Natural radioactivity and radiation hazard assessment of phosphate mining, Quseir-Safaga area, Central Eastern Desert, Egypt

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\textbf{Abstract} Quseir-Safaga area, Central Eastern Desert, Egypt, includes Duwi Formation, which contains uranium-bearing phosphate beds. The present work used the integrated carborne $\gamma$-ray spectrometric data, X-ray analysis and HPGe $\gamma$-ray spectrometer data to investigate the radioactive zones at this area. Carborne $\gamma$-ray spectrometric survey revealed an increase of equivalent uranium, up to 182 ppm. Maps were drawn to show the results of the three radionuclides K, eU, eTh and the eU/eTh ratio. The eU and (eU/eTh) maps reveal that there are twelve important anomalies, associated mainly with phosphate mines. The Hamrawein mines at the northwestern part in the study area seem to represent the highest U-anomalies.

Twelve phosphate samples were collected from the determined twelve anomalies. They were analyzed with HPGe detector gamma-ray spectrometer. It was found that the results of radionuclide concentrations by carborne survey agree well with those obtained by HPGe. Both of them show that phosphate mines effectively contribute to eU anomalies, occurring in the study area. Meanwhile, $^{40}$K, $^{238}$U and $^{232}$Th concentrations in phosphate samples range from 91 to 169, 864 to 3104 and 28.4 to 106 Bq/kg respectively. The highest concentration of $^{238}$U (3104 Bq/kg) occurs in the north of the studied area, close to Hamrawein city. The average concentration of $^{238}$U in the analyzed samples is 1766 Bq/kg, which is 53 times higher than the worldwide average value reaching 33 Bq/kg. The highest $^{232}$Th concentration value reaching 106 Bq/kg is 2.4 times higher than the worldwide value attaining 45 Bq/kg. The absorbed dose rate for the phosphate samples shows the highest value reaching 1468 nGy/h. This is more than 25 times the worldwide value of 58 nGy/h. The annual effective absorbed dose is also high and reached 1.8 mSv/y, which is about twice higher than the permitted value for public exposure of 1.0 mSv/y. The maximum external hazardous index value of 9.2 is more than nine times the unity permitted in ICRP, 2000.

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1. Introduction

Radioactivity is the emission of alpha (α) and beta (β) particles as well as gamma (γ) radiation from the unstable isotopes. Most minerals, that contain potassium, uranium and thorium, are radioactive. Long-live radioactive elements such as uranium (238U), thorium (232Th) and potassium (40K) and any of their decay products, such as radium and radon are examples of naturally occurring radioactive materials (NORM). Study of natural radioactivity is usually done in order to gain information about the present levels of harmful pollutants discharged to the environment itself or in the living creatures (UNSCEAR, 2000). Several studies were performed in different countries to determine the radionuclide concentrations in different food samples and dose assessments from consumption of foodstuff by the population (Shanthi et al., 2010; James et al., 2013).

Phosphates are used extensively, as a source of phosphorous for fertilizers and for manufacturing phosphoric acid and gypsum. Phosphate ores typically contain about 1500 Bq/kg of uranium and radium, although some phosphates contain up to 20,000 Bq/kg of U3O8 (Paschoa and Godoy, 2002). In general, phosphate ores of sedimentary origin have higher concentrations of radionuclides of the uranium family. In 90% of cases, the ore is treated with sulfuric acid. The fertilizers become somewhat enriched in uranium (up to 150% relative to the ore), while 80% of the 226Ra, 30% of 232Th and 5% of uranium are left in phosphogypsum (UNSCEAR, 2008).

The processing of phosphates may generate gaseous and particulate emissions that contain 235U and 226Ra, when discharged to the environment, lead to expose the population to radiation hazards. Local dump sites for phosphogypsum are usually not protected from rainfall and become hydraulically connected to surface water and shallow aquifers (Vandenhove, 2000). The use of phosphate fertilizers in agriculture and of gypsum in building materials is a further source of possible exposure to the public (Paschoa and Godoy, 2002). Elevated radon exposure to the public can further be expected in sites being developed for housing (Vandenhove, 2000).

