Polymerized 4-Fold Coordinated Carbonate Melts in the Deep Mantle

Chrystèle Sanloup, Jessica Hudspeth, Veronika Afonina, Benjamin Cochain, Zuzana Konôpková, Gerald Lelong, Laurent Cormier, Chiara Cavallari

To cite this version:
Chrystèle Sanloup, Jessica Hudspeth, Veronika Afonina, Benjamin Cochain, Zuzana Konôpková, et al.. Polymerized 4-Fold Coordinated Carbonate Melts in the Deep Mantle. Frontiers in Earth Science, Frontiers Media, 2019, 7, pp.72. 10.3389/feart.2019.00072. hal-02147554

HAL Id: hal-02147554
https://hal.archives-ouvertes.fr/hal-02147554
Submitted on 4 Jun 2019

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L’archive ouverte pluridisciplinaire HAL, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d’enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.
Polymerized 4-fold coordinated carbonate melts in the deep mantle

Chrystèle Sanloup 1,*, Jessica M. Hudspeth 2, Veronika Afonina 3, Benjamin Cochain 2, Zuzana Konôpková 1, Gérald Lelong 1, Laurent Cormier 1 and Chiara Cavallari 5

1 Institut de Minéralogie, Physique des matériaux et Cosmochimie, Sorbonne Université, CNRS, 75005 Paris, France
2 Institut des Sciences de la Terre de Paris, Sorbonne Université, CNRS, 75005 Paris, France
3 SUPA, Centre for Science at Extreme Conditions and School of Physics and Astronomy, University of Edinburgh, EH9 3FD, UK
4 DESY Photon Science, Notkestr. 85, 22607 Hamburg, Germany
5 European Synchrotron Radiation Facility, ESRF, 71 Avenue des Martyrs, 38000 Grenoble, France

Correspondence*: Corresponding Author chrystele.sanloup@sorbonne-universite.fr

ABSTRACT

Our understanding of the deep carbon cycle has witnessed amazing advances in the last decade, including the discovery of tetrahedrally coordinated high pressure (P) carbonate phases. However, little is known about the physical properties of their molten counterpart at moderate depths, while their properties at lower mantle conditions remain unexplored. Here, we report the structure and density of FeCO₃ melts and glasses from 44 GPa to 110 GPa by means of in situ x-ray synchrotron diffraction, and ex situ Raman and x-ray Raman spectroscopies. Carbon is fully transformed to 4-fold coordination, a bond change recoverable at ambient P. While low P melts react with silica, resulting in the formation of silico-carbonate glasses, high P melts are not contaminated but still quench as glasses. Carbonate melts are therefore polymerized, highly viscous and poorly reacting with silicates in the lower mantle, in stark opposition with their low P properties.

Keywords: carbonate melts, high pressure, x-ray diffraction, Raman, X-ray Raman, polymerization, deep mantle

1 INTRODUCTION

Although the lower mantle is mostly a reducing environment with the presence of reduced Fe [Frost et al., 2004; Smith et al., 2016], significant amount of subducted carbonates are estimated to be preserved [Litasov and Shatskiy, 2018]. Transition to 4-fold carbon was first predicted for crystalline CaCO₃ [Oganov et al., 2006; Aranan et al., 2007]. This transition strongly depends on the carbonate composition, occuring for CaCO₃ above 105 GPa [Lobanov et al., 2017], 80 GPa for MgCO₃ [Oganov et al., 2008; Boulard et al., 2011], and 50 GPa for FeCO₃ [Liu et al., 2015], while intermediate CaCO₃-MgCO₃-FeCO₃ compositions form a single tetrahedral carbonate phase [Merlini et al., 2017] unlike silicates. This transition induces
polymerization such as sheets or 3-membered rings for MgCO$_3$ (Oganov et al., 2008), and chains for CaCO$_3$ (Oganov et al., 2006). In contrast, our knowledge of carbonate melts structure at depth is scarce and limited to upper mantle pressures. The melting curves of CaCO$_3$, Na$_2$CO$_3$, and FeCO$_3$ have been measured over most of the upper mantle regime (Li et al., 2017; Kang et al., 2015), and viscosity measurements up to 6 GPa span several compositions (K$_2$Ca(CO$_3$)$_2$ and K$_2$Mg(CO$_3$)$_2$) by Dobson et al. 1996, CaCO$_3$ and natural dolomite by Kono et al. 2014, Na$_2$CO$_3$ by Stagno et al. 2018). Structural data instead have only been collected on molten CaCO$_3$ below 10 GPa (Hudspeth et al., 2018) while theoretical investigations of the properties of carbonate melts cover a larger $P$-range but are also limited to the carbon 3-fold stability field (Vuilleumier et al., 2014; Zhang and Liu, 2015; Du et al., 2018; Desmaele et al., 2019). One main question is therefore how this 3-fold to 4-fold transition translates in the molten state, and what are the consequences on the physical and chemical properties of carbonate melts? Of particular interest is the mobility and reactivity of carbonate melts in the lower mantle, knowing that these properties underpin the key role played by carbonate melts in mantle geodynamics through lubrication of plate tectonics, cratonic roots (Foley, 2008) and ascending plumes (Litasov et al., 2013).

