The footprint of cometary dust analogues. - II

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The footprint of cometary dust analogues – II. Morphology as a tracer of tensile strength and application to dust collection by the Rosetta spacecraft

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ABSTRACT
The structure of cometary dust is a tracer of growth processes in the formation of planetesimals. Instrumentation on board the Rosetta mission to comet 67P/Churyumov–Gerasimenko captured dust particles and analysed them in situ. However, these deposits are a product of a collision within the instrument. We conducted laboratory experiments with cometary dust analogues, simulating the collection process by Rosetta instruments (specifically COSIMA, MIDAS). In Paper I, we reported that velocity is a key driver in determining the appearance of deposits. Here in Paper II, we use materials with different monomer sizes, and study the effect of tensile strength on the appearance of deposits. We find that mass transfer efficiency increases from ∼1 up to ∼10 per cent with increasing monomer diameter from 0.3 to 1.5 μm (i.e. tensile strength decreasing from ∼12 to ∼3 kPa), and velocities increasing from 0.5 to 6 m s⁻¹. Also, the relative abundance of small fragments after impact is higher for material with higher tensile strength. The degeneracy between the effects of velocity and material strength may be lifted by performing a closer study of the deposits. This experimental method makes it possible to estimate the mass transfer efficiency in the COSIMA instrument. Extrapolating these results implies that more than half of the dust collected during the Rosetta mission has not been imaged. We analysed two COSIMA targets containing deposits from single collisions. The collision that occurred closest to perihelion passage led to more small fragments on the target.

Key words: space vehicles: instruments – methods: laboratory – comets: individual: 67P/Churyumov–Gerasimenko – interplanetary medium – planets and satellites: formation – dust, extinction.

1 INTRODUCTION
Dust growth is the starting point of planet formation in the early Solar system. However, various barriers exist that inhibit growth from dust to planetesimal, and onwards to larger bodies (Dominik et al. 2007; Johansen et al. 2014; Blum 2018). On Earth, little or no geological remnants can be found from the earliest growth phases, as they have long since disappeared, having been heated and processed. Consequently, the most pristine remnants of the protosolar nebula surviving to this day are found in comets. Being kilometre-sized bodies that have spent most of their existence in cold regions of the Solar system, they contain ‘fossilized’ evidence of early dust growth processes (Blum et al. 2017).

This notion is supported by in situ measurements of cometary dust particles by spacecraft (Levasseur-Regourd et al. 2018). Most recently, the Rosetta mission to comet 67P/Churyumov–Gerasimenko (hereafter 67P) has provided a wealth of information on the dust population in the coma. Two instruments on board Rosetta were able to image particles smaller than 1 mm: COSIMA (10 to a few 100 μm, Kissel et al. 2007) and MIDAS (1 to a few 10 μm, Riedler et al. 2007). These images show that dust particles in the coma consist of aggregates of dust grains with a hierarchical structure down to sub-micron scales (Schulz et al. 2015; Bentley et al. 2016; Mannel et al. 2016). Furthermore, various morphologies of dust aggregates were found (Güttler et al. submitted). Discussion is ongoing whether these morphologies relate to different dust species (Della Corte et al. 2015; Fulle et al. 2015, 2016b; Langevin et al. 2016; Merouane et al. 2016; Fulle & Blum 2017). However, all of these experiments only look at the particles after they interacted...
with the spacecraft during collection, and therefore we do not necessarily have sufficient information on the particles before they were collected by the spacecraft.

One key line of research to interpret cometary dust data from space missions is therefore to perform laboratory experiments with cometary analogue materials. These allow to better understand dust growth processes, and moreover can interpret data from specific instruments where interaction with the equipment naturally influences the outcome of an in situ measurement. Specifically, in the case of the COSIMA and MIDAS instruments, the dust particles imaged are the product of a collision with the instrument target surface, and possibly before that with the collection funnel. Experiments that simulate these collisions provide a tool for interpreting the spacecraft data.

