Epitaxies of Ca sulfates on calcite. II. The main {010}, {001} and {100} forms of bassanite epitaxially deposited on the {10.4} substrate form of calcite

Dino Aquilano, Marco Bruno, Stefano Ghignone, Linda Pastero and Andrea Cotellucci

2D and 3D epitaxies of the main {010}, {001} and {100} forms of deposited bassanite (CaSO\(_4\cdot0.5\)H\(_2\)O) on {10.4} calcite (CaCO\(_3\)) as a substrate are described to provide a theoretical crystallographic background for the replacement of calcite by bassanite both in nature and in the laboratory and by weathering linked to cultural heritage. First, epitaxy in the third dimension, perpendicular to the investigated interfaces, has been verified in order to establish whether adsorption/absorption can occur (as anomalous mixed crystals) at the bassanite/calcite epitaxial contacts. Secondly, and by applying the Hartman–Perdok method, 2D lattice coincidences have been obtained from the physical-geometric matches of bonds running in the common directions within the elementary slices facing the substrate/deposit interfaces. This research represents the second and more detailed part of a wider program extended to the epitaxial interactions between the following pairs: (i) {010}-gypsum/{10.4}-calcite (just published); (ii) bassanite/{10.4}-calcite (the present work); and (iii) anhydrite (CaSO\(_4\))/{10.4}-calcite (coming soon).

1. Introduction

We recently investigated all compatible 3D and 2D lattice coincidences (herein 3D- and/or 2D-LCs) that can occur at the interface between the {10.4} form of cleaved calcite (substrate) and the {010} pinacoid of gypsum (deposit) (Aquilano et al., 2022). Working on calcium sulfates deposited on gypsum, we knew that when gypsum is heated above ~150°C in the dry state (or at 75°C in methanol–water solutions), a part of the crystalline water is removed and CaSO\(_4\cdot0.5\)H\(_2\)O is formed (Maslyk et al., 2022). This mineral, metastable at all temperatures, occurs in nature as bassanite (Bss) (Weiss & Bräu, 2009) and, as a biomineral, in some deep-sea medusae (Tiemann et al., 2002; Becker et al., 2005). In recent times, bassanite has been shown to play an important role in the complex system where gypsum, bassanite and anhydrite replace calcite (Cc) – in both natural and industrial processes. The authors who work in this sector have mainly focused on studies in the thermodynamic and kinetic fields, using the most advanced techniques of characterization. Thus, bassanite has been viewed as a precursor of gypsum (Van Driessche et al., 2012), or as a key product among the Ca sulfates replacing gypsum (Ruiz-Agudo et al., 2015, 2016). During recent years, we have arrived at the point of formulating a tentative general model for Ca sulfate precipitation from solutions and, through nucleation, to explain the occurrence of bassanite on the surface of Mars (Stawski et al., 2020).
Of all this work, the study we consider the most representative was conducted by Ruiz-Agudo et al. (2016), who determined the 3D crystallographic relationship between calcite (parent) and Ca sulfates (products) using X-ray texture analysis. They chemically obtained all the CaSO₄ crystalline phases from the reaction of H₂SO₄ foreign solutions with the replaced calcite and characterized the CaCO₃–CaSO₄ transformations by means of their experimental 2D X-ray diffraction analyses. Through this method, they indicated a clear crystal preferred orientation of the three Ca sulfate phases (both hydrated and anhydrous) formed during the interaction of calcite with sulfate-bearing solutions; briefly, they found that an epitaxial [10.4]Cc/[010]Bss relationship was observed in such a way that ‘...the orientation of the parent calcite determines the disposition of the crystals of the final CaSO₄ phase during transformation. The exact mechanism by which the crystallographic information is transferred in a dissolution–precipitation reaction is not well understood yet’ (Ruiz-Agudo et al., 2016). Using our notations, the geometry of the epitaxy (in Å) was described, at that time, as follows: only [001]Bss = 6.336 is parallel to the 1/3[441]Cc vector = 6.425, the linear misfit between them only reaching 1.45%. No other match was found at the calcite/bassanite epi-contact, and this suggests that only 1D- and not 2D-LCs can exist at this interface, in our opinion.

