Effects of nitrogen deposition on carbon cycle in terrestrial ecosystems of China: A meta-analysis

Hao Chen a, b, Dejun Li b, Geshere A. Gurmesa a, d, f, Guirui Yu c, Linghao Li e, Wei Zhang a, Huajun Fang c, Jiangming Mo a, *

a Key Laboratory of Vegetation Restoration and Management of Degraded Ecosystems, South China Botanical Garden, Chinese Academy of Sciences, Guangzhou 510650, China
b Huanjiang Observation and Research Station for Karst Ecosystem, Key Laboratory of Agro-ecological Processes in Subtropical Region, Institute of Subtropical Agriculture, Chinese Academy of Sciences, Changsha 410125, China
c Institute of Geographical Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 100101, China
d Department of Geosciences and Natural Resource Management, University of Copenhagen, Rolighedsvej 23, DK-1958 Frederiksberg C, Denmark
e State Key Laboratory of Vegetation Environmental Change, Institute of Botany, Chinese Academy of Sciences, Xiangshan, Beijing 100093, China
f Sino-Danish Center for Education and Research, Niels Jensens Vej 2, DK-8000 Aarhus C, Denmark

A R T I C L E   I N F O

Article history:
Received 2 April 2015
Received in revised form 11 July 2015
Accepted 19 July 2015
Available online 30 July 2015

Keywords:
Nitrogen deposition
Carbon cycle
Global change
N-rich and N-limited ecosystems
Meta-analysis
China

A B S T R A C T

Nitrogen (N) deposition in China has increased greatly, but the general impact of elevated N deposition on carbon (C) dynamics in Chinese terrestrial ecosystems is not well documented. In this study, we used a meta-analysis method to compile 88 studies on the effects of N deposition on C cycling on Chinese terrestrial ecosystems. Our results showed that N addition did not change soil C pools but increased above-ground plant C pool. A large decrease in below-ground plant C pool was observed. Our result also showed that the impacts of N addition on ecosystem C dynamics depend on ecosystem type and rate of N addition. Overall, our findings suggest that 1) decreased below-ground plant C pool may limit long-term soil C sequestration; and 2) it is better to treat N-rich and N-limited ecosystems differently in modeling effects of N deposition on ecosystem C cycle.

© 2015 Elsevier Ltd. All rights reserved.

1. Introduction

Global atmospheric nitrogen (N) deposition has dramatically increased due to extensive use of fossil fuels in industries and transportation, heavy application of fertilizers in agriculture, and expansion of animal husbandry (Galloway et al., 2004). Elevated N deposition is likely to change global carbon (C) cycles (Chapin et al., 2009) because N and C cycles are interdependent forming the basis of biogeochemical cycles and energy flows. A number of simulated field studies have been conducted in the past decades to investigate the effects of N deposition on ecosystem C cycles (Hogberg et al., 2006; Hyyonen et al., 2008; Pregitzer et al., 2008). Several meta-analysis studies have synthesized the available information to determine how ecosystem C pools and processes respond to N deposition at a global scale (LeBauer and Treseder, 2008; Lu et al., 2011; Xia and Wan, 2008). These studies have greatly improved our understanding regarding consequences of N deposition on C cycling and indicated a strong regional diversity in the response of ecosystem C to N deposition at a global scale.

China, the second-largest world economy, has undergone rapid economic development with an average annual GDP growth of 9.1 percent during 1989–2014 (NBSC, 2014). This rapid development has caused serious environmental issues. The latest data show that China contributed 29% of the total 36 billion tonnes carbon emitted from all human sources in 2013 (Friedlingstein et al., 2014). N deposition in China has also been among the greatest globally (Liu et al., 2013). The averaged N deposition rate in China has increased from 13.2 kg N ha⁻¹ in the 1980s to 21.1 kg N ha⁻¹ in the 2000s and is projected to increase in the coming decades (Liu et al., 2013). On the other hand, China also has experienced regionally distinct land-use histories and climate trends (Piao et al., 2009). Hence the degree of N deposition and its potential effects on the C cycle in...
Chinese terrestrial ecosystems should be different and are of increasing global concern due to the increasing interest in regional aspects of the global C cycle (Piao et al., 2009). Several N deposition monitoring programs and increased N deposition simulation experiments have been conducted since the late 1990s (Mo et al., 2006; Niu et al., 2010; Tu et al., 2010; Xu et al., 2009). Liu et al. (2011) recently presented a description of the effects of N deposition on ecosystem C cycling in China. However, a synthetic analysis of N deposition impacts on C dynamics in Chinese terrestrial ecosystems remains lacking.

