Diversity of large ornithopod dinosaurs in the upper Hauterivian-lower Barremian (Lower Cretaceous) of Teruel (Spain): a morphometric approach

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ABSTRACT

We report some large ornithopod vertebrae from two upper Hauterivian-lower Barremian (Lower Cretaceous) localities in El Castellar (Maestrazgo Basin, Teruel, Spain). These fossils have been studied systematically as well as morphometrically using a multivariate analysis in order to analyse the diversity of the sample. In fact, principal component analysis has been demonstrated as an useful tool for establishing affinities in isolated iguanodontian vertebrae, at least when size effect is not removed from the analysed dataset. As result of this study, two large indeterminate styracosternans are distinguished in the sample: a large one with platycoelus anterior caudal vertebrae related to the genera Magnamanus and Iguanodon, and a middle-sized one with longer-than-high dorsal vertebrae and amphicoelus anterior caudal vertebrae related to Morelladon. Such diversity of large ornithopods observed in the upper Hauterivian-lower Barremian of the Maestrazgo Basin is similar to that previously observed in western Cameros Basin and it demonstrates the presence of at least two different forms of styracosterman in this stratigraphic range of the Iberian Peninsula.

RESUMEN

En este trabajo describimos y analizamos varias vértebras de grandes ornitópodos procedentes de dos yacimientos situados en el Hauteriviense superior-Barremiense inferior (Cretácico Inferior) de El Castellar (Cuenca del Maestrazgo, Teruel, España). Los fósiles se han estudiado tanto sistemáticamente como morfométricamente empleando un análisis multivariante en el último caso, con el fin de analizar la diversidad de la muestra. De hecho, el análisis de componentes principales realizado ha demostrado ser una herramienta útil para establecer afinidades en vértebras aisladass de iguanodontios, al menos cuando el efecto del tamaño no se elimina del conjunto de datos analizados. Como resultado, dos grandes estiracosternos indeterminados se distinguen en la muestra: uno grande, relacionado con los géneros Magnamanus e Iguanodon, con vértebras caudales anteriores platicélicas, y otro de tamaño mediano, relacionado con Morelladon, con centros vertebrales dorsales más largos que altos y vértebras caudales anteriores anfícelicas. Esta diversidad de grandes ornitópodos en el Hauteriviense superior-Barremiense inferior de la Cuenca del Maestrazgo es similar a la observada en el margen occidental de la Cuenca de Cameros y demuestra la
Keywords: Peñagolosa sub-basin, El Castellar Formation, Ornithopoda, Styracosterna, Principal Components Analysis.

1. INTRODUCTION

The fossils of large ornithopods are the most abundant among those belonging to dinosaurs in the Lower Cretaceous of the Iberian Peninsula (Pereda-Suberbiola et al., 2012; Gasca et al., 2014; Gasulla, 2015; Verdú, 2017; Alcalá et al., 2018). Based on current knowledge, the diversity of this type of dinosaur (all of which are considered Styracostera) in this stratigraphic range is composed of Magnanamus soriaensis Fuentes et al., 2016 in the upper Hauterivian-lower Barremian of Soria province (Fuentes et al., 2016); Iguanodon galvensis Verdú et al., 2015 in the lower Barremian of Teruel province (Verdú et al., 2015); Iguanodon bernissartensis Boulenger in Van Beneden, 1881, Mantellisaurus atherfieldensis (Hooley, 1925), and Morelladon beltrani Gasulla et al., 2015 in the upper Barremian of Cuenca (with the exception of the last taxon) and Castellón provinces (Sanz et al., 1982; Llandres et al., 2013; Gasulla et al., 2014, 2015; Gasulla, 2015; Sanguino & Buscalioni, 2018); and Proa valdearinnoensis McDonald et al., 2012 in the Albain of Teruel province (McDonald et al., 2012). In addition, another styracosternan described in the lower Barremian of Teruel province, ‘Delapparentia turolensis’ Ruiz-Omeñaca, 2011, is generally considered a nomen dubium (Norman, 2015; Verdú et al., 2017).

While the robust basal styracosternan Magnanamus is the unique large ornithopod recognised in the upper Hauterivian-lower Barremian (western part of the Cameros Basin, Soria province), some dorsal vertebrae with elongate neural spines found in deposits of similar age from the nearby Burgos province (Pereda-Suberbiola et al., 2011) potentially represent other taxon, possibly related to the sail-backed styracosternan Morelladon according to their proportions (Gasulla et al., 2015). Additionally, other fossils from the western part of the Cameros Basin have also been related to the robustly built genus Iguanodon in the upper Hauterivian-lower Barremian (Torcida Fernández-Baldor et al., 2006 after Gasca et al., 2014).

In the western Maestrazgo Basin, the upper Hauterivian-lower Barremian outcrops of the Areniscas y Calizas de El Castellar Formation (Salas, 1987; Martín-Closas, 1989) have yielded many diverse dinosaur fossils, including those of large ornithopods (e.g., Luque et al., 2006; Gasca et al., 2009; Gasca, 2011; Cobos et al., 2012; Guerrero & Cobos, 2017). Despite this, ornithopods from this lithostratigraphic unit have been scarcely studied, likely due to the fragmentary conditions of their remains.

In this work, we present the results of a study of an assemblage of vertebral centra recovered from two sites (CT-16 and CT-17) found in the surroundings of El Castellar (southern margin of the Maestrazgo Basin, Teruel, Spain) in outcrops of the El Castellar Formation. In addition to a traditional systematic study of these fossils, a morphometric analysis of the vertebrae has also been conducted to explore vertebral diversity throughout the sample. In ornithopods, these methods have been demonstrated as particularly useful in quantifying morphological differences in functional studies of the axial and appendicular skeleton (e.g., Fearon & Varrichio, 2015) and investigating the evolution of either a particular trait (e.g., VanBuren & Bonnan, 2013) or a specific taxonomic group (e.g., Prieto-Márquez, 2010). Additionally, they are used to identify intraspecific (Verdú et al., 2014) and interspecific variation (e.g., Brink et al., 2014). Finally, morphometric methods have also been employed to test taxonomic affinities of isolated remains (e.g., Barrett et al., 2014) and explore diversity in fossil assemblages (e.g., Blanco et al., 2015), as in the present study.

Hence, the aims of this work are: (a) to anatomically describe the large ornithopod vertebral remains recovered in CT-16 and CT-17, the latter having been previously studied by Cobos et al. (2012); (b) to compare the anatomical features of these fossils with those of other large European ornithopods; (c) to analyse and quantify through morphometric methods the affinities of these fragmentary fossils with other European iguanodontians; and (d) to discuss the implications of these findings for the current knowledge of ornithopod diversity in the Lower Cretaceous of the Iberian Peninsula.