Here, we did not intend to repeat experiments already carried out by others, but only to integrate them and establish a useful tool for comparison and complementary purposes; moreover, we wanted to intervene only when the rules of epitaxy have been clearly violated, especially the crystallographic ones.

Starting from this background and bearing in mind that searching for 2D epitaxy among low-symmetry structures is not always easy, we aim in the present work to investigate all compatible 2D-LCs among the [10.4]Cc and the main [010], [001], [100] morphological forms of bassanite (Becker et al., 2005; Ruiz-Agudo et al., 2016). It is therefore a matter of starting from scratch, slowly moving in crystallographic morphology: to do this, we began with the surface profiles of all the involved forms, obtained through a strict application of the Hartman–Perdok method (Hartman, 1973). As we recently treated the [10.4]-calcite/[010]-gypsum epitaxy (Aquilano et al., 2022) in the same way, we will complete our program in a subsequent study dealing with the [10.4]Cc/anhydrite (CaSO₄) epitaxy, starting from crystallographic experience (Aquilano et al., 1992) acquired many years ago.

2. A short summary on the [10.4] surfaces of calcite
The usual unit cell (in Å) of rhombohedral calcite (space group R̄3c) reads \( a_0 = b_0 = 4.9896; c_0 = 17.06, \alpha = \beta = 90^\circ, \gamma = 120^\circ \). For the rectangular 2D cell of its [10.4] form, the vectors are \([010] = 4.9896\) and \(1/3[421] = 8.103\), this cleavage form being limited by a set of symmetry-equivalent vectors \(1/3[441] = 12.85\), running parallel to the [10.4] edges. Each [10.4] face shows a sharp pseudohexagonal symmetry; in fact, a large supercell occupying an area of 242.58 Å² with multiplicity \((6\times)\) can be drawn (Fig. S1 of the supporting information), and these features are more pronounced when we consider the epi-relationship of calcite/bassanite.

The [10.4]Cc is a flat (F) form growing through the layer-by-layer mechanism (either 2D nucleation or spiral, or both). Actually, four periodic bond chains (PBCs) run within the slice \(d_{[10.4]} = 3.034 \) Å thick. The two main PBCs develop along the \(\langle 441 \rangle\) and \(\langle 48\bar{1} \rangle\) directions, made equivalent through the glide plane ‘c’ [Fig. 1(a)], so building all the edges limiting the six rhombohedron faces. The two other main vectors run along the \(\langle 42\bar{1} \rangle\) and \([010]\) directions. It is fundamental here to

![Figure 1](image)

(a) Cleavage [10.4] calcite rhombohedron (dark-pink colour) with its PBCs running along the vertical glide plane ‘c’ parallel to \(\langle 42\bar{1} \rangle\) and the horizontal \([010]\) directions; the brown parts are related to the \([01.2]\) form. (b) Perpendicular view of the same rhombohedron, where the \(\langle 42\bar{1} \rangle\) and \([010]\) directions are indicated; Ca is blue, C is green and O is red.
recollect the related PBC strength (the end chain energy, erg ion$^{-1} \times 10^{19}$), i.e. the energy released when an ion enters, in a crystallographic position, at one end of each semi-infinite chain: 0.391, 0.359 and 0.333 for the PBCs $(441)$, $(42\bar{1})$ and $(010)$, respectively (Ruiz-Agudo et al., 2016; Stawski et al., 2020). It has also been demonstrated that there is only one way to choose the surface profile of $(10.4)$. Accordingly, the $(10.4)$ profile does not need to be reconstructed, since no atoms can be found on the ideal planes separating two adjacent $d_{10.4}$ slices. In other words, $d_{10.4}$ are ‘self-consistent slices’. Thirty years ago, more or less, we quantified its compactness (Hartman, 1973), i.e. the interaction energy ($E_{\text{slice}}^{(10.4)}$) between the atoms contained within the $d_{10.4}$ slice, and found that $E_{\text{slice}}^{(10.4)} = 0.222$ erg $\times 10^{-10}$ ion$^{-1}$, corresponding to no less than ~94% of the calcite crystallization energy (Aquilano et al., 1992). The shape of growth (or dissolution) of $(10.4)$ patterns (spirals and/or 2D nuclei) is theoretically defined (Hartman, 1973) by the $(441)$ steps limiting the faces, followed by the $(42\bar{1})$ vertical and $(010)$ horizontal directions, as anticipated and demonstrated in Fig. 1.

