Summer temperature—but not growing season length—influences radial growth of *Salix arctica* in coastal Arctic tundra

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Abstract

Arctic climate change is leading to an advance of plant phenology (the timing of life history events) with uncertain impacts on tundra ecosystems. Although the lengthening of the growing season is thought to lead to increased plant growth, we have few studies of how plant phenology change is altering tundra plant productivity. Here, we test the correspondence between 14 years of *Salix arctica* phenology data and radial growth on Qikiqtaruk–Herschel Island, Yukon Territory, Canada. We analysed stems from 28 individuals using dendroecology and linear mixed-effect models to test the statistical power of growing season length and climate variables to individually predict radial growth. We found that summer temperature best explained annual variation in radial growth. We found no strong evidence that leaf emergence date, earlier leaf senescence date, or total growing season length had any direct or lagged effects on radial growth. Radial growth was also not explained by interannual variation in precipitation, MODIS surface greenness (NDVI), or sea ice concentration. Our results demonstrate that at this site, for the widely distributed species *S. arctica*, temperature—but not growing season length—influences radial growth. These findings challenge the assumption that advancing phenology and longer growing seasons will increase the productivity of all plant species in Arctic tundra ecosystems.

Keywords Phenology · Tundra · Shrub · Growth · Dendroecology

Introduction

The Arctic is warming three to four times faster than the rest of the planet (Meredith et al. 2019; You et al. 2021) and tundra plant communities are particularly sensitive to that warming (Elmendorf et al. 2015; Bjorkman et al. 2020). Climate change is resulting in a longer snow- and ice-free season, potentially facilitating longer growing seasons (Cleland et al. 2007; Khorsand Rosa et al. 2015; Prevèy et al. 2021; Frei and Henry 2021). Concurrent with these changes, shifts in distribution and abundance (Sturm et al. 2001; Elmendorf et al. 2012a), biomass (Hudson and Henry 2009), and phenology (timing of life history events) (Oberbauer et al. 2013) have been observed for species across the tundra. Previous research assumes that the altered phenology will correspond directly with increased growth of tundra plants (Myneni et al. 1997; Piao et al. 2007; Ernakovich et al. 2014; Park et al. 2016; Arndt et al. 2019; Kim et al. 2021). Arctic spectral greening trends from satellite vegetation indices, such as the Normalised Differential Vegetation Index (NDVI), are used as proxy metrics of tundra plant phenology (Piao et al. 2007; Zeng et al. 2013; Park et al. 2016) and plant productivity (Myneni et al. 1997; Kim et al. 2021). However, remote sensing studies of tundra phenology may not be capturing on-the-ground plant phenological and growth dynamics and may instead be influenced by other land surface changes such as community compositional shifts and snow cover (Helman 2018) particularly in Arctic systems (Arndt et al. 2019; Myers-Smith et al. 2020; Cao et al. 2020). Satellite (Myers-Smith et al. 2020) and in situ (Oberbauer et al. 2013) studies indicate widespread but variable phenological shifts in the Arctic. Remote sensing studies have linked phenology change with increased plant productivity
in tundra ecosystems (Myneni et al. 1997; Park et al. 2016; Kim et al. 2021); however, in situ studies directly linking plant phenology change to plant growth are rare.

Plant phenology is changing throughout the tundra (Myers-Smith et al. 2019; Bjorkman et al. 2020; Prevéy et al. 2021), but the consequences on plant growth remain unclear. Phenology defines the bounds for plant activity, including photosynthesis, and has shifted around the Arctic due to warming (Assmann et al. 2019; Myers-Smith and Smith 2020). The snow-free season across the Arctic has extended by 2–4 days per decade of warming (Piao et al. 2007; Barichivich et al. 2013; Park et al. 2016; Myers-Smith et al. 2019). Seasons are starting earlier and finishing earlier or later depending on the location and study metrics investigated (Piao et al. 2007; Zeng et al. 2011, 2013; Keenan and Richardson 2015; Park et al. 2016; Myers-Smith et al. 2019). Two key points in deciduous species’ phenology are leaf emergence and leaf senescence: the time between being the entire growing season. Leaf emergence and senescence are both shifting across the Arctic, leading to a longer, earlier growing season at many sites, although changing phenology is not uniform across sites or species (Oberbauer et al. 2013; Assmann et al. 2019; Myers-Smith and Smith 2020). Earlier leaf emergence is associated with earlier snowmelt (Assmann et al. 2019; Myers-Smith and Smith 2019) and declining sea ice (Post et al. 2009; Bhatt et al. 2010; Kerby and Post 2013), although some studies have identified trends towards later phenology in some species and locations usually aligning with later snowmelt (Wipf and Rixen 2010; Bjorkman et al. 2015). Earlier leaf emergence may expose individuals to late spring frost events (Wheeler et al. 2015) or other harsh conditions. Early senescence through deterministic leaf age (Oberbauer et al. 2013; Keenan and Richardson 2015), nutrient availability (Lim et al. 2007), or photoperiod (Arft et al. 1999) may also undermine any growth benefits of earlier leaf emergence. Whether plants can take advantage of an extended growing season to increase productivity and accumulate biomass is therefore uncertain.

