Provenance of upper Permian-lowermost Triassic sandstones, Wutonggou low-order cycle, Bogda Mountains, NW China: implications on the unroofing history of the Eastern North Tianshan Suture

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Abstract

This study investigates the provenance of sedimentary rocks in Bogda Mountains, NW China, and reconstructs the lithology and unroofing history of the Eastern North Tianshan Suture. Petrographic point counting data of sandstones and compositions of conglomerates of upper Permian-lowermost Triassic Wutonggou low-order cycle from Zhaobishan, North Tarlong, Taodonggou, and Dalongkou sections in the southern and northern foothills of Bogda Mountains were used to interpret the temporal and spatial variations of lithology of the Eastern North Tianshan Suture, which is the sediment source area. Three compositional trends were identified. A trend of upward-increasing quartz content and granitic pebbles in Zhaobishan section suggests a change from the undissected volcanic arc, accretionary wedge and trench setting to predominantly transitional volcanic arc and subordinate accretionary wedge and trench, in the eastern part of the Eastern North Tianshan Suture. In North Tarlong and Taodonggou sections, however, the lithic content decreases and the contents of quartz and granitic pebbles increase up sections. These trends indicate that the western part of the Eastern North Tianshan Suture changed from an undissected volcanic arc to the transitional volcanic arc, accretionary wedge and trench. No clear trend in the lithic-rich sandstones of the Dalongkou section indicates that sediments were derived from the undissected volcanic arc in the Eastern North Tianshan Suture and local rift shoulders. Compositional variations of studied rocks suggest that the Eastern North Tianshan Suture was an amalgamated complex with great spatial and temporal heterogeneities in lithology and experienced persistent unroofing during late Permian-earliest Triassic. This study reconstructs a key element of the Chinese Tianshan Suture and serves as an example to understand the unroofing processes of ancient sutures.

Keywords: Provenance, Wutonggou low-order cycle, Bogda Mountains, Unroofing, Eastern North Tianshan Suture, NW China
1 Introduction
The Central Asian Orogenic Belt is one of the accretionary orogens that were the main sites of continental growth since the Phanerozoic, and resulted from the accretion and collision of magmatic arcs, accretionary complexes, trapped oceanic plates, and trailing continental plates (Şengör et al. 1993; Şengör and Natal’ in 1996; Windley et al. 2007; Xiao et al. 2013). The Eastern North Tianshan Suture (ENTS) is located in the southern part of the Central Asian Orogenic Belt (Fig. 1). It marks the closure of the North Tianshan Ocean, which was the major southern segment of the Paleo-Asian Ocean (e.g., Xiao et al., 2009, Xiao et al. 2013). Detailed studies of the ENTS can provide critical information on the final assembly of the southern parts of Central Tianshan Orogenic Belt (e.g., Charvet et al. 2007; Han et al. 2010). Most studies of the ENTS focus on its tectonic evolution based on regional tectonics, structures, and geochemical and geochronological data of the rocks exposed in the ENTS (e.g., Allen et al. 1993; Gao et al., 1998; Xiao et al. 2004, 2013; Wang et al. 2007; Han et al. 2010; Charvet et al. 2011). However, the eroded rocks in the ENTS during late Permian-earliest Triassic are not well understood.

Provenance studies of sandstones are useful to reconstruct the eroded parts of adjacent orogens (e.g., Dickinson and Suczek 1979; Ingersoll and Suczek 1979; Dickinson 1985; Dorsey 1988; Garzanti et al. 1996, 2007; Trop and Ridgway 1997; Ingersoll 2012; Chaudhuri et al. 2018). However, detailed reconstruction of rocks in the ENTS during the late Permian-earliest Triassic cannot be achieved by previous petrographic studies due to out-of-date chronostatigraphy or limited samples (Carroll et al. 1995; Shao et al. 2001; Greene et al. 2005; Guan 2011). This study focuses on the sandstones of upper Permian-lowermost Triassic Wutonggou low-order cycle (WTG-LC) exposed in the northern and southern foothills of Bogda Mountains, the greater Turpan-Junggar intracontinental rift basin (Yang et al. 2010). The main goals are to document the high-resolution temporal and spatial variations of sandstone compositions, to interpret the provenance of these sandstones, and to reconstruct the unroofing history of the ENTS.

2 Geological background
This study focuses on the upper Permian-lowermost Triassic fluvial-lacustrine sandstones of WTG-LC in Zhaobishan, North Tarlong, and Taodonggou sections in the southern and Dalongkou section in the northern foothills of Bogda Mountains, NW China (Figs. 2, 3). The Bogda Mountains is an E-W striking giant anticline with exposures of Devonian to Quaternary sedimentary and igneous rocks. This anticline is located between the Junggar Basin to the north and Turpan-Hami Basin to the south (Fig. 2). It was a part of the greater Turpan-Junggar basin during the late Carboniferous-Jurassic.

Fig. 1 Tectonic map of the Central Asian Orogenic Belt, which is bordered by the East European Craton to the west, Siberia Craton to the east, and North China, Tarim, and Karakum cratons to the south. Modified from Şengör et al. (1993) and Jahn et al. (2000). The Yellow box is the location of Fig. 2.
Fig. 2 Tectonic map of eastern Xinjiang, showing the locations of the Eastern North Tianshan Suture, Central Tianshan Suture, Bogda Mountains, and measured sections. Modified from Xia et al. (2004)

Fig. 3 Geological maps of Zhaobishan (a), North Tarlong-Taodonggou (b), and Dalongkou (c) areas showing names and location (red lines) of measured sections. Modified from Yang et al. (2010), Obrist-Farner and Yang (2015), and Fredericks (2017)
Indian and Asian plates (e.g., Windley et al. 1990; Hen-
the collision between the Central Tianshan
in the western United States (Yang et al. 2010).
former delta plain mudrock, sandstone, conglomerate, and
sedimentary rocks (Allen et al. 1993; Xiao et al. 2004).
Fragmental ophiolite, radiolarian chert, turbidites, and
high-pressure schists mark the subduction of the oceanic
crust between the Junggar Plate and Central Tianshan
Suture (Carroll et al. 1990; Gao et al., 1998; Shu et al.
1999; Xiao et al. 2004; Charvet et al. 2007).

The WTG-LC is an informal cyclostratigraphic unit
defined by Yang et al. (2007, see also Yang et al., 2010)
and approximately correlates with the Wutonggou and
Guodikeng formations (Fig. 4; XBGMR 1993; Yang et al.
2007, 2010). A low-order cycle formed during a period
of long-term stable tectonic and/or climatic conditions
and contains high and intermediate-order cycles. The
high-order cycle is the smallest unit that records the
environmental changes caused by the transgression and
regression of lakes or erosion and deposition of rivers. The
intermediate-order cycle includes several high-order
cycles, representing longer trends of transgression and
regression or erosion and deposition than the high-order
cycles. The WTG-LC records an overall persistently
uplifting history of the source areas and humid to sub-Humid
climate conditions (Yang et al. 2007, 2010; Thomas et al.
2011).