In the present study, we used a meta-analysis technique to synthesize all the available information in China. Our main objectives were to: (1) quantify the responses of C fluxes (C influx and efflux) and C pool sizes (including plant, litter, microbe, soil, and dissolved organic C) to experimental N addition; (2) examine whether ecosystem types, fertilization rates and fertilization forms influence the responses of ecosystem C fluxes and pools to N addition; and (3) reveal the shortage of the current N deposition research in China. Our study provides valuable information to environmental policy- and decision-makers in their attempts to curb N emissions to the atmosphere, and to evaluate the effects of N deposition on terrestrial ecosystems.

2. Materials and methods

2.1. Data collection

Publications that studied C cycle response to N addition (Supplementary Information Table S2) were selected by searching Web of Science (2000–2013) and China National Knowledge Infrastructure (CNKI). In addition, unpublished data from some studies known to us were collected. To avoid publication biases, the following four criteria were applied to select appropriate studies:

(1) the study must contain at least one of our selected variables with a clear record of the ecosystem type, N application rate, N fertilization form, and experimental duration; (2) the N addition and control plots started with the same climatic, soil and vegetation conditions to reduce effects of confounding factors; (3) the means, standard deviations or standard errors and sample sizes of the target variables were directly reported or could be either calculated from data presented in the paper or extracted using Origin 8.0 software (Origin Lab Corporation, Northampton, MA, USA) if the data were graphically presented; (4) the study must not use N deposition gradients as in several other studies (Fang et al., 2011; Huang et al., 2012). Most of the studies used one-time measurement, but latest samplings were used if more than one measurements at different temporal scales were available for the same experiment (Liu and Greaver, 2009; Treseder, 2008). Measurements for different N application rates were considered as independent observations if more than one levels of N addition were applied in the same experiment (Liu and Greaver, 2009).

The compiled database contained 12 variables associated with ecosystem C cycle, including C fluxes (i.e., net photosynthesis rate (NPR), net primary productivity (NPP), litterfall, litter decomposition rate (mass loss or \( k \) value), soil respiration, and ecosystem respiration), ecosystem C pools in above- and below-ground plant biomass, microbial biomass (MBC), dissolved organic C (DOC), organic horizon (OH) and mineral soil C. Data on C fluxes were obtained from studies where these data were directly reported. We used ecosystem respiration to replace soil respiration in grasslands because soil respiration is less measured in grasslands. Above- and below-ground plant C pools were determined by above-ground plant biomass or below-ground root biomass. Soil C pools were determined by soil C content or C storage. In addition, four supporting variables (i.e., soil inorganic N, leaf N, and soil pH) were evaluated (Supplementary Information Table S3) to explain the results of the main variables above.

All the variables were compared among different ecosystem types, N addition forms, and N addition rates. Four ecosystem types (i.e., N-rich subtropical forests, N-limited subtropical forests, temperate forests and grasslands) were identified. These ecosystems were categorized into two main groups as N-rich and N-limited ecosystem based on their initial soil N level and information on plant growth response to experimental N addition. That is, the old-growth subtropical forests (Mo et al., 2008) and forests dominated by N-fixing species (tree age > 30 years) (Zhang et al., 2012) were grouped into N-rich ecosystems as they have high soil N concentrations (Fang et al., 2008) and plant growth in these ecosystems is not limited by N availability (Lu et al., 2010). Early successional forests (Mo et al., 2007; Tu et al., 2010), temperate forests and grasslands, were grouped into N-limited ecosystem because they have low soil N level and plant growth in these ecosystems is limited by N availability. N addition levels were classified into three classes, including low N (LN; \( ≤ 60 \) kg N ha\(^{-1}\) yr\(^{-1}\)), medium N (MN; \( 61–120 \) kg N ha\(^{-1}\) yr\(^{-1}\)), and high N (HN; \( > 120 \) kg N ha\(^{-1}\) yr\(^{-1}\)). In addition, we compared the variables only between two N fertilizer forms; NH\(_4\)NO\(_3\) and urea which were the main fertilization forms in recent China’s N deposition simulated experiments.

