Topographic Constraints on the Origin of the Equatorial Ridge on Iapetus

Erika J. Lopez Garcia\textsuperscript{a}, Edgard G. Rivera-Valentin\textsuperscript{a}, Paul M. Schenk\textsuperscript{b}, Noah P. Hammond\textsuperscript{a}, Amy C. Barr\textsuperscript{a}

\textsuperscript{a}Department of Geological Sciences, Brown University, Providence, RI 02912
\textsuperscript{b}Lunar and Planetary Institute, Houston, TX 77058

Abstract

Saturn’s moon Iapetus has an equatorial ridge system, which may be as high as 20 km, that may have formed by endogenic forces, such as tectonic and convective forces, or exogenic processes such as debris infall. We use high-resolution topographic data to conduct a topographic analysis of the ridge, which suggests a predominantly triangular morphology, with some ridge face slopes reaching 40°, allowing for an exogenic formation mechanism.

Keywords: Iapetus, Satellites, surfaces, Geological processes

1. Introduction

Low resolution \textit{Voyager 2} images show that Saturn’s outermost regular satellite, Iapetus, has unusually high topography (Denk et al., 2000). Later, high-resolution \textit{Cassini} images show several of the peaks identified from \textit{Voyager 2}, and illustrate that those peaks are part of an equatorial ridge system extending over 110° in longitude (Porco et al., 2005). Previous studies report that parts of the ridge reach a height of 20 km (as determined from limb measurements) and could be up to 70 km wide (Porco et al., 2005; Giese et al., 2008). Parts of the ridge are discontinuous and some sections have up to
three nearly parallel ridge sections (Porco et al., 2005; Ip, 2006; Giese et al., 2008). No significant feature pre-dating the ridge system has been discovered (Porco et al., 2005; Neukum et al., 2005; Ip, 2006; Castillo-Rogez et al., 2007; Giese et al., 2008), attesting to its old age.

The location of the ridge at Iapetus’s equator may suggest an endogenic origin. Some suggest tectonism and volcanism as possible endogenic formation mechanisms. Despinning can create complex tectonic patterns on solid planetary bodies (Pechmann and Melosh, 1979; Porco et al., 2005); however, the tectonic features predicted by this mechanism have not been observed (Porco et al., 2005; Castillo-Rogez et al., 2007; Singer and McKinnon, 2011). Giese et al. (2008) suggest that the ridge may be tectonic because some slopes measured at the sides of the ridge are less than the angle of repose for ice, assuming the ridge material had been deposited. The possible tectonic activity could be related to the upwarping of the surface on either side of the ridge (Giese et al., 2008). Others argue the appearance of the ridge is consistent with its formation by material rising from below. Denk et al. (2005) suggest the ridge is a result of extrusive volcanic activity.

Exogenic hypotheses argue that debris infall may have formed the ridge. Ip (2006) proposes that the ridge represents the remains of an early ring system that fell onto the equator. The resultant debris should have a slope close to the angle of repose, which may be between the values for two possible models of ridge material: rounded grains, for which the angle of repose is about ∼25° (Kleinhans et al., 2011), or for snow mixed with particles of hail, for which the angle of repose ∼ 45° (Abe, 2004). The cross-sectional morphology should show a single prominent peak and be uniform as a function
of longitude (Ip, 2006); however, later impact modification would change the morphology of the ridge. Levison et al. (2011) suggest a scenario in which an ancient impact formed a sub-satellite and an impact-generated disk around Iapetus. The remains of the disk fell onto the surface to form the ridge, and the sub-satellite later impacts Iapetus to form one of its large basins during its tidal evolution (Levison et al., 2011). Dombard et al. (2012) also suggest the ridge formed from a sub-satellite. They propose that tidal forces decay and tear the sub-satellite apart, and its debris falls onto the equator. The discovery of a possible ridge system on Rhea (Schenk et al., 2011), may suggest that these types of events are common (Dombard et al., 2012).

Past characterization of the ridge (e.g., Giese et al., 2008) has been limited by the number and quality of digital elevation models (DEMs) available. Characterizing the ridge slopes and morphology is key to deciphering the origin of the ridge. Here, we use DEMs generated from Cassini imagery (Schenk, 2010) to observe how the ridge slopes and cross-sectional morphology varies along its length.

2. Methods

2.1. Profiles

We use eight Cassini Imaging Science Subsystem (ISS) images with corresponding DEMs, which are constructed by a combination of stereo and photoclinometry, of the ridge system within Iapetus' dark terrain to obtain topographic profiles. We focus our morphological study on the dark terrain, where the topography of the ridge is most readily calculated (Schenk, 2010; White and Schenk, 2013). Cassini ISS coverage of the equatorial ridge, as
well as other parts of Iapetus, was used to produce these topographic maps (Schenk, 2010). The images range in size and resolution, allowing us to study the ridge on both a local and regional scale.

