Assessment of local pollution influence on the weather and impact on the air quality over subtropical southern Africa

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Abstract—This study discusses the possible role of air pollution and influence on air quality (AQ) and weather over southern Africa. Although it has long been thought that a lot of pollution occurs in the industrialized regions of the Northern Hemisphere, it is quite clear that a vast amount is also generated in the south, and southern Africa contributes significantly. Since industrial revolution, the sub-continent has become one of the major sources of atmospheric emissions mainly from human-induced activities to meet demand for energy supply and other lifelong needs. The main sources include biomass burning (BB), Aeolian dust and industrial emissions. In the process, trace components generated can negatively impact the atmosphere through aerosol influence on cloud properties and radiation budget, and also affect human health. While anthropogenic emissions are enhanced during winter, natural emissions occur and fluctuate throughout the year. In the case of dust aerosols, they are common over the Kalahari-Namib Desert where dust devils frequently develop overland, keeping the air filled with haze. Common gas pollutants include carbon monoxide (CO), sulphur-and nitrogen oxides (SOx and NOx) from the copper belt (Zambia), Highveld (South Africa), BB-dominated areas and other isolated locations.

1.2 Pollution, weather and air quality
Pollution can have direct and indirect effects (e.g. acidification, eutrophication, stratospheric ozone depletion) with a wide range of impacts on human health, ecosystems, agriculture and air composition. It can be influenced by several factors: weather (or meteorological) components such as wind over/near an emitting source region, chemical transformations in the air, polluting sources (some sources are stationary while others are mobile) and transport mechanisms. To date, industrialization has impacted on the air quality (AQ) through emission of gases and particles and is also attributed to climate change. The weather and AQ are linked through human health as people can be exposed to higher concentrations of air pollutants and suffer disproportionately from effects of deteriorating AQ. For example, people in large populated cities are often exposed to high levels of pollution and can develop respiratory ailments that hinder their comfortable lives (Schwela, 2012). Aerosol pollution may aggravate heart and lung disease and is associated with heart attacks and
cardiac arrhythmias, difficulty in breathing (Jhun et al., 2015, and references therein), and makes people more susceptible to respiratory infections. Just like weather, pollution can change hourly or even daily (USEPA, 2003) from one place to another. AQ pollutants may also have climatic effects; depending on the pollutant, which may be warming or cooling, or a combination of the two (AQEG, 2007). The main cause of air pollution by humans is use of fossil fuels, power generation, industrial and other domestic operations (e.g. burning of firewood, agricultural activities and waste). Because of wide-ranging air pollution problems, governments and other organizations such as the World Health Organization (WHO) put in places some governing standards and regulations to monitor the amount of emissions in the atmosphere. AQ guidelines provide an assessment on health effects due to air pollution; threshold limits are set to ensure that the total average concentrations of pollutants do not exceed certain range or levels within which human can be exposed to.

The purpose of this study is to highlight the role of regional pollution, and possible influence on the weather and AQ over southern Africa. We are motivated by the fact that the sub-continent experiences a significant amount of pollution in addition to its varying climatic conditions. The rest of the paper is arranged as follows: section 2 gives an overview of the climate and pollution over southern Africa; section 3 is the methodology and data acquisition; section 4 focuses on results analysis. Lastly, section 5 is the conclusion on the state of AQ and pollution.

II. CLIMATE OUTLOOK AND POTENTIAL DRIVERS

2.1 Climate pattern

Southern Africa is prone to frequent droughts and uneven rainfall distribution with two distinct seasons (a hot-wet summer and cool dry winter) that significantly determine and control its climate. The most influential factors include the Inter Tropical Convergence Zone (ITCZ), ocean currents, and quasi-stationary high-pressure systems (St. Helena High in the South Atlantic and Mascarene High in the Indian Ocean). Much of the rainfall received between November and February—the period corresponding to the southward displacement of the ITCZ in the Southern Hemisphere (SH) (maximum displacement occurs in January), whereas the dry season occurs when the ITCZ retreats northward (maximum displacement in July). ITCZ annually changes its position on either side of the equator, moving between the Tropics of Cancer and Capricorn (Fig. 1). Its position shifts the belts of planetary winds and pressure systems to the north and to the south annually.

