JUPITER MODELS WITH IMPROVED AB INITIO HYDROGEN EQUATION OF STATE (H-REOS.2)

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ABSTRACT

The amount and distribution of heavy elements in Jupiter gives indications on the process of its formation and evolution. Core mass and metallicity predictions, however, depend on the equations of state (EOSs) used and on model assumptions. We present an improved ab initio hydrogen EOS, H-REOS.2, and compute the internal structure and thermal evolution of Jupiter within the standard three-layer approach. The advance over our previous Jupiter models with H-REOS.1 by Nettelmann et al. is that the new models are also consistent with the observed ≥2 times solar heavy element abundances in Jupiter’s atmosphere. Such models have a rock core mass \( M_\text{c} = 0–8 M_\oplus \), total mass of heavy elements \( M_Z = 28–32 M_\oplus \), a deep internal layer boundary at ≥4 Mbar, and a cooling time of 4.4–5.0 Gyr when assuming homogeneous evolution. We also calculate two-layer models in the manner of Militzer et al. and find a comparable large core of 16–21 \( M_\oplus \), out of which \(~11 M_\oplus \) is helium, but a significantly higher envelope metallicity of 4.5 times solar. According to our preferred three-layer models, neither the characteristic frequency (\( \nu_0 \sim 156 \mu \text{Hz} \)) nor the normalized moment of inertia (\( \lambda \sim 0.276 \)) is sensitive to the core mass but accurate measurements could well help to rule out some classes of models.

Key words: equation of state – planets and satellites: individual (Jupiter)

Online-only material: color figure, machine-readable table

1. INTRODUCTION

Numerous discoveries of unusual hot Jupiters, in particular those that appear to be abnormally inflated (HD 209458b; Baraffe et al. 2003), extremely metal-rich (HAT-P-2b; Leconte et al. 2009), tidally disrupted (WASP-12b; Li et al. 2010), that are thought to have formed by gravitational instability (HR 8799b-e; Baruteau et al. 2011), or that even allow us to study non-equilibrium chemistry in irradiated atmospheres (HD 189733b; Fortney et al. 2010), have drawn some attention away from Jupiter.

While models for exoplanets often span a variety of possible solutions due to the small number of observables characterizing the planet such as radius, mass, age, and their rather large error bars, with Jupiter we face the opposite problem that many models fail to reproduce the available observational constraints (for a review, see Saumon & Guillot 2004). With the upcoming Juno mission, Jupiter will not only be back in focus, but the new data may also invalidate models that cannot cope with the expected accurate measurements for the gravity field, the moment of inertia, or the oxygen abundance in the atmosphere.

As a major source of influence on Jupiter models, Gudkova & Zharkov (1999b) identified the distribution of metals and helium onto certain layers. Saumon & Guillot (2004) demonstrated in addition the sensitivity of the Jupiter models to the equation of state (EOS) of hydrogen. In the past few years, the H EOS in the warm dense matter region where hydrogen metallizes has been improved by applying ab initio simulations (Bonev et al. 2004; Vorberger et al. 2007; Holst et al. 2008; Lorenzen et al. 2010; Morales et al. 2010b, 2010a; Tamblyn & Bonev 2010; Caillabet et al. 2011). The corresponding Jupiter models were consistent with the available observational constraints (Nettelmann et al. 2008, hereafter N08). However, the measured enrichment of Jupiter’s atmosphere with noble gases (∼2 times solar), and with carbon and nitrogen (∼4 times solar) suggests that oxygen is equally abundant, leading to an atmosphere metallicity of two to four times solar. Such high values could not be obtained with the former H EOS (H-REOS.1) used in N08, which instead required solar metallicity atmospheres. This low atmospheric metallicity was confirmed by another group that used the same method of computing the high-pressure EOS data (Militzer et al. 2008). Apart from this similarity however, the Jupiter models suggested by these two groups differed from each other substantially with respect to core mass, heavy element content, and internal layering. In this paper, we present new Jupiter models with ∼2 times solar metallicity atmospheres as a result of applying an improved ab initio H EOS.

In Section 2, we compare the improved H EOS (H-REOS.2) with H-REOS.1 and describe our procedure of structure and evolution modeling. In Section 3, we present the core mass and metallicity of our new Jupiter models and show that the difference in the Jupiter models of Militzer et al. (2008) and our group does not originate from the assumption of two or three layers (Section 3.2). We also present calculations of the moment of inertia (Section 3.3), of the characteristic frequency \( \nu_0 \) of the global oscillations (Section 3.4), and of the thermal evolution (Section 3.5). We discuss Jupiter’s atmospheric metallicity in Section 4 and give our conclusions in Section 5. Appended to this paper are descriptions of our method to calculate the entropy (Appendix A) and the moment of inertia (Appendix B).

2. METHODS

2.1. Improved H EOS

Our new hydrogen EOS (H-REOS.2) is constructed in the same way as the former version H-REOS.1 (N08). The H EOS is assembled from different contributions at the low-density range (\( \rho < 0.1 \text{ g cm}^{-3} \)) and the high-density range (\( 0.2 \text{ g cm}^{-3} \leq \rho \leq 9 \text{ g cm}^{-3} \)). At low densities, pressure dissociation and pressure ionization do not play an important role for the EOS. For these densities, models within the chemical picture agree well with experimental data. As for H-REOS.1, we here apply

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the Fluid Variational theory (FVT+) model (Holst et al. 2007) where ionization is neglected for temperatures below 3000 K.

In the high-density range, our EOS table is obtained from finite-temperature density-functional theory molecular dynamics (FT-DFT-MD) simulations using the code VASP (Kresse & Hafner 1993, 1994; Kresse & Furthmüller 1996). There are three improvements of our new H EOS. First, due to the enormous increase in computing power we were able to perform the simulations with 256 particles in a box compared to 64 particles (Holst et al. 2008). Second, we performed them up to 50,000 K, in order to cover the interior of a warm, young Jupiter, compared to 20,000 K (Holst et al. 2008; N08). In the simulations, a quantum mechanical treatment of the molecular vibrations of the ions is not included. Thus third, after completion of the simulations, we add the energy of a quantum mechanical oscillator, \( \nu_{\text{vib}} = k_B \Theta_{\text{vib}} (0.5 + \exp(\Theta_{\text{vib}}/T) - 1)^{-1} \), to the internal energy of each molecule and subtract its classical energy of vibration, \( k_B T \). The degree of dissociation is estimated via the vibration temperature \( \Theta_{\text{vib}} \) (French & Redmer 2009). For hydrogen, we use the experimental value \( \Theta_{\text{vib}} = 6338.2 \, \text{K} \).

