Unusually variable paleocommunity composition in the oldest metazoan fossil assemblages

Seth Finnegan, James G. Gehling, and Mary L. Droser

Abstract.—Recent excavations of Ediacaran assemblages have revealed striking bed-to-bed variation in diversity–abundance structure, offering potential insight into the ecology and taphonomy of these poorly understood early Metazoan ecosystems. Here we compare faunal variability in Ediacaran assemblages to that of younger benthic assemblages, both fossil and modern. We decompose the diversity of local assemblages into within-collection (α) and among-collection (β) components and show that β diversity in Ediacaran assemblages is unusually high relative to younger assemblages. Average between-bed ecological dissimilarities in the Phanerozoic fossil record are comparable to within-habitat dissimilarities typically observed over meter to kilometer scales in modern benthic marine habitats, but dissimilarities in Ediacaran assemblages are comparable to those typically observed over 10–100 km scales in modern habitats. We suggest that the unusually variable diversity–abundance structure of Ediacaran assemblages is due both to their preservation as near snapshots of benthic communities and to original ecological differences, in particular the paucity of motile taxa and the near lack of predation and infaunalization.

Introduction

Ediacara-type fossil assemblages, known from nearly 40 million years before the Cambrian explosion of skeletonized metazoans, include a diverse suite of globally distributed, exceptionally preserved, macroscopic, morphologically disparate soft-bodied organisms that are largely absent from the post-Ediacaran record (Narbonne 2005; Xiao and Laflamme 2009; Droser and Gehling 2015). Many constituents of Ediacara-type assemblages are interpreted as belonging to whole-group metazoa (Droser and Gehling 2015; Bobrovskiy et al. 2018), although there is considerable uncertainty regarding which specific lineages are represented (MacGabhann 2014). The ecological structure of Ediacaran paleocommunities is enigmatic: they show some structural similarities to younger benthic paleocommunities (Clapham et al. 2003; Droser and Gehling 2015; Laflamme and Darroch 2015), but differ in the absence or near absence of predation and infaunalization and the limited number of motile taxa (Droser and Gehling 2015; Chen et al. 2018). Understanding the ecology of Ediacarans is additionally complicated by the unusual mode of preservation of most Ediacaran assemblages as casts and molds in early-cemented sandstones (Tarhan et al. 2016).

One of the best-studied Ediacaran assemblages is found at the Nilpena National Heritage Site, west of the Flinders Ranges in South Australia. Over the past 16 years, 34 beds in the Ediacara Member of the Rawnsley Quartzite have been excavated and mapped, exposing thousands of fossils, including many new taxa (Droser et al. 2017). One of the most striking observations to emerge from this work is that successive beds exhibit substantial variation in diversity–abundance (Droser and Gehling 2015; Bobrovskiy et al. 2018), although there is considerable uncertainty regarding which specific lineages are represented (MacGabhann 2014). The ecological structure of Ediacaran paleocommunities is enigmatic: they show some structural similarities to younger benthic paleocommunities (Clapham et al. 2003; Droser and Gehling 2015; Laflamme and Darroch 2015), but differ in the absence or near absence of predation and infaunalization and the limited number of motile taxa (Droser and Gehling 2015; Chen et al. 2018). Understanding the ecology of Ediacarans is additionally complicated by the unusual mode of preservation of most Ediacaran assemblages as casts and molds in early-cemented sandstones (Tarhan et al. 2016).

One of the best-studied Ediacaran assemblages is found at the Nilpena National Heritage Site, west of the Flinders Ranges in South Australia. Over the past 16 years, 34 beds in the Ediacara Member of the Rawnsley Quartzite have been excavated and mapped, exposing thousands of fossils, including many new taxa (Droser et al. 2017). One of the most striking observations to emerge from this work is that successive beds exhibit little sedimentological variation but substantial variation in diversity–abundance structure (Droser and Gehling 2015; Droser et al. 2017). If this variability is truly unusual in the context of younger assemblages, it might offer insight into the ecology and/or taphonomy of benthic
ecosystems before the advent of widespread motility, predation, and bioturbation in the Phanerozoic.

Here we use a compilation of fossil and modern marine benthic ecological data sets, each consisting of multiple collections from the same local area, time interval, and habitat, to quantitatively compare the variability of Ediacaran faunal assemblages to that of younger fossil and modern benthic assemblages. We examine both the partitioning of diversity into within-collection (\(\alpha\)) and between-collection (\(\beta\)) components and the pairwise dissimilarities among collections within each data set.