The South Atlantic and Indian Oceans also play an important role in the climate. For example, the east coast is influenced by the southward-flowing Mozambique Current (bringing warm water and humid air from the equator, creating a humid warm climate). Contrastingly, the west coast is influenced by the cold Benguela Current from the Atlantic Ocean—which produces a drier climate. There is a strong rainfall gradient over the interior, running from east to west, mainly occurring during summer in the form of thunderstorms. High daily and
seasonal temperature ranges are also observed as a result of altitude and interior “continental” location (i.e. lack of ocean influences). Total rainfall gradually decreases westward; much of the central and western regions are semi-desert with low and unreliable rainfall. An exception is the southern and western Cape regions of South Africa (this is influenced by maritime conditions and receive winter rainfall as part of a more temperate climate). These parts receive their winter rainfall due to cold fronts rolling from the Atlantic Ocean (Hudson and Jones, 2002) when they are embedded in the westerly wind regimes (Obasi, 2001).

2.2 Transport climatology
Air transport climatology over subtropical southern Africa is classified as daily synoptic situations into predominant circulation types. Percentage zonal transport in easterly and westerly directions; total transport is a function of circulation type and frequency, as well asplume dimensions (Tyson et al., 1996a,b). Two major pathways—the Natal plume (eastwards) and the Angolan plume (westwards)—are the main carriageways of pollution on land.

A typical example is the so-called “river of smoke” outflow from the mainland into the Indian Ocean (Swap et al., 2003). The semi-permanent subtropical continental anticyclones, transient mid-latitude ridging anticyclones and mid latitude westerly disturbances produce major transport into the southwestern Indian Ocean along the Natal plume, whereas the quasi-stationary tropical easterly waves result in appreciable transport into the tropical South Atlantic Ocean through the Angolan plume.

Detailed circulation patterns over southern Africa is given by Tyson et al. (1996a) and Garstang et al. (1996), who described for example, four major circulation types occurring with different frequencies: semi-permanent subtropical anticyclones, transient mid-latitude ridging anticyclones, westerly baroclinic disturbances and barotropic quasi-stationary tropical easterly waves. According to the authors, transport in ridging highs and westerly perturbations are much less and occur throughout the year, with a slight tendency peak in spring.

Monthly, seasonal and annual mass fluxes over and out of southern Africa form transport fields with substantial aerosol concentrations.

2.3 Potential climate drivers

2.3.1 El Niño southern oscillation
El Niño–Southern Oscillation (ENSO) is an irregularly periodical variation in winds and sea surface temperatures (SSTs) over the tropical eastern Pacific Ocean, affecting much of the tropics and subtropics. The Southern Oscillation is an accompanying atmospheric component, coupled with SST change. The warming phase (accompanied with high surface pressure) is known as El Niño, while the cooling phase (accompanied with low air surface pressure in the tropical western Pacific) is known as La Niña. The two phases can last for several months and their effects vary in intensity. El Niño events often begin in the middle of the year with large-scale warming of surface water in the central and eastern equatorial Pacific Ocean and changes in the tropical atmospheric circulation (WMO, 2014). Generally, El Niño events occur every two to seven years and can last up to 18 months; they peak during November–January and decay in the first half of the following year. Strong and moderate El Niño events have a warming effect on the average global surface temperatures (WMO, 2014). Over southern Africa, El Niño usually results in less rain, and normally drought occurs particularly during the critical agricultural period of October to December, and is a major driver to climate variability affecting rainfall (Nicholson and Entekhabi, 1987). For example, taking into account the magnitude of negative impacts in agricultural areas world wide, El Niño of 1992 was classified as the most severe event in 30 years (FOA, 2014). As a result, the economic sector is likely to be affected noting that more than half of the region’s population depends on agriculture.