Silicate aggregates provide a suitable analogue for cometary material after the ices have been sublimated. Their collision physics have been studied extensively both theoretically and experimentally. Experiments have been conducted both with polydisperse and monodisperse material (for an overview, see Blum & Wurm 2008; Gütter et al. 2010; Blum 2018). However, up to this point studies have mostly focused on aggregate growth, and less on the part of the particle left on a solid surface after a collision. We performed laboratory experiments that aim to relate the properties of cometary dust analogues to the deposits they leave on a solid surface after impact. The aim of the first series of experiments, presented in Ellerbroek et al. (2017, hereafter ‘Paper I’), was to relate deposit morphology to particle size and velocity. In Paper I, a single polydispersed material was used as an analogue for cometary dust. A key insight gained from that study is that during impacts, aggregates fragmented and a large fraction of mass was lost in the instrument. Also, the velocity was seen as the main driver of the appearance of deposits. At impact velocities below the breaking or fragmentation barrier (∼2 m s⁻¹ for the material used), particles either stick to or bounce off the target, leaving respectively a single (undamaged) deposit or a shallow footprint of loose monomers on the surface. Above the fragmentation barrier, particles fragment and leave pyramid-shaped deposits. An important open question that remains is: how do material properties (density, packing, monomer size, composition) and the resulting tensile strength influence deposit morphologies?

In this paper, we present a second series of experiments and study the combined effect of tensile strength and velocity on the appearance of deposits. We use three different types of silicate aggregates, consisting of monodisperse monomers of a single size, which directly relates to the aggregate’s tensile strength (Gundlach et al. 2018). This allows us for the first time to quantify the amount of mass transferred to the target plate, as a function of velocity and tensile strength, in a parameter range overlapping with the Rosetta experiments.

The paper is structured as follows. In Section 2, we describe the test material, experimental setup, and the analysis method used. In Section 3, we present the quantitative results relating to mass transfer and deposit characteristics as a function of (pre-collection) dust properties. In Section 4, we discuss our results in the context of cometary dust measurements made by the Rosetta spacecraft.

2 METHODS

In this section, we describe the test material used, the experimental setup, and subsequent data analysis.

2.1 Test material

The particles used in the experiments are aggregates of pure SiO₂, with a material density of ρₘ = 2.0 × 10³ kg m⁻³. Three different samples were used: aggregates of pure SiO₂ spherical monomers, with diameter d₀ = (0.3, 1.0, 1.5) μm, with a standard deviation of 4 percent (Blum et al. 2006). Optical microscope and scanning electron microscope (SEM) images are shown in Fig. 1; the material properties are summarized in Table 1. As the volume filling factor is of order φ ≈ 0.3–0.4, the resulting bulk density of these aggregates is ρₘ ≡ ρₘφ = (0.7 ± 0.1) × 10³ kg m⁻³. Note that the bulk density is not a function of monomer size, but rather scales with φ. The material naturally forms aggregates in their storage canisters; we sieved these aggregates to obtain a particle diameter range of 100–400 μm, which overlaps the range in deposit sizes observed by COSIMA (Hornung et al. 2016).

An important parameter for the collision physics of the aggregates is the tensile strength σ, which decreases as a function of monomer size. Gundlach et al. (2018) performed experiments to measure the tensile strength of the three types of material also used in this paper. Using the scaling relation they provide for different filling factors, we estimate the tensile strength of the materials to vary between ~3 and ~12 kPa for d₀ = 1.5–0.3 μm (Table 1). This is also reflected by the shape of the aggregates that formed in the storage canisters: 1.5 μm aggregates look more spherical (likely due to erosion while sieving) while aggregates of the 0.3 μm material retain a more irregular shape even after sieving. However, the monomers themselves are smooth: Poppe, Blum & Henning (2000) measured surface roughness parameters of 0.5 and 1.2 μm spheres, and find that their surfaces are smooth on a sub-nanometre scale.

2.2 Experimental setup

We use the same experimental setup as described in Paper I. Here, we give a brief summary of the setup and some improvements that were made compared to Paper I.