2. Geologic setting

Quseir-Safaga area is located in the Central Eastern Desert of Egypt, between latitudes 25°58′N and 26°22′N and longitudes 34°03′E and 34°22′E (Fig. 1). The phosphate ore deposits of Egypt represent a part of the Mediterranean Phosphate Province, which extends from Morocco to Turkey. They consist of strata deposited on the shallow Tethyan seafloor, during Late Cretaceous and Early Tertiary times. Over a zone at central eastern desert of Egypt, between 24° and 27°N latitudes, these strata are nearly horizontal, but in the Red Sea area folding and block faulting have resulted in tilting of the mineable beds. The phosphate beds were estimated to contain in excess of 3 billion tons of ore (Notholt, 1985). In the region of the Eastern Desert of Egypt (Red Sea Coast), the phosphate-bearing strata are known as Duwi Formation, in the Nile Valley region as the Sibaiya Formation.

![Figure 1](image_url) Location map of Quseir-Safaga area, Central Eastern Desert, Egypt.
Phosphate Formation and in Western Desert as Phosphate Formation (Awad and Ghobrial, 1966). These beds rest with seeming conformity on the Nubia sandstones and are overlain by deeper-water marls and chalks of the Dakhla Formation (Fig. 2).

The mineralogy of the Egyptian phosphate deposits was studied by many investigators (Dabous, 1981; Sharafeldin, 1999). The major phosphate mineral is carbonate-fluorapatite. The non-phosphatic minerals are represented by dolomite, calcite, quartz, pyrite, goethite, gypsum, smectite, and kaolinite. The clay minerals, associated with Egyptian phosphates, occur mainly as a matrix or a filling of the intergranular pores. Some clay minerals are also present as parts of the phosphatic pellets, or as fillings of the cavity-like microstructures in the bone and teeth fragments. Smectite is the dominant clay mineral, followed by kaolinite, whereas illite and palygorskite are present in trace concentrations (Sharafeldin, 1999). This mode of occurrence was also reported in the Miocene phosphates of Florida (Compton, 1997), the Cretaceous phosphates of Morocco (Prevot, 1990), and the Jordanian Cretaceous phosphates (Khoury et al., 1988). The iron minerals: goethite, hematite, and pyrite were found as coatings on the outer surfaces of phosphate pellets and parts of the cementing material (Dabous, 1981; Sharafeldin, 1999).

3. Carborne gamma-ray spectrometric surveys

The ability of gamma-ray spectrometry to map potassium, uranium, and thorium enrichment or depletion provides powerful exploration guidance in a wide variety of geological settings. Carborne gamma-ray spectrometry was used for source search (Ulvsand et al., 2002), contamination surveying and mapping (Karlsson et al., 2000; Sanderson et al., 2003) and prospecting (Grasty and Cox, 1997). Variations in the radioactivity of rocks may be useful for geological mapping (Aissa and Jubeli, 1997; Gaafar et al., 2016), for acquiring information on the distribution of radiation exposure rates (Raghuwanshi, 1992; Gaafar et al., 2014), and for environmental monitoring (Tauchid and Jubeli, 1991; Gaafar et al., 2012). Total gamma radioactivity can be used by health physicists for acquiring information on the distribution of radiation exposure rates (Raghuwanshi, 1992). Besides, a carborne radiometric system can be effectively used for an emergency
3.1. Spectrometric survey

The carborne gamma-ray spectrometric system that used in the present survey is composed of a GR-320 portable multichannel gamma-ray spectrometer and has typically a 4 l NaI(Tl) scintillation detector. The position for each measurement point was determined by a GPS instrument and recorded with the spectrometric data. The complete system, loaded in a field car, was used to radioactive mineral and environmental exploration after calibration on NMA 4 standard calibration pads. A lot of collected spectrometric data confirm that the system is stable and works reliably, and is a fast and advanced approach for radioactive mineral γ-ray spectrometric exploration. It can be used not only to determine the contents of natural radioactive elements in environment, but also to monitor nuclear pollution and emergency treatment in nuclear accidents significantly.