2.3.2 Sea surface temperatures
SSTs define the predictability of climate in sub-Saharan Africa. This is warm (cool) in the southwest Indian Ocean and cool (warm) in the southeast Indian Ocean; increased (decreased) summer rains may occur over large areas of southeastern Africa. For example, El Niño is the dominant pattern of variability at global scale associated with droughts in the northern and southern hemispheres (Giannini, 2010). SST poles are reversed in sign, decreased precipitation occurs over southeastern Africa as a result of increased low-level divergence or low-level flow and this flow being drier than average.

2.3.3 Monsoon effects

Monsoon is traditionally defined as the seasonal reversing wind accompanied with corresponding changes in precipitation. Two main components of the African monsoons are the West African monsoon (prevail during the Northern Hemisphere (NH) summer (June–September)), and the East African monsoon (common during spring (March–May) and autumn (October–December)). A combined influence of the Indo-Pacific and the Atlantic Oceans drive the inter-annual and the decadal monsoon variability over these regions. Key features of the West African monsoon are the low level southwesterly flow from the Atlantic Ocean and the ITCZ north of the equator. For example, the West African Sahel is well known for its severe droughts that ravaged the region during the 1970s and 1980s (Nicholson, 2013). East African monsoon is associated with the ITCZ.
moving south of the equator. The land-sea contrast forces the summer northern movement of the ITCZ to be displaced further than it would otherwise be (Bigg, 2003). Long-term precipitation changes from both northern and southern regions of Africa are linked to the monsoon circulations and controlled by precessional variation in summer insolation (Schefub et al., 2005, and references therein). The so-called long rains prevail during spring and the short rains during autumn. The transition (equinox) between seasons brings most rainfall to East Africa.

2.4 Sources of air pollution

There are three major air pollution sources defined over Africa: eolian dust, biomass burning and industrial (anthropogenic) emissions (Fig. 2), contributing to the overall global emissions (Piketh and Walton). However, several other common sources also contribute to the overall air pollution over southern Africa, including urbanization, motorization, economic activity, fossil fuel usage and open burning (including vegetative fires and waste burning).

2.4.1 Biomass burning

While domestic fuel (e.g. wood and coal) is widely used in most parts of Africa, the largest source of BB comes from vegetation (Swap et al., 2002; 2003). Human-induced BB is the main source of aerosols in the SH (Eck et al., 2003; Torres et al., 2010) contributing over 86% of the total global emissions of black carbon (soot) - the largest source being the African savannah (Skaeda et al., 2011, and reference therein). BB dominates the chemical atmospheric burden of emissions over most parts of Africa (Crutzen and Andreae, 1990; Liouss et al., 2010; Scholes, et al., 2011), especially during the dry season (Matsueda et al., 2002; Li et al., 2003; Chédin et al., 2005; Magi, 2009). This contributes about 30% of global burning emissions (Smith et al., 2001; Pikethand Walton, 2004), followed by wind-blown dust and industrialemisions respectively. Savannah fires are the global largest source of BB mostly in the tropics (Andreae et al., 1996; Seinfeld and Pandis, 2006), contributing about 75% of all related fire emissions (Piketh and Walton, 2004) with more than 60% of the earth savannah coming from Africa (Li et al., 2003).

The bulk of burning occurs in tropical central Africa (Fig. 3), compared to the much drier north and south extremities. Intense burning reaches its peak around May–August and drops towards summer (September–October) - marking the beginning of the wet season (Piketh et al., 2002; Tummon et al., 2010, and reference therein). Figure 3 shows BB intensity over Africa in 2005, indicating the seasonal fire patterns. The images are based on fires detected by the Moderate Resolution Imaging Spectroradiometer (MODIS) from the National Aeronautics and Space Administration (NASA)’s Terra and Aqua satellites. Each image is a 10-day composite of fire detections (marked in red and yellow); the series include images from every other 10-day period from 1 January through June to 18 August 2005. Red colour indicates locations with few fires detected during the 10-day period whereas yellow indicates many fires detected over an area.