Data

**Ediacaran Data Sets.**—The Ediacara Member of the Rawnsley Quartzite, outcropping in the Flinders Ranges of South Australia, is host to a diverse array of taxa comprising a wide variety of morphologies (Droser and Gehling 2015; Droser et al. 2017). The National Heritage Ediacara fossil site at Nilpena Station provides the opportunity to examine successive bedding planes from two fossiliferous facies representing distinct habitats: (1) oscillation-rippled sandstone facies, which is interpreted to have been deposited between fair-weather and storm wave base, and (2) planar-laminated and rip-up sandstone facies, which is interpreted to represent sub–wave base upper canyon fill deposition (Droser et al. 2017; Tarhan et al. 2017). Beds are laterally discontinuous and range from millimeters to about 10 cm in thickness. We included all beds for which at least 50 individual fossils have been counted (23 from the oscillation-rippled sandstone facies and 4 from the planar-laminated sandstones facies). These two data sets include \(\sim\)28 taxa in total (Supplemental Data File 1). All beds are from within a very small geographic region (maximum distance between samples <500 m) and in many cases are directly superposed. Fronds and associated stalks and holdfasts are an important component of the Nilpena assemblages, but because they are commonly shredded and incomplete, it is not always possible to identify them to genus. All fronds are therefore grouped together in a single “genus.” This is conservative, in that it will tend to reduce, rather than inflate, apparent faunal differences between collections. Macrofossil taxa that have been interpreted as algal, such the informally named “Bundle of Filaments” (Xiao et al. 2013) were excluded.

The only other Ediacaran assemblage that has been comparably studied at the bed scale is the substantially older (\(\sim\)570 Ma) and much deeper-water Mistaken Point biota of Newfoundland. Clapham et al. (2003) censused seven fossiliferous bedding planes in the slope-basinal siliciclastic sediments of the Conception and St. John’s Groups. These formations span a significant range of time and of depositional environments (Narbonne et al. 2003; Ichaso et al. 2007). Consequently, to reduce the potential impact of evolutionary turnover and habitat turnover on bed-to-bed faunal variation, for this analysis we include only the four beds from the Mistaken Point Formation.

**Phanerozoic Fossil Data Sets.**—We downloaded Phanerozoic marine body fossil occurrence records from the Paleobiology Database on March 3, 2018 (Supplementary Data File 2; see Supplementary Appendix 1 for download criteria and metadata). In our analyses we included only collections from level-bottom, soft substrate environments (reefs, bioherms, and buildups excluded). We further limited the analysis to collections made for paleoecological studies in which abundance counts were recorded for all reported taxa. To ensure that our data sets were ecologically comparable to the exclusively benthic Ediacaran data sets, we excluded nektonic and planktonic macrofaunal groups such as ammonites; graptolites; and fish and microfossil groups such as conodonts, radiolarians, and foraminifera. We included only collections consisting of at least 50 individuals representing at least two genera and, to ensure that collecting effort was not focused on a single taxon, at least two classes. We grouped collections into assemblage data sets if they were from the same formation and member, the same stratigraphic stage, the same lithofacies and depositional environment, and within 0.01° latitude and 0.01°longitude of one another (maximum distance between collections of \(\sim\)1.6 km). We excluded collections with poorly resolved lithological descriptors (e.g. “indet. siliciclastic”) to ensure that all
collections in a data set came from genuinely similar lithofacies. The final Phanerozoic fossil database consists of 170 local data sets ranging from middle Cambrian to Pliocene in age. The majority include only skeletonized taxa, but two data sets, from the early Cambrian Maotianshan Shale (Zhao et al. 2013; Supplementary Data File 3) and the middle Cambrian Burgess Shale (Caron and Jackson 2008), include diverse soft-bodied taxa. Fossil data sets are globally distributed, but Europe and North America are disproportionately represented (Fig. 1A).