For H-REOS.1, we obtained convergence of the EOS data within 3% and estimated the overall uncertainty due to both statistical and systematic errors to be below 5%. The new ab initio data in H-REOS.2 have an uncertainty of 1% for most of the density–temperature space. They are within the uncertainty of the former data for \( \rho \geq 0.3 \, \text{g cm}^{-3} \) but shifted to 2% higher pressures in a \( T - \rho \) region where pressure dissociation and ionization occurs. Figure 1 shows selected isotherms. With the onset of pressure dissociation at 0.2 g cm\(^{-3}\), the difference rises to 10%. This is a rather small density for FT-DFT-MD simulations that requires a large computational effort to reach converged results. The displayed region of 5000–10,000 K and 0.2–1 g cm\(^{-3}\) is relevant for the Jupiter adiabat. Here, the SCVH-i hydrogen EOS (Saumon et al. 1995) shows systematically higher pressures. At densities of \( 0.1 \, \text{g cm}^{-3} \leq \rho \leq 0.2 \, \text{g cm}^{-3} \), the difference in pressure is a consequence of a revised interpolation between the low-density regime and the data points at 0.2 g cm\(^{-3}\). This difference is particularly pronounced for the internal energy (not shown). Along the Jupiter adiabat (Figure 2), the revised internal energy requires higher temperatures in order to keep the entropy constant at the level defined by Jupiter’s outer boundary condition of 170 K at 1 bar (see Section 2.3), and also leads to a decrease in the density at low pressures of 0.1–10 GPa. However, due to little mass prevalent at those pressures in Jupiter, this region does not affect our Jupiter interior models. Instead, it is the 10–50 GPa pressure region where the lower densities of H-REOS.2 map onto the new Jupiter models.

Our EOS data tables provide the thermal EOS \( P(\rho, T) \) and the caloric EOS \( u(\rho, T) \), but not the specific entropy \( s(\rho, T) \). In Appendix A, we describe our method to derive an adiabat through a given point \( (P, T) \).

2.2. Code Improvements

Careful inspection of our code used for the calculation of Jupiter’s shape and rotation in N08 revealed that two coefficients (one of second order and the other one of third order) contained errors. Their effect on Jupiter models with H-REOS.1 is twofold: (1) the layer boundary shifts from \( \sim 8 \) to \( 4 \) Mbar when the outer envelope is made to have solar metallicity and (2) the metallicity increases by \( \sim 0.02 \) (in absolute value) in the outer envelope and \( \sim 0.035 \) in the inner one so that the maximum total mass of metals increases from 32 to 41 M\(_{\oplus}\), where the latter value was presented in N08. All other results are almost unchanged. The two errors were found by a comparison of the resulting gravitational coefficients \( J_{2n} \), \( n = 1–4 \), of linear and polytropic density Jupiter and Saturn models with analytic solutions given in Zharkov & Trubitsyn (1978), and with numerical solutions for these density distributions by T. Guillot (2008, personal communication). For details, see Nettelmann (2009). Another improvement of the code is the implementation of a Newton-Raphson scheme to find the central conditions that meet the envelope pressure and mass at the core–mantle boundary. For the current version of our code, we here introduce the label MOGROP-11 (MOdellierungsprogramm für GROße Planeten, 2011).

2.3. Interior Structure and Evolution Modeling

Apart from the code improvements mentioned above, our Jupiter interior and thermal evolution models are generated by...
essentially the same code as introduced in N08, Nettelmann (2011), and Fortney & Nettelmann (2010), where in the latter case the code was applied to the structure and evolution calculations of Uranus and Neptune. We here recall the properties of our three-layer structure model for giant planets, review the observational constraints for the new Jupiter models, and describe the method of structure and evolution modeling.

Three-layer structure model. A sufficient number of parameters to meet the observational constraints and boundary conditions are available if we assume a three-layer structure with two envelopes and a core. The envelopes are adiabatic, homogeneous, and consist of hydrogen, helium, and heavy elements, while the core is made of rocks. Heavy elements in the envelopes are represented by the water EOS H$_2$O-REOS (French et al. 2009; N08), for helium we use He-REOS (Kietzmann et al. 2007; N08), and for rocks the formula given in Hubbard & Marley (1989). The outer envelope has a lower He abundance ($Y_1$) than the inner envelope ($Y_2$). We allow for different heavy element mass fractions in the outer envelope ($Z_1$) and in the inner envelope ($Z_2$). The transition pressure between the envelopes $P_{1-2}$ is a free parameter. The density and entropy are discontinuous at layer boundaries.

Observational constraints. The models are required to meet the observational constraints for the total mass $M_J$, equatorial radius at the 1 bar pressure level $R_1$, 1 bar temperature $T_1$, solid-body rotation rate $\omega$, gravitational moments $J_2$, $J_3$, $J_6$, atmospheric helium abundance $Y_{\text{atm}}$ and have the same He mass fraction Y as the protosolar cloud once had. As long as the oxygen abundance below Jupiter’s water cloud deck is not measured, only the lower limit for the heavy element abundance $Z_{\text{atm}}$ in Jupiter’s atmosphere is known. This value is $1 \times Z$ (see Section 4), where $Z_0$ is the protosolar cloud’s metallicity of 1.5% (Lodders 2003).

These observational constraints for the structure models can be divided into four groups depending on whether they enter the modeling procedure explicitly or need to be fitted by adjusting model parameters, and whether or not they are varied within the 1σ error bars. The first set (explicit, varied) is empty. The second set (explicit, not varied) consists of $M_J = 317.8338 M_\oplus$, $2\pi/\omega = 9h55m30s$ (the period of the magnetic field), the upper limit $T_1 = 170$ K, and $Y_{\text{atm}} = 0.238$. While the small uncertainty in $M_J$ and $\omega$ would not affect the resulting models, lowering $T_1$ by 5 K would reduce $Z_1$ by $\sim 0.3$ $Z_0$ (N08), in the opposite direction to what we are aiming for. Modifying $Y_{\text{atm}}$ by 1σ (±0.007) changes $Z_1$ by about ±0.3 $Z_0$. Because $Y_{\text{atm}}$ was measured by the Galileo entry probe at a depth where vertical convection ensures homogeneous distribution of those elements that do not form clouds (such as the noble gases) deeper inside, we set $Y_1 = Y_{\text{atm}}$. The parameters of the third group (to be fitted, not varied) are $R_1 = 7.1492 \times 10^7$ m and $Y = 0.275$, where $R_1$ is fitted by choosing the mean radius $R_m$ of the equipotential surface that coincides with the 1 bar pressure level and Y by adjusting $Y_2$. We find $R_m = 10.955 R_\oplus$ and $Y_2 = 0.285$–0.325. The gravitational moments $J_2$, $J_3$, and $J_6$ build the fourth group (to be fitted, varied). In particular, $Z_2$ is used to reproduce $J_2/10^{-6}$ within the tight observational bounds 14.697(1), and $Z_3$ is used to adjust $J_3/10^{-6}$ to either $-587$ (the observed central value) or $-589$ (the observed lower limit) (Jacobson 2003). When $J_2$ and $J_3$ are matched, $J_6/10^{-6}$ is found to lie within 34.0 and 37.5, consistent with the observed limits (29.1–39.5; Jacobson 2003). Our computation of Jupiter’s thermal evolution makes use of the observed effective temperature $T_{\text{eff}} = 124.4 \pm 0.3$ K.

