Rare event algorithm study of extreme warm summers and heatwaves over Europe

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Key Points:

- The rare event algorithm increases by several orders of magnitude the number of warm summers and heatwaves sampled by the model.
- Warm summers over either France or Scandinavia are linked to wavenumber 3 hemispheric teleconnection patterns.
- Warm summers in Scandinavia show bimodality due to different distribution of subsequent subseasonal heatwaves.

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Abstract
The analysis of extremes in climate models is hindered by the lack of statistics due to
the computational costs required to run simulations long enough to sample rare events. We
demonstrate how rare event algorithms can improve the statistics of extreme events in state-
of-the-art climate models. We study extreme warm summers and heatwaves over France
and Scandinavia with CESM1.2.2 in present-day climate. The algorithm concentrates the
simulations on events of importance, and shifts the probability distributions of regional tem-
peratures such that warm summers become common. We estimate return times of extremes
orders of magnitude larger than what feasible with direct sampling, and we compute sta-
tistically significant composite maps of dynamical quantities conditional on the occurrence
of the extremes. We show that extreme warm summers are associated to wavenumber 3
hemispheric teleconnection patterns, and that the most extreme summers are related to the
succession of rare subseasonal heatwaves.

Plain Language Summary
The impact of extreme climatic events is often dominated by the rarest events. These
events have return times (a measure of how often they occur on average) of hundreds of
years or more, but they could, and do, happen anytime. These events are poorly understood
because of lack of statistics. Climate models are computationally expensive, and can not be
run for long enough to study events with return times longer than a few decades. We use a
new computational technique that allows simulations to focus only on trajectories leading
to extreme heatwaves over a target region, optimizing the use of computational resources.
We thus gather robust statistics for seasonal heatwaves with return times of hundreds or
thousands of years, and observe even rarer events impossible to observe otherwise. We
find that extreme warm summer over France or Scandinavia are synchronised with extreme
warm summers in specific regions of Asia and North-America by a teleconnection pattern
extending over the entire Northern hemisphere. We also find suggestions that extreme warm
summers over Scandinavia may occur in two distinct ways. This new method can be used
to better study the impact of global warming on the risk of catastrophic events, and to
improve their predictability.

1 Introduction
Recent decades have seen a number of exceptionally warm summers and record breaking
heatwaves at the Northern hemisphere midlatitudes (Luterbacher et al., 2004; García-
Herrera et al., 2010; Barriopedro et al., 2011; Otto et al., 2012; Agha Kouchak, 2012; IPCC,
2012; Russo et al., 2015; Coumou & Rahmstorf, 2012). The Intergovernmental Panel on
Climate Change (IPCC) has concluded that hot days and heavy precipitation events have
become more frequent since 1950 (IPCC, 2013). The precise extent to which heat extremes
will become more common in the future and under the target scenarios of +1.5 °C and
+2 °C global surface temperature with respect to preindustrial levels is an active field of
research (Mueller et al., 2016; Dosio et al., 2018; Suarez-Gutierrez et al., 2018).

A phenomenological theory exists on the causes of heatwaves (Schubert et al., 2014;
Perkins, 2015; Horton et al., 2016), including large scale atmospheric circulation patterns,
(Della-Marta et al., 2007; Cassou et al., 2005; Jezequel et al., 2018), lack of precipitation
and soil moisture (Vautard et al., 2007; Zampieri et al., 2014; D’Andrea et al., 2016), and
sea surface temperature anomalies (Della-Marta et al., 2007; Cassou et al., 2005), with
different levels of importance in different regions (Stéphanon et al., 2012). Long lasting
heatwaves and warm summers are among the most relevant events in terms of impacts
(Camargo & Seth, 2016; Trenberth, 2012). Although they are known to be related to
persistent weather regimes (Lau & Kim, 2012; Teng et al., 2013; Hoskins & Woollings, 2015;
Petoukhov et al., 2016; Boers et al., 2019; Kornhuber et al., 2019), their dynamics is not
fully understood. Some studies show stronger increase with global warming of the frequency
of persistent temperature extremes (Pfleiderer et al., 2019). However, the response of the
global circulation to climate change is extremely complex and not well understood (Hoskins &
Woollings, 2015), in particular for the type of rare dynamics that lead to extreme events.

