foF2 Seasonal Prediction With IRI-2016 Over Quiet Activity at Dakar Station

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Abstract

This study deals with comparison between Dakar station ionospheric F2 layer critical frequency (foF2) data and both subroutines (CCIR and URSI) of IRI-2016 model predictions. Dakar station is located near the crest of the African Equatorial Ionization Anomaly (EIA) region. Comparisons are made for very quiet activity during the four seasons (spring, summer, autumn and winter) over both solar cycles 21 and 22. The quietest days per season are determined by taking the five days with the lowest aa. The relative standard deviation of modeled foF2 values is used to assess the quality of IRI model prediction. Model predictions are suitable with observed data by day than by night. The accuracy is better during spring season and poor during winter season. During all seasons, both model subroutines don’t express the signature of the observed vertical drift $E \times B$. But they express an intense counter electrojet at the place of mean intensity or high electrojet.

Keywords: quiet geomagnetic activity, ionosphere, electrojet current, IRI-2016 model, seasonal prediction

1. Introduction

Experimental data comparison and model estimation is common in ionosphere study (Anderson, Mendillo, & Herniter 1987; P. K. Bhuyan, Chamua, K. Bhuyan, Subrahmanyam & Garg. 2003; Bilitza 1999; Gulyaeva 2012; Hajra et al. 2016; Oyeyemi, Poole, & McKinnell 2005; Zhang, Xiao, Hao, Ridley & Moldwin. 2011). The International Reference Ionosphere (IRI) is a joint undertaking by the Committee on Space Research (COSPAR) and the International Union of Radio Science (URSI), since the 1960s. This international standard's first version parameters of Earth’s ionosphere dates from 1978 (Rawer, Bilitza, & Ramakrishnan 1978). The IRI model has gradually improved every 5 years due to the new data acquired, but also to the advent of better modeling techniques and its latest version is IRI-2016 (Bilitza et al. 2017). IRI model permits several ionospheric parameters determination (e.g. ionosphere layer critical frequencies, ionosphere total electron content (TEC), bottom side thickness (B0) and bottom side shape (B1), electron density (Ne), etc.) (Bilitza et al. 2017; Bilitza et al. 2014; Sethi & Mahajan 2002).

Several authors have worked on the comparison between observed ionospheric parameters and IRI model predictions (Adewale, Oyeyemi, & Ofuase 2010; Bertoni et al. 2006; Chakraborty, Kumar, De & Guha. 2014; Kumar, Tan, Razul, See & Siingh. 2014; Nanema, Gnabahou, Zoundi & Ouattara. 2018; Nanéma & Ouattara 2013; Ouattara & Nanéma 2014; Sandwidi, Gnabahou & Ouattara 2020, Sethi, Dabas, & Das 2007; Zhang, Paxton, Bilitza & Doe. 2010). Like some of these papers, this study is focus in African Equatorial Ionization Anomaly (EIA) region. It concerns the comparison of the F2 layer critical frequency (foF2) variation as measured at Dakar station (Lat: 14.8°N; Long: 342.6°E) with both IRI-2016 subroutines prediction (CCIR and URSI) over quiet geomagnetic activities. The novelty of this paper is that on one hand it concerns the latest version of the IRI model (IRI-2016) and on another hand the comparison is made on seasonal variation.

The outline of the paper is as follows: section 2 concerns materials and methods; section 3 is devoted to results and discussions, and the conclusion ends the paper as its fourth section.
2. Materials and Methods

2.1 Data Used

Three types of data are used in this article: (1) foF2 values measured at Dakar station. This station operated from 1950 to December 1996. Our study covers the period from 1976 to 1995, i.e. solar cycles 21 and 22; (2) Zurich sunspot number (Rz) for the influence of solar cycle phases and (3) Mayaud (1971; 1972; 1973; 1980) geomagnetic index aa for geomagnetic activity impacts. For this study, foF2 diurnal variation are analyzing by considering seasons impacts.

For comparison with IRI model values we use IRI-2016 version that can be ran through the website https://ccmc.gsfc.nasa.gov/modelweb/models/iri2016_vitmo.php.

