The rise in the frequency of anthropogenic earthquakes due to deep fluid injections is posing serious economic, societal, and legal challenges to many geo-energy and waste-disposal projects. Existing tools to assess such problems are still inherently heuristic and mostly based on expert elicitation (so-called clinical judgment). We propose, as a complementary approach, an adaptive traffic light system (ATLS) that is function of a statistical model of induced seismicity. It offers an actuarial judgement of the risk, which is based on a mapping between earthquake magnitude and risk. Using data from six underground reservoir stimulation experiments, mostly from Enhanced Geothermal Systems, we illustrate how such a data-driven adaptive forecasting system could guarantee a risk-based safety target. The proposed model, which includes a linear relationship between seismicity rate and flow rate, as well as a normal diffusion process for post-injection, is first confirmed to be representative of the data. Being integrable, the model yields a closed-form ATLS solution that is both transparent and robust. Although simulations verify that the safety target is consistently ensured when the ATLS is applied, the model from which simulations are generated is validated on a limited dataset, hence still requiring further tests in additional fluid injection environments.

A significant proportion of the world’s global energy production relies on subsurface resources, such as oil, gas and coal production, as well as geothermal energy. In addition, the deep underground is increasingly used for waste storage; typical examples are wastewater from fracking operations and CO₂ sequestration. However, these technologies are not “risk-free,” as shown by the increased frequency of induced seismicity cases around the globe. Recent examples include induced seismicity related to fracking and wastewater disposal, gas extraction, gas storage, CO₂ sequestration and renewable geo-energy. Although some jurisdictions have enforced the use of maximum magnitude thresholds to limit the induced seismicity risk, most of these rules remain heuristic. In this article, we argue that quantitative risk assessment and mitigation strategies rather than trial-and-error methods, should be essential tools to guarantee safety for society. The approach we advocate, which should be seen as a proof-of-concept, allows for an informed risk-cost-benefit analysis involving all stakeholders.

Traffic light systems (TLS) are commonly used to mitigate induced seismicity risk by modifying the fluid injection profile. A TLS is based on a decision variable (earthquake magnitude, peak ground velocity, etc.) and a threshold value above which actions (e.g. stopping the injection or reducing production rates) must be taken. Currently, the definition of this threshold is based on expert judgment and regulations. Here, we propose a data-driven adaptive TLS, termed ATLS, which aims to overcome the limitations of the traditional heuristic methods. Here, the assignment of a magnitude threshold is based on a quantitative risk assessment, subject to a safety criterion imposed by the authorities (e.g., fixed probabilities of unacceptable nuisance, damage or fatalities). As a consequence, the ATLS is an objective and statistically robust mitigation strategy, which facilitates a fair and transparent regulatory process. This approach is in line with the procedures common for most other technological risks, such as in the hydropower, nuclear or chemical industries. Model-based forecasting and alerting are already advocated elsewhere, such as in hurricane data assimilation and forecasting.
Results
Predictive hazard model. A predictive model lies at the heart of any risk assessment. In the case of induced seismicity, a wide range of statistical and physics-based models exists\textsuperscript{16,17}. Undoubtedly, more work is needed to develop, calibrate and validate new models; however, we believe that the missing link is the use of such models for deriving and monitoring quantitative risk thresholds. Here, we use as example a simple and yet robust model that forecasts the piecewise induced seismicity temporal, here daily, rate $\lambda(t, m \geq m_b; \theta)$ as:

$$
\lambda(t, m \geq m_b; \theta) = \begin{cases} 
10^{\phi_b + b m_b V(t)} & ; t \leq t_{\text{shut-in}} \\
10^{\phi_b + b m_b V(t_{\text{shut-in}} - t)} \exp\left(\frac{t - t_{\text{shut-in}}}{\tau}\right) & ; t > t_{\text{shut-in}} 
\end{cases}
$$

(1)

where $V(t)$ is the injection flow rate as a function of time $t$ in m$^3$/day, $\theta = [\phi_b, \alpha_b, \tau]$ a set of model parameters describing the underground characteristics (earthquake size ratio, activation feedback in m$^{-3}$ and mean relaxation time in days, respectively), $m_b$ the minimum magnitude cutoff, and $t_{\text{shut-in}}$ the shut-in time, also in days. Both $\alpha_b$ and $b$ can also be functions of time $t$ (see Decision variable section). In this model, the injection or operation phase is described by a linear relationship between $\lambda(t, m \geq m_b)$ and $V(t)$ in line with previous studies\textsuperscript{17–19}. It derives directly from the linear relationship between $V$ and overpressure\textsuperscript{41}, hence assuming no change of injectivity during any given stimulation. The post-injection phase is described by a pure exponential decay representative of a normal diffusion process\textsuperscript{47}. Although the Modified Omori Law is sometimes used to describe post-injection seismicity\textsuperscript{40}, reasons remain mostly historical\textsuperscript{21,22}. The proposed alternative is verified to be consistent with the tested data (see Methods section) and preferred on analytical grounds, being directly integrable in contrast with the Modified Omori Law, which is conditional on parameter values and may require the definition of an \textit{ad hoc} upper bound\textsuperscript{22}. Finally, no maximum magnitude $M_{\text{max}}$ is imposed. It follows from Eq. (1) that induced seismicity is characterized by the injection profile $V(t)$, and the underground feedback described by the three-parameter set $\theta$. The main limitations of the proposed model are discussed in detail later on.