The wide range of spatial coverage allows observations of the ridge as a semi-continuous system across many kilometers and to closely observe its morphology. We obtained a total of 506 topographic profiles across all eight DEMs.

Figure 1 illustrates the footprints of the study images and the locations of the topographic profiles. Depending on the size of the region covered by each DEM, profiles were taken between $-10^\circ$ to $10^\circ$ latitude, which is about 250 km. There was consistent spacing between the profiles, with a maximum of $0.1^\circ$ latitude. Profiles were taken perpendicular to the ridge and include both the ridge and adjacent plains. Three points on the ridge, including the peak and an endpoint, are used to divide the ridge topography into four segments: northern and southern ridge faces and flanks. For example, in the trapezoidal ridge example in Fig. 2b, the transition from northern flank to face occurs at about $2^\circ$ latitude and the transition from southern face to flank is close to $-4^\circ$ latitude. Least squares fitting was used to determine the best fit line to the topography and corresponding slope angle of each segment with its corresponding standard error of the mean (SEM) to a 95% confidence. Error bars reported here are upper bound values and encompass propagating any errors from lighting or viewing geometry effects. We report positive slopes as north-facing slopes and negative slopes as south-facing slopes.
2.2. Morphology classification

Profile segments that were long enough to cover both the ridge and the adjacent plains are classified based on its morphology. We identify six morphology types defined by the shape of the peak, number of peaks, and the number of different types of peaks found in a single profile: triangular, trapezoidal, crowned, twinned, dissimilar and saddle. An example profile for each morphology is shown in Figure 2.

A triangular peak, as observed by previous work (Castillo-Rogez et al., 2007; Giese et al., 2008), is characterized by its distinct sharp peak with smooth sides on either face of the ridge. The trapezoidal morphological type described by Giese et al. (2008) has a pronounced tabular top, similar to the famous Table Mountain in Cape Town, South Africa (Luckhoff, 1951).

A crowned peak, which has been previously described as a group of three parallel ridges (Porco et al., 2005; Ip, 2006; Giese et al., 2008), is characterized by three triangular peaks, with the middle peak being taller than the two flanking peaks. A twinned morphology has two peaks side by side, where each peak has a similar shape. In some cases, we find two different morphology types in a single profile – often, one triangular peak and one smaller trapezoidal peak – which we name “saddle” peaks. Dissimilar peaks are peaks that have at least three peaks of different sizes and shapes. We also find some profiles where the ridge wasn’t obviously visible or had no distinct morphology.
3. Results and Discussion

We find the triangular peak is the most common morphology, observed in 33% of all of the profiles with a distinguishable shape. For this morphology, the average slopes for the northern and southern faces are similar to previous results, $\sim 15^\circ$ ([Giese et al., 2008]), ranging from $\sim 3^\circ$ to $\sim 39^\circ$ (see Table 1). The trapezoidal peak morphology is the second most frequent, seen in 21% of profiles. Saddle peaks occur in 17% of the profiles and also have face slopes similar to triangular peaks. Twinned peaks occur in 14% of the profiles and crowned peaks are observed in 8% of the profiles. Lastly, the dissimilar ridge morphology is the least common morphology (7% of profiles).

Because we find the triangular shape to have the steepest slopes, and the sharpest peaks, our results suggest the triangular shape could be the least impact-modified ridge morphology on Iapetus. Trapezoidal ridge morphologies could be the result of landslides (e.g., [Singer et al., 2012]), or impact modification. We also suggest that twinned, crowned, and dissimilar peaks formed by further impact modification or landslides ([Singer et al., 2012]). The frequency of trapezoidal and saddle ridge morphologies are similar (21% and 17%), possibly indicating that trapezoidal and saddle peaks represent degraded triangular peaks, which could be modified by landslides ([Singer et al., 2012]). The other less frequent categories have extremely shallow slopes, $< 13^\circ$, suggesting significant erosion. Generally, there is not a typical transition between morphologies.

We find the flanks adjacent to the ridge have a significant slope. Specifically, the southern flank slopes for triangular, trapezoidal, crowned, and saddle peaks are high, $\sim 5^\circ$. This may indicate that the ridge system ex-
tends further out than previously thought (Porco et al., 2005; Giese et al., 2008), may be built upon a pre-existing uplift or bulge, and/or may be more massive than originally estimated. There seems to be no sharp boundary between the faces of the ridge and the plains.

