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Key Points:
- We compare results from an ice sheet model inter-comparison forced using Coupled Model Intercomparison Project phase 6 and phase 5 climate projections
- Projected sea level at 2000 is higher for Greenland under CMIP6 scenarios than CMIP5, but similar for Antarctica under both scenarios
- CMIP6 warmer climate results in increased Greenland surface melt while increased snowfall mitigates loss from ocean warming for Antarctica

Supporting Information:
Supporting Information may be found in the online version of this article.

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Future Sea Level Change Under Coupled Model Intercomparison Project Phase 5 and Phase 6 Scenarios From the Greenland and Antarctic Ice Sheets

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1. Introduction

The overall aim of this paper is to assess whether the stronger future warming shown by many Coupled Model Intercomparison Project phase 6 (CMIP6) models (Forster et al., 2019; Meehl et al., 2020) compared with Coupled Model Intercomparison Project phase 5 (CMIP5) effort. Here we use four CMIP6 models and a selection of CMIP5 models to force multiple ice sheet models as part of the Ice Sheet Model Intercomparison Project for CMIP6 (ISMIP6). We find that the projected sea level contribution at 2100 from the ice sheet model ensemble under the CMIP6 scenarios falls within the CMIP5 range for the Antarctic ice sheet but is significantly increased for Greenland. Warmer atmosphere in CMIP6 models results in higher Greenland mass loss due to surface melt. For Antarctica, CMIP6 forcing is similar to CMIP5 and mass gain from increased snowfall counteracts increased loss due to ocean warming.

2. The CMIP6 Ensemble

We compare a small ensemble of four Earth System Models (ESMs) submitted to the CMIP6 exercise. These models are UKESM1-0-LL, CESM2, CNRM-CM6-1, and CNRM-ESM2-1, which were the only ones available for downscaling at the time. Because the sample is small and based on availability only, it is important to...
understand the difference between the selected models and the larger CMIP6 model ensemble. Effective Climate Sensitivity (ECS) (IPCC, 2013) is a convenient measure of this. ECS estimates the global mean temperature response to doubled atmospheric carbon dioxide concentration (Flato et al., 2013). The four selected models all have ECS at the upper end of the CMIP6 ensemble (CESM2, CNRM-CM6-1, CNRM-ESM2-1 and UKESM1-0-LL have ECS of 5.2°C, 4.8°C, 4.8°C, and 5.3°C, respectively). Roughly half of the CMIP6 ensemble has an ECS of between 4.6°C and 5.6°C, while there is a second similarly sized group with markedly lower ECS in the range 2.3°C–3.2°C (Meehl et al., 2020). In contrast, the CMIP5 ensemble exhibited a fairly continuous range of ECS between 2.1°C and 4.7°C (Flato et al., 2013). The CMIP5 models used in Goelzer, Nowicki, et al. (2020) and Seroussi et al. (2020) were typically drawn from the upper end of this distribution (e.g., MIROC-ESM, HadGEM2-ES, CSIRO-Mk3-6-0 and IPSL-CM5A-LR with ECS of 4.7°C, 4.6°C, 4.1°C, and 4.1°C, respectively) or lay close to the median (e.g., CCSM4, NorESM1-M and MIROC5 with ECS of 2.9°C, 2.8°C and 2.7°C, respectively).

Summaries of the atmospheric and ocean forcing for the two ice sheets are shown in Figures 1 and 2, respectively. Surface warming exhibited over the AIS in CMIP6 lies at or above the high end of the CMIP5 range. A similar pattern is evident in projected changes in Surface Mass Balance (SMB, the annual difference between mass addition, such as snowfall and refrozen rainfall, and mass loss, such as melt and subsequent runoff) over the ice sheet. Neither quantity is, however, significantly higher than the CMIP5 range. For

![Figure 1](image-url)
GrIS, SMB was derived by forcing the MAR regional climate model of Greenland (Fettweis et al., 2013) with CMIP6-derived boundary conditions. In this case, the CMIP6-forced SMB is significantly more negative (i.e., higher GMSLR rise) than is the case for CMIP5 forcing. Indeed, all four SSP585 ESMs fall outside the CMIP5 range and, by 2100, anomalies from UKESM1-0-LL and CESM2 approach twice that of largest CMIP5 ESM. The oceanic forcing of the AIS is described in detail by Jourdain et al. (2020) and for the GrIS by Slater et al. (2020). The thermal forcing derived from the CMIP6 models for both ice sheets lies within the range of the CMIP5 models with the exception of UKESM1-0-LL SSP585, which is occasionally higher. In many cases, the forcing lies towards the center of the CMIP5 range despite the higher ECS of the CMIP6 models. As would be expected thermal forcing from CNRM-CM6-1 SSP126 is less than that from CNRM-CM6-1 SSP585, however the difference is similar to the difference between the four SSP585 models.

