Orbital pacing and secular evolution of the Early Jurassic carbon cycle

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Global perturbations to the Early Jurassic environment (~201 to ~174 Ma), notably during the Triassic–Jurassic transition and Toarcian Oceanic Anoxic Event, are well studied and largely associated with volcanicogenic greenhouse gas emissions released by large igneous provinces. The long-term secular evolution, timing, and pacing of changes in the Early Jurassic carbon cycle that provide context for these events are thus far poorly understood due to a lack of continuous high-resolution δ13C data. Here we present a δ13CTOC record for the uppermost Rhaetian (Triassic) to Pliensbachian (Lower Jurassic), derived from a calcareous mudstone succession of the exceptionally expanded Llanbedr (Mochras Farm) borehole, Cardigan Bay Basin, Wales, United Kingdom. Combined with existing δ13CTOC data from the Toarcian, the compilation covers the entire Lower Jurassic. The dataset reproduces large-amplitude δ13CTOC excursions (>3‰) recognized elsewhere, at the Sinemurian–Pliensbachian transition and in the lower Toarcian serpentinitum zone, as well as several previously identified medium-amplitude (~0.5 to 2‰) shifts in the Hettangian to Pliensbachian interval. In addition, multiple hitherto undiscovered isotope shifts of comparable amplitude and stratigraphic extent are recorded, demonstrating that those similar features described earlier from stratigraphically more limited sections are nonunique in a long-term context. These shifts are identified as long-eccentricity (~405-ky) orbital cycles. Orbital tuning of the δ13CTOC record provides the basis for an astrochronological duration estimate for the Pliensbachian and Sinemurian, giving implications for the duration of the Hettangian Stage. Overall the chronostratigraphy illustrates particular sensitivity of the marine carbon cycle to long-eccentricity orbital forcing.

Significance

Cyclic variations in Earth’s orbit drive periodic changes in the ocean–atmosphere system at a time scale of tens of thousands of years. The Mochras δ13CTOC record illustrates the continued impact of long-eccentricity (405-ky) orbital forcing on the carbon cycle over at least ~18 My of Early Jurassic time and emphasizes orbital forcing as a driving mechanism behind medium-amplitude δ13C fluctuations superimposed on larger-scale trends that are driven by other variables such as tecstratigraphically well-defined δ13CTOC record from uppermost Rhaetian (Triassic) to Pliensbachian (Lower Jurassic) strata, where they are recorded as individual shifts or series of shifts within stratigraphically limited sections. Some of these short-term δ13C excursions have been shown to represent changes in the supraregional to global carbon cycle, marked by synchronous changes in δ13C in marine and terrestrial organic and inorganic substrates and recorded on a wide geographic extent (e.g., refs. 10, 16, 23, and 24). However, due to the previous lack of a continuous dataset capturing and contextualizing all isotopic shifts in a single record, there is no holistic understanding of the global nature, causal mechanisms, and the chronology and pacing of these CIEs. Therefore, these δ13C shifts have largely been interpreted as stand-alone events, linked to a release of 12C from as-yet-undefined sources, reduced organic productivity (leaving more 12C in the ocean–atmosphere system) and/or 13C-depleted carbon sequestration and orbital forced environmental change affecting the carbon cycle on the scale of Milankovitch cyclicity (17, 20, 24, 25). Evidence for the latter is so far limited to the Hettangian to Early Sinemurian and the early Toarcian, where high-resolution isotope records provide the basis for cyclostratigraphic analysis (17–19, 29–31).

The data illustrated herein provide a continuous and biostratigraphically well-defined δ13CTOC record from uppermost Rhaetian (Triassic) to Pliensbachian (Lower Jurassic) strata,

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with a resolution high enough to examine CIEs of varying magnitudes and temporal extent in their stratigraphic context, thereby enabling a distinction between orbital, tectonic, oceanographic, or volcanic forcing mechanisms of the carbon cycle over this time interval.