The impact of climate extremes is often dominated by the rarest events. For instance, the
death toll of the Western European heatwave in 2003, about 70,000 casualties, exceeded
the sum of the fatalities due to all other heatwaves during the last three decades (Barriopedro et al., 2011). Moreover, extreme events with very long return times (also called return
periods) always occur. For instance, an extreme event with a return time of 1000 years, has
a chance 1/1000 to occur next year. Among the huge number of possible extreme events
with a return time of 1000 years, a few of them will certainly occur somewhere sometimes
next year. It is thus crucial to study the most extreme and rare events. However, the study of
such events faces strong scientific limitations: we can not rely on historical data for events
with return times of 100 years or more, because often no similar events have ever been
observed, and it is extremely difficult to study these events using climate models because of
the required computational costs.

In this paper we address the problem of computational cost limitations. In order to
face this problem, we use a rare event algorithm that concentrates ensemble simulations on
trajectories of importance for the extreme events, optimizing the use of the computational
resources. Rare event algorithms have recently been applied to turbulence problems (Grafke et al., 2015; Laurie & Bouchet, 2015; Ebener et al., 2019; Bouchet et al., 2019; Lestang et al., 2020) and climate applications (Ragone et al., 2018; Webber et al., 2019; Ragone & Bouchet, 2020; Plotkin et al., 2019). Based on the phenomenology of the dynamics for each
family of extreme events, an appropriated type of rare event algorithm should be carefully
chosen. Here we use a genealogical algorithm (Ragone et al., 2018; Ragone & Bouchet,
2020) adapted from (Del Moral et al., 2005; Giardina et al., 2011), that is efficient to study
long lasting events. This rare event algorithm has been used to study heatwaves with an
intermediate complexity model (Ragone et al., 2018) in perpetual summer conditions. Here
we use it to sample rare heatwaves in CESM 1.2.2 (Hurrell et al., 2013), a fully realistic
climate model, in presence of daily and seasonal cycles.

The goal of this paper is to demonstrate the applicability of rare event algorithms to
state-of-the-art climate models, and to highlight properties of extreme long lasting heat-
waves that can not be assessed with traditional sampling strategies. We study extreme
warm summers and heatwaves over France and Scandinavia in CESM1.2.2, in present-day
climate. We compute return times orders of magnitude larger than what feasible with direct
sampling, and statistically significant composite maps of dynamical quantities. Our results
show that extreme warm summers are associated to wavenumber 3 teleconnection patterns
in the Northern hemisphere, and suggest that the most extreme summers are related to the
succession of multiple rare subseasonal heatwaves.

2 Data and methods

Model. All simulations are performed with the Community Earth System Model
(CESM) version 1.2.2 (Hurrell et al., 2013). We use an atmosphere and land only setup,
whose active components are the Community Atmospheric Model version 4 (CAM4) and
the Community Land Model version 2 (CLM2). The model is run at statistically sta-
tionary state, with sea surface temperature (SST), sea ice cover, and the concentration of
atmospheric CO2 and other greenhouse gases fixed at values representative of present day
climate (year 2000). Contrary to (Ragone et al., 2018), the model features daily and sea-
sonal cycle, and reproduces a fully realistic climate. See the SI for more details. We study
the statistics of a 1000 years long control run, and the statistics of several sets of simulations
using the rare event algorithm.
Heatwaves and physical observable definition. We consider the surface temperature $T_s(\vec{r}, t)$, where $\vec{r}$ is the space variable and $t$ is time, and define the mean surface temperature $\mathbb{E}(T_s)(\vec{r}, t)$, which varies in time (because of the seasonal cycle) and in space. In practice, $\mathbb{E}(T_s)$ will be approximated by the climatological average of $T_s$ computed from a 1000 years long control run. We study the statistics of the spatially and temporally averaged surface temperature anomaly defined by

$$a(t) = \frac{1}{T} \int_{t_i}^{t_i+T} A(t) \, dt \quad \text{where} \quad A(t) = \frac{1}{|D|} \int_D (T_s - \mathbb{E}(T_s))(\vec{r}, t) \, d\vec{r},$$

(1)

where $A(t)$ is the instantaneous spatial average, $D$ is a spatial domain area, $t_i$ is the starting date of the time average, $T$ is the averaging time, for instance several days up to one season. Heatwaves will be defined as extreme values of the observable $a(t)$, as heatwave indices developed for dynamical studies often use anomalies rather than absolute values (Perkins, 2015). We are specifically interested in long events, for instance warm summers with $T=3$ months. For summer anomalies, we consider $a_{JJA}$ defined by (1) with $t_i$ being June 1st and $t_i+T$ August 29. In the following, we consider $D$ as the area over either France or Scandinavia (see the SI for the definition of the areas).

