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To cite this version:
Elisa Lahcene, Ioanna Ioannou, Anawat Suppasri, Kwanhrai Pakoksung, Ryan Paulik, et al. Characteristics of building fragility curves for seismic and non-seismic tsunamis: case studies of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean tsunamis. Natural Hazards and Earth System Sciences, Copernicus Publ. / European Geosciences Union, 2021, 21 (8), pp.2313-2344. 10.5194/nhess-21-2313-2021. hal-03400302

HAL Id: hal-03400302
https://hal.archives-ouvertes.fr/hal-03400302
Submitted on 31 Oct 2021

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Characteristics of building fragility curves for seismic and non-seismic tsunamis: case studies of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean tsunamis

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Received: 30 November 2020 – Discussion started: 16 December 2020
Revised: 15 June 2021 – Accepted: 28 June 2021 – Published: 6 August 2021

Abstract. Indonesia has experienced several tsunamis triggered by seismic and non-seismic (i.e., landslides) sources. These events damaged or destroyed coastal buildings and infrastructure and caused considerable loss of life. Based on the Global Earthquake Model (GEM) guidelines, this study assesses the empirical tsunami fragility to the buildings inventory of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khao Lak–Phuket, Thailand) tsunamis. Fragility curves represent the impact of tsunami characteristics on structural components and express the likelihood of a structure reaching or exceeding a damage state in response to a tsunami intensity measure. The Sunda Strait and Sulawesi–Palu tsunamis are uncommon events still poorly understood compared to the Indian Ocean tsunami (IOT), and their post-tsunami databases include only flow depth values. Using the TUNAMI two-layer model, we thus reproduce the flow depth, the flow velocity, and the hydrodynamic force of these two tsunamis for the first time. The flow depth is found to be the best descriptor of tsunami damage for both events. Accordingly, the building fragility curves for complete damage reveal that (i) in Khao Lak–Phuket, the buildings affected by the IOT sustained more damage than the Sunda Strait tsunami, characterized by shorter wave periods, and (ii) the buildings performed better in Khao Lak–Phuket than in Banda Aceh (Indonesia). Although the IOT affected both locations, ground motions were recorded in the city of Banda Aceh, and buildings could have been seismically damaged prior to the tsunami’s arrival, and (iii) the buildings of Palu City exposed to the Sulawesi–Palu tsunami were more susceptible to complete damage than the ones affected by the IOT, in Banda Aceh, between 0 and 2 m flow depth. Similar to the Banda Aceh case, the Sulawesi–Palu tsunami load may not be the only cause of structural destruction. The buildings’ susceptibility to tsunami damage in the waterfront of Palu City could have been enhanced by liquefaction events triggered by the 2018 Sulawesi earthquake.

1 Introduction

Indonesia regularly faces natural disasters such as earthquakes, volcanic eruptions, and tsunamis because of its geographic location in a subduction zone of three tectonic plates (Eurasian, Indo-Australian, and Pacific plates) (Marfai et al., 2008; Sutikno, 2016). The Sunda Arc extends for 6000 km from the north of Sumatra to Sumbawa Island (Lauterjung et al., 2010) (Fig. 1a). Megathrust earthquakes regularly occur in this region, causing horizontal and vertical movement of the ocean floor which tends to be tsunamigenic (McCloskey et al., 2008; Nalbant et al., 2005; Rastogi, 2007). These tsunamis are likely to cause greater destruction as they can follow prior damaging earthquake ground shaking and/or liquefaction (Sumer et al., 2007; Sutikno, 2016). Earthquake-
generated tsunamis also tend to have longer wave periods affecting the coast than non-seismic ones (Day, 2015; Grezio et al., 2017). On 26 December 2004, the Sumatra–Andaman earthquake \( (M_w = 9.0–9.3) \) hit the north of Sumatra, Indonesia (Fig. 1b). The rupture of the seafloor is estimated at 1200 km length and around 200 km width (Ammon et al., 2005; Krüger and Ohrnberger, 2005; Lay et al., 2005). In the city of Banda Aceh, a strong ground shaking was recorded (Lavigne et al., 2009). This megathrust earthquake was the second largest ever recorded (Løvholt et al., 2006) and caused the deadliest tsunami in the world. Overall, a dozen Asian and African countries were devastated, with around 280,000 casualties (Asian Disaster Preparedness Center, 2007; Suppasri et al., 2011). Although earthquakes represent the main cause of tsunamis, non-seismic events such as landslides can also initiate tsunami waves (Grezio et al., 2017; Ward, 2001). After a few months of volcanic activity in the Sunda Strait, Indonesia, the Anak Krakatau volcano erupted on 22 December 2018, leading to its southwestern flank failure (Fig. 1c). It triggered a relatively short wave period tsunami (\( \sim 7 \) min) (Muhari et al., 2019), which devastated the western coast of Banten and the southern coast of Lampung with a death toll of 437 (Heidarzadeh et al., 2020; Muhari et al., 2019; National Agency for Disaster Management (BNPB), 2018; Syamsidik et al., 2020). The tsunami generation process is unclear. The subaerial and submarine landslide volume is still being investigated and ranges between 0.10 and 0.30 km\(^3\) according to recent studies (Dogan et al., 2021; Grilli et al., 2019; Omira and Ramalho, 2020; Paris et al., 2020; Williams et al., 2019). Almost 2 months before this event, an unexpected tsunami struck Palu Bay, on Sulawesi Island, claiming 2000 lives and considerable loss to property (Association of Southeast Asian Nations (ASEAN)-Coordinating Centre for Humanitarian Assistance on disaster, 2018; Omira et al., 2019). The Sulawesi earthquake \( (M_w = 7.5) \) occurred near the Palu-Koro strike-slip fault, 50 km northwest of Palu Bay (Fig. 1d) (Socquet et al., 2019). Ground shaking led to significant liquefaction along the coast.
The term “tsunami fragility” is a measure recently proposed to estimate structural damage and casualties caused by a tsunami, as mentioned by Koshimura et al. (2009b). Tsunami fragility curves are functions expressing the damage probability of structures (or death ratio) based on the hydrodynamic characteristics of the tsunami inundation flow (Koshimura et al., 2009a, b). These functions have been widely developed after tsunami events such as the 2004 Indian Ocean tsunami (IOT; Koshimura et al., 2009a, b; Murao and Nakazato, 2010; Suppasri et al., 2011), the 2006 Java tsunami (Reese et al., 2007), the 2010 Chilean tsunami (Mas et al., 2012), or the 2011 great eastern Japan tsunami (Suppasri et al., 2012, 2013). Several methods aim to develop building fragility curves based on (i) a statistical analysis of on-site observations during field surveys of damage and flow depth data (empirical methods) (Suppasri et al., 2015, 2020), (ii) the interpretation of damage data from remote sensing coupled with tsunami inundation modelling (hybrid methods) (Koshimura et al., 2009a; Mas et al., 2020; Suppasri et al., 2011), or (iii) structural modelling and response simulations (analytical methods) (Attary et al., 2017; Macabuag et al., 2014).