3. Bassanite 001: the sharp stacking difference of its elementary $d_{020}$ layers with respect to those of the [001] and [100] forms. A comparison with the $d_{10.4}$ layers of the cleavage calcite rhombohedron

Table 1 and Fig. 2 show the profound differences between the behaviour of the main forms [010], [001] and [100] of bassanite with respect to the same [10.4]$_{\text{Cc}}$ substrate. In other words, there is no compatibility between the thickness of the elementary slices $d_{10.4} = 3.043$ Å (calcite) and $d_{020} = 3.4635$ Å (bassanite). To find an acceptable correspondence in the thickness $(10.4)_{\text{Cc}}/(010)_{\text{Bss}}$ one has to rise up to the thickness of $(6-9) \times d_{020}$ layers of bassanite; after this, the misfit starts to rise again (from $+2.43\%$) and adsorption/absorption of 2D-[010]$_{\text{Bss}}$ layers into the bulk of the [10.4] form of calcite becomes improbable in comparison with the cases for other basic bassanite forms. This means that bassanite adsorption can only occur at the $(10.4)_{\text{Cc}}/(010)_{\text{Bss}}$ interface. Fig. 2 provides evidence showing that, for the other two interfaces $(10.4)_{\text{Cc}}/(001)_{\text{Bss}}$ and $(10.4)_{\text{Cc}}/(100)_{\text{Bss}}$, the thickness correspondences are everywhere very close to each other, in such a way that the
misfit reaches a maximum of \(-1.16\) and \(+4.1\%\) for \(d_{200}\) and \(d_{002}\), respectively. Underlining this difference is useful, as it highlights the pseudo-quadratic 2D symmetry of (010)\(_{\text{Bss}}\) with respect to the pseudo-hexagonality of both the (100) and the (001) planes.

3.1. The pseudo-quadratic hexagonal symmetry of bassanite viewed along [010] and the pseudo-hexagonal symmetry along both [100] and [001] directions

As described in the Introduction, we adopted the bassanite structure proposed by Ballirano et al. (2001) and Hildyard et al. (2011), who determined the monoclinic space group \(I\_2\) and the cell parameters (in Å) \(a_0 = 12.032\), \(b_0 = 6.927\), \(c_0 = 12.671\) and \(\beta = 90.27^\circ\). Fig. 3 roughly describes the sub-symmetry of bassanite:

The right side shows that the [010] form looks ‘pseudo-quadratic’; in fact, the vector [200]\(_{\text{Bss}}\) = 24.064 and its perpendicular [002]\(_{\text{Bss}}\) = 25.342 differ by a misfit of 5.31\%. The comparison between the supercell made by these two vectors and that made by calcite [42\(\bar{1}\)] = 24.309 and [010] = 24.948, building the \(\{10.4\}\)\(_{\text{Cc}}\) supercell, points out the striking 2D-LCs (see cases 4a and 4b in Table 2) occurring between [010]\(_{\text{Bss}}\) and \(\{10.4\}\)\(_{\text{Cc}}\).

The left side outlines that the [001] and [100] forms are ‘pseudohexagonal’. In fact, (i) the vertical side is common and has a length of 13.854; (ii) the diagonal ones have lengths of 13.883 and 14.441 in the forms [001] and [100], respectively; and (iii) the six internal angles range from 119.93 to 122.64\(^\circ\) in the [001] form, and from 118.67 to 122.64\(^\circ\) in the [100] form. The areas (in Å\(^2\)) of these 2D-LCs [multiplicity (6\times)] vary from 500.074 to 526.63, going from the [001] to [100] forms. The resulting \(\Delta\%\) reaches 5.31.

To summarize, through the exposed surfaces, we can obtain further proof that it is reasonable to treat [010]\(_{\text{Bss}}\) separately from [001] and [100], when the epi-contact with \(\{10.4\}\)\(_{\text{Cc}}\) is made.

