A Study on Monte Carlo Simulation of the Radiation Environment above GeV at the DAMPE Orbit

S. X. Wang¹,², J. J. Zang³,¹, W. Jiang¹, S. J. Lei¹,², C. N. Luo¹,², Z. L. Xu¹,², and J. Chang¹,²

¹Key Laboratory of Dark Matter and Space Astronomy, Purple Mountain Observatory, Chinese Academy of Sciences, Nanjing 210023, China
zangjingjing@lyu.edu.cn, xuzl@pmo.ac.cn
²School of Astronomy and Space Science, University of Science and Technology of China, Hefei 230026, China
³School of Physics and Electronic Engineering, Linyi University, Linyi 276000, China

Received 2022 January 10; accepted 2022 February 10; published 2022 March 17

Abstract

The Dark Matter Particle Explorer (DAMPE) has been undergoing a stable on-orbit operation for more than 6 yr and acquired observations of over 11 billion events. A better understanding of the overall radiation environment of the DAMPE orbit is crucial for both simulation data production and flight data analysis. In this work, we study the radiation environment at low Earth orbit and develop a simulation software package using the framework of ATMNC3, in which state-of-the-art full 3D models of the Earth’s atmospheric and magnetic-field configurations are integrated. We consider in our Monte Carlo procedure event-by-event propagation of cosmic rays in the geomagnetic field and their interaction with the Earth’s atmosphere, focusing on the particles above GeV that are able to trigger the DAMPE data acquisition system. We compare the simulation results with the cosmic-ray electron and positron (CRE) flux measurements made by DAMPE. The overall agreement on both the spectral and angular distribution of the CRE flux demonstrates that our simulation is well established. Our software package could be of more general usage for simulation of the radiation environment of low Earth orbit at various altitudes.

Key words: methods: data analysis – instrumentation: detectors – astroparticle physics

1. Introduction

The Dark Matter Particle Explorer (DAMPE), launched on 2015 December 17 is a space-borne high-energy particle detector covering an energy range from GeV to TeV (Chang et al. 2017; Xu et al. 2018; Huang et al. 2020). The observation results of DAMPE can be used for scientific research on the origin of cosmic rays (Yuan & Feng 2018; An et al. 2019; Yuan et al. 2021), indirect detection of dark matter (Ambrosi et al. 2017; Yuan et al. 2017; Alemanno et al. 2021; Xu et al. 2022) and gamma-ray astronomy (Duan et al. 2019). DAMPE has been operating on a solar synchronized orbit of 500 km for more than 6 yr and has acquired observational data of over 11 billion cosmic-ray events. At this low Earth orbit (LEO⁴), massive data are collected in a complex cosmic-ray radiation environment mainly formed by the interaction of galactic cosmic rays (GCRs) with the geomagnetic field and Earth’s atmosphere. The charged GCRs propagating in interplanetary space are modulated by solar activities, showing a flux fluctuation with a period of about 11 yr. The transportation of charged GCRs in Earth’s magnetosphere region is mainly affected by the geomagnetic field, forming several well-known phenomena, such as the radiation belt, the South Atlantic Anomaly area and the east–west effect. At even closer distance to the Earth, GCRs may interact with atoms in the atmosphere, generating a large amount of secondary particles. As DAMPE keeps operating in such a complex radiation environment, accurate on-orbit simulation of the radiation environment is the basis for fine calibrations of the detector that are crucial to understanding the flight data, such as the minimum ionized particles calibration, the geomagnetic rigidity cutoff measurement, the event rate and data volume estimation of the data acquisition (DAQ) system, and so on.

Given the critical role played by a correct understanding of the radiation environment at the DAMPE orbit in both the simulation data production and flight data analysis, we study in this work the Monte Carlo (MC) simulation of the radiation environment at the LEO, and the DAMPE orbit in particular. As the inputs of our simulation, we make use of the fluxes of primary cosmic-ray protons, helium nuclei, and electrons and positrons (CREs) above GeV measured by the AMS02 (Aguilar et al. 2014, 2015a, 2015b). We then back trace the propagation of these charged particles in the geomagnetic field, obtaining the distribution of primary cosmic-ray particles on the satellite orbital plane. The interaction between these primary cosmic-ray particles and the Earth’s upper atmosphere

