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A Discrete Fracture Network Model With Stress-Driven Nucleation: Impact on Clustering, Connectivity, and Topology

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The realism of Discrete Fracture Network (DFN) models relies on the spatial organization of fractures, which is not issued by purely stochastic DFN models. In this study, we introduce correlations between fractures by enhancing the genetic model (UFM) of Davy et al. [1] based on simplified concepts of nucleation, growth and arrest with hierarchical rules. To do so, the nucleation of new fractures is correlated with the elastic strain energy of distortion stored in the matrix, which is a function of preexisting fractures. Discrete Fracture Networks so generated show multi-scale clustering effects with fractal dimensions below the topological dimension over a broad range of scales. The fractal dimension depends on the way one correlates the nucleation occurrence to the strain energy. Fracture clustering entails a spatial variability of the fracture density, which increases with the intensity of the coupling between stress and nucleation. The analysis of connected clusters density and of fracture intersections also highlights the differences between the UFM models and its equivalent Poisson model. We show that our stress-dependent nucleation model introduces some new fracture size-positions correlations, with small fractures tending to connect to the largest ones.

Keywords: discrete fracture networks (DFNs), nucleation, clustering, connectivity, topology

INTRODUCTION

Fractures are ubiquitous structures controlling both flows and rock mechanical strength in geological environments. Modeling the fracture network is thus a key prerequisite of forecasting modeling in many industrial applications such as managing groundwater/petroleum resources, assessing risks associated with geotechnical constructions or deep waste disposal, among others. In most of the cases, fractures cannot be modeled deterministically, because they cannot be observed in three-dimensions with sufficient resolution at all scales. Hence, the modeling must be stochastic, which consists of generating a 3D fractured medium statistically equivalent to measures and observations. Discrete Fracture Network modeling is one of the most convenient and used of the stochastic methods; it describes fractured rocks as a population of individual fractures, whose parameters (size, shape, orientation, aperture, and position) are drawn from statistical probability distributions derived from observation maps (mainly 2D outcrop and tunnels, 1D fracture intensity along wells, or 3D geophysics imagery) and models [see [2, 3] for reviews]. In its simplest form, the model (which we will refer as the Poisson model) consists of positioning fractures...
at random in the generation volume with a given density, and of assigning other fracture parameters independently by bootstrapping the parameter distributions of observations [4–9]. The method is easy to implement, but it neglects most of the complexity of the underlying fracturing process, in particular the correlations induced by fracture-to-fracture mechanical interaction [10–12]. This so-called Poisson model is thus a crude representation of geological fractures that can lead to large discrepancy between modeled and natural network in terms of network topology [13, 14], having dramatical impact on the estimated hydrological and mechanical behavior of the fractured rock mass [15–17]. A way to improve the realism of DFN models is the use of genetic models, in which the fracture hierarchy reproduce correlations between their different geometrical attributes. Nevertheless, a full mechanical description of the fracturing process [18–20] is not feasible when dealing with dense networks made of fractures having sizes from centimeter to tens of kilometers. The broad range of natural fracture size distributions and their footprint extends to several kilometers. Recent papers [1, 24] suggest that all scales matters. This is even more important for industries such as nuclear deep waste disposal where small fractures may have an important role at the nearfield of the repository, whose footprint extends to several kilometers. Recent papers [1, 24] have proposed a genetic model of DFN, called “Universal Fracture Model” (UFM model), using simplified fracturing-relevant rules for nucleation, growth and arrest of fractures to draw complex and dense networks. With simple kinematic rules that mimic the main mechanical processes, the model produces fracture size distributions and fracture intersections that are consistent with observations [24]. It results in less connected networks than the Poisson model and changes in the network topology [25], which has been proven to have an impact on flow properties, particularly decreasing effective permeability and increasing flow channeling [26]. Nonetheless, the random positioning of new fractures (nuclei) in this model is still too simplified to reproduce spatial variability and clustering effect observed in natural fracture network patterns [11, 23].