Modeling procedures. Internal profiles of $m$, $P$, $T$, and $\rho$ which depend on the radial coordinate $l$ are obtained by numerically integrating the structure equations

$$\rho^{-1} dP/dl = -d(V + Q)/dl$$

(1)

and

$$dm/dl = 4\pi l^2 \rho.$$  (2)

Equation (1) states that the force acting on a mass element with density $\rho$ at radius $l$ is balanced by a pressure gradient, where the force $-d(V + Q)/dl$ arises from the gravity field $V$ and the centrifugal force with potential $Q$ due to rigid-body rotation. Equation (2) expresses the amount of mass $dm$ included in a spherical shell of density $\rho$ and thickness $dl$ at radius $l$. These two ordinary, first-order differential equations are integrated inward starting at the surface $l = R_m$ with the two boundary conditions $P(R_m) = 1$ bar and $m(R_m) = M_1$. In order to satisfy in addition the inner boundary condition $m(0) = 0$, a further parameter is needed. For this we choose the core mass $M_\bullet$. Thus, for given values of the observational constraints and of the transition pressure $P_{1-2}$, the result is one single Jupiter model, obtained in terms of the internal profile and values for $Z_1$, $Z_2$, and $M_\bullet$. To calculate the planetary shape and the gravitational moments, we apply the theory of figures (Zharkov & Trubitsyn 1978) up to third order as in N08 and also to fourth order. Such structures describe Jupiter at present time $t_0 = 4.56$ Gyr.

For modeling Jupiter’s thermal evolution, we apply a procedure similar to that of Saumon et al. (1992). To describe Jupiter’s interior at earlier times $t < t_0$, we select a few representative structure models and generate for each of them a sequence of ~80 warmer interior profiles by increasing $T_1$ up to 800 K while $m(P_{1-2})$, $M_\bullet$, and the element abundances are kept at constant values. Further models with intermediate $T_1$ values are then generated by interpolation. While for the structure models at the present time with known $R_1(t_0)$ the core mass was chosen to ensure mass conservation, we here invert the problem to find the planet radius $R_1(t)$ for the given core mass. Unlike Saumon et al. (1992), rotation is included in the approximation of spherical symmetry where the centrifugal force reduces to the zero-order term $-dQ/dl = (2/3)\omega^2 l$. We calculate the cooling curves in the usual approximation of constant angular velocity (Saumon & Guillot 2004; Fortney et al. 2011), and then also by including angular momentum conservation (Hubbard 1970), which requires just one to three additional iterations for a given value of $T_1$, and the corresponding change in the energy of rotation, $dE_{\text{rot}}$. The time $d t$ passed between profiles with different surface temperatures and hence different internal entropies is given by the equation of energy balance,

$$L_{\text{eff}} - L_{\text{eq}} = -dE_{\text{int}}/dt,$$  (3)

where $L_{\text{eff}}$ is the planet’s observable luminosity in the infrared and $L_{\text{eq}}$ is the luminosity the planet would have in equilibrium with the insolation. Using the Stefan–Boltzmann law for the energy emitted by a blackbody with uniform temperature, these luminosities can be used to define the commonly used temperatures $T_{\text{eff}}$, $T_{\text{eq}}$, and $T_{\text{int}}$. This way, $T_{\text{eq}}(t_0)$ is derived from the $L_{\text{eq}}$ value given in Guillot (2005). It is set constant over time or, alternatively, varied with time according to a linear increase of the irradiation flux $F_{\text{eq}} := L_{\text{eq}}/4\pi R_1^2$ starting

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2 Throughout this paper, the label $Y$ denotes an He mass fraction with respect to the H–He subsystem.
with \( F_{\text{eq}}(0) = 0.7 F_{\text{eq}}(t_0) \) as predicted by theoretical standard models for the Sun (Bahcall & Pinsonneault 1995; Sackmann & Boothroyd 2003). Note that the solar standard evolution model predicts an increasing bolometric luminosity with time, whereas the short-wavelength (chromospheric and coronal) emissions of the young Sun may have been up to 1000 times stronger than those of the present Sun as indicated by the measured X-ray and ultraviolet (XUV) fluxes of young solar-like stars (Ribas et al. 2005). Absorbed XUV irradiation can lead to heating of and mass loss from the upper planetary atmosphere (Lammer et al. 2003) but absorption at high altitudes is not suspected to influence the energy balance of the interior (Guillot & Showman 2002).

Therefore, we ignore the activity-related, time-dependent solar XUV flux in our evolution calculations for Jupiter.

The difference between the luminosities radiated away and absorbed equals the loss of the planet’s intrinsic energy \( dE_{\text{int}} \) per time, see Equation (3). In the envelope, the energy lost per mass shell simply is the heat \( \delta q = T(m) ds(m) \). Guillot et al. (1995) offer a convenient closure relation by relating \( T_{\text{eff}} \) to \( T_1 \), namely, \( T_1 = K T_{\text{eff}}^{1.244} g^{-0.467} \) according to the model atmosphere grid by Graboske et al. (1975), where the parameter \( K \) can be used to reproduce the observed \( T_{\text{eff}}(t_0) \) and \( g \) is the surface gravity. Also like Guillot et al. (1995), we take into account the core’s contribution to the planet’s intrinsic luminosity. The heat loss of the core is \( M_c c_v dT_{\text{core}} \), where we use \( c_v = 1 \ J K^{-1} \ g^{-1} \) as in Guillot et al. (1995) and \( dT_{\text{core}} \) is the temperature difference of the isothermal cores of subsequent interior profiles with different \( T_1 \) values, as is \( ds(m) \) the entropy difference at mass shell \( m \) between those interior profiles. Although the abundances of radioactive elements decrease exponentially with time so that the energy production from radioactive decay in rock with meteoric composition may have been an order of magnitude larger in the past than it is today, we here use the present Earth’s value \( L_{\text{radio}} = 2 \times 10^{13} \ J \ s^{-1} \ M_{\oplus}^{-1} \) to account for this contribution. This treatment seems justified for Jupiter because we find the prolongation of the cooling time due to the core’s heat loss to be 0.01 Gyr only. Thus, the cooling equation reads

\[
\frac{dt}{\pi R^2} = -\frac{1}{4} \frac{M_0}{dE_{\text{int}}} \int M_0 \ T ds + M_c c_v dT_{\text{core}} + dE_{\text{rot}}.
\]  

After integrating Equation (4) backward in time, we obtain the cooling time \( t \) that the particular Jupiter model needs to cool down from an arbitrarily hot initial state to the present state.

