New evidence of Neoarchean crustal growth in southern São Francisco Craton: the Carmópolis de Minas Layered Suite, Minas Gerais, Brazil

Luís Emanoel Alexandre Goulart¹*, Mauricio Antonio Carneiro², Issamu Endo³, Marcos Tadeu de Freitas Suita³

ABSTRACT: The Carmópolis de Minas Layered Suite is situated in the Campo Belo Metamorphic Complex domain, Minas Gerais, SE Brazil, which, together with other metamorphic complexes in the surroundings of the Quadrilátero Ferrífero, composes the Archean sialic crust of the southern part of the São Francisco Craton. This suite encompasses a remaining segment of a magmatic arc, which is deformed and metamorphosed to upper amphibolite-granulite facies. It shows layered and massive strata, constituted by rock types with lithogeochemical signature equivalent to low-K₂O tholeiitic series. Both tholeiitic and calc-alkaline magmatism phases were identified. Amphibolite and metarhyolite samples were used for zircon U-Pb dating using MC-LA-ICP-MS. The amphibolite zircons yield 232Th/238U ratios ≈ 0.17 – 0.83 and a Concordia age of 2,752 ± 18 Ma. This interval yields an approximation to minimum age of the calc-alkaline phase. The metarhyolite zircons yielded 232Th/238U ratios ≈ 0.25 – 0.66 and ages of 2,713 ± 9.9 and 2,710 ± 31 Ma. This variation was interpreted as the minimum age interval of the calc-alkaline phase. One of the metarhyolite samples yielded two inherited zircon populations. These populations were dated at 3,374 ± 30 and 2,859 ± 23 Ma, suggesting the contribution of older crusts in the genesis of the suite. The results showed that the genesis of this suite represents an episode of juvenile arc crust formation during the Rio das Velhas Tectonomteral Event (2,780 Ma). This syn- to late-orogenic episode involved arc development, tholeiitic to calc-alkaline magmatism, metamorphism, and crustal recycling.

KEYWORDS: magmatic arc; U-Pb geochronology; calc-alkaline; tholeiitic; Archean.

RESUMO: A Suíte Acamadada Carmópolis de Minas está situada no domínio do Complexo Metamórfico Campo Belo, Minas Gerais, SE Brasil, que, junto a outros complexos metamórficos nos arredores do Quadrilátero Ferrífero, compõem a crosta sialica arqueana da parte meridional do Craton São Francisco. Essa suíte compreende um segmento remanescente de arco magmático, deformado e metamorfoizado em fácies anfibolito-superior a granulito. Existe estratos acamadados e maciços, constituídos por litotipos com assinatura litogeoquímica equivalente a uma série tholeiítica de baixo-K. Uma fase de magmatismo tholeiítico e outra de magmatismo calcio-alcalino foram identificadas. Uma amostra de anfibolito e duas de metarriolito foram utilizadas para datação U-Pb em zírconos usando MC-LA-ICP-MS. O anfibolito apresentou zírconos com razões 232Th/238U ≈ 0,17 – 0,83 e idade concórdia de 2,752 ± 18 Ma. Esse intervalo foi interpretado como idade mínima da fase tholeiítica. As amostras de metarriolito apresentaram zírconos com razões 232Th/238U ≈ 0,25 – 0,66 e foram datadas em 2,713 ± 9,9 e 2,710 ± 31 Ma. Esse intervalo fornece uma aproximação para a idade mínima da fase calcio-alcalina. Uma das amostras de metarriolito apresentou duas populações de zírconos herdados. Essas populações foram datadas em 3,374 ± 30 e 2,859 ± 23 Ma, sugerindo a contribuição de crostas mais antigas na gênese da suíte. Os resultados demonstraram que a gênese dessa suíte representa um episódio de formação de crosta juvenil durante o Evento Tectonomteral Rio das Velhas (2.780 Ma). Esse episódio sin a tardi-orogênico envolveu desenvolvimento de arco, magmatismo tholeiítico e calcio-alcalino, metamorfismo e reciclagem crustal.

PALAVRAS-CHAVE: arco magmático; geocronologia U-Pb; calcio-alcalino; tholeiítico; Arqueano.
INTRODUCTION

The Archean tectonic evolution of the southern portion of the São Francisco Craton (SFC) was marked by episodes of formation, thickening, and consumption of crust, as proposed by Carneiro (1992), Machado & Carneiro (1992), Machado et al. (1992, 1996), Noce et al. (1998), Silva et al. (2000), Teixeira (1985), and Teixeira et al. (2000), among others. Among this collection of papers, those of Teixeira et al. (1996, 1998) stand out, because they point to a polyphase history with three periods of intense tectono-magmatic activity between the Paleo- and Neoarchean Eras. According to these authors, the first period incorporates ages older than 3.1 – 2.9 Ga and contributes to the set of oldest U-Pb ages of the South American Platform. The ages in this range represent a fruitful period of crustal formation and thickening (e.g. Teixeira et al. 1996, 1998, Carneiro et al. 1998; Hartmann et al. 2006). U-Pb ages with this order of magnitude were obtained in zircons from the gneisses of Campo Belo Metamorphic Complex (CBMC; Teixeira et al. 1998), the Belo Horizonte Metamorphic Complex (CMBH; Teixeira et al. 1996), the mafic-ultramafic volcanic sequence of the Piumhi greenstone belt (Machado & Schrank 1985), and the metasedimentary rocks of the Rio das Velhas Supergroup (SGRV; Machado et al. 1996; Noce et al. 2005; Hartmann et al. 2006). This period is also recorded in Sm-Nd TDM model ages, and Pb/Pb and Rb-Sr isochron ages (Carneiro 1992; Machado & Carneiro 1992; Machado et al. 1992; Teixeira et al. 1996, 1998), obtained in the metamorphic complexes that surround the Quadrilátero Ferrífero mining district. The second event occurred between 2.9 – 2.8 Ga. Several studies in zircons from the migmatites neosome of the CBMC (Teixeira et al. 1998) and BHMC (Noce 1995; Noce et al. 1998) record an episode of high-grade metamorphism and migmatization between 2860 and 2839 Ma. In addition, processes of oceanic crust formation during this period are suggested by Sm-Nd isochron age of 2863 ± 65 Ma, obtained in komatiitic rocks from the Morro do Ferro greenstone belt (Pimentel & Ferreira Filho 2002). There is no mention in the literature of records on the 2860 – 2839 Ma event in the Bonfim Metamorphic Complex. This fact indicates that the tectonic evolution of this unit was distinct from the other metamorphic complexes in this region. However, the CMB records evidence of an event of magmatism and crustal reworking between 2780 and 2700 Ma (e.g. Carneiro 1992; Machado & Carneiro 1992). This event was defined by Carneiro (1992) as the Ribeirão do Velhas Tectono-thermal Event (RVTE) and comprises the third period of tectonic activity described by Teixeira et al. (1996, 1998), also referred to as Ribeirão dos Velhas Orogeny. RVTE involved reworking of older TTG gneisses during the evolution of a continental margin, generating metamorphism (Carneiro 1992) and producing andesitic to calc-alkaline magmatism (Carneiro 1992; Machado et al. 1992, 1996; Noce et al. 2005). The apex of this event was marked by a collisional stage between 2752 and 2700 Ma (e.g. Carneiro 1992; Noce et al. 1998), which included the Nova Lima Group sedimentation (Hartmann et al. 2006), and, at the final stages of stabilization of the orogen, an extensional event, responsible for the emplacement of mafic-ultramafic bodies in the crust (Pinese 1997; Carneiro et al. 2004) and a syn- to late-orogenic granitogenesis (Carneiro 1992; Noce 1995; Teixeira et al. 1996; Romano 1989).