Geological Setting

The Llanbedr (Mochras Farm) borehole (hereafter referred to as Mochras) cored the Lower Jurassic of the Cardigan Bay Basin (Wales, United Kingdom), an extensional structure related to the breakup of Pangaea (32). In the Early Jurassic, the basin was located at a midpaleolatitude in the Laurasian Seaway on the northwest fringes of the European shelf (Fig. 1 and refs. 33 and 34). The uppermost Pliensbachian and lower Toarcian strata are regarded as having been deposited in an unrestricted, open-marine setting (35).

The recovered sedimentary succession at Mochras comprises 32.05 m of continental Upper Triassic (Rhaetian) deposits (1,938.83 to 1,906.78 m below surface, mbs), ~1,305 m of Lower Jurassic Hettangian to Toarcian marine strata (1,906.78 to 601.83 mbs), and is unconformably overlain by Paleogene–Neogene sandstones and glaciogenic sediments (601.83 to 0 mbs, ref. 36). Ammonite biostratigraphy of the core was defined to a zonal and even subzonal level, and all ammonite zones of the Lower Jurassic have been identified with the exception of the lowermost Hettangian tilmanni zone (37, 38). Due to the lack of the base-Jurassic biostratigraphic marker Psiloceras speleae, the Triassic–Jurassic boundary in the Mochras borehole is placed at a lithological change from calcitic dolostone to calcareous mudstone at ~1,906.78 mbs (36, 38). About 1.7 m of biostratigraphically undefined strata lying between the base Jurassic and the base of the planorbis zone are referred to as “pre-planorbis beds,” likely equivalent to the basal Jurassic tilmanni zone (38).

The relative thinness of the pre-planorbis beds suggests a base-Jurassic hiatus at the sharp lithological change at ~1,906.78 mbs. A calcite-veined interval in the mid-Sinemurian oxynotum zone may be marked by a fault which, if present at all, cuts out less than one ammonite subzone (36). A small hiatus may also be present at the level of intraformational conglomerate at 627.38 mbs (36) within the upper Toarcian pseudoradiosus zone, and a further unconformity is present at the top of the Lower Jurassic (at 601.85 mbs), where sediments of the uppermost Toarcian aalenis zone are overlain by Paleogene strata (7, 36). In all other respects, the Lower Jurassic succession appears to be stratigraphically complete. However, core preservation below ~1,290 mbs is largely limited to reserve collection samples, each of which aggregate ~1.4 m intervals of broken core, with consequential reduction of stratigraphic resolution (see Materials and Methods and SI Appendix).

The Jurassic succession at Mochras is markedly expanded, with relatively uniform lithology compared to coeval strata elsewhere (38, 39). The strata primarily comprise calcareous mudstone, with varying silt and clay content, alternating with strongly bioturbated calcareous siltstone and silty limestone (36). Average Rock-Eval thermal maturation parameter (Tmax = ~430 °C) and vitrinite reflectance (Ro = 0.38 to 0.63) from previous studies indicate the presence of immature to early mature sedimentary organic matter (7, 22, 40). δ13C data from total organic carbon (δ13C_TOC) and carbonates (δ13C_carb) generated in previous studies suggest that the Mochras sedimentary archive records the long-term pattern of global carbon-cycle change (7, 14, 22, 41, 42).

Results

The high-resolution δ13C_TOC and Rock-Eval data from Mochras presented here for the uppermost Rhaetian to Pliensbachian are combined with published data for the Toarcian derived from the same core (ref. 7 and Fig. 2). The compiled δ13C_TOC record illustrates significant long- and short-term fluctuations in δ13C_TOC through the Lower Jurassic of the Mochras core. At a longer time scale, the record shows a long-term ~5‰ positive shift in δ13C_TOC from the lowermost Hettangian to upper Sinemurian. The Sinemurian–Pliensbachian boundary is characterized by a symmetrically shaped ~5‰ negative long-term trend and subsequent “recovery” (upper oxynotum to upper ibex zones, ~1,360 to ~1,060 mbs), reaching the lowest values in the lower Jamesoni zone. The mid-Pliensbachian interval presents a stable plateau in δ13C_TOC, followed by the upper margaritatus zone (subnodosus and gibbosus subzones) where δ13C_TOC values rise gradually and culminate in an abrupt ~2‰ positive excursion in the upper margaritatus zone (~930 to ~926 mbs). The margaritatus–spinatum zone boundary is marked by a sharp ~4‰ drop in organic carbon-isotope ratios, followed by a gradual positive shift throughout the spinatum zone. The Toarcian record comprises a lower Toarcian