In contrast with complex fluid modelling\textsuperscript{16}, the closed-form Eq. (1) can be computed on-the-fly; moreover, it includes the mean relaxation time, $\tau$, hence taking into account the long-term underground feedback after shut-in. Finally, being integrable, it leads in turn to a closed-form ATLS, as demonstrated in the next section. Although the physical process governing the rate of induced seismicity is more complex than what is represented by Eq. (1), this rate model is proven to be valid in a Poissonian probabilistic setting (see Methods section). Moreover, the physical processes are either not still clear (in fact, there is not an unanimous consensus among scientists), or computationally expensive. Therefore, pragmatism imposes the use of statistical models until both an agreement is found on the physics of induced seismicity and computational time of complex physical modelling is reduced.

The model (Eq. 1) was fitted to six induced seismicity sequences observed in fluid injection experiments from enhanced geothermal systems (EGS)\textsuperscript{13,23–26}, the initial stage of a long-term brine sequestration\textsuperscript{27}, and one fracking at an oil field\textsuperscript{28} (Table 1). The model succeeds to describe most of the data as shown in Figure 1 (see results of statistical tests in the Methods section – Note that the rare outliers above $3\sigma$ may be due to missing on-site data that may affect seismicity, such as unknown technical operations on wells, or to second-order physical processes missing in Eq. (1), as discussed below). The parameters are found to range over $0.8 \leq b \leq 1.6$, $-2.8 \leq \phi_b \leq 0.1$ m$^{-3}$ and $0.2 \leq \tau \leq 20$ days and show a relatively large scattering between sites and between different stimulations at a same site (Table 2). It is important to note that we specifically chose those six datasets, as they had been made publicly available. Those cases are characterized by high pressures and flow rates, and are rich in induced earthquakes. This represents a selection bias and the $\phi_b$ parameter could decrease to much lower values elsewhere (Fig. 2). For instance, most injection wells in the U.S. do not cause felt earthquakes\textsuperscript{29}; large $\phi_b$ variations between regions and sites might be explained by different regional crustal stresses\textsuperscript{30}. The present sites provide however natural laboratories to test our model and the associated ATLS, without generalizing or inferring any high level of risk for all existing deep fluid injections. Figure 2 illustrates the parameters’ scattering, including results from past studies\textsuperscript{18} for EGSs and other hard rock settings enlarging the range of values to $0.7 \leq b \leq 3$ and $-4.2 \leq \phi_b \leq 0.4$ m$^{-3}$. It has been shown that in fracking environments, the activation feedback can be as low as $\phi_b = -9.25$ m$^{-3}$\textsuperscript{30}.

| Site (country ISO code), year | Injection profile | Earthquake catalog |
|-----------------------------|------------------|-------------------|
| Basel (CH), 2006            | Digitized from (13) | (23) |
| Garvin (US), 2011           | Digitized from (28) | (28) |
| KTB (DE), 1994              | Digitized from (24) | (24) |
| Paradox Valley (US), 1994   | http://www.usbr.gov/uc/wcao/progact/paradox/RI.html | http://www.usbr.gov/uc/wcao/progact/paradox/RI.html |
| Newberry (US), 2012         | Digitized from (25) | http://fracture.lbl.gov/cgi-bin/Web_CatalogSearch.py |
| Newberry (US), 2014         | Digitized from (26) | http://fracture.lbl.gov/cgi-bin/Web_CatalogSearch.py |

Table 1. Source” of stimulation experiment datasets. *Online material last assessed in June 2017.
Safety criterion. A safety criterion is a probability of exceedance that can be fixed with respect to different safety metrics, such as fatalities, economic loss, building damage or level of nuisance. Given the selected metric, the corresponding safety criterion can be converted in the magnitude space into the probability of exceedance \( \Pr(m \geq m_{\text{saf}}) = Y \), which will ensure that the acceptable level of risk is preserved at all time:

**Figure 1.** Induced seismicity model fitting of six fluid injection experiments: 1994 German Continental Deep Drilling Program (KTB); 1994 Paradox Valley, United States; 2006 Basel, Switzerland; 2011 Garvin, United States; and 2012–2014 Newberry, United States. For both KTB and 2014 Newberry, experiments are broken down into two separate stimulations, each with its own post-injection tail. The model (Eq. 1) is represented by the red curves on the induced seismicity time series with the ±3σ uncertainty envelope shown in light red. Dashed and dotted vertical lines indicate the shut-in times and sub-stimulation periods, respectively.

| Experiment | \( m_b \) | \( b \) | \( a_b \) (m/s) | \( \Sigma^\ast \) | \( \tau \) (days) |
|------------|--------|------|----------------|--------|--------|
| B06        | 0.8    | 1.58 | 0.10           | 0.4    | 1.12   |
| G11        | 1.0    | 0.77 | -1.35          | N/A    | 0.28   |
| KTB94a     | -1.5   | 0.98 | -1.35          | -1.65  | 0.03†  |
| KTB94b     | -1.4   | 0.87 | -1.65          | -1.65  | 0.22   |
| PV94       | 0.6    | 1.08 | -2.40          | -2.6   | 14.13  |
| NB12       | 0.2    | 0.80 | -2.80          | N/A    | 12.59  |
| NB14a      | 0.0    | 0.98 | -1.60          | N/A    | 3.55   |
| NB14b      | 0.2    | 1.05 | -1.60          | N/A    | 3.16   |

