THE USE OF DIGITAL ELEVATION MODELS FOR DETAILED MAPPING OF SLOPE SOILS

A. B. Achasov, A. O. Achasova, A. O. Soved. ВИКОРИСТАННЯ ЦИФРОВИХ МОДЕЛЕЙ РЕЛЬЄФУ ДЛЯ ДЕТАЛЬНОГО КАРТОГРАФУВАННЯ СХИЛОВИХ ГРУНТІВ. Останнім часом у світі тривають активні пошуки нових підходів до картографування грунтів. Традиційні методики картографування, які спирались на результати польових обстежень, не відповідають вимогам сьогодення. Стаття присвячена питанню цифрового картографування грунтів на основі використання цифрових моделей рельєфу (ЦМР). На дослідженої поверхні площою 5 га було проведено детальне польове дослідження з метою визначення зносу грунтів від антропогенної деградації. За результатами обстежень, проведених на основі ЦМР, визначено кількісні характеристики деградації, які знаходять застосування в наукових та практичних дослідженнях.

Andriy Borisovich Achasov, Doctor of Sciences (Agriculture), Associate Professor, Head of the Department of Geodesy, Cartography and Geoinformatics, V. V. Dokuchaev Kharkiv National Agrarian University, «Dokuchaevske-2», Kharkiv region, Kharkiv district, 62483, e-mail: achasov.ab@gmail.com, https://orcid.org/0000-0003-2446-3707;

Alla Oleksandrivna Achasova, PhD (Biology), Associate Professor, National Scientific Center «A. N. Sokolovsky Institute of Soil Science and Agrochemistry», 4 Chaykovskaya str., Kharkiv, 61024, e-mail: alsisa971@gmail.com, https://orcid.org/0000-0002-6294-2445;

Arkady Oleksandrivich Sived, Senior Lecturer of the Department of Geodesy, Cartography and Geoinformatics, V. V. Dokuchaev Kharkiv National Agrarian University, e-mail: shakhmet1985@gmail.com, https://orcid.org/0000-0003-0604-4015

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© Achasov A. B., Achasova A. O., Sived A. O. https://doi.org/10.26565/2410-7360-2019-50-06

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Formulation of the problem. Rapid development of agrotechnologies requires their respective information support. First of all, it concerns available detailed and up-to-date soil cover information, which is the basis of agricultural production. This is especially true for Ukraine because, on the one hand, the agrarian sector is one of the main budget-forming branches of the country, and on the other hand, the mapping database of soils of Ukraine is very outdated. A solid large-scale survey of soils in Ukraine was carried out in the 1950’s and 1960’s, and the latest occasional soil map corrections date back to the 90’s. Consequently, even the latest generation of ground maps in Ukraine has become outdated for at least 20 years.

In addition, it should be noted that even the second round of large-scale soil survey, using the old methods, does not fundamentally change the situation. A significant (albeit objective, at that time) drawback of the classic method of soil mapping, which was used in the USSR [1] and was at one time inherited by Ukraine [2], was a weak substantiation of the boundaries of the earth’s habitats. The contours of the ground plots were applied to maps based on expert decisions of soil explorercartographers who used the materials of point field studies of the soil profiles, topographic maps and, sometimes, aerial photographs. Quantitative methods for deciphering aerial photographs and relief analysis, as a rule, were not used.

Modern agrarian production requires digital soil and agrochemical maps, representig the corresponding geospatial data bases. At the same time, the main importance of these maps should be the possibility to represent continuum of surface characteristics required by the user with the necessary detail and accuracy. Continuity of soil characteristics on such digital surface models (DSM) must be quantified. There are not so many ways to provide such a justification. Two spatial translators that can reliably transmit certain characteristics of the ground cover have already been named –remote sensing data in all of their modern manifolds, and landforms. Geostatistical capabilities for the construction of geopoles and vegetation as surface indicato are added.

Among all these sources of additional information on the soil cover, landforms are distinguished, which, unlike other factors, do not passive-ly reflect the spatial diversity of soils, but is a factor in the formation of the surface structure by regulating the arrival of solar radiation and precipitation on the soil surface. It has been shown previously [3] that the topography factor can be considered as a determining one for the description of the variegated soil cover for the automorphic zonal soils in Ukraine in the scale of the administrative region.

Consequently, having a digital surface model (DEM) of the territory and established mathematical relationships between the parameters of the relief and soil characteristics, it is possible to create predictive TGKs. Subsequently, these maps are adapted to local conditions by adding auxiliary information (remote sensing data, archival and field data), and field verification.