Fig. 1. Early Jurassic paleogeography showing the location of the Mochras borehole (red star) within the northern Eurasian Seaway (red rectangle). Reprinted from ref. 39. Copyright (2019) with permission from Elsevier.
overarching positive CIE interrupted by the large negative CIE associated with the T-OAE, as described in previous studies (7, 14, 22, 41). All larger-scale trends in the $^{13}$C$_\text{TOC}$ data are reproduced in $^{13}$C$_\text{wood}$ presented for the upper Sinemurian to Toarcian, although the latter dataset shows a larger degree of variability (Fig. 2).

At a decameter scale, the $^{13}$C$_\text{TOC}$ record is characterized by consecutive alternating positive and negative shifts of ~0.5 to 2‰ magnitude, superimposed on the observed long-term isotopic trends. These fluctuations in $^{13}$C$_\text{TOC}$, hereafter referred to as medium-amplitude shifts, are particularly well-defined in the Hettangian to uppermost Sinemurian (between 1,906.78 and 1,340 mbs) and the mid-Pliensbachian ibex to lower margaritatus zones (between 1,120 and 980 mbs). The individual shifts appear larger in magnitude and more stratigraphically extensive in the Hettangian and Sinemurian compared with those in the Pliensbachian. Superimposed on these medium-amplitude shifts, fluctuations in $^{87}$Sr/$^{86}$Sr of up to 2‰ on a meter to centimeter scale occur, with larger magnitudes in the upper Sinemurian and Pliensbachian likely being an artifact of differing sample resolution. The calcium carbonate ($\text{CaCO}_3$) content of the Mochras strata is highly variable (~0.6 to 95%; Fig. 2). The long-term shifts in $\text{CaCO}_3$ appear to negatively correlate with the broad $^{13}$C$_\text{TOC}$ trends, with the exception of the upper Pliensbachian and lower
Toarcian successions. On a decameter scale, the CaCO₃ shows a clear fluctuation in the Hettangian and Sinemurian interval, but the pattern does not correspond to the medium-scale shifts in δ¹³C_TOC. The relatively higher variability in CaCO₃ on a meter to decameter scale over the Sinemurian–Pliensbachian transition and the upper Pliensbachian interval is associated with the higher data resolution obtained in these intervals.

The total organic carbon (TOC) content and hydrogen index (HI) values are generally low throughout the Hettangian to Pliensbachian of the Mochras core (Fig. 2). TOC and HI values in the Hettangian and most of the Sinemurian (~1.5 wt % and ~80 mg HC/g TOC on average, respectively) notably increase in the upper Sinemurian to lower Pliensbachian (2.6 wt %, up to 380 mg HC/g TOC, respectively) and are moderately elevated through the Pliensbachian (~1.4 wt % and ~170 mg HC/g TOC on average, respectively).

The increase in both TOC and HI accompanies the downward limb of the Sinemurian–Pliensbachian negative CIE, and some stratigraphic intervals with distinctly enhanced TOC and HI values also occur in the Hettangian and Sinemurian, coinciding with minimum values in δ¹³C_TOC (for example in the angulata, bucklandi, and lower rarasolitata zones). Similarly, the lowermost Jurassic sediments of the pre-planorbis beds and planorbis zone are also marked by distinctly elevated TOC and HI values (of 3.2 wt % and up to 660 mg HC/g TOC, respectively), coinciding with negative δ¹³C_TOC values of ~28‰ (Fig. 2). Overall, there is no clear correlation between TOC and δ¹³C_TOC (SI Appendix, Fig. S5).

Throughout the Toarcian, TOC values fluctuate between 0.5 and 1.5 wt %, with HI values of ~100 mg HC/g TOC (7). Slightly higher TOC and HI values are recorded in the bifrons and variabilis zones, and elevated TOC and HI values (up to 2.5 wt % up to 339 mg HC/g TOC, respectively) are associated with the negative CIE interval in the serpentinum zone (7).