At mid-latitudes of the Arctic, shrub growth can be particularly sensitive to climate (Myers-Smith et al. 2015a) and there is ground based and satellite evidence for rapid shrub expansion in the region (Fraser et al. 2011; Moffat et al. 2016; Myers-Smith et al. 2019). Dendroecology allows for the exploration of the growth history of shrubs based on the width of rings formed during seasonal woody tissue deposition (Myers-Smith et al. 2015b). Individual annual growth ring chronologies can be compared with environmental variables to reveal the climate sensitivity of radial growth over time. Through dendroecology, we can directly observe how changing conditions affect shrub growth, validating assumptions and models. Individual growth is a key element in our understanding of shrub expansion throughout the Arctic (Tape et al. 2006; Myers-Smith et al. 2011a; Elmendorf et al. 2012b; Myers-Smith et al. 2019; García Criado et al. 2020). Increasing shrub cover and canopy height alters ecosystem processes and species interactions (Myers-Smith et al. 2011a; Tape et al. 2016, 2018; Way and Lapalme 2021) through snow trapping (Myers-Smith and Hik 2013), shading (Bloq et al. 2010), hydrology and albedo (Sturm et al. 2005; Pomeroy et al. 2006), food webs (Ravolainen et al. 2014)—including soil assemblages (DeMarco et al. 2014)—and habitat provision for wildlife (Boelman et al. 2015). Shrub encroachment has been linked to warming in studies using dendrochronology (Forbes et al. 2010), remote sensing (Myneni et al. 1997; Myers-Smith et al. 2020), field observations (Hudson and Henry 2009; Myers-Smith et al. 2011b), and experiments (Elmendorf et al. 2012a, 2015; Khorsand Rosa et al. 2015; Frei and Henry 2021). To accurately predict the future structure and function of northern ecosystems, we must understand how plant growth is changing (Myers-Smith et al. 2020), especially the role of phenology as ecological dynamics change under warming (Keenan and Richardson 2015; Myers-Smith et al. 2019; Bjorkman et al. 2020; Samplonius et al. 2020).

Arctic vegetation change plays a key role in regional and global feedback loops (Liston et al. 2002; Sturm et al. 2005; Pearson et al. 2013; Grosse et al. 2016) and carbon budgets (Piao et al. 2007; McGuire et al. 2009; Parker et al. 2021). As phenology changes, we expect compositional shifts and increased growth during longer growing seasons (Myneni et al. 1997; Ernakovich et al. 2014; Panchen and Gorelick 2017). Biome-wide shifts in growth rates and community composition could have profound implications for global carbon budgets through biomass accumulation (Piao et al. 2007) and decomposition (DeMarco et al. 2014). Warming drives earlier leaf emergence (Ernakovich et al. 2014; Park et al. 2016), which has been linked with increased plant productivity using remote sensing observations (Myneni et al. 1997). And thus, studies of satellite-derived spectral greening trends have linked changes in phenology to changes in plant productivity (Myneni et al. 1997; Park et al. 2016; Kim et al. 2021). Furthermore, accurate Earth system models depend on our understanding of plant growth–climate relationships and ecosystem–climate feedbacks (Sturm et al. 2005; Loranty and Goetz 2012; Richardson et al. 2013; Pearson et al. 2013; Fisher et al. 2018; Bonan and Doney 2018). Despite underpinning global models, uncertainty remains in the expected association between phenology and growth of Arctic plants and whether warmer temperatures or longer growing seasons are the primary drivers of increasing tundra plant productivity.