Analysis of recent research and publications. Influence of the relief factor on the surface formation process has been known for a long time. P.A. Kostychev in his time, studying virgin steppe soils of Ukraine, showed that soils of the northern slopes are more humus-like, compared to the soils of southern exposures [4]. Specificity of soil formation on sloping lands was noted in the classical works of G.N. Vysotsky [5], D. Demek [6], A. Ger- rard [7], P. Furley [8] and many others.

Development of computer and geoinformation technology has given this topic a powerful new impetus. According to Bishop and Minasny [9], almost 80% of digital soil mapping projects are performed using DEM as the most powerful source of data for forecasting local soil specificities.

DEM and its derivative geomorphological parameters can be used as the basic material for surface mapping in the following four ways:

1) To update the existing topographic maps.

The analysis of the literature provides numerous examples of the DEM use and derivative models for medium-sized and small-scale soil mapping [10, 11, 12, 13, 14]. A typical example is the work of Biggs and Slater [15] who carried out a medium-scale soil survey, using DEM and compared the results with an existing ground maps created by the traditional method. Using such topographic parameters as slope, surface curvature, topographic humidity index, relative height and slope position, a scale map of 1: 100000 was constructed, which allowed updating the existing map and increasing its reliability. Hammer et al. [16] have showed that the use of DEM with a spatial resolution of 10 m makes it possible to construct maps of slopes that have a large potential application for soil survey and land use planning.

2) For the allocation of soil-landscape units.

Two potential approaches to obtain ground land-
scapes can be used. First, it is an automatic approach based on clustering [17, 18], when there are no definite criteria for the landscape classification. In these studies, an automated clustering procedure is used to determine meaningful clusters of the area, using a set of DEM derivatives. The second approach is based on compliance with the existing classification of landscapes, using expert assessment to determine the naming of selected landscapes according to the classification used [19]. Dobos et al. [20] identified the soil units according to the SOTER (World Soil Database) classification based on the analysis of the terrain by parameters such as height, relief and tilt. Hengl and Rossiter [21] applied photointerpretation in the reference regions for extrapolation of geomorphological units to the entire study area, using nine parameters of the relief.

3) For direct evaluation of soil parameters [22, 23, 14, 24, 25].

The geomorphological parameters of the territory, as shown by the research [26, 27], can successfully be used to forecast the soil cover characteristics, as there is a dependable relationship between soil parameters and landforms characteristics. Geomorphological characteristics derived from DEM are increasingly used as auxiliary covariates for spatial prediction of soil properties [28, 29].

10 derivatives from DTM (STRM) were used in work [30] to assess organic carbon stocks in the soils of the Tibetan Highlands. It has been established that parameters such as altitude and slope length are closely related to the depth of the soil profile. M. Menezes and colleagues successfully used the DIM, derived from a topographic map of scale 1: 50000, and obtained a model for digital ground mapping of soils in Brazil from it [31]. A detailed review of the possible methods for forecasting soil parameters for DEM is presented by T. Bishop and V. Minasny in [9].

4) To optimize the soil sampling strategy. Analysis of DEM with the allocation of homogeneous sections and key structural elements of the topography optimizes sampling schemes [35, 32] and minimizes the forecast error of individual soil parameters [36].

It should be noted that the use of DEM in soil research becomes much more effective if additional information is included in the analysis. Thus, in [37], it is indicated that using DEM together with information on the geological structure, makes it possible to forecast the types of soils in the studied territory. The expediency of the integral use of DEM with the data of remote sensing and other additional materials to improve the efficiency of soil mapping is confirmed in publications [12, 20, 33, 38].

**Selection of previously unsettled parts of the general problem that the article focuses on.** The literary review has showed that, despite a large number of publications, a lot of issues connected with surface digital mapping remain relevant.

In particular, individual geomorphological parameters are used in assessing the influence of topography on soil formation processes: altitude, exposition, steepness, slope length, topographic humidity index, curvature, topographic position index [39, 40, 41, 10, 42] or their combination as a set of regression equations.

In our opinion, this is a logical but not optimal way of formalizing the effect of relief on soil formation. Such an approach does not allow the transition from empirical observations to a truly effective analysis of landscape-soil dependencies. The systematic approach is an alternative, taking into account the impact of relief not as a set of individual indicators, but as an integral function, which largely determines soil development [3].

It has also been noted that most of the well-known publications refer to medium- and small-scale soil mapping. The use of geomorphological characteristics in large-scale and detailed mapping is given much less attention to. Although for today's agrarian production the most important are maps of scale 1: 5000 and larger.