1.1. Institutional abbreviations

GPIT – Institut für Geowissenschaften, Tübingen, Germany.
IWCM – Dinosaur Isle, Sandown, United Kingdom.
MAP – Museo Aragonés de Paleontología, Fundación Conjunto Paleontológico de Teruel-Dinópolis, Teruel, Spain.
MDS – Museo de Dinosaurios de Salas de los Infantes, Salas de los Infantes, Spain.
2. GEOGRAPHIC AND GEOLOGICAL SETTINGS

Dinosaur sites CT-16 and CT-17 were found in the southern section of El Castellar village. Geographically, this town is located in the southeast region of the Iberian Range and 40 km east of the city of Teruel (Fig. 1a).

Geologically, the area of El Castellar lies in the Peñagolosa sub-basin (southern margin of the Maestrazgo Basin) (Salas & Guimerá, 1996). Both CT-16 and C-17 fossils were found ex situ in two separate and clearly differentiated concentrations where the El Castellar Formation outcrops (Figs 1b-1c). This lithostratigraphic unit was defined in the Peñagolosa sub-basin by Salas (1987) and uncomfortably overlies either the Villar del Arzobispo Formation (Kimmeridgian-Tithonian) or the Mora de Rubielos Formation (upper Berriasian-lower Valanginian) (Campos-Soto et al., 2017). In this area, the El Castellar Formation is overlain by the Camarillas Formation (lower Barremian) (Campos-Soto et al., 2017). Broadly, the El Castellar Formation is divided into two stages in the Peñagolosa sub-basin (Cobos et al., 2012). The lower stage generally consists of clays intercalated with channelled bodies of sandstones, which are interpreted as alluvial sediments in an alluvial plain crossed by channels. The upper stage represents lacustrine deposits composed of a marl level with gypsum and, after that, marls alternating with limestones containing numerous fossils of freshwater flora and fauna. The location of CT-16 and CT-17 and the clayish matrix partially covering some fossils indicates that they come from a grey or pinkish clay layer that belongs to the top part of the lower stage of the El Castellar Formation (Fig. 1d). Biostratigraphic studies based on charophytes from this lithostratigraphic unit point to an uppermost Hauterivian-basal Barremian age to the El Castellar Formation (Martín-Closas, 1989).
3. MATERIAL AND METHODS

3.1. Material

According to the size of the fossils studied, the absence of repetition of anatomical elements in fossils of similar size, and some other characteristics, it has been determined that at least three individual large ornithopods, representing two different morphotypes (MOR1 and MOR2), were present in the set of vertebrae from CT-16 and CT-17. Site CT-16 yielded fossils corresponding to a huge ornithopod (MOR1), which was mainly represented by platycerolus caudal vertebrae, and fossils belonging to a middle-sized ornithopod (MOR2), which was represented by a longer-than-high platycerolus dorsal and an amphiocerolus caudal vertebra (the latter unlike MOR1) (MPA collections; see catalogue numbers in Table 1). The size of the remains of MOR2 is consistent with belonging to the same individual, but such identity is nevertheless uncertain. Appendicular remains were also recognised in CT-16 (e.g., a proximal fragment of a femur which likely belonged to MOR2 because of its size), but they were not considered in this work due to their scarcity. Additionally, a big partial centrum (MAP-803) was also ruled out from this study because it was unclear if it belonged to the dorsal or caudal series. Site CT-17 (Cobos et al., 2012) provided the fossils of an individual similar to and comparable in size with MOR2 of CT-16. This specimen was represented by longer-than-high platycerolus dorsal vertebrae (like MOR2 in CT-16), some caudal remains, and appendicular fragments (e.g., a fibular fragment), but only the dorsals were studied here (Table 1) because of the poor preservation of the other fossils. All of these specimens are here considered ‘adults’ for reasons explored in the ‘Discussion’. Note that CPT collection numbers in Cobos et al. (2012) are the same as those designated here as MAP. All these fossils are currently housed in the Museo Aragonés de Paleontología.

Table 1. Fossil vertebrae from CT-16 and CT-17 (housed in the Museo Aragonés de Paleontología, Teruel, Spain).

| Site | Collection numbers | Element         | Assignation       |
|------|--------------------|-----------------|-------------------|
| CT-16| MAP-793, MAP-794,  | Caudal vertebras| Morphotype 1      |
|      | MAP-796, MAP-797,  |                 |                   |
|      | MAP-799, MAP-4548, |                 |                   |
|      | MAP-7730, MAP-7731 |                 |                   |
|      | MAP-3907           | Dorsal vertebra |                   |
|      | MAP-3906           | Caudal vertebra | Morphotype 2      |
| CT-17| MAP-812, MAP-814,  | Dorsal vertebra |                   |
|      | MAP-815, MAP-816,  |                 |                   |
|      | MAP-929            |                 |                   |
the analysis. Posteriorly, this dataset was analysed by principal components analysis (PCA) using free PAST v.3.0 software (Hammer et al., 2001) to determine which European taxa were more morphologically related to fossils from CT-16 and CT-17. PCA is a multivariate exploratory technique that transforms original variables into a new uncorrelated set of variables called principal components (PC) (Blanco et al., 2015). In fact, these new variables are linear combinations of the originals (Hammer, 2018). In this way, PCA allows researchers to quantify the variance between species as well as specimens (Fearon & Varricchio, 2015). Two different PC analyses were performed for each dorsal and caudal centra: one analysed raw data without applying any transformation, and the other removed the size component from the dataset to minimize its impact on the variance (Fearon & Varricchio, 2015). In the second analysis, measurements were transformed with the ‘Remove size from distances’ tool, applying the option ‘Allometric vs. standard’ prior to PCA. For both analyses, PCA was done using a variance-covariance matrix because all data had the same units (millimetres) (Hammer, 2018). Moreover, options for groups were set at ‘Disregard’. Additionally, estimation of missing values was made through the recommended option of ‘Iterative imputation’ (Hammer, 2018). Finally, 1000 bootstrapping replicates were performed for each analysis.

As result, the output data of the PCA included: a distribution plot of the calculated PC scores of each dorsal/caudal vertebra for each taxa with respect to the two most important eigenvalues. Additionally, R-squared Pearson index for correlation between the original variable and the PC scores (Hammer, 2018).