Table 2. Maximum likelihood estimates per stimulation experiment. *Seismogenic index obtained by ref.18; †unreliable, part of the tail being likely hidden by the KTB94b sequence.
with $m_{\text{safe}}$ the magnitude at which the given safety limit (e.g., damage, fatality) is reached. Note that Eq. (2) is a closed-form expression where $V(t)$ is the cumulative injected fluid volume, and $m_{\text{safe}}$ and $Y$ are derived from the safety criterion (see Methods section). Moreover, the set of parameters $\theta$ is updatable at any given time. The mapping from risk to earthquake magnitude is required to control injection operations based on short-term observations (see Decision Variable section). While peak ground velocity (PGV) is a more direct measure, conversion to magnitude is in any case inevitable to estimate the risk potential of larger earthquakes from the $b$-value. It is also a unique measure, while PGV requires a location that is not trivial to assign. It should be added that no maximum magnitude $M_{\text{max}}$ is imposed in Eqs (1–2). This implicitly assumes that both small to medium-size induced events and large triggered earthquakes on existing faults are treated the same way. This remains debated, although a recent study demonstrated that the observed $M_{\text{max}}$ in fluid injections is compatible with the null-hypothesis of the Gutenberg-Richter law with no upper limit. The role of $M_{\text{max}}$ (and therefore of triggered earthquakes) is however only critical when the risk of fatalities (e.g., individual risk) is evaluated. For nuisance or minor damage thresholds, risk is more likely dominated by medium-size induced events. This important discussion has no significant impact on the method proposed, as proved in the Methods section.

Before an ATLS is set, the likelihood of failure of the planned project with respect to a specified limit state function defined by the safety criterion in magnitude space (Eq. 2) can approximately be determined. In this study, we select as main metric the annual individual risk ($\text{IR}$) over the entire project period, and as safety criterion $\text{IR} \leq 10^{-6}$ (i.e. the probability that a statistically representative individual dies for the introduced hazard),

$$\Pr\left(m \geq m_{\text{safe}}\right) = 1 - \exp\left\{-10^{b_m - b_m_{\text{safe}}} [V(t_{\text{shut-in}}) + \tau V(t_{\text{shut-in}})]\right\} = Y$$

Figure 2. Three-parameter set $\theta = [b, a_{\text{fb}}, \tau]$ scattering in fluid injection experiments & impact on project validation for fixed safety threshold ($\text{IR} \leq 10^{-6}$) and different fluid injection scenarios. Dots represent $\theta$ estimates obtained in the present study (Fig. 1), squares the ones obtained by ref. (with no information on $\tau$) and curves (made of successive dots) time-dependent estimates obtained where the earthquake catalogue is large enough (see Methods section). The hypothetical project injects a total volume $V = 10,000$ m$^3$ of fluids with constant flow rate $\dot{V}$ at a distance $d$ from the nearest building: a. $\dot{V} = 1$ m$^3$/min, $d = 0$ km; b. $\dot{V} = 1$ m$^3$/min, $d = 50$ km; c. $\dot{V} = 10$ m$^3$/min, $d = 0$ km and d. $\dot{V} = 10$ m$^3$/min, $d = 50$ km. Lines represent the safety threshold for different values of $\tau$ (in days) (Eq. 2). Colour green or red indicates if the safety target is respected or not, respectively (darker colours are used in the case of time-dependent estimates). Letters represent sites: Basel (B), Cooper Basin (CB), Garvin (G), KTB, Newberry (NB), Ogachi (O), Paradox Valley (PV) and Soultz (S), followed by the experiment year’s last two digits.
which is a threshold commonly enforced for hazardous installations\(^{14}\). Figure 2 shows the acceptable domain for a fixed limit state function, tested for different injection scenarios in which a hypothetical project plan is to inject a total volume \(V = 10,000\) m\(^3\) of fluids at a depth of 4 km\(^{-1}\). The injection profile is assumed to be flat with a constant flow rate \(\dot{V} = 1\) or 10 m\(^3\)/min, having an impact on injection duration and tail behaviour (Eq. 2; Fig. 2). The project is considered to be located at a distance \(d = 0.1\) or 50 km from the nearest building. For a given site, with no knowledge of the underground feedback to fluid injection, project operators and regulators in an EGS setting (where high seismicity rates are common during stimulation) could use the known \(\theta\) scattering for an \textit{a priori} parameterization. This preliminary assessment shows the likelihood of the project to pass or fail the safety threshold. As shown in Figure 2, the results can be ambiguous, due to the large uncertainties associated with subsurface characteristics. Nevertheless, it provides a preliminary assessment of the risk reflecting the limited knowledge of the induced seismicity process. Future estimations of \(\theta\) at additional injection sites will likely refine the results and improve the decision process. In addition, rules of decision-making under uncertainty can account for that ambiguity\(^{16,39}\). Decisions become more obvious in cases in which the diagram would be entirely green (clear go) or red (clear no-go). Underground stimulation activities in areas with low exposure (e.g. remote EGS plant locations with large distance \(d\) from the nearest habitations) would evidently have a lower induced-seismic risk and, thus, shrink the red area. The termination of the 2006 Basel EGS project was due to the high induced-seismic risk emerged from the high exposure of the urban built environment\(^{27}\). Note that Eq. (2) can be used to predict a distance \(d\) for which the induced-seismic risk would become acceptable—conditional to a given injection profile \(\dot{V}(t)\) and parameter set \(\theta\) (since \(m_{th}\) is a function of \(d\); see Methods section).

**Decision variable.** The ATLS decision variable must be selected and updated as new data allows estimating \(\theta\) more accurately, or if the planned injection scheme is changed. Here, a threshold earthquake magnitude \(m_{th}\) is used as decision variable. In particular, \(m_{th}\) is defined as the magnitude value for which mitigating actions must be taken, here corresponding to stopping injection, i.e.:

\[
m_{th} = \frac{1}{k} \log_{10}(Y - 10^{\theta_{fb saf}} \dot{V}\tildet_{dus-in}) + m_{mf}
\]

(3)

(see Methods section). If \(m_{th}\) is updated “on-the-fly”, the project is guaranteed to meet the defined safety criterion.