3. Summary of ISMIP6 Experimental Procedure

The procedures used to convert the climate information summarized in Figures 1 and 2 into forcing imposed on ice sheet models are summarized in a series of papers for Antarctic ocean (Favier et al., 2019; Jourdain et al., 2020), Greenland ocean (Slater et al., 2019, 2020) and Greenland atmosphere (Fettweis et al., 2020; Jourdain et al., 2020; Slater et al., 2020).
et al., 2013; Goelzer, Noël, et al., 2020). Details of the experimental protocols employed can be found in Nowicki et al. (2016) and Nowicki et al. (2020) and employed a carefully chosen sub-sample of six CMIP5 models for each ice sheet.

These protocols were primarily employed by ice sheet modeling groups to generate projections using forcing from the CMIP5 ensemble, which are reported in Goelzer, Nowicki, et al. (2020) for GrIS and Seroussi et al. (2020) for AIS, however groups also conducted experiments using forcing from the CMIP6 ensemble as summarized in Tables 1 and 2. Both tables refer to experiments using the following numbering: (a) The CNRM-CM6-1 model run with scenario SSP585 (roughly equivalent to RCP8.5 of CMIP5), (b) CNRM-CM6-1 with SSP126 (roughly equivalent to RCP2.6 of CMIP5), and SSP585 with (c) UKESM1-0-LL, (d) CESM2, (e) CNRM-ESM2-1. Within the ISMIP6 design, experiments could be performed under “standard” or “open” configurations (see Nowicki et al., 2020). The former refers to the full implementation of ISMIP6 protocols for converting climate forcing into the mass fluxes experienced by the ice sheets, while in the latter individual groups used their own previously existing methods to do this.

### 4. GMSLR Projections

Figure 3 shows projections for the AIS from the seven participating ice sheet models for each CMIP6-forced experiment along with ranges from the equivalent CMIP5-forced experiments (Seroussi et al., 2020). Figure 3b–3d compares these projections with ranges derived for the CMIP5 ensemble at 2100. The equivalent ranges for the whole AIS are −14–155 mm for RCP2.6, and −76–300 mm for RCP8.5. The regional contributions from West and East AIS are within or below the ranges reported for CMIP5 forcing. In many cases, they sit in the lower half of this range. This, however, is likely to reflect the high GMSLR associated with one ESM in CMIP5 ensemble of six (HadGEM2-ES), whose projected GMSLR was typically much higher (roughly twice that of the other ESMS for West AIS and positive rather than negative for East AIS). The projected GMSLR for all three AIS regions for CMIP6 and CMIP5 is very compatible if HadGEM2-ES is excluded from the latter.

Comparing projections for SSP126 (one ESM only) and SSP585 (four ESMS) suggests that there is little impact of emission scenario on projected GMSLR for AIS. This is, again, most likely to be related to the contrasting impacts for global warming on the ice sheet’s mass budget through increases in both mass loss by ice-sheet discharge and gain by snow accumulation.

The relationship between forcing and GMSLR for each CMIP6 ESM is complicated. For instance, ocean thermal forcing (Figure 2), air temperature anomalies (Figure 2) tend to be larger for UKESM1-0-LL; however, this is not reflected in their projected GMSLR. This is most likely to be associated with the compensatory effect of increased precipitation (Figure 2) in these ESMS.

Figure 4 shows projections for the GrIS from the 14 participating ice-sheet models for each CMIP6-forced experiment along with ranges from the equivalent CMIP5-forced experiments (Goelzer, Nowicki, et al., 2020). Projected GMSLR is either at the upper end of the CMIP5-forced range

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**Table 1**

Overview of Experiments and Modeling Groups Participating in the CMIP6-Forced Exercise for AIS

| Group   | Model   | Open | Standard | Symbol |
|---------|---------|------|----------|--------|
| AWI     | PISM    | 1–5  | 1–5      | o      |
| ILTS_PIK| SICOPOLIS| 1–5  | <        |        |
| JPL     | ISSM    | 1–5  | △        |        |
| NCAR    | CISM    | 1–5  | △        |        |
| LSCE    | GRISLI  | 1–5  | □+       |        |
| UC1IPL  | ISSM    | 1–5  | □−       |        |
| VUB     | AISPALAO| 1–3  |          |        |

*Total: 2  7*

**Note.** Please refer to Seroussi et al. (2020) for model and group details. Symbols are those used in Figure 3. Abbreviations: AIS, Antarctic ice sheet; CMIP6, Coupled Model Intercomparison Project phase 6.