A schematic drawing of the experimental setup is displayed in Fig. 2. For each experiment, a collection of 20 aggregates with diameter 100–400 μm were loaded on to a piston in a vacuum chamber. The chamber was depressurized to ~0.03 mbar, so as to minimize the influence of air drag. Although this pressure is much lower than that experienced during a comet impact, we assume a similar effect on the deposit properties.
higher than the outer space environment where Rosetta operates, it suffices the purpose of minimizing the influence of air drag on the collision time-scales in the experiment. Similarly, the collision dynamics are not affected significantly by gravity (see Paper I, fig. 4c). Our experiments thus approximate the circumstances in space, where contact forces are dominant in determining the collision dynamics.

A current pulse was applied to a lifting magnet, thereby launching the aggregates vertically on to a COSIMA target (polished silver) placed 8 cm above the piston. By tuning the voltage, impact velocities in the range 0.5–6.5 m s\(^{-1}\) were reached, with a spread in velocity of \(\sim 0.5\) m s\(^{-1}\) within a single shot. Depending on its properties, a particle may stick, bounce or breakup on the target surface, resulting in a collection of fragments left on the target surface. A collection of fragments resulting from the collision of a single particle is referred to as a ‘deposit’.

The particle collisions were recorded by a high-speed camera, placed outside of the vacuum chamber. The camera was inclined at an angle of 6.7°, and was focused on the centre of the target (focal depth 0.5 cm). The spatial resolution of the images was 20.7 \pm 0.3 \mu m pixel\(^{-1}\). Simultaneously with launching, exposures of 0.05 ms were taken at a rate of 20 000 frames per second. In this way we fully monitored the particles’ approach to and collisions on the target, allowing to measure their size, shape, and velocity, and to monitor the collisions with the target surface.

As described in Paper I, the average pressure exerted on an aggregate through acceleration during launch scales as \(P \propto v^2\), and is approximately 1 kPa for \(v \sim 1\) m s\(^{-1}\). Considering the tensile strength of the different samples, particles can be expected to fragment upon launch when accelerated to a velocity above a fragmentation barrier of \((7.7, 2.4, 1.5)\) m s\(^{-1}\) for monomer sizes \(d_0 = (0.3, 1.0, 1.5)\) \mu m. This expected breakup on launch as a function of monomer strength was indeed observed empirically (see Section 3), corroborating our earlier estimates of tensile strength.

After launch and on-target collisions, the vacuum chamber was slowly pressurized. The target, containing deposits sticking to it, was dismounted with plastic tweezers only touching its sides, and photographed with a photo camera before it was stored for further analysis. This picture, taken immediately after the experiment, was used in later stages of the post-impact analysis as a reference to exclude any fragments added to or removed from the target plate during transfer, storage, and further imaging (see Section 2.3).

### 2.3 Data analysis

In this sub-section, we describe the methods used to analyse the experimental data. The analysis was split into two parts: analysis of camera images (referred to as pre-impact) and analysis of the resulting deposits on the targets (post-impact).

Our method differs somewhat from Paper I, where we matched individual pre-impact and post-impact particles one-on-one. That method only allows studying a limited number of individual collisions. In the current study, we compare statistics of all particles in a single experiment pre-impact and deposits post-impact. This allows a quantitative measure of the mass transfer function, and bulk properties of the deposits. This situation also likely resembles the dust collection on the COSIMA plates, where a collection of fragments on target are often seen to be caused by a single pre-impact ‘parent’ particle, which may already have been fragmented on the funnel wall shortly before hitting the target (Merouane et al. 2017).

We comment on the different post-impact deposit morphologies in Section 3.

#### 2.3.1 Pre-impact analysis

Measuring pre-impact particle size distribution and mass was done by a careful examination of individual movie stills centred on the same location, at different times. A montage was made of 5–10 stills preceding the impact on target (see Fig. 3 for an example). The time difference between selected stills ranges from \(\sim 0.5\) ms (for high-velocity experiments) to \(\sim 4\) ms (for low-velocity experiments). The montage was constructed so that it displays all particles above the detection limit \(d_{\text{eq, pre}} \gtrsim 24\) \mu m that eventually hit the target. Red ellipses are drawn around these particles (Fig. 3). Unmarked particles either miss the target eventually, or are downward-moving rebounds off the target.