The role of Fe in the deep carbon cycle is emphasized by the predominance of Fe-rich ferropericlase in diamond inclusions from the lower mantle (Kaminsky, 2012). The lowest transition $P$ from 3-fold to 4-fold C in FeCO$_3$ amongst carbonates justifies its choice as the first composition to investigate. Not only this transition occurs at less challenging experimental conditions, but it might be driven by Fe high spin to low spin transition at 40.4 GPa (Weis et al., 2017), a consequence of which being the large enrichment in Fe of (Mg,Fe)-carbonates coexisting with bridgmanite to almost pure FeCO$_3$ (Lobanov et al., 2015). Besides, high Fe concentration stabilizes (Ca,Mg,Fe)$_{IV}$CO$_3$ with respect to single cation 3-fold carbonates at mid mantle conditions (30-50 GPa) (Solomatova and Asimow, 2018). Formation of Fe-carbonates in the lower mantle might also result from carbonation of Fe-oxides ((Mg,Fe)O, FeOOH) (Boulard et al., 2012, 2018). Last but not least, FeCO$_3$ is a technical choice as it can be laser heated, which is required to reach lower mantle conditions without the need for additional laser coupler.

2 MATERIAL AND METHODS

Materials and chemical analyses

The starting natural crystalline siderite sample (mineralogical collection at Sorbonne Université) was loaded in the sample chamber laser-drilled in a rhenium gasket as approximately 20 µm-thick platelet between two equally thick platelets of compressed SiO$_2$ powder. The SiO$_2$ platelets act as thermal insulators and $P$-transmitting medium. Only one sample was used per $P$ point (Fig. 1) to avoid repeated laser-heatings, and preserve the chemical integrity of the sample. Six samples could be recovered after the experiments, embedded in epoxy and polished for analysis. Samples 8, 9 and 15 were carbon-coated for SEM imaging (Fig. 2), samples 8 and 15 were then repolished and gold-coated along with samples 13, 14 and 20 for electron microprobe analysis using a CAMECA SX-FIVE analyzer (EMPA) at the Camparis centre of Sorbonne Université (Table 1), using the following operating conditions: 15 keV, 10 nA. We used a defocussed beam size of 10 µm to get an average composition at the laser-heated spot.

$P$-$T$ conditions

We used diamond-anvil cells and a double-sided infra-red laser focussed down to 20 µm to generate high $T$ and $P$. For each $P$ point, targeted power was increased in 2 W increments from 20 to 50 W of power on each laser depending on $P$ until complete melting of the sample. Melting was identified by disappearance of diffraction peaks apart from SiO$_2$ peaks, and by the appearance of diffuse scattering. As we used the
off-axis heating system to avoid using carbon mirrors that would add to the x-ray background signal and
compromise processing of the scattered signal. $T$ could not be measured by pyrometric techniques. FeCO$_3$
melting curve has only been measured up to 20 GPa (Kang et al., 2015), where it reaches 1865 K. The
stishovite to CaCl$_2$ SiO$_2$ transition has been investigated up to 90 GPa (Fischer et al., 2018), this constrains
$T$ to a maximum of 2300 K at 79 GPa and 2500 K at 83 GPa as CaCl$_2$ is the observed SiO$_2$ structure for
the three highest $P$ runs, while stishovite is observed below. We therefore consider that x-ray diffraction
patterns were collected on molten FeCO$_3$ within the 2000 K-2500 K interval except for the highest $P$
point that is only constrained to below 3500 K from extrapolation of the stishovite-CaCl$_2$ Clapeyron slope
(Fischer et al., 2018).

$P$ is measured at room $T$ using fluorescence of a ruby sphere added in the sample
chamber (Mao et al., 1986) and SiO$_2$ equations of state (Andrault et al., 1998; Nishihara et al., 2005)
for quenched samples, and using only SiO$_2$ equations of state for molten samples with error bars on $P$
including the effect of a 2000 K-2500 K $T$-range, and up to 3500 K for the 110 GPa data point.