To avoid reaching \(m_{th}\) before the planned stop of the fluid injection one may be inclined to reduce the flow rate \(\dot{V}\), but here it would only delay the time at which the injection must stop, as the risk is mostly controlled by total volume injected \(V\) (Eq. 2; the secondary role of \(\dot{V}\) is highlighted in Figure 2 by the change of width of the yellow band, for different \(\dot{V}\) and \(\tau\)). It is common practice to reduce \(\dot{V}\) however\(^{13}\), but results of such action remain unclear since verification of a safety threshold requests a large number of experiments (see simulations below); indeed: (i) It is plausible that such action has no overall effect, the risk remaining the same in average over a fixed \(V\); (ii) If such action has an effect, it would indicate that the model of Eq. (1) does not properly describe the role of different injection strategies. Despite the proposed model being classified as a pure statistical method, it is based on physical considerations. Its first term, for example, builds on the linear relationship between volume change and pressure\(^{37}\). Other relationships can be envisioned such as a bilinear relationship indicative of a change of injectivity\(^{26}\), which may explain some second-order relationship observed between flow rate and \(M_{max}\)\(^{35}\). The linear relationship could also be shifted in time by including a minimum pressure threshold\(^{37}\) below which no induced seismicity is triggered. Adding such processes would likely allow for smarter mitigation strategies in which the shape of the injection profile would play a role. Since any model change would require the inclusion of additional parameters (which have yet to be constrained), and since Eq. (1) is verified to be consistent with most of the data tested, we consider Eq. (1) to be a reasonable first-order model for the proposed ATLS.

To validate the ATLS in a realistic time-dependent setting, we simulate the Basel induced seismicity sequence using Eq. (1) with \(\dot{V}\tildet_{dus-in}(t)\) (Figs 1; 3a) and \(\theta_{tildet_{dus-in}}(t) = [b(t), a(t), \tau = 1.12\) day\) (Figs 2; 3b; Table 2). We use the safety threshold \(IR \leq 10^{-6}\) with the nearest building above the borehole (\(d = 0\) km), yielding \(m_{mf} = 5.8\) (see Methods section). The value \(m_{th}\) as a function of time is shown in Figure 3c with two examples of induced seismicity time series. The first one, in grey, is a reproduction of the 2006 Basel experiment; the second one, coloured, is the case where the ATLS is used. Figure 3d finally shows that the safety target, while not respected for synthetic versions of the Basel experiment, is reached once the newly proposed ATLS (Eq. 3) is considered. Due to the stochastic nature of the earthquake process, the operational safety target is only reached on average over multiple sequences. It is plausible that an improved model with a non-linear relationship between flow rate and overpressure would open the possibility for mitigation strategies based on different injection profiles, hence potentially avoiding prematurely stopping the stimulation and the project.

**Discussion**

The main purpose of this study was to present a statistical-based first order ATLS, which verifies that a quantitative safety criterion is ensured. An important benefit of the outlined ATLS approach to both operators and regulators is its transparency and execution speed (being a suite of simple closed-form expressions). By principle, its adoption would make any project in compliance with the safety threshold whatever the response of the underground. This does not ensure that a project is financially successful, but it gives the operator the maximum allowable chance to reach success, based on a quantitative risk-based method.

It is important to note that existing traffic light systems based on heuristics can already provide reasonable results (for instance, the magnitude threshold \(m_{th} = 2.9\) fixed during the Basel experiment\(^{14}\) is not far off the \(m_{th}\) computed with the new ATLS; Fig. 3c). However, magnitude thresholds enforced by different jurisdictions vary significantly (with \(0.5 \leq m_{th} \leq 4^{19}\) with no clear link with the standard risk-based safety criteria used in other hazardous industries\(^{14}\). Although the application of the ATLS in the decision process might appear at first
complex, there is substantial evidence that the algorithmic (or actuarial) approach is superior to the so-called clinical approach of informal judgement\textsuperscript{36,37}. This must apply too to induced seismicity prognostics, experts necessarily basing their judgements, consciously or intuitively, on past observations shown here to be reasonably well described by Eq. (1) (Fig. 1 and Methods section). Generic parameter values are here updated after the first 100 events are observed. Once injection stops at \( t_{\text{shut-in}} \), \( b \) is assumed to remain constant since no control on induced seismicity is possible after that time; (c) Simulated version of the 2006 Basel time series (in grey) and shorten (coloured) when injection is stopped by having an event with magnitude \( m > m_{\text{th}} \) (in green); (d) ATLS validation showing that the frequency of events with \( m \geq m_{\text{th}} \) tends to \( Y \) over many simulations (fluctuations around \( Y \) represent the inherent uncertainty of the earthquake Poisson process, which decrease with increasing sample size).

Figure 3. ATLS validation on synthetics of the 2006 Basel fluid injection experiment. (a) Observed time series and flow rate \( \dot{V}(t) \) during the 2006 Basel experiment\textsuperscript{13,23}; (b) Maximum likelihood estimates (MLE) of ground parameters \( a_b \) and \( b \) over time determined from the Basel time series using a moving window of 100 events (see Methods section). Generic parameter values are here updated after the first 100 events are observed. Once injection stops at \( t_{\text{shut-in}} \), \( b \) is assumed to remain constant since no control on induced seismicity is possible after that time; (c) Simulated version of the 2006 Basel time series (in grey) and shorten (coloured) when injection is stopped by having an event with magnitude \( m > m_{\text{th}} \) (in green); (d) ATLS validation showing that the frequency of events with \( m \geq m_{\text{th}} \) tends to \( Y \) over many simulations (fluctuations around \( Y \) represent the inherent uncertainty of the earthquake Poisson process, which decrease with increasing sample size).