⁴ https://www.sciencedirect.com/topics/earth-and-planetary-sciences/low-earth-orbit
is simulated to generate secondary particles. All these primary and secondary particles are then connected to the DAMPE official Geant4 package to simulate their interaction with the detectors, and finally produce reconstructed simulation data in the same format as the flight data. Our simulation of the radiation environment focuses on the high-energy GCRs above GeV, as the solar wind particles below hundreds of MeV cannot trigger the DAQ system (Zhang et al. 2019). To investigate the accuracy of our on-orbit simulation, several validations on the spectral and spatial distribution are performed in several different radiation regions. Our results show good agreement between flight data and simulation.

### 2. Simulation Methods

The ultimate goal of the software is to generate simulation data in the same format as the flight data by simulating the cosmic-ray environment in the DAMPE orbit and the process of flight DAQ. For this purpose, the on-orbit simulation is separated into four steps as illustrated in Figure 1: the event generation, the primary particle determination, the secondary particle generation and the official Geant4 simulation. The first three steps are developed based on the ATMNC3 (Honda et al. 2004) software framework that was originally developed to calculate atmospheric neutrino flux, but which can also be used to simulate cosmic-ray flux near the Earth with proper modification. In our simulation, the GCR events are uniformly generated on an Earth-centered spherical surface (10 km above the DAMPE orbit, as displayed in Figure 2) with an isotropic flux, and then only inward-going particles are considered to ensure that they can reach the DAMPE orbit. A so-called back tracing process is then employed to determine if the particle is reasonable. For each event, the charge sign and the momentum direction are reversed, and the particle will propagate along its time-reversed trajectory in the geomagnetic field. If the particle reaches a distance of 10 Earth-radii, it is considered reasonable, otherwise a forbidden trajectory or a primary particle has been counted. During the backtracing, the tracing step length is optimized to balance CPU time consumption and trajectory accuracy. As the inputs of our simulation, we use the measurements of the AMS02 for the incoming fluxes of various GCR components, and only three GCR components of the most interest are considered, namely, protons (Aguilar et al. 2015b), helium nuclei (Aguilar et al. 2015a) and CREs (Aguilar et al. 2014). We consider in our simulation a full 3D description of the Earth’s atmospheric (NRLMSISE00 atmospheric model, Picone et al. 2002) and magnetic-field (IGRF-12 geomagnetic field model, Beggan et al. 2015) configurations using the latest available models. The NRLMSISE00 model describes the atmospheric density, composition and changes over time from sea level to 1000 km altitude. The IGRF12 is an international geomagnetic reference field model provided by the V-MOD International Cooperative Group of the International Association of Geomagnetism and Aeronomy (IAGA), covering the period of 2015–2020. The relative error of IGRF12 is about 1 nT while the typical field strength is about 40,000 nT. Event-by-event simulation of GCR propagation in the geomagnetic field and their interaction with the Earth’s atmosphere is implemented in the ATMNC3 framework also with the help of the DMPJET3 (Ranft et al. 2003) and PHITS (Niita et al. 2006) particle interaction models. The DMPJET3 is a new version of the MC event generator DPMJET for air shower simulation, and the PHITS is an MC particle propagation model that applies to particles in the energy range from 1 to 200 GeV.

After the ATMNC3 simulation, for the validation of our simulation results by comparing to flight data, all the primary and secondary particles generated are imported into the DAMPE official Geant4 simulation and reconstruction package (Jiang et al. 2020), where the relativistic inelastic and elastic interaction between a particle and the DAMPE detector is simulated. In the end, the MC data are reconstructed using the same algorithm for the flight data reconstruction, and a simulation data sample in the same format as flight data are generated.
3. Validation of On-orbit Simulation

In this section, we validate our on-orbit simulation by comparing the MC results with the CRE flux measurement of DAMPE. The primary CREs nearly isotropically propagating in interplanetary space, however, are deflected by geomagnetic force when entering the Earth’s magnetic field. As a consequence, the energy spectrum of primary CREs near the Earth is no longer a single power-law but rather a broken power-law with a magnetic latitude dependent cutoff, the so-called geomagnetic rigidity cutoff. The isotropic angular distribution is also broken by the rigidity cutoff and the so-called east–west asymmetry in the angular distribution is formed. Meanwhile, some secondary CREs generated in the top-layer of the atmosphere can also trigger the DAQ, so the CREs observed by DAMPE have two origins, one population is from the primary GCRs and the other comes from the interaction between GCRs and the atmosphere. By comparing the spectrum of primary and secondary combined CREs, we can validate both the determination method of the primary particles and their interaction with the Earth’s atmosphere.