**Table 2**

Overview of Experiments and Modeling Groups Participating in the CMIP6-Forced Exercise for GrIS

| Group   | Model   | Open | Standard | Symbol |
|---------|---------|------|----------|--------|
| AWI     | ISSM1   | 1–5  |          |        |
| AWI     | ISSM2   | 1–5  |          |        |
| AWI     | ISSM3   | 1–5  |          |        |
| BGC     | BISICLES| 1–3  | □+       |        |
| GSFC    | ISSM    | 1–2  | □−       |        |
| ILTS_PIK| SICOPOLIS1| 1–5  | △        |        |
| ILTS_PIK| SICOPOLIS2| 1–5  | □−       |        |
| IMAU    | IMAUICE2| 1–3,5| □+       |        |
| JPL     | ISSM    | 1–5  | □−       |        |
| JPL     | ISSM2   | 1–5  | □−       |        |
| NCAR    | CISM    | 1–5  | □−       |        |
| UAF     | PISM1   | 1–3,5| □−       |        |
| UAF     | PISM2   | 1–3,5| □−       |        |
| UC1IPL  | ISSM1   | 1–3  | □+       |        |
| VUB     | GISM    | 1–5  | □−       |        |

*Total: 2  14*

**Note.** Please refer to Goelzer, Nowicki, et al. (2020) for model and group details. Symbols are those used in Figure 4. “f” refers to filled symbol. Abbreviations: CMIP5, Coupled Model Intercomparison Project phase 6; GrIS, Greenland ice sheet.
or well above it. Indeed, both CESM2 and UKESM1-0-LL-based projections do not overlap with the CMIP5 range at all and, in the latter case, are almost double. In contrast to the AIS, projections for SSP126 (one ESM) are considerably lower than SSP585 (four ESMs) such that the ranges for CMIP6 SSP126 and SSP585 do not overlap. The trajectory of GMSLR associated with SSP126 starts to become distinct from SSP585 around 2060 but is not entirely separate until 2090. There is also a suggestion that GMSLR may stabilize (or at least increase at a far reduced rate) beyond 2100 for SSP126, which is certainly not the case for SSP585.

5. Discussion

We present the first comparison between CMIP5 and CMIP6-based projections of the contribution of ice sheets to future GMSLR up to 2100. This comparison is particularly interesting because many CMIP6 ESMs have higher climate sensitivity than their CMIP5 counterparts (Forster et al., 2019; Meehl et al., 2020) and their projections of future global warming are therefore higher. The comparison is hampered by the use of a relatively small ensemble of available CMIP6 ESMs, which are all at the upper end of CMIP6's range of climate sensitivity.

The comparison between CMIP5 and CMIP6 is markedly different for the two ice sheets, reflecting the very different ways in which the ice sheets are impacted by and respond to changes in the global climate system.
For the GrIS, our results suggest that GMSLR contributions under CMIP6 are much higher than for CMIP5 perhaps by a factor of two. They also suggest a significant difference between SSP585 and SSP126, with the former experiencing accelerating rates of mass loss in marked contrast to the tendency towards stabilization of the latter.

Goelzer, Nowicki, et al. (2020) demonstrate that in excess of 80% of GrIS’ contribution to GMSLR can be explained by changing SMB (primarily by surface melt and subsequent runoff), which is mostly controlled by atmospheric processes. The link between global warming and mass loss from the ice sheet is therefore fairly direct and a strong relationship between the two should be expected. The higher climate sensitivity of the sampled CMIP6 ESMs will therefore manifest itself as a larger GMSLR contribution in comparison to CMIP5. It should also be noted that for GrIS (in contrast to AIS), global warming is likely to favor increased mass loss by both atmospheric (i.e., SMB) and ocean forcing (i.e., discharge). However it appears that, at least within the ISMIP6 experimental design, ocean forcing plays a secondary role to the atmosphere.