In this study, we use dendroecology to test the correspondence between in situ phenology observations, environmental factors, and radial growth of Salix arctica Pall. (Salicaceae) on Qikiqtaruk–Herschel Island in the Western Canadian Arctic. Salix arctica has a circum-Arctic
distribution, woody tissues which enable dendrology, and exceptionally closely monitored phenology at the site (Myers-Smith et al. 2019), offering a suitable species for this study. At this site, S. arctica phenology has advanced in both spring and autumn, although autumn only marginally, overall lengthening the growing season by 2 days per decade (Myers-Smith et al. 2019). We test three questions: (1) Do longer growing seasons facilitate greater shrub radial growth? (2) Of phenological metrics, does leaf emergence date, senescence date, or growing season length best explain shrub radial growth? (3) Do climatic factors (air temperature, precipitation, sea ice concentration, or snowmelt), or maximum plant productivity (estimated through the spectral Normalised Difference Vegetation Index, NDVI), explain shrub radial growth better than phenology? We hypothesise that (1) a longer realised period of growth will increase radial growth; (2) growing season length will explain shrub radial growth better than leaf emergence or senescence date as it encompasses the cumulative change in the growth period; and (3) growing season length will best explain shrub radial growth relative to temperatures or other variables as short growing seasons in Arctic systems are limiting plant growth (Myers-Smith et al. 2019).

**Methods**

**Study site**

We studied S. arctica phenology and radial growth on Qikiqtaruk–Herschel Island, YT, Canada (69.57 °N, 138.90 °W) (Fig. 1). The island is approximately 100 km² in area, with soils formed of marine and glacial deposits atop ice-rich permafrost (Burn and Zhang 2009). Qikiqtaruk sits at the northerly extent of tall shrubs, particularly Salix richardsonii Hook. (Salicaceae) (Myers-Smith et al. 2011b), which feature heavily in its flora alongside Eriophorum vaginatum L. (Cyperaceae) tussock tundra and dwarf shrub heath rich in S. arctica (Myers-Smith et al. 2019).

**Salix arctica sampling**

As woody perennials, shrubs grow annual rings of wood which record radial growth over time (Myers-Smith et al. 2015b). Here, we focus on S. arctica, a prostrate willow with a circum-Arctic range which reaches as far north as the north coast of Greenland (Argus 2007). We collected 38 shrub samples on a coastal floodplain on the east side of the island in the Ice Creek watershed, a site of known vegetation change (Myers-Smith et al. 2019). The vegetation is dominated by Salix spp., defined by patches of S. richardsonii, a canopy-forming willow, with S. arctica at ground level. Cross-sections of 3–5 cm in length were taken from the base of the thickest stem of 6–8 individuals each along five parallel transects on the Qikiqtaruk floodplain in July 2016. Individuals were at least 10 m apart and transects were 50–100 m apart to lower the risk of sampling clones, particularly given the sprawling prostrate growth form of S. arctica (Argus 2007).

**Dendroecological methods**

We sliced thin sections (~25 µm) of each sample with a sledge microtome, then photographed (Online Resource 1), and measured the rings along four radii per sample using the ObjectJ package (1.04a) for ImageJ (2.0.0-rc-59/1.51j).
We visually crossdated samples to check for partly missing rings, then averaged the radii for each individual, and crossdated again to check for entirely missing rings between samples. We removed the first 2 years of radial growth data for each individual to account for age-related growth effects and the data from the year of sampling, as the growth for the season was not completed (Myers-Smith et al. 2015b). To maintain a minimum of 4 years of radial growth data after removing the data from 2016 and the first 2 years, the sample size was reduced to 28 individuals, running from 2002 to 2015 (Online Resource 1). Individuals with fewer than 7 years of radial growth were removed and we calculated the basal area increments from the ring width data. We then detrended the basal area increment data, fitting a smoothing spline (dplR package for R, \( f = 0.5, \text{nyrs} = 0.67 \)), which removed variation in radial growth beyond interannual variation. We visually assessed different detrending methods, testing spline, negative exponential, and no detrending and found that spline detrending best removed individual growth trends to help focus our analyses on interannual variation in radial growth (Myers-Smith et al. 2015b). We therefore used the detrended basal area increment data for the main text analysis. We also conducted the same analysis using ring widths and the negative exponential detrending method and found similar results to our main text analysis, and these are presented in the supplementary materials (Online Resource 3). For the statistical analysis, all variables were normalised between 1 and −1, so that effect sizes could be compared.