**Modern Data Sets.**—We downloaded benthic data sets from a variety of data sources, with the majority coming from the Ocean Biographic Information System (OBIS) and its regional affiliates (Supplementary Data File 4) and from the U.S. Environmental Protection Agency’s Environmental Monitoring and Assessment (EMAP) program and National Coastal Assessment (Supplementary Data File 5). We included only collections gathered by grab sampling or box-coring of level-bottom, soft-substrate environments in which all or most benthic macrofauna and megafauna were identified at least to genus and counted. We used two criteria, substrate type and depth, to assign collections to habitat types roughly comparable to facies in the sedimentary record. Substrate data are available for all EMAP collections but not for all OBIS collections, so for European data sets we intersected collection coordinates with a benthic substrate map of European waters (Populus et al. 2017). We grouped collections into local data sets if they were from the same source and collected during the same year using the same protocols, from the same substrate type, the same 5 m depth bin, and from within 0.01° latitude and 0.01° longitude of one another. As with the fossil data, we only included data sets consisting of two or more collections with at least 50 individuals in each collection. The final modern database consisted of 265 local data sets, primarily from northern Europe and the United States (Fig. 1B). Whereas in the fossil record there is an almost unavoidable trade-off between increasing geographic scale of analysis and decreasing temporal resolution (Jablonski 2000), modern collections provide the opportunity to examine the relationship between increasing geographic distance and degree of faunal dissimilarity in a single geological instant. Quantifying this relationship while holding depth and substrate type constant can contextualize within-facies bed-to-bed faunal dissimilarities in the fossil record. For these comparisons, modern data sets were defined as above, but without imposing geographic limitations. All data sets used are available for download at the Harvard Dataverse Repository (https://dataverse.harvard.edu/dataset.xhtml?persistentId=doi:10.7910/DVN/JLQRXP).

**Methods**

All data sets were analyzed at the genus level due to inconsistent reporting of species identities in both fossil and modern databases and variation in taxonomic practices across different groups and time intervals. This is conservative, in that it can only act to reduce, rather than inflate, apparent faunal differences among the collections in a data set. Diversity partitioning and pairwise sample dissimilarity analyses were carried out using the ‘vegan’ R package. Because all diversity metrics are strongly influenced by variation in sample size (Beck et al. 2013), we standardized sample sizes by subsampling, without replacement, 50 individuals from each collection. We calculated mean α and multiplicative β diversity (Jost 2007; Baselga 2010) for each subsampled data set with the multipart function, using a Hill’s number (q) of 1 so that diversity estimates are equivalent to linearized Shannon-Wiener diversity and reflect both the evenness and richness components of diversity (Jost 2007). We determined the mean “expected” β diversity of each subsampled data set by taking the mean β diversity from 100 random draws of N samples of 50 individuals from the pooled (γ) distribution, where N is the number of collections in the data set (random draws were done by setting the “nsimul” parameter in multipart to 100). We then calculated “excess” β diversity by subtracting expected from observed β diversity values in each subsampled data set. Because there is considerable debate over the properties of different β-diversity metrics, we also calculated β diversity for each subsampled data set...
as the mean Bray-Curtis dissimilarity between all subsampled collections and the data set centroid. All of the above procedures were repeated 100 times for each data set to place 95% confidence intervals on diversity estimates.

To quantify the relationship between β diversity and geographic distance within modern benthic habitats and compare it with bed-to-bed β diversity in fossil assemblages, we calculated pairwise Bray-Curtis distances among all collections within both modern and fossil data sets. In the modern data sets, we calculated the median dissimilarity between collection pairs in each log2 distance bin. Annotated code for reproducing our analyses is in Supplemental File 1.

Results

As is expected given the absence of soft-bodied taxa in most fossil assemblages, Phanerozoic fossil data sets have significantly lower α diversity than modern data sets (two-sided Wilcox test \( p << 0.001 \), two-sided
Kolmogorov-Smirnov test $p << 0.001$ (Fig. 2). However, excess $\beta$ distributions of fossil and modern data sets do not differ significantly (two-sided Wilcoxon test $p = 0.22$, two-sided Kolmogorov-Smirnov test $p = 0.47$). Ediacaran data sets exhibit relatively low $\alpha$ diversity but high excess $\beta$ compared with both the Phanerozoic and modern fossil data sets (Fig. 2). Ediacaran data sets have $\alpha$ diversities near or below the modal value for Phanerozoic and modern data sets, but excess $\beta$ values that are very high relative to Phanerozoic and modern data sets (97th to 100th percentile). Within the Phanerozoic, there is a trend toward higher $\alpha$ diversity, consistent with previous analyses (Bambach 1977; Bush and Bambach 2004), but no consistent trend in excess $\beta$ (Fig. 3).