Methods

Time series analysis. The induced seismicity temporal rate model \( \lambda(t) \) of Eq. (1) is fitted to the induced seismicity data by using the maximum likelihood estimation (MLE) method\textsuperscript{38}. The probability \( p_i \) that the observed number \( n_i \) of induced earthquakes results from a Poisson process\textsuperscript{39} with rate \( \lambda_i \) is

\[
P_i = \frac{\lambda_i^{n_i} \exp(-\lambda_i)}{n_i!}
\]

which yields the log-likelihood function.
where $X = \{n_1, \ldots, n_n, \ldots, n_{imax}\}$ is the observation set and $\theta = [a, b, \tau]$ is the parameter set of Eq. (1). The maximum likelihood estimate of $\theta$ is finally $\hat{\theta}_{MLE} = \arg \max \LL(\theta, X)$, $n_i$ is fixed to $M_1$, the completeness magnitude defined as the magnitude bin with the highest number of events $b$ is estimated independently of $\lambda$, also based on the MLE method. The model is fitted to 8 datasets (from 6 stimulations in various injection settings; Table 1); the resulting maximum likelihood estimates are listed in Table 2.

**Sensitivity analysis.** Temporal changes in $a$ and $b$ are evaluated for induced seismicity sequences that are large enough (i.e., made of hundreds of events, such as in 1994 Paradox Valley, 2006 Basel and 2014 Newberry). The parameters are estimated using a moving window with constant event number $n = 100$. Before $n$ is reached, MLE estimates obtained in retrospect (Table 2) are used, as shown in Figure 3b. In a prospective case, generic values should be used, e.g. the median or mean of values taken by the parameters in previous experiments. Since the post-injection phase is not considered in the sensitivity analysis, $a_0$ is directly obtained from Eq. (1) so that

$$a_{\beta, i} = b m_0 + \log_{10} n - \log_{10}(V_{i+1} - V)$$

with $i$ the incremental window step. Noteworthy, $n \geq 100$ is reached only for high $a_0$ and/or low completeness $M_1$. Depending on the underground $a$ priori knowledge, the injection profile plan, and the safety criterion (Eq. 3), one may estimate what a reasonable $M_1$ would be for parameter updating over time. This may require further seismic network planning, with the number and spatial arrangement of seismic stations potentially derived from the simple function $M_1(d, k) = c_d d(k)^2 + c_p$, where $d$ is the distance to the $k^{th}$ nearest seismic station and $c_d, c_p$ empirical parameters (although there are additional theoretical limits on event detection that must be taken into account).

**Post-injection data analysis.** While the linear relationship between $\lambda$ and $V$ is well established, the pure exponential behaviour of the induced seismicity post-injection tail has only been demonstrated for the 2006 Basel case. Here, we compare three relaxation models: pure power law $\lambda(t) \propto t^{-\beta}$, pure exponential $\lambda(t) \propto \exp(-t/\tau)$ and stretched exponential $\lambda(t) \propto t^{-2} \exp(-(t/\tau)^\beta)$ by using the Akaike Information Criterion. We find that, of the 7 datasets (discarding KTB94a, which tail is likely cut), 5 are best described by the pure exponential function. The 2 others are best described by a stretched exponential function with stretching parameter $\beta = 0.9$ and 0.7, for PV94 and NB14b, respectively ($\beta = 1$ representing the pure exponential). This justifies the use of a pure exponential in Eq. (1) hence limiting $\theta$ to a simple three-parameter set.

**Model goodness-of-fit.** Figure 1 shows the general agreement between model and data by visual inspection. Additionally we test Eq. (1) against our datasets using the Kolmogorov-Smirnov (KS) confidence bounds. We first convert the dataset $D(t)$ into a transformed dataset $\tilde{D}(T)$ as follows:

$$D(t) = \{t_1, \ldots, t_n, \ldots, t_{i=1}\} \rightarrow \tilde{D}(T) = \{T_1, \ldots, T_i, \ldots, T_n\}$$

$$T_i = \int_{0}^{T_i} \lambda(t, \theta)dt$$

The two datasets are equivalent but with the distribution of $\tilde{D}(T)$ being the one of a uniform Poisson process with unit rate. Then using the KS confidence bounds, we estimate graphically whether the empirical cumulative distribution function (CDF) of $\tilde{D}(T)$, $F_{\tilde{D}}(T)$, deviates significantly from the CDF of the uniform distribution, $F_{U}(T)$. Here, we use the confidence bounds not to perform a KS statistical test but rather to examine the physical-engineering evidence against the proposed model. When $F_{\tilde{D}}(T)$ is within the 95–99% KS bounds, we classify the model as performing well; when $F_{\tilde{D}}(T)$ falls locally outside the 95–99% KS bounds, we classify the model as performing poorly. Results are shown in Figure 4 with the dashed lines representing the two-sided 95% and 99% confidence intervals. The following conclusions can be drawn: The model performs well for the datasets KTB94b, B06 and G11; fairly well for the datasets KTB94a, PV94, NB12 and NB14b; and poorly for dataset NB14a. Given the range of different datasets, we conclude that the rate model expressed in Eq. (1) describes fairly well the relationship between injection flow rate and fluid-induced seismicity rate.