X-ray diffraction methods

We collected in situ high $P$-$T$ x-ray diffraction data in laser-heated diamond anvil cells at the extreme
conditions beamline P02.2 at the PETRAIII synchrotron. We used symmetric diamond-anvil cells equipped
with 70° opening Boehler-Almax seats in order to access a wider $q$-range up to 10 Å$^{-1}$, and reduce the
diamond Compton contribution as Boehler-Almax anvils are only 1.5 mm thick. The x-ray monochromatic
beam (42.7 keV) was focussed down to a size of $4 \times 6$ µm$^2$, allowing high spatial resolution in direct
space. To limit iron migration away from the laser heating spot due to Soret effect, the laser shutters were
opened only once the targeted power was reached, and held open for 10 s during which 10 x-ray diffraction
patterns of 1 s acquisition time were recorded on a Perkin-Elmer 2-D detector. 2-D patterns were integrated
using the Fit2D software (Hammersley et al., 1996). In order to isolate the scattered intensity from the
molten FeCO$_3$ only, each sample was removed from the gasket, and the gasket put back in place to collect
x-ray data on the empty cell. Obtained patterns were then scaled vertically to match the baseline of x-ray
patterns collected on the starting crystalline sample under $P$ (Sanloup and de Grouchy, 2018). This last
step ensures that any $P$ effect on the background is corrected for. Amongst eight successful runs (Table 1)
for which full melting was observed, intensity from molten FeCO$_3$ could only be processed for the highest
$P$ run for which the sample vs SiO$_2$ platelets thickness ratio was slightly higher, the scattered intensity
being too weak for the lower $P$ points. All glass patterns could be processed. The x-ray diffracted intensity
data are converted into the structure factor, $S(q)$ (Fig. ?? and Fig.4), using the Ashcroft-Langreth formalism.
The radial distribution function $g(r)$ (Fig 3B), that describes ion-ion contributions in real space, is obtained
by Fourier transforming of $S(q)$,

$$g(r) = \frac{1}{2\pi^2 n} \int_{0}^{\infty} q S(q) \sin(qr) dq$$  

(1)

where $n = \frac{aN_A}{M}$, $N_A$ is the Avogadro number, $M$ the mean atomic molar mass, and $\rho$ the density.

Density measurements

The method to derive density from x-ray diffraction data on melts compressed in diamond-anvil cell
experiments (Eggert et al., 2002; Sanloup et al., 2013) consists in minimizing the oscillations in $g(r)$ where
there should not be any signal, i.e. below the minimum interatomic distance ($r < 0.95$ Å here). This method
requires that the background, essentially the Compton signal from the diamond anvils that dominates the
total diffracted intensity, is perfectly subtracted.
As the C-O contribution is distinct on \( g(r) \) of quenched glasses up to 83 GPa, we also ran consistency checks by fixing the C-O coordination number to 4 as indicated by x-ray Raman spectra (cf. Results section), and simulating the C-O contribution using the obtained density values against a gaussian with the following equation:

\[
g(r) = \frac{1}{nS_\infty} \frac{A}{\sigma\sqrt{2\pi}} \exp\left(-\frac{(r - d)^2}{2\sigma^2}\right)
\]

where

\[
S_\infty = \frac{\Sigma p K_p^2}{Z_{\text{tot}}^2}
\]

and

\[
A = \int \frac{4\pi r^2}{\sigma\sqrt{2\pi}} \exp\left(-\frac{(r - d)^2}{2\sigma^2}\right) dr
\]

with \( K_p \), the effective atomic number \(^{\text{Eggert et al., 2002}}\), \( Z_{\text{tot}} \) the total atomic number of the compositional unit \( \text{e.g. FeCO}_3 \), \( CN \) the C-O coordination number fixed to 4, \( d \) the C-O inter-atomic distance, and \( \sigma \) a parameter depending on structural disorder, \( \sigma = k\sqrt{d} \) where \( k \) is an adjustable parameter \(^{\text{Hosemann and Bagchi, 1962}}\) with a value of 0.11 here. The C-O contribution to \( g(r) \) thus calculated adequately fits the experimental ion-ion contribution (dashed lines on Fig.3B), hence comforting the obtained density values.