Recent work by Romano et al. (2013) and Lana et al. (2013), corroborating the previous papers, proposed three main tectono-magmatic events of TTG crust formation. The authors named these events as Santa Bárbara (3220 – 3200 Ma), Rio das Velhas I (RV I; 2930 – 2900 Ma), and Rio das Velhas II (RV II; 2800 – 2770 Ma). After the RV II Event, the coalescence of the crustal segments that currently represent the metamorphic complexes in the vicinity of the Quadrilátero Ferrífero was recorded, where an extensive 2760 – 2700 Ma potassic magmatic event took place, marking the Archean cratonic core stabilization (e.g. Lana et al. 2013).

The isotopic records of these events are also identified at the Campo Belo Metamorphic Complex domain. The current tectonic models (e.g. Teixeira et al. 1998; Carneiro et al. 1998; Campos & Carneiro 2008) indicate that the crustal segment represented by CMCB collided with the N/NE sialic blocks, over the probable Neoarchean continental margin, corresponding to SW limit of the Quadrilátero Ferrífero mining district (Carneiro 1992; Machado & Carneiro 1992; Campos & Carneiro 2008). In contrast, either there is no information about the style of margin in the block opposite (CMCB) or available information is more generalist, addressing only the collisional orogen as a whole.

In this sense, within the Campo Belo Metamorphic Complex and close to the inferred limit with the Bonfim Metamorphic Complex crops out the Carmópolis de Minas Layered Suite (CMLS – e.g. Goulart & Carneiro 2013). Previous work by Carneiro et al. (2007) originally correlated the CMLS rocks with the Neoarchean Ribeirão dos Motas Layered Sequence (RMLS), which crops out at the SW studied area. This unit was interpreted by Carvalho Júnior (2001) as a stratiform layered intrusion, correlated to the mafic-ultramafic volcanic rocks of the Rio das Velhas Supergroup. Whole-rock Sm-Nd isotope constraints (Carneiro et al. 1997, 2004) yielded an isochron age of ca. 2.79 Ga for the Ribeirão dos Motas Layered Sequence, supporting this interpretation. Nevertheless, later works (e.g. Goulart & Carneiro 2013; in press) observed evidence of a tardy calc-alkaline magmatic phase in the Carmópolis de Minas Layered Suite, not observed...
in the Ribeirão dos Motas Layered Suite. These evidences consist of the emplacement of intermediate to acidic rocks, posterior to the tholeiitic magmatic phase. Given the new geological constraints, Goulart et al. (2013, unpublished) and Goulart & Carneiro (2013, in press) characterized the SACM as Neoarchean juvenile magmatic arc-related suite, correlated to the late RV II. The present work shows the preliminary results of U-Pb (LA-ICP-MS) geochronological dating in zircons from rocks of the Carmópolis de Minas Layered Suite, which discusses: 1) ages of the phases of tholeiitic and calc-alkaline magmatism; and 2) possible implications of this unit in the crustal evolution of the Southern São Francisco Craton.

**LOCATION AND GEOLOGIC CONTEXT OF CARMÓPOLIS DE MINAS LAYERED SUITE**

Carmópolis de Minas Layered Suite comprises a crustal segment of the Campo Belo Metamorphic Complex, located southwest of the Quadrilátero Ferrífero (QF), in the Carmópolis de Minas municipality (State of Minas Gerais, Figs. 1 and 2). This suite is positioned in a sinformal structure approximately 30 km long (Carmópolis de Minas Synclinal – Fig. 2). To the north, the structure axis trends NE-SW and inflects NW-SE to the south.

Contacts between the Carmópolis de Minas Layered Suite and the CBMC TTG gneisses are inferred. Along the unit, granitoids and three mafic dyke systems (two NW-SE and one NE-SW direction), designated as Lençóis I / II and Timboré (e.g. Carneiro et al. 2007) suites, intrude into this stratigraphic sequence. Remnants of detrital and chemical sedimentary units crop out subordinately and in small volumes. These rocks consist of metachert, gondite, magnetitite layers, and quartzite, which occasionally occur interspersed with amphibolite lenses.

The suite is composed of layered and massive, plutonic or sub-volcanic, mafic-ultramafic to felsic lithotypes, metamorphosed at amphibolite to granulite facies conditions (Figs. 3A to C). The typical outcrops of the CMLS show NE-SW-vertical to sub-vertical layering. In the syncline closing, the magmatic layering shows expressive deformation, exhibiting well-defined NE-SW-oriented M-fold axes (Fig. 3B). Geochemical and isotopic constraints (Goulart et al. 2013, unpublished; Goulart & Carneiro 2013, in press) indicate that low Mg number (Mg# < 57.87, except metaultramafic cumulates), low K2O contents (0.1 – 1.45 wt.%), and positive to slightly negative εNd(t) values (+13.14 to -0.25) are ubiquitous features of these rocks. According to the cited authors, the predominant low Mg# values in association with vestiges of metasedimentary rocks indicate that the SACM protolith may have been derived from more evolved magmas, which in turn were fractionated from primary magmas, to then be housed at shallower depths in the crust. The low-K2O rocks, as opposed to high-K2O and TTG suites currently known in the sialic substrate of Southern CSF (e. g. Alkmin & Noce 2006; Campos & Carneiro 2008; Romano et al. 2013, Lana et al. 2013), indicate the SACM to be an Archean metamorphic equivalent of modern arc-related tholeiitic suites of the boninite-basalt-andesite-rhyolite type, suggesting that their protoliths were originated in a juvenile oceanic arc setting (see Gill 1981; Leat et al. 2003 Tamura et al. 2009).