Stratigraphic correlations have been largely based on
lithostratigraphy, biostratigraphy, and cyclostratigraphy
(Zhang 1981; Liao et al. 1987; Wartes et al. 2002; Yang et al.
2007, 2010). The chronostratigraphy in the greater
Turpan-Junggar basin is not well constrained. Yang et al.
(2010) placed the Permo-Triassic boundary in a 90 m-

thick interval in North Tarlong section. Based on strati-
graphic correlation and petrographic studies, the strata
in Tarlong and Taodonggou areas were interpreted as
being deposited within one half-graben, termed the
Tarlong-Taodonggou half-graben (Yang et al. 2010; see
also Guan 2011; Peng 2016; Obrist-Farner and Yang
2017; Fredericks 2017). The basin geometry of the
Zhaobishan and Dalongkou areas is not clear due to a
limited number of measured sections and is speculated
to be similar to that of Tarlong-Taodonggou half-
graben.

Sandstones of WTG-LC can be divided into fluvial, del-
taic, and littoral-lakeplain facies (Yang et al. 2010; Fig. 5).
The fluvial facies includes meandering stream and braided
stream deposits. The former is characterized by a high-
relief erosional base with upward-fining succession of
channel-fill conglomerate, point bar sandstone, and over-
bank siltstone and shale; and the latter by a low-relief ero-
sional base with upward-fining association of channel-fill
conglomerate, bar sandstone and absence of overbank de-
posit. The deltaic facies includes the upward-coarsening and
thickening successions of prodeltaic shale and silt-
stone, delta front sandstone and/or conglomerate in the
lower part, and distributary channel-fill and interdistribu-
tary delta plain mudrock, sandstone, conglomerate, and

(Shao et al. 1999, 2001; Greene et al. 2001, 2005; Yang
et al. 2007, 2010). The greater Turpan-Junggar basin is
speculated to be a back-arc basin (Hsu 1988), a transi-
tional basin from rift to foreland basin (Carroll et al.
1990; Hendrix et al. 1992; Shao et al. 1999, 2001; Greene
et al. 2001, 2005), or a rift basin (Allen et al. 1991; Allen
et al. 1993; Shu et al. 2005, 2011; Yang et al. 2010, 2013).
The seismic profiles (Yang et al. 2010), mixed tholeiitic
volcanism and marine sedimentation in the uppermost
basement (Yang et al. 2013), the bimodal volcanic rocks
(Shu et al. 2005, 2011), continental rift-type geochemical
signatures (Allen et al. 1991; Shu et al. 2005, 2011), and
regional scale strike-slip shear zones (Laurent-Charvet
et al. 2002, 2003; Shu et al. 2005, 2011) support that the
greater Turpan-Junggar basin was an intracontinental
rift basin over a Carboniferous volcanic arc or back-arc
basement (Shu et al. 2005; Yang et al. 2010; Yang et al.
2013). Regional dextral strike-slip movement triggered
the rifting that starting in the latest Carboniferous
(Shao et al. 1999, 2001; Greene et al. 2001, 2005; Yang
et al. 2010, 2013). The WTG-LC consists of Ordovician to Devonian-Carboniferous
rock units (Allen et al. 1993; Xiao et al. 2004; Charvet et al.
2007).

The Chinese Tianshan separates the Junggar Basin to
the north from the Tarim Basin to the south and has
been created since the Cenozoic collision between the
Indian and Asian plates (e.g., Windley et al. 1990; Hen-
drix et al. 1994; Yin et al. 1998). Before the Cenozoic
collision, a series of suture zones were formed in the
Chinese Tianshan area during the Paleozoic, one of
which is the North Tianshan Suture. The North Tianshan Suture is further divided into the western and eastern
segments in terms of their relative locations to the
city of Urumqi (Fig. 2). This study focuses on the eastern
segment of the North Tianshan Suture, which is situated
about 100 km south of Bogda Mountains and north of the
Central Tianshan Suture. The origin of the ENTS is
not fully understood, but the southward subduction of
North Tianshan Ocean beneath the Central Tianshan
Suture, and the collision between the Central Tianshan
Suture and the trailing Junggar Plate are widely accepted
(Windley et al. 1990; Gao et al., 1998; Xiao et al. 2004,
2013; Charvet et al. 2011). The timing of the collision is
in debate, varying from Middle Ordovician (Gao et al.,
1998), Devonian-early Carboniferous (Xiao et al. 2004,
2013) to Late Devonian-Carboniferous (Charvet et al.
2011). Similarly, the proposed closure time varies from
the end of Early Carboniferous (Gao et al., 1998), Late
Carboniferous (Windley et al. 1990; Xiao et al. 2004,
2013; Han et al. 2010), to Late Carboniferous-Early Per-
mian (Allen et al. 1993; Carroll et al. 1995). The current
ENTS consists of Ordovician to Devonian-Carboniferous
volcanic-arc rocks and associated submarine volcanic-
paleosol in the upper part. The littoral-lakeplain facies is characterized by the upward-coarsening succession of sublittoral shale and littoral well-washed sandstone and conglomerate in the lower part, and muddy to sandy paleosol in the upper part.

3 Data and methodology
Sixty sandstones from four sections were studied to observe their compositional and textural characteristics (Fig. 6; see Yang et al. 2007, 2010 for detailed measured sections). Three hundred framework grains in each thin section were counted using both Suttner’s (1974) and Gazzi-Dickinson’s methods (Gazzi 1966; Dickinson 1970). The Gazzi-Dickinson method counts sand-size (0.063 mm) mineral crystals within large rock fragments as individual grains, whereas the Suttner’s method does not count the mineral crystals, only as the rock fragments (Ingersoll et al. 1984). As most of the counted lithic grains in the studied sandstones contain minerals smaller than sand, the results of these two methods are similar. Definitions of raw and recalculated parameters of point counting categories are tabulated in Table 1. The interpretations of volcanic lithic fragments follow the descriptions of Dickinson (1970) and Marsaglia and Ingersoll (1992). The interpretation of polycrystalline quartz grains follows those of Basu et al. (1975), Young (1976), and Blatt et al. (2006). Point counting data in recalculated parameters are presented in Table 2 in terms of Gazzi-Dickinson method so that petrofacies can be defined by composition and compared with the tectonic fields in templates from previous studies (Dickinson and Suczek 1979; Dickinson et al. 1983; Dickinson 1985; Marsaglia and Ingersoll 1992; Critelli and Ingersoll 1995).

One thousand nine hundred fourteen gravels in 17 conglomeratic beds were counted and described in the field to obtain the spatial and temporal trends of clast composition. Fresh surfaces were used to identify lithologies. A rectangular grid on the outcrop surface was laid out as a guide for counting. About 100 clasts were counted in each location. In addition, 168 attitudes of nine tabular cross beds of fluvial sandstones were measured in the field. They were later corrected using the method of Davis et al. (2011), pp. 710–714). The correction, rose diagrams, and mean vectors were performed using the software StereoNet of Allmendinger (2005).