### 3 | RESULTS

All samples are systematically quenched as a glass. Chemical integrity of FeCO\(_3\) molten spheres is observed for runs conducted above 40 GPa, apart from a marginal fraction at the glass-SiO\(_2\) interface in one sample showing enrichment of the \( P \)-transmitting medium in Fe and C. Instead, the lowest \( P \) samples, \textit{i.e.} 11 GPa and 15 GPa, have reacted with the SiO\(_2\) \( P \)-transmitting medium. This is shown by SEM imaging (Fig.2) and EMPA analysis on sample 8 (Table 1). High \( P \) carbonate melts are thus much less reactive than low \( P \)
melts. This might not contradict the observed reactivity of high $P$ crystalline MgCO$_3$ with SiO$_2$ (Seto et al. 2008; Maeda et al. 2017) due to the much longer heating durations (20-240 minutes against 10 seconds heating duration in this work); alternatively, Fe stabilizing effect on high $P$ carbonates could be at stake. We observe no disproportionation of Fe as was reported in the crystalline state in some studies (Boulard et al., 2011; Cerantola et al., 2017) but not in others (Liu et al., 2015). This might be due to different $P$-$T$ paths followed, i.e. flash heating here instead of continuous $T$ increase (Boulard et al., 2011; Cerantola et al., 2017).

A striking characteristic of glassy FeCO$_3$ is its strong first sharp diffraction peak (FSDP) that persists in the structure factor up to the highest $P$ investigated (Fig.3A), indicative of a strong medium-range order. This is in stark contrast to silicate glasses that lose their medium-range order with increased $P$ (Sato and Funamori, 2008), but consistent with $ab$ initio calculations on carbon-bearing silicate melts reporting $P$-induced polymerisation of carbonate species into dimers and with the silicate network (Ghosh et al., 2017; Solomatova and Asimow, 2019). A second noticeable feature is the decrease of the contribution at 4 Å$^{-1}$ attributed in molten carbonates to the O-O bond (Wilding et al., 2016). On radial distribution functions, $g(r)$ (Fig.3B), the C-O contribution is clearly visible at 1.2-1.3 Å with none or little overlap with the second contribution (Fe-O and O-O) at ~2 Å in the glass, and with some overlap in the melt. No significant structural changes are observed between molten and quenched glassy state at 110 GPa, apart from a generally lower intensity in the melt due to the high $T$ and consequent higher degree of disorder. For $g(r)$, this weaker intensity translates into broader C-O and Fe-O/O-O contributions in the molten state.

For glasses quenched at 11 GPa and 15 GPa, the x-ray structure factor, $S(q)$, is intermediate between that of pure SiO$_2$ glass (Sato and Funamori, 2008) and high-$P$ FeCO$_3$ glasses (Fig.4). SEM image of sample 8 (15 GPa, Fig.2) shows heterogeneities in the quenched glass, which indicates that the x-ray structure factor likely averages at least two types of glass structure and therefore data cannot be interpreted quantitatively.

The x-ray Raman C K-edge spectrum of quenched FeCO$_3$ glass shows no presence of sp2 3-fold carbon characterized by an intense $\pi^*$ peak at 290 eV (Fig.5 $\pi^*$ peak). Only the $\sigma^*$ peak of tetrahedrally coordinated carbon (Shieh et al., 2013) is visible (Fig5 $\sigma^*$ peak). The totally missing $\pi^*$ peak is indicative of a fully sp3 state of carbon atoms in the siderite glass. $P$-induced coordination changes of major cations in silicate melts (e.g. Si, Al) were first reported from the study of glasses quenched from high $P$ (Meade, Hemley and Mao, 1992; Yarger et al., 1995), and later confirmed by $in$ situ studies in the molten phase (Sanloup et al., 2013; Drewitt, 2015). However, the opposite, i.e. coordination change occurring only in the quenched glass, not in the high $P$ melt, have not been reported nor been theoretically predicted. The 3-fold to 4-fold transition therefore occurs in molten Fe-carbonates at $P$ less or equal to 51 GPa. This transition is preserved upon quenching to the glassy state, and is recoverable at ambient conditions, opening the way to the synthesis of a new class of glassy materials. Two broad bands are observed in the Raman spectra (Fig.9), very different from those of the only two carbonate systems that quench as glasses at room $P$, MgCO$_3$-K$_2$CO$_3$ and La(OH)$_3$-Ca(OH)$_2$-CaCO$_3$-CaF$_2$BaSO$_4$ (Sharma and Simons, 1979), that are essential dominated by the strong CO$_3^{2-}$ stretching mode at ~1080 cm$^{-1}$. Instead, present Raman spectra are reminiscent of those reported for calcium silicate glasses (Fig.6) albeit at higher Raman shift values for the broadest band (1200-1600 cm$^{-1}$ for glassy FeCO$_3$ vs 850-1100 cm$^{-1}$ for calcium silicate glasses).

