the influence of \( n \) variables like rainfall intensity, rainfall duration, and antecedent rainfall on the initiation of debris flows, it has limitations of data scarcity. Moreover, multi-dimensional data visualization itself has several limitations. Therefore, this research is limited to the two-dimensional situation.

The two-dimensional Bayesian method considers the influence of rainfall and temperature work together on triggering debris flows and expresses the probability of its occurrence under the conditional probability of all the considered conditions. It is difficult to draw a clear dividing line between rainfall and temperature as the cause of debris flows by using this method. However, it is possible to identify a reasonable threshold based on the probability of the occurrence of debris flows. Moreover, such methods can be dynamically updated with the addition of new data (Berti et al. 2012). For actual forecasting and early warning, the cost of missing an event and raising a false alarm must be considered simultaneously.
4 Results

4.1 Major triggers of debris flow

The values of hydrometeorological variables at the event days were chosen from the observed and calculated time series. There were 6 days (DF 1, 3, 5, 7, 8, 9) with debris flows observed rainfall exceeded $R = 43.2$ mm d$^{-1}$, equivalent to a rainfall event with $P_e < 0.01$ over the research duration. Besides, rainfall recorded of 4 events (DF 2, 4, 10, 14) reached moderate exceedance probability $0.01 < P_e < 0.1$ (~23.4 mm d$^{-1}$). And 3 event days (DF 6, 11, 12) rainfall recorded with $0.1 < P_e < 0.5$, while only 1 day (DF13) nearly, no rainfall was recorded. For all calculated antecedent rainfall, most variables of antecedent rainfall were $0.1 < P_e < 0.5$ and $P_e > 0.5$, and only the exceedance probability of DF7’s antecedent rainfalls reached $P_e < 0.01$.

About variables of daily mean air temperature, there were only 2 event days (DF4, DF13) with a high probability of $0.01 < P_e < 0.1$ (14.6–16.4 °C). Other exceedance probabilities of daily mean air temperature $T$ not reached 0.01, especially DF5 occurred with $P_e > 0.5$. However, it can be found that all the exceedance probabilities of $MT3$ were $P_e < 0.5$; among them, 3 events (DF 2, 6, 14) reached $0.01 < P_e < 0.1$, and DF13 reached $P_e < 0.01$ from the calculated temperature variables. Other calculated variables of temperature like $MT7$, $MT14$, and $MT30$ corresponded to 3, 1, 2, respectively, times high exceedance probability events with $0.01 < P_e < 0.1$.

The above exceedance probabilities of variables on the event days allowed us to identify the different relevance to different variables of rainfall and temperature for triggering the recorded events on the 14 event days and to classify the initiation class base on the variable which plays the most significant role in triggering debris flows (Table 1). The highest correlation between the observed and calculated variable of rainfall and temperature was taken as the main contributor, and the contribution degrees were expressed according to the different probability grades in the variable. Among them, $P_e < 0.01$ represents extremely high correlation (+ + +), $0.01 < P_e < 0.1$ represents high correlation (+ +), $0.1 < P_e < 0.5$ represents medium correlation (+), and $0.5 < P_e < 1$ represents a low correlation.

By comparing the degree of contribution to different variables, it is found that the debris flow events dominated by rainfall accounted for 7 of the 14 event days (50%), while the temperature as the leading cause accounted for 2 event days (14%), 5 event days (36%) caused by the same contribution degree of rainfall and temperature. On the 6 of 7 event days, rainfalls caused most with daily rainfall $R > 43.2$ mm d$^{-1}$ observed probably to have an extremely high relevance to triggering the debris flows. Most of the exceedance probabilities of the values of $T$ for these events reached $0.1 < P_e < 0.5$, indicating that there are some moderately relevant additional contributions to runoff and increase the initiate water of debris flows. For example, the rainfall on July 12, 1990, reached extremely 64.5 mm with falling temperature for a few days, a dramatic runoff was observed, and debris flow occurred (Fig. 2a). Another one of 7 event days caused by high rainfall $R = 41.8$ mm d$^{-1}$ $(0.01 < P_e < 0.1)$ and moderate temperatures $(0.1 < P_e < 0.5)$. High-intensity rainfall played the dominant role in those 7 events, while temperature played a minor role.