Indeed, nucleation is a complex process both controlled by the initial nuclei size and the parameters (that can be a power-law, exponential, etc...) and a rate $\dot{n}_N = dN_N/dt$ (with $N_N$ the number of nuclei introduced in the system). Nuclei positions are assumed to be uniformly distributed in space here, we will refer to this model as the stress-independent UFM model. Once created, fractures grow following a power law relationship to describe the crack tip velocity in the subcritical regime [36]:

$$v(l) = Cl^a$$

with $l$ the fracture length, $C$ the growth rate and $a$ the growth exponent. If not arrested, nuclei size increases non-linearly with time and becomes infinite for a finite time $t_{\infty}$ dependent on the initial nuclei size and the parameters $C$ and $a$. For constant nucleation rate and no fracture arrest, even if fractures grow, there is a stationary solution for the fracture size distribution [1]:

$$n_G(l) = \frac{\dot{n}_N}{C} l^{-a}$$

The arrest rule is assumed to reflect the mechanical interaction between fractures. In this model, we consider these interactions as a binary law where fractures can only abut on larger ones, but the reverse is not likely to occur. It results in a large proportion of T-shape intersections that are consistent with field observation, and in a quasi-universal self-similar fracture size distribution:

$$n_A(l) = Dl^{D-1}$$

THE MODEL

The Discrete Fracture Network approach for modeling fractured rock masses refers to numerical models explicitly representing the geometry of each fractures forming the network. Generally, this geometry (positions, orientations, size...) is generated stochastically from data statistics. The simplest model considers fractures independent of each other; it will be referred to as the Poisson DFN model. The DFN model developed by Davy et al. [1] is based on basic mechanical concepts described in Davy et al. [24]. The fracturing process is divided in three main stages in a time-wise approach: nucleation, propagation and arrest of fractures. We first develop the model and how in its simplest form—i.e., with a Poisson distribution of nuclei—it controls the network size and intersection distributions. Then, we introduce a more complex nucleation model based on the stress perturbation of pre-existing fractures.
Stress-Driven Nucleation

In this section, we further develop the stress-independent UFM model by making nucleation dependent on stress redistributions caused by existing fractures. For this, we introduce a stress field based probabilistic sampling of nuclei locations. The stress field evolves over the whole domain as the fracturing process. Considering the system to be linear elastic, which may not be the case for highly damaged materials, we define the stress field \( \overline{\sigma}(x, t) \) at any position \( x \) in the domain at time \( t \) as the superposition of every fracture \( \sigma_f(x, t) \): 

\[
\overline{\sigma}(x, t) = \sigma^{\infty}(t) + \sum_f \sigma_f(x, t)
\]

New nuclei are progressively introduced in the system following a probability field \( P(x, t) \) derived of this stress field: 

\[
P(x, t) \sim [\sigma_{VM}(x, t)]^m
\]

where \( \sigma_{VM}(x, t) \) is the Von Mises stress [37] and \( m \) a parameter that quantifies the coupling between nucleation occurrence and stress (thereafter called the selectivity parameter). The Von Mises stress is a scalar invariant measure of the deviatoric stress intensity and a measure of the elastic strain energy of distortion stored in the matrix. We then generate a scalar stress-intensity field that will serve as a basis to construct a discrete probability distribution for nuclei position sampling, without using any strength criteria. This stress-driven nucleation process is thus defined as an annealed disorder process [35] where nuclei positioning is directly a function of the system evolution. The model then needs two more parameters: the remote stress field tensor \( \sigma^{\infty} \) and the selectivity parameter \( m \). The latter quantifies the influence of the stress field heterogeneity on nucleation. For large \( m \), nuclei tend to concentrate in regions with high stress intensity. In the following, we will refer to this model as the stress-driven UFM model. The case \( m = 0 \), where the nucleation is uniformly distributed in space, corresponds to the stress-independent UFM model.