Gamma-ray spectrometry provides a direct measurement of the surface distribution of naturally occurring radioelements (K, U, and Th). Potassium is a major constituent of most rocks, while uranium and thorium are present in trace amounts, as mobile and immobile elements, respectively. As the concentrations of these radioelements vary between different rock types, the measured radioelement distribution can be reliably used to map and distinguish different lithologies (IAEA, 2010). The concentrations of K, eU and eTh and eU/eTh ratio are illustrated along the image maps as shown in Figs. 3–6. There are twelve eU anomalies distributed along the car traverses, from south to north direction, having the values of: 66, 35, 57, 98, 80, 77, 97, 143, 182, 152, 116 and 85 ppm respectively (Fig. 4 and Table 1). The eU/eTh ratio shows strong positive correlation with eU, suggesting uranium mobilization in the phosphate that located in the northwest part of the studied area. The strong positive correlation between eU/eTh ratio and eU and the negative correlation between eU/eTh ratio and eTh coincide with phosphate mines (Figs. 4–6). The prominent high eU/eTh ratios are associated with phosphates without any contribution to the other rock units. The strong increase of eU/eTh ratio with eU and not with eTh and K suggests post magmatic redistribution of uranium. This could be a favorable economic criterion for uranium mineralization, because uranium is likely to be remobilized and reconcentrated.
4. Results of X-ray analysis

This study deals with the estimation of 14 trace elements: (Cr, Cd, Co, Ni, Cu, Zn, Zr, Rb, Y, Ba, Pb, Sr, V and Nb) in the 12 phosphate samples using X-ray fluorescence technique. Some of them, such as phosphate, contain abundant trace elements (Bech et al., 2009). Determination of trace elements, present in environmental samples, is of considerable interest because of the increased awareness about environmental pollution and human health since some elements which are nutrients, can also act as toxin, if they are present above certain limits (Hayumbu et al., 1995). Phosphate ores of different geographical provenances can vary substantially in trace element contents. Therefore, in order to minimize environmental and health risks, it is important to know the phosphate trace element assemblages, and use suitable phosphates of known provenance with low potential toxic trace element contents.

The concentrations of trace elements in phosphate samples at Quseir-Safaga area were measured at NMA and are listed in Table 2. It is clear that the average concentrations (in ppm) of the trace elements for phosphate samples, in a decreasing order, are as follows: Sr (751) > Zn (297) > Cr (121) > Ba (99) > V (90) > Y (56) > Zr (15.1) > Ni (14.7) > Cu (13.4) > Pb (6.6). The average concentrations of trace elements in Quseir-Safaga samples are also shown in Table 2 and Fig. 7. It is clear that all samples contain high concentration of Strontium (Sr) and Zinc (Zn).

Table 2 shows the concentrations of Cr from 91 to 160 ppm, Ni from 5 to 27 ppm, Cu from 11 to 16 ppm, Zn from 47 to 1358 ppm, Zr from 9 to 24 ppm, Y from 40 to 121 ppm, Ba from 31 to 340 ppm, Pb from <2 to 24 ppm, Sr from 451 to 1047 ppm and V from 42 to 173 ppm. The lowest values of Ni, Cu, Zr and Ba are in sample No. Q12, Zr, Y, Pb and Sr are in Q7, Cr and Cu are in Q1 and V in Q3 and Q5. Meanwhile, the highest values of Cr, Ni and V are in sample Q8, Zr, Y, Ba and Sr in sample Q9, Cu in sample Q7, Zn in sample Q2 and Pb in sample Q1.

5. Gamma-ray spectroscopic analysis

Each phosphate sample was placed in a 500 g airtight PVC container. The inner lid was placed in and closed tightly with outer cap. The container was sealed hermetically and externally using cellophane tape and kept aside for about one
month to ensure equilibrium between Ra and its daughter products, before being taken for gamma-ray spectrometric analysis.