The composition of the lithological assemblage is in agreement with the expected average proportions between tholeiitic and calc-alkaline rocks from juvenile magmatic arc-related suites (e.g. Gill 1981). Tholeiitic rocks are volumetrically predominant and represent several metamorphic lithotypes as arc-related amphibolite-metaultramafic cumulates and metamorphosed arc-related basic to intermediate rocks.

The metamafites correspond to several varieties of amphibolite and leucocambriolite. Garnet-amphibolite was described and occurs in small amounts. These rocks are composed of plagioclase, hornblende, pyroxene, Fe-Ti-oxides, garnet, and quartz. Metasomatized rocks contain actinolite, Na-plagioclase, biotite, epidote, scapolite, quartz, and carbonat. The meta-metaultramafic rocks are orthopyroxenite, hornblendite, plagioclase-metapiroxenite, meta-olivine-websterite, and meta-harzburgite. Garnet-bearing amphibolite and garnet-bearing websterite have been described, but they occur in small amounts. The lithotypes are composed of orthopyroxene, clinopyroxene, hornblende, olivine, garnet, spinel, Fe-Ti-oxides, and sulfides. The secondary paragenesis is formed by serpentinite, talc, tremolite/Mg-hornblende, and chlorite in smaller quantities. Metadunite and metaherzolite are unusual and present original mineralogy intensely replaced by secondary paragenesis. These rocks keep relic primary textures such as euhedral twinned clinopyroxene crystals showing rim recrystallization (mantled texture), euhedral tabular plagioclase crystals with recrystallized cores (ghost crystals), and several relic cumulatic textures constituted by plagioclase, clinopyroxene, orthopyroxene, and olivine in lesser importance. The calc-alkaline rocks are less abundant and consist of metamorphosed pyroxenitic/ryholitic rocks. These lithotypes occur as felsic injections, parallel to sub-parallel to the metaultramafic-amphibolitic bedding. Granitoids crop out in some portions of the unit, but the relationship of these with the CMLS is not known. These rocks are formed by Na-plagioclase, K-feldspar, and quartz. Mafic minerals are biotite, pyroxene, and/or hornblende and Fe-Ti-oxides but occur in small amounts.
Figure 1. Simplified geologic map of southern São Francisco Craton showing the location of the study area and part of the SW limit of the Quadrilátero Ferrífero. The Carmópolis de Minas Layered Suite boundaries are still represented as in Carneiro et al. (2007). In the upper left corner the location of the study area is shown in relation to the Southern São Francisco Craton (dark gray area). Sources: Leite et al. (2004), Carneiro et al. (2007).
Figure 2. Geologic map of Carmópolis de Minas Layered Suite.

1 and 2: undifferentiated gneisses of the CBMC; 3: metamafic rocks of the CMLS; 4: metaultramafic rocks of the CMLS; 5: Chemical metasedimentary rocks; 6: Metagranitoid; 7: regional-scale fold axis (inverted syncline); 8: second-order fold axis; 9: roads and highways; 10: rivers; 11: town; 12: sampling site.
Figure 3. Carmópolis de Minas Layered Suite typical structures. (A) Magmatic layering constituted by leucoamphibolite and amphibolite. (B) “M”-folded magmatic layering constituted by leucoamphibolite and amphibolite, with stratification formed by the accumulation of poikilitic hornblende crystals. (C) Magmatic layering in meta-websterite. (D) and (E) Amphibolite enclave in metarhyolite. These structures are associated with the presence of poikilitic hornblende (Hbl). (F) Thin section showing poikilitic hornblende crystals (Hbl) hosting plagioclase (Pl) and magnetite (Mag). Observe the euhedral to subhedral habit of the crystals (PPL, enlargement 5X). (G) and (H) Outcrop showing the sampled metarhyolite and amphibolite. The acidic injections are parallel to sub-parallel to the amphibolite strata.
Mafic enclaves, microgranular and/or orbicular aggregates have been observed in the stratified domains, poikilitic crystals of pyroxene are present, allowing for the analysis of the rock's isotopic ratios. The selected crystals were mounted in an araldite circular section 2.5 cm in diameter and then analyzed by LA-MC-ICP-MS (Laser Ablation Mass Spectrometry) equipment at the Geochronology Laboratory of the Geosciences Institute of Brasília University. Zircon crystals were separated from the non-magnetic fraction, and the samples were analyzed for their isotopic compositions.

In press) obtained point to two magmatic sources for the SACM rocks. Positive εNd(t) values (> +0.41) point out that the tholeiitic rocks are mantle-derived, whereas the calc-alkaline rocks, presenting slightly negative εNd(t) values (-0.25), show that they may have been generated in a sin- to late- orogenic phase by sub-arc crust anatexis. Several field evidences confirm the two magmatic phases during SACM formation:

- Mafic enclaves, microgranular and/or orbicular aggregates, and poikilitic euhedral crystals of hornblende, ortho- and clinopyroxene (Figs. 3D to F), hosting countless Fe-Ti-oxides and phenocrystal of plagioclase, are observed in the contact between amphibolite layering and felsic injections. These structures suggest assimilation of sub-arc mafic crust by felsic magmas.

- Amphibolite enclaves and lenses showing low-grade metamorphic paragenesis with Na-plagioclase, epidote, actinolite, quartz, and minor biotite also occur at the felsic injections contact. According to Morgan & London (1987), the contact between acidic melts and mafic rocks may cause metasomatism and low-grade paragenesis formation. This point of view implies that the tholeiitic and calc-alkaline magmatic phase cannot be contemporaneous, but suggests that the calc-alkaline magmatism may have involved several pulses, where not all have yet been identified.

- In the stratified domains, poikilitic crystals of pyroxene and hornblende were also observed. According to several authors (e.g. Storey et al. 1989; Seaman et al. 1995; Lindberg & Eklund 1988), similar textures can be generated by the contact reaction of melts with distinct rheological characteristics. Regarding CMLS, these textures suggest that tholeiitic and calc-alkaline magmatic phases could have coexisted, interacted, and generated hybrid magmas.

### RESULTS

Two metarhyolite samples (MR1 and MR2) and one amphibolite sample (AE19) were analyzed. These samples were obtained from the same outcrop, located in the northernmost part of CMLS, at the left bank of the Pail River, which corresponds to the closing of the Carmópolis de Minas Synclinal (UTM 540888/7733485; Figs. 2, 3G, and 3H). This place was chosen for the large lithotype diversity, with peculiar structures, such as abundant poikilitic hornblende crystals and mafic enclaves observed in metarhyolite. The selected grains showing low metamictization degree and sizes compatible with the laser spot diameter were analyzed. Despite the fact that the zircons had shown low contrast in the BSE imaging, the analyses were directed preferentially to the crystal cores, because the sizes of these domains are more compatible with the laser spot. Morphology and isotopic characteristics of zircon grains are presented in Figs. 4 and 5. Geochronological results are presented in Tab. 1 and Fig. 5.

