Soil N retention and nitrate leaching in three types of dunes in the Mu Us desert of China

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A large reservoir of soil nitrate in desert subsoil zones has been demonstrated in previous studies; however, information on the subsoil nitrate reservoir and its distribution characteristics in the deserts of China is still limited. This study investigated the distribution patterns of soil total nitrogen (N), nitrate, ammonium, and stable isotopic ratios of 15N (δ15N) in shallow (1 m) and subsoil (5 m) profiles in three types of dunes in the Mu Us desert of China. We found that soil N retention of the fixed and semi-fixed dunes followed a progressive nutrient depletion pattern in shallow soil profiles, whereas the subsoil nitrate of the fixed, semi-fixed and mobile dunes maintained a conservative accumulation pattern. The results indicate that the subsoil of the Mu Us desert may act as a reservoir of available nitrate. Furthermore, a soil δ15N analysis indicates that the nitrate content of the fixed dune is likely derived from soil nitrification, whereas the nitrate content in the mobile dune is derived from atmospheric nitrate deposition. Within the context of looming climate change and intensifying human activities, the subsoil nitrate content in the deserts of northern China could become mobilized and increase environmental risks to groundwater.

Deserts cover one-third of the land surface worldwide and play an important role in the function of global ecosystems and biogeochemical cycling1–4. Desert soils are generally thought to be nutrient poor and low in total nitrogen (N)1,5. However, recent studies have demonstrated that desert subsoil represents a large reservoir of bioavailable N in the form of nitrate, suggesting that this N pool has been previously overlooked6–9. Investigations of this subsoil N storage could increase estimates of vadose-zone N content by 14 to 71% for warm deserts and arid shrublands worldwide6. Moreover, subsoil nitrate may contaminate groundwater and exert further negative effects after land-use or climate change in deserts6. Therefore, additional investigations of subsoil nitrate reservoirs are required in the field of desert environmental research.

China contains several of the largest areas of desert and desertified land in the world. The total area of desert in China is estimated at approximately 1.53 × 106 km², and deserts occupy approximately 15.9% of the total national land area10,11. Numerous studies have suggested that desert subsoil in China could accumulate a large amount of bioavailable N from massive atmospheric nitrate depositions and active N fixation by biological soil crusts8,12–14. However, to our knowledge, few studies have been conducted to characterize the subsoil nitrate distribution and dynamics in Chinese deserts. Moreover, numerous studies have demonstrated that the climate of northwest China has experienced an increasing warming and wetting trend over the past several decades and suggest that this trend will continue throughout the 21st century15–17. Increases in precipitation would lead to additional leaching and mobilization of subsoil nitrate in deserts and heighten the environmental risk to groundwater. Therefore, investigations of the distribution characteristics of subsoil nitrate reservoirs in Chinese deserts are urgently required, and the resulting information will broaden our understanding of desert N cycling.

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The Mu Us desert is located in the southeastern region of the Ordos Plateau in northern China and covers an area of 40,000 km². Over the past decade, the region has been the most active area of economic growth in China because of its rich coal, oil and natural gas resources. Groundwater is the main water source in the Mu Us desert, and soil nitrate is readily mobilized to the groundwater because of the shallow groundwater table. Moreover, the rapid development of petroleum and coal industries in the region could significantly increase the atmospheric deposition of nitrate and pose an additional potential threat to the groundwater. However, little is known of the distribution patterns of subsoil nitrate in the region, which limits our understanding of whether a subsoil nitrate reservoir occurs in the area and how it might be characterized. In this study, three typical landforms in the Mu Us desert – fixed dune, semi-fixed dune and mobile dune – were selected, and soil samples from shallow profiles (1 m) and subsoil drilling cores (5 m) were analysed to determine the total N, nitrate, ammonium, and δ¹⁵N (Fig. 1). The objectives of this study were to (1) investigate the characteristics of soil N retention in shallow soil profiles and (2) characterize the distribution patterns of soil nitrate in the subsoil horizon.
Results

Plant and soil properties among the three types of dunes. The three types of dunes exhibited considerable differences in plant and soil properties (Table 1). The fixed dune presented the highest vegetation cover, plant species composition, silt and clay content, and bulk density, and *Artemisia ordosica* and *Hedysarum fruticosum* were the dominant species. The semi-fixed dune presented lower vegetation cover, plant species composition, silt and clay content, and bulk density compared with that of the fixed dune, and *A. ordosica* was the dominant species. The mobile dune was barren and presented the highest sand content and lowest soil bulk density.