Because it is expressed as a multiple of mean $\alpha$ diversity, $\beta$ diversity (and excess $\beta$) is more likely to reach very high values when mean $\alpha$ diversity is low (Veech and Crist 2010), as it generally is for Ediacaran data sets. However, excess $\beta$ values of Ediacaran data sets are high, even relative to Phanerozoic and modern data sets with comparable mean $\alpha$ diversity (Fig. 2). The $\beta$ diversity of a given data set is also partially dependent on the number of collections examined (Marion et al. 2017). The number of collections in Ediacaran data sets ranges from 4 (Nilpena planar-laminated sandstone facies and Mistaken Point Formation) to 19 (Nilpena oscillation-rippled sandstone facies), but excess $\beta$ values in all three cases are high relative to other fossil data sets with comparable numbers of collections (Fig. 4), individuals (Fig. 5), and genera (Fig. 6). This result is not method dependent, as Ediacaran $\beta$ values are also very high when measured as mean Bray-Curtis dissimilarity from the data set centroid (Fig. 7).

We aggregated modern collections into local data sets at the same geographic scale as fossil collections (within $\sim$1 km of each other). However, this choice of spatial scale is somewhat arbitrary, given that the fossil data sets represent far greater time spans. Plotting the distance-decay of faunal similarity within modern habitats, as measured by pairwise Bray-Curtis distance, provides a richer framework for comparing the spatial variation observed in a single time slice with the temporal variation observed in the fossil record (Fig. 8). The distribution of mean pairwise distances between beds in Phanerozoic fossil data sets suggests that they typically capture a range of faunal variation comparable to that seen over meters to kilometers in modern shallow, benthic ecosystems. In contrast, Ediacaran data sets exhibit mean pairwise distances more comparable to those observed over tens to hundreds of kilometers in modern ecosystems.

**Discussion**

Unrecognized environmental or temporal gradients are a potential source of error in our analyses. The Ediacaran data sets were field collected with associated sedimentological data, making it possible to confidently differentiate facies and examine only within-habitat $\beta$ diversity. We were as stringent as possible in limiting habitat variation in fossil and modern data sets.
with the data available, but it is possible that some data sets cover significant gradients in depth and/or substrate type. In addition, although we restricted the temporal coverage of fossil data sets as much as possible given available chronostratigraphic constraints, in some data sets there may still be some temporal turnover in the composition of the species pool. However, this cannot explain the low $\beta$ diversity of Phanerozoic and modern data sets, as

**FIGURE 3.** A, $\alpha$-diversity component of fossil data sets through time. B, Excess $\beta$-diversity component of fossil data sets through time (Nilpena points are overplotted). Confidence intervals as in Fig. 1. E, Ediacaran; Cm, Cambrian; O, Ordovician; S, Silurian; D, Devonian; C, Carboniferous; P, Permian; T, Triassic; J, Jurassic; K, Cretaceous; Pg, Paleogene; Ng, Neogene.

**FIGURE 4.** Excess $\beta$ diversity versus log number of collections in each data set. Confidence intervals and marginal frequency distributions as in Fig. 2.

**FIGURE 5.** Excess $\beta$ diversity versus log number of individuals in each data set. Confidence intervals and marginal frequency distributions as in Fig. 2.

**FIGURE 6.** Excess $\beta$ diversity versus log number of genera in each data set. Confidence intervals and marginal frequency distributions as in Fig. 2.
Environmental or temporal gradients can only inflate apparent \( \beta \) diversity by mixing samples from different habitats or time intervals. Ediacaran organisms are less well understood than younger taxa, and so it is possible that they are taxonomically undersplit, but this would tend to reduce Ediacaran \( \beta \)-diversity estimates rather than inflate them.

Accepting that the pattern is genuine and not an artifact of how collections are aggregated into data sets, three non–mutually exclusive hypotheses can be formulated to explain the exceptional bed-to-bed variability of Ediacaran assemblages. First, it is possible that Ediacaran assemblages represent different depositional environments, with different gradients and disturbance dynamics, than other assemblages. Second, Ediacaran assemblages may be less spatiotemporally averaged than younger fossil assemblages due to the absence of skeletonization and bioturbation. Finally, Ediacara-type faunas may have been genuinely more variable in space and time than younger benthic ecosystems.