Rare event algorithm experiments and importance sampling. Rare event algorithms allows a numerical model to produce very efficiently rare trajectories of importance for the study of extremes. We simulate an ensemble of $N$ model trajectories, and at constant intervals of a resampling time $\tau$ we perform trajectory selection, pruning and cloning trajectories in order to favour those leading to the extreme events. See the SI for a more detailed description and (Ragone et al., 2018). The algorithm also allows to compute probabilities for the obtained trajectories. Let $X(t)$ be the vector of values of all the model variables at time $t$ (which includes but is not limited to the temperature $T_s$ defined above). We consider the trajectory $\{\vec{X}(t)\}$, with all values of $X$ along all the interval $t_a \leq t \leq t_a + T_a$, where $t_a$ is the starting date and $T_a$ the trajectory duration in the algorithm. We denote $P_k(\{\vec{X}(t)\})$ the probability distribution function of the trajectories obtained in the algorithm. Then the rare event algorithm produce trajectories distributed according to

$$P_k(\{\vec{X}(t)\}) = \frac{e^{\int_{t_a}^{t_a+T_a} A(u) \, du}}{Z} P_0(\{\vec{X}(t)\}),$$

(2)

where $P_0(\{\vec{X}(t)\})$ is the probability distribution function of the trajectories in the model climate, $Z$ is a normalization term computed by the algorithm, $A$ is the selection function (with the dependance on the trajectory implicit for simplicity, $A(t) = A(\{\vec{X}(t)\})$), and $k$ controls the strength of the algorithm selection. Equation 2 is called an importance ratio formula. Using equation 2 we can compute the actual probabilities of the rare trajectories obtained in the simulations with the algorithm. See the SI and (Ragone et al., 2018) for more details.

We use as selection function $A$ the same function used to define a heatwave in formula (1). We see from equation (2) that if the parameter $k$ is positive, then trajectories with large values of the time average of the surface temperature anomaly will be much more probable in simulations obtained with the algorithm rather than in simple simulations with the model. The larger $k$, the stronger the selection, and the more probable trajectories with extreme values of the time averaged surface temperature. We analyse extremely warm summers over France and Scandinavia. For each case, we perform $K=10$ ensemble simulations with the algorithm, each with $N=100$ trajectories, biasing parameter $k=30$, with $T_a=90$ days from June 1st to August 29th, and resampling time $\tau=5$ days (see the SI). A 1000 years long control run is used to provide initial conditions for the experiments with the algorithm, and as a benchmark for the statistics. The computational cost of the experiments with the algorithm is equivalent to simulating 1000 summers in the control run, but they allow to gather a much richer statistics for the extreme events of interest.
3 Results

3.1 Importance sampling of extreme warm summers

The main goal of the algorithm is to perform importance sampling for the distribution of the summer temperature anomalies over the target region. Figures 1a and 1b show the probability distribution functions of the summer temperature anomalies $a_{JJA}$ for the control run and the rare event algorithm. The control distribution shows a similar variance of about 1 °K for both France and Scandinavia. The algorithm is very effective in performing importance sampling and populating the upper tails of the distributions. The typical value of the seasonal anomaly $a_{JJA}$ in the rare event algorithm experiments is around 4 °K for both cases, which are values never observed in the control run. Thanks to the algorithm most of the computational power is indeed used to simulate extremely warm summers, rather than trajectories belonging to the bulk of the distribution. In the case of Scandinavia the algorithm statistics in figure 1b shows a bimodality that we will discuss in section 3.3.