As the particles cross multiple pixels during a single exposure, their silhouettes are ‘smeared out’ in the vertical direction on the movie image. The size of each ellipse encompassing a particle was corrected for this blurring effect by reducing the vertical ellipse axis by the particle crossing distance (ranging from 2 to 14 pixels in the range \(v = 0.5–6.0\) m s\(^{-1}\)). Subsequently, the particle volume was approximated as

\[
V_{\text{pre}} = \frac{\pi}{6} d_{\text{eq, pre}}^3 ,
\]

where \(d_{\text{eq, pre}}\) is the average diameter of the deconvolved ellipse. By inspection of multiple frames, we measured velocities and excluded particles that were moving downwards after rebounding.

This method results in the measurements, for every experiment, of the volumes of particles with \(d_{\text{eq, pre}} \gtrsim 24\) \mu m. We adopt a uniform error of 20 per cent in the measure of particle volume, mainly driven by smearing, defocus, and irregular shape. Also, in some cases the field was crowded, resulting in the wrongful inclusion of some particles that in fact missed the target, passing behind it. This results in the bulk volume being an overestimate of the actual volume hitting the target.

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Table 1. Characteristics of the three analogue particles used in this study. Here, \(d_0, \phi, \) and \(\sigma\) are the diameter of the monomer grains, the packing density of the aggregates, and the tensile strength of the aggregates, respectively. The latter are derived from Gundlach et al. (2018).

| \(d_0 (\mu m)\) | \(\phi\) | \(\rho_0 (10^3\) kg m\(^{-3}\)) | \(\sigma\) (kPa) |
|----------------|---------|----------------|---------|
| 1.5            | 0.35 ± 0.03 | 0.70 ± 0.06   | 2.5 ± 0.4 |
| 1.0            | 0.42 ± 0.03 | 0.84 ± 0.06   | 4.6 ± 1.0 |
| 0.3            | 0.32 ± 0.02 | 0.64 ± 0.04   | 11.6 ± 1.5 |

Figure 2. Schematic drawing of the experimental setup (side view). Taken from Paper I.
Coverage of the target was obtained by taking the per-pixel maximum of the COSISCOPE M and P frames. This image was manually cleaned of artefacts and the area containing the screws fixing the target to the target holder were masked. The resulting image was thresholded at 1σ above the average illumination level of an empty section of the target on the combined M and P frames. The area of individual fragments was calculated; every interconnected particle is considered as one fragment. From this, the size dimension $d_{eq,post}$ for fragments was calculated as being the equivalent diameter of a circle with the same area as the fragment.

The length of the shadow on the separate P and M images provide a measure of the particle height, $h_{post}$. This value is used in calculating the post-impact deposit volume of a single fragment as

$$V_{post} = \frac{\pi}{4} \epsilon d_{eq,post}^2 h_{post}. \quad (2)$$

The parameter $\epsilon$ parametrizes the deposit geometry (see Hornung et al. 2016 and Paper I). We adopt $\epsilon = 0.33$ (a pyramid shape) for all deposits.

We calculated height for selected fragments (around 25 per cent of the total number) across the deposit size and morphology range, to arrive at an estimate for the typical height-to-base ratio as a function of velocity, monomer size, and deposit size. We assume no compaction of the material has taken place during the collisions with the target; see the discussion in Paper I (Section 4.3). As the bulk density of monodisperse material is comparable to or higher than the polydisperse material used in Paper I, and the bulk density is also similar, we assume compaction can also be neglected in the current study. Combining the total volume of material pre- and post-impact, we calculate the mass transfer function $TF$, being the fraction of the total pre-impact dust mass deposited on to the target in one experiment, as

$$TF = \frac{\sum_{i} V_{post,i}}{\sum_{i} V_{pre,i}}, \quad (3)$$

where $V_{pre,i}$ and $V_{post,i}$ are the values of individual particles and fragments, respectively.

3 RESULTS

We present the results of nine experiments, one for every combination of monomer size $d_0 (0.3, 1.0, 1.5) \, \mu m$ and $v$ (low: $\sim 1 \, m \, s^{-1}$, medium: $\sim 2.5 \, m \, s^{-1}$, high: $\sim 6 \, m \, s^{-1}$). See Table 2 for a summary of the analysis.