Available data provide mixed support for the first hypothesis. Inferred depositional environments of Ediacaran data sets analyzed here range from nearshore (Nilpena) to outer slope (Mistaken Point), and nothing in the sedimentology of these units suggests that they represent unusual environments. Although it has been suggested that the Ediacara Member includes a range of terrestrial environments (Retallack 2013), this interpretation has been widely rejected (Callow et al. 2013; Xiao et al. 2013; Tarhan et al. 2015a). The Ediacara Member may represent a lagoonal or estuarine setting, but Ediacaran data sets exhibit elevated variability.
even when compared with fossil data sets from lagoonal and estuarine settings (Fig. 9). Dissolved oxygen levels in shallow waters were probably lower in the Ediacaran than during most of the Phanerozoic (Lyons et al. 2014; Sperling et al. 2015), but the typically low faunal variability of modern low-oxygen settings (Levin 2003) suggests that low oxygen is unlikely to explain the unusual variability of Ediacaran assemblages. However, there is growing evidence for strongly fluctuating redox conditions in some Ediacaran-aged successions (Bowyer et al. 2017), and such fluctuations would be expected to generate highly variable faunal assemblages within a single facies.

Mode of preservation and degree of time averaging are almost certainly an important component of the exceptional variability of Ediacaran assemblages. Phanerozoic data sets consist primarily of shelly fossils in bioturbated sediments and are consequently relatively time-averaged (Kidwell and Daniel 1991). Because they predate the advent of extensive bioturbation, Ediacaran assemblages are expected to have higher temporal resolution. Time averaging of shelly fossils acts to transfer diversity from \( \beta \) to \( \alpha \) levels (Finnegan and Droser 2008; Tomašových and Kidwell 2009), potentially obscuring the original patchiness of benthic communities. The faunal variability of Ediacaran assemblages compared with that of the Phanerozoic fossil record may thus be, at least in part, a consequence of their high temporal resolution—they may offer snapshots of larval settlement patterns (Droser and Gehling 2008; Mitchell et al. 2015) and successional dynamics (Clapham et al. 2003; Coutts et al. 2016; Reid et al. 2017) that occur on timescales too short to be captured by the time-averaged Phanerozoic shelly record.

Some support for this hypothesis comes from comparison with the Cambrian Burgess Shale and Maotianshan Shale, the only other soft-bodied Lagerstätte for which sufficient paleoecological data are available (Caron and Jackson 2008; Zhao et al. 2013). Both of these data sets exhibit relatively high \( \beta \) diversity (Fig. 10). Although direct comparison is difficult, because the Cambrian assemblages formed in relatively deep-water habitats, the generally high \( \beta \) diversity of Ediacaran and Cambrian Lagerstätten suggests that data sets with high temporal resolution do indeed tend to preserve greater bed-to-bed variation in diversity–abundance structure. Thus, in addition to providing unusual windows into organismal morphology and paleocommunity composition, some Lagerstätten may also preserve exceptional records of paleoecological dynamics. We also note that although there is little trend in the mean excess \( \beta \) value of Phanerozoic shelly fossil data sets, some of the highest observed values occur in Ordovician–Silurian data sets (Fig. 3). This is consistent with suggestions that, because they precede the post-Silurian advent of modern bioturbation intensities (Tarhan et al. 2015b), some early Paleozoic benthic shelly fossil assemblages may preserve a higher-resolution paleoecological record than younger assemblages (Bambach et al. 1991).

It is likely that the unusual variability of Ediacaran assemblages also reflects some key ecological differences compared with younger benthic paleocommunities. The most important of these may be the rarity of motile taxa in Ediacaran assemblages, which would be expected to lead to patchier communities strongly influenced by dispersal limitations. One of the few motile taxa in the Nilpena

**Figure 9.** Mean \( \alpha \) versus excess \( \beta \) diversity for all fossil data sets broken out by broad depth zone. Confidence intervals and marginal frequency distributions as in Fig. 2. Points correspond to Ediacaran and Phanerozoic points in Fig. 2.
assemblages, *Dickinsonia*, occurs on nearly all beds in both facies (Droser et al. 2017). Only the holdfast form genus *Aspidella* shares a similar distinction (Droser et al. 2017), and it is unknown how many different frond taxa *Aspidella* might in fact represent (Gehling et al. 2000).

Another factor that may help to explain the high faunal patchiness of Ediacaran assemblages is the inferred presence of multiple reproductive and dispersal modes. A number of Ediacaran taxa exhibit evidence of planktonic dispersal (Clapham et al. 2003; Droser and Gehling 2008; Darroch et al. 2013; Mitchell et al. 2015), but some may also have engaged in asexual clonal reproduction via stolons (Mitchell et al. 2015), a trait that would lead to patchy species distributions. Patchiness may also reflect successful colonization related to organism height and dispersal (Mitchell and Kenchington 2018).