In figures 1c,d we compare return times of $a_{JJA}$ in the control run and in the experiments with the rare event algorithm. A description of the computation of return times both for direct sampling and importance sampling using formula (2) is presented in (Lestang et al., 2018). The black curves are obtained from the control run using 1000 years of data. In order to estimate uncertainty ranges, we compute also an estimate dividing the control run in $K = 10$ samples of 100 years each, computing 10 estimates of the return times curve. We then take their average (blue curves), and compute the empirical standard deviation (blue shaded areas). The red line and shaded areas are obtained in the same way, but using the estimates from $K = 10$ independent experiments with the algorithm.

With the rare event algorithm we reach return times up to $10^5$-10$^6$ years with uncertainty ranges comparable with the ones with the control run for return times of order $10^2$ years. Given the large value of $k$ we chose, we have a small range of overlap for estimates of the return times from the control run and the rare event algorithm. When they do overlap, the values are consistent with each other within the uncertainty ranges. After about $10^6$-10$^7$ years the return time curves reach a plateau. Such plateaux are due to undersampling, as discussed in the SI and in (Lestang et al., 2018). A full demonstration of the reliability of the results obtained with the rare event algorithm in similar simulations can be found in (Ragone et al., 2018; Ragone & Bouchet, 2020).

3.2 Teleconnection patterns for extreme warm summers

We study the dynamical properties of extremely warm summers with return time larger than 100 years (called 100-year warm summers or seasonal heatwaves from now on). We compute composite maps of anomalies of the local JJA surface temperature and 500 hPa geopotential height conditional on the occurrence of 100-year warm summers, for the control run and the rare event algorithm experiments. One of the key advantage of the rare event algorithm is that it gives much better results than the control run for composite statistics for large return times. Moreover the rare event algorithm gives access to composite statistics for return times larger than 1000 years, which is impossible with the control run.

Figure 2a shows composite statistics of surface temperature and 500 hPa geopotential height of 100-year warm summers over France in the control run. The map hints at the presence of a wavenumber 3 teleconnection pattern for both observables. However, how much of these patterns, which were obtained by averaging over only 10 warm summers, is statistically significant? Figure 2b shows the $t$-value map for the geopotential height (see the SI for details). Only the geopotential height anomalies over Europe pass a statistical significance test with $|t| > 2$, which means then we can not really assess the reality of the teleconnection pattern using a 1000 years control run: we just do not have enough data. Figures 2c,d show the same composite maps and statistical significance analysis for 100-year warm summers over France, but computed using the rare event algorithm. The algorithm
results are globally significant, except over the northern part of the Pacific area. Transition areas between cyclonic and anticyclonic anomalies have a $|t|$ value smaller than 2, because they have a low value of the conditional average. They are however rather narrow, so that their location can be still considered rather precise. The rare event algorithm thus allows to properly assess the existence of teleconnections of warm summers in Europe, North America and Central Asia.

The algorithm also gives a much better estimate of the amplitude of the anomalies, since it gives access to large amplitude events unavailable in the control run, because they are too rare. In the same way, the rare event algorithm gives also access to composite statistics that totally unavailable with the control run. For instance figure 3a shows composite statistics of surface temperature and 500 hPa geopotential height for 1000-year warm summers over Scandinavia, which also show a teleconnection pattern with a different spatial pattern but similar broad stroke features.

Based on the rare event algorithm results, we obtain a better overview of the characteristics of the synoptic dynamics occurring during extremely warm summers in the considered regions. A warming pattern centered over the target region is present, encompassing a larger area on a spatial scale of a few thousands km, which coincide with central-Western Europe for France and Northern-Eastern Europe for Scandinavia. Persistent anticyclonic synoptic scale structures centered over the target area are associated to this warming pattern. Locally these structures are consistent with the observed synoptic conditions for the occurrence of European heatwaves. In particular, for the North Atlantic and European area, the local patterns are very close to the observed patterns for the Western-European and the Scandinavian heatwave clusters obtained from reanalysis data in (Stéfanon et al., 2012).