**Sample AE19**

This sample is fine to medium grained and composed of hornblende, plagioclase, and quartz. Accessory minerals are titanite, zircon, apatite, and Ti-magnetite. This sample constitutes zircons D1, D2, D3, D11, D13, and D16. The grains present short and elongated prismatic shapes, occasionally with sub-rounded rims, as a function of metamorphic overgrowth. Internal domains are homogeneous with some fracturing.

Samples and standards were corrected after Pb and Hg blanks; 207Pb/206Pb and 206Pb/238U ratios were corrected after common Pb; common Pb assuming as concordant age 206Pb/238U * 207Pb/235U; 206Pb = 1/137.88 * U * U * 0.992743. Errors were calculated with 1 sigma (% for isotopic ratios, absolute ages). 7. Discordance % = 100 * (1 - 206Pb/238U t/207Pb/206Pb).
Despite the weak contrast of the images, it was possible to distinguish a blurred, oscillatory fine zoning in some crystals. Some zircons present metamorphic overgrowth separated from the internal domains by fractures with dissolution-resorption structures and inclusions of silicate or phosphate crystals. Some grains exhibit peripheral domains more enriched in inclusions. The analyses revealed moderate to high $^{232}$Th/$^{238}$U ratios (0.17 – 0.83; Tab. 1). The obtained results are concordant to strongly discordant (2 – 65% discordance) and consist of a large interval of $^{206}$Pb/$^{207}$Pb ages with range from 2749 ± 22 Ma to 2448 ± 22. The analytical points defined a Pb-loss line that intercepts the Concordia in the upper part of the diagram at 2744 ± 11 Ma and the lower part at 252 ± 12 Ma (MSWD = 0.38), the latter with no apparent geologic meaning. These discordant results are probably correlated with the presence of internal fractures in the analyzed grains. Nevertheless, the grouping of analysis defined by zircon grains D3, D11, and D16 yielded an apparent Concordia age of 2752 ± 18 Ma (MSWD = 1.6; Fig. 5).

Sample MR1

Sample MR1 is composed of plagioclase, K-feldspar, quartz, and subordinate biotite and orthopyroxene. Titanite and zircon are accessory minerals. The analyzed zircons were labeled with prefixes H and I. The grains are predominantly prismatic, short, subhedral, and/or equidimensional, occasionally sub-rounded and double-terminated. Internal domains are ample and homogeneous. Occasionally the subtle contrasts

Figure 4. Backscattering electron probe images of some zircons of this study. The circles represent the position of the laser spots.
in BSE images suggest the presence of blurred oscillatory zoning. The development of fractures is insignificant. In general, grains present peripheral domains with homogeneous texture, in some cases (zircon H25) with inexpressive concentric fractures delimiting the internal domains. Occasionally they are associated with dissolution structures, suggesting fluid percolation. Some grains present mineral inclusions. Results yielded moderate to high $^{235}$Th/$^{238}$U ratios in the 0.21 – 0.66 interval. Analysis from zircons H9, H15, H25, H28, and H18 presented concordant analytical points (discordance < 1%). These results yielded a weighted mean $^{206}$Pb/$^{207}$Pb age of 2718 ± 15 Ma and an apparent Concordia age of 2713 ± 9.9 Ma (MSWD = 1.6; Fig. 5).

### Sample MR2

The essential paragenesis of this sample is plagioclase, quartz, and K-feldspar. The mafic mineral is biotite and the accessory minerals are Ti-magnetite, titanite, apatite, and zircon.

The zircons of this sample define three populations with well-defined $^{206}$Pb/$^{207}$Pb ages. The first population is represented by zircons E6, E8, E11, E21, and E28. The crystals are prismatic, elongated, with sub-rounded rims, and internal domains with homogeneous texture or blurred oscillatory zoning. Mineral and fluid inclusions, fracturing, and botryoidal structures are present in all grains. It is possible to distinguish with some difficulty pervasive transgressive recrystallization in weakly contrasting BSE images. These structures indicate that these zircons were partially recrystallized during medium- to high-grade metamorphism (e.g. Vavra et al. 1999; Connelly 2000; Corfu et al. 2003).

The zircon analysis yielded high $^{232}$Th/$^{238}$U ratios (> 0.42 – 0.57). Grains E21 and E28 showed Th and U enrichment relative to Pb, when compared to E6, E8, and E11, resulting in slightly lower $^{235}$Th/$^{238}$U ratios (0.26 and 0.32). Zircons E6, E8, E11, E21, and E28 presented concordant to discordant results (1 – 10% disc.) and $^{206}$Pb/$^{207}$Pb ages with range in ages from 2717 ± 47 to 2183 ± 97 Ma. Discordant results probably reflect the presence of fractures in the zircons. Particularly, the analyses of zircons E21 and E28 resulted in apparent $^{207}$Pb/$^{206}$Pb ages, inconsistent with the other samples of this group (2183 ± 97 and 2210 ± 31 Ma), but aligned with these in a radiogenic Pb-loss line that intercepts the Concordia at 2710 ± 31 Ma and 773 ± 64 Ma (MSWD = 1.06). Zircons E08 and E11 (discordance <1%) yielded an apparent Concordia age of 2713 ± 15 Ma (MSWD = 0.42). The other zircon populations yielded well-defined $^{207}$Pb/$^{206}$Pb ages that are distinct