2.2 Methods

2.2.1 Geomagnetic Activity Determination

Based on the strong correlation between geomagnetic index aa and solar wind velocity established by Svalgaard (1977), it is well-known, according to Legrand and Simon (1989), Ouattara and Amory-Mazaudier (2009), Ouattara and Zerbo (2011), Zerbo, Amory-Mazaudier, Ouattara, and Richardson. (2012), that disturbed activities are characterized by \( aa > 20 \ nT \) and are classified into three groups: (1) recurrent activities caused by solar high wind stream and Corotating interaction Region (CIRs). (2) shock activities caused by Coronal Mass Ejections (CMEs) and magnetic clouds and (3) fluctuating activities, consequences of solar heliosheet fluctuation.

To better study the impact of quiet geomagnetic activity on the season; very quiet days are not selected according to the reference limit previously defined by Zerbo, Ouattara, Zoundi, and Gyébré (2011) and Gnabahou and Ouattara (2012) but rather with the objective of selecting the quietest days of the considered month. Indeed, for a given month, Mayaud (1980) specifies that the quietest days are obtained by taking \( aa < 10 \ nT \). But here, our quietest days per month are determined by taking the five days with the lowest aa. Daily aa values are obtained from the following website: http://isgi.unistra.fr/data_download.php.

2.2.2 Seasons Determination

There are four seasons: (1) spring (March, April and May); (2) summer (June, July and August); (3) autumn (September, October and November) and (4) winter (December, January and February). But referring to the present paper, the Equinoctial month’s concern March and September and that of solstice are June and December.

The retained days per solar cycle phases over selected months are mentioned in Table 1.

2.2.3 Data Analysis Method

This article aims to study foF2 hourly variations over seasons. To appreciate IRI model predictions, we will compare foF2 values from the URSI and CCIR subroutines of IRI-2016 model with those of Dakar station measured data. For this we will use error bars in the graphs given by the following equation:

\[
\sigma = \sqrt{V}
\]

where \( V \) is the variance defined by \( \frac{1}{N} \sum_{i=1}^{N} (x_i - \bar{x})^2 \) with \( x_i \) the foF2’s hourly average value of considered days, \( \bar{x} \) the average of the foF2’s hourly average values and \( N \) the total number of observations.

Two types of analyzes will be carried out: (1) comparison between foF2 profiles and the five standard profiles established by Faynot and Vila (1979) for African equatorial ionization anomaly (EIA) region and (2) quantitative analysis based on comparison between foF2 observed data profiles with those of both IRI-2016 subroutines.
The five standard profiles established by Faynot and Vila are linked to the nature, the strength or the absence of electric currents in the E layer of ionosphere (Dunford, 1967; Acharya, Roy, Sivaraman, & Dasgupta. 2010; Acharya, Roy, Sivaraman, & Dasgupta, 2011; Vassal, 1982a; Vassal, 1982b). Those are: "Noon bite out" or “B” profile characterized by a double peak (morning and evening) with a trough around midday; "Reversed" or “R” profile characterized by a single peak at evening; "Morning pic” or “M” profile characterized by a single maximum at morning; "Plateau” or “P” profile characterized by an ionization plateau during daytime and "Dome profile” or “D” profile characterized by a single maximum around noon. These five profiles express respectively a high electrojet, an important afternoon conter-electrojet, moderate electrojet, a weak electrojet and absence of electrojet (Vassal 1982a; Vassal 1982b).

Assessment of foF2 measured data prediction’s quality by both IRI-2016’s subroutines will be done by analyzing the relative deviation of foF2 defined by:

$$\sigma_{foF2} = \frac{foF2_{m} - foF2_{exp}}{foF2_{exp}} \times 100$$

(2)

where $foF2_{m}$ and $foF2_{exp}$ are respectively modeled and experimental foF2 values. $\sigma_{foF2}$ is the relative deviation with the following appreciation:

- $\sigma_{foF2} > 10\%$ model overestimates foF2’s measured values;
- $\sigma_{foF2} < -10\%$ model underestimates foF2’s measured values;
- $-10\% < \sigma_{foF2} < 10\%$ model predictions are suitable.