4 Results
Sandstone compositions are used to classify petrofacies, from which the lithology and tectonic settings of the source areas can be interpreted. Clast composition of
conglomerates and paleocurrent directions supplement the classification of petrofacies and provenance interpretation. The stratigraphic distribution of petrofacies, clast composition, and paleocurrent directions in individual sections show the temporal changes of sandstone composition. Finally, the spatial variations of sandstone composition are interpreted on the basis of correlations among sections in the study areas.
4.1 Framework grains

The major framework grains in sandstones of WTG-LC include quartz, feldspar, and lithic fragment. They are further differentiated on the basis of optical mineralogic characteristics, such as extinction, twinning, and relict textures (Tables 1, 2). Accessory minerals, including micas and heavy and opaque minerals, are scarce.

4.1.1 Quartz

Quartz occurs as monocrystalline crystals (Qm) and polycrystalline (Qp) and microcrystalline aggregates (Cht). Qm grains are clear, inclusion-free, and subangular, and are subdivided into nonundulatory (Qnu; Fig. 7a) and undulatory types (Qu). Qnu grains exhibit straight extinctions, whereas Qu grains are strained with undulose extinction at an angle between 5 and 10 degrees. Qp grains are subdivided into polycrystalline ones with metamorphic deformed texture (Qpt; Fig. 7b) and polycrystalline ones without such texture (Qpw). Qpt grains contain more than five sutured, elongate quartz crystals. In contrast, Qpw grains contain two to five monocrystalline quartz grains without sutured contacts. Cht grains are aggregates of microcrystalline quartz and interpreted as fragments of chert (Fig. 7c).

4.1.2 Feldspars

This group includes plagioclase (P; Fig. 7a) and potassium feldspar (K; Fig. 7a). Plagioclase grains usually exhibit polysynthetic twinning. The plagioclase grains in North Tarlong and Taodonggou sections are common with albite twinning. They occur as discrete angular-subangular grains and lath or mosaic phenocrysts in volcanic lithic fragments. Potassium feldspar includes microcline and orthoclase. The microclines exhibit twinning and only occur in Zhaobishan and Taodonggou sections. The orthoclase is usually Carlsbad twinned or untwinned. Clear, inclusion-free orthoclase is common in the studied thin sections.

4.1.3 Lithic fragments

Lithic fragments include volcanic (Lv), sedimentary (Ls), and metamorphic (Lm). Lv and Ls grains dominate and
account for over 99% of the total lithics. Lm grains are rare.

The Lv grains are subdivided into four types based on their textures, including felsic (LvF; Fig. 7d), microlitic (LvMi; Fig. 7e, f), lathwork (LvL; Fig. 7e), and vitric (LvV; Fig. 7f). LvF, LvMi, LvL, and LvV grains are interpreted to be derived from felsic (LvF), intermediate (LvMi), and mafic (LvL) igneous rocks and volcanic glass (LvV), respectively (Dickinson 1970). The Ls grains are subdivided into three types based on their textures, including mudrock (Lmd; Fig. 7e), siltstone (Lslt), and sandstone (Lsd). The Lmd grains account for more than 95% of the total sedimentary lithics. Finally, a trace amount of Lm grains, mainly schist fragments, are identified on the basis of their foliations.

4.1.4 Accessory minerals
Accessory mineral grains are the minor framework grains in WTG-LC sandstones and account for 2% of the total detrital grains. Muscovite, biotite, zircon, tourmaline, amphibole, and opaque minerals are observed.

4.2 Matrix, cement, and sandstone classification
WTG-LC sandstones contain 1% of matrix. Based on the sandstone classification of Dott (1964) 46 litharenites, 11 feldspathic arenites, and three lithic wackes (Table 2) were recognized. The cements in the sandstones are predominantly calcite, clays, and iron oxides and sulfides. Zeolite and silica are rare. As the matrix is not further studied and cements are largely controlled by diagenesis rather than provenance lithology (Dickinson and Suczek 1979), they are not included in the classification of petrofacies.

4.3 Composition of gravels and paleocurrent measurements
Conglomerates in the WTG-LC are typically polymictic, either clast- or matrix-supported. The composition of individual conglomeratic beds is summarized in Section 5 to facilitate the interpretations of source lithology.

Gravels are igneous, sedimentary, or metamorphic in composition (Fig. 6; Table 3). Igneous gravels are volcanic and plutonic clasts, including white or gray rhyolite, dark green or dark purple andesite, and dark green, dark purple or black basalt, and reddish granite. Sedimentary gravels are mudrock and chert, including green, purple, and brown soft clasts of mudrock and massive or laminated gray chert. Metamorphic gravels include white quartzite, of which the boundaries of single quartz crystals are interlocking with each other.

The paleocurrent directions were only measured from cross-beddings of a decimeter scale within point bar sandstones, in order to remove the possible errors induced by bed-form hierarchies (Allen 1968; Miall 1974). The mean vectors of paleocurrent directions provide qualitative estimates of the locations of the surrounding highs. These results are further discussed in Section 4.5 to aid the interpretations of source locations (Fig. 6).

4.4 Petrofacies and implications on lithology and tectonic settings of source areas
Three petrofacies were recognized in sandstones of the WTG-LC on the basis of relative abundance of quartz, feldspar, and lithic fragment. The distributions of these petrofacies are shown in the QFL and QmFLt ternary

| Table 1 Raw and recalculated grain types and categories for point-counting and petrofacies classification |
| --- |
| Symbol | Definition |
| Raw |  |
| Qnu | Nonundulose monocrystalline quartz |
| Qu | Undulose monocrystalline quartz |
| Qpt | Polycrystalline quartz with metamorphic textures |
| Qpw | Polycrystalline quartz without metamorphic textures |
| Cht | Chert and chalcedony |
| Lvf | Volcanic lithic with felsic texture |
| Lvmi | Volcanic lithic with microlitic texture |
| LvL | Volcanic lithic with lathwork texture |
| Lvv | Volcanic lithic with vitric texture |
| Lvun | Unidentified volcanic lithic |
| Lmd | Mudrock fragment |
| Lslt | Siltstone fragment |
| Lsd | Sandstone fragment |
| Lm | Metamorphic lithic |
| Carb | Carbonate |
| AM | Accessory transparent minerals |
| OM | Opaque minerals |
| Bio | Bioclastic grains |
| Uni | Unidentified grain |
| Recalculated |  |
| Qm | Qnu + Qu |
| Qp | Qpt + Qpw + Cht |
| Q | Qm + Qp |
| F | K + P |
| Lv | Lvf + Lvmi + LvL + Lvv + Lvun |
| Ls | Lmd + Lslt + Lsd |
| Lmt | Lm + Qpt |
| Lst | Ls + Cht |
| L | Lv + Ls + Lm |
| Lt | Lv + Lmt + Lst |
Table 2: Recalculated point-counting data