**Time series simulation.** We simulate induced seismicity time series with time-dependent flow rates $V(t)$ and $\theta(t) = [a(t), b(t), \tau]$ by using the thinning method for the injection phase. For the post-injection phase for which $\tau$ is assumed constant, as no control over post-injection seismicity is possible, the inversion method is used instead to simulate event occurrence times $t^{\text{post}}$. Magnitudes $m$ are also simulated using the inversion method (for both injection and post-injection phases with $b$ time-dependent in the first phase).

**Risk-to-magnitude mapping.** We translate the safety threshold $IR = 10^{-6}$ as $Pr(I = 9) = 10^{-5}$ assuming that a statistically average “poor” building collapse is reached for a ground intensity $I = 9$ (i.e., most buildings of class A in the EMS98 code suffer collapse) and that once a building collapses, there is a 10% chance of individual fatality. The magnitude $m_{\text{safe}}$ is then estimated from an intensity prediction equation (IPE). For induced
seismicity, we use the IPE derived from the U.S. Geological Survey “Did You Feel It?” rich database, corrected for induced seismicity:

$$I(d_{hyp}) = c_1 + c_2(m_{tecto} - 6) + c_3(m_{tecto} - 6)^2 + c_4\log_{10}d_{hyp} + c_5d_{hyp} + c_6m_{tecto}\log_{10}d_{hyp} + 3\sigma$$

with $m_{tecto}$ the earthquake magnitude in the tectonic context, $d_{hyp} = \sqrt{d^2 + h^2}$ the hypocentral distance in km (point source hypothesis at depth $h = 4$ km), $c_1 = 11.72$, $c_2 = 2.36$, $c_3 = 0.1155$, $c_4 = -0.44$, $c_5 = -0.002044$, $c_6 = -0.479 \times 38$, $\sigma = 0.4$, $m_{saf} = m_{tecto} + m_{corr}$ and $m_{corr} = 0.82$. The correction is based on the observation that induced earthquakes would seem less severe in average than tectonic ones. We assume that the Modified Mercalli Intensity (MMI) and the USGS Community Internet Intensity (CII) are equivalent for sake of simplicity. For a specific project, a fully probabilistic risk approach is recommended to derive the parameters $Y$ and $m_{saf}$ of the ATLS closed-form expressions (Eqs 2–3). Here, the safety thresholds shown in Figure 2 use $Y = 10^{-5}$ with $m_{saf}(d=0\text{km}) = 5.8$ and $m_{saf}(d=50\text{km}) = 7.9$ (equivalent in the tectonic case to 5.0 and 7.1, respectively). This simplified approach is used to illustrate our ATLS proof-of-concept for a general case with no site-specific conditions. A detailed risk analysis, which would integrate risk over the full magnitude range, remains outside the scope of the present study. The high values of $m_{saf}$ are due to using IR as the safety metric. Some experts may disagree that such high magnitudes can be reached in the induced seismicity context, although it is statistically plausible. The choice of the safety metric has however no impact on the validity of the proposed method. For example, using the minor

Figure 4. Goodness-of-fit of the model represented by Eq. (1). Using the Kolmogorov–Smirnov (KS) confidence bounds, we find that the model performs well for the datasets KTB94b, B06 and G11; fairly well for the datasets KTB94a, PV94, NB12 and NB14b; and poorly for dataset NB14a (see the Methods section for details).
damage threshold \( \Pr(t = 0) = Y \) (i.e., most buildings of class A in the EMS98 code suffer negligible to slight damage) would yield \( m_{\text{saf}}(d = 0\, \text{km}) = 4.0 \). From Eq. (3), the same ATLS \( m_{\text{saf}} \) would be computed for both \( \Pr(t = 0) \approx 10^{-3} \) assuming e.g. \( b = 1 \) in \( \log_{10}(Y) = \log_{10}(Y) + b(m - m_{\text{saf}}) \).

**Decision variable definition.** The threshold \( m_{\text{saf}} \) is fixed such that \( \Pr(m \geq m_{\text{saf}}) = Y \). Assuming that earthquakes follow a non-homogeneous Poisson process,

\[
\Pr\left( m \geq m_{\text{saf}} \right) = 1 - \exp\left(-\Lambda(T, m \geq m_{\text{saf}})\right) = Y
\]  

(9)

where \( \Lambda(T, m \geq m_{\text{saf}}) = \int_0^T \lambda(t, m \geq m_{\text{saf}}; 0) dt \) is the mean cumulative number of events and \( T \) is the observation time. For sufficiently large \( T \) (i.e. \( T \gg t_{\text{shut-in}} + \tau \)),

\[
\Lambda(T, m \geq m_{\text{saf}}) = 10^{b \left[-m_{\text{saf}}\right] t_{\text{shut-in}}} + \tau \Rightarrow \bar{V}(t_{\text{shut-in}}) \approx Y
\]  

(10)

with \( V(t) \) the cumulative injected volume and \( \Lambda(T, m \geq m_{\text{saf}}) \approx Y \) for \( \Pr(m \geq m_{\text{saf}}) < 1 \). Eq. (2) is then obtained by injecting Eq. (10) into Eq. (9) (note that one could replace \( Y = -\ln(1-Y) \) in Eq. (2) with an impact on the results only if \( Y \) was tending to 1, which is unlikely in safety norms). We then define the ATLS as the operational magnitude threshold \( m_{\text{sh}} \) at which the injection is stopped in order to meet the safety target. Note that \( m_{\text{sh}} \) also provides the completeness magnitude to attain in the region and thus a basis for seismic network monitoring planning.32,33 We thus obtain the following system of equations (for \( Y < 1 \)):

\[
\begin{align}
&\left(10^{b \left[-m_{\text{saf}}\right] t_{\text{shut-in}}} + \tau \Rightarrow \bar{V}(t_{\text{shut-in}}) \approx Y
\right) \\
&\Rightarrow \left(10^{a \left[-m_{\text{sh}}\right] t_{\text{shut-in}}} V(t_{\text{shut-in}}) + \tau \bar{V}(t_{\text{shut-in}}) \right) \\
&\Rightarrow \left(10^{a \left[-m_{\text{sh}}\right] t_{\text{shut-in}}} V(t_{\text{shut-in}}) = 1 \right)
\end{align}
\]  