The application of the rare event algorithm allows to assess in a statistically robust way that these local dynamics are part of hemispheric structures of approximately wavenumber 3, which induce a temperature teleconnection pattern, with persistent regional temperature anomalies of alternating signs around the hemisphere. The patterns for the two regions are in broad strokes similar, although they differ in the exact location of positive and negative anomalies. Warm summers over France occur systematically with warm summers over Siberia and North-East America. The corresponding tripolar structure of anticyclonic anomalies is accompanied by a localized low over Central-Asia and a general lower pressure over the Arctic, with minimum over Greenland. The structure related to Scandinavian warm summers is similar on the North-Atlantic sector, but over Asia it is quite different, with a negative temperature anomaly (and related low pressure) extending from Southern Europe to the whole central Asia, with the positive anomaly constrained over the far East. The circulation over the Arctic is also different, with a strong low over the Pole.

### 3.3 Bimodality of warm Scandinavian summers and subseasonal heatwaves

Figure 1b a bimodality of the Scandinavia summer temperatures in the rare event algorithm data. This suggests that two types of distinct dynamical events might lead to Scandinavian extreme summers. We note that, because of the nonlinear relation (2), a bimodality in the algorithm distribution does not necessarily imply a bimodality in the tail of the model distribution. Still it is interesting to test the hypothesis of two types of dynamics.

Figure 3a shows composite statistics of surface temperature and 500 hPa geopotential height for 1000-year warm summers over Scandinavia. Figures 3b and 3c show the composite maps from the rare event algorithm computed separately for $a_{JJA} < 4.2^\circ K$ and $a_{JJA} > 4.2^\circ K$, where $4.2^\circ K$ corresponds to the local minimum between the two peaks of the distribution in red in figure 1b. While the overall structure is the same as the total composite map in figure 3a, the map of the first range (figure 3b) shows a weaker teleconnection pattern, and an anticyclonic anomaly over the North-Atlantic that is not present in the map of the second range (figure 3c). The Scandinavian warm summer dynamics is compatible with a
northward shift of the jet stream along the entire hemisphere. In the case of the first range (figure 3b), this shift is less clear over the North-Atlantic, where a different dynamics seems to be in place. It is however likely that, if two distinct types of dynamics are occurring, a selection based only on the range of seasonal regional temperature fluctuations is not enough to differentiate between different dynamics, and therefore simple composite maps show from that point of view mixed information.

The study of relation between subseasonal fluctuations of surface temperature and extreme warm summers can give an explanation of this bimodality. To visualize this relation, we look at the genealogical structure of the trajectories in an experiment with the rare event algorithm for Scandinavia in figure 4a. The black lines correspond to trajectories belonging to the first range ($a_{JJA} < 4.2^\circ$K), while the red lines to trajectories belonging to the second range ($a_{JJA} > 4.2^\circ$K). In the experiment represented in figure 4a, trajectories belonging to each of the two ranges of the bimodal distribution coexist. In figure 4b, we show the same genealogical tree, but indicate with different colors the values of the 5-days temperature anomaly in each segment. These values are computed using equation 1 with $T = 5$ days and $t_f$ corresponding to the date of each resampling event. In particular red segments indicate a 5-day anomaly above $4.5^\circ$K. The chosen threshold of $4.5^\circ$K is the median of the distribution of the 5-day temperature anomalies rare event algorithm experiments for Scandinavia (shown in the SI). This value corresponds to the 96.6th percentile of the distribution of the 5-day temperature anomalies in the control run. We can thus consider 5-days periods with temperature anomaly larger than this threshold as heatwave periods, and a succession of 2 or more consecutive heatwave periods as a subseasonal heatwave.

With this definitions, during warmer summers belonging to the second range (red in figure 4a), we see a first subseasonal heatwave that lasts about 20 days in late June-early July, and then a second subseasonal heatwave that lasts again about 15 days in August. Less extreme summers (black in figure 4a) have instead only one subseasonal heatwave in June-July. The bimodality of the algorithm statistics therefore corresponds to two qualitatively different types of warm summer, with either one or two subseasonal heatwaves. We note that the 2003 warm summer in France was characterized by two separate subseasonal heatwaves in June and August (García-Herrera et al., 2010).