**Formulating the purpose of the article (statement of the task).** The purpose of the article is to study the use of topographic indicative models in creation of detailed digital surface mapping materials.

**Presentation of the main research material.**

**Object and research methods.** The research was carried out in 2016-2017 on the territory of the testing ground located on the lands of V.V. Dokuchaev Kharkiv National Agrarian University. The experimental site is located on the slope of the south-eastern exposition. The ravine divides the territory of the landfill into two sections - "fields". Field №1 is northeastern with the area of 5 hectares, field number 2 is southwestern with the area of 4.3 hectares (Fig. 1). Major studies were focused on field number 1.

Soil cover of the landfill was studied according to generally accepted methods [10] by laying out and describing 7 soil sections, as well as by additional selection of 64 soil samples from the surface.

Soil samples were selected according to two schemes: a regular network with a distance of approximately 50 m, and an irregular one - for visually detected drainage basins. The sampling points were fixed (tied to the geographical coordinate system) using the GPS device.

In all samples, the content of organic matter (C) was determined for [43]. It is an important invariant characteristic of the soil and can serve as a reliable indicator of its condition. The soil samples were ana-
To analyze the influence of landforms on soil formation at the local level and the use of its characteristics in the digital surface mapping, a detailed digital model of the landfill site topography was used. The DEM spatial resolution is 20 x 20 cm per pixel.

The digital surface model was built on the photography from an unmanned aerial vehicle DJI "Phantom 3Advanced". The shooting took place on June 22, 2016, at a height of 25 m. With the help of the specialized mobile application "Pix4D Capture", an autopilot with overlapping of 80% of photos and average speed was created. The resolution of the pictures was 5 x 5cm. The received photos and construction of DEM were processed in the program "AgisoftPhotoScan" [44]. Subsequent analysis of the DEM was performed, using the ArcGIS software package.

**Research results.** The best description of any phenomenon or a process is a mathematical model based on well-defined physical laws. The first step towards formalization of the soil formation process in automorphic zonal non-aerated soils of Ukraine can be considered parametrization of two main substance-energy agents in soil formation: solar energy and moisture, incoming to the ground. According to [3], should the granulometric composition of the soil layer varies slightly, hydrothermal conditions of soil formation are completely controlled by the topography of the investigated area.

According to the results of the morphological description of soil sections and analysis of soil samples, it has been established that the soil cover of the landfill is represented by chernozem typical with different degrees of erosion. The depth of the humus horizon (H) of the soil gradually decreases down the slope from 51 to 30 cm, which may be due to both the soil erosion and their xeromorphism, caused by specific sloping conditions of soil formation. The soil-forming rock for the territory of the landfill is heavy-sandy carbonate loess-like loam (the average content of particles is <0.01 mm is 47.7%). Thus, the process of soil formation before the beginning of active human activity was mainly controlled by topography.

To quantify this effect, it is suggested to use xeromorphism coefficient (Kc) [40], which characterizes the change of hydrothermal conditions for a specific landforms site in comparison with the horizontal surface:

\[ K_k = K_i / K_z \] (1),

where Ki is insolation coefficient, determining direct solar radiation on the real slope compared with the horizontal surface;

Kz is the relative humidity coefficient, characterizing the flow of moisture to the soil on this slope compared with the horizontal surface.

The calculating method of the reduced coefficients is presented in works [45, 46, 34]. The coefficients can be used to calculate the absolute values of solar radiation and precipitation for specific slopes, as well as the relative parameters that characterize the spatial soil-forming “potential” of the terrain.

It should be noted that KK is by its very nature close to the radiation index of dryness, widely used in geography to assess the climatic conditions of the
terrain formation. The xeromorphity coefficient obtains its physical meaning if it is modified as a ratio with numerator equal to the product of the annual radiation balance on Ki, and denominator is the product of the Kz per year of the amount of precipitation and the latent energy of steam generation. As a result, the analogue of the Grigoriev-Budiko dryness index for local landscapes is obtained.

It is clear that Kk cannot fully reflect the influence of topography on soil formation, but in the first approximation it allows to quantify hydrothermal conditions at each point of the relief. The effectiveness of Kc use in the study of soil cover has been confirmed for the conditions of the Steppe and Forest-steppe of Ukraine. It has been established that Qk in non-eroded soils is closely related to the content of humus in black earths and the depth of their profile [45], and, accordingly, can act as a predictor in ground mapping [43].