### Table 1. LA-MC-ICPMS analytical results

| Zircon | $^{206}$Pb/$^{207}$U | $^{207}$Pb/$^{206}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U | $^{206}$Pb/$^{207}$U |
|--------|----------------|----------------|----------------|----------------|----------------|----------------|----------------|----------------|----------------|
| D05    | 14.1096       | 2.84           | 0.5384         | 2.45           | 0.86           | 0.1900         | 1.43           | 2777           | 68             |
| D11    | 14.1099       | 1.33           | 0.5363         | 1.06           | 0.80           | 0.1908         | 0.81           | 2768           | 28             |
| D12    | 14.2281       | 2.23           | 0.5407         | 1.97           | 0.88           | 0.1900         | 1.05           | 2796           | 55             |
| D01    | 5.7651        | 3.46           | 0.2453         | 0.61           | 0.18           | 0.1705         | 3.41           | 1414           | 9              |
| D02    | 8.4533        | 1.56           | 0.5334         | 1.13           | 0.85           | 0.1859         | 0.76           | 1855           | 21             |
| D13    | 3.1016        | 1.18           | 0.1412         | 0.83           | 0.70           | 0.1593         | 0.84           | 852            | 17             |
| H09    | 13.453        | 2.87           | 0.520          | 2.45           | 0.85           | 0.1888         | 1.49           | 2700           | 66             |
| H15    | 13.323        | 2.08           | 0.518          | 1.36           | 0.65           | 0.187          | 1.58           | 2689           | 56             |
| H25    | 13.597        | 2.50           | 0.527          | 2.09           | 0.84           | 0.187          | 1.37           | 2730           | 57             |
| H28    | 13.416        | 1.88           | 0.522          | 1.57           | 0.84           | 0.186          | 1.03           | 2709           | 43             |
| I18    | 13.687        | 2.69           | 0.526          | 2.14           | 0.80           | 0.189          | 1.63           | 2723           | 58             |
| E01    | 15.2065       | 1.22           | 0.4483         | 1.02           | 0.85           | 0.2460         | 0.68           | 2587           | 24             |
| E27    | 26.0351       | 2.24           | 0.6662         | 2.01           | 1.02           | 0.2854         | 0.99           | 3291           | 66             |
| E04    | 13.2531       | 1.34           | 0.4663         | 0.52           | 0.39           | 0.2059         | 1.24           | 2467           | 13             |
| E17    | 8.5267        | 16.45          | 0.3044         | 14.62          | 0.89           | 0.2031         | 7.53           | 1713           | 250            |
| E06    | 12.8810       | 2.56           | 0.5035         | 2.27           | 0.88           | 0.1853         | 1.19           | 2629           | 60             |
| E08    | 13.6823       | 4.43           | 0.5304         | 1.08           | 0.92           | 0.1871         | 1.72           | 2743           | 112            |
| E11    | 13.4722       | 1.92           | 0.5249         | 1.61           | 0.84           | 0.1862         | 1.04           | 2720           | 44             |
| E21    | 3.8727        | 5.32           | 0.2057         | 2.94           | 0.55           | 0.1565         | 4.43           | 1206           | 55             |
| E28    | 4.5974        | 2.18           | 0.2405         | 1.66           | 0.76           | 0.1386         | 1.41           | 1390           | 23             |

**Geochemistry**

- **Th (ppm)**: 20.6
- **U (ppm)**: 238.2
- **Pb (ppm)**: 2718

Luís Emanoel Alexandre Goulart et al.

Braz. J. Geol., São Paulo, 43(3): 445-459, September 2013
Figure 5. Wetherill Concordia diagrams for Carmópolis de Minas Layered Suite rocks.

DISCUSSION AND CONCLUSIONS

Although the BSE images do not allow for a clear view of the internal features of the zircon grains, the external morphology in association with $^{232}\text{Th}/^{238}\text{U}$ ratios can lead to some interpretations. The BSE imaging of zircon grains, as previously presented, revealed stubby (short prismatic) to prismatic habits (preferentially), with sub-rounded overgrowths and possible internal euhedral zoning. This morphology is consistent with zircons that underwent metamorphism in the upper-amphibolite to granulite facies conditions (see Vavra et al. 1999, Corfu et al. 2003, Hoskin & Schaltegger 2003). Several studies have interpreted higher $^{232}\text{Th}/^{238}\text{U}$ ratios in high-grade metamorphic zircons from felsic (e.g. Schaltegger et al. 1999; Hoskin & Black 2000) and mafic protolith (e.g. Vavra et al. 1999; Liati & Gebauer 1999; Rubatto et al. 1999) as a magmatic signature. Although there is no consensus on the interpretations of this feature, the $^{232}\text{Th}/^{238}\text{U}$ ratios tend to be higher for granulite facies zircons than for eclogite and amphibolite facies zircons (Vavra et al. 1999). Work by Rubato (2002) showed that zircons from metamorphic eclogite have chemical compositions similar to igneous zircons; however, they have lower $^{232}\text{Th}/^{238}\text{U}$ ratios (< 0.1). The low $^{232}\text{Th}/^{238}\text{U}$ ratios may indicate a complete recrystallization of zircon and consequently reset of the isotopic systems (Hoskin & Schaltegger 2003). In this sense,
studies on eclogite and mafic-ultramafic granulate (e.g., Lati & Gebauer 1999) have considered high $^{232}\text{Th}/^{238}\text{U}$ ratios from zircons with little or no preserved magmatic domains, as a pre-metamorphic heritage, and correlated ages are interpreted as the minimum age for crystallization of the protolith. This is the case of the U-Pb zircon dating from CMLS rocks.

The obtained ages between $= 3374$ and $2710$ Ma are recurrent in the Southern CSF (e.g., Teixeira et al. 1998; Noce et al. 2005; Hartmann et al. 2006). In accordance with the studies of Teixeira (1996, 1998), Hartmann et al. (2006), and Lana et al. (2013), this age range records four magmatic events that occurred between the Paleo- and Neoarchean at the South São Francisco Craton.

Analyses of the amphibolite AE19 yielded a discordant upper intercept age of $2744 \pm 11$ Ma (MSWD = 1.5 – Fig. 5), which we interpret as a reasonable approximation to the age of the tholeiitic magmatic phase in the Carmópolis de Minas Layered Suite. However, because of restrictions regarding the areas not clearly preserved in the magmatic zircons sampled, we interpret the apparent Concordia age of $2752 \pm 18$ Ma as being the most likely crystallization age of the amphibolite protolith and the minimum age of the tholeiitic magmatism of the CMLS. This interpretation is also supported by the predominant low Mg# ($< 57.87$) of this suite, indicating that the protoliths of the studied tholeiitic rocks were not crystallized from primary magmas, but from evolved magmas that were previously fractionated into the crust (Goulart et al. 2013 in preparation). The presented results point that the CMLS tholeiitic magmatic phase represents a period of emersion of ocean arcs (Fig. 6A) during ETRV, which contributed to the formation of the Campo Belo Metamorphic Complex. According to Hartmann et al. (2006), the age of $= 2.75$ Ga represents the Rio das Velhas Orogeny apex ($= 2759 – 2751$ Ma) in the southern São Francisco Craton.

The sample MR1 yielded a weighted mean $^{206}\text{Pb}/^{207}\text{Pb}$ age of $2718 \pm 19$ Ma. This result overlaps within error with the upper intercept age of $2710 \pm 31$ Ma (MSWD = 1.06) that was obtained in sample MR2 and constrains a likely age range for the development of the calc-alkaline magmatism recorded in the CMLS. Some field correlations that we presented in this work suggest the possibility of the fact that the calcium-alkaline magmatic phase might have occurred in several pulses. Therefore, we interpret the Concordia ages of $= 2713$ Ma, which the samples MR1 and MR2 yielded, as the crystallization age of the metarhyolite protolith.