| Sample Number | Q | F | L | Qm | F | L | Q | K | Qp | L | Ls | Lm | Lv | Lm | Lvf | Lvmi | Lvl | Matrix Classification | Petrofacies |
|---------------|---|---|---|----|---|---|---|---|----|---|----|----|----|----|----|-----|----|---------------------|-------------|
| S15–31        | 55| 32| 13| 29 | 32| 39| 47| 43| 9  | 27| 32| 41| 22| 34| 44| 90 | 10 | 0  | 2       | FA           | PF1         |
| S15–33        | 48| 13| 39| 20 | 13| 67| 61| 28| 11 | 16| 45| 39| 15| 45| 40| 69 | 24 | 6  | 0       | LA           | PF1         |
| S15–35        | 49| 32| 19| 15 | 32| 53| 34| 39| 27 | 55| 21| 23| 42| 28| 30| 77 | 19 | 3  | 0       | FA           | PF1         |
| S15–36        | 56| 24| 19| 19 | 24| 56| 44| 43| 13 | 27| 23| 50| 21| 25| 54| 65 | 26 | 9  | 3       | LA           | PF1         |
| S15–37        | 63| 17| 20| 22 | 17| 60| 56| 18| 26 | 16| 25| 58| 12| 27| 62| 71 | 29 | 0  | 3       | LA           | PF1         |
| S15–38        | 58| 31| 10| 32 | 31| 37| 50| 17| 33 | 14| 23| 63| 9  | 24| 66| 78 | 22 | 0  | 3       | LA           | PF1         |
| S15–40        | 53| 37| 10| 34 | 37| 29| 48| 18| 34 | 25| 25| 49| 21| 27| 52| 65 | 29 | 6  | 2       | FA           | PF1         |
| S15–41        | 44| 45| 11| 17 | 45| 38| 29| 43| 28 | 36| 23| 41| 28| 25| 46| 96 | 0   | 4  | 0       | FA           | PF1         |
| S15–45        | 64| 21| 15| 30 | 21| 49| 59| 17| 24 | 33| 18| 50| 27| 19| 54| 59 | 41 | 0  | 0       | LA           | PF1         |
| S15–47        | 53| 46| 1  | 32 | 46| 21| 44| 24| 31 | 16| 4  | 80| 6  | 4  | 90| 100| 0  | 0  | 0       | FA           | PF1         |
| S15–48        | 51| 7  | 42| 25 | 7 | 68| 77| 10| 13 | 14| 24| 62| 13| 25| 63| 47 | 37 | 16 | 0       | LA           | PF1         |
| S15–49        | 49| 41| 10| 21 | 41| 38| 36| 47| 17 | 45| 18| 37| 29| 23| 48| 75 | 25 | 0  | 1       | FA           | PF1         |
| S15–51        | 42| 37| 21| 16 | 37| 47| 30| 45| 24 | 16| 38| 46| 6  | 43| 51| 43 | 18 | 39 | 0       | FA           | PF1         |
| S15–52        | 38| 23| 39| 21 | 23| 57| 48| 42| 10 | 15| 51| 34| 11| 54| 36| 56 | 36 | 9  | 0       | LA           | PF1         |
| S15–53        | 20| 43| 37| 7  | 43| 50| 15| 53| 32 | 14| 71| 15| 6  | 77| 17| 69 | 18 | 13 | 0       | LA           | PF2         |
| NTR1 Section  | gw9–10* | 7  | 32| 62| 4  | 32| 64| 14| 79 | 6  | 0  | 87| 13 | 0  | 87| 13 | 55 | 15 | 0       | LA           | PF3         |
| NTR1–9*       | 12| 14| 74 | 2  | 14| 84| 16| 80 | 5  | 4  | 46| 51 | 0  | 47| 53 | 52 | 28 | 20      | 0       | LA           | PF3         |
| NTR1–12*      | 10| 3  | 87 | 0  | 3  | 97| 25| 75 | 0  | 1  | 67| 31 | 0  | 68| 32 | 47 | 48 | 0       | 4       | 0       | FA           | PF3         |
| NTR1–7*       | 18| 1  | 80 | 1  | 1  | 98| 33| 67 | 0  | 1  | 25| 74 | 0  | 25| 75 | 67 | 33 | 0       | 4       | 0       | FA           | PF3         |
| NTR36–17      | 6  | 3  | 91 | 3  | 3  | 94| 44| 44 | 13 | 0  | 57| 43 | 0  | 57| 43 | 68 | 17 | 15      | 0       | LA           | PF3         |
| NTR3–17       | 13| 16| 70 | 1  | 17| 83| 4 | 92 | 4  | 4  | 45| 51 | 0  | 47| 53 | 50 | 25 | 25      | 2       | LA           | PF3         |
| TD140         | 6  | 3  | 91 | 0  | 3  | 97| 13| 63 | 25 | 0  | 35| 65 | 0  | 35| 65 | 46 | 41 | 13      | 3       | LA           | PF3         |
| TDG Section   | TD2–17 | 1  | 2 | 96 | 0  | 2  | 98| 0  | 100| 0  | 0  | 70| 30 | 6  | 87| 7  | 0   | LA       | PF3         |
| TD3–17        | 2  | 4  | 94 | 1  | 4  | 95| 17| 58 | 25 | 0  | 84| 16 | 0  | 84| 16 | 53 | 63 | 23      | 0       | LA           | PF3         |
diagrams (Fig. 8; Table 2). All quartz grains are grouped together in the QFL diagram to emphasize the variations of grain stability among quartz, feldspar, and lithic fragments. In contrast, polycrystalline and microcrystalline quartz grains are counted as lithic fragments in the QmFLt diagram to emphasize the grain size of the source rocks, because fine-grained source rocks produce more lithic fragments than monocrystalline grains (Dickinson and Suczek 1979). Moreover, ternary diagrams of QmPK, QpLvLs, LmtLvLst, and LvfLvmiLvl use subgroups of QFL to show the characteristics of monocrystalline, polycrystalline, lithic, and volcanic lithic grains, respectively (Fig. 8; Dickinson and Suczek 1979; Dickinson et al. 1983). These diagrams are used to further classify the petrofacies for detailed interpretation of source lithology. Finally, ternary diagrams of QFL, QmFLt, and QpLvLs are used to interpret the tectonic settings of source areas using the templates of Dickinson and Suczek (1979) and Dickinson et al. (1983).

4.4.1 Petrofacies 1

Petrofacies 1 has a mean composition of Q 51 F30 L 19 and Qm21F30Lt49 and occurs in 11 lithic arenites, nine feldspathic arenites, and one lithic wacke (Fig. 8; Table 2). Quartz grains dominate and consist of monocrystalline, polycrystalline, and chert grains, which account for 21%, 13%, and 17% of the total grains, respectively. Feldspars account for 30% of the total grains and are slightly more enriched with plagioclase than potassium feldspars, with an average plagioclase/feldspar (P/F) ratio of 0.55 and a mean composition of Qm43P33K24 (Fig. 8). Polycrystalline and lithic grains are mainly composed of polycrystalline quartz and chert with mean compositions of Qp 61 Lv 25 Ls 14 and Lmt 19 Lv 29 Lst 52 (Fig. 8). Other lithic fragments include volcanic and mudrock, which account for 12% and 7% of the total grains, respectively. Volcanic lithic fragments are the minor component and are mainly composed of felsic volcanic lithics with a mean composition of Lvf 66 Lvmi 25 Lvl 19 (Fig. 8).

