Data Article

Dataset for modeling the hydrological patterns and salinity fluctuations of a Mediterranean confined coastal lagoon system (La Pletera salt marsh area, NE Spain)

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\textbf{A R T I C L E  I N F O}

\textbf{Article history:}
Received 27 July 2022
Revised 2 September 2022
Accepted 8 September 2022
Available online 13 September 2022

\textbf{Keywords:}
General lake model
Groundwater surface-water interaction
Constructed coastal lagoon
Water balance
Hydrology

\textbf{A B S T R A C T}

This dataset is related to the research article entitled “Effects of morphology and sediment permeability on coastal lagoons’ hydrological patterns” (W. Meredith, X. Casamitjana, X. D. Quintana, A. Menció) \cite{1}, and was obtained in the La Pletera salt marshes in Catalunya between 2016 and 2019, to model the water balance and salinity fluctuations of 6 permanent lagoons using the General Lake Model (GLM). As no inflow and outflow data were available, water level and bathymetric data were used to calculate the net balance of inflows and outflows according to the observed daily volume fluctuations. Meteorological data were obtained from the L’Estartit Meteorological station north of the lagoons. Daily solar radiation was measured in Mas Badia (La Tallada, \textasciitilde10km from the La Pletera) in 2016 and 2017 and in situ with radiation sensors in 2018 and 2019. Together with the bathymetry and water levels of the lagoons, calculated inflows and calibrated salinity and temperature data are pro-

\DOI{10.1016/j.jhydrol.2022.128259}

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https://doi.org/10.1016/j.dib.2022.108593

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provided to further confined coastal lagoons system modeling where inflow and outflow data are not available. Meteorological data and observed lagoon salinity and temperature are provided for comparison. As this is one of the few datasets that have modeled coastal water bodies less than 3m in depth using the GLM, the data presented here can be useful in stress testing the General Lake Model to other coastal lagoon systems, as well as to other global aquatic ecosystems.

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### Specifications Table

| Subject: | Environmental Science |
| Specific subject area: | Hydrological modelling of coastal lagoons, Groundwater surface-water interaction, Lagoon morphometry, Restoration |
| Type of data: | Table, Figure, Excel File |
| How the data were acquired: | Procedures of collection are described in paragraph 3.1, 3.2 and 3.3 in Methods of Meredith et al., 2022 [1]. |
| Data format: | Raw, Analyzed |
| Description of data collection: | Daily water levels from November 2014 to September 2017 were obtained from Schlumberger water level data loggers (accuracy ± 0.02 m). Water levels in 2018 and 2019 were obtained biweekly from depth gauge boards installed in the lagoons. Biweekly values for temperature and salinity were obtained from a CTD profiler (Sea & Sun Technology). Bathymetric data and water levels were used to calculate lagoon volume and corresponding water inflow quantities. Meteorological data was obtained from the L'Estartit meteorological station, 2 km north of the lagoons. Daily solar radiation was measured in Mas Ràbia (La Tallada, ~10km from the La Pletera) in 2016 and 2017 and in situ with radiation sensors in 2018 and 2019. |
| Data source location: | University of Girona, L’Estartit, Girona, Catalonia, Spain |
| | 42.0513° N, 3.1905° E |
| Data accessibility: | Repository name: Figshare |
| | Direct URL to data: https://doi.org/10.6084/m9.figshare.20198804.v3 |
| Related research article | W. Meredith, X. Casamitjana, X. D. Quintana, A. Menció, Effects of morphology and sediment permeability on coastal lagoons’ hydrological patterns, J. Hydrol. 612 (2022) 128259. https://doi.org/10.1016/j.jhydrol.2022.128259. |

### Value of the Data

- This dataset provides information of the hydrology of a confined coastal lagoon system that is strongly influenced by irregular and unpredictable Mediterranean climatic events.
- This dataset is one of a few studies that have modeled coastal systems at a water depth lower than 3m.
- This dataset provides the basis to anyone who wants to use the General Lake Model in modeling confined coastal lagoons systems where inflow and outflow data are not available.
- The data can be useful in stress testing the General Lake Model to other coastal lagoon systems, as well as to global aquatic ecosystems.
1. Data Description