3. RESULTS

3.1. Core Mass and Metallicity of Three-layer Models

Results for the core mass, the outer envelope metallicity, and the inner envelope metallicity as functions of the depth of the layer boundary between the envelopes are shown in Figure 3. If the layer boundary is located higher in the planet (lower transition pressures), \( Z_1 \) and \( Z_2 \) become smaller. The lighter envelopes then require a larger core mass to conserve the total mass. When \( Z_1 \) decreases down to zero, the core mass cannot grow any further. However, we do not consider models with \( Z_1 < Z_{\text{core}} \) acceptable as this is inconsistent with Jupiter’s observed atmospheric abundances. We find \( Z_2 > Z_1 \) and \( M_c < 10 M_{\oplus} \) for all our acceptable three-layer models. As the gravitational moments are most sensitive to the density in the outer part of the planet at pressures of a few Mbars \( (J_2) \) or even below 1 Mbar \( (J_4) \) and increase with the local density, \( Z_1 \) must decrease when the denser inner envelope extends farther out. Otherwise, \( |J_4| \) would become too large. On the other hand, as we go to higher pressures deeper inside the planet, the sensitivity of \( J_4 \) and later also of \( J_2 \) approaches zero. Therefore, \( Z_1 \) changes only weakly with \( P_{1-2} \) for deep internal layer boundaries. However, since \( J_2 \) is adjusted by \( Z_2 \), \( Z_2 \) must rise strongly with \( P_{1-2} \) in order to provide a sufficiently high mass density that is able to keep \( J_2 \) on the high level of the observed value. With rising metallicity the envelope becomes denser, so the core mass must decrease. When \( M_c = 0 \) is reached, the layer boundary is as deep as it can be, and \( Z_1 \) adopts a maximum value. For models with H-REOS.1, this maximum is 1.0 \( Z_{\text{core}} \) if \( J_4/10^{-5} \) is adjusted to \(-587 \), and 1.5 \( Z_{\text{core}} \) if \( J_4/10^{-6} \) is adjusted to \(-589 \). For models with our improved H EOS, this maximum is 2.5 \( Z_{\text{core}} \) \((J_4/10^{-6} = -587) \) and 2.7 \( Z_{\text{core}} \) \((J_4/10^{-6} = -589) \) respectively. This enhancement in \( Z_1 \) arises...
from the ∼2% higher pressures at given hydrogen mass density (Figure 1), implying ∼2% lower hydrogen mass density at given pressure. The lower partial density of hydrogen in the mixture requires it to be compensated for by a similar amount of metals. Including fourth-order coefficients to the calculations of Jupiter’s shape and rotation systematically shifts the solutions to ∼0.8 \( Z_\oplus \) lower outer envelope metallicities, inducing an additional uncertainty on \( Z_1 \) that is of the same size as the error bars of single observables such as \( T_1 \) and \( J_1 \); see below. For \( M_c \) and \( Z_2 \), the inclusion of higher-order terms has an even greater effect than the improvement in the H EOS. Obviously, a convergence check of the solution with increasing order of the theory is highly important but beyond the scope of this paper. The accuracy of our current computer code is insufficient to reliably study the effect of higher than fourth-order coefficients. From preliminary calculations, solutions with fifth-order coefficients are found to lie between those of third and fourth order.

Figure 3 also shows the observed element abundances in Jupiter’s atmosphere. Assuming that O is as equally enriched as C, these abundances indicate an average enrichment of 2 (noble gases) to 4 (C, N, O). Taking two times solar as the lower limit for \( Z_1 \), models with \( P_{1-2} < 4 \text{ Mbar} \) drop out of the realm of acceptable models, reducing the upper limit for the core mass down to 8 \( M_\oplus \). The total mass of metals is \( M_Z = 29–32 \ M_\oplus \).

As representatives for our acceptable models we recommend models J11-4a and J11-8a, highlighted in Figure 3. Model J11-4a is the one with the outermost possible layer boundary, i.e., at 4 \( \text{Mbar} \). Model J11-8a has \( P_{1-2} = 8 \text{ Mbar} \) for which \( Z_1 \) is \( \geq 2 \ Z_\oplus \) for all of the above considered uncertainties. A machine-readable table for model J11-4a is provided as supplemental material; see Table 1. For these two models, we also vary \( Y, Y_1, T_1 \), and \( J_4 \) within their 1σ error bars. Modifying \( Y \) by \( \Delta Y = \pm 0.05 \) influences \( Y_2 \) in the same direction, and thus leads to a change in \( Z_2 \) of \( \pm 8\% \). Modifying \( Y_1 \) by \( \Delta Y_1 = \pm 0.05 \) causes a response in \( Z_1 \) of \( \pm 10\% \). Similarly, a 5 K colder 1 bar level leads to a 50% decrease in \( Z_1 \). In both cases, \( Z_2 \) consequently changes by \( \pm 3\% – 4\% \) in the opposite direction to \( Z_1 \), and \( M_c \) changes by \( \pm 2\% – 3\% \), again opposite to \( Z_2 \). The colder outer envelope adiabat also cools the interior, resulting in a core that is 100 K colder. The slightly denser H–He subsystem that results reduces the amount of metals in the envelope required to match \( J_2 \), but the net effect of \( \Delta T_1 \) on \( Z_2 \) remains positive. Setting \( |J_4|/10^{-6} \text{ down to } 585 \) lowers \( Z_2 \) by \( \sim 15\% \). Clearly, the response of \( Z_2 \) (increase) and thus of \( M_c \) (decrease) in this case is stronger than in case of \( \Delta T_1 \) and \( \Delta Y_1 \). Including these uncertainties, the total mass of metals changes but insignificantly to 28–32 \( M_\oplus \).