Predominant components of sedimentary organic matter identified by maceral analysis are liptinites, most of which are represented by liptodetrinite (up to 96.7 vol %; SI Appendix, Fig. S7), a product of aerobic and mechanic degradation of lipinitic macerals (43). Markedly smaller but variable amounts of less-degraded liptinite macerals are algal in origin (alginite). Bituminite, also known as amorphous organic matter (AOM), which originates from anaerobic decomposition of algae and faunal plankton under anoxic conditions (44, 45), is primarily present in Pliensbachian samples (up to 24 vol %). Terrestrial organic matter comprising coal clasts, vitrinite, inertinite, sporinite, and cutinite accounts for variable relative amounts (3.3 to 58.8 vol %) of the total organic matter. Notably, the Pliensbachian samples contain a larger relative amount of terrestrially derived organic matter and bituminite compared to the Hettangian and Sinemurian samples. This stratigraphic trend is also reflected in the comparatively high abundance of macrofossil wood in the Pliensbachian and Toarcian part of the core, contrasting with very rare occurrences in the Hettangian and Sinemurian.

Discussion

Source and Preservation of Bulk Organic Matter in the Mochras Core.

The formation of liptodetrinite, the predominant organic component in the bulk organic matter assemblage, is associated with physical disintegration of lipinitic macerals in the water column and is indicative of extensive water-column circulation and high oxygen availability (45). The precursors of liptodetrinite can be of marine-aquatic origin or derive from terrestrial liptinites, such as sporinite and cutinite (46). The highly degraded shape of liptodetrinite macerals observed in Mochras (SI Appendix, Fig. S6) suggests fragile marine organic matter such as algae as primary precursor. The marine origin of the degraded particles is further supported by its fluorescence. The fluorescence of terrestrial organic matter should decrease with greater biodegradation (47) but in the studied samples appears higher compared to intact terrestrial liptinite. The HI values recorded in Mochras are likely highly compromised as a result of aerobic bacterial degradation of initially hydrogen-rich marine organic matter and are therefore not indicative of the primary source of the organic matter. Comparably larger relative amounts of bituminite (AOM) in the upper Pliensbachian strata indicate more anaerobic bottom-water conditions, resulting in preservation of lipid and hydrogen-rich organic matter. An increase in AOM and foraminifera organic inner wall linings was previously reported from Mochras, concomitant with the increase in TOC in the raricostatum to davoei zones (22). The elevated TOC and HI values around the Sinemurian–Pliensbachian transition were interpreted to result from both the increase in organic flux to the seafloor and low-oxygen bottom waters (22). Elevated TOC and HI values manifested on a smaller (decameter) scale and stratigraphically coincident with negative CIEs (well-expressed, for example, in the angulata, bucklandi, and lower obtusum zones) may similarly be linked to enhanced preservation of organic matter as a response to redox conditions.

The Lower Jurassic δ¹³C_TOC Record in the Mochras Core. The large-magnitude CIEs (~3‰) recorded in the Mochras core are the same as isotope events that have been previously recorded elsewhere, such as the early Toarcian negative CIE, punctuating an overarching positive excursion, in the tenonicostatum-serpentinum zones (7, 14, 22, 41) and the Sinemurian–Pliensbachian boundary negative CIE, both of which have been interpreted as due to increased release of isotopically light carbon into the ocean–atmosphere system (4, 5, 39). Markedly well-expressed in the Mochras δ¹³C_TOC record is an upper Pliensbachian (uppermost margaritatus zone) negative CIE, revealing a sharp ~4‰ downward shift following the upper margaritatus positive CIE. Both these CIEs have been recognized in multiple European basins (10, 15, 25, 48), as well as in the North American realm (16). The positive margaritatus zone CIE has been linked to widespread deposition of isotopically light organic matter during high sea levels and warm climates (6, 25, 49), with no apparent temporal link to large-scale volcanism (16). Subsequent cooling, associated sea-level fall, and restored water-column mixing have been suggested to have released accumulated light carbon through sediment reworking and oxidative and heterotrophic remobilization, causing the negative CIE in the uppermost margaritatus zone (8, 10, 26, 50, 51). Similarly, the influx (upwelling/recycling) of ¹⁴C-rich deep waters associated with a climatic cooling trend has been suggested as an alternative driving mechanism for the Sinemurian–Pliensbachian boundary negative CIE (11).