Table 2. Measurements (in mm) of vertebral specimens listed by site and morphotype. Abbreviations and symbols: V: Type of vertebra (C, Caudal; D, Dorsal); L: Length; W: Width; H: Height; a: anterior; p: posterior; m: minimum; r1: L/(Wa + Wp)/2; r2: L/(Ha + Hp)/2; r3: Ha/Wa; r4: Hp/Wp; >: incomplete; ~: estimated. Note that specimens are listed in presumably anatomical order.

| Museum number | V | L   | Wa  | Ha  | Wp  | Hp  | Wm  | r1  | r2  | r3  | r4  |
|---------------|---|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|
| **CT-16 SITE** |   |     |     |     |     |     |     |     |     |     |     |
| **Morphotype 1** |   |     |     |     |     |     |     |     |     |     |     |
| MAP-797       | C | 120 | 147.9| >113.4| 144.9| >124.4| -   | 0.82| -   | -   | -   |
| MAP-794       | C | 120.1| 139.6| >105.9| 129.3| 121.3| 106.5| 0.89| 0.99| -   | 0.94|
| MAP-796       | C | 107.2| >114.3| >109.4| 125.3| 121.4| -   | 0.86| 0.88| -   | 0.97|
| MAP-793       | C | 97  | 107 | >116.2| 104.5| >111.4| ~78.8| 0.92| 0.85| -   | -   |
| MAP-7731      | C | 89.8| 103.1| >104.1| 98.5| 105  | ~71 | 0.89| 0.86| -   | 1.07|
| MAP-7730      | C | 93.7| 80.2| 90.2 | 84.2 | ~82.9| -   | 1.14| 1.04| 1.13| -   |
| MAP-4548      | C | 100.8| 87.6| 86.1 | 90  | 79.2 | 74.3| 1.14| 1.22| 0.98| 0.88|
| MAP-799       | C | 63.1| >46.6| 46.5| >41.1| >42.4| 27.7| -   | 1.36| -   | -   |
| **Morphotype 2** |   |     |     |     |     |     |     |     |     |     |     |
| MAP-3907      | D | 85  | ~54.6| 66.9 | 53.7 | 65.6| >33.3| -   | 1.28| -   | 1.22|
| MAP-3906      | C | 76.6| 64  | 67  | 60.6| 62.8| 44  | 1.23| 1.18| 1.05| 1.04|
| **CT-17 SITE** |   |     |     |     |     |     |     |     |     |     |     |
| **Morphotype 2** |   |     |     |     |     |     |     |     |     |     |     |
| MAP-929       | D | 81.6| -   | -   | -   | -   | -   | -   | -   | -   | -   |
| MAP-812       | D | -   | -   | -   | 60.4| 76  | -   | -   | -   | 1.26| -   |
| MAP-816       | D | 90.8| 54.9| 71.4| 61.7| 75  | 39.6| 1.56| 1.24| 1.30| 1.22|
| MAP-814       | D | 85.6| 61.3| 72  | 59.6| 71.2| 40.2| 1.42| 1.2  | 1.17| 1.19|
| MAP-815       | D | 80.8| 64.6| 80.7| 64.2| 74.4| 47.3| 1.25| 1.04| 1.25| 1.16|
| MAP-930       | D | >63.5| -   | -   | -   | -   | -   | -   | -   | -   | -   |
4. SYSTEMATIC PALAEONTOLOGY

Superorder DINOSAURIA Owen, 1842
Order ORNITHISCHIA Seeley, 1887 (*sensu* Sereno, 2005)
Suborder ORNITHOPODA Marsh, 1882
Clade IGUANODONTIA Dollo, 1888 (*sensu* Sereno, 1998)
Clade EUIGUANODONTIA Coria & Salgado, 1996 (*sensu* Verdú et al., 2018)
Clade DRYOMORPHA Sereno, 1986 (*sensu* Sereno, 2005)
Clade ANKYLOPOLLEXIA Sereno, 1986 (*sensu* Sereno, 2005)
Clade STYRACOSTERNA Sereno, 1986

Styracosterna indet. 1 (MOR1)

**Material.** Eight anterior, middle, and posterior caudal vertebral centra (MAP-793, MAP-794, MAP-796, MAP-797, MAP-799, MAP-4548, MAP-7730, and MAP-7731) from CT-16 were analysed (Figs 2-3).

**Description.** Caudal vertebrae. The most anterior caudal centra, such as MAP-797 (Fig. 2a) and MAP-794 (Fig. 2b), have a block-like shape with broad (i.e., wider than high) subquadrangular anterior and posterior articular faces. However, articular faces become progressively and proportionally higher towards the posterior vertebrae of this part of the tail, so height and width are similar in MAP-796 (Fig. 2c), MAP-793 (Fig. 2d), and MAP-7731 (Fig. 3a). Even though MAP-797 can be described as amphiplatyan (slightly platycoelous), the articular surfaces are fully platycoelous in the successive anterior caudal centra (MAP-794, MAP-796, MAP-793, and MAP-7731). In lateral view, the first anterior caudal centra are as long as they are high (MAP-797 [Fig. 2a], MAP-794 [Fig. 2b], and MAP-796 [Fig. 2c]), but they become higher than long toward the middle region of the tail (MAP-793 [Fig. 2d] and MAP-7731 [Fig. 3a]). Although the ventral surface is poorly preserved in all examples, chevron facets seem to be small (or absent) in MAP-797, similar to the structure observed in the first caudal vertebra of *I. bernissartensis* (Norman, 1980). In lateral view, the ventral margin descends from the anterior articular surface to the posterior because of the drop of the tail from the sacrum (Norman, 1980). The neurocentral suture is closed in all examples, and part of the transverse process is still appreciable laterally in some centra (e.g., MAP-793 [Fig. 2d]).

Moving backwards along the caudal series, the middle caudal vertebrae are recognised for the loss of the transverse processes in the neural arch. Unfortunately, MAP-7730 (Fig. 3b) and MAP-4548 (Fig. 3c) are the only middle centra preserved in CT-16. These are rectangular and longer than high in lateral view. Particularly, articular faces adopt a hexagonal outline in MAP-4548 (Fig. 3c) as consequence of a ridge, which extends longitudinally on the lateral side of the centrum. This is characteristic of Iguanodontia (Norman, 2004). MAP-7730 is platycoelous, but MAP-4548, posterior to the former, is slightly amphicoelous. Additionally, chevron facets are more modest than in anterior caudal vertebrae, with the posterior ones being the largest. A ventral sulcus runs along the ventral surface in both MAP-7730 (Fig. 3b) and MAP-4548 (Fig. 3c). As in the anterior caudal vertebrae, the neurocentral suture is closed in the middle caudal vertebrae.