2.2. Analysis of data

The data were analyzed using meta-analysis method as described in Hedges et al. (1999). The effects of N addition on terrestrial ecosystem C pools and fluxes were estimated by response ratio (RR), which was calculated as:

\[
RR = \ln \left( \frac{X_t}{X_c} \right) = \ln(X_t) - \ln(X_c)
\]

where \( (X_t) \) and \( (X_c) \) are the mean values of a given variable in the treatment and the control group, respectively. The natural log was used for meta-analyses because its bias is small and its sampling distribution is approximately normal (Luo et al., 2006). The variance \( (v) \) of RR was estimated as:

\[
v = \frac{s_t^2}{n_t X_t^2} + \frac{s_c^2}{n_c X_c^2}
\]

where \( n_t \) and \( n_c \) are the sample sizes for the treatment and control groups, respectively; and \( s_t \) and \( s_c \) are the standard deviations for the treatment and control groups, respectively.

To test whether the experimental conditions alter the response magnitude to N addition, each observation was categorized into three groups: ecosystem type, N addition rate, and forms of N addition. The data were sub-divided and the mean of response ratio (RR\(_{+,+}\)) and standard error \( [s(RR_{+,+})] \) were calculated as:

\[
RR_{+,+} = \frac{\sum_{i=1}^{m} \sum_{j=1}^{k} W_{ij} RR_{ij}}{\sum_{i=1}^{m} \sum_{j=1}^{k} W_{ij}}
\]

\[
s(RR_{+,+}) = \sqrt{\frac{1}{\sum_{i=1}^{m} \sum_{j=1}^{k} W_{ij}}}
\]

where \( m \) is the number of groups (e.g., different N addition levels, N forms or ecosystem types), \( k \) is the number of comparisons in the \( i \)th group, and \( W_{ij} \) is the weighting factor and was estimated as:
The 95% confidence interval (CI) for the response ratio was given as:

$$95\% \text{ CI} = RR_{++} \pm 1.96s(\Delta RR_{++})$$  \hspace{1cm} (6)

If the 95% CI values of $RR_{++}$ for a variable did not cover zero, the effects of N addition on the variable were considered to differ significantly between treatments.

The meta-analysis was conducted using the METAWIN (Sinauer Associates, Inc. Sunderland, MA, USA). We also used a one-way ANOVA to examine whether the $RR_{++}$ of a variable differed significantly between ecosystem types, and among different N addition rates or forms. For a better explanation, mean of the response ratio ($RR_{++}$) was transformed back to the percentage change calculated by the formula: $\left(e^{RR_{++}} - 1\right) \times 100\%$.

3. Results

3.1. Data overview

Our meta-analysis study compiled data from 88 studies that included 607 observations at 33 sites (Fig. 1), and all were natural ecosystems. Four types of ecosystems were included in the analysis: N-rich subtropical forest ($n = 98$), N-limited subtropical forest ($n = 217$), temperate forest ($n = 88$), and grassland ($n = 204$) (Table 1). In total 12 response variables were extracted from the selected studies. However, not all variables were measured at all the sites. The missing variables include NPR and litterfall in grassland;

\[ W_j = \frac{1}{V} \hspace{1cm} (5) \]
NPR, NPP, above-ground plant C and OH soil C pools in temperate forest. Similarly, ecosystem respiration was measured only in grasslands (Figs. 2 and 3). The fertilization duration in all selected studies lasted up to 10 years and fertilization durations in 85.5% of studies were ≤ 4 years (Table 1). The N addition rates ranged from 10 to 640 kg N ha\(^{-1}\) year\(^{-1}\), and the N fertilizer forms included NH\(_4\)NO\(_3\) (counting for 73.3%), urea (counting for 21.6%) and others (such as NaNO\(_3\), and (NH\(_4\))\(_2\)SO\(_4\) and mixed fertilizer) (Table 1).

Fig. 2. Weighted response ratios (RR\(_{w+}\)) of different ecosystem types to N additions for C fluxes. (a) net photosynthesis rate; (b) net primary productivity; (c) litterfall; (d) litter decomposition; (e) ecosystem respiration; and (f) soil respiration. Bars represent RR\(_{w+}\) and 95% confidence intervals. The vertical lines are drawn at RR\(_{w+}\) = 0. The sample size for each variable is shown next to the bar.
Fig. 3. Weighted response ratios (RR\textsubscript{+}) of different ecosystem types to N additions for C pools. (a) Above-ground plant C pool; (b) Below-ground plant C pool; (c) MBC; (d) DOC; (e) OH soil C pool; and (f) Mineral soil C pool. Bars represent RR\textsubscript{+} and 95% confidence intervals. The vertical lines are drawn at RR\textsubscript{+} = 0. The sample size for each variable is shown next to the bar.
3.2. N addition effects on C fluxes and C pools