The clear parallelism between fossil plant matter and bulk organic δ¹³C records demonstrates, however, that the ocean–atmosphere and biosphere carbon reservoirs were simultaneously affected and the shifts in δ¹³C_TOC can thus not solely be explained with oceanographic changes such as redox-related preservation or upwelling of ¹⁴C-enriched deep waters. Despite possible circulation-related redox changes throughout the strata, notably around the Sinemurian–Pliensbachian transition and the upper Pliensbachian (22, 26, 48), the δ¹³C_wood record presented here signifies that Mochras appears to record the global δ¹³C evolution.

The most dominant feature illustrated by the Mochras record is, however, the periodic appearance of medium-amplitude (~0.5 to 2‰) fluctuations on a decameter scale throughout the interval studied, superimposed on the longer-term isotopic shifts discussed above. The medium-amplitude δ¹³C_TOC fluctuations appear in a sequence of excursions that are comparable in magnitude and stratigraphic extent. They are expressed throughout the section but appear larger in magnitude in the Hettangian and Sinemurian stratigraphic interval, the deposition of which was associated with more oxygenated bottom waters.
A comparison of the Sinemurian and Pliensbachian $\delta^{13}\text{C}_{\text{TOC}}$ record of the Sancerre-Couy core (Paris Basin, France; ref. 15), and the Hettangian to lower Sinemurian record from St Audries Bay/East Quantoxhead/Kilve (Bristol Channel Basin, England, United Kingdom; refs. 4 and 17–19), which together represent the hitherto longest high-resolution $\delta^{13}\text{C}_{\text{TOC}}$ record of the Lower Jurassic, demonstrates that numerous medium-amplitude $\delta^{13}\text{C}_{\text{TOC}}$ shifts recorded in Mochras have previously been merged into what appears as stratigraphically more expanded shifts, or were missed entirely (Fig. 3). The dominant expression of the medium-amplitude CIEs in Mochras likely results from the comparatively high sedimentation rate as well as the high data resolution compared to the Sancerre-Couy record.

Several medium-amplitude CIEs observed in Mochras have previously also been identified in other $\delta^{13}\text{C}$ records covering shorter stratigraphic intervals from geographically widespread sections, and obtained from different carbon substrates, demonstrating that these shifts do not represent a local phenomenon restricted to the Cardigan Bay Basin. In the Sinemurian, for example, a positive CIE in the tumeri zone is also recorded in $\delta^{13}\text{C}_{\text{TOC}}$ on the Dorset coast of the United Kingdom (20) and in western North America (23). A negative CIE in the obtusum–oxynotum zones is also present in previously published $\delta^{13}\text{C}$ records of bulk organic matter, belemnites, and terrestrially derived palynomorphs from Lincolnshire, United Kingdom, and the $\delta^{13}\text{C}_{\text{TOC}}$ records from Robin Hood’s Bay, United Kingdom and Sancerre-Couy, France (15, 21, 24, 42). Although biostratigraphically less well constrained, a similar pattern in $\delta^{13}\text{C}_{\text{TOC}}$ has been observed in records from Italy and Morocco (12, 52). Likewise, multiple likely coeval medium-amplitude shifts in $\delta^{13}\text{C}$ within the Pliensbachian ibex and davoei zones are recorded on a supraregional scale, including the European and African Tethyan margin (9–11, 14, 15, 25, 27, 28, 48, 50, 53). Exceptionally well-expressed are multiple consecutive shifts recorded in $\delta^{13}\text{C}_{\text{carb}}$ derived from belemnites.