Finally, the most distal region of the tail is only represented by MAP-799 (Fig. 3d), a small and amphiplatyan reel-like centrum that is highly compressed between both articular faces. In lateral view, this centrum has a markedly arched ventral surface.

Styracosterna indet. 2 (MOR2)

**Material.** One dorsal vertebra (MAP-3907) and one caudal vertebra (MAP-3906) from CT-16 (Fig. 4), along with several dorsal vertebrae (MAP-812, MAP-814, MAP-815, MAP-816, and MAP-929) and a caudal vertebra (MAP-930) from CT-17 were analysed (Fig. 5).

**Description.** Fossils from CT-17 have already been described in detail by Cobos et al. (2012), so only a general description of them is provided in the following paragraph.

**Dorsal vertebrae.** MAP-3907 from CT-16 (Fig. 4a) is tentatively identified as part of the anterior-to-middle region of the dorsal series, whereas MAP-929 (Fig. 5a), MAP-816 (Fig. 5b), MAP-814 (Fig. 5c), and MAP-815 (Fig. 5d) are likely from the middle-to-posterior region. MAP-812 represents part of a vertebral centrum (possibly its posterior articular face), and its anatomical position is uncertain. Overall, the centra are spool-like, longer than high, and moderately compressed between the articular faces. Their articular surfaces are platycoelous and elliptical (higher than wide), with no everted margins. Additionally, they have a sharp longitudinal keel extending along the ventral side of the centrum (e.g., Fig. 5c). The sinuous neurocentral suture is closed in all the vertebrae studied, except for MAP-816, whose suture is partially closed. Finally, the unique changes noticed along the dorsal series are that the centra become slightly anteroposteriorly compressed, more robust, and acquire a more modest ventral keel (MAP-815, Fig. 5d) toward the sacrum.

**Caudal vertebrae.** MAP-3906 (Fig. 4b) is the best-preserved caudal vertebrae for MOR2 (MAP-930 from CT-17 is highly eroded). It can be identified as a member
Figure 2. Caudal vertebrae of Styracosterna indet. 1 (MOR1) from CT-16. At top: the estimated position of the vertebrae along the caudal series (in grey). Both the outline of the caudal series and the ornithopod shade have been redrawn and modified from *Iguanodon bernissartensis* drawing in Paul (2008; Fig. 1c). a) Anterior caudal centrum MAP-797: (a1) anterior view; (a2) left lateral view; (a3) posterior view. b) Anterior caudal centrum MAP-794: (b1) anterior view; (b2) left lateral view; (b3) posterior view. c) Anterior caudal centrum MAP-796: (c1) anterior view; (c2) right lateral view (reversed); (c3) posterior view. d) Anterior caudal centrum MAP-793: (d1) anterior view; (d2) right lateral view (reversed); (d3) posterior view; (d4) ventral view.
Figure 3. Caudal vertebrae of Styracosterna indet. 1 (MOR1) from CT-16. At top: the estimated position of the vertebrae along the caudal series (in grey). Both the outline of the caudal series and the ornithopod shade have been redrawn and modified from *Iguanodon bernissartensis* drawing in Paul (2008; Fig. 1C). 

a) Anterior caudal centrum MAP-7731: (a1) anterior view; (a2) left lateral view; (a3) posterior view. 

b) Middle caudal centrum MAP-7730: (b1) anterior view; (b2) left lateral view; (b3) posterior view; (b4) ventral view. 

c) Middle caudal centrum MAP-4548: (c1) anterior view; (c2) right lateral view (reversed); (c3) posterior view; (c4) ventral view. 

d) Posterior caudal centrum MAP-799: (d1) anterior view; (d2) right lateral view (reversed); (d3) posterior view; (d4) ventral view.
of the anterior series (likely one of the last in this region) since it still preserves the base of the transverse processes. MAP-3906 is a bit larger than high, slightly amphicoelous (unlike platycoelus anterior caudal vertebrae of MOR1), and has inverted triangular articular faces. Additionally, it bears chevron facets both anteriorly and posteriorly on its ventral surface. The neurocentral suture is closed in this specimen.

5. MULTIVARIATE ANALYSIS RESULTS

It should be noted that specimens included in the analysis do not necessarily represent the entire dorsal or caudal series of a particular taxon, so their morphospaces are not completely represented in several cases. Additionally, it should be considered when interpreting the results that some vertebral centra exhibit similar proportions in different regions of dorsal/caudal series for different taxa. For instance, the length/height ratio in the distal caudal vertebrae of the styracosternan I. bernissartensis (1.82 in the 43rd caudal vertebra of RBINS R51 [Norman, 1986]) is similar to that of the middle caudal vertebrae of the dryosaurid Valdosaurus (1.81 in the 20th caudal vertebra of IWCMS 2013.175 [Barrett, 2016]). This could explain some of the phylogenetically inconsistent overlapping observed among different taxa in the results (Fig. 6).

Caudal vertebrae. When the size effect is not removed from the dataset, the first component (PC1) contributes to the variance by 94.8%, whereas the second component (PC2) contributes by only 4.5%. Eigenvalues of both components are significant (Table 3). PCA loading scores indicate that correlation with the length, the height, and the width is high with PC1, but not with PC2 (Table 3). The analysis of caudal centra does not resolve the separated morphospace in shape among the different European iguanodontians (Fig. 6a), except for the dryosaurid Eousdryosaurus, which is characterized for having larger than high anterior caudal vertebrae (Escaso et al., 2014). MOR1 from CT-16 occupies an area, which overlaps significantly with that of Magnan anus and I. bernissartensis, and to a lesser extent with that of Mantellisaurus. Additionally, the caudal centrum of MOR2 from CT-16 falls within the morphospace of MOR1, Mantellisaurus, and Magnan anus. When size effect is removed from the dataset (Fig. 6b), the differences among the caudal series of European iguanodontians become even more attenuated so that the majority are agglomerated around central scores of both components with an extensive overlap of all of them, including MOR1 and MOR2. As in the previous analysis, the unique iguanodontian, which occupies an (almost) exclusive area in the negative scores of the first component, is Eousdryosaurus. In this analysis, PC1 explains 88% of the variance whereas PC2 explains 12%. Loading scores indicate that height and width correlates with PC1, especially the latter, and that height partially correlates with PC2 (Table 3). Note that length has been equalled in every specimen when the size component is removed from the dataset so that it does not produce variation.