(11)

The second equation is always true since the expected number of events with \( m \geq m_{\text{sh}} \) is one, given the assumption that injection is stopped for \( t = t_{\text{shut-in}} \) as soon as \( m \geq m_{\text{sh}} \) is first observed. Substituting \( V(t_{\text{shut-in}}) \) in the first equation of this system yields Eq. (3). It is worth noting that omitting the post-injection tail effect with \( \tau = 0 \) yields the basic frequency-magnitude distribution threshold \( m_{\text{sh}} = \log_{10}(Y)/b + m_{\text{saf}} \).

**Data availability.** All the data used in this study are publicly available. For more information, please contact arnaud.mignan@sed.ethz.ch.

**References.**

1. Ellsworth, W. L. Injection-Induced Earthquakes. Science 341, 1225942 (2013).
2. Keranen, K. M., Weingarten, M., Abers, G. A., Bekins, B. A. & Ge, S. Sharp increase in central Oklahoma seismicity since 2008 induced by massive wastewater injection. Science 345, 448–451 (2014).
3. Shrizai, M., Ellsworth, W. L., Tampio, K. F., Gonzales, P. J. & Manga, M. Surface uplift and time-dependent seismic hazard due to fluid injection in eastern Texas. Science 353, 1416–1419 (2016).
4. van Thienen-Visser, K. & Breunese, J. N. Induced seismicity of the Groningen gas field: History and recent developments. The Leading Edge 34, 664–671 (2015).
5. Cesca, S. et al. The 2013 September-October seismic sequence offshore Spain: a case of seismicity triggered by gas injection? Geophys. J. Int. 198, 941–953 (2014).
6. White, J. A. & Foxhall, W. Assessing induced seismicity risk at CO2 storage projects: Recent progress and remaining challenges. Int. J. Greenhouse Gas Control 49, 413–424 (2016).
7. Giardini, D. Geothermal quake risks must be faced. Nature 462, 848–849 (2009).
8. Guglielmi, Y., Cappa, F., Avouac, J.-P., Henry, P. & Elsworth, D. Seismicity triggered by fluid injection-induced aseismic slip. Science 348, 1224–1226 (2015).
9. Mignan, A., Landtwing, D., Kästli, P., Mena, B. & Wiemer, S. Induced seismicity risk analysis of the 2006 Basel, Switzerland, Enhanced Geothermal System project: Influence of uncertainties on risk mitigation. Geothermics 53, 133–146 (2015).
10. Bosman, K., Baig, A., Vegas, G. & Urbanic, T. Towards an improved understanding of induced seismicity associated with hydraulic fracturing. First Break 34, 61–66 (2016).
11. Fischhoff, B. The realities of risk-cost-benefit analysis. Science 350, 527 (2015).
12. Bommer, J. J. et al. Control of hazard due to seismicity induced by a hot fractured rock geothermal project. Eng. Geol. 83, 287–306 (2006).
13. Häring, M. O., Schanz, U., Ladner, F. & Dyer, B. C. Characterisation of the Basel 1 enhanced geothermal system. Geothermics 37, 469–495 (2008).
14. Jonkman, S. N., van Gelder, P. H. A. J. M. & Vrijling, J. K. An overview of quantitative risk measures for loss of life and economic damage. J. Hazardous Materials A99, 1–30 (2003).
15. Barker, D. M., Huang, W., Guo, Y.-R., Bourgeois, A. J. & Xiao, Q. N. A Three-Dimensional Variational Data Assimilation System for MMS: Implementation and Initial Results. Mon. Weather Rev. 132, 897–914 (2004).
16. Rutqvist, J. Status of the TOUGH-FLAC simulator and recent applications related to coupled fluid flow and crustal deformations. Computers & Geosciences 37, 739–750 (2011).
17. Mignan, A. Static behaviour of induced seismicity. Nordin. Processes Geophys. 23, 107–113 (2016).
18. Dinske, C. & Shapiro, S. A. Seismotectonic state of reservoirs inferred from magnitude distributions of fluid-induced seismicity. J. Seismol. 17, 13–25 (2013).
19. van der Elst, N., Page, M. T., Weiser, D. A., Goebel, T. H. W. & Hosseini, S. M. Induced earthquake magnitudes are as large as (statistically) expected. J. Geophys. Res. 121, 4575–4590 (2016).
20. Langenbruch, C. & Shapiro, S. A. Decay rate of fluid-induced seismicity after termination of reservoir stimulations. Geophysics 75, MA53–MA62 (2010).
21. Mignan, A. Revisiting the 1894 Omori Aftershock Dataset with the Stretched Exponential Function. Seismol. Res. Lett. 87, 107–113 (2016).
22. Mignan, A. Reply to “Comment on ‘Revisiting the 1894 Omori Aftershock Dataset with the Stretched Exponential Function’ by A. Mignan” by S. Hainzl and A. Christophersen. Seismol. Res. Lett. 87, 1134–1137 (2016).
23. Kraft, T. & Deichmann, N. High-precision relocation and focal mechanism of the injection-induced seismicity at the Basel EGS. Geothermics 52, 59–73 (2014).
