A polarization filter for surface wave analysis: statistical approach

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ABSTRACT
Polarization properties of seismic surface waves have been used for the identification and interpretation of small seismic events. In the present method of polarization filtering, normally elliptically polarizing noise is discriminated against specially polarizing surface waves. The statistical random nature of noise is taken into account in the determination of filter weights.

RIASSUNTO
Sono state usate le proprietà di polarizzazione delle onde sismiche superficiali per l'identificazione e l'interpretazione di piccoli eventi sismici. Nella presente tecnica di filtraggio della polarizzazione, il rumore, che generalmente si polarizza ellitticamente, è ben discriminato rispetto alle onde superficiali polarizzatamente. Viene presa in considerazione, per la determinazione dei pesi del filtro, la natura statisticamente casuale del rumore.

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REFERENCES
Nella stesura delle proprietà di polarizzazione delle onde sismiche superficiali sono state coinvolte tecniche avanzate di elaborazione del segnale. Le tecniche di elaborazione del segnale sono state sviluppate dall'Institute for the Earth and Man, Dallas, Texas. Per ulteriori dettagli si rinvia all'originale articolo di D.J. JIN e ai lavori successivi pubblicati in riviste scientifiche specializzate.
Polarization filters have been developed and used in the analysis of seismic data. In this study, polarization properties of seismic waves relative to noise are described and utilized. Characteristics of noise are then incorporated in designing filters to be used simultaneously for the analysis of seismic surface waves. Methods similar to the present one have been used by various authors (Shimshoni and Smith, 1964; Aki, 1965; Flinn, 1966; Griffin, 1966; Simons, 1968; Montalbetti and Kanasewich, 1970; Strauss, 1976). Kanasewich (1973) allocates a chapter of his text book for the discussion of polarization analysis. The difference of the present method from the representative methods used by previous investigators is briefly discussed in the following paragraphs.

In this paper the medium through which surface waves propagate is assumed to be perfectly elastic and only fundamental modes of seismic surface waves are considered.

Shimshoni and Smith's method is for the discrimination of elliptically polarized noise against rectilinearly polarized body waves. Aki's method is for the discrimination of generally elliptically polarized noise against rectilinearly polarized surface waves. Noise is generally elliptically polarized since, due to its nature of randomness, the horizontal component may either lead or lag the vertical component by an arbitrary phase difference and the ellipticity, which is the ratio of semiaxes, is completely random and arbitrary. Rayleigh waves are specially elliptically polarized since the horizontal component may either lead or lag the vertical component, with the latter case being possible on a layered medium, by a certain phase difference of $\pi/2$. They are characterized by their ellipticity of about $2/3$.

The present method has an analogy to the method devised by Aki (1965) in the sense that the phase difference between components of surface waves is exploited, but they are quite different in their nature and objective. Aki uses the phase difference between Love and Rayleigh waves in order to discriminate
the earthquake source phase difference in studying earthquake mechanisms, while in this paper the phase difference between two components of Rayleigh waves is used in order to suppress the coexisting noise.

Flinn (1965) has utilized the rectilinearity and directionality of waves in order to separate compressional wave motion from the shear or surface wave motion, with emphasis on the isolation and identification of the \( pP \) phase. Although the present method is philosophically same as that of Simons (1968), utilization of statistical properties of noise characteristics in designing the filters distinguishes the present method from that of Simons.

The following symbols are employed in the discussion which follow:

- \( V \): vertical component of Rayleigh waves,
- \( T \): transverse component of surface waves or Love waves,
- \( R \): radial component of Rayleigh waves,
- \( V, T, A \): amplitude of \( V, T, R \), respectively,
- \( \phi, \theta, \chi \): phase of \( V, T, R \), respectively,
- \( v, t, a \): real parts of \( V, T, R \), respectively,
- \( v, t, a \): imaginary parts of \( V, T, R \), respectively.