Based on the results of the UAV-shooting of the DEM landfill (Fig. 2), derivative models were built by ArcGIS program of the following relief parameters: slope steepness, exposure, insolation coefficient (Ki), hydromorphic coefficient (Kg), xeromorphic coefficient (Kc). Table 1 shows statistical characteristics of these parameters for sites where samples of soil were selected.

The model of soil xeromorphism became the

**Fig. 2. DEM of experimental testing ground**

| Variable | Arithmetic mean | Minimum value | Maximum value | Standard deviation | Variation coefficient |
|----------|----------------|---------------|---------------|--------------------|----------------------|
| Steepness | 3,1 | 0,7 | 5,8 | 1 | 32 |
| Exposition | 125 | 55 | 161 | 25 | 20 |
| Hight | 155 | 144 | 161 | 3 | 2 |
| Ki | 1,028 | 0,958 | 1,073 | 0,021 | 2 |
| Kk | 0,891 | 0,802 | 0,975 | 0,033 | 3 |
| Kc | 1,156 | 1,026 | 1,337 | 0,058 | 5 |
| C1 | 2,1 | 1,3 | 2,9 | 0,3 | 13 |
| C2 | 2,7 | 2,2 | 3,1 | 0,2 | 6 |

Note: C1 - the content of organic carbon in a layer of soil 0-10 cm, obtained analytically; C2 - the content of organic carbon in a layer of soil 0-10 cm, calculated on the model of xeromorphism.
spatial base for the creation of cartograms of organic carbon content in the upper layer of the landfill (Fig. 3). The calculation was based on logic-statistical models established earlier for a large sample (164 sections) of non-aerated soils [45, 46].

Since most of the samples on the territory of the landfill (56 pcs.) was selected on field number 1, all the illustrations will characterize it in future.

Correlation analysis has showed that for the landfill territory there is no significant link between values of C1 content and relief parameters (Table 2). There is no link between C1 and the calculated carbon content (C2): r = -0.04, either.

Comparing C1 and C2 values by 56 points, it

has been proved that the cartogram obtained during the geoinformation modeling, demonstrates the value of organic carbon content on average by 0.51% higher than the values established analytically in the soil survey. Moreover, the difference between the maximum values in the samples of the predicted and actually existing content C was only 0.2%, while the minimum values of C for the two samples differed by 0.9%.

This is absolutely logical because the cartographic materials obtained by the proposed algorithm for calculating the content of organic carbon reflect the situation of natural soil formation, without taking into account modern anthropogenic degradation of the soil. That is, they reflect natural potential of the soil. Loss of the upper layer of soil as a result of erosion processes leads to a decrease in C in the layer of soils in proportion to the degree of their erosion.

Having studied changes in the depth profile of the soil under the transect located along the slope (Figure 4), it has been established that the depth of the humused part of the profile (horizons H + Hr) of the arable soil gradually reduces due to decrease in the absolute height of the surface, indicating the influence of water erosion processes.

At the same time, section №2, located 10 m below the margin of the field in the area under virgin steppe vegetation, on the contrary, is characterized by super-high value of the humus part in the profile. Studies have shown that a solid layer of dark-colored humus soil is revealed at a depth of 200 cm. The reason for this is a sharp decrease in the rate of water flow in the area covered with herbaceous ve-

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Fig. 3. The cartogram of organic carbon potential content in the upper layer of soil (C2) is constructed by method of geoinformation analysis of the landforms.

Symbols: 1) C2 <1.7%; 2) C2 = 1.7-2.4%; 3) C2 = 2.4-4.7%; 4) C2> 2.7

Table 2

|       | C1   | C2   | Kk  | Ki  | Kr  | Exposition | Steepness | Height |
|-------|------|------|-----|-----|-----|-------------|-----------|--------|
| C1    | 1    | -0.04| -0.94| 0.23| 0.08| 0.27        | -0.08     | 0.11   |
| C2    | -0.04| 1    | -0.99| -0.71| 0.91| -0.52       | -0.91     | 0.69   |
getation, which contributes to the deposition of soil particles taken from the above cultivated slope.

Another confirmation of the erosion transformation of the landfill sites is the fact that the variation coefficient for the sample of calculated values C is half the coefficient of variation for the analytical values (Table 1). Indeed, the studied fields have practically the same constant exposure, which determines the weak spatial variability of solar radiation arrival on the surface. Accordingly, the process of organic matter formation in soils should be more homogeneous for the entire study area.