Finally, competitive interactions may have had a stronger influence on paleocommunity composition in the Ediacaran than in younger ecosystems. Relative abundance distributions suggest that at least some Ediacaran assemblages included taxa that competed with one another for resources (Darroch et al. 2018). Ediacaran ecosystems also lacked widespread predation and bioturbation—processes that exert a powerful influence on the structure of younger benthic ecosystems (Erwin and Tweedt 2012) and may act to reduce the importance of competition (Stanley 2008). In their absence, then, competitive interactions could have had a stronger influence on community structure and composition. Competition-dominated ecosystems are generally predicted to be characterized by low α diversity and high β diversity (Hautmann 2014), a prediction that is consistent with the patterns we observe.

**Conclusions**

Here we have demonstrated that, compared with the Phanerozoic fossil record, Ediacaran assemblages are characterized by unusual bed-to-bed variability in faunal diversity and abundance. Comparisons between fossil and modern benthic assemblages are complicated by the limited spatial scale of the former and the limited temporal scale of the latter, but whereas the faunal variability of Phanerozoic shelly fossil data sets is consistent with relatively local-scale variation in modern benthic habitats, Ediacaran assemblages exhibit variability typically observed only over much greater distances in modern habitats. Additional high-resolution field studies of Ediacaran and Phanerozoic fossil assemblages and modeling exercises are required to determine whether this unusual variability primarily reflects differences in taphonomy, community ecology, or a combination of the two.

**Acknowledgments**

We thank R. and J. Fargher for access to the National Heritage Ediacara fossil site Nilpena on their property, acknowledging that this land lies within the Adnyamathanha Traditional Lands. We thank M. Clapham, L. Tarhan, and A. Tomášových for helpful discussions, and M. Laflamme and an anonymous reviewer for constructive reviews. This work was funded by the NASA Exobiology Program (NASA grant NNX14AJ86G to M.L.D.); the Australian Research Council (grant DP0453393 to J.G.G.); the National Geographic Society (grant 9476-14 to M.L.D.); the Ediacaran Foundation; the South Australian Museum Waterhouse Club;
the University of California, Riverside; Beach Energy (M.L.D. and J.G.G.); and a David and Lucile Packard Fellowship (S.F.). Some of the data described in this article were produced by the U.S. Environmental Protection Agency through its Environmental Monitoring and Assessment Program (EMAP), http://www.epa.gov/emap. Information contained here has been derived from data made available under the European Marine Observation Data Network (EMODNet) Seabed Habitats project (www.emodnet-seabedhabitats.eu), funded by the European Commission’s Directorate-General for Maritime Affairs and Fisheries (DG MARE). This is Palaeobiology Database contribution number 327.

**Literature Cited**

Bambach, R. K. 1977. Species richness in marine benthic habitats through the Phanerozoic. Paleobiology 3:152–167.

Bambach, R. K., M. L. Droser, and J. J. Sepkoski Jr. 1991. Secular changes in Phanerozoic event bedding and the biological overprint. Pp. 299–312 in G. Einsele, W. Ricken, and A. Seilacher, eds. Cycles and events in stratigraphy. Springer-Verlag, Berlin.

Baselga, A. 2010. Multiplicative partition of true diversity yields independent alpha and beta components; additive partition does not. Ecology 91:1974–1981.

Beck, J., J. D. Holloway, and W. Schwanghart. 2013. Undersampling and the measurement of beta diversity. Methods in Ecology and Evolution 4:370–382.

Bobrovskiy, I., J. M. Hope, A. Ivantsov, B. J. Nettserhems, C. Hallmann, and J. J. Brooks. 2018. Ancient steroids establish the Ediacaran fossil *Dickinsonia* as one of the earliest animals. Science 361:1246–1249.

Bowyer, F., R. A. Wood, and S. W. Poulton. 2017. Controls on the evolution of Ediacaran metazoan ecosystems: a redox perspective. Geobiology 15:516–551.

Bush, A. M., and R. K. Bambach. 2004. Did alpha diversity increase during the Phanerozoic? Lifting the veils of taphonomic, latitudinal, and environmental biases. Journal of Geology 112:625–642.