Density values are reported in Fig.7 along with predictions for lower $P$ melt properties (Kang et al., 2015), $P$-evolution of crystalline siderite, and with the Earth’s seismological PREM model (Dziewonski and Anderson, 1981). Density profile below 40 GPa is calculated using $K_T$, value of 80.23 GPa (Kang et al., 2015), consistent with that reported for molten calcite (Hudspeth et al., 2018), and density at room $P$ of 2500 kg·m$^{-3}$ by assuming a similar density jump upon melting as for other carbonates for which
Density of non-crystalline FeCO$_3$ remains considerably lower than that of its crystalline counter parts, even at the highest investigated $P$, by approximately 15%. The situation is thus very different from that of molten and crystalline silicates which density converge at deep mantle conditions [Petitgirard et al., 2015; Sanloup, 2016], and such difference could be attributed to the very strong medium-range order preserved in tetrahedral high $P$ carbonate melts while it is mostly collapsed by 5 GPa in silicate melts. That high $P$ FeCO$_3$ melts quench as glasses contrasts with the behaviour observed at lower $P$, and suggests an important increase of carbonate melt viscosity consistent with the observation of a very strong medium-range order. It is also opposite to the behaviour of molten basalt that systematically quenches as crystalline phases above 11 GPa [Sanloup et al., 2013] and as a glass below. The strongly reduced chemical reactivity of high $P$ FeCO$_3$ melts with silica along with their glass-forming ability suggest that unlike at lower $P$, tetrahedral carbonate melts are not pervasive, which could contribute to the longevity of carbonates in the deep mantle where allowed by oxydizing conditions or slow reduction kinetics [Litascov and Shatskiy, 2018].
FUNDING

The research leading to these results has received funding from the European Community’s Seventh Framework Programme (FP7/2007-2013) under grant agreements no. 312284 and 259649 (European Research Council starting grant to C.S.). Portions of this research were carried out at the light source PETRA III at DESY, a member of the Helmholtz Association (HGF). The laser heating system on beamline P02.2 is funded by the German BMBF (project number 05K10RFA).

ACKNOWLEDGMENTS

We acknowledge E. Boulard for providing the starting siderite sample, K. Glazyrin for his help with ex situ diamond-anvil cell laser-heating synthesis in PETRA III, L. Rémuat at Museum National d’Histoire Naturelle (Paris, France) for gold coating of the recovered samples, the ESRF (Grenoble, France) and PETRAIII (Hamburg, Germany) for provision of synchrotron radiation facilities.

REFERENCES

Andrault, D., Fiquet, G., Guyot, F., and Hanfland, M. (1998). Pressure-induced Landau-type transition in stishovite. Science 282, 720–724. doi:10.1126/science.282.5389.720.

Arapan, S., De Almeida, J. S., and Ahuja, R. (2007). Formation of sp(3) hybridized bonds and stability of CaCO$_3$ at very high pressure. Phys. Rev. Lett. 98, 268501. doi:10.1103/PhysRevLett.98.268501.

Boehler, R., Ross, M., and Boercker D. B. (1997). Melting of LiF and NaCl to 1 Mbar: Systematics of Ionic Solids at Extreme Conditions Phys. Rev. Lett. 78, 4589–4592.

Boulard, E., Gloter, A., Corgne, A., Antonangeli, D., Auzende, A.-L., Perrillat, J.-P., et al. (2011). New host for carbon in the deep Earth. Proc. Natl Acad. Sci. USA 108, 5184–5187. doi:10.1073/pnas.1016934108.

Boulard, E., Guyot, F., and Fiquet, G. (2012). The influence on Fe content on Raman spectra and unit cell parameters of magnesite-siderite solid solutions. Phys. Chem. Miner. 39, 239–246. doi:10.1007/s00269-011-0479-3.

Boulard, E., Guyot, F., Menguy, N., Corgne, A., Auzende, A.-L., Perrillat, J.-P., and Fiquet, G. (2018). CO$_2$-induced destabilization of pyrite-structured FeO$_2$H$_x$ in the lower mantle. Natl. Sci. Rev. 5, 870–877. doi:10.1093/nsr/nwy032.

Cerantola, V., Bykova, E., Kupenko, I., Merlini, M., Ismailova, L., McCammon, C., et al. (2017). Stability of iron-bearing carbonates in the deep Earth’s interior. Nat. Commun. 8. doi:10.1038/ncomms15960.

Desmaele, E., Sator, N., Vuilleumier, R. and Guillot, B. (2019). Atomic simulations of molten carbonates: Thermodynamic and transport properties of the Li$_2$CO$_3$-Na$_2$CO$_3$-K$_2$CO$_3$ system. J. Chem. Phys. 150, 094504.