Table 1

| \( m \) (\( M_\oplus \)) | \( P \) (GPa) | \( l \) (\( R_\oplus \)) | \( T \) (K) | \( \rho \) (g cm\(^{-3} \)) | \( s_2 \) | \( s_4 \) | \( s_6 \) |
|-----------------|-----------|---------|------|-------------|-----|-----|-----|
| 317.833802      | 1.0000E–04| 1.000000| 170.0| 1.6970E–04  | –4.4964E–02| 1.9858E2E–03| –1.01995E–04 |
| ...             | ...       | ...     | ...  | ...          | ...  | ...  | ...  |
| 227.158225      | 3.9999E+02| 0.755537| 8995.9| 1.3253E+00  | –3.5329E–02| 1.0943E8E–03| –3.8205E–05 |
| 227.155808      | 4.0001E+02| 0.755532| 8996.0| 1.4332E+00  | –3.5329E–02| 1.0943E7E–03| –3.8204E–05 |
| ...             | ...       | ...     | ...  | ...          | ...  | ...  | ...  |
| 7.567754        | 4.2729E+03| 0.114019| 20041.2| 4.3140E+00  | –1.0080E+02| 8.0544E7E–05| –6.7238E–07 |
| 7.567666        | 4.2729E+03| 0.114518| 20041.2| 1.881E+01   | –1.0080E+02| 8.0544E7E–05| –6.7238E–07 |

Notes. The columns show the internal profile of the Jupiter model J11-4a in terms of the mass coordinate \( m \), the pressure \( P \), the radial coordinate \( l \) scaled by Jupiter’s calculated mean radius \( R_n = 10.95517 \ R_\oplus \), the temperature \( T \), the density \( \rho \), and the figure functions \( s_2 \)–\( s_6 \); see Appendix B. (This table is available in its entirety in a machine-readable form in the online journal. A portion is shown here for guidance regarding its form and content.)

3.2. Core Mass and Metallicity of Two-layer Models

The striking difference between the Militzer et al. (2008, hereafter M08) predictions for the core mass of Jupiter (16 ± 2 \( M_\oplus \)) and that of our group (0–8 \( M_\oplus \)), where both groups applied similar ab initio simulations to generate the EOS data, has prompted speculations about causes for that difference (Militzer & Hubbard 2009). We here investigate how this difference relates to the modeling assumptions of two-layer (M08) or three-layer (N08; this work) structures. For that purpose, we have calculated two-layer models following M08. As in M08, the homogeneous envelope has \( Y_1 = 0.238 \), and the missing helium to obtain an average \( Y \) of 0.275 is not explicitly accounted for but may be part of the core mass. The core is either pure rocks or rock–ice, where we switch from ice to rock at 70 Mbar. As a result, only the innermost 4 \( M_\oplus \) are rocks. Because these models do not have the degree of freedom \( Z_2 \), we can only match either \( J_2 \) or \( J_4 \). As in M08 we choose to match \( J_2 \) by varying the envelope metallicity \( Z \). We apply our usual numerical procedure (see Nettelmann 2011 for details) and vary \( Z \), starting from 0.01, until \( J_2(Z) \) meets the observed value. The core mass must decrease with increasing envelope metallicity (denser envelopes) in order to ensure total mass conservation. If the number of layers assumed for the structure of the model were responsible for the different Jupiter models, we would expect to obtain the same two-layer model as in M08. This is not the case, as illustrated in Figure 4. While the M08 models require 4 ± 2 \( M_\oplus \) heavy elements in the envelope to meet \( J_2 \), corresponding to \( Z = 1.0 ± 0.5 \ Z_\oplus \) for a 16 \( M_\oplus \) core, our two-layer models require \( Z = 4.3 \ Z_\oplus \).

This significantly larger envelope metallicity of our two-layer models enables us to make a transition to three-layer models where metals are shuffled from the outer part of the envelope to the inner part in order to get \( |J_4|/10^{-6} \) down to the observed value (from 614 to 587). The M08 models do not have this degree of freedom as the already low envelope metallicity would become zero long before \( J_4 \) is matched. Hence, the M08 models stay at \( |J_4|/10^{-6} = −614 \) and a large core, while our two-layer models can undergo a transition to three-layer models where \( Z_2 \) becomes smaller and \( Z_2 \) becomes larger than in the homogeneous envelope case. As a consequence of large \( Z_2 \) values, the core mass decreases further. In contrast, the core mass of the
two-layer models is large because this mass also contains the missing mass of helium $dM_{\text{He}} = (0.275-0.238)(M_7 - M_4) \sim 11 M_\oplus$ needed for an overall abundance $Y = 0.275$. Thus, the ice–rock core mass of two-layer models is about 5–10 $M_\oplus$, just at the upper limit of our three-layer models. This estimate would also hold for the two-layer model of M08. However, we cannot reproduce their Jupiter model, the reason for which remains unexplained at the moment. We conclude that the difference between the M08 two-layer models and our three-layer models does not originate from assumptions about the number of layers, but instead is a symptom of already different models within the two-layer frame.

### 3.3. Moment of Inertia

We have calculated the moment of inertia $I$ as described in Appendix B. The non-dimensional form ($\lambda$) is obtained as $\lambda = I M_\odot^{-1} R_\odot^{-2}$. For all of our three-layer models, we find an almost invariant value $\lambda = 0.27605 \pm 0.03\%$. If we ignore Jupiter’s shape deformation, the resulting $\lambda$ decreases by 4% down to 0.26539 $\pm$ 0.03%. This is close to the predictions of 0.2629–0.2645 of Helled et al. (2011), who considered a broad set of interior models with core masses between 0 and 40 $M_\oplus$ as allowed by the use of six-order polynomials to represent the pressure–density relation in Jupiter’s envelope. From our calculations, we conclude that $\lambda$ is not an appropriate parameter to constrain Jupiter’s core mass further. In contrast, Helled et al. (2011) found a significant variation of 0.6% over the full range of their models and still a 0.2% variation when excluding models with $M_c$ larger than 10 $M_\oplus$. Juno’s measurement of Jupiter’s $\lambda$ is expected to have an accuracy of 0.2% (Helled et al. 2011), hence sufficient to distinguish between the different predictions.