Characteristics of soil N retention in the upper 1 m of the soil. In the upper 1 m of the soil, the content of soil total N, nitrate and ammonium presented minimal changes with increasing soil depth for the mobile dune, whereas for the fixed and semi-fixed dunes, the soil N content decreased with increasing depth and showed a progressive nutrient depletion pattern (Fig. 2a–c). A statistical analysis demonstrated that the fixed dune has the highest content of soil total N and ammonium among the three types of dunes (total N: \( P = 0.015 \); ammonium: \( P = 0.012 \)), whereas soil nitrate does not show significant differences (nitrate: \( P = 0.151 \)). Ammonium was the dominant form of soil inorganic N preserved in the upper 1 m of the soil, and the ratios of soil nitrate to ammonium averaged 0.24 for the three types of dunes (Fig. 2d).

Distribution patterns of soil nitrate in the 5 m soil layer. In the 5 m soil layer, soil nitrate exhibited obvious pattern of leaching for all of the three types of dunes (Fig. 3). In the mobile dune, the soil nitrate content presented minimal changes within the depth range of 0–3.5 m, whereas below 3.5 m, the soil nitrate content significantly increased and showed an obvious pattern of leaching. In the fixed and semi-fixed dunes, nitrate in the shallow soil profiles showed a progressive nutrient depletion pattern, whereas nitrate in the deep soil horizon exhibited an increasing pattern, although the distribution patterns varied.

Soil \( ^{15} \)N values among the three types of dunes. The values of \( ^{15} \)N in the upper 1 m of the soil exhibited considerable variation among the three types of dunes, averaging 1.74%, 0.80% and −0.15% for the fixed, semi-fixed and mobile dunes, respectively (Fig. 4). A statistical analysis demonstrated that the fixed dune had the highest values of \( ^{15} \)N, whereas the mobile dune had the lowest (\( P < 0.001 \)).

Discussion

The three types of Chinese desert dunes evaluated in this study were found to have large differences in their levels of vegetative cover, which is the primary causal factor for the corresponding differences in the soil N retention and nitrate leaching. In the fixed dune, the vegetation cover was approximately 85%, with a large number of biological soil crusts observed to cover the soil surface. Dense vegetation cover constitutes a powerful system for biological cycling of soil N and usually leads to progressive nutrient depletion patterns in shallow soil profiles. Moreover, the dominant species of the fixed dune were *A. ordosica* and *H. fruticosum*. *H. fruticosum* is a species of the legume family, and biological N fixation by *H. fruticosum* and soil crusts also promotes soil N increases at the soil surface. Moreover, a high degree of vegetation cover and biological soil crust formation can interfere with the atmospheric deposition of N. These factors contributed to the current situation in which the fixed dunes showed

| Plant properties | Soil properties |
|------------------|-----------------|
| **Vegetation cover (%)** | **Soil texture (0–20 cm, %, mean ± s.d.)** | **Bulk density (0–100 cm, g cm\(^{-3}\), mean ± s.d.)** | **Soil pH (0–100 cm, mean ± s.d.)** |
| **Plant species composition** | **Sand** | **Silt** | **Clay** | **Sand** | **Silt** | **Clay** |
| **Fixed dune** | 85 | Artemisia ordosica, Hedysarum fruticosum, *Hedysarum scoparium*, *Stipa bungeana*, *Cleistogenes squarrosa*, *Oxytropis psammochans*, *Lespedeza davurica*, *Pylgosa tehuifolia* | Artemisia ordosica, *Hedysarum fruticosum* | 67.85 ± 10.10* | 27.97 ± 9.24* | 4.10 ± 1.02* | 1.66 ± 0.02* | 7.62 ± 0.17* |
| **Semi-fixed dune** | 25 | Artemisia ordosica, *Psammochloa villosa*, *Cynanchum komarovii*, *Chenopodium aristatum*, *Hedysarum fruticosum* | Artemisia ordosica | 80.54 ± 3.49* | 16.94 ± 3.28* | 2.52 ± 0.42* | 1.60 ± 0.03* | 7.67 ± 0.22* |
| **Mobile dune** | 0 | none | none | 94.50 ± 1.56* | 4.53 ± 1.11* | 0.97 ± 0.47* | 1.53 ± 0.02* | 7.63 ± 0.09* |

Table 1. Plant and soil properties in the fixed, semi-fixed and mobile dunes. Notes: *with different letters within a variable indicates a significant difference.*
Figure 2. Distribution patterns of total N (a), nitrate (b), ammonium (c) and ratios of nitrate to ammonium (d) in the upper 1 m of the soil.