Here, we empirically developed building fragility curves for the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khaok Lak–Phuket, Thailand) tsunamis based on the Global Earthquake Model (GEM) guidelines (Rocsetto et al., 2014). From the field surveys conducted after the 2018 Sunda Strait (Syamsidik et al., 2019), 2018 Sulawesi–Palu (Paulik et al., 2019), and 2004 Indian Ocean (Khaok Lak–Phuket, Thailand) (Foytong and Ruangrassamee, 2007; Ruangrassamee et al., 2006) events, we utilize three databases called DB_Sunda2018, DB_Palu2018, and DB_Thailand2004, respectively. In the literature, tsunami inundation modelling has been performed many times to better understand the tsunami hydrodynamics, especially for earthquake-generated tsunamis (Charvet et al., 2014; Gokon et al., 2011; Koshimura et al., 2009a; Macabuag et al., 2016; De Risi et al., 2017; Suppasri et al., 2011). Compared to the 2004 IOT, the 2018 Indonesian tsunamis are uncommon events remaining less understood. Therefore, to improve our understanding of the structural damage caused by the Sunda Strait and Sulawesi–Palu tsunamis and to discuss the impact of wave period, ground shaking, and liquefaction events, we reproduce their tsunami intensity measures (i.e., flow depth, flow velocity, and hydrodynamic force) based on two-layer modelling (TUNAMI two-layer). We then compared the fragility curves of the Sunda Strait, Sulawesi–Palu, and Indian Ocean (Khaok Lak–Phuket) tsunamis to those derived for the 2004 IOT in Banda Aceh (Indonesia), produced by Koshimura et al. (2009a). In this study, we explore the characteristics of building fragility curves for the 2018 Sunda Strait event and 2004 IOT in Khaok Lak–Phuket, as well as for complex events, such as the 2018 Sulawesi–Palu tsunami in Palu City and the 2004 IOT in Banda Aceh, where the tsunamis may not be the only cause of structural destruction. Studying the impact of the wave period, ground shaking, and liquefaction events on the structural performance of buildings aids to improve our knowledge on the relationship between local vulnerability and tsunami hazard in Indonesia.

2 Post-tsunami databases

A post-tsunami database has been established for the Sunda Strait area by Syamsidik et al. (2019), Palu Bay by Paulik et al. (2019), and Khaok Lak–Phuket by Ruangrassamee et al. (2006) and Foytong and Ruangrassamee (2007) in urban areas strongly affected by these events. These databases include 98, 371, and 120 observed flow depth traces at buildings, respectively. Here, the tsunami fragility analysis stands on subsets of the original databases of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khaok Lak–Phuket) tsunamis, as explained in Sects. 3.2.2, 3.2.3, and 2.2, respectively. We define these subsets as “new” databases, and we call them DB_Sunda2018, DB_Palu2008, and DB_Thailand2004, respectively. We note that the use of smaller databases for the fragility assessment is expected to increase the uncertainty in the exact shape of the fragility curves. Each database gathers exclusive information regarding the degree of damage, the building characteristics, and the flow depth traces (Tables 1 and 2). A brief analysis of the key variables (i.e., damage scale, building class, and tsunami intensity) are presented below.

2.1 Damage state

Each field survey adopted a different scale to record the degree of structural damage. In DB_Sunda2018, the five-state damage scale proposed by Macabuag et al. (2016) and Suppasri et al. (2020) is adopted, ranging from no damage to complete damage or washed away. In DB_Palu2018, the observed damage was classified into four states: no damage, partial damage repairable, partial damage unrepairable, and complete damage, as proposed by Paulik et al. (2019). Finally, in DB_Thailand2004, a four-state damage scale is...
defined by Ruangrassamee et al. (2006). To simplify the comparison between the fragility curves, a harmonization of damage scales is proposed (Table 1). In this study, a four-state damage scale ranging from $ds_0$ to $ds_3$ is used.

### 2.2 Building characteristics

Each survey also recorded the building construction type, which influences the damage probability (Suppasri et al., 2013). In Table 2, among the 94 buildings included in DB_Sunda2018, 67 are confined masonry, 26 are timber, and 1 is a steel frame building. In DB_Palu2018, most of the buildings are confined masonry with unreinforced clay bricks ($\sim 95\%$). The database also includes reinforced concrete and timber buildings. Finally, DB_Thailand2004 contains only reinforced concrete buildings. We note that after the 2004 IOT, 120 flow depth traces were recorded at reinforced concrete structures (e.g., residence, hotel, school, shop, bridge) in the Khao Lak–Phuket area. As we are not considering the data regarding the surveyed bridges, DB_Thailand2004 includes only 117 reinforced concrete buildings.

### 2.3 Tsunami intensity

The tsunami intensity has been measured in terms of flow depth level. Table 2 also presents the number of flow depth traces at surveyed buildings and the range of flow depth levels for each database.

### 3 Tsunami intensity simulations

#### 3.1 Tsunami numerical modelling with a landslide source

**3.1.1 Tsunami inundation model**

The TUNAMI two-layer tsunami model used in the Sunda Strait and Palu areas relies on a two-layer numerical model solving non-linear shallow water equations. It considers two interfacing layers and appropriate kinematic and dynamic boundary conditions at the seafloor, interface, and water surface (Imamura and Imteaz, 1995; Pakoksung et al., 2019). To reproduce the landslide-generated tsunami, we model the interactions between tsunami generation and submarine landslides as upper and lower layers. The mathematical model performed in the landslide-tsunami code is obtained from a stratified medium with two layers. The first layer, composed of a homogeneous inviscid fluid with constant density, $\rho_1$, represents the seawater, and the second layer is composed of a fluidized granular material with a density, $\rho_s$, and porosity, $\varphi$. As assumed by Macías et al. (2015), the mean density of the fluidized sliding mass is constant and equals $\rho_2 = (1 - \varphi)\rho_s + \varphi\rho_1$. We consider the two layers immiscible. The governing equations are written as follows.

The continuity equation of the seawater (first layer) is

$$\frac{\partial Z_1}{\partial t} + \frac{\partial Q_{1x}}{\partial x} + \frac{\partial Q_{1y}}{\partial y} = 0.$$  \hspace{1cm} (1)

The momentum equations of the seawater in the $x$ and $y$ directions are

$$\frac{\partial Q_{1x}}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q_{1x}^2}{D_1} \right) + \frac{\partial}{\partial y} \left( \frac{Q_{1x}Q_{1y}}{D_1} \right) + gD_1 \frac{\partial Z_1}{\partial x} + gD_1 \frac{\partial Z_2}{\partial x} + \tau_{1x} = 0,$$  \hspace{1cm} (2)

$$\frac{\partial Q_{1y}}{\partial t} + \frac{\partial}{\partial x} \left( \frac{Q_{1x}Q_{1y}}{D_1} \right) + \frac{\partial}{\partial y} \left( \frac{Q_{1y}^2}{D_1} \right) + gD_1 \frac{\partial Z_1}{\partial y} + gD_1 \frac{\partial Z_2}{\partial y} + \tau_{1y} = 0.$$  \hspace{1cm} (3)

The continuity equation of the landslide (second layer) is

$$\frac{\partial Z_2}{\partial t} + \frac{\partial Q_{2x}}{\partial x} + \frac{\partial Q_{2y}}{\partial y} = 0.$$  \hspace{1cm} (4)

The momentum equations of the landslide in the $x$ and $y$ directions are
Table 2. Observed flow depth traces at buildings, range of flow depth levels, and building characteristics in DB_Sunda2018, DB_Palu2018, and DB_Thailand2004.

| Building Characteristics | DB_Sunda2018 | DB_Palu2018 | DB_Thailand2004 |
|--------------------------|--------------|-------------|-----------------|
| Observed flow depth traces at buildings | 94 | 124 | 117 |
| Range of observed flow depth levels at buildings (m) | (0.20, 6.60) | (0.10, 3.65) | (0.15, 10.00) |
| Number of buildings per construction type | 67 confined masonry | 119 confined masonry | 117 reinforced concrete |

The tsunami inundation model is calibrated using two performance parameters: $K$ and $\kappa$ proposed by AIDA (1978), as defined below:

$$\log K = \frac{1}{n} \sum_{i=1}^{n} \log K_i,$$

where $K_i$ represents the Manning’s roughness coefficient through the term $D$, which represents the simulated flow depth (m). In the urban areas of Sunda Strait and Palu, the average occupation ratios are 24% and 84%, respectively (Kotani, 1998).