4. Conclusions

Using high-resolution imagery data, we conduct a topographic survey of the equatorial ridge of Saturn’s moon Iapetus. We obtained 506 topographic profiles across eight Cassini ISS images with corresponding DEMs to calculate the slope of the northern and southern faces and flanks of the ridge.

In addition to the three morphological types already classified - triangular, trapezoidal, and crowned peaks (Porco et al., 2005; Castillo-Rogez et al., 2007; Giese et al., 2008), we find three additional ridge shapes: a twinned morphology, a dissimilar morphology, and a saddle morphology. Previous work had ruled out exogenic formation processes due to shallow ridge slopes and varied morphology (Giese et al., 2008); however, we find that the most common morphology of the ridge is a triangular peak with face slopes reaching $\sim 40^\circ$. Slopes on triangular ridge faces are close to the angle of repose, which can vary between $\sim 45^\circ$ for snow with small particles of hail (Abe, 2004) or $\sim 25^\circ$ for rounded grains (Kleinhans et al., 2011), for $\sim 10\%$ of the topographic profiles. This may point to an exogenic origin for the ridge, such as formation by debris infall.

Iapetus is among the most unusual planetary bodies in the Solar System. Although the formation of its equatorial ridge remains mysterious, the most common triangular ridge morphology and the evidence of slope angles close
to the angle of repose make the case for an exogenic origin more plausible.

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References

Abe, Osamu. Shear strength and angle of repose of snow layers including graupel. Annal. Glaciol. 38 1, 305-308.

Castillo-Rogez, J. C., Matson, D. L., Sotin, C., Johnson, T. V., Lunine, J. I., Thomas, P. C., 2007. Iapetus’ geophysics: Rotation rate, shape, and equatorial ridge. Icarus 190, 179–202.

Denk, T., Matz, K., Roatsch, T., Wolf, U., Wagner, R., Neukum, G., Jaumann, R., 2000. Iapetus (1): Size, topography, surface structures, craters abstract no. 1596. In: Lunar and Planet. Sci. Conf. Abstracts XXXI. Lunar and Planetary Institute, Houston, Tex.

Denk, T., Neukum, G., Roatsch, T., McEwen, A., Turtle, E., Thomas, P., Helfenstein, P., Wagner, R., Porco, C., Perry, J., 2005. First Imaging Results from the Iapetus B/C Flyby of the Cassini Spacecraft abstract no. 2268. In: Lunar and Planet. Sci. Conf. Abstracts XXXVI. Lunar and Planetary Institute, Houston, Tex.
Dombard, A. J., Cheng, A. F., McKinnon, W. B., Kay, J. P., 2012. Delayed formation of the equatorial ridge on Iapetus from a subsatellite created in a giant impact. Journal of Geophysical Research 117, E03002.

Giese, B., Denk, T., Neukum, G., Roatsch, T., Helfenstein, P., Thomas, P., Turtle, E., McEwen, A., Porco, C., 2008. The topography of Iapetus’ leading side. Icarus 193, 359–371.

Ip, W.-H., 2006. On a ring origin of the equatorial ridge of Iapetus. Geophysical Research Letters 33, L16203.

Kleinhans, M. G., Markies, H., de Vet, S. J., in ’t Veld, A. C., Postema, F. N., 2011. Static and dynamic angles of repose in loose granular materials under reduced gravity. Journal of Geophysical Research 116, E11004.

Levison, H., Walsh, K., Barr, A., Dones, L., 2011. Ridge formation and despinning of Iapetus via an impact-generated satellite. Icarus 214, 773–778.

Luckhoff, C. A., 1951. Table Mountain: Our national heritage after three hundred years. AA Balkema, Cape Town.

Neukum, G., Wagner, R., Denk, T., Porco, C. C., Team, t. C. I., 2005. The cratering record of the Saturnian satellites Phoebe, Tethys, Dione, and Iapetus in comparison: First results from analysis of the Cassini ISS imaging abstract no. 2034. In: Lunar and Planet. Sci. Conf. Abstracts XXXVI. Lunar and Planetary Institute, Houston, Tex.

Pechmann, J. B., Melosh, H. J., 1979. Global fracture patterns of a despun planet: Application to Mercury. Icarus 38 (2), 243–250.
Porco, C., Baker, E., Barbara, J., Beurle, K., Brahic, A., Burns, J., Charnoz, S., Cooper, N., Dawson, D., Del Genio, A., 2005. Cassini imaging science: Initial results on Phoebe and Iapetus. Science 307, 1237–1242.