Upon analysis of the pre-impact movies, the breakup of particles upon launch caused the material with low tensile strength (large monomer size) to contain relatively more small fragments compared to the high-strength material (small monomer size). This is schematically displayed in Fig. 4. Note that in all experiments, over 90 per cent of the pre-impact mass is contained in particles larger than 80 $\mu m$ (see below).

Post-impact images of the nine targets are displayed in Fig. 5 (top panel). The bottom panel of this figure displays thresholded images of these targets, which highlight the coverage with dust fragments (in black). These were used to calculate the area (and hence $d_{eq}$) of individual fragments. The height of $\sim 25$ per cent of the on-target fragments was measured, and averaged over fragments smaller and larger than 80 $\mu m$. These values were used accordingly to calculate the total post-impact volume. The values lie in the range $h/d_{eq,post}$ $\sim 0.1–0.2$ and decrease with size, tensile strength, and velocity, similar to what was measured for the deposits in Paper I.
Table 2. Summary of experiments and results. Columns, from left to right: experiment label; monomer diameter; velocity range of particles; number of particles identified on pre-impact movie stills, above threshold (see the text); power-law index of particle size distribution; number of post-impact fragments on target above threshold (see the text); power-law index of fragment size distribution; mass transfer function.

| Experiment | $d_0$ (μm) | Pre-impact Velocity range (m s$^{-1}$) | # Particles ($d_{eq>24}$ μm) | $\alpha_{pre}$ | Post-impact # Fragments ($d_{eq>16}$ μm) | $\alpha_{post}$ | TF (percent) |
|------------|------------|--------------------------------------|----------------------------|---------------|----------------------------------------|---------------|--------------|
| A1         | 1.5        | 0.4–1.2                              | 225                        | $-1.2 \pm 0.2$ | 159                                    | $-1.9 \pm 0.2$ | 1.4 ± 0.3    |
| A2         | 1.5        | 2.7–3.3                              | 330                        | $-1.3 \pm 0.2$ | 410                                    | $-1.6 \pm 0.2$ | 4.8 ± 0.9    |
| A3         | 1.5        | 5.7–6.5                              | 298                        | $-0.4 \pm 0.1$ | 698                                    | $-1.6 \pm 0.2$ | 7.4 ± 1.5    |
| B1         | 1.0        | 0.5–1.1                              | 95                         | $-0.7 \pm 0.2$ | 395                                    | $-2.6 \pm 0.2$ | 0.35 ± 0.07  |
| B2         | 1.0        | 2.1–2.9                              | 82                         | $-0.9 \pm 0.2$ | 299                                    | $-1.8 \pm 0.2$ | 0.83 ± 0.17  |
| B3         | 1.0        | 5.5–6.3                              | 221                        | $-2.0 \pm 0.2$ | 469                                    | $-1.6 \pm 0.2$ | 2.0 ± 0.4    |
| C1         | 0.3        | 0.5–1.1                              | 154                        | $-2.3 \pm 0.2$ | 161                                    | $-2.9 \pm 0.2$ | 0.02 ± 0.02  |
| C2         | 0.3        | 2.6–3.2                              | 221                        | $-1.8 \pm 0.2$ | 209                                    | $-2.3 \pm 0.2$ | 0.21 ± 0.04  |
| C3         | 0.3        | 5.4–5.6                              | 261                        | $-3.4 \pm 0.3$ | 713                                    | $-2.0 \pm 0.2$ | 0.25 ± 0.05  |

The pre- and post-impact analyses are summarized in Figs 6 and 7. Fig. 6 displays pre- and post-impact fragment size distributions (normalized to 1). The top two graphs in Fig. 7 summarize the trends observed in the mass distributions. The histogram on the top panel shows the fraction of the pre- and post-impact mass contained in large (>80 μm) particles (pre-impact) and fragments (post-impact), respectively. In all experiments, more than 90 percent of the measured pre-impact mass is contained in particles larger than 80 μm (4 pixels across on the movie image). In the post-impact size distribution, however, a clear trend is visible. Material with lower tensile strength leaves relatively more large fragments on the target. Velocity has a similar effect: at higher velocities, larger deposits are made.