**Phenology data**

Phenological observations have been collected since 2001 by Yukon Parks Rangers on Qikiqtaruk every 2–3 days from April (snowmelt) until September (leaf senescence) along established transects (~250 m from our study site). The rangers record phenology throughout the growing season (Myers-Smith et al. 2019) in line with ITEX protocol. The data from the year of sampling, as the growth for the season was not completed (Myers-Smith et al. 2015b). To maintain a minimum of 4 years of radial growth data after removing the data from 2016 and the first 2 years, the sample size was reduced to 28 individuals, running from 2002 to 2015 (Online Resource 1). Individuals with fewer than 7 years of radial growth were removed and we calculated the basal area increments from the ring width data. We then detrended the basal area increment data, fitting a smoothing spline (dplR package for R, \( f = 0.5, \text{nyrs} = 0.67 \)), which removed variation in radial growth beyond interannual variation. We visually assessed different detrending methods, testing spline, negative exponential, and no detrending and found that spline detrending best removed individual growth trends to help focus our analyses on interannual variation in radial growth (Myers-Smith et al. 2015b). We therefore used the detrended basal area increment data for the main text analysis. We also conducted the same analysis using ring widths and the negative exponential detrending method and found similar results to our main text analysis, and these are presented in the supplementary materials (Online Resource 3). For the statistical analysis, all variables were normalised between 1 and −1, so that effect sizes could be compared.

Other environmental data

We also compared radial growth to seasonal air temperatures and precipitation, snowmelt, sea ice concentration, and productivity. The temperature data came from Environment Canada Qikiqtaruk–Herschel Island weather station (ID 1560) and precipitation data came from the ERA5-gridded dataset produced by Copernicus Climate Change Service and the European Centre for Medium-Range Weather Forecasts (Hersbach et al. 2020). We collated data into seasons (spring: April–May, summer: June–July, autumn: August–September, winter: October–March) including the lagged data for the preceding summer and autumn, as monthly resolution was higher than useful for this study. Snowmelt data are from the Qikiqtaruk phenology dataset, where the date at which transects are free of snow is recorded (Myers-Smith et al. 2019). We used the onset of sea ice melt data (Assmann et al. 2019), determined using the NOAA/NSIDC Climate Data Record (CDR) v3 Passive Microwave Sea Ice Concentrations (Meier et al. 2017). Productivity data are the annual maxima (estimated by smoothing trends in the data with a generalised additive model) of the MODIS MOD13A1v6 NDVI satellite dataset (Myers-Smith et al. 2020).

**Statistical analysis**

We selected our models a priori, using single-predictor models to compare individual variables and assess their predictive power on shrub radial growth. We used a Bayesian framework for our analyses including weakly informative priors of a half Student t prior with three degrees of freedom. The effect sizes of models were assessed relative to their credible intervals (95%). If the credible intervals for the estimated slope did not cross zero, we considered an effect to be significant. If the credible interval of the model slope sits at zero or fluctuates between overlapping zero and not overlapping zero, we consider the effect to be marginally significant.

To test the relationships between variables we used hierarchical linear mixed-effect models, with year, transect, and individual as random effects and individuals nested within transects. We used mixed models due the hierarchical structure of our data, caused by non-independence of individuals within transects and within a given year. Shrubs also shared conditions in each year, making them non-independent. The variability in absolute growth amongst individuals was high, as seen during crossdating, but due to sampling a single species within a relatively small area, similar relative growth...
responses were expected across groups. Thus, we chose not to use random slopes, only random intercepts. As a secondary analysis presented in the supplementary materials (Online Resource 2, Online Resource 3), we used a frequentist framework and compared models using AIC to see whether the models fit the data better than a null model using the conventional threshold ($\Delta AIC_{null} \geq 2$) (Akaike 1974). We also calculated conditional and marginal pseudo-$R^2$ to test the absolute model fit including and excluding random effects, respectively (Nakagawa and Schielzeth 2013). We used models with maximum likelihood estimation for AIC comparisons and models with restricted maximum likelihood estimation for pseudo-$R^2$ and effect size values. Residuals of models were visually assessed for normality with fitted residual plots and temporal autocorrelation (first- or second-order) with correlograms (Online Resource 2). We did not detect a signal of temporal autocorrelation, and residuals were similar across all models (Online Resource 2). We tested the correlation amongst all environmental and phenological variables and correlation coefficients varied between −0.76 and 0.75 (Online Resource 1).

All statistical tests were carried out in R (3.6.3), via RStudio (1.2.1335), including the brms package for Bayesian analysis: Code and data are available at the following GitHub repository: https://github.com/ShrubHub/ShrubRingPhenoHub.