2.4.2 Industrial/Anthropogenic emissions

The most significant industrial trace gases emitted over southern Africa include sulphur- and nitrogen oxides (SOx and NOx), carbon monoxide (CO) and other gas compounds. Most of the gas emissions are concentrated over the Mpumalanga highveld (South Africa), the copper-belt (Zambia) and other isolated points such as the Bamangwato Concessions Limited (BCL) mine in Botswana. However, BCL is currently not operating, after being placed under liquidation in 2016. Also, there is a large amount of aerosol pollution generated from various source sectors such as cement and timber processing, construction/manufacturing sites, mining and other operations across the sub-continent.
Some studies (e.g., Benkovitz et al., 1996; Meter et al., 1999) had proposed that emissions from industrial sources contributed up to 2.24 Mt per annum. For example, it was estimated that the spatial distribution contributed a total of 1.1 Mt of sulphur emitted into the regional atmosphere annually (Sivertsen et al. (1995). Out of this, 66% was said to come from South Africa alone (Piketh et al., 1999), from which about 90% was from the Mpumalanga’s highveld area (Sivertsen et al., 1995; Wells et al., 1996). South Africa is the most industrialized country in southern Africa (Piketh and Walton, 2004), and one of the world’s most carbon intensive countries with per capita emissions higher than most European countries (DEA, 2014). The Cu-Ni and smelter plant in Botswana has been one significant source of sulphur emissions (before its closure), accounting for 0.19 Mty⁻¹. During winter (when anthropogenic emissions increase) it would be expected that industrial and BB emissions become well mixed over land.

2.4.3 Aeolian dust

Desert dust is particularly the main aerosol component in many arid and semi-arid regions (e.g., Sahara and sub-Saharan Africa: Zakey et al., 2006). These emissions (emitted through suspension, saltation and creeping processes) occur mostly in episodic events determined by threshold near-surface wind conditions. Naturally, dust aerosols originate from the local bare soils (devoid of vegetation or grass). Dust can be a regional scale climatic forcing agent when vast dust plumes are suspended into the atmosphere. This is particularly true for regions where dust plumes are suspended into the atmosphere and transported over hundreds to thousands of kilometers. Large amounts of mineral dust can be transported across, which in combination with other emissions lead to persistent haze because of lack of wet removal (Knippertz et al., 2015). The fine dust particles can be lifted up to higher altitudes and transported over long distances away from their source regions. Their effects can be felt not only locally, but also at places/regions far away from their source regions (Zakey et al., 2006).

Africa has vast tracks of desert dust aerosol emissions, mainly from the Sahara Desert (North Africa) and the southwest coast of Namib Desert (southern Africa). Strong winds known as “berg winds” (mountain winds) can loft plumes of dust directly out into the Atlantic Ocean depending on the wind speed and direction (Piketh and Walton, 2004).

2.4.4 Automobiles and othersources

The rapid, unplanned and uncoordinated town/city growth has negatively impacted population centers in different regions, resulting in many people moving between the
city centers to urban area outskirts. High population increases together with pollution levels seriously compromise existing transportation systems and significantly increase the challenge of creating future transportation systems (Road management, 1998). Population growth naturally influences car ownership (privately owned). For example, motor vehicle fleets have recently increased in countries such as Botswana and Zimbabwe (Simukanga et al., 2003), due to imported second-hand vehicles no longer meeting certain standards in their countries of origin (Wiston, 2017). These imported vehicles are not properly maintained or checked on how they emit pollutants. This increase in number of vehicles results in an increase in fuel consumption. For example, in city centers where traffic congestion can be responsible for 90–95% of CO and 60–70% of NO\(_x\) (Schwela, 2004). NO\(_x\) (where NO\(_x\) = NO + NO\(_2\)) and hydrocarbons pose a major threat to human health and natural resources (Chanda, 2014). These emissions can contribute to photochemical smog, especially in areas experiencing high traffic density such as central business districts (CBDs). The escalating emission rates are also exacerbated by road congestions, poor vehicle maintenance and high average age of the vehicle fleet.

