Mineralogy and Chemistry of Coronadite from Middle Cambrian Manganese Deposits at Wadi Dana, Southern Jordan

1Ahmad Al-Malabeh and 2Tayel El-Hasan

1Department of Earth and Environmental Sciences, Hashemite University, Zarqa 13115, P. O. Box 150459, Jordan
2Faculty of Science, Mu’tah University, 61710, P.O. Box (7), Al-Karak, Jordan

Abstract: Coronadite was reported from the upper-most horizons of the Middle Cambrian sediments at two locations in Wadi Dana, central Wadi Araba and Jordan. The unit is composed of dolomite, limestone and shale. Geochemical investigations show appreciable variations in Mn, Pb, Fe, K and Ba contents in the coronadite between the two studied sites. Pb was found to increase downward in both sites in spite of the lateral distance between them. Fe does not vary vertically, but its concentration decreases eastward. This can be attributed to the nature and mobilization direction, duration of the process, as well as the mechanism of both Fe and Pb bearing solutions. Mineralogical analysis has revealed the presence of hollandite-coronadite, cryptomelane-coronadite and psilomelane-coronadite in solid solutions. These petrographical and geochemical characteristics of the coronadite-bearing samples indicated that they formed epigenetically.

Key words: Middle Cambrian, coronadite, epigenetic genesis, lateralization, solid solution

INTRODUCTION

Coronadite is a Pb-Mn hydrous oxide that belongs to the tetravalent Mn-oxide-rich mineral that belongs to cryptomelane group, which is usually associated with supergene manganese ores[1-3]. The tetravalent Mn-oxides indicating oxidizing conditions during formation[4]. However, coronadite is reported also to form during lateralization[5]. Ehrlich et al.,[6] found manganese nodules that were composed mainly of coronadite formed by epigenetic solution.

Coronadite is fairly rare mineral of the oxidation zone, is in most cases the product of simultaneous weathering of Mn and Pb-minerals; Pb-replacement deposits in Mn-rich limestones[2]. Originally discovered and described by[7] on ore vein in the state of Arizona, USA. Later was discovered associated with cryptomelane, hollandite, limonite, quartz and kaolinite in Morocco[8].

Manganese deposits crop out in central Wadi Araba, mainly at Wadi Dana and Wadi Um Ghadah (Bir Madkour)[9]. The manganese in Wadi Um Ghadah is hosted within the bedded arkosic sandstone Lower Cambrian Saleb Formation, while the manganese of Wadi Dana occurs mainly within the Middle Cambrian Burj Formation, Numeric member composed of dolomite, limestone and shale (Fig. 1). Both deposits occur in three zones arranged from bottom to top: the primary sedimentary ore; the supergene stage, and finally the epigenetic stage[11]. Each stage has its own chemical and mineralogical characteristics.

Problem statements: This study is concerned with the Wadi Dana occurrence and mainly with the upper epigenic horizons stage. Coronadite is confined to this particular horizon, which shows higher Pb concentrations. These high Pb concentrations of the Burj Formation at Wadi Khaled and Wadi Mahjoob was the main reason for further investigations. El-Hasan et al.,[11] reported higher Ba, Pb and Zn values within two manganese ore samples from Wadi Khaled. Moreover, coronadite was mentioned in the assemblage of the Timna manganese deposit, Israel[13,14].

The mineralogical analysis results confirmed the presence of coronadite within this horizon. This mineral phase was identified for the first time as a major mineral phase within the Jordanian manganese deposits. In the study area, the mineral was identified at two sites: in Wadi Khaled in the west and in Wadi Mahjoob in the east. It is confined to the upper part of these profiles represents the Middle Cambrian Burj Formation (Figs. 1 and 2). The presence of coronadite is important to understand the genesis of the mineral paragenesis within the Cambrian manganese deposits as
Geologic setting: The coronadite-hosting Burj Formation is underlain by the lower Cambrian Saleb Formation, (Fig. 2). The Burj Formation is a succession of thin layers of sandstone, dolomitic sandstone - sandy dolomite and red shale layers. The thickness is changing along the course of Wadi Dana ranging from 25-50 m, due to the irregular Precambrian paleo-terrain. The lithology of Saleb Formation starts with a dolomitic layer at the bottom and then changes into arkosic sandstone faces with clay alternations. This profile was very rich in manganese that occurs as thin bands, concretions, irregular bodies and as cross cutting veins at the top of the formation.

The varying thickness is accompanied by lateral lithological changes. This is visible specifically within the coronadite-hosting horizon, which consists of thin alternations of fine-grained sandstone with black claystone and carbonaceous layers in the west (Wadi Khaled). This horizon is only 4 m thick, divided into three parts. These are from bottom to top: (i) 1 m thick fine-grained sandstone with very thin alternations of clay, followed by (ii) 2 m thick very thin alternations of varve-like carbonaceous sandstone, with black claystone and siltstone. Mineral determination confirms the presence of a cryptocrystalline fluor-apatite as one of the main mineral constituents of this layer. Finally, the sequence is overlain by (iii) 1 m of massive, hard sandstone cemented by manganese minerals and secondary copper mineral (Chrysocolla) encrustations.