3.2. 2D coincidence lattices between bassanite [010] and calcite [10.4]

According to the preceding sections, the best fit between a vector in the (010) plane of bassanite and a chain in the (10.4) plane of calcite is that between the most important edge \(1/3\langle 441\rangle_{\text{calcite}} = 12.85\) Å and the most important axis of bassanite [001]\(_{\text{Bss}}\) = 12.671 Å, the linear misfit being 1.45\%. Accordingly, the statement (Ruiz-Agudo et al., 2016) that no 2D-LC can be found at this interface (see the Introduction) is rather pessimistic. In fact, from Table 2 and Fig. 4, one can find the following:

(i) The [441]\(_{\text{Cc}}\) chain is fundamental to build up an epitaxy on (010)\(_{\text{Bss}}\) and the corresponding side of the shared 2D-LC is
[100]_{\text{Bss}}. The linear misfit between the two vectors is \(-6.84\)%, which is compensated by the opposite misfit (+7.82%) occurring between the two other sides of the 2D-LC. The linear compensation is reflected in the low value (\(-6\))% of the maximum area misfit which, together with the minimum multiplicity (1\times) of the 2D-LC of bassanite and its tolerable (4.44\%) obliquity value, allows us to say that the epitaxy constraints in case 1 are well satisfied.

(ii) Another reasonable condition that could be found for a 2D epitaxy occurs with case 2 (Table 2). The epitaxy constraints are also fulfilled in this case, but the linear misfits are coherent and hence the error propagates towards the long-range interactions, and the multiplicity of the 2D common area is twice the preceding value. Accordingly, the probability of epitaxy exists, but is lowered.

(iii) In cases 3a and 3b, the 2D common areas or angular misfits (or both) exceed the geometrical constraints in order for an epitaxy to occur.

(iv) In cases 4a and 4b, one obtains the largest 2D common areas but the best of the angular misfits. Furthermore, the linear misfits are in opposition everywhere.

To summarize, [010]_{\text{Bss}} has two opportunities to form good epitaxies with [10.4]_{\text{Cc}}. Case (1) illustrates the ‘short-range’ 2D-LC, owing to the lowest multiplicity (1\times) of the [010]_{\text{Bss}} lattice. Cases 4a and 4b describe the ‘long-range’ 2D-LC, as it ensues from the multiplicity (4\times). By now, only the values of the adhesion energy between [010]_{\text{Bss}} and [10.4]_{\text{Cc}} could indicate which one of the epitaxies will be the preferred one. Anyway, it is worth remembering that the small 2D epi-nuclei of bassanite could form at medium–high supersaturation, whereas the larger ones are stable even at low supersaturation (with respect to bassanite). In the first case, short-range 2D-LCs are coupled with small nuclei, whereas the long-range ones will be coupled with the larger nuclei. Accordingly, this is the best evidence that epitaxy [010]_{\text{Bss}}/[10.4]_{\text{Cc}} has good probability to occur.

From the occurrence frequency expressed in Table 2, one obtains these observed rules:

(i) [100]_{\text{Bss}} is parallel to [441]_{\text{Cc}}, [42T]_{\text{Cc}}, [010]_{\text{Cc}} and, less frequently, to [419]_{\text{Cc}}, [42T]_{\text{Cc}}, [010]_{\text{Cc}} and, less frequently, to [417]_{\text{Cc}}.

(ii) [001]_{\text{Bss}} is parallel to [42T]_{\text{Cc}}, [010]_{\text{Cc}} and, less frequently, to [451]_{\text{Cc}} and [411]_{\text{Cc}}.

(iii) [101]_{\text{Bss}} is parallel to [42T]_{\text{Cc}} and, less frequently, to [451]_{\text{Cc}} and [411]_{\text{Cc}}.

(iv) [201]_{\text{Bss}} is parallel to [441]_{\text{Cc}}.

(v) [101]_{\text{Bss}} is parallel to [411]_{\text{Cc}}.