2.5 Weather, pollution and air quality

Air pollution has a significant impact on climate and air composition in many ways. For example, GHGs in the atmosphere can trap heat and contribute to global warming. These gases (e.g., carbon dioxide (CO\(_2\)), ozone (O\(_3\)), methane (CH\(_4\))) once emitted can remain in the atmosphere for decades to centuries. They absorb and emit radiation within the thermal infrared range. Their concentrations can become well-mixed throughout the global atmosphere depending on the prevailing meteorology, regardless of location of emission source and their effects on climate could be long lasting (USEPA, 2009). Also, particles such as black carbon (BC) and nitrates absorb and scatter radiation, therefore affecting the radiative balance. Similarly, the weather can affect air pollution; it can either raise the development of pollution or stop its development depending on the environmental conditions such as in tropical areas where precipitation often occurs. It is more likely that pollutants would be washed out by precipitation (Pachauri, 2007). Changes in atmospheric gas- and aerosol concentrations, land cover and solar radiation alter the energy balance of the climate system (IPCC, 2007) and are drivers of climate change. The changes in energy balance due to these factors are expressed as radiative forcing—a process used to compare the warming or cooling influence on global climate.

In principle, any pollutant that contributes to local/regional pollution and act as a radiative forcing agent or changes solar distribution may potentially produce a linkage between AQ and climate change (AQEG, 2007). Changes in weather and climate can affect AQ as different areas experience varying climatic conditions and different pollutant concentrations. For example sunshine, rain, precipitation and wind can affect the amount of air pollution in an area, leading pollution washout, heat absorption or pressure variation in the atmosphere. Photo-energy makes some pollutants undergo chemical reactions in the atmosphere; higher temperatures and strong winds can also speed up chemical reactions in the atmosphere, rain washes out water-soluble pollutants. The wind can carry air contaminants away from their source, and distribute them to other areas. Higher (strong) winds result in fast dispersal of contaminants and increase their mixing in the atmosphere; the boundary layer may also change due to mixing.

III. METHODOLOGY

3.1 Data acquisition and sources

Meteorological and chemical data is obtained from several databases: National Oceanic and Atmospheric Administration (NOAA) (www.esrl.noaa.gov/psd/data/composites/hour/), the geostationary Dundee Satellite (http://www.sat.dundee.ac.uk/satellites.html) and the NASA’s Giovanni (www.giovanni.gsfc.nasa.gov/Giovanni/service) for wind field, temperature, precipitation, clouds and chemical emissions over the study domain. NASA’s Giovanni (Geospatial Interactive Online Visualization and Analysis Infrastructure) is a web-based application developed by the Goddard Earth Sciences Data and Information Services Center (GES-DISC) to provide simple and intuitive way to visualize, analyze, and access vast amounts of Earth science remote sensing data. Satellites provide quantitative data, images (maps) showing spatial distribution and pollutant concentration over an area.

We focus on the summer (Nov–Jan) and winter (May–July) seasons over the 10-year period (2000–2010) looking at the trend of pollutants and climatology over the sub-continent. Specifically, we look at emission scenarios in 2000, 2005 and 2010 at 5-day intervals - both surface and temporal distributions. Pollution patterns are then discussed in relation to the meteorological parameters over the same period and locations in order to understand how the weather interacts with the chemistry. For example, the period 2000–2010 experienced more constructions and industrial activities over southern Africa. This period also falls within the time during which many African countries were preparing for the world
societal tournament (2010 World Cup), which took place between June and July 2010 in South Africa. Most activities were aimed at attracting tourists and investors to the African continent.

IV. RESULTS AND DISCUSSION
We present and discuss outputs for meteorological and chemical variables over the domain as highlighted above. First we start with the meteorology before the chemistry and later look at their interactions. Because the chemistry is driven by the meteorology one needs to first understand and appreciate the dynamics in the meteorology.