Wadi Mahjoob is located 4 km east of Wadi Jamal (Fig. 1). The Saleb Formation starts with 4 m basal conglomerate, followed by 25 m thick massive, vary-colored, cross-beded, arkosic sandstone. Some Mn mineralizations were found in this formation disseminated in the sandstone layers, in addition to thin, mineralized beds. Above the Saleb Formation the 40 m thick Burj Formation starts with dolomite interbedded by thin clay layers, followed by a succession of fine- to medium-grained sandstone, rose-brown colour and massive dolomite-dolomitic sandstone, with thin claystone intercalations. The upper part is a sandy facies and comprises the main ore horizon. It is intercalated with red claystone layers overlain by the Upper Cambrian sandstone Umm Ishrin Formation. The manganese in this profile occurs as bands, lenses, irregular bodies and concretions. In the upper portion Mn is filling the joints and forms stock-work veins.

MATERIALS AND METHODS

The studied samples have been collected from the horizons with the higher concentrations of Pb among the all studied ore samples from Wadi Dana profiles.
They consist of four samples two from Wadi Khaled (Kh2 and Kh4) and two from Wadi Majoob (Mh13 and Mh14). The samples were analyzed for their bulk composition using XRF machine (RIGAKU model 3270). The operation conditions were: tube voltage 50 kV and tube current 50 mA. The mineral identification was carried out using XRD micro-diffract meter machine with CuKα radiation. The samples were studied under the microscope and undergone SEM analysis using the machine (SEM-EDS - JOEL - JSM 5400).

The quantitative mineral chemistry was carried out using the EPMA (JOEL-JXA 8621) EPMA Superprobe. The microprobe analytical conditions were 1×10⁻⁸ mA, 25-50 nA specimen current potential, 10 second integration time and 20 kV acceleration potential, the SPI mineral standards were used for the calibration process.

RESULTS AND DISCUSSION

Petrography: The coronadite found within the profiles of Wadi Khaled occur as secondary veinlets associated within the secondary chrysocolla and apatite, as well as, thin laminations alternating with apatite and siltstone layers (Fig. 3). The second profile of Wadi Mahjoob is located about 14 km to the east of Wadi Khaled. The coronadite-bearing horizon is the upper layers, where the coronadite is confined to these veinlets (Fig. 4).

Frodel, C. and E. Heinrich[15] has described the coronadite in polish sections as rhythmic structures alternating with cryptomelane, psilomelane. Moreover, it belongs to tetragonal system and its XRD prime at 3.10Å (100). Moreover, the qualitative SEM photographs and chemical charts results for samples from both studied profiles reveal obviously the presence of coronadite as the peaks of Mn and Pb beside small peaks of Ba are evident as secondary veinlets (Fig. 5) and as curved alternating lamellae as in (Fig. 6). Furthermore, the lithology of the studied profiles is similar to Hakalil Formation, Nimra Formation, Nehustan Formation and Mikhrot Formation (from the bottom to the top) of Timna manganese outcrop at southern Israel[12,15,18].

Geochemistry of the host rocks: The EPMA analysis has shown high Pb concentrations within the upper mineralized horizons (up to 8%) (Table 1). The SEM, XRD and EPMA results confirmed the presence of coronadite. Previously high Pb and Ba concentrations were reported by[12]. He gave the following concentrations (5.4, 0.6, 0.7 and 0.8%) for MnO, Cu, Ba and Pb, respectively for one sample from Wadi Khaled.

In spite of this high Pb content, coronadite was not detected at that time, which might be due to the lack of the instruments.
Fig. 6: SEM photo (Magn. X 150) of coronadite phase in the veins of Fig. 4 showing the curved lamellae texture; down is SEM chart showing the quantitative composition of the above texture at the point (x).

Table 1: Average of 15 EPMA analytical results for each studied sample. All in wt%.

|       | Kh2     | Kh4     | Mh13    | Mh14    |
|-------|---------|---------|---------|---------|
| Mn    | 33.28±4.68 | 31.63±7.29 | 39.38±4.13 | 46.31±6.30 |
| Fe    | 5.16±2.59  | 0.81±0.56  | 0.32±0.2  | 0.17±0.07  |
| K     | 0.64±0.55  | 0.75±0.26  | 0.21±0.05 | 2.24±0.97  |
| Ba    | 3.98±1.14  | 4.98±0.65  | 4.79±1.14 | 4.48±1.56  |
| Pb    | 9.82±3.77  | 3.87±0.83  | 10.8±2.22 | 3.26±1.08  |

The XRF analysis of whole rock test was carried out on samples from all Wadi Dana mineralized areas. The results show that among 43 ore samples there are 4 samples that contain high Pb concentration. These 4 samples belong to Wadi Khaled and Wadi Mahjoob profiles that were mentioned in the analytical section. Bar-Matthews, M. [13] have used the Pb plus Fe versus Mn discriminatory plot in order to discriminate between the various types of Timna Mn ore. However, when the discriminatory diagram is applied on the studied samples. It can be seen that the samples have nearly the same ore types (Fig. 7). Nevertheless, the studied samples were plotted within the epigenetic field or B-type after [13] because of their high Pb contents. Moreover, the same samples were found to bear relatively higher Zn concentrations, therefore, they were found to scatter within the field of dubhite [11].