3.2 Calibration and validation of the tsunami inundation model

3.2.1 Performance parameters

The tsunami inundation model is validated using two performance parameters: $K$ and $\kappa$ proposed by AIDA (1978), as defined below:

$$\log K = \frac{1}{n} \sum_{i=1}^{n} \log K_i,$$
\begin{equation}
\log \kappa = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (\log K_i)^2 - (\log K)^2},
\end{equation}

where $x_i$ and $y_i$ are the recorded and simulated tsunami flow depths at location $i$. $K$ is defined as the geometrical mean of $K_i$, and $\kappa$ is defined as deviation or variance from $K$. The Japan Society of Civil Engineers (JSCE) (2002) recommends $0.95 < K < 1.05$ and $\kappa < 1.45$ for the model results to achieve “good agreement” in the tsunami source model and propagation and inundation model evaluation (Otake et al., 2020; Pakoksung et al., 2018).

### 3.2.2 The 2018 Sunda Strait tsunami inundation model

To correct the digital surface model (DSM), we removed the vegetation, building, and infrastructure elevations based on the linear smoothing method and used the resulting digital elevation model (1st DEM) as topography in the tsunami inundation model (Fig. 3). The vertical accuracy of the DSM and DEM is about 4 m. The 2018 Sunda Strait tsunami model depends on the density of the landslide ($\rho_2$), its stable slope ($\alpha$), its volume ($V_L$), and its sliding time ($t_S$). As proposed by Paris et al. (2020), the low sensitivity parameters are set as follows: $\rho_2 = 1500 \text{ kg m}^{-3}$, $\alpha = 5^\circ$, and $V_S = 0.15 \text{ km}^3$. We reach the best fit between the simulated and observed flow depths at buildings for 10 min sliding time. Nevertheless, most of the simulated flow depths are underestimated compared to the observed ones, with a mean difference of $0.28 \pm 1$ m. Using quantum GIS (QGIS) software, we smoothed the 1st DEM to remove these mean differences in elevation at buildings where the flow depth is underestimated. The resulting DEM (2nd DEM) provides a topography more reliable at buildings (Fig. 3). We completed three cross sections along the Sunda Strait coasts to show the different corrections applied to the DSM (Fig. 3). The $K$ and $\kappa$ values for damaged buildings are 0.99 and 1.11, respectively, which means that we achieve “good agreement” for the Sunda Strait tsunami model, displayed in Fig. 5a–f. We note that the simulated inundation zone overlays 94 buildings out of 98. In Sect. 4.1, the Sunda Strait tsunami fragility assessment is based on these 94 buildings (DB_Sunda2018). Simulation snapshots of the Sunda Strait tsunami propagation are shown in Fig. B1 10, 20, 60, and 120 s after the tsunami generation. In Fig. B2, the simulated tsunami height based on the best-fitting parameters is also displayed. Figure B3 illustrates the maximum simulated flow velocity of the 2018 Sunda Strait tsunami inundation model.
Figure 3. Topographic corrections performed on the DSM and the 1st DEM. The 2nd DEM is used as new topography in the TUNAMI two-layer model.

Figure 4. (a) Cross sections along Sunda Strait coasts. One cross section is realized in the computational areas (b and e) 1, (c and f) 2, and (d and g) 3 to illustrate the topographic corrections applied to the DSM at buildings using QGIS (a triangle represents a building) (background ESRI and © Google Maps).
Figure 5. (a, c, and e) Sunda Strait final tsunami inundation model with the maximum simulated flow depth overlaid on the damaged building data in the computational areas 1 to 3, and (b, d, and f) magnified views of the maximum simulated flow depth in the Rajabasa, Pejamben, and Sumur areas (background ESRI and © Google Maps).
Table 3. Hypothesized landslide parameters (location and volume) in Palu Bay.

| No. | Location (latitude; longitude) | Volume (10^6 m^3) |
|-----|--------------------------------|------------------|
| L1^a| −0.655; 119.749                | 37.54            |
| L2^a| −0.670; 119.801                | 31.93            |
| S1^b| −0.680; 119.821                | 0.60             |
| S2^b| −0.703; 119.842                | 0.18             |
| S3^b| −0.737; 119.851                | 0.25             |
| S4^b| −0.789; 119.862                | 0.75             |
| S5^b| −0.852; 119.878                | 0.22             |
| S6^b| −0.879; 119.871                | 0.60             |
| S7^b| −0.885; 119.858                | 2.44             |
| S8^b| −0.846; 119.822                | 4.45             |
| S9^b| −0.832; 119.813                | 0.83             |
| S10^b| −0.804; 119.808               | 2.17             |
| S11^b| −0.774; 119.792              | 0.55             |
| S12^b| −0.754; 119.788              | 0.83             |

a Based on our assumption from Arikawa et al. (2018) and Heidarzadeh et al. (2019). b Based on observations from satellite imagery, field surveys, and video footage (Arikawa et al., 2018; Carvajal et al., 2019).

3.2.3 The 2018 Sulawesi–Palu tsunami inundation model

We increased the mean sea level (MSL) by 2.3 m to reproduce the high tide during the 2018 Sulawesi–Palu tsunami. As shown by Pakoksung et al. (2019), the observed waveform at Pantoloan tidal gauge does not fit the simulated one with the finite fault model of TUNAMI-N2. Although recent studies show that seismic seafloor deformation may be the primary cause of the tsunami (Gusman et al., 2019; Ulrich et al., 2019), in this study, the main assumption is that the 2018 Sulawesi–Palu event was triggered by subaerial and submarine landslides. According to Heidarzadeh et al. (2019), a large landslide to the north or the south of Pantoloan tidal gauge is responsible for the significant height wave recorded. Arikawa et al. (2018) also identified several sites of potential subsidence in the northern part of Palu Bay. Based on these previous studies, we assume two large landslides: L1 and L2. Small landslides (S1–S12) also occurred in the bay; their location is known from observations from satellite imagery, field surveys, and video footage (Arikawa et al., 2018; Carvajal et al., 2019) (Fig. 6). From the trial and error method and the topographic and bathymetric data provided by the Geospatial Information Agency (BIG), we determined the soil property and achieved the volume of the landslides (Table 3). In Fig. 7, the submarine landslides model reproduces well the tsunami observations at Pantoloan.

The calibration of the model depends on the landslide S8 because (i) as a small landslide, its volume is too small to distort the simulated wave height at the Pantoloan tidal gauge, (ii) it has the largest volume among the other small landslides, and (iii) it is close and ideally oriented to Palu City; the slide direction, captured by an aircraft pilot, is perpendicular to the bay (Carvajal et al., 2019). The density of the landslides (ρ2), their stable slope (α), and their sliding time (t2) are set as follows: ρ2 = 2000 kg m^−3 (Palu Bay receives a large amount of fine continental deposits such as clay-sized sediments; Frederik et al., 2019), α = 14° (Chakrabarti, 2005), and t2 = 10 min. For a landslide ratio of 1.2 (i.e., S8 volume is multiplied by 1.2), the tsunami model shows a great similarity between observed and simulated flow depths (α = 1.027). The simulated tsunami inundation zone overlays 175 traces out of 371 because (i) 151 buildings with flow depth traces are not included in our computational area (Fig. 2c) and (ii) 45 buildings are outside the simulated tsunami envelope, which is shorter than the surveyed one (Fig. 8). The geometric mean is near the recommended values (K = 0.93), while the standard deviation and the root mean square error (RMSE) are high (κ = 2.18, RMSE = 0.92 m). Therefore, to develop accurate and reliable curves, we set a 1 m confidence interval including 124 flow depth traces at buildings out of 175 (Fig. 9). In Sect. 4.2, the Sulawesi–Palu tsunami fragility assessment is based on these 124 buildings (DB_Palu2018). K and κ values for damaged buildings are 0.93 and 2.14, respectively, with a root mean square error of 0.26 m. The validity of the model is mainly based on the geometric mean K, close to 0.95, so we consider the tsunami inundation model accurate enough (Fig. 8). In Fig. C1, the simulation snapshots of the Sulawesi–Palu tsunami propagation are shown 2, 10, 30, and 60 s after the tsunami generation. The simulated tsunami height based on the best-fitting parameters is also displayed in Fig. C2. Figure C3 illustrates the maximum simulated flow velocity of the 2018 Sulawesi–Palu tsunami inundation model.