The seismogram from three-component seismometers given:

\[
\begin{align*}
V \exp(i \phi) &= V + i T, \\
T \exp(i \theta) &= T + i T, \\
R \exp(i \chi) &= R + i R.
\end{align*}
\]
where both polar and Cartesian coordinate representations are shown. We set a coordinate system such that:

\[ V' \exp (i (0V - 0v)) = V, \]

\[ T' \exp (i (0T - 0v)) = T + iT', \]

\[ R' \exp (i (0R - 0v)) = R + iR'. \]

where \( V', T', R', R' \) are the new coordinate versions of \( V, T, R, R \) respectively. That is, by taking the difference of \( v \) between the real and imaginary parts of the complex Rayleigh waves becomes purely real but Figs 1a and 1b.

Since the real parts of Eq. (2) are of our utility, we shall consider only the real parts of Eq. (2). Since \( V', T', \) and \( R' \) are all real quantities, they all in phase with each other. Therefore, by definition, the radial component of Rayleigh waves.

![Diagram](image-url)
After rotation of coordinate axes in the complex domain or correspondingly shift of reference or origin time in seismogram, so that Eq. [2] holds, whole of radial component of Rayleigh waves should be purely imaginary so that it is out of phase with $V$ by $\frac{\pi}{2}$ since $\mathbf{T}$ and $\mathbf{R}'$ are, however, horizontal component of surface waves, but not radial component of Rayleigh waves, it should be love waves whether they are desired ones or not.

It is important to note that if there existed only pure signal of surface waves, $\mathbf{R}'$ would vanish as is shown in Fig. 1b. But due to ambient noise, $\mathbf{R}'$ is not usually zero. That is, we have not only the real part of desired love waves but also part of noise which is observed through our seismometers as if it were part of desired love waves. The interpretation in the real language of what is said above is this: Due to the existing noise $\mathbf{R}'$, whose presence is observed in the Fourier component of the
seismogram, the particle motion of Love waves has been deviated from the direction perpendicular to the radial direction. In other words, the propagation direction of Love waves has apparently been changed from its expected direction. Fig. 2 depicts this concept in a schematic manner.

Fig. 2 - Components of the real part of Love waves shown in the real space domain: 1: Radial direction. 2: Transverse direction, which is the expected direction of the particle motion of Love waves. 3: Vertical direction.

1A: Love-wave propagation direction deviated from the radial direction due to noise.
2A: Love-wave particle-motion direction deviated from the transverse direction due to noise.
From Fig. 2, the deviation angle $\beta$ of propagation direction of Love waves from the radial direction, can be computed by

$$\beta = \arctan \left( \frac{R'}{r'} \right)$$

This angle $\beta$ is computed for each frequency component. We want to design a filter, by taking advantage of $\beta$, such that full weight would be assigned to the case where $\beta$ is zero or no noise exists and reduced or null weight as the noise increases. As it is discussed above, the present method utilizes the phase information from all three components of the surface waves. This is another aspect which makes the present method distinguished from that of Shimshoni and Smith (Shimshoni and Smith, 1964, p. 665). Strauss (1976) used similar method to the present one, but he tried to utilize an angle which corresponds to $\beta$ of this paper by shifting the observed transverse component back to the expected direction by rotating the axis by $\beta$. In the present author's opinion, his method adds the observed noise to the actual transverse signal. The present method will give proper adjustment of the observed information according to the amount of the existing noise.

In designing a filter by making use of $\beta$, it should be borne in mind that the noise which yields the angle $\beta$ is inherently random in phase and consequently in its real part, which is the observed magnitude on a seismogram. Therefore an intelligent way in designing a filter which deals with something random may be to find out the statistical probability distribution of $\beta$, which we will do in a later section.

Since $R'$, discussed above is due to noise, we will adopt, simply or consistently, a new symbol $N_{Rr}$ for $R'$, by $N_{Rr}$ implying the real part of $R'$ in the radial direction. One point we wish to note is that as long as there exists the real part of noise in the radial direction, it is reasonable...
to assume the existence of the real part of the noise in the transverse direction. Therefore, if scattered data should be resolved into two components - the real part of desired Love waves, \( L \), and the real part of the noise in that direction, \( N \). The breakdown of the observed Love waves is shown in Fig. 3.