Consequently, the difference between the estimated and actual values of the C content in the upper layer of the landfill and absence of significant correlations of the actual carbon content with the relief parameters are associated with water erosion processes. This leads to a dynamic complex of washed away slope soils whose genesis is caused by the simultaneous influence of water erosion and sloping xeromorphism.

Furthermore, an actual cartogram of content C was constructed according to the laboratory measurements with the help of the ArcGIS program (Fig. 5).

The basis for cartograms was the value of C for 56 points located on field number 1. Interpolation was carried out using the method of weighted distances. Spatial structure of the content C in the upper layer of soils of the landfill also confirms a significant degree of their transformation under the influence of erosion processes.

Comparison of the obtained cartograms (Figures 3 and 5) allows to construct a detailed map of the landfill grounds. The algorithm for constructing such a map is as follows:

1) To establish the correspondence between C contents in a layer 0-10 cm and the degree of soil wash off. For this purpose, a standard characteristic of the C content distribution on the depth in the immutable soil is used. In our studies it was section number 20, which was laid outside the testing ground in the watershed and represented a typical modal chernozem. The diagnostic feature of weakly washed off soils is the erosion occupying half thickness of the humus accumulative horizon N. Accordingly, we assume that for our landfill value C in the upper layer of immutable soils should be more than 2.7%, for weakly washed off soils it should fluctuate in the range of 2.4-2.7%.

Average washed soils are characterized by washing off from half to the entire horizon N. Content C in the upper layer of such soils varies within 1.7-2.4%. Heavily washed soils, in which there is no horizon H and part of the upper transition horizon, are characterized by value of C <1.7%. We used these ranges in the construction of cartograms C (Figures 3 and 5).

2) Overlaying of cartograms 3 and 5 in order to select all soil habitats formed on the crossing of the initial data.

3) Ground identification of the received map-diagram (Fig. 6). It is based on the following logical considerations:
The cartogram of the actual content of organic carbon (Fig. 5) can be interpreted as a map of soil erosion, where area № 1 corresponds to the heavily washed off soils, area № 2 – medium washed off soils and area № 3- weakly washed off chernozem.

The cartogram of the estimated potential content C (Fig. 3) reflects the quasi-solid state of soils, in which the erosion processes are a priori absent. Thus, the four distinct areas characterize no erosion but specific sloping hydrothermal conditions, in which it is impossible to form soil similar to the ground of the plain. Such conditions are typical for, for example, southern slopes of great steepness. Such underdeveloped soils have recently been named xenomorphic [47]. Accordingly, area № 1 corresponds to heavily xenomorphic, area № 2 – average xenomorphic, area № 3 – weakly xenomorphic and area № 4 – modal chernozem.

The resultant forecasting map of soils is given in Fig. 6. It describes the soil cover of the field in more details, taking into account both its present condition and the soil formation potential. The map looks rather unusual. This is due to the detailed received information and the variability of the ground cover of the landfill.

Let us summarize the obtained results. The cartogram of potential content C (Fig. 5), constructed on the basis of the DEM analysis, reflects the initial state of the soil, or in other words, shows the "potential of the soil". It is, in fact, the standard to compare the parameters of modern soils to assess the degree of their degradation. Thus, the average loss of the organic carbon by the landfill soil due to anthropogenic degradation can be estimated at 0.51% by weight, that is, 5.1 kg per 1 ton of soil.

Combination of such a "quasi-virgin" ground map with the results of modern field soil surveys can significantly increase the informative resulting map.

In particular, analysis of Fig. 3 allows us to conclude that the initial conditions of soil formation for the studied field varied slightly. Weak- and medium-xenomorphic chernozems were formed only in the lower part of the slope due to significant slopes. Anthropogenic erosion degradation of soils greatly complicated the structure of the soil, making it stochastic. Thus, the southern part of the field adjacent to the bottom of the slope, being rather steep, is characterized by a high content of organic carbon, which contradicts the logic of soil formation and indicates the redeposition of soil with high hums content that has been taken out of the upper part of the field.

The obtained map, based on the research, (Fig. 6) not only reflects the current condition of the soil, but also allows to approach the process of managing soil resources in a new way. As a result of such mapping we obtain spatial information. Based on it, we can distinguish between degraded but potentially fertile soils from soils whose low fertility is due to natural factors - unfavorable hydrothermal conditions. Accordingly, different approaches should be applied to the use of these soils and taking soil protection measures.
Fig. 6. Map of the landfill soils.