Table 3. Output data of the PCA. Abbreviations: L: Length; H: Height; PC: Principal component; W: Width.

| PC    | Eigenvalue | % variance | Confidence interval | Correlation |
|-------|------------|------------|---------------------|-------------|
|       |            |            | 2.5%                | 97.5%       | L   | H   | W   |
| Caudal – including size | | | | |
| 1     | 2894.84    | 94.779     | 92.545              | 96.652      | 0.91396 | 0.99336 | 0.98468 |
| 2     | 137.218    | 4.4926     | 2.7536              | 6.709       | 0.40462 | -0.060178 | -0.14848 |
| 3     | 22.2537    | 0.7286     | 0.27864             | 1.2545      | 0.030898 | -0.098078 | 0.091442 |
| Caudal – excluding size | | | | |
| 1     | 1043.49    | 88.006     | 81.458              | 92.907      | 0     | 0.80212 | 0.98365 |
| 2     | 142.217    | 11.994     | 7.0931              | 18.542      | 0     | 0.59717 | -0.18007 |
| 3     | 0          | 0          | 0                   | 0           | 0     | 0     | 0     |
| Dorsal – including size | | | | |
| 1     | 2101.56    | 88.539     | 82.908              | 92.937      | 0.79232 | 0.98004 | 0.94765 |
| 2     | 203.292    | 8.5647     | 5.0305              | 13.118      | 0.55156 | 0.10225 | -0.2967 |
| 3     | 68.735     | 2.8958     | 1.4793              | 4.4882      | 0.2608 | -0.1705 | 0.118 |
| Dorsal – excluding size | | | | |
| 1     | 1048.77    | 89.407     | 6.47 x 10^28        | 92.917      | 0     | 0.94977 | 0.94102 |
| 2     | 124.255    | 10.593     | 6.2708              | 17.778      | 0     | -0.31296 | 0.33835 |
| 3     | 1.01 x 10^{-21} | 8.58 x 10^{-24} | 6.53 x 10^{-28} | 93.37      | 0     | 0     | -9.27 x 10^{-45} |
**Dorsal vertebrae.** When size effect is not removed from the dataset, PC1 explains the 88.4% variance, while PC2 only contributes to it by 8.6%. According to loading correlation values (Table 3), the length, height, and width contribute highly to the variation of PC1. If the resulting plot of the PCA is observed (Fig. 6c), two areas of distribution are clearly differentiated along the morphospace. Dorsal centra of MOR2 from CT-16 and CT-17 are placed into the negative scores of PC2 along with *Valdosaurus, Morellodon, Mantellisaurus*, and the indeterminate tall-spined iguanodontian (GPIT 1802/1-7). However, dorsal centra of *Magnamanus, I. bernissartensis*, and *I. cf. galvensis* are distributed over the positive scores of PC2. Particularly, the area occupied by the dorsal centra from CT-17 widely overlaps with *Morellodon*, whereas the unique vertebra from CT-16 does not fall in the morphospace of any iguanodontian. In contrast, distinctions among dorsal vertebrae of different iguanodontians are not as clear as in the previous analysis when size effect is removed from the dataset (Fig. 6d). Not only the CT-16 morphospace overlaps with that of *Morellodon*, it also overlaps broadly with those of *Valdosaurus, Magnamanus, I. bernissartensis*, or *I. cf. galvensis*. In this case, PC1 explains 89.4% of the variation, while PC2 only explains 10.6%. The two variables of height and width widely explain PC1 (Table 3).
Figure 5. Dorsal vertebrae of Styracostrum indet. 2 (MOR2) from CT-17. At top: the estimated position of the vertebrae along the dorsal series (in grey). The outline of the dorsal series has been redrawn from Norman (1986; Figs 29b and 31b). The ornithopod profile has been modified from Mantellisaurus atherfieldensis drawing in Paul (2008; Fig. 1c). a) Dorsal centrum MAP-929: (a1) anterior view; (a2) right lateral view (reversed); (a3) posterior view; (a4) ventral view. b) Dorsal centrum MAP-816: (b1) anterior view; (b2) left lateral view; (b3) posterior view; (b4) ventral view. c) Dorsal centrum MAP-814: (c1) anterior view; (c2) right lateral view (reversed); (c3) posterior view; (c4) ventral view. d) Dorsal centrum MAP-815: (d1) anterior view; (d2) right lateral view (reversed); (d3) posterior view; (d4) ventral view.
6. DISCUSSION

The features of MOR1 and MOR2 are similar to those of ornithopods described in Europe (as will be discussed in the following paragraphs). In fact, MOR1 and MOR2 are strongly different from other European and Barremian dinosaurs, despite their fragmentary nature. These Barremian faunas are composed of brachiosaurids and rebbachisaurids (among the sauropods), ceratosaurs, baryonychine spinosaurids, carcharodontosaurids, ornithomimosaurs, other small coelurosaurs (among the non-avian theropods), and ankylosaurs (among the thyreophorans), apart from small and large ornithopods (Pereda-Suberbiola et al., 2012; Alcalá et al., 2018 and references therein). Iberian and Barremian sauropods resembling either the upper Hauterivian-lower Barremian *Soriatitan* or the Barremian-Aptian *Tastavinsaurus* have bigger adult vertebrae, pleurocoelic dorsals, and more rounded articular faces in the proximal and distal caudals (Canudo et al., 2008; Royo-Torres, 2009; Royo-Torres et al., 2017) than MOR1 and MOR2. Theropods similar in size to MOR1 and MOR2, such as either the upper Barremian carcharodontosaurid *Concavenator* from the Iberian Peninsula or the baryonychine spinosaurid *Baryonyx*, exhibit extremely compressed dorsal and caudal centra between their articular faces (Charig & Milner, 1997; Ortega et al., 2010; Cuesta et al., 2019), strikingly different to the condition observed in MOR1 and MOR2. Finally, thyreophorans such as the Barremian ankylosaur *Polacanthus*, whose remains have been identified in the Iberian Peninsula (e.g., Gasulla et al., 2011), have very different wider-than-high dorsal vertebrae which are excavated dorsally forming a V-shape pit, anterior caudals with hearth-shaped articular faces, and middle caudal vertebrae with rounded articular faces and transverse processes (Hulke, 1881; Pereda-Suberbiola, 1994), unlike MOR1 and MOR2.