4 Conclusions

We have shown how simulations with the algorithm produce hundreds of times more extremes than a control run for the same computational cost, and allow us to estimate return times of extreme events orders of magnitude larger. This allows to compute precisely statistically significant composite maps of dynamical quantities for warm summers with extremely strong seasonal surface temperature anomalies. In this way we are able to identify rigorously the occurrence of persistent teleconnection patterns of wavenumber 3 during very warm summers. These patterns and the related dynamics can not be properly studied with direct sampling, as the corresponding local anomalies of temperature and geopotential height are statistically significant only over the target region. The improved statistics given by the rare event algorithm instead allows to obtain statistically significant maps and a very clear signal also away from the target region.

Teleconnection patterns similar to what we have obtained here were observed in the first application of the rare event algorithm to a climate model (Ragone et al., 2018). In that case the model was an intermediate complexity model, run in perpetual summer setup, and the target area was the whole Europe. In this case we use a much more realistic model with seasonal cycle, calibrated to reproduce present day climate, and we target areas consistent with observed heatwave clusters (Stéfanon et al., 2012) and recent cases of very intense heatwaves or warm summer, like the heatwave over France of 2003 and the warm summer over Northern Europe of 2018. It is striking that we obtain similar qualitative results. This supports the idea that these teleconnections at seasonal scale and the corresponding
wavenumber 3 dynamics are robust features, mainly dynamical and little affected by the details of the physical parameterizations.

Comparison with observations is not straightforward, due to the lack of data and the fact that most studies have focused on events at subseasonal scale. Several authors have recently highlighted the role of Rossby waves in determining teleconnection of extreme events (Schubert et al., 2011; Lau & Kim, 2012; Petoukhov et al., 2013, 2016), with particular emphasis on strong heatwaves, e.g. cases over the central USA (Teng et al., 2013), Alberta (Petoukhov et al., 2018) and Western Europe (Kornhuber et al., 2019). These studies typically find patterns with wavenumber 5 to 7 and analyse subseasonal temperature fluctuations. Several different detection methods are used in the literature to identify the dynamics leading to heatwaves, including projections of heatwave states on typical variability patterns obtained with cluster analysis (Cassou et al., 2005) or empirical orthogonal functions (Teng et al., 2013), spectral analysis (Petoukhov et al., 2018; Kornhuber et al., 2019) and/or indicators based on resonance models (Petoukhov et al., 2013, 2016). Merging these approaches with rare event simulations should lead to very promising future studies, aimed at investigating the relation between the seasonal scale structures we find here and the possible role of higher wavenumber stationary Rossby waves proposed in the literature, as well as the possible multimodality in the way of occurrence of some extreme events.

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Figure 1. Distribution of seasonal JJA temperatures anomalies averaged over France (a) and Scandinavia (b) from the control run (black) and the rare event algorithm (red). Return times for the seasonal temperature anomalies for France (c) and Scandinavia (d), from the control run (black) and the rare event algorithm (red). The shaded areas, light blue for the control run and light red for the algorithm, correspond to 1 standard deviation of the sample used to compute the estimate (see SI).
Figure 2. Northern hemisphere composite maps (conditional statistics) for the JJA anomalies of the surface temperature (colors) and 500 hPa geopotential height (contours) for 100-year warm summers over France, in the control run (a) and in the rare event algorithm statistics (c). Panels b) and d) show for the control run and the rare event algorithm respectively the corresponding maps of the $t$ values of the 500 hPa geopotential height anomalies (see the SI).
Figure 3. Composite maps for the anomalies of the surface temperature (colors) and 500 hPa geopotential height (contours) for warm summers over Scandinavia with return times larger than 1000 years (a). Panel b) and c) show respectively composite maps for the same variables with the condition of JJA temperature anomaly smaller or higher that 4.2°C.
Figure 4. Panel a) shows the genealogical tree of a rare event algorithm experiment, with $N = 100$ ensemble members, for selection of Scandinavia seasonal heatwaves. Each broken line, composed by one trunk and its branches up to the last leave, represents a trajectory. The horizontal axis represents time in blocks of $\tau = 5$ days. The genealogical trees have been created with the help of the ETE Python toolkit (Huerta-Cepas et al., 2016). Black lines represent trajectories with $a_{JJA} < 4.2^\circ$K, red lines trajectories with $a_{JJA} > 4.2^\circ$K. Panel b): the same, but with different colors for each 5 day period: blue corresponds to negative 5-day time averaged anomalies, black between 0 and 4.5$^\circ$K, and red above 4.5$^\circ$K.
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