Callow, R. H. T., M. D. Brasier, and D. McNy. 2013. Discussion: “Were the Ediacaran siliciclastics of South Australia coastal or deep marine?” by Retallack et al., *Sedimentology*, 59, 1208–1236. *Sedimentology* 60:624–627.

Caron, J.-B., and D. A. Jackson. 2008. Paleocology of the greater phyllopod bed community, Burgess Shale. Palaeogeography, Palaeoclimatology, Palaeoecology 258:222–256.

Chen, Z., X. Chen, C. Zhou, X. Yuan, and S. Xiao. 2018. Late Ediacaran trackways produced by bilaterian animals with paired appendages. Science Advances 4eaao6691.

Clapham, M. E., G. M. Narbonne, and J. G. Gehling. 2003. Paleocology of the oldest known animal communities: Ediacaran assemblages at Mistaken Point, Newfoundland. Palaeobiology 29:527–544.

Coifs, F., J. J. Gehling, and D. C. García-Bellido. 2016. How diverse were early animal communities? An example from Ediacara Conservation Park, Flinders Ranges, South Australia. Alcheringa: An Australasian Journal of Palaeontology 40:407–421.

Darroch, S. A. F., M. Laflamme, and M. E. Clapham. 2013. Population structure of the oldest known macroscopic communities from Mistaken Point, Newfoundland. Palaeobiology 39:591–608.

Darroch, S. A. F., M. Laflamme, and P. J. Wagner. 2018. High ecological complexity in benthic Ediacaran communities. Nature Ecology and Evolution 2:1541–1547.

Droser, M. L., and J. G. Gehling. 2008. Synchronous aggregate growth in an abundant new Ediacaran tubular organism. Science 319:1660–1662.

Droser, M. L., J. G. Gehling, L. G. Tarhan, S. D. Evans, C. M. S. Hall, J. V. Hughes, E. B. Hughes, M. E. Dzaugis, M. P. Dzaugis, P. W. Dzaugis, and D. Rice. 2017. Piecing together the puzzle of the Ediacara Biota: excavation and reconstruction at the Ediacara National Heritage site Nilpena (South Australia). Palaeogeography, Palaeoclimatology, Palaeoecology 513:132–145.

Erwin, D. H., and S. Tweedt. 2012. Ecological drivers of the Ediacaran–Cambrian diversification of Metazoa. Evolutionary Ecology 26:417–433.

Finnegan, S., and M. L. Droser. 2008. Reworking diversity: effects of storm deposition on evenness and sampled richness, Ordovician of the Basin And Range, Utah and Nevada, USA. Palaeo. 25:87–96.

Gehling, J. G., G. M. Narbonne, and M. M. Anderson. 2000. The first named Ediacaran body fossil, *Aspidella terranova*. Palaeontology 43:427–442.

Hautmann, M. 2014. Diversification and diversity partitioning. Paleobiology 40:162–176.

Ichaso, A., R. Dalrymple, and G. M. Narbonne. 2007. Paleoenvironmental and basin analysis of the late Neoproterozoic (Ediacaran) upper Conception and St. John’s groups, west Conception Bay, Newfoundland. Canadian Journal of Earth Sciences 44:25–41.

Jablonski, D. 2000. Micro- and macroevolution: scale and hierarchy in evolutionary biology and paleobiology. Paleobiology 26:15–52.

Jost, L. 2007. Partitioning diversity into independent alpha and beta components. Ecology 88:2427–2439.

Kidwell, S. M. B., and W. J. Daniel. 1991. Taphonomy and time-averaging of marine shale faunas. Pp. 115–209 in P. A. Allison and D. E. G. Briggs, eds. Taphonomy: releasing the data locked in the fossil record. Plenum, New York.

Laflamme, M., and S. A. F. Darroch. 2015. Palaeobiology: ecological revelations in Ediacaran reproduction. Current Biology 25: R1047–R1050.

Levin, L. A. 2003. Oxygen minimum zone benthos: adaptation and community response to hypoxia. Oceanography and Marine Biology 41:1–45.

Lyons, T. W., C. T. Reinhard, and N. J. Planavsky. 2014. The rise of oxygen in Earth’s early ocean and atmosphere. Nature 506:307–315.

MacGabhann, B. A. 2014. There is no such thing as the “Ediacara Biota.” Geoscience Frontiers 5:53–62.

Marion, Z. H., J. A. Fordyce, and B. M. Fitzpatrick. 2017. Pairwise beta diversity resolves an underappreciated source of confusion in calculating species turnover. Ecology 98:933–939.