4.1 Meteorology (weather pattern)
(a) Wind speed and direction

High wind speeds occur in summer especially over the Atlantic and southern part of the domain. For example, summer speed reach over 11 m/s in 2000 and 9 m/s in winter, south of Madagascar (Fig. 4(a) and (b)). The winds were mostly northeasterly coming from the Indian Ocean during summer and strong southerly over the west coast of Namibia, whereas in winter they are mostly easterly and northeasterly in the south. The winter of 2005 had relatively low winds than in summer. Summer maxima reach over 12 m/s while winter maxima was
below 11 m/s. The pattern is almost similar but more clearly defined in 2010 and the summer of 2005 (Fig. 4(c–f)). This indicates the possible direction of pollution transport, especially in the south where winds are westerly. The pattern also shows that the center of strong winds over the Atlantic Ocean slightly moved away from the sub-continent (southeasterly) between 2000 and 2010 during summer, and a new center developed from the east.

(b) Temperature
Next we show mean temperature distribution (Fig. 5). Again there are similar patterns in temperature as observed for wind. Summer temperatures were higher over the Kalahari, extending to the northeastern part of Namibia and central Zambia-Malawi, reaching 309 K while most of the landmass experienced temperatures between 290 K and 306 K. Winter had the temperature maximum of just over 300 K over central Angola while the rest of the sub-continent remained cooler, especially over the oceans. Higher temperatures were observed in the summer of 2005 (Fig. 5(c)) reaching over 307 K over Zambia and the southern part of Kalahari, while winter temperatures (Fig. 5(d)) had a maximum of just above 300 K over Angola and average of 250 K over much of the southern part. The summer of 2010 (Fig. 5(e)) had its temperature maxima over Botswana-Namibia, while winter had maximum temperature over northern Angola/southern Democratic Republic of Congo (DRC).

c) Precipitation and clouds
We also show accumulated precipitation (Fig. 6), where we observe higher precipitation during summer (maximum of 40 kg/m²) over southern and northeastern Mozambique and west of Angola. Lowest rainfall (below 15 kg/m²) was experienced over the western part of South Africa and southwestern tip of Namibia. The highest rainfall in winter was 32 kg/m² over eastern Mozambique coast with low rainfall across Lesotho and southern parts of South Africa (Fig. 6(b)). More precipitation occurred over the Mozambique coast (Fig. 6(c)), reaching well over 45 kg/m² whereas winter was mostly dry, with the highest rainfall around 17 kg/m² over the Indian Ocean and northern part of Mozambique coast.

Fig 5: 1000-mb temperature during summer and winter in 2000 ((a), (b); 2005 ((c), (d)); 2010 ((e), (f)) respectively. Contours indicate temperature variation and scales are colour-coded in terms of temperature range. Output between 05–10 December and June. Source: NCEP/NCAR reanalysis observation.
The summer of 2010 also shows precipitation almost over the same location as in 2005 (about 40 kg/m$^2$ between Mozambique-Madagascar), while the winter was drier over much of the landmass. Overall, it appears that summers accumulate more precipitation over the Indian Ocean and west of Angola.

Clouds are one of the important factors in weather, because they are directly linked to pollution through aerosols. To better understand aerosol–cloud effects, one needs to know whether there are clouds over the domain and where they are located. Figure 7 shows satellite imagery (from the Dundee satellite station) in 2005 and 2010.

Satellites provide estimates of area coverage of clouds by the cloud-top temperature and the column-integrated optical thickness. They maintain an up-to-date archive of images covering the globe. The summer of 2005 (Fig.7(a)) had substantial cloud cover, mostly over the central-towards-east and some cloud patches over the Atlantic Ocean. Much of the domain was clear during winter, with few cloud bands over the oceans. Similarly, the summer of 2010 (Fig.7(c)) had more clouds concentrated mostly over the northern part of southern Africa and marine environments. While the winter was similar to that of 2005, few cloud patches appear over the South Atlantic Ocean.
Fig 7: Accumulated column integrated clouds over Africa between summer and winter in 2005 ((a), (b)) and 2010 ((c), (d)) respectively. Output between 05–10 December and June. Source: Meteosat infrared satellite images (http://www.sat.dundee.ac.uk).