Fig. 3. $\delta^{13}\text{C}_{\text{TOC}}$ record of the Hettangian to Pliensbachian of the Mochras core, upper Hettangian to Pliensbachian of the Sancerre-Couy core, Paris Basin (15) and the composite Hettangian to lower Sinemurian record from the Bristol Channel Basin (4, 17–19). Key for identified ammonite subzones in Mochras given in Fig. 2. Identified ammonite subzones in the Sancerre-Couy stratigraphy: a = valdensi, b = stokesi, c = subnodosus, d = gibbosus, and e = solare. The blue line represents seven-point moving average. Carbon-isotope events identified in Sancerre Couy core are correlated with the Mochras $\delta^{13}\text{C}_{\text{TOC}}$ record. Note the differences in stratigraphic resolution and stratigraphic completeness, data resolution, and how some individual medium-amplitude CIEs in the Paris Basin record appear merged into a single event when correlated to Mochras.
from Dorset, United Kingdom, throughout the Pliensbachian *jamesoni*, *ibex*, and lower *davoei* zones (11). Furthermore, multiple shifts in $\delta^{13}$C$_{TOC}$ are recorded in the upper Pliensbachian *kauea* and *carlottense* zones of eastern Oregon in the United States, correlative to the mid-*margaritatus* to *spinatum* zones of the northwestern European realm (16).

These CIEs have thus far been discussed as single or episodic events, but as aggregated here in a continuous record and viewed collectively in the overall stratigraphic context they appear as elements in a regular series. Based on the occurrence in multiple carbon substrates and in various sedimentary basins, it seems apparent that most, if not all, of the medium-amplitude CIEs recorded in the Mochras core reflect changes in the $\delta^{13}$C composition of an at least supraregional marine dissolved inorganic carbon pool.

**Pacing of Early Jurassic Carbon-Cycle Fluctuations.** The $\delta^{13}$C$_{TOC}$ data presented here point to a common and strongly repetitive, supraregionally to globally acting driving force pacing the observed fluctuations in $\delta^{13}$C$_{TOC}$ and concomitant shifts in $\delta^{13}$C$_{Carb}$. The most plausible driving mechanism acting continuously over an extended time interval is orbital forcing. Compared to shorter Milankovitch periodicities, the long-eccentricity (405-ky) orbital signal can be well expressed in $\delta^{13}$C records due to the long residence time of carbon in the ocean–atmosphere system and the associated “memory effect” of carbon in the oceans (54, 55).

Spectral analysis of the Mochras $\delta^{13}$C$_{TOC}$ dataset identified dominant spectral peaks, which are changing to slightly higher frequencies up-sequence (*SI Appendix, Figs. S2 and S3*). Average spectral misfit (ASM) testing of these dominant spectral peaks and orbital target frequencies signifies orbital influence as the likely driving mechanism behind the dominant spectral components. The periodicities corresponding to the long-eccentricity (405-ky) cycles visually match the observed medium-amplitude CIEs (*SI Appendix, Fig. S2*) and are in a similar stratigraphic range compared to dominant spectral peaks identified in elemental calcium concentrations and gamma-ray logs identified in the upper Sinemurian to Pliensbachian strata of the same core, which also have been interpreted to represent 405-ky cycles (39, 56).

Thus far, it has not been resolved how eccentricity forcing impacts the carbon cycle and the $\delta^{13}$C signature. Orbital eccentricity forcing modulates the precessional amplitude of Earth’s insolation, leading to cyclic changes in seasonal contrasts, with eccentricity maxima marked by high seasonal contrasts and short but intense “monsoon-like” wet intervals followed by prolonged dry periods, whereas eccentricity minima are characterized by more uniform precipitation patterns (57–59). High precipitation and weathering rates during the high-eccentricity wet season is associated with increased continental runoff and fluvial freshwater inputs, as well as increased nutrient and terrestrial organic and inorganic carbon transfer into the oceans, resulting in productivity blooms, a stratified water column, and bottom-water anoxia (58).

During the dry season, oxidation of terrestrial organic matter is favored on land, restored water-column mixing oxidizes marine organic matter, and regional carbonate production increases. More stable conditions during eccentricity minima lead to a constant input of freshwater, nutrients, and carbon, resulting in constant productivity rates, persistent watermass stratification, and continuous accumulation of organic-rich deposits in the ocean, as well as increased net production of terrestrial biomass and its storage in stable tropical soils, wetlands, and peats (58, 60, 61). These eccentricity-paced variations in the accumulation and remineralization of marine and terrestrial organic carbon, and the ratio between burial flux of organic carbon and the accumulation rate of inorganic (carbonate) carbon, impact the oceanic carbon pool and are sufficient to drive $\delta^{13}$C fluctuations (58, 60–63).