As mentioned above, the group of dinosaurs with which MOR1 and MOR2 share further characters are the ornithopods. Unfortunately, synapomorphy characters of
Described in the lower Barremian (Camarillas Formation) of the nearby locality of Galve (Teruel, Spain), is an immature individual (Ruiz-Omeñaca et al., 2012) but can be totally differentiated from MOR1 and MOR2 for the reasons explained regarding Hypsilophodon.

Both MOR1 and MOR2 share more features with dorsal and caudal vertebrae of large non-hadrosaurid styracosternans described in the Hauterivian-Barremian of the Iberian Peninsula than with any other group discussed. Next, taxonomic affinities of both morphotypes with other taxa will be discussed independently. Overall, the caudal centra of MOR1 closely resemble those of several iguanodontians (but unlike basal ornithopods such as Hypsilophodon [Galton, 1974]) in having higher-than-long anterior centra that become enlarged in the middle vertebra of the series. This feature is not only extensively present in European genera of large non-hadrosaurid styracosternans such as Barilium (Norman, 2011), Hypselospinus (Norman, 2015), Magnamanus (Fuentes et al., 2016), Iguanodon (Norman, 1980), and Mantellisaurus (Norman, 1986), but also in several other iguanodontians, e.g., Zalmoxes (Weishampel et al., 2003; Godefroit et al., 2009), Tenontosaurus (Forster, 1990), Camptosaurus (Gilmore, 1909; Carpenter & Wilson, 2008), Ouranosaurus (Taquet, 1976; Bertozzo et al., 2017), Eolambia (McDonald et al., 2012), and Bactrosaurus (Godefroit et al., 1998). This consistency of the lateral shape in the caudal vertebrae along iguanodontian evolution might explain why the multivariate analysis concentrates the caudal morphospace for almost all the taxa analysed when the size effect is removed (Fig. 6b). The English Valanginian genus Barilium is characterized by having autapomorphic wider-than-high sub-rectangular articular faces in the first caudal (Norman, 2011). Although the first caudal in MOR1 is incomplete (Fig. 2a), it does not seem to be proportionally as wide as that in Barilium and it closely resembles the squared shape described in Iguanodon (Norman, 1980; Verdú, 2017; Verdú et al., 2017) or Mantellisaurus (Norman, 1986). MOR1 have platycoelous caudal vertebrae such as those of the Iberian taxa Magnamanus (Fuentes et al., 2016) and I. bernissartensis (Norman, 1980). Evidence for the affinity of MOR1 with Magnamanus and Iguanodon is also present in the morphometric study when size is considered (Fig. 6a), even though the type of articular faces is not analysed. Additionally, the platycoelous caudal vertebrae of MOR1 are identical to those of MOR2 in the most posterior dorsal centra (Valdosaurus: IWCMS 2013.175, L/H: 1.18-1.87), they exhibit sub-circular anterior and posterior faces (Barrett, 2016). Similarly, proximal caudals of Hypsilophodon, basal rhabdodontomorphans and Valdosaurus have a L/H ratio (Hypsilophodon: 1.22-1.33 in Fig. 30 in Galton [1974]; basal rhabdodontomorphans: 1.41-1.93 in MDS. VG 72 and 101 [Dieudonné et al., 2016]; Valdosaurus: 1.18 in the 6th of IWCM 2013.175 [Barrett, 2016]) greater than those of MOR1 (L/H: 0.85-0.99). Moreover, the middle caudal centra of Hypsilophodon do not show the typical hexagonal articular faces of iguanodontians (Norman, 2004), in contrast with MOR1. It should be noted that the unique specimen of the basal ornithopod Gideonmantellia amosanjuanae Ruiz-Omeñaca et al., 2012, which was described in the lower Barremian (Camarillas Formation) of the nearby locality of Galve (Teruel, Spain), is an immature individual (Ruiz-Omeñaca et al., 2012) but can be totally differentiated from MOR1 and MOR2 for the reasons explained regarding Hypsilophodon.

Ornithopoda, which are primarily restricted to the cranial and appendicular skeleton (Norman, 2004; Butler et al., 2008; Boyd, 2015), are not identified in MOR1 and MOR2. However, attending to the differences stated above with other contemporaneous dinosaurs, the following combination of characters suggests an affinity with Ornithopoda: the dorsal centra of MOR2 is moderately compressed between the narrow articular faces, and, according to Knoll (2009), without pleurocoeli and with a sinuous neurocentral suture; and caudal centra of MOR1 and MOR2 are slightly compressed between the articular faces, which range from sub-quadrangular to triangular, and without transverse processes in the middle ones. Four groups of ornithopods can be distinguished in the Lower Cretaceous of the Iberian Peninsula: small basal ornithopods related to the English taxon Hypsilophodon (e.g., Sanz et al., 1983; Ruiz-Omeñaca et al., 2012), small basal iguanodontians related to the Upper Cretaceous rhabdodontids (Dieudonné et al., 2016), dryosaurids related to the English taxon Valdosaurus (e.g., Galton, 2009), and diverse large non-hadrosaurid styracosternans (e.g., Gasca et al., 2014; Gasulla et al., 2014, 2015; Verdú et al., 2015; Fuentes et al., 2016), the latter being the group most frequently found in the fossil record. It should be noted that some Lower Cretaceous Iberian fossils identified as basal ornithopods, indeed, might be basal ornithischians (Pereda-Suberbiola et al., 2012). In this context of ornithopod diversity, the vertebral centra of both MOR1 and MOR2 are much too big to belong to mature specimens of basal ornithopods such as Hypsilophodon (Galton, 1974), basal rhabdodontomorphans (Dieudonné et al., 2016) or dryosaurids such as Valdosaurus (Barrett, 2016). Additionally, dorsal centra in Hypsilophodon and basal rhabdodontomorphans exhibit length/height ratio (L/H) (Hypsilophodon: L/H > 1.8 [Cobos et al., 2012]; basal rhabdodontomorphans: MDS. VG 66 and 69: 1.46-1.76 [Dieudonné et al., 2016]) greater than dorsal vertebrae of MOR2 (L/H: 1.04-1.28) and have more rounded and proportionally wider articular faces. Even though the ratio L/H of Valdosaurus certainly overlaps with those of MOR2 in the most posterior dorsal centra (Valdosaurus: IWCM 2013.175, L/H: 1.18-1.87), they exhibit sub-circular anterior and posterior faces (Barrett, 2016). Similarly, proximal caudals of Hypsilophodon, basal rhabdodontomorphans and Valdosaurus have a L/H ratio (Hypsilophodon: 1.22-1.33 in Fig. 30 in Galton [1974]; basal rhabdodontomorphans: 1.41-1.93 in MDS. VG 72 and 101 [Dieudonné et al., 2016]; Valdosaurus: 1.18 in the 6th of IWCM 2013.175 [Barrett, 2016]) greater than those of MOR1 (L/H: 0.85-0.99). Moreover, the middle caudal centra of Hypsilophodon do not show the typical hexagonal articular faces of iguanodontians (Norman, 2004), in contrast with MOR1. It should be noted that the unique specimen of the basal ornithopod Gideonmantellia amosanjuanae Ruiz-Omeñaca et al., 2012, which was
basal iguanodontians such as Zalmoxes (Weishampel et al., 2003) and Tenontosaurus (Forster, 1990), but centra in MOR1 are significantly bigger. In contrast, anterior caudal centra are amphiplatyan in the middle-sized taxon Mantellisaurus (Norman, 1986), of which fossils have also been described in the upper Barremian of the Iberian Peninsula (Llandres et al., 2013; Gasulla et al., 2014), and in the robust English and Valanginian genus Barilium (Norman, 2011). The deep ventral sulcus present in some middle caudal centra of MOR1 (e.g., MAP-4548, Fig. 3c) does not have systematic value because it was reported as individually variable (Verdú et al., 2017). Unfortunately, caudal vertebrae have not been described in the large Iberian styracosternans Morelladon (Gasulla et al., 2015) and Proa (McDonald et al., 2012).