For AIS, our results up to 2100 suggest little difference between CMIP6 and CMIP5-forced projections. This reflects the more complex interactions between this ice sheet and the global climate system. Global warming is likely to favor mass loss through changes in discharge resulting from increased ocean thermal forcing; however, the opposite is expected of the atmospheric forcing where warming is likely to favor mass gain (as a consequence of increased snow accumulation). The higher climate sensitivity of the sampled CMIP6 ESMs will therefore manifest itself as a larger GMSLR contribution in comparison to CMIP5. It should also be noted that for GrIS (in contrast to AIS), global warming is likely to favor increased mass loss by both atmospheric (i.e., SMB) and ocean forcing (i.e., discharge). However it appears that, at least within the ISMIP6 experimental design, ocean forcing plays a secondary role to the atmosphere.

The experimental design of the CMIP6-forced experiments reported here does not include the fracture and collapse of AIS’ floating ice shelves resulting from meltwater ponding due to significant atmospheric warming (Trusel et al., 2015). This process has been cited as a necessary precursor to rapid ice loss by the retreat of marine ice cliffs (DeConto & Pollard, 2016). As ice shelf fracture was included in the CMIP5-forced experiments, an initial assessment can be made by comparing the amount of atmospheric warming projected to occur under CMIP5 and CMIP6. Figure 2 suggests that CMIP6 ESMs lie close to or above the maximum CMIP5 surface temperature warming for AIS. For CMIP5 forcing, this process is limited to the Antarctic Peninsula and areas around George VI ice shelf and Totten glacier and its impact on GMSLR is ∼28 mm (Seroussi et al., 2020). Ice-shelf fracture and associated processes may therefore be important under some CMIP6 forcing, in particular for CESM2 and UKESM1-0-LL, and likely be enhanced beyond 2100.
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References

Barthel, A., Agosta, C., Little, C. M., Hattermann, T., Jourdain, N. C., Goelzer, H., et al. (2020). CMIP5 model selection for ISMIP6 ice sheet model forcing: Greenland and Antarctica. The Cryosphere, 14, 855–879. https://doi.org/10.5194/tc-14-855-2020

DeConto, R. M., & Pollard, D. (2016). Contribution of Antarctica to past and future sea-level rise. Nature, 531(7596), 591–597. https://doi.org/10.1038/nature17145

Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., & Taylor, K. E. (2016). Overview of the coupled model intercomparison project phase 6 (cmip6) experimental design and organization. Geoscientific Model Development, 9(5), 1937–1958. https://doi.org/10.5194/gmd-9-1937-2016

Favier, L., Jourdain, N. C., Jenkins, A., Merino, N., Durand, G., Gagliardini, O., et al. (2019). Assessment of sub-shelf melting parameterizations using the ocean-ice-sheet coupled model NEMO/ICE/MARS. Geoscientific Model Development, 12, 2255–2283. https://doi.org/10.5194/gmd-12-2255-2019

Fettweis, X., Franco, B., Tesedo, M., van Angelen, J. H., Lenaerts, J. T. M., van den Broeke, M. R., & Gallée, H. (2013). Estimating the Greenland ice sheet surface mass balance contribution to future sea level rise using the regional atmospheric climate model mar. The Cryosphere, 7(2), 469–489. https://doi.org/10.5194/tc-7-469-2013

Flato, G., Marotzke, J., Abiodun, B., Braconnot, P., Chou, S., Collins, W., & Rummukainen, M. (2013). Evaluation of climate models. In Stocker, T., Qin, D., Plattner, M., Tignor, S. K., Allen, J., Boschung, A., Nauels, Y., & Michel, C. (Eds.). Climate change 2013: The physical science basis. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change (pp. 741–866). Cambridge and New York, NY: Cambridge University Press. https://doi.org/10.1038/9781107415324.020

Forster, P., Maycock, A., McKenna, C., & Smith, C. (2019). Latest climate models confirm need for urgent mitigation. Nature Climate Change, 10, 7–10. https://doi.org/10.1038/s41558-019-0660-0

Goelzer, H., Noël, B. P. Y., Edwards, T. L., Fettweis, X., Gregory, J. M., Lipscomb, W. H., et al. (2020). Remapping of Greenland ice sheet surface mass balance anomalies for large ensemble sea-level change projections. The Cryosphere, 14(6), 1747–1762. https://doi.org/10.5194/tc-14-1747-2020