To estimate the activity levels of $^{238}\text{U}$, $^{232}\text{Th}$ and $^{40}\text{K}$ in the 12 samples, a high resolution gamma-ray detection system in the Laboratories of the Nuclear Physical Division, Faculty of Girls, Ain Shams University, was used for gamma-ray analysis. The detection system is basically a hyper pure germanium (HPGe), model No. GEM-15190, coaxial type detector with a serial No. 27-p-1876A recommended operating bias, negative 3 kV. The individual samples were placed on the detector manually during the work and each sample was analyzed for a time of 70,000 s. The $\gamma$-ray emitting radionuclides, specifically recorded, were as follows: $^{238}\text{U}$, $^{232}\text{Th}$, $^{226}\text{Ra}$ and $^{40}\text{K}$.

Natural radionuclides of relevance for this work are mainly gamma-ray emitting nuclei in the decay series of $^{238}\text{U}$ and $^{232}\text{Th}$, and single occurring $^{40}\text{K}$. Meanwhile, $^{40}\text{K}$ can be measured directly by its own gamma-rays, $^{238}\text{U}$ and $^{232}\text{Th}$ are not directly gamma-ray emitters, but they are possibly measured by the gamma-ray of their decay products. Decay products for $^{238}\text{U}$ ($^{214}\text{Pb}$: 295 & 352 keV; and $^{214}\text{Bi}$: 609 keV) and $^{232}\text{Th}$ ($^{228}\text{Ac}$: 209, 338 and 911 keV; $^{212}\text{Pb}$: 239 keV; $^{212}\text{Bi}$: 727 keV; and $^{208}\text{TI}$: 583 keV) were used for the measurements of their parents, assuming the decay series to be in secular equilibrium (Firestone and Shirley, 1998). Weighted averages of several decay products were used to estimate activity concentrations of $^{238}\text{U}$ and $^{232}\text{Th}$. The natural abundance of $^{232}\text{U}$ is only 0.72% of the total uranium content and hence was not considered in the present study.

5.1. Radiation health hazards

Radium equivalent activity ($\text{Ra}_{\text{eq}}$) is a widely used index in radiation health hazards. It can be obtained from the following equation (Ravisankar et al., 2012):

$$\text{Ra}_{\text{eq}} = A_{\text{Ra}} + 1.43A_{\text{Th}} + 0.077A_{\text{K}}$$  \hspace{1cm} (1)

where $A_{\text{Ra}}$, $A_{\text{Th}}$ and $A_{\text{K}}$ are the activities of $^{226}\text{Ra}$, $^{232}\text{Th}$ and $^{40}\text{K}$, respectively, in Bq/kg. The maximum value of $\text{Ra}_{\text{eq}}$ must be less than 370 Bq/kg, in order to keep the annual external absorbed dose safe and less than 1.0 mSv/y (UNSCEAR, 2000; USEPA, 2004).
The absorbed dose rate due to $\gamma$-radiation in air at 1 m above the ground surface was also calculated. The conversion factors used for ($D_R$) calculations are 0.462 nGy/h for $^{226}$Ra, 0.621 nGy/h for $^{232}$Th, and 0.0417 nGy/h for $^{40}$K. Accordingly, $D_R$ can be calculated as follows (UNSCEAR, 2000):

$$D_R = 0.462A_{Ra} + 0.621A_{Th} + 0.0417A_{K}$$  

(2)

where $A_{Ra}$, $A_{Th}$ and $A_{K}$ are the activities of $^{226}$Ra, $^{232}$Th and $^{40}$K, respectively, in Bq/kg.