For numerical implementation of the model, we use the same basic assumptions as Davy et al. [1]: fractures are modeled as interacting growing disks in a time-wise process. At each time step \( \Delta t \), \( n_{\Delta t} \Delta t \) nuclei are introduced in the cubic system. Those who are not intersecting any existing fractures are kept in the system and grow following equation [1], until they cross a larger fracture or reach an infinite size, that we set to be twice the system size. Nuclei are not allowed to intersect pre-existing fractures so that the available space for new nuclei and the effective nucleation rate decrease with simulation time. The stress perturbation associated to each fracture can be approximated using the 3D tensorial analytical solutions of Fabrikant [38] for uniformly loaded freely-slipping penny-shaped cracks, considering traction and/or shearing. If the system is under compression, then only the shearing part is considered. Each fracture is assumed to be uniformly loaded by the remote stress field and generate a stress perturbation that depends on the fracture size, the input remote stress field intensity, and their relative orientations. To fasten calculations, we do not calculate the interaction terms between fractures and set them to zero. Although these terms may be non-negligible when fractures get close with each other [39–41], in particular when the fracture density increases, we have estimated that this approximation is consistent with the degree of simplification used for the different stages of the model. A more elaborate version is under development.

The stress field is computed over the whole domain, on a regular cartesian stress grid \( S_{x} \) of resolution \( r_{\text{stress}} \). In order to obtain a dimensionless stress-intensity scalar field (Figure 1), each cell value is divided by the remote stress Von Mises value.

For each nucleation step, a cell is chosen from a discrete probability sampling over the whole stress grid \( S_{x} \), where the probability for each cell \( c \) is defined by:

\[
P(x_c) = \frac{[\sigma_{VM}(x_c)]^m}{\sum_{c \in S} [\sigma_{VM}(x_c)]^m}
\]

Once the nucleus cell has been determined, the nucleus center position is randomly taken inside this cell. The number of

**FIGURE 1** | Stress-intensity field generated on a regular grid of size \( L = 1 \) and resolution \( r = 0.01 \).
computations is directly related to the nucleation rate, the stress grid resolution, and the increasing number of fractures already present in the medium, which can be large. Since the addition of a single small nucleus does not affect the stress field at the system size, a single stress grid can be used for several nucleation steps $n_{\text{step}}$ in order to accelerate computations. This heterogeneous probabilistic point process of nuclei positions constitutes an improvement of the stress-independent UFM model, while keeping constant nucleation rate.

RESULTS

In this section, we focus on the spatial and topological analysis of fracture networks generated by this stress-driven UFM model. We analyze the evolution of the pattern complexity of this model with the selectivity parameter $m$, and compare the results with the stress-independent UFM ($m = 0$) and equivalent Poisson model (i.e., same population of fractures with random positions in the domain).

Numerical Simulations

For all models, we seed and let fractures grow in a domain of size $L = 1$ with a growth exponent $a = 3$, so that the power-law exponent of the dilute regime is $-3$, which is consistent with field data [24]. Nuclei appears in the system with a constant nucleation rate $\dot{n} = 20$, growth rate $C = 1$, and a size drawn from a narrow-ranged power-law distribution:

$$p_N(l) = \frac{(b-1)}{l_N} \left( \frac{l}{l_N} \right)^{-5}$$

$l_N$ is the minimum nuclei size; its value has been set at 0.01 in order to cover two orders of magnitude in the resulting fracture size distribution. For each timestep $\Delta t$, we introduce $n_{\text{step}} = 200$

![FIGURE 2](A) Poisson, (B) stress-independent UFM ($m = 0$), and (C) stress-driven UFM ($m = 3$) models, and (D) associated size distributions ($m \in [0, 5]$).
new nuclei. Nuclei intersecting existing fractures are rejected, the effective nucleation rate is thus decreasing with time, since the available space for new fractures to form is also decreasing. We stop simulations when the time is close to the \( t_{\infty}(l_N) \), i.e., the time necessary for the smallest nuclei to become infinite [1]. Networks are generated for different values of selectivity parameters \( m \in \{0,1,2,3,4,5\} \) in order to quantify its impact on fracture clustering. The equivalent Poisson model is obtained by moving the fracture centers randomly in space, so that the size distribution remain identical but the correlations between fracture size and position are destroyed. We only show the Poisson model derived from the stress-independent UFM \((m = 0)\). All stress-driven UFM are generated from the same constant and compressive remote stress field \( \sigma \), so that \( \sigma_1 = \sigma_{xx} = -4, \sigma_2 = \sigma_{yy} = -2, \sigma_3 = \sigma_{zz} = -1, \) and \( \sigma_{xy} = \sigma_{xz} = \sigma_{yz} = 0 \). The intensity and orientations of the principal stress components do not matter in this set of simulations since we consider that, for all generated DFNs, the orientation distribution is stress-correlated and is assumed uniform in order to minimize asymmetry due to the remote stress field. By doing so, we aim at focusing on the consequences of the stress field heterogeneity on spatial correlations only, but not on its spatialization. For this set of parameters, the simulation time for a stress-driven UFM model is about 60 times larger than for the stress-independent UFM model for which there is no stress. For all the models, we perform 50 realizations of each model for statistical analysis.