The mean composition of Petrofacies 1 suggests that its source rocks were probably a suite of felsic igneous rocks, quartzite, chert, and mudrocks. The common monocrystalline quartz and feldspar grains and the occurrence of felsic volcanic lithic sandy and gravelly clasts (see also Section 5) indicate a felsic plutonic and
volcanic origin for this petrofacies. The undulatory quartz grains and polycrystalline quartz with deformed texture indicate a metamorphic origin. Finally, chert and mudrock fragments were originated from sedimentary rocks.

The mean composition of Petrofacies 1 also provides clues to understand the tectonic setting of source areas. The mean compositions of QFL, QmFLt, QpLvLs groups fall within the tectonic fields of the recycled orogen, transitional arc, and subduction complex (Fig. 8; Dickinson and Suczek 1979; Dickinson et al. 1983). The tectonic settings interpreted from QFL and QmFLt distributions are different, because polycrystalline quartz and chert grains are incorporated as quartz in QFL diagram but as lithic fragments in QmFLt diagram. The chert-rich sandstones from Klamath Mountains within the North American Cordilleran known magmatic arc and associated accretionary wedge and trench sources fall in the recycled orogen in QFL and lithic recycled field in QmFLt (Dickinson and Suczek 1979; Dickinson et al. 1983). Similarly, Petrofacies 1 of WTG-LC sandstones is also enriched in polycrystalline quartz and chert and falls in the recycled orogen in QFL and transitional arc in QmFLt plots. Hence, Petrofacies 1 was likely derived from transitional

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**Fig. 7** Photomicrographs of sandstones in the Wutonggou low-order cycle. 

- **a** A nonundulatory monocrystalline quartz grain (Qnu) with embayment, a potassium feldspar with Carlsbad twinning, and a plagioclase with albite twinning. Sample TD110, lower Taodonggou section; 
- **b** A polycrystalline quartz grain with sutured quartz crystals, indicating its metamorphic origin. The blue arrow points to an elongate muscovite grain. S15–36, lower Zhaobishan section; 
- **c** A slightly clay coated chert grain. NTR39–17, upper North Tarlong section; 
- **d** A volcanic lithic fragment with felsic texture. The phenocrysts are mainly feldspar grains. TD110, lower Taodonggou section; 
- **e** A volcanic lithic fragment with microlitic texture, an angular mudrock fragment, and a volcanic lithic with lathwork texture showing large feldspar laths. These grains suggest sedimentary and volcanic origins. TD108, lower Taodonggou section; 
- **f** A volcanic lithic fragment with microlitic texture and a volcanic lithic with vitric texture, suggesting a volcanic origin. TD168, upper Taodonggou section. All micrographs are taken under cross-polarized light. Scale bar is 1 mm long in all photos. See Table 1 for the abbreviations of grain categories.
arc sources and associated accretionary wedge and trench, which, as a whole, is termed as a subduction complex by Dickinson and Suczek (1979). The QpLvLs plot substantiates the interpretation that Petrofacies 1 falls in the subduction complex field (Fig. 8). Overall, these three diagrams indicate that the sources of Petrofacies 1 include felsic volcanic and plutonic rocks from a transitional volcanic arc, and quartzite, chert and mudrock from the associated accretionary wedge and trench.

4.4.2 Petrofacies 2

Petrofacies 2 has a mean composition of Q_{28}F_{36}L_{36} and Qm_{14}F_{36}Lt_{50} and consists of seven litharenites and two feldspathic arenites (Fig. 8; Table 2). Quartz is still the major component but less enriched than Petrofacies 1. The monocrystalline and polycrystalline quartz and chert account for 14%, 7%, and 7% of the total grains, respectively. Feldspars are slightly enriched than Petrofacies 1 and account for 36% of the total grains. Plagioclase is dominant with an average P/F ratio of 0.65 and a mean composition of Qm_{29}P_{47}K_{24} (Fig. 8). The polycrystalline grains and lithic fragments include mainly volcanic lithic fragments and subordinate polycrystalline quartz, chert and mudrock fragments with mean compositions of Qp_{28}Lv_{65}Ls_{7} and Lmt_{6}Lv_{68}Lst_{23}. Similar to Petrofacies 1, the felsic volcanic lithic fragments are the major components of volcanic lithic fragments with a mean composition of Lv_{28}Lv_{m15}Lv_{7}.

The composition of Petrofacies 2 suggests that the sources include predominant felsic volcanic and plutonic rocks and subordinate quartzite, chert, and mudrocks. The occurrence of large amounts of volcanic lithic fragments, especially felsic volcanic ones, is indicative of felsic volcanic and plutonic source rocks. The common occurrence of monocrystalline quartz and feldspar grains and rhyolitic and granitic gravels (see Section 5) support this interpretation. The content of polycrystalline quartz, chert, and mudrock fragments is markedly lower than that of volcanic lithic fragments, suggesting that metamorphic and sedimentary rocks are subordinate sources.

The mean compositions of samples of Petrofacies 2 on the QFL, QmF Lt, and QpLvLs ternary diagrams fall within the tectonic fields of transitional arc and the mixed zone between subduction complex and arc orogen (Fig. 8). Mean composition falls in the field of transitional arc in both QFL and QmF Lt diagrams, indicating the transitional volcanic arc origin of these sandstones. In addition, mean composition falls in the