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**Table 1** The concentrations of the three radioelements and eU/eTh ratio for the 12 representative samples from the 12 phosphate mining of Quseir-Safaga area, Central Eastern Desert, Egypt.

| Sample | Long. (E) | Lat. (N) | K (%) | eU (ppm) | eTh (ppm) | eU/eTh |
|--------|----------|----------|-------|----------|-----------|--------|
| Q1     | 34°10'22" 25°59'15" | 0.5 | 66    | 7.7   | 8.6       |
| Q2     | 34°16'20" 26°03'31"  | 0.2 | 35    | 3.2   | 10.9      |
| Q3     | 34°06'45" 26°05'20"  | 0.3 | 57    | 3.7   | 15.4      |
| Q4     | 34°07'40" 26°05'17"  | 0.3 | 98    | 5.4   | 18.1      |
| Q5     | 34°10'52" 26°11'17"  | 0.3 | 80    | 4.8   | 16.7      |
| Q6     | 34°10'33" 26°12'40"  | 0.6 | 77    | 8.6   | 9.0       |
| Q7     | 34°06'11" 26°17'34"  | 0.8 | 97    | 9.1   | 10.7      |
| Q8     | 34°06'12" 26°16'57"  | 0.3 | 143   | 7.3   | 19.6      |
| Q9     | 34°05'58" 26°17'42"  | 0.3 | 182   | 9.7   | 18.8      |
| Q10    | 34°05'33" 26°18'15"  | 0.3 | 152   | 8.3   | 18.3      |
| Q11    | 34°05'59" 26°19'26"  | 0.2 | 116   | 7.8   | 14.9      |
| Q12    | 34°05'43" 26°20'16"  | 0.3 | 85    | 6.1   | 13.9      |
| Q13    | 34°05'43" 26°20'16"  | 0.3 | 85    | 6.1   | 13.9      |
| Min    | 34°10'33" 25°59'15"  | 0.2 | 35    | 3.2   | 8.6       |
| Max    | 34°17'02" 26°20'16"  | 0.8 | 182   | 9.7   | 19.6      |
| Average| 34°11'17" 26°10'30" 0.37 | 90 | 6.8   | 14.6   |
The annual effective dose (Eff. Dose) is a concept that allows the radiation doses from different radionuclides and from different types and sources of radioactivity to be added. The annual effective dose rate (in mSv/y) can be calculated using the following equation (UNSCEAR, 2000):

\[
\text{Eff. Dose} = \frac{DR}{8760/C2} \times \frac{0.7}{C2} \times \frac{0.2}{C10} \times \frac{3}{C3}
\]

where 0.7 Sv/Gy is the conversion coefficient from absorbed dose to effective dose and 0.2 is the fraction of time spent outdoors.

The external hazard index \( (H_{ex}) \) considers only the external exposure risk due to \( \gamma \)-rays and can be obtained by the following relation (Bereka and Mathew, 1985):

\[
H_{ex} = \frac{A_{Ra}}{185} + \frac{A_{Th}}{259} + A_{K}/4810
\]

where \( A_{Ra}, A_{Th} \) and \( A_{K} \) are the activities of \( ^{226} \text{Ra}, ^{232} \text{Th} \) and \( ^{40} \text{K} \), respectively, in Bq/kg.

The internal exposure to \( ^{222} \text{Rn} \) and its daughter products is controlled by an internal hazard index \( H_{in} \) which is defined as

\[
H_{in} = \frac{A_{Ra}}{185} + \frac{A_{Th}}{259} + A_{K}/4810
\]

The mass exhalation rate \( E_{Rn} \) (Bq kg\(^{-1}\) s) of \( ^{222} \text{Rn} \) was determined through the following equation:

\[
E_{Rn} = \frac{A_{Rn}}{370} + \frac{A_{Th}}{259} + A_{K}/4810
\]

If the maximum concentration of radium is half that of the normal acceptable limit, then \( H_{in} \) will be ≤ 1 (Abdullah, 2010).

The representative level index can be calculated using the following relation (Ebaid and Bakr, 2012):

\[
I_{c} = \frac{1}{150} A_{Ra} + \frac{1}{100} A_{Th} + \frac{1}{1500} A_{K}
\]

where \( A_{Ra}, A_{Th} \) and \( A_{K} \) are the activities of \( ^{226} \text{Ra}, ^{232} \text{Th} \) and \( ^{40} \text{K} \), respectively, in Bq/kg.