This effect is also reflected by the graph in the middle panel, which shows the power-law index $\alpha_{post}$ of the post-impact size distribution, defined as

$$N \propto d_{eq}^{\alpha_{post}},$$

measured for particles/fragments between $30 > d_{eq} > 300$ μm, binned with logarithmic intervals of 0.15, similar to Merouane et al. (2016, 2017, and references therein). In general, a steep power law indicates the presence of more small fragments. Fig. 7 (middle panel) shows that in our experiments, low velocity and high tensile strength lead to relatively more smaller fragments on the target. This reiterates the trend that was observed from the histogram in the top panel of this figure.

The bottom graph shows that the mass transfer function TF increases with increasing impact velocity and increasing monomer size, hence decreases with increasing material strength. TF does not exceed 10 percent in any of the experiments, and is less than 1 percent in some of the low-velocity and high-strength experiments. We observe that higher values of $\alpha_{post}$ are observed in experiments with higher values of TF, which emphasizes that the mass is dominated by large deposit fragments.

The trends observed in TF and $\alpha_{post}$ imply that when attempting to retrieve pre-impact characteristics from post-impact mass distributions, a degeneracy exists between velocity and tensile strength. We will discuss this further in Section 4.1.

4 DISCUSSION

4.1 Factors that influence mass transfer

The main finding of this series of experiments is that, within the parameter space considered, mass transfer efficiency increases...
Figure 5. **Top:** Targets after experiments, illuminated from right and left (co-added, see the text). **Bottom:** Same as above, with thresholded images highlighting target coverage by dust in black.
with impact velocity and decreases with tensile strength. This is consistent with earlier work by Gütterl et al. (2010); see also the discussion in Paper I (Section 4.1). At the values considered in this study, even at the softest material (∼3 kPa) at highest velocities (∼6 m s⁻¹), TF does not exceed 10 per cent. Furthermore, our experiments show that multiple parameters influence the appearance of deposits in a similar way. In this sub-section we discuss these factors.

The influence of impact velocity and tensile strength on mass transfer efficiency is similar. This degeneracy is not lifted by considering other observables, like the relative fragment size distribution, as this has a similar scaling relation with both parameters. A closer examination of the individual deposits, however, may prove fruitful in determining its tensile strength or impact velocity, for example by relating this to the height-to-base ratio of individual fragments. Such an in-depth analysis is beyond the scope of this paper.

In the imaging data acquired by the Rosetta mission, just the post-impact situation is known to us. If one were to retrieve the strength and velocity of the particle(s) prior to their impact on the target, one is met with various complications. First of all, the degeneracy described above makes it difficult to isolate the effect of impact velocity. Another complication that inhibits the retrieval of the properties of a parent particle is the unknown pre-impact size distribution. It is expected that some of the particles that entered the instrumentation were broken up upon collision with the instrument funnel (Merouane et al. 2016). The effects of these funnel

Figure 6. Cumulative size distribution for pre-impact (left-hand column, particles) and post-impact (right-hand column, fragments) sizes, normalized to the total cross-sectional area. Tensile strength decreases per row from top to bottom, as monomer size increases; capital letters A, B, and C correspond to experiments listed in Table 2. Note that the absolute value of \( V_{\text{post,total}} \) (right-hand column) is up to a factor 10–100 lower than \( V_{\text{pre,total}} \) (left-hand column; also see middle right panel).
collisions may be studied by simulating them in a future series of experiments.

4.2 Comparison to cometary dust

Fig. 8 displays the relation between mass transfer and tensile strength found in our experiments. It can be seen that mass transfer steeply increases towards strengths below 4 kPa. While measurements at lower strengths are not available, it is likely that of the dust that initially hit the COSIMA targets during the Rosetta mission, more than half has not ended up there following a bouncing and/or fragmenting collision.

Some considerations should be taken into account when extrapolating our experimental results to cometary dust. First, the monodisperse material used in our experiments is not expected to occur in nature. However, it has a clear dependence of strength on monomer size, and allows us to study this parameter. Langevin et al. (2017) show that the albedo of cometary dust is much lower than our sample material. As the contrast with the background of our material is higher than it would have been with darker material, this may have resulted in a systematic overestimate of post-impact fragment sizes.