| Section | Lithology (%) |
|----------|---------------|
|          | Rhyolite | Andesite | Basalt | Chert | Quartzite | Granite | Limestone | Sandstone | Mudstone |
| ZBS 26   | 24       | 30       | 25     | 10    | 11        | 0       | 0         | 0         | 0        |
| ZBS 204  | 8        | 32       | 21     | 25    | 14        | 0       | 0         | 0         | 0        |
| ZBS 501  | 32       | 26       | 30     | 6     | 6         | 0       | 0         | 0         | 0        |
| ZBS 865  | 23       | 29       | 29     | 0     | 10        | 6       | 3         | 0         | 0        |
| Ave      | 22       | 29       | 26     | 9     | 10        | 3       | 3         | 1         |          |
| TDG 5    | 17       | 35       | 43     | 0     | 0         | 0       | 0         | 4         | 0        |
| TDG 29   | 43       | 26       | 26     | 4     | 0         | 0       | 0         | 0         | 0        |
| TDG 38   | 17       | 30       | 50     | 0     | 0         | 0       | 0         | 3         | 0        |
| TDG 80   | 13       | 35       | 48     | 4     | 0         | 0       | 0         | 0         | 0        |
| TDG 113  | 35       | 10       | 50     | 0     | 0         | 5       | 0         | 0         | 0        |
| TDG 139  | 29       | 7        | 64     | 0     | 0         | 0       | 0         | 0         | 0        |
| TDG 281  | 27       | 18       | 50     | 0     | 0         | 5       | 0         | 0         | 0        |
| Ave      | 26       | 23       | 47     | 1     | 1         | 1       | 1         | 1         |          |
| NTRL 35  | 16       | 72       | 72     | 6     | 3         | 3       |           |           |          |
| NTRL 160 | 28       | 41       | 28     | 2     |           | 3       |           |           |          |
| NTRL 232 | 72       | 16       | 10     | 0     |           | 2       |           |           |          |
| NTRL 247 | 56       | 32       | 9      | 0     |           | 3       |           |           |          |
| NTRL 322 | 48       | 22       | 26     | 0     |           | 4       |           |           |          |
| NTRL 523 | 39       | 13       | 35     | 0     | 9         | 4       |           |           |          |
| Ave      | 43       | 33       | 19     | 1     | 9         | 3       | 4         |           |          |
mixed zone rather than the arc orogen field in the QpLvLs diagram, which indicates that the sources are mixed rocks from volcanic arc, accretionary wedge, and trench. The sources of accretionary wedge and trench are not reflected in the QFL and QmFLt diagrams due to the relatively low content of chert and polycrystalline quartz grains. Overall, the three diagrams indicate that the transitional or dissected volcanic arc rocks are the major sources, and the accretionary wedge and trench metamorphic and sedimentary rocks are the secondary sources.

4.4.3 Petrofacies 3

Petrofacies 3 has a mean composition of $Q_{6}F_{13}L_{79}$ and $Qm_{2}F_{13}L_{85}$ and occurs in 28 litharenites and two lithic wackes (Fig. 8; Table 2). Quartz is no longer the major component. Monocrystalline, polycrystalline, and chert grains only account for 2%, 1%, and 5% of the total grains, respectively. Similarly, feldspars decrease significantly and only account for 13% of the total grains. Plagioclase is still more than K-feldspars, with an average P/F ratio of 0.69 and a mean composition of $Qm_{16}P_{64}K_{20}$ (Fig. 8). In contrast to Petrofacies 1 and Petrofacies 2, polycrystalline grains and lithic fragments are the major components and largely consist of volcanic and sedimentary fragments with mean compositions of $Qp_{6}Lv_{60}L_{s34}$ and $Lmt_{0}Lv_{60}L_{st40}$ (Fig. 8d, e). Finally, although felsic volcanic lithic fragments still dominate in Petrofacies 3, the proportion of microlitic volcanic lithic fragments increases significantly, as indicated by the mean composition of volcanic lithic fragments $Lv_{f46}Lv_{mi39}Lv_{15}$ (Fig. 8).

The mean composition of Petrofacies 3 indicates that the sources are rhyolites, andesites, and mudrocks. Moreover, the mean compositions of Petrofacies 3 fall within the fields of undissected volcanic arc in QFL and QmFLt diagrams and arc orogen in the QpLvLs diagram, indicating the presence of undissected volcanic arc rocks in the sources. However, although the mean composition is in the undissected arc field in QmFLt diagram, 15 sandstones of Petrofacies 3 fall within the fields of undissected volcanic arc in QFL and QmFLt diagrams and arc orogen in the QpLvLs diagram, indicating the presence of undissected volcanic arc rocks in the sources. However, although the mean composition is in the undissected arc field in QmFLt diagram, 15 sandstones of Petrofacies 3 fall within the lithic recycled field (Fig. 8). The recycled lithics may be derived from uplifted older sedimentary strata (Dickinson et al. 1983; Dickinson 1985). Thus, these sandstones may have mixed sources of sedimentary and volcanic rocks. The sedimentary sources are unlikely the trench-fill sedimentary rocks, because the mudrocks fragments of Petrofacies 3 include abundant angular rip-up clasts and the concurrent chert fragments are rare. As a result, mudrock fragments were more likely derived from a
nearby source area, such as the rift shoulders, the uplifted hanging wall of the graben (Yang et al. 2010; Guan 2011; Obrist-Farner and Yang 2017).

4.5 Provenance of sedimentary rocks in Bogda Mountains

The stratigraphic distribution of petrofacies along each section provides clues to understand the evolution of provenance through time. Clast composition of conglomerates and paleocurrent directions are used to substantiate provenance interpretations. Finally, the correlation of the four studied sections demonstrates the spatial variations of sandstone composition and source areas.

4.5.1 Provenance of sedimentary rocks in the Zhaobishan section

Petrofacies of sandstones of the WTG-LC in Zhaobishan section change upsection, suggesting that the lithology and tectonic setting of the source areas of the lower and upper WTG-LC sandstones are different. Petrofacies 1 occurs in 14 sandstones in the lower 420 m, whereas Petrofacies 2 occurs in nine sandstones in the upper 350 m of Zhaobishan section (Figs. 6, 9; Table 2). The occurrence of Petrofacies 1 suggests that the sandstones in the lower WTG-LC were derived from a transitional volcanic arc and associated accretionary wedge and trench. As the ENTS was the collisional product of oceanic plate, volcanic arc, and continental plates (Allen et al. 1991; Gao et al., 1998; Xiao et al. 2004, 2013; Charvet et al. 2011), it should contain volcanic, metamorphic, and sedimentary rocks. Rhyolites, fragmental radiolarian chert and high-pressure metamorphic rocks are still exposed in current ENTS (Xiao et al. 2004; Wang et al. 2006), indicating that the ENTS was the available source area of Zhaobishan section during late Permian-earliest Triassic.

Clast composition of WTG-LC conglomerates and paleocurrent directions support the provenance interpretation. Two conglomeratic beds at the 26 and 204 m thickness points in the lower Zhaobishan section consist of abundant volcanic and a few quartzite and chert gravels (Fig. 6; Table 3). The paleocurrent direction is northward at the bottom of Zhaobishan section (Fig. 6). These data indicate that the ENTS, located ~100 km south of the section, served as the volcanic, metamorphic, and sedimentary sources to the lower WTG-LC sandstones in Zhaobishan section. Finally, the absence of plutonic gravels in the conglomerates suggests that the volcanic arc might be undissected or slightly dissected without major exposure of plutons.
In contrast, the occurrence of Petrofacies 2 in the upper WTG-LC sandstones indicates that the sources are mainly transitional volcanic arc and subordinately accretionary wedge and trench. The ENTS is still interpreted as the available source areas containing igneous, metamorphic and sedimentary rocks. The northward paleocurrent direction identified in the bed at 450 m in the upper half of the section, supports the interpretation (Fig. 6). The significant decrease of polycrystalline quartz and chert fragments and the occurrence of granitic gravels in conglomeratic beds at the thickness levels of 501 and 765 m (Fig. 6; Table 3) indicate that the source lithology contains a significant amount of granites along with rhyolites from the transitional volcanic arc and diminishing quartzite and chert from the accretionary wedge and trench.