4 Tsunami fragility assessment

The proposed fragility assessment framework has two main steps. In the first step, an exploratory analysis aims to (i) assess the trends that the available data follow and (ii) determine the main explanatory variables that need to be included in the statistical model and their influence on the slope and intercept of the fragility curves. Then, we select a statistical model and examine its goodness-of-fit to the data based on the observations of the exploratory analysis. We note that the development of the computed fragility curves for the 2018 Sunda Strait and 2018 Sulawesi–Palu tsunamis is directly based on DB_Sunda2018 and DB_Palu2018, in which each building has both observed and simulated flow depth values (Table 4).

To explore the relationship between the tsunami intensity and the probability of damage, we fit a generalized linear model (GLM) to the data of each database, as proposed by the GEM guidelines (Rossetto et al., 2014). A GLM assumes that the response variable y_ij is assigned 1 if the building

https://doi.org/10.5194/nhess-21-2313-2021
Table 4. Number of buildings used for the tsunami fragility analysis of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 IOT (Khao Lak–Phuket) events.

| Database            | Observed flow depth | Simulated flow depth | Simulated flow velocity | Simulated hydrodynamic force |
|---------------------|---------------------|----------------------|-------------------------|-----------------------------|
| DB_Sunda2018\(^a\)  | 94                  | 94                   | 94                      | 94                          |
| DB_Palu2018\(^b\)   | 124                 | 124                  | 124                     | 124                         |
| DB_Thailand2004     | 117                 | –                    | –                       | –                           |

\(^a\) Surveyed buildings included in the Sunda Strait simulated tsunami inundation zone. \(^b\) Surveyed buildings included in the Palu simulated tsunami inundation zone and in the 1 m confidence interval.

Figure 6. Location of the hypothesized landslides (S: small; L: large) in Palu Bay (background ESRI).

Figure 7. Comparison between observed and simulated wave heights at Pantoloan tidal gauge in Palu Bay, Sulawesi, Indonesia.

Based on the aforementioned observations, we construct parametric statistical models for the three databases to (i) identify the simulated tsunami measure type that fits the data best and (ii) construct fragility curves for the tsunami intensity type that fits the data best.

Ideally, the response variable \( y_{ij} \) of an appropriate statistical model is the damage state \( DS \geq ds_j \) and 0 otherwise. The variable follows a Bernoulli distribution:

\[
y_{ij} \sim \text{Bernoulli}(\pi_i(\tilde{x}_j)),
\]

where \( \pi_i(\tilde{x}_j) \) is the probability that a building \( j \) will reach or exceed the “true” damage state \( ds_j \) given estimated tsunami intensity level \( \tilde{x}_j \). The Bernoulli distribution is characterized by its mean,

\[
\mu_{ij} = \pi_i(\tilde{x}_j),
\]

which is expressed here in terms of a probit model, commonly used to express the mean in the empirical fragility assessment field (Rossetto et al., 2013), defined in terms of \( \Phi[.], \) the cumulative distribution function of a standard normal distribution:

\[
\Phi^{-1}[\pi_i(\tilde{x}_j)] = \eta_{ij},
\]

where \( \eta_{ij} \) is the linear predictor, which can be written in the form

\[
\eta_{ij} = \theta_0 + \theta_1 \ln(\tilde{x}_j),
\]

where \( \theta_1 \) and \( \theta_0 \) are the two regression coefficients representing the slope and the intercept, respectively, of the fragility curve corresponding to damage state \( ds_j \). For the exploratory analysis, the tsunami intensity is measured in terms of observed flow depth levels. We also fit the GLM models to subsets of data of each database to explore the importance of the construction type to the shape of the fragility curves. The confidence in the exact shape of the mean curves is estimated and presented in terms of the 90% confidence intervals around the best-estimate curves.
which describes the possible levels of damage \( i = \{0, 1, 2, 3\} \) sustained by a given building (Table 1). The random component of this model can be written as

\[
y_{ij} \sim \text{Categorical} \left( P(\text{DS} = d_{s_i} | \tilde{x}_j) \right), \tag{16}
\]

where \( P(\text{DS} = d_{s_i} | \tilde{x}_j) \) is the probability that a building \( j \) will reach the “true” damage state \( d_{s_i} \) given estimated tsunami intensity level \( \tilde{x}_j \).

\[
P(\text{DS} = d_{s_i} | \tilde{x}_j) = \begin{cases} 
1 - \pi_i (\tilde{x}_j), & i = 0 \\
\pi_i (\tilde{x}_j) - \pi_{i+1} (\tilde{x}_j), & 0 < i < i_{\text{max}} \\
\pi_{i_{\text{max}}} (\tilde{x}_j), & i = i_{\text{max}}
\end{cases} \tag{17}
\]

Multiple expressions of the systematic component are constructed to test their goodness of fit. With regard to the link function, apart from the commonly used probit function, two alternative expressions found in the GEM guidelines for empirical vulnerability assessment (Rossetto et al., 2013), namely the logit and complementary loglog (termed here “cloglog”), are considered in the form

\[
\eta_{ij} = \begin{cases} 
\Phi^{-1} \left[ \pi_i (\tilde{x}_j) \right], & \text{probit} \\
\ln \left( \frac{\pi_i (\tilde{x}_j)}{1 - \pi_i (\tilde{x}_j)} \right), & \text{logit} \\
\ln \left( -\ln \left( 1 - \pi_i (\tilde{x}_j) \right) \right), & \text{cloglog}
\end{cases} \tag{18}
\]

The linear predictor is also expressed in various forms of increasing complexity, as depicted in Eq. (19).

\[
\eta_{ij} = \begin{cases} 
\theta_0 + \theta_1 \tilde{x}_j & \text{(19a)} \\
\theta_0 + \theta_1 \tilde{x}_j + \theta_2 \text{ class} & \text{(19b)} \\
\theta_0 + \theta_1 \tilde{x}_j + \theta_2 \text{ class} & \text{(19c)} \\
\theta_0 + \theta_1 \tilde{x}_j + \theta_2 \text{ class} + \theta_3 \tilde{x}_j \text{ class} & \text{(19d)}
\end{cases}
\]

Class is a categorical unordered variable which expresses here the construction type. \( \theta_{0-3} \) are the unknown regression coefficients of the model. Equations (19a) and (19b) assume that the fragility curves are only influenced by the tsunami intensity. Equation (19a) assumes that the slope of the fragility curves is the same for all damage states. In contrast, Eq. (19b) allows the slope of each curve to vary for each damage state; the slope varies for each fragility curve. The following three equations account for the influence of the building class (i.e. the construction type) in the shape of the fragility curves. All three equations assume that the construction type affects the intercept of the fragility curves, and only Eq. (19e) assumes that the construction type affects both the intercept and the slope of the curves. Finally, Eqs. (19c) and (19e) assume identical slopes for all fragility curves irrespective of the damage state. In contrast, Eq. (19d) relaxes this assumption and considers that the slope changes for each damage state. The combinations of random and systematic components result in five distinct models (Table 5).