This dataset was used to model the water balance and salinity fluctuations of the La Pletera salt marshes in Catalunya. It forms the basis of core parameters required by the General Lake Model (GLM) to model water volumes, salinity and temperature profiles of the six lagoons labeled; BPI, FRA, G02, L01, L04, M03. As this is a confined coastal system, it was not possible to measure inflows or outflows and they had to be calculated manually from the fluctuating volume levels of each lagoon. Table 1 is the bathymetry data that calculate the volumes of the lagoons at any single depth using a polynomial fit described in [1]. Table 2 are the GLM physical parameters used in modeling the La Pletera salt marshes. Because of the small volume and shallow depth of the lagoons, the minimum values for volume and thickness of the Lagrangian layers are established one order of magnitude smaller than the typical values. The physical parameters \( C_K, \eta, C_T, C_S \) and \( C_{HYP} \) are related to the individual mixing process efficiencies and

| Height above sea level (m) | Surface Area (m²) |
|---------------------------|-------------------|
|                           | BPI   | FRA   | G02  | L01  | L04  | M03  |
| 1.5                       | 17290 | 2991  |
| 1.25                      |       | 2700  |
| 1                         | 5387  | 6733  |
| 0.9                       |       | 2500  | 8657 |
| 0.8                       |       | 7732  | 5000 |
| 0.7                       | 1383.72| 14478 |
| 0.6                       |       | 2000  | 3492.85| 9878|
| 0.5                       | 435.83 | 13605 |
| 0.4                       |       | 1550  | 9060.63| 2810.47|
| 0.3                       |       | 2684.69| 7372 |
| 0.25                      | 288.58 | 12154 |
| 0.2                       |       | 1500  | 2322.05| 6399.4| 2154.33|
| 0.1                       |       | 1827.29| 4428.9| 1906.65|
| 0                         | 213.31 | 2530.12|
| -0.1                      |       | 1419  | 1512.18| 2950.08| 1630.66|
| -0.2                      | 138.12 |       | 0     | 1606.16| 500 |
| -0.3                      |       |       | 883.58 | 0 |
| -0.4                      |       |       | 585.51 | 0 |
| -0.5                      | 60.12  |       | 1516.16| 500 | 217.5|
| -0.6                      |       | 300   | 217.5 |
| -0.75                     | 1056.78|       |
| -1                        |       | 250   | 0     |
| -1.5                      | 259.39 |       |
| -1.6                      |       | 0     |       |

Table 2
The values of the GLM physical parameters used in modelling the La Pletera salt marshes.

| Mixing and thermodynamic parameters |
|-------------------------------------|
| \( C_K \)                         | Mixing efficiency-convective overturn |
| \( \eta \)                         | Mixing efficiency-wind stirring vs convection |
| \( C_S \)                         | Mixing efficiency-shear production |
| \( C_T \)                         | Mixing efficiency-kinetic requirement |
| \( C_{HYP} \)                     | Mixing efficiency-hypolimnetic mixing |

| Model structure                  |
|----------------------------------|
| Maximum Lagrangian layers        | 200 m³ |
| Minimum layer volume             | 0.025 m³ |
| Minimum layer thickness          | 0.005 m |
| Maximum layer thickness          | 0.05 m |
don’t require calibration; their values are based on observations, experiments in the laboratory and theoretical deliberation [2], and are set to the usual values.

In the excel files “InputDataGLM.xlsx” there are 9 excel sheets:

- The matrices of the first 6 sheets are the input data of inflow, temperature and salinity of each of the lagoons required by the GLM. These have been calibrated to fit the observed volumes, temperatures and salinity. It is important to note these files need to be converted .csv to run in the GLM.
- The seventh sheet is the observed water levels of the lagoons in centimeters above sea level (note that the bottom of the lagoons are below sea level).
- The last two sheets are the biweekly observed salinity and temperature values between 2016 and 2019 of the six lagoons.