3. Results and Discussion

3.1 Results

Figure 1 shows foF2 diurnal time variation in left column and relative deviation percentage ($\sigma_{foF2}$) in right
column for spring (panel “a”), summer (panel “b”), autumn (panel “c”) and winter (panel “d”) seasons of both solar cycles involved. Observed data from Dakar station are represented by solid line while IRI-2016 predictions are in dotted (for URSI) and in dash lines (for CCIR).

At spring (panel "a") model curves present "R" profiles while experimental data curve presents a "B" profile with a trough observed at local noon and an afternoon peak higher than the morning one. Experimental data afternoon peak is observed at 1700 LT, simultaneously with that of CCIR and one hour after that of URSI. Observed data profile shows the presence of a high electrojet (signature of the $E \times B$ drift) associated with an intense counter-electrojet. But only the latter, characterized by the afternoon peak, is observed for model predictions. Thus, IRI-2016’s both subroutines don’t show the upward vertical drift $E \times B$ in the spring. By day, URSI's curve is above that of CCIR and it's the opposite by night. $\sigmafoF2$’s curves show that the model overestimate experimental data during night for CCIR (1900-0500 LT) and for URSI (from 1900 to 0200 LT). Positive peaks are observed at 2300 LT for URSI (+ 42.4%) and for CCIR (+ 50.2%). URSI’s negative peak (−3.9%) is observed at 1000 LT and that of CCIR, (−5.6%) at 0700 LT. Thus, at spring URSI overestimates and CCIR underestimates the observed data more than CCIR. The previsions are better during the day than at night. Indeed, $\sigmafoF2$ curves present values within ±10% for CCIR (from 0500 to 1900 LT) and for URSI (from 0200 to 2000 LT).

At summer (panel "b") all profiles are type "R". Experimental data curve’s peak is observed at 1600 LT, simultaneously with that of URSI and one hour before that of CCIR. Model’s profiles show the intense counter-electrojet observed in summer. $\sigmafoF2$ curves show that both model subroutines overestimate observed data by night (2000-0500 LT) while they underestimate them in the morning, precisely for CCIR (0600 -0800 LT) and for URSI (from 0600 to 0700 LT). Positive peaks are observed at 0100 LT for URSI (+ 86%) and for CCIR (+ 105.4%). URSI’s negative peak (−13.3%) is observed at 0600 LT and that of CCIR (−26.4%) at 0700 LT. At summer, CCIR underestimates and overestimates observed data more than URSI. Nevertheless, one notes suitable predictions by the day than by night. Indeed, $\sigmafoF2$’s curves present values within ±10% from 0500 to 2000 LT for both subroutines, except from 0600 to 0700 LT for URSI and from 0600 to 0800 LT for CCIR.

At autumn (panel "c") model's curves show "R" profiles and the one of observed data shows a "B" profile with a late trough around 1400 LT and an afternoon peak predominance. This afternoon peak is observed at 1700 LT, simultaneously with that of CCIR and one hour after that of URSI. Experimental data profile highlights an intense counter-electrojet associated with a strong electrojet around 1400 LT while model’s profiles only show an intense counter-electrojet. Thus, IRI-2016 doesn’t reproduce the upward vertical drift $E \times B$ during the fall. $\sigmafoF2$’s curves show that the model overestimate observed data during night, precisely from 1900 to 0500 LT for CCIR and from 2000 to 0500 LT for URSI and also during the day for URSI (1300-1500 LT). Negative peaks are observed at 0700 LT for CCIR (−20.8%) and for URSI (−2.7%) and positive peaks at 2300 LT for CCIR (+ 107.8%) and for URSI (84%). In autumn, CCIR underestimates and overestimates experimental data more than URSI. Nevertheless, the predictions are suitable the day than by night with values of $\sigmafoF2$ within ±10% for URSI (from 0600 to 1300 LT and 1500 to 2000 LT) and for CCIR (from 0500 to 1900 LT, except between 0600 and 0800 LT).