It is found that DF6 and DF13 occurred with little antecedent, which exceedance probabilities were $0.1 < P_e < 0.5$; however, the triggered water is also common rainfall, DF6 occurred with recorded daily rainfall $R = 8.7$ mm d$^{-1}$ and DF13 with $R = 0.2$ mm d$^{-1}$. In particular, the daily rainfall of DF13 nearly no measured rainfall but debris flow also triggered (Fig. 2b). According to the recorded temperature variables of them, we found that
Table 1  Fourteen recorded debris flow events in Hailuogou since 1988

| Event no | Date     | Rainfall/mm d<sup>-1</sup> | Temperature/°C | Contribution degree |
|----------|----------|----------------------------|----------------|---------------------|
|          |          | rain_d | rain_3d | rain_7d | rain_14d | rain_30d | temp_d | temp_3d | temp_7d | temp_14d | temp_30d | R     | T     |
| 1        | 1989/7/8 | 57.6   | 7.8     | 15.2    | 35.6     | 38.3     | 13.1   | 13.5   | 10.9    | 9.6      | 10.3     | +    | +    |
| 3        | 1990/7/12| 64.5   | 3.6     | 5.4     | 17.5     | 20.5     | 13.2   | 14.1   | 14.1    | 13.6     | 12.3     | +    | ++   |
| 5        | 1995/8/11| 44.7   | 19.0    | 41.1    | 56.6     | 59.8     | 11.8   | 13.9   | 13.5    | 13.6     | 13.6     | ++   | ++   |
| 7        | 1997/7/4 | 43.3   | 81.6    | 97.8    | 103.8    | 110.4    | 13.5   | 13.7   | 12.6    | 11.8     | 10.2     | ++   | +    |
| 8        | 1997/8/1 | 48.0   | 19.6    | 37.9    | 44.3     | 48.6     | 12.8   | 12.9   | 12.8    | 12.4     | 11.9     | ++   | +    |
| 9        | 1997/8/15| 91.0   | 4.1     | 15.3    | 27.3     | 31.3     | 12.4   | 13.4   | 12.5    | 13.6     | 12.9     | ++   | +    |
| 10       | 2003/7/26| 41.8   | 25.0    | 37.6    | 42.0     | 47.5     | 13.5   | 14.0   | 12.3    | 13.5     | 12.6     | +    | +    |
| 6        | 1996/7/17| 8.7    | 17.6    | 34.1    | 47.3     | 52.9     | 14.5   | 14.7   | 12.8    | 12.2     | 11.8     | +    | ++   |
| 13       | 2010/7/31| 0.2    | 23.4    | 48.9    | 59.8     | 63.8     | 16.3   | 16.1   | 14.7    | 13.7     | 13.7     | +    | +++  |
| 2        | 1989/7/26| 42.8   | 21.5    | 36.8    | 54.4     | 60.4     | 12.4   | 14.4   | 15.0    | 14.0     | 11.9     | ++   | ++   |
| 4        | 1995/7/28| 31.6   | 0.9     | 20.4    | 36.3     | 39.2     | 15.2   | 14.2   | 13.8    | 13.6     | 12.3     | ++   | ++   |
| 11       | 2005/8/11| 17.4   | 28.3    | 50.7    | 63.6     | 66.0     | 12.8   | 12.3   | 11.9    | 12.9     | 12.8     | +    | +    |
| 12       | 2010/7/17| 15.6   | 13.2    | 31.3    | 41.2     | 49.3     | 12.7   | 14.0   | 12.9    | 13.5     | 12.6     | +    | +    |
| 14       | 2019/7/29| 25.8   | 5.9     | 18.3    | 31.1     | 36.1     | 13.5   | 14.5   | 13.4    | 13.5     | 12.0     | +    | ++   |

For each individual variable for a given event, bold and italic values indicate an extremely low occurrence probability ($P_e \leq 0.01$), bold values as a low occurrence probability ($0.01 < P_e \leq 0.1$); italic values as moderate occurrence probability ($0.1 < P_e \leq 0.5$) and normal values as a high occurrence probability ($0.5 < P_e \leq 1$).
the 3 days mean temperature $MT_3$ of DF6 was 14.7 °C ($0.01 < P_e \leq 0.1$) and of DF13 was 16.1 °C ($P_e \leq 0.01$). Their recorded daily temperature was a medium and high correlation, respectively. Therefore, those two events were triggered by the meltwater caused by rising high temperatures, and the rainfall played only a minor role.