Du, X., Wu, M., Tse, J. S., and Pan, Y. (2018). Structures and Transport Properties of CaCO3 Melts under Earth’s Mantle Conditions ACS Earth Space Chem 2, 1–8. doi:10.1021/acsearthspacechem.7b00100.

Dziewonski, A. M. and Anderson, D. L. (1981). Preliminary reference Earth model. Phys. Earth Planet. Int. 25, 297–356.

Drewitt, J. W. E., Jahn, S., Sanloup, C., de Grouchy, C., Garbarino, G. and Hennet L. (2015). Development of chemical and topological structure in aluminosilicate liquids and glasses at high pressure. J. Phys.: Cond. Matt. 27, 105103.

Eggert, J. H., Weck, G., Loubeyre, P., and Mezouar, M. (2002). Quantitative structure factor and density measurements of high-pressure in diamond anvil cells by x-ray diffraction: Argon and water. Phys. Rev. B 65, 174105.
Fischer, R. A., Campbell, A. J., Chidester, B. A., Reaman, D. M., Thompson, E. C., Pigott, J. S., et al. (2018). Equations of state and phase boundary for stishovite and CaCl₂-type SiO₂. *Am. Mineral.* 103, 792–802.

Foley, S. F. (2008). Rejuvenation and erosion of the cratonic lithosphere. *Nature Geosci.* 1, 503–510. doi:10.1038/ngeo261.

Frost, D., Liebske, C., Langenhorst, F., McCammon, C., Tronnes, R., and Rubie, D. (2004). Experimental evidence for the existence of iron-rich metal in the Earth’s lower mantle. *Nature* 428, 409–412. doi:10.1038/nature02413.

Ghosh, D. B., Bajgain, S. K., Mookherjee, M., and Karki, B. B. (2017). Carbon-bearing silicate melt at deep mantle conditions. *Sci. Rep.* 7. doi:10.1038/s41598-017-00918-x.

Hammersley, A. P., Svensson, S. O., Hanfland, M., Fitch, A. N., and Hausermann, D. (1996). Two-dimensional detector software: From real detector to idealised image or two-theta scan. *High Press. Res.* 14, 235–248.

Hosemann, R. and Bagchi, S. N. (1962). *Direct Analysis of Diffraction by Matter* (Amsterdam: North-Holland).

Huang, F., Wu, Z., Huang, S., and Wu, F. (2014). First-principles calculations of equilibrium silicon isotope fractionation among mantle minerals. *Geochem. Cosmochim. Acta* 140, 509–520. doi:10.1016/j.gca.2014.05.035.

Kaminsky, F. (2012). Mineralogy of the lower mantle: A review of ‘super-deep’ mineral inclusions in diamond. *Earth Sci. Rev.* 110, 127–147. doi:10.1016/j.earthsci.2011.10.005.

Litasov, K. D., Shatskiy, A., Ohtani, E., and Yaxley, G. M. (2013). Solidus of alkaline carbonatite in the deep mantle. *Geology* 41, 79–82. doi:10.1130/G33488.1.

Litasov, K. D., Shatskiy, A., Ohtani, E., and Yaxley, G. M. (2013). Solidus of alkaline carbonatite in the deep mantle. *Geology* 41, 79–82. doi:10.1130/G33488.1.

Li, Z., Li, J., Lange, R., Liu, J., and Militzer, B. (2017). Determination of calcium carbonate and sodium carbonate melting curves up to Earth’s transition zone pressures with implications for the deep carbon cycle. *Earth Planet. Sci. Lett.* 457, 395–402.

Lobanov, S. S., Goncharov, A. F., and Litasov, K. D. (2015). Optical properties of siderite (FeCO₃) across the spin transition: Crossover to iron-rich carbonates in the lower mantle. *Am. Mineral.* 100, 1059–1064. doi:10.2138/am-2015-5053.
Lobanov, S. S., Stevanovic, V., Gavryushkin, P. N., Litasov, K. D., Greenberg, E., Prakapenka, V. B., Oganov, A. R. and Goncharov, A. F. (2017). Raman spectroscopy and x-ray diffraction of sp(3) CaCO$_3$ at lower mantle pressures. *Phys Rev B* 96, 104101. doi:10.1103/PhysRevB.96.104101.

Maeda, F., Ohtani, E., Kamada, S., Sakamaki, T., Hirao, N., and Ohishi, Y. (2017). Diamond formation in the deep lower mantle: a high-pressure reaction of MgCO$_3$ and SiO$_2$. *Sci. Rep.* 7. doi:10.1038/srep40602.