In what follows, we fit multiple models to each database based on the observations of the exploratory analysis. We examine the goodness of fit of these models for a given tsunami intensity measure and link function with two formal tests, as proposed in the GEM guidelines (Rossetto et al., 2014).
Table 5. Statistical models examined for each database.

| Model | Component   |
|-------|-------------|
|       | Random      | Systematic |
| M1    | Eq. (19a)   |             |
| M2    | Eq. (19b)   |             |
| M3    | Eq. (16)    | Eq. (19c)   |
| M4    | Eq. (19d)   |             |
| M5    | Eq. (19e)   |             |

Figure 10. Probit functions fitted for each individual damage state to DB_Sunda2018 (a) to assess whether the observed flow depth is an efficient descriptor of damage and (b) to assess whether the construction type affected the shape of fragility curves for $ds_2$ and $ds_3$. In both cases, the 90% confidence interval is plotted.

Firstly, we compare the Akaike information criterion (AIC) values, which estimates the prediction error of the examined models (Akaike, 1974). The model with the lowest value fits the data best. The alternative models used in this study are nested, which means that the more complex model includes all the terms of the simpler ones plus an additional term. For this reason, we also perform a series of likelihood ratio tests to examine whether the fit provided by the model with the lowest AIC value is statistically significant over its alternative nested models, which relaxes its assumptions (Rossetto et al., 2014). We also use the AIC value to determine which of these simulated intensity measures fits the data best. Furthermore, the 90% confidence intervals of the best-estimate fragility curves are constructed using bootstrap analysis. According to the latter analysis, 1000 samples of the database are obtained with a replacement, and the selected model is refitted to each sample.

4.1 DB_Sunda2018

We fit the GLM models to the data in DB_Sunda2018 (irrespective of their structural characteristics), and we plot the obtained probit functions against the natural logarithm of the observed flow depth to explore how the slope and the intercept of the models change for each damage state (Fig. 10a). The 90% confidence intervals around the best-estimate curves are also included. All three curves have positive slopes, which indicates that the flow depth is an adequate descriptor of the damage caused by a tsunami as the probability of a given damage state being reached or exceeded increases with the increase in the flow depth. The slope of each function is similar for $ds_2$ and $ds_3$ and different for $ds_1$. Nonetheless, the curve corresponding to $ds_1$ is also associated with substantial uncertainty. In Fig. 10b, we fit probit models to subsets of the available data for the two main construction types. One of the drawbacks of the small database is that not all damage states were observed for each building class. Therefore, the comparison of probit models is limited for damage states $ds_2$ and $ds_3$. The curves for the two construction types appear to be substantially different. As expected, the timber buildings are more vulnerable than the confined masonry buildings. Their intercept is responsible for the difference as the two curves are parallel. It indicates the need to develop a statistical model which allows only the intercept to change with the construction type, and the slope should be identical.

Following the main observations of the exploratory analysis, we consider that M3 is an acceptable model with two explanatory variables: the tsunami intensity and the construction type. To assess its goodness of fit, we consider each link function with three alternatives for the linear predictor (i.e., M4, M5, and M1) which relax some of its assumptions. In Table 6, we compare the AIC values of the three models to assess the fit of the different models for the observed flow depth levels assuming the probit link function. M3 has the smallest AIC value compared to its alternatives, which indicates that it fits the data better than the remaining three models. Nonetheless, some of these differences are rather small, and it raises the question of whether the improvement in the fit provided by M3 is statistically significant over its alternatives. To address this, we perform likelihood ratio tests, and the results are reported in Table 7. We note that the $p$ values vary for the three comparisons. The $p$ value is significantly above the 0.05 threshold when the identical slope for each fragility curve assumption (i.e. comparison of M3 and M4) is tested. This means that M4 (which assumes varying slopes for each damage state) does not provide a statistically significant improvement than its alternative. Therefore, the fit
of M3 is the best. Similarly, the $p$ value is well above the threshold for M3 vs. M5, highlighting that the construction type does not affect the slope of the fragility curves. In contrast, the $p$ value is well below the threshold for the comparison of M3 and M1, indicating that the construction type is an important variable and affects only the intercept. Having concluded that M3 based on the observed flow depth data fits the data better than its alternatives (i.e. M4, M5, and M1), we repeat the procedure to identify which simulated intensity type fits the data best. Table 6 also shows the comparison of the AIC values for the three simulated tsunami intensity types. For all simulated intensity types, M3 is identified as the model which fits the data better than its alternatives, and this conclusion is further reinforced by the likelihood ratio tests presented in Table 7. By comparing the AIC values for M3 for all three simulated intensity types, we note that the simulated flow depth is the tsunami intensity that fits the data best. The aforementioned observations can also be made if instead of the probit link function, the two alternative functions (i.e. logit and cloglog) are considered, as depicted in Table D1. The comparison of the AIC values of M3 for the three link functions identifies the probit link function as the one that fits the data best.

The regression coefficients of the 2018 Sunda Strait fragility curves based on the best-fitted M3 model with a probit link function are listed in Table E1. An advantage of constructing a complex model that accounts for the ordinal nature of the damage and for the two main construction types in the systematic component is that fragility curves for timber buildings can be obtained even for the states for which there are available data. A timber building is found to sustain more damage than confined masonry buildings for the more intense damage states. Nonetheless, there is substantially more uncertainty in the prediction of the likelihood of damage, and this can be attributed to the rather small sample size.

### Table 6. AIC values for the three models assuming probit link function fitted to the observed and simulated tsunami intensity measures of DB_Sunda2018.

| Model | Observed flow depth | Simulated flow depth | Simulated flow velocity | Simulated hydrodynamic force |
|-------|---------------------|----------------------|------------------------|-----------------------------|
| M3    | 129.9               | 138.5                | 224.2                  | 194.9                       |
| M4    | 137.7               | 148.4                | 227.7                  | 210.3                       |
| M5    | 131.6               | 139.8                | 225.3                  | 196.1                       |
| M1    | 162.0               | 169.0                | 246.5                  | 216.9                       |

### Table 7. Likelihood ratio test summary for all available observed and simulated tsunami intensity measures of DB_Sunda2018.

| Model | Observed simulated flow depth | $p$ value | Simulated simulated flow depth | $p$ value | Simulated simulated flow hydrodynamic force | $p$ value |
|-------|-------------------------------|-----------|--------------------------------|-----------|---------------------------------------------|-----------|
| M3    | ~0.41                         |           | ~0.72                         |           | ~0.08                                       |           |
| M4    | ~0.56                         |           | ~0.39                         |           | ~0.35                                       |           |
| M5    | ~0.00                         |           | ~0.00                         |           | ~0.00                                       |           |

### Table 8. AIC values for model M1 fitted to the simulated tsunami intensity measures of DB_Palu2018.

| Link function | Model | Simulated flow depth | Simulated flow velocity | Simulated hydrodynamic force |
|---------------|-------|----------------------|------------------------|-----------------------------|
| probit        | M1    | 276.8                | 286.3                  | 283.3                       |
| logit         | M1    | 276.2                | 286.3                  | 283.1                       |
| cloglog       | M1    | 280.3                | 286.5                  | 284.7                       |

### 4.2 DB_Palu2018

We also fit GLM models to the data in DB_Palu2018 using the observed tsunami flow depth to express the tsunami intensity and then to construct fragility curves and their 90% confidence intervals for the three individual damage states (Fig. 11). The data seem to produce fragility curves with positive slopes for $d_{S1}$ and $d_{S2}$ and a negative slope for $d_{S3}$. This latter observation is counter-intuitive as it is expected that the likelihood of collapse will grow with the increase in the tsunami depth. This outcome could be attributed to the collected sample, which includes very few collapsed buildings observed at low flow depth levels.