Symbols: 1 - chernozem typical weakly washed off, 2 - chernozem typical weakly xenomorphic, 3 - chernozem typical weakly xenomorphic and slightly washed off, 4 - chernozem typical weakly xenomorphic, medium washed off, 5 - chernozem typical medium xenomorphic, alluvial; 6 - chernozem typical medium washed off, 8 - chernozem typical heavily xenomorphic alluvial

Conclusion.
1. Soils on slopes are formed in essentially different conditions in comparison with soils on the plain territories. Sloping soil formation is the result of a complex spatial combination of natural processes of xeromorphism, separation, transit and resedimentation of soil particles due to soil erosion. Therefore, it is necessary to use a fundamentally different approach in comparison with the soils of plain areas for correct mapping of slope soils. The combined effect of these processes should be taken into account.
2. Such an approach is used in geoinformation analysis of the terrain, making it possible to quantitatively formalize hydrothermal conditions of soil formation on the slopes and to spatially assess its xeromorphism.
3. Mapping materials constructed on this methodical approach characterize the potential of the landscape in terms of soil formation and reflect the quasicyclical condition of the soil or "soil potential".
4. Detailed surface maps created by combination of the field and laboratory soil surveys results with the results of soil potential assessment for the analysis of DEM in the article make it possible to adequately estimate and objectively reflect distribution of eroded and xeromorphic soils and their complexes on sloping lands.

Acknowledgements. The authors are grateful for the assistance in carrying out UAV shooting and processing of the results to the company InteticsGeo and personally O. Seliverstov.

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Authors Contribution: All authors have contributed equally to this work.
Andriy Achasov,
Doctor of Sciences (Agriculture), Associate Professor,
Head of the Department of Geodesy, Cartography and Geoinformatics,
V. V. Dokuchaiev Kharkiv National Agrarian University,
«Dokuchaevskye-2», Kharkiv region, Kharkiv district, 62483,
e-mail: achasov.ab@gmail.com, https://orcid.org/0000-0003-2446-3707;

Alla Achasova,
PhD (Biology), Associate Professor,
National Scientific Center «A. N. Sokolovsky Institute of Soil Science and Agrochemistry»,
4 Chaykovskaya str., Kharkiv, 61024,
e-mail: alsisa971@gmail.com, https://orcid.org/0000-0002-6294-2445;

Arkadiy Siedov,
Senior Lecturer of the Department of Geodesy, Cartography and Geoinformatics,
V. V. Dokuchaiev Kharkiv National Agrarian University,
e-mail: shakhmet1985@gmail.com, https://orcid.org/0000-0003-0604-4015

THE USE OF DIGITAL ELEVATION MODELS FOR DETAILED MAPPING OF SLOPE SOILS

Formulation of the problem. The current stage of agricultural development in Ukraine requires highly informative soil maps.
One of the directions for creating such maps is the use of digital elevation models (DEM) as a spatial translator of soil characteristics. The literary review has showed that despite the large volume of publications on this topic, a number of issues of soils digital mapping remain relevant.
The purpose of the article is to study the possibilities of using relief-ground indicative models in creation of detailed digital ground maps.
Methods. The research was conducted on the territory of the testing ground. During the field study of the landfill, it has been established that the soils are represented by chernozem typical of varying erosion degrees.
It is proposed to use a xeromorphism coefficient for the quantitative account of the landforms influence on soil formation, characterizing changes in hydrothermal conditions for a particular section of the topography, compared with the horizontal surface.
A detailed DTM was obtained, using the "Phantom 3" UAV. The derivative models of a number of topographic parameters were built on its basis later. A digital model of xeromorphism of the territory was built, transformed into a model of organic carbon (C) content.
The specified map shows initial conditions of the soil cover. It can be used as a standard to compare parameters of modern soils for assessment of extent of their degradation.
Comparison of this map with the map of actual C content has shown that average loss of C by soils of the studied area owing to anthropogenic degradation can be estimated at 5.1 kg on 1 t of the soil.
Results. Investigations have proved that the geoinformation analysis of landforms allows to quantitatively shape hydrothermal conditions of soil formation for a certain territory. The cartographic materials constructed on such a methodical approach characterize landscape potential on soil formations and reflect quasi-virgin land condition of the soil.
Scientific novelty and practical significance. The detailed soil maps, based on the results of field and laboratory soil studies, leading to potential soil assessment for DEM analysis in the article, allow to adequately estimate and objectively represent distribution of eroded and xeromorphic soils and their complexes on sloping lands.
Keywords: soil, map, geographic information systems, digital surface model, organic carbon, chernozem soil, erosion.

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