4.5.2 Provenance of sedimentary rocks in the North Tarlong section

The distribution of petrofacies in sandstones of the WTG-LC in North Tarlong section also indicates different provenances for lower and upper WTG-LC sandstones. Petrofacies 3 occurs in 11 sandstones in the lower 600 m, whereas Petrofacies 1 occurs in four sandstones in the upper 220 m of North Tarlong section. The occurrence of Petrofacies 3 in the lower WTG-LC sandstones indicates that the sources were an undissected volcanic arc and rift shoulders. The overall northward paleocurrent directions documented in the beds at 5, 29, and 38 m suggest the presence of a highland to the south, likely the ENTS (Fig. 6; Table 3). Abundant volcanic clasts are present in the lower five conglomeratic beds (Fig. 6). Thus, rhyolites and andesites probably covered a large area in the ENTS as the main source. In addition, the large number of mudrock fragments in the lower sandstones may have been derived from local rift shoulders.

In contrast, the occurrence of Petrofacies 1 in the upper WTG-LC sandstones suggests that the source changed to the transitional volcanic arc and associated accretionary wedge and trench of the ENTS. This change is also evidenced by the presence of granitic gravels in the uppermost conglomeratic bed in North Tarlong section (Fig. 6; Table 3).

4.5.3 Provenance of sedimentary rocks in the Taodonggou section

Sandstone petrofacies of the WTG-LC also vary in the Taodonggou section, indicating changes of provenance. Petrofacies 3 occurs in six sandstones in the lower 200 m and six sandstones in the upper 90 m of the section, whereas Petrofacies 1 occurs in three sandstones in the middle 50 m of Taodonggou section (Figs. 6, 9). The occurrence of Petrofacies 3 in the lower and upper parts of Taodonggou section suggests volcanic and sedimentary sources. The rift shoulders are likely the sources for mudrock fragments. In addition, the basement in the Taodonggou-Tarlong area contains upper Carboniferous basaltic, andesitic, and sedimentary rocks (Yang et al. 2010; Yang et al. 2013). Thus, the rift shoulders, if the basement rocks were exposed, might also supply basaltic and andesitic fragments. The abundant volcanic clasts in WTG-LC conglomerates and southward paleocurrent directions documented in Taodonggou section (Fig. 6; Table 3) support a rift-shoulder source. However, the rhyolitic fragments might come from some other sources. Therefore, the ENTS might have provided fragments of rhyolites and likely, andesites.

In contrast, the occurrence of Petrofacies 1 in the middle part of Taodonggou section suggests that sediments were derived from rocks in a transitional volcanic arc and the associated accretionary wedge and trench. The ENTS might have been the likely source to supply the felsic volcanic and plutonic, metamorphic, and sedimentary rocks in the sandstones of the middle Taodonggou section.

4.5.4 Provenance of sedimentary rocks in the Dalongkou section

Only Petrofacies 3 occurs in the seven sandstones in Dalongkou section, suggesting a persistent provenance. Petrofacies 3 indicates volcanic and sedimentary sources from rift shoulders and the ENTS, as discussed above. In addition, the paleocurrent directions are either northward or southward (Fig. 6), which suggest a complex dispersal pattern, probably originated from surrounding rift shoulders. No basement rocks are exposed in the Dalongkou area at the present time. Thus, the volcanic clasts may have been derived from either ancient rift shoulders and/or undissected volcanic arc in the ENTS. A dominant rift-shoulder source conforms to that for the sandstones of the underlying Quanzijjie low-order cycle (Obrist-Farner and Yang 2017).

4.5.5 Spatial correlations of petrofacies

The distribution of petrofacies in the four studied sections was correlated to identify spatial variations of sandstone compositions and provenance. Petrofacies of North Tarlong section is correlative with those of South Tarlong section (Fig. 9; Guan 2011). These sections are located at the northern and southern limbs of a syncline and both converge toward the axis and were deposited in the same graben (Fig. 3; Yang et al. 2010; Guan 2011). Twenty-three WTG-LC sandstones were divided into the lower lithic-rich and upper quartz- and feldspar-rich petrofacies, which resemble the distribution of petrofacies in North Tarlong section (Guan 2011; Figs. 9, 10). The good correlation between these two sections
suggests that they shared the same provenance during the deposition of the WTG-LC.

Correlation between petrofacies of North Tarlong and South Tarlong sections and those of Taodonggou section shows a slight difference. The petrofacies of the three sections shift from Petrofacies 3 to Petrofacies 1 upsection, suggesting a similar trend of provenance evolution. However, Petrofacies 1 occurs only in a thin interval in the middle part of the Taodonggou section, about one-seventh of the total thickness, and abruptly changes to Petrofacies 3 again. Taodonggou section is 6 km away from the North Tarlong-South Tarlong sections and located in the same half-graben as North Tarlong-South Tarlong sections (Yang et al. 2010). However, the thickness and types of high-order cycles change significantly. The Taodonggou section was at the ramp of the half-graben, while North Tarlong-South Tarlong sections were at the depocenters (Yang et al. 2010). Thus, they do not have the same depositional environments and even might not have the same drainage areas (e.g., Soreghan and Cohen 1993). These factors may have caused the abrupt shift of petrofacies in the upper Taodonggou section.

Additionally, the correlation between petrofacies of Tarlong-Taodonggou areas and that of the northwestern Turpan Basin has variations. The northwestern Turpan Basin is about 30 km south of Tarlong-Taodonggou area (Fig. 2), where two Wutonggou Formation sandstones are lithic-rich, resembling the lower WTG-LC sandstones in North Tarlong, South Tarlong, and Taodonggou sections (Shao et al. 2001; Fig. 10). However, the upper quartz- and feldspar-rich sandstones are not documented in the northwestern Turpan Basin.

The trend of petrofacies evolution of Zhaobishan section cannot be correlated with those of North Tarlong, South Tarlong, and Taodonggou sections, because Pet- rofacies 3 is absent in Zhaobishan section. On the other hand, the petrofacies shifts in these sections all occur in the middle parts of the sections (Fig. 9), which suggests an approximately coeval tectonic event in both source areas of Zhaobishan and Tarlong-Taodonggou areas in the ENTS. The shift in North Tarlong section occurred at a bed 10 m below a bentonite with an age of 253.11 ± 0.05 Ma (Yang et al. 2010; Fig. 9). Therefore, the petrofacies shift of the studied sections occurred probably during Wuchiapingian-early Changhsingian transition.

Moreover, the petrofacies of Zhaobishan section may be similar to those of sandstones of the Wutonggou Formation in the northeastern part of the Turpan Basin, as documented by Shao et al. (2001). Eleven sandstones of the Wutonggou Formation from Xishan and Kulai sections in northeastern Turpan Basin are quartz- and feldspar-rich, similar to Petrofacies 2 in this study (see Fig. 2 for the section locations; Fig. 10). This suggests that the two areas may share a similar provenance.