The ages of the samples MR1 and MR2 show that the CMLS calc-alkaline magmatic phase is slightly younger than the felsic magmatism of the Rio das Velhas Supergroup ($= 2746$ Ma) in the Quadrilátero Ferrífero domain (Machado et al. 1992; Noce et al. 2005; Hartmann et al. 2006). However, the magmatic protolith ages of the metamorphosed acidic rocks are contemporaneous to several granitoids, commonly described in metamorphic complexes present in the outskirts of QF (Romano 1989; Carneiro 1992; Noce 1995; Romano et al. 2013) that now are associated with the Neoarchean potassic magmatic event defined by Romano et al. (2013) and Lana et al. (2013). Other granitoids mark the development of continental margin, at the SW limit of the QF, in the region currently represented by the Bonfim Metamorphic Complex (e.g., Carneiro 1992; Machado & Carneiro 1992; Campos & Carneiro 2008; Lana et al. 2013). In contrast, the spatial and temporal correlations between the SACM and these rocks suggest that to the southwest, the convergence margin of the segment currently represented by Campo Belo Metamorphic Complex may have been characterized by the development of oceanic arcs parallel to the continental edge, defining one accretionary edge (Fig. 6C). Inside the continental block and in response to the evolution of a convergent margin, the development of an extensional environment supports the intrusion of the Ribeirão dos Motas Layered Sequence (Fig. 6C), whose whole-rock isochron age of ca. $2790$ Ma (Carneiro et al. 2004) overlaps within error with the $2752 \pm 18$ Ma age obtained for the amphibolite AE19.

At the moment, the difference that we observe between the ages of the amphibolite AE19 and metarhyolite MR1 and MR2 rules out the possibility of magmatic hybridization as an explanation to the structures observed in the contact between the tholeiitic and calc-alkaline phases; on the other hand, it strengthens the thesis of assimilation and metasomatism of the preexisting amphibolite crust by acidic melts. However, the presence of reliquiar compositional banding with the accumulation of poikilitic hornblende/orthopyroxene or microgranular mafic aggregates indicates that physical-chemical instability occurred in the contact between magmas with distinct rheological characteristics (see Lindberg & Eklund 1988). This observation still suggests that possibly other phases of calc-alkaline magmatism, not yet identified, could have coexisted with the tholeiitic magmatism along the CMLS evolution.

Tectonic models correlate the calc-alkaline magmatism genesis with the anatexis of the mafic sub-arc crust (e.g., Stern 2010), where the crustal emplacement of intermediate to felsic magmas records the beginning of the intermediate maturity stages of arc evolution, well defined by the formation of a felsic middle crust (Fig. 6B). The mafic lower crust was formed during infant and juvenile stages of the subduction and arc-related magmatism by the gradual replacement of the preexisting mafic-ultramafic oceanic crust. Presumably, the formation age of the initial mafic
arc crust is constrained by the age of the CMLS tholeiitic magmatic phase and therefore must be equal to or older than 2752 Ma (Fig. 6B). Development of juvenile arc-related magmatism would have involved the consumption and recycling of preexisting oceanic crusts and assimilation of crust-derived components.

Although the amphibolite AE19 has not presented direct evidence of crust-derived component assimilation, the sample MR2 records the presence of inherited zircons, with U-Pb ages ranging between ≈ 3374 and 2859 Ma. These zircons represent a crustal inheritance that was passed from the lower arc crust to acidic melts during anatexis. Ages with this magnitude record two tectonic events that occurred between Paleo- and Mesoarchean, involving the development of continental nuclei and probable juvenile crust formation (see Machado & Schrank 1989; Pimentel & Ferreira Filho 2002; Hartmann et al. 2006; Lana et al. 2013). These events correspond to the Santa Barbara and Rio das Velhas I Events, according to that which was proposed by Lana et al. (2013), and reaffirm the formation of primitive (oceanic and continental) crustal nuclei at that period.

Figure 6. Schematic model for the evolution of the Carmópolis de Minas Layered Suite arc magmatism (based on Stern 2010). (A) Juvenile arc. Tholeiitic magmatism (amphibolite, mafic, and ultramafic cumulates). (B) Late to mature arc. Crustal thickening and beginning of calc-alkaline magmatism by sub-arc crust anatexis. (C) Tectonic model showing the genesis of the SACM. Observe correlation with Ribeirão dos Motas Layered Sequence.
Unpublished work by Goulart & Carneiro (2013, submitted), which supports the proposition of this model, was based on Nd-Sr isotope data from a representative part of the lithological framework that composes the Carmópolis de Minas Layered Suite. The authors report a significant isotopic data set that includes several positive to slightly negative $\varepsilon_{Nd}$ parameters, variable $\varepsilon_{Sr}$ parameters, a pseudo-isochron age for CMLS of 2736 ± 300 Ma with $\varepsilon_{Nd,i} = +0.4$, and six Sm-Nd TDM model ages ranging from 3266 to 2860 Ma. From these results we can highlight the following:

- The pseudo-isochron age is consistent with the ages that we have presented in this paper, whereas the slightly positive value of $\varepsilon_{Nd,i}$ reflects the affinity mantle of the primitive protolith of the CMLS.

- The Sm-Nd TDM model age range constrains, with a best approximation, the age of the magma sources and characterizes an isotopic inheritance from protolith that has been involved in the arc-magmatism or in sub-arc crust formation.

- Whole-rock analysis carried out in amphibolite, similar to sample AE19, yielded positive $\varepsilon_{Nd}(2713)$ and $\varepsilon_{Sr}(2713)$ parameter values, with an Sm-Nd TDM model age of 2985 Ma, indicating that the tholeiitic primordial magmas were originated in depleted mantle sources and later, during emplacement in the crust, they assimilated older crustal components.

- The single metahyolite sample (MR1) yielded a Sm-Nd TDM model age of 2860 Ma, older than the suite, although newer than Sm-Nd TDM model ages of the other samples. Thus, the genesis of the protolith of this rock can be linked to sub-arc crust anatexis, a fact corroborated by the slightly negative $\varepsilon_{Nd}$ (2713) values (-0.25) and positive $\varepsilon_{Sr}$ (2713) values (+12), suggesting a protolith correlated to juvenile crustal reservoir, with short crustal residence time. The Sm-Nd TDM model age of 2860 Ma replicates within error the upper intercept age that was obtained in the ≈ 2859 Ma inherited zircons (E04, E17, and E24) of sample MR2, confirming the probable age of one of the crusts involved in the sub-arc crust formation.