The negative shifts in $\delta^{13}$C$_{TOC}$ in Mochras may thus reflect enhanced preservation of isotopically light organic matter in response to orbitally paced redox conditions in the water column and/or bottom waters and sediments. The more distinct expression and larger magnitude of $\delta^{13}$C$_{TOC}$ shifts in the Hettangian to upper Sinemurian strata may be linked to the more oxygenated bottom-water conditions that were likely more susceptible to the orbitally paced redox changes. Increased drawdown of $\delta^{13}$C-enriched organic matter is, however, generally associated with positive shifts in $\delta^{13}$C. Conversely, in Mochras, increased TOC values correspond to negative isotope shifts instead (for example, in the *angularata* and *bucklandi* zones), but overall no correlation between TOC and $\delta^{13}$C$_{TOC}$ is apparent. Other Lower Jurassic geochemical records show that medium-amplitude shifts are preceded by, rather than comcomitant with, increased TOC intervals (11, 26). This offset may suggest that the accumulation of organic matter took place elsewhere, outside the Cardigan Bay Basin. Furthermore, the long-eccentricity (405-ky) cycles in $\delta^{13}$C$_{TOC}$ appear independent of the overall climatic background, as demonstrated by the isotope shifts expressed during the late Pliensbachian, which represents a climatically cold interval (8, 51, 64) and may be less conducive for the development of high seasonal contrasts.

**Early Jurassic Astronomical Time Scale.** Due to the stability of the orbital long-eccentricity cycle over the past 215 My (65), orbital tuning of datasets encoding this astronomical meteronome can provide reliable time constraints. The 405-ky tuned $\delta^{13}$C$_{TOC}$ record from Mochras shows dominant spectral peaks corresponding to amplitude modulation, short eccentricity, and obliquity Milankovitch cycles (Fig. 4). A floating time scale based on the tuned dataset implies a duration of 8.8 My for the Pliensbachian (Fig. 4), which is closely comparable to the cyclostratigraphic duration estimate obtained from elemental calcium concentrations from the same core (8.7 My; ref. 39). The tuned $\delta^{13}$C$_{TOC}$ record of the preceding stage provides a direct cyclostratigraphic duration estimate for the Sinemurian and its constituent ammonite zones (Fig. 4). If all cycles were identified correctly, the stage duration of ~6.6 My defined herein is shorter compared with the previous estimate of 7.6 My, which is based on the assumed linear decrease in $87^{\text{Sr}}/86^{\text{Sr}}$ of belemnite rostra extracted from the Belemnite Marls, Dorset, United Kingdom (66).

Some uncertainty regarding cycle allocation associated with a comparably weak spectral peak in the original and tuned dataset appears in the upper Sinemurian *osynotum* zone (Fig. 4 and *SI Appendix, Fig. S2*). In this stratigraphic interval, the potential stratigraphic break associated with the possible occurrence of a fault and/or the transition between available sample types and data resolution (transition between samples taken from preserved core slabs and reserve bags) may have led to imprecise cycle allocation. Assuming a relatively uniform sedimentation rate, and considering that the fault, if present, cuts out less than one cycle, the results most likely indicate that, if at all, not more than one cycle may be missing in this stratigraphic interval.

Controversial astronomical duration estimates are currently debated for the Hettangian Stage, ranging between 1.7/1.9 My and 4.1 My (17–19, 67, 68). The shorter estimates are derived from multiple astrochronological studies (17–19), with duration estimates from the Hartford Basin, United States, and the biostratigraphically well-constrained St Audries Bay succession, Somerset, United Kingdom, being effectively indistinguishable and supported by correlation to the geomagnetic polarity time scale, as well as radioisotopic ages for the Triassic–Jurassic boundary and earliest Sinemurian from the Pucará Basin, Peru (18, 67, 69–71). The longer estimate of 4.1 My is based on cyclostratigraphic interpretation of spliced magnetic-susceptibility
records from multiple UK locations, the individual sections of which suggest durations of 2.9 to 3.2 My (68).