N addition-induced changes in C fluxes showed large variation. Compared with control groups, NPR and NPP were significantly increased by 5.2% and 10.3%, respectively under experimental N addition (mean effect size = 0.051 and 0.154; CI: 0.01 to 0.09, and 0.04 to 0.26; Table 2, Fig. 2a,b). Litterfall was increased by 4.4% (mean effect size = 0.043; CI: −0.02 to 0.11; Table 2, Fig. 2c) but the change was statistically not significant. N addition significantly decreased litter decomposition (mean effect size = −0.066; CI: −0.10 to −0.02; Table 2, Fig. 2d) and soil respiration (mean effect size = −0.059; CI: −0.16 to −0.07; Table 2, Fig. 2f) by 6.4% and 5.3%, respectively, across all ecosystems, but it significantly increased ecosystem respiration by 19.7% in grasslands (mean effect size = 0.15; CI: 0.05 to 0.24; Table 2, Fig. 2e).

Changes in C stock due to N addition varied among ecosystem pools. N addition significantly increased above-ground plant C pool (mean effect size = 0.198; CI: 0.12 to 0.27; Table 2, Fig. 3a) and DOC (mean effect size = 0.146; CI: 0.10 to 0.19; Table 2, Fig. 3d) by 21.9% and 15.8%, respectively. However, N addition significantly decreased below-ground plant C (mean effect size = −0.287; CI: −0.46 to −0.11; Table 2, Fig. 3b) and MBC (mean effect size = 0.123; CI: −0.17 to −0.07; Table 2, Fig. 3c) by 25.0% and 11.6%, respectively. OH soil C pool exhibited only a minor increase (mean effect size = 0.039; CI: −0.02 to 0.08; Table 2, Fig. 3e), and no effect of N addition on mineral soil C pool was observed across all ecosystems (mean effect size = 0.019; CI: −0.01 to 0.03; Table 2, Fig. 3f).

3.3. Variation between N-rich and N-limited ecosystems

Despite the general trends observed above, responses of some measured variables to N addition greatly varied between N-rich and N-limited ecosystems. For example, changes in above-ground plant C, litterfall, litter decomposition, soil respiration, below-ground plant C and MBC after N addition were significantly different between these two groups (Fig. 3a–f; P < 0.01 for all), showing decreasing trends in N-rich subtropical forest but increased trends in other N-limited ecosystems. However, changes in some C fluxes (e.g. NPR, NPP) and C pools (e.g. OH soil C pool, mineral soil C pool and DOC) for these two ecosystem types did not differ from the general trends observed.

3.4. Effects of different fertilization rates and fertilizer forms

Fertilization rate influenced the responses of C fluxes and C pools to N addition. For example, litter decomposition and soil respiration were significantly different between N addition levels (Fig. 2d, f; both P < 0.01), which decreased under MN and HN treatments but increased under LN treatment. Below-ground plant C pool, MBC and DOC were found to decrease under N addition, but their decreasing extents were significantly larger under MN treatment and HN treatment than those under LN treatment (Fig. 3b, d; P < 0.01 for both). Other variables showed a similar response to N addition under different fertilization rate. However, except for MBC, we did not find any differences between N fertilization forms in all variables.

3.5. Effects of N additions on other variables

Soil inorganic N and leaf N concentrations were significantly increased by 85.9% and 35.6% (Table 2, mean effect size = 0.62 and 0.30; CI: 0.53 to 0.71, and 0.23 to 0.38, respectively), respectively. Soil pH was significantly decreased by 6.4% compared to control treatment (Table 2, mean effect size = −0.28; CI: −0.08 to −0.48). Correlation analysis showed that there was a significant relationship between leaf N and net photosynthesis rate (Fig. 4, P = 0.01).