### 3.4. Characteristic Frequency

Jovian seismology has long been recognized as a unique opportunity to infer interior structure properties (Vorontsov et al. 1976). In particular, the predicted low-frequency acoustic free oscillations depend sensitively on the core mass and internal layer boundaries of theoretical Jupiter models (Gudkova & Zharkov 1999a). However, the expected small amplitudes ($\sim10$ cm) and Jupiter’s rapid rotation make it difficult to measure acoustic modes. Recently, Gualme et al. (2011) used the SYMPA spectrometer to detect the lowest frequency modes ($\sim1000 \mu$Hz) of Jupiter from spatially resolved radial velocity measurements. In the approximation of low degrees $l$ and overtones of high radial order $n \gg l$, the acoustic modes $\nu_{n,l}$ are proportional to the characteristic frequency $v_0$ (also called equidistance) and are equally spaced by $v_0$ for a given $l$: $\nu_{n,l} \simeq (n+1/2)v_0$. Gualme et al. (2011) observed $v_0 = 155.3 \pm 2.2 \mu$Hz. The inverse of $v_0$ is twice the time a sound wave needs to travel from the center to the surface (Jupiter’s troposphere),

$$v_0 = \left[ 2 \int_0^{R_\odot} dr \ c_s^{-1} \right]^{-1},$$

a useful parameter that interior models can easily be compared to. In Equation (5), $c_s$ is the adiabatic sound velocity. Gudkova & Zharkov (1999a) calculate $v_0 = 152–155 \mu$Hz for SCvH-i EOS-based Jupiter models with five layers and core masses of 3–10 $M_\oplus$. Gualme et al. (2011) obtain the same range of values for $v_0$ with three-layer models of Jupiter that have a small internal He discontinuity and core masses of 0–6 $M_\oplus$. For our three-layer models with $M_c = 0–8 M_\oplus$, we find $v_0 = 155.7–156.3 \mu$Hz. Since this range is rather narrow and slightly above that of former Jupiter models, a more accurate determination of the global mode spacing may help to rule out some Jupiter models but not constrain Jupiter’s core mass further within the current uncertainty of $\sim0–10 M_\oplus$.

### 3.5. Cooling Curves

For a better comparison with published calculations of the cooling of Jupiter, we first assume a constant irradiation flux and constant angular velocity over time. The latter assumption is justified by the rather small change of the rotational energy during evolution compared to Jupiter’s intrinsic energy loss (Hubbard 1977) and the small change in planet radius over most of the contraction timescale (Guillot et al. 1995). Our H-REOS.2-based Jupiter model J11-4a has a calculated cooling time of 4.66 $\pm$ 0.04 Gyr, see Figure 5, and model J11-8a of 4.68 $\pm$ 0.04 Gyr, where the uncertainty mostly arises from the observational error bar of $T_{\text{eq}}$. This result can be considered in good agreement with the age of the solar system ($T_{\odot} = 4.56$ Gyr) given that none of the alternative proposed Jupiter models that rely on free-energy models for the EOS (Saumon & Guillot 2004), such as LM-H4 (4.0 Gyr), SCvH-i (4.7 Gyr), and LM-SCOP (4.8 Gyr), gives a better agreement. On the other hand, Fortney et al. (2011) investigated the uncertainty arising from the use of the Graboske et al. (1975) model atmosphere grid for Jupiter’s evolution. If instead a self-consistent model atmosphere grid was applied as is the common approach for exoplanets, Jupiter’s cooling time increases by $\sim0.5$ Gyr. Sinking He droplets from H–He phase separation in Jupiter would also prolong the cooling time further. A process working in the opposite direction could be rising material from core erosion. During the first 500 Myr, Jupiter has shrunk from...
1.4–1.5 $R_J$ down to 1.1 $R_J$ (Figure 5). Because the initial conditions for the long-term evolution of a 1 $M_J$ planet are forgotten after 0.01 Gyr (Marley et al. 2007), we consider the displayed radius evolution at young ages realistic.

In a second step, we keep the angular momentum conserved to Jupiter’s current value (within a numerical accuracy of 0.4% which gives sufficiently smooth evolution curves) and include the subsequent change of rotational energy $dE_{\text{rot}}(t)$, as also done in Hubbard (1970). Such second-order effects are important for estimates of the size of the correction factors that are necessary to bring the cooling time in agreement with $\tau_\odot$. It is clear that both effects ($\lambda$ and $dE_{\text{rot}}$) will prolong the cooling time: (1) for the same total mass, a larger planet (the young Jupiter) rotates slower, implying a weaker centrifugal force that pushes matter outward, and the consequently smaller radius will allow less energy to be radiated away from the surface and (2) angular momentum conservation ($L = I\omega$) then implies a lower energy of rotation $E_{\text{rot}} = 1/2 L\omega$ at young ages. The increase in $E_{\text{rot}}$ with time must be compensated for by a reduced luminosity, implying again a longer cooling time. For model J11-4a, the first effect increases $\tau$ by 0.1 Gyr, and the second one by 0.2 Gyr, so that we find $\tau = 4.96$ Gyr. In a third step, we consider a time-dependent solar irradiation by a linear approximation of the Sun’s luminosity evolution, which is assumed to have started with 70% of the current value. We find the lower irradiation to speed up the cooling by $\sim$0.6 Gyr, so that we end up with $\tau = 4.41 \pm 0.04$ Gyr. Figure 6 shows the relations between the structure parameter $T_1$ that determines the planet radius, $P = 2\pi/\omega$ and $\lambda$. Their evolution with time can be read from the cooling curve $T_1(t)$.

The resulting cooling times of 4.4 Gyr (linearly increasing insolation) to 5.0 Gyr (constant insolation) suggest a reasonable understanding of Jupiter’s interior and evolution, where remaining uncertainties seem to be attributable to uncertainties in the luminosity evolution of the Sun. This may be counterintuitive as the present Sun is, by means of accurate helioseismology, far better constrained than Jupiter. For instance, uncertainties in the sound velocity profile of the solar standard model are of the order of 0.1% only (Boothroyd & Sackmann 2003). On the other hand, non-standard solar evolution models that predict a bright and more massive young Sun (Sackmann & Boothroyd 2003) cannot completely be ruled out by observationally derived stellar mass-loss rates; see Güdel (2007) for a review.

However, homogeneous, adiabatic evolution models for Saturn and Uranus that are equivalent to those presented here for Jupiter keep failing to reproduce the observed luminosities (Fortney et al. 2011), indicating that the standard three-layer model assumption may be too simplistic for some giant planets. Therefore, Leconte & Chabrier (2012) investigated the possibility of layered convection in Jupiter, where heat is transported inefficiently by diffusion across thin stable layers, separated vertically by convective cells. They determined the maximum superadiabaticity that would give density distributions in agreement with the gravity field data. The deduced presence of roughly $10^4$ diffusive interfaces in Jupiter will qualitatively reduce the heat flux out of the deep interior and shorten the cooling time (Stevenson 1985). Quantitative estimates of (at least) this effect are necessary before one can conclude a reasonable understanding of Jupiter’s evolution.