Figures 2A–C show 2D slices of generated three-dimensional Poisson, stress-independent UFM and stress-driven UFM \((m = 3)\), respectively. Visually, the stress-driven nucleation process seems to increase the clustering effect of fractures positions. Simulations are stopped when \( t = t_{\infty}(l_N) \), so that both the dilute and dense regime can be observed on the fracture size distribution (Figure 2D) and are consistent with equations [2] and [3], with a transitional length \( l_c = 0.15 \):

\[
n(l) = \begin{cases} 20l^{-3} & \text{if } l < l_c \\ 3l^{-4} & \text{if } l > l_c \end{cases}
\]

For all models, fracture size distributions are almost the same.

When increasing the selectivity parameter, new nuclei to form should be attracted to the tip of the largest existing fractures since the stress is high here, increasing the probability of rejection. Hence, for the same simulation time, the fracture density should decrease when increasing selectivity parameter \( m \). We consider three-dimensional fracture densities here, such as defined by Dershowitz and Herda [42]: fracture number density \( P_{30} \) (number of fractures per unit volume), fracture intensity \( P_{32} \) (total fracture surface per unit volume), and percolation parameter \( p \) (total excluded volume around fractures per unit volume) that quantifies the network connectivity [43]. Considering the disk-shape assumption we made in our DFN models, we define these densities as:

\[
P_{30} = \int n(l) \Pi(l,L) dl
\]

\[
P_{32} = \frac{\pi}{4} \int n(l) \Pi(l,L) l^2 dl
\]

\[
\Pi(l,L) = \frac{\pi^2}{8} \int n(l) \Pi(l,L) l^3 dl
\]

\(\Pi(l,L)\) is surface ratio of the fracture \( l \) included in the domain of size \( L \). Figure 3 summarize the density statistics of the generated networks.

One can notice a decreasing of both \( P_{30}, P_{32}, \) and \( p \) with selectivity parameter \( m \) due the increasing number of rejected nuclei. Moreover, the fracture density of the stress-independent UFM model is slightly slower than its equivalent Poisson model, because both models are subjected to different finite size effects.

**Impact on Clustering**

The spatial organization and topology of fractures in a network may have dramatical impact on its connectivity and hydraulic behavior. Quantifying the fractures organization in DFN models and comparing with natural networks is a key challenge to qualify the relevance of simplified models. Natural fracture networks have shown complex clustered patterns [11, 44] that has consequences on connectivity [21, 45, 46].

Considering the multiscale nature of fractures, and thus of the subsequent stress fluctuations, we expect fractal correlations to develop in our model. The full multifractal spectrum of fracture organization may be quantified using the box-counting method [47], computing the number of boxes to cover the network at different scales. Nevertheless, this technique has been shown to be strongly affected by finite size effects [11, 23, 48]. We then compute the 2-point correlation integral (or correlation pair function) to describe the spatial correlations of fractures positions. This method gives the probability for two fractures to belong to the same cluster. For a population of \( N_f \) fractures, the associated correlation pair function is defined by:

\[
C_2(r) = \frac{2N(r)}{N_f(N_f - 1)}
\]