In other words, one has to search for these alignments, having remembered that only adsorption of bassanite on calcite can occur, according to the last row of Table 1. In these cases, (for adsorption alone) the action of screw dislocation cannot be foreseen at the outcropping calcite/bassanite interface: this means that periodic polysynthetic twins cannot be obtained on the growing surfaces (Boistelle & Aquilano, 1977; Aquilano, 1977), although they could easily occur when adsorption/absorption mixes one or more complex interfaces.

The situation shown by 2D-LCs (cases 4a, 4b) is quite interesting. In fact, in both cases, the linear and 2D-area misfit are very low or negligible; the obliquity is nil; the linear misfits are opposite. Finally, and this is amazing, both 2D-LCs are practically quadratic, the directions of their sides being parallel to the cell axes \(a_0\) and \(c_0\) of bassanite. It is not by chance that we suppose [in Fig. 4(b)] the reasonable existence

| Table 2 | 2D-LCs between the [10.4]_{\text{Cc}} and the [010]_{\text{Bss}} pinacoid. |
|---------|---------------------------------------------------------------------------|
| Ranking | [10.4]_{\text{Cc}}, lattice vectors (\(\text{Å}\)) | [010]_{\text{Bss}}, lattice vectors (\(\text{Å}\)) | Maximum linear and area misfit (\(\Delta\%\)) | Obliquity (\(\text{°}\)) | Notes |
| Case 1  | 1/3[441] = 12.85 | [100] = 12.032 | -6.84 | 2D twin law axis [10\(\overline{1}\)]_{\text{Bss}} |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | -2/3[42T] = 16.206 | [10\(\overline{1}\)] = 17.5145 | 7.82 | Opposite misfits |
| Case 2  | 1/3[4.17,1] = 26.231 | [100] = 25.342 | -3.50 | Coherent misfits |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | -2/3[457] = 19.137 | [10\(\overline{1}\)] = 17.5145 | 9.26 | Opposite misfits |
| Case 3a | 1/3[4.19,1] = 25.7 | [120] = 27.196 | 5.82 | Coherent misfits |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | -1/3[4.11,1] = 17.021 | [10\(\overline{1}\)] = 17.5145 | 2.90 | Coherent misfits |
| Case 3b | 1/3[4.19,1] = 35.855 | [3100] = 36.096 | 0.67 | Coherent misfits |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | 1/3[4.11,1] = 17.021 | [10\(\overline{1}\)] = 17.5145 | 2.90 | Coherent misfits |
| Case 4a | 5[010] = 24.948 | [100] = 24.064 | -3.67 | Opposite misfits |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | [42T] = 24.309 | [2001] = 25.342 | 4.25 | Coherent misfits |
| Case 4b | 5[010] = 24.948 | [100] = 24.064 | -3.67 | Opposite misfits |
| 2D cell area (\(\text{Å}^2\)) and multiplicity | [42T] = 24.309 | [2001] = 25.342 | 4.25 | Opposite misfits |
of \(\{10\bar{1}\}\) ledges in bassanite, to avoid the superposition of the original parent embryos.

3.3. The surface structures of the \(\{001\}\) form of bassanite

In cases 1, 2a and 2b described in Table 3, the sides \([010]_{\text{Bss}}\) and \([110]_{\text{Bss}}\) exactly coincide with the sides of the pseudo-hexagonal 2D supercell described in Fig. 3 (left). In Figs. 5(a) and 5(b) the 2D twin law is the same: \([120]\) is the twin axis,

Figure 4

(a) Drawing of case 1 from Table 2. The side (blue) of the parent nucleus, \([100]_{\text{Bss}}\), is parallel to one main side \([441]\) of the calcite substrate, whereas the other side, \([101]_{\text{Bss}}\), is nearly parallel to one of the other main PBCs \([421]_{\text{Cc}}\). The angular misfit between \([101]_{\text{Bss}}\) and \([441]_{\text{Cc}}\) is very low at 1.34°. This means that we are dealing with a new 2D twin law generated by the \([010]_{\text{Bss}}/\{10.4\}_{\text{Cc}}\) epitaxy. This 2D twin law has the \([10\bar{1}]\) axis of bassanite. (b) In case 4a (Table 2), we observe \([100]_{\text{Bss}}/\{010\}_{\text{Cc}}\). (c) In case 4b, we obtain \([2001]_{\text{Bss}}/\{010\}_{\text{Cc}}\). Note, we can obtain two new 2D twin laws for bassanite: twin axes \([100]_{\text{Bss}}\) and \([001]_{\text{Bss}}\) for cases 4a and 4b, respectively. Finally, all the cases illustrated in (b) and (c) show that the linear misfits run in the opposite sense, which would be the ideal situation to calculate \((010)_{\text{Bss}}/(10.4)_{\text{Cc}}\) adhesion energy.