Figure 3. Distribution patterns of soil nitrate in the 5 m soil layer.
the highest soil total N content and the most obvious pattern of soil N decreases with depth among the three types of dunes (Fig. 2a–c). In contrast, the soil total N, nitrate and ammonium content showed little changes with increasing soil depth for the mobile dune.

Previous studies have demonstrated that the high soil-surface temperatures of deserts (greater than 50 °C), which are driven by solar radiation, can cause abiotic losses of N in the form of NO\textsubscript{y} (all forms of oxidized gaseous N) and NH\textsubscript{3}\textsuperscript{29}. Moreover, the soil pH tends to be high in extremely dry areas, which is also a key driver of ammonium loss associated with volatilization\textsuperscript{30}. Therefore, ammonium is generally considered difficult to preserve in the desert soil. In this study, we found that ammonium was the dominant form of inorganic N preserved in the upper 1 m of the soil and that the level of soil ammonium was nearly three times higher than that of soil nitrate (Fig. 2d). We conclude that the relatively low temperature and soil pH played a critical role in regulating the soil ammonium retention at the study site. Hu et al. (2008) reported that soil ammonium retention is both negatively and significantly correlated with the air temperature and soil pH in the drylands of central East Asia, where low temperature and soil pH are found to correspond to high soil ammonium content\textsuperscript{31}. In this study, the annual mean temperature of the area was 6.7 °C and the monthly mean temperatures from April to October range from 7.4–21.9 °C\textsuperscript{32}. Moreover, the soil pH of the study area is not too high due to a relatively high rate of precipitation (with an annual average of 345 mm) (Table 1). Numerous studies have reported that low temperature and soil pH could inhibit the ammonium oxidation rate and suppress ammonia volatilization\textsuperscript{30,33–35}, which are both beneficial to the soil ammonium storage\textsuperscript{31}. Moreover, the fixed dune had the highest content of soil ammonium measured among the three types of dunes (P = 0.012), indicating that the net soil N accumulation dominates the ecosystem N cycling process as the biological N fixation of the legume species and soil crusts.

In deserts, the storage of available soil N is low\textsuperscript{36–38}; as such, the resulting nitrate leaching from desert soils is also expected to be low. We found, however, that the soil nitrate content increased significantly in the subsoil horizon, which indicates that high levels of available soil N in the presence of nitrate have leached into the deeper soil. This result is similar to the findings obtained by Walvoord et al. (2003)\textsuperscript{6}. In 2003, Walvoord and colleagues discovered the first large reservoir of nitrate beneath desert soils in a western region of the United States. However, Jackson et al. (2004) questioned the generality of their results and found no significant increase in the soil nitrate measured in 16 desert soil profiles at depths of 10 m in the Jornada and Sevilleta desert regions of the United States\textsuperscript{39}. Recently, Graham et al. (2008) reported that large near-surface nitrate pools were found in the soils capped by desert pavement in the Mojave Desert\textsuperscript{7}. Therefore, we speculate that subsoil nitrate reservoirs may exist in deserts. In this study, the increase seen in the nitrate content of subsoil indicates the likely presence of a nitrate reservoir in the subsoil zones of the Mu Us desert. However, because of the large differences in the vegetation cover, the three types of dunes showed different patterns of subsoil nitrate distribution. The mobile dune is barren and does not provide physical buffering and biological regulation of vegetation. Consequently, the mobile dune showed a very clear process of nitrate leaching (Fig. 3). In the fixed and semi-fixed dunes, the presence of vegetation was seen to affect the soil N cycling (Figs 2 and 3). As a result, the fixed and semi-fixed dunes had a fluctuating pattern of nitrate leaching (Fig. 3).