with \( N(r) \) the number of points whose mutual distance is less than \( r \) [47]. For a large population of points, this quantity...
tends to scale a power-law $r^{D_2}$, where $D_2$ is the correlation dimension of fracture centers. The correlation dimension can thus be obtained by computing the slope of the correlation pair function in a log-log plot. Figure 4A shows the evolution of the correlation pair function and its derivative with scale $r$. This function is affected by finite-size and resolution effects when the mutual distance $r$ tends to domain finite size and nuclei size, respectively. We calculate a correlation dimension $D_2$ in the interval $[0.03, 0.08]$ that is not affected by size effects. Indeed, below the lower bound, the derivative of the correlation pair function increases dramatically because few fractures are so close to each other. This resolution effect is related to the no-intersecting assumption when introducing new nuclei in the UFM model. On the other hand, finite size effect due to the system size are already dramatic below the upper bound, as we should obtain a correlation dimension $D_2 = 3$, for the Poisson model. As expected, the correlation dimension of the 3D Poisson and stress-independent UFM models is $D_2 \sim 3$, which is consistent with a uniform distribution of fracture centers in space for both models. $D_2$ is smaller than 3 for the stress-driven UFM model and decreases when increasing the selectivity parameter $m$. For large values of $m$, nuclei concentrate in zones of high stress, mainly near the tips of the largest fractures, leading to a clustering of fractures. Figure 4B shows that correlations exist even at early simulation times. For large $m$ values, the correlation dimension increases slightly with time, as the available space around fractures tips decreases.

The correlation dimension of fracture centers indicates how much a fracture network occupies its underlying metric space. Nevertheless, two networks can have the same correlation dimension but very different patterns. In order to describe the texture associated to a network, we use the concept of lacunarity [49]. Fundamentally, lacunarity is a dimensionless representation of the variance to mean ratio [50] defined here as:

$$\lambda_M(s) = \left[ \frac{\sigma_M(s)}{\mu_M(s)} \right]^2$$

with $\sigma_M(s)$ and $\mu_M(s)$ the standard deviation and mean of a measure $M$ at scale $s$. For any density measure $M$, if $\lambda_M(s) \to 0$, the pattern is perfectly homogeneous at scale $s$. Lacunarity is a scale dependent measure, whose analysis quantify the degree of clustering and anti-clustering [51], and potentially on different regimes [50, 52, 53], when analyzing lacunarity curves, showing $\lambda_M(s)$ evolution with scale $s$. Lacunarity analysis can then be used to analyze textural heterogeneity of fracture densities [54]. For three-dimensional fracture networks, we can define various measures $M$ quantifying fracture density as defined by Dershowitz and Herda [42], or in section Numerical Simulations.

Figure 5 shows the lacunarity curves of the three-dimensional fracture densities defined in section Numerical Simulations. As expected, the $p_{30}$ lacunarity curve is the same for the stress-independent UFM model and its equivalent Poisson
Consequences on Connectivity and Topology

Fracture correlation is likely changing the connectivity of the overall network. Maillot et al. [26] show that stress-independent UFM networks (which they refer as a "kinematic" model) have permeabilities 1.5–10 times smaller than the equivalent Poisson model, and a higher channeling (a higher portion of the total fracture surface where the flow is significant). Some studies [46, 55] show that, for networks with a power-law size distribution, the evolution of connectivity with scale is strongly dependent on the power-law exponent $a$ and on the fractal dimension of fracture centers $D_2$. In our case, because fracture centers tend to concentrate in clusters around large fracture tips, the fracture interconnectivity may also increase. Increasing the connectivity between fractures should increase the backbone density, defined as the structure carrying flow in the network [56]. We here define the backbone by removing iteratively fractures having only one connection with the network or the boundary, keeping only the connected clusters without flow dead-ends. Figure 6 shows that the percentage of fractures in number and in total area ($p_{30}$ and $p_{32}$) involved in the backbone is smaller for any kind of UFM model (stress-independent or stress-driven) than for corresponding Poisson model. Moreover, for the UFM models, even if $<25\%$ of fractures are part of the backbone, this represents more than $65\%$ of the backbone surface, which shows that connectivity is mostly ensured by large structures [21, 45]. Finally, as the number of fractures involved in the backbone structure increases more than the total area with $m$, we can conclude that we tend to connect mostly small fractures to the backbone with this stress-driven nucleation process.