4.2 Emissions (pollution)

(a) Dust

Now we look at chemical emissions, focusing on dust (aerosols), NO\textsubscript{x}, SO\textsubscript{2} and CO (gases). These are primary pollutants (i.e. directly from the sources) and are shown for the same time periods as the meteorology discussed above for consistency. Figure 8 shows dust concentration obtained from the NASA’s Giovanni database. More dust appears during summer along the southwest coast (Namib Desert) throughout the years, and a significant amount over DRC in 2000. One also notices an increase in the east along the Kenyan coast during the winter of 2005 (Fig. 8(d)). The winter of 2010 shows dust over the Atlantic, along Namibia-Angola coasts. Dust emissions are generally common along the west (Kalahari-Namib Desert) than other geographical locations. However, that is so expected for this part of the globe because region is drier than its eastern counterpart—the so-called Kalahari Transect (Shughart et al., 2004).
Fig. 8: Dust concentrations between (a) summer and (b) winter respectively. Contours indicate increasing particle concentration and scales colour-coded in terms of concentration maxima. Source: MODIS onboard the Aqua satellite measurement (http://giovanni.gsfc.nasa.gov/giovanni/).

(b) Sulphur and nitrogen oxides
Next we show the spatial distribution for SO$_2$ (Fig. 9), from the same database. As expected, more SO$_2$ is observed over/around industrial locations highlighted earlier (see Sect. 2.4). SO$_2$ is an air pollutant often associated with combustion in power generation, fossil fuel refining and ore smelting (Zunckel, et al., 2000) such as the mines. About 1.8 million tons or more of SO$_2$ is emitted from electrical generation annually. For example, the annual ambient SO$_2$ concentration from South Africa, on average, is said to approach 20 ppb guideline (WHO; Kgabi, 2012). These are quite clear in the southern part—with more industrial operations. Typical locations include the copper-belt, highveld and other isolated small plants.
Similarly, we can show another common gas pollutant NO\textsubscript{x} (herein NO\textsubscript{2}) to compare with the SO\textsubscript{2} (Fig. 10). NO\textsubscript{x} is an important primary pollutant produced from anthropogenic and natural sources (e.g. fossil fuel combustion, BB, soil bacteria, lightning, oxidation of ammonia, mobile transport and other heavy machinery). While SO\textsubscript{2} is from the MODIS-on-board the Aqua satellite measurement, NO\textsubscript{x} is a snap shot from satellite imagery in 2006. Again hot spots (column density order of 10\textsuperscript{16} mol/cm\textsuperscript{2}) for NO\textsubscript{2} are observed over the Highveld and Angola-DRC; concentration patterns show how emissions spread to other areas from the source locations.

(c) Carbon monoxide

Next we show the spatial distribution of CO (Fig. 11) also from the NASA’s Giovanni database. The concentration increases (more reddish) in the top-left corner (west) into the South Atlantic Ocean, but declines towards the east. Again concentration decreases towards the south. While the pattern is similar throughout, the summer of 2005 (Fig. 11(a)) has low CO concentration than other panels. However, the summer of 2010 (Fig. 11(c)) shows more wide emission coverage and some isolated hotspots over Madagascar.

It is also worth noting that concentrations increase mostly over the BB areas (tropical central) than anthropogenic-dominated regions. One distinct difference between the spatial distribution of the emissions in dust and COs that gases are more spread than aerosols.
4.3 Meteorology-chemistry interactions and influences

The atmosphere over southern Africa has unique meteorological and chemical characteristics that influence pollution transport and climate forcing (Garstang et al., 1996). Emission distribution is strongly influenced by the transport dynamics and air masses that prevail over most of the time (Kirkman et al., 2000). These circulation systems and transport pathways play vital roles in the local meteorology. For example, the frequent high-pressure systems act as ‘accumulation’ mechanisms for trace gas and aerosols associated with haze layers and the westerly wave passages are responsible for the ‘clean out’ signatures (Stein et al., 2003). There is an indication that emissions have increased over the years, more pronounced during summer, although sometimes during winter (e.g. in 2005 there was a substantial increase in wind speeds).