4. Discussion

4.1. Responses of C storage and fluxes to N addition

Nitrogen addition led to a significant increase in above-ground plant C pool and a slight increase in litter production (Fig. 2c and Fig. 3a) which is likely due to stimulated plant growth. Our results are consistent with previous findings in meta-analysis studies on effects of N addition on above-ground plant growth (LeBauer and Treseder, 2008; Xia and Wan, 2008); gradient analysis (Laubhan et al., 2009; Solberg et al., 2009; Thomas et al., 2010), and model simulation (Wamelink et al., 2009a, 2009b). The increased above-ground plant C pool could be attributed to the enhanced NPP (Fig. 2) which could, in turn, be attributed to improved soil N availability to support plant growth as suggested by several studies (Pregitzer et al., 2008; Vitousek and Howarth, 1991). Increased N
availability to plants due to N addition is evident in our analysis as indicated by the significant increase in soil inorganic N (NH$_4^+$ plus NO$_3^-$ ) and leaf N content by 85.9% and 35.6%, respectively (Table 2). The significant positive relationship between the leaf N concentration and plant photosynthesis we found (Fig. 4) supports the above explanation.

However, N addition has no effect on soil C pools (in both OH and mineral soil C) in this study (Fig. 2, Table 1). This is consistent with a previous suggestion from a meta-analysis study (Lu et al., 2011) that N addition had minor influence on C pools of OH or mineral horizons across global forests and grasslands. However, other reviews and meta-analyses showed that N fertilization increased soil C storage (Hyvonen et al., 2008; Janssens et al., 2010; Nave et al., 2009), and increased C pool in OH but not in mineral horizons were reported elsewhere (Liu and Greaver, 2010).

In terrestrial ecosystems, any changes in soil C pool are generally related to the changes in C fluxes. In this study, we found that adding N increased litterfall and decreased both litter decomposition and soil respiration (Fig. 5), which all suggest potentially increase in soil C sequestration. The increase in DOC concentration we observed (Fig. 3d) did not contribute to changes in soil C pool because DOC is low in soil and is generally suggested to have minor contribution for C sequestration.

The reason why soil C pools were not changed significantly by N addition is far from conclusion based on our data. It could be partly explained by the decrease in below-ground plant C (Fig. 3b) because dead roots are the main mechanism for soil C inputs. Our result may suggest that reduced below-ground plant C under N addition may be an important factor for limiting long-term soil C sequestration. However, this needs further investigation because it contradicts results from previous meta-analysis studies (Xia and Wan, 2008; Lu et al., 2011) who reported that plant root biomass was increased by 23%.

There are two potential explanations for the decrease in below-ground C pool. First, plants invest less C to roots because less effort is required to acquire this resource from the soil (Boxman et al., 1998; Haynes and Gower, 1995). Thus the reduction in below-ground plant C by N addition may be the result of a C tradeoff between above-ground plant tissues (leaves) and below-ground plant tissues. In support of our explanation, previous studies suggested that N addition tended to reduce root: shoot ratio (Johnson and Thornley, 1987). We did also find that in grasslands above-ground plant C increased but the below-ground plant C decreased under N addition, which supports above explanation. However, this explanation does not seems to fit to other ecosystems, especially the N-rich subtropical forest where above- and below-ground plant both decreased. N-addition induced soil acidification may be a better explanation for our finding. It has been suggested that N deposition will lead to soil acidification (Godbold et al., 2003) which would reduce root biomass through directly causing root death (Johnson and Thornley, 1987). In this study, we have found that N addition significantly decreased soil pH by 6.4% across all ecosystems (Table 2). In addition, the similar response trends in soil pH, MBC, and root biomass (Supplementary Information Fig. S1) suggest that N-addition induced soil acidification did have direct damage to plant and microbe.

4.2. Effects of ecosystem type and N addition rate

In this study, very different responses between N-rich subtropical forests and other N-limited ecosystems were shown under N addition. In a previous meta-analysis study, the impact of N addition on ecosystem C dynamics has been found to largely depend on ecosystem types (Liu and Greaver, 2010); however, the difference between N-rich and N-limited ecosystems has not been presented in any previous meta-analysis studies. Our observation can be explained by N saturation hypotheses (Aber et al., 1998) because the two N-rich forest types in our study, i.e., old-growth
forests and N-fixing tropical plantation are both N-saturated systems as reported by previous studies (Mo et al., 2008; Zhang et al., 2012). N saturation hypotheses suggest that once an ecosystem reaches N saturation, further increase in N addition will not increase NPP of the ecosystem (Aber et al., 1998). This is in agreement with our meta-analysis results showing that experimental N addition significantly increased above-ground plant C pools in N-limited systems in subtropical forests, temperate forests and grasslands, but not in N-rich subtropical forests. Our results are also supported by recent field experiments in other regions, which suggested that plant growth may not respond to increased N deposition in subtropical forests with N-rich soils and high rates of soil N cycling (Cusack et al., 2010; Harrington et al., 2001; Kaspari et al., 2008; Miramanto et al., 1999). The N saturation hypotheses also suggest that further N addition to an N-saturated system will lead to reduced plant performance (fine roots) and soil functions such as N retention and nutrient cycling (Aber et al., 1998, 1989). Decreased fine root biomass is evident from our result (Fig. 2b).