The number of intersections per fracture is a good indicator of fracture connectivity. Maillot et al. [26] showed that the number of intersections per fracture is a function of fracture size for both Poisson and UFM models. Moreover, they showed that whatever the fracture size, the number of intersections is about two times lower for UFM model than for equivalent Poisson model. We here push further this topological analysis computing the fracture intersection matrix $P_i$ so that for $n_i$ and $n_j$ fractures of size $l_i$ and $l_j$, respectively, $P[i,j]$ gives the number of fractures of size $l_i$ intersecting fractures of size $l_j$, divided by $n_i n_j$. Figures 7A–C show the mean intersection matrix for all generated Poisson, stress-independent UFM ($m = 0$), and stress-driven UFM ($m = 5$) models, respectively. As expected, in any case, the probability of intersection increases with fractures sizes. Figures 7D,E show the mean intersection matrix for stress-independent UFM ($m = 0$), and stress-driven UFM ($m = 5$) models, normalized by the mean equivalent Poisson’s model intersection matrix, in order to highlight differences between UFM and Poisson models. Our analysis shows that the number of fractures intersections is smaller for UFM models than their equivalent Poisson. We can also notice that this number is much smaller for fractures of the same size, which is a consequence of the UFM rule assuming that a fracture cannot cross a larger one. Finally, Figure 7F shows the stress-dependent UFM ($m = 5$) model intersection matrix, normalized by the one of the stress-independent UFM ($m = 0$) model, showing that our stress-driven nucleation process tend to increase the connectivity of small fractures with the smallest and the largest fractures. Indeed, new fractures tend to develop at the tip of the largest existing fractures, increasing connectivity between both.
CONCLUSION

The genetic UFM model developed by Davy et al. [1], describing the fracturing process as a combination of simplified nucleation, growth and arrest laws, introduces a fracture size-position correlation in DFN modeling, that does not exist in equivalent Poisson model. It results in two distinct power-law fracture size distributions and a number of T-intersections that are consistent with field data [24]. Nevertheless, the model does not take into account the mechanical feedback loop between fracture growth and birth, therefore neglecting fracture-to-fracture positioning correlations. In this paper, we pushed further the model by improving the nucleation process, conditioning the position of newly created fractures by the stress perturbation induced by preexisting ones, in a timewise process. This stress perturbation is a function of fractures geometry (size and orientation), and the applied remote stress (orientation and intensity). This results in more correlated networks, showing fractal positioning, and a higher variability of fracture densities. We introduce a selectivity parameter \( m \) that quantifies the dependency of nucleation with the stress field. When nucleation is stress-independent \( (m = 0) \), uniform positions, we show that fracture density variability associated to UFM networks is much smaller than equivalent Poisson model. This means that the UFM rule, imposing that a fracture cannot cross a larger one, tends to organize networks into more homogeneous patterns than if fractures were positioned at random. Nonetheless, when nucleation is stress-driven \( (m \neq 0) \), the higher \( m \), the lower the correlation dimension of fracture positions, and the higher the spatial variability of fracture densities. This effect is dependent on the density measure, i.e., on the dependency of the density with fracture size. It is higher for the number of fractures per unit volume than for the percolation parameter. Moreover, our connectivity analysis brings up that the UFM rule tends to create a hierarchy between fractures, so that fractures of the same size order are less likely to cross one each other. The stress-driven nucleation process we propose tends to connect small fractures all together with the largest ones, that are responsible for the main stress perturbations. We also show that UFM models have lower percentage of fractures involved in the backbone than their equivalent Poisson model [26], whatever the selectivity parameter.

Finally, constraining fractures orientation according to the computed local stress field, in order to account for fracture position-orientation correlations, would constitute a huge improvement of the model. Once a full simplified mechanical description of the fracturing process is performed, this would allow us to perform real case studies, and compare our analysis results (clustering, connectivity...) between 2D numerical outcrops from generated DFNs, with real ones.

DATA AVAILABILITY STATEMENT

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation, to any qualified researcher.
AUTHOR CONTRIBUTIONS

EL, PD, CD, and RM conceived the idea of this study. EL and PD contributed in developing the model. EL ran network analysis and took the lead in writing manuscript. All authors provided critical feedback.