Consequently there is a decline in precipitation with very little or no rainfall over the coastal areas. Temperatures along the west coast (between 2005 and 2010) appear to be cool in summer and become cooler in winters. It would of interest to try to understand all potential climate controls as highlighted earlier. Satellite images indicate more cloud coverage in summer, correlating with the rise in dust aerosols (more aerosols enhance the chance of cloud formation because they act as sites for developments). However, more clouds (precipitation) may result in pollution washout from the atmosphere, leading to pollution removal (wet deposition), whereas strong winds can also carry out pollution (dry deposition) from the atmosphere or away from the source. Urban pollution can enhance down welling radiation during clear nights and lead to increase in night-time minimum temperatures when warm air is mixed from aloft because of radiative destabilization (Knippertz et al., 2015). However, these kinds of interactions between weather parameters and chemistry are cyclic between the seasons.

We also show temporal emission distribution for dust, \( \text{SO}_2 \) and \( \text{CO}_2 \) concentrations from 2000 through 2010 (Fig. 12). The trends show fluctuating patterns in pollutant concentrations across all the species. As expected, CO has the highest concentration than \( \text{SO}_2 \) and dust. Global pollutants’ emissions increase mainly driven by the demand for energy and transportation. As highlighted earlier, the changing meteorological patterns and seasonal human activities play vital roles in pollution levels. For instance the summer winds in 2000 show wind divergence over the Mpumalanga Highveld while the winter \( \text{SO}_2 \) concentration appear to be spread in all directions but mostly directed towards the Mozambique coastal areas.
Common air circulations during winter (mostly westerly) also have the potential to carry the SO$_2$ emission observed over the highland and further distribute it across. Between 2005 and 2010, wind patterns become more intense in summer, this could possibly scatter emissions further. Sulphate aerosols can also act as sites for cloud formation (cloud condensation nuclei [CCN]), although clouds formed around these compounds have varying radiative properties than other CCN. Increases in sulphate abundance may lead to more clouds or increase cloud albedo, and cause a net reduction in solar radiation reaching the earth surface (AQRS, 2001). An increase in the rainfall intensity in summer would imply more pollution washout from the atmosphere, especially in the proximity of polluting areas. On the other hand, frequent winds in winter increase likelihood of pollution transport and/or result in the build up of pollution (haze) above. This can alter radiation budget (depending on the intensity of absorbing or scattering aerosols present), or impact on cloud microphysical properties.

V. CONCLUSION

Southern Africa is characterized by the tropical and subtropical climatic conditions with great influence from the adjacent oceans, as well as the common air circulations. Since the late 20$^{\text{th}}$ Century, the sub-continent has transformed from a rural to a complex society and has made great strides towards industrialisation, urbanisation, and economic development (Simukanga et al., 2003). These changes, compounded with population growth, brought about environmental problems virtually non-existent in the past. Owing to different major pollutant sources, the region has experienced a considerable change recently and is recognised as a major source of pollutants (Wiston, 2016). Emissions exhibit a mixture of large quantities of aerosols and gases from BB, domestic fires, fossil fuel sources [automobiles and industries], construction, manufacturing and Aeolian dust. While the SH summer is often characterized with more cloud coverage and wetness, winter is characterized with common air masses blowing over land. It is quite clear that pollution can influence the AQ composition over a given location. Consequently, some of the areas experiencing high levels of air pollution are the most populous and/or capital cities. Although there might not be national AQ problems on record, a number of air pollution conditions exist where severe AQ problems are likely to occur; examples being locations within active BB and high anthropogenic emissions. Vast amount of pollutants can be transported over long distances, impacting other areas downwind or those initially not in the vicinity of emitting sources. Under these conditions, humans are often exposed to high levels of pollution, which is not always monitored and sometimes exceed.
threshold limits (Wiston, 2017). Bad AQ cannegatively impact human health and may lead to respiratory problems such as heart failure, asthma, cardiac arrests, congestive heart failure and eye irritations. Anthropogenic and human-induced BB emissions increase mostly in winter when industrial activities are in operation and more vegetation burnt for various purposes. On the other hand, natural emissions are generated and fluctuate throughout the year. It is worth noting that southern African pollution plays a significant role on the weather and regional air quality composition.

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