### Results

Contrary to our first hypothesis that a longer realised period of growth will increase radial growth, we found that growing season length had no effect on *S. arctica* basal area increments, nor was there a lagged effect from the previous growing season (Table 1, Fig. 2). Contrary to our second hypothesis that growing season length will explain shrub radial growth better than leaf emergence or senescence date, it was leaf senescence date that was the best predictor of radial growth amongst these variables. Radial growth was significantly greater in years with an earlier leaf senescence date (negative effect), and we detected no relationship between radial growth and leaf emergence date, growing season length, or previous growing season length (Table 1, Fig. 2). Contrary to our third hypothesis that growing season length will best explain shrub radial growth relative to temperatures and other variables, we found that rather than growing season length, summer temperature was the best predictor

### Table 1

| Predictor variable                  | Estimate | Standard Error | Lower 95% CI | Lower 95% CI | R-hat | Bulk ESS | Tail ESS | Effect |
|-------------------------------------|----------|----------------|--------------|--------------|-------|----------|----------|--------|
| Leaf emergence                      | −0.04    | 0.06           | −0.16        | 0.08         | 1.00  | 2494     | 2063     | Fixed  |
| Leaf senescence                     | −0.10    | 0.04           | −0.18        | −0.01        | 1.00  | 2613     | 1760     | Fixed  |
| Growing season length               | −0.05    | 0.06           | −0.18        | 0.07         | 1.00  | 2268     | 1700     | Fixed  |
| Previous growing season length      | −0.02    | 0.06           | −0.13        | 0.09         | 1.00  | 3014     | 2128     | Fixed  |
| Previous summer temperature         | 0.00     | 0.06           | −0.12        | 0.13         | 1.00  | 2905     | 1892     | Fixed  |
| Previous autumn temperature         | 0.09     | 0.05           | −0.01        | 0.20         | 1.00  | 2743     | 2005     | Fixed  |
| Winter temperature                  | −0.02    | 0.05           | −0.12        | 0.07         | 1.00  | 1887     | 1917     | Fixed  |
| Spring temperature                  | 0.02     | 0.05           | −0.09        | 0.12         | 1.00  | 2618     | 2475     | Fixed  |
| Summer temperature                  | 0.12     | 0.05           | 0.01         | 0.22         | 1.00  | 3546     | 2036     | Fixed  |
| Autumn temperature*                 | 0.07     | 0.06           | −0.04        | 0.18         | 1.00  | 2968     | 2152     | Fixed  |
| Previous summer precipitation       | −0.03    | 0.06           | −0.14        | 0.08         | 1.00  | 2698     | 2094     | Fixed  |
| Previous autumn precipitation       | −0.01    | 0.05           | −0.11        | 0.09         | 1.00  | 2389     | 2016     | Fixed  |
| Winter precipitation                | −0.01    | 0.06           | −0.13        | 0.12         | 1.00  | 2999     | 2697     | Fixed  |
| Spring precipitation                | −0.01    | 0.05           | −0.10        | 0.08         | 1.00  | 2559     | 1856     | Fixed  |
| Summer precipitation                | 0.02     | 0.06           | −0.10        | 0.13         | 1.00  | 2681     | 2159     | Fixed  |
| Autumn precipitation                | −0.01    | 0.06           | −0.12        | 0.10         | 1.00  | 2781     | 2215     | Fixed  |
| MODIS NDVI                          | 0.01     | 0.06           | −0.10        | 0.12         | 1.00  | 2164     | 1501     | Fixed  |
| Minimum sea ice extent              | −0.14    | 0.12           | −0.38        | 0.11         | 1.00  | 3588     | 2755     | Fixed  |
| Sea ice melt onset date             | −0.10    | 0.06           | −0.22        | 0.03         | 1.00  | 2373     | 2286     | Fixed  |
| Date snow free                      | −0.04    | 0.09           | −0.21        | 0.13         | 1.00  | 1493     | 724      | Fixed  |

All models span the period 2002–2015. Asterisk (*) indicates a model where the 95% credible intervals do not overlap zero and which we therefore consider to be significant. Sample depth per year as follows: 2002, n = 5; 2003, n = 6; 2004, n = 8; 2005, n = 10; 2006, n = 14; 2007, n = 16; 2008, n = 20; 2009, n = 23; 2010, n = 25; 2011, n = 27; and 2012–5, n = 28 (Online Resource 1). Only the fixed effects are shown here, the results for the random effects are available in Online Resource 2.
amongst all variables. Summer temperature explained 2.8% of the variation in radial growth (Fig. 3, marginal pseudo-$R^2$, Online Resource 2), with higher temperature coinciding with higher annual radial growth. However, the overall explanatory power of the models was low with no model explaining more than 2.8% of variation without including random effects (marginal pseudo-$R^2$, Online Resource 2).