The mass exhalation rate \( E_{Rn} \) (Bq kg\(^{-1}\) s) of \( ^{222} \text{Rn} \) was determined through the following equation:

\[
E_{Rn} = \frac{A_{Rn} \times A_{Th} \times \lambda_{Rn}}{1500}
\]

where \( \lambda_{Rn} \) is the decay constant \( ^{222} \text{Rn} \) (2.1 \times 10\(^{-6}\) per s) (Addo et al., 2013).

5.2. Results of natural radioactivity

The concentrations and distribution of radionuclides in the 12 studied samples from Quseir-Safaga area were determined.
using HPGe spectrometer to evaluate their environmental radioactivity.

The activity concentrations of $^{238}$U were measured from $^{234}$Pa (1001 keV), $^{226}$Ra from (186 keV), $^{214}$Pb from (352, 295 keV) and $^{214}$Bi from (609, 1120, 1765 keV); meanwhile, the activity concentration of $^{228}$Ac was measured from 911, 338.4 keV and Tl from 583, 2614 keV. They were used to evaluate the specific activity of $^{232}$Th series radionuclides. $^{40}$K was determined from 1460 keV photo peak. For the actinium series $^{235}$U -energies of 144 keV and 163 keV were taken to represent the $^{235}$U activity (Merdanoglu and Altinsoy, 2006).

Table 3 shows the activity concentrations of the radionuclides of the 12 Quseir-Safaga phosphate samples. The radionuclides of $^{238}$U series are $^{234}$Pa, $^{226}$Ra, $^{214}$Bi and $^{214}$Pb while for $^{232}$Th series are $^{228}$Ac, $^{208}$Tl and $^{212}$Pb. The activity concentrations of $^{238}$U series are included in $^{234}$Pa that varies from 864 to 3104 Bq/kg, $^{226}$Ra varies from 963 to 3309 Bq/kg, $^{214}$Bi from 484 to 1584 Bq/kg, and $^{214}$Pb from 477 to 1671 Bq/kg. Meanwhile, $^{232}$Th series has values of $^{228}$Ac that ranges from 19 to 198 Bq/kg, $^{208}$Tl from 58 to 162 Bq/kg, and $^{212}$Pb from 2.04 to 51 Bq/kg. $^{235}$U was found to vary from 26 to 102 Bq/kg.

The activity concentrations of $^{238}$U range from 864 to 3104 Bq/kg, $^{232}$Th from 28.4 to 106 Bq/kg, $^{226}$Ra from 963 to 3309 Bq/kg, $^{40}$K from 91 to 169 Bq/kg and $^{235}$U from 26 to 102 Bq/kg (Table 4 and Fig. 8).

The radioelements worldwide average (W.A.) values are 33 Bq/kg for $^{238}$U, 32 Bq/kg for $^{226}$Ra, 45 Bq/kg for $^{232}$Th and 412 Bq/kg for $^{40}$K (UNSCEAR, 2008). The data on worldwide average outdoor dose rates value is 58 nGy/h (UNSCEAR, 2008). The concentrations of all Quseir-Safaga phosphate samples are higher than the worldwide average (33 and 32 Bq/kg) for $^{238}$U and $^{226}$Ra respectively (Table 4 and Fig. 9). Meanwhile, the concentrations of six phosphate
samples are lower than the worldwide average (W.A.) of 45 Bq/kg for thorium while the samples (Q4, Q5, Q6, Q10, Q11 and Q12) have concentrations higher than the worldwide average. On the other hand, the concentrations of all Quseir-Safaga phosphate samples are lower than the worldwide level for potassium (412 Bq/kg) (Table 4 and Fig. 9).