Walvoord et al. (2003) explained that available soil nitrate is not completely consumed by plants or returned to the atmosphere\textsuperscript{6}, which is a possible cause of nitrate leaching in deserts. Moreover, Gebauer and Ehleringer (2000) discovered that desert plants do not necessarily consume water and nutrients simultaneously\textsuperscript{40}, which could also contribute to nitrate loss. In this study, the relatively high precipitation rate was found to also promote nitrate leaching, while the soil δ\textsuperscript{15}N analysis showed that different dunes may have different sources of nitrate. The fixed dune had the highest values of δ\textsuperscript{15}N (1.74%), while
the mobile dune had the lowest values of $\delta^{15}$N (−0.15%). Nitrate is known to have a more negative $\delta^{15}$N value than ammonium$^{41-43}$, indicating a source with increased nitrate content, whereas a more positive value of $\delta^{15}$N indicates a source with increased ammonium content$^{44}$. In this study, the highest value of soil $\delta^{15}$N was observed in the fixed dune, thus indicating that the biological N fixation of legume species and soil crusts dominates the surface soil N process, producing ammonium and maintaining the nutrient demands of the desert ecosystem. Therefore, nitrate in the fixed dune was most likely to have been derived from soil nitrification. Correspondingly, the lowest soil $\delta^{15}$N content was found in the mobile dune, indicating that the soil N in this dune was primarily derived from the atmospheric nitrate deposition.

A number of studies have indicated that changes in climate or land use could mobilize the subsoil nitrate reservoirs that have accumulated over thousands of years in deserts$^{6,7}$. The activated subsoil nitrate can contaminate groundwater and adversely affect public water supplies$^{45}$. In the Mu Us desert, the risk of nitrate pollution to the groundwater is high because of the large-scale coal mining operations in the area. Moreover, Hong et al. (2014) demonstrated that the rates of summer precipitation in arid eastern Central Asia (to include northwestern China) have increased steadily over the past 8,500 years$^{45}$ and suggested that the trend of a wetter climate would continue in the future$^{13,16,46}$. Increasing precipitation in the arid regions of China has the potential to mobilize the subsoil nitrate in deserts and increase the environmental risks to groundwater. Therefore, the adverse effects of subsoil nitrate on the groundwater quality should be considered when managing water resources in the arid and semiarid regions of China.

**Methods**

**Study area.** This study was conducted at the Chinese Academy of Sciences’ Ordos Sandy Grassland Research Station, which is located in the Mu Us desert of Inner Mongolia, China (see Supplementary Fig. S1 online). The area has a typically semiarid climate with marked seasonal and diurnal temperature variations and low precipitation. The annual mean temperature is 6.7°C, with monthly mean temperatures falling below 5°C from November to March and ranging from 7.4°C to 21.9°C from April to October. The annual mean precipitation is 345 mm, with an annual mean evaporation of 2535 mm. From April to October, the mean precipitation is 322 mm, which accounts for approximately 93% of the annual precipitation$^{32}$. The topography of the area is characterized by sand dunes and desert shrub vegetation, and only a small area of grasslands is distributed in the lowland and upland areas$^{18}$.

In the vicinity of the Ordos Sandy Grassland Research Station, three typical landforms – fixed, semi-fixed, and mobile dunes – were selected as sampling sites. The fixed dune sampling site lies at 111° 11′ 39″ E and 39° 29′ 43″N and is 1313 m above sea level. Vegetation cover accounts for approximately 85% of the area, and the main plant species are *A. ordosica*, *H. fruticosum*, *Hedyarum scoparium* and *Stipa bungeana*$^{21,46,47}$, with *A. ordosica* and *H. fruticosum* identified as the dominant species (see Supplementary photographs of common plant species in the Ordos Sandy Grassland Research Station online, Figs S2–S66, Zhu Yajuan took and assembled in 2007). The percentages of sand (2–0.02 mm), silt (0.02–0.002 mm) and clay (<0.002 mm) measured at depths of 0–20 cm in the soil profile are 67.85%, 27.97% and 4.10%, respectively$^{48}$. The semi-fixed dune sampling site lies at 111° 8′ 38″ E and 39° 28′ 51″N and is 1289 m above sea level. Vegetation cover accounts for approximately 25% of the area, and the main plant species are *A. ordosica*, *P. villosa* and *C. komarovii*$^{21,46,47}$, with *A. ordosica* as the dominant species. The percentages of sand, silt and clay measured at depths of 0–20 cm in the soil profile are 80.54%, 16.94% and 2.52%, respectively$^{48}$. The mobile dune sampling site lies at 111° 15′ 50″ E and 39° 28′ 39″N and is 1317 m above sea level. The mobile dune is barren, and the percentages of sand, silt and clay measured at depths of 0–20 cm in the soil profile are 94.50%, 4.53% and 0.97%, respectively$^{48}$.