All other variables aside from leaf senescence date and summer temperature were non-significant predictors of radial growth (Fig. 3). Leaf emergence date; current and previous years’ growing season length; temperatures from the winter, spring, autumn, and the previous year; all precipitation models; and snowmelt date did not explain variation in radial growth (Fig. 3). We used year as a random effect in our models (individual and transect level growth variation are accounted for during detrending and scaling of the radial growth index) and its effect was significant, indicating variation in radial growth amongst years beyond the effect of the bioclimatic variable of interest in that year (Online Resource 2, Online Resource 3). The model results agreed between the
different detrending approaches, and although exact effect sizes differed slightly, the results were similar when models used detrended ring width data or used detrended basal area increment (Online Resource 3).

Discussion

Through a unique study of long-term in situ phenology monitoring and dendroecology, we compared interannual variation in phenology, environmental conditions, and NDVI to interannual variation in radial growth of *S. arctica* on Qikiqtaruk. We found that summer temperatures and leaf senescence—but not leaf emergence or growing season length—explained variation in radial growth for the widespread Arctic shrub *S. arctica* (Table 1, Fig. 3). Precipitation, sea ice, snowmelt, and NDVI did not correspond with variability in interannual radial growth in our study. Thus, we did not find support for the hypothesised relationship between phenology and radial growth. Our results suggest that factors other than the timing of the growing period from leaf emergence through senescence, such as temperature, can exert a larger influence on shrub growth in this tundra ecosystem. These findings ultimately have implications for how tundra shrub growth is modelled and thus the projection of Arctic carbon budgets.

Phenology

We found no evidence that earlier leaf emergence and longer growing seasons corresponded with increased radial growth in *S. arctica*, including growth in the following year (Table 1, Fig. 3). Results for preliminary analyses including other willow species from this and other sites have reached similar conclusions (Angers-Blondin 2019). Earlier leaf emergence did not result in greater *S. arctica* radial growth; however, we did find evidence of greater radial growth in years with earlier leaf senescence. Although we cannot identify a particular biological mechanism linking earlier leaf senescence with enhanced radial growth, early leaf senescence was correlated with warmer summer temperatures (Pearson’s product–moment correlation, *df*= 263, *p* < 0.001, *ρ* = −0.60, Online Resource 1). Individuals may reach a threshold after intense early growth (Rumpf et al. 2014), allowing for early cessation of growth above ground. Alternatively, growth and leaf age could be deterministic (Oberbauer et al. 2013; Keenan and Richardson 2015; Semenchuk et al. 2016; Parker et al. 2017), with growth...
ending at a fixed time after growth begins each year. Or earlier leaf senescence could occur in years with warmer temperatures without a mechanistic link between the two variables. Our findings are in line with previous evidence that the timing of tundra plant senescence is driven at least in part by non-climatic factors (Arft et al. 1999; Oberbauer et al. 2013). Taken together, our results suggest that shifts to earlier shrub leaf emergence and longer growing seasons are not necessarily driving changes in tundra shrub growth, contrary to interpretations of satellite remote sensing data (Myneni et al. 1997; Zeng et al. 2011, 2013; Arndt et al. 2019) and reviews (Ernakovich et al. 2014).