Fig. 2 Examples of the observed hydrometeorological variables for the 3000 m station, including station rainfall, air temperature, and observed runoff. a the debris flow event on July 12, 1990, which may be interpreted as high-intensity rainfall; b on July 31, 2010, which suggests being intense snowmelt induced by high temperature; c–d events may be triggered by the interaction of long-lasting (or short-medium) rainfall and rising temperature.
Besides, there were 5 events triggered by the same degree of rainfall and temperature, including DF2, 4, 11, 12, 14. Among them, DF11 and 12 occurred with medium rainfall and temperature. From Fig. 2c, we found that this event was triggered by the combination of long-lasting rainfall and rising temperature; variables with moderate exceedance probability (0.1 < \( P_e \leq 0.5 \)) can also cause debris flow. DF2, 4, 14 were triggered by the combination of high rainfall and temperature with low exceedance probability (0.01 < \( P_e \leq 0.1 \)). The antecedent rainfall of DF4 showed a low correlation of debris flow, but the daily rainfall and temperature, which reached 0.01 < \( P_e \leq 0.1 \), caused debris flow (Fig. 2d).

It can be concluded that it is more difficult to trigger periglacial debris flows by the influence of high temperature than rainfall in this area, and debris flows triggered by the combination of rainfall and temperature tend to have a greater impact and faster speed. At the same time, we plot the distribution of variables to compare the difference between three triggering modes (Fig. 3). It is found that there are obvious distribution differences in daily rainfall (\( R \)), daily temperature (\( T \)), and 3 days mean temperature (\( MT3 \)) of 3 different triggering modes.

### 4.2 One-dimensional Bayesian probability

Our study has shown that periglacial debris flows are more sensitive to rainfall than temperature.

The results of the one-dimensional Bayesian analysis are shown in Figs. 4, 5. The left side of the two figures (a, c, e, g, and i) compares the initial rainfall/temperature of the debris flows and the overall frequency distribution, namely \( P(B|A) \) and \( P(B) \), and the right side of the figures (b, d, f, h, and j) shows the probability of debris flow occurrence \( P(A|B) \). The difference between \( P(B|A) \) and \( P(B) \) is large, and \( P(A|B) \) is higher, indicating that the significance of the variable under consideration is higher. The two figures show that the probability of debris flows increased with the event’s severity, and the rainfall variable is more significant than temperature. The triggers for debris flows can be ranked as follows: \( R > AR3 > AR14 > AR7 > AR30 > MT3 > MT30 > T > MT7 > MT14 \). Event daily rainfall \( R \) is
the most significant explanatory variable among these ten variables capable of triggering a debris flow event as a single variable. When $R > 30 \text{ mm d}^{-1}$, $P(A|B)$ gradually increases. When $R > 40 \text{ mm d}^{-1}$, $P(A|B)$ reaches 0.2, and when $R > 60 \text{ mm d}^{-1}$, $P(A|B)$ is as high as 0.33. Here, the $P(A|B)$ reaches 100% when $R > 90 \text{ mm d}^{-1}$, because there was only 1 rainfall event (DF7’s daily rainfall) recorded more than 90 mm d$^{-1}$ and no other rainfall events recorded between 70 to 90 mm d$^{-1}$. Among the temperature variables, $MT3$ is more significant in triggering debris flows confirming the finding of Paranunzio et al. (2015). When $MT3 > 16 ^{\circ}C$, the $P(A|B)$ value reaches the highest value of 0.06.

4.3 Two-dimensional Bayesian probability

The rainfall–3 days mean temperature threshold of periglacial debris flows—is dynamic, and $P(A|R, MT3) \approx 0.1$ can be selected as the early warning threshold of debris flows.

Two-dimensional Bayesian analyses refers to the conditional probability analyses of debris flow events given by two control variables. The variables should be selected from the variables with the highest explanatory in the one-dimensional Bayesian analysis. In this paper, it should be the event rainfall and antecedent rainfall in the previous 3 days, but these two variables have a high degree of coincidence, and therefore, we chose the parameters of event rainfall and 3 days mean temperature for analysis.
We connected the points with equal posterior probability $P(A|R, MT3)$ of debris flows under different parameters into contour lines and plotted them in Fig. 6. It can be seen that when the 3 days mean temperature is lower than $11 \, ^{\circ}C$ and no rainfall event occurs, the probability of debris flows is 0. The probability of debris flows is also 0 when the rainfall is less than $8.7 \, \text{mm d}^{-1}$, and the 3 days mean temperature is $8 \, ^{\circ}C$. The debris-free area shown in Fig. 6 includes all the points where $P(A|R, MT3) = 0$ because there is no rainfall in this area in the recorded meteorological conditions of debris flow outbreaks when $MT3 > 8 \, ^{\circ}C$ and rainfall $R > 50 \, \text{mm d}^{-1}$, and the probability of a debris flow outbreak reaches the highest value of 0.4.