The petrofacies trend of Dalongkou section cannot be correlated with those in the other sections, because the section contains only Petrofacies 3. The Dalongkou section was interpreted to be related to a separated drainage system (see above). This interpretation fits the tectonic setting of the greater Turpan-Junggar basin as a highly-partitioned rift basin (Yang et al. 2010), where the abundant rift shoulders might have hampered the transport of sediments from the ENTS. Alternatively, rivers originating from the ENTS might have been persistently draining an area rich in volcanic and sedimentary rocks during the deposition of the entire WTG-LC.

Previous studies suggest the WTG-LC sandstones in the southern Junggar Basin and Northern Turpan Basin are uniformly volcanic-rich with slight variations in compositions based on limited numbers of point-counting data (Carroll et al. 1995; Hendrix 2000; Greene et al. 2005). The petrofacies distribution and correlation among multiple detailed stratigraphic sections present a clear spatial and temporal pattern of provenance evolution during the deposition of WTG-LC sandstones.

5 Discussion: the unroofing history of the Eastern North Tianshan Suture

The interpreted provenance lithology and tectonic setting, and the evolution of WTG-LC sandstones can be used to reconstruct the unroofing history of the ENTS. Overall, the ENTS had been persistently unroofed during late Permian-earliest Triassic to provide a large
amount of siliciclastic sediments northward into the greater Turpan-Junggar basin. Uplifting of the ENTS is likely, but the rate of uplifting cannot be confirmed. During approximately the Wuchiapingian Stage when the lower WTG-LC sandstones in Zhaobishan section were deposited, the source area in the eastern part of the ENTS was composed of rocks of the undissected volcanic arc, accretionary wedge, and trench (Fig. 11). During approximately Changhsingian-early Induan stages, when the upper WTG-LC sandstones were deposited, the source area was covered by rocks mainly generated in a transitional volcanic arc and subordinately the accretionary wedge and trench, where deep-seated granitic plutons started to expose (Fig. 11). The unroofing trend indicates the amalgamation of accretionary wedge, trench, and volcanic arc rocks was caused by the collision between the Junggar Plate and Central Tianshan Suture (Allen et al. 1993; Gao et al., 1998; Xiao et al. 2004, 2013; Charvet et al. 2011). The southward subduction of the North Tianshan Ocean formed the accretionary wedge and trench, which were accreted together by later continuous movement. A similar trend was reported from the Eocene-Middle Miocene sandstones in the collisional zone between Izu Arcs and the Honshu Arc in central Japan (Okuzawa and Hisada 2008), where the older sources are volcaniclasts, and the younger sources are accretionary wedge and trench rocks. The unroofing trend of the ENTS continued through the Triassic as indicated by increasingly quartzose compositions of Triassic sandstones in the Turpan-Hami Basin (Shao et al. 2001; Greene et al. 2005), as a consequence of progressive dissection of the volcanic arc in the ENTS.

In contrast to the source area of Zhaobishan section in the eastern part of the ENTS, the source areas of Tarlong-Taodonggou sections are located in the western part of the ENTS. These areas are probably 90 km west of the Zhaobishan source area, which is the present-day distance between Zhaobishan and Tarlong-Taodonggou areas. The Tarlong-Taodonoggou source area in the ENTS has a different unroofing history. It was covered by an undissected volcanic arc during the Wuchiapingian at the time of deposition of the lower WTG-LC sandstones. However, the area was covered with complex lithologies of the amalgamated transitional volcanic arc, accretionary wedge, and trench rocks during the Changhsingian-early Induan at the time of deposition of the upper WTG-LC sandstones. This unroofing trend is similar to that of the continental arc within the Turan Plate, Western Turkmenistan (Garzanti and Gaetani 2002) and the Sierra Nevada in North America (Ingersoll 2012), where the contents of quartz and feldspar in the studied sandstones increase at the expense of volcanic lithic fragments. During continuous plate consumption, sandstones may also show an increased increase of polycrystalline quartz and chert.

Fig. 11 Unroofing trends of source areas in Eastern North Tianshan Suture, as interpreted from petrofacies of Wutonggou low-order cycle sandstones in Bogda Mountains. The trend for the Zhaobishan section shows that the sources changed from the undissected volcanic arc, accretionary wedge and trench rocks to the transitional volcanic arc rocks. The trend for North Tarlong, South Tarlong, and Taodonggou sections shows that the sources shifted from undissected volcanic arc rocks to transitional volcanic arc, accretionary wedge and trench rocks. See Fig. 8 for the explanations of the tectonic fields.
fragments derived from the accretionary wedge and trench (Dickinson and Suczek 1979; Dickinson et al. 1983; Garzanti et al. 2007).

The two different unroofing trends between the source areas for Zhaobishan and Tarlong-Taodonggou sections indicate that the ENTS was an amalgamated complex with spatial and temporal variations in lithology during the late Permian-earliest Triassic. During the Wuchiapingian, the source areas of Zhaobishan section in the eastern part of the ENTS contained amalgamated rocks of the undissected volcanic arc, accretionary wedge and trench, whereas the source areas of Tarlong-Taodonggou areas in the western part of the ENTS contained assemblages of undissected volcanic arcs. During
6 Conclusions
Sandstones from the upper Permian-lowermost Triassic WTG-LC in Bogda Mountains, NW China, provide critical information to understand the provenance and unroofing history of the ENTS. The source of the Zhaobishan section (east part) changed from the rocks of the undissected volcanic arc, accretionary wedge and trench to those of transitional volcanic arc with subordinate accretionary wedge and trench. The source of the North Tarlong and Taodonggou sections (west part) shifted from an undissected volcanic arc and sedimentary rocks from the ENTS and rift shoulders to rocks in the transitional volcanic arc, accretionary wedge and trench. The sources of the Dalongkou section, located 70 km north of the North Tarlong and Taodonggou sections, are undissected volcanic arc and sedimentary rocks from the ENTS and rift shoulders. Unroofing history differs between the source areas for Zhaobishan and North Tarlong-Taodonggou sections, indicating that the eastern ENTS changed from the trinity of volcanic arc, accretionary wedge, and trench to the transitional volcanic arc, whereas the western ENTS shifted from the volcanic arc to the trinity of transitional volcanic arc, accretionary wedge and trench. This study provides sedimentological evidence to support that the ENTS was an amalgamated complex of volcanic arc, accretionary wedge, and trench with great spatial and temporal heterogeneity in lithology and experienced persistent unroofing during the late Permian-earliest Triassic. This study reconstructs a key element of the Chinese Tianshan Suture and serves as an analog for future studies to understand the lithology and unroofing processes of ancient sutures.

Abbreviations
ENTS: Eastern North Tianshan Suture; WTG-LC: Wutonggou lower-order cycle

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Authors’ contributions
DYZ performed the data collection, analysis, interpretation, and drafted the manuscript. WY collected data, and revised the manuscript. The author(s) read and approved the final manuscript.

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Availability of data and materials
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Competing interests
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