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\[
\phi = \arctan \left( \frac{N_R}{r + N_T} \right)
\]

Radial direction

Transverse direction

Assumption 1: The phase angles, \( \phi_R \), \( \phi_T \), and \( \phi_N \), of the radial, transverse and vertical components of the noise are random and uncorrelated to each other.

Assumption 2: The amplitudes of the radial, transverse and vertical components of the noise are the same. From assumption 2,

\[
N_R = N \cos \phi_R \quad N_T = N \cos \phi_T
\]

\[
N_N = N \cos \phi_N \quad N \cos \phi_N
\]
where \( N_r \) and \( N_t \) represent the amplitudes of the radial and transverse components of the noise and are identical to a quantity \( N \).

From Eqs. [4] and [5],

\[
\beta = \arctan \left( \frac{N_r}{N_t} \right) \quad \text{and} \quad \phi = \arctan \left( \frac{N_r}{N_t} \right).
\]

We assign a pair of random values ranging from zero to \( 2\pi \) to \( \phi \) and \( \beta \) in Eq. [6] for a given signal-to-noise ratio, \( r_j/N \), and obtain a corresponding value for \( \beta \). From the probability distribution function for \( \beta \), a filter for Love waves is designed.

Although the phase difference between the vertical and radial components of the Rayleigh waves should be \( \pi/2 \), we expect that the noise would distort it to some extent, again requiring some form of filtering. We approach this problem with a statistical modeling similar to that for \( \beta \) in Love waves.

We wish to construct a probability distribution function for the apparent phase angle between the vertical and radial components of the Rayleigh waves under the influence of the noise. A summary of the symbols used in this part of the discussion and in Fig. 4 is as follows:

- \( c \): vertical component of desired Rayleigh waves, and it is purely real after rotation of the coordinate axes;
- \( N_v \): vertical component of noise;
- \( N_v \), \( N_v \) : amplitude and phase of \( N_v \);
- \( N_v \), \( N_v \) : real and imaginary parts of \( N_v \).

\( \phi \) : vertical component of apparent Rayleigh waves under the influence of the noise.
— $\psi$: radial component of desired Rayleigh waves, and it is purely imaginary after rotation of the coordinate axes.

— $N_{v}, N_{w}, N_{d}, N_{r}$: radial component counter parts of $N_{v}, N_{w}, N_{d}, N_{r}$ respectively.

— $\gamma$: radial component of apparent Rayleigh waves under the influence of the noise, and has its phase different from that of $\psi$ by $\gamma$.

— $\delta$: the phase difference between the apparent vertical and radial components, i.e., $\psi$ and $\gamma$ of Rayleigh waves.

From Fig. 4,

\[ \sin \left( \frac{\pi}{2} \right) = \sin \left( \frac{\pi}{2} - \gamma \right) \]

Hence,

\[ \tau = \arctan \frac{\sin \delta}{\cos \delta} \]

Similarly

\[ \sin \gamma = \sin \left( \gamma - \frac{\pi}{2} \right) \]

Hence,

\[ \phi = \arctan \frac{\cos \gamma}{\sin \gamma} \]

From the relation

\[ \delta = \pi/2 - \gamma \]

and the two assumptions we made on the characteristics of the noise,

\[ \delta = \pi/2 - \arctan \frac{\cos \gamma}{\sin \gamma} = \arctan \frac{\sin \gamma}{\cos \gamma} \]

\[ \delta = \pi/2 - \gamma \]

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In Eq. (9), we let $N = N_{VI}$ and employed 2/3 for the ellipticity of the elliptic polarization of the Rayleigh waves without losing generality. From Eq. (9), we construct a probability distribution curve for $\theta$ in a similar manner to that for $\delta$, and based on this a filter for the Rayleigh waves is designed.

![Diagram](image)

**Fig. 4** - Apparent phase differences between the vertical and radial components of Rayleigh waves, shown in the complex domain.

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