Goelzer, H., Nowicki, S., Payne, A., Larour, E., Seroussi, H., Lipscomb, W., & van den Broeke, M. (2020). The future sea-level contribution of the Greenland ice sheet: A multi-model ensemble study of ISMIP6. The Cryosphere, 14(9), 3071–3096. https://doi.org/10.5194/tc-14-3071-2020

IPCC (2013). Annex III: Glossary. In Stocker, T., Qin, D., Plattner, M., Tignor, S. K., Allen, J., Boschung, A., Nauels, Y., & Michel, C. (Eds.). Climate change 2013: The physical science basis. Contribution of working group I to the fifth assessment report of the intergovernmental panel on climate change (p. 1447–1466). Cambridge and New York, NY: Cambridge University Press. https://doi.org/10.1017/CBO9781107415324.031

Jenkins, A., Shosmith, D., Dutrieux, P., Jacobs, S., Kim, T. W., Lee, S. H., et al. (2018). West Antarctic ice sheet retreat in the Amundsen Sea driven by decadal oceanic variability. Nature Geoscience, 11(10), 733–738. https://doi.org/10.1038/s41561-018-0207-4

Jourdain, N. C., Assay-Davis, X., Hattermann, T., Straneo, F., Seroussi, H., Little, C. M., & Nowicki, S. (2020). A protocol for calculating basal melt rates in the ISMIP6 Antarctic ice sheet projections. The Cryosphere, 14(9), 3111–3134. https://doi.org/10.5194/tc-14-3111-2020

Meehi, G. A., Senior, C. A., Eyring, V., Flato, G., Lamarque, J.-F., Stouffer, R. J., et al. (2020). Context for interpreting equilibrium climate sensitivity and transient climate response from the CMIP6 earth system models. Science Advances, 6(26), eaba1981. https://doi.org/10.1126/sciadv.aba1981

Nowicki, S., Goelzer, H., Seroussi, H., Payne, A. J., Lipscomb, W. H., Abe-Ouchi, A., et al. (2020). Experimental protocol for sea level projections from ISMIP6 stand-alone ice sheet models. The Cryosphere, 14(7), 2331–2368. https://doi.org/10.5194/tc-14-2331-2020

Nowicki, S. M. J., Payne, A., Larour, E., Seroussi, H., Goelzer, H., Lipscomb, W., et al. (2016). Ice sheet model intercomparison project (ISMIP6) contribution to CMIP5. Geoscientific Model Development, 9(12), 4521–4545. https://doi.org/10.5194/gmd-9-4521-2016

Schoof, C. (2007). Ice sheet grounding line dynamics: Steady states, stability, and hysteresis. Journal of Geophysical Research, 112(F3), F03S28. https://doi.org/10.1029/2006JF000664

Seroussi, H., Nowicki, S., Payne, A. J., Goelzer, H., Lipscomb, W. H., & Abe-Ouchi, A., & Zwinger, T. (2020). ISMIP6 Antarctica: A multi-model ensemble of the Antarctic ice sheet evolution over the 21st century. The Cryosphere, 14(9), 3033–3070. https://doi.org/10.5194/tc-14-3033-2020

Slater, D. A., Feliksion, D., Straneo, F., Goelzer, H., Little, C. M., Morlighem, M., et al. (2020). Twenty-first century ocean forcing of the Greenland ice sheet for modeling of sea level contribution. The Cryosphere, 14(3), 985–1008. https://doi.org/10.5194/tc-14-985-2020

Slater, D. A., Straneo, F., Feliksion, D., Little, C. M., Goelzer, H., Fettweis, X., & Holte, J. (2019). Estimating Greenland tidewater glacier retreat driven by submarine melting. The Cryosphere, 13(9), 2489–2509. https://doi.org/10.5194/tc-13-2489-2019

Taylor, K. E., Stouffer, R. J., & Meehl, G. A. (2012). An overview of CMIP5 and the experiment design. Bulletin of the American Meteorological Society, 93(4), 485–498. https://doi.org/10.1175/BAMS-D-11-00094.1

Trusel, L. D., Frey, K. E., Das, S. B., Karnauskas, K. B., Kuipers Munneke, P., van Meijgaard, E., & van den Broeke, M. R. (2015). Divergent trajectories of Antarctic surface melt under two twenty-first-century climate scenarios. Nature Geoscience, 8(12), 927–932. https://doi.org/10.1038/ngeo2563