3.1. Primary and Secondary CRE Flux with Geomagnetic Rigidity Cutoff

In the analysis, we select primary and secondary CRE flux calculated in the magnetic equator region where the geomagnetic field lines are almost parallel to the surface of the Earth. The charged particles are deflected by the Lorentz force, while in the polar region, the magnetic field lines are open and charged cosmic rays can reach the Earth’s surface along field lines. As a consequence, the geomagnetic rigidity cutoff is much higher in the equatorial region, where the typical value is about a few tens of GeV, well in the sensitive energy range of DAMPE. For comparison, we use the CRE fluxes near the geomagnetic equator in DAMPE orbit (Zang et al. 2017), and the DAMPE measurement of the CRE fluxes above 1.2 times the rigidity cutoff is normalized to the fluxes published by AMS02 (Aguilar et al. 2014).

As depicted in Figure 3, the CRE flux is broken into two parts due to the rigidity cutoff. Above the cutoff, primary CREs dominate the population, while below the cutoff they are almost all secondary ones since few primary CREs with that energy can reach the equator region due to the existence of the Earth and the geomagnetic field. By comparing the spectra of the flight data and MC data, spectral validation has been performed in all regions of DAMPE’s orbit. The primary and secondary flux of simulated CREs agrees well with flight data, especially near the rigidity cutoff region where the contributions from neither primary nor secondary are ignorable. The overall agreement demonstrates that the spectra of both primary and secondary CREs have been described well.
3.2. Angular Distribution of Primary and Secondary CREs

As mentioned above, the charged GCRs propagating in interplanetary space are highly isotropic in terms of their direction of arrival. Deflected by the Lorentz force while propagating in the geomagnetic field, they are no longer isotropic on LEO, but show an asymmetric angular distribution, i.e., the so-called east–west effect. That is, more positively charged particles are seen from west moving eastward, and vice versa for negatively charged particles. To describe the asymmetric angular distribution, the azimuth and zenith angle are defined in an Earth-centered coordinate system. In the system, the coordinate origin is at the center of the Earth, the azimuth angle is defined as geographical azimuth, particles from North have an azimuth of 0°, East is at 90°, South is at 180° and West is at 270°. Figure 4 displays the azimuthal distributions of CREs extracted from flight data of DAMPE and on-orbit simulation in four typical energy bins in an L interval of 1–1.14. Since DAMPE has no ability to distinguish the origins of CREs, flight data only give the entire azimuth distribution including both primary and secondary.

Figure 4. The CRE azimuthal distribution of MC and flight data are compared for different energy ranges of (a) [8.0, 10.3] GeV, (b) [12.1, 13.2] GeV, (c) [14.4, 15.7] GeV and (d) [17.1, 18.6] GeV in the McIlwain L value range of [1, 1.14]. The CRE azimuthal distribution is averaged over all zenith angles. The flight data CRE azimuthal distribution is from Zang et al. (2017). Quite good agreement between flight data and our simulation could be seen in the figures for the CRE azimuthal distribution in different energy ranges.

4. Research in Astronomy and Astrophysics, 22:045011 (5pp), 2022 April Wang et al.

Figure 5. The CRE zenithal distributions of MC and flight data on-orbit are compared. The flight data result here takes account of all the CREs observed by DAMPE, and the MC result is convolved with the DAMPE acceptance. An overall good agreement between flight data and simulation is also seen for the CRE zenithal simulation.

Our on-orbit simulation can trace and distinguish primary and secondary CREs. The overall good agreement of azimuth between flight data and primary+secondary demonstrates that the angular distribution has been described well. Benefiting
from this structured azimuth distribution, the method of estimating the fraction of the primary part is developed, and that is the main idea of simultaneous measurement flux of primary and secondary CREs. In the Earth-centered coordinate system, zenith is defined as the angle between the particle motion direction and the line pointing from the satellite to the Earth’s center. Figure 5 shows the zenithal direction for CREs collected on DAMPE orbit with an L range of 1–1.14. The distribution is convolved with the DAMPE acceptance. There is no significant discrepancy between real data and simulation, demonstrating zenith direction has been described well too.