Fig. 7: Scatter diagram of studied samples showing clear variation in their Pb and Mn concentrations.

Table 2: A comparison between the Pb-enriched manganese samples from Wadi Dana and hollandite-coronadite solid solution from Timna Mn deposits/ south Israel. All in wt% only Zn in ppm.

|       | 1     | 2     | 3     | 4     | 5     |
|-------|-------|-------|-------|-------|-------|
| Mn    | 18.74±4.02 | 22.67±39.8 | 32.13±46.03 | 45.75±50.99 | 30-52 |
| Ba    | 1.89±3.68  | 3.9±5.83  | 2.16±6.16  | 2.76±6.91  | 2-10  |
| Pb    | 5.86±13.5  | 2.19±5.48  | 6.3±16.94  | 1.6±4.35  | 2-8   |
| Cu    | 1.27±5.77  | 2.91±3.83  | 1.55±1.92  | 1-3      |
| Fe    | 3.48±8.38  | 0.44±4.94  | 0.25±2.0   | 0.14±0.56 | 0.2-2.5|
| K     | 0.24±8.38  | 0.46±2.16  | 0.22±0.6   | 1.13±3.74 | 0.07-1 |
| Zn    | 723-2330   | 560-2030   | n.d.       | n.d.     | 2000-5000|

Nicholson, K. [5] was the first to introduce the term dubhite that described the supergene terrestrial oxides formed by weathering of Mn rich sequences. This result is in consistency with the results given by [19], in which he reported a cryptomelane-hollandite-coronadite series in Vani manganese deposits, Milos, in Greece to contain high concentrations K, Ba, Pb and Zn. Furthermore, this comparison shows the different behavior and mode of formation between Fe and Pb. On one hand, Fe is enriched in the samples of Wadi Khaled and decreases toward the east the samples of Wadi Mahjoob. On the other hand, Pb shows a vertical or stratiform mode of occurrence. The behavior of Fe and Pb and their lateral spreading might be attributed to parameters of the environment such as Eh and pH [20,14].

Mineral chemistry and genesis: Coronadite was described as Pb-Mn oxide of variable composition [21]. However, the EPMA obtained data in Table 2 shows that the studied samples resemble those of hollandite-coronadite solid solution mentioned by [13]. Moreover, when comparing these analytical results with coronadite of [22] (Table 3), it is obvious that studied samples cannot
Table 3: Selected EPMA data from the studied samples compared with coronadite analysis after [2]

|            | 1    | 2    | 3    | 4    | 5    |
|------------|------|------|------|------|------|
| Coronadite | n.d  | 0.21 | 2.47 | 0.03 | 0.33 |
| Mn         | 43.94| 43.88| 56.18| 50.59| 54.17|
| Ba         | 26.37| 24.55| 0.21 | 0.90 | 13.56|
| Pb         | 0.04 | 0.11 | 0.71 | 0.67 | 1.64 |
| Cu         | 0.39 | 0.21 | 0.05 | 0.04 | 0.03 |
| K          | n.d  | n.d  | 4.28 | 0.11 | 2.83 |
| Zn         | 0.08 | n.d  | 6760 | 1870 | n.d  |

n.d: not analyzed

CONCLUSION

Mineral chemistry was shown an appreciable variation in Mn, Pb, Fe, K and Ba contents in the coronadite of the studied samples. The petrography and geochemistry of the coronadite bearing samples reveals the epigenetic mode of formation. The Pb was found to increase downward in the two sites in spite of the wide lateral distance between them. Meanwhile, Fe shows almost horizontal trend, where higher Fe at the west and it decreases toward the east. This might be attributed to the nature, direction, time, as well as the mechanism of both Fe and Pb bearing solutions. The mineralogical and EPMA analysis revealed the presence of hollandite, cryptomelane and psilomelane-coronadite solid solution phases. The mineral chemistry of coronadite reveals a variation in the coronadite chemistry that is better indicated from the recalculated formula: 

\[
(Ba_{0.03-1.06} K_{0.58-1.52} Pb_{0.13-0.4} Cu_{0.01-0.26}) \text{Fe}_{(0.02-0.28)} \text{Mn}_{(5.47-7.67)} O_{16-x}H_2O
\]

The upper position of coronadite-containing horizon and the nature of ore types, beside the microscopic and mineral chemistry of the solid solution phases all together would indicate the epigenetic nature of the forming solution. These four samples were found to belong to dubhites ore type. Such ore type is usually formed by lateralization process, which might accompany with descending meteoric water. Mineral of cryptomelane group are varied widely over short distance due to the strong influence of local variation in water chemistry [22]. This lateralization process affected the upper manganese supergene ores causing the formation of these Pb-rich solutions. Many previous studies [5,6,21], have indicated the epigenetic nature of coronadite. However, in our case it’s obvious that it is not pure coronadite phase but cryptomelane, psilomelane and hollandite-coronadite solid solution phases. Such solid solution is an additional evidence of epigenetic mode of formation.

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