The excel file “MeteoData.csv” is the meteorological data of the La Pletera salt marshes during the study period and is required by the GLM. Data includes short wave radiation, cloud cover, air temperature, relative humidity, wind speed and rain. The last column is snowfall and is not applicable in this coastal system. This file is applicable to all the lagoons listed here. The resulting modeled vs. observed values for volume, salinity and temperature are depicted in Figs. 1 –3, respectively. Rainfall between 2016 and 2019 is shown in Fig. 4 and can be used to compare with volume fluctuations of the lagoons. The lagoons also receive surface inflow from seawater inputs during winter cyclonic storms, as well as subterranean water flows during the summer period, when there is usually no rainfall. These inflows are modeled separately from rainfall.
Fig. 2. Modeled vs observed values for salinity of all the lagoons. Salinity is measured in parts per thousand.

Fig. 3. Modeled vs observed values for temperature of all the lagoons.
Fig. 4. Rainfall in the La Pletera salt marsh between 2016 and 2019.

2. Experimental Design, Materials and Methods

The measurements for water levels were taken daily by Schlumberger water level data loggers (accuracy $\pm 0.02$ m) in 2016 and 2017 and biweekly from depth gauge boards installed in the lagoons in 2018 and 2019. Biweekly data for temperature and salinity were measured by a sonde (CTD) profiler (Sea & Sun Technology). The L’Estarit meteorological station, 2 km north of the lagoons, provided daily-average measurements of relative humidity, precipitation, and maximum and minimum temperatures. Daily solar radiation was measured in Mas Badia (La Tallada, $\sim 10$ km from the La Pletera) in 2016 and 2017 and in situ with radiation sensors in 2018 and 2019. Inflow and outflow measurements were estimated from the water levels of the lagoons. As there were no inflow and outflow measurements, the volumes were estimated using a polynomial fit [1], and the net daily inflows and outflows were calculated according to the volume fluctuations observed. The modeled inflows and outflows were then set from the calculated net daily inflow and outflow calculations. Rain and evaporation fluxes are modeled separately from the inflows and outflows. As a result, the inflows and outflows were adjusted until the modeled and real volume, temperature, and salinity values showed the smallest possible differences through an iterative process.

When modeling the water balance, salinity and temperature fluctuations, the General Lake Model incorporates inflows/outflows, mixing, and surface heating and cooling, and calculates vertical profiles of temperature, salinity, and density with a flexible Lagrangian layer structure [3,4]. These layers contract or expand according to inflows, outflows, surface mass fluxes and mixing. From the total daily inflow and outflow data and daily-average meteorological data, the surface momentum, sensible heat, and latent heat fluxes are computed from the following bulk aerodynamic formulae for the stress $\tau$ ($\text{N m}^{-2}$), the sensible heat transfer $H$ ($\text{W m}^{-2}$), and the evaporative heat transfer $E$ ($\text{W m}^{-2}$)

$$\tau = \rho_A C_D U^2 \quad (1)$$

$$H = -\rho_A C_p C_H (T_A - T_S) \quad (2)$$

$$E = -\rho_A L_V C_W U (q_A - q_S) \quad (3)$$

where $\rho_A$ = air density; $U$ = wind speed; $T$ = air temperature; $q$ = specific humidity (all daily averaged); and subscripts A for air and S for water surface values. $C_H$, $C_W$ and $C_D$ are bulk aerodynamic transfer coefficients, and are determined by the height where the data were taken. $C_p$ is the specific heat of water at constant pressure and $L_V$ is the latent heat of evaporation of water. The water mass evaporation can be calculated from $E/L_V$ of the lagoons and is in $\text{kg m}^{-2} \text{s}^{-1}$. 