**Temperature**

We found that higher summer temperatures increased the radial growth of *S. arctica* at our site. The summer is the peak season for growth and individuals are sensitive to warming in this period (Andreu-Hayles et al. 2020), as observed across the biome (Myers-Smith et al. 2015a; Myers-Smith and Hik 2018) from dendrochronology (Forbes et al. 2010; Blok et al. 2011; Myers-Smith et al. 2011a; Li et al. 2016; Weijers et al. 2018; Le Moullec et al. 2019; Prendin et al. 2022), repeat photography (Sturm et al. 2001; Tape et al. 2006), and experiments (Elmendorf et al. 2012a, 2015; Khorsand Rosa et al. 2015; Frei and Henry 2021). Temperature–growth relationships are heterogeneous across the tundra biome, with relatively low climate sensitivity observed on Qikiqtaruk compared with other mid-latitude tundra locations (Myers-Smith et al. 2015a). Growth response to temperature has decreased over time in another dwarf willow species, *Salix polaris* Wahlenb. (Salicaceae), at Bjørnøya, Svalbard (Owczarek et al. 2021), suggesting that growth responses may not be fixed over long timescales. The growth response to early leaf senescence suggests the importance of resource accumulation for growth in the following season, shrubs may senesce above ground but remain active below ground for longer periods. There is increasing evidence that above-ground phenology may be asynchronous with below-ground root growth (Blume-Werry et al. 2016, 2017; Ögren 2017; Liu et al. 2022), though root phenology itself may not respond to autumn warming (Schwieger et al. 2018). Snow cover insulates shrubs from winter and spring temperatures (Kelsey et al. 2020; Rixen et al. 2022), and Krab et al. (2018) found diverging shrub radial growth responses to winter temperature, spring warming, and snowmelt amongst species. *Vaccinium vitis-idaea* L. (Ericaceae) grew more with delayed snowmelt with a contrasting reduction in growth in *Empetrum nigrum* L. (Ericaceae). We, however, found no association between temperatures in the previous year and radial growth, and no relationship for winter, spring, and autumn temperatures and radial growth (Table 1, Fig. 3).

**Hydrology**

We did not find a strong influence of summer precipitation, sea ice, or snowmelt on interannual variation in radial growth *S. arctica* in this study. Growth of Arctic shrubs can be moisture limited (Keuper et al. 2012; Ackerman et al. 2017; Buchwal et al. 2020; Weijers 2022). Moisture sensitivity of growth can depend on temperature (Li et al. 2016) and can vary within (Thompson and Koenig 2018) and between sites (Myers-Smith et al. 2015a). Soils on Qikiqtaruk are frequently saturated, likely reducing the impacts of drought locally (Myers-Smith et al. 2019). We did not detect any influence of precipitation from summer rain, snowmelt, or cloud cover (Table 1, Fig. 3). Decreasing snow cover reduces soil insulation in winter and limits productivity increase under warming and earlier phenology in Alaska (Kelsey et al. 2020). The lack of a precipitation signal detected in our study could be influenced by our use of gridded climate datasets due to a lack of a complete local record for precipitation at this site. Gridded climate datasets poorly capture spatially variable precipitation, due to the paucity of Arctic meteorological stations and the high spatial variability of precipitation (Macias-Fauria et al. 2014; Myers-Smith and Myers 2018). For sea ice, we found that lower annual minima and earlier melt are weakly associated with increased radial growth of *S. arctica*, although phenology for this species was not found to vary with sea ice extent (Assmann et al. 2019). Sea ice could influence plant growth and phenology through interactions with local climate (Post et al. 2009; Bhatt et al. 2010; Kerby and Post 2013; Macias-Fauria et al. 2017; Assmann et al. 2019) and drought-stress (Forchhammer 2017; Buchwal et al. 2020). We found no relationship between snowmelt date and radial growth, which is consistent with the primary mechanism of snowmelt controlling phenology and so influencing plant growth (Assmann et al. 2019; Myers-Smith et al. 2019). Taken together, our results suggest that temperature rather than growing season length, precipitation, or sea ice dynamics was the primary factor controlling *S. arctica* radial growth on Qikiqtaruk.

**NDVI**

We observed no correlation between NDVI and interannual variation in *S. arctica* radial growth, consistent with results for other shrub species at the site (Myers-Smith et al. 2019). Whilst satellite datasets do not entirely correspond with each other (Guay et al. 2014), and shrub biomass cannot be directly estimated from NDVI alone (Cunliffe et al. 2020), NDVI is easily scaled, well studied, and part of a broader picture of complex Arctic tundra vegetation change (Myers-Smith et al. 2020). Arctic shrubification has been linked with satellite-derived Arctic greening trends (Macias-Fauria et al. 2012), and comparing ground observations to spectral
greening observed by satellites improves broad-scale interpretation of these trends (Myers-Smith et al. 2020). Correlation of NDVI and shrub growth has been found in some studies (Forbes et al. 2010; Macías-Fauria et al. 2012), but is not universal and varies with site and the time of year (Blok et al. 2011; Brehaut and Danby 2018; Andreu-Hayles et al. 2020). Taken together, these results suggest that satellite spectral greening indices are not capturing all of the variation in plant productivity, including the length of the snow-free season, indicated by analyses with shrub radial growth (Angers-Blondin 2019; Berner et al. 2020).