It is possible that the spliced records contain unrecognized stratigraphic overlaps or wrongly identified lithological cycles and thus overestimate the duration of the Hettangian Stage. Equally, however, the individual sections studied for astrochronological constraints may be stratigraphically incomplete and therefore underestimate the time involved in their deposition (68). An additional complication is the possibility that the radiosotopically dated volcanic ashes from the lower Sinemurian in the Pucará Basin in Peru are reworked or correlated inaccurately to the European base-Sinemurian stratotype (68). Temporal constraints on the Hettangian–Sinemurian boundary are thus not entirely resolved.

Although the Hettangian δ¹³C_TOC record in Mochras shows consistent medium-amplitude fluctuations similar to the overlying strata, which can be interpreted to represent long-eccentricity forcing, tuning of the Hettangian δ¹³C_TOC record in Mochras cannot in itself resolve the duration of the Hettangian Stage as the strata are most likely stratigraphically incomplete at the Triassic–Jurassic boundary.

Biostratigraphically calibrated temporal constraints on the Triassic–Jurassic boundary (201.36 ± 0.17 Ma, recalculated from refs. 70 and 72) are supported by absolute age constraints on the preceding end-Triassic extinction event (73) and the Newark–Hartford astrochronology and geomagnetic polarity time scale (69). The Pliensbachian–Toarcian boundary projected age of 183.7 ± 0.5 Ma is supported by biostratigraphically constrained U-Pb ages from the lower Toarcian (ref. 75, corrected in refs. 76 and 77) and upper Pliensbachian (16, 78). Based on these absolute ages bracketing the studied time interval, the combined Hettangian, Sinemurian, and Pliensbachian Stages cover some 17.7 ± 0.5 My. According to the cyclostratigraphic duration estimates for the Sinemurian and Pliensbachian (~6.6 and 8.8 My, respectively) given here, the Hettangian Stage is constrained to a duration of ~2.3 ± 0.5 My. Despite uncertainties, this time scale strongly supports the nonspliced cyclostratigraphic duration...
estimates for the Hettangian Stage from individual UK sections (17–19, 68).

**Conclusions**

The Mochras δ13C TOC data from the uppermost Rhaetian to Pliensbachian interval, combined with available data from the Toarcian from the same core, provides a continuous, biostratigraphically well-defined, high-resolution chemostратigraphic record for the Lower Jurassic. Beside large-scale CIEs (>3‰) known from isotopic records elsewhere, CIEs of generally smaller magnitude (0.5 to 2‰) occur throughout the Hettangian to Pliensbachian interval. These medium-amplitude CIEs, including shifts that have previously been recorded in stratigraphically shorter intervals, appear more singular in the context of a continuous record. Spectral and ASM analysis of the data reveals that these medium-amplitude CIEs are paced by long-eccentricity (405-ky) cycles, implying the impact of orbital forcing on the ocean-atmosphere carbon reservoir. Orbital tuning of the isotope record provides a duration estimate of 8.8 My for the Pliensbachian and offers an estimate for the Sinemurian Stage (6.6 My). Combined with published biostratigraphically defined radiocarbon age constraints for the Triassic–Jurassic and Pliensbachian–Toarcian boundaries, the data presented herein suggest a duration for the Hettangian Stage of ~2.3 ± 0.5 My.

**Materials and Methods**

Preserved core slabs, bagged core fragments known as the "reserve collection," and registered specimens of the Mochras drill core are housed at the British Geological Survey National Geological Repository at Keyworth, United Kingdom. For this study, bulk rock samples between 1,290 and 863 mbs were collected from well-preserved core slabs at a 30-cm resolution. Bulk-rock samples below the ~1,290 mbs level were largely sampled from reserve collections, each of which aggregates ~1.4 m intervals of broken core. A single sample was taken from each bag and referred to the depth of the midpoint of the sampled interval (reserve bag samples marked as white squares in Fig. 2). Macroscopic fossil plant material was extracted from reserve bags only. The sample resolution, ammonite and foraminiferal biostratigraphy, and a lithological log of the Mochras drill core are shown in **SI Appendix**, Fig. S1.

**Data Availability Statement**

All data discussed in the paper will be made available in the **SI Appendix and Dataset S1**.

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