Our study also showed that the response of C dynamics to N addition is affected by rates of N addition. Interestingly the effects of N-addition at lower rates on the measure C dynamics (including below-ground plant C pool, MBC, litter decomposition, and soil respiration) are similar to the trends observed in N-limited ecosystems while that of N addition at higher rates were similar to that observed in N-rich ecosystems. These results showed that N availability, regardless of its source, can affect C dynamics in an ecosystem. Different N addition rates indicate increased N availability from exogenous N, but the N-rich vs. N-limited comparison indicates differences in N availability from endogenous N sources. They support nitrogen saturation hypotheses from two different sides. These results suggest that even though N addition at lower rates has a positive effect on C dynamics, the effect could turn negative when an optimal threshold level (i.e., N saturation) is exceeded. Because below-ground plant C pool, MBC, litter decomposition, and soil respiration might be affected differently by N addition rates, further studies focusing on these specific pools and processes are needed to avoid uncertainties in making conclusion about the overall effects of increased N deposition on C cycling in terrestrial ecosystems on regional scales.

4.3. Implications and future research

In the present study, we used meta-analysis to synthesize the effects of N addition on ecosystem C fluxes and pools. These processes and pools were quantified across ecosystems and summarized in Fig. 5. However, our finding based on the currently available data is hampered by some uncertainties to be considered in future research on this topic.

85.5% of the selected studies lasted less than four years in terms of N treatment duration, and all the studies lasted less than 10 years. Due to the short experiment time scales in combination with the inherently high spatial variability, it is difficult to accurately assess the changes in soil C dynamics, especially in the mineral soil (Chapin et al., 2002). In addition, high N fertilization experiments (>120 N kg ha⁻¹ yr⁻¹) for a quit short period were commonly used in recent field experiments to simulate future N deposition. Such high N loads may not be realistic especially in natural ecosystems, and effects from such studies could be too “artificial”. Thus, we suggest that follow-up studies should focus on long-term effects of N deposition using more realistic N deposition levels.

Our analysis showed a large difference between the N-rich and N-limited ecosystems. However, the compiled studies in our meta-analysis were mainly conducted in N-limited ecosystems (i.e., young subtropical forests, temperate forests and grasslands), and only few N-rich subtropical forests were included. Lack of more studies from N-rich ecosystems may influence our evaluation of the integrated response of terrestrial ecosystems to N addition in China. This highlights the need for additional studies in tropical old-growth forests, N-fixing forests, and N-saturated temperate forests and grasslands with different N deposition gradient for a better understanding of how C cycling is affected by increased N deposition under future air pollution scenarios.

Acknowledgments

We thank Drs. Xiaoming Zhu, Qinggong Mao for providing unpublished data, and Wantong Wang for drawing the Fig. 1. This study was supported by the National Basic Research Program of China (2010CB833502) and the National Natural Science Foundation of China (41473112).

Appendix A. Supplementary data

Supplementary data related to this article can be found at http://dx.doi.org/10.1016/j.envpol.2015.07.033.

References

Aber, J., McDowell, W., Nadelhoff, K., Magill, A., Berntson, G., Kamakea, M., McNulty, S., Currie, W., Rustad, L., Fernandez, I., 1998. Nitrogen saturation in temperate forest ecosystems - hypotheses revisited. Bioscience 48, 921–934. Aber, J.D., Nadelhoff, K.J., Steudler, P., Melillo, J.M., 1989. Nitrogen saturation in northern forest ecosystems. Bioscience 39, 378–386.

Boxman, A.W., Blanck, C., Brandrud, T.-E., Emmitt, B.A., Gundersen, P., Hoekstra, R., Kjosa, O.J., Persson, H., Timmermann, V., 1998. Vegetation and soil biota response to experimentally-changed nitrogen inputs in coniferous forest ecosystems of the NITREX project. For. Ecol. Manag. 101, 65–79.

Chapin III, F.S., Matson, P.A., Mooney, H.A., 2002. Principles of Terrestrial Ecosystem Ecology. New York.

Chapin III, F.S., McFarland, J., McGuire, A.D., Euskirchen, E.S., Esser, R.W., Kieland, K., 2009. The changing global carbon cycle: linking plant-soil carbon dynamics to global consequences. J. Ecol. 97, 840–850.