The short wave radiant flux that is distributed through the water column is calculated using Beer’s law formulation

$$Q(z) = Q_0 e^{-\eta z}$$  \hspace{1cm} (4)

where $Q_0$ = measured radiation at the surface; $Q(z)$ = the intensity at depth $z$, and $\eta$ = the light attenuation coefficient. The constant light attenuation factor was set to 1.7 m$^{-1}$, which is a typical value for eutrophic waters [5]

An integral turbulent kinetic energy model forms the basis for surface layer dynamics [2] and is divided into four discrete processes: wind stirring, convective overturn, interfacial shear production, and Kelvin-Helmholtz billowing. The model calculates the energy available through each of these processes and is a function of the nature of the stratification and the strength of the forcing. This is then compared with the potential energy required to combine the mixed layer with the layer immediately below. The layers are mixed by averaging their properties if sufficient energy is available. This is repeated until not enough energy remains within the present time step to continue the deepening process. This residual energy is then added to the available energy in the next time step. The parameterization for the available turbulent kinetic energy ($KE_A$) is:

$$KE_A = \frac{C_k}{2} \left( w_*^3 + \eta^3 u_*^3 \right) \Delta t + \frac{C_S}{2} \left( u_1^2 + \frac{u_1^2 d\delta}{\delta h} + \frac{u_1 \delta du_1}{3 \delta h} \right) \delta h$$  \hspace{1cm} (6)

and for the required potential energy ($PE_R$) is:

$$PE_R = \frac{C_T}{2} \left[ \left( w_*^3 + \eta^3 u_*^3 \right) \hat{z}^2 + \frac{\Delta \rho gh}{\rho_0} + \frac{g \delta^2}{24 \rho_0} \frac{d \Delta \rho}{dh} + \frac{g \Delta \rho \delta}{12 \rho_0} \frac{d \delta}{dh} \right] \delta h$$  \hspace{1cm} (7)

where $u_*$ and $w_*$ = velocity scales for wind shear and penetrative convection, respectively; $u_1$ = shear velocity at the surface; $\Delta \rho$ = density jump between the surface layer (with depth $h$) and the layer immediately below it (with depth $dh$); $\rho_0$ is a reference density; $\delta$ is the Kelvin-Helmholtz billow thickness scale; $\Delta t$ = the time step; and $g$ = the acceleration due to gravity.

The turbulent diffusivity coefficient, $D_z$, is used to model hypolimnetic mixing. This depends directly on the dissipation of the turbulent kinetic energy and inversely on the stratification. The formula is based on Weinstock (1981), and is given by the expression

$$D_z = \frac{C_{HYP} \varepsilon}{N^2 + \nu^2 \delta^2}$$  \hspace{1cm} (8)

where $k_0$ is the wave number of the largest eddies; $ut^*$ is the turbulent velocity scale; $\varepsilon$ is dissipation; and $C_{HYP}$ is a constant related to the mixing efficiency of the turbulence.

**Ethics Statements**

The authors declare there are no ethical issues with the data presented or methods used.

**Declaration of Competing Interest**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

**Data Availability**

Dataset for modelling the hydrological patterns and salinity fluctuations of a Mediterranean confined coastal lagoon system. (Original Data) (Figshare).
CRediT Author Statement

Warren Meredith: Conceptualization, Formal analysis, Visualization, Writing – original draft; Xavier Casamitjana: Methodology, Conceptualization, Writing – review & editing; Xavier D. Quintana: Conceptualization, Writing – review & editing; Anna Menció: Conceptualization, Supervision, Writing – review & editing.

Acknowledgments

The authors would like to thank Mònica Martinoy and Josep Pascual, who provided the water level, salinity measurements, and meteorological data for the La Pletera lagoons. WM benefited from a Ph.D. grant (IFUdG 2019) from the University of Girona. Financial support was provided by the Life + Program from the European Commission (Life Pletera; LIFE13NAT/ES/001001), the Ministerio de Economía y Competitividad (CGL 2016-76024-R AEI/FEDER/UE) and the Ministerio de Ciencia e Innovacion (PID2021-123860OB-I00) and the European Commission-Horizon 2020 (H2020-LCLLA-2019-2). Part of this research has received funding from the European Union’s Horizon 2020 research and innovation program under grant agreement No 869296 – The PONDERFUL Project.

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