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REFERENCES

1. Davy P, Le Goc R, Darcel C. A model of fracture nucleation, growth and arrest, and consequences for fracture density and scaling. J Geophys Res. (2013) 118:1393–407. doi: 10.1002/jgrb.50120
2. Jing L. A review of techniques, advances and outstanding issues in numerical modelling for rock mechanics and rock engineering. Int J Rock Mech Mining Sci. (2003) 40:283–353. doi: 10.1016/S1569-8994(03)00013-3
3. Lei Q, Latham J-P, Tsang C-F. The use of discrete fracture networks for modelling coupled geomechanical and hydrological behaviour of fractured rocks. Comput Geotech. (2017) 85:151–76. doi: 10.1016/j.compgeo.2016.12.024
4. Long J, Remer J, Wilson C, Witherspoon P. Porous media equivalents for networks of discontinuous fractures. Water Resour Res. (1982) 18:645–58. doi: 10.1029/WR018i003p00645
5. Baecher GB. Statistical analysis of rock mass fracturing. J Int Assoc Math Geol. (1983) 15:329–48. doi: 10.1007/BF01036074
6. Andersson J, Shapiro AM, Bear J. A stochastic model of a fractured rock conditioned by measured information. Water Resour Res. (1984) 20:79–88. doi: 10.1029/WR020i001p00079
7. Endo H, Long J, Wilson C, Witherspoon P. A model for investigating mechanical transport in fracture networks. Water Resour Res. (1984) 20:1390–400. doi: 10.1029/WR020i010p01390
8. Long J, Gilmour P, Witherspoon PA. A model for steady fluid flow in random three-dimensional networks of disc-shaped fractures. Water Resour Res. (1985) 21:1105–15. doi: 10.1029/WR021i008p01105
9. Andersson J, Dverstorp B. Conditional simulations of fluid flow in three-dimensional networks of discrete fractures. Water Resour Res. (1987) 23:1876–86. doi: 10.1029/WA023i010p01876
10. Ackermann RV, Schlische RW. Anticlustering of small normal fractures defined around faults. J Geol. (1997) 25:1127–30. doi: 10.1130/0091-7613(1997)025<0176:AOAF>2.3.CO;2
11. Bonneau F, Caumon G, Renard P. Impact of a stochastic sequential initiation of fractures on the spatial correlations and connectivity of discrete fracture networks. Water Resour Res. (2016) 52:8526–45. doi: 10.1002/2016WR018973
12. Atkinson BK, Meredith PG. The theory of subcritical crack growth with applications to minerals and rocks. Fracture mechanics of rock. (1987) 2:111–66. doi: 10.1016/B978-0-12-66266-1.50009-0
13. Engelder T. Joints and shear fractures in rock. Fract Mech Rock. (1987) 27–69.
14. Pollard DD, Aydin A. Progress in understanding jointing over the past century. Geol Soc Am Bull. (1988) 100:1181–204. doi: 10.1130/0016-7606(1988)100<1181:TJOT>2.3.CO;2
15. Tapponnier P, Brace W. Development of stress-induced microcracks in Westerly granite. Int J Rock Mech Min Sci Geomech Abstr. (1976) 13:103–12. doi: 10.1016/0148-9626(76)91937-9
16. Ingraffea AR. Theory of crack initiation and propagation in rock. Fract Mechanics of Rock. (1987) 10:93–4. doi: 10.1016/B978-0-12-66266-1.50008-9
17. Betekhtin V, Kadomtsev A. Evolution of microcracks and pores in solids under loading. Phys Solid State. (2005) 47:825–31. doi: 10.1134/1.1924839
18. Renhaw CE, Pollard DD. Numerical simulation of fracture set formation: a fracture mechanics model consistent with experimental observations. J Geophys Res. (1994) 99:9359–72. doi: 10.1029/94JB00139
19. Olson JE. Predicting fracture swarms—The influence of subcritical crack growth and the crack-tip process zone on joint spacing in rock. Geol Soc Lond Spec Publ. (2004) 231:73–88. doi: 10.1144/GSL.SP.2004.2.31.01.05
20. Bonneau F, Caumon G, Renard P. Impact of a stochastic sequential initiation of fractures on the spatial correlations and connectivity of discrete fracture networks. J Geophys Res. (2016) 121:5641–58. doi: 10.1002/2015JB012451
21. Charlot R. Dynamic fatigue of glass. J Appl Phys. (1958) 29:1657–62. doi: 10.1063/1.1723019
22. Mises RV. Mechanik der festen Körper im plastisch-deformablen Zustand. Nachrichten von der Gesellschaft der Wissenschaften zu Göttingen, Mathematisch-Physikalische Klasse. (1913) 1:582–92.
38. Fabrikant VI. Complete solutions to some mixed boundary value problems in elasticity. *Adv Appl Mech.* (1989) 27:153–223. doi: 10.1016/S0065-2156(08)70196-0