Correlations between $^{238}$U and $^{226}$Ra, $^{238}$U and $^{40}$K, $^{232}$Th and $^{40}$K, $^{238}$U and $^{235}$U are displayed in Fig. 10A–E. The resultant correlations are weak in Fig. 10B–D and strong in Fig. 10A and E. There is a high and positive correlation between $^{238}$U and $^{226}$Ra ($R^2 = 0.8519$) and also a high and positive correlation between $^{238}$U and $^{235}$U ($R^2 = 0.7571$) because the two elements accompany each other. Meanwhile, Fig. 10B shows a negative correlation between $^{238}$U and $^{40}$K ($R^2 = 0.0884$), which could be indication of high U-enrichment. Fig. 10C shows weak correlation between $^{238}$U and $^{232}$Th ($R^2 = 0.0065$) due to the lower activity concentration of $^{232}$Th compared with the high re-concentrated $^{238}$U. Fig. 10D shows a poor negative correlation between $^{232}$Th and $^{40}$K ($R^2 = 0.1534$), which indicates moderate activity concentration level of $^{232}$Th compared to very low $^{40}$K activity concentration.

The specific activity of a sample containing 1.0 ppm, by weight, of $^{238}$U is 12.35 Bq/kg, 1.0 ppm of Ra is 11.1 Bq/kg, 1 ppm of $^{232}$Th is 4.06 Bq/kg and 1% of $^{40}$K is 313 Bq/kg (El Galy et al., 2008). For $^{235}$U the specific activity concentration also given in ppm is considered as 1 ppm = 80 Bq/kg (Barton, 1995).

The radionuclide concentration of $^{238}$U series, $^{232}$Th series and $^{40}$K varies between 70 and 251 ppm, 7 and 26.2 ppm and 0.29% and 0.54% respectively (Table 5). The arithmetic mean of all studied phosphate samples (0.11) of $^{232}$Th/$^{238}$U ratio is much lower than the Clark’s value (3.5), which indicates U-enrichment in the phosphates in the studied area. $^{238}$U/$^{226}$Ra ratios for most of the 12 phosphate samples are lower than unity, reflecting a state of radioactive disequilib-
Table 6  Values of radium equivalent (Ra eq), external hazard (H ex), internal hazard (H in), external gamma radiation (I c), dose rate (D R) and the effective dose (Eff. Dose) for Quseir-Safaga phosphate samples, Central Eastern Desert, Egypt.

| Samples | Ra eq (Bq/kg) | H ex (nGy/y) | H in (nGy/y) | I c | D R (nGy/h) | Eff. Dose (mSv/y) |
|---------|---------------|--------------|--------------|-----|-------------|------------------|
| Q1      | 1110          | 3.0          | 5.8          | 7.5 | 468         | 0.57             |
| Q2      | 1030          | 2.8          | 5.4          | 6.9 | 429         | 0.53             |
| Q3      | 1589          | 4.3          | 8.4          | 10.7| 646         | 0.79             |
| Q4      | 2823          | 7.6          | 15.0         | 19.0| 1062        | 1.30             |
| Q5      | 2674          | 7.2          | 14.1         | 18.0| 1248        | 1.53             |
| Q6      | 2011          | 5.4          | 10.6         | 13.5| 608         | 0.75             |
| Q7      | 1105          | 3.0          | 5.8          | 7.4 | 588         | 0.72             |
| Q8      | 3332          | 9.0          | 17.8         | 22.4| 1102        | 1.35             |
| Q9      | 2938          | 7.9          | 15.7         | 19.7| 1286        | 1.58             |
| Q10     | 3387          | 9.2          | 18.1         | 22.7| 1468        | 1.80             |
| Q11     | 1832          | 5.0          | 9.7          | 12.3| 692         | 0.85             |
| Q12     | 2088          | 5.6          | 10.9         | 14.1| 663         | 0.81             |
| Min     | 1030          | 2.8          | 5.4          | 6.9 | 429         | 0.53             |
| Max     | 3387          | 9.2          | 18.1         | 22.7| 1468        | 1.8              |
| Average | 2160          | 5.8          | 11.4         | 14.5| 855         | 1.05             |
| P.L.    | 370           | 1.0          | 1.0          | 1.0 | 58          | 1.0              |

Figure 11  Values of radium equivalent (Ra eq) in (Bq/kg), external and internal hazard (H ex & H in) in (mGy/y), external gamma radiation (I c), dose rate (D R) in (nGy/h) and effective dose (Eff. Dose) in (mSv/y) for Quseir-Safaga phosphate samples, Central Eastern Desert, Egypt.
Correlation between radon mass exhalation rate (E_{Rn}) and activity of $^{226}$Ra ($A_{Ra}$) for Quseir-Safaga phosphate samples, Central Eastern Desert, Egypt.