Figure 6 provides a series of rainfall—3 days mean temperature probability values. The key task is selecting a conditional probability value (if any) as the debris flow warning threshold. In our observation, we can define $P(A|R, MT3) \approx 0.1$ as the threshold for debris flows warning because at least a single variable has reached the critical value to cause extreme events on this line (rainfall). For example, when $MT3 \approx 8$, $R$ is already higher than $23.4 \, \text{mm d}^{-1}$ of low exceedance probability $0.01 < P_e < 0.1$ on this line.
5 Discussion and conclusions

By analyzing the relationship between eight debris flows activities in the Hailuogou periglacial region and associated meteorological conditions, the following conclusions can be drawn from this study:

(1) In the periglacial areas with few glaciers and snow-covered watersheds, debris flow in the area is affected by rain and temperature. Of the 14 debris flow event days, seven are attributable to high-intensity rainfall, two are mainly triggered by rising high temperature, and five are affected by high rainfall and temperature.

(2) Periglacial debris flow is more sensitive to rainfall. The overall probability of the occurrence of a debris flow event depends by the parameters in the order of daily rainfall > antecedent rainfall 3 days > antecedent rainfall 14 days > antecedent rainfall 7 days > antecedent rainfall 30 days > 3 days mean temperature > 30 days mean temperature > daily temperature > 7 days mean temperature > 14 days mean temperature. If we consider only one parameter as the trigger, the probability of occurrence of debris flows is determined by these conditions: when the daily rainfall $R > 40 \text{ mm d}^{-1}$, the probability of a debris flow outbreak in this area is 20%, when $R > 60 \text{ mm d}^{-1}$, the probability reaches 33.3%; when the $MT3 > 16 \text{ °C}$, the probability of a debris flow outbreak is 6%.

(3) Based on the analysis of the data set, a contour map of the conditional probability of the periglacial debris flows caused by the rainfall—3 days mean temperature—was plotted,
and the conditional probability when \( P(A|E,M_{T3}) \approx 0.1 \) was defined as the threshold value, which can be used as the reference value for debris flows early warning in this region. For the high-intensity rainfall triggered debris flows, which are caused by heavy rainfall on the day of its occurrence, it is easier to discover and provide early warning, while in the cases of temperature triggered debris flows, a lower daily rainfall can also induce debris flows, and thus, early warning is difficult. Therefore, one should be alerted to identify the outbreak of debris flows caused by both rainfall and temperature to be able to provide early warning during actual monitoring.

Although some progress has been made in understanding the meteorological causes and prediction and early warning of debris flow outbreaks in the Mount Gongga periglacial area, some problems persist:

(1) Our study is limited to only 14 days of occurrences of historical debris flows, and, likely, some events have not been recorded. As a result, the numbers of recorded debris flow events are incomplete, leading to unavoidable errors in the probability estimations. Therefore, more debris flows data must be added in future research to optimize the warning threshold of periglacial debris flows in this area.

(2) The triggering of periglacial debris flows is often the result of combined effects of meteorology, hydrology, and geomorphology (Jomelli et al. 2007; Stoffel et al., 2011; Borga et al. 2014). Our research focuses on the relationship between the outbreak of debris flows and rainfall and temperature without considering the impact of geomorphological conditions on the susceptibility to debris flows. In future research, the influence of channel characteristics will be added to the analysis.

(3) Earthquakes can trigger debris flows, especially those accompanying extreme drought years, causing large-scale debris flow events (Chen et al. 2014). According to data available, an earthquake with a magnitude of 5.2 Mw occurred in Shimian county, approximately 50 km away from Hailuogou, on June 9, 1989, which may be one of the reasons for the two consecutive debris flows in July of that year. Therefore, further studies shall consider combining seismic activity with climate change.

(4) As natural systems are affected by land use, land to cover, rainfall patterns, and human activities, changes in these will also affect the frequency of debris flows.

Therefore, our research is only appropriate for predicting, and early warning of debris flows by using hydrometeorological conditions within the study area.

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Declarations

Conflict of interest The authors declare no conflict of interest.
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