Thus, the development of the Carmópolis de Minas Layered Suite occurred in a restricted range time (2752 – 2710 Ma), during the last 40 Ma of the Rio das Velhas Orogeny. In this period, the crustal segments that currently compose the Campo Belo Metamorphic Complex consisted of isolated young crustal nuclei that were agglutinated between Rio das Velhas Orogeny and potassic magmatism after this. At the end of the Neoarchean events, the isolated continental nuclei, including the younger protocontinental segments and the last magmatic arcs (e.g. CMLS), were amalgamated, starting the architecture and stabilization of the Archean platform that preceded the Proterozoic events and presently represents the sialic substrate of this portion of the São Francisco Craton.

ACKNOWLEDGMENTS

We wish to thank FAPEMIG for the financial support for the projects (MAC; CRA - RDP-00067-10), Capes for the doctoral scholarship (LEAG), and CNPq for the productivity scholarship (MAC). We would also like to thank Editor Cordani, U. and all anonymous reviewers for constructive comments.

REFERENCES

Campos J.C.S & Carneiro M.A. 2008. Neoarchean and Paleoproterozoic granitoids marginal to the Jecéa-BomSucesso lineament (SE border of the southern São Francisco craton): Genesis and tectonic evolution. Journal of South American Earth Sciences, 26:463-484.

Carneiro M.A. 1992. O Complexo Metamórfico Bonfim Setentrional - Quadrilátero Ferrífero, Minas Gerais: Itaconstratigrafia e evolução geológica de um segmento de continental do arqueano. Tese de Doutorado, Instituto de Geociências, Universidade de São Paulo, São Paulo, 235 p.

Carneiro M.A., Jord-Evangelista H. & Teixeira W. 1997. Eventos magmáticos arqueanos de natureza cálcio-alcalina e tholeítica no Quadrilátero Ferrífero e suas implicações tectônicas. Revista Brasileira de Geociências, 27:121-128.

Carneiro M.A., Teixeira W., Carvalho Júnior I.M., Oliveira A.H. & Fernandes R.A. 1997. Archean Sm/Nd isochron age from the Ribeirão dos Motas layered rocks sequence, Southern São Francisco Craton, Brazil. In: South-American Symposium on Isotope Geology, Campos do Jordão. Extended Abstracts. p. 63-64.

Carneiro M.A., Teixeira W., Carvalho Júnior I.M. & Fernandes R.A. 1998. Ensilic Tectonic Setting of the Archean Rio das Velhas Greenstone Belt: Nd and Pb Isotopic Evidence from the Bonfim Metamorphic Complex, Quadrilátero Ferrífero, Brazil. Revista Brasileira de Geociências, 28(2):189-200.

Carneiro M.A., Teixeira W., Carvalho Jr. I.M., Pimentel M.M. & Oliveira A.H. 2004. Comportamento dos Sistemas Sm-Nd e Rb-Sr da Seqüência Acamadada Máfico-Ultramáfica Ribeirão dos Motas (Arqueano), Cratão São Francisco Meridional: Evidências de Enriquecimento Mantelício e Fracionamento Isotópico. Revista do Instituto de Geociências, Geologia USP Série Científica, São Paulo, 4(2):13-26.

Carneiro M.A., Endo I, Nalini Jr. H.A, Sales J.C.C., Goulart L.E.A., Silva E.F., Pereira A.A., Tavares T.D., Jamelaro F, Carneiro J.M., Mariano
Carvalho J.R. 2008. Orogenic Andesites and Plate Tectonics. Reviews in Mineralogy and Geochemistry, 53:423-439.

Corfu F., Hanchar J.M., Hoskin P.W.O., Kinny P. 2003. Atlas of zircon textures. In: Hanchar J. M. & Hoskin W. O. P. (eds). Zircon. Reviews in Mineralogy and Geochemistry 53:468-500.

Gill J. 1981. Orogenic Andesites and Plate Tectonics. Springer, 390 p.

Goulart L.E.A. & Carneiro M.A. 2013 (in press). Evolution of arc magmatism in the Carmópolis de Minas Layered Suite, Minas Gerais, Brazil. Sm-Nd and Rb-Sr isotope geochemistry. Revista Escola de Minas (no date to be published).

Hartmann L.A., Endo I., Suita M.D.F., Santos J.O.S., Frantz J.C., Carneiro M.A., McNaughton N.J. & Bailey M.E. 2006. Provenance and age delimitation of Quadrilátero Ferrífero sandstones based on zircon U-Pb isotopes. Journal of South American Earth Sciences, 20:273-285.

Hoskin P.W.O. & Black L.P. 2000. Metamorphic zircon formation by solid-state recrystallization of protolith igneous zircon. Journal of Metamorphic Geology, 18:425-459.

Hoskin P.W.O. & Schaltegger U. 2003. The Composition of Zircon and Igneous and Metamorphic Petrogenesis. In: Hanchar J.M. & Hoskin P.W.O. (eds). Zircon. Reviews in Mineralogy and Geochemistry, 53:57-62.

Hawkins J.W. & Ishizuca O. 2009. Petrologic evolution of Palau, a volcanic arc. Geological Society of American Bulletin, 121:259-276.

Liati A. & Gebauer D. 1999. Constraining the prograde and retrograde P-T-t path of Eocene HP rocks by SHRIMP dating different zircon domains; inferred rates of heating, burial, cooling and exhumation for central Rhodope, northern Greece. Contributions to Mineralogy and Petrology, 135:340-354.

Lindberg B. & Eklund O. 1998. Interactions between basaltic and granitic magmas in a Svecofennian postorogenic granitoid intrusion, Aland, southwest Finland. Lithos, 22:13-23.

Ludwig K.R. 2000. Isoplot 3.0: A Geochronological Toolkit for Microsoft Excel, vol. 4. Berkeley Geochronology Center Special Publication, 71 p. (CD).

Machado N. & Carneiro M.A. 1992. U-Pb evidence of late Archean tectonothermal activity in the southern São Francisco shield, Brazil. Canadian Journal of Earth Science, 29:2341-2346.

Machado N. & Schrank A. 1989. Geochronology U-Pb do Maciço de Piúmba: resultados preliminares. In: Simpósio de Geologia de Minas Gerais, V. 1989. Belo Horizonte. Anais, Belo Horizonte: SBG, p. 45-49.