At winter (panel "d") model curves present "R" profiles with an afternoon peak observed at 1700 LT for CCIR and at 1600 LT for URSI while experimental data curve shows an "M" profile associated with two troughs (the first at 1200 LT up to 1300 LT and the second at 1600 LT). Experimental data profile expresses a moderate electrojet’s signature associated with two successive "strong electrojet", the first lasting from 1200 to 1300 LT and the second at 1600 LT. On other hand, model’s predictions show the signature of an intense counter-electrojet. Thus, IRI-2016 doesn’t reproduce the upward vertical drift $E \times B$ during winter. $\sigmafoF2$’s curves show that both model subroutines overestimate observed data by night (from 2100 to 2300)
Figure 1. Seasonal diurnal variations profiles of foF2’s measured data at Dakar station and those of IRI-2016 model (left column). Right column presents the relative deviations percentage curves of both model’s subroutines (CCIR and URSI).

LT, around 0100 LT and from 0300 to 0600 LT) and also by day for URSI (1300-1700 LT) and for CCIR
(1500-1700 LT). URSI’s negative peak (−3.2%) is observed at 0800 LT and its positive peak (+38.9%) at 0600 LT. CCIR’s positive peak (+33.3%) is observed at 2300 LT and its negative peak (−24.7%) at 0700 LT. CCIR underestimates observed data more than URSI while URSI overestimates them more than CCIR. Nevertheless, predictions are suitable the day than by night. $\sigma$foF2’s curves present values within ±10% during night for both subroutines (around 0000 and 0200 LT and from 1800 to 2000 LT) and also by day for URSI (0700-1200 LT) and for CCIR (0900-1400 LT).

3.2 Discussion

Generally, experimental data’s afternoon peak appears simultaneously with that of CCIR at 1700 LT and one hour after that of URSI during equinox seasons (spring and autumn). On other hand, during summer URSI’s afternoon peak appears simultaneously at 1600 LT with that of experimental data and one hour before that of CCIR. But in winter, model curves show an afternoon peak while experimental data curve shows a morning peak. Also, model doesn’t reproduce the upward vertical drift $E \times B$. This result is the same as that found by Ouattara (2013) and Ouattara and Nanema (2014) by comparing experimental data diurnal profiles from Ouagadougou station with those of IRI-2007 and IRI-2012 models predictions respectively.

Quantitative analysis of model predictions proves that predictions are better by day than by night (precisely from morning to 1900 LT for all seasons and around 2000 LT in summer). During all seasons, URSI’s predictions are suitable for a longer time in the day than for CCIR, except in winter. This indicates as already demonstrated by Guibula, Zerbo, Kaboré, Ouattara (2019) at Korhogo station that URSI’s predictions are better than CCIR’s.

These results prove not only the importance to include the Equatorial Ionization Anomaly region ionosphere layer's electrodynamic mechanisms in the IRI model's configuration parameters. Indeed, the upward drift being equatorial zone characteristic phenomenon, it goes without saying that IRI model, parameterized with characteristic phenomena in particular in the mid-latitude zone, cannot reproduce the $E \times B$ drift as well as possible.

In addition, it is well known that at night, the predominant recombination processes in the F2 layer contribute to the drop in electron density; on other hand in high latitude regions this electron loss is compensated by electron precipitation phenomena in auroral zone. These high latitude regions electrodynamic phenomena could explain IRI-2016's poor prediction by night in equatorial region.

4. Conclusion

This paper shows that IRI-2016 model gives good prediction of foF2 variation at Dakar station, especially during day time. Model's peaks match those of observed foF2 data. In addition, model's subroutines don't express the signature of $E \times B$. They show intense counter-electrojet at the place of high or moderate electrojet. Predictions are better during spring and they are poor during winter. It would be important for next studies to quantify the variation of foF2 in relation with electrodynamic phenomena characteristic of low latitudes such as the vertical drift velocity. This will allow us to be able to integrate those low latitudes region phenomena ($E \times B$ drift or electrojet, recombination processes, etc.) in the IRI model configuration parameters to improve its accuracy.

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