Table 7 Values of activity of $^{226}$Ra ($A_{Ra}$), emanation coefficient of radon ($C_{Ra}$), emanation factor ($F$) and mass exhalation rate of radon ($E_{Rn}$) for Quseir-Safaga 12 phosphate samples, Central Eastern Desert, Egypt.

| Sample No. | $A_{Ra}$ (Bq/kg) | $C_{Ra}$ (Bq/kg) | $F$ | $E_{Rn}$ (μBq/kg) |
|------------|-----------------|-----------------|-----|-----------------|
| Q1         | 1039            | 518             | 0.50 | 1.1             |
| Q2         | 963             | 479             | 0.50 | 1.0             |
| Q3         | 1527            | 753             | 0.49 | 2.4             |
| Q4         | 2711            | 1439            | 0.53 | 8.2             |
| Q5         | 2528            | 1076            | 0.43 | 5.7             |
| Q6         | 1892            | 1028            | 0.54 | 4.1             |
| Q7         | 1056            | 315             | 0.30 | 0.7             |
| Q8         | 3265            | 1886            | 0.58 | 12.9            |
| Q9         | 2875            | 1480            | 0.52 | 8.9             |
| Q10        | 3309            | 1725            | 0.52 | 12.0            |
| Q11        | 1742            | 870             | 0.50 | 3.2             |
| Q12        | 1928            | 1045            | 0.54 | 4.2             |
| Min        | 963             | 315             | 0.30 | 0.7             |
| Max        | 3309            | 1886            | 0.58 | 12.9            |
| Average    | 2070            | 1051            | 0.50 | 5.4             |

The activity concentration of U-238 was measured from the activity of $^{226}$Ra ($A_{Ra}$) and mass exhalation rate of radon ($E_{Rn}$), both with a correlation coefficient $R^2 = 0.9513$, which means that $E_{Rn}$ and $A_{Ra}$ accompany with each other.

6. Conclusions

High-sensitivity carbone γ-ray spectrometric survey proves high concentrations of eU ranging from 34 ppm to 182 ppm associated with locations of phosphate mines which are considered high risk for populations. The northwestern part of the study area possesses the highest content of eU. Meanwhile, all phosphate mines represent the lowest level of both eTh and K concentrations in the study area. The remaining rock types of the study area, which include older granites, metavolcanics and ophiolites as well as sedimentary rocks, are saved where they are free from any high radioelement concentrations.

The concentrations of trace elements in phosphate samples at Quseir-Safaga area reveal high concentration of both Strontium (Sr) and Zinc (Zn) with an average 51% and 20% respectively.

The current study estimated the activity concentrations of the three radionuclides: $^{226}$Ra, $^{238}$U, $^{232}$Th and $^{40}$K by means of gamma-ray spectrometry in twelve phosphate samples in Quseir-Safaga area, Central Eastern Desert, Egypt. The average values of activity concentrations of $^{226}$Ra, $^{238}$U, $^{232}$Th and $^{40}$K in these 12 samples were 2070 Bq/kg, 1766 Bq/kg, 57 Bq/kg and 117 Bq/kg respectively. The maximum concentrations of $^{226}$Ra, $^{238}$U, $^{232}$Th and $^{40}$K reached 3309 Bq/kg, 3104 Bq/kg, 106 Bq/kg and 169 Bq/kg respectively. Though, the area of large concentration of $^{238}$U coincides with slightly high values of $^{232}$Th but does not coincide with any high values of $^{40}$K. The total annual effective doses which could be received by members of the public from exposition to phosphate mines ranged from 0.53 mSv/y to 1.8 mSv/y, with an average value of 1.05 mSv/y, which is higher than the worldwide value of 1.0 mSv/y for the permitted public exposure.

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