Figure 5

Three examples of 2D-LCs between \(\{001\}_{\text{Bss}}\) and \(\{10.4\}_{\text{Cc}}\). It is coincidental that the interpenetration twins, described in cases 1 and 2b, and drawn in (a) and (c), lead to the same swallow angle, even though they are formed by different bassanite sides in the two cases.
Table 3 2D-LCs between [001]_{Bss} and [10.4]_{Cc}.

The form [001]_{Bss} should have slices of thickness \( d_{200} \), as required by the constraint 00l \( \rightarrow l = 2n \). Cases 4a and 4b are provided in Table S1, as the bassanite sides are not parallel to the (441) sides of calcite.

| Ranking | [10.4]_{Cc} lattice vectors (\( \AA \)) | [001]_{Bss} lattice vectors (\( \AA \)) | Maximum linear and area misfit (\( \Delta \% \)) | Obliquity (\( ^\circ \)) | Notes |
|---------|--------------------------------------|-------------------------------------|---------------------------------------------|----------------|-------|
| Case 1  | 1/3[441] = 12.855 4[010] = 19.958 2[010] = 13.860 1[20] = 18.349 | [2010] = 13.860 [210] = 18.349 | +7.82 | -8.79 | [150]_{Bss} twin axis |
| 2D cell area (\( \AA^2 \)) and multiplicity | 161.733 (4\( \times \)) | 166.583 (2\( \times \)) | 2.99 | 1.85 | Opposite linear misfits |
| Case 2a | 1/3[441] = 12.855 1/3[4.14]1 = 21.540 | [110] = 13.883 5[010] = 20.780 | +7.99 | -3.61 | [150]_{Bss} twin axis |
| 2D cell area (\( \AA^2 \)) and multiplicity | 242.60 (6\( \times \)) | 249.87 (3\( \times \)) | +2.99 | 4.34 | Opposite linear misfits |
| Case 2b | -1/3[441] = 12.855 6[010] = 29.938 | -[110] = 13.883 [140] = 30.208 | +7.99 | +0.90 | [140]_{Bss} |
| 2D cell area (\( \AA^2 \)) and multiplicity | 242.60 (6\( \times \)) | 249.87 (3\( \times \)) | +2.99 | 2.48 | Coherent linear misfits |
| Case (3) | 2/3[441] = 19.032 5[010] = 24.948 | [150] = 18.349 [210] = 25.017 | -3.72 | +0.276 | [210]_{Bss} twin axis |
| 2D cell area (\( \AA^2 \)) and multiplicity | 404.33 (10\( \times \)) | 416.46 (5\( \times \)) | +2.99 | 6.71 | Opposite linear misfits |

Table 4 2D-LCs between [100]_{Bss} and [10.4]_{Cc}.

The [100] slices have to be of thickness \( d_{200} \), as required by the constraint 0k0 \( \rightarrow k = 2n \).

| Ranking | [10.4]_{Cc} lattice vectors (\( \AA \)) | [100]_{Bss} lattice vectors (\( \AA \)) | Maximum linear and area misfit (\( \Delta \% \)) | Obliquity (\( ^\circ \)) | Notes |
|---------|--------------------------------------|-------------------------------------|---------------------------------------------|----------------|-------|
| Case 1  | 1/3[441] = 12.855 4[010] = 19.958 2[010] = 13.860 1[20] = 18.349 | [2010] = 13.860 [210] = 18.778 [220] = 20.780 | +7.82 | -8.79 | [021]_{Bss} twin axis |
| 2D cell area (\( \AA^2 \)) and multiplicity | 161.733 (4\( \times \)) | 175.606 (2\( \times \)) | +8.58 | -3.38 | Opposite linear misfits |
| Case 2  | 1/3[441] = 12.855 1/3[8.13.2] = 22.061 | [001] = 12.67 3[010] = 20.79 | -1.46 | -6.11 | [032]_{Bss} twin axis |
| 2D cell area (\( \AA^2 \)) and multiplicity | 283.015 (7\( \times \)) | 263.409 (3\( \times \)) | -7.44 | 3.65 | Coherent linear misfits |