Furthermore, we observe that the SiO₂ material breaks up into tiny fragments with sizes below the COSISCOPE resolution limit, brightening up neighbouring pixels in the images. For cometary dust, in situ ‘experiments’ of the breakup of cometary dust by impact or electrostatic forces due to charging show that the dust particles do breakup into elements in the 10⁻⁵⁻⁻⁵₀ μm size range with dark areas in between, without a hint for blurring or an increase in reflection due to partial filling of the instrument pixel with reflections of tiny particles below the resolution limit (Langevin et al. 2016; Hilchenbach et al. 2017).

These differences should be considered when directly comparing experimental results with COSIMA data. For a further detailed discussion on the similarity to cometary material, we refer to the discussion in Paper I (Section 4.2).

Analysis of Rosetta data has suggested the presence of either two different types of dust populations: a more compact component and a fluffy component, or a more uniform composition with one component. In the two different type of hypothesis, the compact component constitutes with mass density \( \rho_b \sim 10^3 \text{ kg m}^{-3} \) (Rotundi et al. 2015; Fulle et al. 2016b, 2017), and a filling factor \( \phi > 0.1 \). A second, low-density component (\( \rho_b \sim 1 \text{ kg m}^{-3} \)) is detected by GIADA (Fulle et al. 2015, 2016a). This component is hypothesized to consist of the fractal dust particles detected by MIDAS (Bentley et al. 2016; Mann et al. 2016), but a limited amount of data is available to confirm this hypothesis. The comet formation model proposed by Blum et al. (2017) also includes the fluffy component.
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Figure 9. Two different COSIMA targets: 2D1 (left) was exposed for 7 weeks before perihelion and 1CD (right) was exposed for 19 weeks including perihelion.

Table 3. Summary of properties of COSIMA targets displayed in Fig. 9.

| Target | Exposure dates | Spacecraft altitude (km) | Heliocentric distance (km) | Collision event date | Percent of on-target volume collected during event | Spacecraft altitude (km) | Heliocentric distance (au) |
|--------|----------------|--------------------------|---------------------------|---------------------|-----------------------------------------------|--------------------------|--------------------------|
| 2D1    | 10/04/2015–27/05/2015 | 91–321                   | 1.89                      | 11–12/05            | 73                                            | 146–162                  | 1.66 (in)                |
|        | (inbound)       |                          |                           |                     |                                               |                          |                          |
| 1CD    | 30/05/2015–07/10/2015 | 153–1502                 | 1.53                      | 16–17/05            | 24                                            | 126–133                  | 1.62 (in)                |
|        | (inbound)       |                          |                           |                     |                                               |                          |                          |

to be present in between the more compact particles. Fulle & Blum (2017) conclude that the presence of this fluffy component suggests that no collisions at velocities higher than 1 m s$^{-1}$ have taken place during formation of the comet.

The COSIMA dust collection observations can be modelled with one dust particle population with a tensile strength due to van der Waals forces of $\sigma \sim 1$ kPa and mass density $\rho_b \sim 1\sim 10^2$ kg m$^{-3}$ for particles in the 60–300 μm size range (Hornung et al. 2016), a filling factor down to $\phi \sim 0.1$, based on optical observations (Langevin et al. 2017) and mechanical parameter analysis (Hornung et al. 2016). The rationale for the later dust particle model is that particles classified prior as ‘compact’ after impact can be fragmented by Lorentz forces due to in situ charging in the instrument and therefore do not constitute a separate dust particle class (Langevin et al. 2016; Merouane et al. 2016; Hilchenbach et al. 2017).

It is intriguing to find out whether an application of our results to COSIMA can shed light on the strength of cometary dust, or the existence of the different dust species.