**Soil sampling and laboratory analysis.** Soil sampling was performed in July, 2013. In the study area, three soil profiles with depths of 0–100 cm were established in each of the fixed, semi-fixed, and mobile dune sampling sites. The distance among the three soil profiles of the fixed dune was approximately 150 m. At each fixed-dune profile, we used a soil corer with a diameter of 5 cm to collect the soil samples. After removing the biological soil crust and litterfall, soil samples were collected at 10-cm intervals to a depth of 100 cm below the soil surface. Soil samples were stored using polyethylene bags. Soil samples were then analysed to measure soil nitrate content.

All of the collected soil samples were air-dried in the laboratory, and any remaining roots were carefully removed. The soil pH was measured in a deionized water suspension (soil:water, 1:2.5) using a DMP-2 mV/pH detector (Quark Ltd., Nanjing, China). For the soil N analysis, the air-dried soil samples were ground in an agate mortar and passed through a 0.15-mm sieve. The soil total N content was measured using the Kjeldahl method$^{11}$, and then the nitrate was determined using ion chromatography$^{49}$. The nitrite concentration was lower than 0.005 mmol/kg$^{49}$. The soil total N, nitrate, ammonium, $\delta^{15}$N and pH were measured in the fixed, semi-fixed, and mobile dune sampling sites. The distance among the three soil profiles of the fixed dune was approximately 200 m. Soils were sampled by the core method and brought to laboratory for further analysis. Soil samples were collected at 10-cm intervals to a depth of 100 cm below the soil surface. In the sampling sites of the semi-fixed and mobile dunes, three replicate samples were collected from each profile at intervals of 10 cm to conduct a soil carbon analysis. In total, 30 soil samples were collected from each sampling site. After removing the biological soil crust and litterfall, soil samples were collected at 10-cm intervals to a depth of 100 cm below the soil surface. Soil samples were stored using polyethylene bags. Soil samples were then analysed to measure soil nitrate content.
δ15N = \left( \frac{R_{\text{sample}}}{R_{\text{standard}}} - 1 \right) \times 1000 \tag{1}

where R represents the isotope ratio (15N/14N) and R_{\text{standard}} is the 15N/14N ratio for atmospheric N2. The analytical precision for δ15N was 0.2%.

Data analysis. Statistically significant differences in the soil total N, nitrate, ammonium and δ15N among the three sampling sites were identified using a one-factor analysis of variance (ANOVA) and least significant difference calculations at an alpha level of 0.05 (α < 0.05). All statistical analyses were performed with the Statistical Program for Social Sciences (SPSS 11.0; SPSS Inc., Chicago, IL, USA).

