Clockwise rotation of the Tarim basin driven by the Indian plate impact

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Abstract: Tarim basin is usually regarded as a rigid tectonic block that resists the northward expansion of the Tibetan plateau that is associated with ongoing northward motion of the Indian plate. The structural complexity of its tectonic boundaries with adjacent units inferred from seismological studies suggests, however, that this simple view of a static block under north-south compression needs revision. Here, we demonstrate (from Global Position System data) that Tarim basin is rotating clockwise (at a rate of 0.461°/Myr) around a (virtual) vertical axis within the unit. This rotation logically results from the asymmetric nature of the India-Asia collision zone as inferred from seismological studies and can explain the otherwise puzzling pattern of lithospheric interactions along the Tarim basin borders as well as the diachronous seawater retreat and the closing of the westerly moisture pathway that results in aridification and desertification of Tarim.

Keywords: Tarim Basin, block rotation, Indian plate, Eurasian plate, continental collision, paleoenvironment, westerly moisture pathway

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исследований, дает основание считать, что этот упрощенный взгляд, предполагающий простое движение Таримского бассейна в результате меридионального сжатия статического блока, нуждается в пересмотре. Наши данные позволяют продемонстрировать в Глобальной системе координат, что Таримский бассейн вращается по часовой стрелке со скоростью 0.461° в миллион лет вокруг виртуальной оси внутри структуры. Это вращение логически вытекает из асимметричной природы зоны Индо-Азийского столкновения, которая установлена в процессе сейсмологических исследований, призванных объяснить загадочную картину петро-сферных взаимодействий вдоль границ Тарима, а также изучения диахронного отступления границы моря и перекрытия западного канала доступа морской воды, что привело к аридизации и опустыниванию Таримского бассейна.

Ключевые слова: Таримский бассейн, вращение блока, Индийская плита, континентальная коллизия, палеогеографические реконструкции, западный морской канал

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The Tarim basin is a cratonic block located between the Siberian craton and the Tibetan / Pamir plateau, bounded on all sides by major faults and mountain belts (Fig. 1). Its Precambrian basement is covered by thick Paleozoic-Mesozoic-Cenozoic marine sediments and post-Miocene Taklimakan Desert eolian dune sands [1, 2]. In the classical view the basin (backstopped by Siberia) resists the northward motion of the Tibetan plateau caused by the India-Asia collision [3], with minor rotations of the basin around Euler poles outside the unit [4, 5]. Seismological studies of Tarim basin boundaries suggest, however, that lithospheric interactions with surrounding units are more complicated than expected from this canonical model. Changes in environment, such as the Paleogene aridification and desertification, attributed to a combination of Tibetan Plateau uplift and seawater retreat from the area [6], also suggest higher order complexity. Here we report evidence from Global Position System (GPS) data for clockwise rotation of the Tarim block around a virtual axis within the unit, yielding larger strain gradients than predicted previously from paleomagnetic data. This rotation explains the observed structural complexity along the boundaries, as well as changes in paleoenvironment, and is itself a logical consequence of the asymmetry and diachronous nature of the India-Asia collision zone. Over the past two decades a number of seismic experiments have been conducted across the borders of the Tarim basin (Fig. 1) with the purpose of studying the lithospheric structure of the basin with respect to the surrounding mountain belts. Details about the experiments are given in the Online Supplement; here we summarize the results insofar they are relevant for the discussion of regional variations in deformation style. In the north, the observation of a double Moho beneath Tien Shan and joint inversion of seismic, gravity, and geomagnetic data suggest that Tarim basin crust underthrusts central Tien Shan [7] (Line XB; Fig. 2, a). In the northeast, the seismic data reveal southward underthrusting of the Junggar basin lower crust beneath Tien Shan and a gap between the Moho of Tarim basin and eastern Tien Shan (Line KJ; Fig. 2, b). In the southeast, wide-angle reflection/refraction and joint inversion with gravity data reveal southward underthrusting of the Tarim lithosphere beneath the Altyn Shan and Tibet [8] (Line BD; Fig. 2, c). In the southwest, pure shear deformation is accommodated by deep faults and receiver function profiles reveal a Moho step of ~20 km, from ~80 km beneath Tibet to ~60 km beneath Tarim [9, 10] (Line ANTILOPE-I; Fig. 2, d).

Starting at the northern part of the Tarim Basin and ending at the northern margin of the Junggar Basin, the XB Line crosses northern Tarim Basin, the Tien Shan, and the Junggar Basin, with a length of 900 km. Along the profile, wide-angle seismic reflection/refraction and a joint inversion of gravity
Zhao Junmeng, Zhang Peizhen, Yuan Xiaohui, et al. Clockwise rotation of the Tarim Basin and a double Moho structure beneath the Tien Shan has been observed (Fig. 2, a).

Fig. 1. Map of Tarim Basin and surrounding Tibetan and Tien Shan orogenic belts

Solid lines show locations of the profiles. Dashed lines indicate the front of the Indian mantle lithosphere seen by seismic tomography and receiver functions (green line denotes [11] and blue line denotes [10]. Thin red lines denote main faults around the Tarim Basin. Black triangle pairs indicate compression or extension.