4.3 Application to Rosetta/COSIMA measurements

As an illustration of the application of these experiments, we applied our method to Rosetta data. We compare two COSIMA targets that have captured multiple particles. Fig. 9 shows stray-light corrected images (P illumination) of the two targets. Some properties of the targets and their collection periods are listed in Table 3. The targets have been exposed for 7 weeks (2D1) and 19 weeks (1CD). During the second collection period, at least two major outbursts (2015 May 6 and August 15) and several smaller events have taken place (Feldman et al. 2016; Knollenberg et al. 2016; Vincent et al. 2016; Pajola et al. 2017). However, the bulk of the fragments on these targets were collected during single collision events outside the time window of the major outburst events, as was identified by Merouane et al. (2017, table A1). From this analysis, it can be inferred that these collision events took place within $\sim$24 h time windows 3 months (2D1: 2015 May 11–12 and 16–17) and 2 weeks (1CD, 2015 July 31–August 1) before perihelion. The spacecraft speed relative to the comet was less than 0.1 m s$^{-1}$ at all times during these collection events. The major differences in the circumstances were that during the second collection period the spacecraft was both 1.5 times further away from the comet surface, and 1.3 times closer to the sun.

Merouane et al. (2017) find that the power-law index of the fragment size distribution increases during outburst events, due to either higher velocities or material with lower tensile strength. While there is no evidence that the collision events have coincided with outburst events, a possible effect contributing to the abundance of small fragments on target 1CD is that in the period around an outburst event, higher pressure may lead to breakup into smaller aggregates. These can be lifted more easily because of the inverse tensile strength relation of aggregate packings (see Skorov & Blum 2012). When we apply the same method as described in Section 3, the power-law index for 2D1 is $\alpha \sim -1.9$, and for 1CD $\alpha \sim -3.6$; see Fig. 10. In other words, target 1CD contains a relative high amount of small fragments. If the tensile strength of both dust particles is equal, this implies a higher impact velocity for period 1CD. Conversely, if velocities...
were similar, the 1CD particle may have been of a stronger dust species.

A more likely scenario is that in both cases, a particle collided with the funnel wall and broke up into smaller parts. For 1CD, however, the particle may have hit at higher velocity, causing breakup into more and smaller fragments (consistent with our pre-impact analysis). These fragments subsequently impacted at lower velocities. Our result thus shows that the dust on the two COSIMA targets presented are consistent with originating from particles with a single strength. The significant difference in the distribution of particles on their surface may thus be explained by different impact velocities.

5 SUMMARY AND CONCLUSIONS

Aiming to simulate the circumstances of dust collection in the COSIMA and MIDAS instruments on the Rosetta spacecraft, we performed nine experiments, in which we collided SiO$_2$ aggregates in the size range 30–400 μm with a target surface at impact velocities of 0.5–6.5 m s$^{-1}$. Three different monodisperse monomer sizes were used, resulting in tensile strengths in the range 3–12 kPa. We used COSIMA targets, and also COSISCOPE imaging to obtain a like-for-like comparison with COSIMA data. Our main conclusions are as follows:

(i) The transfer function increases with increasing velocity and increasing monomer size (hence, decreases with increasing tensile strength).

(ii) No more than 10 per cent of the impacting material is transferred during collisions on the target surface. This fraction drops to below 1 per cent for low velocity and high tensile strength.

(iii) Material with lower tensile strength leaves relatively more deposits exceeding 80 μm in size.

(iv) Extrapolating these results to the strength parameters of cometary dust imply that more than half of the dust that enters COSIMA and MIDAS was lost in the instrumentation.

While differences in material properties, unknowns pre-impact and degeneracies between parameters inhibit a direct comparison to COSIMA data, the qualitative results of the experiments may be used to correlate trends found in the fragment distribution patterns. In future space missions to comets (e.g. CAESAR sample return mission to 67P; Squyres et al. 2018), we suggest to use a different technique (e.g. combination of laser curtain, enhanced imaging capabilities, multiple resolution modes for scanning) that avoids the uncertainty introduced by the impact in the instruments.

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Figure A1. Examples of individual deposit morphologies. The images displayed are height maps obtained with a 3D laser scanning confocal microscope. The height scale is normalized to the maximum height of a deposit; numbers within boxes correspond to the maximum height level (in μm) of the deposit. Frame colours indicate the type of collision that led to the deposit. All scale bars indicate 100 μm.