4. DISCUSSION

According to our new Jupiter models with the improved H EOS (H-REOS.2), metals are enriched by a factor of at most 2.7 in the outer envelope and atmosphere. That value can be increased to 3.0 if the atmospheric He abundance is lowered by $\sigma$, but remains close to the lower limit of the measured C and N abundances. However, Jupiter’s atmospheric metallicity is not directly observable. What can be measured are the abundances of single chemical species from which the atomic ratios with respect to the number of H atoms can then be inferred. Such measurements have been obtained for Jupiter by remote sensing with the Voyager IR and UV spectrometers, IR spectrometry with the Earth Orimer Infrared Space Observatory, Galileo orbiter measurements in the near-IR, the Galileo Probe Mass Spectrometer data, and ground-based remote sensing from IR to radio wavelengths. Among the detected species, the measured abundances of C, N, S, P, Ar, Kr, and Xe (see Figure 3) are
taken to be representative for the convective region below the 1 bar level, the outer boundary of our models, as they are not subject to non-equilibrium chemistry, cloud formation, or other processes that could significantly affect their abundance at such low altitudes (Atreya et al. 2003). Therefore, the enrichment of these elements is assumed to be representative also for those elements that are still subject to such processes, such as oxygen. The Juno orbiter, which is en route to Jupiter, is designed to measure the O abundance below the cloud level. In Figure 7, we vary the O:H ratio and calculate the corresponding metallicity of Jupiter’s atmosphere with the help of Equation (6),

\[ Z_{\text{atm}} = \frac{\sum_i \mu_i (N_i:H)}{\mu_\text{H} + \mu_\text{He}(\text{He:H}) + \sum_i \mu_i (N_i:H)}, \]

where \( \mu_i \) is the atomic weight of species \( i \). By taking into account the measured abundances \( N_i:H \) with their error bars (mean, minimum, and maximum values) and including the non-detected elements \{Mg, Al, Ca\}, because they are abundant in the solar system, with enrichment factors of zero, one, and three times solar, we aim to cover the uncertainty in Jupiter’s metallicity for a given O:H ratio. The unknown abundances of Mg, Al, and Ca contribute an uncertainty to Jupiter’s \( Z_{\text{atm}} \) of \( \pm 0.2 \times Z_\odot \), and the observational error bars of the measured species contribute an additional \( \pm 0.3 \times Z_\odot \).

Assuming Jupiter accreted its volatiles as a result of infall of icy planetesimals from cold, outer regions of the protosolar cloud which resembled the interstellar medium as we observe it today (Owen & Bar-Nun 1995), we would expect O:H to be similarly enriched in Jupiter as C:H (Owen et al. 1999), i.e., O:H \( \gtrsim 3 \times \) solar. For this ratio, we derive from Figure 7 a minimum atmospheric metallicity of 2 \( Z_\odot \) that Jupiter models should satisfy. A \( \sim 2.5 \) times solar metallicity (this work) is consistent with an O:H of 2–4 times solar. A measured O:H ratio greater than 4.5 times solar would not be consistent with the Jupiter models presented here. According to the uncertainties discussed above, the measured O:H ratio at the 19 bar level sets a lower limit to Jupiter’s envelope metallicity of 0.8–1.9 \( Z_\odot \) (based on Figure 7).

The metallicity of the envelope is directly tied to the density of the H–He mixture for fixed \((T, P)\). Our models are always warmer \((T_c \sim 20,000 \text{ K})\) than the M08 models \((T_c \sim 14,000 \text{ K})\), implying a less dense H–He mixture that allows us to add more metals compared to the M08 Jupiter adiabat. This is surprising as nonlinear H–He mixing effects, which have been shown to decrease the density of the mixed phase (Vorberger et al. 2007) compared to the linear mixture that we apply, are already included in the H–He EOS of the M08 Jupiter models. Nonlinear H–He mixing effects remain an appealing possibility (see N08) to enhance \( Z_1 \) in our three-layer Jupiter model calculations further.

We favor the three-layer models over the two-layer models as they allow us to reproduce \( J_4 \) without invoking the hypothesis of deep-zonal winds because the penetration depth of such winds might have to be much deeper (\(< 0.96 \ R_J \)) than physically allowed in the presence of convective flow motions (Liu et al. 2008). If deep-zonal winds indeed exist and require a correction of \( J_4 \) as proposed in M08, the atmospheric metallicity of our Jupiter models would then rise up to 4.5 \( Z_\odot \) and the ice–rock core mass up to 10 \( M_\oplus \).

The three-layer model framework is amenable to additional variations. Fortney & Nettelmann (2010) modified the composition of the core by adding envelope material to the core region, mimicking a diluted core. For extremely diluted cores with rock mass fraction below 20\% in the central region, they found a 50\% enhancement of the resulting atmospheric metallicity of Jupiter. Thus, the relatively low outer envelope metallicities of our models suggest that diluted cores, possibly from core erosion, may be a better assumption for Jupiter than pure rock cores.

5. CONCLUSIONS

We have aimed to push three-layer Jupiter models with our improved H EOS (H-REOS.2) in the direction of the largest possible outer envelope metallicity \( Z_1 \). We find \( Z_1 \lesssim 3.0 \ Z_\odot \), corresponding to an O:H enrichment factor \( \lesssim 5 \). However, we prefer our models with \( Z_1 = 2.0–2.5 \ Z_\odot \), corresponding to O:H = 1–4 \( \odot \) H:solar, because they do not require us to push all constraints to their 1\% uncertainty limits. This increase in the envelope metal content compared to our earlier Jupiter models based on the H-REOS.1 hydrogen EOS arises from about 2\% higher pressures in the \( 0.2–1 \) g cm\(^{-3}\) and 5000–10,000 K region of the H EOS, where the ab initio simulations are challenging since dissociation and ionization occurs. The resulting \( Z_1 \) values in our Jupiter models depend on the depth of the assumed transition to a helium-rich inner envelope. The imposed constraint \( Z_1 > 2 \ Z_\odot \) requires a transition at \( P_{1-2} \geq 4 \) Mbar. This threshold rises to \( P_{1-2} \geq 7 \) Mbar if fourth-order terms in the rotational perturbation of Jupiter’s gravity field and shape are included. Future investigations of the internal structure of Jupiter, and also of Saturn, should include a convergence check of the models with respect to the treatment of rotation.

Our two-layer model calculated with our EOS and our code differs from the Militzer et al. (2008) Jupiter model. We find a larger envelope metallicity (4.5 \( Z_\odot \)) that enables us to switch over to a three-layer model where the envelope is divided into a metal-poor outer and a metal-rich inner part to the expense of the core mass. In a low-metallicity envelope (the M08 Jupiter model), this degree of freedom does not exist. The origin of the differences between these two-layer Jupiter models remains unexplained.
For our three-layer models, we calculate a cooling time of 4.4–5.0 Gyr, an asymptotic frequency spacing \(\nu = 156 \pm 0.03 \mu\text{Hz}\) for global oscillation modes, and a normalized moment of inertia \(\lambda = 0.276 \pm 0.04\%\). These models are consistent with the available observational constraints. The \textit{Juno} mission will be very helpful for further constraining Jupiter’s interior by measuring the deep atmospheric oxygen abundance, higher-order gravitational moments, and the moment of inertia.