Blue start denotes the Euler pole, around which the Tarim block is rotating clockwise.

Abbreviations for tectonic units: TB – Tarim Basin; JB – Junggar Basin; CTS – Central Tien Shan; ETS – Eastern Tien Shan; AS – Altyn Shan; KS – Kunlun Shan; KKS – Karakoram Shan; QB – Qaidam Basin; TP – Tibetan plateau

Numbered faults: 1 – Talas-Fergana Fault; 2 – West Kunlun Piedmont Fault; 3 – Kangxiwar Fault; 4 – Miya-Tianshendaban Fault; 5 – East Kunlun Fault; 6 – Altyn Fault; 7 – Tarim Northern Margin Fault

Рис. 1. Карта Таримского бассейна и прилегающих к нему территорий Тибета и Тяньшаньских орогенических поясов

Сплошными линиями показано положение профилей. Пунктирными линиями отмечена передняя часть литосферной мантии по данным сейсмической томографии и функций сейсмоприемника (зеленая линия – по данным источника [11] и синяя – по данным источника [10]. Тонкими красными линиями отмечены главные разломы вокруг Таримского бассейна. Черные стрелки указывают направление сжатия или растяжения. Синей звездой отмечен полюс Эйлера, вокруг которого Таримский блок вращается по часовой стрелке

Сокращения для тектонических единиц: TB – Таримский бассейн; JB – Джунгарский бассейн; CTS – Центральный Тянь-Шань; ETS – Восточный Тянь-Шань; AS – Алтын Шань; KS – Кунлунь Шань; KKS – Каракорум Шань; QB – бассейн Кайдам; TP – Тибетское плато

Пронумерованные разломы: 1 – Таласо-Ферганский; 2 – Западно-Куньлуньский; 3 – Разлом Канксиар; 4 – Ми-Таньшэндабан; 5 – Восточно-Куньлуньский; 6 – Разлом Алтын; 7 – Разлом северной границы Тарима
Fig. 2. Cartoons showing interaction of crust and mantle lithosphere between Tarim block and the surroundings orogenic belts

Along four profiles shown in Fig. 1

The crust is homogeneously indicated by yellow color, while mantle lithospheres of different units are denoted by different colors: green – Tarim Basin; cyan – Jungar Basin; red – Tien Shan; blue – Tibetan Plateau

Abbreviations are the same in Fig. 1

The recent 600-km long seismic profile (KJ line) crosses the northeastern margin of the Tarim Basin, the east Tien Shan and the southern part of the Junggar Basin. This comprehensive geophysical profile reveals a double-Moho structure underneath the northern margin of the east Tien Shan with underthrusting of the Junggar Basin lower crust beneath the Tien Shan (Fig. 2, b). Beneath the southern Tien Shan margin a gap exists between the Moho of the Tarim Basin and that of the east Tien Shan (Fig. 2, b). Along with previous seismic profiles this reveals a complex configuration along the basin boundaries. The NW-SE striking, 1400-km long seismic profile (BD Line) traverses the eastern part of the Tarim Basin, the Altyn Shan, and the Qaidam Basin. Wide-angle reflection/refraction and joint inversion of gravity and geomagnetism revealed an underthrusting of the Tarim lithosphere beneath the Altyn Shan to a depth of 90 km [8] (Fig. 2, c). The ANTILOPE-I is a passive-source seismic profile, located in...
Zhao Junmeng, Zhang Peizhen, Yuan Xiaohui, et al. Clockwise rotation of the Tarim Basin, extending from Himalaya to the southern Tarim Basin, with a total length of ~800 km. The Moho is seen at a depth of ~80 km underneath Tibet and suddenly shallows to ~60 km underneath the Tarim Basin with a step of ~20 km [10]. Close to the profile a Sino-French experiment revealed a similar Moho step beneath the Altyn Shan [9]. The crustal shortening here is accommodated by pure shear in contrast to simple shear at the southern Tarim margin (along BD line) (Fig. 2, d).

The lithospheric interactions between bounding units implied by these observations are inconsistent with the canonical framework of predominantly NS compression but are readily explained if the strain gradients due to clockwise rotation of the Tarim Block are larger than previously suggested. Paleomagnetic data suggest negligible post-Cretaceous rotation of Tarim basin [4], whereas a kinematic model derived from modeling slip rates along major thrust and strike-slip faults revealed a Cenozoic clockwise rotation of the Tarim basin with respect to Siberia at a rate of 0.65±0.3°/Mys and around a pole at (43.5°N, 95.7°E) [5]. Euler poles outside Tarim produce rotations that are too small, however, to explain the different deformation styles inferred from the seismological studies.