even if in Fig. 5(a) [010]_{Bss} is parallel to [441]_{Cc}, whereas in Fig. 5(b), [110]_{Bss}=(441)_{Cc}. In Fig. 5(c) the 2D twin law changes too: [140] is the new twin axis. Note that the angle of 78.15° is the same in the ‘swallow tail’ of different laws (upper side left and lower side), because in both cases the twin axis runs parallel to the main [010] calcite PBC.

3.4. The surface structures of the [100] form of bassanite

In both cases of Table 4, the [100] bassanite nucleus has its sides parallel to the most important (441) PBC of the substrate. In the first case [Fig. 6(a)], a new 2D twin axis [021]_{Bss} is obtained. The penetration twin has a swallow angle of 84.91° determined by the (010) directions of the bassanite parent (P) and ‘c’ twinned (T) individuals. The lateral sides of the penetration twin are both parallel to the other sides of the nucleus and coincide with the 2D twin axis [021] of bassanite. Concerning case 2 [Fig. 6(b)], another 2D twin axis [032]_{Bss} works: the angle formed by the ‘c’ equivalent (010) bassanite directions is 101.85°.

4. Conclusions

Starting from the premise cited in the Introduction, it could be assumed that bassanite is a valuable replacement for CaSO₄ to make 2D and/or 3D epitaxy with the (1.04) cleaved form of calcite. By closely observing the interface between bassanite and [10.4]_{Cc} with deeper crystallographic insight, we instead realized in the present work that all the main bassanite forms can produce new 2D twin laws, when in epi-contact with the basic [10.4]_{Cc} rhombohedron. By taking into account the point established in Section 3 on the properties of the interfaces, we can summarize the following interactions between the calcite substrate and the new twin laws determined by bassanite deposition:

(i) [010]_{Bss}: three twin laws were determined, [010]_{Bss} [100]_{Bss} and [001]_{Bss}. In the first 2D-[010]_{Bss} twin law, the swallow-tail angle (92.68°) is formed by the (100)_{Bss} steps and can be attributed to the ‘c’ glide plane that is invariably present in the (10.4) cleaved calcite. In the other two laws ([100]_{Bss} and [001]_{Bss}), the swallow angle that originated between the (101) directions varies between 84.07° (2D-[100]_{Bss} twin law) and 86.78° (2D-[001]_{Bss} twin law). Readers will notice that the maximal variation in these three swallow angles is minimal (<9°) and that visually the triplets look alike, so that they can be easily confused; actually, they differ from each other both in physics and in geometry.

(ii) [001]_{Bss}: first, it is remarkable that in all three main twin laws the bassanite [001] nuclei are always perfectly aligned along the (441)_{Cc} sides. When the alignment is parallel to [010]_{Bss} or [110]_{Bss}, a new twin law arises, with [120]_{Bss} as a new
2D twin axis (cases 1, 2a); instead, in case 2b, where [110]Bss is again aligned along the \(h_{441}\) side, \([1\bar{4}0]_{Bss}\) becomes the new 2D twin axis. Three other cases are given in Table S1 of the supporting information, either because some coincidence has been found with the \(h_{441}\) sides of calcite or because the 2D-Cl area misfit is too high. In two of them, the new 2D twin axes are \([210]_{Bss}\) and \([010]_{Bss}\). Finally, also for this bassanite form, cases 1 and 2b show the same bassanite swallow angle (81.89°/C14), which differs only by a total obliquity of 3.74°/C14 with respect to the theoretical one (calcite), as drawn in Fig. 5 of the text.