Cusack, D., Silver, W., Tom, M., McDowell, W., 2010. Effects of nitrogen additions on above- and below-ground carbon dynamics in two tropical forests. Biogeochemistry 1–23.

Fang, Y.T., Yoh, M., Koba, K., Zhu, W.X., Takebayashi, Y., Xiao, Y.H., Lei, C.Y., Mo, J.M., Zhang, W., Lu, X.K., 2011. Nitrogen deposition and forest nitrogen cycling along an urban-rural transect in southern China. Glob. Change Biol. 17, 872–885.

Fang, Y.T., Gundersen, P., Mo, J.M., Zhu, W.X., 2008. Input and output of dissolved organic and inorganic nitrogen in subtropical forests of South China under high air pollution. Biogeosciences 5, 339–352.

Friedlingstein, P., Andrew, R.M., Rogelj, J., Peters, G.P., Canadell, J.G., Knutti, R., Luderer, G., Raupach, M.R., Schaeffer, M., van Vuuren, D.P., Le Quere, C., 2014. Persistent growth of CO₂ emissions and implications for reaching climate targets. Nat. Geosci. 7, 709–715.

Galloway, J.N., Dentener, F.J., Capone, D.G., Boyer, E.W., Howarth, R.W., Seitzinger, S.P., Asner, G.P., Cleveland, C.C., Green, P.A., Holland, E.A., Karl, D.M., Michaels, A.F., Porter, J.H., Townsend, A.R., Vos痛苦, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.

Godbold, D.L., Fritz, H.W., Jentschke, G., Meesenburg, H., Rademacher, P., 2003. Root turnover and root necromass accumulation of Norway spruce (Picea abies) are affected by soil acidity. Tree Physiol. 23, 915–921.

Harrington, R.A., Fownes, J.H., Vitousek, P.M., 2001. Production and resource use efficiencies in N- and P-limited tropical forests: a comparison of responses to long-term fertilization. Ecosystems 4, 646–657.

Haynes, B.E., Gower, S.T., 1995. Below-ground carbon allocation in unfertilized and fertilized red pine plantations in northern Wisconsin. Tree Physiol. 15, 317–325.

Hedges, L.V., Gurevitch, J., Curtis, P.S., 1999. The meta-analysis of response ratios in experimental ecology. New York.

Hogervorst, R.F., Kjonaas, O.J., Persson, H., Timmermann, V., 1998. Vegetation and soil biota response to experimentally-changed nitrogen inputs in coniferous forest ecosystems of the NITREX project. For. Ecol. Manag. 101, 65–79.

Kielland, K., 2009. The changing global carbon cycle: linking plant-soil carbon dynamics to global consequences. J. Ecol. 97, 840–850.

McNulty, S., Currie, W., Rustad, L., Fernandez, I., 1998. Nitrogen saturation in northern forest ecosystems. Bioscience 39, 378–386.

Michaelas, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.

Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.

Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.

Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.

Michaels, A.F., Porter, J.H., Townsend, A.R., Vorosmarty, C.J., 2004. Nitrogen cycles: past, present, and future. Biogeochemistry 70, 151–226.
acidification in response to 30 years of experimental nitrogen loading on boreal forest. Glob. Change Biol. 12, 489—499.

Huang, L., Zhu, W., Ren, H., Chen, H., Wang, J., 2012. Impact of atmospheric nitrogen deposition on soil properties and herb-layer diversity in remnant forests along an urban-rural gradient in Guangzhou, southern China. Plant Ecol. 213, 1187—1202.

Hyvonen, R., Persson, T., Andersson, S., Olsson, B., Agren, G., Linder, S., 2008. Impact of long-term nitrogen addition on carbon stocks in trees and soils in northern Europe. Biogeochemistry 89, 121—137.

Janssens, I.A., Leemans, W., Luyssaert, S., Subke, J.A., Reichstein, M., Ceulemans, R., Ciais, P., Dolman, A.J., Grace, J., Matteucci, G., Popale, D., Piao, S.L., Schulze, E.D., Tang, J., Law, B.E., 2010. Reduction of forest soil respiration in response to nitrogen deposition. Nat. Geosci. 3, 315—322.

Johnk, I.R., Thornley, J.H.M., 1987. A model of shoot — root partitioning with optimal-growth. Ann. Bot. 60, 133—142.