**Study limitations**

Whilst our findings bring together phenology and dendrochronology, two important fields of study of Arctic change, there are limitations. Sampling stem elongation (primary growth) and root collars rather than stems alone would improve the capture of interannual variation in shrub productivity. Primary and secondary (radial) growth can be driven by different controls (Bret-Harte et al. 2002; Campioli et al. 2012a, b), so study of annual stem increments or other measures of shrub growth in addition to radial growth would more robustly address questions of shrub growth (Myers-Smith et al. 2015b). Root collars show greater climate sensitivity (Ropars et al. 2017) and less response to individual conditions than stems (Sonesson and Callaghan 1991; Sadras and Denison 2009; Myers-Smith et al. 2015b), yet root collars are challenging to find and excavate in clonal species, such as *S. arctica*, and are more destructive to sample. We were not easily able to locate root collars consistently at this site (Angers-Blondin 2019). The destructive nature of dendrological sampling also prevented us from sampling the individuals in the long-term phenology transect directly, so we sampled nearby individuals as phenology is consistent across the site (Myers-Smith et al. 2019). Future research across different sites and species using localised climate and microenvironmental variables may shed more light on the relationships between plant phenology and growth.

**Future study**

Modern techniques facilitate below-ground monitoring of tundra plant phenology and root growth (Iversen et al. 2015; Sloan et al. 2016; Blume-Werry et al. 2016, 2017), exposing an overlooked dimension of tundra dynamics. An increasing number of studies indicate phenological asynchrony above and below ground at sub-Arctic sites (Ögren 2017; Blume-Werry et al. 2017), with below-ground root growth extending into the late summer and autumn in now thawed soils. Fungal symbiotes such as mycorrhizae can influence plant growth and carbon exchange in tundra shrubs which could be altering growth–climate interactions (Clemmensen et al. 2006; Compan et al. 2010; Deslippe et al. 2011). Iler et al. (2013) suggest that phenology responses to warming are reaching physiological limits in some Arctic and alpine species, potentially reducing the magnitude of future change. Collins et al. (2021) found that reproductive and vegetative phenologies are affected differently by experimental warming, which could alter ecosystem dynamics via trophic mismatches and resource allocation (Post and Forchhammer 2008; Clausen and Clausen 2013; Kerby 2015; Wheeler et al. 2015). There has been relatively little investigation of plant senescence and the drivers of the end of the growing season, creating uncertainty in our understanding of plant responses to warming across the growing season.

Innovative techniques such as drone-derived biomass estimates could also help with scaling up to landscape-wide analyses (Cunliffe et al. 2020). Newer approaches to studying tundra plant phenology such as time lapse cameras (a.k.a. phenocams) are overcoming inherent challenges of data collection in the Arctic (Westergaard-Nielsen et al. 2017; Richardson et al. 2018; Parmentier et al. 2021). Local observations of plant phenology and growth can be scaled up using drone and satellite data to bridge scale gaps and form a landscape perspective on tundra productivity change (Riihimäki et al. 2019; Assmann et al. 2019, 2020; Cunliffe et al. 2020). Challenges of scaling and data collection are being met by technological solutions, allowing us to see Arctic change from new angles and more clearly than ever before. Although further research is required, particularly for Arctic systems (Diepstraten et al. 2018), the increasing scope of monitoring of above- and below-ground plant responses encompassing phenology and growth allows for the investigation of key knowledge gaps about tundra ecosystem responses to global change.

**Conclusion**

Our findings demonstrate that plant phenology does not necessarily predict growth in an Arctic shrub, but that warmer temperatures in the summer are associated with increased annual radial growth. Interannual variation in precipitation, sea ice, snow cover, and MODIS NDVI for the landscape were not strongly related to radial growth. Our results indicate that future Arctic warming will likely enhance shrub growth and encroachment (Tape et al. 2006; Myers-Smith et al. 2011a; García Criado et al. 2020). Where this growth is not limited by water or nutrients (Mack et al. 2004; Myers-Smith et al. 2015a; Ackerman et al. 2017), there may be significant consequences for water, energy, and carbon fluxes (Loranty and Goetz 2012; Pearson et al. 2013; Parker et al. 2021). Taller shrub canopies could influence soil temperatures, litter decomposition rates, nutrient cycling, and ultimately the tundra carbon cycle (Sturm et al. 2005;
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