Mao, H. K., Xu, J., and Bell, P. M. (1986). Calibration of the ruby pressure gauge to 800 kbar under quasi-hydrostatic conditions. *J. Geophys. Res.* 91, 4673–4676.

Mao, H. K., Xu, J., and Bell, P. M. (1986). Calibration of the ruby pressure gauge to 800 kbar under quasi-hydrostatic conditions. *J. Geophys. Res.* 91, 4673–4676.

Merlini, M., Cerantola, V., Gatta, G. D., Gemmi, M., Hanfland, M., Kupenko, I., et al. (2017). Dolomite-IV: Candidate structure for a carbonate in the Earth’s lower mantle. *Am. Mineral.* 102, 1763–1766. doi:10.2138/am-2017-6161.

Nishihara, Y., Nakayama, K., Takahashi, E., Iguchi, T., and Funakoshi, K. (2005). P-V-T equation of state of stishovite to the mantle transition zone conditions. *Phys. Chem. Min.* 31, 660–670.

Oganov, A., Glass, C., and Ono, S. (2006). High-pressure phases of CaCO$_3$: Crystal structure prediction and experiment. *Earth Planet. Sci. Lett.* 241, 95–103. doi:10.1016/j.epsl.2005.10.014.

Oganov, A. R., Ono, S., Ma, Y., Glass, C. W., and Garcia, A. (2008). Novel high-pressure structures of MgCO$_3$, CaCO$_3$ and CO$_2$ and their role in Earth’s lower mantle. *Earth Planet. Sci. Lett.* 273, 38–47.

Petitgirard, S., Malfait, W. J., Sinmyo, R., Kupenko, I., Hennet, L., Harries, D., et al. (2015). Fate of MgSiO$_3$ melts at core-mantle boundary conditions. *P. Natl. Acad. Sci. USA* 112, 14186–14190. doi:10.1073/pnas.1512386112.

Sanloup, C. (2016). Density of magmas at depth. *Chem. Geol.* 429, 51–59. doi:10.1016/j.chemgeo.2016.03.002.

Sanloup, C. and de Grouchy, C. J. L. (2018). X-ray diffraction structure measurements (Amsterdam, The Netherlands: Elsevier), chap. 5. 137–154.

Sanloup, C., Drewitt, J. W. E., Konôpková, Z., Dalladay-Simpson, P., Morton, D. M., Rai, N., et al. (2013). Structural change in molten basalt at deep mantle conditions. *Nature* 503, 104–107.

Sato, T. and Funamori, N. (2008). Sixfold-coordinated amorphous polymorph of SiO$_2$ under high pressure. *Phys. Rev. Lett.* 101, 255502.

Sharma, S. and Simons, B. (1979). *Raman study of K$_2$CO$_3$-MgCO$_3$ glasses* (Carnegie Institute), vol. 79. 322–326.

Shieh, S. R., Jarrige, I., Wu, M., Hiraoka, N., Tse, J. S., Mi, Z., et al. (2013). Electronic structure of carbon dioxide under pressure and insights into the molecular-to-nonmolecular transition. *Proc. Natl. Acad. Sci. U. S. A.* 110, 18402–18406. doi:10.1073/pnas.1305161110.

Smith, E. M., Shirey, S. B., Nestola, F., Bullock, E. S., Wang, J., Richardson, S. H., et al. (2016). Large gem diamonds from metallic liquid in earth’s deep mantle. *Science* 354, 1403–1405.
Sanloup et al.

High Pressure Carbonate Melts

Solomatova, N. V., Caracas, R., and Manning, C. E. (2019). Carbon sequestration during core formation implied by complex carbon polymerization. Nat. Comm. 10, 789. doi:10.1038/s41467-019-08742-9.

Stagno, V., Stopponi, V., Kono, Y., Manning, C. E. and Tetsuo, I. (2018). Experimental determination of the viscosity of Na$_2$CO$_3$ melt between 1.7 and 4.6 GPa at 1200-1700 degrees C: Implications for the rheology of carbonatite magmas in the Earth’s upper mantle. Chem. Geol. 501,19–25. doi:10.1016/j.chemgeo.2018.09.036.

Vuilleumier, R., Seitsonen, A., Sator, N., and Guillot, B. (2014). Structure, equation of state and transport properties of molten calcium carbonate (CaCO$_3$) by atomistic simulations. Geochim. Cosmchim. Acta 141, 547–566. doi:10.1016/j.gca.2014.06.037.