39. Kachanov M. Three-dimensional problems of strongly interacting arbitrarily located penny-shaped cracks. *Int J Fract.* (1989) 41:289–313. doi: 10.1007/BF00188661

40. Kachanov M. On the problems of crack interactions and crack coalescence. *Int J Fract.* (2003) 120:537–43. doi: 10.1023/A:1025448314409

41. Thomas RN, Paluszny A, Zimmerman RW. Quantification of fracture interaction using stress intensity factor variation maps. *J Geophys Res.* (2017) 122:7698–717. doi: 10.1002/2017JB014234

42. Dershowitz WS, Herda HH. Interpretation of fracture spacing and intensity. In: *The 33th US Symposium on Rock Mechanics (USRMS)* Santa Fe, NM: American Rock Mechanics Association (1992).

43. De Dreuzy JR, Davy P, Bour O. Percolation parameter and percolation-threshold estimates for three-dimensional random ellipses with widely scattered distributions of eccentricity and size. *Phys Rev E.* (2000) 62:5948–52. doi: 10.1103/PhysRevE.62.5948

44. Davy P. On the frequency-length distribution of the San Andreas fault system. *J Geophys Res.* (1993) 98:12141–51. doi: 10.1029/93JB00372

45. Bour O, Davy P. On the connectivity of three-dimensional fault networks. *Water Resour Res.* (1998) 34:2611–22. doi: 10.1029/98WR01861

46. Darcel BO, Davy P, de Dreuzy JR. Connectivity properties of two-dimensional fracture networks with stochastic fractal correlation. *Water Resour Res.* (2003) 39:1272. doi: 10.1029/2002WR001628

47. Hentschel H, Procaccia I. The infinite number of generalized dimensions of fractals and strange attractors. *Phys D.* (1983) 8:435–44. doi: 10.1016/0167-2789(83)90235-X

48. Moein MJA, Valley B, Evans KF. Scaling of fracture patterns in three deep boreholes and implications for constraining fractal discrete fracture network models. *Rock Mech Rock Eng.* (2019) 52:1723–43. doi: 10.1007/s00603-019-1739-7

49. Mandelbrot BB, Pignoni R. *The Fractal Geometry of Nature.* New York, NY: WH Freeman (1983). doi: 10.1119/1.13295

50. Plotnick RE, Gardner RH, Hargrove WW, Prestegaard K, Perlmutter M. Lacunarity analysis: a general technique for the analysis of spatial patterns. *Phys Rev E.* (1996) 53:5461. doi: 10.1103/PhysRevE.53.5461

51. Kaye BH. *A Random Walk Through Fractal Dimensions.* John Wiley & Sons (2008).

52. Roy A, Perfect E, Dunne WM, Odling N, Kim J-W. Lacunarity analysis of fracture networks: evidence for scale-dependent clustering. *J Struct Geol.* (2010) 32:1444–9. doi: 10.1016/j.jsg.2010.08.010

53. Roy A, Perfect E, Dunne WM, McKay LD. A technique for revealing scale-dependent patterns in fracture spacing data. *J Geophys Res.* (2014) 119:5979–83. doi: 10.1002/2013JB010647

54. Lavoine E, Davy P, Darcel C, Le Goc R. On the density variability of poissonian discrete fracture networks, with application to power-law fracture size distributions. *Adv Geosci.* (2019) 49:77–83. doi: 10.5194/adgeo-49-77-2019

55. Bour O, Davy P. Connectivity of random fault networks following a power law fault length distribution. *Water Resour Res.* (1997) 33:1567–83. doi: 10.1029/96WR00433

56. Stauffer D, Aharony A. *Introduction to Percolation Theory: Revised.* 2nd Edition. CRC Press (2014).

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