References

1. Noy-Meir, I. Desert ecosystems. I. Environment and producers. Annu. Rev. Ecol. Evol. Syst. 4, 25–52 (1973).
2. Jickells, T. D. Global iron connections between desert dust, ocean biogeochemistry, and climate. Science 308, 67–71 (2005).
3. Austin, A. T. Has water limited our imagination for aridland biogeochemistry? Trends Ecol. Evol. 26, 229–235 (2011).
4. Maestre, F. T. et al. Plant species richness and ecosystem multifunctionality in global drylands. Science 335, 214–218 (2012).
5. Post, W. L., Pastor, J., Zinke, P. J. & Statenberger, A. Global patterns of nitrogen storage. Nature 317, 613–616 (1985).
6. Walvoord, M. A. et al. A reservoir of nitrate beneath desert soils. Science 302, 1021–1024 (2003).
7. Graham, R. C., Hirnass, D. R., Wood, Y. A. & Amrhein, C. Large near-surface nitrate pools in soils capped by desert pavement in the Mojave Desert, California. Geology 36, 259–262 (2008).
8. Ma, J. Z. et al. Spatial distribution of chloride and nitrate within an unsaturated dune sand of a cold-arid desert: Implications for paleoenvironmental records. Catena 96, 68–75 (2012).
9. Al-Tairi, A. A. & Al-Qudah, K. A. Investigation of desert subsurface nitrate in Northeastern Badia of Jordan. Sci. Total Environ. 442, 111–115 (2013).
10. Zhu, Z. D., Liu, S. & Ding, X. M. Retrospect of the history of desert research and a few problems in the field in China. J. Desert Res. 4, 3–7 (1984).
11. Dong, G. R., Li, S., Li, B. S., Wang, Y. & Yan, M. C. A preliminary Study on the formation and evolution of deserts in China. J. Desert Res. 11, 23–32 (1991).
12. Zhang, Y. et al. Nitrate Deposits in the Xinjiang Province (Xinjiang University Press, Urumqi, China, 2000).
13. Qin, Y. et al. Massive atmospheric nitrate accumulation in a continental interior desert, northwestern China. Geology 40, 623–626 (2012).
14. Li, K. et al. Atmospheric nitrogen deposition at two sites in an arid environment of central Asia. PLoS ONE 8, e67018, doi: 10.1371/journal.pone.0067018 (2013).
15. Shi, Y. F., Shen, Y. P. & Hu, R. J. Preliminary study on signal, impact and foreground of climatic shift from warm dry to warm humid in northwest China. J. Glaciol. Geocryol. 24, 219–226 (2002).
16. Shi, Y. et al. Discussion on the present climate change from warm dry to warm wet in northwest China. Quaternary Sci. 23, 152–164 (2003).
17. Zhai, P. M., Zhang, X. B., Wan, H. & Pan, X. H. Trends in total precipitation and frequency of daily precipitation extremes over China. J. Climate 18, 1096–1108 (2005).
18. Zhang, Y. S. The optimization pattern and principle of grassland construction and ecological background of Maowusu sandland. Acta Phytocenol. Sin. 18, 1–16 (1994).
19. Cheng, D. H. et al. Relationship between vegetation and groundwater in Mu Us desert. J. Jilin Univ. (Earth Sci. Ed.) 42, 184–189 (2012).
20. Wei, Y. et al. Atmospheric dry and wet nitrogen deposition in typical agricultural areas of North Shaanxi. Chin. J. App. Ecol. 21, 255–259 (2010).
21. Wang, Q. S., Dong X. J., Chen X. D. & Wang, B. Y. Study on some features of Artemisia ordosica community at the different successional stages. Acta Phytocenol. Sin. 21, 531–538 (1997).
22. Zhang, J. H. et al. Biological soil crust distribution in Artemisia ordosica communities along a grazing pressure gradient in Mu Us Sandy Land, Northern China. J. Arid Land 5, 172–179 (2013).
23. Schlesinger, W. H. & Pilmanis, A. M. Plant-soil Interactions in Deserts. Biogeochemistry 42, 169–187 (1998).
24. Jobbágy, E. G. & Jackson, R. B. The distribution of soil nutrients with depth: Global patterns and the imprint of plants. Biogeochemistry 53, 51–77 (2001).
25. Belnap, J. Nitrogen fixation in biological soil crusts from southeast Utah, USA. Biol. Fertil. Soils 35, 128–135 (2002).
26. Wang, S. L. & Ma, C. The spatial heterogeneity of organic carbon and total nitrogen contents in sand soil under different densities of Hedysarum Monglicum communities. Chin. J. Soil Sci. 45, 851–856 (2014).
27. Shachak, M. & Lovett, G. M. Atmospheric deposition to a desert ecosystem and its implications for management. Ecol. Appl. 8, 455–463 (1998).
28. Zaady, E. Seasonal change and nitrogen cycling in a patchy Negev desert: a review. Arid Land Res. Manag. 19, 111–124 (2005).
29. McCall, C. K. & Sparks, J. P. Abiotic gas formation drives nitrogen loss from a desert ecosystem. Science 326, 837–839 (2009).