Based on the observations of the exploratory analysis, we use identical slopes for the fragility curves for all three damage states ($d_{S1} - d_{S3}$) to tackle the negative slope for $d_{S3}$ and three link functions. Therefore, model M1 is fitted to DB_Palu2018 assuming that the tsunami intensity is expressed in terms of simulated flow depth, flow velocity, and hydrodynamic force. Table 8 depicts the AIC values for each model. We note that for all cases the flow depth fits the data the best. Table 8 also shows that the logit function fits the data best. The regression coefficients of the 2018 Sulawesi–Palu fragility curves for the logit function are depicted in Table E2.

https://doi.org/10.5194/nhess-21-2313-2021

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5.1 Building fragility curves of the 2018 Sunda Strait tsunami

The fragility curves determine conditional damage probabilities according to the tsunami intensity measures of the 2018 Sunda Strait event for both confined masonry concrete (Fig. 13a–c) and timber (Fig. 14a–c) buildings of DB_Sunda2018. In Fig. 14a and b, there are no data to predict the shape of the curves between 0–1 m flow depth and 0–1 m s\(^{-1}\) flow velocity. The curves as a function of the observed flow depth reveal a great similarity with the ones based on the simulated flow depth from the TUNAMI two-layer model (Figs. 13a and 14a). For instance, when the observed and simulated flow depths reach 3 m, the likelihood of minor to major damage (i.e., \(\geq d_{s1}\) and \(\geq d_{s2}\)) for both timber and confined masonry buildings is approximately 99% (Fig. 14a and b). In contrast, the likelihood of complete damage (i.e., \(\geq d_{s3}\)) is 70% for timber buildings and only 10% for confined masonry buildings. Consequently, the tsunami functions based on observation and simulation are highly similar, which illustrates the accuracy and the reliability of the tsunami inundation model. The curves show that confined masonry-type buildings have higher performance than timber structures. When the flow depth is greater than 5 m and 2.5 m, the probability of complete damage is around 99% for confined masonry and timber buildings, respectively. We also compare the completely damaged or washed away fragility curve for confined-masonry buildings to Syamsidik et al. (2020), who developed the curve as a function of observed flow depth for these buildings, as depicted in Fig. 13a. Fragility curves representing complete damage or washed away are similar up to 4.5 m flow depth. Each curve estimates a 15% building damage probability at 3.5 m flow depth. However, a few data points are available beyond 5 m in the Sunda Strait area. Therefore, the damage probability uncertainty is greater for this value, hence the difference between our \(d_{s3}\)-curve and the one produced by Syamsidik et al. (2020). The curves as functions of the maximum simulated flow velocity and the hydrodynamic force are displayed in Figs. 13b and 14b and Figs. 13c and 14c for confined masonry concrete and timber buildings, respectively.

5.2 Building fragility curves of the 2018 Sulawesi–Palu tsunami

The 2018 Sulawesi–Palu tsunami curves are developed for confined masonry buildings with unreinforced clay brick of DB_Palu2018. The computed and surveyed curves show a similar damage trend. When the observed and simulated flow depths reach 1.5 m, the building damage probabilities for partial damage repairable (i.e., \(\geq d_{s1}\)), partial damage unreparable (i.e., \(\geq d_{s2}\)), and complete damage (i.e., \(\geq d_{s3}\)) are around 90%, 40%, and 15%, respectively (Fig. 15a). The fragility curves based on the observed and simulated flow depths are relatively similar, especially for \(d_{s1}\) and \(d_{s3}\). The curves based on the flow velocity and the hydrodynamic force are displayed in Fig. 15b and c.

5.3 Comparison between the 2018 and 2004 building fragility curves

In Fig. 16, we compare (i) the Sunda Strait and Sulawesi–Palu \(d_{s3}\)-curves based on the simulated tsunami intensity measures for confined masonry-type buildings, (ii) the 2004 Indian Ocean (Khao Lak–Phuket, Thailand) \(d_{s3}\)-curve based on the observed flow depth for reinforced-concrete infilled frames buildings (Foytong and Ruangrassamee, 2007; Ros-
Figure 11. Probit functions fitted for each individual damage state to DB_Palu2018 to assess whether the observed flow depth is an efficient descriptor of damage. The 90% confidence interval is plotted.

Figure 12. Probit functions fitted for each individual damage state to DB_Thailand2004 to assess whether the observed flow depth is an efficient descriptor of damage. The 90% confidence interval is plotted.
Figure 13. The 2018 Sunda Strait curves for confined masonry concrete buildings. Best-estimate fragility curves, with their 90% confidence intervals, as functions of (a) the observed and the maximum simulated flow depths, (b) the maximum simulated flow velocity, and (c) the simulated hydrodynamic force for confined masonry concrete buildings of DB_Sunda2018 sustaining minor or moderate damage ($d_{S1}$), major damage ($d_{S2}$), and complete damage or washed away ($d_{S3}$) in Sunda Strait area.

setto et al., 2007; Ruangrassamee et al., 2006), and (iii) the 2004 Indian Ocean (Banda Aceh, Indonesia) $d_{S3}$-curves produced by Koshimura et al. (2009a). The curves are based on a visual damage interpretation of remaining roofs using the pre- and post-tsunami satellite data (IKONOS) and are thus developed for mixed buildings (low-rise wooden, timber-framed, and non-engineered reinforced concrete constructions; Koshimura et al., 2009a; Saatcioglu et al., 2006). For 1 m flow depth, the likelihood of complete damage is greater in Palu (10%) than in Banda Aceh, Khao Lak–Phuket, and Sunda Strait (Fig. 16a, Table 10). However, when the flow depth reaches 3 m, the damage probability is about 50% in Banda Aceh, 25% in Palu City, and less than 20% in Khao Lak–Phuket. We also note that the likelihood of completely damaged or washed away buildings is higher in Sunda Strait than in Khao Lak–Phuket above 4 m flow depth. However, the data points in Thailand are mostly ranging from 0 to 5 m, and the 90% confidence interval upon this value is constantly increasing with the flow depth. Below 1 m s$^{-1}$, the flow velocity has a low impact on the damage probability in Banda Aceh (< 1%). However, beyond this value, the probability of damage becomes very sensitive to the current velocity (Fig. 16b, Table 10). As an example, when the flow velocity attains 6 m s$^{-1}$, the curve estimates 99% building damage probability in Banda Aceh. The hydrodynamic force also contributes to increase the probability of complete damage in Banda Aceh. For example, when the force reaches 25 kN m$^{-1}$, the damage probability is around 99% in Banda Aceh (Fig. 16c, Table 10).

6 Discussion

6.1 Reliability of the building fragility curves

The reliability of the curves depends mainly on (i) the quality and the quantity of post-tsunami data and (ii) whether the tsunami intensity measures are efficient predictors of damage. With regard to the first factor, DB_Sunda2018, DB_Palu2018, and DB_Thailand2004 include relatively little data (Table 2). For each database, the relatively broad confidence intervals around the best-estimate fragility curves reflect the small sample size. Moreover, the complexity of each studied event also plays a role in how well the selected tsunami intensity measure can represent the tsunami damage. In particular, in DB_Sunda2018 and DB_Thailand2004, only the tsunami load is responsible for the building dam-
Figure 14. The 2018 Sunda Strait curves for timber buildings. Best-estimate fragility curves, with their 90% confidence intervals, as functions of (a) the observed and the maximum simulated flow depths, (b) the maximum simulated flow velocity, and (c) the simulated hydrodynamic force for timber buildings of DB_Sunda2018 sustaining minor or moderate damage ($ds_1$), major damage ($ds_2$), and complete damage or washed away ($ds_3$) in Sunda Strait area.

6.2 Impact of the wave period, ground shaking, and liquefaction events on the building damage probability

The curve comparison illustrates well the relationship between the 2004 Indian Ocean, the 2018 Sunda Strait, and the...
2018 Sulawesi–Palu tsunamis characteristics, summarized in Table 11, and the structural performance of buildings.