Dorsal vertebrae of MOR2 are like those of Morelladon (Gasulla et al., 2015) and Mantellisaurus (Norman, 1986) in having longer-than-high middle-to-posterior dorsal centra, but they are different from those having higher-than-long middle-to-posterior dorsal centra, such as Magnanamus (Fuentes et al., 2016), I. galvensis (Verdú, 2017), and I. bernissartensis (Norman, 1980). This observation agrees with the morphometric study of the dorsal vertebrae when size effect is considered. Here, MOR2 occupies a similar morphospace as Mantellisaurus and Morelladon, being particularly overlapped with the latter (Fig. 6c). Longer-than-high cylindrical middle-to-posterior centra is present in several other iguanodontian genera, including Zalmoxes (Weishampel et al., 2003; Godfrey et al., 2009), Valdosaurus (Barrett, 2016), Camptosaurus (Carpenter & Wilson, 2008), Hippodraco (McDonald et al., 2010), and Ouranosaurus (Taquet, 1976; Bertozzo et al., 2017). Nevertheless, middle-to-posterior dorsal vertebrae are amphiplatyan in Mantellisaurus (Norman, 1986), unlike those of MOR2, which are platycoelous, similar to those of Morelladon (Gasulla et al., 2015). In contrast, Magnanamus (Fuentes et al., 2016), I. cf. galvensis (Verdú, 2017), and I. bernissartensis (Norman, 1980) have amphiplatyan dorsal vertebrae. Platycoelous dorsal vertebrae have been only described in the non-hadrosaurid iguanodontians Tenontosaurus (Forster, 1990; Winkler et al., 1997) and Hippodraco (McDonald et al., 2010), apart from Morelladon. Additionally, MOR2 exhibits a ventral keel in the dorsal vertebrae similar to Mantellisaurus (Norman, 1986), Morelladon (Gasulla et al., 2015), and I. bernissartensis (Norman, 1980). In contrast, this feature is absent in the dorsal vertebrae of Magnanamus (Fuentes et al., 2016) and in I. cf. galvensis (Verdú, 2017). Moreover, the ventral keel feature in the dorsal vertebrae is broadly extended among iguanodontians, e.g., Zalmoxes (Weishampel et al., 2003; Godfrey et al., 2009), Tenontosaurus (Forster, 1990), Cumnoria (Galton & Powell, 1980), Barilium (Norman, 2011), Hypselospinus (Norman, 2015), Ouranosaurus (Taquet, 1976; Bertozzo et al., 2017), and Bactrosaurus (Godefroit et al., 1998). MOR2 dorsal centra are also very similar to the platycoelous vertebrae of GPIT 1802/1-7 found in beds of similar age from Salas de los Infantes (Burgos, Spain) (Pereda-Suberbiola et al., 2011) and which closely resemble those of Morelladon (Gasulla et al., 2015). Dorsal centra of GPIT 1802/1-7 are resolved in the same sector as those of MOR2 in the morphometric study but they do not overlap it, perhaps either due to intraspecific variation or because they belong to different parts of the dorsal series.

The caudal vertebra of MOR2 is resolved in the area occupied by Mantellisaurus, as well as MOR1 and Magnanamus. Similarity of the caudal vertebra of MOR2 with those of Mantellisaurus is not surprising as it is also observed for the dorsal vertebrae. In contrast, recovered relation with Magnanamus and MOR1 is striking and might have been due to the similar proportion among some elements of the caudal series in these taxa. The anterior caudal centrum of MOR2 is amphicoelous and it differs from those amphiplatyan in Morelladon (Norman, 1986) or those platycoelous in Magnanamus (Fuentes et al., 2016) and I. bernissartensis (Norman, 1980). Unfortunately, caudal vertebrae are unknown in Morelladon (Gasulla et al., 2015), but amphicoelous caudal vertebrae might be one of its features. Camptosaurus (Gilmore, 1909) and Ouranosaurus (Taquet, 1976; Bertozzo et al., 2017) are unique non-hadrosaurid iguanodontian genera with amphicoelous anterior caudals.