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**APPENDIX A**

**METHOD OF CALCULATING THE ENTROPY**

The functions \(P(\rho, T)\) and \(u(\rho, T)\) contain the full thermodynamic information. We first reconstruct the free energy \(F\) and then use the Gibbs–Duhem relation to convert to entropy. It allows us to derive the specific entropy \(s(\rho, T)\) with an offset \(s_0\) with respect to a reference state \((\rho_0, T_0)\). The offset is unknown but constant for a given reference state. First, \(F\) is derived with an unspecified offset \(F_0 = F(\rho_0, T_0)\) by integration of the total differential \(dF/\rho\) along some path in \(\rho\) space. \(F\) is related to the entropy by definition, \(F = U - TS\), where \(U = Mu\) and \(S = Ms\) are the extensive internal energy and the extensive entropy, respectively, and \(M = V\rho\) is the mass contained in the volume \(V\). With the reference state \((T_0, V_0)\) and the unknown entropy offset \(s_0\), we can write

\[
S(T, V) = \frac{U(T, V)}{T} - \left(\frac{F(T, V)}{T} - \frac{F(T_0, V_0)}{T_0}\right) + s_0 M. \tag{A1}
\]

The term in parentheses in Equation (A1) is the solution to the line integral,

\[
\int_{T_0, V_0}^{T, V} dF(T', \rho') \frac{T'}{T}, \tag{A2}
\]

and independent of the chosen path of integration if the EOS data are thermodynamically consistent. We choose lines of constant density and of constant temperature as our path of integration. While this is a choice of convenience, it has the drawback of requiring a large-scale EOS in \(T - \rho\) space off the planetary adiabat. With \(dU = -P\,dV + T\,dS\), we have

\[
\begin{align*}
\frac{d(U - TS)}{dT} &= -\frac{U}{T^2} \, dT, \\
\frac{d(U - TS)}{dV} &= -\frac{P}{T} \, dV. \tag{A3}
\end{align*}
\]

After switching back to \((\rho, T)\)-space, we can write

\[
\frac{1}{M} \int_{\rho_0}^{\rho} \int_{T_0}^{T} dF(T', \rho') \frac{T'}{T} = \int_{\rho_0}^{\rho} d\rho' \frac{1}{\rho'^2} P(T_0, \rho') - \int_{T_0}^{T} dT' \frac{u(T', \rho')}{T'^2}. \tag{A4}
\]

After integrating the right-hand side of Equation (A4), we can numerically calculate the entropy

\[
s(\rho, T) = \frac{u(T, \rho)}{T} - \frac{1}{M} \int_{\rho_0}^{\rho} \int_{T_0}^{T} dF(T', \rho') \frac{T'}{T} + s_0. \tag{A5}
\]

If the underlying EOS is a mixture of different components, the entropy derived from Equation (A5) implicitly includes the mixing terms, for instance the ideal entropy of mixing if the EOS is an ideal mixture. The constant of integration \(s_0\) may depend on composition, but not on the initial condition of the planetary adiabat. If a mixture is kept constant during a planet’s evolution, the offsets \(s_0\) of planetary adiabats through different outer boundaries cancel each other, so that the real entropy difference between two adiabats is then known.

**APPENDIX B**

**MOMENT OF INERTIA**

The moment of inertia \(I\) of a rigid body with spin-angular velocity \(\omega\) and axis of rotation \(\vec{e}_\omega\) through the body’s center of mass is

\[
i = \int_V d^3r \; \rho(\vec{r}) (\vec{e}_\omega \times \vec{r})^2, \tag{B1}
\]

where the integral is taken over the volume \(V\) of the body. If we choose the orientation of the body such that \(\vec{e}_\omega\) runs along the \(z\)-axis, then \((\vec{e}_\omega \times \vec{r})^2 = r^2 \sin^2 \vartheta\) in spherical coordinates. Because for an isolated fluid planet the assumption of hydrostatic equilibrium implies symmetry around the axis of rotation, Equation (B1) becomes

\[
i = 2\pi \int_0^\pi \int_0^R d\vartheta \, dr \, \rho(r, \vartheta) r^4 \sin^3 \vartheta, \tag{B2}
\]

where \(R\) is the planet’s radius. If the planet rotates rapidly such as Jupiter, \(R_{\text{eq}}\) depends on the latitude \(\vartheta\); in particular, the equatorial radius \(R_{\text{eq}} = R(\vartheta = \pi/2)\) is larger than the polar radius \(R(\vartheta = 0)\). Therefore, in order to calculate the two-dimensional integral, the figure of the planet must be known. The figure is the shape at an iso-surface, equivalently to an equipotential surface. For the solar system giant planets, this surface is taken to be the 1 bar pressure level by convention. We have used the \textit{theory of figures} (Zharkov & Trubitsyn 1978) to determine Jupiter’s figure in terms of the figure functions \(s_n(l)\) which give the shape of equipotential surfaces

\[
\mathrm{r}_l(l, \vartheta) = l \left(1 + \sum_{n=0}^{\infty} s_n(l) P_{2n}(\cos \vartheta)\right). \tag{B3}
\]

If the new radial coordinate \(l\) is fixed, the total potential remains unchanged under variation of latitude \(\vartheta\). The functions \(P_{2n}(\cos \vartheta)\) are the Legendre polynomials. Only even expansion coefficients occur because of the assumed symmetry between northern and southern hemisphere. The sum in Equation (B3) is truncated after the third-order term \((n = 3)\) because that is sufficient to calculate the gravitational moment \(J_2\). We abbreviate the right-hand term of the definition \(\mathrm{r}_l(l, \vartheta)\) by \(r_l = l(1 + \sum)\). By inserting Equation (B3) into Equation (B2) and changing the independent variable from \(r\) to \(l\) using

\[
dr/dl = \left[\left(1 + \sum s_{2n} P_{2n}\right) + l \left(\sum_0 \frac{ds_{2n}}{dl} P_{2n}\right)\right], \tag{B4}
\]
and $l(P = 1\text{bar}) = R_m$, the mean radius, we can finally compute $I$, 

\[
I = 2\pi \int_0^\pi d\vartheta \sin^3 \vartheta \int_0^{R_m} dl \left[ (1 + \sum_i + l \left( \frac{ds_{2n}}{dl} P_{2n} \right) \right] \\
\times \rho(l) l^4 (1 + \sum_i)^4.
\]

(B5)

In case of spherical symmetry, $\sum_i = 0$ and $ds_{2n}/dl = 0$.

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