Impact of the wave period

The 2018 Sunda Strait tsunami and the 2004 IOT (Khao Lak–Phuket, Thailand) are characterized by dominant wave periods of about 7 min (Muhari et al., 2019) and 40 min (Karlsson et al., 2009; Puspito and Gunawan, 2005; Tsuji et al., 2006), respectively (Table 11). Damage from ground shaking or liquefaction episodes was not reported, so the tsunami is the main cause of building damage. We compare the Sunda Strait and the Indian Ocean (Khao Lak–Phuket) curves based on the flow depth to investigate the impact of the tsunami wave period on buildings. In Fig. 16a, the curves showed that the short wave period tsunami in the Sunda Strait is less damaging than the 2004 IOT below 5 m flow depth. For instance, for 3 m flow depth, the likelihood of complete damage is around 20% in Khao Lak–Phuket against only 10% in the Sunda Strait area (Table 10). On the other hand, above 5 m flow depth, the structures in Khao Lak–Phuket reveal a better performance than the ones in the Sunda Strait area. As few data points are available beyond this value for completely damaged buildings, the Sunda Strait and the Indian Ocean (Khao Lak–Phuket) curve reliability is insufficient. Even though the long wave periods of the IOT seem to increase the likelihood of building damage, the sample size of collapsed buildings beyond 5 m flow depth is too small to validate this assumption.

Impact of ground shaking and liquefaction events

The city of Banda Aceh and the Khao Lak–Phuket area were damaged by the 2004 IOT. Along Banda Aceh shores, the simulated tsunami wave period ranges from 40 to 45 min (Prasetya et al., 2011; Puspito and Gunawan, 2005), and the one simulated in Khao Lak–Phuket is estimated at approximately 40 min (Karlsson et al., 2009; Puspito and Gunawan, 2005; Tsuji et al., 2006). Although the tsunami wave periods are similar at both locations, the 2004 Indian Ocean earthquake was strongly felt in the city of Banda Aceh, where it lasted about 10 min (Table 11). The earthquake intensity is estimated at VII to VIII on the Modified Mercalli Scale (Ghobarah et al., 2006; Saatcioglu et al., 2006). Despite the
Table 10. Damage probabilities of buildings reaching complete damage according to the intensity measures of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khao Lak–Phuket and Banda Aceh) tsunamis.

| Tsunami intensity measure                  | Building damage probability (%) |
|-------------------------------------------|---------------------------------|
|                                           | Sunda Strait | Sulawesi–Palu | Khao Lak–Phuket | Banda Aceh |
| Observed and simulated flow depths (m)    | 1            | < 1           | 10              | < 1       |
|                                           | 3            | 8             | 25              | 17        |
|                                           | 6            | 62            | –               | 43        |
| Simulated flow velocity (m s\(^{-1}\))    | 1            | 13            | 10              | –         |
|                                           | 3            | 19            | 11              | –         |
|                                           | 6            | 25            | 13              | –         |
| Simulated hydrodynamic force (kN m\(^{-1}\)) | 25           | 35            | 17              | –         |
|                                           | 50           | 48            | 19              | –         |
|                                           | 100          | –             | 22              | –         |

Other reasons include the high tsunami wave height and the high ground motion intensity. The high tsunami wave height and the high ground motion intensity could have triggered liquefaction events and enhanced the building susceptibility to tsunami damage in the waterfront of Palu City. This assumption cannot be verified through satellite images; it needs direct and close observations, which might be erased by the tsunami.

7 Conclusions

According to the GEM guidelines, building fragility curves of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khao Lak–Phuket, Thailand) tsunamis are empirically developed from post-tsunami databases respectively called DB_Sunda2018, DB_Palu2018, and DB_Thailand2004. To improve our understanding of the

https://doi.org/10.5194/nhess-21-2313-2021

Nat. Hazards Earth Syst. Sci., 21, 2313–2344, 2021
Figure 16. Best-estimate fragility curves for the 2018 Sunda Strait tsunami, 2018 Sulawesi–Palu tsunami, and 2004 IOT in Khao Lak–Phuket (Thailand) and Banda Aceh (Indonesia) as functions of (a) the observed and the maximum simulated flow depths, (b) the maximum simulated flow velocity, and (c) the simulated hydrodynamic force. These fragility functions are developed only for completely damaged or washed away buildings with their 90% confidence intervals.

Table 11. Characteristics of the 2004 Indian Ocean tsunami in Banda Aceh (Indonesia) and Khao Lak–Phuket (Thailand), as well as the 2018 Sulawesi–Palu and 2018 Sunda Strait tsunamis.

| Tsunami event | Indian Ocean | Indian Ocean | Sulawesi–Palu | Sunda Strait |
|---------------|--------------|--------------|---------------|-------------|
| Database      | Koshimura et al. (2009a) | DB_Thailand2004 | DB_Palu2008 | DB_Sunda2018 |
| Location      | Banda Aceh, Indonesia | Khao Lak–Phuket, Thailand | Palu City, Indonesia | Sunda Strait, Indonesia |
| Tsunami source| Earthquake   | Earthquake   | Landslides    | Landslide   |
| Ground shaking| +           | –            | +             | –           |
| Liquefaction  | +           | –            | +             | –           |
| Wave period   | Long (~ 40–45 min) | Long (~ 40 min) | Short (~ 3.5 min) | Short (~ 7 min) |
| Construction type | Mixed (e.g., reinforced concrete, timber) | Reinforced concrete | Confined masonry | Confined masonry, timber |

+ : recorded; – : not recorded.
Figure 17. (a) Liquefaction areas surveyed inland near Palu City and (b) magnified view of the maximum simulated flow depth of the 2018 Sulawesi–Palu tsunami overlaid on the masonry-type buildings completely damaged ($d_{33}$) and location of the coastal retreats surveyed in the waterfront of Palu City (background ESRI).

structural damage caused by the Sunda Strait and Sulawesi–Palu tsunamis, we reproduce their tsunami intensity measures (i.e., flow depth, flow velocity, and hydrodynamic force) with the TUNAMI two-layer model for the first time. The flow depth is constantly the best descriptor of tsunami damage for each event. The building fragility curves for complete damage reveal the following. (i) The buildings affected by the Sunda Strait tsunami sustained less damage than the ones in Khao Lak–Phuket (IOT). For example, for 3 m flow depth, the building damage probability is around 20% in Khao Lak–Phuket against 10% in the Sunda Strait area hit by a short wave period tsunami (landslide source). Considering that the tsunami was the main cause of structural damage (i.e., damage related to ground shaking and/or liquefaction was not recorded), the longer wave period of the 2004 IOT may have increased the likelihood of complete damage, and (ii) the building resilience is weaker in Banda Aceh than in Khao Lak–Phuket. For 3 m flow depth, the likelihood of complete damage is about 50% in Banda Aceh and 20% in Khao Lak–Phuket. Although both locations were hit by the 2004 IOT, Banda Aceh was strongly affected by ground shaking before the tsunami’s arrival, and (iii) the buildings affected by the Sulawesi–Palu tsunami were more susceptible to be completely damaged than the ones affected by the IOT in Banda Aceh (i.e., ≤2 m). As an example, for 1 m flow depth, the building damage probability of complete damage is about 10% in Palu and 5% in Banda Aceh. The Sulawesi–Palu tsunami is a complex event as it may not be the only cause of structural destruction. The 2018 Sulawesi earthquake caused minor damage to buildings and most importantly could have triggered liquefaction events in the waterfront of Palu City where coastal retreats have been observed, increasing the susceptibility of buildings to tsunami damage.
Appendix A: Two-layer modelling of a subaerial and submarine landslide

Figure A1. Two-layer modelling of a subaerial and submarine landslide (from the original sketch of Pakoksung et al., 2019): (a) pre-failure, (b) generation of negative and positive waves due to the landslide, and (c) landslide in progress and wave propagation.
Appendix B: The 2018 Sunda Strait tsunami generation, propagation, and inundation modelling

Figure B1. Temporal evolution of the 2018 Sunda Strait tsunami wave (a) 10, (b) 20, (c) 60, and (d) 120 s after the volcano flank collapse. The red line is the topography and bathymetry before the landslide (Pakoksung et al., 2020).