GPS measurements constrain present-day surface motions around the Tarim basin. In a reference frame of stable Eurasian plate (Fig. 3, a) GPS vectors in the western Tarim regions are dominated by N-S direction while eastern Tarim moves toward the NE. From GPS data [12] we calculated the relative motion of Tarim basin with respect of its surrounding areas (Fig. 3, b). First, we solved for the Euler vector that minimized the RMS velocity of all these surrounding stations. Then, through reverse rotation of the GPS velocity field around the Euler vector thus obtained we removed the overall rigid rotation of the surrounding units. The resulting “Tarim fixed reference frame” highlights the clockwise rotation of the Tarim basin with respect to its surroundings. The rotation rate (0.461±0.036°/Myr) is close to the inference from kinematic data by Avouac et al. [5], but a pole within the Tarim basin (at 39.2°± 0.27°N, 82.0°± 0.34°E) produces strain gradients that can explain morphological changes along the Tien Shan [3] as well as the otherwise puzzling styles of deformation in the boundary zones, that is, compression in NW with central Tien Shan, trans-tension in NE with eastern Tien Shan, compression with Altyn Tagh Shan in the SE, and left lateral shear motions in the SW (Fig. 1).

The clockwise rotation of the Tarim block can explain variations in the topography of the Tien Shan. Geographically, the central Tien Shan is north-south symmetric (Fig. 2, a), whereas the eastern Tien Shan is asymmetric (Fig. 2, b). The geometry of the crustal interplay may result in an extension and strike sliding underneath the southern Tien Shan margin. The elevation of the southern Tien Shan is reduced by weathering and erosion, causing asymmetric topography of the Tien Shan along the KJ line. The tectonic difference of the southern and northern Tien Shan, and the deformation of the basins at its two sides are different. At the southern margin of the Junggar Basin, the deformation is strong, causing steep structures and intense folding. The northern margin of the Tarim Basin is relatively flat, because it has experienced less tectonic deformation.

Block rotations are not expected in an overall regime of (pure shear) north-south convergence, but we propose that the rotation of the Tarim basin is driven by the asymmetric indentation of the northward moving Indian plate and facilitated by lateral variations in geomechanical properties of plate boundary zone. First, the lithospheric boundary between the Indian plate and the Eurasian plate beneath Tibet roughly follows a line from the western Tarim basin to the eastern Himalayan syntaxis (Fig. 1) [10, 11]. Second, increased seismic anisotropy [10, 13–15] and high attenuation in the uppermost mantle [16, 17] suggest that in central and east Tibet hot and more easily deformable mantle material fills the gap between
Fig. 3. Observed GPS velocity vectors (red arrows) in and around the Tarim block in a reference frame of Siberia representing the stable Eurasia plate (a); GPS velocity vectors (red arrows) in and around the Tarim block in the “Tarim surrounding vicinity fixed reference frame” (b).

The blue star indicates the Euler pole location of the clockwise rotation of the Tarim block with respect to its surroundings. The blue vectors are modeled GPS velocity (see text for explanation).

Рис. 3. Наблюдаемые векторы скорости GPS (красные стрелки) внутри и вокруг блока Тарим в системе отсчета Сибири, представляющей стабильную Евразийскую плиту (a); векторы скорости GPS (красные стрелки) внутри и вокруг блока Тарим в “Таримском окружающем вицините фиксированной системы отсчета” (b).

Синяя звезда показывает расположение полюса Эйлера для вращения блока Тарим по часовой стрелке относительно его окружения. Синие стрелки показывают смоделированные скорости GPS (пояснения даны в тексте).
Indian and Asian plates. In western Tibet, to-
toward Pamir, indentation of the Indian plate
across the entire width of the Plateau exerts
a northward push on the western Tarim ba-
sin and the weak zone further east allows
the Tarim basin to rotate clockwise in re-
sponse to the northward push further west.

The rotation of the Tarim block had
significant consequences for the paleoenvi-
ronment of the Asian interior. Previous stud-
ies have indicated that the Northern Pamir
has been displaced northward 400–300 km
with respect to the rest of Eurasia [18], lead-
ing to collision with the southeastern Tien
Shan. Traditional views attribute all crustal
shortening in this region simply to the north-
ward motion of Pamir in response to the In-
dia-Eurasia collision [18]. Our new geophys-
ical evidence refines this view and suggests
that the collision of the Northern Pamir with
the southern Tien Shan is so localized be-
cause the asymmetric indentation of India and
the concomitant clockwise rotation of
the Tarim Block enhanced the northward
displacement of Pamir compared to other
parts of the collision zone. The localized
convergence influenced the retreat of sea
water from the Tarim Basin in the late Eo-
cene [19, 20] and accelerated the final colli-
sion in the latest Miocene [21–24], which
blocked the transport of water-vapor by the
prevailing westerlies from the Atlantic and
the Mediterranean Sea to the downwind Ta-
rim Basin, leading to the extreme arid cli-
mate and the formation of the world second-
largest shifting-sand desert (Taklimakan De-
sert) in the Tarim Basin in the latest Miocene
[1, 25].

Block rotation is common in plate
boundary zones controlled by strike slip or
transpression, but our study indicates that
such rotations can also result from asym-
metry in the structure and mechanical prop-
herties of colliding tectonic units and produce
lateral changes in paleoenvironment in pre-
dominantly pure shear regimes.

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**Responsibility for plagiarism / Конфликт интересов**

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