Kaspari, M., Garcia, M.N., Harms, K.E., Santana, M., Wright, S.J., Yavitt, J.B., 2008. Multiple nutrients limit litterfall and decomposition in a tropical forest. Ecol. Lett. 11, 35—43.

Laubhammer, D., Sterba, H., Reinds, G.J., de Vries, W., 2009. The impact of atmospheric deposition and climate on forest growth in European monitoring plots: an individual tree growth model. For. Ecol. Manag. 258, 1751—1761.

LeBauer, D.S., Treseder, K.K., 2008. Nitrogen limitation of net primary productivity in terrestrial ecosystems is globally distributed. Ecology 89, 371—379.

Liu, L.L., Greaver, T.L., 2009. A review of nitrogen enrichment effects on three biogenic GHGs: the CO2 sink may be largely offset by stimulated N2O and CH4 emission. Ecol. Lett. 11, 35.

Liu, L.L., Greaver, T.L., 2010. A global perspective on belowground carbon dynamics effects of atmospheric nitrogen deposition. Nat. Geosci. 3, 315—322.

Liu, X.J., Duan, L., Mo, J.M., Du, E.Z., Shen, J.L., Lu, X.K., Zhang, Y., Yang, H.J., Wan, S.Q., 2010. Nitrogen effects on net ecosystem carbon exchange in a temperate steppe. Glob. Change Biol. 16, 144—155.

Piao, S.L., Fang, J.Y., Ciais, P., Peylin, P., Huang, Y., Sitch, S., Wang, T., 2009. The carbon balance of terrestrial ecosystems in China. Nature 458, 1099—U1082.

Pregitzer, K.S., Burton, A.J., Zak, D.R., Talhelm, A.F., 2008. Simulated chronic nitrogen deposition increases carbon storage in northern temperate forests. Glob. Change Biol. 14, 142—153.

Solberg, S., Dobbertin, M., Reinds, G.J., Lange, H., Andreassen, K., Fernandez, P.G., Hildingsson, A., de Vries, W., 2009. Analyses of the impact of changes in atmospheric deposition and climate on forest growth in European monitoring plots: a stand growth approach. For. Ecol. Manag. 258, 1753—1758.

Thomas, R.Q., Canham, C.D., Weathers, K.C., Gourley, C.D., 2010. Increased tree carbon storage in response to nitrogen deposition in the US. Nat. Geosci. 3, 13—17.

Treseder, K.K., 2008. Nitrogen additions and microbial biomass: a meta-analysis of ecosystem studies. Ecol. Lett. 11, 1111—1120.

Tu, L.-H., Hu, T.-X., Zhang, J., Li, R.-H., Dai, H.-Z., Luo, S.-H., 2010. Short-term simulated nitrogen deposition increases carbon sequestration in a Pinus thunbergii var. mongolica plantation. Plant Soil 1—14.

Vitousek, P.M., Howarth, R.W., 1991. Nitrogen limitation on land and in the sea — how can it occur. Biogeochemistry 13, 87—115.

Wamelink, G.W.W., van Dobben, H.F., Mol-Dijkstra, J.P., van Oijen, M., de Vries, W., 2009. Modelling impacts of changes in carbon dioxide concentration, climate and nitrogen deposition on carbon sequestration by European forests and forest soils. For. Ecol. Manag. 258, 1774—1779.

Wamelink, G.W.W., Wieggers, H.J.J., Reinds, G.J., Kros, J., Mol-Dijkstra, J.P., van Oijen, M., de Vries, W., 2009b. Modelling impacts of changes in carbon dioxide concentration, climate and nitrogen deposition on carbon sequestration by European forests and forest soils. For. Ecol. Manag. 258, 1794—1805.

Xia, J.Y., Wan, S.Q., 2008. Global response patterns of terrestrial plant species to nitrogen addition. New Phytol. 179, 428—439.

Xu, X.K., Han, L., Luo, X.B., Liu, Z.R., Han, S.J., 2009. Effect of nitrogen deposition reduction on biodiversity and carbon sequestration. For. Ecol. Manag. 258, 1774—1779.

Xia, J.Y., Wan, S.Q., 2009b. Modelling impacts of changes in carbon dioxide concentration, climate and nitrogen deposition on carbon sequestration by European forests and forest soils. For. Ecol. Manag. 258, 1794—1805.

Yamaguchi, K., 2002. Global response patterns of terrestrial plant species to nitrogen addition. New Phytol. 155, 428—439.