24. Jost, M. L., Büßelberg, T., Jost, O. & Harjes, H.-P. Source Parameters of Injection-Induced Microearthquakes at 9 km Depth at the KTB Deep Drilling Site, Germany. Bull. Seismol. Soc. Am. 88, 815–832 (1998).
25. Petty, S., Nordin, Y., Glassley, W., Cladouhos, T. T. & Swyer, M. Improving geothermal project economics with multi-zone stimulation: Results from the Newberry volcano EGS demonstration. Proc. 38th Workshop Geotherm. Res. Eng. SGPr-TR-198 (2013).
26. Cladouhos, T. T., Petty, S., Swyer, M. W., Uddenberg, M. E. & Nordin, Y. Results from Newberry Volcano EGS Demonstration. Proc. 40th Workshop Geotherm. Res. Eng. SGPr-TR-204 (2015).
27. Ake, J., Mahrer, K., O’Connell, D. & Block, L. Deep-Injection and Closely Monitored Induced Seismicity at Paradox Valley, Colorado. Bull. Seismol. Soc. Am. 95, 664–683 (2005).
28. Holland, A. A. Earthquakes Triggered by Hydraulic Fracturing in South-Central Oklahoma. Bull. Seismol. Soc. Am. 103, 1784–1792 (2013).
29. Rubinstein, J. L. & Mahani, A. B. Myths and Facts on Wastewater Injection, Hydraulic Fracturing, Enhanced Oil Recovery, and Induced Seismicity. Seismol. Res. Lett. 86, 1060–1067 (2015).
30. Gobel, T. A comparison of seismicity rates and fluid-injection operations in Oklahoma and California: Implications for crustal stresses. The Leading Edge 34, 640–645 (2015).
31. McGarr, A. Maximum magnitude earthquakes induced by fluid injection. J. Geophys. Res. 119, 1008–1019 (2014).
32. Evans, K. F., Zappone, A., Kraft, T., Deichmann, N. R. & Moia, F. A survey of the induced seismic responses to fluid injection in geothermal and C2 reservoirs in Europe. Geothermics 41, 30–54 (2012).
33. Camerer, C. & Weber, M. Recent Developments in Modeling Preferences: Uncertainty and Ambiguity. J. Risk and Uncertainty 5, 325–370 (1992).
34. Maxwell, S. C., Zhang, F. T. & Damjanac, B. Geomechanical modeling of induced seismicity resulting from hydraulic fracturing. The Leading Edge 34, 678–683 (2015).
35. Kwiatek, G. et al. Effects of long-term fluid injection on induced seismic parameters and maximum magnitude in northwestern part of The Geysers geothermal field. J. Geophys. Res. 120, 7085–7101 (2015).
36. Dawes, R. M., Faust, D. & Meehl, P. E. Clinical Versus Actuarial Judgment. Science 243, 1668–1674 (1989).
37. Grove, W. M. Clinical Versus Statistical Prediction: The Contribution of Paul E. Meehl. J. Clinical Psychology 61, 1233–1243 (2005).
38. Lewis, P. A. W. & Shedler, G. S. Statistical Analysis of Non-stationary Series of Events in a Data Base System. IBM J. Res. Dev. 20, 465–482 (1976).
39. Langenbruch, C., Dinske, C. & Shaprio, S. A. Inter event times of fluid-induced earthquakes suggest their Poisson nature. Geophys. Res. Lett. 38, L21302 (2011).
40. Mignan, A. Functional shape of the earthquake frequency-magnitude distribution and completeness magnitude. J. Geophys. Res. 117, B08302 (2012).
41. Aki, K. Maximum Likelihood Estimate of b in the Formula log N = a−bM and its Confidence Limits. Bull. Earthquake Res. Inst. 43, 237–239 (1965).
42. Mignan, A., Werner, M. J., Wiemer, S., Chen, C.-C. & Wu, Y.-M. Bayesian Estimation of the Spatially Varying Completeness Magnitude of Earthquake Catalogs. Bull. Seismol. Soc. Am. 101, 1371–1385 (2011).
43. Kwiatek, G. & Ben-Zion, Y. Theoretical limits on detection and analysis of small earthquakes. J. Geophys. Res. 121, 5898–5916 (2016).
44. Akaeke, H. & Nakada, H. A New Look at the Statistical Model Identification. IEEE Trans. Automatic Control AC 19, 716–723 (1974).
45. Mignan, A. Modeling aftershocks as a stretched exponential relaxation. Geophys. Res. Lett. 42, 9726–9732 (2015).
46. Ogata, Y. Statistical Models for Earthquake Occurrences and Residual Analysis for Point Processes. J. Am. Stat. Assoc. 83, 9–27 (1988).
47. Lewis, P. A. W. & Shedler, G. S. Simulation of nonhomogeneous Poisson processes by thinning. Nav. Res. Logistics Q. 26, 403–413 (1979).
48. Devroye, L. Non-Uniform Random Variate Generation (Springer, 1986).
49. Grünthal, G. European Macroseismic Scale 1998 (Imprimerie J. Beffort, 1998).
50. FEMA. HAZUS MR4 Technical Manual (FEMA, 2003).
51. Atkinson, G. M. & Wald, D. J. "Did You Feel It?" Intensity Data: A Surprisingly Good Measure of Earthquake Ground Motion. Seismol. Res. Lett. 78, 362–368 (2007).
52. Hough, S. E. Shaking from Injection-Induced Earthquakes in the Central and Eastern United States. Bull. Seismol. Soc. Am. 104, 1086–1095 (2014).
53. Kraft, T., Mignan, A. & Giardini, D. Optimization of a large-scale microseismic monitoring network in northern Switzerland. Geophys. J. Int. 195, 474–490 (2013).