Schenk, P., Hamilton, D. P., Johnson, R. E., McKinnon, W. B., Paranicas, C., Schmidt, J., Showalter, M. R., 2011. Plasma, plumes and rings: Saturn system dynamics as recorded in global color patterns on its midsize icy satellites. Icarus 211, 740–757.

Schenk, P. M., 2010. Global Topographic Mapping Of Saturn’s Midsize Icy Satellites: System-wide Thermal And Impact Effects, abs. no. 09.16. In: AAS/Division for Planetary Sciences Meeting Abstracts #42.

Singer, K. N., McKinnon, W. B., 2011. Tectonics on Iapetus: Despinning, respinning, or something completely different? Icarus 216, 198–211.

Singer, K. N., McKinnon, W. B., Schenk, P. M., Moore, J. M., 2012. Massive ice avalanches on Iapetus mobilized by friction reduction during flash heating. Nature Geoscience 5, 574–578.

White, O. L., Schenk, P. M., 2013. Impact basin relaxation on Rhea and Iapetus and relation to past heat flow. Icarus 223, 699–709.
|                   | North Flank | North Face | South Face | South Flank |
|-------------------|-------------|------------|------------|-------------|
| **Triangular**    |             |            |            |             |
| average           | 3.0° ± 1.1° | 15.2° ± 1.4° | −16.5° ± 1.7° | −4.0° ± 1.5° |
| range             | −9.9° to 13.8° | 3.8° to 29.4° | −38.8° to −2.7° | −19.4° to 6.9° |
| **Trapezoidal**   |             |            |            |             |
| average           | 3.8° ± 1.0° | 15.9° ± 1.5° | −16.8° ± 1.8° | −4.0° ± 1.4° |
| range             | −9.7° to 11.3° | 3.3° to 29.2° | −39.9° to −5.7° | −16.2° to 5.7° |
| **Crowned**       |             |            |            |             |
| average           | 3.8° ± 1.1° | 13.7° ± 3.3° | −13.1° ± 3.6° | 6.3° ± 2.2° |
| range             | 0.7° to 10.7° | −9.0° to 26.5° | −36.1° to 7.1° | −13.9° to 4.3° |
| **Twinned**       |             |            |            |             |
| average           | 4.5° ± 1.4° | 8.9° ± 3.1° | −8.0° ± 3.2° | −2.7° ± 1.3° |
| range             | −5.5° to 14.5° | −10.0° to 23.7° | −29.7° to 10° | −13.3° to 7.2° |
| **Dissimilar**    |             |            |            |             |
| average           | 3.1° ± 2.7° | 11.9° ± 4.8° | −10.1° ± 3.9° | −2.8° ± 1.7° |
| range             | −12.2° to 12.9° | −16.0° to 32.3° | −22.3° to 8.2° | −9.2° to 5.7° |
| **Saddle**        |             |            |            |             |
| average           | 4.3° ± 1.2° | 14.8° ± 1.5° | 14.6° ± 1.9° | −5.6° ± 1.7° |
| range             | −5.3° to 14.6° | 1.3° to 22.3° | −33.9° to −3.2° | −16.3° to 4.4° |

Table 1: Average and range of slope angles for the six ridge morphology types.
Figure 1: A global map of Iapetus (image PIA11116 resolution of 0.8 km/pxl) with the location of each image (except images 1, 7 and 8) and the types of ridge morphology observed. Red is triangular peak, pink is trapezoidal peak, green is crowned peak, blue is twinned peak, cyan is dissimilar peak, yellow is saddled peak, and white is no discernible peak. These are the image names that are on the global map: (2) W1568130503 (resolution 0.41 km/pxl); (3) W1568125439 (resolution 0.32 km/pxl); (4) N1483151512 (resolution 0.90 km/pxl); (5) W1568133373 (resolution 0.79 km/pxl); and (6) N1482859934 (resolution 6.7 km/pxl). Images (1) W1568127165 (resolution 0.09 km/pxl, latitude and longitude $-4.48083^\circ$, 160.992°, 5.21678°, 167.332°), (7) N1568127472 (resolution 0.009 km/pxl, latitude and longitude $-0.502969^\circ$, 165.907°, 0.65417°, 166.711°), and (8) N1568127660 (resolution 0.809 km/pxl, latitude and longitude $-0.36639^\circ$, 165.015°, 0.65908°, 165.939°) are too small to be depicted on this map.
Figure 2: Representative examples of the six ridge morphological types observed in the topographic profiles. Vertical exaggeration $\sim 10$ times.