30. Wang, C. C. et al. Aridity threshold in controlling ecosystem nitrogen cycling in arid and semi-arid grasslands. Nat. Commun. 5, 4799, doi: 10.1038/ncomms4799 (2014).
31. Hu, L., Lee, X. Q., Huang D. K. & Cheng, J. Z. Ammonium nitrogen in surface soil of arid and semi-arid Central East Asia. Geochimica 37, 572–580 (2008).
32. Jin, Z., Dong, Y. S., Qi, T. C. & An, Z. S. Soil respiration and net primary productivity in perennial grass and desert shrub ecosystems at the Ordos Plateau of Inner Mongolia, China. J. Arid Environ. 74, 1248–1256 (2010).
33. Zhang, J. B., Bai, Z. C., Zhu, T. B., Yang, W. Y. & Müller, C. Mechanisms for the retention of inorganic N in acidic forest soils of southern China. Sci. Rep. 3, 2342, doi: 10.1038/srep02342 (2013).
34. Zhang, J. B., Zhu, T. B., Cai, Z. C., & Müller, C. Nitrogen cycling in forest soils across climate gradients in Eastern China. *Plant Soil* **342**, 419–432 (2011).
35. Delgado-Baquerizo, M. *et al.* Decoupling of soil nutrient cycles as a function of aridity in global drylands. *Nature* **502**, 672–676 (2013).
36. Schlesinger, W. H. *et al.* Biological feedbacks in global desertification. *Science* **247**, 1043–1048 (1990).
37. Billings, S. A., Schaeffer, S. M. & Evans, R. D. Soil microbial activity and N availability with elevated CO₂ in Mojave Desert soils. *Global Biogeochem. Cy.* **18**, GB1011, doi: 10.1029/2003GB002137 (2004).
38. Austin, A. T. *et al.* Water pulses and biogeochemical cycles in arid and semi-arid ecosystems. *Oecologia* **141**, 221–235 (2004).
39. Jackson, R. B., Berthrong, S. T., Cook, C. W. & Jobbágy, E. G. Comment on "A Reservoir of Nitrate Beneath Desert Soils". *Science* **304**, 51 (2004).
40. Gebauer, R. L. & Ehleringer, J. R. Water and nitrogen uptake patterns following moisture pulses in a cold desert community. *Ecology* **81**, 1415–1424 (2000).
41. Högberg, T. 15N natural abundance in soil-plant systems. *New Phytol.* **137**, 179–203 (1997).
42. Houlton, B. Z. & Bai, E. Imprint of denitrifying bacteria on the global terrestrial biosphere. *Proc. Natl. Acad. Sci. USA* **106**, 21713–21716 (2009).
43. Liu, X. Y. *et al.* Ammonium first: natural mosses prefer atmospheric ammonium but vary utilization of dissolved organic nitrogen depending on habitat and nitrogen deposition. *New Phytol.* **199**, 407–419 (2013).
44. Scanlon, B. R., Reedy, R. C., Stonestrom, D. A., Prudic, D. E. & Dennehy, K. F. Impact of land use and land cover change on groundwater recharge and quality in the southwestern US. *Global Change Biol.* **11**, 1577–1593 (2005).
45. Hong, B. *et al.* Increasing summer rainfall in arid eastern-Central Asia over the past 8500 years. *Sci. Rep.* **4**, 5279, doi: 10.1038/srep05279 (2014).
46. Cheng, X. L. *et al.* The correlation between the desertification of grassland and the change of vegetation characteristics in Eerduosi. *J. Nanjing Univ. (Nat. Sci.)* **37**, 232–239 (2001).
47. Li, X. R. Study on shrub community diversity of Ordos Plateau, Inner Mongolia, Northern China. *J. Arid Environ.* **47**, 271–279 (2001).
48. Jin, Z., Dong, Y. S., Qi, Y. C., Liu, W. G. & An, Z. S. Characterizing variations in soil particle size distribution along grass-desert shrub transition in the Ordos Plateau of Inner Mongolia, China. *Land Degrad. Dev.* **24**, 141–146 (2013).
49. Bao, S. T. *Chemical Analysis for Agricultural Soil* (China Agriculture Press, Beijing, 1999).
50. Li, M., Zhou, X., Zhang, Q. & Cheng, X. Consequences of afforestation for soil nitrogen dynamics in central China. *Agr. Ecosyst. Environ.* **183**, 40–46 (2014).

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Author Contributions
Z.J. conducted the field soil sampling; Y.Z. took and assembled the photographs of common plant species in the Ordos Sandy Grassland Research Station; X.L. conducted the soil sample chemical analysis; Z.J., X.L., Y.D. and Z.A. conducted the data analysis; and Z.J. took the photographs of the three types of dunes and prepared the sampling location map. The manuscript was written by Z.J., Y.D. and Z.A. The project was planned by Z.J. All of the authors discussed the results and commented on the manuscript.

Additional Information
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