4. Conclusion

A software package for the DAMPE on-orbit radiation environment simulation is developed using the ATMNC3 framework, within which state-of-the-art models of the geomagnetic field and the atmosphere are also integrated. By taking delicate consideration of the GCR propagation in the geomagnetic field and their interaction with the atmosphere, the software package is able to simulate the cosmic-ray radiation environment on the DAMPE orbit with sufficient accuracy and reasonable computational consumption. The simulation results are validated by comparing to the real observations for the CREs flux in several energies and geomagnetic latitude ranges. The overall agreements on the spectral and angular distributions demonstrate that the CRE radiation environment is simulated well. Specifically, the structured azimuthal distribution can serve as a tool to estimate the fraction of primary or secondary CREs. Our simulation tool thus could be helpful to the scientists and engineers working on the data analysis of space detectors after launch as well as in their design of a science satellite before launch. The software package can also be used to simulate the radiation environment in other LEOs by simply changing the orbital altitude, so it can also be applied in possible space science projects in the future, such as HERD (Zhang et al. 2014) and VLAST.5

Acknowledgments

This work is supported in part by the National Key R&D Program of China (2021YFA0718404), the National Natural Science Foundation of China (Grant Nos. 11773085, U1738207, and 12173098), the Youth Innovation Promotion Association CAS and the Scientific Instrument Developing Project of the Chinese Academy of Sciences, Grant No. GJSTD20210009.

References

Aguilar, M., Aisa, D., Alpat, B., et al. 2014, PhRvL, 113, 221102
Aguilar, M., Aisa, D., Alpat, B., et al. 2015a, PhRvL, 115, 211101
Aguilar, M., Aisa, D., Alpat, B., et al. 2015b, PhRvL, 114, 171103
Alemany, F. (DAMPE collaboration), et al. 2021, arXiv:2112.08860
Ambrosi, G. (DAMPE collaboration), et al. 2017, Natur, 552, 63
An, Q. (DAMPE collaboration), et al. 2019, SciA, 5, eaax3793
Beggan, C. D., Hamilton, B., Taylor, V., Macmillan, S., & Thomson, A. 2015, in Magnetic Interactions 2015, http://nora.nerc.ac.uk/id/eprint/511731/
Chang, J. (DAMPE collaboration), et al. 2017, ApJ, 95, 6
Duan, K.-K., Jiang, W., Liang, Y.-F., et al. 2019, RAA, 19, 132
Hilton, H. 1971, JGR, 28, 6952
Honda, M., Kajita, T., Kasahara, K., & Midorikawa, S. 2004, PhRvD, 70, 043008
Huang, Y.-Y., Ma, T., Yue, C., et al. 2020, RAA, 20, 153
Jiang, W., Yue, C., Cui, M.-Y., et al. 2020, ChPhL, 37, 119601
Niita, K., Sato, T., Iwase, H., et al. 2006, RadM, 41, 1080
Picone, J. M., Hedin, A. E., Drob, D. P., & Aikin, A. C. 2002, JGRA, 107, 15
Ranft, J., Engel, R., & Roesler, S. 2003, NuPhS, 122, 392
Xu, Z.-L., Duan, K.-K., Jiang, W., et al. 2022, FrP, 17, 34501
Xu, Z.-L., Duan, K.-K., Shen, Z.-Q., et al. 2018, RAA, 18, 027
Yuan, Q., & Feng, L. 2018, SCPMA, 61, 101002
Yuan, Q., Feng, L., Yin, P.-F., et al. 2017, arXiv:1711.10989
Yuan, Q., Qiao, B.-Q., Guo, Y.-Q., Fan, Y.-Z., & Bi, X.-J. 2021, FrP, 16, 24501
Zang, J., Yue, C., & Li, X. 2017, PoS, ICRC2017, 197
Zhang, S. N., Adriani, O., Albergo, S., et al. 2014, Proc. SPIE, 9144, 91440X
Zhang, Y.-Q., Guo, J.-H., Liu, Y., et al. 2019, RAA, 19, 023

5 chinaXiv:202203.00033