Weis, C., Sternemann, C., Cerantola, V., Sahle, C. J., Spiekermann, G., Harder, M., et al. (2017). Pressure driven spin transition in siderite and magnesiosiderite single crystals. Sci. Rep. 7. doi:10.1038/s41598-017-16733-3.

Wilding, M. C., Wilson, M., Alderman, O. L. G., Benmore, C., Weber, J. K. R., Parise, J. B., et al. (2016). Low-dimensional network formation in molten sodium carbonate. Sci. Reports 6, 24415.

Yarger, J. L., Smith, K. H., Nieman, R. A., Diefenbacher, J., Wolf, G. H., Poe, B. T., McMillan, P. F. (1995). Al Coordination Changes in High-Pressure Aluminosilicate Liquids. Science 270, 1964–1967.

Zhang, Z., and Liu, Z. (2015). High pressure equation of state for molten CaCO$_3$ from first principles simulations. Chin. J. Geochem. 34, 13–20.
FIGURE CAPTIONS

Figure 1. Microphotograph of the sample after laser heating at 110 GPa. Single shot laser heating resulted in the formation of a quasi-spherical pure carbonate glass that was removed from the gasket for EPMA and/or SEM analyses.

Figure 2. SEM images of recovered samples. Low $P$ sample 8 (a) shows pervasive contamination of carbonate sample with SiO$_2$ $P$-transmitting medium. High $P$ samples 9 (b) and 15 (c) show that chemical integrity of carbonate melt (homogeneous light gray zone) was preserved.
**Figure 3.** Structure of non-crystalline FeCO$_3$ at high pressures. (A) Structure factor, $S(q)$, for all quenched glasses (black) and the highest $P$ melt (red). (B) Corresponding radial distribution functions, $g(r)$. Dashed lines are fits to the C-O contribution at 1.2-1.3 Å where there is no overlap with farther contributions.

**Figure 4.** Structure factor, $S(q)$, for low $P$ reacted FeCO$_3$+SiO$_2$ glass (black), compared to SiO$_2$ glass at 20 GPa (Sato and Funamori, 2008) (brown) and FeCO$_3$ glass at 44 GPa (red). Low $P$ sample 8 (15 GPa) shows intermediate structure between SiO$_2$ glass and high $P$ FeCO$_3$ glasses.
Figure 5. X-ray Raman spectra collected at the carbon K-edge on crystalline siderite and high $P$-quenched FeCO$_3$ glasses at ambient conditions. The disappearance of the $\pi^*$ feature, which is solely related to the three-fold coordinated carbon, is a spectroscopic evidence of a full four-fold coordination state in the glassy structure of FeCO$_3$.

Table 1. Run conditions, quenched products and their chemical composition in wt% obtained from EMPA. One standard deviations are given in parentheses. Starting natural siderite sample also contained less than 0.1 wt% CaO and MnO.

| #  | $P$ melt/glass (GPa) | CO$_2$   | FeO       | MgO       | SiO$_2$   | Total     |
|----|----------------------|----------|-----------|-----------|-----------|-----------|
| 6  | 11.6/ –              | not recovered, reaction confirmed by XRD (Fig.4) |
| 8  | 15/14                | 25.7(9.2) | 42.2(5.8) | 0.1(0.1)  | 24.2(6.7) | 92.2      |
| 15 | 51/44                | 40.6(0.5) | 58.9(9.3) | 0.3(0.1)  | 0.3(0.2)  | 100.0     |
| 13 | 55/ –                | 41.2(2.6) | 54.4(1.3) | 0.3(0.2)  | 2.0(1.9)  | 98.0      |
| 20 | 63/57                | 36.6(6.8) | 57.7(1.3) | 0.4(0.4)  | 0.7(0.8)  | 95.4      |
| 9  | 79/72                | not analyzed, C-coated for SEM (Fig.2) |
| 12 | 83/77                | not recovered |
| 14 | 110/108              | 37.8(8.7) | 58.5(1.1) | 0.2(0.1)  | 0.7(0.8)  | 97.3      |

sample for x-ray Raman 59 not analyzed, only glass sphere preserved
Figure 6. Raman spectra collected on high $P$-quenched FeCO$_3$ glasses (runs 9 and 14) at ambient conditions.
Figure 7. Density evolution of glassy, molten and crystalline siderite with pressure. Molten low $P$ siderite (plain curve), high $P$ data on glass (black points) and the highest $P$ melt (red point), crystalline equation of state (dashed curve) includes the transition from high spin siderite I to low spin siderite II at 50 GPa (Liu et al., 2015).