Mitchell, E. G., and C. G. Kenchington. 2018. The utility of height for the Ediacaran organisms of Mistaken Point. Nature Ecology and Evolution 2:1218–1222.

Mitchell, E. G., C. G. Kenchington, A. G. Liu, J. J. Matthews, and N. J. Butterfield. 2015. Reconstructing the reproductive mode of an Ediacaran macro-organism. Nature 524:343–346.

Narbonne, G. M. 2005. The Ediacara Biota: Neoproterozoic origin of animals and their ecosystems. Annual Review of Earth and Planetary Sciences 33:421–442.

Narbonne, G. M., R. W. Dalrymple, D. A. Wood, J. G. Gehling, and M. E. Clapham. 2003. Paleoenvironmental analysis of the late Neoproterozoic Mistaken Point and Trepassay formations, south-eastern Newfoundland. Canadian Journal of Earth Sciences 40:1375–1391.
Populus, J., M. Vasquez, J. Albrecht, S. Agnesi, Z. Al Hamdani, J. Andersen, A. Annunziatellis, T. Bekkby, A. Bruschi, V. Doncheva, V. Drakopoulou, G. Duncan, R. Inghilse, C. Kyriakidou, F. Lalli, H. Lillis, G. Mo, M. Muresan, M. Salomidi, D. Sakellariou, M. Simboura, A. Teaca, D. Tezcan, V. Todorova, and L. Tunesi. 2017. EUSeaMap. A European broad-scale seabed habitat map. http://archimer.ifremer.fr/doc/00388/49975, accessed 13 March 2018.

Reid, L. M., D. C. García-Bellido, J. L. Payne, B. Runnegar, and J. G. Gehling. 2017. Possible evidence of primary succession in a juvenile-dominated Ediacara fossil surface from the Flinders Ranges, South Australia. Palaeogeography, Palaeoclimatology, Palaeoecology 476:68–76.

Retallack, G. J. 2013. Ediacaran life on land. Nature 493:89–92.

Sperling, E. A., C. J. Wolock, A. S. Morgan, B. C. Gill, M. Kunzmann, G. P. Halverson, F. A. Macdonald, A. H. Knoll, and D. T. Johnston. 2015. Statistical analysis of iron geochemical data suggests limited late Proterozoic oxygenation. Nature 523:451–454.

Stanley, S. M. 2008. Predation defeats competition on the seafloor. Paleobiology 34:1–21.

Tarhan, L. G., M. L. Droser, and J. G. Gehling. 2015a. Depositional and preservation environments of the Ediacara Member, Rawnsley Quartzite (South Australia): assessment of paleoenvironmental proxies and the timing of “ferruginization.” Palaeogeography, Palaeoclimatology, Palaeoecology 434:4–13.

Tarhan, L. G., M. L. Droser, N. J. Planavsky, and D. T. Johnston. 2015b. Protracted development of bioturbation through the early Palaeozoic Era. Nature Geoscience 8:865–869.

Tarhan, L. G., A. v S. Hood, M. L. Droser, J. G. Gehling, and D. E. G. Briggs. 2016. Exceptional preservation of soft-bodied Ediacara Biota promoted by silica-rich oceans. Geology 44:951–954.

Tarhan, L. G., M. L. Droser, J. G. Gehling, and M. P. Dzaugis. 2017. Microbial mat sandwiches and other anactualistic sedimentary features of the Ediacara Member (Rawnsley Quartzite, South Australia): implications for interpretation of the Ediacaran sedimentary record. Palaios 32:181–194.

Tomašových, A., and S. M. Kidwell. 2009. Fidelity of variation in species composition and diversity partitioning by death assemblages: time-averaging transfers diversity from beta to alpha levels. Paleobiology 35:94–118.

Veech, J. A., and T. O. Crist. 2010. Diversity partitioning without statistical independence of alpha and beta. Ecology 91:1964–1969.

Xiao, S., and M. Laflamme. 2009. On the eve of animal radiation: phylogeny, ecology and evolution of the Ediacara biota. Trends in Ecology and Evolution 24:31–40.

Zhao, F., J.-B. Caron, D. J. Botter, S. Hu, Z. Yin, and M. Zhu. 2013. Diversity and species abundance patterns of the Early Cambrian (Series 2, Stage 3) Chengjiang Biota from China. Paleobiology 40:50–69.