(iii) \([100]_{Bss}/[10.4]_{Cc}\) epitaxy. Only the sides of the 2D-[100] cell can be found parallel to the lateral \((4\bar{4}1)_{Cc}\) sides of calcite substrate. In both cases the bassanite twin axes are parallel to \([010]_{Cc}\), one of the main PBCs of the substrate.

84.91°. Once again, this highlights how deceptive first appearances can be when observing ‘swallow-tail twins’.

Accurate research, based on the lattice epi-correspondence between the main bassanite and \((10.4)_{Cc}\) PBCs, allowed us to identify nine unexpected 2D twin laws, generated by the intrinsic \((10.4)_{Cc}\) symmetry coupled with the surface symmetry of bassanite, which increasingly works like a transition compound. Further, swallow-tail twins do theoretically occur in all three cases of the epitaxies of bassanite on \((10.4)_{Cc}\). In brief, these new 2D twin laws have been found for bassanite, promoted by the calcite substrate; together with the \((10.4)_{Cc}/(010)_{Gypsum}\) coupling examined earlier (Aquilano et al., 2022), a new way of thinking is being developed in detail about the epitaxy between different species. This can be particularly useful when a new mineralogical species tends to replace another, as in the case of Ca sulfates (gypsum, bassanite, anhydrite) replacing calcite, in nature and/or the laboratory (Ruiz-Agudo et al., 2016). According to our planning, the next step will be the CaSO₄ anhydrite/(10.4)Cc epitaxy.

Acknowledgements

Open Access Funding provided by Universita degli Studi di Torino within the CRUI-CARE Agreement.

References

Aquilano, D. (1977). J. Cryst. Growth, 37, 215–218.
Aquilano, D., Bruno, M., Cotellucci, A., Pastero, L. & Ghignone, S. (2022). CrystEngComm, 24, 5120–5127.
Aquilano, D., Rubbo, M., Catti, M., Pavese, A. & Ugliengo, P. (1992). J. Cryst. Growth, 125, 519–532.
Ballirano, P., Maras, A., Meloni, S. & Caminiti, R. (2001). Eur. J. Mineral. 13, 985–993.
Becker, A., Söte, I., Paulmann, C., Beckmann, F., Donath, T., Boese, R., Prymak, O., Tiemann, H. & Epple, M. (2005). Dalton Trans. pp. 1545–1550.
Boistelle, R. & Aquilano, D. (1977). Acta Cryst. A33, 642–648.
Hartman, P. (1973). Editor. Crystal Growth: an Introduction. Amsterdam: North-Holland.
Hildyard, R. C., Llana-Fuñe, S., Wheeler, J., Faulkner, D. R. & Prior, D. J. (2011). J. Petrol. 52, 839–856.
Maslyk, M., Dallos, Z., Koziol, M., Seiffert, S., Hieke, T., Petrović, K., Kolb, U., Mondeshki, M. & Tremel, W. (2022). Adv. Funct. Mater. 32, 2111852.
Ruiz-Agudo, E., Álvarez-Lloret, P., Ibáñez-Velasco, A. & Ortega-Huertas, M. (2016). Cryst. Growth Des. 16, 4950–4959.
Ruiz-Agudo, E., Putnis, C. V., Hövelmann, J., Álvarez-Lloret, P., Ibáñez-Velasco, A. & Putnis, A. (2015). Geochim. Cosmochim. Acta, 156, 75–93.
Stawski, T. M., Besselink, R., Chatzipanagis, K., Hövelmann, J., Benning, L. G. & Van Driessche, A. E. S. (2020). J. Phys. Chem. C, 124, 8411–8422.
Tiemann, H., Söte, I., Jarms, G., Paulmann, C., Epple, M. & Hasse, B. (2002). Dalton Trans. pp. 1266–1268.
Van Driessche, A. E. S., Benning, L. G., Rodrigo-Blanco, J. D., Ossorio, M., Bots, P. & García-Ruíz, J. M. (2012). Science, 336, 69–72.
Weiss, H. & Bräu, M. F. (2009). Angew. Chem. Int. Ed. 48, 3520–3524.