Machado N. & Schrank A.N., Noce C.M. & Gauthier G. 1996. Ages of detrital zircon from archean paleoproterozoic sequences: implications for greenhouse belt setting and evolution of a transamazonian foreland basin in Quadrilátero Ferrifero, southeast Brazil. Earth Planetary Science Letters, 141:259-276.

Machado N., Noce C.M., Ladeira E.A., Oliveira O.A.B. 1992. U-Pb geochronology of Archean magmatism and Proterozoic metamorphism in the Quadrilátero Ferrifero, Southern São Francisco Craton, Brazil. Geological Society of American Bulletin, 104:1221-1227.

Morgan G.B. & London D. 1987. Alteration of amphibolitic wallrocks around the Tanco rare-element pegmatite, Bennic Lake, Manitoba, American Mineralogist, 72:1057-1121.

Noce C.M. 1995. Geocronologia dos eventos magmáticos, sedimentares e metamoricos na regiao do Quadrilátero Ferrifero, Minas Gerais. Tese de Doutoramento, Instituto de Geociências, Universidade de São Paulo, São Paulo, 128 p.

Noce C.M., Machado N. & Teixeira W. 1998. U-Pb Geochronology of gnaisses and granitoids in the Quadrilátero Ferrifero (Southern São Francisco Craton): age constraints for archean and paleo-proterozoic magmatism and metamorphism. Revista Brasileira de Geociências, 1(28):95-102.

Noce C.M., Zucchi M., Baltazar O.F., Armstrong R.C., Dantas E., Renger F.E.A. & Lobato L.M. 2005. Age of felsic volcanism and the role of ancient continental crust in the evolution of the Neoarchean Rio das Velhas Greenstone belt (Quadrilátero Ferrifero, Brazil): U-Pb zircon dating of volcaniclastic granitic rocks. Precambrian Research, 141: 67-82.

Pimentel M. & Ferreira Filho C.F. 2002. Idade Sm-Nd de komatiitos do greenstone belt do Morro do Ferro, Fortaleza de Minas (MG). Revista Brasileira de Geociências, 32(1):147-148.

Pinse J.P.P. 1997. Geoquímica, geologia isotópica e aspectos petrográficos dos diques maficos pré-cambrianos da região de Lavras (MG). Porção Sul do Craton São Francisco. Tese de Doutorado, Instituto de Geociências, Universidade de São Paulo, São Paulo, 178 p.

Romano A.W. 1989. Evolution tectonique de la region NW du Quadrilatere Ferrifere - Minas Gerais, Brazil. Nancy. These, Docteur, Université de Nancy, France, 259 p.
Romano R., Lana C., Alkmim F.F., Stevens G. & Armstrong R. 2013. Stabilization of the southern portion of the São Francisco craton, SE Brazil, through a long-lived period of potassic magmatism. Precambrian Research, 224:143-159.

Rubatto D. 2002. Zircon trace element geochemistry: partitioning with garnet and the link between U-Pb ages and metamorphism. Chemical Geology, 184:123-138.

Rubatto D., Gebauer D. & Compagnoni R. 1999. Dating of eclogite-facies zircons: the age of Alpine metamorphism in the Sesia-Lanzo Zone (Western Alps). Earth and Planetary Science Letters, 167:141-158.

Schaltegger U., Fanning C.M., Gunther D., Maurin J.C., Schulmann K. & Gebauer D. 1999. Growth, annealing and recrystallization of zircon and preservation of monazite in high-grade metamorphism: conventional and in situ U-Pb isotope, cathodoluminescence and microchemical evidence. Contributions to Mineralogy and Petrology, 134:186-201.

Seaman S.J., Scherer E.E. & Standish J.J. 1995. Multistage magma mingling and the origin of flow banding in the Aliso lava dome, Tumacacori Mountains, southern Arizona. Journal of Geophysical Research, 100:8381-8398.

Silva L.C., Noce C.M. & Lobato L.M. 2000. Dacitic volcanism in the course of the Rio das Velhas (2800–2960 Ma) Orogeny: a Brazilian Archean analogue (TTD) to the modern adakites. Revista Brasileira de Geociências, 30:384-387.

Stacey J.S. & Kramers J.D. 1975. Approximation of terrestrial lead isotope evolution by a two-stage model. Earth and Planetary Science Letters, 26:207-221.

Stem R.J. 2010. The anatomy and ontogeny of modern intra-oceanic arc systems. In: Kusky T.M., Zhai M.G. & Xiao W. (eds), The Evolving Continents: Understanding Processes of Continental Growth, Geological Society, London, Special Publications, 358:7-54.

Storey M., Wolff J.A. & Narr M.J. & Marriner G.F 1989. Origin of hybrid lavas from Água de Pau volcano, São Miguel, Azores. In: Sanders A.D. & Norry M.J. (Eds), Magmatism in Ocean Basins. Geological Society of London, Special Publications, Classics, p. 161-180.

Tamura Y. 2003. Some geochemical constraints on hot fingers in the mantle wedge: evidence from NE Japan. In: Larter R.D. & Leat P.T. (eds), Intra-Oceanic Subduction Systems: Tectonic and Magmatic Processes. Geological Society, London, Special Publications, 219:221-237.

Teixeira W. 1985. A evolução geotectônica da porção meridional do Craton do São Francisco, com base em interpretações geocronológicas. Tese de Doutorado, Instituto Geociências, Universidade de São Paulo, São Paulo, 207 p.

Teixeira W., Carneiro M.A., Noce C.A., Machado N., Sato K. & Taylor P.N. & Nagasawa T. 1996. Pb, Sr and Nd isotope constraints on the Archean evolution of gneissic granitoid complexes in the southern São Francisco Craton, Brazil. Precambrian Research, 78:151-164.

Teixeira W., Cordani U.G., Nutman A.P. & Sato K. 1998. Polyphase Archean evolution in the Campo Belo Metamorphic Complex, Southern São Francisco Craton, Brazil: shrimp and U-Pb zircon evidence. Journal of South American Earth Science, Oxford, 3(11): 279-289.

Teixeira W., Sabate P., Barbosa J.S.F., Noce C.M. & Carneiro M.A., 2000. Archean and Paleoproterozoic Tectonic evolution of the São Francisco Craton, Brazil. In: Cordani, U.G., Milani, E.J., Thomas Filho, A., Campos, D.A. (Eds.), Tectonic Evolution of the South America, v.31. International Geologic Congress, Rio de Janeiro, Brazil, p. 101-157.

Vavra G., Schmid R. & Gebauer D. 1999. Internal morphology, habit and U-Th-Pb microanalysis of amphibolites to granulite facies zircons: geochronology of the Ivrea zone (Southern Alps). Contribution to Mineralogy and Petrology, 134:380-404.