Hence, MOR1 and MOR2 cannot be classified with more precision than that of Iguanodontia indet. when considering only anatomical features. A summary of those characters which suggest iguanodontian affinity include the following: caudal centrum of MOR1, which are anteroposteriorly compressed in the anterior vertebrae but which elongate toward the middle part of the tail; middle caudals, according to Norman (2004), which acquire hexagonal articular faces; and dorsal centra of MOR2, tall in lateral view (more than in basal ornithopods), with higher-than-wide elliptical articular faces and with the presence of a ventral keel. However, their Lower Cretaceous age, when the abundance and diversity of styracosternans was high in the Iberian Peninsula, and the similarity of MOR1 and MOR2 with sympatric styracosternan genera allow us to classify them tentatively as Styracosterna indet. Note that Cobos et al. (2012) classified the fossils from CT-16 (here MOR2) as Hadrosauriformes indet. However, we consider it more appropriate to include the fossils in a more inclusive clade with further support by phylogenetic studies. As such, the classification is Styracosterna (e.g., Verdú et al., 2018). Even though MOR1 and MOR2 share proportions and configuration of articular faces with Magnanamus/Iguanodon and Morelladon respectively, it is better to be conservative when working with fossils that are so fragmentary and not propose genera/specific assignations to avoid the introduction of taxonomic noise.
Several authors have reported relevant ontogenetic changes that affected the axial skeleton during growth in some small and large-sized iguanodontians (e.g., Dewaele et al., 2015; Verdú, 2017; Hübner, 2018). Since vertebrae of MOR2 are smaller than those of MOR1, they might belong to an immature specimen of the latter rather than to a different taxon. For this reason, the determination of the approximate ontogenetic stage of studied fossils is crucial for discriminating between systematic informative characters and those that are ontogenetically variable. In this sense, the size of the fossils and the pattern of closure of the neurocentral suture (i.e., the degree of fusion of the neural arch with its centrum) are good qualitative indicators of the degree of development in ornithopods (Verdú, 2017 and references therein). All caudal vertebrae of MOR1 have their neurocentral sutures totally closed, including those in the most anterior region of the series. The fusion of the neural arch to the centra in these vertebrae occurred between the juvenile and sub-adult stages in Iguanodon (Verdú, 2017). However, the size of the caudal centra of MOR1 reveals that it might be even older than a sub-adult since they are larger (e.g., MAP-797, 120 mm of length) than those of big ‘adult’ specimens of I. bernissartensis (e.g., RBNS R51, 104 mm [Norman, 1986]) or Magnamanus (e.g., MNS 2000/132.2, 112.8 mm [Fuentes et al., 2016]). Thus, vertebrae of MOR1 likely belonged to a fully-growth adult.

In the case of MOR2, specimens from CT-16 and CT-17 are similar in size, so it is assumed that they had roughly the same somatic age. All the dorsal vertebrae of MOR2 have the neurocentral sutures closed (except for MAP-816, which has the sutures partially closed), including MAP-3907, which is tentatively identified as an anterior-to-middle dorsal vertebra. Closure of the neurocentral suture in the anterior dorsal vertebrae occurred late in Iguanodon development, specifically between the sub-adult and adult stages (Verdú, 2017). Moreover, the length of the dorsal (e.g., MAP-816, 90.8 mm) and caudal vertebrae (MAP-3906, 76.6 mm) of MOR2 is compatible with that of supposed ‘adult’ specimens of the middle-sized iguanodontians Mantellisaurus (e.g., RBINS R57, dorsal: 80 mm, caudal: 79 mm [Norman, 1986]) and Morelladon (e.g., CMP-MS-03-09, dorsal: 93.5 mm). Thus, the middle-sized specimen MOR2 can be confidently attributed to a mature specimen as MOR1, and therefore it does not represent an immature specimen of the later morphotype. Consequently, all anatomical differences observed between the caudal vertebrae of MOR1 and MOR2 can be considered systematically significant, including the size of the centra, and similarities reported with other taxa are not a product of ontogeny.

Finally, evidence from the southern region of the Maestrazgo Basin provided in this work further supports to the hypothesis of coexistence of two forms of large ornithopods during the late Hauterivian-early Barremian in the Iberian Peninsula as was previously observed in the western part of the Cameros Basin (Pereda-Suberbiola et al., 2011, 2012; Fuentes et al., 2016): a robust styracosternan related to Magnamanus/Iguanodon and a slender one related to Morelladon. However, it should be noted that Magnamanus is anatomically very similar to Iguanodon, and future reviews of the holotype might shed light on a closer relationship between them than previously thought (Verdú, 2017). According to evidence derived from the study of different iguanodontian dentary teeth, the diversity of styracosternans might ascend to three taxa in the lower Barremian (Gasca et al., 2014). Similarly, younger upper Barremian deposits in Spain have also yielded fossils of the large and robust I. bernissartensis (Gasulla et al., 2014) and two slender taxa, Mantellisaurus (Llandres et al., 2013; Gasulla et al., 2014) and Morelladon (Gasulla et al., 2015). As for Magnamanus, fragmentary remains from the Iberian Peninsula assigned to Mantellisaurus might be related to Morelladon in future studies given the close phylogenetic relationship between both taxa. In the same way as this study, Rey et al. (2018) also identified two different styracosternan morphotypes (one robust and other slender) in the Albian of Teruel (Spain), based on fragmentary vertebral remains. In fact, Norman (2012) already observed the recurrent record of both a robust and a slender large styracosternan associated in the Wealden Group of the southern England: Barilium and Hypselospinus in the Valanginian, and I. bernissartensis and Mantellisaurus in the Barremian-lower Aptian. Similar association of styracosternans is also found in the Sainte-Barbe Clays Formation (upper Barremian-lower Aptian) of Belgium: I. bernissartensis and Mantellisaurus (Norman, 2012).

6. CONCLUSIONS

We have systematically and morphometrically studied a sample of large-ornithopod dorsal and caudal vertebral centra found in the upper Hauterivian-lower Barremian (Lower Cretaceous) deposits of the southern margin of the Maestrazgo Basin (northeastern Teruel, Spain). They correspond to at least three mature (or almost mature) individuals of two different but indeterminate iguanodontians, MOR1 and MOR2, which are ascribed tentatively to the more inclusive clade of Styracosternia. First, MOR1 is represented by several caudal centra, which can be related to the robust and large Iberian styracosternans Magnamanus (from the upper Hauterivian-lower Barremian) and Iguanodon (from the Barremian-Aptian) because of the values of their L/H ratio and their platycoelous anterior caudals. Morphometric study with a multivariate PCA also shows morphological affinity of...
these two taxa with MOR1, but not when size effect is removed, perhaps because proportions of caudal vertebrae are consistent along iguanodontian phylogeny.

The second styracosternan, MOR2, is represented by dorsal and caudal centra. The dorsals of MOR2 resemble those of the sail-backed styracosternan from the upper Barremian *Morelladon* in their L/H ratio and their platycoelus articular faces. Morphometric study with a multivariate PCA also supports this observation for the dorsal vertebrae of MOR2, at least when size is considered. Anterior caudal of MOR2 is amphicoelus unlike MOR1.

Thus, this study provides further support, using evidence recorded in the El Castellar Formation, for the hypothesis that two forms of large styracosternans coexisted in sympatry during the late Hauterivian-early Barremian in what is today the Iberian Peninsula. In future research, incorporation of either additional lineal measurements taken along vertebrae (including those in the neural arch) or 3D geometric morphometric methods might convert multivariate analysis into a useful tool for identifying the taxonomical affinities of isolated ornithopod vertebrae.

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