Figure B2. Temporal evolution of the 2018 Sunda Strait tsunami wave (a) 10, (b) 20, (c) 60, and (d) 120 s after the volcano flank collapse (Pakoksung et al., 2020).

https://doi.org/10.5194/nhess-21-2313-2021
Figure B3. (a–c) Magnified views of the maximum simulated flow velocity of the 2018 Sunda Strait tsunami overlaid on the damaged building data in the Rajabasa, Pejamben, and Sumur areas (background ESRI and © Google Maps).
Appendix C: The 2018 Sulawesi–Palu tsunami
generation, propagation, and inundation modelling

Figure C1. Temporal evolution of the 2018 Sulawesi–Palu tsunami wave (a) 2, (b) 10, (c) 30, and (d) 60 s after the S8 landslide. The red line is the topography and bathymetry before S8 landslide (Pakoksung et al., 2019).
Figure C2. Temporal evolution of the 2018 Sulawesi–Palu tsunami wave (a) 2, (b) 10, (c) 30, and (d) 60 s after the S8 landslide (Pakoksung et al., 2019).
**Figure C3.** Sulawesi–Palu final tsunami inundation model with the maximum simulated flow velocity overlaid on the damaged building data (background ESRI).

**Appendix D: Statistical model selection: comparison of AIC values for logit and cloglog link functions (DB_Sunda2018)**

**Table D1.** AIC values for the three models assuming logit and cloglog link functions fitted to the observed and simulated tsunami intensity measures of DB_Sunda2018.

| Model | Observed flow depth | Simulated flow depth | Simulated flow velocity | Simulated hydrodynamic force |
|-------|---------------------|----------------------|------------------------|-----------------------------|
|       | logit               | cloglog              |                        |                             |
| M3    | 132.5               | 134.8                | 224.3                  | 196.0                       |
| M4    | 146.4               | 144.5                | 229.0                  | 220.3                       |
| M5    | 134.2               | 136.1                | 225.5                  | 197.4                       |
| M1    | 163.6               | 168.8                | 247.0                  | 217.9                       |

https://doi.org/10.5194/nhess-21-2313-2021 Nat. Hazards Earth Syst. Sci., 21, 2313–2344, 2021
Appendix E: Regression coefficients for the building fragility curves of the 2018 Sunda Strait, 2018 Sulawesi–Palu, and 2004 Indian Ocean (Khao Lak–Phuket) tsunamis

**Table E1.** Regression coefficients for the 2018 Sunda Strait tsunami fragility curves based on DB_Sunda2018.

| Tsunami intensity measure                  | Regression coefficients (best estimate, standard error) |
|-------------------------------------------|--------------------------------------------------------|
|                                           | $\theta_1$     | $\theta_2$     | $\theta_3$     | $\theta_1$     | $\theta_2$ (class=Timber) |
| Observed flow depth                       | −0.29, 0.415   | −1.99, 0.402   | −4.52, 0.639   | 2.76, 0.408     | 2.08, 0.416              |
| Simulated flow depth                      | −0.26, 0.377   | −1.69, 0.355   | −4.03, 0.545   | 2.40, 0.346     | 1.96, 0.390              |
| Simulated flow velocity                   | 0.80, 0.300    | 0.14, 0.293    | −1.17, 0.307   | 0.27, 0.276     | 1.40, 0.296              |
| Simulated hydrodynamic force              | −4.07, 1.016   | −4.95, 1.058   | −6.50, 1.116   | 0.61, 0.118     | 1.45, 0.311              |

**Table E2.** Regression coefficients for the 2018 Sulawesi–Palu tsunami fragility curves based on DB_Palu2018.

| Tsunami intensity measure                  | Regression coefficients (best estimate, standard error) |
|-------------------------------------------|--------------------------------------------------------|
|                                           | $\theta_1$     | $\theta_2$     | $\theta_3$     | $\theta_1$     |
| Observed flow depth                       | 2.33, 0.315    | −0.71, 0.193   | −2.09, 0.286   | 0.57, 0.272     |
| Simulated flow depth                      | 2.37, 0.319    | −0.79, 0.199   | −2.20, 0.293   | 0.91, 0.286     |
| Simulated flow velocity                   | 2.07, 0.428    | −0.87, 0.370   | −2.23, 0.428   | 0.18, 0.335     |
| Simulated hydrodynamic force              | 0.35, 1.034    | −2.65, 1.061   | −4.03, 1.096   | 0.24, 0.127     |

**Table E3.** Regression coefficients for the 2004 IOT in Khao Lak–Phuket (Thailand) based on DB_Thailand2004.

| Tsunami intensity measure                  | Regression coefficients (best estimate, standard error) |
|-------------------------------------------|--------------------------------------------------------|
|                                           | $\theta_1$     | $\theta_2$     | $\theta_3$     | $\theta_1$     |
| Observed flow depth                       | 0.71, 0.377    | −1.59, 0.361   | −3.84, 0.481   | 2.00, 0.342     |
Code and data availability. Post-tsunami field survey data are available from references cited in the text. The bathymetric and topographic data for the Sunda Strait area were provided by BATNAS and DEMNAS, Indonesia, respectively (http://tides.big.go.id/DEMNAS/index.html, last access: 1 February 2020) (DEMNAS, 2020). The Geospatial Information Agency (BIG), Indonesia, provided the bathymetric and topographic data for Palu Bay. The tidal gauge records were supplied by the Coastal Disaster Mitigation Division, Ministry of Marine Affairs and Fisheries, Jakarta, Indonesia. Spatial data in this study are depicted through QGIS software.

Author contributions. FI, AS, KP, EL, II, and FB designed and coordinated this research. AS, KP, and EL performed the tsunami simulations and participated in the calibration of the inundation models in Palu Bay and in Sunda Strait. SS and RP contributed to the tsunami data collection in the Sunda Strait and Palu areas, respectively. II developed the fragility functions through advanced statistical analysis. All authors contributed to the drafting of the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We greatly acknowledge the three reviewers for their constructive comments and recommendations that helped to improve the quality of this manuscript. This research was funded and supported by the Japan Society for the Promotion of Science (JSPS) Grant-in-Aid for Young Scientists, the JSPS-NRCT Bilateral Research grant, the World Class Professor (WCP) programme 2018–2020 promoted by the Ministry of Education and Culture of the Republic of Indonesia, the Pacific Consultants Co., Ltd., the Willis Research Network (WRN), the Tokio Marine & Nichido Fire Insurance Co., Ltd., the National Institute of Water and Atmospheric Research (Project: CARH2106), UKRI GCRF Urban Disaster Risk Hub, and GLADYS.

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Financial support. This research has been supported by the Japan Society for the Promotion of Science (Grant-in-Aid for Young Scientists (B)), the National Research Council of Thailand (Bilateral Research grant, fiscal year 2017–2018), the Kementerian Riset Teknologi Dan Pendidikan Tinggi Republik Indonesia (World Class Professor, WCP, programme 2018–2020), the Université de Montpellier (grant GLADYS), the Global Challenges Research Fund (grant no. Urban Disaster Risk Hub NE/S009000/1), the National Institute of Water and Atmospheric Research (project: CARH2106), Pacific Consultants Co., Ltd., Willis Research Network (WRN), and Tokio Marine & Nichido Fire Insurance Co., Ltd.

Review statement. This paper was edited by Maria Ana Baptista and reviewed by three anonymous referees.

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